

1                   **Triple-Dip La Niña in 2020-22: Impact of Annual Cycle**  
2   **in Tropical Pacific SST**

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14  
15   **Key Points:**

- 16   • The triple-dip La Niña event in 2020-2022 is associated with the persistent  
17   easterly and southeasterly wind anomalies over the tropical Pacific.
- 18   • The persistent negative phase of the symmetric mode of annual cycle SST in the  
19   tropical Pacific maintains the easterly wind anomalies.
- 20   • The contrast of SST between the northern and southern hemispheres provides a  
21   favorable background for the southeasterly wind anomalies.

22       **Abstract**

23       The triple-dip La Niña in 2020-22 is characterized by persisting easterly and  
24       southeasterly wind anomalies over the tropical central and southeastern Pacific. Our  
25       results show that the wind anomalies are associated with two leading modes of the  
26       annual cycle (antisymmetric and symmetric modes) of sea surface temperature (SST).  
27       The first two modes account for 82.2% and 13.5% of total variances, linking to the  
28       seasonal swing of SST distribution in the northern and southern hemispheres and the  
29       temporal evolution of El Niño-Southern Oscillation (ENSO), respectively. The  
30       persistent negative phase of the symmetric mode enhances easterly wind, while the  
31       antisymmetric mode strengthens the southeasterly wind. The negative phase of the  
32       antisymmetric mode is regulated by the contrast of SST anomalies between the  
33       northern and southern hemispheres. Therefore, both the zonal and the meridional  
34       process associated with annual cycle anomalies may play an important role in the  
35       evolution of the triple-dip La Niña in 2020-22.

36       **Plain Language Summary**

37       El Niño-Southern Oscillation (ENSO) can be depicted by the interannual  
38       variation of the annual cycle sea surface temperature (SST) anomaly, which alternates  
39       between El Niño and La Niña. The La Niña phase usually persists for a longer period  
40       than the El Niño. The triple-dip La Niña in 2020-22 would be the first three-year La  
41       Niña since the 1998-2001 event. Previous studies suggested that the southeasterly  
42       wind over the central Pacific in spring would trigger the third-year La Niña in 2022.

43 In this study, we further point out that the persistent easterly and southeasterly wind  
44 anomalies over the tropical central and southeastern Pacific associated with the  
45 negative phase of annual cycle modes of SST are conducive to this event. Our results  
46 highlight the importance of the inconsistent of SST warming between the north and  
47 south Pacific to prolonged La Niña events.

48 **Keywords:** Triple-dip La Niña, Annual cycle anomaly, Sea surface temperature

## 49 **1. Introduction**

50 El Niño-Southern Oscillation (ENSO) is the most prominent interannual signal  
51 of the tropical Pacific and has significant socioeconomic impacts (Timmermann et al.,  
52 2018). ENSO is characterized by sea surface temperature (SST) anomalies over the  
53 central and eastern equatorial Pacific, which develops during boreal spring and  
54 summer, reaches maturity during autumn and winter, and decays in the following  
55 spring. Such seasonal evolution is referred to as the ENSO and annual cycle  
56 phase-lock phenomena (Chen & Jin, 2020; Galanti E 2000; Jin et al., 1994;  
57 Rasmusson & Carpenter, 1982). The characteristics of the annual cycle in the central  
58 and eastern equatorial Pacific have been widely discussed (Horel, 1982; T. M. Li &  
59 Philander, 1996; Song et al., 2020; Tozuka & Yamagata, 2003; B. Wang, 1994; X. L.  
60 Wang, 1994; Shang-Ping Xie, 1994). The annual cycle mode of tropical Pacific SST  
61 can be decomposed into symmetric and antisymmetric modes (X. L. Wang, 1994).  
62 The symmetric mode directly affects ENSO by changing the zonal SST gradient,  
63 while the antisymmetric mode reflects the meridional gradient of SST, which affects

64 ENSO by modifying the cold tongue (B. Wang, 1994). Tozuka and Yamagata (2003)  
65 also investigated the annual ENSO, where the ENSO is interpreted as the interaction  
66 between two distinct modes of air-sea interaction: the annual ENSO mode and the  
67 interannual ENSO mode. Therefore, the influence of the annual cycle should be  
68 considered in the variability and predictability of ENSO (Shin et al., 2021).

