

3 **Arctic Tropospheric Ozone Trends**

4 K. S. Law¹, J. Liengard Hjorth², J. B. Perno^{2,3}, C. H. Whaley⁴, H. Skov², M. Collaud Coen⁵, J.
5 Langner⁶, S. R. Arnold⁷, D. Tarasick⁸, J. Christensen², M. Deushi⁹, P. Effertz^{10,14}, G. Faluvegi^{11,12},
6 M. Gauss¹³, U. Im², N. Oshima⁹, I. Petropavlovskikh^{10,14}, D. Plummer⁴, K. Tsigaridis^{11,12}, S.
7 Tsyro¹³, S. Solberg¹⁵ and S.T. Turnock^{16,7}

8 ¹Sorbonne Université, LATMOS/IPSL, UVSQ, CNRS, Paris, France

9 ²Department of Environmental Science, Interdisciplinary Centre for Climate Change, Aarhus University,
10 Frederiksborgvej 4000, Roskilde, Denmark

11 ³Extreme Environments Research Laboratory, École Polytechnique fédérale de Lausanne,
12 1951 Sion, Switzerland

13 ⁴Canadian Centre for Climate modeling and analysis, Environment and Climate Change Canada, Victoria,
14 BC, Canada

15 ⁵Federal Office of Meteorology and Climatology, MeteoSwiss, Payerne, Switzerland

16 ⁶Swedish Meteorological and Hydrological Institute, Norrköping, Sweden

17 ⁷Institute of Climate and Atmospheric Science, School of Earth and Environment, University of Leeds,
18 Leeds, United Kingdom

19 ⁸Air Quality Research Division, Environment and Climate Change Canada, Toronto, ON, Canada

20 ⁹Meteorological Research Institute, Japan Meteorological Agency, Tsukuba, Japan

21 ¹⁰Cooperative Institute for Research in Environmental Sciences (CIRES), University of Colorado, Boulder,
22 CO, USA.

23 ¹¹NASA Goddard Institute for Space Studies, New York, NY, USA

24 ¹²Center for Climate Systems Research, Columbia University; New York, USA

25 ¹³Norwegian Meteorological Institute, Oslo, Norway

26 ¹⁴National Oceanic and Atmospheric Administration (NOAA) ESRL Global Monitoring Laboratory,
27 Boulder, CO, USA.

28 ¹⁵Norwegian Institute for Air Research (NILU), Kjeller, Norway.

29 ¹⁶Met Office Hadley Centre, Exeter, UK

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37 **Introduction**

38 The supporting information gives details about surface and ozonesonde data coverage (Text S1,
39 Figs. S1 and S2), the trend analysis method (Text S2), and the models (Text S3) used in this
40 study, and discusses trends of tropospheric ozone (O₃) precursors in the Arctic (Text S4).
41 Additional figures show O₃ trends at Arctic locations not shown in the main text, and trends over
42 different time periods for both the observations and models (Figs. S3-S9). Trends in surface
43 carbon monoxide (CO) (Fig. S10), and in low springtime O₃ concentrations at Canadian Arctic

44 sites (Fig. S11) are also shown. Tables S1 and S2 provide additional details about the models and
45 surface annual O₃ trends, respectively. See Fig. 1 in main text for measurement site locations.

46 **Text S1. Data coverage**

47 To ensure proper statistical representation of the data coverage thresholds at least 50% of
48 available data is required for the calculation of monthly and annual trends. Data coverage for the
49 surface sites is shown in Fig. S1. It is also verified that at least one complete year of data is
50 available at the beginning and end of the time series, since incomplete years at the beginning and
51 end of a time series can have a larger influence on trend analysis results (Collaud Coen et al.,
52 2020). Esrange has full data coverage, with less than five months having coverage lower than
53 50%. Pallas, Summit, Tustervatn and Zeppelin have one to three periods of 1-2 years without data
54 that do not preclude long-term trend analysis, although it should be noted that the Summit record
55 only started in June 2000. Alert and Villum have a high monthly data coverage but suffer from
56 missing data over a 4-5 year period in the middle of the time series. For Alert, there are two
57 missing periods throughout the time series. For Villum there is only one missing period although
58 the period 1996-2002 suffers from poor data coverage. Therefore, trends at Villum and Alert can
59 be considered as valid but should be interpreted with caution (Collaud Coen et al., 2020).

60
61 Ozonesondes were launched at least once per week at the six Arctic stations used in this analysis.
62 There are periods with up to three soundings a week, mostly in winter and spring (Fig. S2).
63 Periods without measurements do not exceed 1 month except at Eureka, where there are five
64 periods with missing data of three to eight months mostly in spring (2000, 2003, 2005, 2006, and
65 2016) and a 3-month period missing in 2006 at Sodankyla. The mean yearly number of soundings
66 varies between 38 (Resolute) and 90 (Ny Ålesund). At Alert and Sodankyla a lower number of
67 soundings were performed in 2016-2020. A visual inspection of the O₃ time series at the 20
68 pressure levels used in this analysis does not show evidence for any rupture apart from at the 4
69 lowest levels before 1995 at Eureka, potentially resulting in lower O₃ concentrations.

70 **Text S2. Trend analysis methodology**

71 Long-term trend determination needs to be applied to homogeneous time series in order to
72 analyze real variations or changes in the observations as opposed to artifacts. For the surface O₃
73 concentrations, while data is taken from quality-controlled repositories, this does not take into
74 account technical changes such as modification or changes in instrumentation, station position,
75 changes to calibration procedures or instrumental drifts. Visual inspection allowed detection of
76 potential ruptures in the time series.

77 Trends in surface O₃ concentrations at different sites are determined using the non-parametric
78 Mann-Kendall test and Sen's slope methodology (Theil, 1950; Sen, 1968), and calculated using
79 daily medians of the O₃ volume mixing ratios using hourly observations. These tests require the
80 data to be serially independent and homogeneously distributed. However, the O₃ measurements
81 are significantly lag-1 auto-correlated and exhibit high amounts of seasonality. Therefore, the
82 data is prewhitened to remove the lag-1 autocorrelation. The seasonal Mann-Kendall test is
83 applied to address the seasonality present in the data (Hirsch et al., 1982) using the Matlab
84 version (Collaud Coen and Vogt, 2021) of the 3PW algorithm from Collaud Coen et al. (2020).

85 The 3PW method uses two prewhitening methods prior to testing for statistical significance. The
86 first method from Kulkarni and von Storch (1995) simply removes the lag-1 autocorrelation
87 (referred to as PW). This method has a low rate of false positives but lowers the power of the
88 Mann-Kendall test. The second prewhitening method from Yue and Wang (2002), detrends the
89 data before prewhitening (referred to as TFPW-Y). This method returns the power of the Mann-
90 Kendall test although it increases the rate of false positives. Trends must be statistically
91 significant using both methods to be considered significant in the 3PW algorithm. If a significant
92 trend is present, then the slope is calculated using the variance corrected trend free prewhitening
93 (VCTFPW) method from Wang et al. (2015), which gives an unbiased estimate of the Theil Sen
94 slope. This method maximizes the advantages of these prewhitening methods and minimizes their
95 disadvantages. We use the seasonal Mann-Kendall test on a monthly temporal segmentation (i.e.,
96 a trend analysis is performed on each value in a respective month over the period analyzed). For
97 the O₃ soundings weekly resolution is applied since regular daily measurements are not available.
98 Relative monthly trends (% per year) for the 20 pressure levels of the ozonesonde data are
99 calculated as absolute monthly trends divided by monthly median concentrations x 100 %.
100 Statistically significant (ss) trends are determined at the 90th and 95th % confidence levels (CLs).
101 The Theil-Sen estimator is the median of all possible pairwise slopes. Upper CL and lower CL,
102 on the 95th % CL, contain the middle 95% of the pairwise slopes, which are normally distributed
103 (Gilbert, 1987). To test the robustness of the results, trends are also compared for different
104 periods of available records (1993-2019, 1993-2013, 1999-2019), and comparison with the model
105 results (1995-2015). Monthly medians are used to compare model results with the observations
106 since higher temporal resolution results are not available from the models. In all cases, model
107 results are averaged to produce multi-model median (MMM) results.

108 For trends in the yearly median, a simple Mann-Kendall test and Theil Sen Slope without
109 prewhitening are employed, using the MAKESENS application (Salmi et al., 2002), since there is
110 no seasonality present, and the data are not autocorrelated. The minimum requirement for
111 including observation series in this evaluation is data available from 1996 to 2018 in order to
112 have comparable trends. Nearly all of the stations have years with less than 50% data coverage
113 which are not included in the calculations of the annual trends. Relative trends are given as
114 percentages of median concentrations.

