

# **Surface sampling of wetlands in a tropical dry forest after a large fire: Wind direction dominates the transport and deposition of fire proxies**

Nithin Kumar<sup>1</sup>, Prabhakaran Ramya Bala<sup>1\*</sup>, Diptimayee Behera<sup>2</sup>, Ambili Anoop<sup>2</sup>, Raman Sukumar<sup>3</sup>

<sup>1</sup>National Institute of Advanced Studies, Bangalore, India

<sup>2</sup>Indian Institute of Science Education and Research, Mohali, India

<sup>3</sup>Indian Institute of Science, Bangalore, India

\* Corresponding author - Prabhakaran Ramya Bala – [pramyabala@gmail.com](mailto:pramyabala@gmail.com), [pramyabala@nias.res.in](mailto:pramyabala@nias.res.in)

## Abstract

A major fire raged through the dry tropical forests of Bandipur Tiger Reserve, a 'biodiversity hotspot' in southern India, in February 2019. The fire occurred in patches, burning >10,000 acres of forest, becoming one of the largest forest fires in India in recent times. Very few studies have been able to capture the dynamics of fire proxies from an active surface fire, especially in the dry tropics. We opportunistically sampled two wetlands roughly sandwiched between the largest and the second-largest burnt patches a week after the fire was extinguished. We collected surface samples from each wetland and looked at popular fire proxies – macrocharcoal, microcharcoal, microcharcoal/pollen (C/P) ratio and the abundance and distribution of Polycyclic Aromatic Hydrocarbons (PAHs). Macrocharcoal counts were low (mean ~5), while the C/P ratio was ~1. Low Molecular Weight (LMW) PAH molecules Phenanthrene, Anthracene, Fluoranthene and Pyrene were found in both sites, while High Molecular Weight (HMW) PAHs were only found in one. None of the proxies is particularly indicative of the large surface fire that occurred ~15 km away. Analysis of wind speed and direction from weather station data and forward and backward HYSPLIT model trajectories tell us that both wetlands were not downstream of the smoke plume. There was also no recorded precipitation between the fire event and the sampling date. Through this opportunistic study of fire proxies, we show that wind direction and wet scavenging are essential factors determining the transport and deposition of fire proxies in this environment. Hence reconstruction of fire histories should be done using multi-site data since the absence of fire proxies does not equate to the absence of fire.

# 1 Introduction

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3 Approximately 420 million years ago ( late Silurian Period) the emergence of land vascular plants  
4 (Beerling, 2007) and the rise in the photosynthetic atmospheric oxygen (>13%), made fire ignition  
5 possible on Earth (Bowman et al., 2009; Glasspool et al., 2004; Scott, 2000a; Scott and Glasspool,  
6 2006). By Late Tertiary, increase in oxygen (Berner, 2006) and burnable biomass, fire became a  
7 global phenomenon capable of changing and shaping every terrestrial ecosystem across the globe  
8 (Herring, 2013; Jia et al., 2003; Pausas and Keeley, 2009). In addition to being a global terrestrial  
9 modifier, fire also played a critical role in human evolution as well. Cooking meat aided in the  
10 development of brain size, bipedalism, and utilization of stone tools (Dunbar, 2009). The knowledge  
11 of ignition and management of fire came out much later, out of necessity, due to the temperature  
12 drop from the warm Pliocene to Pleistocene ice ages (deMenocal, 2004; Glikson, 2013). From 790-  
13 690 kyr, humans started using fire more extensively (deMenocal, 2004) and subsequently, they  
14 started to make appreciable effects on the forest ecology (Archibald et al., 2012; Bowman et al.,  
15 2011; D. Burton, 2009; Marlon et al., 2009). With these increased anthropogenic pressures and the  
16 changes induced due to climate change and global warming on fire-prone ecosystems, several  
17 concerns have been raised about their fate and resilience (Bowman, 2015; Flannigan et al., 2009;  
18 Foley et al., 2013; Lawson et al., 2013; Schumacher and Bugmann, 2006). The information that is  
19 required to answer such concerns is achieved through the study of the long-term trends of vegetation  
20 and fire histories.

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22 In the last 50 years, several palaeoecological investigation techniques have been developed to aid  
23 this process. This includes the investigation through charcoal particles (Clark, 1988; Leys et al.,  
24 2013; Scott, 2000b; Tinner and Hu, 2003; Vachula, 2021; Whitlock and Larsen, 2002) Polycyclic  
25 Aromatic Hydrocarbons (PAHs) (Denis et al., 2012; Edwards, 1983; Kong et al., 2021; Musa  
26 Bandowe et al., 2014; Vachula et al., 2022; Yunker et al., 2002) in archeological material and natural  
27 paleoenvironmental archives, magnetic susceptibility of soil (Gedye et al., 2000; Rummery, 1983),  
28 fire scars in dendrochronology (Stephens et al., 2003), and also from historical documents. Of all  
29 these methods, charcoal-based reconstruction is the most widely used due to several advantages.  
30 Charcoal is inert (Scott, 2010), which results in its being well-preserved in a wide variety of  
31 environmental conditions. Charcoal has the potential to predict various characteristics of fire like

32 intensity (Duffin et al., 2008), severity (Whitlock and Larsen, 2002), extent, and frequency (Leys et  
33 al., 2013). Charcoal morphologies have also been used to trace the type of vegetation burnt (Enache  
34 and Cumming, 2006; Frank-DePue et al., 2022; Jensen et al., 2007; Mustaphi and Pisaric, 2014).  
35 The latter has developed an elaborate classification of sedimentary charcoal particles to identify fuel  
36 sources. Particles were categorized into seven major morphological classes (A to G) based on the  
37 overall shape, subdivided into 27 subclasses based on dominant surface textures or major features.  
38 With this classification scheme, they introduced an identification flow chart to identify fuel sources.  
39 Charcoal has also been combined with other proxies such as pollen (MacDonald et al., 1991),  
40 phytolith (Gu et al., 2007), molecular fire proxies (Argiriadis et al., 2018; Conedera et al., 2009; Denis  
41 et al., 2012; Musa Bandowe et al., 2014; Vachula et al., 2022) waxes (Lerch et al., 2022) and, sterols  
42 (Argiriadis et al., 2018). Since several other proxies can be simultaneously analyzed and creatively  
43 combined with charcoal, it provides robust ways of examining linkages between forest fires,  
44 vegetation, and anthropogenic activities.

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46 However, these reconstruction techniques with charcoal are not without problems. The lack of  
47 standardization in extraction procedures, identification, and quantification of macrocharcoal  
48 (Conedera et al., 2009) is of great concern. The uncertainties in relationships between particle size,  
49 burnt area, and dispersal distance (Clark, 1988; Higuera et al., 2007; Marlon et al., 2009; Vachula,  
50 2021; Vachula et al., 2018; Vachula and Richter, 2018) as well as the use of charcoal morphology  
51 to trace fuel sources are yet to be resolved (Enache and Cumming, 2006; Frank-DePue et al., 2022;  
52 Mustaphi and Pisaric, 2014). There are also variability associated with the production rate of charcoal  
53 depending on the vegetation burned (Feurdean, 2021; Pereboom et al., 2020; Vachula et al., 2022).  
54 Additionally, factors like wind speed and direction, morphology and size of the airborne particles,  
55 geographic location of sampling and source sites are also likely to affect the distribution, deposition,  
56 and retrieval of these proxies.

