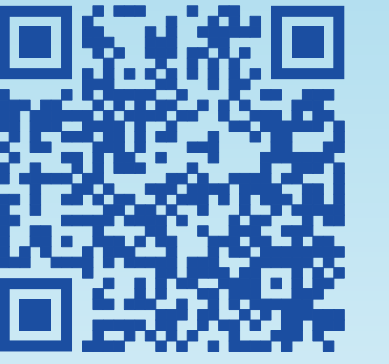


# Dynamics of the Global Energy Budget with a time dependant Climate Feedback Parameter



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The 0-dimensional linearised energy balance model (EBM) introduced by Budyko (1968) allows us to study the response of the climate system to a radiative forcing such that an increase of atmospheric CO<sub>2</sub>. This EBM reads:  $C \frac{dT_s}{dt} = N = F - \lambda T_s$ , where  $C$  is the ocean heat capacity,  $T_s$  is the global surface temperature,  $N$  is the Earth Energy Imbalance and  $\lambda$  is the constant climate feedback parameter.

However, recent studies show that a constant climate feedback parameter cannot represent accurately the long term dynamics of the climate response, notably due to the dependence of  $\lambda$  on the pattern of warming (Armour et al. 2013).

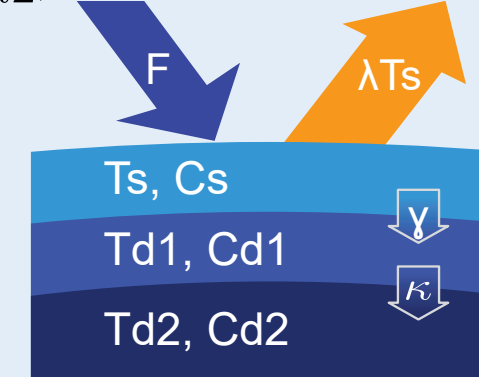
**Here, we introduce the time dependence of  $\lambda$  in a simple energy balance model and develop the consequences on climate sensitivity.**

## Theoretical framework

### 1. Energy Balance Model

$$\begin{aligned} ① \quad C_s \frac{dT_s}{dt} &= F - \lambda T_s - \gamma (T_s - T_{d1}) \\ ② \quad C_{d1} \frac{dT_{d1}}{dt} &= \gamma (T_s - T_{d1}) - \kappa (T_{d1} - T_{d2}) \\ ③ \quad C_{d2} \frac{dT_{d2}}{dt} &= \kappa (T_{d1} - T_{d2}) \end{aligned}$$

Three layers energy balance model, adapted from Geoffroy et al. (2013)



### 2. Hypotheses

- The climate system is a forced dynamical system. We assume the existence of steady states variables:  $T_{s0}$ ,  $F_0$  and  $\lambda_0$ .
- When  $\delta F$  is applied, the system deviates from its steady state and tends to reach a new equilibrium.
- The perturbation theory allows us to study this new dynamical system. **We hypothesise that our study is in the scope of the perturbation theory.**

### 3. Applying perturbation theory

We apply perturbation theory to the surface equation to derive the perturbed anomaly system.

**With a constant  $\lambda_0$**

$$C_s \frac{d}{dt} (\delta T_s) = \delta F - \lambda_0 \delta T_s - \gamma (\delta T_s - \delta T_{d1})$$

**With forced variation of  $\lambda$**

$$C_s \frac{d}{dt} (\delta T_s) = \delta F - \lambda_0 \delta T_s - \delta \lambda(t) T_{s0} - \gamma (\delta T_s - \delta T_{d1})$$

