

Accuracy and reproducibility of coral Sr/Ca SIMS timeseries in modern and fossil corals

Hussein R. Sayani¹, Kim M. Cobb¹, Brian Monteleone², and Heather Bridges¹

¹School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, Georgia, USA

²Department of Geology and Geophysics, Woods Hole Oceanographic Institute, Woods Hole, MA, USA

Corresponding author: Hussein R. Sayani (hsayani@gatech.edu)

Key Points:

- Diagenesis can impart large artifacts to coral Sr/Ca-based climate reconstructions even with minor levels of alteration
- SIMS and other microscale analytical techniques can be used to bypass diagenesis and recover Sr/Ca values from primary fossil coral material
- Transects of SIMS Sr/Ca analyses from an altered section of fossil coral are significantly correlated to SST in the early 20th century

Abstract

Coral Sr/Ca ratios provide quantitative estimates of past sea surface temperatures (SST) that allow for the reconstruction of changes in the mean state and climate variations, such as the El Niño-Southern Oscillation, through time. However, coral Sr/Ca ratios are highly susceptible to diagenesis, which can impart artifacts of 1-2 °C that are typically on par with the tropical climate signals of interest. Microscale sampling via Secondary Ion Mass Spectrometry (SIMS) for the sampling of primary skeletal material in altered fossil corals, providing much-needed checks on fossil coral Sr/Ca-based paleotemperature estimates. In this study, we employ a set modern and fossil corals from Palmyra Atoll, in the central tropical Pacific, to quantify the accuracy and reproducibility of SIMS Sr/Ca analyses relative to bulk Sr/Ca analyses. In three overlapping modern coral samples, we reproduce bulk Sr/Ca estimates within $\pm 0.3\%$ (1 σ). We demonstrate high fidelity between 3-month smoothed SIMS coral Sr/Ca timeseries and SST ($R = -0.5$ to -0.8 ; $p < 0.5$). For lightly-altered sections of a young fossil coral from the early-20th century, SIMS Sr/Ca timeseries reproduce bulk Sr/Ca timeseries, in line with our results from modern corals. Across a moderately-altered section of the same fossil coral, where diagenesis yields bulk Sr/Ca estimates that are 0.6mmol too high (roughly equivalent to -6 °C artifacts in SST), SIMS Sr/Ca timeseries track instrumental SST timeseries. We conclude that 3-4 SIMS analyses per month of coral growth can provide a much-needed quantitative check on the accuracy of fossil coral Sr/Ca-derived estimates of paleotemperature, even in moderately altered samples.

1. Introduction

Surface coral skeletons are one of the few marine archives capable of providing absolutely-dated, sub-annually resolved records of past tropical climate and oceanic variability. Cores from living coral colonies are increasingly used to reconstruct past variability in surface temperature (SST), salinity, and other oceanic parameters across recent centuries (e.g. *Nurhati et al.*, 2009; 2011; *Vásquez-Bedoya et al.*, 2012; *Sanchez et al.*, 2016; *Murty et al.*, 2017; *Jimenez et al.*, 2018; *Rodriguez et al.*, 2019), and cores from fossil corals can extend such reconstructions through the Holocene and beyond (e.g. *Beck et al.*, 1997; *Linsley*, 2000; *Tudhope et al.*, 2001; *Cobb et al.*, 2003; *DeLong et al.*, 2010; *Cobb et al.*, 2013; *McGregor et al.*, 2013; *Felis et al.*, 2014; *Toth et al.*, 2015; *Grothe et al.*, 2019; *Felis*, 2020). While fossil corals provide unique snapshots into past seasonal-to-decadal climate variability, estimates of mean temperature change from glacial-aged or older fossil corals are often $\sim 2\text{-}3^\circ\text{C}$ cooler than those from other marine archives (*Gagan et al.*, 2012). Analyses of similarly-aged corals reveal that post-depositional alteration of the coral skeleton, or diagenesis can easily introduce cool artifacts in coral-based reconstructions (e.g. *Cohen and Hart*, 2004; *Allison*, 2005). As the number of fossil coral-based mean climate reconstructions increases, there is a pressing need to better constrain their accuracy, especially if they are to be included in large multi-proxy reconstruction and data assimilation efforts (e.g. *Emile-Geay et al.*, 2013a; 2013b; *Hakim et al.*, 2016; *Tardif et al.*, 2019; *Sanchez et al.*, 2021).

Diagenesis typically occurs as dissolution and/or precipitation of secondary cements, both of which can produce significant artifacts in coral-based paleoclimate reconstructions. A majority of existing coral-based climate reconstructions utilize either oxygen isotope ratios (^{18}O), which tracks both changes in SST and the hydrologic budget (*Epstein et al.*, 1953; *Weber and Woodhead*, 1972; *Hasson et al.*, 2013; *Conroy et al.*, 2014; 2017), and/or strontium-to-calcium ratios (Sr/Ca), which primarily reflect changes in SST (*Weber*, 1973; *Smith et al.*, 1979; *Beck et al.*, 1992). Secondary aragonite cements, which have more enriched ^{18}O and higher Sr/Ca than coral aragonite, produce cool artifacts in paleo-SST reconstructions (*Enmar et al.*, 2000; *Müller et al.*, 2001; *Ribaudo-Laurenti et al.*, 2001; *Quinn and Taylor*, 2006; *Hendy et al.*, 2007). On the other hand, secondary calcite cements are typically lower in ^{18}O and Sr/Ca relative to coral aragonite, producing warm artifacts in paleo-SST reconstructions (e.g. *McGregor and Gagan*, 2003). The effects of skeletal dissolution on coral-based paleoclimate records remain poorly characterized. However, evidence suggests that some skeletal elements, such as the centers of calcification (COCs), which are higher in Sr/Ca (*Cohen et al.*, 2001), dissolve more easily, producing warm artifacts in paleo-SST reconstructions (e.g. *Hendy et al.*, 2007).

A variety of tools such as x-ray imaging, x-ray diffraction (XRD), scanning electron microscopy (SEM), and light microscopy, are used to screen coral samples for diagenetic alteration. With the exception of x-ray imaging, these techniques are destructive, and as such, screening is conducted a few millimeters to centimeters away from the section of the core used for geochemical analysis. This strategy works well in heavily altered corals, where the alteration is more consis-

tent across large sections of a core. In light to moderately altered corals, however, alteration is more heterogeneously distributed (e.g. *Hendy et al.*, 2007; *Sayani et al.*, 2011) and can escape detection, leading to diagenetic material being inadvertently included in coral powders drilled for conventional bulk or “mm-scale” Sr/Ca or ^{18}O analyses. The inclusion of even 1% of secondary cements in coral powders can introduce warm/cool artifacts of 1-2 ° C in coral-based temperature reconstructions (*Allison et al.*, 2007; *Sayani et al.*, 2011). While fossil corals are routinely screened for diagenesis, modern corals seldom are even though alteration can occur while colonies are still living (e.g. *Nothdurft et al.*, 2005; *Hendy et al.*, 2007; *Nothdurft et al.*, 2007; *Nothdurft and Webb*, 2008). The prevalence of diagenesis in both modern and fossil corals highlights the need for both rigorous screening and the development of techniques capable of extracting reliable constraints on mean climate from altered corals.

