

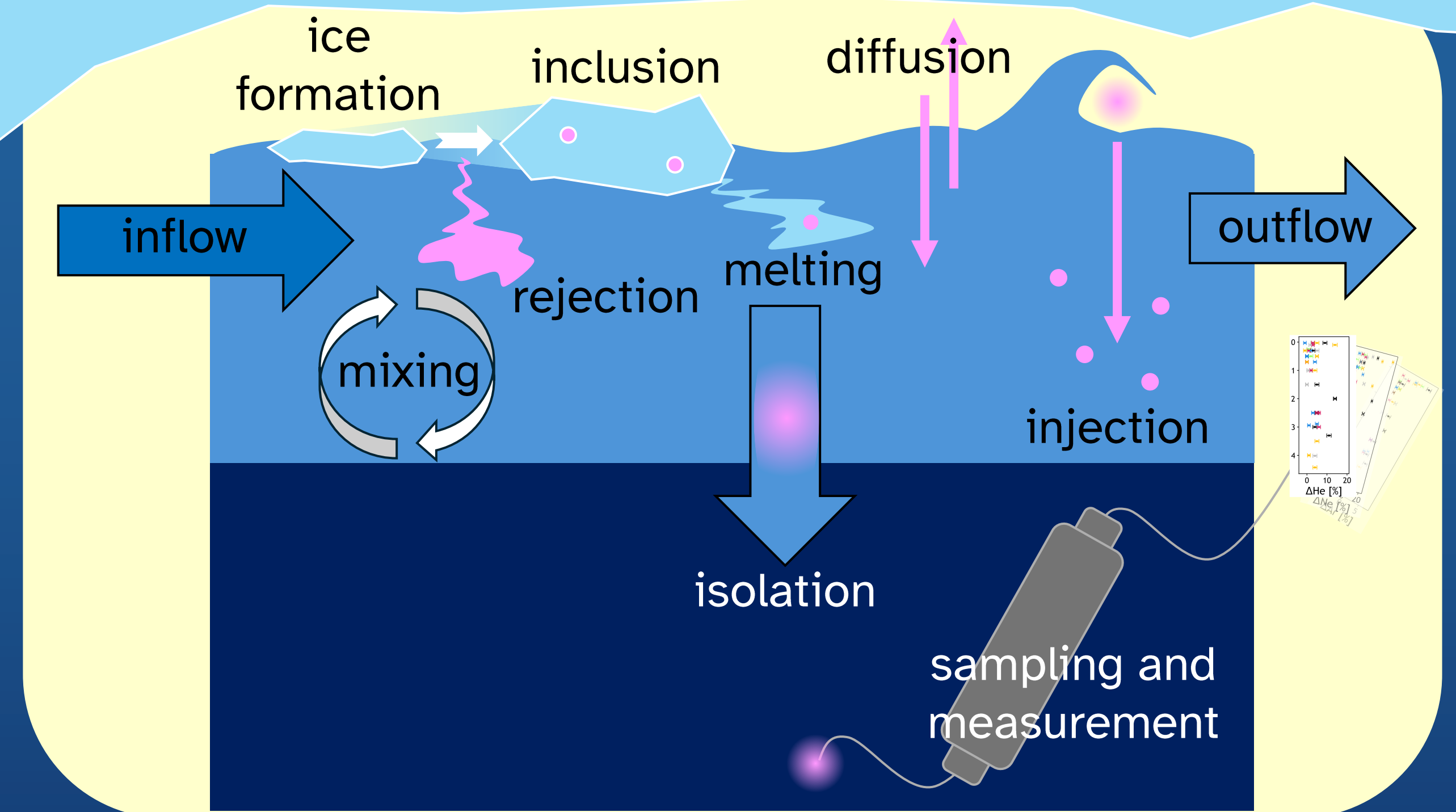
Background

- Project: Ventilation and Anthropogenic Carbon in the Arctic Ocean (VACAO), Synoptic Arctic Survey 2021 (SAS21) [1]
- Goals: understand timescales of ventilation and carbon uptake in the Arctic Ocean
- Physical processes at the surface control rates of gas uptake
- Noble gas concentration anomalies act as non-chemical tracers for physical gas exchange processes. [2]
- Formation/melting of sea ice, injection of air and rapid cooling can be modelled and constrained using noble gas data.
- Results are used to quantify the primary surface exchange processes and to correct age-tracer data which assumes diffusion as the only form of gas exchange