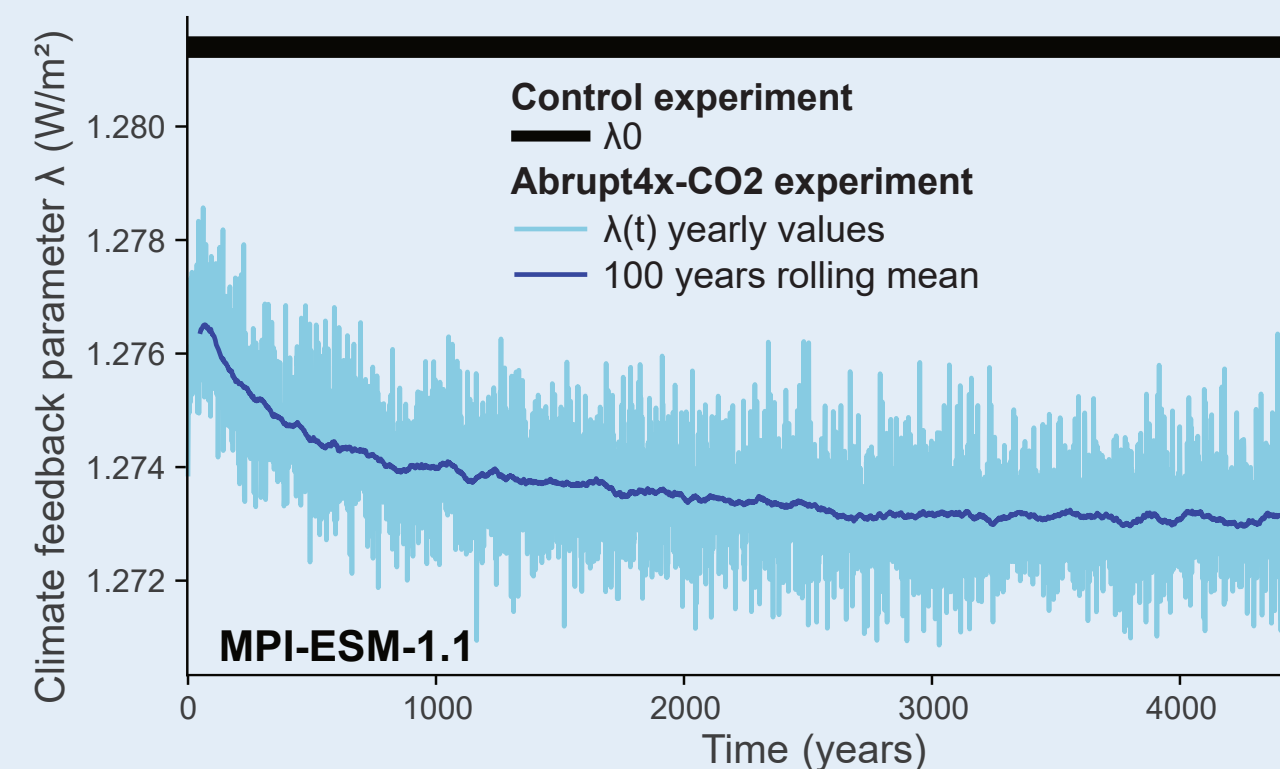
**Assuming a variable  $\lambda$  leads to the emergence of a new term in the anomaly energy budget**

## Climate Feedback Parameter time series

With  $N = F - \lambda T_s$ , we can write

$$\lambda(t) = \frac{F_0 + \delta F - N(t)}{T_{s0} + \delta T_s(t)}$$

Which gives a times series of the climate feedback parameter.