69 ENSO has become more complex during the last few decades, including the  
70 intensity difference, asymmetric evolution, and flavor diversity (Z.-Z. Hu et al., 2020;  
71 Timmermann et al., 2018; Yeh et al., 2009). For instance, there have been fewer El  
72 Niño events and frequent La Niña events since the 1990s (R.-H. Zhang et al., 2022),  
73 and some La Niña events are usually followed by another La Niña (Z.-Z. Hu et al.,  
74 2014; Luo et al., 2017) or even persisted three years, such as 1973-76 and 1998-2001.  
75 Recently, persistent cold SST anomalies (SSTAs) have been occurring in the central  
76 and eastern equatorial Pacific since 2020, and it is another triple-dip La Niña event  
77 after 1998-2001 (Cates et al., 2022; Fang et al., 2023; Zheng et al., 2022). It has been  
78 argued that multi-year La Niña might be caused by westward propagation of reflected  
79 cold ocean temperature anomaly in the off-equatorial Pacific, the asymmetric  
80 response of the atmosphere due to climatological seasonal cycle, and inter-basin  
81 interaction with the Atlantic and Indian Oceans (DiNezio et al., 2017; Gao & Zhang,  
82 2016; Z.-Z. Hu et al., 2014; Luo et al., 2017; C. Zhang et al., 2019; R.-H. Zhang et al.,  
83 2022). Both X. Li et al. (2022) and Fang et al. (2023) emphasized the role of  
84 persistent easterly wind anomalies in the triple-dip La Niña in 2020-22, while Hasan

85 et al. (2022) indicated the contribution of inter-basin interaction. But it is unclear for  
86 the role of the annual cycle anomalies in the La Niña in 2020-22, in the context of  
87 heterogenic warming trends in the tropical Pacific (Gao et al., 2022).

88 The evolution of ENSO is related to the zonal SST gradient across the tropical  
89 Pacific, however, due to the differences in land and sea distribution and heat content,  
90 the warming trend in the northern hemisphere is stronger than in the southern  
91 hemisphere (Cavaleri et al., 1997; Flato & Boer, 2001), and increasing meridional  
92 temperature gradient may strengthen cross-equatorial winds. The cross-equatorial  
93 winds can affect ENSO variation (Shang-Ping Xie et al., 2018), leading to a “La  
94 Niña-like” change in the tropical Pacific (S. Hu & Fedorov, 2018). Recently, Fang et  
95 al. (2023) argued that the meridional processes in the eastern Pacific are important to  
96 trigger a third-year La Niña event in 2022-23. The meridional processes seem  
97 associated with the antisymmetric mode of annual cycle SST which could be an  
98 important factor in the occurrence of the triple-dip La Niña in 2020-22. In the present  
99 study, we intend to examine the interannual variability of annual cycle SST in the  
100 tropical Pacific associated with ENSO, and try to discuss its role in the triple-dip La  
101 Niña event in 2020-22, as well as the contribution of the meridional SST gradient.

## 102 **2. Data and Methods**

103 We use the daily Optimum Interpolation SST (OISST.v2) records from the National  
104 Oceanic and Atmospheric Administration (NOAA) with a  $0.25^{\circ} \times 0.25^{\circ}$  grid in a  
105 period of 1982-2022 (Reynolds et al., 2007). The atmospheric components are from