115 **Text S3. Models**

116 Four global chemistry-climate (CMAM, GISS-E2.1, MRI-ESM2, UKESM1) and two chemistry-
117 transport (DEHM, EMEP MSC-W) models were run using the same (ECLIPSEv6b)
118 anthropogenic emissions for 1990 (1995 for GISS model) to 2015 as part of the Arctic
119 Monitoring and Assessment Programme (AMAP) assessment (AMAP, 2021). They were run
120 with different biogenic emissions and meteorology and nudged using different reanalysis products
121 (see Table S1). The models also vary in their representation of the stratosphere with only a subset
122 of having a fully simulated stratosphere. CMAM, MRI-ESM2, GISS-E2.1, and UKESM1 contain
123 relatively complete stratospheric O₃ chemistry (involving nitrogen oxides (NO_x), chlorine and
124 bromine chemistry), while the other models have no stratospheric chemistry (DEHM, EMEP
125 MSC-W). Model results from 1995-2015 are used in this study.
126

127 Present-day model simulations for 2014-2015, including the models contributing to this study,
128 were evaluated first against a limited set of tropospheric Arctic O₃ observations (Whaley et al.,
129 2022), and also in more detail, including tropospheric Arctic O₃ seasonal cycles, in our
130 companion paper Whaley et al. (2023). These evaluations show large variability in model
131 performance. The MMMs capture surface Arctic O₃ observations quite well, except for low
132 concentrations observed during high Arctic polar spring (Whaley et al., 2023). Most atmospheric
133 models, including all of the models in this study, do not yet contain Arctic tropospheric halogen
134 chemistry, and thus cannot simulate the surface-level bromine and iodine-driven O₃ depletion
135 events that occur during spring at some Arctic locations (Whaley et al., 2023). Model
136 performance in the free troposphere is better (within +/- 10%) compared to satellite, aircraft and
137 ozonesonde data, but upper tropospheric O₃ is overestimated (Whaley et al., 2022, 2023).

138 **Text S4. Trends in tropospheric Arctic O₃ precursors**

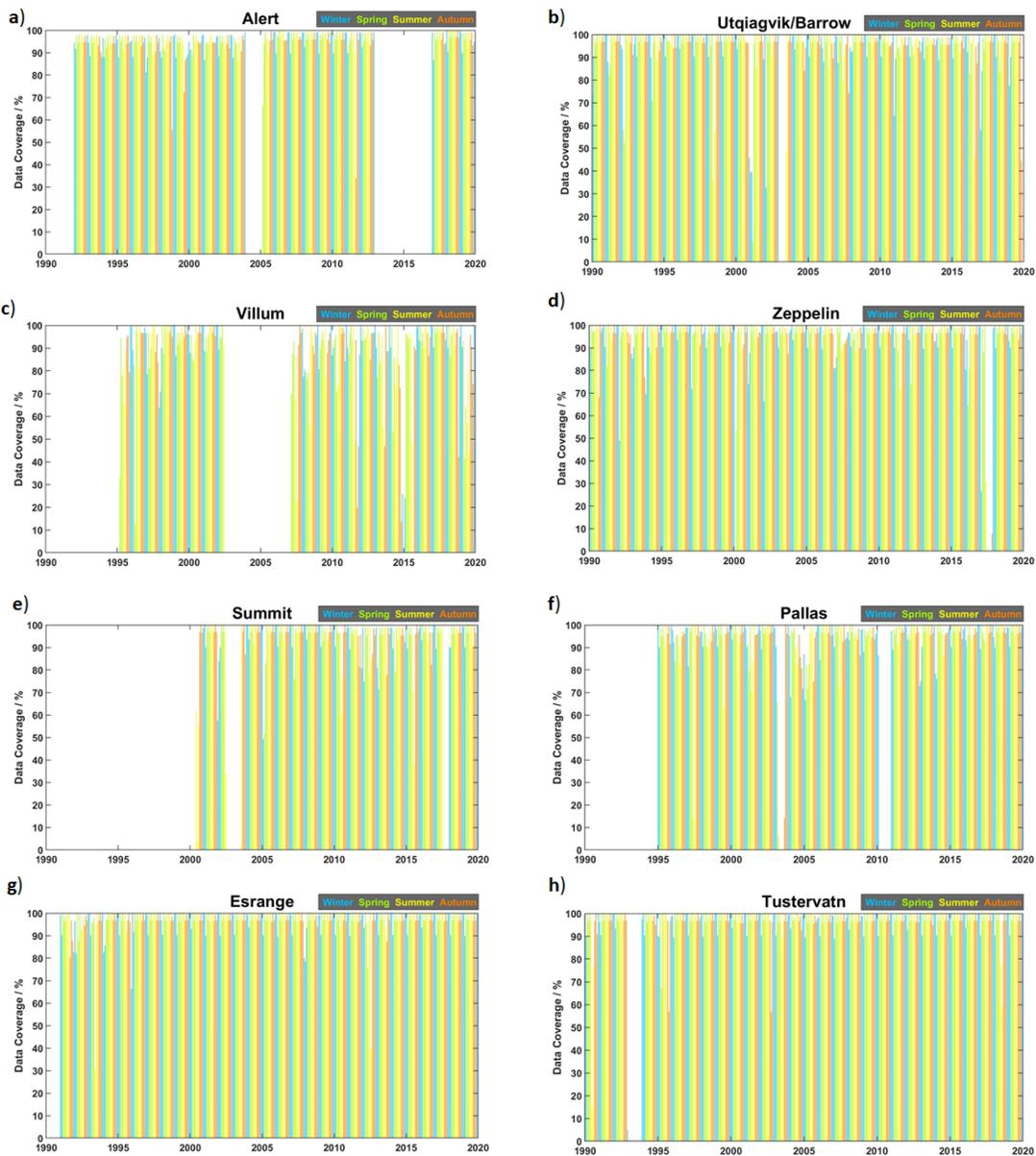
139 To understand Arctic tropospheric O₃ trends, it is important to understand trends in O₃ precursors
140 as well as other factors such as changing transport patterns. O₃ precursors include methane (CH₄),
141 CO, NO_x and non-methane volatile organic compounds (NMVOCs). While, as pointed out in the
142 main text, precursor trends over Northern Hemisphere (NH) mid-latitude emission regions are
143 likely contributing to changes in Arctic tropospheric O₃, it is also useful to examine precursor
144 trends in the Arctic, when long-term measurements are available, and to evaluate MMM trends.
145 The contribution of different sources, including precursor emissions, and sinks of Arctic O₃ is
146 discussed in Whaley et al. (2023) and references therein.

147 Arctic CH₄ at the surface rose from 1.75 ppmv in 1984 to 1.95 ppmv in 2020 (AMAP, 2021). The
148 increasing trend was interrupted by a plateau between 2000 and 2005 that accelerated after 2015
149 based on data from Pallas, Zeppelin and Utqiagvik (formerly known as Barrow). AMAP (2021)
150 derived an ss (95th % CL) observed annual trend of +2.29+/-0.55 ppbv per year, whereas modeled
151 CH₄, which was prescribed, had an annual trend of +2.79 ppbv per year, ss only at 90th % CL.
152 Thus, modeled and measured trends are reasonably comparable. Around 40% of Arctic O₃
153 response to precursor emission changes may be due to increasing CH₄ (AMAP, 2015).

154 Mackie et al. (2016) showed that CO decreased from 1989-2012 at Utqiagvik with the largest ss
155 decreases in winter and spring attributed to decreasing anthropogenic CO emissions in Europe
156 and North America. Figure S10 compares seasonal CO trends from observations and MMM
157 results at the few sites with long time series (Alert, Utqiagvik and Zeppelin). These trend
158 calculations are based on monthly averages, applying the 3PW algorithm, as discussed in 2.3 and
159 S2, for periods with available data. MMM results are in general agreement with the observations
160 showing ss negative trends in winter and autumn. However, it can be noted that the models
161 underestimate surface NH CO in winter (Whaley et al, 2023). This discrepancy is larger earlier in
162 the time series, while the models agree better with observed CO later in the time series (not
163 shown).