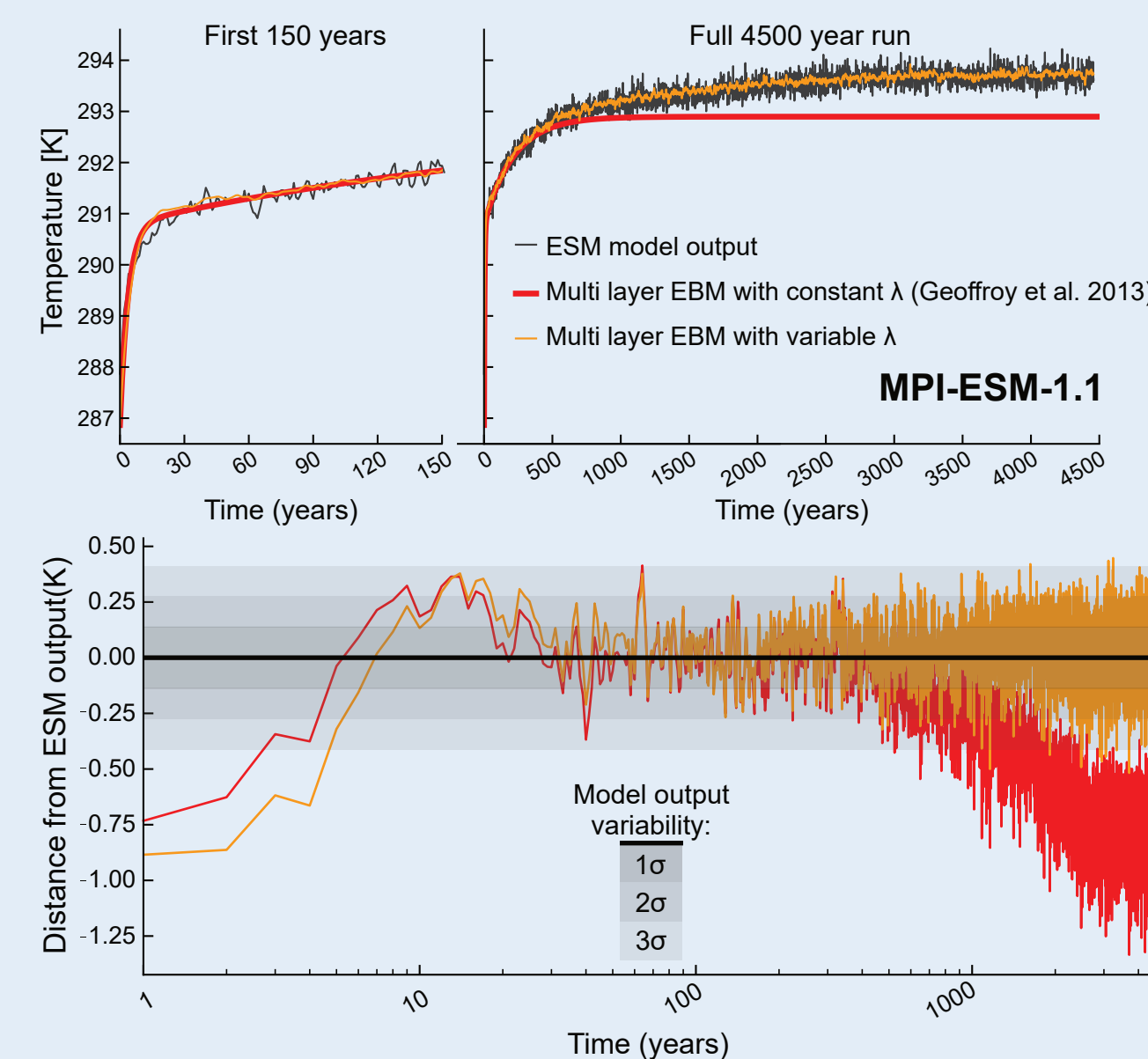


**We verify  $\delta \lambda \ll \lambda_0$**

which validates the perturbation theory hypothesis.

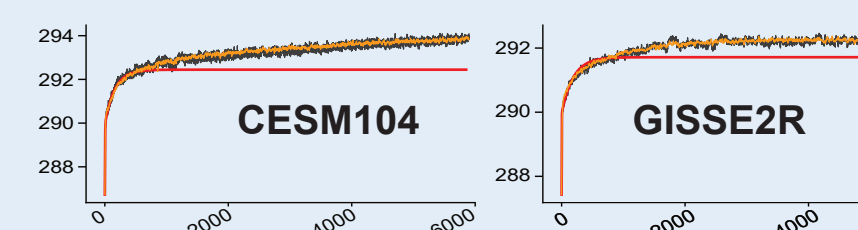
With such a climate feedback parameter, we expect to reproduce the dynamics of the global surface temperature in the model with the numerical integration of the system.

## Numerical integration



We reproduce the dynamics of the global average surface temperature in the MPI-ESM1.1 abrupt4x-CO<sub>2</sub> run from the longrunMIP experiment (Rugenstein et al. 2019) **at all time scales.**

How about other models?



