

1 **Effect of shape and size on the transport of floating**
2 **particles on the free surface in a meandering stream**

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9 **Key Points:**

- 10 • The velocity of floating particles in turbulent streams is weakly affected by their shape
11 and size.
- 12 • Larger particles disperse faster on the free surface due to their ability to filter out
13 small-scale turbulent fluctuations.
- 14 • Rods re-orient following the mean shear of the surface flow and rotate according to
15 the integral scales of the free surface turbulence.

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Abstract

Understanding how floating particles are transported by streaming waters is crucial in predicting the transport of plastic pollution, which is dramatically abundant in rivers, lakes, and oceans. Using particle tracking velocimetry, we investigate the motion of floating particles of different shape and size on the turbulent free surface of a field-scale meandering stream. We consider two different locations, in both of which the role of surface waves on transport is deemed negligible. Millimetre-sized spheres are used as tracers to characterize the surface flow. These are compared with centimetre-sized discs and rods, approximating typical-sized pieces of floating litter. The larger particles exhibit similar mean and fluctuating velocities as the tracers but filter out the extreme turbulent accelerations. Consequently, their motion is more time-correlated and their spreading rate is larger. This behaviour is also confirmed by complementary laboratory measurements in an open channel flow. The rotation of the rods, affected by a range of turbulent scales, reduces the correlation time scale of their translational motion, and leads to a slower dispersion compared to the discs, despite the rods' length being larger than the discs' diameter. Taken together, these results indicate that the motion of finite-sized objects floating on the surface of weakly wavy turbulent waters is consistent with the behaviour of inertial particles in three-dimensional turbulence. These results can be valuable when constructing predictive models of floating plastics.

Plain Language Summary

Plastic debris is a rising global issue severely affecting the state of our rivers, lakes and oceans. Understanding how pieces of litter, often floating, travel in streaming waters is crucial for predicting and ultimately limiting plastic pollution. The main goal of this research is to investigate how the shape and size of small floating objects may affect their journey on the surface of water. To this end, we use high-speed imaging to track floating objects of different shape and size in an outdoor stream laboratory. The motion of centimetre-sized discs and rods, approximating typical pieces of plastics found in rivers, is directly compared to the motion of millimetre-sized spheres that follow the surface flow. We find that the larger discs and rods spread faster on the surface of water. Not only can these results be used to devise effective sequestration strategies, but they can be important to inform computer models that predict the abundance, and fate, of plastic litter in natural waters.

1 Introduction

Plastic debris is ubiquitous in our lakes, oceans, and coastal waters, posing a serious threat to human health and the environment (Eriksen et al., 2013; van Sebille et al., 2015; Lebreton et al., 2018). Recent findings demonstrate that about 1000 rivers account for 80 % of the global annual emissions of 0.8 to 2.7 million tons of plastics into the oceans per year, with small urban rivers among the most polluting (Meijer et al., 2021). Riverine ecosystems themselves are also affected by such pollution (van Emmerik & Schwarz, 2020). Plastic objects enter such systems in a wide range of compositions, shapes, and sizes before degrading into so-called microplastics (typically defined as pieces smaller than 5 mm). Significant efforts have been made to characterize the transport of microplastics throughout the water column (Ballent et al., 2012; H. Zhang, 2017). Several studies have explored different approaches to mitigate plastic pollution with different remediation strategies (Helinski et al., 2021; E. Zhang et al., 2022). Still, a large proportion of plastic waste in the U.S. is comprised of polyethylene and polypropylene, which are less dense than fresh water (Jambeck et al., 2015), and in general, it is estimated that more than half of all plastics produced are positively buoyant (Geyer et al., 2017). The question that motivates the present study is at which rate floating meso- and macroplastics (particles in the size range of 5 mm and larger) spread over the surface of turbulent streaming waters.

65 The transport of floating particles has been mainly investigated in terms of its depen-
66 dence on surface waves. These impart a net drift velocity in the wave propagation direction,
67 known as Stokes drift (van den Bremer & Breivik, 2018). While this is typically much smaller
68 than the mean advective velocity, its magnitude increases with wave steepness and can play
69 a role in the long-term dispersion (van Sebille et al., 2020). De Leo and Stocchino (2022)
70 found that the wave-induced transport of negatively buoyant plastic particles is confined to
71 a ballistic regime and a diffusive regime is rarely observed. However, the particle-to-fluid
72 density ratio has been shown to affect the total transport by waves (Stocchino et al., 2019).
73 While these studies have considered microplastics, mesoplastics have been shown to dwell
74 in the near-shore regions until they degrade into microplastics which then spread offshore
75 (Isobe et al., 2014). DiBenedetto et al. (2018, 2019) showed that non-spherical particles in
76 wavy waters tend to follow a preferred orientation, which affects their settling velocity if
77 those are negatively buoyant. For buoyant particles, DiBenedetto (2020) found that waves
78 result in non-uniform particle concentration. Ultimately, to obtain a global perspective
79 of the transport of plastics, one must also consider the effects of wind mixing, boundary
80 currents and meteorological conditions (Ourmieres et al., 2018; Kukulka et al., 2012).

81 The nature of turbulence of the free surface is still debated. Pan and Banerjee (1995)
82 identified hallmark features such as upwelling and downwelling motions and long-lived vor-
83 tices. Kumar et al. (1998) measured a k^{-3} decay of the velocity spectra (k being the
84 wavenumber), consistent with the expectation for two-dimensional (2D) turbulence. On the
85 other hand, the field measurements of Chickadel et al. (2011) displayed a $k^{-5/3}$ behaviour
86 typical of three-dimensional (3D) turbulence. In the riverine environment, the shallowness
87 of the flow plays a significant role in determining the nature of the turbulence: in partic-
88 ular, in the presence of strong lateral shear, the limited depth inhibits vortex stretching
89 and may result in vortex dynamics akin to 2D-turbulence, especially at low wavenumbers
90 (Uijttewaal & Booij, 2000). Most previous studies focused on free surface turbulence have
91 been concerned with the topological features of the flow, often in relation to air-water gas
92 fluxes (Shen et al., 1999; Shen & Yue, 2001; McKenna & McGillis, 2004; Turney & Baner-
93 jee, 2013; Herlina & Wissink, 2014), with only a few studies concerned with the transport
94 of particles on it. Particularly, Cressman et al. (2004) and Lovecchio et al. (2013) found
95 that tracer particles floating on the free surface cluster into string-like structures with long
96 lifetimes. Characteristic features of shallow flows, such as transitional macro-vortices, have
97 been found to greatly affect the single-particle and particle-pair dispersion (Stocchino et al.,
98 2011).

99 Several field studies have been concerned with natural free surface flows, focusing on
100 the effectiveness of free surface velocity measurements (e.g., for discharge estimation as well
101 as flow monitoring during flood events). The methods include acoustic Doppler velocimetry
102 (ADV) but also imaging techniques originally developed for laboratory flow studies, such
103 as particle image velocimetry (PIV) and particle tracking velocimetry (PTV) (Raffel et al.,
104 2018; Adrian & Westerweel, 2011). Free surface PIV and PTV present technical challenges
105 even in laboratory studies: the choice of appropriate tracers and their successful imaging in
106 spite of surface reflections (Weitbrecht et al., 2002; Miozzi et al., 2010; Miozzi & Romano,
107 2020; Gomit et al., 2022). These difficulties are exacerbated in field studies due to uneven
108 natural illumination and scarcity of detectable floating tracers. Nevertheless, these tech-
109 niques have gained favour in riverine flow investigations due to the richness of the data they
110 can provide (Jin & Liao, 2019; Tauro et al., 2016, 2019). Recent studies regarding floating
111 debris have shown the importance of surface tension and how it can play a key role in the
112 transport of partially submerged floating macroplastics, as its effects can be of the same
113 order or magnitude as buoyancy and turbulence (Valero et al., 2022).

114 Here we investigate experimentally the motion of floating particles on the turbulent
115 free surface of a meandering stream in an outdoor facility which offers laboratory-quality
116 measurements, and control, in a field-scale setting. We focus on regimes in the absence
117 of wind where the amplitudes of, and the drift induced by, surface waves are too small to



Figure 1. (a) The OSL facility, with the locations of the two ROIs (meander and pool) indicated by arrows. The tent, shown here at a downstream location, is deployed over the ROIs in the present experiments. (b) The traversing system holding the camera used for free surface imaging, indicating the approximate location of the FOV in the meander and the 2D coordinate system.

118 appreciably modify the advective transport. The main goal of the study is to explore the
 119 influence of the shape and size of floating particles along their trajectories when driven by
 120 the multi-scale fluctuations of the free surface flow. The focus is on size ranges relevant to
 121 meso- and macroplastics (≥ 5 mm) which are highly relevant to but largely understudied
 122 in river flows (van Emmerik & Schwarz, 2020). Applying time-resolved PTV to millimetre-
 123 sized spheres, we obtain surface velocity fields at two different locations along the stream.
 124 We then characterize the transport of centimetre-sized discs and rods and directly compare
 125 them to the behaviour of the spheres. In particular, we examine the floating particles’
 126 response to the free surface turbulent fluctuations which in turn affects their spreading
 127 rate. The observed behaviours are confirmed in well-controlled laboratory experiments,
 128 indicating the findings hold beyond the specific field settings. The rotational dynamics of
 129 the rods is considered to gain insight on their dispersion as compared to the discs. As we
 130 will discuss, the sensitivity of the particle dispersion to small-scale turbulence may have
 131 important consequences for modelling approaches based on flow velocity data (which are
 132 necessarily coarse-grained in space and time).

133 2 Materials and Methods

134 2.1 Field-scale Stream Facility and Hydrodynamic Characterization

135 Measurements are performed in the Outdoor StreamLab (OSL), an outdoor field-scale
 136 experimental stream facility at the Saint Anthony Falls Laboratory, University of Minnesota
 137 (Figure 1a). Water is drawn from the Mississippi River, flows through a meandering channel
 138 and discharges back into the river. The flow rate is controlled via a valve at the inlet, and
 139 the incoming water flows into a headbox and over a weir before entering the channel. The
 140 monitoring of the water height at the weir allows real-time calculation of the flow rate Q .
 141 Two flow rates are considered, $Q_1 = 32.1 \text{ L s}^{-1}$ and $Q_2 = 53.7 \text{ L s}^{-1}$, for which transport of
 142 sediment is negligible and the river bed is static. A Massa M300 ultrasonic distance probe
 143 and a sonar transducer are mounted on a programmable measurement carriage, performing
 144 2D elevation scans of the water surface and channel bed. Measurements are acquired in two
 145 regions of interest (ROI): one located at one of the meanders in the stream, and the other
 146 over a scour pool downstream of a riffle. We will refer to these measurement locations as the
 147 meander and the pool, respectively. The riverbanks can affect significantly the transport of
 148 particles (van Emmerik & Schwarz, 2020). In the present experiments, we have not focused
 149 on this aspect and rather investigated the interaction of floating particles on the turbulent
 150 free surface. Therefore, to minimize interaction with the bank, vegetation along it was either
 151 trimmed away or pinned down to avoid particle entrapment.

