

520

521

Geophysical Research Letters

522

Supporting Information for

523

**Major modes of climate variability dominate nonlinear Antarctic ice-sheet elevation changes 2002-
2020**

524

2020

525

Matt A. King^{1,2}, Poul Christoffersen^{2,3}

526

¹School of Geography, Planning, and Spatial Sciences, University of Tasmania, Hobart, Tasmania
7001, Australia

527

528

²The Australian Centre for Excellence in Antarctic Science, University of Tasmania, Hobart,
Tasmania 7001, Australia

529

530

³Institute for Marine and Antarctic Studies, University of Tasmania, Hobart, Tasmania 7001,
Australia

531

532

533

534

Contents of this file

535

536

Text S1

537

Figures S1 to S9

538

Table S1

539

540

541 **Text S1**

542

543 To make a first-order estimate of plausible velocity-driven changes on surface elevation on Pine
 544 Island and Thwaites Glacier, we used annual Measures v2 ice velocities (Rignot et al., 2017) to
 545 derive year-on-year velocity changes at the glacier point positions where we extracted time series
 546 of surface elevation changes (shown in Fig. 3). The annual velocity changes were extracted in
 547 polar stereographic coordinates (x, y) and averaged over a 25×25 km box centred around each
 548 of the two geographical locations. The corresponding changes in surface height were estimated
 549 from the conservation of mass:

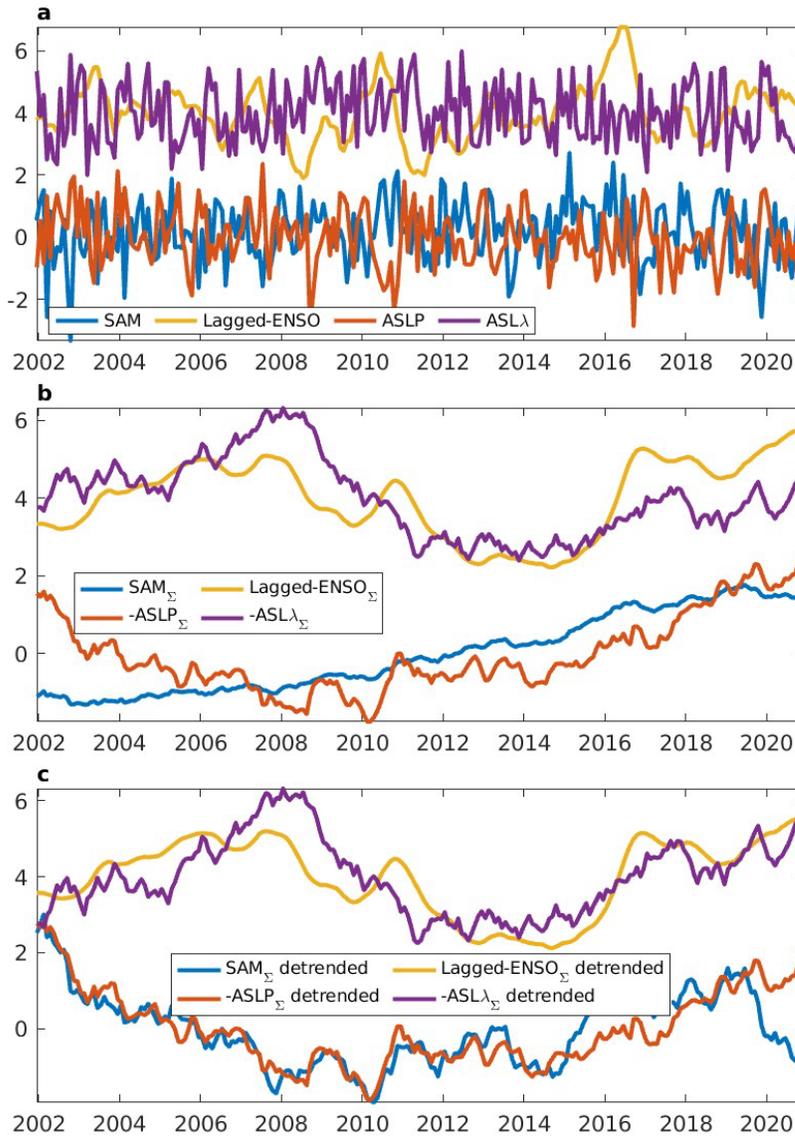
$$550 \quad \frac{\partial h}{\partial t} = \dot{a} + \dot{b} + \bar{v} \cdot \nabla H + H \cdot \dot{\epsilon}_z$$