164 There are no reported trends in NO_x or NMVOCs in the Arctic due to a lack of long-term
165 measurements (AMAP, 2021). In general, observed NMVOC concentrations are low at Arctic
166 sites monitoring background air masses, and away from local sources (Pernov et al. 2021). Thus,

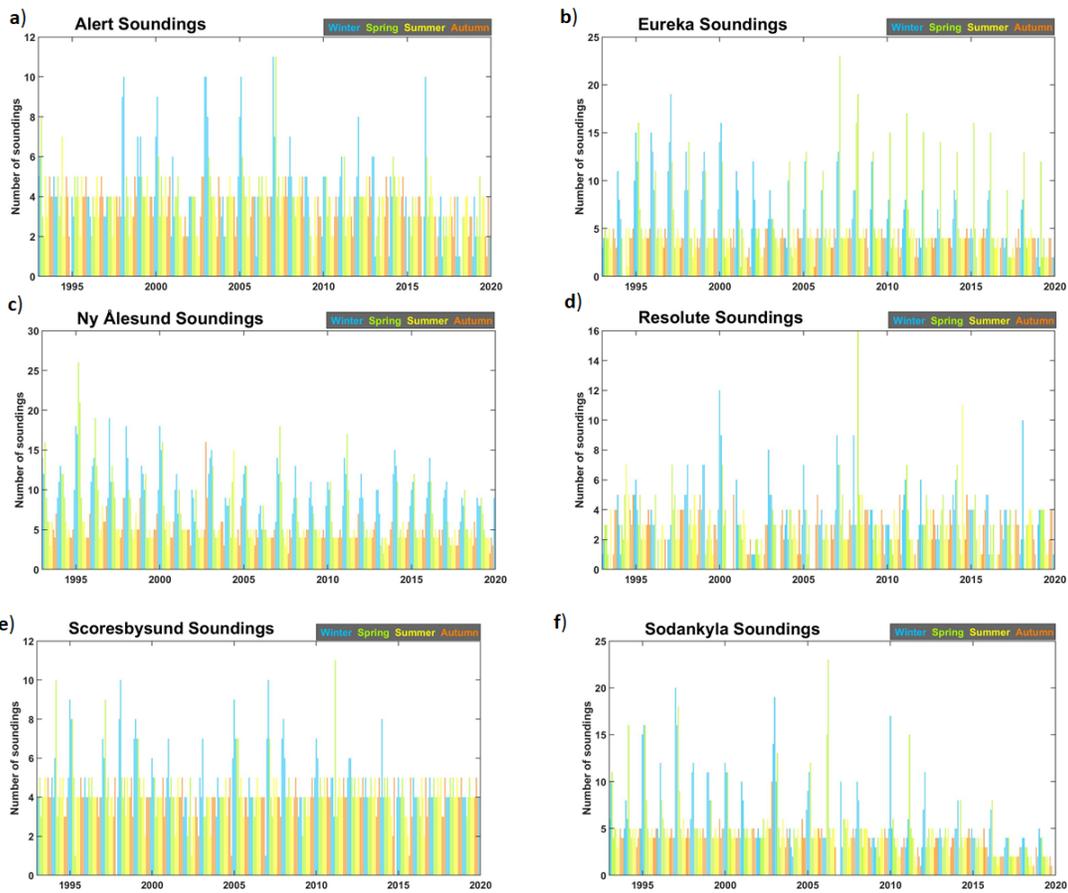
167 local NMVOC photo-oxidation only has limited effect on Arctic tropospheric O₃ as discussed by
168 Helmig et al. (2014) and Gautrois et al. (2003). However, long-range transport of air masses,
169 influenced by mid-latitude anthropogenic and natural NMVOC emissions, can influence Arctic
170 tropospheric O₃ (Whaley et al., 2023). NO_x concentrations are also generally low in the
171 background Arctic troposphere (Whaley et al., 2023), but can be elevated near local sources, for
172 example due to shipping emissions (e.g. Marelle et al., 2016). Also, long-range transport of NO_x
173 reservoir species such as peroxy-acetyl nitrate (PAN) or nitric acid can decompose in the warmer
174 spring and summer months producing NO_x (Law et al. (2014) and refs therein). Walker et al.
175 (2012) estimated that more than 50% of O₃ in the Arctic during summer is formed from local
176 PAN decomposition. Whaley et al (2023) showed that MMMs underestimate CO and NO_x
177 throughout the troposphere, but overestimates PAN compared to observed aircraft profiles in the
178 Arctic. Despite these differences, models match observed O₃ profiles in the troposphere, but
179 significantly overestimates O₃ near the tropopause, as noted earlier.



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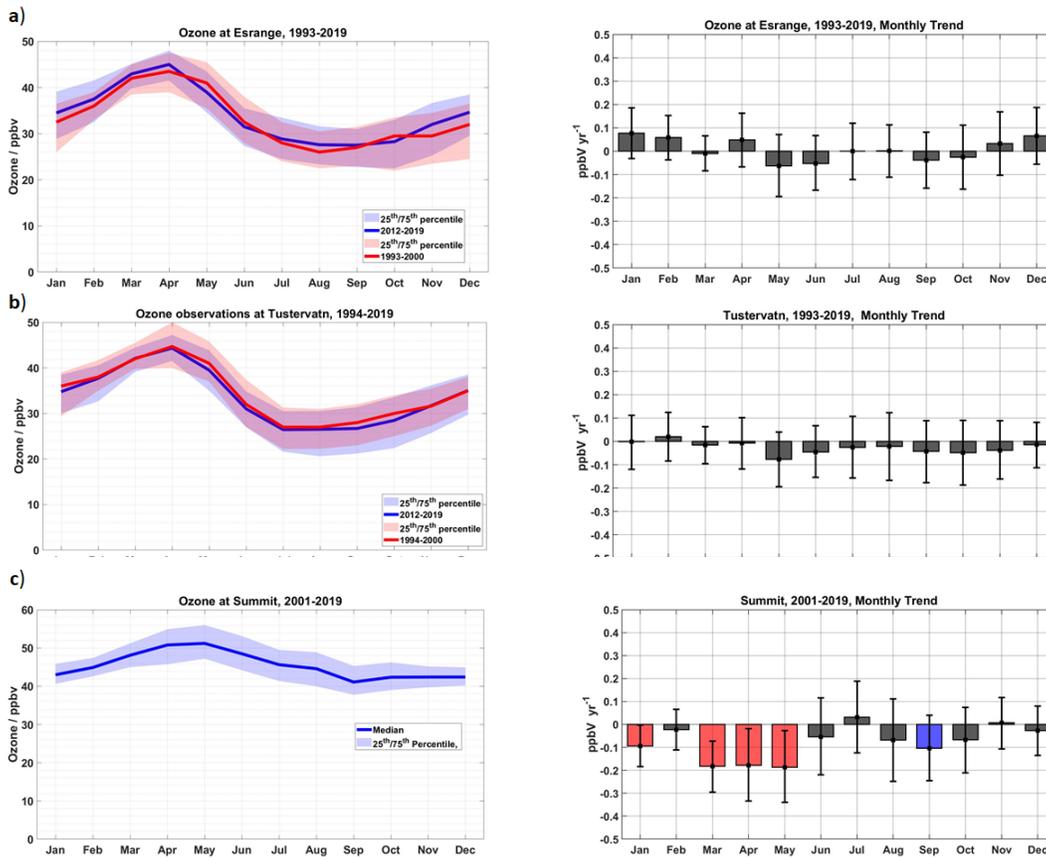
182 **Figure S1.** Data coverage for surface O_3 measurements at a) Alert, b) Utqiagvik, c) Villum, d)
 183 Zeppelin, e) Summit, f) Pallas, g) Esrange, and h) Tustervatn. Coverage is given as a percentage
 184 of hours with measurements compared to the total number of hours for each month. Monthly
 185 coverage is color-coded by the meteorological seasons. See Figure 1 for station locations.



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187

188 **Figure S2.** Ozonesonde data coverage at a) Alert, b) Eureka, c) Ny Alesund, d) Resolute, e)
 189 Scoresbysund, and f) Sodankyla. The number of O_3 soundings in each month is shown color-
 190 coded by the meteorological seasons. See Figure 1 for station locations.