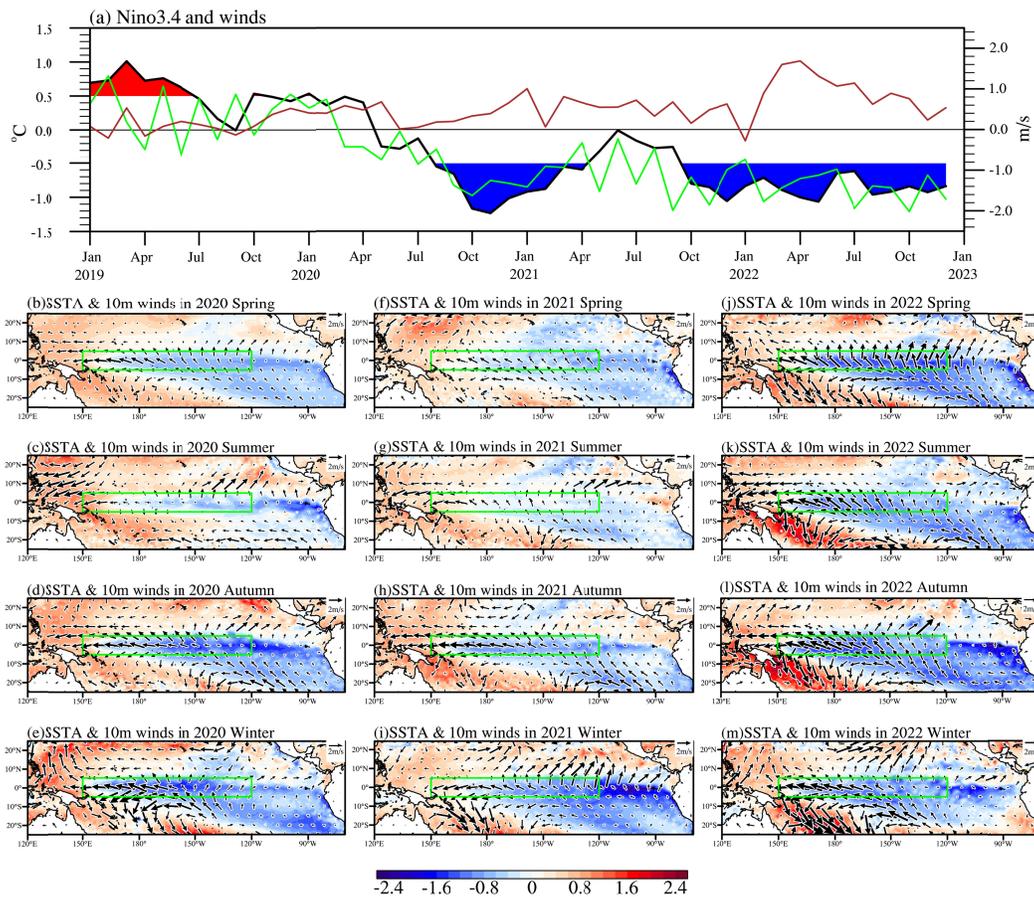
106 the daily JRA-55 Reanalysis dataset, released by the Japanese Meteorological Agency,  
107 with a horizontal resolution of  $1.25^{\circ} \times 1.25^{\circ}$  for the period of 1982-2022 (Kobayashi  
108 et al., 2015). The monthly mean outgoing longwave radiation (OLR) is provided by  
109 NOAA, with a horizontal resolution of  $2.5^{\circ} \times 2.5^{\circ}$  (Liebmann & Smith, 1996).

110 Here, the period of the annual cycle is defined as from July to June of the next  
111 year. Empirical orthogonal function (EOF) analysis is applied to show the annual  
112 cycle modes of climatological SST in the tropical Pacific. The climatology is defined  
113 as the 30-yr average between 1991 and 2020. The statistical methods used include  
114 Student's t-test, partial correlation, and linear regression. The ENSO events are  
115 defined when the 3-month running-mean Niño 3.4 index (averaged SSTA in the  
116 region of  $5^{\circ}\text{S}$ - $5^{\circ}\text{N}$ ,  $120^{\circ}\text{W}$ - $170^{\circ}\text{W}$ ) is above the threshold of  $0.5^{\circ}\text{C}$  for 5 consecutive  
117 months, consistent with the NOAA Climate Prediction Center's definition. Here,  
118 winter is defined from December to February of the next year (DJF). The results are  
119 not sensitive to the selections of different data sets and different periods of  
120 climatology.

### 121 **3. Time evolution of triple-dip La Niña in 2020-22**

122 Figure 1a shows that the Niño 3.4 index has been in a negative phase since the  
123 middle of 2020. Except for the summer of 2021, the index value was lower than the  
124 threshold of  $-0.5^{\circ}\text{C}$  from the mid of 2020 until the end of 2022, implying a triple-dip  
125 La Niña event. The fluctuations of the Niño 3.4 index and SSTAs along the equatorial  
126 Pacific are coherent with the variations of zonal wind anomalies along the equator

134 (Figure 1). For example, the strengthening of the easterly wind over the central  
 135 tropical Pacific in the autumn of 2020 is linked to the decline of the Niño 3.4 index,  
 136 and SST cooling in the central and eastern tropical Pacific, sustaining the cold SST  
 137 condition by the Bjerknes positive feedback (Bjerknes, 1969). On the other hand,  
 138 there have also been meridional wind anomalies over the central tropical Pacific since  
 139 the winter of 2019 (Figure 1a). It seems also important in triggering and evolution of a  
 140 La Niña event (Fang et al., 2023).