551 where \dot{a} is the surface accumulation from snowfall, \dot{b} is the basal accumulation when water at
 552 the bed freezes on (negative for melting), \bar{v} is the velocity vector, ∇ is the gradient operator, H is
 553 ice thickness and $\dot{\epsilon}_z$ is the vertical strain rate averaged for the ice column. We used the third and
 554 fourth term on the RHS to derive estimates of the elevation change stemming specifically from
 555 the change in velocity, ignoring the accumulation terms \dot{a} and \dot{b} . If the velocity change is $\delta\bar{v}$, the
 556 change in height is $-\delta\bar{v} \cdot \nabla H + H \cdot \delta\dot{\epsilon}_z$, where the first term denotes elevation gain from
 557 advection of thicker ice (or the opposite) and the second term is elevation loss due to dynamic
 558 thinning when the ice velocity increases (or the opposite). The change in velocity, $\delta\bar{v} =$
 559 $(\Delta v_x, \Delta v_y)$ was the annual difference from one year to the next. The advection term was
 560 calculated as $\Delta v_x dh/dx + \Delta v_y dh/dy$, while strain thinning was calculated as
 561 $H(-\Delta(dv_x/dx) - \Delta(dv_y/dy))$ with the assumption that ice is incompressible ($\delta\dot{\epsilon}_z = -\delta\dot{\epsilon}_x -$
 562 $\delta\dot{\epsilon}_y$) and vertical strain in the ice column is uniform. Fig. S10 shows the estimated elevation
 563 change in terms of advection and strain and their total.

564

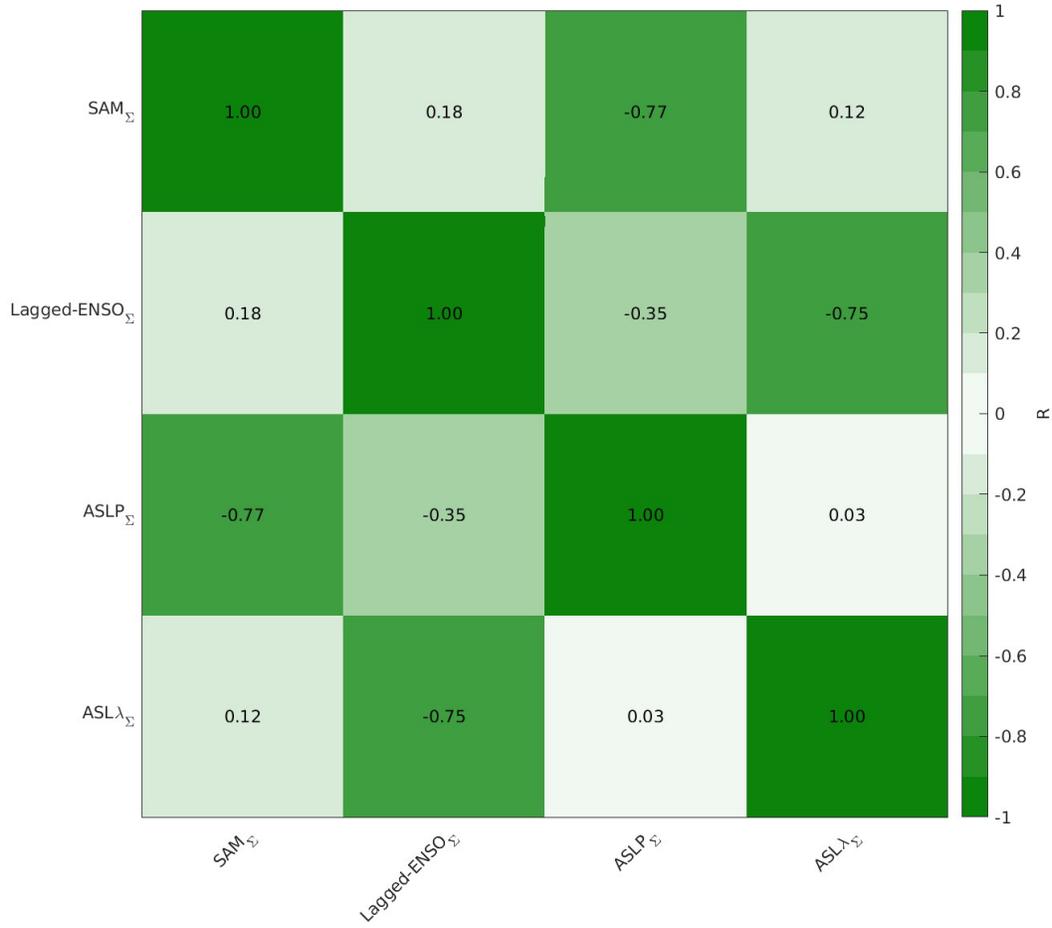
565

566
 567 Supplementary Figures
 568
 569



570
 571 **Fig S1.** Cumulatively summed and detrended climate indices. The panels show the normalized
 572 climate indices (a), their cumulative sum renormalized (b), after further detrending and
 573 renormalization (c). ENSO and ASL λ terms are shown offset by 4 units for clarity.
 574
 575

576



577
578
579
580

Fig S2. Correlation coefficients of summed and detrended climate indices.

