

1 **Characterization of Seismicity from Different Glacial Bed Types: Machine Learning**
2 **Classification of Laboratory Stick-Slip Acoustic Emissions**

3 **S. Saltiel¹†, N. Groebner², T. Sawi¹, C. McCarthy¹, B.K. Holtzman¹**

4 ¹ Lamont-Doherty Earth Observatory, The Earth Institute, Columbia University of New York,
5 NY, USA

6 ² Strabo Analytics, Inc, New York, NY, USA.

7 Corresponding author: Seth Saltiel (ssaltiel@unr.edu)

8 † now at Nevada Seismological Laboratory, University of Nevada, Reno, NV, USA

9 **Key Points:**

- 10 • Ice slip on frozen till or rock at high velocity produce stick-slip stress-drops with AEs
11 recorded on transducers frozen into the ice
- 12 • Supervised learning can predict whether an event waveform originated from frozen till or
13 rock, but spectral features are not predictive
- 14 • Feature importance shows that till events are more impulsive, they generally have higher
15 steady-state friction and stress-drops

16 **Abstract**

17 Subglacial seismicity provides the opportunity to monitor inaccessible glacial beds in high
18 resolution. There are different types of glacial beds, which determine the mechanics of slip and, if
19 unstable, characteristics of resulting seismicity. Utilizing a double direct shear apparatus, we found
20 conditions for instability at freezing temperatures and high slip rates for both rock and till beds,
21 although with very different frictional evolution. During stick-slip stress-drops, we recorded
22 acoustic emissions with piezoelectric transducers frozen into the ice. Supervised machine learning
23 can classify recorded waveforms as coming from rock or till, while spectral information is not
24 predictive. The Random Forest Classifier is interpretable, with the prediction based on the first
25 three oscillation peaks. Till events are generally higher stress-drop, with more impulsive first
26 arrivals compared to rock waveforms. These seismic signatures of mechanical slip processes and
27 associated bed conditions can potentially greatly enhance interpretation of subglacial seismic data.

28 **Plain Language Summary**

29 A glacier can lurch forward while slipping on its base, like an earthquake, releasing seismic waves
30 which are monitored from the surface. Just like in a tectonic setting, only certain conditions allow
31 for this type of motion, and aspects of the bed conditions affect the mechanics of slip and resulting
32 waveforms. We replicate realistic glacial bed conditions in the lab of two very different types, soft
33 (sediment) and hard (rock), and measure lurching behavior and resulting waves from each. Using
34 a variety of data science techniques, we decipher subtle differences between the two bed types
35 from ‘remotely-sensed’ waves. This suggests that seismicity can provide important information on
36 glacial bed conditions and how they differ in time and space.

37 **1 Introduction**

38 Future sea-level rise will largely be determined by fast-slipping polar glaciers, known as ice
39 streams [Cuffey & Paterson 2010]. Since motion is mostly concentrated at their beds, conditions
40 in this region have an outsized effect on the entire system’s mass-balance and evolution. Glacial
41 beds are separated, to first order, into hard bedrock or soft sediment (till), and then as either wet
42 (melting temperature) or dry (frozen or drained) [Clarke 2005]. Water and sediment can move and
43 change on much shorter time scales than ice deforms, so the bed is one of the most dynamic parts
44 of the ice sheet system, assumed to be responsible for recent changes in ice flow configurations
45 [Bougamont et al., 2015] and ongoing responses to the changing climate [Parizek et al., 2013].

46 Although the basal system is difficult to directly access, growing observations of subglacial
47 seismicity offer the opportunity to monitor changes with high temporal and spatial resolution
48 [Aster & Winberry 2017]. Recent studies have used subglacial seismicity observations to infer
49 differences in bed strength [Guerin et al., 2021], failure mechanism [Kufner et al., 2021], fine-
50 scale asperity interactions [Gräff et al., 2021], basal water pressure [Gräff & Walter 2021], as well
51 as local basal shear-stresses and slip-rates [Hudson et al., 2022].

52 Seismic observations are particularly useful since there are limited glacial bed conditions that have
53 been shown to exhibit the requisite conditions for seismic failure [Iverson 2010, Lipovsky et al.,
54 2019]. Classically, ice deformation, and thus slip due to regelation and viscous creep, is assumed
55 to be rate-strengthening [Schoof 2005]. Till deformation was also first treated as viscous but later
56 shown to be Coulomb plastic, essentially rate-neutral [Iverson 2010, Zoet & Iverson 2020]. But
57 nucleation of seismic instability requires rate-weakening resistance, described by the rate-state
58 stability parameter ($b - a$), which allows acceleration due to feedback with decreasing friction, as

59 has been shown for fault rocks and gouge [Marone 1998]. This situation provides the opportunity
60 for seismic observations to present a strong constraint on the conditions at their epicentral location
61 and origin time, but each potential stick-slip mechanism and resulting seismicity characteristics
62 must be thoroughly understood to determine what recorded seismic events represent.

63 Laboratory simulations provide the opportunity to directly observe slip behavior under controlled
64 subglacial conditions. To date, seismically required rate-weakening has been reported for debris-
65 laden ice on impermeable rock at sub-freezing temperature and permeable rock at the pressure
66 melting point [Zoet et al., 2013], pure ice on impermeable rock at sub-freezing temperature
67 [McCarthy et al., 2017], and pure ice on till at sub-freezing temperature [Saltiel et al., 2021], with
68 stick-slip stress-drops only reported for debris-laden ice on impermeable rock at sub-freezing
69 temperature [Zoet et al., 2020]. These findings suggest that seismicity is largely associated with
70 dry (frozen or drained) conditions, but experiments have also shown rate-weakening is possible
71 due to cavity formation behind hard bed obstacles [Zoet & Iverson 2016] and pore-pressure
72 feedback from clast ploughing [Thomason & Iverson 2008]. Although each of these mechanisms,
73 and the bed conditions which enable them, show rate-weakening drag, their frictional evolution
74 can differ dramatically. For example, the critical slip distance (D_c) over which friction evolves to
75 a new steady-state after a change in slip rate varies by more than an order of magnitude between
76 rock and till beds under similar conditions in the same apparatus [McCarthy et al., 2017, Saltiel et
77 al., 2021]. These mechanisms' different frictional characteristics and applicable scales likely
78 contribute to aspects of the resulting seismicity, which could further constrain epicentral bed
79 conditions.