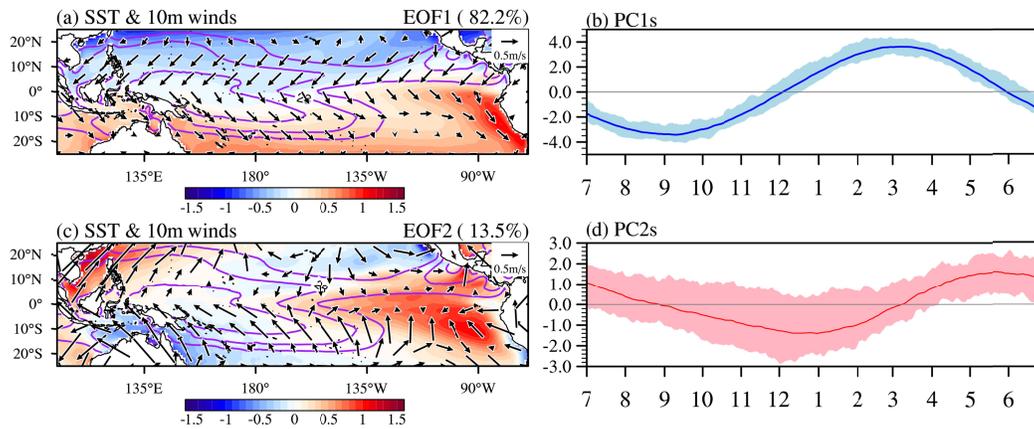


135 Figure 1. (a) Time series of Niño3.4 index (black line), averaged zonal wind anomaly (green line)  
 139 and meridional wind anomaly (brown line) in the region of green box (5°S-5°N, 150°E-120°W)  
 140 during 2019-2022. Horizontal distributions of seasonal SSTA (shading; °C) and 10m winds  
 141 anomalies (vector; m/s) for (b-e) 2020, (f-i) 2021, (j-m) 2022.  
 142

139        In autumn 2020, a La Niña pattern developed, characterized by strong trade wind  
140        over the central equatorial Pacific and southeasterly wind over the southeastern  
141        Pacific (Figure 1d). A similar anomaly pattern occurred in autumn 2021 and autumn  
142        2022 (Figure 1h, l). There was a turning point in the summer of 2021, and the cold  
143        SSTA weakened, accompanied by weak southeasterly wind over the central  
144        equatorial Pacific (Figure 1g). The southeasterly wind further strengthens in the next  
145        season (Figure 1h). Thus, the easterly over the central equatorial Pacific and  
146        southeasterly wind anomalies over the southeastern Pacific could be the important  
147        factors for the evolution of the triple-dip La Niña in 2020-22.

#### 148        **4. Impacts of annual cycle SST modes**

149        The seasonal mean SSTAs in the central and eastern tropical Pacific are negative  
150        during 2020-22. Spatially, in addition to the zonal SSTA gradient, there are  
151        appreciable meridional SSTA gradients in the tropical Pacific (Figure 1), which may  
152        be related to the two leading modes of the annual cycle in the tropical Pacific. Thus,  
153        there could be a linkage between the annual cycle of SSTA in the eastern Pacific and  
154        prolonged La Niña in 2020-22.



156

161 Figure 2. The EOF analysis of tropical Pacific daily mean climatological SST in the region of 100°  
 162 E-70°W, 25°S-25°N. (a, c) The two leading EOF modes of SST (shading; °C) and the  
 163 PCs-regressed surface wind (vectors; m/s). The purple lines in (a, c) indicate the annual mean SST.  
 164 (b, d) shows the first (PC1, solid blue line) and the second (PC2, solid red line) PC of the annual  
 165 cycle modes and their interannual variability (blue and red shades) from 1982-2022.