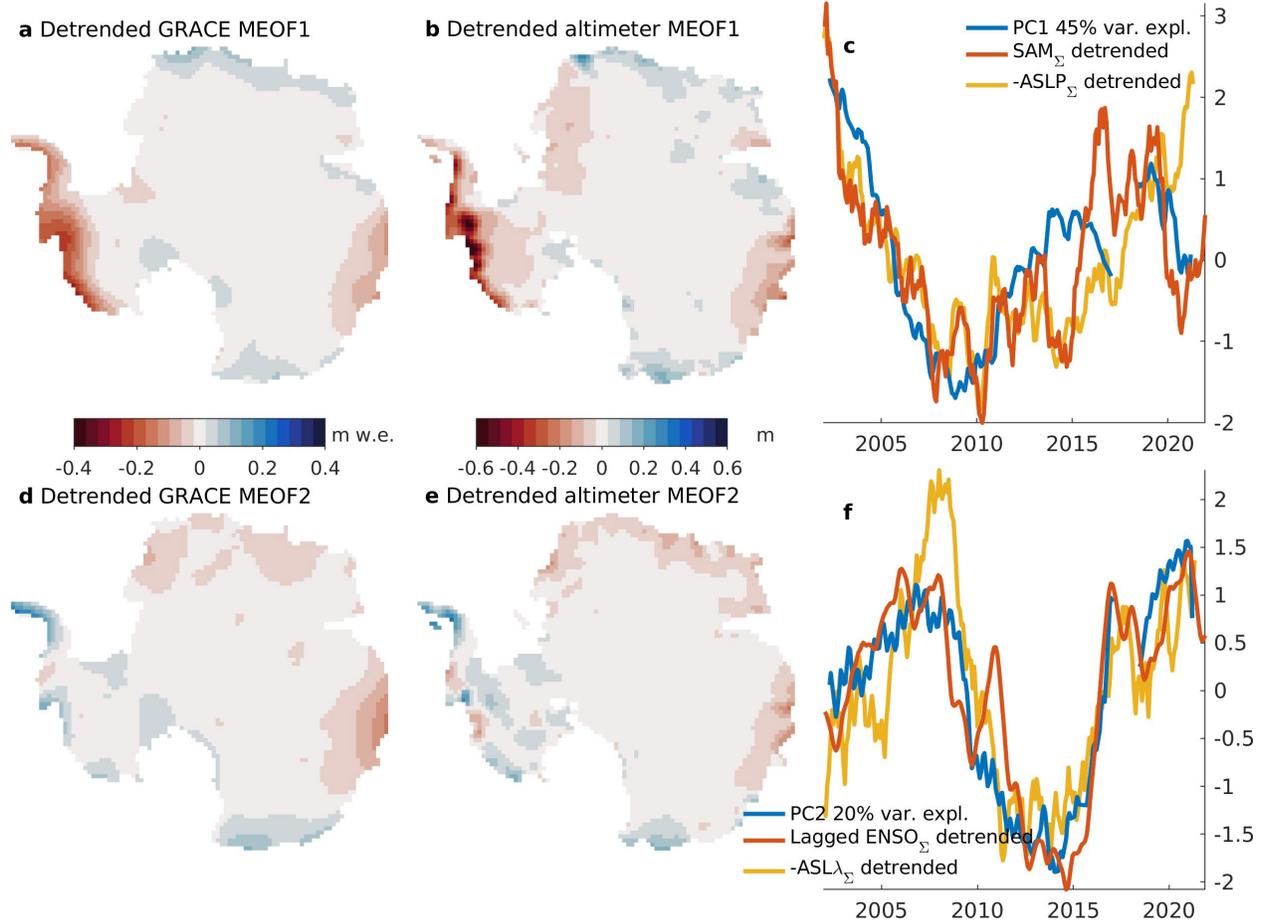
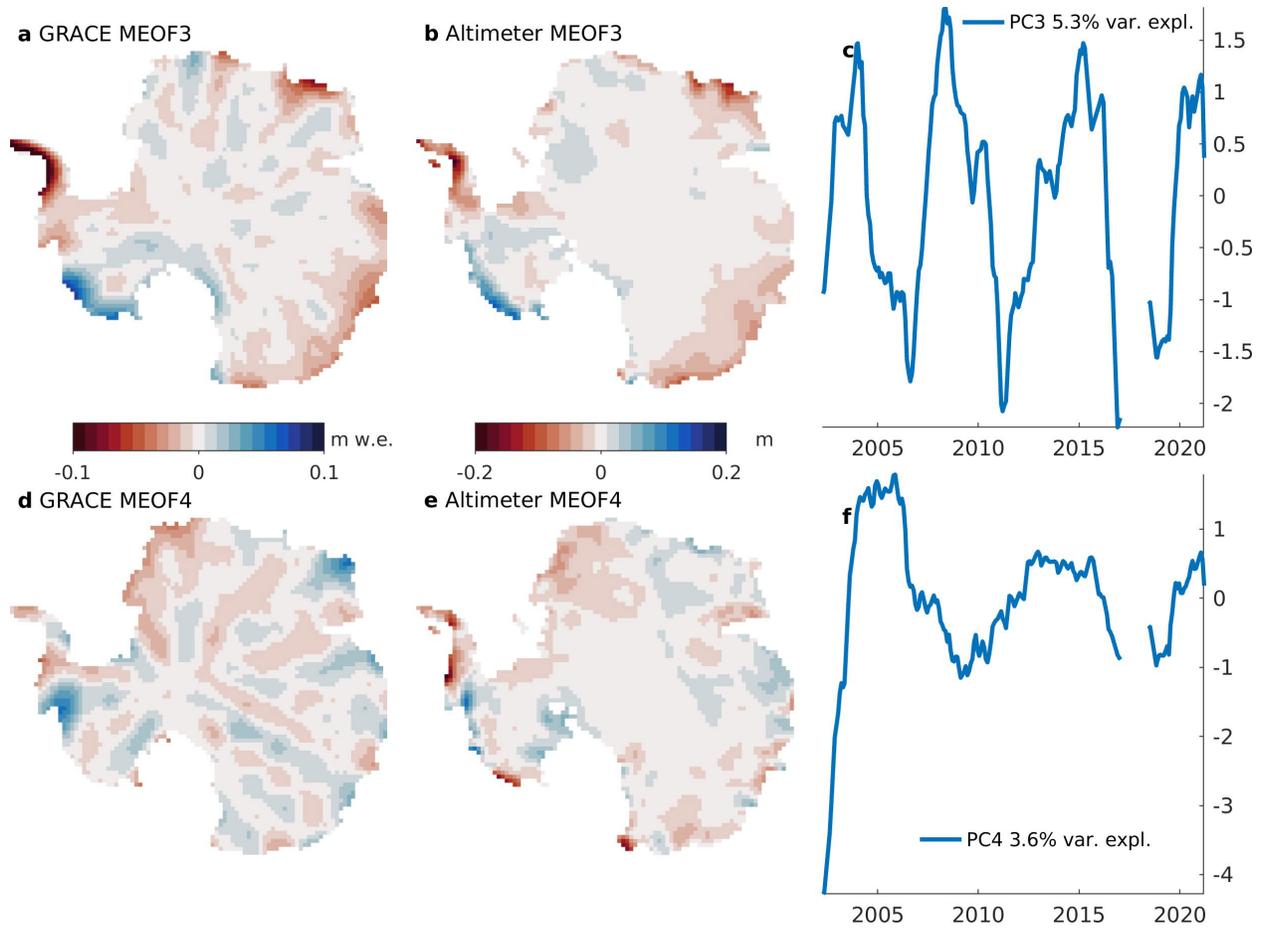