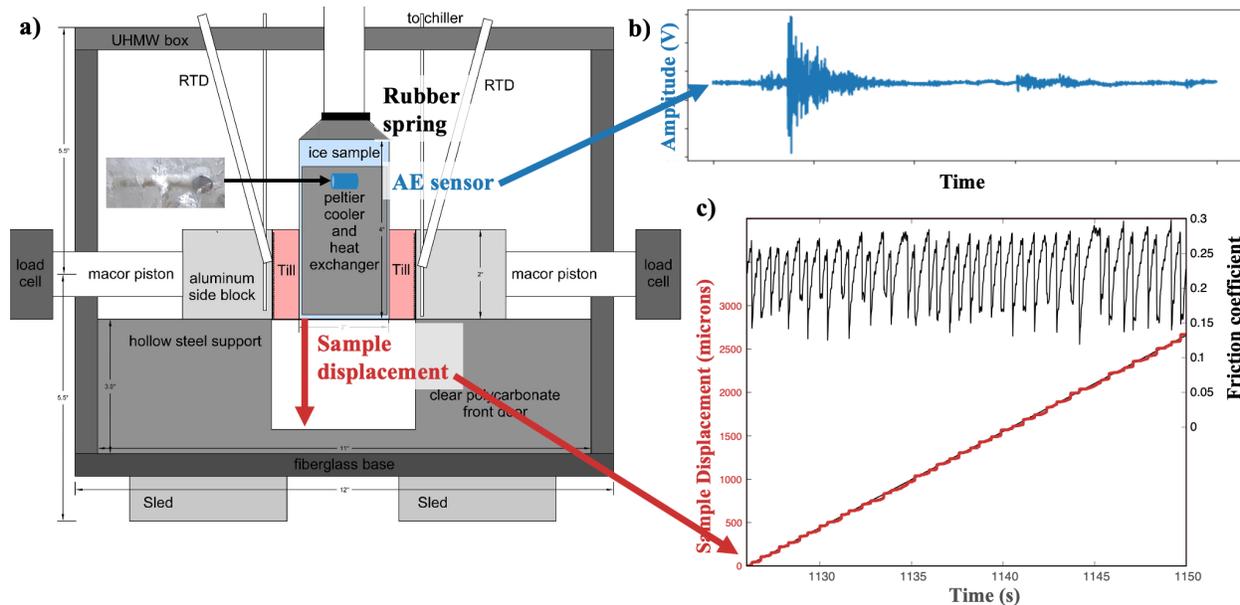
80 We report here, for the first time, experimental stick-slip stress-drops for pure ice on impermeable
81 rock and till at sub-freezing temperatures. In addition, we measured Acoustic Emissions (AEs)
82 from these settings and analyze the measured waveforms using Machine Learning (ML)
83 classification algorithms to find the characteristics associated with each bed type and its resulting
84 mechanics. By improving our understanding of the mechanisms of unstable slip in glacial settings
85 and their expression in seismic emissions, these experiments and analysis techniques provide the
86 opportunity to extract more information of conditions / source mechanics of subglacial or other
87 seismic settings.

88 **2 Experimental Methods and Materials**

89 Experiments were conducted using an ambient pressure, cryogenic temperature, servo-hydraulic
90 biaxial friction apparatus [McCarthy et al., 2016], with modifications to the insulating cryostat and
91 loading procedure to allow measurement of till [Saltiel et al., 2021]. In this double-direct-shear
92 configuration, a central ice block slides against two stationary side blocks, with layers of pre-
93 compacted and frozen till or rock, on opposite sides of the ice, such that applied horizontal load is
94 resolved as normal stress and vertical load as shear stress on the sliding interfaces (Figure 1a).
95 Additional experimental details are described in supporting text S1.

96 We made three additional modifications to the apparatus from that past study. An additional Linear
97 Variable Inductance Transducer (LVIT) position sensor measures the sample displacement
98 separate from the loading point's preset displacement. This allowed measurement of displacement
99 in each stress-drop 'slip' event as well as how much slip occurs during 'stuck' periods and the
100 timing of both relative to stress-drops (Figure 1c). Here we refer only to mechanical or bulk stress-
101 drops, the stress change during a slip event as measured by our vertical load cell, not to be confused
102 with seismologically derived stress-drops. A rubber gasket material was also inserted into the

103 loading geometry that effectively reduced the stiffness of the apparatus, reaching critical stiffness
 104 and allowing stick-slip instability. We estimate the effective apparatus stiffness using the
 105 mechanical data's reloading slope between stress-drops, relative to the compression of the loading
 106 train including rubber, the load point displacement minus sample displacement (Figure 1c). We
 107 estimate the apparatus stiffness after adding the rubber to be ~ 0.1 kPa/ μm or $\sim 5 \times 10^5$ N/m,
 108 significantly less stiff than was estimated without the rubber ~ 1 kPa/ μm [Saltiel et al., 2021].
 109 Additionally, commercial piezoelectric transducers were frozen into the central ice block, facing
 110 one of the ice-bed interfaces, to measure AEs. After experimenting with four different types of
 111 transducers of varying sizes and frequency sensitivities, we settled on Physical Acoustic's Nano-
 112 30™ miniature AE sensor due to its small size and 125-750 kHz response, covering the major
 113 frequency content of the events. All AEs analyzed here were recorded with a single Nano-30.