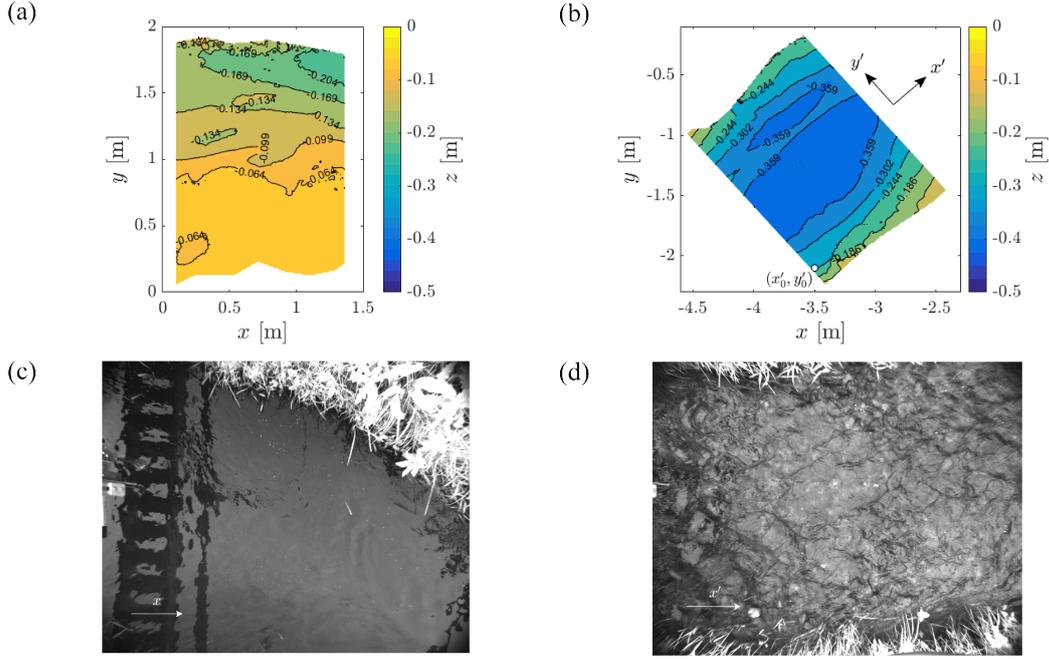


Figure 2. Bathymetry of the meander (a) and the pool (b) for Q_2 . The origin of the alternative coordinate system for the pool (x'_0, y'_0) and the directions of the abscissa and ordinate are indicated. Instantaneous photographs of the free surface at the meander (c) and the pool (d) for the same flow rate, indicating the streamwise direction x and x' , respectively.

152 The bathymetries of both ROIs are shown in Figure 2a-b. The origin of the global
 153 coordinate system is chosen to be on the bank of the meander, with x approximately in the
 154 streamwise direction, y pointing from the inner to the outer bank, and $z = 0$ m corresponding
 155 to the water surface. At the pool we also define an additional coordinate system $x'-y'$, with
 156 origin $(x'_0, y'_0) = [-3.5, -2.2]$ m and x' approximately aligned with the local flow direction.

157 The near- and sub-surface flow velocity $\vec{u}(\vec{x}, t)$ is also characterized by a Nortek Vectrino
 158 ADV probe, traversed along the cross-sections at $x = 1$ m and $x' = 1$ m in the meander and
 159 the pool, respectively. The phase-space thresholding technique described in Parsheh et
 160 al. (2010) is used to remove occasional spurious velocity spikes (e.g. due to air bubbles).
 161 Measurements are acquired at 100 Hz for 120 s. In the meander 24 and 27 locations are
 162 sampled along the cross-section for Q_1 and Q_2 , respectively. Correspondingly, 21 locations
 163 are sampled in the pool for both flow rates.

164 The hydrodynamic conditions of both ROIs are summarized in Table 1. The Reynolds
 165 number $Re = HU_b/\nu$ and the Froude number $Fr = U_b/\sqrt{gH}$ are based on the water depth
 166 H and the bulk flow velocity U_b , both spatially averaged over the respective ROIs. Here, U_b
 167 $= Q/A$ is calculated from the cross-sectional area A inferred from the bathymetry, g is the
 168 gravitational acceleration, and ν is the fluid kinematic viscosity. Despite the meander being
 169 shallower and associated with a larger Fr , the pool displays a wavier surface (Figure 2c-d)
 170 which is attributed to the turbulence induced by the rocky bed of the riffle upstream of
 171 this region (Brocchini & Peregrine, 2001). In both ROIs $Fr \ll 1$, and indeed the ultrasonic
 172 probe data indicates limited deformation of the free surface: the root mean square (RMS)
 173 fluctuations of the water surface level are approximately 1 mm and 2 mm in the meander and
 174 the pool, respectively, which provide an estimate of the wave amplitude a . Instantaneous
 175 images (acquired as described below) indicate wavelengths λ of 3 to 6 cm in the meander
 176 and 4 to 8 cm in the pool. To obtain first-order estimates of the wave effect on the floating

Table 1. Main hydrodynamic parameters of the meander and the pool for both flow rates Q : mean depth of the channel H , mean width of the channel B , mean cross-sectional area A , bulk fluid velocity U_b , Reynolds number Re , and Froude number Fr .

Meander	H [m]	B [m]	A [m ²]	U_b [m s ⁻¹]	Re	Fr
$Q_1 = 32.1 \text{ L s}^{-1}$	0.08	1.72	0.143	0.225	18,480	0.25
$Q_2 = 53.7 \text{ L s}^{-1}$	0.10	—	0.177	0.303	30,910	0.30
Pool	H [m]	B [m]	A [m ²]	U_b [m s ⁻¹]	Re	Fr
$Q_1 = 32.1 \text{ L s}^{-1}$	0.29	1.68	0.492	0.065	18,977	0.04
$Q_2 = 53.7 \text{ L s}^{-1}$	0.31	—	0.525	0.102	31,750	0.06

Table 2. Main properties of the floating particles: material, particle-to-fluid density ratio ρ_p/ρ , equatorial radius a , polar radius c , and aspect ratio λ .

Particle Type	Material	ρ_p/ρ	a [mm]	c [mm]	λ
Spheres	Polypropylene	0.9	2.5	2.5	1.00
Discs	Softwood	0.75	19.1	1.6	0.08
Rods	Softwood	0.75	0.9	31.8	35.28

177 particle transport, we use relations for monochromatic surface waves (Lighthill, 2001). The
 178 maximum horizontal velocity of a floating particle due to the wave field is the maximum
 179 orbital velocity Sc_p , where $S = ak$ is the wave slope, $k = 2\pi/\lambda$ is the wavenumber and
 180 $c_p = \sqrt{gk}$ is the deep-water phase velocity. This yields around 0.03 m s^{-1} and 0.05 m s^{-1}
 181 for the meander and pool, respectively, which are small compared to the measured free surface
 182 velocities. Consistent with this estimate, Del Grosso et al. (2019) reported RMS free surface
 183 velocities induced by gravity-capillary waves of a few cm s^{-1} , but for waves with much larger
 184 amplitude and similar slope. In conclusion, while the wave-induced surface motion may
 185 participate to the transport, it is not expected to majorly affect our conclusions.

186 2.2 Floating Particles

187 Three types of floating particles are used in the present experiments. White polypropy-
 188 lene beanbag filler pellets, approximately spherical with a 5 mm diameter, are used to char-
 189 acterize the surface flow velocity. These are sufficiently large to be accurately detected by
 190 imaging and can be recaptured downstream of the ROIs. To explore the effect of shape and
 191 size on particle transport, larger centimetre-sized discs and rods are utilized. The discs con-
 192 sist of wooden craft circles and the rods are wooden toothpicks, both spray-painted white
 193 to increase their visibility and to reduce the absorption of water.

194 When describing non-spherical particles in turbulence such as discs and rods, it is
 195 common to idealize their shape as spheroids. Any spheroid can be specified by its aspect
 196 ratio $\lambda = c/a$, defined as the ratio between the polar radius c (i.e., the length of the semi-axis
 197 perpendicular to the plane of symmetry) and the equatorial radius a (i.e., the length of the
 198 semi-axis along the plane of symmetry): $\lambda = 1$ is a sphere, $\lambda < 1$ is an oblate spheroid (disc),
 199 and $\lambda > 1$ is a prolate spheroid (rod). The different particle properties are summarized in
 200 Table 2.

2.3 Particle Response Time

When assessing the ability of particles to follow the fluid velocity fluctuations it is customary to quantify the Stokes number $St = \tau_p/\tau_f$, where τ_p is the particle response time and τ_f is a relevant time scale of the flow. The latter is usually taken as the Kolmogorov time scale τ_η (the time scale of the smallest turbulent eddies), due to the significance of the particle interaction with the microscale structure of the turbulence (Wang & Maxey, 1993; Balachandar & Eaton, 2010; Brandt & Coletti, 2022). While the dynamics of free surface turbulence is not fully understood and the applicability of Kolmogorov theory is debated (Hunt & Graham, 1978; Magnaudet, 2003), experimental and numerical studies have documented a $k^{-5/3}$ scaling of the energy spectra at or near the free surface (Chickadel et al., 2011; Flores et al., 2017). As we will show, both the near-surface ADV measurements and the PTV measurements confirm such a scaling in the present setting, allowing us to estimate τ_η (see Section 3.2). Evaluating τ_p of floating particles, however, is especially challenging. This is usually defined as the characteristic time over which a particle responds to changes in the surrounding fluid velocity through the drag force. The latter depends on the level of submergence (Beron-Vera et al., 2019), which is not accurately known for particles floating in turbulent flows and cannot be accurately measured here. Alternatively, τ_p can be defined as the integral, over time, of the particle acceleration autocorrelation. However, measuring this reliably requires a spatio-temporal resolution hardly achievable in a large-scale outdoor setting, and beyond the capability of the present imaging system. Therefore, in the following, leveraging previous studies of finite-size particles in turbulence, we opt for an estimate of St based on the size of the particles compared to the Kolmogorov scale η (the size of the smallest free surface turbulent eddies, see Section 3.2).

