



PATH EFFECTS ON SURFACE-WAVE AMPLIFICATION IN SEDIMENTARY BASINS (S23C-0546)

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MOTIVATIONS

- Surface waves propagating in sedimentary basins strongly affect earthquake ground motions and cause strong damage (*Kawase, 1996*)
- *Bowden & Tsai (2017)* proposed a 1-D semi-analytical method to predict surface-wave basin amplification between two sites
- 1-D approximation of the near-surface geologic structure does not account for path effects (reflections, conversions)

⇒ **The current study aims to provide quantitative estimates of the importance of these various path effects** on surface-wave amplification and also extend the current 1-D theory to more complex multi-dimensional basin structures.

1 - SEMI-ANALYTIC MODELS

Pure 1D theory

- **Conservation of energy flux**
⇒ relative amplitude between two sites:

$$\frac{A_n}{A_n^R} = \frac{u_n(0)}{u_n^R(0)} \left(\frac{U I_0}{U^R I_0^R} \right)^{-1/2},$$

where A_n wave amplitude at the surface,
 $u_n(0)$ surface-wave eigenfunction amp. at the surface,
 U group velocity, $I_0 = \int_0^\infty \rho(z)(u_1(z)^2 + u_2(z)^2)dz$

- Neglect path effects (reflections and mode conversions)

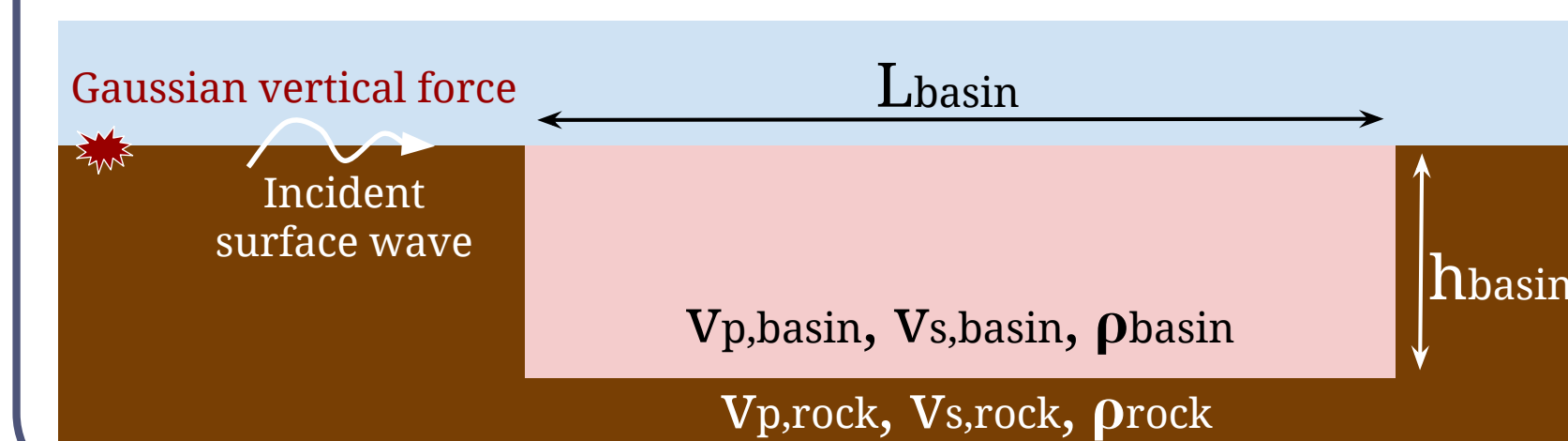
Eigenfunctions: computed using Computer Program in Seismology (*Herrmann, R. B., 2013*).

Reference solutions: high-order numerical axisymmetric solutions from SPECFEM package (*Komatitsch & Vilotte 1998*).