173 The essential characteristic of the annual cycle SST and ENSO seasonality can  
 174 be described by the first two EOF modes of the annual cycle of the climatological  
 175 SST in the tropical Pacific (Figure 2). EOF1 shows an antisymmetric mode, which  
 176 accounts for 82.2% of the total variance of SST. This mode is associated with the  
 177 annual cycle of solar radiation and represents the reverse change of SST between the  
 178 northern and southern hemispheres (Kim & Chung, 2001; B. Wang, 1994; X. L. Wang,  
 179 1994). Due to the huge thermal inertia of the ocean, the corresponding principal  
 180 component (PC1) lags solar radiation for about 2-3 months (Figure 2b) and reaches its  
 181 peak in March (the winter solstice occurs in late December). At this time, the SST  
 182 gradient between the northern and southern hemispheres is the largest, and the  
 183 trans-equatorial northerly wind prevails in the tropical Pacific (Figure 2a). Under the  
 184 influence of the Coriolis force, the northeast winds turn to northwest winds from the

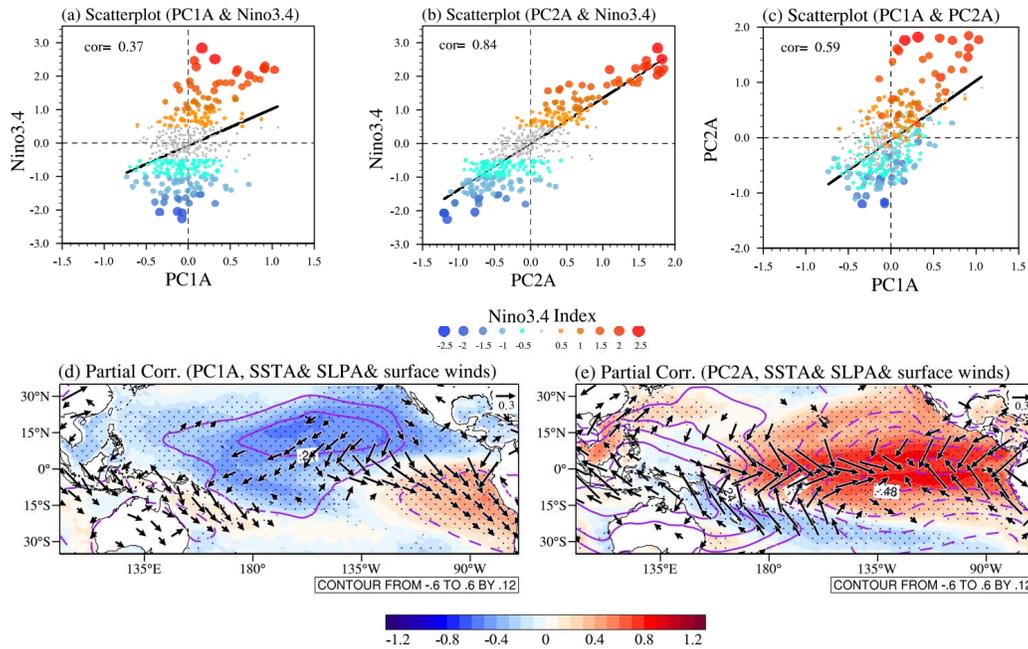
173 northern hemisphere to the southern hemisphere. In the positive phase of PC1, the  
174 weakened southeasterly trade winds off the coast of Peru can reduce the wind  
175 evaporation feedback cooling due to reduced wind speed (Shang-Ping Xie &  
176 Philander, 1994) and suppress the upwelling of the eastern tropical Pacific. On the  
177 other hand, the loading of the annual cycle mode is small in the warm pool region in  
178 the western tropical Pacific where SSTs exceed 28°C and the seasonal variation is  
179 weak.

180 EOF2 represents a symmetric variation to the equator in the central and eastern  
181 tropical Pacific (Figure 2c), with main loadings in the eastern equatorial Pacific which  
182 is the main region associated with ENSO. This mode reaches its peak in May, lagging  
183 PC1 by two months. The center of the warming is around 95°W along the equator,  
184 where the thermocline is shallowest (Meyers, 1979).