Laboratory experiments from Fiabane et al. (2012) and particle-resolved simulations from Homann and Bec (2010) and Uhlmann and Chouippe (2017) indicate that spherical particles in turbulence behave as tracers up to $d_p \sim 5\eta$; while larger particles have a Stokes number that approximately scales as $St = \tau_p/\tau_\eta \sim 1 + 0.08d_p/\eta$. Homann and Bec (2010) argued for a power-law dependence $\propto (d_p/\eta)^{2/3}$, but the quantitative outcome is similar. As we will see, $\eta \approx 0.5$ mm in the meander, thus, $d_p/\eta \sim 10$ (i.e., $St \sim 2$) for the spheres. While the spheres may not respond faithfully to the smallest-scale fluctuations, they are expected to capture most of the turbulent kinetic energy, to first-order accuracy, being 30 to 50 times smaller than the energy-containing turbulent eddies of size L . As a result, we will regard them as tracers of large-scale motions. This is consistent with Nikora et al. (2007), where 3 mm floating particles were deemed suitable tracers for free surface turbulence in a laboratory flume. For non-spherical particles, the estimation of τ_p is even more complex due to their geometry. Considering the length of maximum extension, $2a$ for the discs and $2c$ for the rods, we estimate $St \approx 9$ and 14, respectively. The elongated shape of the rods, however, suggests that alternative measures of their effective size (e.g., the volume-equivalent diameter) may be more suitable. In general, the larger particles are expected to have significantly longer response times than the mm-sized spheres.

In the pool, we lack precise estimates of the Kolmogorov scales; as the turbulence intensity and so the dissipation rate is higher, η is expected to be somewhat smaller, hence St may be accordingly larger yet comparable to the levels in the meander. As we will discuss, the observed behaviour of the larger particles is consistent with such estimates. We shall remark that St is not a sufficient parameter to characterize the behaviour of finite-sized particles in turbulence (Lucci et al., 2010). The present estimates are solely meant to guide in the later phenomenological interpretation of the results.

2.4 Particle Imaging and Tracking

A 1 Mpx CMOS camera (Allied Vision Mako U-130B) with a 3 mm wide-angle lens is mounted on a cantilever arm attached to a traversing system composed of aluminum beams (Figure 1b). The camera is suspended 1.5 m above the water surface, imaging a $2.2 \text{ m} \times 1.7 \text{ m}$

252 field of view (FOV). As it will be shown, for both considered locations, this is much larger
 253 than the integral scale of turbulence L . To minimize reflections on the water surface, a large
 254 tent is set up to enclose the camera and the FOV, blocking direct sunlight that would cause
 255 reflections and any wind that may affect the free surface.

256 The particles are dispensed using a bin spanning the width of the channel, ensuring a
 257 nearly homogeneous particle distribution, and retrieved via a nylon seine net at the down-
 258 stream end of the stream. The camera records at a frame rate of 30 to 50 Hz depending on
 259 the ROI and Q , keeping the inter-frame particle displacement to about 6 pixels. For each
 260 flow rate case and each ROI, measurements are performed over four separate runs to prevent
 261 the net from filling with particles and obstructing the water flow. Each run contains about
 262 15 000 to 20 000 images. In total, this yields approximately 16 000 particle trajectories for
 263 the spheres and 1000 trajectories for the discs and rods. We verify that each of the four
 264 runs per particle type yields the same quantitative results for each flow rate, thus statistical
 265 uncertainty due to finite sample size does not affect the conclusions.

266 The wide-angle camera lens introduces some image distortion. To correct it, a $0.9\text{ m} \times$
 267 1.2 m checkerboard pattern is imaged at the same distance as the water surface, and the
 268 appropriate de-warping transform is determined (Z. Zhang, 2000). Despite the tent blocking
 269 direct sunlight, some glare off the water surface from the diffused ambient light is still
 270 present. This time-dependent background noise is removed using the proper orthogonal
 271 decomposition (POD)-based method by Mendez et al. (2017), which isolates the modes
 272 mostly contributing to the intensity variance of the images. We subtract the first two
 273 modes, which successfully removes most of the glare while preserving the particles in the
 274 images.

275 Particles are identified by employing threshold-based image segmentation (i.e., finding
 276 continuous groups of pixels exceeding an intensity threshold). The probability distribution
 277 function (*p.d.f.*) of the areas of these groups of pixels is considered, and a rejection criterion
 278 is set at ± 2 standard deviations from the expected value based on the pixel/mm ratio. Par-
 279 ticle trajectories $\vec{x}_p(t)$ are formed using a custom-written nearest-neighbour PTV algorithm
 280 (Baker & Coletti, 2019, 2021, 2022), and their velocities $\vec{u}(\vec{x}_p(t))$ and accelerations $\vec{a}(\vec{x}_p(t))$
 281 are obtained by convolution with the first and second derivative of a Gaussian kernel in the
 282 time domain, respectively. A temporal kernel $t_k = 16$ frames is chosen as the smallest value
 283 beyond which the total acceleration variance σ_a^2 decays exponentially (Figure 3a). This
 284 approach has been used in several previous laboratory and field studies (Voth et al., 2002;
 285 Nemes et al., 2017; Li et al., 2022; Berk & Coletti, 2021; Baker & Coletti, 2021, 2022). We
 286 also characterize the rods' orientation and rotation rate along their trajectory. The orienta-
 287 tion is defined by the unit vector \hat{p} aligned with the rod's symmetry axis, obtained from an
 288 ellipse best-fit to the valid pixel groups. The angular velocity $\Omega(t)$ is obtained by convolving
 289 $\hat{p}(t)$ with the first derivative of a Gaussian kernel, analogous to the particle velocity using
 290 the same t_k .

291 2.5 Laboratory Water Channel

292 To complement the study in the OSL, laboratory experiments on particles floating in
 293 turbulent water are conducted at ETH Zürich (Figure 4). A recirculating open channel is
 294 used, with a $0.5\text{ m} \times 0.5\text{ m} \times 2.0\text{ m}$ test section. The water depth is 0.4 m and a bulk velocity
 295 $U_b = 0.24\text{ m s}^{-1}$ is imposed by a centrifugal pump. Turbulence is generated by a square-
 296 mesh grid inserted at the inlet of the test section, with mesh size $M = 35\text{ mm}$ yielding a
 297 Reynolds number $\text{Re}_M = MU_b/\nu = 8400$. With a Froude number $\text{Fr} = 0.12$, the water
 298 surface is weakly deformed by waves with amplitude smaller than 0.5 mm. To characterize
 299 the free surface turbulence 2 mm polyethylene spheres (Cospheric LLC) are used. These are
 300 around 5 times larger than η thus sufficiently small to be considered effective tracers. To
 301 investigate the effect of particle size, discs of 5 mm and 10 mm in diameter, laser cut out of
 302 1 mm thick polypropylene sheets, are utilized. A nylon net is placed at the outlet of the test

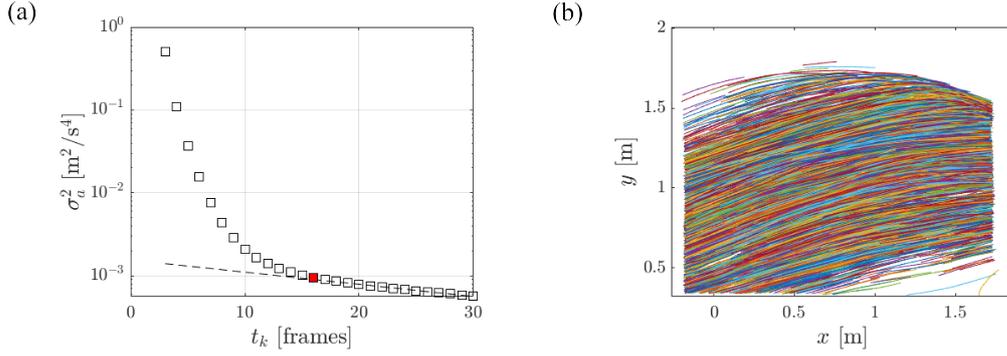


Figure 3. (a) Total particle acceleration variance plotted against the Gaussian smoothing kernel size for the spheres in the meander for Q_1 . The filled data point corresponds to the chosen kernel size for this data set and the dashed line represents the exponential decay of the acceleration variance. (b) 1% of the respective smoothed particle trajectories, drawn with different colours for visualization purposes.

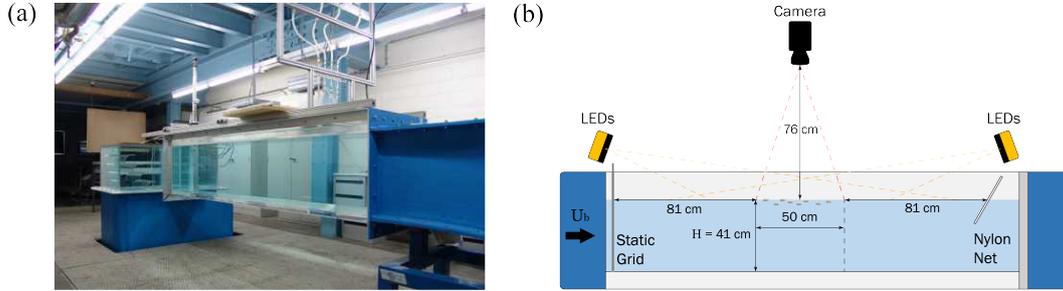


Figure 4. (a) The large recirculating water channel operated at ETH Zürich. (b) The experimental imaging setup.

303 section to recapture the particles. A $0.45 \text{ m} \times 0.5 \text{ m}$ FOV is imaged via a 12 Mpx CMOS
 304 camera (Baumer VQXT-120C.HS) operated at 90 Hz with a 35 mm lens. The upstream edge
 305 of the FOV is located 0.81 m from the grid, which is sufficient for the turbulence to have
 306 reached equilibrium conditions (Hearst & Lavoie, 2014). The particles are illuminated by
 307 a pair of continuous LED lights. Their centroids are obtained via threshold-based image
 308 segmentation followed by a circle-finder routine. The trajectories are reconstructed using
 309 the same PTV algorithm as described above for the outdoor stream measurements. At least
 310 10 000 trajectories per particle type are acquired.