## Consequences on climate sensitivity

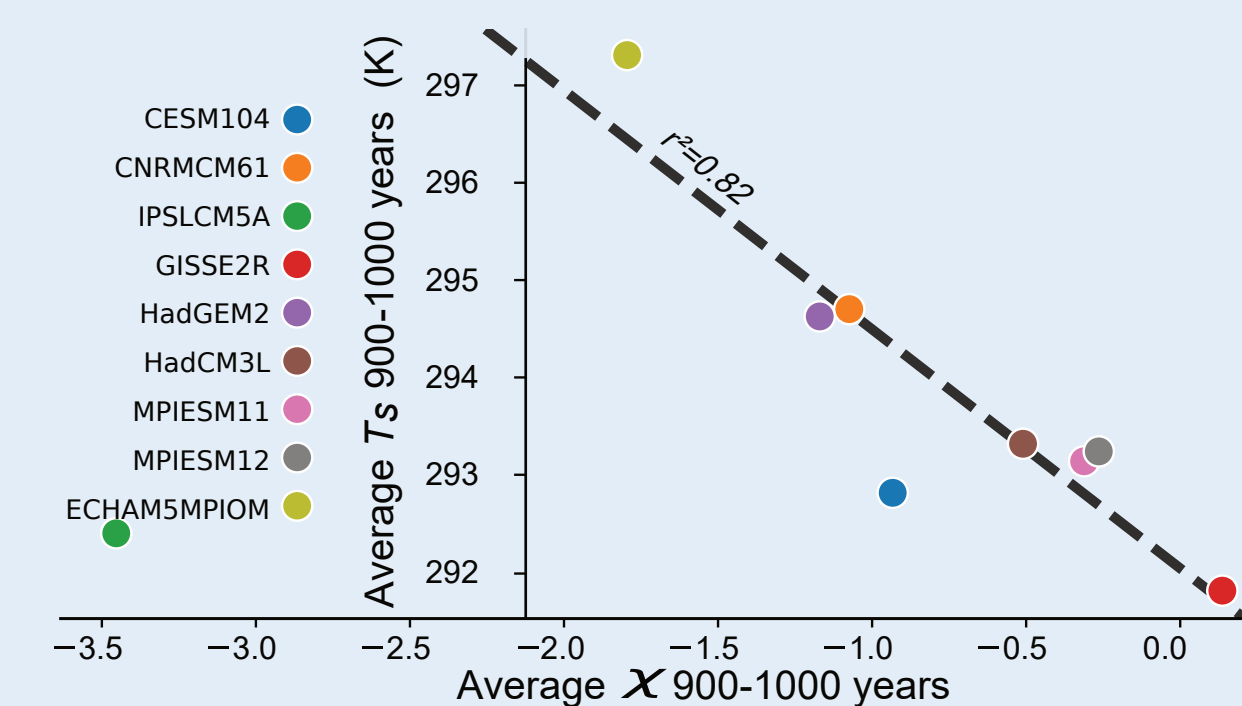
Getting  $\lambda$  from  $N = F - \lambda T_s$  and developing the new EBM to equilibrium leads to a new expression of the climate sensitivity:

$$S = S_0 (1 - \chi)$$

$$\chi = \frac{F_0 \delta \lambda}{\lambda_0 \delta F} \quad S_0 = \frac{\delta F}{\lambda_0}$$

Where  $\chi$  is the **climate susceptibility** to forcing

- Explicit dependance on the initial climate state
- Explicit dependence on  $\lambda$  variations



**The intermodel spread in climate sensitivity in the LongRunMIP experiment is due to different variations of  $\lambda$  among models.**

## Conclusions

- A simple theory is developed to account for the time dependency of  $\lambda$  in the global energy budget.
- The resulting differential equation accurately reproduces the response of the climate under abrupt changes in CO<sub>2</sub> concentrations **at all time scales** as simulated in a multitemillenia earth system model.
- Analysis of the asymptotic form of the differential equation yields a new expression of the climate sensitivity which explicitly depends on the temporal variations of the climate feedback parameter.
- We find that **the spread in climate sensitivity among climate models of the LongRunMIP experiment is essentially due to different temporal changes in  $\lambda$**  (and thus different pattern effect) among models.

## References

- Budyko (1969). The effect of solar radiation variations on the climate of the Earth. *Tellus*, 21(5), 611–619.
- Armour et al. (2013). Time-varying climate sensitivity from regional feedbacks. *Journal of Climate*, 26(13), 4518–4534. <https://doi.org/10.1175/JCLI-D-12-00544.1>
- Rugenstein et al. (2019). LongRunMIP: Motivation and design for a large collection of millennial-length AOGCM simulations. *Bulletin of the American Meteorological Society*, 100(12), 2551–2570. <https://doi.org/10.1175/BAMS-D-19-0068.1>