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58 Another recently developing fire indicator is the polycyclic aromatic hydrocarbons (PAHs). They are  
59 produced either by the combustion of fossil fuels or by the burning of organic materials during a  
60 forest fire (Edwards, 1983). The source of the PAH can be distinguished by certain ratios of specific  
61 PAHs (Denis et al., 2012; Vachula et al., 2022; Yunker et al., 2002). There are several advantages  
62 associated with this proxy. The PAH produced can vary depending on the combustion temperature

63 (Blumer, 1976; Laflamme and Hites, 1978) and they have a broader temperature range than charcoal  
64 (Conedera et al., 2009; Lu et al., 2009). Moreover, PAH assemblages also vary with fuel type, fuel  
65 moisture and oxygen availability (Burns et al., 1997; Lu et al., 2009; Yang et al., 2007). Thus, PAH  
66 analysis can give a lot of information about the fire event and the fuel sources, similar to charcoal. A  
67 few studies have in fact used PAH as indicators of fire along with charcoal (Denis et al., 2012; Kong  
68 et al., 2021; Vachula et al., 2022; Yunker et al., 2002) but there are problems associated with this as  
69 well. The significant ones being lack of understanding of the factors that control their production,  
70 spatial distribution, and deposition mechanism. The PAH in the sediment can also undergo microbial  
71 degradation, especially the low molecular ones, leading to its depletion in the sediment (Yan et al.,  
72 2022, 2021). Further, the low molecular ones are also subjected to degradation during transport  
73 (Vachula et al., 2022).

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75 Thus, certain aspects of these proxies (charcoal and PAH) are not very well understood and thus  
76 require a modern fire-proxy calibration to see how these proxies are captured and preserved in  
77 environmental archives. It also gives an idea on how they behave during after the fire event and also  
78 about the transport mechanism. Studies have been conducted on various fire-prone ecosystems like  
79 temperate, boreal, subalpine forests, grasslands, and savannahs across the world, making use of  
80 atmospheric and lake sediment traps, lake sediment cores, and lake surface sampling (Aleman et  
81 al., 2013; Clark et al., 1998; Leys et al., 2015; Lynch et al., 2004; Pisaric, 2002; Tinner et al., 2006).  
82 But none have been reported so far from the frequently burning tropical dry forests, especially from  
83 India, and this would be the first study reporting it.

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85 In this study, we look at the modern fire-proxy relationship in the dry tropical forests of southern India,  
86 through an opportunistic sampling of two wetlands in the Mudumalai Tiger Reserve (MTR), after a  
87 fire broke out on 22 February 2019. The fire occurred in patches, in the forests of Bandipur,  
88 Wayanad, and Mudumalai Tiger reserves of the Nilgiris Biosphere Reserve in the Western Ghats.  
89 According to a newspaper report, this was purported to be a man-made disaster caused by  
90 miscreants deliberately setting elephant dung on fire (Chennabasaveshwar, 2019) . The initial fires  
91 were brought under control within hours, but the dominant easterly winds of > 9km/hr speed (Kargudi  
92 weather station data), aided its spread, burning an area of more than 10,920 acres in just three days.  
93 The fire was brought under control by 26 February by the collective effort of forest officials, local

94 people, volunteers from the public and the Indian Air Force. A satellite study reported that the land  
95 temperature during the fire reached up to 140°C (Iyer et al., 2021). Even a week after the main fire  
96 event, complete extinguishment of the residual fire was not achieved.

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## 98 **Methods**

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### 100 **Sample location**

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102 The samples were collected from two wetlands, Ombetta (110 34'36.73"N, 76036'29.24"E) and  
103 Imbrella (11034'36.73"N, 76031'29.24"E), one week after the fire, on 7 March 2019 (Fig 1). The fire  
104 occurred in patches with the largest area burnt in the Bandipur Tiger reserve (BTR), while a second,  
105 smaller patch burnt in the adjacent Mudumalai Tiger reserve (MTR). Ombetta is ~5 km north of the  
106 MTR patch and 15 km south of the BTR patch. Imbrella is located ~7 km north of the MTR patch and  
107 10 km south of the BTR fire patch. These two wetlands are separated by a distance of about three  
108 kilometers.

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### 110 **Sample Collection**

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112 A total of 20 samples, 10 from each wetland were collected by walking along the edges. The ten  
113 samples comprise of five surface samples collected by scooping 2 cm of the surface and five  
114 bottle core samples collected by pressing a 5 cm tall plastic bottle into the sediment and collecting  
115 the contents. The collected samples were stored in a freezer in the field base at -20° C. They  
116 were transported in frozen state to the lab at the Indian Institute of Science, Bengaluru, where  
117 they were freeze-dried (Labconco bulk drier), crushed gently using pestle and mortar and stored  
118 in room temperature until further analysis.

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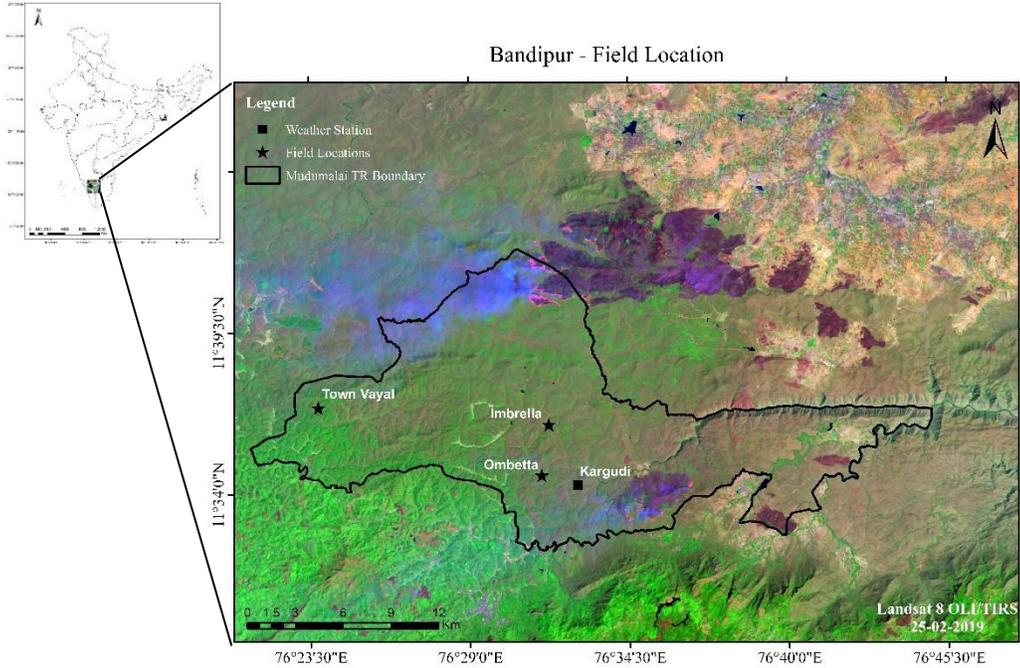


Fig 1: Landsat Image of Mudumalai Tiger Reserve (MTR) during the final day of the fire. The black line demarcates MTR boundaries. Imbrella and Ombetta (black stars) are the sampling wetlands, black square is the Kargudi weather station from which the wind data was collected. Two dark violet patches on the top of MTR and southeast of Kargudi are the major fire patches. The fire started in east and spread towards the west. You can see smoke arising moving west.