1D Theory w/ transmission coef. (SWRT)

- *Levshin (1989)* At a vertical boundary
⇒ wavefield = **incident, reflected and transmitted waves**
- Approximation of reflection/transmission coef. by **Green's function method** of *Its and Yanovskaya (1985)*
- Numerical code for transmission coef. by *Datta (2018)* named Surface Wave Reflec. Trans. (SWRT)
- Neglect body-wave diffraction at the basin edge

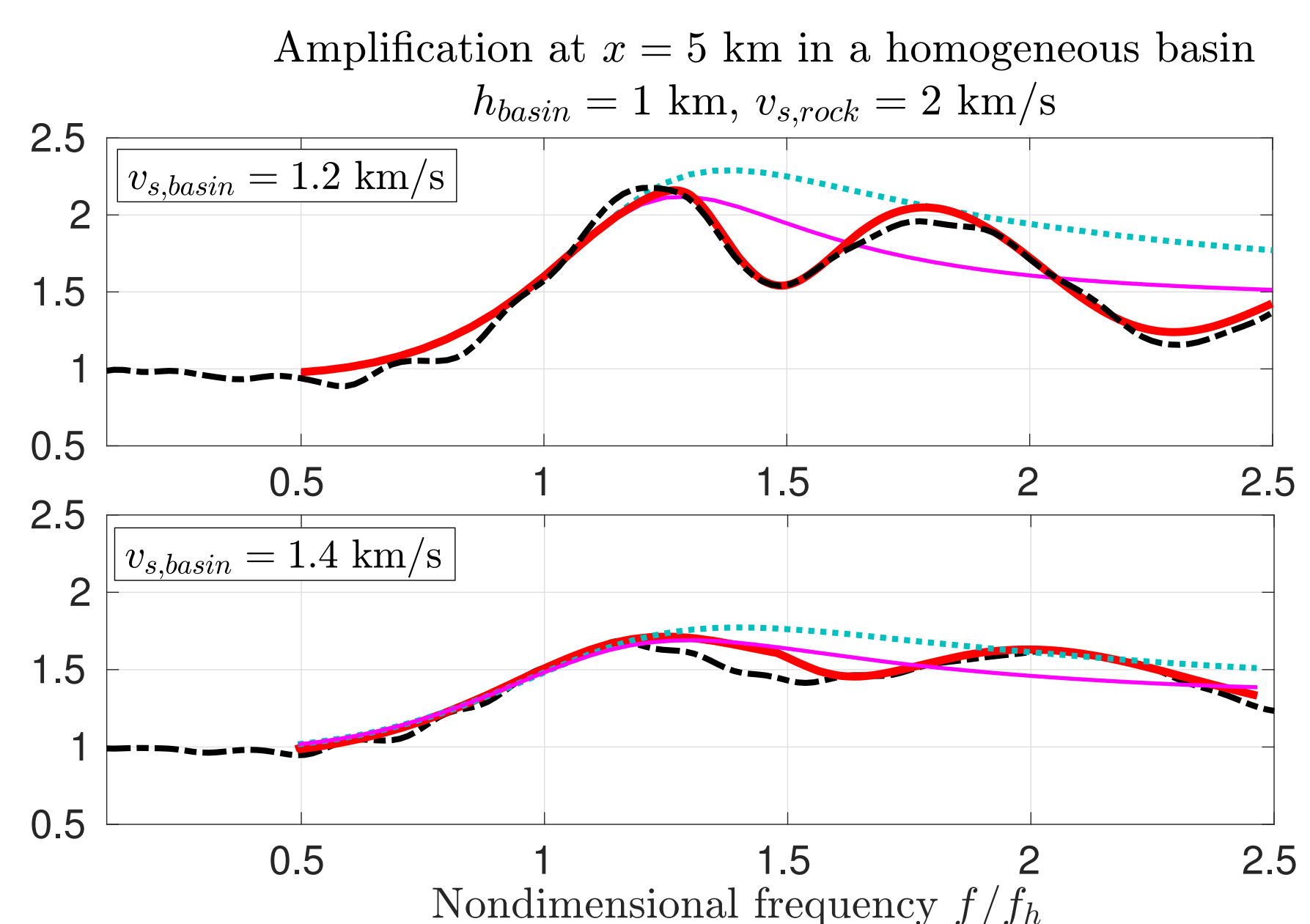
