

Effect of Shape and Size on the Transport of Floating Particles on the Free-surface of a Meandering Stream

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Key Points:

- Velocity of particles floating on turbulent streams are weakly affected by their shape and size.
- Larger particles disperse faster on the free-surface due to their ability to filter out the small-scale turbulent fluctuations.
- Rods re-orient following the mean shear of the surface flow and rotate according to the integral scales of the free-surface turbulence.

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Abstract

Understanding how floating particles are transported by streaming waters is crucial in predicting the transport of plastic pollution, which is dramatically abundant in rivers, lakes, and oceans. Using particle tracking velocimetry, we investigate the motion of floating particles of different shape and size on the turbulent free-surface of a field-scale meandering stream. We consider two locations with different turbulence levels where the role of surface waves on the transport is deemed negligible. Millimetre-sized spheres are used as tracers to characterise the surface flow. These are compared with centimetre-sized discs and rods, much larger than the tracers, approximating typical-sized pieces of litter found in rivers. These larger particles exhibit similar velocities as the small tracers but filter out the extreme accelerations. Consequently, their motion is more time-correlated and their spreading rate is larger. This notion is confirmed by the mean-square displacement of single particles and mean-square separation of particle pairs, which grow faster in time compared to the tracers. The rotation of the rods, affected by a range of turbulent scales, reduces the correlation time scale of their translational motion, and leads to a slower dispersion compared to the discs, despite the rods' length being larger than the discs' diameter. Taken together, these results indicate that the motion of finite-size objects floating on non-wavy turbulent water is consistent with the behaviour of inertial particles in three-dimensional turbulence. These results can be valuable when constructing predictive models of floating plastics in natural waters.

Plain Language Summary

Plastic debris is a rising global issue severely affecting the state of our rivers, lakes and oceans. Understanding how pieces of litter, often floating, travel in streaming waters is crucial for predicting and ultimately limiting plastic pollution. The main goal of this research is to investigate how the shape and size of small floating objects may affect their journey on the surface of water. To this end, we use high-speed video recordings to track floating objects of different shape and size in an outdoor stream laboratory. The motion of centimetre-sized discs and rods, approximating typical pieces of plastics found in rivers, is directly compared to the motion of millimetre-sized spheres that follow the surface flow. We find that the larger discs and rods spread faster on the surface of water. Not only can these results be used to devise effective sequestration strategies, but they can be important to inform computer models that predict the abundance and fate of plastic litter in natural waters.

1 Introduction

Plastic debris is ubiquitous in our lakes, oceans, and coastal waters, posing serious threats to human health and the environment (Eriksen et al., 2013; van Sebille et al., 2015; Lebreton et al., 2018). Recent findings demonstrate that about 1,000 rivers account for 80 % of the global annual emissions of 0.8 to 2.7 million tons of plastics into the oceans per year, with small urban rivers among the most polluting (Meijer et al., 2021). Riverine ecosystems themselves are affected by such pollution (van Emmerik & Schwarz, 2020). Plastic objects enter such systems in a wide range of compositions, shapes, and sizes before degrading into so-called microplastics (typically defined as smaller than 5 mm). A large proportion of the plastic waste in the U.S. is comprised of polyethylene and polypropylene, which are less dense than fresh water (Jambeck et al., 2015). In general, it is estimated that more than half of all plastics produced are positively buoyant (Geyer et al., 2017). The question that motivates the present study is at which rate floating mesoplastics (particles in the size range of 5 to 50 mm) spread over the surface of turbulent streaming waters.

The transport of floating particles has been mainly investigated in terms of its dependence on surface waves. These impart a net drift velocity in the direction of wave propagation, known as Stokes drift (van den Bremer & Breivik, 2018). While this is typically much

66 smaller than the mean advective velocity, its magnitude increases with wave steepness and
67 it can play a role in the long-term dispersion (van Sebille et al., 2020). DiBenedetto et al.
68 (2018, 2019) showed that non-spherical particles in wavy waters tend to follow a preferred
69 orientation, which affects their settling velocity if those are negatively buoyant. For buoyant
70 particles, DiBenedetto (2020) found that waves result in non-uniform particle concentration.

71 Here we focus on regimes and scales where surface waves do not appreciably modify
72 the dispersion along the free-surface, which is instead dictated by turbulence. The nature of
73 turbulence on the vicinity of the surface is still debated. Pan and Banerjee (1995) identified
74 hallmark features such as upwelling and downwelling motions and long-lived vortices. Kumar
75 et al. (1998) measured a k^{-3} decay of the velocity spectra (k being the wavenumber),
76 consistent with the expectation for two-dimensional (2D) turbulence. On the other hand,
77 the field measurements of Chickadel et al. (2011) displayed a $k^{-5/3}$ behaviour typical of
78 three-dimensional (3D) turbulence. In the riverine environment, the shallowness of the flow
79 plays a significant role in determining the nature of the turbulence: in particular, in the
80 presence of strong lateral shear, the limited depth inhibits vortex stretching and may result
81 in vortex dynamics akin to 2D-turbulence, especially at low wavenumbers (Uijttewaal &
82 Booij, 2000). Most previous studies focused on free-surface turbulence have been concerned
83 with the topological features of the flow, often in relation to air-water gas fluxes (Shen
84 et al., 1999; Shen & Yue, 2001; McKenna & McGillis, 2004; Turney & Banerjee, 2013;
85 Herlina & Wissink, 2014), with only few studies concerned with the transport of particles
86 on it. Particularly, Cressman et al. (2004) and Lovechio et al. (2013) found that tracer
87 particles floating on the free-surface cluster into string-like structures with long lifetimes.
88 Characteristic features of shallow flows, such as transitional macro-vortices, have been found
89 to greatly affect the single-particle and particle-pair dispersion (Stocchino et al., 2011).

90 Several field studies have been concerned with natural free-surface flows, focusing on
91 the effectiveness of free-surface velocity measurements, e.g., for discharge estimation as well
92 as flow monitoring during flood events. The methods include acoustic Doppler velocimetry
93 (ADV) but also imaging techniques originally developed for laboratory flow studies, such
94 as particle image velocimetry (PIV) and particle tracking velocimetry (PTV) (Raffel et al.,
95 2018; Adrian & Westerweel, 2011). Free-surface PIV and PTV present technical challenges
96 even in laboratory studies, e.g., the choice of appropriate tracers and their successful imag-
97 ing in spite of surface reflections (Weitbrecht et al., 2002; Miozzi et al., 2010; Miozzi &
98 Romano, 2020; Gomit et al., 2022). The difficulties are exacerbated in field studies due to
99 uneven natural illumination and scarcity of detectable floating tracers. Nevertheless, these
100 techniques have gained favor in riverine flow investigations due to the richness of the data
101 they can provide (Jin & Liao, 2019; Tauro et al., 2016, 2019).

102 Here we investigate experimentally the motion of floating particles on the turbulent
103 free-surface of a meandering stream in an outdoor facility which offers laboratory-quality
104 measurements and control in a field-scale setting. The main goal of the study is to explore
105 the influence of the shape and size of floating particles along their trajectories when driven
106 by the multi-scale fluctuations of the free-surface flow. The focus is on size ranges relevant
107 to meso- and macroplastics (> 5 mm) which are highly relevant to but largely understudied
108 in river flows (van Emmerik & Schwarz, 2020). Applying time-resolved PTV to millimetre-
109 sized tracers, we obtain surface velocity fields at two different locations along the stream. We
110 then characterize the transport of centimetre-sized discs and rods and directly compare them
111 to the behaviour of the tracers. In particular, we examine the floating particles response
112 to the free-surface turbulent fluctuations which in turn affects their spreading rate. The
113 rotational dynamics of the rods are then considered to gain insight on their dispersion as
114 compared to the discs. As we will discuss, the sensitivity of the particle dispersion to small-
115 scale turbulence may have important consequences for modeling approaches based on flow
116 velocity data (which are necessarily coarse-grained in space and time).



Figure 1. (a) The OSL facility, with the locations of the two ROIs (Meander and Riffle) indicated by arrows. The tent, shown here at a downstream location, is deployed over the ROIs in the present experiments. (b) The traversing system holding the camera used for free-surface imaging, indicating the approximate location of the field of view in the Meander and the coordinate system.

117 2 Materials and Methods

118 2.1 Field-scale Stream Facility and Hydrodynamic Characterization

119 Measurements are performed in the Outdoor StreamLab (OSL), an outdoor field-scale
 120 stream facility at the Saint Anthony Falls Laboratory, University of Minnesota (Figure
 121 1a). Water is drawn from the Mississippi River, flows through a meandering channel and
 122 discharges back into the river. The flow rate is controlled via a valve at the inlet, and the
 123 incoming water flows into a headbox and over a weir before entering the channel. A Massa
 124 M300 ultrasonic distance probe and a sonar transducer are mounted on a programmable
 125 measurement carriage, performing 2D elevation scans of the water surface and channel bed.
 126 The monitoring of the water height at the weir allows real time calculation of the flow rate
 127 Q . Two regimes are considered, $Q = 32.1 \text{ L s}^{-1}$ and $Q = 53.7 \text{ L s}^{-1}$, for which transport
 128 of sediment is negligible and the bed is static. Measurements are acquired in two regions
 129 of interest (ROIs): one located at one of the meanders in the stream, and the other at the
 130 straight section upstream of a riffle. We will refer to these measurement locations as the
 131 Meander and the Riffle, respectively.

