

Dynamics of the Global Energy Budget with a time dependant Climate Feedback Parameter



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The 0-dimensional linearised energy balance model (EBM) introduced by Budyko (1968) allows us to study the response of the climate system to a radiative forcing such that an increase of atmospheric CO₂. This EBM reads: $C_s \frac{dT_s}{dt} = N = F - \lambda T_s$, where C is the ocean heat capacity, T_s is the global surface temperature, N is the Earth Energy Imbalance and λ is the constant climate feedback parameter.

However, recent studies show that a constant climate feedback parameter cannot represent accurately the long term dynamics of the climate response, notably due to the dependence of λ on the pattern of warming (Armour et al. 2013).

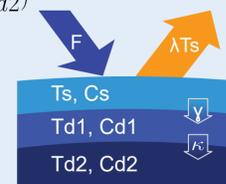
Here, we introduce the time dependence of λ in a simple energy balance model and develop the consequences on climate sensitivity.

1. Energy Balance Model

$$1 \quad C_s \frac{dT_s}{dt} = F - \lambda T_s - \gamma (T_s - T_{d1})$$

$$2 \quad C_{d1} \frac{dT_{d1}}{dt} = \gamma (T_s - T_{d1}) - \kappa (T_{d1} - T_{d2})$$

$$3 \quad C_{d2} \frac{dT_{d2}}{dt} = \kappa (T_{d1} - T_{d2})$$



Three layers energy balance model, adapted from Geoffroy et al. (2013)

Theoretical framework

2. Hypotheses

- The climate system is a forced dynamical system. We assume the existence of steady states variables: T_{s0}, F₀ and λ₀.
- When δF is applied, the system deviates from its steady state and tends to reach a new equilibrium.
- The perturbation theory allows us to study this new dynamical system system. We hypothesise that our study is in the scope of the perturbation theory.

3. Applying perturbation theory

We apply perturbation theory to the surface equation to derive the perturbed anomaly system.

With a constant λ₀

$$C_s \frac{d}{dt} (\delta T_s) = \delta F - \lambda_0 \delta T_s - \gamma (\delta T_s - \delta T_{d1})$$

With forced variation of λ

$$C_s \frac{d}{dt} (\delta T_s) = \delta F - \lambda_0 \delta T_s - \delta \lambda(t) T_{s0} - \gamma (\delta T_s - \delta T_{d1})$$

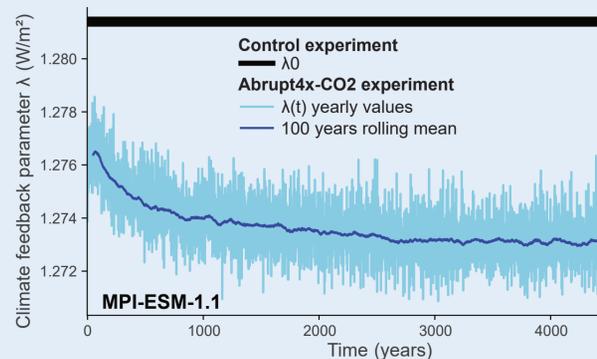
Assuming a variable λ leads to the emergence of a new term in the anomaly energy budget

Climate Feedback Parameter time series

With $N = F - \lambda T_s$, we can write

$$\lambda(t) = \frac{F_0 + \delta F - N(t)}{T_{s0} + \delta T_s(t)}$$

Which gives a times series of the climate feedback parameter.

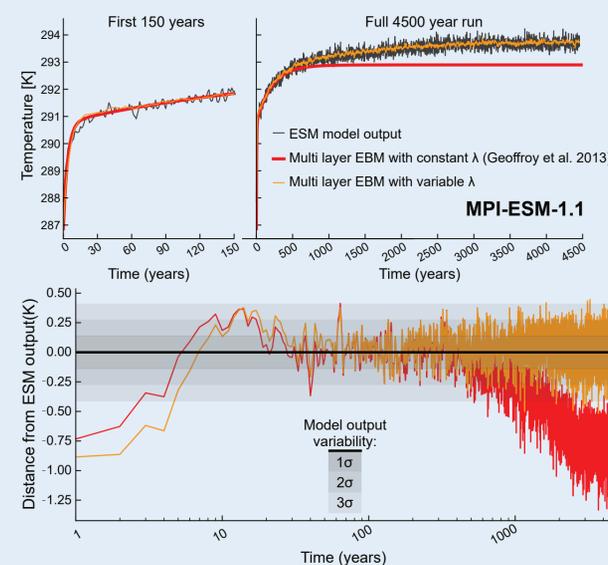


We verify $\delta \lambda \ll \lambda_0$

which validates the perturbation theory hypothesis.

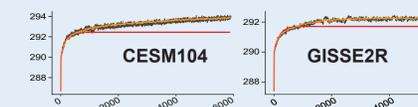
With such a climate feedback parameter, we expect to reproduce the dynamics of the global surface temperature in the model with the numerical integration of the system.

Numerical integration



We reproduce the dynamics of the global average surface temperature in the MPI-ESM1.1 abrupt4x-CO₂ run from the longrunMIP experiment (Rugenstein et al. 2019) at all time scales.

How about other models?



Consequences on climate sensitivity

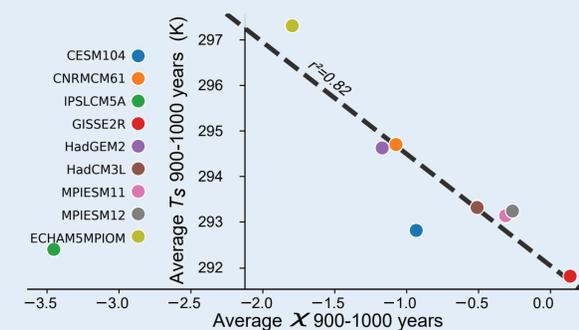
Getting λ from $N = F - \lambda T_s$ and developing the new EBM to equilibrium leads to a new expression of the climate sensitivity:

$$S = S_0 (1 - \chi)$$

$$\chi = \frac{F_0 \delta \lambda}{\lambda_0 \delta F} \quad S_0 = \frac{\delta F}{\lambda_0}$$

Where χ is the climate susceptibility to forcing

- Explicit dependance on the initial climate state
- Explicit dependance on λ variations



The intermodel spread in climate sensitivity in the LongRunMIP experiment is due to different variations of λ among models.

Conclusions

- A simple theory is developed to account for the time dependency of λ in the global energy budget.
- The resulting differential equation accurately reproduces the response of the climate under abrupt changes in CO₂ concentrations at all time scales as simulated in a multitemillenia earth system model.
- Analysis of the asymptotic form of the differential equation yields a new expression of the climate sensitivity which explicitly depends on the temporal variations of the climate feedback parameter.
- We find that the spread in climate sensitivity among climate models of the LongRunMIP experiment is essentially due to different temporal changes in λ (and thus different pattern effect) among models.

References

- Budyko (1969). The effect of solar radiation variations on the climate of the Earth. *Tellus*, 21(5), 611–619.
- Armour et al. (2013). Time-varying climate sensitivity from regional feedbacks. *Journal of Climate*, 26(13), 4518–4534. <https://doi.org/10.1175/JCLI-D-12-00544.1>
- Rugenstein et al. (2019). LongRunMIP: Motivation and design for a large collection of millennial-length AOGCM simulations. *Bulletin of the American Meteorological Society*, 100(12), 2551–2570. <https://doi.org/10.1175/BAMS-D-19-0068.1>