311 3 Results and discussion

312 3.1 ADV Measurements

313 We first consider the ADV measurements to assess the near- and sub-surface flow in
 314 the ROIs. Figure 5 shows the temporal mean and RMS fluctuations of the streamwise
 315 velocities for the meander (a-b) and the pool (c-d). Measurements are shown for one flow
 316 rate in each ROI, the trends being analogous for both considered flow rates. In the meander,
 317 horizontal near-surface velocity averaged over the ROI is $\langle \vec{u} \rangle = [0.30, 0.07] \text{ m s}^{-1}$ and $\langle \vec{u} \rangle =$
 318 $[0.34, 0.05] \text{ m s}^{-1}$ for Q_1 and Q_2 , respectively. The nonzero spanwise velocity is expected for
 319 a curved section of a natural stream. Conversely, the pool displays weak spanwise velocity

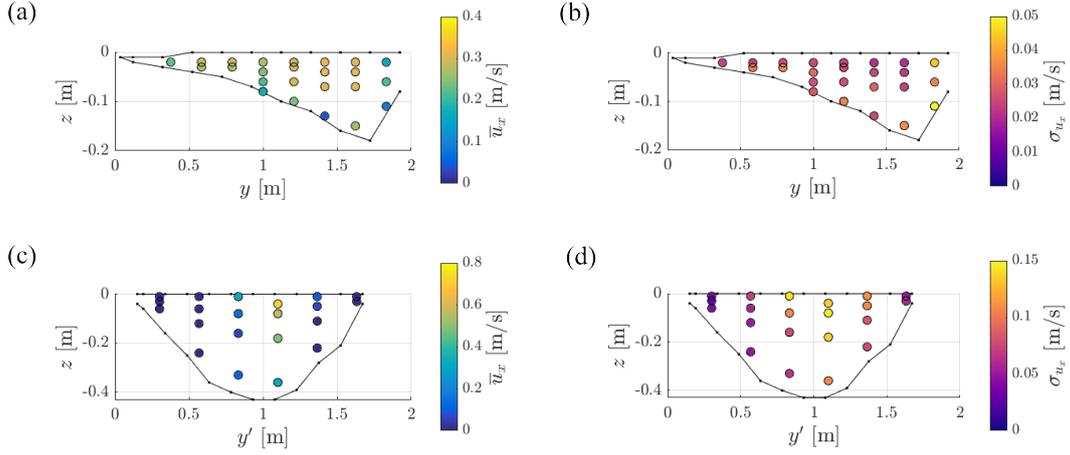


Figure 5. The temporal mean (a) and RMS fluctuations (b) of the streamwise velocity in the meander, measured by ADV, at various points along the cross-section at $x = 1$ m for Q_1 . The temporal mean (c) and RMS fluctuations (d) of the streamwise velocity in the pool, measured by ADV, at various points along the cross-section at $x' = 1$ m for Q_2 .

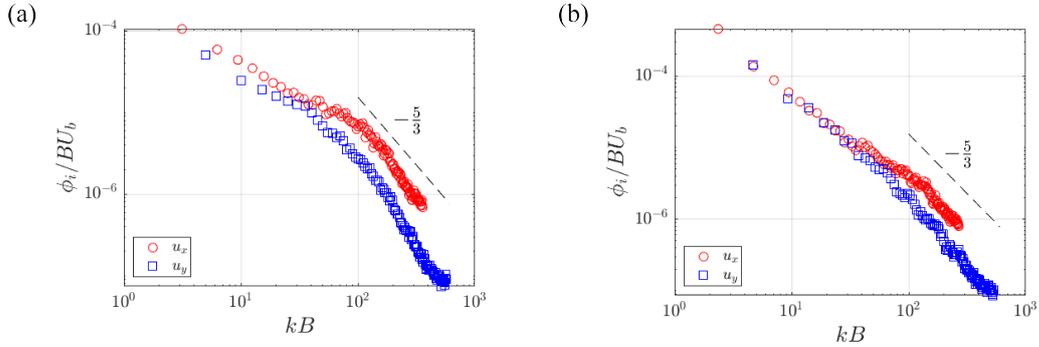


Figure 6. Normalized power spectral densities of the near-surface ADV measurements taken in the meander at $\vec{x} = [1, 1, -0.02]$ m for Q_1 (a) and Q_2 (b). The dashed line corresponds to $k^{-5/3}$ scaling.

320 and high streamwise velocity along the midline. This indicates a jet-like flow structure,
 321 bounded by shear layers which are associated with high-velocity fluctuations. This view
 322 will be confirmed by the free surface flow imaging. Furthermore, the flow in the meander
 323 displays significant turbulence intensity throughout the water column, with streamwise RMS
 324 fluctuations exceeding 10% of U_b . Figure 6 shows the normalized power spectral density of
 325 the near-surface velocity fluctuations measured 2 cm below the water surface. We recover
 326 the classic $k^{-5/3}$ scaling for the streamwise and spanwise components of the free surface
 327 velocity over a sizeable range of wavenumbers.

3.2 Free Surface Flow Characterization

328
 329 We consider the Eulerian fields of the mean velocity $\tilde{U}(\vec{x})$, and the RMS fluctuations
 330 $\sigma_U(\vec{x})$, where $U = \sqrt{\vec{u}(t) \cdot \vec{u}(t)}$ is the norm of the particle velocity vector. The Eulerian
 331 data is obtained by binning the trajectories into fixed interrogation windows of 5 cm \times 5 cm.
 332 This allows for a temporal averaging of at least 25 instantaneous vectors in each window and
 333 is indicated by $\tilde{\cdot}$. The results for both measurement locations for Q_1 are shown in Figure

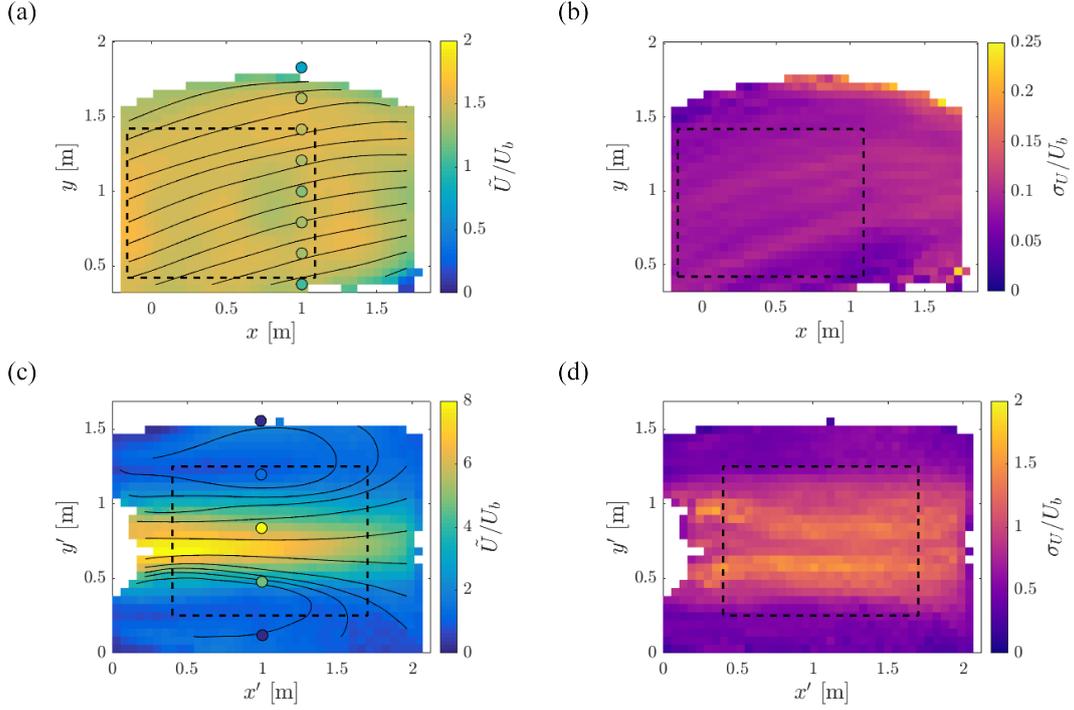


Figure 7. Eulerian mean velocity (a-c) and RMS velocity fluctuation (b-d) fields of the tracers for Q_1 , normalized by the bulk velocity; meander (a-b) and pool (c-d). The black lines indicate streamlines and the dashed boxes indicate the sub-regions where Lagrangian quantities are evaluated. The coloured circles (a-c) correspond to the near-surface ADV measurements of \bar{u}_x along the cross-section $x = x' = 1$.

334 7. Also displayed are the near-surface (2 cm depth) ADV measurements. These reasonably
 335 agree with the Eulerian fields obtained by PTV, except for the regions near the shallow
 336 banks. As anticipated, the meander displays a remarkably homogeneous surface flow. In
 337 particular, we define a 1.25×1 m sub-region in it (highlighted in the figure) where \tilde{U} and
 338 σ_U remain within $\pm 2.5\%$ and 9.3% of their respective spatial mean and the streamlines
 339 are relatively straight. In this sub-region we investigate unbiased single-point and two-
 340 point flow statistics, characterizing the spatio-temporal flow scales, using the framework
 341 of homogeneous turbulence (presented in the next section), and examine the Lagrangian
 342 particle transport. On the other hand, the jet-like flow structure in the pool is clearly visible
 343 with two shear layers associated with large velocity fluctuations and flanked by recirculation
 344 zones (Figure 7c). Because of the significant spatial inhomogeneity, the scales of the free
 345 surface turbulence in the pool are not carried out, as this would require spatial averaging
 346 and the evaluation of velocity fluctuations around a local mean. The Lagrangian particle
 347 transport in this ROI is quantified in a 1.1×1 m sub-region. For both ROIs, the choice of
 348 the sub-region avoids statistics being strongly influenced by the proximity of the banks and
 349 reduces potential bias from short trajectories as the particles exit the FOV.

350 3.3 Free Surface Turbulence in the Meander

351 For statistical analysis of the free surface turbulence, we are particularly interested in
 352 the instantaneous velocity fluctuations. For this purpose, the particle velocity fluctuations
 353 $u'_i(t)$ are calculated by subtracting from the measured velocity $\vec{u}(t)$ the global mean $\langle \vec{u} \rangle$,
 354 hence $u_i(t) = \langle u_i \rangle + u'_i(t)$, known as Reynolds decomposition. The mean flow velocity vector

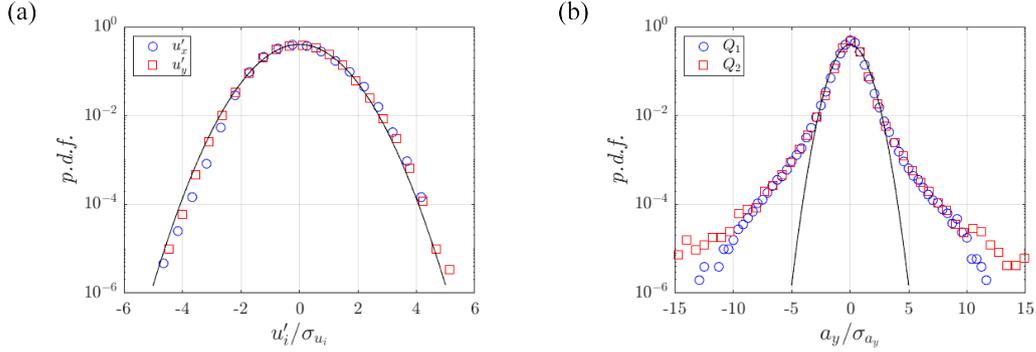


Figure 8. (a) Streamwise and spanwise velocity fluctuation *p.d.f.* of the spheres in the meander for Q_2 . (b) Spanwise acceleration *p.d.f.* for both flow rates. The distributions are normalized by their respective RMS quantities. The continuous line represents the normalized Gaussian distribution.