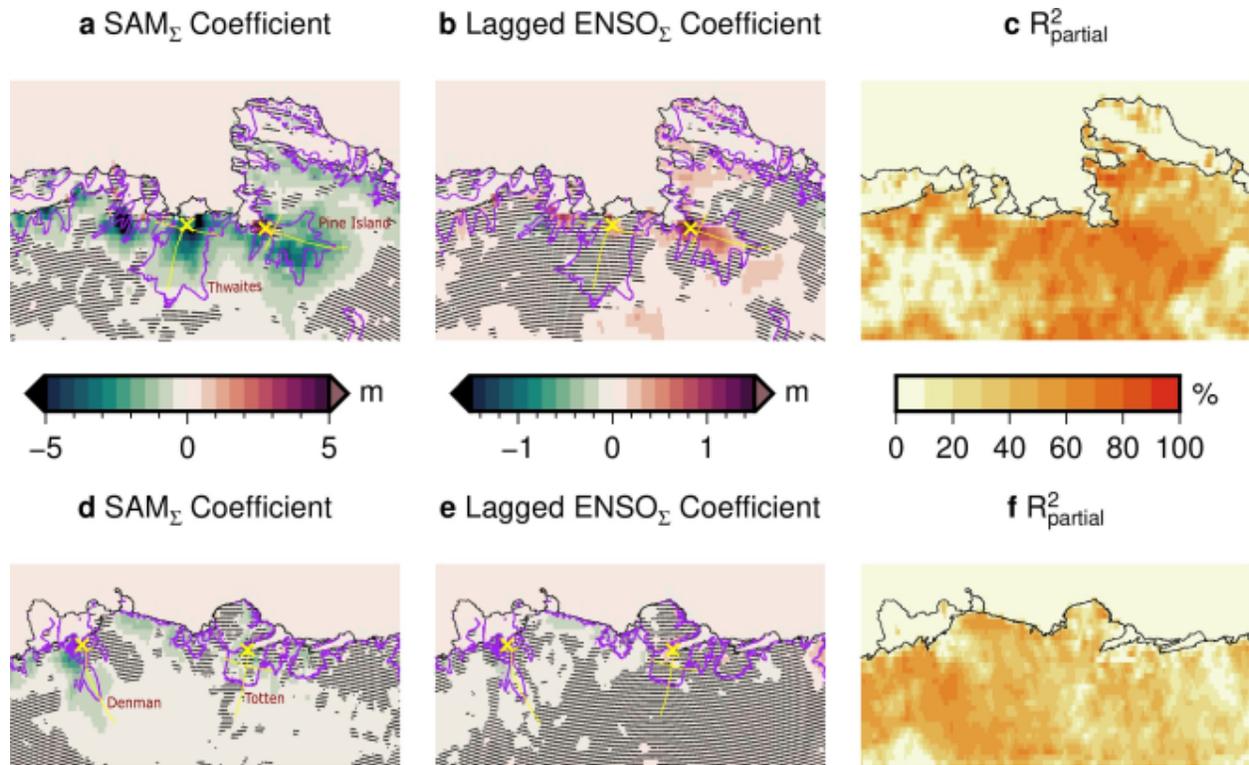
581
582583
584
585
586
587
588
589