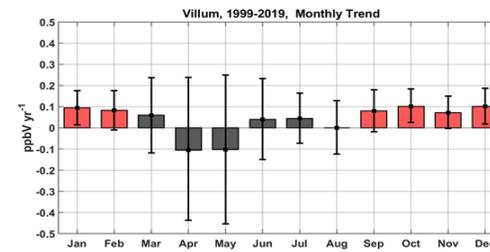
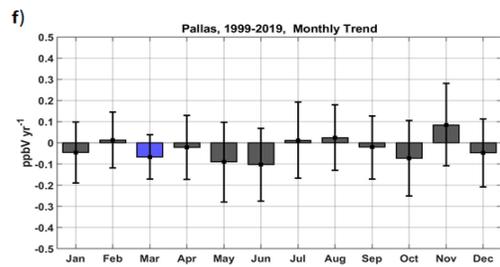
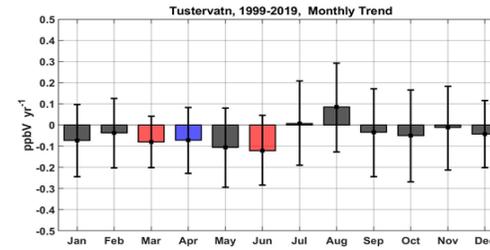
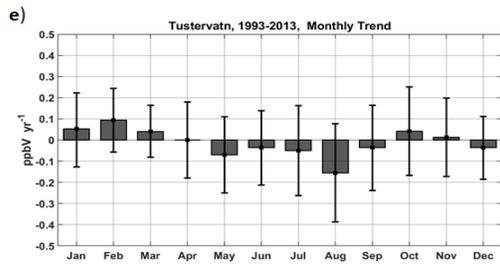
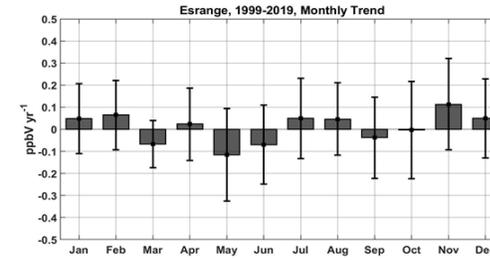
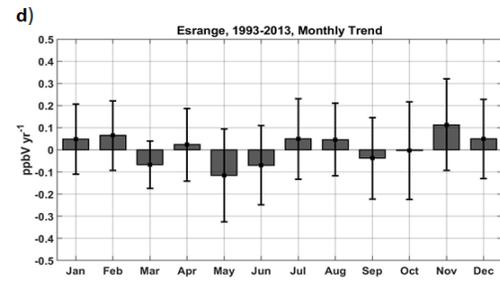
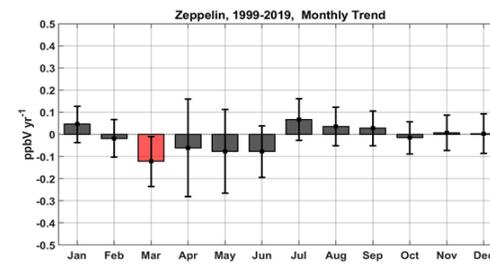
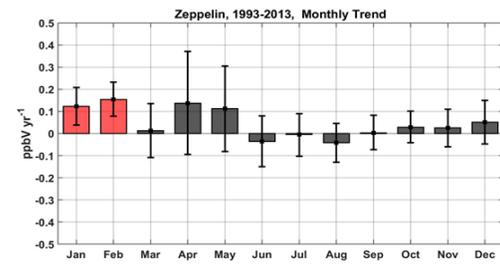
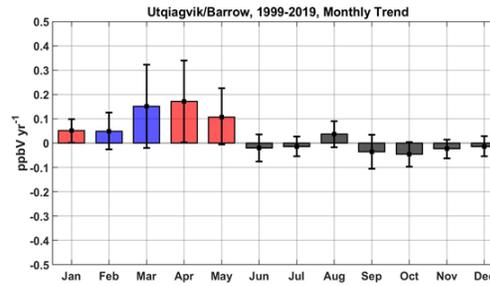
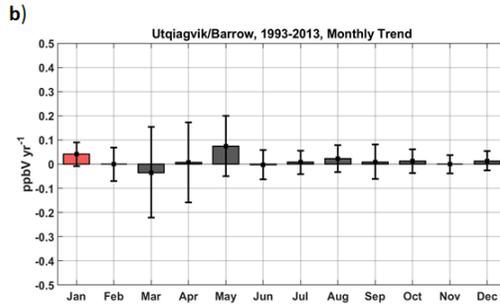
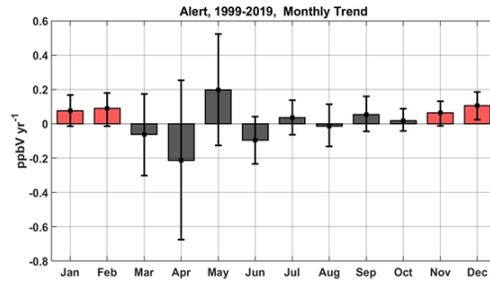
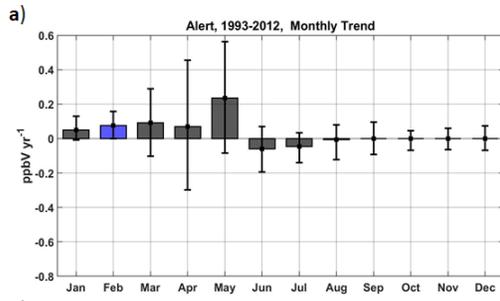


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193 **Figure S3.** Observed surface O_3 trends and seasonal cycles. Left panels: seasonal cycles of
 194 monthly median O_3 (ppbv) at a) Esrange, and b) Tustervatn for 1993-2000 (blue lines) vs 2012-
 195 2019 (red lines), and c) Summit for 2001-2019 only. Shaded areas show upper and lower
 196 quartiles of hourly values. Right panels: monthly trends over 1993-2019, or for shorter periods
 197 depending on data availability. Boxes represent the slope of the trend in ppbv per year with red
 198 boxes significant at 95th% CL, blue boxes at 90th% CL, and black boxes not statistically
 199 significant. Error bars show 95th% CLs.

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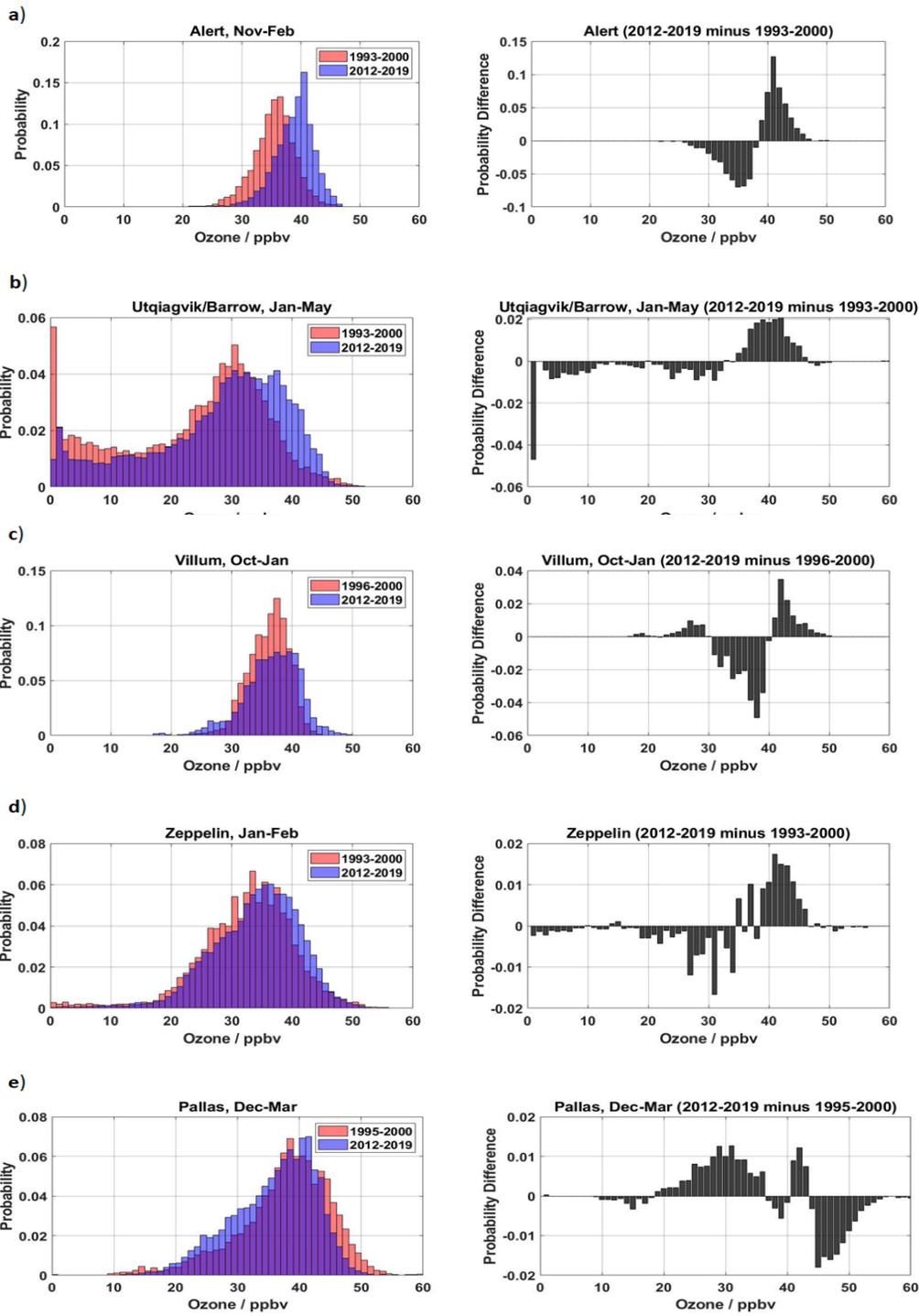
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204 **Figure S4.** Comparison of observed surface O_3 monthly median trends in ppbv per year for
205 different time periods at a) Alert, b) Utqiagvik, c) Zeppelin, d) Esrange, e) Tustervatn for 1993-
206 2012 and 1999-2019, depending on data availability, and f) Pallas (left) and Villum (right) for
207 1999-2019, when good data coverage is available. Box coloring and error bars are the same as
208 Fig. S3 (right panels).