355 $\langle \vec{u} \rangle$ is evaluated by ensemble averaging the free surface velocity vector $\vec{u}(t)$ of the spheres
 356 obtained by PTV and the subscript i specifies the 2D velocity component. Figure 8 displays
 357 the *p.d.f.* of the streamwise (u'_x) and spanwise (u'_y) velocity fluctuations for Q_2 , as well as
 358 the *p.d.f.* of the spanwise accelerations (a_y) for both flow rates; all quantities are normal-
 359 ized by their respective RMS quantities. Both components of the velocity fluctuations are
 360 normally distributed. Conversely, the acceleration *p.d.f.* possesses long exponential tails, in-
 361 dicating strong intermittency (i.e., a relatively large probability of extreme events occurring,
 362 especially for the higher Reynolds number). This behaviour of Lagrangian accelerations has
 363 been well documented in 3D turbulence (Voth et al., 2002; Mordant et al., 2004; Toschi &
 364 Bodenschatz, 2009). While the kurtosis of the velocity fluctuations approximately equals
 365 the Gaussian value of 3, the acceleration kurtosis is 8.1 and 15.9 for Q_1 and Q_2 , respectively.
 366 These levels of intermittency are typical of fully developed 3D turbulence (Voth et al., 2002;
 367 Ishihara et al., 2007).

368 To characterize how the turbulent energy is distributed across the scales of the flow, we
 369 consider the Eulerian second-order velocity structure function $S_2^E(\vec{r})$ (Kolmogorov, 1941;
 370 Pope, 2000). This is defined as the second moment of the velocity difference $\delta^E \vec{u}(\vec{r}) =$
 371 $\vec{u}(\vec{x}, t) - \vec{u}(\vec{x} + \vec{r}, t)$, where $\vec{u}(\vec{x}, t)$ and $\vec{u}(\vec{x} + \vec{r}, t)$ are the velocities of two particles separated
 372 by a distance \vec{r} at a given time t

$$S_2^E(\vec{r}) = \langle \delta^E \vec{u}(\vec{r})^2 \rangle \quad (1)$$

373 Leveraging spatial homogeneity and isotropy, we ensemble-average over all particle pairs
 374 at a distance $r = \sqrt{\vec{r} \cdot \vec{r}}$. The ensemble-averaging requires binning the data over ranges of
 375 separation $r \pm \Delta r$, where we take $\Delta r = 1$ mm as a trade-off between resolution in scale-space
 376 and statistical convergence. Here we focus on the longitudinal structure function, in which
 377 the velocity component parallel to the separation vector \vec{r} is considered. Figure 9a shows
 378 that this exhibits an approximate $r^{2/3}$ scaling over separations from about 3 cm to 10 cm
 379 which is equivalent to the $k^{-5/3}$ scaling of the velocity spectra (Figure 6). This suggests the
 380 validity of the Kolmogorov (1941) ansatz in the inertial sub-range, hence

$$S_2^E(r) = C_2(\epsilon r)^{2/3} \quad (2)$$

381 where ϵ is the dissipation rate of the turbulent kinetic energy, and C_2 is a constant. Fur-
 382 thermore, Flores et al. (2017) report that even though the mechanism underlying the $k^{-5/3}$

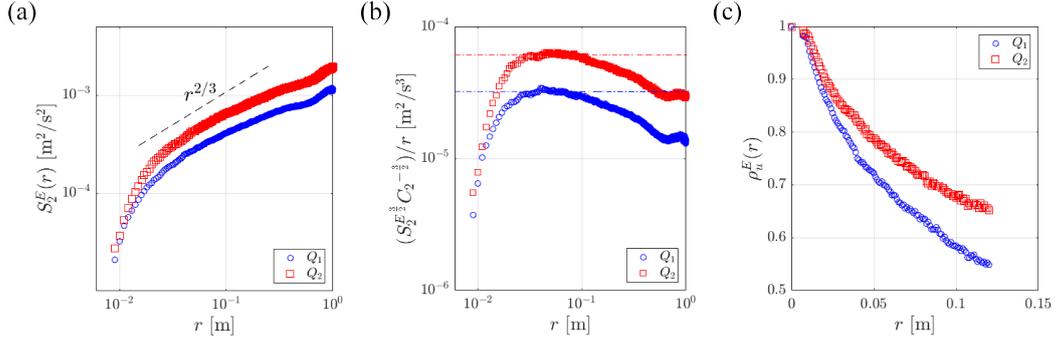


Figure 9. (a) Eulerian longitudinal second-order structure function, (b) compensated structure function and (c) Eulerian velocity autocorrelation function of the tracers for both flow rates. The dashed line in (a) corresponds to $r^{2/3}$ scaling. The dashed-dotted horizontal lines in (b) show the plateau of the compensated structure function which corresponds to the turbulent dissipation rate.

383 spectral slope at the near-surface may differ from the 3D turbulence dynamics in the bulk,
 384 the proportionality constants are roughly the same. Therefore, we assume $C_2 = 2.1$ as in
 385 3D turbulence (Pope, 2000; Saddoughi & Veeravalli, 1994) and use Equation 2 to estimate
 386 ϵ from the plateau of the compensated structure functions in Figure 9b. We then estimate
 387 the dissipative scales of the free surface turbulence: the Kolmogorov length and time scale,
 388 respectively

$$\eta = \left(\frac{\nu^3}{\epsilon} \right)^{1/4} \quad (3)$$

$$\tau_\eta = \left(\frac{\nu}{\epsilon} \right)^{1/2} \quad (4)$$

389 To determine the integral scales of the free surface turbulence, we make use of the
 390 Eulerian velocity autocorrelation function, which for homogeneous turbulence can be easily
 391 derived from the second-order structure function

$$\rho_i^E(r) = \frac{\langle u'_i(\vec{x}, t) u'_i(\vec{x} + \vec{r}, t) \rangle}{\sigma_i^2} = 1 - \frac{S_2^E(r)}{2\sigma_i^2} \quad (5)$$

392 where σ_i^2 is the variance of u'_i . The obtained velocity correlation exhibits an approximately
 393 exponential decay (Figure 9c), and the integral length scale L is evaluated by least-square
 394 fitting to it a function $Ae^{-r/L}$ where A is a constant of order unity. The estimates for
 395 the dissipative and integral scales, summarized in Table 3, support the notion that the $r^{2/3}$
 396 scaling of the structure function applies over an inertial sub-range $\eta \ll r \ll L$. Additionally,
 397 an alternative estimate of the dissipation rate can be obtained from the classic scaling
 398 (Tennekes & Lumley, 1972)

$$\epsilon \approx C \frac{\sigma_u^3}{L} \quad (6)$$

399 Taking the typical proportionality constant $C = 0.5$ as for 3D turbulence in the high-
 400 Reynolds number limit (Burattini et al., 2005; Carter et al., 2016), we obtain dissipation
 401 estimates consistent with those found from the second-order structure function.

Table 3. Main physical quantities characterizing the free surface turbulence for both flow rates in the Meander: RMS of the velocity fluctuations σ_u , dissipation rate of turbulent kinetic energy ϵ , integral length scales L , integral time scale T_L , Kolmogorov length scale η , and Kolmogorov time scale τ_η .

Meander	σ_u [m s^{-1}]	ϵ [$\text{m}^2 \text{s}^{-3}$]	L [m]	T_L [s]	η [mm]	τ_η [s]
$Q_1 = 32.1 \text{ L s}^{-1}$	0.022	$3.2 \cdot 10^{-5}$	0.175	1.02	0.4	0.18
$Q_2 = 53.7 \text{ L s}^{-1}$	0.032	$6.1 \cdot 10^{-5}$	0.243	1.01	0.4	0.13

3.4 Effect of Particle Shape and Size in the Meander

In this section, we compare the motion of the larger particles (discs and rods) against the spheres. We start by considering the meander where the flow homogeneity allows for a comprehensive statistical description of the transport.

The Eulerian velocity fields of all particle types are found to be quantitatively similar. This is evident from Figure 10a-b, where the velocities of the larger particles are normalized by those of the spheres. For both considered flow rates, the RMS difference between the three particle types and the near-surface ADV measurements is less than 2% of \tilde{U} and less than 17% of σ_U . Also displayed are the particle velocity and acceleration *p.d.f.* for selected components and flow rates (Figure 10c-d). To highlight the difference between the different particle types, the Kolmogorov velocity scale $u_\eta = \eta/\tau_\eta$ and acceleration scale $a_\eta = u_\eta/\tau_\eta$ are used for normalization. The velocity fluctuations are similar between all particle types, closely approximating a Gaussian distribution (Figure 10c). Contrarily, the acceleration intermittency shown by the spheres is significantly reduced for the larger particles (Figure 10d). Moreover, at Q_1 , the RMS acceleration of the discs and rods is 9% and 21% lower than that of the tracers, respectively, while at Q_2 the reduction becomes 8% and 20%, respectively.

To characterize the spreading rate of the floating particles, we consider their Lagrangian motion characterized by single-particle dispersion; examining how far, on average, a single particle migrates from its origin over time. Leveraging the homogeneity of the flow in the meander and following the classic framework of Taylor (1921), the single-particle diffusivity can be derived from the Lagrangian velocity autocorrelation

$$\rho_u^L(\tau) = \left\langle \frac{\sum \vec{u}'(t) \cdot \vec{u}'(t + \tau)}{\sum \vec{u}'(t)^2} \right\rangle \quad (7)$$

Here, the summation extends to all values of τ along each trajectory (i.e., the autocorrelation is first calculated along each trajectory and normalized by its velocity variance, before ensemble-averaging over all trajectories). This ensures that each trajectory has the same weight when contributing to the global autocorrelation coefficient (Guala et al., 2007). Additionally, we only consider trajectories whose duration is longer than the time delay τ (Mordant et al., 2004). Figure 11a-b display the Lagrangian velocity autocorrelation of each particle type for both considered flow rates, showing that the motion of the discs and rods is more time-correlated than that of the spheres. This is consistent with the trend reported by numerical simulations of inertial particles (Squires & Eaton, 1991; Jung et al., 2008) and laboratory observations of finite-size particles (Machicoane & Volk, 2016) in 3D turbulence.