Fig S3. The leading two modes of the detrended altimeter and GRACE data based on MVEOF. The EOFs are shown after scaling to reverse the effects of normalization. GRACE modes (a, d) are shown in units of meters of water equivalent (w.e.), and altimeter modes (b, e) shown in units of meters of ice elevation. Overlain on PC1 is the detrended SAM_Σ and -ASLP_Σ indices and overlain on PC2 is the lagged ENSO_Σ and -ASLλ_Σ. The variances explained by each mode are indicated in the legends of panels c and f.



590
591
592
593
594
595
596

Fig S4. Modes 3 and 4 of the MVEOF. Note the changes in color scale from Fig S3.



597
 598
 599
 600
 601
 602
 603
 604
 605

Fig S5. Detail of Figure 1 in the Amundsen Sea Embayment (top) and Denman-Totten glacier (bottom) regions. Thick black lines show the grounding line and thin black lines the ice shelf limits (Haran et al., 2014, updated 2019). Pink lines define the 100 m/yr speed contour (Rignot et al., 2017). Yellow lines mark the locations of ice stream profiles and yellow cross marks the location of the ice stream time series.

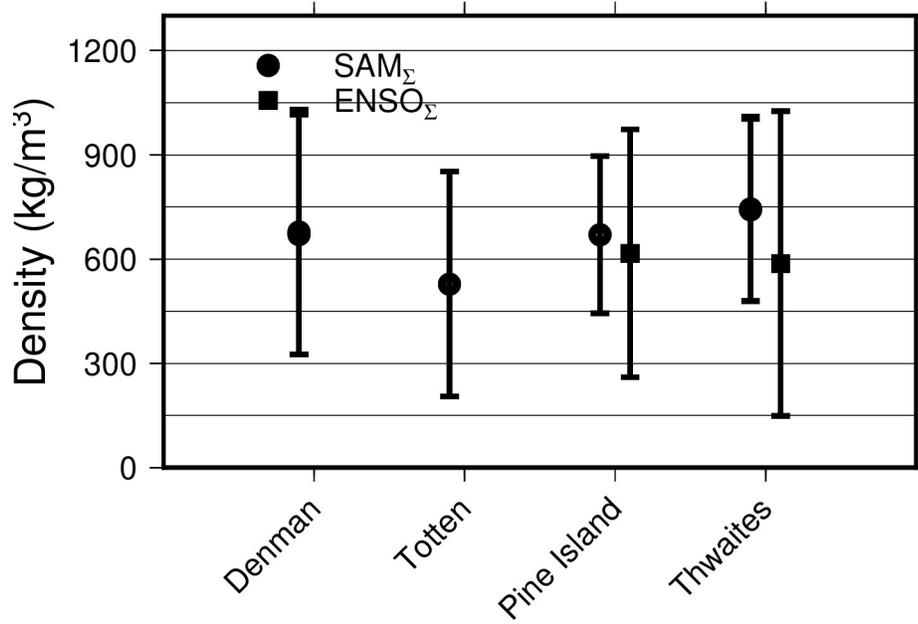
606
607608
609
610
611
612
613

Fig S6. Density of SAM Σ and ENSO Σ variability from the ratio of GRACE and smoothed-altimetry coefficients. Error bars reflect 1-sigma uncertainties. ENSO Σ values are not shown for Denman and Totten as they are too uncertain.

