

Supporting Information for ”Structure and thermal evolution of exoplanetary cores”

Irene Bonati¹, Marine Lasbleis^{1,2}, and Lena Noack³

¹Earth-Life Science Institute, Tokyo Institute of Technology, Tokyo 152-8550, Japan

²Laboratoire de Planétologie et Géodynamique, LPG, UMR 6112, CNRS, Université de Nantes, Université d’Angers, France

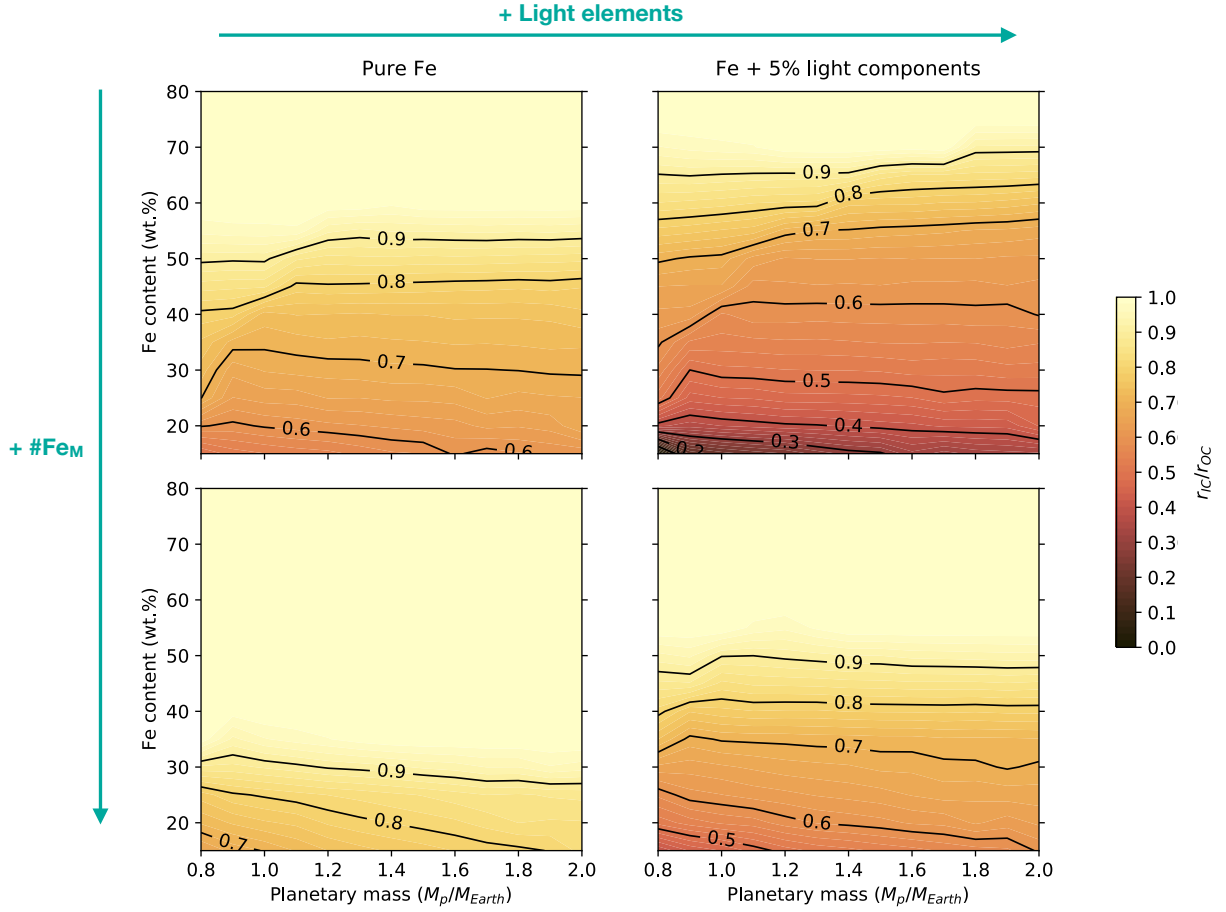
³Department for Earth Sciences, Freie Universität Berlin, Malteserstr. 74-100, D-12249 Berlin, Germany

Contents of this file

1. Figures S1 to S3

Figure S1.

Inner core radius fractions after 5 Gyr of evolution



Radial fraction of the inner core (r_{IC}/r_{OC}) after 5 Gyr of evolution as a function of planetary mass, bulk iron content, mantle iron number (upper row: $\#Fe_M = 0$, lower row: $\#Fe_M = 0.1$), and core composition (left column: pure iron, right column: iron and 5% light elements).

Figure S2.

L_ρ (left) and A_ρ (right) for a range of planetary masses and iron contents. Values for Earth are $L_\rho = 7683$ km and $A_\rho = 0.484$ (Labrosse, 2015).