2 - SIMULATION SETUP



- Axisym. basin with length L_{basin} , depth h_{basin} and shear vel. $v_{s,basin}$
- Relationships between v_p , v_s and ρ are extracted from (*Brocher, 2005*)
- We use a **nondimensional freq.** $f_h \approx \frac{f}{v_{s,basin}/3h_{basin}}$ (*Colombero, 2018*)

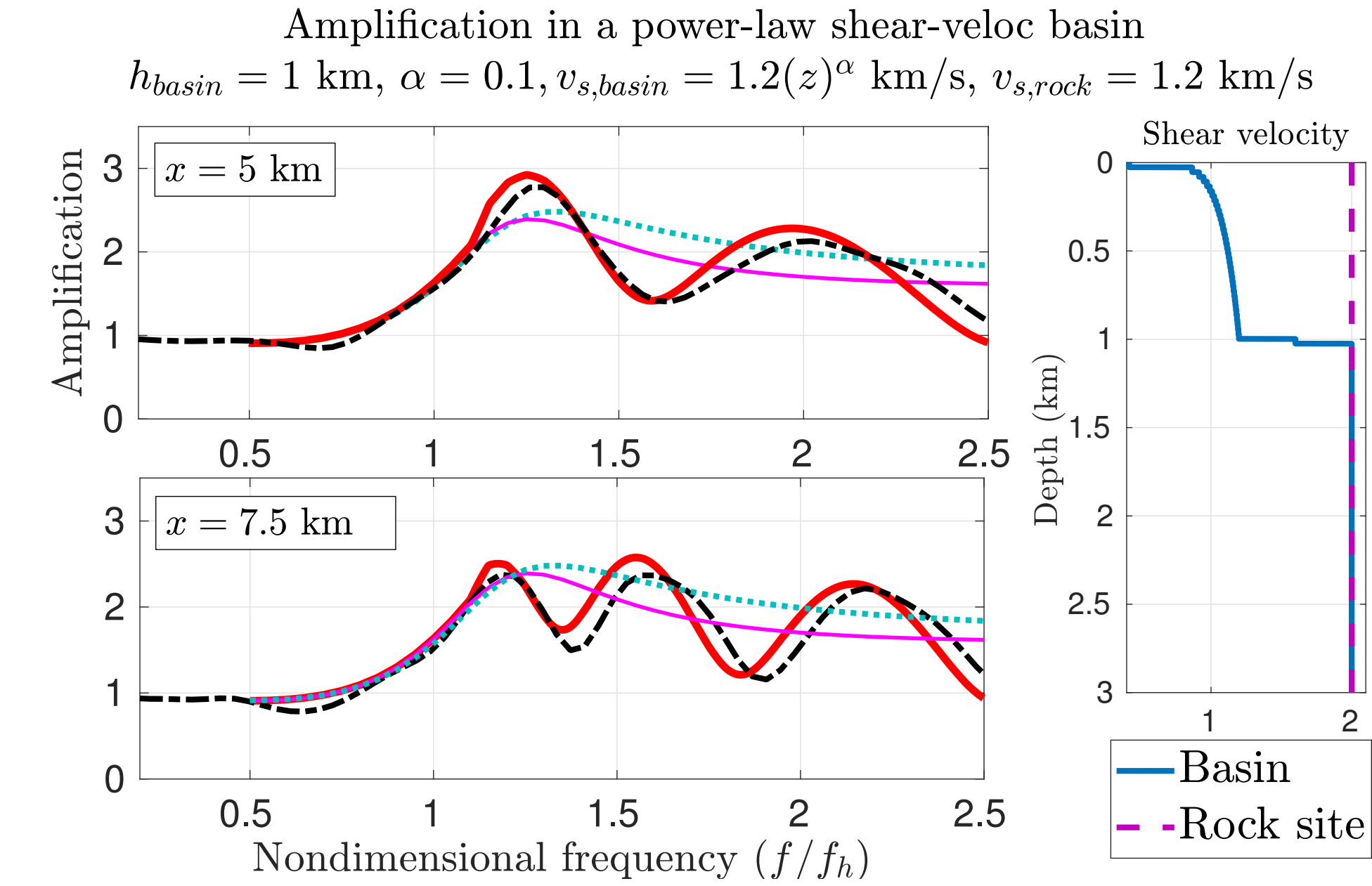
3 - TRANSMISSION AND CONVERSION IN SEMI-INFINITE BASINS ($L_{basin} \rightarrow \infty$)