The noble gases as exchange process tracers

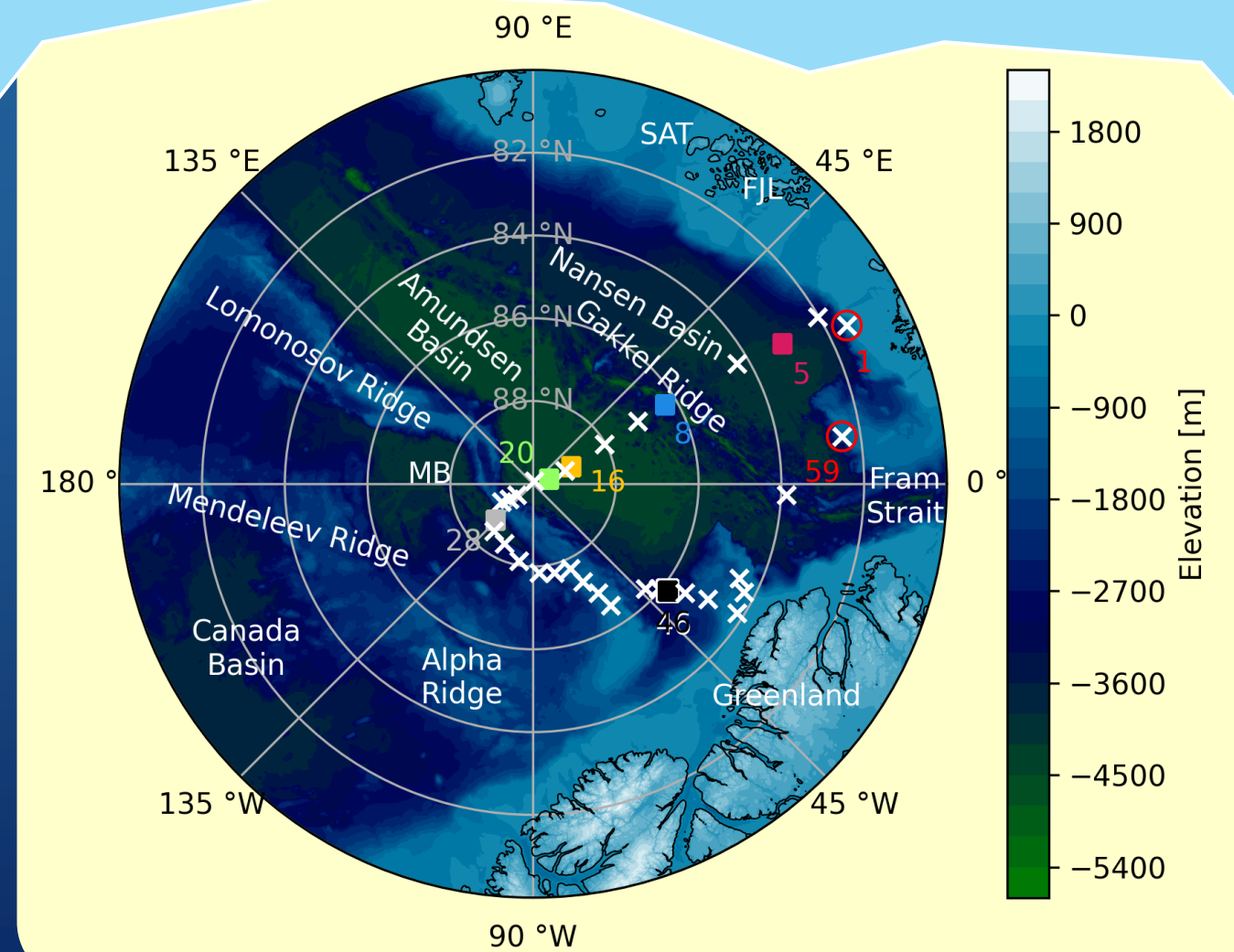
- Five stable noble gases: He, Ne, Ar, Kr, Xe
- (Bio-)Chemically inert and have constant atmospheric histories
- Practically no degradation in subducted waters: effectively conservative
- Concentrations only dependent on *physical* exchange parameters
- Atmospheric concentrations, solubilities, solubility-temperature gradients and behaviour in ice varies greatly from species to species of noble gas
- These properties make them excellent environmental tracers for surface gas exchange parameters
- Concentration/Anomaly profiles can be used with an appropriate model and fitting algorithm to parameterise surface gas exchange [2]