The diffusivity K is obtained by integrating the decaying Lagrangian velocity autocorrelation (Taylor, 1921)

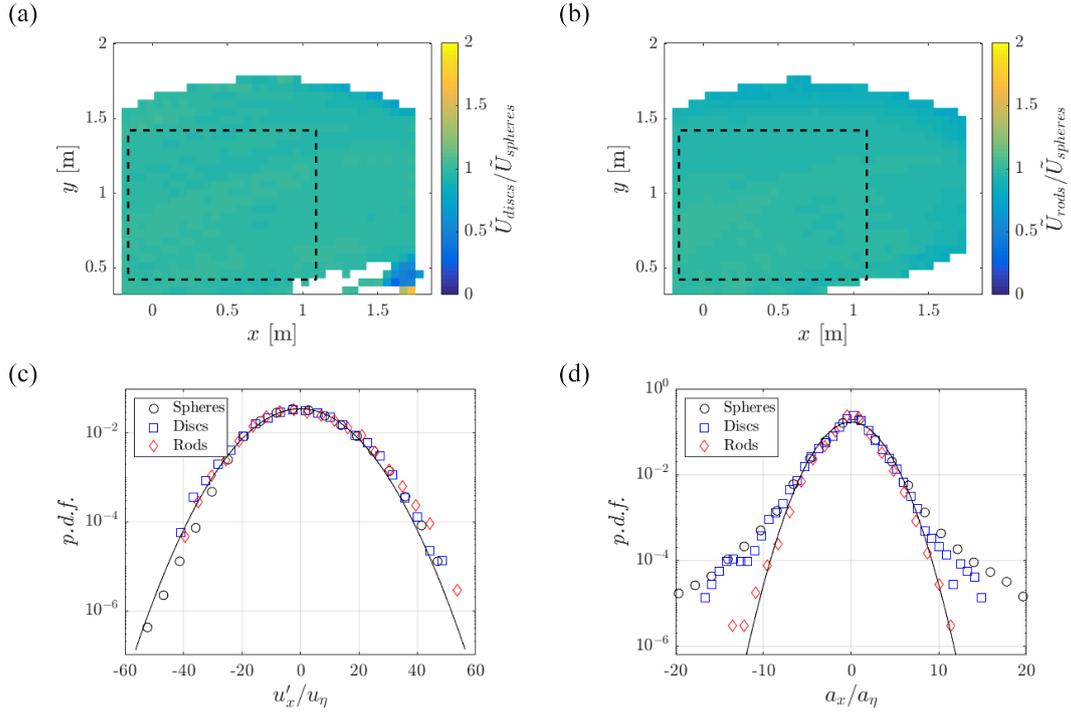


Figure 10. Eulerian mean velocity fields of the discs (a) and rods (b) for Q_1 . (c) Streamwise velocity fluctuation $p.d.f.$ of the different particle types in the meander for Q_2 . (d) Streamwise acceleration $p.d.f.$ of the different particle types in the same location for Q_1 . The distributions are normalized by Kolmogorov scaling. The continuous line represents the normalized Gaussian distribution.

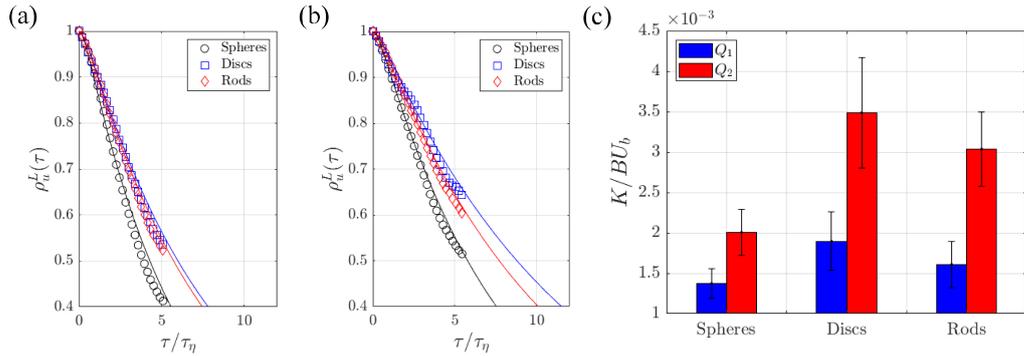


Figure 11. Lagrangian velocity autocorrelation function of each particle type for Q_1 (a) and Q_2 (b). The solid lines are the autocorrelation functions computed using Equation 9 which are integrated to obtain diffusion coefficients. (c) Normalized diffusivity of the different particle types for both flow rates. The error bars represent the standard deviation of the diffusion coefficients from separate runs.

$$K = \sigma_u^2 \int_0^\infty \rho_u^L(\tau) d\tau \quad (8)$$

436 As the extreme of integration grows, the autocorrelation is expected to decay to negligibly
 437 small values and correspondingly the diffusivity will asymptote to a value independent of
 438 time. Due to the finite length of the recorded trajectories, we extrapolate the autocorrelation
 439 using the stochastic model proposed by Sawford (1991)

$$\rho_u^L(\tau) = \frac{T_L e^{-\tau/T_L} - T_2 e^{-\tau/T_2}}{T_L - T_2} \quad (9)$$

440 Here, two time scales are required: the integral time scale of the turbulence T_L , and a
 441 characteristic time scale related to the dissipation T_2 . The former is defined as the charac-
 442 teristic decay time of the Lagrangian velocity autocorrelation function of the spheres and
 443 is estimated by least-square fitting $\rho_u^L(\tau)$ to an exponential function of the form e^{-t/T_L} ,
 444 and reported in Table 3. The value of T_2 is estimated by fitting the experimental curve to
 445 Equation 9 and found to be approximately $0.3\tau_\eta$; this is the same order of magnitude as
 446 in 3D turbulence studies (Voth et al., 2002; Mordant et al., 2004). The diffusivity is then
 447 determined by the long-time asymptote of K using Equations 8 and 9. For the spheres we
 448 obtain normalized diffusivities $K/u_\tau d_p \approx 0.5$ for both flow rates, where we estimate the
 449 friction velocity u_τ from its relationship with the dissipation rate, $\epsilon = u_\tau^3/d_p$ (assumed
 450 to be mainly driven by bed friction (Raymond et al., 2012)). This falls well in the range
 451 $K/u_\tau d_p = 0.3$ to 0.9 reported for meandering channels (Fischer et al., 1979; Rutherford,
 452 1994). The diffusivity is plotted in Figure 11c for the different particle types and for both
 453 considered flow rates. One clearly sees an increase in K with increasing flow rate, hence
 454 with Reynolds number. Most importantly, the larger particles exhibit larger diffusivity than
 455 the spheres, with the discs spreading faster than the rods. We remark that the extrapola-
 456 tion using Sawford (1991) model adds quantitative uncertainty to the estimated value of K .
 457 Therefore below, we report measures of the Lagrangian transport that are not affected by
 458 such extrapolation.

459 Next, we consider the mean square displacement (MSD) of recorded PTV trajectories
 460 due to turbulent fluctuations

$$\langle X(t)^2 \rangle = \langle \|\vec{x}_p(t) - \vec{x}_p(t_0) - \langle \vec{u} \rangle \Delta t\|^2 \rangle \quad (10)$$

461 where $\vec{x}_p(t)$ is the particle position at time t and $\vec{x}_p(t_0)$ is the reference position at the
 462 temporal origin of the trajectory t_0 . The advective displacement $\langle \vec{u} \rangle \Delta t$, due to the mean
 463 flow during the time interval $\Delta t = t - t_0$, is subtracted to isolate the contribution of the
 464 turbulent fluctuations. Leveraging spatial homogeneity, the advective flow is taken to be a
 465 uniform motion, which avoids the ambiguities associated with subtracting different advective
 466 displacements at different points along the same trajectory. The MSD of each particle type
 467 for both flow rates is plotted in Figure 12a-b and confirms that the discs spread faster than
 468 the rods, which spread faster than the spheres. Calculating the diffusivity from a least-
 469 square fit to the linear part of the MSD returns a value of diffusivity in agreement with
 470 those reported above. Although not shown, we also note that the MSD can alternatively be
 471 computed by integrating the autocorrelation twice (Taylor, 1921; Pope, 2000)

$$\langle X(t)^2 \rangle = 2\sigma_u^2 \int_0^t \int_0^{t'} \rho_u^L(\tau) d\tau dt' \quad (11)$$

472 where t' is a second integration variable. This yields analogous trends when compared to
 473 Equation 10.

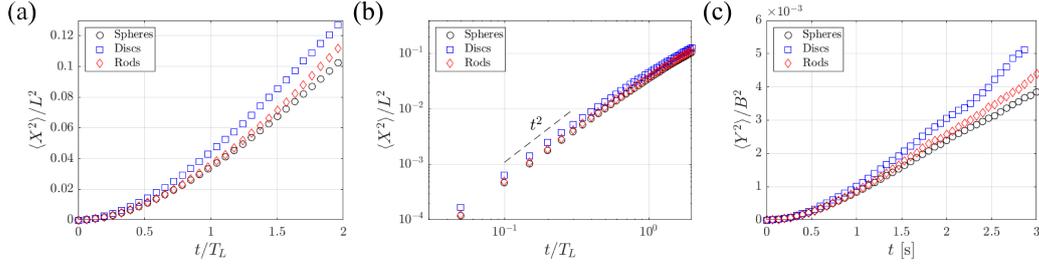


Figure 12. Normalized MSD due to turbulent velocity fluctuations of each particle type in the meander for Q_1 (a) and Q_2 (b). (c) Spanwise MSD of the different particles in the pool for Q_1 .

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3.5 Effect of Particle Shape and Size in the Pool

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In this section, we verify that the trends observed in the meander also apply to the significantly different flow conditions found in the pool. Also here, the Eulerian fields of \tilde{U} and σ_U for the discs and rods (not shown) are close to those measured for the spheres, shown in Figure 7c-d, with RMS difference between the three particle types and the near-surface ADV measurements less than 12% for \tilde{U} and less than 16% for σ_U . Nevertheless, as for the meander, we shall see that the particle shape and size influences the Lagrangian dispersion.