Homogeneous semi-inf. basin w/ various shear velocities



- Fund.-mode transmission coefficient **captures well the average amplification spectrum**
- **Higher modes introduce strong oscillations** in the spectrum that can be reproduced by considering fund.-to-higher modes conversions

Heterogeneous semi-inf. basin at various distances



- Amplification spectra for **vertically heterogeneous basin velocity structures can be well approximated by transmission coefficients**
- Main amplification peak amplitude can be increased by higher modes

4 - 1D FUND.-MODE AMPLIFICATION

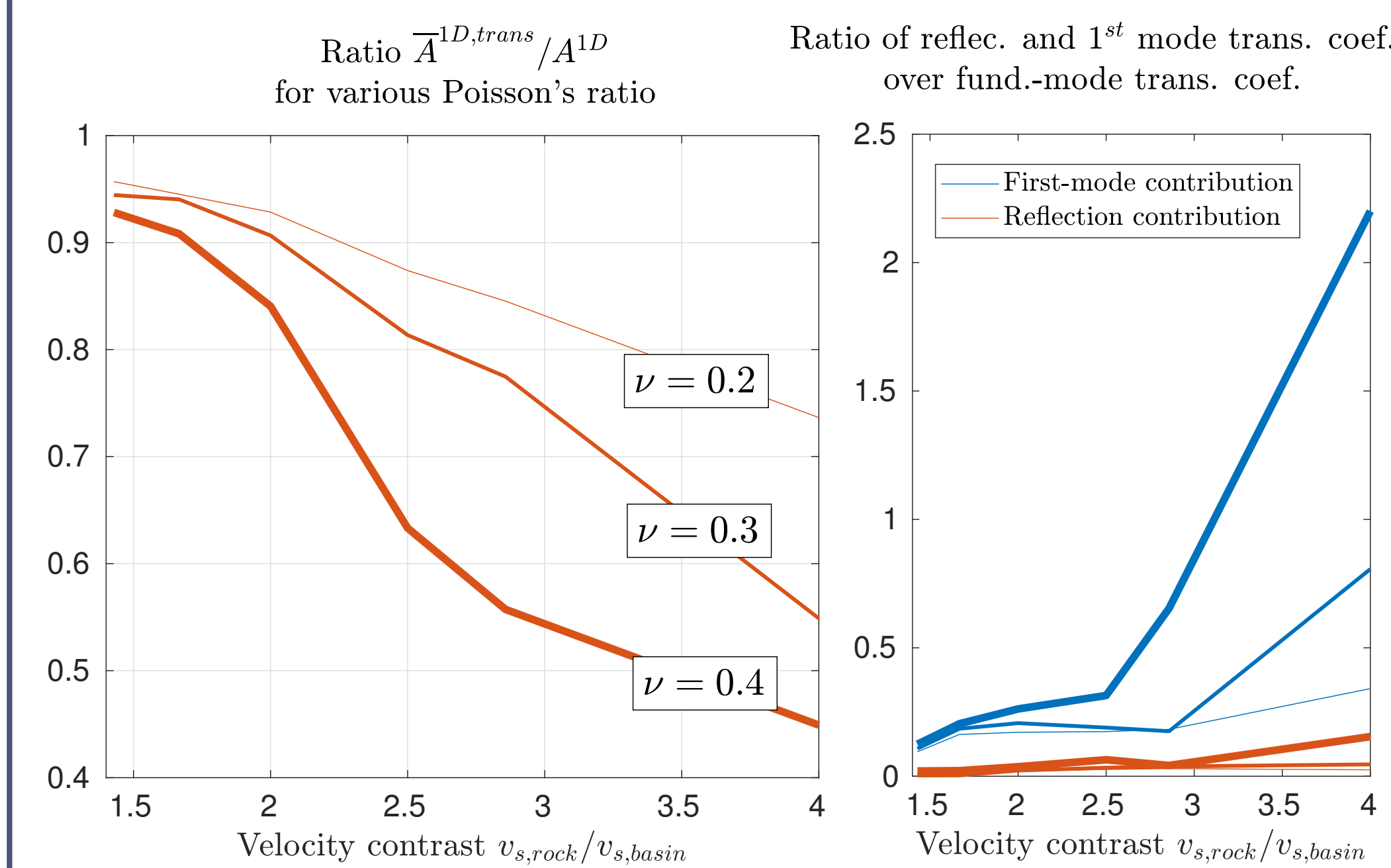


Figure: Ratio of max. amplification from the 1D theory A^{1D} and the mean transmission coefficients over the 10 first km from basin edge $\bar{A}^{1D,trans}$ against velocity contrasts for various Poisson's ratio

- Discrepancies between pure 1D theory and trans. coef. come from mode conversions and reflection at the basin boundary
- **Using a nondimensional freq. f_h we can assess the accuracy of the 1D theory** to predict the surface-wave amplification