Physical gas exchange processes



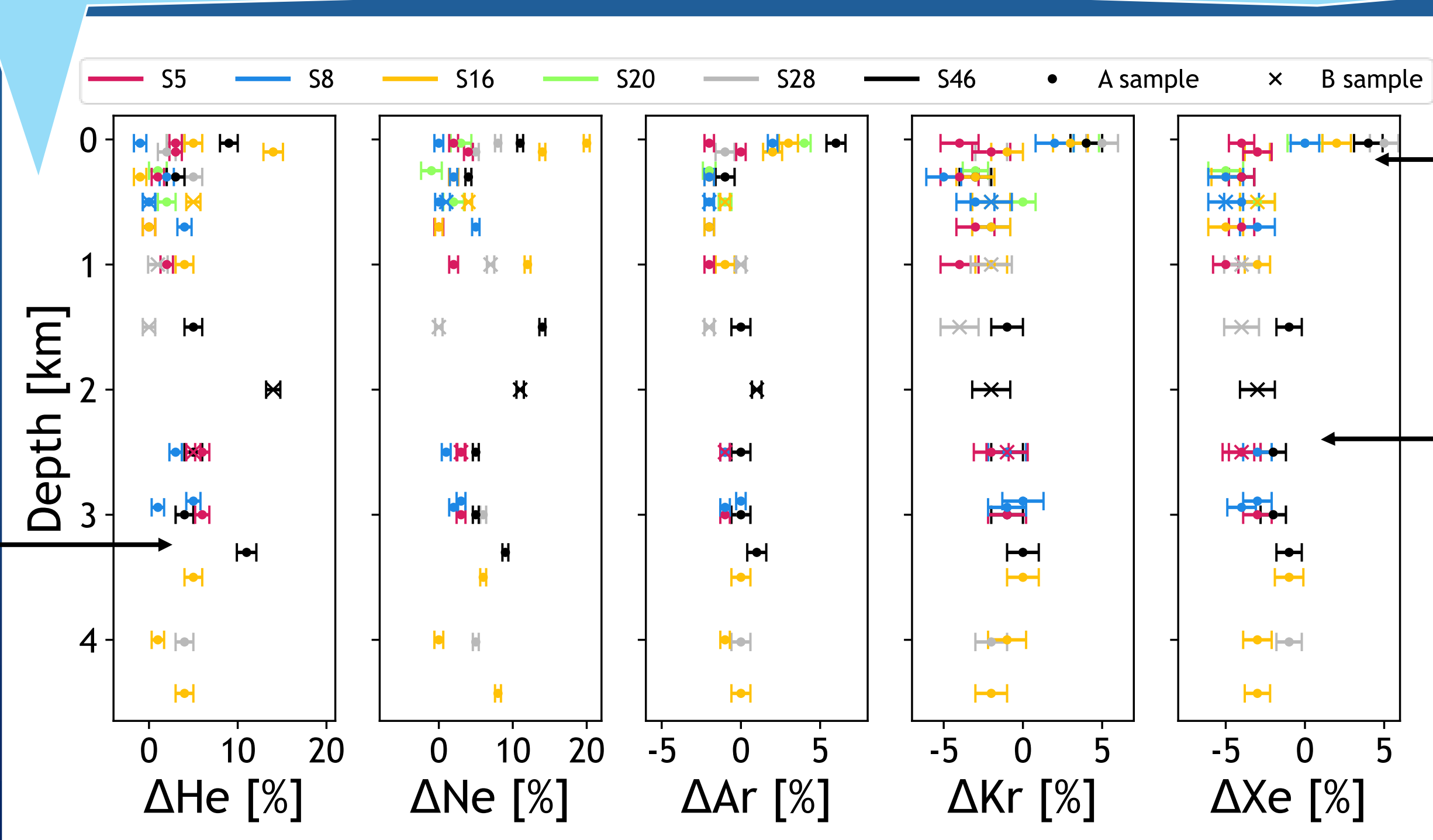
- Both models we present here consider only the interaction between the atmosphere and elements in the Polar Mixed Layer (PML)
- PML is assumed well-mixed and coupled to the atmosphere
 - Diffusive exchange: acts to push gas concentrations toward equilibrium
 - Excess air injection: wave breaking forces bubbles into the water which release gases – diffusivities and abundances of gases control how much their concentrations are affected
- Ice formation: all noble gases (except He) are squeezed out of the ice matrix upon freezing, causing fractionation [3]
- Ice melting: gases trapped in ice are released into the PML
- Water below the PML is considered isolated and noble gases conserved