185 The two leading modes explain 95.7% of the total variance of annual cycle SST,  
186 which can capture the basic characteristics of the climatological seasonal evolution of  
187 SST in the tropical Pacific. To obtain the interannual variation of the annual cycle, the  
188 observed SSTs in each year are projected into the EOF1 and EOF2 of climate state  
189 (shading in Figure 2b, d). Here, PC1A (PC2A) is defined as the difference between  
190 PC1 (PC2) for each year and climatic PC1 (PC2). Since PC2A shows large  
191 interannual fluctuations in the winter and the largest loading of EOF2 is in the central  
192 and eastern equatorial Pacific, thus this mode is related to ENSO. The correlation  
193 coefficient between PC2A and Niño 3.4 index is 0.84 (Figure 3b), implying that the

194 interannual variation of EOF2 is associated with the ENSO evolution. Interestingly,  
195 PC1A is positively correlated with PC2A (Figure 3c) which is significant at a 99%  
196 level and suggests some non-orthogonality between the interannual modulation of  
197 symmetric mode and antisymmetric mode of annual cycle anomalies.

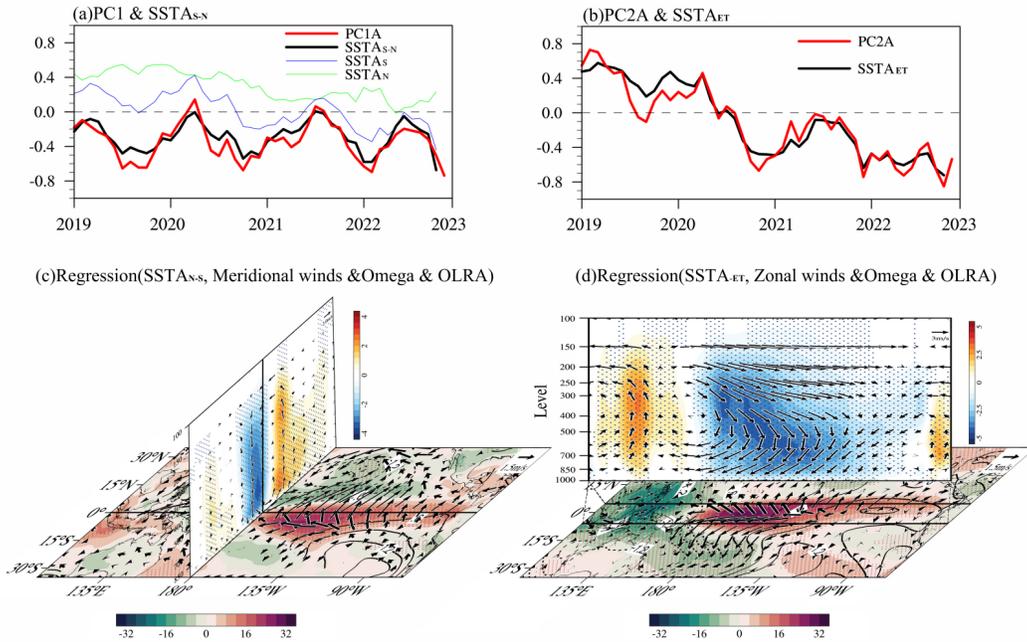
198       We compute partial correlation to investigate the relationship of the PC1A and  
199 PC2A with ENSO events. When excluding the PC1A component, PC2A shows an  
200 ENSO pattern (Figure 3e). The positive (negative) phase of symmetric mode (EOF2)  
201 is associated with the decrease (increase) in the zonal pressure gradient, westerly  
202 (easterly) wind anomaly, and positive (negative) SSTA in the central and eastern  
203 tropical Pacific. During the last three years, PC2A was in a negative phase and  
204 showed a downward trend just like the Niño 3.4 index (Figures 1a, 4b), meaning that  
205 the development of La Niña is related to the strengthening of the trade wind over the  
206 central tropical Pacific.