As a microscale analytical technique, secondary ion mass spectrometry (SIMS) has the potential to bypass diagenetic phases by targeting only better-preserved areas of the coral skeleton for geochemical analyses. These techniques capitalize on the fact that secondary cements and/or low levels of dissolution leave much of the original coral skeleton geochemically intact (e.g. *Cohen and Hart*, 2004; *Allison*, 2005; *Sayani et al.*, 2011). Early application of SIMS technique to altered fossil corals from the western tropical Pacific demonstrate that targeted ion-microprobe analyses of pristine material in altered fossil corals yield mean temperature estimates that are more consistent with other reconstructions from the region (*Cohen and Hart*, 2004; *Allison*, 2005). Despite these early successes, microscale analytical techniques like SIMS have remained underutilized in paleoclimate reconstruction for several reasons. For one, few SIMS facilities exist, given the cost and complexity associated with the equipment. Second, SIMS coral Sr/Ca analyses are roughly 100 times more time-consuming and expensive than standard ICP-OES analyses. Moreover, the composition of biogenic carbonates is highly variability at microscales (*Allison*, 1996; *Hart and Cohen*, 1996; *Sinclair et al.*, 1998; *Meibom*, 2003; *Rollion-Bard et al.*, 2003; *Gagnon et al.*, 2007; *Allison and Finch*, 2010b), potentially obscuring the climate signals of interest. In corals, these large microscale fluctuations in composition likely reflect biological processes rather than environmental variability (e.g. *Meibom*, 2003; *Meibom et al.*, 2008), limiting the utility of SIMS for reconstructing daily- or weekly-resolved climate records. Averaging or smoothing multiple SIMS measurements can resolve monthly SST variability (e.g. *Allison and Finch*, 2004; *Sinclair*, 2005; *Allison and Finch*, 2009), however, additional work is required to establish best practices for applying SIMS to fossil corals and quantify the uncertainties of the resulting reconstructions.

In this study, we measure SIMS Sr/Ca in three modern corals that overlap in time and one ~100yr-old fossil coral from Palmyra Atoll (6 ° N, 162 ° W) in the central tropical Pacific. We use the modern corals to devise a SIMS sampling strategy to yield monthly-resolved timeseries, and then test their reproducibility across all three modern corals and their fidelity to SST. Guided by SIMS Sr/Ca sampling protocols developed in our modern corals, we analyze SIMS Sr/Ca data

from three altered sections of a relatively young fossil coral that grew during the early-20th century. We compare the fossil coral SIMS Sr/Ca timeseries to both bulk Sr/Ca analyzed in the same coral, and to instrumental SST, to assess the accuracy of SIMS coral Sr/Ca across a gradient of diagenesis. Lastly, we synthesize our findings into a discussion of the strengths and weaknesses of employing SIMS-based coral Sr/Ca analyses for paleo-SST reconstruction in fossil corals.

2. Methods

We present new SIMS Sr/Ca records from three *Porites* spp. modern corals, as well as bulk and SIMS Sr/Ca records from a *Porites* spp. fossil coral, collected from Palmyra Atoll (5°53'N, 162°5'W). Modern corals PM, PM1, and PM5 are used to test and refine our SIMS methodology as they've been previously vetted for diagenesis and used to develop monthly coral Sr/Ca and oxygen isotope records (Cobb *et al.*, 2001; Nurhati *et al.*, 2009; 2011; Sayani *et al.*, 2019). We also present new bulk Sr/Ca timeseries and micro-scale SIMS Sr/Ca timeseries from Palmyra fossil coral NB9, which extends from 1908 to 1935 (Cobb *et al.*, 2003).

1. Bulk Sr/Ca analyses and diagenesis screening

Coral cores were cut into ~1-cm thick slices and prepared for geochemical analyses using standard procedures. For bulk Sr/Ca analyses, coral powders were drilled using a 1mm bit at 1mm intervals along a transect parallel to the primary growth axis. Coral powders were digested in 2% trace-metal grade HNO₃ and analyzed using a Horiba Jobin-Yvon Ultima 2C Inductively Coupled Plasma Optical Emission Spectrometer (ICP-OES) located at Georgia Tech using procedures outlined in Sayani *et al.*, (2019). Analytical precisions for bulk Sr/Ca measurements are < ±0.1% or ±0.09mmol/mol (1).

Potentially altered horizons in core NB9 were identified through visual screening of coral xrays for density anomalies and the existence of exceptionally noisy intervals in the bulk Sr/Ca record, and flagged for subsequent analysis via Scanning Electron Microscopy (SEM). Samples for SEM imaging were prepared by extracting small (2-4mm) coral chunks near the bulk Sr/Ca transect and gold coating. SEM images were obtained using either a Hitachi S-800 field emission gun SEM or a LEO 1530 thermally-assisted field emission SEM located at Georgia Tech.

1. Coral sample preparation for SIMS

Ion microprobe measurements were made on a series of continuous, doubly-polished, thin-section rounds. Thin sections were prepared by cutting 4-6cm long segments from each coral slab, typically within 1-3 cm from the transect used for bulk Sr/Ca analyses. Each segment was lightly scored on the back and snapped by hand into ~1cm pieces to preserve continuity across the sampling surface. Coral segments were then embedded in epoxy, sectioned at similar depths, polished and mounted onto 1-inch rounds, and then polished again to

a final thickness of at least ~ 30 μm . Prior to SIMS analyses, each thin section was gold coated and then imaged using both a transmitted and a reflected light microscope. Transmitted light microscope images were used to identify the location of COCs, secondary cements, and any imperfections in the thin section. Paired reflected light images were used to map continuous sampling paths and navigate across the thin section when loaded onto the SIMS.

1. SIMS coral Sr/Ca analyses

SIMS Sr/Ca measurements were made using a Cameca IMS-1280 ion microprobe, located at the Northeast National Ion Microprobe Facility (NENIMF) at the Woods Hole Oceanographic Institute, equipped with a $\sim 6\text{nA}$ O^- ion beam, accelerated at 10keV. SIMS analyses were made using an ~ 20 m diameter beam, targeting pristine sections spots on aragonite fibers located between the COCs and the edge of the trabeculae. Following a pre-analysis burn time of 90s to remove gold coating, ^{42}Ca and ^{88}Sr were measured using a -80eV energy filter across 10 cycles lasting 10s and 15s each, respectively. As measurements were made using a single collector, we use dead-time interpolation scheme to correct for drift or loss in signal intensity across the 10 cycles. Analytical uncertainty for SIMS Sr/Ca analyses, estimated as the standard error of $^{88}\text{Sr}/^{42}\text{Ca}$ measurements across 9 dead-time corrected cycles, ranges from ± 0.01 to $\pm 0.05\text{mmol/mol}$ (1). Mass ^{42}Ca and ^{88}Sr intensities were converted to concentration using calibration curves built by measuring three well-characterized standards (*Gaetani and Cohen*, 2006; *Holcomb et al.*, 2009; *Gaetani et al.*, 2011) at the beginning and end of each day. Standards include an OKA Carbonatite crystal, a calcite crystal (Blue-0875) and an aragonite crystal (AG1) whose Sr/Ca ratios span 0.56-19.3mmol/mol. The average reproducibility of Sr/Ca measurements on each standard, across four sessions between 2012 and 2016, are $\pm 1.4\%$ for OKA, $\pm 2.4\%$ for AG1, and $\pm 2.3\%$ for Blue-0875 per session.