5 - LATERAL RESONANCE

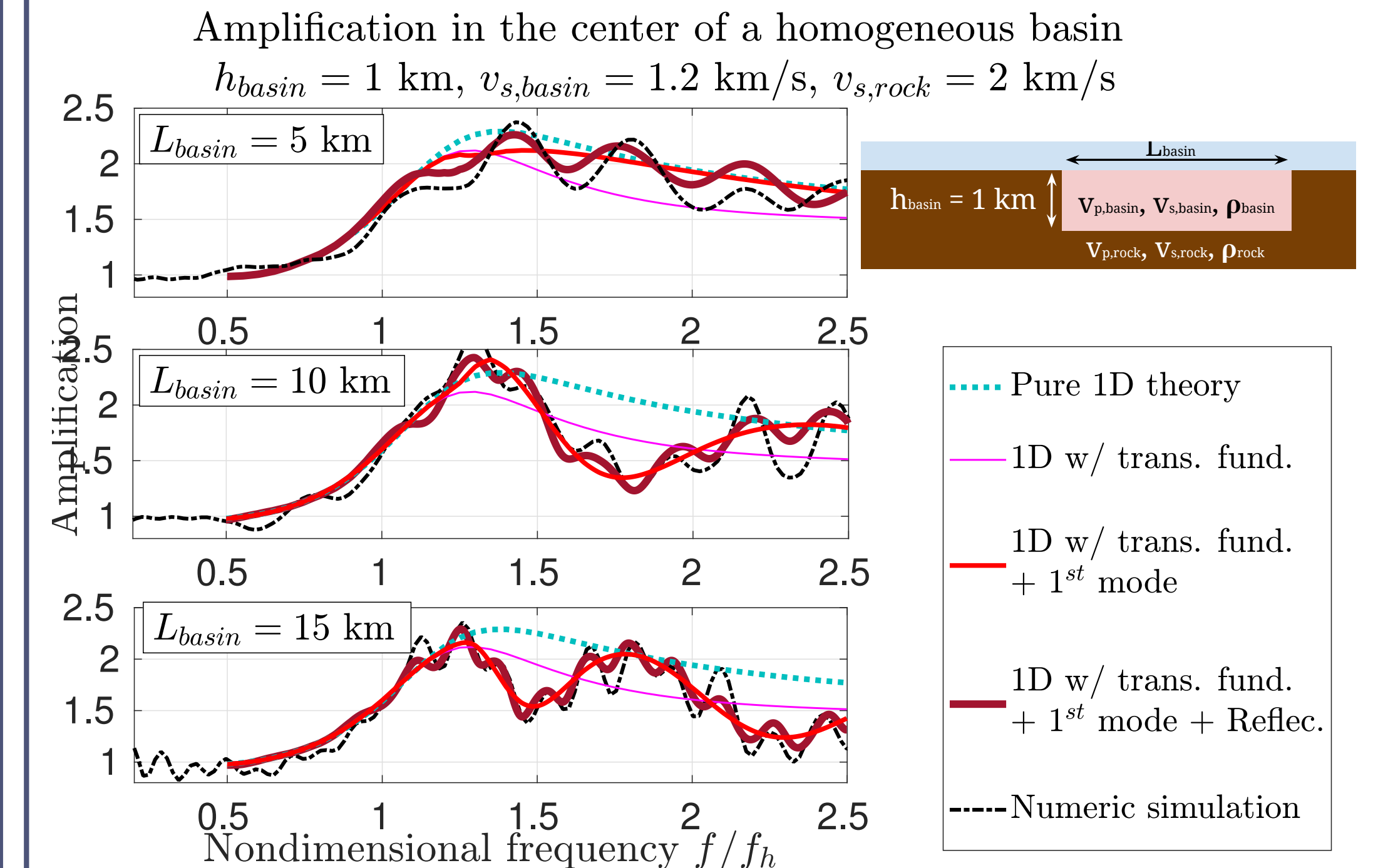


Figure: Top to bottom, Spectral amplification against nondimensional frequency f_h for basin length $L_{basin} = 5, 10, 15$ km and basin depth $h_{basin} = 1$ km.

- Lateral boundaries introduce **extra oscillations in the amplification spectrum** due to back and forth reflections within the basin
- Close to the basin edges and/or as the surface-wave wavelength range tends to the basin length, the **maximum amplitude can be significantly altered**

6 - LOS ANGELES BASIN AMPLIFICATION

- **2D velocity profile in the LA basin** extracted from SCEC Community Velocity Model (CVM-S4.26, *Lee (2014)*) w/ sharp velocity jump
- Basin edge location is chosen at the largest horizontal shear-velocity jump ($d \approx 66$ km)
- Transmission coefficients are computed from the 1D profiles beneath the stations