Because the mean velocity in the pool is predominantly aligned with x' , we can isolate the turbulent dispersion by considering the lateral displacement (i.e., the MSD of particle trajectories along the spanwise direction y')

$$\langle Y(t)^2 \rangle = \langle [y'(t) - y'(t_0)]^2 \rangle \quad (12)$$

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and is plotted in Figure 12c for Q_1 (Q_2 displaying analogous results). This indicates again that the larger particles spread faster than the spheres, with the discs spreading faster than the rods. An estimate of the lateral diffusion coefficient can be derived from the relation

$$K_{y'} = \frac{1}{2} \frac{d\langle Y(t)^2 \rangle}{dt} \quad (13)$$

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A linear least-square fit to the data over the range $t > 1.5$ s (where the MSD is approximately linear with time) yields $K_{y'} = 0.002 \text{ m}^2 \text{ s}^{-1}$, $0.003 \text{ m}^2 \text{ s}^{-1}$ and $0.0025 \text{ m}^2 \text{ s}^{-1}$ for the tracers, discs, and rods, respectively.

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3.6 Rotational dynamics

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The translational and rotational motion of anisotropic particles in turbulence are strongly coupled to each other (Voth & Soldati, 2017). Therefore, we consider the rotational dynamics of the rods, as it can provide insight into the transport behaviour presented in the previous section. We present results for Q_2 , with Q_1 showing analogous trends.

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We first consider the alignment of the rods defined by the orientation vector $\hat{p}(t)$. Figure 13a shows the *p.d.f.* of $|\hat{p}(t) \cdot \hat{u}(t)|$, where $\hat{u}(t)$ is the unit vector parallel to the particle velocity. For both ROIs, the rods display a preference to align with the direction of motion. Considering the close similarity between the velocity fields of the spheres and those of the rods, this can be interpreted as a preferential alignment with the flow direction.

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The intermittent nature of the free surface turbulence, displayed in the acceleration *p.d.f.* in Figure 8b, is also reflected in the distribution of the angular velocity Ω shown in Figure 13b. The kurtosis of these distributions are 5.9 and 8.8 for the meander and the

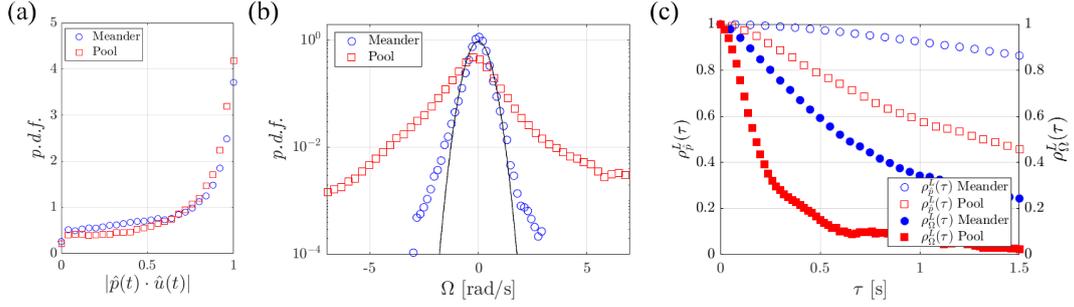


Figure 13. (a) The *p.d.f.* of the absolute value of the cosine of the orientation angle of the rods in both ROIs. (b) The *p.d.f.* of the angular velocities in both ROIs. (c) Lagrangian autocorrelation functions of the rods' orientation and angular velocity in both ROIs.

500 pool, respectively, indicating a relatively large probability of extreme events with angular
 501 velocities of several rad s^{-1} , especially with a higher turbulence intensity of the free surface.
 502 Such sudden changes in orientation are expected to alter the Lagrangian transport by the
 503 underlying flow.

504 The curvature of the streamlines in the ROIs is small, but the rods' orientation varies
 505 in time due to flow fluctuations. We characterize the time scales associated with the rods'
 506 re-orientation by the Lagrangian autocorrelation of the orientation vector $\rho_p^L(\tau)$, calculated
 507 analogously to the velocity autocorrelation function in Equation 10 and shown in Figure
 508 13c. In the meander the particle orientation is remarkably stable, which is consistent with
 509 its moderate turbulence intensity: the fluid velocity, with which the rods tend to be aligned,
 510 remains mostly oriented in the streamwise direction. The orientation autocorrelation in
 511 the pool shows a faster decay with a characteristic time of approximately 1.5 s. Given the
 512 jet-like flow structure, a candidate time scale dictating the rod reorientation is provided by
 513 the intense shear layers (Figure 7c-d). Indeed, visual observation confirms that the rods'
 514 rotation in those regions follows the direction of the mean shear. The associated time scale
 515 can be estimated from the jet half-width $d_{1/2} \approx 0.5$ m and the velocity difference across it
 516 $\Delta\tilde{U} \approx 0.4$ m s^{-1} such that $d_{1/2}/\Delta\tilde{U} \approx 1.25$ s, which approximately agrees with the observed
 517 correlation time scale. The fact that the time scale of re-orientation is attributed to the
 518 mean shear of the surface flow is consistent with the observation that the rods' orientation
 519 is very stable in the meander, where the flow is highly homogeneous and lateral shear is
 520 weak.

521 Figure 13c also shows the autocorrelation of the angular velocity $\rho_\Omega^L(\tau)$, which as ex-
 522 pected decays significantly faster than $\rho_p^L(\tau)$. For the meander, the correlation time scale of
 523 $\rho_\Omega^L(\tau)$ is approximately 1 s, matching the integral time scale of the free surface turbulence
 524 T_L . In the pool, the same quantity decays with a characteristic time scale around 0.25 s.
 525 While a single value of T_L can hardly be defined in the pool due to spatial inhomogeneity,
 526 we note that σ_U is roughly 4 times larger than in the meander. This suggests that, in both
 527 ROIs, the correlation time scale of $\rho_\Omega^L(\tau)$ is dictated by the energetic eddies that determine
 528 the integral scales of the turbulence. Since the rods' length is two orders of magnitude
 529 larger than η and a fraction of L , this finding is in line with the view that rods' rotation is
 530 controlled by eddies of size comparable to or larger than their length (Parsa & Voth, 2014;
 531 Voth & Soldati, 2017).

532 3.7 Laboratory Results

533 The analysis of the laboratory measurements is analogous to the outdoor stream study.
 534 The trajectory trajectories are first binned into $4 \text{ mm} \times 4 \text{ mm}$ interrogation windows to

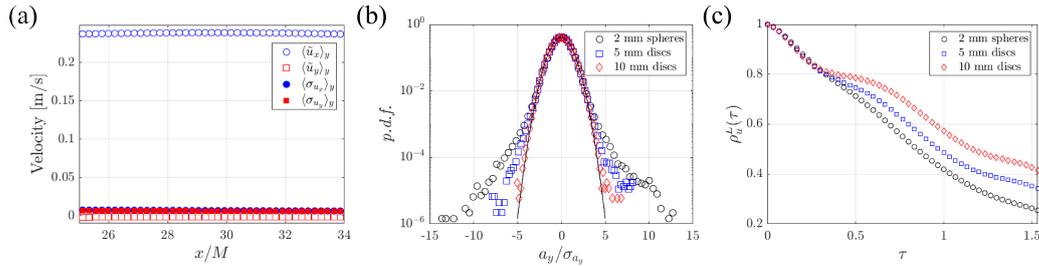


Figure 14. (a) The spanwise average of both components of the mean velocity and RMS fluctuations for the tracers. (b) Spanwise acceleration *p.d.f.* of the different particles normalized by their respective standard deviations. (c) Lagrangian velocity autocorrelation for the different particle types (c).

535 generate Eulerian mean fields. In this case, as expected, the degree of homogeneity is much
 536 higher, and therefore the presented data is obtained from an ROI that coincides with the
 537 FOV. A similar analysis to the one presented for the meander indicates that $\epsilon = 2 \cdot 10^{-5}$
 538 $\text{m}^2 \text{s}^{-3}$ on the free surface, for $\eta = 0.5 \text{ mm}$. Figure 14a displays the spanwise average of
 539 both components of the mean velocity $\langle \tilde{u}_i \rangle_y$ and RMS fluctuations $\langle \sigma_i \rangle_y$ for the tracers,
 540 indicating nearly isotropic turbulence along the free surface. Figure 14b displays the *p.d.f.*
 541 of the spanwise acceleration a_y for each particle type normalized by their respective standard
 542 deviation. As in the outdoor stream, there is an apparent reduction in intermittency for
 543 the larger particles when compared to the tracers, with the larger discs displaying almost
 544 Gaussian accelerations. Finally, Figure 14c shows the Lagrangian velocity autocorrelation
 545 for the different particle types. Clearly, the motion of the larger particles has a degree of
 546 temporal correlation that increases with particle size.

547 4 Discussion

548 The PTV measurements of the small spheres, especially in the spatially homogeneous
 549 sub-region of the meander, inform us of the nature of the free surface flow in the considered
 550 riverine environment. We remark that, for fundamental reasons, free surface turbulence
 551 is not expected to be equivalent either to 2D or 3D turbulence: the surface exchanges
 552 energy and enstrophy with the flow underneath, hence neither quantity can be regarded as
 553 invariant and dimensional scaling arguments do not strictly apply (Cressman et al., 2004).
 554 However, the present measurements do indicate a strong similarity with the phenomenology
 555 of 3D turbulence. In particular, the behaviour of the second-order structure function is
 556 consistent with Kolmogorov (1941) scaling in the inertial sub-range. While a similar scaling
 557 is also expected in the inverse-cascade range of 2D turbulence (Kraichnan, 1967), the latter
 558 framework is inconsistent with the observed intermittency of the acceleration (Boffetta &
 559 Ecke, 2012). The close agreement between the dissipation estimates from Equations 2 and
 560 9 further supports the applicability of a 3D turbulence framework. The similarity between
 561 3D and free surface turbulence is possibly due to the surface carrying the prominent imprint
 562 of sub-surface vortices connected to it. These evolve by diffusion and stretching, as vortex
 563 tilting is annihilated at the surface (Shen et al., 1999; C. Zhang et al., 1999; Shen & Yue,
 564 2001). In other words, unlike in 2D turbulence, the free surface boundary condition affects
 565 but does not suppress vortex stretching, which is essential to the energy cascade in 3D
 566 turbulence (Davidson, 2015; Carbone & Bragg, 2020; Johnson, 2020).

567 Our results are specific to a particular riverine flow configuration, therefore further
 568 studies are needed to assess the generality of the observations, especially as a function of
 569 the water depth, which is known to influence the turbulence dynamics (Nezu et al., 1994).
 570 Additionally, water depth influences the respective role of water-column turbulence and

571 bed friction in setting the dissipation rate at the surface (Raymond et al., 2012; Ulseth et
 572 al., 2019). Moreover, vortex stretching is hindered in shallow flows, which can trigger the
 573 emergence of features peculiar to 2D turbulence (Uijttewaal & Booij, 2000; Stocchino et al.,
 574 2011).