132 The bathymetries of both ROIs are shown in Figure 2a-b. The origin of the global
 133 coordinate system is chosen to be on the bank of the Meander, with x approximately in
 134 the streamwise direction, y pointing from the inner to the outer bank, and $z = 0 \text{ m}$ cor-
 135 responding to the water surface. At the Riffle we also define an additional coordinate system
 136 $x' - y'$, with x' approximately aligned with the local flow direction.

137 The sub-surface flow velocity in the Meander is also characterized by a Nortek Vectrino
 138 ADV probe traversed along the cross-section at $x = 1 \text{ m}$. The phase-space thresholding
 139 technique described in Parsheh et al. (2010) is used to remove occasional spurious velocity
 140 spikes due to air bubbles. Measurements are acquired at 100 Hz for 120 s, in 23 and 28
 141 locations along the cross-section for $Q = 32.1 \text{ L s}^{-1}$ and $Q = 53.7 \text{ L s}^{-1}$, respectively. Besides
 142 confirming the flow rates measured at the weir, the ADV measurements indicate significant
 143 turbulence intensity throughout the water depth, with root mean square (RMS) fluctuations
 144 exceeding 10% of the bulk velocity.

145 The hydrodynamic conditions at both ROIs are summarized in Table 1. The Reynolds
 146 number $\text{Re} = DU_{bulk}/\nu$ and the Froude number $\text{Fr} = U_{bulk}/\sqrt{gD}$ are based on the water
 147 depth D and the bulk flow velocity U_{bulk} , both spatially averaged over the respective ROIs.
 148 $U_{bulk} = Q/A$ is calculated from the cross-sectional area A inferred from the bathymetry, g is
 149 the gravitational acceleration, and ν is the kinematic viscosity. Despite the Meander being

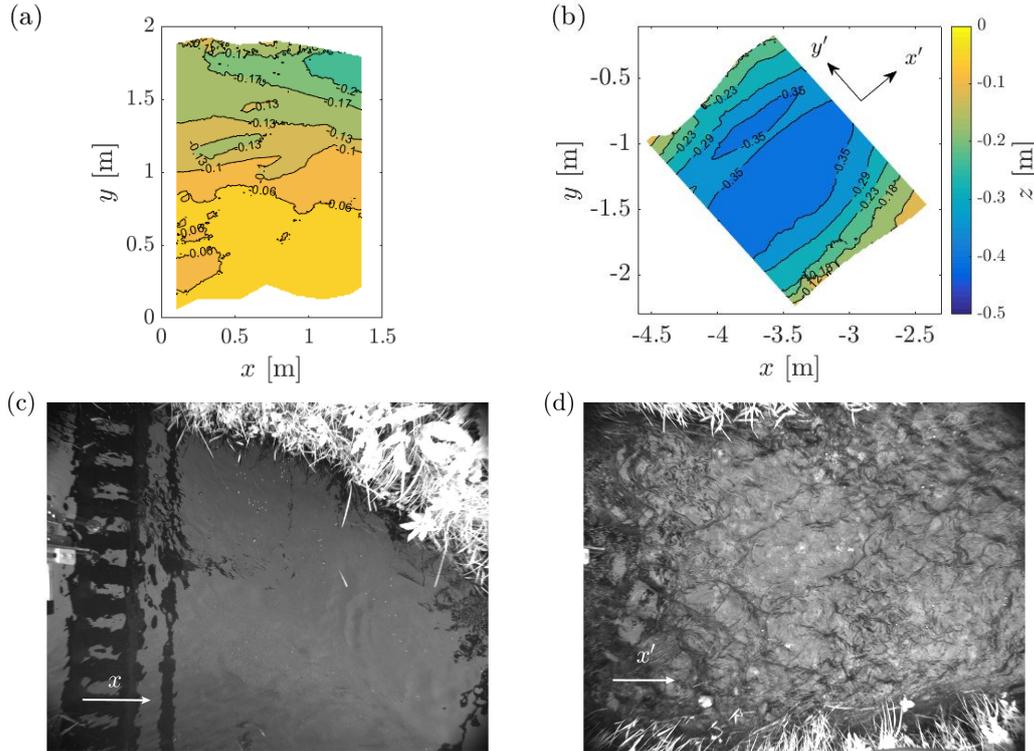


Figure 2. Bathymetry of the Meander (a) and Riffle (b) for $Q = 53.7 \text{ L s}^{-1}$. Instantaneous photographs of the free-surface at the Meander (c) and Riffle (d) for the same flow rate, indicating the streamwise direction x and x' , respectively.

150 shallower and associated to a larger Fr , the Riffle displays a wavier surface (Figure 2c-d)
 151 which is attributed to the turbulence induced by the rocky bed in this region (Brocchini
 152 & Peregrine, 2001). In both ROIs $Fr \ll 1$, and indeed the ultrasonic probe data indicates
 153 limited deformation of the free-surface: the RMS fluctuations of the water surface level
 154 are approximately 1 mm and 2 mm in the Meander and Riffle, respectively. Instantaneous
 155 images (acquired as described below) indicate wavelengths of 3 to 6 cm in the Meander and
 156 4 to 8 cm in the Riffle. Thus, the estimated Stokes drift velocity is at least one order of
 157 magnitude smaller than the RMS velocity fluctuations measured by PTV, confirming that
 158 the waves do not significantly alter the transport of the floating particles in the considered
 159 flow.

160 2.2 Floating Particles

161 Three types of floating particles are used in the present experiments. White polypropy-
 162 lene bean-bag filler pellets, approximately spherical with a 5 mm diameter, are used to
 163 characterize the surface flow velocity. These are sufficiently large to be accurately detected
 164 by imaging and can be recaptured downstream of the ROIs. Systematic studies assessing
 165 the ability of floating particles to faithfully follow the surface turbulent fluctuations are
 166 scarce. In Nikora et al. (2007), 3 mm floating particles were deemed suitable as tracers for
 167 free-surface turbulence in a laboratory flume. The largest eddies in homogeneous turbu-
 168 lence have diameters several times larger than the Kolmogorov scale η (Jiménez et al., 1993;
 169 Carter et al., 2016). Thus, neutrally buoyant particles up to 5η are typically considered
 170 faithful tracers (Fiabane et al., 2012), and even for particle sizes of $\sim 10\eta$ the response time

Table 1. Main hydrodynamic parameters of the Meander (a) and Riffle (b) for both volumetric flow rates Q : mean depth of the channel D , mean width of the channel W , mean cross-sectional area A , bulk fluid velocity U_{bulk} , Reynolds number Re , and Froude number Fr .

(a) Meander	D [m]	W [m]	A [m ²]	U_{bulk} [m s ⁻¹]	Re	Fr
$Q = 32.1 \text{ L s}^{-1}$	0.08	1.72	0.143	0.225	18,480	0.25
$Q = 53.7 \text{ L s}^{-1}$	0.10	1.72	0.177	0.303	30,910	0.30
(b) Riffle	D [m]	W [m]	A [m ²]	U_{bulk} [m s ⁻¹]	Re	Fr
$Q = 32.1 \text{ L s}^{-1}$	0.29	1.68	0.492	0.065	18,977	0.04
$Q = 53.7 \text{ L s}^{-1}$	0.31	1.68	0.525	0.102	31,750	0.06

Table 2. Summary of the main properties of the floating particles.

Particle Type	Spheres (Tracers)	Discs	Rods
Material	Polypropylene	Wood	Wood
Density [g cm ⁻³]	0.9	0.7	0.7
Major Axis [mm]	5.0	38.1	63.5
Minor Axis [mm]	5.0	38.1	1.8
Thickness [mm]	—	3.2	—

171 is about 1.5 times that of fluid tracers (Qureshi et al., 2007; Volk et al., 2011). As it will be
 172 shown, the estimated Kolmogorov length scale of the free-surface turbulence in the Meander
 173 is roughly 10 times smaller than the spherical pellets, which are 35 to 50 times smaller than
 174 the integral scale of the turbulence. Therefore, while these particles may not respond faith-
 175 fully to the smallest-scale fluctuations, they capture most of the turbulent kinetic energy
 176 and will be regarded as tracers.

177 To explore the effect of shape and size on particle transport, larger discs and rods are
 178 utilized. The discs consist of wooden craft circles and the rods are wooden toothpicks,
 179 both spray-painted white to increase their visibility and to reduce their absorption of water.
 180 The different particle properties are summarized in Table 2. An estimate of the particle
 181 Stokes number St (the ratio of the particle response time to a relevant flow time scale)
 182 is not attempted. Directly measuring the response time to the fluid velocity fluctuations
 183 would require the accurate measurement of the particle acceleration autocorrelation, which
 184 is beyond the capability of the present imaging system. Theoretical expressions require
 185 assumptions on the drag exerted on the particles; this depends on their level of submergence
 186 (Beron-Vera et al., 2019), which cannot be accurately measured here.

187 2.3 Particle Imaging and Tracking

188 A 1-megapixel CMOS camera (Allied Vision Mako U-130B) with a 3 mm wide-angle
 189 lens is mounted on a cantilever arm attached to a traversing system composed of aluminum
 190 beams (Figure 1b). The camera is suspended 1.5 m above the water surface, imaging a
 191 2.2 m \times 1.7 m field of view (FOV). To minimize reflections on the water surface, a large tent
 192 is set up to enclose the camera and the FOV, blocking direct sunlight that would cause
 193 reflections and any wind that may affect the free-surface.