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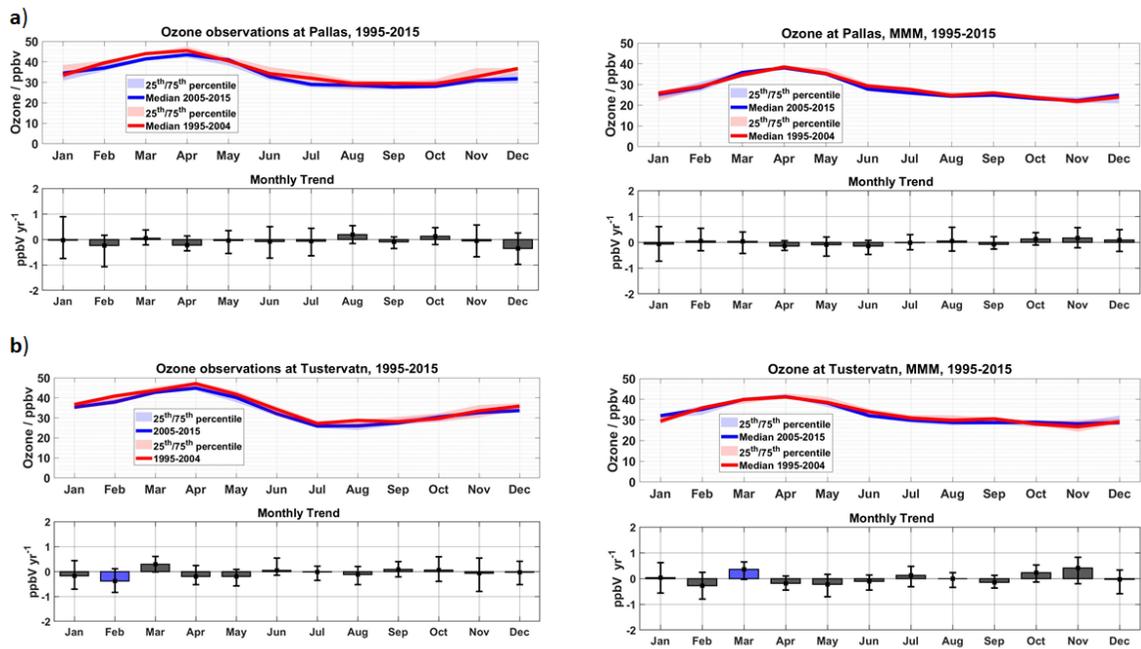


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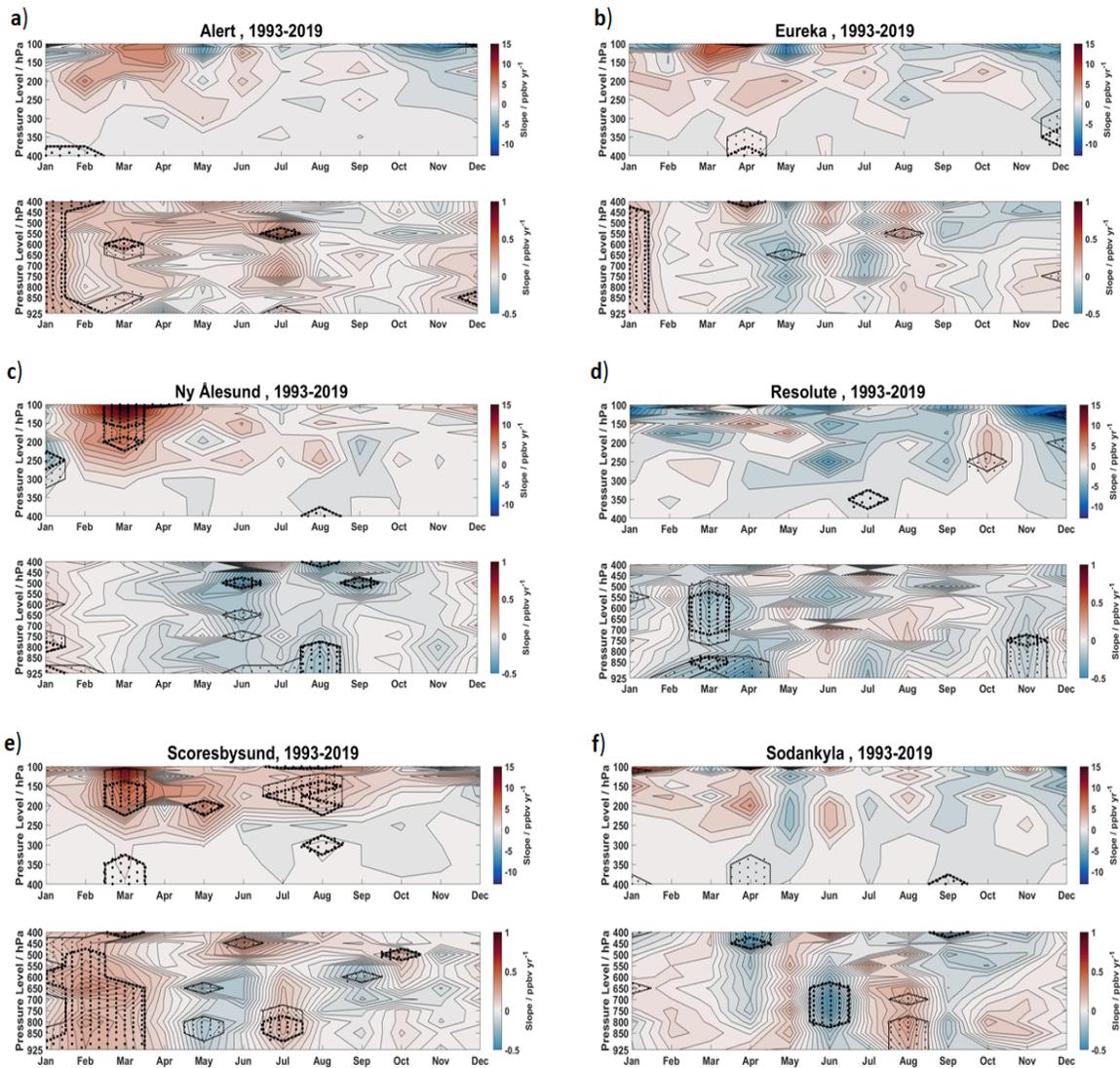
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213 **Figure S5.** Changes in probability distributions of surface O_3 concentrations in ppbv between
 214 earlier (1993-2000) and later (2012-2019) periods for selected months with statistically
 215 significant trends (see Figs. 2 and S3) at a) Alert, b) Utqiagvik, c) Villum, d) Zeppelin, and e)
 216 Pallas. Periods shown vary depending on data availability.



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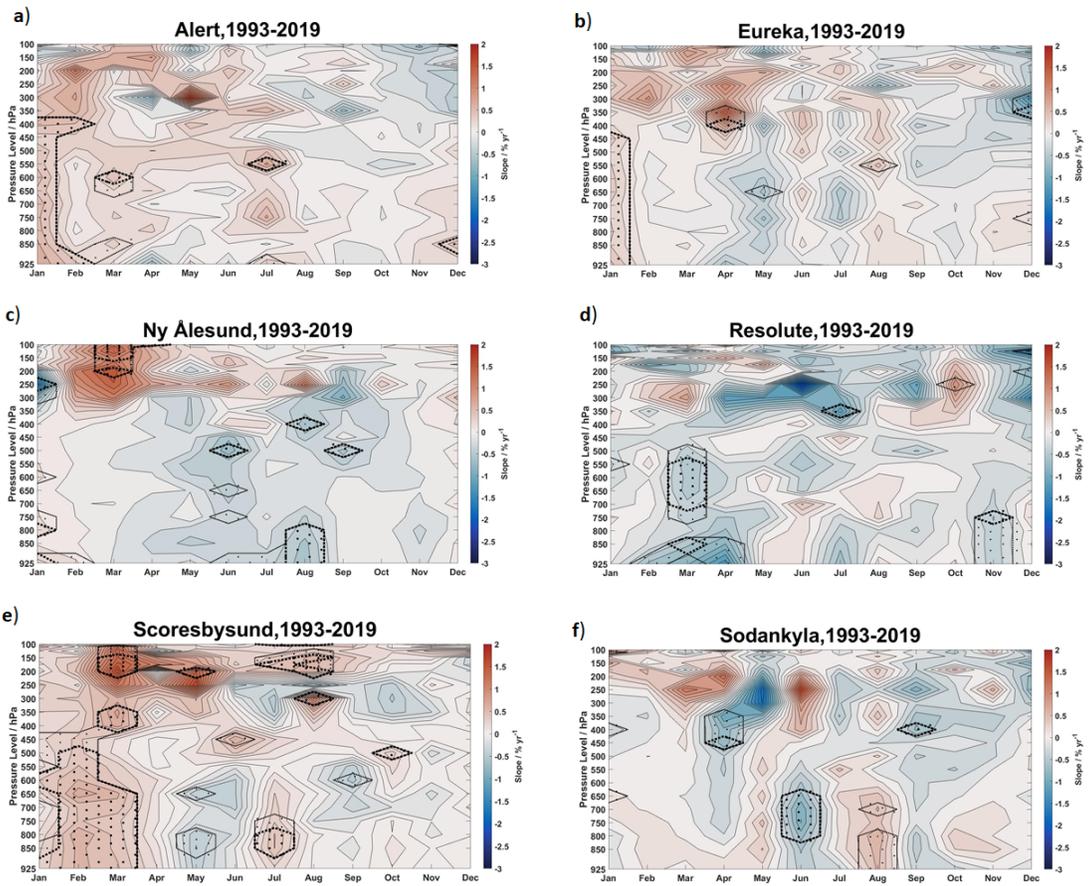
218 **Figure S6.** Comparison of observed (left) and MMM (right) surface O_3 trends and seasonal
 219 cycles at a) Pallas and b) Tustervatn. Upper panels: seasonal cycles for 1995-2004 (red lines) vs
 220 2005-2015 (blue lines). Shaded areas show upper and lower quartiles of hourly values
 221 (observations only). Lower panels: monthly median trends in ppbv per year for 1995-2015. Box
 222 coloring and error bars are the same as Fig. S3. Shorter periods are shown depending on data
 223 availability.