SAS21 Cruise track



Vertical gas concentration profiles taken at each station shown [1, 4]

NOBLE GAS LAB ANALYSIS



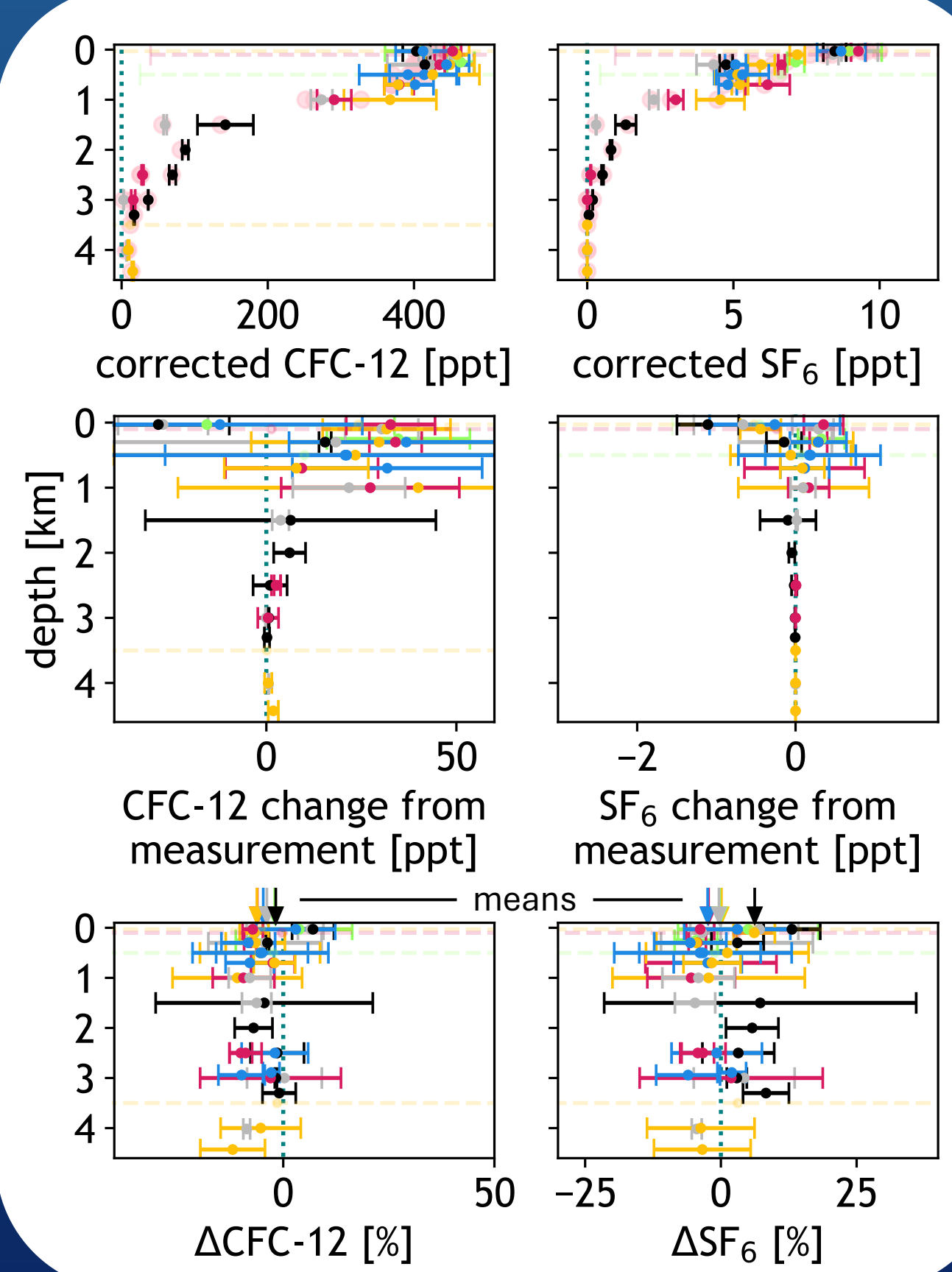
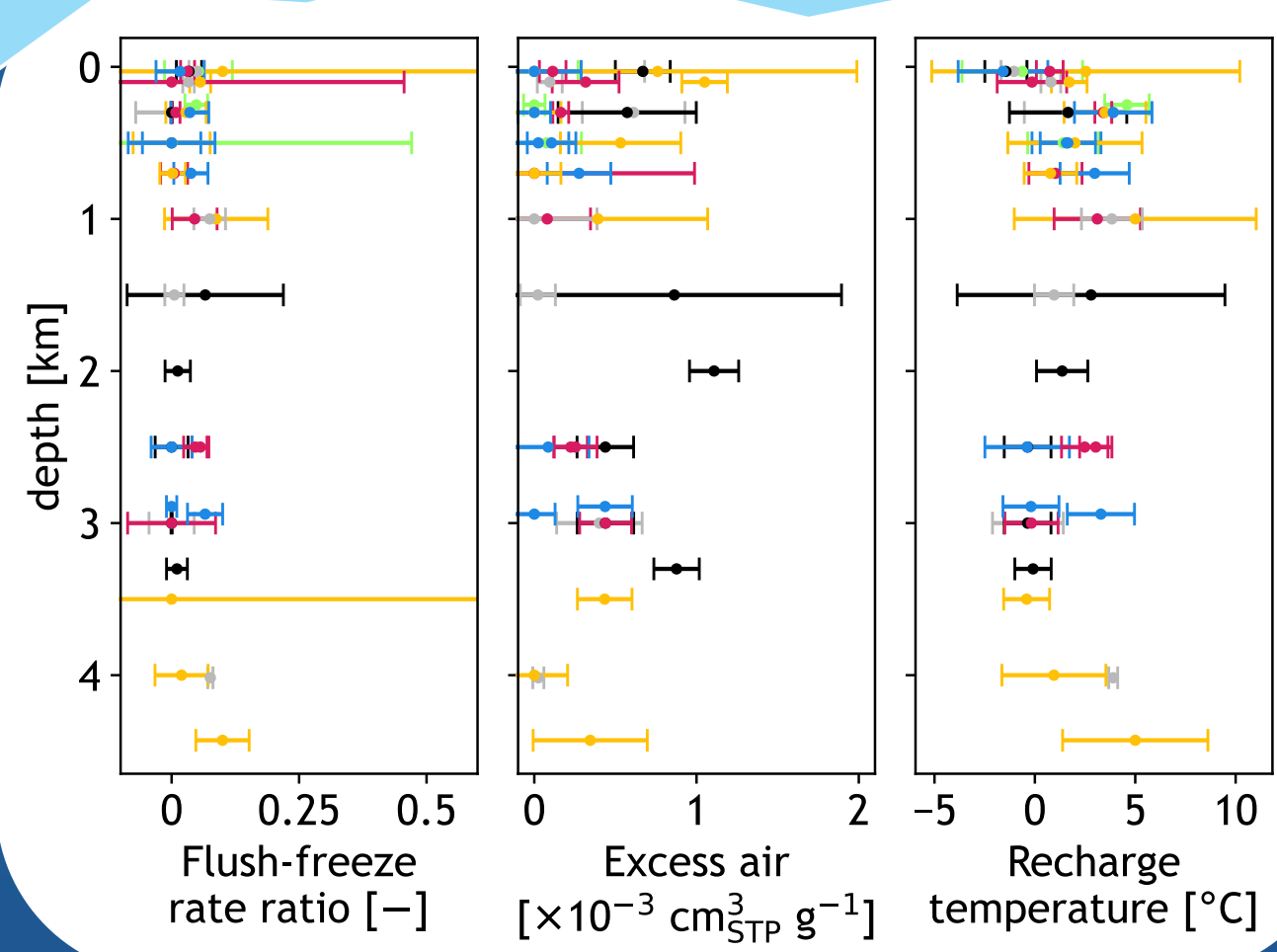
Surface oversaturation of all gases: ice formation?