194 To recapture the floating particles, a nylon seine net is suspended from a PVC pipe
 195 frame across the channel 3 m downstream of the Meander. Particles are manually seeded
 196 onto the water surface by gently shaking a wide bin spanning the channel width upstream of

197 the ROIs. The camera records at a frame rate of 30 to 50 Hz depending on the ROI and flow
 198 rate, keeping the inter-frame particle displacement to about 6 pixels. The measurements
 199 are performed over four runs for each case, to prevent the net from filling with particles
 200 and obstructing the water flow. This results in about 15,000 to 20,000 images, yielding
 201 approximately 16,000 particle trajectories for the tracers and 1,000 trajectories for the discs
 202 and rods, per each case. We verify that each run yields the same quantitative results for each
 203 case, thus statistical uncertainty due to finite sample size does not affect the conclusions.

204 The wide-angle camera lens introduces some image distortion. To correct it, a $0.9\text{ m} \times$
 205 1.2 m checkerboard pattern is imaged at the same distance as the water surface, and the
 206 appropriate de-warping transform is determined using MATLAB. Despite the tent blocking
 207 direct sunlight, some glare off the water surface from the diffused ambient light is still
 208 present. This time-dependent background noise is removed using the proper orthogonal
 209 decomposition (POD)-based method by Mendez et al. (2017), which isolates the modes
 210 mostly contributing to the intensity variance of the images. We subtract the first two
 211 modes, which successfully removes most of the glare while preserving the particle images.

212 Particles are identified as continuous groups of pixels exceeding an intensity threshold.
 213 Considering the probability distribution function (PDF) of the areas of these groups of pixels,
 214 a rejection criterion is set at ± 2 standard deviations from the expected value (based on the
 215 pixel/mm ratio). Particle centroids are tracked using a custom-written nearest-neighbour
 216 PTV algorithm (Baker & Coletti, 2019, 2021), and their velocities and accelerations are
 217 obtained from the trajectories by convolution with the first and second derivative of a
 218 Gaussian kernel in the time domain, respectively. A temporal kernel of 16 frames is chosen
 219 as the shortest interval beyond which the acceleration variance decays exponentially. This
 220 approach has been used in several previous laboratory and field studies (Voth et al., 2002;
 221 Nemes et al., 2017; Li et al., 2022; Berk & Coletti, 2021; Baker & Coletti, 2021, 2022).
 222 We also characterize the rods' orientation and rotation rate along their trajectory. The
 223 orientation is defined by the unit vector $\hat{\mathbf{p}}$ aligned with the rod axis, obtained from an
 224 ellipse best-fit to the object image. The angular velocity ω is obtained convolving $\hat{\mathbf{p}}(t)$ with
 225 the first derivative of a Gaussian kernel, analogously to the particle velocity and using the
 226 same temporal kernel.

227 3 Results and discussion

228 3.1 Free-Surface Flow: Eulerian Fields

229 We first consider the Eulerian fields of the mean velocity \tilde{U} , and the RMS fluctuations
 230 $\tilde{\sigma}_u$. Here $U = \sqrt{U_x^2 + U_y^2}$ is the norm of the free-surface velocity vector \mathbf{U} obtained by 2D
 231 imaging. The Eulerian data is obtained by binning the PTV trajectories in fixed interroga-
 232 tion windows of $5\text{ cm} \times 5\text{ cm}$ and indicated with a tilde. This allows for a temporal averaging
 233 of at least 25 instantaneous vectors in each window. The results for both measurement loca-
 234 tions and flow rates are shown in Figure 3. The Meander displays a remarkably homogeneous
 235 surface flow. In particular, we define it in a $1.25 \times 1\text{ m}$ sub-region (highlighted in the figure)
 236 where \tilde{U} and $\tilde{\sigma}_u$ remain within $\pm 2.5\%$ and 9.3% of their respective spatial mean and the
 237 streamlines are relatively straight. This allows us to investigate unbiased single-point and
 238 two-point flow statistics, characterizing the spatio-temporal flow scales using the framework
 239 of homogeneous turbulence (presented in the next section). Lagrangian particle transport
 240 is investigated in the same sub-region.

241 The mean flow in the Riffle, on the other hand, resembles a jet-like flow with two shear
 242 layers associated to high velocity fluctuations and flanked by recirculation zones. Because
 243 of the large spatial inhomogeneity, the scales of the free-surface turbulence in the Riffle are
 244 not carried out using two-point statistics, as this would require spatial averaging and the
 245 evaluation of velocity fluctuations around a global mean. The Lagrangian particle transport
 246 in this ROI is quantified in a $1.1 \times 1\text{ m}$ sub-region. For both ROIs, the choice of the sub-

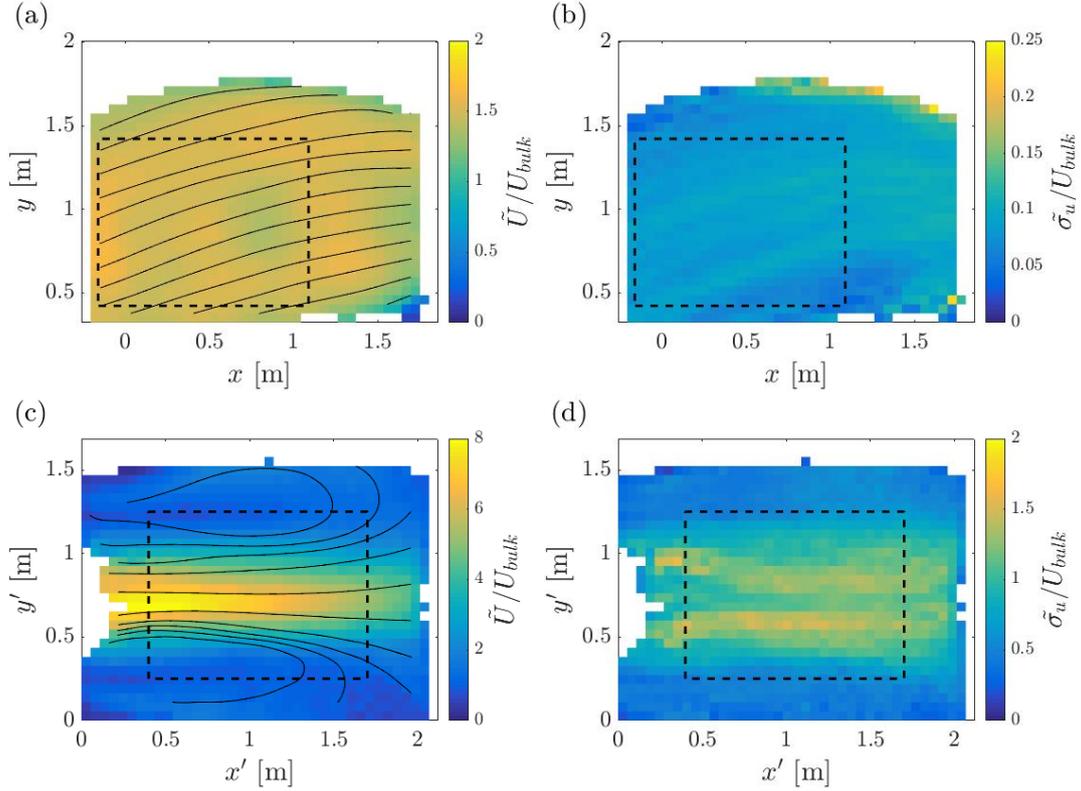


Figure 3. (a-c) Eulerian mean velocity and (b-d) RMS velocity fluctuation fields of the tracers for $Q = 32.1 \text{ L s}^{-1}$, normalized by the bulk velocity; Meander (a-b) and Riffle (c-d). The black lines indicate streamlines and the dashed boxes indicate the sub-regions where Lagrangian quantities are evaluated.

247 region avoids that the statistics be strongly influenced by the proximity of the banks and
 248 reduces potential bias from short trajectories as the particles exit the FOV.

249 3.2 Free-surface Turbulence in the Meander

250 For a statistical analysis of the turbulence structure, we are particularly interested
 251 in the instantaneous velocity fluctuations. For this purpose, the velocity fluctuation $\mathbf{u}(t)$
 252 is calculated by subtracting from the measured velocity \mathbf{U} the global mean $\langle \mathbf{U} \rangle$, $\mathbf{U}(t) =$
 253 $\langle \mathbf{U} \rangle + \mathbf{u}(t)$. Here t indicates time and angled brackets indicate ensemble-averaging over all
 254 realizations and all spatial locations within the homogeneous sub-region. Figure 4 displays
 255 the PDF of the streamwise (u_x) and spanwise (u_y) velocity fluctuations for $Q = 53.7 \text{ L s}^{-1}$,
 256 as well as the PDF of the spanwise accelerations (a_y) for both flow rates; all quantities
 257 are normalized by their respective standard deviation. Both components of the velocity
 258 fluctuations appear normally distributed. On the other hand, the acceleration PDFs possess
 259 long exponential tails, indicating strong intermittency, i.e., relatively large probability of
 260 extreme events, especially for the higher Reynolds number. This behavior of Lagrangian
 261 accelerations has been well documented in 3D turbulence (Voth et al., 2002; Mordant et
 262 al., 2004; Toschi & Bodenschatz, 2009). While the kurtosis of the velocity fluctuations
 263 approximately equals the Gaussian value of 3, the acceleration kurtosis is 8.1 and 15.9 for
 264 $Q = 32.1 \text{ L s}^{-1}$ and $Q = 53.7 \text{ L s}^{-1}$, respectively. These levels of intermittency are typical
 265 of fully developed 3D turbulence (Voth et al., 2002; Ishihara et al., 2007).