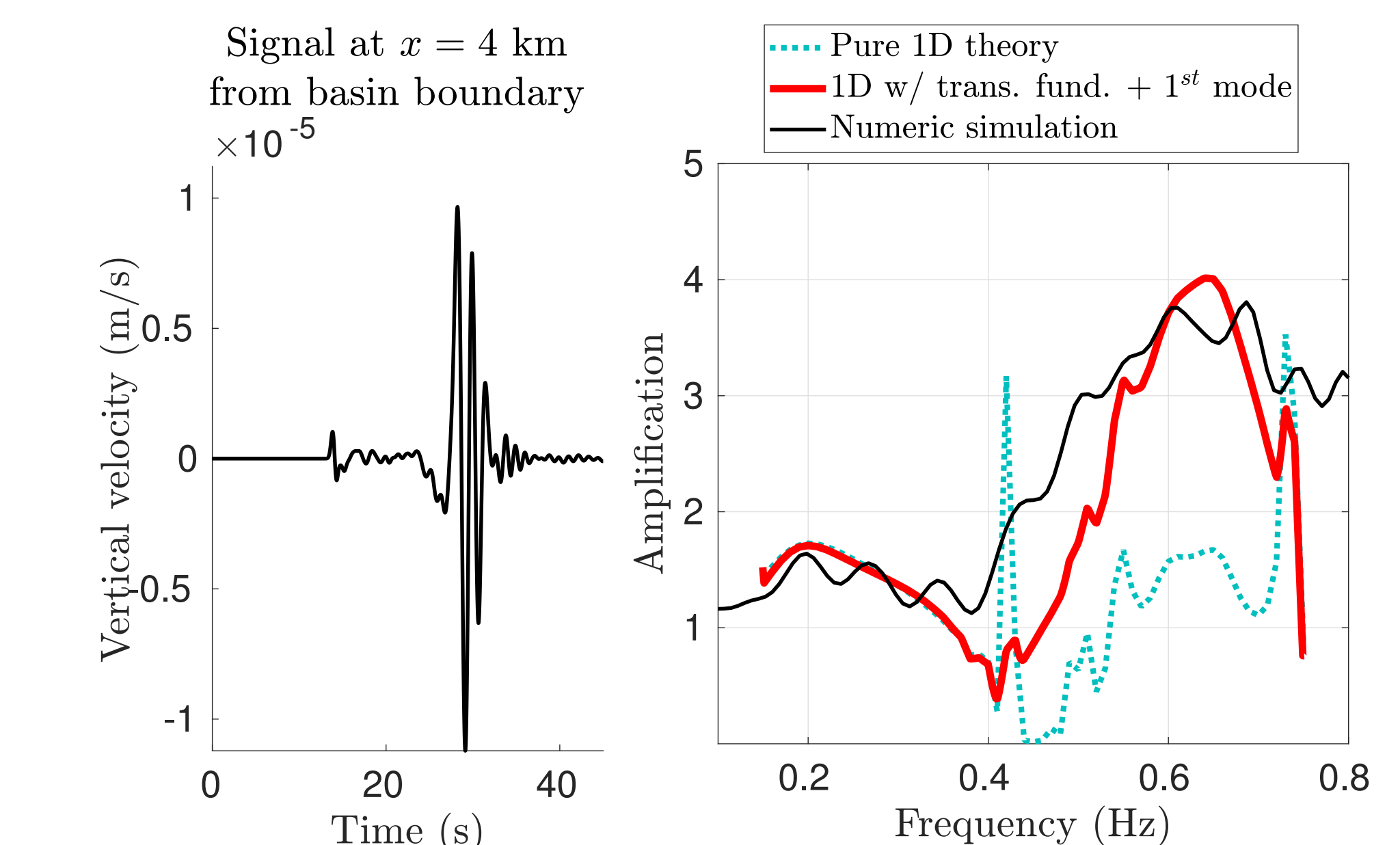
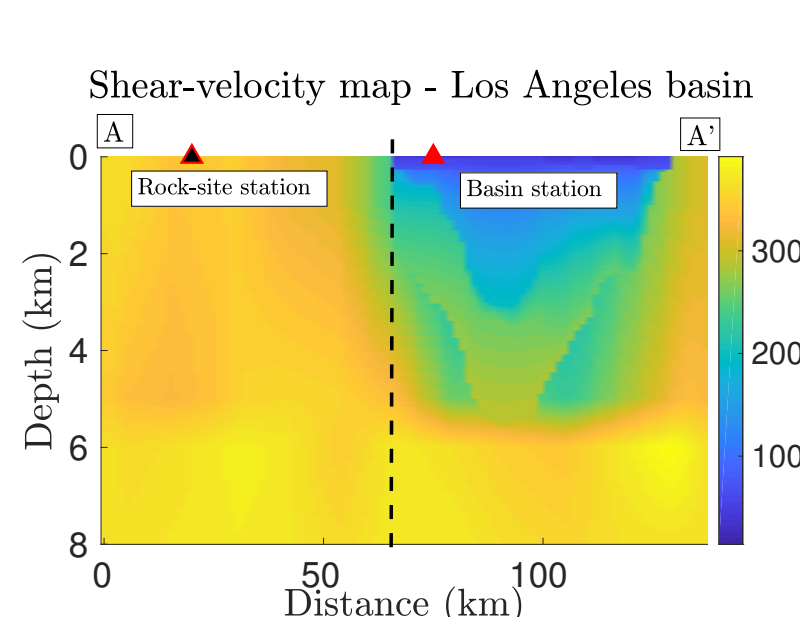
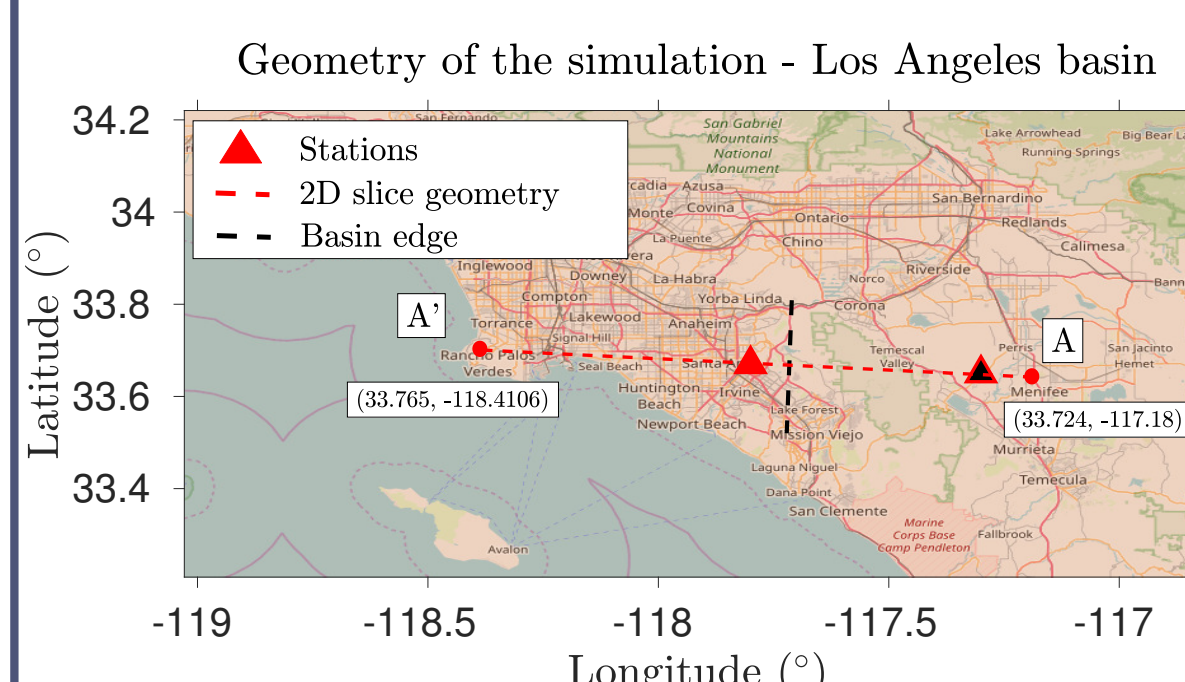


Figure: Left, Vertical velocity against time at the basin station. Right, amplification spectrum against frequency at the basin station

CONCLUSIONS AND FUTURE WORK

Conclusions

- 1D Theory
⇒ **good estimate of the amplification** (over-predict. < 30% of max. amp.) **for low velocity contrast** $\frac{v_{s,rock}}{v_{s,basin}} < 2.5$
- Approximate trans./reflec. coefficients:
⇒ can very well reproduce amplitude and variations of the amp. in axisym. basins
⇒ **good estimate of amplification in laterally heterogeneous axisym. sedimentary basins** w/ sharp vertical boundaries

Future work will include

- **More complex axisym. basin geometries**
- **Love-wave amplification** and Rayleigh-to-Love conversions
- **Full 3D basins structures** and subsequent path effects