**Macrocharcoal analysis**

One gram of sample was taken and kept in 10ml of 10% sodium hexametaphosphate ( $\text{NaPO}_3)_6$  solution overnight to deflocculate the sediments. We found Sodium hexametaphosphate was better in dispersing the clay than KOH (Andreola et al., 2004). This was followed by Sodium hypochlorite ( $\text{NaOCl}$ ) treatment to bleach the organics (Stevenson and Haberle, 2005). The samples were then sieved through a 125  $\mu\text{m}$  mesh and residue  $>125 \mu\text{m}$  was collected for identification. Macrocharcoal particles are black, opaque and usually planar and were counted under a stereomicroscope (LeicaS4E) at 10X magnification. For each sample, three extractions were done and the charcoal particles counted, its mean, median, maximum and minimum values are given in fig 2.

157 **Microcharcoal and Pollen counts**

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159 Microcharcoal extraction was done using the same method used for pollen (Bennett, 1990), the  
160 standard procedure followed at the French Institute of Pondicherry (FIP). In this method, samples  
161 are treated with a dispersing agent and coarse particles (>150 µm) and clay particles are  
162 removed. Then it is treated with 10% hydrochloric acid (HCl) to remove carbonates and (48%)  
163 hydrofluoric Acid (HF) to remove silicates. This is followed by acetolysis (9:1 mixture of acetic  
164 anhydride (CH<sub>3</sub>CO)<sub>2</sub>O and sulfuric acid (H<sub>2</sub>SO<sub>4</sub>) to remove polysaccharides (Gunnar, 1960). The  
165 sample is then centrifuged and the supernatant is removed. The extracts were mounted on glass  
166 slides using glycerin. Both microcharcoal and pollen grains were counted (minimum of 1000  
167 pollen) under a microscope (Olympus CX43). Pollen counting was done using the Thanikaimoni  
168 reference collection at (FIP) (Anupama et al., 2014;).

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170 **Polycyclic Aromatic Hydrocarbons (PAHs)**

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172 Sediment samples, in powdered form (~6 g), were combined homogeneously with unsaturated  
173 silica. This blend was subsequently packed into stainless steel enclosures and subjected to lipid  
174 extraction in a Buchi Speed Extractor E-914. The solvent mixture used was dichloromethane  
175 (DCM) and methanol in a ratio of 93:7. The extractor parameters were set to conduct two cycles  
176 at a temperature of 100 °C and a pressure of 70 bar, facilitating the maximal extraction of organic  
177 material. Post extraction, the substance obtained was concentrated carefully at 30 °C using a  
178 Buchi P-6 Multivapor. The total lipid extract (TLE) dissolved in 2 ml of hexane, was then  
179 transferred into a glass column filled with 2 cm of glass wool and 17 cm of silica gel. 20 ml of  
180 hexane was used to isolate the saturated hydrocarbon component from the TLE. The aromatic  
181 fraction was eluted using 80 ml of a solution of hexane-DCM in the ratio of 4:1. This fraction was  
182 then completely condensed via dry nitrogen gas and dissolved in a final volume of 0.5 ml DCM  
183 for further analysis (Ajay et al., 2021; Behera et al., 2022)

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185 The fractionated aliquots were subjected to analytical testing using gas chromatography-mass  
186 spectrometry (GC-MS; Agilent 7890B/5977 MSD), adhering to the procedure detailed in (Behera  
187 et al., 2022). A non-polar capillary column (HP5-MS, with dimensions of 30 m × 250 µm × 0.25

188  $\mu\text{m}$ ) was utilized for this analysis, with helium as carrier gas. The injection method used was  
189 spitless, with the inlet temperature set at 320 °C and a maximum sample volume of 1  $\mu\text{L}$ . The  
190 GC oven's programming commenced with a base temperature of 40 °C, held steady for 2  
191 minutes, followed by a progressive increase to a final temperature of 320°C at a gradient of 4°C  
192 per minute. The rate of flow for the helium gas within the column was sustained at 1.4  $\text{cm}^2/\text{sec}$ .  
193 As for the mass spectrometer detector (MSD) conditions, an EI ionization source of 70 eV was  
194 used, with a mass range spanning from 45–600 atomic mass units (amu). The multiplier voltage  
195 was set to 2341 V, with the ion-source temperature maintained at 230 °C.

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197 The polycyclic aromatic hydrocarbons (PAHs) were recognized by aligning the distinct mass  
198 spectra with existing literature and the National Institute of Standards and Technology (NIST)  
199 database. In order to evaluate the quantity of PAHs present within the samples, an external  
200 calibration curve was generated. This was based on analytical standards with a range of  
201 concentrations ( $n=8$ ), incorporating the Sigma CRM47930 PAHs mix with concentrations of 50,  
202 100, 200, 300, 400, 600, 800, and 1200 ng/mL. Several indices related to PAHs were determined  
203 to aid in the quantification and understanding of the varied origins contributing to the organic  
204 content in the system.

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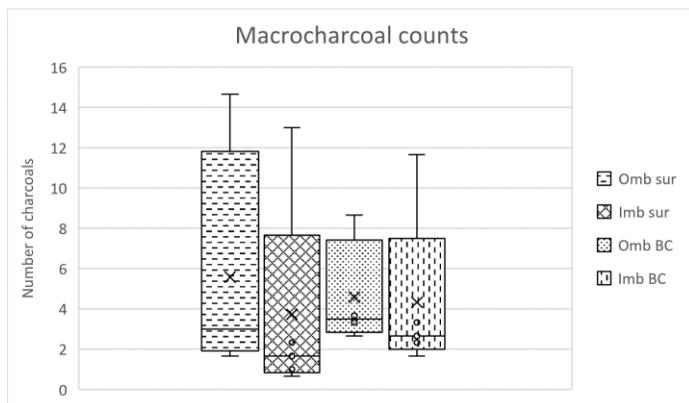
## 206 **Wind Trajectory Analysis**