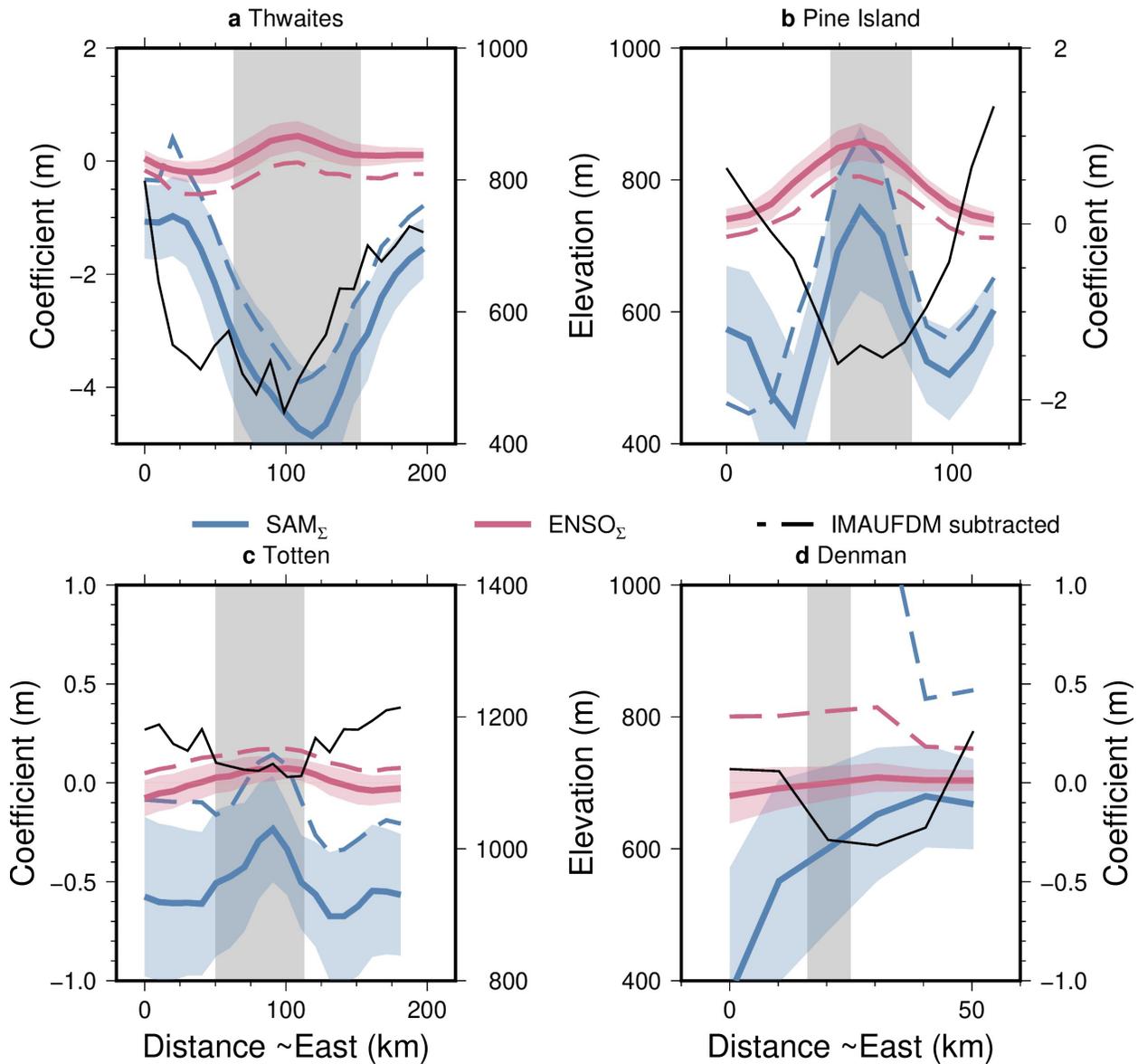
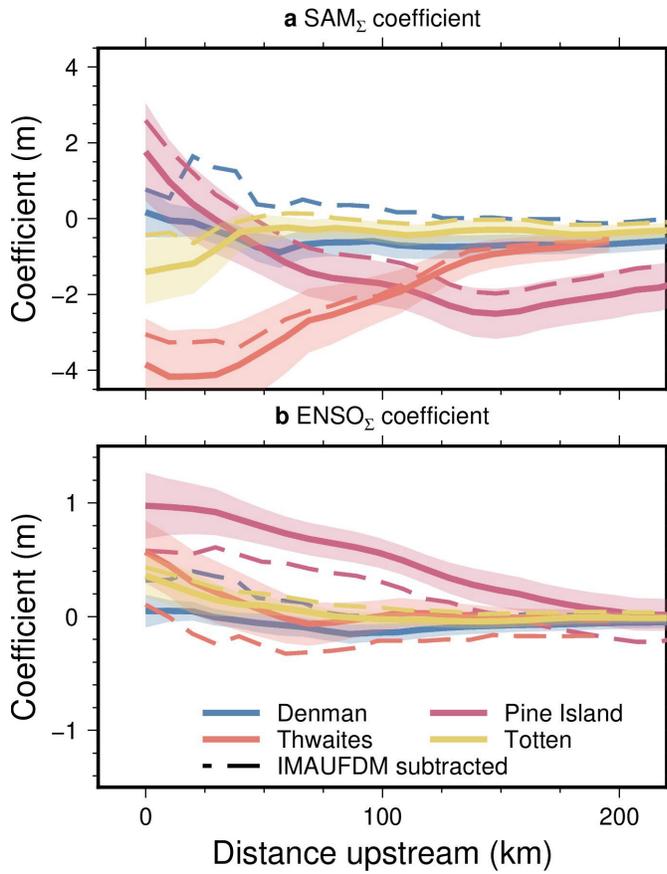
614
615616
617