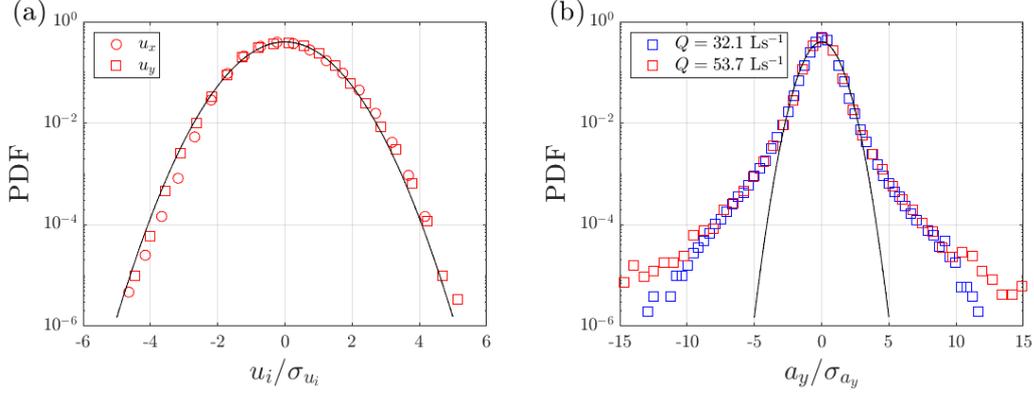


Figure 4. PDF of the streamwise and spanwise velocity fluctuations of the tracers in the Meander for $Q = 53.7 \text{ L s}^{-1}$. (b) PDFs of the spanwise accelerations for both flow rates. The distributions are normalized by their respective standard deviations. The continuous line represents the normalized Gaussian distribution.

266 To characterize how the turbulent energy is distributed across the scales of the flow,
 267 we consider the Eulerian second-order structure function of the velocity fluctuations $S_2^E(\mathbf{r})$
 268 (Kolmogorov, 1941; Pope, 2000). This is defined as the second moment of the velocity
 269 difference $\delta^E \mathbf{u}(\mathbf{r}) = \mathbf{u}(\mathbf{x}, t) - \mathbf{u}(\mathbf{x} + \mathbf{r}, t)$, where $\mathbf{u}(\mathbf{x}, t)$ and $\mathbf{u}(\mathbf{x} + \mathbf{r}, t)$ are the velocities of
 270 two particles separated by a distance \mathbf{r} at a given time t . Leveraging spatial homogeneity,
 271 we ensemble-average over all particle pairs at a distance r :

$$272 \quad S_2^E(\mathbf{r}) = \langle \delta^E \mathbf{u}(\mathbf{r})^2 \rangle \quad (1)$$

273 The ensemble-averaging requires binning the data over ranges of separation $r \pm \Delta r$, where we
 274 take $\Delta r = 1 \text{ mm}$ as a trade-off between resolution in scale-space and statistical convergence.
 275 Here we focus on the longitudinal structure function, in which the velocity component
 276 parallel to the separation vector \mathbf{r} is considered. Figure 5a shows that this exhibits an
 277 approximate $r^{2/3}$ scaling over separations from about 3 mm to 10 cm. This suggests the
 278 validity of the Kolmogorov (1941) ansatz in the inertial sub-range:

$$279 \quad S_2^E(r) = C_2(\epsilon r)^{2/3} \quad (2)$$

280 where ϵ is the dissipation rate of the turbulent kinetic energy, and C_2 is a constant. While the
 281 applicability of Kolmogorov theory to free-surface turbulence is debated (Hunt & Graham,
 282 1978; Magnaudet, 2003), experimental and numerical studies have documented a $k^{-5/3}$
 283 scaling of the energy spectra at or near the free-surface (Chickadel et al., 2011; Flores et
 284 al., 2017), equivalent to the $r^{2/3}$ scaling of the second-order structure function. Flores et al.
 285 (2017) report that, even though the mechanism underlying the spectral slope at the surface
 286 may differ from the 3D turbulence dynamics in the bulk, the proportionality constants are
 287 roughly the same. Therefore, here we assume $C_2 = 2.1$ as in 3D turbulence (Pope, 2000;
 288 Saddoughi & Veeravalli, 1994) and use Equation 2 to estimate ϵ from the plateau of the
 289 compensated structure functions in Figure 5b. We then estimate the dissipative scales of
 290 the free-surface turbulence, i.e., the Kolmogorov length and time scales, respectively:

$$291 \quad \eta = \left(\frac{\nu^3}{\epsilon} \right)^{1/4} \quad (3)$$

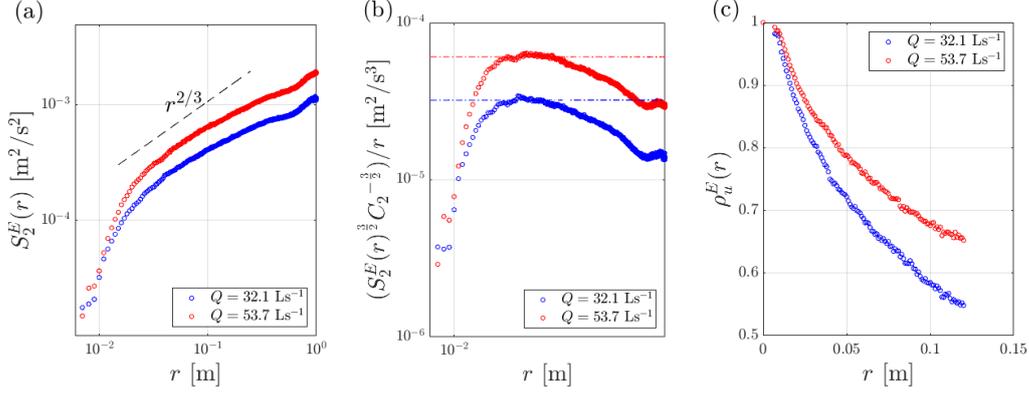


Figure 5. (a) Eulerian longitudinal second-order structure function, (b) compensated structure function and (c) Eulerian velocity autocorrelation function of the tracers for both flow rates. The dashed line in (a) corresponds to $r^{2/3}$ scaling. The dashed-dotted horizontal lines in (b) show the plateau of the compensated structure function which corresponds to the turbulent dissipation rate.

$$\tau_\eta = \left(\frac{\nu}{\epsilon}\right)^{1/2} \quad (4)$$

To determine the integral scales of the free-surface turbulence, we make use of the Eulerian velocity autocorrelation function, which in homogeneous isotropic turbulence can be calculated from the second-order structure function:

$$\rho_u^E(r) = \frac{\langle u(\mathbf{x}, t)u(\mathbf{x} + \mathbf{r}, t) \rangle}{\sigma_u^2} = 1 - \frac{S_2^E(r)}{2\sigma_u^2} \quad (5)$$

where σ_u^2 is the velocity variance. The velocity correlation exhibits an approximately exponential decay (Figure 5c), and the integral length scale L is evaluated by least-square fitting to it a function $Ae^{-r/L}$, where A is a constant of order unity. The estimates for the dissipative and integral scales, summarized in Table 3, support the notion that the $r^{2/3}$ scaling of the structure function applies over an inertial sub-range $\eta \ll r \ll L$. Additionally, an alternative estimate of the dissipation rate can be obtained from the classic scaling (Tennekes & Lumley, 1972):

$$\epsilon \approx C \frac{\sigma_u^3}{L} \quad (6)$$

Taking the proportionality constant $C = 0.5$ as typical for 3D turbulence in the high-Reynolds number limit (Burattini et al., 2005; Carter et al., 2016), we find dissipation estimates very close to those found from the second-order structure function (Table 3).

3.3 Effect of Particle Shape and Size in the Meander

In this section we compare the motion of the larger particles (discs and rods) against the tracers. We start by considering the Meander where the flow homogeneity allows for a comprehensive statistical description of the transport.

The Eulerian velocity fields of all particle types are found to be quantitatively similar. This is evident by comparing Figure 6a-b, where the mean velocity fields of the discs and

Table 3. Main physical quantities characterizing the free-surface turbulence for both flow rates in the Meander: RMS of the velocity fluctuations σ_u , integral length scales L , dissipation rate of turbulent kinetic energy ϵ , Kolmogorov length scale η , and Kolmogorov time scale τ_η .

Meander	$Q = 32.1 \text{ L s}^{-1}$	$Q = 53.7 \text{ L s}^{-1}$
σ_u [m s^{-1}]	0.022	0.032
L [m]	0.175	0.243
ϵ from (2) [$\text{m}^2 \text{ s}^{-3}$]	$3.2 \cdot 10^{-5}$	$6.1 \cdot 10^{-5}$
ϵ from (6) [$\text{m}^2 \text{ s}^{-3}$]	$3.1 \cdot 10^{-5}$	$6.6 \cdot 10^{-5}$
η [mm]	0.4	0.4
τ_η [s]	0.18	0.13

314 rods are displayed for $Q = 32.1 \text{ L s}^{-1}$, with the tracers in Figure 3a. For both considered
 315 flow rates, the RMS difference between the three particle types is less than 2% of \tilde{U} and
 316 less than 17% of $\tilde{\sigma}_u$.