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208 To track the wind direction, we used the Hybrid Single-Particle Lagrangian Integrated Trajectory  
209 model (HYSPLIT) (Draxler and Hess, 1998). This computer model is widely used to forecast  
210 wildfire smoke trajectories as it can impact air quality and human health. It also has archival data  
211 of the movement of air parcels. This archival information was combined with two functions of the  
212 model – 1. Forward trajectory that gives the direction the air parcel moved from a given point of  
213 interest and 2. Backward trajectory that gives the direction from which air parcels came to the  
214 point of interest. In our case, we used forward trajectories for the fire locations to see the direction  
215 of potential charcoal transport and backward trajectories on sampling locations to see the  
216 potential sources of charcoal particles deposited. The model was run on all days starting from  
217 22<sup>nd</sup> February to the date of sample collection. It was run for 24 hours each day with new  
218 trajectories starting every 2 hours. Since this was a large fire, we considered wind parcels at

219 different altitudes ranging from 100 meters to 1 kilometer from the ground level. We also collected  
 220 wind speed and wind direction data from Kargudi weather station, about 2 km and 4 km southeast  
 221 of Ombetta and Imbrella, respectively. The Kargudi weather Station is close to the fire patch in  
 222 MTR (~5 km) (fig 1).

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Fig 2: Box plot of macrocharcoal counts per gram of soil. Three extracts were done for each sample. The mean, median, maximum and minimum values of charcoal obtained are given in this plot. For all the samples, average number of charcoals lies between 4 and 6.

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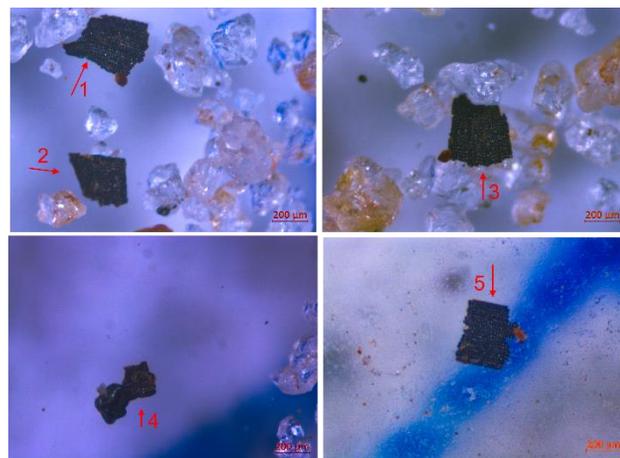


Fig 3 macrocharcoal particles (red arrows) found from the wetland samples. Only macrocharcoal of this type was found from the samples. This according to Mustaphi and Pisaric, 2014's classification scheme is of wood origin. The W/L ratio of these charcoals also show that they are of wood origin. The transparent crystals surrounding the charcoals are quartz. The W/L ratio of 1, 2,3 ,4 and 5 are 0.95, 0.80,

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## 240 Results

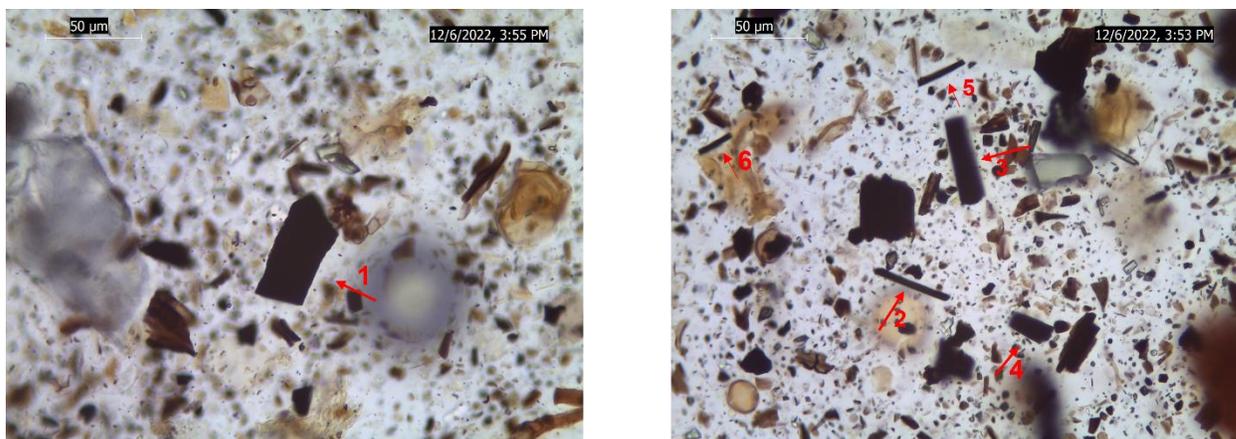
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### 242 Macro and Microcharcoal

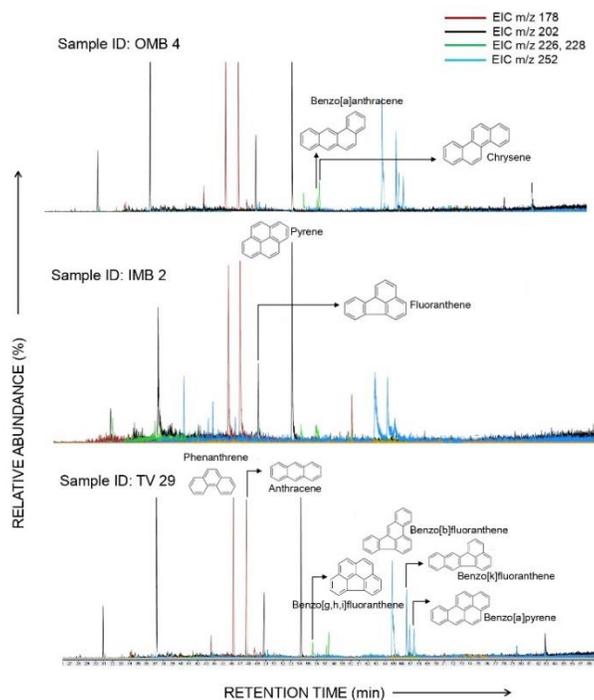
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244 Macrocharcoal counts of both sets of samples, surface and bottle core, were disproportionately  
 245 low compared to the range and extent of the forest fire (see fig 2). The maximum number of  
 246 macrocharcoal counted was 15 from one Ombetta surface sample, while the average number of  
 247 particles was ~5 for all samples from both the wetlands. All the charcoal particles observed have

248 a textured broad shape indicative of wood burning (Mustaphi and Pisaric, 2014) (fig 3). The width  
 249 to length ratio of all these particles are  $>0.5$  which again indicates a woody origin (Aleman et al.,  
 250 2013) ( fig 3). Macrocharcoal counting of the top 5 cm of Town Vayal (TV), another wetland from  
 251 a very low fire frequency area of Mudumalai was also done. This top layer sample had 13  
 252 charcoals per gram of sample. This sample also had morphologically similar type of  
 253 macrocharcoal in them. Microcharcoal to pollen (C/P) ratio was 1.2 and 1.1 for Ombetta and  
 254 Imbrella respectively. Length to width ratio of the microcharcoals gave values ranging from 1.2  
 255 (leaf) to  $>2.13$  (wood) to  $>3.62$  (grass) (Cui et al., 2009) (fig 4). Thus, vegetation of all types is  
 256 represented in the pollen slides while only wood charcoals are represented in the macrocharcoal.