For SIMS Sr/Ca measurements, we primarily targeted skeletal material between the centers of calcification and the edge of each trabecula. At m-scales, coral skeletal elements (e.g. the thecal wall, septa, dissepiments) are composed of dense bundles of crystalline aragonite fibers which radiate out from granular, sub-micron “centers of calcification” (COCs) (see review in *Rabier et al.*, 2008). While both the COCs and aragonite fibers likely track environmental variability (*Allison and Finch*, 2004), they have distinct geochemical compositions with COCs generally being more enriched in Sr/Ca and other trace elements (e.g. *Cohen et al.*, 2001; *Allison and Finch*, 2004; *Meibom et al.*, 2008). As the COCs are too small to be discretely measured with ion beam sizes >10 m, we focus our analyses on the aragonite fibers $>95\%$ of the coral skeleton.

To generate smoothed SIMS Sr/Ca time series, we remove any data points that are ± 2 standard deviations outside the mean of all analyses from each sample, and then apply an 11-point running mean filter. The 2 filter primarily serves to remove measurements that may have inadvertently included material from COCs (visually confirmed by examining thin sections under a light microscope following SIMS analyses), which are systematically higher in Sr/Ca and

can bias the resulting time series. The 11-month running mean filter averages SIMS Sr/Ca measurements across a ~2.5month time period that consistently provides the best match with bulk Sr/Ca time series from each coral. Uncertainty in smoothed SIMS Sr/Ca is estimated by calculating the standard error of all points used to calculate the running mean at each time step. Initial age models were derived by mapping the SIMS measurements to core depth and then interpolating ages using the age-depth relationship previously developed for Sr/Ca records from each core. To account for changes in extension rate changes between the SIMS and ICP-OES transections, and sample loss between contiguous thin sections, we slightly finetuned age models for SIMS data from PM and PM1 by shifting some points by up to 1mm (equivalent to 2-3 weeks in time).

3. Results

3.1. Sampling strategy for deriving SIMS Sr/Ca records

Centers of calcification (COCs) contain significantly more Sr/Ca than aragonite fibers, which make up >90% of the coral skeleton, and are difficult to detect on gold-coated thin sections. To develop a filter for removing any SIMS datapoints that might include a COC, we measured spots with COCs and adjacent aragonite fibers at 7 locations in core PM. SIMS measurements on aragonite fibers yield a mean Sr/Ca value of 8.95 ± 0.02 mmol/mol (1, n=14), consistent with the average bulk Sr/Ca values across this horizon (8.95 ± 0.04 mmol/mol, 1, n=3). In contrast, SIMS analyses that include COCs yield an average Sr/Ca value of 9.15 ± 0.03 mmol/mol (1; n=7), two standard deviations greater than the mean Sr/Ca of the adjacent aragonite fibers. As such, we exclude any SIMS Sr/Ca measurements that are ± 2 standard deviations outside the mean Sr/Ca value of each coral from subsequent analyses.

To determine the minimum number of SIMS analyses needed to resolve a seasonal cycle, we measured SIMS Sr/Ca at ~200 m intervals across an 18mm-long transect in core PM, targeting skeletal structures where the location of COCs were more apparent in thin sections. We observe large point-to-point variability of ~0.6mmol/mol across this initial SIMS time series, consistent with results from previous studies measuring coral Sr/Ca at microscales (*Sinclair et al.*, 1998; *Cohen et al.*, 2001; *Meibom et al.*, 2008; *Allison and Finch*, 2009). However, the average of these SIMS Sr/Ca measurements (8.99 ± 0.02 mmol/mol, n=80) is indistinguishable from that of bulk Sr/Ca measurements across the same interval (8.97 ± 0.02 mmol/mol, n=20). Applying an 11-point running mean filter, equivalent to ~2.5 months, yields a timeseries with seasonal variability similar to that of bulk Sr/Ca. By subsampling this initial dataset at coarser resolutions, we find that a minimum sampling resolution of ~400 m is needed to resolve seasonal variability in our SIMS data. Given core PM's extension rate of 16-18mm/year, this translates to roughly one SIMS Sr/Ca analysis for every week of coral growth.

3.2. Reproducibility of SIMS Sr/Ca analyses across overlapping modern corals

To assess the reproducibility of SIMS analyses between corals, we measured Sr/Ca between 1985-1989 in three modern corals from Palmyra atoll. Similar to observations from core PM, SIMS Sr/Ca measurements from the 1986-1989 sections of cores PM5 (Figure 3B) and PM1 (Figure 3C) are also characterized by high-amplitude fluctuations of ~ 0.6 mmol/mol to ~ 0.7 mmol/mol, respectively. No significant correlations are observed between raw SIMS Sr/Ca timeseries from cores PM, PM1, and PM5. However, SIMS analyses do yield an average Sr/Ca value that is within error of the average bulk Sr/Ca from each core. SIMS Sr/Ca measurements from the 1986-1989 section of core PM1 yield an average value of 9.02 ± 0.01 mmol/mol ($n=114$), consistent with the 9.04 ± 0.01 mmol/mol ($n=45$) average of bulk Sr/Ca from across the same time period. Similarly, SIMS analyses from the 1986-1989 horizon of PM5 yield an average Sr/Ca value (8.89 ± 0.01 mmol/mol; $n=84$) that is consistent with bulk Sr/Ca from this core (8.86 ± 0.01 mmol/mol; $n=46$). The average SIMS Sr/Ca values of these cores are offset by ± 0.07 mmol/mol (1 σ), similar to the observed ± 0.08 mmol/mol (1 σ) offsets in bulk Sr/Ca records among these cores (Sayani *et al.*, 2019).

Smoothed SIMS Sr/Ca timeseries resolve seasonal variability observed in both bulk Sr/Ca from each core and instrumental SST at our site. Smoothed SIMS Sr/Ca timeseries from cores the 1985-1989 sections of PM and PM1 (Figure 3) are well correlated with monthly SST ($R = -0.5$ and -0.7 , $P < 0.05$, respectively) and their bulk Sr/Ca timeseries ($R = 0.4$ and 0.6 , $P < 0.05$, respectively). In contrast, smoothed SIMS Sr/Ca from core PM5 timeseries diverges considerably from both bulk Sr/Ca ($R = -0.4$, $P < 0.05$) and SST ($R = 0.5$, $P < 0.05$) between 1986-1987 (Figure 3B). Inspection of thin sections from PM5 after the SIMS session reveals that several spots sampled across this horizon included COCs, which bias the smoothed timeseries towards higher Sr/Ca values. However, smoothed PM5 SIMS data from the year before and after this interval (1986-1987 and 1988-1989) reproduce variability observed in bulk Sr/Ca and SST bulk Sr/Ca. We also tested out this methodology to a second section of PM1 that covers 1995-1997 (Figure 3D), just before the 1997/98 El Niño event. Across this horizon, smoothed SIMS Sr/Ca reproduces both bulk Sr/Ca ($R = 0.5$, $P < 0.05$) and SST ($R = -0.8$, $P < 0.05$).