317 Figure 6 also displays PDFs of particle velocities and accelerations for selected com-
 318 ponents and flow rates. To highlight the difference between the different particle types,
 319 the Kolmogorov velocity scale $u_\eta = \eta/\tau_\eta$ and acceleration scale $a_\eta = u_\eta/\tau_\eta$ are used for
 320 normalization. The velocity fluctuations are similar between all particle types, closely ap-
 321 proximating a Gaussian distribution (Figure 6c). On the other hand, the acceleration inter-
 322 mittency shown by the tracers is significantly reduced for the larger particles (Figure 6d).
 323 At $Q = 32.1 \text{ L s}^{-1}$, the RMS acceleration of the discs and rods is 9% and 21% lower than
 324 that of the tracers, respectively; while at $Q = 53.7 \text{ L s}^{-1}$ the reduction becomes 8% and
 325 20%, respectively.

326 To characterize the spreading rate of the floating particles, we consider their Lagrangian
 327 motion characterized by single-particle and particle-pair dispersion. The former examines
 328 how far, on average, a single particle migrates from its origin over time. Leveraging the
 329 homogeneity of the flow in the Meander and following the classic framework of Taylor (1921),
 330 the single-particle diffusivity can be derived from the Lagrangian velocity autocorrelation:

$$331 \quad \rho_u^L(\tau) = \left\langle \frac{\sum u(t)u(t+\tau)}{\sum u(t)^2} \right\rangle \quad (7)$$

332 Here the summation extends to all values of τ along each trajectory, i.e., the autocorrela-
 333 tion is first calculated along each trajectory and normalized by its velocity variance, before
 334 ensemble-averaging over all trajectories. This ensures that each trajectory has the same
 335 weight when contributing to the global autocorrelation coefficient (Guala et al., 2007). Ad-
 336 ditionally, we only consider trajectories whose duration is longer than the time delay τ
 337 (Mordant et al., 2004). Figure 7a-b display the Lagrangian velocity autocorrelation of each
 338 particle type for both considered flow rates, showing that the motion of the discs and rods
 339 is more time-correlated than that of the tracers. This is consistent with the trend reported
 340 by numerical simulations of inertial particles (Squires & Eaton, 1991; Jung et al., 2008) and
 341 laboratory observations of finite-size particles (Machicoane & Volk, 2016) in 3D turbulence.

342 The diffusivity K is obtained by integrating the decaying Lagrangian velocity autocor-
 343 relation (Taylor, 1921):

$$344 \quad K = \sigma_u^2 \int_0^\infty \rho_u^L(\tau) d\tau \quad (8)$$

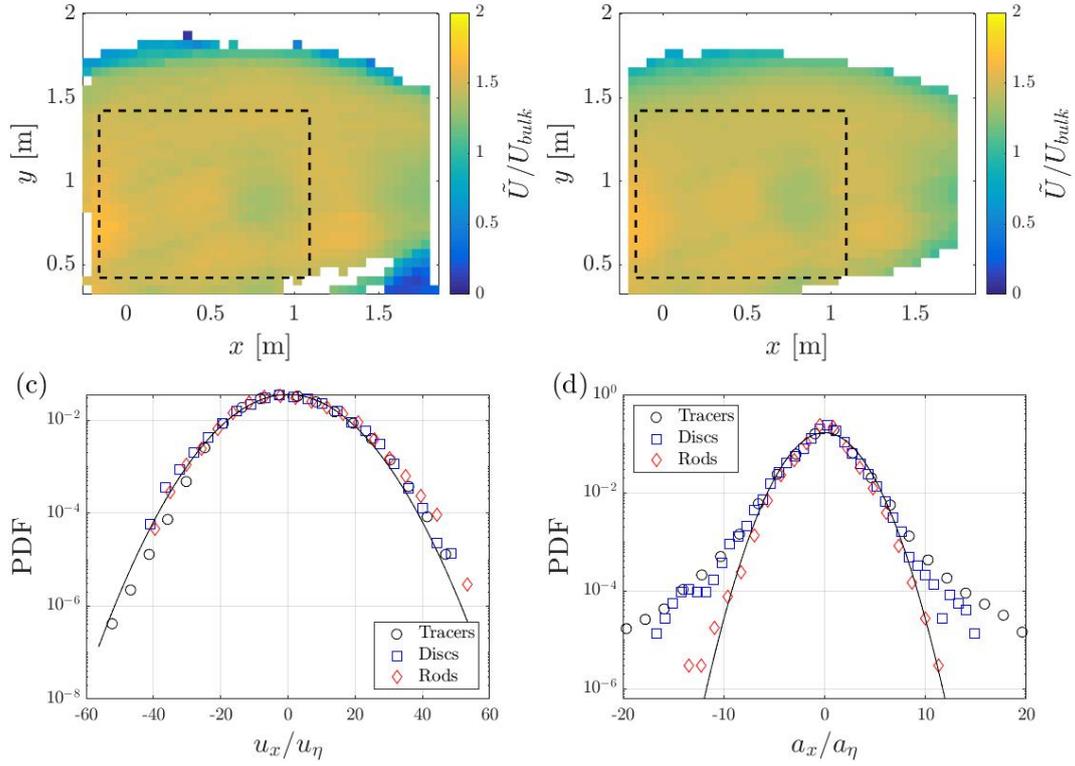


Figure 6. Eulerian mean velocity fields of the discs (a) and rods (b) for $Q = 32.1 \text{ L s}^{-1}$. (c) PDFs of the streamwise velocity fluctuations of the different particle types in the Meander for $Q = 53.7 \text{ L s}^{-1}$. (d) PDFs of the streamwise accelerations of the different particle types in the same location for $Q = 32.1 \text{ L s}^{-1}$. The distributions are normalized by Kolmogorov scaling. The continuous line represents the normalized Gaussian distribution.

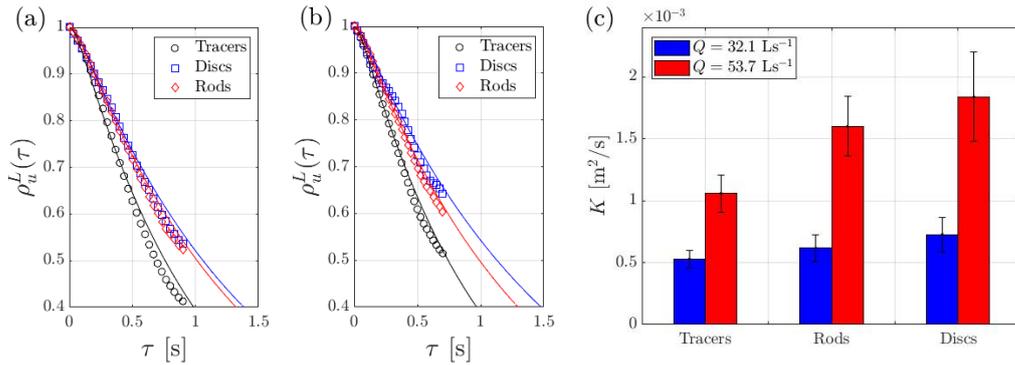


Figure 7. Lagrangian velocity autocorrelation function of each particle type for $Q = 32.1 \text{ L s}^{-1}$ (a) and $Q = 53.7 \text{ L s}^{-1}$ (b). The solid lines are the autocorrelation functions computed using Equation 9 which are integrated to obtain diffusion coefficients. (c) Diffusivity of the different particle types for both flow rates. The error bars represent the standard deviation of the diffusion coefficients from separate runs.

345 As the extreme of integration grows, the autocorrelation is expected to decay to negligibly
 346 small values and correspondently the diffusivity will asymptote to a value independent of
 347 time. Due to the finite length of the recorded trajectories, we extrapolate the autocorrelation
 348 using the stochastic model proposed by Sawford (1991):

$$349 \quad \rho_u^L(\tau) = \frac{T_L e^{-\tau/T_L} - T_2 e^{-\tau/T_2}}{T_L - T_2} \quad (9)$$

350 Two time scales are required: the integral time scale of the turbulence T_L , and a charac-
 351 teristic time scale related to the dissipation T_2 . The former is defined as the characteristic
 352 decay time of the Lagrangian velocity autocorrelation function and is estimated by least-
 353 square fitting $\rho_u^L(\tau)$ to an exponential function. The value of T_2 is estimated by fitting the
 354 experimental curve to Equation 9 and found to be approximately $0.3\tau_\eta$; this is the same
 355 order of magnitude as in 3D turbulence studies (Voth et al., 2002; Mordant et al., 2004).
 356 The diffusivity is then determined by the long-time asymptote of K . For the tracers we
 357 obtain normalized diffusivities $K/(u_\tau D) \approx 0.5$ for both flow rates, where we estimate the
 358 friction velocity u_τ from its relationship with the dissipation rate, $\epsilon = u_\tau^3/D$ (assumed
 359 to be mainly driven by bed friction (Raymond et al., 2012)). This falls well in the range
 360 $K/(u_\tau D) = 0.3$ to 0.9 reported for meandering channels (Fischer et al., 1979; Rutherford,
 361 1994). The diffusivity is plotted in Figure 7c for the different particle types and for both
 362 considered flow rates. One clearly sees an increase in K with increasing flow rate, hence
 363 with Reynolds number. Most importantly, the larger particles exhibit larger diffusivity than
 364 the tracers, with the discs dispersing faster than the rods.