207  
 208 Figure 3. (a-c) are the scatterplots of PC1A and Niño 3.4 index, PC2A and Niño 3.4 index, and  
 209 PC1A and PC2A, respectively. Their linear relationship was shown by the black line. The color  
 210 and size of the dots represent the value of the Niño 3.4 index (greater than 0.5 for red, less than  
 211 -0.5 for blue, and others for gray). (d-e) The partial correlation of PC1A and PC2A with SSTA  
 212 (shading), SLPA (contour), and surface winds (vectors) from 1982-2022. The black vector arrows,  
 213 contour, and stippling indicate the region where the correlation coefficient is statistically  
 214 significant at the 95% confidence level.

215       When excluding the PC2A component, PC1A has a significant partial correlation  
 216 with the SSTA in the southeastern and northern Pacific. The positive phase of PC1A is  
 217 associated with an anomalous northwestern-southeastern gradient of SLP in the east  
 218 of the dateline (Figure 3d). The anomalous SLP gradient is associated with the  
 219 northwesterly wind anomalies over the southeastern Pacific, which reduces the  
 220 southeasterly trade winds and suppresses the upwelling along the coast of Peru.  
 221 Previous research has emphasized the contribution of SSTA in the southeastern  
 222 Pacific in the evolution of ENSO (H. Zhang et al., 2014; Zhu et al., 2016).  
 223 Interestingly, the PC1A has been in the negative phase since early 2019 (Figure 4a),

224 which is conducive to the persistent southeasterly wind prevailing over the  
 225 southeastern Pacific. The southeasterly wind anomalies over the southeastern Pacific  
 226 favor the persistence of La Niña condition. This suggests that the antisymmetric  
 227 (EOF1) mode may also play a role in the prolonged La Niña event in 2020-22.  
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229  
 230 Figure 4. (a) The time series of SSTA in the northern tropical Pacific ( $SSTA_N$ , averaged in  $0^\circ$ - $25^\circ$   
 231 N,  $110^\circ$ E- $80^\circ$ W) and SSTA in the southern tropical Pacific ( $SSTA_S$ , averaged in  $25^\circ$ S- $0^\circ$ ,  $110^\circ$ E- $80^\circ$   
 232 W) and the difference between them ( $SSTA_{S-N}$ ) and PC1A during 2019-2022; (b) The time series  
 233 of SSTA in eastern tropical Pacific ( $SSTA_{ET}$ , averaged in  $15^\circ$ S- $15^\circ$ N,  $180^\circ$ E- $80^\circ$ W) and PC2A  
 234 during 2019-2022. The  $SSTA_{N-S}$  regressed (c) pressure velocity (vertical shading;  
 235  $\omega \times 100$ Pa/s), meridional circulation (vectors; m/s) in the eastern Pacific (averaged in  $180^\circ$   
 236 E- $80^\circ$ W), SLPA (contour; hPa) and OLRA (horizontal shading;  $W/m^2$ ) from 1982 to 2022; the  
 237  $SSTA_{ET}$  regressed (d) pressure velocity, zonal circulation in the tropical Pacific (averaged in and  
 238  $5^\circ$ S- $5^\circ$ N), SLPA and OLRA from 1982 to 2022. The stippling and vector and contour in (c-d)  
 239 indicate the region where the regression coefficient is statistically significant at the 95%

240 confidence level.

241       The antisymmetric mode is associated with the seasonal evolution of the SST  
242 between the northern and southern hemispheres, which is represented by the SST  
243 difference between the southern (25°S-0,110°E-80°W) and northern (0-25°N, 110°  
244 E-80°W) tropical Pacific ( $SSTA_{S-N}$ ). The correlation coefficient between  $SSTA_{S-N}$   
245 and PC1A during 1982-2022 is 0.82 and significant at a 99% level. The index is  
246 negative in the recent three years (Figure 4a), which is consistent with the triple-dip  
247 La Niña in 2020-22.

248       In addition to that the overall warming trend is larger in the northern hemisphere  
249 than in the southern hemisphere (IPCC, 2021), the SSTA in the southern tropical  
250 Pacific ( $SSTA_S$ ) has even shown a cooling trend in recent years (Figure 4a), as a result,  
251 the negative  $SSTA_{S-N}$  persists and induces an anomalous meridional gradient of SLP  
252 in the east of the dateline (Figure 4c). The anomalous SLP gradient is favorable for  
253 maintaining the anomalous southeasterly over the eastern tropical Pacific in the recent  
254 three years. The southeasterly wind anomalies over the southeastern Pacific enhance  
255 the wind speed and cool down SST (Figure not shown) which (Figure 3d) can in turn  
256 further strengthen southeasterly wind, a wind-evaporation-SST feedback (Shang-Ping  
257 Xie et al., 2009). Thus, the spatially heterogeneous warming trends may also have an  
258 impact on the triple-dip La Niña in 2020-22. The anomalous SLP gradient and the  
259 cross-equatorial meridional wind associated with negative  $SSTA_{S-N}$  also produce an  
260 anomalous vertical circular circulation, which causes a weakening of the Hadley cell.