3.3. Bulk vs. targeted SIMS Sr/Ca analyses in an altered fossil coral

Bulk Sr/Ca from fossil coral NB9 generally tracks instrumental SST ($R = -0.6$, $p < 0.05$) and overlapping bulk Sr/Ca from modern core PM ($R = 0.5$, $p < 0.05$) between 1917-1935. We do not observe any large departures of coral Sr/Ca from SST variations across this horizon of NB9, however, there are several intervals (e.g. around 1918, 1922, 1928, and 1931) where individual bulk Sr/Ca datapoints from both NB9 and PM deviate from each other and observed SST by up to ~ 1 °C. SEM images from these sections of NB9 reveal smooth skeletal surfaces, such as those seen in modern coral PM (Figure 5A), with a few patches of small (< 5 m-long) secondary aragonite cements and minor dissolution (Figure 5B, C). In contrast, the bulk NB9 Sr/Ca record from 1911-1917 is characterized by large excursions of up to -0.6 mmol/mol away from the baseline defined by

SST variations during this time. Across this ~6yr period, bulk NB9 Sr/Ca suggests temperatures were 1-6 °C cooler than instrumental observations. SEM images from this horizon of the core reveal more continuous coverage of larger (~10 m-long) secondary aragonite cements (Figure 5D), which are ~30% more enriched in Sr/Ca relative to coral aragonite (*Sayani et al.*, 2011) and produce cool artifacts in Sr/Ca-based temperature reconstructions.

We measured SIMS Sr/Ca across three horizons of NB9 with increasing levels of alteration. We first applied the methodology developed using the modern corals above to two lightly-altered sections of NB9: (i) from 1931-1933, where the skeleton is well preserved and bulk Sr/Ca matches SST, and (ii) from 1920-1922, where the skeleton contains trace alteration and bulk Sr/Ca contains cool excursions of up to ~1 °C. SIMS Sr/Ca analyses across the 1931-1933 section of NB9 produce an average value of 9.10 ± 0.02 mmol/mol (n=78), equivalent to the average of bulk NB9 Sr/Ca (9.08 ± 0.01 mmol/mol; n=33). Both bulk and SIMS Sr/Ca from NB9 yield similar temperature estimates of 27.4 ± 0.05 °C and 27.2 ± 0.2 °C, respectively, consistent with the instrumental SST average of 27.4 °C across this interval. Similarly, SIMS analyses from the 1920-1922 section of NB9 yield an average value of 9.09 ± 0.02 mmol/mol (n=73), consistent with bulk NB9 Sr/Ca (9.07 ± 0.01 mmol/mol; n=31). Once again, bulk and SIMS Sr/Ca yield temperature estimates (27.3 ± 0.2 °C and 27.5 ± 0.1 °C, respectively) that are consistent with average instrumental SST across this period (27.7 °C).

Smoothed SIMS coral Sr/Ca time series across 1931-1933 and 1920-1922 from NB9 generally track overlapping bulk Sr/Ca and instrumental SST, albeit with a few discrepancies. We measured Sr/Ca across these two sections of NB9 using a sampling resolution of ~400 m, which worked well for our faster growing modern corals (16-18mm/yr), but did not provide enough analyses per month from the slower (10-14mm/yr) growing NB9 to fully capture seasonal variability. For subsequent SIMS analyses from NB9, we measure 3-4 spots per month of coral growth (i.e. between consecutive dissepiments; DeCarlo & Cohen, 2017) instead of using a fixed sampling interval to account for the slower extension rate. Nonetheless, SIMS and bulk Sr/Ca from NB9, as well as bulk Sr/Ca from PM, suggest temperatures in early-1922 that are much cooler than instrumental observations. However, agreement between average SIMS and bulk Sr/Ca measurements across both of these sections of NB9 suggests that the low levels of diagenetic alteration observed here do not have a noticeable impact on bulk Sr/Ca measurements.

Finally, we apply SIMS to the moderately-altered 1911-1914 section of NB9, where SEM images show more continuous diagenesis (Figure 6) and bulk Sr/Ca contains excursions of up to -6 °C. SIMS Sr/Ca analyses across this horizon yield an average value of 9.06 ± 0.02 mmol/mol (n=119) or 27.3 °C, consistent with the average instrumental SST of 27.6 °C observed at Palmyra across this interval. In contrast, bulk Sr/Ca measurements yield an average value of 9.27 ± 0.02 mmol/mol (n=46), ~3 °C cooler than both SIMS Sr/Ca-derived and instrumental SST, likely reflecting contamination by high-Sr/Ca secondary pre-

cipitates. Moreover, by sampling only preserved sections of the coral skeleton between 1911-1914, we derive a smoothed SIMS Sr/Ca timeseries that reproduces seasonal SST variability during this interval ($R=-0.9$, $p<0.05$; Figure 4).

4. Discussion

4.1. Reproducibility of SIMS analyses in modern corals

Early application of SIMS to altered fossil corals highlighted that microscale analytical techniques could circumvent analysis of diagenetic phases (e.g. *Cohen and Hart*, 2004; *Allison*, 2005; *Sayani et al.*, 2011), but none of these studies thoroughly benchmarked modern coral SIMS Sr/Ca measurements against SST. Here, SIMS coral Sr/Ca timeseries from three time-overlapping modern corals demonstrate that SIMS provides mean Sr/Ca values and monthly-resolved timeseries that are reproducible across coral colonies and coherent with instrumental SST. Consistent with previous studies featuring microscale trace-element and isotope measurements in corals, our SIMS Sr/Ca measurements are characterized by large, high-amplitude fluctuations of 0.6 to 0.7mmol/mol in SIMS Sr/Ca that are not correlated between time-overlapping samples. These fluctuations exceed diurnal variability at our site (1.5°C or 0.1 mmol/mol) and are an order of magnitude too large to be explained by analytical error (± 0.01 to ± 0.03 ; 1). As such, it is likely that they reflect either non-linear skeletal accumulation (*Barnes et al.*, 1995; *Cohen and Sohn*, 2004; *Sinclair*, 2005) or non-environmental factors that impact the biomineralization process (e.g. *Juillet-Leclerc et al.*, 2009; *Allison and Finch*, 2010a). Notably, we also find that a 2.5-month smoothing of the SIMS coral Sr/Ca data isolates a seasonal cycle that mirrors seasonal SST variability at the site, consistent with findings from other studies (e.g. *Sinclair et al.*, 1998; *Allison and Finch*, 2004; 2009). Thus, while individual SIMS Sr/Ca analyses may not yield meaningful daily- or weekly-resolved SST records, smoothing SIMS measurements across several months of coral growth can provide robust sub-seasonal SST reconstructions.