$$\Delta \text{Gas} = \left(\frac{C_{\text{Gas}}^{\text{meas}}}{C_{\text{Gas}}^{\text{eq}}} - 1 \right) \times 100\%$$

$$\Delta \text{Gas} = 0\% \Rightarrow \text{equilibrium/saturation}$$

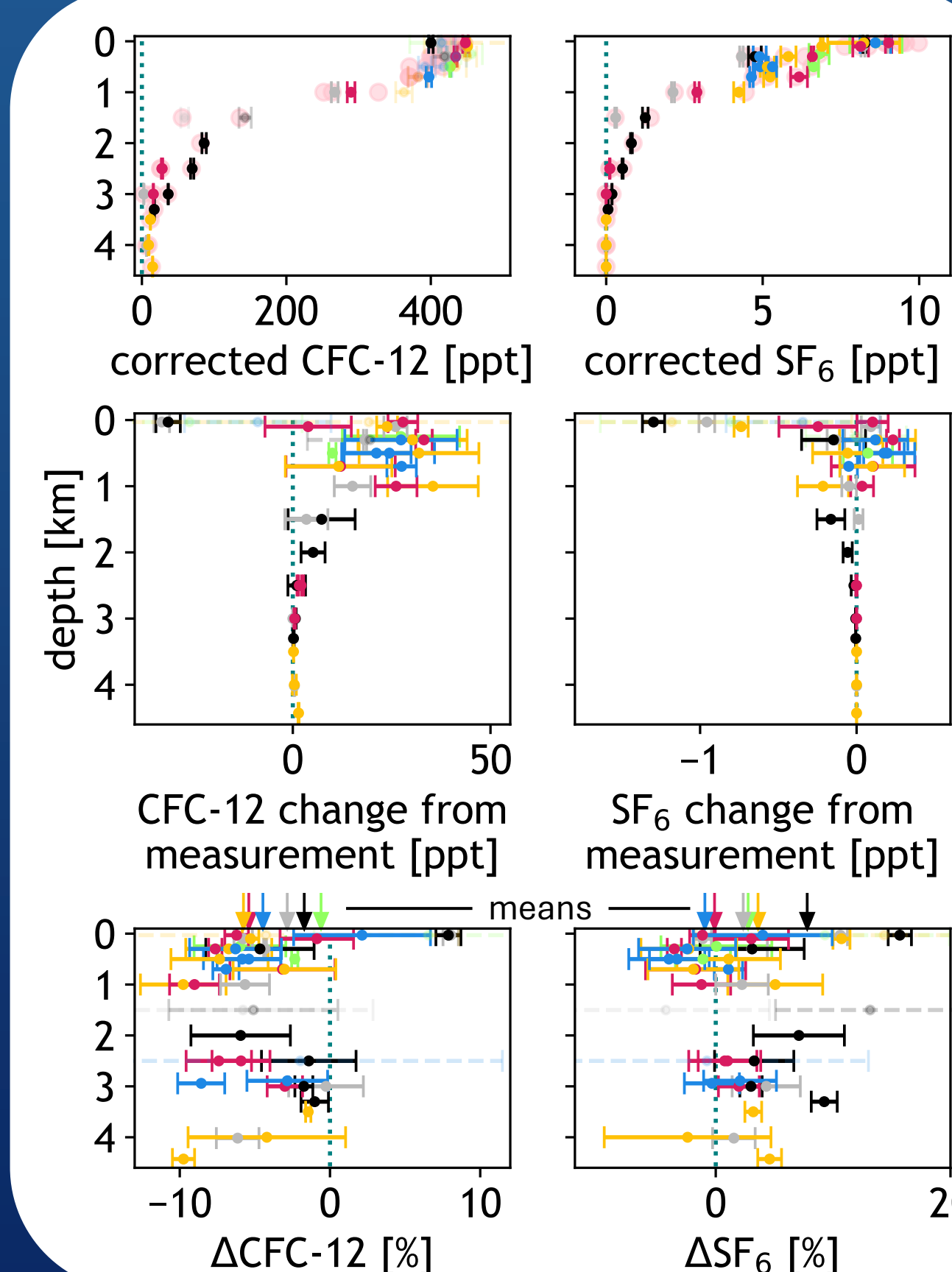
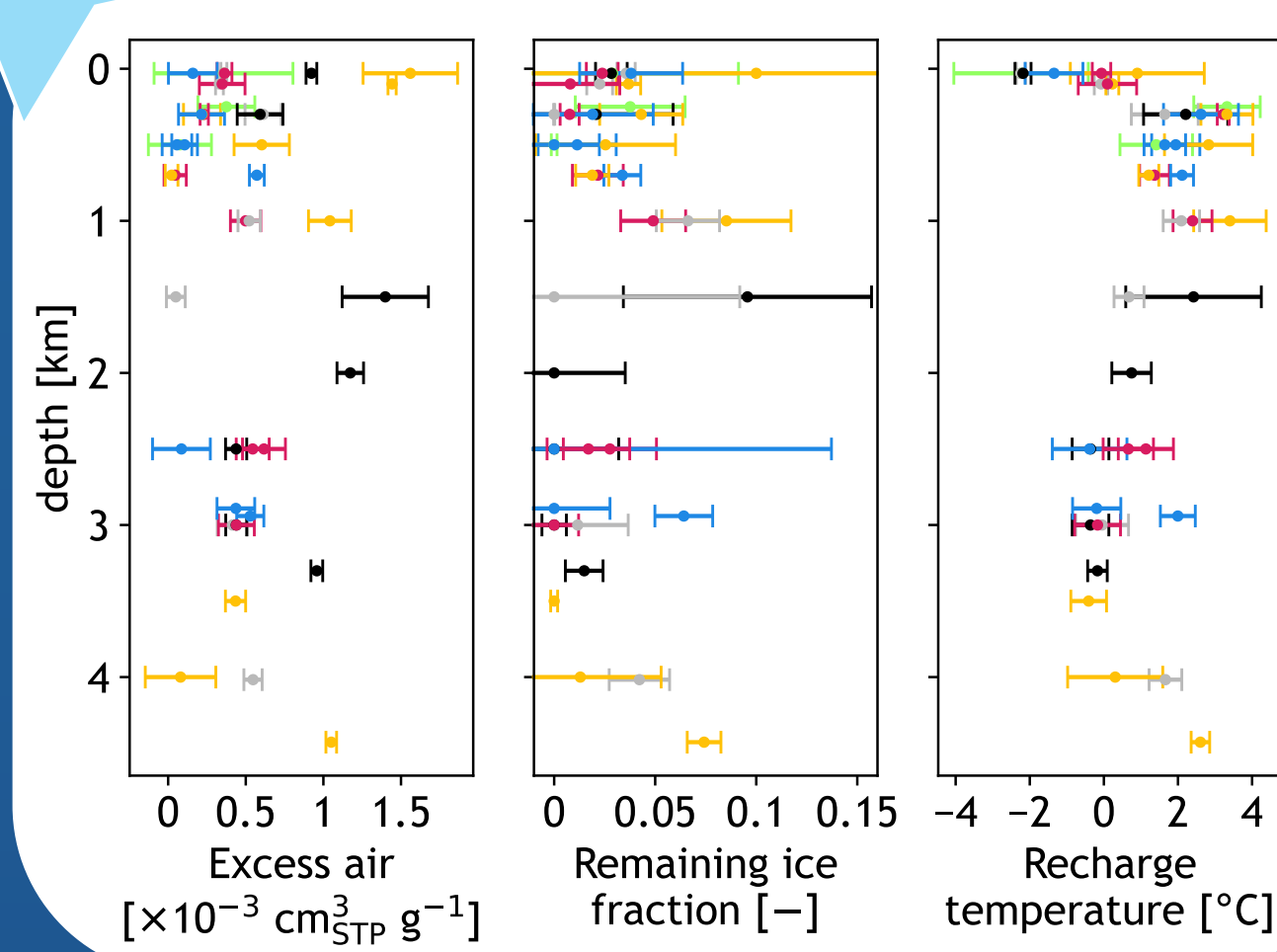
Undersaturated heavy gases: rapid cooling?