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226 **Figure S7a.** Absolute vertical trends in observed monthly O_3 for 1993-2019 in ppbv per year at a)
 227 Alert, b) Eureka, c) Ny Alesund, d) Resolute, e) Scoresbysund and f) Sodankyla. Upper panels:
 228 upper troposphere and lower stratosphere (400-100 hPa). Lower panels: mid- and lower
 229 troposphere (925-400 hPa). Note the different color scales. Stippled lines/areas show statistical
 230 significance at the 90th % CL (smaller marker size) and 95th % CL (larger marker size).



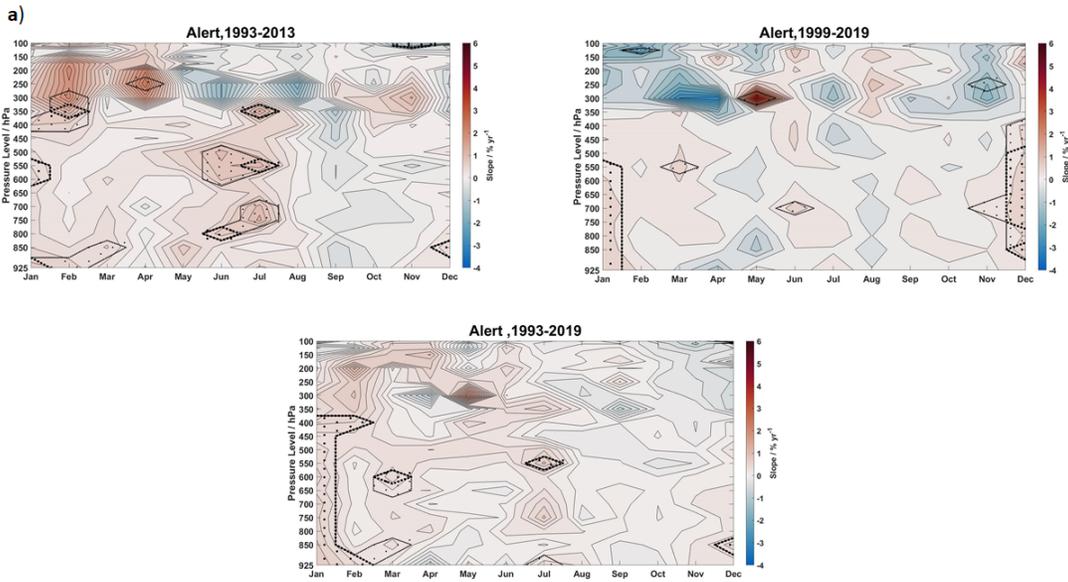
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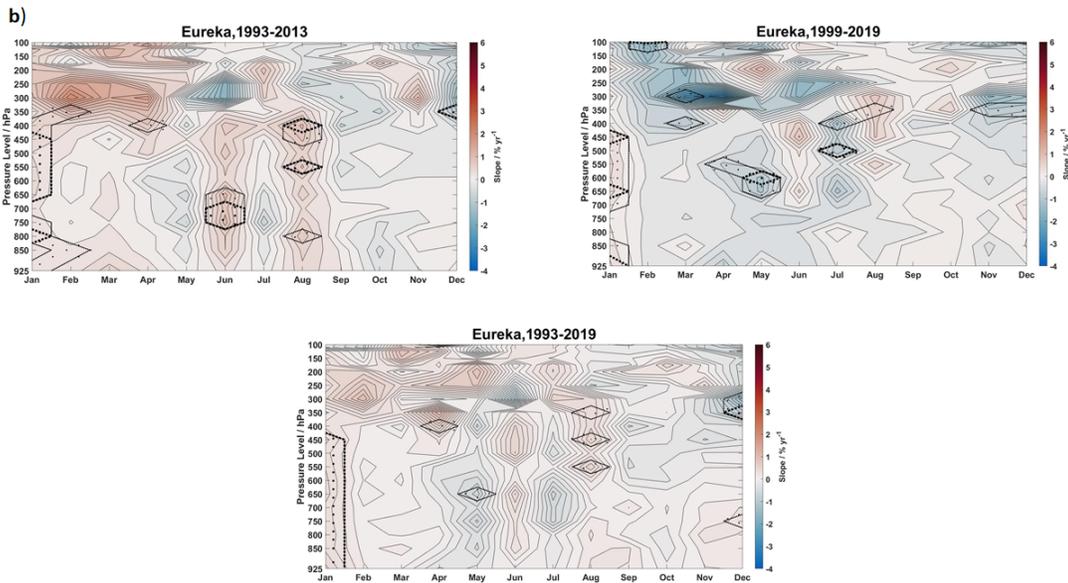
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234 **Figure S7b:** Same as Figure S7a, but vertical trends from 925-100 hPa in observed monthly O_3
 235 for 1993-2019, relative to monthly median concentrations, in % per year. Shading and symbols
 236 are the same as Fig S7a.