4.2. Application of SIMS in altered fossil corals

We apply the SIMS methodology developed using modern corals to a fossil coral with varying degrees of alteration to assess the impact of low alteration levels on the fidelity of bulk Sr/Ca analyses. Targeted SIMS analyses of preserved coral aragonite across two lightly-altered sections of fossil coral NB9 demonstrate that low-levels of alteration have a negligible impact on the fidelity of bulk Sr/Ca records. Across two lightly-altered horizons of fossil coral NB9 (1931-1933 and 1920-1922), targeted SIMS Sr/Ca analyses reproduce the mean and seasonal variability observed in bulk Sr/Ca from both NB9 and Palmyra modern coral PM (Figure 4). This consistency among timeseries generated by two different Sr/Ca analytical techniques in two different corals implies that the low levels of secondary cement coverage observed in one of the corals have no impact on bulk Sr/Ca across this horizon. As such, the relatively small $\sim 1^{\circ}\text{C}$ discrepancies between bulk Sr/Ca-derived SST and instrumental SST in 1922 may reflect vital effects. Where both bulk Sr/Ca records show similar deviations from SST,

the discrepancies may reflect the fact that gridded SST products contain large uncertainties in the early-20th century due to limited SST observations across much of the tropical Pacific (e.g. *Deser et al.*, 2010; *Tokinaga et al.*, 2012).

Targeted SIMS analyses across the moderately-altered horizon of NB9 allow us to completely bypass secondary cements present in this section of the coral, yielding Sr/Ca-derived temperatures that agree with instrumental observations. These results are consistent with previous studies that have used SIMS to derive more realistic mean temperature estimates from ancient altered fossil corals (e.g. *Cohen and Hart*, 2004; *Allison*, 2005; *Sayani et al.*, 2011). However, building on this previous work, we use SIMS to derive the first, realistic monthly-resolved SST timeseries from an altered section of a fossil coral. Smoothed SIMS Sr/Ca data from this more altered horizon of NB9 agrees with both published modern coral Sr/Ca data from Palmyra (*Nurhati et al.*, 2011) and instrumental SST. Thus, while diagenesis can produce significant artifacts in temperature reconstructions based on bulk Sr/Ca measurements, our results highlight that SIMS is a powerful tool that can be used to verify the accuracy of bulk Sr/Ca measurements and derive more accurate paleo-SST reconstructions.

4.3. Practical implications for coral-based SST reconstructions

Most if not all fossil corals are subject to some level of diagenetic alteration, motivating the development of an evidence-based protocol for acquiring accurate paleotemperature estimates from this archive. This is especially true because, except in the most extreme cases, diagenesis can escape detection via standard screening protocols. While alteration is often associated with older fossil corals, diagenesis routinely occurs in relatively young corals, such as the <100yr-old fossil coral featured in this study, and is also observed in cores collected from living colonies (e.g. *Nothdurft et al.*, 2007; *Nothdurft and Webb*, 2008; *Rabier et al.*, 2008; *Nurhati et al.*, 2011; *Sayani et al.*, 2011). As the composition of secondary cements is significantly different from that of coral aragonite, even light to moderate levels of alteration can produce significant errors reconstructions that rely on bulk analytical techniques. Moreover, given the scarcity of fossil corals, researchers often attempt to recover reconstructions from corals that grew during a given time period of interest, even if they have experienced moderate to significant alteration. Thus, given the pervasiveness of diagenesis and its substantial impact on accuracy of fossil coral Sr/Ca-based paleotemperature estimates, rigorous screening for diagenesis would ideally be combined with micro-scale SIMS Sr/Ca as an independent constraint on paleotemperature.

SIMS Sr/Ca of coral skeletons allows for the selective analysis of less altered skeletal material, but the cost- and labor-intensive nature of these analyses requires a sampling strategy that maximizes the accuracy of the resulting paleotemperature reconstructions with as few analyses as possible. Our results suggest that the most strategic use of SIMS Sr/Ca involves analysis of discrete multi-year windows to check the accuracy of longer bulk coral Sr/Ca timeseries generated by ICP-OES. In doing so, it is important to consider how to achieve the following goals with a minimum number of SIMS analyses: collecting enough

measurements in a given time interval to overcome microscale compositional heterogeneity in coral aragonite, and collecting timeseries of SIMS measurements that will provide an apples-to-apples comparison to bulk Sr/Ca measurements. By sampling ~3-4 samples per month of coral growth, and applying a ~2.5 month running mean filter, we obtained monthly-resolved Sr/Ca with an analytical uncertainty of ± 0.04 mmol/mol or ± 0.6 °C (1). While this is adequate for validating bulk Sr/Ca, it may not be sufficient for exploring paleotemperature changes without accompanying longer, bulk Sr/Ca timeseries. Increasing the number of SIMS Sr/Ca analyses per month of coral growth would indeed reduce the analytical uncertainty of SIMS-based Sr/Ca timeseries. For example, ~20 analyses/month would yield uncertainties as low as ± 0.022 mmol/mol (1sigma) or ± 0.3 °C in smoothed SIMS Sr/Ca, assuming a Gaussian distribution, however this may not be practical in most cases. Additionally, paleo-SST reconstructions from fossil corals must also contend with intercolony offsets, such as the ± 0.08 mmol/mol or ± 1 °C (1) offset observed in SIMS data from the three modern corals studied here, that make it difficult attribute differences in mean Sr/Ca values between two corals to climate change (e.g. *Sayani et al.*, 2019). As such, given that multiple corals are needed from a desired time interval to derive robust estimates of mean paleo-SST change, the most practical use of SIMS coral Sr/Ca is to validate the accuracy of bulk coral Sr/Ca timeseries.

5. Conclusion

Our results demonstrate that SIMS is a powerful tool that can be used to generate accurate, multi-year Sr/Ca time series from altered fossils. As most coral proxies are impacted by diagenesis to some extent, the sampling and data processing strategies presented here serve as a “roadmap” for recovering accurate paleoenvironmental reconstructions from altered corals using a wide variety of micro-scale analytical techniques, e.g. SIMS-based oxygen isotopes or Mg/Ca using laser ablation ICP-MS. Given the sheer volume of analyses needed to overcome fine-scale heterogeneity in coral geochemistry (e.g. the 730 individual SIMS analyses presented here cost ~\$15,000), microscale geochemical analyses are not practical in all circumstances. However, microscale analyses can and should be used to verify the accuracy of conventional, bulk Sr/Ca records from fossil corals, given that (i) no fossil coral is devoid of diagenesis, and (ii) even trace diagenesis can impart a significant artifact to coral Sr/Ca-based SST reconstructions, given the sensitivity of the coral Sr/Ca-SST relationship.