114 **Figure 1: a)** Schematic of biaxial cryostat with additions of rubber spring to decrease loading
 115 stiffness, AE sensor frozen into central ice block (pictured within ice in inset on left), and sample
 116 displacement measurement, modified from Saltiel et al., [2021]. **b)** An example AE waveform
 117 before processing, from a single stress-drop / slip event, and **c)** an example experiment of measured
 118 friction drops (in black on top) and stick-slip sample displacement (in red on the bottom) with the
 119 steady load point displacement (in black) for reference, due to instability induced by apparatus
 120 reaching subcritical stiffness.
 121

122 AEs were recorded using a preamplifier and TiePie™ HS6 differential digital oscilloscope. To
 123 ensure we recorded all relevant spectral content in the waveforms, they were recorded at a very
 124 high sample rate of 100 MHz for 2 ms time windows around each triggered event. These
 125 oscilloscope settings provided the optimum real-time viewing of triggered waveforms as they were
 126 being recorded (Figure 1b), but subsequent analysis showed most of the energy was under 1 MHz,
 127 and waveforms were subsequently down sampled to 10 MHz and windowed to 15 μs , lowering
 128 file size. Recordings of continuous acoustic signal without applied shear found electrical noise
 129 above 3 MHz, so filtering also helped remove persistent noise sources. The oscilloscope was set
 130 in rising-limb trigger mode with trigger amplitude set just above the noise level before the
 131 deformation program started, such that it did not trigger without an audible stress-drop. Since
 132 electrical and other sources of noise can vary, this trigger level was adjusted throughout the

133 experiment to maximize the number of captured events and minimize waveforms of purely noise,
134 but some events were missed, and many events triggered by noise were saved.

135 **3 Data Processing and Machine Learning Analysis**

136 To remove noisy events, non-events triggered by noise, and to normalize the waveform in a way
137 that focuses on the initial wave arrivals, we implemented a data cleaning and normalization
138 approach based on that implemented by Nolte & Pyrak-Nolte [2022], described in supporting text
139 S2.

140 After removing noisy waveforms, we end up with 2817 total events, including 1547 waveforms
141 from 6 till experiments and 1270 waveforms from 6 rock experiments. With this labeled catalog
142 (Figure 2), we systematically explored the predictive performance of numerous supervised ML
143 algorithms on the waveforms as well as on spectra and spectrograms creating from the waveforms.
144 We found that none of the spectral-based algorithms were substantially more predictive than
145 random (at best ~55%), so here we focus on waveform-based results.

146 We divide the data into training and test sets based on experiment, i.e., for a given model training
147 run the waveforms from 5 till and 5 rock experiments are used for the training set, and the
148 remaining 1 till and 1 rock experiment are used for testing. By separating training and test sets by
149 experiment, any experiment-dependent features of the waveforms would be irrelevant for
150 classification. As experiments vary in number of events (between 94 – 465), we calculated
151 balanced prediction accuracy for each set of test data. The prediction accuracy is summarized by
152 a 6 till by 6 rock experiments matrix, giving the accuracy for 36 models with each combination
153 used as the testing data (Figure 3).

154 We focus our analysis on the Random Forest Classifier model [Breiman 2001] applied to processed
155 waveforms, since it obtained some of our highest prediction accuracies (68% mean accuracy),
156 independent of which experiments were used for testing, and it gives the feature importance for
157 interpretability. The feature importance shows the weighting of each waveform sample in making
158 predictions (Figure 4a). The feature importance is key for interpreting how the prediction is made
159 and visually highlighting the subtle differences between different waveform sources.