365 To verify this notion, we consider the mean square displacement (MSD) of the particles
 366 due to turbulent fluctuations:

$$367 \quad \langle X(t)^2 \rangle = \langle \|\mathbf{x}(t) - \mathbf{x}(t_0) - \langle \mathbf{U} \rangle \Delta t\|^2 \rangle \quad (10)$$

368 where $\mathbf{x}(t)$ is the particle position at time t and $\mathbf{x}(t_0)$ is the reference position at the
 369 temporal origin of the trajectory t_0 . The advective displacement $\langle \mathbf{U} \rangle \Delta t$, due to the mean
 370 flow during the time interval $\Delta t = t - t_0$, is subtracted to isolate the contribution of the
 371 turbulent fluctuations. Leveraging spatial homogeneity, the advective flow is taken to be a
 372 uniform motion, which avoids the ambiguities associated to subtracting different advective
 373 displacements at different points along a same trajectory. The MSD of each particle type
 374 for both flow rates is plotted in Figure 8a-b and confirms that the discs spread faster than
 375 the rods, which spread faster than the tracers. We also note that MSD can alternatively be
 376 compute integrating the autocorrelation twice (Taylor, 1921; Pope, 2000):

$$377 \quad \langle X(t)^2 \rangle = 2\sigma_u^2 \int_0^t \int_0^{t'} \rho_u^L(\tau) d\tau dt' \quad (11)$$

378 where t' is a second integration variable. This yields analogous trends, not shown for brevity.

379 We then turn to pair-dispersion, quantified via the mean square separation (MSS):

$$380 \quad \langle [D(t) - D(t_0)]^2 \rangle = \langle [\|\mathbf{x}_i(t) - \mathbf{x}_j(t)\| - D(t_0)]^2 \rangle \quad (12)$$

381 Here $\mathbf{x}_i(t)$ and $\mathbf{x}_j(t)$ are the positions of the i -th and j -th particles at time t , respectively.
 382 The initial separation $D(t_0) = \|\mathbf{x}_i(t_0) - \mathbf{x}_j(t_0)\|$ is subtracted to account for possible cor-
 383 relation between initial separation and relative velocity of the particle pair (Ouellette et
 384 al., 2006). The minimum distance between particle edges is typically 0.01 m, thus inter-
 385 molecular effects as well as inter-particle capillary forces are negligible. Figure 8c-d shows
 386 the MSS of each particle type for $Q = 53.7 \text{ L s}^{-1}$, in both linear and logarithmic axes. The

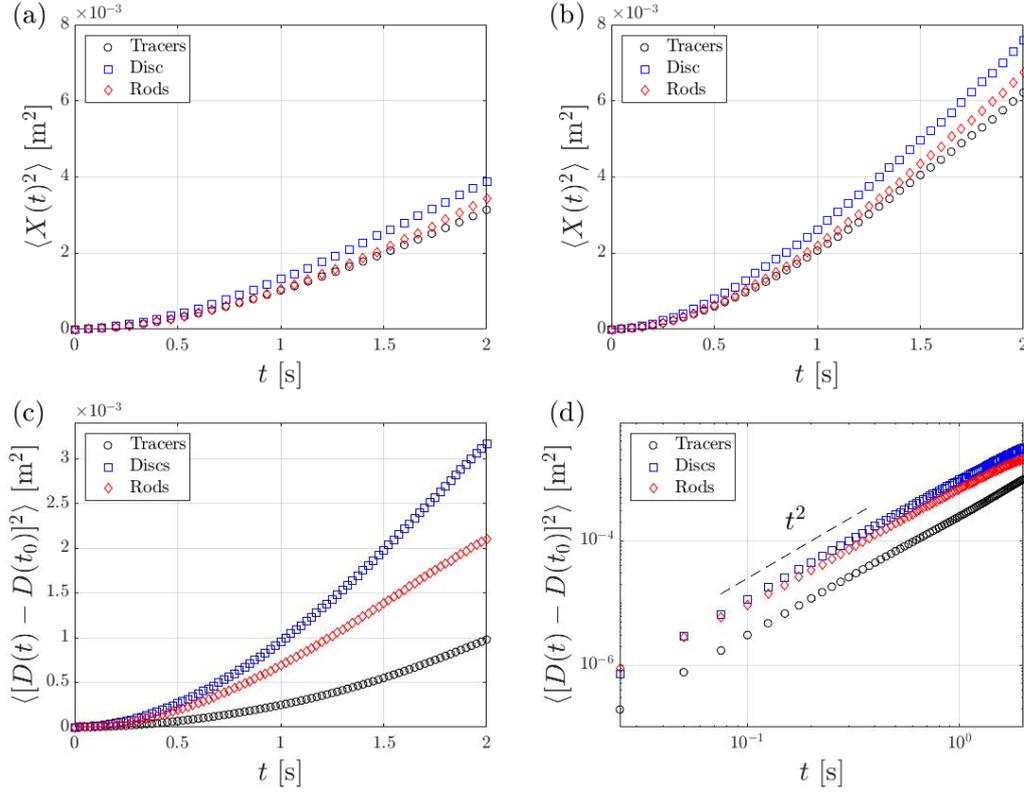


Figure 8. Mean square displacement due to turbulent velocity fluctuations of each particle type for $Q = 32.1 \text{ L s}^{-1}$ (a) and $Q = 53.7 \text{ L s}^{-1}$ (b). Mean square separation of each particle type for $Q = 53.7 \text{ L s}^{-1}$, plotted in linear axes (c) and logarithmic axes (d). The dashed line in (d) corresponds to the t^2 (ballistic) scaling.

387 t^2 scaling predicted by Batchelor (1950) for particles in turbulence in the initial phase of
 388 separation is recovered, indicating ballistic separation in the considered regime and time
 389 scales. Consistently with the picture gained from the single-particle dispersion analysis, the
 390 discs separate from each other faster than the rods, and the latter separate faster than the
 391 tracers. We note that, for delays longer than the turnover time of eddies of size $D(t_0)$ (with
 392 $D(t_0)$ in the inertial sub-range), Richardson (1921) predicted super-diffusive dispersion with
 393 $D^2 \sim t^3$. However, such regime has been historically difficult to retrieve even in laboratory
 394 experiments (Bourgoin, 2015; Tan & Ni, 2022; Elsinga et al., 2022).

395 3.4 Effect of Particle Shape and Size in the Riffle

396 In this section we verify that the trends observed in the Meander also applies to the
 397 significantly different flow conditions found in the Riffle. Also here, the Eulerian fields of \tilde{U}
 398 and $\tilde{\sigma}_u$ for the discs and rods (not shown) are close to those measured for the tracers and
 399 shown in Figure 5c-d, with RMS difference between the three particle types less than 12 %
 400 for \tilde{U} and less than 16 % for $\tilde{\sigma}_u$. Nevertheless, as for the Meander, we shall see that the
 401 particle shape and size influences the Lagrangian dispersion.

402 Because the mean velocity in the Riffle is predominantly aligned with x' , we can isolate
 403 the turbulent dispersion by considering the lateral displacement, i.e., the MSD along the
 404 spanwise direction y' :

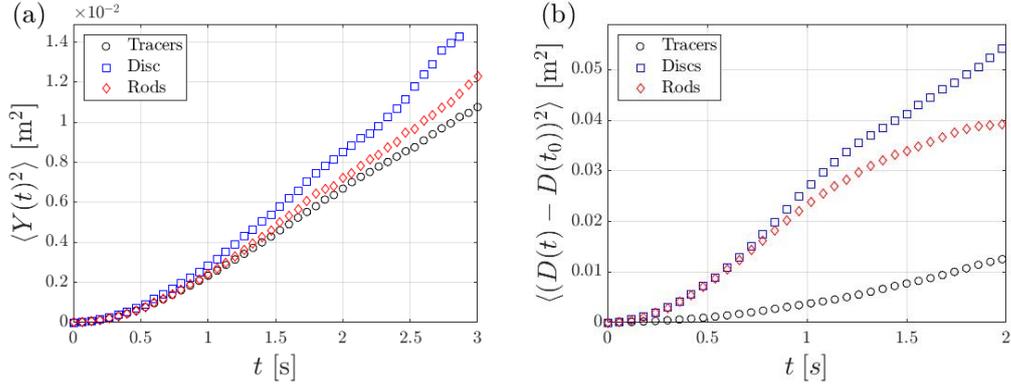


Figure 9. (a) Spanwise MSD of the different particles in the Riffle for $Q = 32.1 \text{ L s}^{-1}$. (b) MSS of each particle type in the same location for $Q = 53.7 \text{ L s}^{-1}$.

$$\langle Y(t)^2 \rangle = \langle [y'(t) - y'(t_0)]^2 \rangle \quad (13)$$

This is plotted in Figure 9a for $Q = 32.1 \text{ L s}^{-1}$ ($Q = 53.7 \text{ L s}^{-1}$ displaying analogous results). This indicates again that larger particles spread faster than the tracers, with the discs spreading faster than the rods. An estimate of the lateral diffusion coefficient can be derived from the relation:

$$K_{y'} = \frac{1}{2} \frac{d\langle Y(t)^2 \rangle}{dt} \quad (14)$$

A linear least-square fit to the data over the range $t > 1.5 \text{ s}$ (where the MSD is approximately linear with time) yields $K_{y'} = 0.002 \text{ m}^2 \text{ s}^{-1}$, $0.0025 \text{ m}^2 \text{ s}^{-1}$ and $0.003 \text{ m}^2 \text{ s}^{-1}$ for the tracers, rods, and discs, respectively. The same trend is retrieved for the particle-pair dispersion quantified by the MSS (Figure 9b), which is calculated as per Equation 12 for $Q = 53.7 \text{ L s}^{-1}$. Both MSD and MSS indicate faster spreading rates in the Riffle compared to the Meander, which is expected due to the higher turbulence intensity σ_u/U_{bulk} in the former.