Acknowledgements

We’d like to thank Anne L. Cohen for providing laboratory resources and feedback that helped shape this work, and the phenomenal staff of the Northeast National Ion Microprobe Facility, including Nobumichi Shimizu, Glenn Gaetani, Kathryn Rose-Pietro, and Peter Landry, for their invaluable assistance in making this work possible. We’d also like to thank Yolande Berta and Georgia Tech’s Center for Nanostructure Characterization for providing access to their SEM facilities, and the Khaled bin Sultan Living Ocean Foundation and The Nature Conservancy for financial and logistical support for field excursions to

Palmyra. Funding for this work was provided by the National Science Foundation (Award #s 1502832 and 2002458 to K.M.C) and the National Oceanic and Atmospheric Administration (Award#: NA11OAR4310165 to K.M.C). ERSST V3b data provided by the NOAA PSD, Boulder, Colorado, USA, from their web site (<http://www.cdc.noaa.gov/>). New records presented in this study are archived on the Paleoclimatology section of the NOAA National Centers for Environmental Information data repository (link pending)

References

- Allison, N. (1996), Comparative determinations of trace and minor elements in coral aragonite by ion microprobe analysis, with preliminary results from Phuket, southern Thailand, *Geochimica et Cosmochimica Acta*, doi:10.1016/0016-7037(96)00171-8.
- Allison, N. (2005), Reconstruction of deglacial sea surface temperatures in the tropical Pacific from selective analysis of a fossil coral, *Geophys. Res. Lett.*, *32*(17), L17609, doi:10.1029/2005GL023183.
- Allison, N., A. A. Finch, J. M. Webster, and D. A. Clague (2007), Palaeoenvironmental records from fossil corals: The effects of submarine diagenesis on temperature and climate estimates, *Geochimica et Cosmochimica Acta*, *71*(19), 4693–4703, doi:10.1016/j.gca.2007.07.026.
- Allison, N., and A. A. Finch (2004), High-resolution Sr/Ca records in modern *Porites lobata* corals: Effects of skeletal extension rate and architecture, *Geochem. Geophys. Geosyst.*, *5*(5), doi:10.1029/2004GC000696.
- Allison, N., and A. A. Finch (2009), Reproducibility of minor and trace element determinations in *Porites* coral skeletons by secondary ion mass spectrometry, *Geochem. Geophys. Geosyst.*, *10*(4), n/a–n/a, doi:10.1029/2008GC002239.
- Allison, N., and A. A. Finch (2010a), The potential origins and palaeoenvironmental implications of high temporal resolution ^{18}O heterogeneity in coral skeletons, *Geochimica et Cosmochimica Acta*, *74*(19), 5537–5548, doi:10.1016/j.gca.2010.06.032.
- Allison, N., and A. A. Finch (2010b), ^{11}B , Sr, Mg and B in a modern *Porites* coral: the relationship between calcification site pH and skeletal chemistry, *Geochimica et Cosmochimica Acta*, *74*(6), 1790–1800, doi:10.1016/j.gca.2009.12.030.
- Barnes, D. J., R. B. Taylor, and J. M. Lough (1995), On the inclusion of trace materials into massive coral skeletons. Part II: distortions in skeletal records of annual climate cycles due to growth processes, *Journal of Experimental Marine Biology and Ecology*, *194*(2), 251–275.
- Beck, J. W., J. Recy, F. Taylor, R. L. Edwards, and G. Cabioch (1997), Abrupt changes in early Holocene tropical sea surface temperature derived from coral records, *Nature*, *385*(6618), 705–707.

- Beck, J. W., R. L. Edwards, E. Ito, F. W. Taylor, and J. Recy (1992), Sea-surface temperature from coral skeletal strontium/calcium ratios, *Science*, doi:10.1126/science.257.5070.644.
- Cobb, K. M., C. D. Charles, and D. E. Hunter (2001), A central tropical Pacific coral demonstrates Pacific, Indian, and Atlantic decadal climate connections, *Geophys. Res. Lett.*, *28*(11), 2209–2212, doi:10.1029/2001GL012919.
- Cobb, K. M., C. D. Charles, H. Cheng, and R. L. Edwards (2003), El Nino/Southern Oscillation and tropical Pacific climate during the last millennium, *Nature*, *424*(6946), 271–276, doi:10.1038/nature01779.
- Cobb, K. M., N. Westphal, H. R. Sayani, J. T. Watson, E. Di Lorenzo, H. Cheng, R. L. Edwards, and C. D. Charles (2013), Highly Variable El Nino-Southern Oscillation Throughout the Holocene, *Science*, *339*(6115), 67–70, doi:10.1126/science.1228246.
- Cohen, A. L., and R. A. Sohn (2004), Tidal modulation of Sr/Ca ratios in a Pacific reef coral, *Geophys. Res. Lett.*, *31*(16), doi:10.1029/2004gl020600.
- Cohen, A. L., and S. R. Hart (2004), Deglacial sea surface temperatures of the western tropical Pacific: A new look at old coral, *Paleoceanography*, *19*(4), doi:10.1029/2004PA001084.
- Cohen, A. L., G. D. Layne, S. R. Hart, and P. S. Lobel (2001), Kinetic control of skeletal Sr/Ca in a symbiotic coral: Implications for the paleotemperature proxy, *Paleoceanography*, *16*(1), 20–26, doi:10.1029/1999PA000478.
- Conroy, J. L., D. M. Thompson, K. M. Cobb, D. Noone, S. Rea, and A. N. LeGrande (2017), Spatiotemporal variability in the 18O-salinity relationship of seawater across the tropical Pacific Ocean, *Paleoceanography*, *32*(5), 484–497, doi:10.1002/2016pa003073.
- Conroy, J. L., K. M. Cobb, J. Lynch-Stieglitz, and P. J. Polissar (2014), Constraints on the salinity–oxygen isotope relationship in the central tropical Pacific Ocean, *Marine Chemistry*, *161*, 26–33, doi:10.1016/j.marchem.2014.02.001.
- DeLong, K. L., T. M. Quinn, C.-C. Shen, and K. Lin (2010), A snapshot of climate variability at Tahiti at 9.5 ka using a fossil coral from IODP Expedition 310, *Geochem. Geophys. Geosyst.*, *11*(6), n/a–n/a, doi:10.1029/2009GC002758.
- Deser, C., A. S. Phillips, and M. A. Alexander (2010), Twentieth century tropical sea surface temperature trends revisited, *Geophys. Res. Lett.*, *37*(10), doi:10.1029/2010GL043321.
- Emile-Geay, J., K. M. Cobb, M. E. Mann, and A. T. Wittenberg (2013a), Estimating Central Equatorial Pacific SST Variability over the Past Millennium. Part I: Methodology and Validation, *J. Climate*, *26*(7), 2302–2328, doi:10.1175/JCLI-D-11-00510.1.
- Emile-Geay, J., K. M. Cobb, M. E. Mann, and A. T. Wittenberg (2013b), Estimating Central Equatorial Pacific SST Variability over the Past Millen-