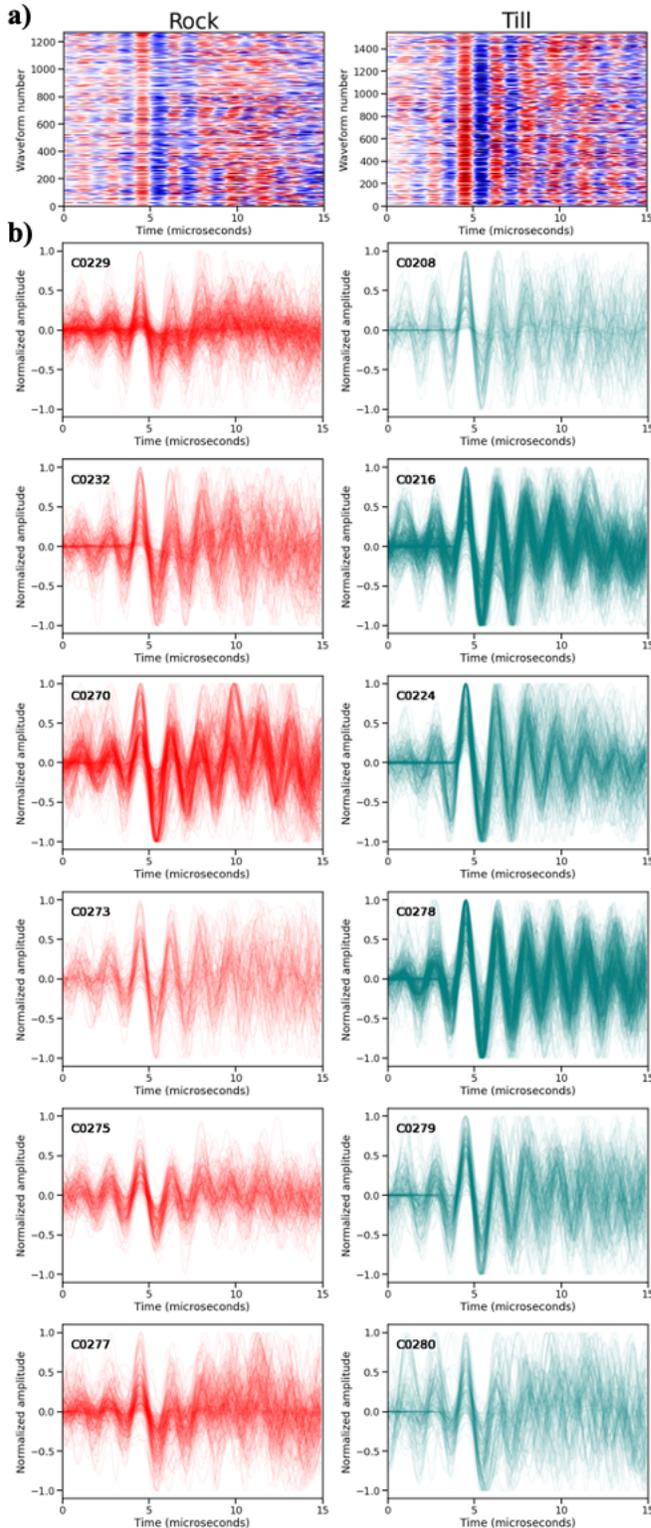
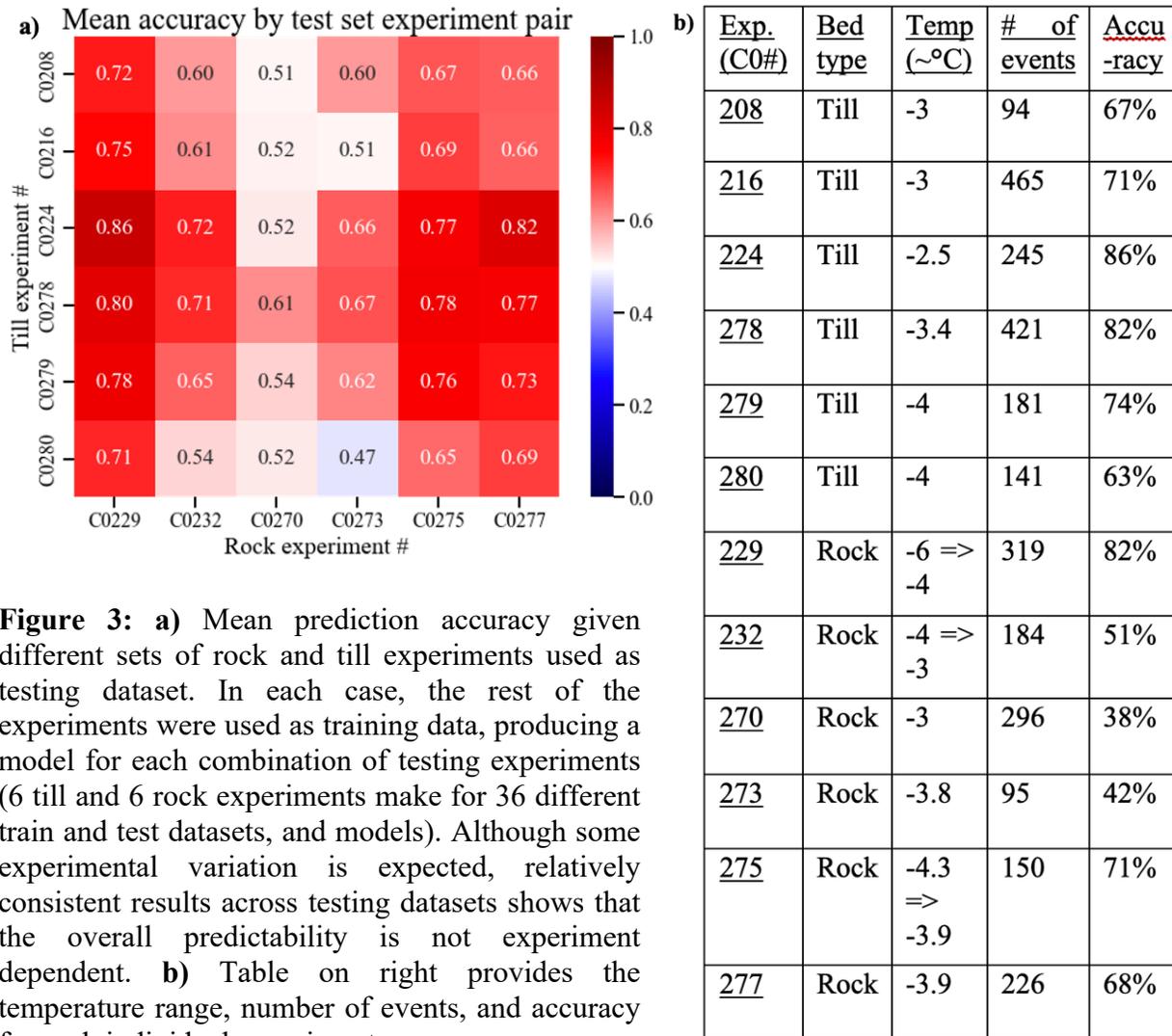


Figure 2: a) Waveforms plotted in chronological order along y-axis, oldest experiments, lower number, on bottom, colored by amplitude and normalized by maximum amplitude (red is positive and blue negative) with rock events plotted on left and till on the right. b) Waveforms plotted together for each experiment (labelled on upper left). Rock experiments are plotted in red, while the till events are teal. Each waveform is plotted with a thin, light line, so the dark parts show many waveforms aligned on top of each other, and broader lines show less alignment. Since experiments vary significantly by number of events (94 – 465), that also contributes to the appearance of each experiment plot. Although there are subtle differences, it is not visually clear that the two beds can be deciphered, making it a useful dataset to explore ML-based classification.