266 Fig 4: microcharcoal images of the wetland samples. The red arrows show the microcharcoal. Unlike macrocharcoal,  
 267 microcharcoal of all sizes are found in the sample. The *L/W* values of 1,2,3,4,5, and 6 are 1.97, 12.3, 3.90, 2.19, 1.22,  
 268 and 6.55



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 Fig 5: Chromatograms of the 10 PAHs obtained. These samples are representative of all the  
 samples analysed. OMB and IMB are the wetland samples and TV29 is a surface sample  
 from a low fire-frequency area of MTR. The PAHs found include phenanthrene, anthracene,  
 fluoranthene, pyrene, benzo[g,h,i]fluoranthene, benzo[a]anthracene, chrysene,  
 benzo[b]fluoranthene, and benzo[a]pyrene, shown in the figures with representative peak  
 positions.

## 279 PAH Analysis

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281 Out of the 32 PAH compounds reported in literature (Dzepina et al., 2007), 10 compounds were  
282 detected from the nine surface samples of the wetlands (fig 5) . This includes four Low Molecular  
283 Weight (LMW) ones, phenanthrene, anthracene, fluoranthene and pyrene and six High Molecular  
284 Weight (HMW) ones, benzo[g,h,i]fluoranthene, benz[a]anthracene, Chrysene,  
285 benzo[b]fluoranthene, and benzo[a]pyrene. All the Imbrella samples are missing  
286 benzo[g,h,i]fluoranthene and benz[a]anthracene except for IMB 4 and IMB 5 respectively. LMW/  
287 total PAH value shows that Ombetta has low LMW/ $\Sigma$ PAHs (mean  $\sim$ 0.37) while Imbrella has  
288 comparatively high values (mean  $\sim$ 0.74). However, Imbrella is generally low in PAHs (mean of  
289 total PAHs  $\sim$ 0.76 ng/g) compared to Ombetta (mean of total PAHs  $\sim$ 3.5 ng/g). Imbrella also lacks  
290 many of the higher end compounds. The  $\Sigma$ LMW PAHs of Imbrella is also low compared to  
291 Ombetta. Ombetta is rich in HMW PAHs (avg  $\sim$ 0.63). Furthermore, Town Vayal surface sample,  
292 which is from a low fire-frequency area, also yielded the same 10 compounds and is rich in LMW  
293 compounds (>80% of total PAH detected). The  $\Sigma$ PAHs for town vayal is comparable to Ombetta  
294 (2.94 ng/g) but is generally rich in LMW PAHs than HMW unlike Ombetta.

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296 Ratios of different PAHs were used to determine the source of PAHs. Yunker et al., 2002 show  
297 that samples with anthracene / (anthracene (Ant)+ phenanthrene (Ph)) ratio  $>0.1$  are pyrogenic  
298 in origin. This value for Ombetta, Imbrella and Town Vayal lies between 0.75 - 0.81, 0.38-0.61  
299 and 0.69 respectively. Similarly, for the Fluoranthene (Fla) / (fluoranthene + pyrene (Py)) ratio  
300  $>0.5$  indicate pyrogenic source. This value lies between 0.53-0.61, 0.4-0.63, and 0.58 for  
301 Ombetta, Imbrella and town vayal respectively. One sample of Imbrella (IMB 1) shows petrogenic  
302 signature, while rest of the samples including the town vayal are pyrogenic in origin (Fig 6a).  
303 Ratios of Fla/Py vs Ph/Ant (Fig 6b) and Fla/ Py vs LMW/HMW (Fig 6c) (Ranjbar Jafarabadi et al.,  
304 2017) also suggest pyrogenic source for all the samples except for town vayal which has a high  
305 LMW/HMW value (Fig 6c).

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311 Table 1 PAH analysis summary

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Site	$\Sigma$ PAH	$\Sigma$ LMW	$\Sigma$ HMW	a/(a+p)	f/(f+p)	LMW/HMW	f/p
IMB	0.764637	100277	41398	0.5614	0.56	0.863	1.34
OMB	3.5893	242044	432627	0.7802	0.58	0.592	1.39
TV	2.9468	427586	125852	0.6925	0.58	3.397	1.40

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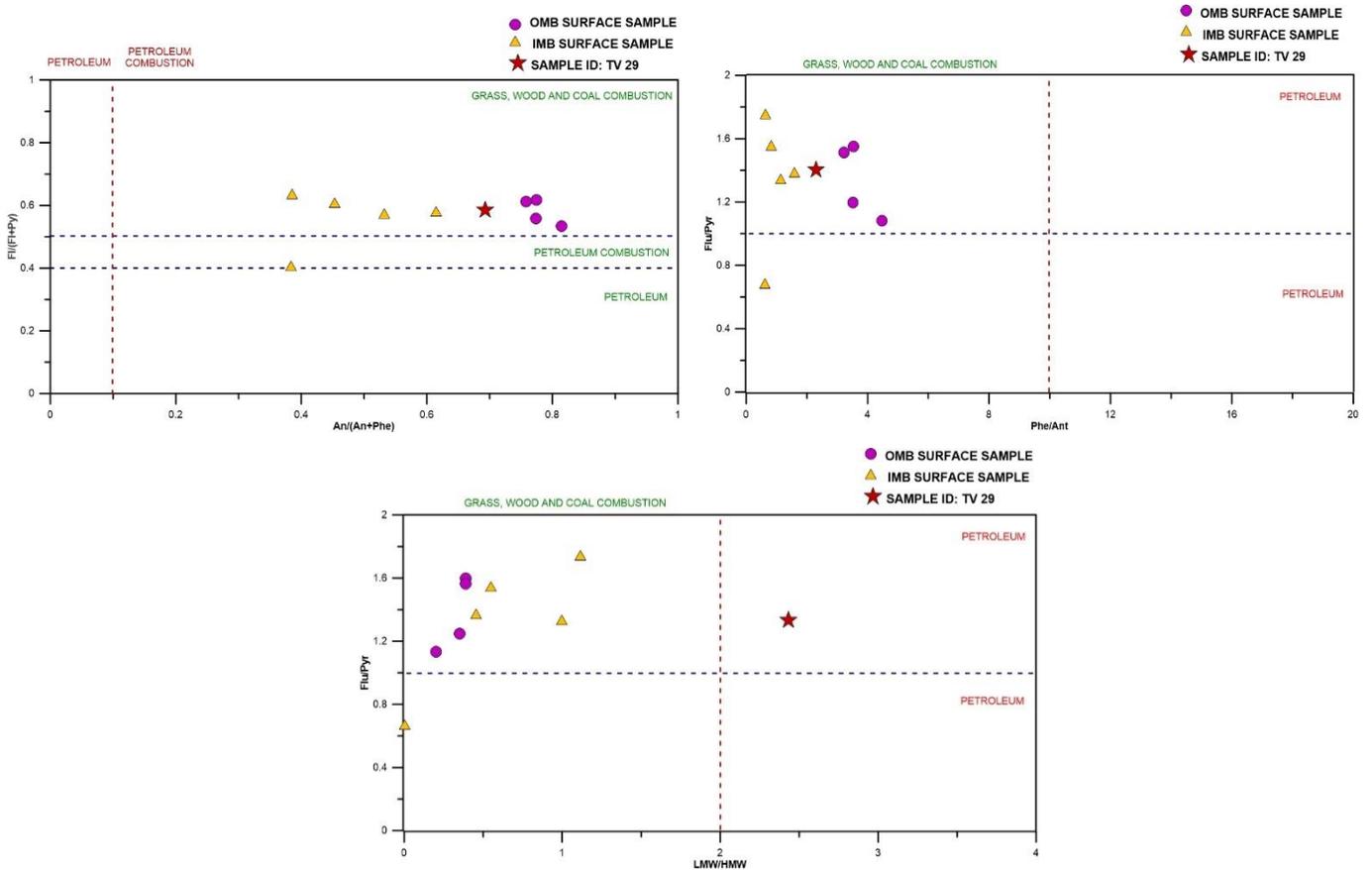
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Fig 6: PAH source determination plots