- nium. Part II: Reconstructions and Implications, *J. Climate*, *26*(7), 2329–2352, doi:10.1175/JCLI-D-11-00511.1.
- Enmar, R., M. Stein, M. Bar-Matthews, E. Sass, A. Katz, and B. Lazar (2000), Diagenesis in live corals from the Gulf of Aqaba. I. The effect on paleo-oceanography tracers, *Geochimica et Cosmochimica Acta*, *64*(18), 3123–3132, doi:10.1016/S0016-7037(00)00417-8.
- Epstein, S., R. Buchsbaum, H. A. Lowenstam, and H. C. Urey (1953), Revised Carbonate-Water Isotopic Temperature Scale, *Geological Society of America Bulletin*, *64*(11), 1315–1325, doi:10.1130/0016-7606(1953)64[1315:RCITS]2.0.CO;2.
- Felis, T. (2020), Extending the instrumental record of ocean-atmosphere variability into the last interglacial using tropical corals, *Oceanography*, *33*(2), 69–79, doi:10.5670/oceanog.2020.209.
- Felis, T. et al. (2014), Intensification of the meridional temperature gradient in the Great Barrier Reef following the Last Glacial Maximum, *Nature Communications*, *5*, ncomms5102.
- Gaetani, G. A., A. L. Cohen, Z. Wang, and J. Crusius (2011), Rayleigh-based, multi-element coral thermometry: A biomineralization approach to developing climate proxies, *Geochimica et Cosmochimica Acta*, *75*(7), 1920–1932, doi:10.1016/j.gca.2011.01.010.
- Gaetani, G. A., and A. L. Cohen (2006), Element partitioning during precipitation of aragonite from seawater: A framework for understanding paleoproxies, *Geochimica et Cosmochimica Acta*, *70*(18), 4617–4634, doi:10.1016/j.gca.2006.07.008.
- Gagan, M. K., G. B. Dunbar, and A. Suzuki (2012), The effect of skeletal mass accumulation in Porites coral Sr/Ca and ^{18}O paleothermometry, *Paleoceanography*, *27*(1), n/a–n/a, doi:10.1029/2011PA002215.
- Gagnon, A. C., J. F. Adkins, D. P. Fernandez, and L. F. Robinson (2007), Sr/Ca and Mg/Ca vital effects correlated with skeletal architecture in a scleractinian deep-sea coral and the role of Rayleigh fractionation, *Earth and Planetary Science Letters*, *261*(1-2), 280–295, doi:10.1016/j.epsl.2007.07.013.
- Grothe, P. R. et al. (2019), Enhanced El Niño-Southern Oscillation variability in recent decades, *Geophys. Res. Lett.*, 2019GL083906, doi:10.1029/2019GL083906.
- Hakim, G. J., J. Emile-Geay, E. J. Steig, D. Noone, D. M. Anderson, R. Tardif, N. Steiger, and W. A. Perkins (2016), The Last Millennium Climate Reanalysis Project: Framework and First Results, *Journal of Geophysical Research: Atmospheres*, *121*(12), 6745–6764, doi:10.1002/2016JD024751.
- Hart, S. R., and A. L. Cohen (1996), An ion probe study of annual cycles of Sr/Ca and other trace elements in corals, *Geochimica et Cosmochimica Acta*, *60*(16), 3075–3084, doi:10.1016/0016-7037(96)00154-8.

- Hasson, A. E. A., T. Delcroix, and R. Dussin (2013), An assessment of the mixed layer salinity budget in the tropical Pacific Ocean. Observations and modelling (1990–2009), *Ocean Dynamics*, *63*(2-3), 179–194, doi:10.1007/s10236-013-0596-2.
- Hendy, E. J., M. K. Gagan, J. M. Lough, M. McCulloch, and P. B. deMenocal (2007), Impact of skeletal dissolution and secondary aragonite on trace element and isotopic climate proxies in Porites corals, *Paleoceanography*, *22*(4), n/a–n/a, doi:10.1029/2007PA001462.
- Holcomb, M., A. L. Cohen, R. I. Gabbitov, and J. L. Hutter (2009), Compositional and morphological features of aragonite precipitated experimentally from seawater and biogenically by corals, *Geochimica et Cosmochimica Acta*, *73*(14), 4166–4179, doi:10.1016/j.gca.2009.04.015.
- Jimenez, G., J. E. Cole, D. M. Thompson, and A. W. Tudhope (2018), Northern Galápagos Corals Reveal Twentieth Century Warming in the Eastern Tropical Pacific, *Geophys. Res. Lett.*, *45*(4), 1981–1988, doi:10.1002/2017GL075323.
- Juillet-Leclerc, A., S. Reynaud, C. Rollion-Bard, J. P. Cuif, Y. Dauphin, D. Blamart, C. Ferrier-Pagès, and D. Allemand (2009), Oxygen isotopic signature of the skeletal microstructures in cultured corals: Identification of vital effects, *Geochimica et Cosmochimica Acta*, *73*(18), 5320–5332, doi:10.1016/j.gca.2009.05.068.
- Linsley, B. K. (2000), Decadal Sea Surface Temperature Variability in the Subtropical South Pacific from 1726 to 1997 A.D, *Science*, *290*(5494), 1145–1148, doi:10.1126/science.290.5494.1145.
- McGregor, H. V., and M. K. Gagan (2003), Diagenesis and geochemistry of porites corals from Papua New Guinea, *Geochimica et Cosmochimica Acta*, *67*(12), 2147–2156, doi:10.1016/S0016-7037(02)01050-5.
- McGregor, H. V., M. J. Fischer, M. K. Gagan, D. Fink, S. J. Phipps, H. Wong, and C. D. Woodroffe (2013), A weak El Niño/Southern Oscillation with delayed seasonal growth around 4,300 years ago, *Nature Geoscience*, *6*(11), 949–953, doi:10.1038/ngeo1936.
- Meibom, A. (2003), Monthly Strontium/Calcium oscillations in symbiotic coral aragonite: Biological effects limiting the precision of the paleotemperature proxy, *Geophys. Res. Lett.*, *30*(7), 1418, doi:10.1029/2002GL016864.
- Meibom, A., J.-P. Cuif, F. Houlbrèque, S. Mostefaoui, Y. Dauphin, K. L. Meibom, and R. Dunbar (2008), Compositional variations at ultra-structure length scales in coral skeleton, *Geochimica et Cosmochimica Acta*, *72*(6), 1555–1569, doi:10.1016/j.gca.2008.01.009.
- Murty, S. A., N. F. Goodkin, H. Halide, D. Natawidjaja, B. Suwargadi, I. Suprihanto, D. Prayudi, A. D. Switzer, and A. L. Gordon (2017), Climatic Influences on Southern Makassar Strait Salinity Over the Past Century, *Geophys. Res. Lett.*, *44*(23), 11,967–11,975, doi:10.1002/2017GL075504.