182 **4 Results and Discussion**

183 Using a wide range of classification algorithms, we consistently find prediction accuracy above
 184 50%, mostly between 60% and 85%, showing it is possible to tell if waveforms were emitted by
 185 till or rock beds. This is not clear by visually examining the waveforms (Figure 2), showing
 186 algorithms successfully extract subtle waveform features correspond to the different bed labels.
 187 Our preliminary explorations found almost no predictive power in the spectral data, including
 188 spectra, spectrograms, and further extracted. In contrast, every method of processing the waveform
 189 data and every algorithm we tested, found some overall predictive capability in the waveforms.

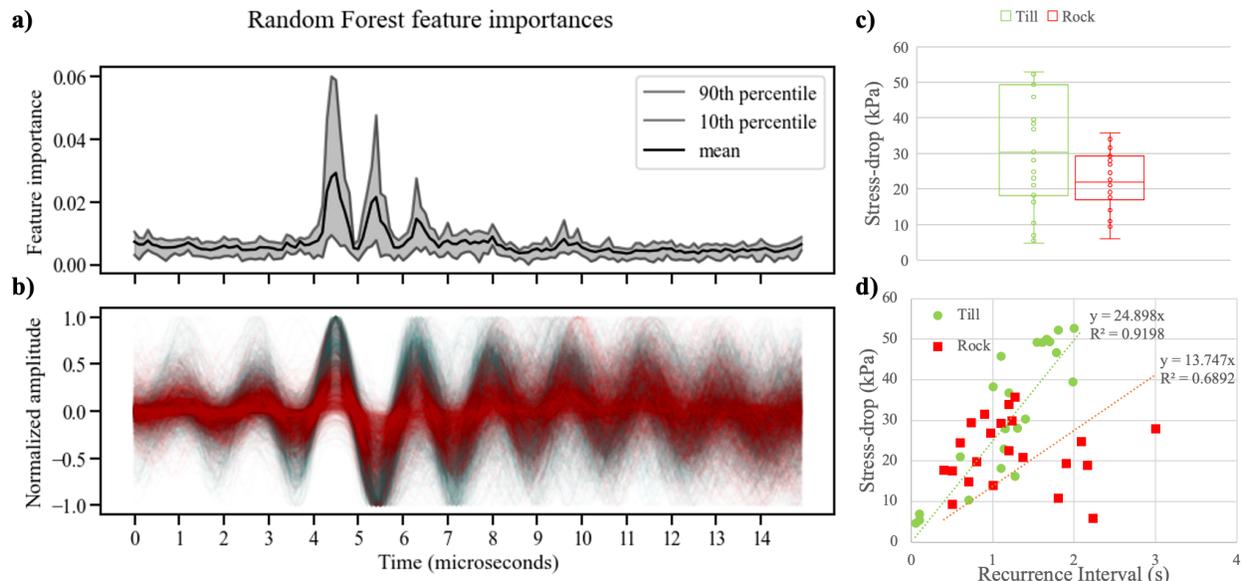


190 **Figure 3:** a) Mean prediction accuracy given
 191 different sets of rock and till experiments used as
 192 testing dataset. In each case, the rest of the
 193 experiments were used as training data, producing a
 194 model for each combination of testing experiments
 195 (6 till and 6 rock experiments make for 36 different
 196 train and test datasets, and models). Although some
 197 experimental variation is expected, relatively
 198 consistent results across testing datasets shows that
 199 the overall predictability is not experiment
 200 dependent. b) Table on right provides the
 201 temperature range, number of events, and accuracy
 202 for each individual experiment.

203 This prediction accuracy calculates how often the model could correctly classify individual
 204 waveforms as coming from till or rock beds, but we envision a tool whereby a collection of seismic
 205 events recorded from a given location would be analyzed to determine the probability it came from
 206 a till or rock-based glacier. So, the more relevant accuracy is if a single experiment can be
 207 accurately predicted to be till or rock, and how many events would be needed to make such a
 208 prediction accurate. By this metric, all 6 till experiments would be correctly predicted (with total

209 test accuracies well above 50%), while only 3 rock experiments are robustly predicted as rock.
 210 This shows that the model is very sensitive to till features, able to detect them in all the till
 211 experiments, but more specific for rock, in that every experiment predicted as rock was correct.
 212 Even though the algorithms correct for class imbalances there is a minimum number of events that
 213 are needed to give accurate results, which can be analyzed using a receiver operating characteristic
 214 (ROC) curves. This minimum catalog size, along with seismicity rates, would set the maximum
 215 temporal and spatial resolution of bed identification through this method.