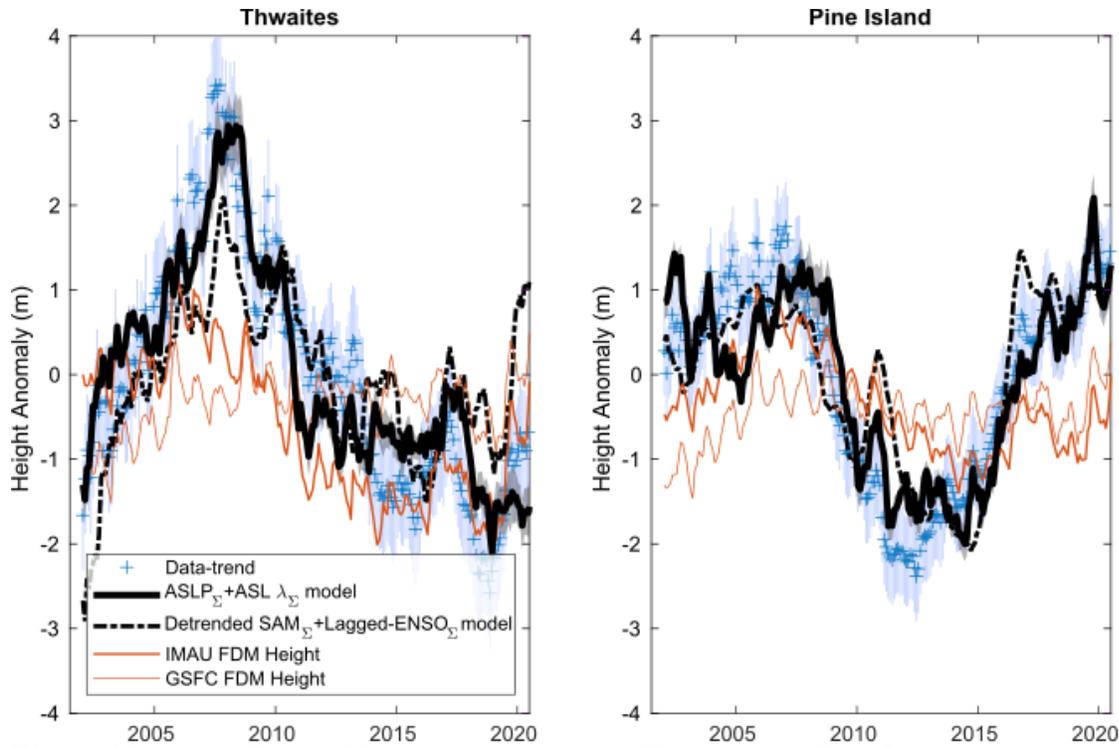
Fig S7. Cross-section profiles of SAM_{Σ} (a) and $ENSO_{\Sigma}$ (b) coefficients as a function of distance across Thwaites (a), Pine Island (b), Totten (c) and Denman (d) glaciers. Coefficients are shown as estimated before (colored solid line) and after (colored dashed line) subtraction of the IMAU FDM. The ice elevation extracted from the REMA v2.0 1 km mosaic (Howat, 2022) is shown as a black line (central axes). Colored shading indicates the 1-sigma confidence limits. The vertical grey box indicates the 100 m/yr limits of ice flow based on Measures v2. The locations of the profiles are shown in Figure S5 as yellow lines.

626
627



628
629
630
631
632
633
634
635

Fig S8. Centre-line profiles of SAM_{Σ} (a) and $ENSO_{\Sigma}$ (b) coefficients as a function of distance upstream of the grounding line of major ice streams. Coefficients are shown before (solid line) and after (dashed line) subtraction of the IMAU FDM. Shading indicates the 1-sigma confidence limits. The locations of the profiles are shown in Figure S5 as yellow lines.