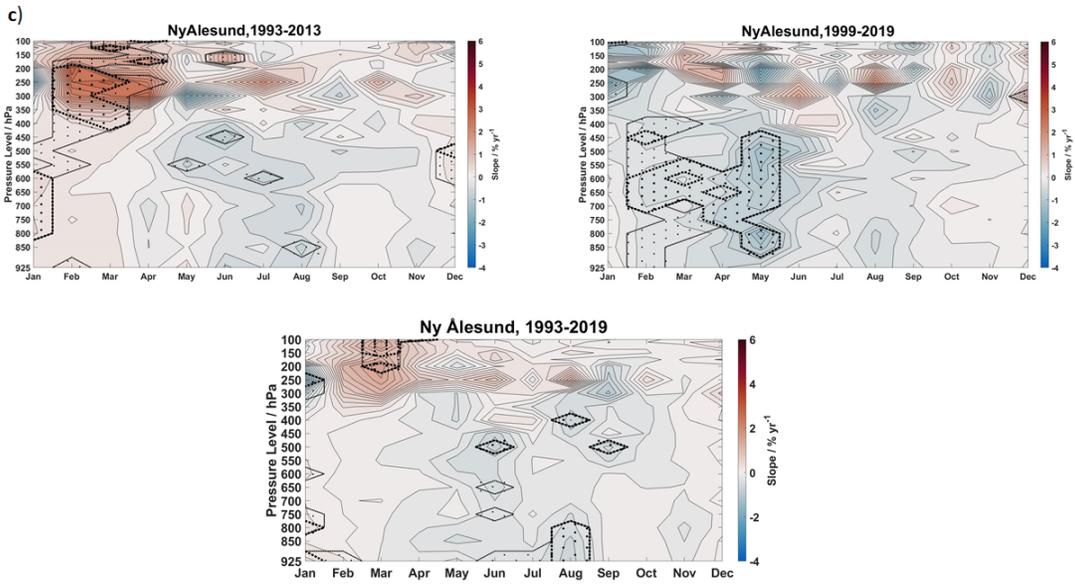
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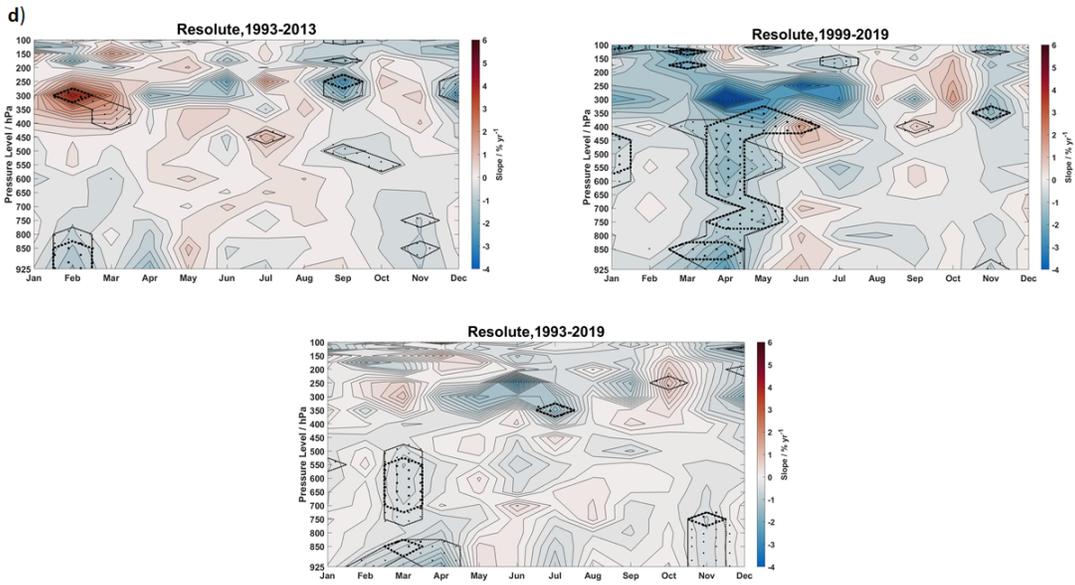
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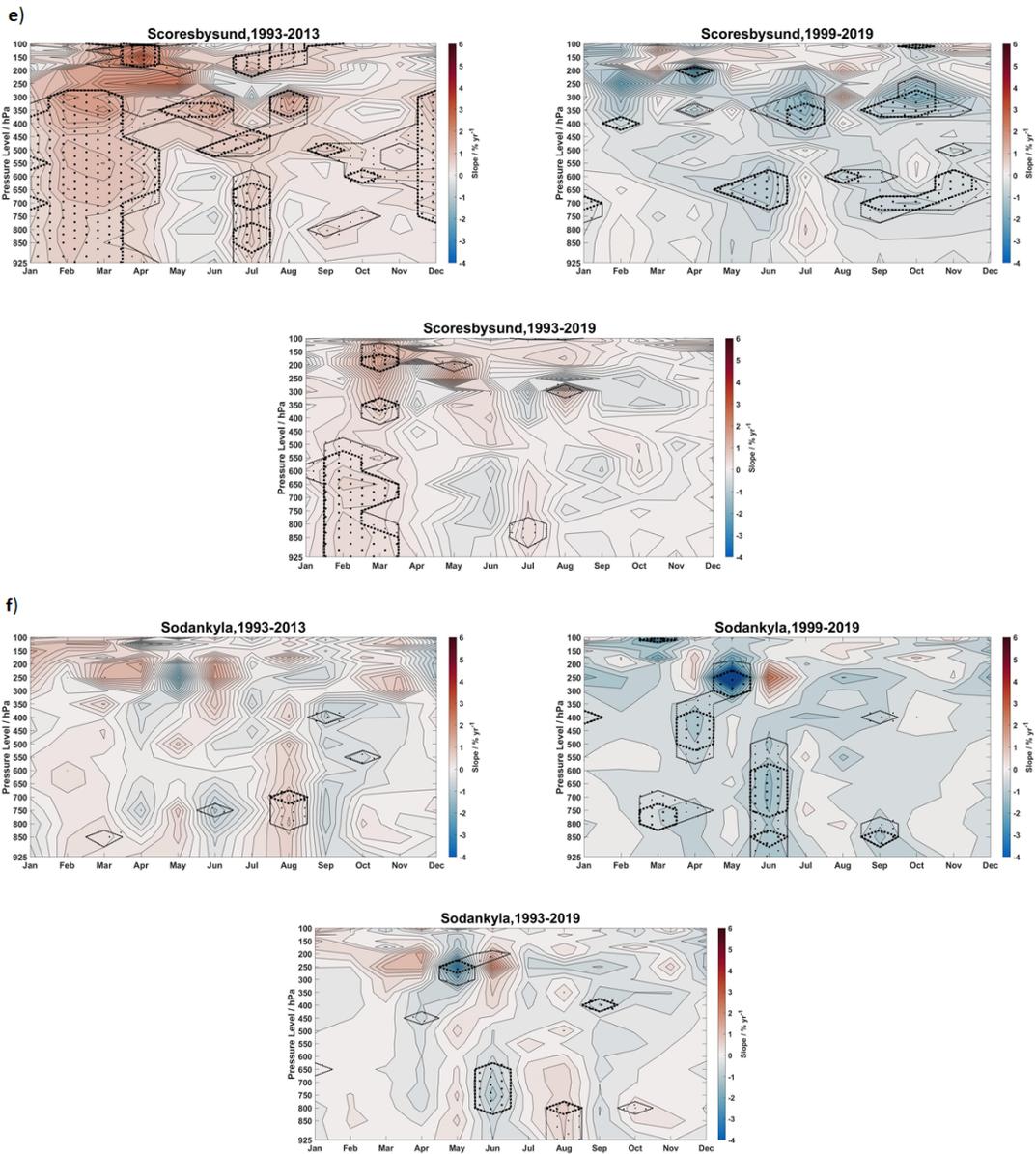
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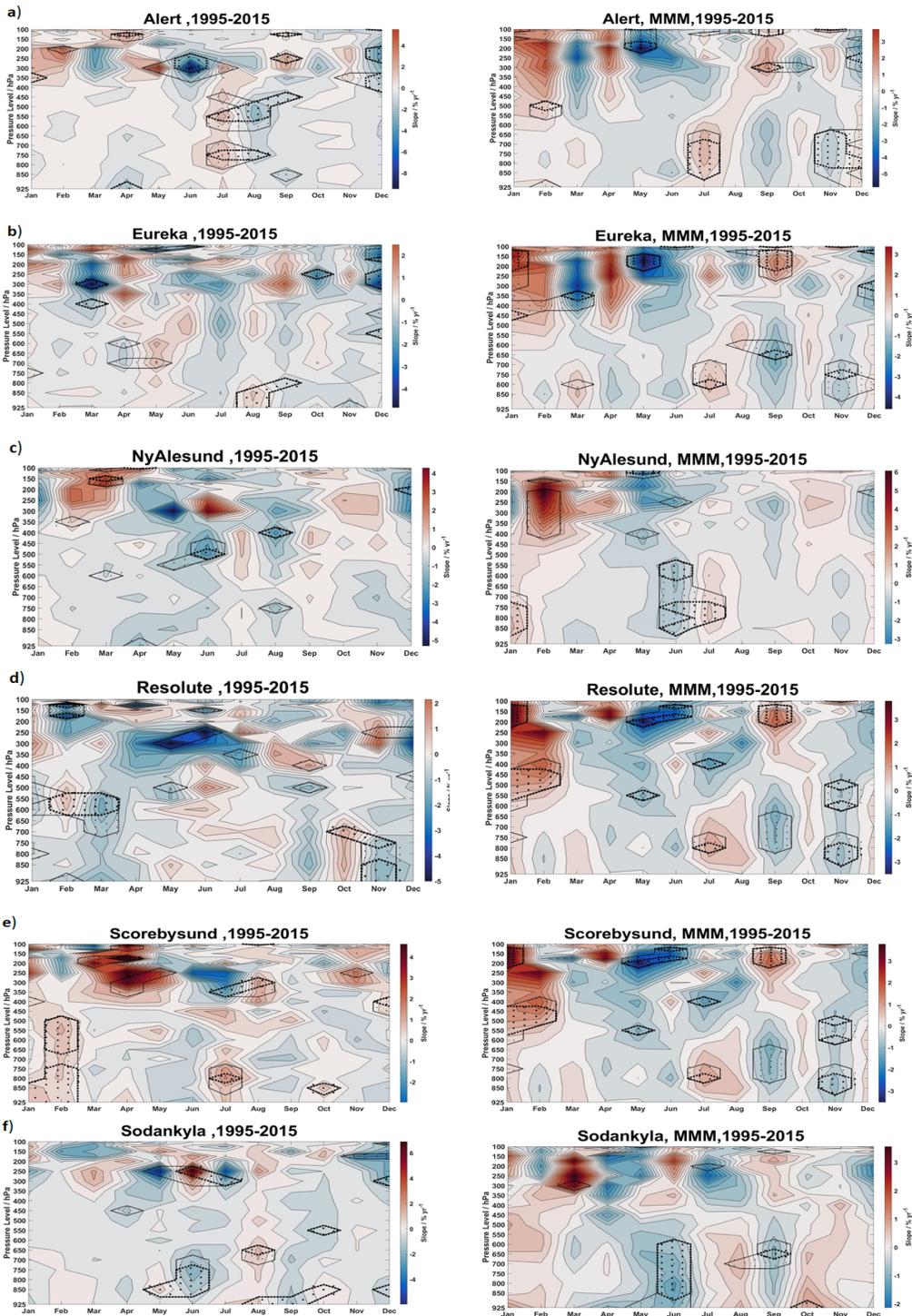
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245 **Figure S8.** Vertical trends in observed monthly O_3 , relative to monthly median concentrations, in
 246 % per year, for different time periods (1993-2013, 1999-2019, and 1993-2019) from 925-100 hPa
 247 at a) Alert, b) Eureka, c) Ny Alesund, d) Resolute, e) Scoresbysund, and f) Sodankyla. Shading
 248 and symbols are the same as Fig S7a.