3.5 Rotational dynamics

The translational and rotational motions of anisotropic particles in turbulence are strongly coupled to each other (Voth & Soldati, 2017). We therefore consider the rotational dynamics of the rods, as they can provide insight in the transport behaviour presented in the previous section. We present results for $Q = 53.7 \text{ L s}^{-1}$, with $Q = 32.1 \text{ L s}^{-1}$ showing analogous trends. We first consider the rod alignment defined by the orientation vector $\hat{\mathbf{p}}(t)$. Figure 10a shows the PDF of $|\hat{\mathbf{p}}(t) \cdot \hat{\mathbf{U}}(t)|$, where $\hat{\mathbf{U}}(t)$ is the unit vector parallel to the particle velocity. For both ROIs, the rods display a preference to align with the direction of motion. Considering the close similarity between the velocity fields of the tracers and those of the rods, this can be interpreted as a preferential alignment with the flow direction.

The curvature of the streamlines in the ROIs is small, but the rods orientation varies in time due to the flow fluctuations. We characterize the time scales associated to the rods' re-orientation by the Lagrangian autocorrelation of the particle orientation vector $\rho_{\hat{\mathbf{p}}}^L(\tau)$, calculated analogously to the velocity autocorrelation function in Equation 11 and shown in Figure 10b. In the Meander the particle orientation is remarkably stable, which is consistent with its moderate turbulence intensity: the fluid velocity, with which the rods tend to be aligned, remains mostly oriented in the streamwise direction. The orientation

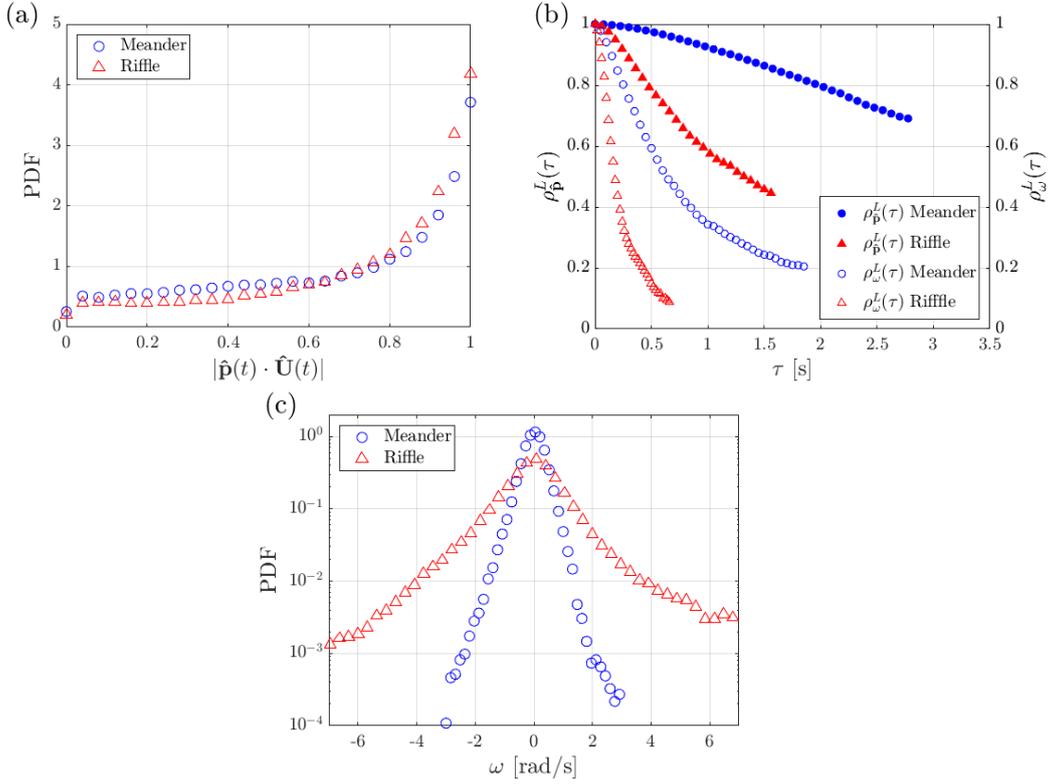


Figure 10. (a) PDF of the cosine of the orientation angle of the rods in both ROIs. (b) Lagrangian autocorrelation functions of the rods' orientation and angular velocity in both ROIs. (c) PDFs of the angular velocities in both ROIs.

434 autocorrelation in the Riffle shows a faster decay with a characteristic time of approximately
 435 1.5 s. Given the jet-like flow structure, a candidate time scale dictating the rod reorientation
 436 is provided by the intense shear layers (Figure 5c-d). Indeed, visual observation confirms
 437 that the rods' rotation in those regions follows the direction of the mean flow shear. The
 438 associated time scale can be estimated from the jet half width, $d_{\frac{1}{2}} \approx 0.5$ m, and the velocity
 439 difference across it, $\Delta\tilde{U} \approx 0.4$ m s⁻¹: $d_{\frac{1}{2}}/\Delta\tilde{U} \approx 1.25$ s, which approximately agrees with the
 440 observed correlation time scale. The fact that the time scale of re-orientation be attributed
 441 to the mean shear of the surface flow is consistent with the observation that the rods'
 442 orientation is very stable in the Meander, where the flow is highly homogeneous and lateral
 443 shear is weak.

444 Figure 10b also shows the autocorrelation of the angular velocity $\rho_{\omega}^L(\tau)$, which as expected
 445 decays significantly faster than $\rho_{\mathbf{p}}^L(\tau)$. For the Meander, the correlation time scale
 446 of $\rho_{\omega}^L(\tau)$ is approximately 1 s, matching the integral time scale of the free-surface turbulence
 447 T_L . In the Riffle, the same quantity decays with a characteristic time scale around 0.25 s.
 448 While a single value of T_L can hardly be defined in the Riffle due to spatial inhomogeneity,
 449 we note that σ_u is roughly 4 times larger than in the Meander. This suggests that, in both
 450 ROIs, the correlation time scale of $\rho_{\omega}^L(\tau)$ is dictated by the energetic eddies that determine
 451 the integral scales of the turbulence. Since the rods' length is two orders of magnitude
 452 larger than η and a fraction of L , this finding is in line with the view that rods' rotation is
 453 controlled by eddies of size comparable to or larger than their length (Parsa & Voth, 2014;
 454 Voth & Soldati, 2017).

455 The intermittent nature of the free-surface turbulence, displayed in the acceleration
 456 PDF in Figure 6d, is also reflected in the distribution of the angular velocity PDF shown
 457 in Figure 10c. The kurtosis of these distribution is 5.9 and 8.8 for the Meander and Riffle,
 458 respectively, indicating relatively large probability of extreme events with angular velocities
 459 of several rads^{-1} , especially with higher turbulence intensity of the free-surface. Such sud-
 460 den changes in orientation are expected to alter the Lagrangian transport by the underlying
 461 flow.

462 4 Discussion

463 The PTV measurements of the small tracer particles, especially in the spatially ho-
 464 mogeneous sub-region of the Meander, informs us on the nature of the free-surface flow in
 465 the considered riverine environment. We remark that, for fundamental reasons, free-surface
 466 turbulence is not expected to be equivalent to either 2D or to 3D turbulence: the surface
 467 exchanges energy and enstrophy with the flow underneath, hence neither quantity can be
 468 regarded as invariant and dimensional scaling arguments do not strictly apply (Cressman et
 469 al., 2004). However, the present measurements do indicate strong similarity with the phe-
 470 nomenology of 3D turbulence. In particular, the behaviour of the second-order structure
 471 function is consistent with Kolmogorov (1941) scaling in the inertial sub-range. While a
 472 similar scaling is also expected in the inverse-cascade range of 2D turbulence (Kraichnan,
 473 1967), the latter framework is inconsistent with the observed intermittency of the accelera-
 474 tion (Boffetta & Ecke, 2012). The close agreement between the dissipation estimates from
 475 Equations 2 and 6 further supports the applicability of the 3D turbulence framework. The
 476 similarity between 3D and free-surface turbulence is possibly due to the surface carrying
 477 the prominent fingerprints of sub-surface vortices connected to it. These evolve by diffusion
 478 and stretching, as vortex tilting is annihilated at the surface (Shen et al., 1999; Zhang et
 479 al., 1999; Shen & Yue, 2001). In other words, unlike in 2D turbulence, the free-surface
 480 boundary condition affects but does not suppress vortex stretching, which is essential to the
 481 energy cascade in 3D turbulence (Davidson, 2015; Carbone & Bragg, 2020; Johnson, 2020).
 482 Our results are specific to a particular riverine flow configuration, therefore further studies
 483 are needed to assess the generality of the observations, especially for different water depths.
 484 Indeed, vortex stretching is hindered in shallow flows, which can trigger the emergence of
 485 features peculiar to 2D turbulence (Uijttewaalt & Booij, 2000; Stocchino et al., 2011). More-
 486 over, water depth influences the respective role of water-column turbulence and bed friction
 487 in setting the dissipation rate at the surface (Raymond et al., 2012; Ulseth et al., 2019).