575 Our main finding is that, in both investigated ROIs, larger floating particles disperse
 576 faster than smaller tracer-like particles. This result can be interpreted based on our un-
 577 derstanding of the behaviour of inertial particles in turbulence. We remind that the term
 578 “inertial” indicates objects too heavy and/or too large to faithfully follow the fluid flow
 579 (Brandt & Coletti, 2022). Indeed, both discs and rods display weaker and less intermittent
 580 accelerations than the spheres. This behaviour is well known from the investigation of 3D
 581 turbulence laden with inertial particles and is attributed to two concurring mechanisms:
 582 preferential sampling of high-strain/low-vorticity regions, prevalent for small St , and iner-
 583 tial filtering of the small-scale/high-frequency fluctuations, prevalent for $St \gg 1$ (Bec et al.,
 584 2006; Toschi & Bodenschatz, 2009). Here we have estimated $St = \mathcal{O}(10)$ for the cm-sized
 585 particles. This supports inertial filtering as the likely cause of the observed behaviour. These
 586 relatively large particles respond to a spatial average of the fluid velocity, making them less
 587 sensitive to the smaller and faster-decaying eddies. This is consistent with the increasingly
 588 time-correlated motion of the larger particles. The trends found in the outdoor stream have
 589 also been confirmed by a dedicated laboratory study of similar regimes, indicating that the
 590 conclusions possess a degree of generality. The slower decay of the velocity autocorrelation
 591 is consistent with the simulations of Shin and Koch (2005) for rods in 3D turbulence, who
 592 found the correlation time scale T_L to increase with the rods’ length. However, in such
 593 a study, the RMS velocity fluctuations of the rods σ_u were found to decrease with their
 594 length, and the diffusivity $K = \sigma_u^2 T_L$ ultimately decreased. We remark that our estimate
 595 of St is consistent with previous studies on inertial particles in turbulence. In particular,
 596 based on the direct numerical simulations of Jung et al. (2008), particles with $St = \mathcal{O}(10)$ in
 597 homogeneous turbulence have a Lagrangian integral time scale ~ 2.0 times larger than the
 598 one of tracers (see their Figure 6a); this is consistent with our observations for the discs. In
 599 their study, however, the inertial particles also showed a significant reduction in fluctuating
 600 velocity, hence the increase of the inertial particle diffusivity was milder. In the present
 601 case, the RMS velocity fluctuations of the particles are not significantly affected by their
 602 size and shape, and thus the diffusivity follows the same trend as T_L . The fact that the
 603 particles are relatively large (as opposed to material points, as in Jung et al. (2008)) may
 604 be the cause of the difference. Also, their simulations spanned a limited range of scales,
 605 $L/\eta < 30$. In such a situation, particles with $St = \mathcal{O}(10)$ based on the Kolmogorov time
 606 scale have a response time comparable to the integral time scale of the turbulence, which
 607 may result in the significant reduction of the fluctuating energy of the particles.

608 Despite the rods’ length being almost twice the discs’ diameter, the latter disperse
 609 faster than the former. This may be due to the discs possessing a larger wetted area, thus
 610 more effective filtering of the small-scale fluctuations. However, the object shape is also
 611 likely to have a profound influence on the Lagrangian transport. The characteristic time
 612 scale of Ω and its intermittent nature indicate that the instantaneous orientation of the
 613 rods is affected by a range of turbulent scales. These may contribute to decorrelating their
 614 translational motion, which for anisotropic particles is strongly coupled with the rotational
 615 motion (Voth & Soldati, 2017). Moreover, the rods’ tendency to align with the flow direction
 616 suggests that the large scales of the turbulence (at least those larger than the rods’ length)
 617 are not isotropic and likely populated by streamwise-oriented structures. Indeed, already
 618 early studies of open channel flows highlighted the connection between near-wall bursts in
 619 the bottom-wall boundary layer and the coherent motions that transfer mass to and from
 620 the free surface (Nakagawa & Nezu, 1981; Rashidi & Banerjee, 1988). The complex bed
 621 topography of a natural channel is likely to enhance this connection by generating energetic
 622 eddies that can travel up to the surface, as indicated by the fact that bed roughness in
 623 shallow streams strongly correlates with gas transfer velocity (Ulseth et al., 2019).

Besides shape and size, other properties of floating particles may be influential towards their free surface transport; in particular, bulk density and surface characteristics. Particles of higher density and mass may be more effective in filtering small-scale turbulent fluctuations, which could further enhance their diffusivity. However, depending on the size, this effect could be counteracted by a lack of responsiveness to some of the energetic scales responsible for dispersion. Moreover, density and surface characteristics, in particular hydrophobicity, will affect the balance between surface tension and gravity, determining the submerged fraction of the floating object (Koh et al., 2009; Ji et al., 2018). In turn, submergence will determine the amount of windage, i.e., the drag exerted by the airflow on objects partly protruding out of the water (Zambianchi et al., 2014; Beron-Vera et al., 2019). Finally, while we have limited our study to sparse objects that do not significantly interact with each other, compressibility of the free surface flow is known to produce intense clustering that can bring floaters into close contact (Cressman et al., 2004; Lovecchio et al., 2013). Again, the material properties of the particles are then expected to affect the short-range interactions and possibly lead to aggregation (Vella & Mahadevan, 2005). The impact of such particle properties, which is outside the scope of the present work, clearly warrants further systematic investigations using different particle materials.

The observed influence of the particles' properties on dispersion, once confirmed for a wider range of particle types and flow conditions, may have profound implications for the transport of floating particles; in particular, the transport of meso- and macroplastics in small streams and turbulent waters in general. The diffusivity, which we find to roughly double from mm-sized to cm-sized objects, is a crucial quantity to incorporate the effect of unresolved spatio-temporal scales in Lagrangian transport models for rivers, lakes, and the oceans (Liu et al., 2011; Park et al., 2017; van Sebille et al., 2018; Daily & Hoffman, 2020; McDonald & Nelson, 2021). Our results indicate that such a parameter varies significantly not only with the flow conditions but also with the particle properties. Parameterizations that also include the latter appear necessary to obtain accurate predictions from such models.

5 Conclusion

Motivated by the need of understanding the transport of plastic litter in river flows, we have used time-resolved PTV to characterize the motion of particles of different shape and size floating on the surface of a field-scale meandering stream. We have considered two locations with different turbulence levels, in which the role of surface waves on the transport is deemed negligible. We have measured the position, velocity, and acceleration along the trajectories of thousands of millimetre-sized spherical pellets and centimetre-sized discs and rods, as well as the orientation and rotation of the latter, and evaluated the spatio-temporal scales associated with such quantities. At the meander, the homogeneity of the flow properties allows us to identify both dissipative and integral scales of the free surface turbulence, providing essential terms of comparison for the size of the particles and the scales of their motion. The spheres are small enough to capture most if not all scales of the free surface motion and are regarded as flow tracers; while the length of the rods and the diameter of the discs are $\mathcal{O}(100)$ times larger than the dissipative scales and several times smaller than the integral scales of the turbulence. The analysis of the particles' motion leads to the following observations:

- I. All considered particles display almost indistinguishable mean velocities and RMS velocity fluctuations. These are determined by the largest scales of the surface flow, to which the particles respond faithfully.
- II. While the velocity fluctuations follow normal distributions unaffected by the particle shape and size, the accelerations show a sizeable degree of intermittency which decreases for larger particles. This is attributed to the finite size of the particles, filtering out the smallest scales of the turbulence associated with the most intense gradients.

- 675 III. Consequently, the larger particles spread more rapidly on the turbulent free surface,
676 with diffusivity coefficients roughly doubling for centimetre-sized particles as compared
677 to millimetre-sized tracers. This is due to the motion of the larger particles being more
678 time-correlated, which in turn is rooted in their impaired response to the small-scale
679 turbulent fluctuations.
- 680 IV. The rods tend to align with the flow direction, but their instantaneous orientation is
681 influenced by a range of scales: they re-orient following the mean shear, rotate ac-
682 cording to the turnover time of the energetic eddies, and exhibit intermittency in their
683 angular velocities. This leads to less time-correlated motions and slower dispersion
684 than the discs, despite the rods' length being larger than the discs' diameter.

685 Overall, the behaviour of the free surface turbulence and the motion of particles floating
686 on it appears to be consistent with the phenomenology of inertial finite-sized particles in
687 3D turbulence. This similarity, to be confirmed in a wider range of flow conditions and
688 particle types, may allow leveraging of established results and recent advances in the field of
689 particle-laden turbulence (Balachandar & Eaton, 2010; Brandt & Coletti, 2022), furthering
690 the predictive understanding of the transport of floating plastics in natural waters. We
691 observe that the shape and size of floating particles in turbulent streaming waters only
692 affects the higher-order statistics (which in turn influences the Lagrangian transport), while
693 the mean velocity and RMS fluctuations are not measurably affected. This may be valuable
694 for modelling the transport of non-spherical floating particles in rivers.

695 Future studies shall expand the present work in several directions. Our experiments
696 have been carried out in a relatively small stream; studies in larger and deeper rivers,
697 in which the dissipation mechanisms in the water column are inherently different (Moog &
698 Jirka, 1999), are needed to expand and generalize the results. In such cases, particle imaging
699 may require the use of uncrewed aerial vehicles, which have been successfully utilized to
700 characterize natural flows (Blois et al., 2016; Liu et al., 2021). Given the variety of debris
701 types found in water streams, the range of particle properties should be expanded beyond
702 shape, size, and density: deformability and brittleness have recently been investigated in
703 laboratory studies and are especially relevant to plastic pollution (Brouzet et al., 2014, 2021).
704 Finally, high-Froude streams and/or streams under the action of wind where breaking and
705 non-breaking waves occur may play a major role in the transport of floating particles. The
706 recent laboratory experiments of Lenain et al. (2019), confirming computational results by
707 Deike et al. (2017), found that breaking waves induce much stronger transport of cm-sized
708 spherical particles compared to Stokes drift. Moreover, Ruth et al. (2022) showed that
709 bubbles entrained during wave-breaking events travel downstream faster than the Stokes
710 drift associated to buoyant particles in non-breaking waves. Overall, studies investigating
711 the effect of particle properties in wave-breaking conditions are warranted.

712 **Open Research**

713 Data - Pre-processed background-subtracted images for the different regions of interest and
714 flow rates and particles are available at <https://doi.org/10.3929/ethz-b-000572787>.

715 Data - Smoothed particle trajectories for the different regions of interest and flow rates and
716 particles are available at <https://doi.org/10.3929/ethz-b-000573259>.

717 Software - Analysis and figures were done with MATLAB version R2020a, available under
718 the MATLAB license at <https://mathworks.com/>.

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