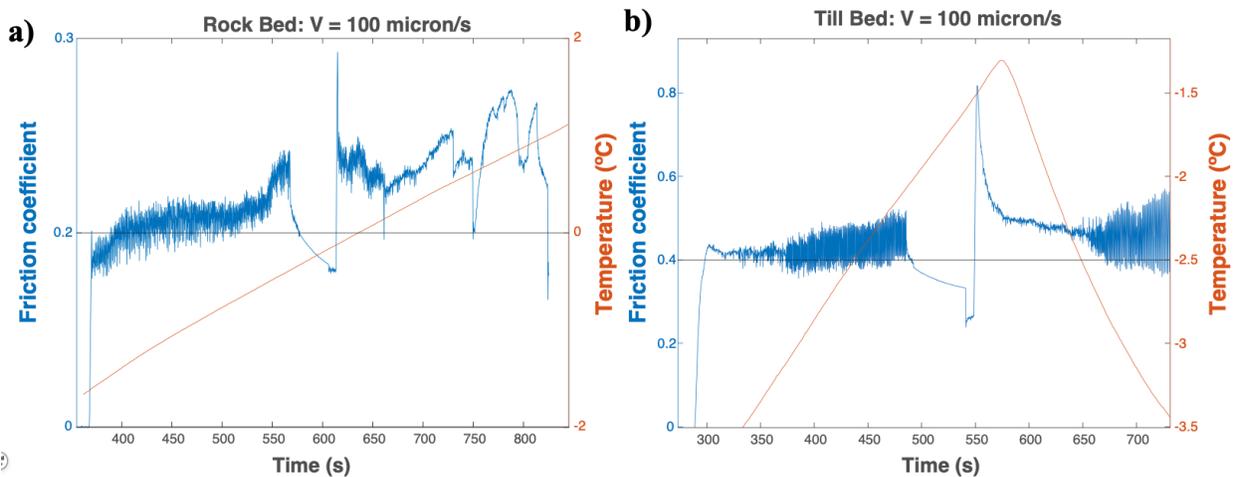
216 To be able to apply our findings from laboratory AEs to field-scale seismicity, it is vital that we
 217 can interpret how the algorithms make their prediction. Although transfer learning methods offer
 218 the potential to train with labelled laboratory or modelled datasets and ‘transfer’ the model to more
 219 limited field or laboratory data [Wang et al., 2021], clear differences in the spectral content, path
 220 effects, and scale of field seismic data make this a daunting task. But by isolating and interpreting
 221 the features the algorithms are using to make their successful predictions, we can understand the
 222 differences in the waveforms to look for and interpret in field data. The feature importance for all
 223 36 Random Forest Classifier models show similar parts of the waveform are used to make the
 224 prediction, focusing on the first three peaks (positive and negative) of the initial wave arrivals
 225 (Figure 4a). Plotting all the normalized waveforms (color coded by bed type) on top of each other,
 226 we can see that the till (teal) waves tend to have higher amplitude in these first peaks (Figure 4b).
 227 Analyzing the mechanical data from 23 till and 22 rock experiments (including other experiments
 228 without recorded AEs, in addition to those described above), we find that the stress-drops of stick-
 229 slip events on till beds are generally higher (Figure 4c). This difference in stress-drop is consistent
 230 with till’s higher healing rates (Figure 4d) and higher average friction (Figure 5), but given the
 231 range of values for each and subtlety of waveform differences, ML techniques were needed to
 232 quantify how accurately bed-type can be deciphered from waveforms. Along with the fact that till
 233 events are generally predicted more accurately, we interpret the algorithm’s attention to initial
 234 arrivals as evidence that this till feature is the main identification and the rest are classified as rock.



235 **Figure 4:** **a)** Mean feature importance, showing the relevance of each waveform sample to the
 236 models’ prediction, with 10th and 90th percentile error bands for all 36 Random Forest Classifier
 237 models, and **b)** plot of all rock (red) and till (teal) waveforms. Results highlight the importance of
 238

239 the initial three wave arrivals in all models. Plotting the superimposed normalized waveforms,
 240 shows that the till (teal) events are higher amplitude in these first few oscillations. **c)** Box and
 241 whisker plots of largest repeated mechanical stress-drop amplitude from 23 till and 22 rock
 242 experiments at similar stress and temperature conditions show till also has higher stress-drops,
 243 although with overlap. **d)** Stress-drops vs their recurrence interval for till and rock experiments,
 244 the till beds' greater healing (higher slope) contributes to their higher stress-drops.
 245

246 These experiments also show the temperature dependence of instability, as both rock and till
 247 experiments were over a range of temperatures. Although analyzing the temperature dependence
 248 of AEs is outside of the scope of this letter, we find stick-slip instability is limited to frozen
 249 temperatures ($< \sim 0$ °C for rock and $< \sim -2.5$ °C for till beds in Figure 5. Given temperatures are
 250 approximate, since they are measured behind the till/rock, there is some lag time before the
 251 temperature on the sliding interface reached those recorded. This finding is consistent with that of
 252 rate-weakening friction in till beds at ~ -3 °C using the same apparatus [Saltiel et al., 2021]. We
 253 estimate the apparatus stiffness with rubber to be ~ 0.1 kPa/ μm or $\sim 5 \times 10^5$ N/m, which is the
 254 same order of magnitude as the critical stiffness of estimated from velocity-step experiments \sim
 255 0.02 kPa/ μm or 1×10^5 N/m at similar conditions (see section on critical stiffness in Saltiel et al.,
 256 [2021]). This factor of five difference is consistent with the error inherit in applying estimations
 257 of rate-state friction parameters ($b - a$, D_c) from few experiments, as well as in our rough
 258 estimation of apparatus stiffness. Past studies of ice-on-rock friction did not find rate-weakening
 259 until lower temperatures, $< \sim -18$ °C for McCarthy et al., [2017] using this apparatus. In that study,
 260 experiments above -18 °C which exhibited slight rate-strengthening were undertaken at less than
 261 half the slip rate, which could affect the rate-dependence of friction as well as stability more
 262 broadly [Schulson & Fortt 2012]. It is also possible to reach instability at nominally stable
 263 conditions given the strong elastic contrast between ice and rock beds [Rice et al., 2001]. This
 264 highlights the range of factors that contribute to seismic instability. Further experiments and
 265 analysis are needed to fully map subglacial stability.