488 Our main finding is that, in both investigated ROIs, larger floating particles disperse
 489 faster than smaller tracers. This result can be interpreted based on our understanding
 490 of the behaviour of inertial particles in turbulence. We remind that the term “inertial”
 491 indicates objects too heavy and/or too large to faithfully follow the fluid flow (Brandt &
 492 Coletti, 2022). Indeed, both discs and rods display weaker and less intermittent accelerations
 493 than the tracers. This behaviour is well known from the investigation of 3D turbulence
 494 laden with inertial particles and is attributed to two concurring mechanisms: preferential
 495 sampling of high-strain/low-vorticity regions, prevalent for small St ; and inertial filtering of
 496 the small-scale/high-frequency fluctuations, prevalent for large St (Bec et al., 2006; Toschi
 497 & Bodenschatz, 2009). Although we do not evaluate St , the large size of the discs and rods
 498 compared to the Kolmogorov scales suggest that inertial filtering is the likely cause: the
 499 larger particles respond to a spatial average of the fluid velocity, making them less sensitive
 500 to the smaller and faster-decaying eddies. This is consistent with the increasingly time-
 501 correlated motion of the larger particles. The slower decay of the velocity autocorrelation
 502 is consistent with the simulations of Shin and Koch (2005) for rods in 3D turbulence, who
 503 found the correlation time scale T_L to increase with the rods’ length. However, in such study
 504 the RMS velocity fluctuations of the rods σ_u was found to decrease with their length, and
 505 the diffusivity $K = \sigma_u^2 T_L$ ultimately decreased. Vice versa, in the present case, the RMS
 506 velocity fluctuations of the particles are not significantly affected by their size and shape,

507 and thus the diffusivity follows the same trend as T_L . This different outcome with respect
508 to Shin and Koch (2005) could be rooted in qualitative differences between free-surface and
509 3D turbulence. Another possible explanation is the range of scales at play: their simulations
510 reached relatively low Reynolds numbers, with a range of scales $L/\eta < 30$, and were focused
511 on rods of length comparable to or larger than the integral scales. On the other hand, in
512 the present study L/η exceeds 600, and the particle size is at most one third of the integral
513 scale. Therefore, the floating discs and rods filter out only a fraction of the fluctuating
514 energy contained in the large eddies, and therefore their velocity variance remains close to
515 that of tracers.

516 Despite the rods' length being almost twice the discs' diameter, the latter disperse faster
517 than the former. This may be due to the discs possessing a larger wetted area, thus more
518 effective filtering of the small-scale fluctuations. However, the object shape is also likely to
519 have a profound influence on the Lagrangian transport. The characteristic time scale of the
520 angular velocity and its intermittent nature indicate that the instantaneous orientation of the
521 rods is affected by a range of turbulent scales. These may contribute to decorrelating their
522 translational motion, which for anisotropic particles is strongly coupled with the rotational
523 motion (Voth & Soldati, 2017). Moreover, the rods' tendency to align with the flow direction
524 suggests that the large scales of the turbulence (at least those larger than the rods' length)
525 are not isotropic but likely populated by streamwise-oriented structures. Indeed, already
526 early studies of open channel flows highlighted the connection between near-wall bursts in
527 the bottom-wall boundary layer and the coherent motions that transfer mass to and from
528 the free-surface (Nakagawa & Nezu, 1981; Rashidi & Banerjee, 1988). The complex bed
529 topography of a natural channel is likely to enhance this connection by generating energetic
530 eddies that can travel up to the surface, as indicated by the fact that bed roughness in
531 shallow streams strongly correlates with gas transfer velocity (Ulseth et al., 2019).

532 Besides shape and size, other properties of floating particles may be influential towards
533 their free-surface transport; in particular, bulk density and surface characteristics. Par-
534 ticles of higher density and mass may be more effective in filtering small-scale turbulent
535 fluctuations, which could further enhance their diffusivity. However, depending on the size,
536 this effect could be counteracted by a lack of responsiveness to some of the energetic scales
537 responsible for the dispersion. Moreover, density and surface characteristics, in particular
538 hydrophobicity, will affect the balance between surface tension and gravity, determining
539 the submerged fraction of the floating object (Koh et al., 2009; Ji et al., 2018). In turn,
540 submergence will determine the amount of windage, i.e., the drag exerted by the airflow
541 on objects partly protruding out of the water (Zambianchi et al., 2014; Beron-Vera et al.,
542 2019). Finally, while we have limited our study to sparse objects that do not significantly
543 interact with each other, compressibility of the free-surface flow is known to produce intense
544 clustering that can bring the floaters in close contact (Cressman et al., 2004; Lovecchio et
545 al., 2013). Again, the material properties of the particles are then expected to affect the
546 short-range interactions and possibly lead to aggregation (Vella & Mahadevan, 2005). The
547 impact of such particle properties, which is outside the scope of the present work, clearly
548 warrants further systematic investigations using different particle materials.

549 The observed influence of the particles' properties on dispersion, once confirmed for a
550 wider range of particle types and flow conditions, may have profound implications for the
551 transport of floating particles; in particular, the transport of meso- and macroplastics in
552 small streams and turbulent waters in general. The diffusivity, which we find to roughly
553 double from mm-sized to cm-sized objects, is a crucial quantity to incorporate the effect of
554 unresolved spatio-temporal scales in Lagrangian transport models for rivers, lakes, and the
555 oceans (Liu et al., 2011; Park et al., 2017; van Sebille et al., 2018; Daily & Hoffman, 2020;
556 McDonald & Nelson, 2021). Our results indicate that such parameter varies significantly not
557 only with the flow conditions, but also with the particle properties. Parameterizations that
558 include also the latter appear necessary to obtain accurate predictions from such models.

559 5 Conclusion

560 Motivated by the need of understanding the transport of plastic litter in river flows,
 561 we have used time-resolved PTV to characterize the motion of particles of different shape
 562 and size floating on the surface of a field-scale meandering stream. We have considered
 563 two locations with different turbulence levels, in which the role of surface waves on the
 564 transport is deemed negligible. We have measured the position, velocity, and acceleration
 565 along the trajectories of thousands of millimetre-sized spherical pellets and centimetre-sized
 566 discs and rods, as well as the orientation and rotation of the latter, and evaluated the spatio-
 567 temporal scales associated with such quantities. At the Meander, the homogeneity of the
 568 flow properties allows us to identify both dissipative and integral scales of the free-surface
 569 turbulence, providing essential terms of comparison for the size of the particles and the
 570 scales of their motion. The spheres are small enough to capture most if not all scales of the
 571 free-surface motion and are regarded as flow tracers; while the length of the rods and the
 572 diameter of the discs are $\mathcal{O}(100)$ times larger than the dissipative scales and several time
 573 smaller than the integral scales of the turbulence. The analysis of the particles' motion leads
 574 to the following observations:

- 575 I. All considered particles displays almost indistinguishable mean velocities and RMS
 576 velocity fluctuations. These are determined by the largest scales of the surface flow,
 577 to which the particles respond faithfully.
- 578 II. While the velocity fluctuations follow normal distributions unaffected by the particle
 579 shape and size, the accelerations show a sizeable degree of intermittency which de-
 580 creases for larger particles. This is attributed to the finite-size of the particles, filtering
 581 out the smallest scales of the turbulence associated to the most intense gradients.
- 582 III. Consequently, the larger particles spread more rapidly on the turbulent free-surface,
 583 with diffusivity coefficients roughly doubling for centimetre-sized particles as compared
 584 to millimetre-sized tracers. This is due to the motion of the larger particles being more
 585 time-correlated, which in turn is rooted in their impaired response to the small-scale
 586 turbulent fluctuations.
- 587 IV. The rods tend to align with the flow direction, but their instantaneous orientation is
 588 influenced by a range of scales: they re-orient following the mean shear, rotate ac-
 589 cording to the turnover time of the energetic eddies, and exhibit intermittency in their
 590 angular velocities. This leads to less time-correlated motions and slower dispersion
 591 than the discs, despite the rods' length being larger than the discs' diameter.

592 Overall, the behaviour of the free-surface turbulence and the motion of particles floating on it
 593 appears consistent with the phenomenology of inertial finite-sized particles in 3D turbulence.
 594 This similarity, to be confirmed in a wider range of flow conditions and particle types, may
 595 allow leveraging of established results and recent advances in the field of particle-laden
 596 turbulence (Balachandar & Eaton, 2010; Brandt & Coletti, 2022), furthering the predictive
 597 understanding of the transport of floating plastics in natural waters.

598 Future studies shall expand the present work in several directions. Our experiments
 599 have been carried out in a relatively small stream; studies in larger and deeper rivers,
 600 in which the dissipation mechanisms in the water column are inherently different (Moog &
 601 Jirka, 1999), are needed to expand and generalize the results. In such cases, particle imaging
 602 may require the use of uncrewed aerial vehicles, which have been successfully utilized to
 603 characterize natural flows (Blois et al., 2016; Liu et al., 2021). Given the variety of debris
 604 types found in water streams, the range of particle properties should be expanded beyond
 605 shape, size, and density: deformability and brittleness have recently been investigated in
 606 laboratory studies and are especially relevant to plastic pollution (Brouzet et al., 2014, 2021).
 607 Finally, high-Froude streams and/or under the action of wind, breaking and non-breaking
 608 waves may play a major role in the transport of floating particles. The recent laboratory
 609 experiments of Lenain et al. (2019), confirming computational results by Deike et al. (2017),
 610 found that breaking waves induce much stronger transport of cm-sized spherical particles

611 compared to Stokes drift. Studies investigating the effect of particle properties in similar
 612 situations are warranted.

613 **Data Availability Statement**

614 The pre-processed background-subtracted images for the different regions of interest, flow
 615 rates and particles are available at <https://doi.org/10.3929/ethz-b-000572787>.

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617

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