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## 338 Wind Trajectory Analysis

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340 From the Kargudi weather station data, we created a wind rose diagram to observe the general  
341 wind pattern during and after the fire events (fig 7). During the fire event (22 – 26 February 2019),  
342 the weather station recorded a wind speed ranging from 0.3 to 18.3 km/hr. The wind rose diagram  
343 (fig 7) showed that winds with speed greater than 9.3 km/hr dominated these days. More than  
344 60% of the winds were from the Northeast, East or Southeast and thus in general, easterly in  
345 direction. This seems to have aided the spread of the fire towards the west. The westerly winds  
346 start to get stronger towards the end of the fire event. After the fire event, from 27 February 2019,  
347 wind speed ranged between 0.3 to 12.3 km/hr, more than 65% of the wind was westerly.

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349 HYSPLIT model was used to look at wind trajectories every hour of the day at different heights  
350 from the surface. The forward trajectory models from the major fire locations were created. The  
351 results show that wind from the larger Bandipur fire patch did not pass through any of the sample  
352 locations both during and after the fire event except on 6 March, 2019, one day before the sample  
353 collection. Forward trajectory models on the smaller MTR fire patch show that the wind was  
354 passing through the sample locations both during and after the fire days. This implies that  
355 whatever fire proxies were captured by the wetlands would be dominated by the MTR fire source.  
356 A backward trajectory of the wind parcels at the sample locations was also analyzed that  
357 confirmed the conclusions from the forward trajectory models (REF Supplementary material (III)).

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Fig 7: Wind rose diagram created from the Kargudi weather station data. The left side image is after fire (27<sup>th</sup> feb to 7 of march, 2019). Right side image is from 22<sup>nd</sup> feb to 26<sup>th</sup> feb, 2019, when the fire was prominent. During the fire days the wind was mostly westerly. More than 40% of them are high speed wind (15.3-18.3 km/hr). Easterly winds start to develop towards the end of fire. After the fire event, the wind is mostly from the east with speed ranging 12.3-6.3 km/hr. The change in wind direction has in fact helped in curbing the fire.

## 367 Discussion

368

### 369 **Macrocharcoal and Microcharcoal Count**

370

371 Macrocharcoal counts from both wetlands are low (mean ~5) for all the samples. Whitlock and  
372 Larsen (2002), in their study, show that macrocharcoal count >50 was indicative of fire events.  
373 However, we must note that tropical forests are not included in this study and hence this arbitrary  
374 value may not be applicable to our study from tropical India. Since ours is the first study to try  
375 and evaluate fire-proxy relationships in the dry tropical forests of India, we do not yet have  
376 baseline values for a fire peak representation in surface sediments in this landscape. Hence we  
377 chose to compare our results to investigations on sediment cores from the larger region of the  
378 Western Ghats (Bhagwat et al., 2012, 2014; Kulkarni et al., 2021; Nogué et al., 2018) and a  
379 surface sample (top 5cm) from Town vayal. These investigations on fire-vegetation-human inter-  
380 relationships have been done in the Kodagu district of Karnataka, central Western Ghats in  
381 anthropogenically modified landscapes, albeit in tropical wet evergreen forests and very low  
382 burning part of the Mudumalai Tiger reserve, respectively. On one of the sediment cores, the  
383 authors report an increase of fire in two time periods: ~1800-1400 yr BP and 400-0 yr BP, was  
384 represented by 61 charcoal particles per cm<sup>3</sup> of sample (Bhagwat et al., 2012; Kulkarni et al.,  
385 2021) respectively. This shows that >50 charcoal count for fire indication may hold for tropical  
386 India as well. However further study has to be done to confirm this. The surface of these sediment  
387 core had about 18 charcoals per cm<sup>3</sup> of sample. This low count is not very surprising as the region  
388 is very unlikely to burn as it is a plantation area. This strongly implies that the surface of our  
389 wetlands is not representative of the large surface fire. The understanding that macrocharcoal as  
390 a result of their larger size and limited mobility would only be able to travel small distances and  
391 represent a local fire event aligns well with our findings as both the wetlands are 5-7 kms away  
392 from either of the fire patches (Clark, 1988; Mooney and Tinner, 2011).

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## 398 **Macrocharcoal: Morphology**

399

400 The Bandipur fire was one of the largest recent forest fires and burnt large swathes of the forest,  
401 including canopy trees and understory vegetation – both grass and shrub. Yet, only  
402 macrocharcoal, possibly of woody plants origin, were found (irregular, opaque with textured  
403 surface), while charcoal from the other vegetation types were not found. However, microcharcoal  
404 indicative of diverse vegetation types were found in the samples (based on the large range of  
405 sizes and shapes (Fig 4). Macrocharcoal of similar morphology is also found from the low burning,  
406 800-year, old Town Vayal sample (unpublished data) as well from its surface. Thus, the presence  
407 of this macrocharcoal in both low burning location and >7km from the fire edge (OMB and IMB)  
408 shows that it can travel larger distances. It could also be distinctive of the tropical dry deciduous  
409 forest.

410

## 411 **Polycyclic Aromatic Hydrocarbons (PAHs)**

412

413 All of our wetland samples show pyrogenic source for the PAHs (fig 6). TV is more proximate to  
414 human occupation; it is likely that the signatures of PAHs there are of a mixed pyro-petrogenic  
415 origin. The higher LMW/HMW value for TV sample would be indicative of that. However, since  
416 the sample was collected in what is well within Tiger Reserve boundaries, we expect that most  
417 of the PAHs would be of pyrogenic origin for the purposes of this study.