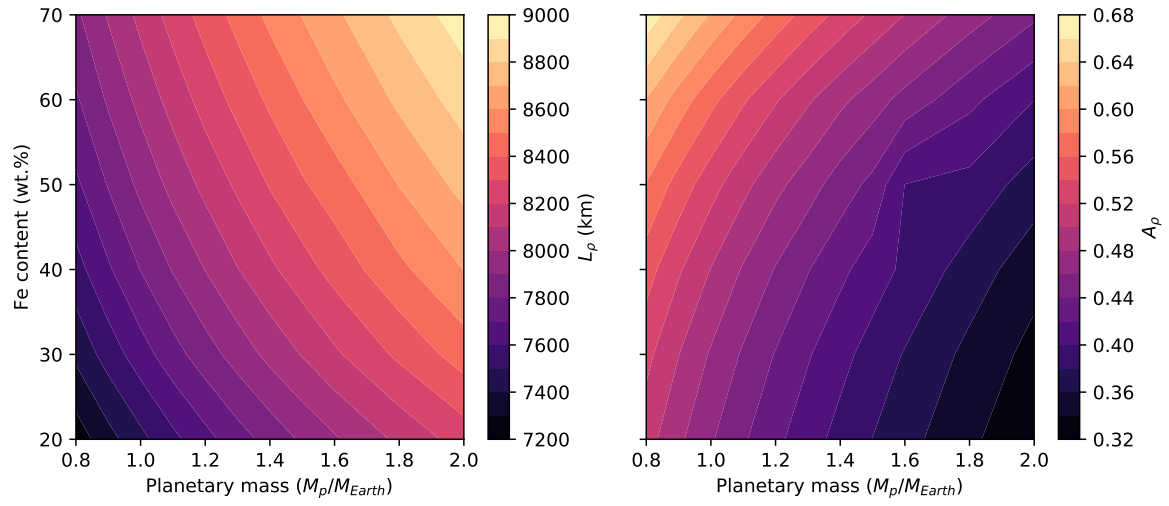
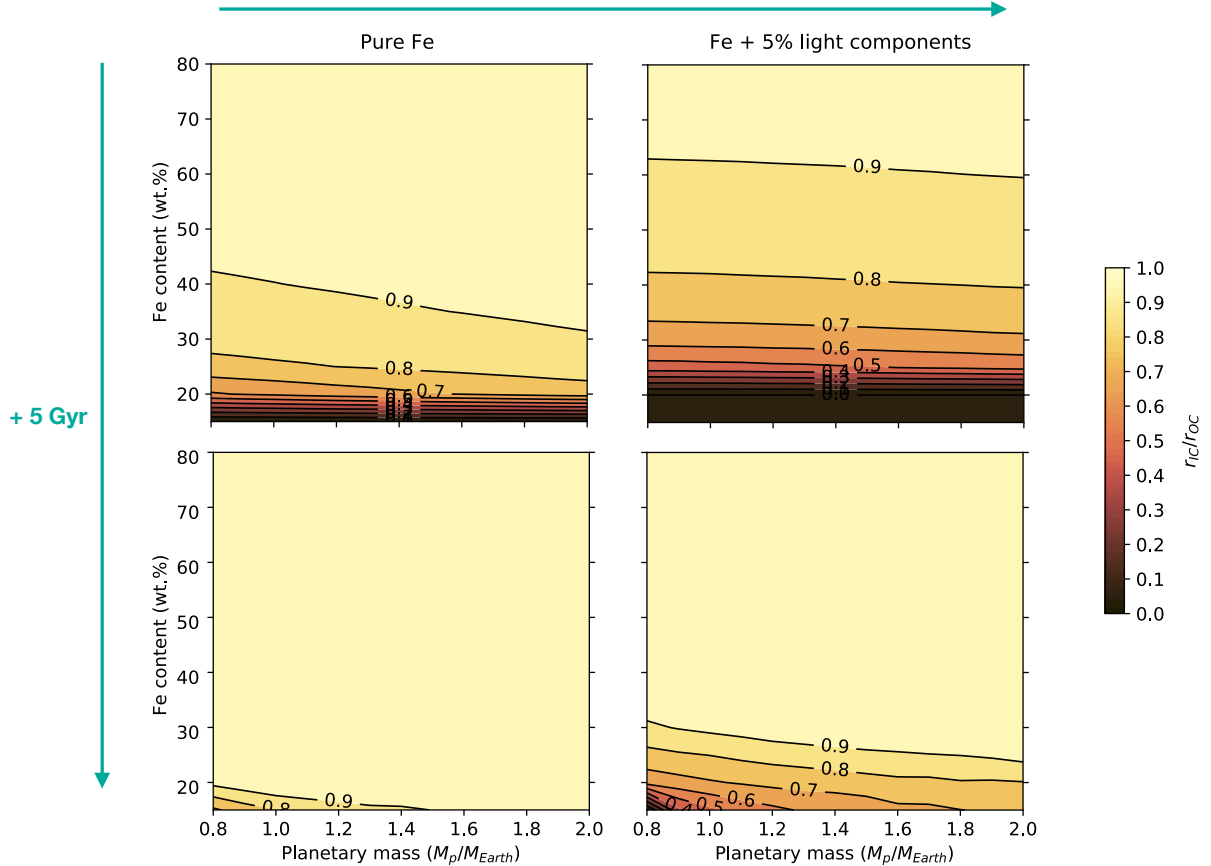


Figure S3.

Inner core radius fractions at the beginning (end of accretion) and at the end (5 Gyr later) for planets with a mantle iron number $\#Fe_M = 0.2$. The core is

made of pure iron (left panel) and iron with 5 % of light elements (right panel)



References

Labrosse, S. (2015). Thermal evolution of the core with a high thermal conductivity.

Physics of the Earth and Planetary Interiors, 247, 36–55.