- Müller, A., M. K. Gagan, and M. T. McCulloch (2001), Early marine diagenesis in corals and geochemical consequences for paleoceanographic reconstructions, *Geophys. Res. Lett.*, *28*(23), 4471–4474, doi:10.1029/2001GL013577.
- Nothdurft, L. D., and G. E. Webb (2008), Earliest diagenesis in scleractinian coral skeletons: implications for palaeoclimate-sensitive geochemical archives, *Facies*, *55*(2), 161–201, doi:10.1007/s10347-008-0167-z.
- Nothdurft, L. D., G. E. Webb, N. A. Buster, C. W. Holmes, J. E. Sorauf, and J. T. Kloprogge (2005), Brucite microbialites in living coral skeletons: Indicators of extreme microenvironments in shallow-marine settings, *Geol*, *33*(3), 169, doi:10.1130/G20932.1.
- Nothdurft, L. D., G. E. Webb, T. Bostrom, and L. Rintoul (2007), Calcite-filled borings in the most recently deposited skeleton in live-collected Porites (Scleractinia): Implications for trace element archives, *Geochimica et Cosmochimica Acta*, *71*(22), 5423–5438, doi:10.1016/j.gca.2007.09.025.
- Nurhati, I. S., K. M. Cobb, and E. Di Lorenzo (2011), Decadal-Scale SST and Salinity Variations in the Central Tropical Pacific: Signatures of Natural and Anthropogenic Climate Change, *J. Climate*, *24*(13), 3294–3308, doi:10.1175/2011JCLI3852.1.
- Nurhati, I. S., K. M. Cobb, C. D. Charles, and R. B. Dunbar (2009), Late 20th century warming and freshening in the central tropical Pacific, *Geophys. Res. Lett.*, *36*(21), L21606, doi:10.1029/2009GL040270.
- Quinn, T. M., and F. W. Taylor (2006), SST artifacts in coral proxy records produced by early marine diagenesis in a modern coral from Rabaul, Papua New Guinea, *Geophys. Res. Lett.*, *33*(4), L04601, doi:10.1029/2005GL024972.
- Rabier, C., Y. Anguy, G. Cabioch, and P. Genthon (2008), Characterization of various stages of calcitization in Porites sp corals from uplifted reefs — Case studies from New Caledonia, Vanuatu, and Futuna (South-West Pacific), *Sedimentary Geology*, *211*(3-4), 73–86, doi:10.1016/j.sedgeo.2008.08.005.
- Ribaud-Laurenti, A., B. Hamelin, L. Montaggioni, and D. Cardinal (2001), Diagenesis and its impact on Sr/Ca ratio in Holocene Acropora corals, *Int J Earth Sci*, *90*(2), 438–451, doi:10.1007/s005310000168.
- Rodriguez, L. G., A. L. Cohen, W. Ramirez, D. W. Oppo, A. Pourmand, R. L. Edwards, A. E. Alpert, and N. Mollica (2019), Mid-Holocene, Coral-Based Sea Surface Temperatures in the Western Tropical Atlantic, *Paleoceanography and Paleoclimatology*, *34*(7), 1234–1245, doi:10.1029/2019PA003571.
- Rollion-Bard, C., D. Blamart, J. P. Cuif, and A. Juillet-Leclerc (2003), Microanalysis of C and O isotopes of azooxanthellate and zooxanthellate corals by ion microprobe, *Coral Reefs*, *22*(4), 405–415, doi:10.1007/s00338-003-0347-9.
- Sanchez, S. C., C. D. Charles, and J. D. Carriquiry (2016), Two centuries of coherent decadal climate variability across the Pacific North American region,

Geophysical ..., 43(17), 9208–9216, doi:10.1002/2016GL069037.

Sanchez, S. C., G. J. Hakim, and C. P. Saenger (2021), Climate model teleconnection patterns govern the Niño-3.4 response to early nineteenth-century volcanism in coral-based data assimilation reconstructions, *J. Climate*, 34(5), 1863–1880, doi:10.1175/JCLI-D-20-0549.1.

Sayani, H. R., K. M. Cobb, A. L. Cohen, W. C. Elliott, I. S. Nurhati, R. B. Dunbar, K. A. Rose, and L. K. Zaunbrecher (2011), Effects of diagenesis on paleoclimate reconstructions from modern and young fossil corals, *Geochimica et Cosmochimica Acta*, 75(21), 6361–6373, doi:10.1016/j.gca.2011.08.026.

Sayani, H. R., K. M. Cobb, K. DeLong, N. T. Hitt, and E. R. M. Druffel (2019), Intercolony ^{18}O and Sr/Ca variability among *Porites* spp. corals at Palmyra Atoll: Toward more robust coral-based estimates of climate, *Geochem. Geophys. Geosyst.*, 20(11), 5270–5284, doi:10.1029/2019GC008420.

Sinclair, D. J. (2005), Correlated trace element “vital effects” in tropical corals: A new geochemical tool for probing biomineralization, *Geochimica et Cosmochimica Acta*, 69(13), 3265–3284, doi:10.1016/j.gca.2005.02.030.

Sinclair, D. J., L. P. J. Kinsley, and M. T. McCulloch (1998), High resolution analysis of trace elements in corals by laser ablation ICP-MS, *Geochimica et Cosmochimica Acta*, 62(11), 1889–1901, doi:10.1016/S0016-7037(98)00112-4.

Smith, S. V., R. W. Buddemeier, R. C. Redalje, and J. E. Houck (1979), Strontium-Calcium Thermometry in Coral Skeletons, *Science*, 204(4391), 404–407, doi:10.1126/science.204.4391.404.

Tardif, R., G. J. Hakim, W. A. Perkins, K. A. Horlick, M. P. Erb, J. Emile-Geay, D. M. Anderson, E. J. Steig, and D. Noone (2019), Last Millennium Reanalysis with an expanded proxy database and seasonal proxy modeling, *Clim. Past*, 15(4), 1251–1273, doi:10.5194/cp-15-1251-2019.

Tokinaga, H., S.-P. Xie, A. Timmermann, S. McGregor, T. Ogata, H. Kubota, and Y. M. Okumura (2012), Regional Patterns of Tropical Indo-Pacific Climate Change: Evidence of the Walker Circulation Weakening*, <http://dx.doi.org/10.1175/JCLI-D-11-00263.1>, 25(5), 1689–1710, doi:10.1175/JCLI-D-11-00263.1.

Toth, L. T., R. B. Aronson, K. M. Cobb, H. Cheng, R. L. Edwards, P. R. Grothe, and H. R. Sayani (2015), Climatic and biotic thresholds of coral-reef shutdown, *Nature Climate Change*, 5(4), 369–374, doi:10.1038/nclimate2541.

Tudhope, A. W., C. P. Chilcott, M. T. McCulloch, E. R. Cook, J. Chapell, R. M. Ellam, D. W. Lea, J. M. Lough, and G. B. Shimmield (2001), Variability in the El Niño-Southern Oscillation Through a Glacial-Interglacial Cycle, *Science*, 291(5508), 1511–1517, doi:10.1126/science.1057969.

Vásquez-Bedoya, L. F., A. L. Cohen, D. W. Oppo, and P. Blanchon (2012), Corals record persistent multidecadal SST variability in the Atlantic Warm Pool

since 1775 AD, *Paleoceanography*, 27(3), n/a–n/a, doi:10.1029/2012PA002313.

Weber, J. N. (1973), Incorporation of strontium into reef coral skeletal carbonate, *Geochimica et Cosmochimica Acta*, 37(9), 2173–2190, doi:10.1016/0016-7037(73)90015-X.

Weber, J. N., and P. Woodhead (1972), Temperature dependence of oxygen-18 concentration in reef coral carbonates, *J. Geophys. Res.*, doi:10.1029/JC077i003p00463.