266 **Figure 5:** Example experiments of the temperature effect on slip stability for **a)** rock and **b)** till
 267 beds. Each experiment shows stress-drops in the beginning of the experiment but, after an
 268 experimentally induced hold (described in supporting text S1), with increasing temperature the ice
 269 starts to stably slide without sudden drops in friction or audible stick-slips. The transition to stable
 270 sliding occurs around ~ 0 °C for the rock experiment. In the till experiment, stability is reached
 271 during the hold, but as the temperature is lowered again stress-drops resume after reaching ~ -2.5
 272

273 °C. Each estimated transition temperature is highlighted with a solid black horizontal line. The till
 274 experiment has higher friction and healing rate (as the friction coefficient rose significantly more
 275 after hold times of similar duration).

276 **5 Conclusions**

277 This study presents stick-slip stress-drops and resultant AE waveforms for ice on rock and till beds
 278 at sub-freezing temperatures, a labeled dataset in which we explore how ML can decipher which
 279 bed produced the events. We found that instability, and thus seismicity, only occurs for each bed
 280 below a certain temperature (~ 0 °C for rock and ~ -2.5 °C for till), sliding stably as the temperature
 281 warms above and stick-slipping again when frozen below that temperature. Although the different
 282 bed types exhibit stick-slip behaviors at similar conditions, the mechanics of their drag are very
 283 different, demonstrated by friction that evolves over an order of magnitude more distance (D_c),
 284 significantly more rate-weakening ($b - a$), higher friction, and healing rates in frozen till compared
 285 to rock beds [Saltiel et al., 2021]. The higher healing rates contribute to the generally higher stress-
 286 drops (Figure 4d). Resultant emissions have subtle differences, difficult to decipher visually, but
 287 which ML-based classification was able to identify; successfully predicting the bed type of a given
 288 waveform about 60% to 85% of the time. Given the events from an entire experiment, all 6 till
 289 experiments were correctly identified, but only half of the rock experiments were robustly
 290 predicted. In contrast, spectral data was not predictive. The Random Forest Classifier was
 291 particularly successful and interpretable, since it provides feature importance of each waveform
 292 sample, showing the models focus on the first three wave arrivals, where till waveforms are higher
 293 amplitude. This is consistent with more impulsive failure, higher stress-drops, friction, and healing.

294 These findings are counter to our original hypothesis based on the longer frictional evolution
 295 distances (D_c) found in velocity-step experiments, which suggest less impulsive, lower frequency
 296 emissions. It is possible that different aspects of the frictional mechanics counter each other, for
 297 example more healing has been associated with higher frequency emissions in laboratory and
 298 natural faults [McLaskey et al., 2012], which could cancel the spectral effect of longer D_c . In a
 299 similar way, till experiments' higher D_c and $b - a$ balance each other to produce a critical
 300 rheological stiffness of the same order as rock [Saltiel et al., 2021]. In the end, our findings suggest
 301 that ML-based classification and correlation studies could find unknown and non-intuitive
 302 relationships between seismic emission characteristics and the mechanics / conditions of rupture
 303 in subglacial, as well as tectonic, volcanic, induced seismicity settings. Laboratory experiments
 304 offer the opportunity to obtain well-controlled, labeled datasets, but results need to be interpretable.
 305 Although it will be difficult to transfer models trained in the lab directly to field-scale data, the
 306 understanding gained can be used to infer characteristics of natural or induced seismic sources.

307 **Acknowledgments, Samples, and Data**

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313 The datasets generated for this study are available on figshare.com at doi:
 314 [10.6084/m9.figshare.21257730](https://doi.org/10.6084/m9.figshare.21257730), and Jupyter notebook for processing data is available at
 315 <https://github.com/StraboAI/IcesAEs>.