Mixed reactor



- Useful water age-tracers like CFC-12 and SF₆ are also affected by physical gas exchange
- Modelled anomalies can be used to “correct” their measured concentrations: $C^{\text{corr}} = C^{\text{meas}} \div (1 + \Delta^{\text{model}})$ [6]
- CFC12 undersaturated in both models, SF6 differs between the two

AIFM



Model details and fitting

- Model parameters are constrained in a least-squares sense [2] using Levenberg-Marquardt nonlinear minimisation of:

$$\chi^2 = \sum_{\text{Gas} \in \{\text{He}, \dots, \text{Xe}\}} \left(\frac{C_{\text{Gas}}^{\text{meas}} - C_{\text{Gas}}^{\text{model}}(\text{fitted parameters})}{\sigma_{C_{\text{Gas}}^{\text{meas}}}} \right)^2$$

- Mixed reactor model:**

$$C_{\text{Gas}}^{\text{meas}} = (1 + (2 - \kappa_{\text{Gas}}) \cdot R_{\text{ff}}) \cdot C_{\text{Gas}}^{\text{eq}} + A \cdot z_{\text{Gas}}$$

- κ = ice fractionation coefficient, z = atmospheric abundance, A = excess air, R_{ff} = flush-freeze rate ratio
- Steady-state mixed reactor where fluxes due to ice formation balance those due to flushing rate of PML
- Excess air added *after* as transient term, melting ignored

- Air injection, freezing and melt model (AIFM):**

$$C_{\text{Gas}}^{\text{meas}} = (C_{\text{Gas}}^{\text{eq}} + A \cdot z_{\text{Gas}}) \cdot (1 + (1 - \kappa_{\text{Gas}}^2) \cdot f_{\text{ri}})$$

- f_{ri} = remaining proportion of ice as a fraction of PML after melting
- Excess air injected *before* freezing occurs
- Freezing modelled as a single Rayleigh-fractionation process [5], melting as a single rapid event

Questions?
Ask me here
or email me!



Conclusions and outlook

- Decision still must be made as to if a model should over- or under-saturate SF₆
- Excess air should be divided into diffusive and non-diffusive regimes
- Glacial meltwater contribution could be modelled with a water-fraction model [7]
- Corrected profiles should be smoothed and used with transit time distributions of SF₆/CFC-12 tracer pairs to better constrain timescales of Arctic water circulation
- Application to the novel tracer ³⁹Ar is also a possibility

References and acknowledgements

- [1] Snoeijs-Leijonhalm, P., Expedition Report SWEDARCTIC Synoptic Arctic Survey 2021 with icebreaker Oden, Luleå: Swedish Polar Research Secretariat, 2022.
- [2] Jung, M. and Aeschbach, W., “A new software tool for the analysis of noble gas data sets from (ground)water,” *Environmental Modelling & Software*, vol. 103, pp. 120–130, 2018.
- [3] Top, Z., Martin, S. and Becker, P., “A laboratory study of dissolved noble gas anomaly due to ice formation,” *Geophysical Research Letters*, vol. 15, no. 8, pp. 796–799, 1988.
- [4] GEBCO Bathymetric Compilation Group, “The GEBCO 2022 Grid - a continuous terrain model of the global oceans and land,” 2022.
- [5] Kluge, T., Marx, T., Aeschbach, W., Spötl, C. and Richter, D. K., “Noble gas concentrations in fluid inclusions as tracer for the origin of coarse-crystalline cryogenic cave carbonates,” *Chem. Geol.*, vol. 368, pp. 54–62, 2014.
- [6] Hamme, R. C., Emerson, S. R., Severinghaus, J. P., Long, M. C. and Yashayaev, I., “Using noble gas measurements to derive air-sea process information and predict physical gas saturations,” *Geophys. Res. Lett.*, vol. 44, no. 19, pp. 9981–9989, 2017.
- [7] Loose, B., Jenkins, W. J., Moriarty, R., Brown, P., Jullian, L., Garabato, A. C. N., Valdes, S. T., Hoppema, M., Ballentine, C. J. and Meredith, M. P., “Estimating the recharge properties of the deep ocean using noble gases and helium isotopes,” *Journal of Geophysical Research: Oceans*, vol. 121, pp. 5959–5979, 2016.