418

419 The individual PAH compounds can give an idea of the spatial distribution of the fire (Vachula et  
420 al., 2022). According to Vachula et al., 2022, naphthalene, anthracene, acenaphthylene, and  
421 fluorine record local fire histories (within 40km) whereas phenanthrene, benzo[g,h,i] perylene,  
422 chrysene, benzo[k]fluoranthene, benzo[b]fluoranthene record both local and regional fires. Our  
423 samples are rich in compounds that can record both local and regional fires and thus indicate  
424 that the big near fire event was not well captured by the wetland.

425

426 LMW/ $\Sigma$ PAH is found to show a positive correlation with accumulation rate of sedimentary  
427 macrocharcoal of all size classes (Vachula et al., 2022). They thus postulate that the adsorption  
428 of PAHs (LMWs) to charcoal can be an important mode of transport and deposition for them.

429 However, (Kong et al., 2021) in their study of the peat samples from Malaysia observed a positive  
430 correlation of the HMW PAHs with charcoal abundance. The LMW was in fact decreasing with  
431 increasing charcoal abundance. In our case, both the wetlands and Town Vayal core top have  
432 low charcoal counts. While Town Vayal shows a high LMW/ $\Sigma$ PAH, Ombetta is rich in HMW PAHs.  
433 This difference in concentration of LMW and HMW in low charcoal scenario highlights the  
434 possibility for more research in linking either of these PAHs with charcoal abundance. Based on  
435 our results neither of the PAHs show any relationship with macrocharcoal. This indicates that the  
436 increase in concentration is independent of adsorption to charcoal as a transportation  
437 mechanism.

438  
439 HMW PAHs are produced more in a high temperature scenario (McGrath et al., 2003). Forest  
440 fires, depending on the fuel that is burned and the climate, can be highly variable in terms of  
441 temperature of burning. Even though both Ombetta and Imbrella are close to each other (~3km),  
442 there is a huge distinction between them in terms of the PAH that were captured (LMW vs HMW).  
443 Imbrella is closer to the bigger fire patch. However, the  $\Sigma$  PAHs that got captured by the wetland  
444 is low compared to Ombetta. This shows that the signals from the bigger BTR patch were not  
445 captured by the wetland. HYSPLIT model also confirms that none of the wind parcel from the  
446 BTR passed through our sample locations. The high value of HMWs in Ombetta (Ombetta is  
447 closer to MTR patch) indicates that the signal might be from the MTR patch. HYSPLIT model  
448 results show that some of the winds did go through our sample location (Ombetta).

449  
450 High percentages of LMW PAHs compared to HMW PAHs in the wet forest sediments of Town  
451 vayal shows that they can travel farther and get deposited. Their presence in low amounts even  
452 in Imbrella shows that they can get deposited in less transport-favorable environments as well.  
453 The LMW PAHs have low gas-particle partition coefficient (Alam et al., 2014; Xie et al., 2014)  
454 and so they are mostly associated with gas phase(Wang et al., 2016). This property of the LMW  
455 PAHs makes sense with their long range of transport. The opposite is true for HMW PAHs. Thus,  
456 high percentages of HMW PAHs in Ombetta and absence/ low concentration in both Imbrella and  
457 Town Vayal shows that they are more local fire indicators and are particle or transport medium  
458 dependent. However, more modern-day calibration has tinbe done on this regard before using  
459 them as reliable fire proxies.

460 According to Wang et al., 2016, PAH flux deposition of LMW PAHs (2 + 3 rings) are high with wet  
461 condition. This is because the LMWs are more water soluble (Mackay and Shiu, 1977) and tend  
462 to be associated more with gaseous phase. The HMW PAHs (4+5+6 rings) on the other hand,  
463 favor dry deposition. The PAH flux deposition was also found to be high on a calm day (Wang et  
464 al., 2016). In our case, from the fire start day to the day of sample collection, the climate was dry  
465 and windy. So, most of the LMW PAHs might have been carried away by the wind as they are  
466 more associated with the gaseous phase. Thus, the dry deposition mode should have been more  
467 favored in such conditions. However, the absence of HMW PAHs in Imbrella shows that wind  
468 does determine their transport and deposition just as is the case for macrocharcoal.

469

470

## 471 Conclusion

472

473 From our opportunistic sampling of wetland surfaces following a large fire event, we report the  
474 following important findings. First, we demonstrate the importance of wind direction in proxy  
475 transport and deposition. We see that all fire proxies considered in both wetlands –  
476 macrocharcoal, microcharcoal/pollen ratio and PAHs is disproportionately low and not  
477 representative of one of the largest forest fire events in the recent past in India. This is ascribed  
478 to the lack of favorable wind for deposition of these proxies. Second, the distance travelled by  
479 macrocharcoal (blocky texture) in our study site. For the smaller burnt area whose downwind fire  
480 proxies we seem to have captured through our sampling of these two wetland surfaces, charcoal  
481 particles of size 120-200  $\mu$  has been transported up to 7 km which to our knowledge is reported  
482 for the first time in India. Third, we report ambiguity in the use of PAH as fire proxy. HMW PAHs  
483 seem to capture local and high intensity fire well while LMW PAHs seem to be regional in nature.  
484 We also report ambiguity with respect to transportation and preservation of PAHs and their  
485 relationship with wet and dry deposition events. More modern-day fire calibration is required to  
486 have a more nuanced understanding of the factors that affect fire proxy transport and deposition  
487 and to reconstruct past fire events with more confidence.