636

637

638

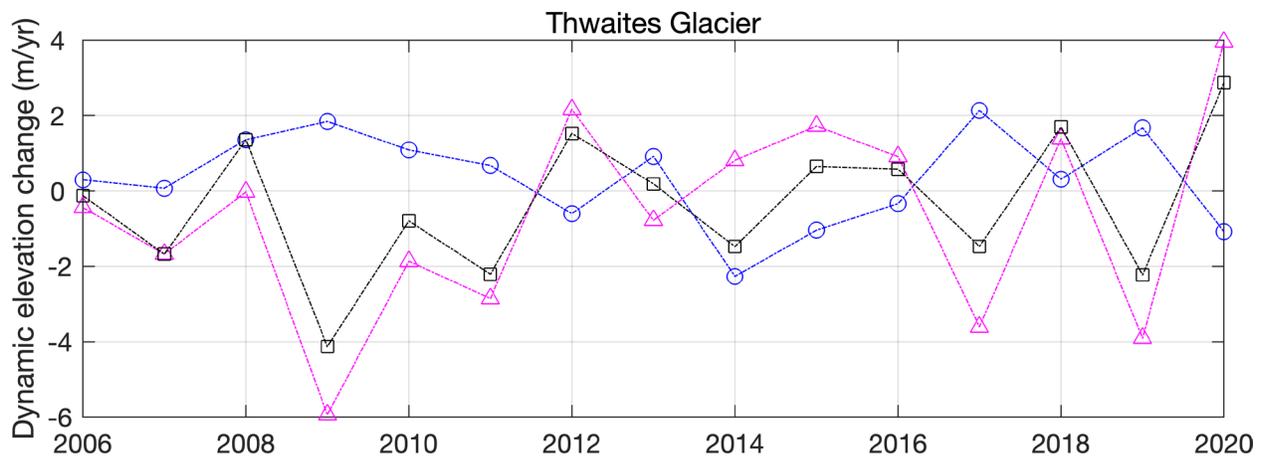
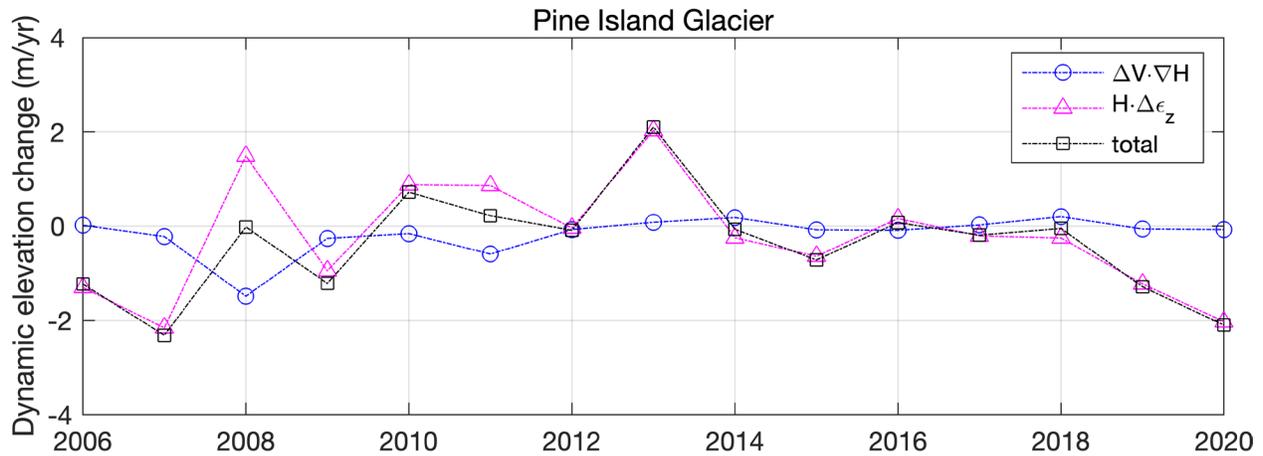
639

640

641

642

Figure S9. As for Figure 3 but showing only the Thwaites and Pine Island time series. Panels show (left axes) the detrended data and their 2-sigma uncertainties (blue pluses and error bars), outputs of two models of firn densification (FDM; brown lines), the best fitting models based on ASL (black solid line with 2-sigma uncertainty in gray shading) and SAM+ENSO (black dashed lines) terms.



643
 644 **Figure S10.** Dynamic elevation change anomalies (positive for gain) estimated from annual
 645 changes in satellite derived velocities for Pine Island Glacier (top) and Thwaites Glacier
 646 (bottom). These elevation change anomalies are calculated from the advective thickness change
 647 (blue circles) and strain (magenta triangle) tied to the observed change in velocity from one year
 648 to the next. The total elevation change (black squares) is the sum. Estimates are based on gridded
 649 surface velocities (v_x, v_y) in a 25 x 25 km box centered around glacier points used to show
 650 observed elevation changes in Figure 3 of main text. Other sources of elevation change are
 651 ignored. See Text S1 for details.

652

653 Table S1

654

655 Table S1. Location in Antarctic Polar Stereographic coordinates (EPSG:3031) of sites in Fig. 3
656 and Fig S9 and discussed in the main text.

| Glacier | X(m) | Y(m) |
|----------------|-------------|-------------|
| Thwaites | -1511702 | -463473 |
| Pine Island | -1583860 | -232513 |
| Totten | 2274795 | -988293 |
| Denman | 2496950 | -423352 |

657