261 The anomalous downward (upward) motion over the eastern south (north) Pacific is in  
262 agreement with the OLR anomalies (Figure 4c).

263 On the other hand, the symmetric mode is associated with the seasonal evolution  
264 of SST in the eastern tropical Pacific ( $15^{\circ}\text{S}$ - $15^{\circ}\text{N}$ ,  $180^{\circ}\text{E}$ - $80^{\circ}\text{W}$ ). The correlation  
265 coefficient between  $\text{SSTA}_{ET}$  and PC2A is 0.92 and significant at a 99% level. The  
266 index is negative since the summer of 2020 (Figure 4b), linked to the strengthened  
267 trade wind over the central tropical Pacific via increasing the zonal SLP gradient  
268 (Figure 4d). The anomalous easterly wind and SLP gradient also strengthen the  
269 Walker circulation, accompanied by increased convection over the Maritime  
270 Continent. Thus, both the zonal wind associated with the symmetric and the  
271 meridional wind associated with the antisymmetric mode are conducive to the  
272 triple-dip La Niña in 2020-22.

## 273 **5. Summary and Discussion**

274 It has been pointed out that the persistent southeasterly winds over the tropical  
275 Pacific played a crucial role in the triple-dip La Niña in 2020-22 (Fang et al., 2022,  
276 2023). From the annual cycle SST perspective, we further analyzed the connection of  
277 the triple-dip La Niña event in 2020-22 with the anomalies of leading annual cycle  
278 modes. We argued that the easterly and southeasterly wind anomalies over the central  
279 and southeastern Pacific were associated with the anomalies of two leading annual  
280 cycle modes of SST.

281 The annual cycle SST in the tropical Pacific can be largely described by

282 antisymmetric and symmetric modes. The symmetric mode with loading mainly in the  
283 central and eastern tropical Pacific links to the temporal evolution of ENSO. Its  
284 positive (negative) phase is associated with weakened (enhanced) zonal SLP gradient  
285 and reduced (strengthened) trade winds over the central and eastern equatorial Pacific,  
286 and El Niño (La Niña) condition. While the antisymmetric mode with loading mainly  
287 in the off-equatorial Pacific reflects the seasonal meridional migration of SST  
288 between the northern and southern hemispheres. In its positive (negative) phase, the  
289 north-to-south SST gradient decreases (increases), which is associated with the  
290 weakening (strengthening) of the southeasterly wind over the southeastern and central  
291 and eastern equatorial Pacific, favorable for the growth of positive (negative) SSTA  
292 and El Niño (La Niña) in the central and eastern equatorial Pacific. The modes are  
293 both in negative phases since 2019 and are favorable for the development of the  
294 triple-dip La Niña event.

295       The spatial heterogeneity of warming trends may affect the SST contrast between  
296 the southern and northern hemispheres. That may modulate ENSO evolution (S. Hu &  
297 Fedorov, 2018). In the scenario of the increase in the SST contrast between the north  
298 and south Pacific in the future warming, more frequent prolonged La Niña events  
299 might be expected.

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314 **Data Availability Statement**

315 The daily SST of the NOAA OIv2.1 SST high-resolution dataset is downloaded from

316 <https://psl.noaa.gov/data/gridded/data.noaa.oisst.v2.highres.html>. The monthly mean

317 outgoing longwave radiation (OLR) from the U.S. NOAA is downloaded from

318 <https://psl.noaa.gov/data/gridded/data.olrcdr.interp.html>. The JRA - 55 reanalysis

319 dataset is obtained from the National Center for Atmospheric Research,

320 Computational and Information Systems Laboratory

321 (<https://doi.org/10.5065/D6HH6H41>).

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