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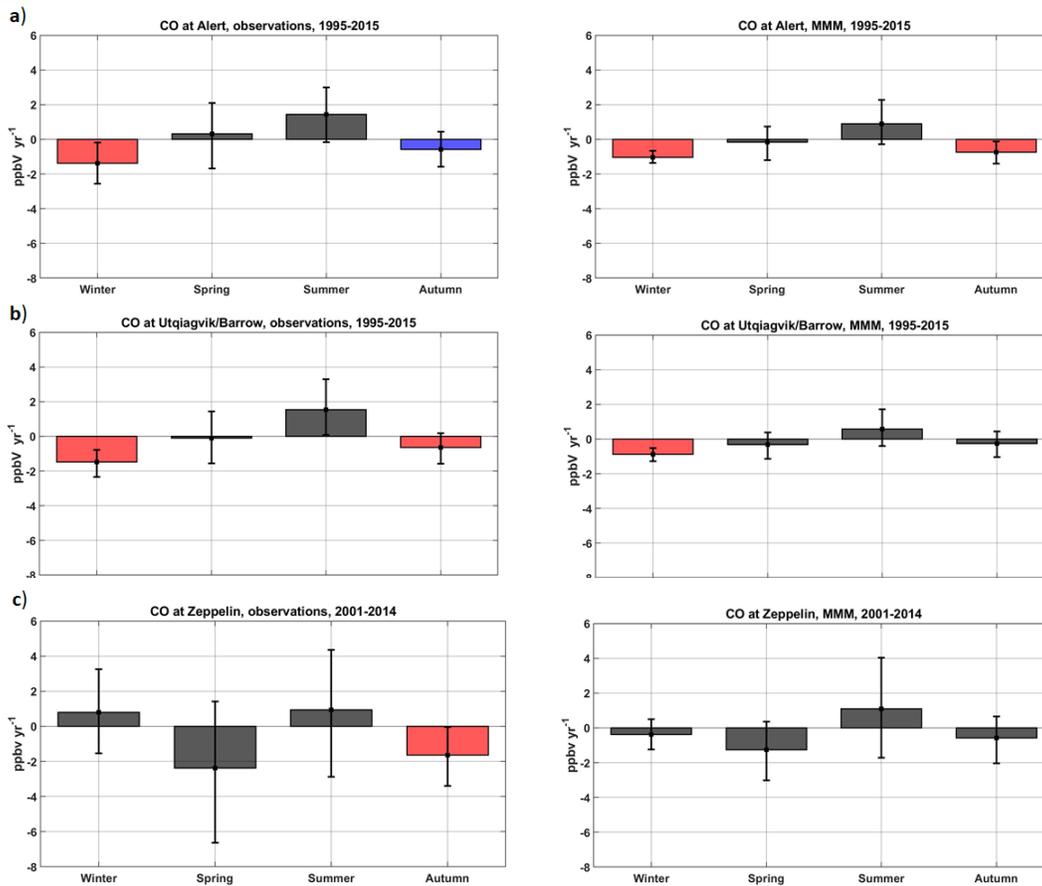
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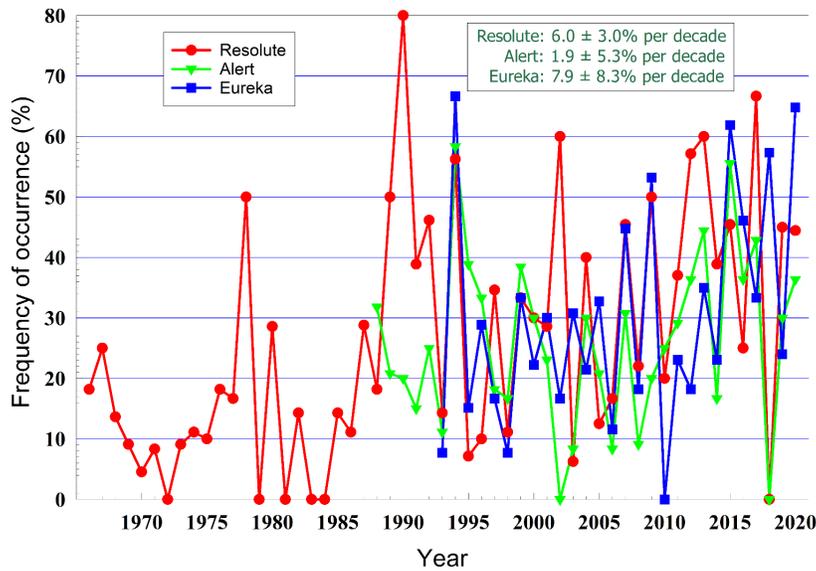
Figure S9: Comparison of observed (left) and MMM (right) vertical trends in monthly O_3 , relative to monthly medians, in % per year, from 925-100 hPa over 1995-2015 at a) Alert, b) Eureka, c) Ny Alesund, d) Resolute, e) Scoresbysund, and f) Sodankyla. Shading/symbols are the same as Fig. S7a.



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258 **Figure S10.** Observed (left) and MMM (right) seasonal surface CO trends at a) Alert, b)
 259 Utqiagvik, and c) Zeppelin. Boxes represent the slope of the trend in ppbv per year with red boxes
 260 significant at 95th % CL, blue boxes at 90th % CL, and black boxes not statistically significant.
 261 Error bars show 95th % CLs.



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Figure S11. Frequency of occurrence in % of March to May low O_3 concentrations, defined as <10 ppbv and indicative of O_3 depletion events, using ozonesonde data at the surface, and the first measurement level after balloon release, at Canadian Arctic sites from 1966–2020. Data are adjusted for the effects of occasional variation in sounding frequency. Details about the methodology are given in Tarrasick and Bottenheim (2002).

269 **Table S1.** Emissions and meteorology used in the models. See Whaley et al. (2022) for further
 270 details.

Model name	Biogenic emissions	Forest fire emissions	Meteorology
CMAM	None	CMIP6	Nudged to ERA-Interim reanalysis
DEHM	MEGANv2	GFAS	Nudged to ERA-Interim reanalysis
EMEP MSC-W	EMEP scheme (Simpson et al., 2012)	FINN (based on Wiedinmyer et al., 2011)	Driven by 3-hourly data from the Integrated Forecast System (IFS) at ECMWF
GISS-E2.1	Isoprene:Guenther et al. (2012); Terpenes: ORCHIDEE; Online DMS, Sea-salt and dust	CMIP6	Nudged to NCEP reanalysis
MRI-ESM2	Biogenic VOCs emissions are taken from Horowitz et al. (2003)	CMIP6	Nudged to the Japanese 55-year Reanalysis (JRA55)
UKESM1	Isoprene and monoterpenes interactive with land surface vegetation scheme	Prescribed from CMIP6 dataset	Nudged to ERA-Interim reanalysis

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272 **Table S2.** Annual surface O₃ trends in % per year based at Arctic sites. ‘Significance level’ is
 273 the probability that the observed trend is not the result of random variations. Lower and upper 95
 274 % CLs are also shown together with the significance (confidence) level for the annual trends
 275 calculated over the periods indicated. Statistically significant trends (above 90% CL) are in bold.
 276 The far right column displays the number of years included in the trend calculation compared to
 277 the maximum number of years over the period considered.

Site	Annual trend (%)	Lower 95% CL	Upper 95% CL	Significance level	Period	Number yrs included/ max. yrs
Alert	0.29	0.00	0.75	95%	1999-2019	16/21
	0.24	0.00	0.60	95%	1993-2019	22/27
Utqiaġvik	0.53	-0.32	1.18	<90%	1999-2019	20/21
	0.26	-0.21	0.76	<90%	1993-2019	26/27
Villum	1.98	0.09	3.11	95%	1999-2019	16/21
	0.68	-0.28	2.30	<90%	1996-2019	18/24
Zeppelin	-0.19	-0.45	0.15	<90%	1999-2019	20/21
	0.18	-0.03	0.46	90%	1993-2019	26/27
Summit	-0.28	-0.77	0.12	<90%	2001-2019	19/19
Esrange	0.08	-0.37	0.64	<90%	1999-2019	21/21
	0.00	-0.24	0.38	<90%	1993-2019	27/27
Pallas	-0.30	0.85	-0.06	90%	1999-2019	21/21
	-0.40	-0.84	0.00	90%	1995-2019	25/25
Tustervatn	-0.52	-1.14	-0.08	99%	1999-2019	21/21
	-0.18	-0.51	0.08	<90%	1994-2019	26/26

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