316 **References**

- 317 Aster, R. C., & Winberry, J. P. 2017. Glacial seismology. *Reports on Progress in Physics*, **80**,
 318 126801. doi:10.1088/1361-6633/aa8473
- 319 Bougamont, M., Christoffersen, P., Price, S. F., Fricker, H. A., Tulaczyk, S., & Carter, S. P.
 320 2015. Reactivation of Kamb Ice Stream tributaries triggers century-scale reorganization
 321 of Siple Coast ice flow in West Antarctica. *Geophysical Research Letters*, **42**(20), 8471-
 322 8480.
- 323 Breiman, L. 2001. Random Forests. *Machine Learning* **45**, 5–32 doi:10.1023/A:1010933404324
- 324 Clarke, G. K. 2005. Subglacial processes. *Annual Review of Earth and Planetary Sciences*, **33**(1),
 325 247-276.
- 326 Cole, D. M. 1979. Preparation of polycrystalline ice specimens for laboratory experiments, *Cold*
 327 *Reg. Sci. Technol.*, 1, 153–159, doi: 10.1016/0165-232X(79)90007-7.
- 328 Cuffey, K.M. & Paterson, W.S.B., 2010. *The Physics of Glaciers*, 4th ed, Elsevier
- 329 Gräff, D., Köpfl, M., Lipovsky, B. P., Selvadurai, P. A., Farinotti, D., & Walter, F., 2021. Fine
 330 structure of microseismic glacial stick-slip. *Geophysical Research Letters*, **48**,
 331 e2021GL096043. <https://doi.org/10.1029/2021GL096043>
- 332 Gräff, D., & Walter, F., 2021. Changing friction at the base of an Alpine glacier. *Sci. Rep.* **11**,
 333 10872. <https://doi.org/10.1038/s41598-021-90176-9>
- 334 Guerin, G., Mordret, A., Rivet, D., Lipovsky, B. P., & Minchew, B. M., 2021. Frictional origin of
 335 slip events of the Whillans Ice Stream, Antarctica. *Geophysical Research Letters*, **48**(11),
 336 e2021GL092950.
- 337 Iverson, N.R., 2010. Shear resistance and continuity of subglacial till: hydrology rules. *Journal of*
 338 *Glaciology*, **56** (200) 1104-1114. doi:10.3189/002214311796406220.
- 339 Hudson, T., Kufner, S. K., Brisbourne, A., Kendall, M., Smith, A., Alley, R., ... & Murray, T.
 340 2022. Friction and slip measured at the bed of an Antarctic ice stream., Preprint from
 341 *Research Square*, DOI: 10.21203/rs.3.rs-1214097/v1
- 342 Kufner, S.-K., Brisbourne, A. M., Smith, A. M., Hudson, T. S., Murray, T., Schlegel, R., et al.
 343 2021. Not all icequakes are created equal: Basal icequakes suggest diverse bed deformation
 344 mechanisms at Rutford Ice Stream, West Antarctica. *Journal of Geophysical Research:*
 345 *Earth Surface*, 126, e2020JF006001. Doi:10.1029/2020JF006001
- 346 Lipovsky, B.P., Meyer, C.R., Zoet, L.K., McCarthy, C., Hansen, D.D., Rempel, A.W. & Gimbert,
 347 F., 2019. Glacier sliding, seismicity and sediment entrainment. *Annals of Glaciology*,
 348 **60**(79), pp.182-192.
- 349 Marone, C., 1998. Laboratory-derived friction laws and their application to seismic faulting, *Ann.*
 350 *Rev. Earth Planet. Sci.*, **26**, 1, 643–696.
- 351 McCarthy, C., Savage, H.M., Koczyński, T., & Nielson, M.A., 2016. An apparatus to measure
 352 frictional, anelastic, and viscous behavior in ice at temperate and planetary conditions.
 353 *Review of Scientific Instruments*, **87**.
- 354 McCarthy, C., Savage, H.M., & Nettles, M., 2017. Temperature dependence of ice-on-rock friction
 355 at realistic glacier conditions, *Phil. Trans. R. Soc. A*, **375** (2086), 20150348.
- 356 McLaskey, G. C., Thomas, A. M., Glaser, S. D., & Nadeau, R. M., 2012. Fault healing promotes
 357 high-frequency earthquakes in laboratory experiments and on natural faults. *Nature*,
 358 **491**(7422), 101-104.
- 359 Nolte, D. D., & Pyrak-Nolte, L. J., 2022. Monitoring fracture saturation with internal seismic
 360 sources and twin neural networks. *Journal of Geophysical Research: Solid Earth*, **127**(2),
 361 e2021JB023005.

- 362 Parizek, B. R., Christianson, K., Anandakrishnan, S., Alley, R. B., Walker, R. T., Edwards, R. A.,
 363 Wolfe, D. S., Bertini, G. T., Rinehart, S.K., Bindschandler, R. A., & Nowicki, S. M. J.,
 364 2013. Dynamic (in) stability of Thwaites Glacier, West Antarctica. *Journal of Geophysical*
 365 *Research: Earth Surface*, **118**(2), 638-655.
- 366 Rice, J. R., N. Lapusta, & Ranjith K., 2001. Rate and state dependent friction and the stability of
 367 sliding between elastically deformable solids, *J. Mech. Phys. Solids*, **49**, no. 9, 1865–1898.
- 368 Saltiel, S., McCarthy, C., Creyts, T. T., & Savage, H. M., 2021. Experimental evidence of velocity-
 369 weakening friction during ice slip over frozen till: Implications for basal seismicity in fast
 370 moving, soft-bed glaciers and ice streams. *Seismological Research Letters*, **92**(5), 2793-
 371 2810.
- 372 Schoof, C. 2005. The effect of cavitation on glacier sliding. *Proceedings of the Royal Society A:*
 373 *Mathematical, Physical and Engineering Sciences*. 461.2055. 609-627.
- 374 Schulson, E. M., & Fortt, A. L. 2012. Friction of ice on ice. *Journal of Geophysical Research:*
 375 *Solid Earth*, **117**(B12).
- 376 Thomason, J.F. & Iverson, N.R., 2008. A laboratory study of particle ploughing and pore-pressure
 377 feedback: a velocity-weakening mechanism for soft glacier beds. *Journal of Glaciology*,
 378 **54**(184), pp.169-181.
- 379 Wang, K., Johnson, C. W., Bennett, K. C., & Johnson, P. A., 2021. Predicting fault slip via transfer
 380 learning. *Nature communications*, **12**(1), 1-11. Doi: 10.1038/s41467-021-27553-5
- 381 Zoet, L.K., Carpenter, B., Scuderi, M., Alley, R.B., Anandakrishnan, S., Marone, C., Jackson, M.,
 382 2013. The effects of entrained debris on the basal sliding stability of a glacier. *Journal of*
 383 *Geophysical Research: Earth Surface*, **118**, 656-666.
- 384 Zoet, L.K., & Iverson, N.R., 2016. Rate-weakening drag during glacier sliding. *Journal of*
 385 *Geophysical Research: Earth Surface*, **121**(7), 1206-1217.
- 386 Zoet, L.K. & Iverson, N.R., 2018. A healing mechanism for stick-slip of glaciers. *Geology*, **46**(9),
 387 pp.807-810.
- 388 Zoet, L.K. & Iverson, N.R., 2020. A slip law for glaciers on deformable beds. *Science*, **368**(6486),
 389 pp.76-78.
- 390 Zoet, L.K., Ikari, M.J., Alley, R.B., Marone, C., Anandakrishnan, S., Carpenter, B.M., & Scuderi,
 391 M.M., 2020. Application of constitutive friction laws to glacier seismicity. *Geophysical*
 392 *Research Letters*, **47**(21), e2020GL088964.