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## 490 References

- 491 Ajay K, Behera D, Bhattacharya S, Mishra PK, Ankit Y and Anoop A (2021) Distribution and  
492 characteristics of microplastics and phthalate esters from a freshwater lake system in Lesser  
493 Himalayas. *Chemosphere* 283: 131132:  
494 doi:10.1016/j.chemosphere.2021.131132.
- 495 Alam MS, Delgado-Saborit JM, Stark C and Harrison RM (2014) Investigating PAH relative  
496 reactivity using congener profiles, quinone measurements and back trajectories. *Atmospheric*  
497 *Chemistry and Physics* 14(5): 2467–2477:  
498 doi:10.5194/acp-14-2467-2014.
- 499 Aleman JC, Blarquez O, Bentaleb I, Bonté P, Brossier B, Carcaillet C, et al. (2013) Tracking  
500 land-cover changes with sedimentary charcoal in the Afrotropics. *The Holocene* 23(12): 1853–  
501 1862:  
502 doi:10.1177/0959683613508159.
- 503 Andreola F, Castellini E, Manfredini T and Romagnoli M (2004) The role of sodium  
504 hexametaphosphate in the dissolution process of kaolinite and kaolin. *Journal of the European*  
505 *Ceramic Society* 24(7): 2113–2124:  
506 doi:10.1016/S0955-2219(03)00366-2.
- 507 Anupama K, Prasad S and Reddy CS (2014) Vegetation, land cover and land use changes of  
508 the last 200 years in the Eastern Ghats (southern India) inferred from pollen analysis of  
509 sediments from a rain-fed tank and remote sensing. *Quaternary International* 325: 93–104:  
510 doi:10.1016/j.quaint.2014.02.003.
- 511 Archibald S, Staver AC and Levin SA (2012) Evolution of human-driven fire regimes in Africa.  
512 *Proceedings of the National Academy of Sciences* 109(3): 847–852:  
513 doi:10.1073/pnas.1118648109.
- 514 Argiriadis E, Battistel D, McWethy DB, Vecchiato M, Kirchgeorg T, Kehrwald NM, et al. (2018)  
515 Lake sediment fecal and biomass burning biomarkers provide direct evidence for prehistoric  
516 human-lit fires in New Zealand. *Scientific Reports* 8(1): 12113: doi:10.1038/s41598-018-30606-  
517 3.
- 518 Beerling D (2007) *The Emerald Planet: How Plants Changed Earth's History*. Oxford University  
519 Press.
- 520 Behera D, Bhattacharya S, Rahman A, Kumar S and Anoop A (2022) Molecular tracers for  
521 characterization and distribution of organic matter in a freshwater lake system from the Lesser  
522 Himalaya. *Biogeochemistry* 161(3): 315–334: doi:10.1007/s10533-022-00984-y.
- 523 Bennett KD (1990) Textbook of pollen analysis. K. FAEGRI, J. IVERSEN (4th edn by K.  
524 FAEGRI, P. E. KALAND, K. KRZYWINSKI), Publisher John Wiley and Sons, Chichester 1989  
525 (328 pp) £51.00 ISBN 0 471 92178 5. *Journal of Quaternary Science* 5(3): 254–255:  
526 doi:10.1002/jqs.3390050310.
- 527 Berner RA (2006) GEOCARBSULF: A combined model for Phanerozoic atmospheric O<sub>2</sub> and  
528 CO<sub>2</sub>. *Geochimica et Cosmochimica Acta* 70(23): 5653–5664: doi:10.1016/j.gca.2005.11.032.

- 529 Bhagwat SA, Nogué S and Willis KJ (2012) Resilience of an ancient tropical forest landscape to  
530 7500years of environmental change. *Biological Conservation* 153: 108–117:  
531 doi:10.1016/j.biocon.2012.05.002.
- 532 Bhagwat SA, Nogué S and Willis KJ (2014) Cultural drivers of reforestation in tropical forest  
533 groves of the Western Ghats of India. *Forest Ecology and Management* 329: 393–400:  
534 doi:10.1016/j.foreco.2013.11.017.
- 535 Blumer M (1976) Polycyclic Aromatic Compounds in Nature. *Scientific American* 234(3): 34–45:  
536 doi:10.1038/scientificamerican0376-34.
- 537 Bowman DM (2015) What is the relevance of pyrogeography to the Anthropocene? *The*  
538 *Anthropocene Review* 2(1): 73–76: doi:10.1177/2053019614547742.
- 539 Bowman DMJS, Balch J, Artaxo P, Bond WJ, Cochrane MA, D'Antonio CM, et al. (2011) The  
540 human dimension of fire regimes on Earth: The human dimension of fire regimes on Earth.  
541 *Journal of Biogeography* 38(12): 2223–2236: doi:10.1111/j.1365-2699.2011.02595.x.
- 542 Bowman DMJS, Balch JK, Artaxo P, Bond WJ, Carlson JM, Cochrane MA, et al. (2009) Fire in  
543 the Earth System. *Science* 324(5926): 481–484: doi:10.1126/science.1163886.
- 544 Burns WA, Mankiewicz PJ, Bence AE, Page DS and Parker KR (1997) A principal-component  
545 and least-squares method for allocating polycyclic aromatic hydrocarbons in sediment to  
546 multiple sources. *Environmental Toxicology and Chemistry* 16(6): 1119–1131:  
547 doi:10.1002/etc.5620160605.
- 548 Chennabasaveshwar P (2019) What Caused the Fire Bandipur Forest? , 25 February. Available  
549 at: <https://www.oneindia.com/india/what-caused-fire-in-bandipur-forest-2856666.html?story=2;>
- 550 Clark JS (1988) Particle Motion and the Theory of Charcoal Analysis: Source Area, Transport,  
551 Deposition, and Sampling. *Quaternary Research* 30(1): 67–80: doi:10.1016/0033-  
552 5894(88)90088-9.
- 553 Clark JS, Lynch J, Stocks BJ and Goldammer JG (1998) Relationships between charcoal  
554 particles in air and sediments in west-central Siberia. *The Holocene* 8(1): 19–29:  
555 doi:10.1191/095968398672501165.
- 556 Conedera M, Tinner W, Neff C, Meurer M, Dickens AF and Krebs P (2009) Reconstructing past  
557 fire regimes: methods, applications, and relevance to fire management and conservation.  
558 *Quaternary Science Reviews* 28(5–6): 555–576: doi:10.1016/j.quascirev.2008.11.005.
- 559 Cui Q, Marquer L, Arzarello M and Lebreton V (2009) An attempt to separate anthropic and  
560 natural fire signals in an archaeological context-The case of the Mousterian site Grotta Reali  
561 (Rocchetta a Volturno Molise, Central Italy). *Frontiers of Earth Science in China* 3(2): 171–174:  
562 doi:10.1007/s11707-009-0014-8.
- 563 D. Burton F (2009) *Fire: The Spark That Ignited Human Evolution*. .
- 564 deMenocal PB (2004) African climate change and faunal evolution during the Pliocene–  
565 Pleistocene. *Earth and Planetary Science Letters* 220(1–2): 3–24: doi:10.1016/S0012-  
566 821X(04)00003-2.

- 567 Denis EH, Toney JL, Tarozo R, Scott Anderson R, Roach LD and Huang Y (2012) Polycyclic  
568 aromatic hydrocarbons (PAHs) in lake sediments record historic fire events: Validation using  
569 HPLC-fluorescence detection. *Organic Geochemistry* 45: 7–17:  
570 doi:10.1016/j.orggeochem.2012.01.005.
- 571 Draxler RR and Hess GD (1998) NOAA Technical Memorandum ERL ARL-224. .
- 572 Duffin KI, Gillson L and Willis KJ (2008) Testing the sensitivity of charcoal as an indicator of fire  
573 events in savanna environments: quantitative predictions of fire proximity, area and intensity.  
574 *The Holocene* 18(2): 279–291: doi:10.1177/0959683607086766.
- 575 Dunbar RIM (2009) Richard Wrangham, *Catching Fire: How Cooking Made Us Human*: Basic  
576 Books, New York, 2009, \$26.95. *Human Nature* 20(4): 447–449: doi:10.1007/s12110-009-9075-  
577 3.
- 578 Dzepina K, Arey J, Marr LC, Worsnop DR, Salcedo D, Zhang Q, et al. (2007) Detection of  
579 particle-phase polycyclic aromatic hydrocarbons in Mexico City using an aerosol mass  
580 spectrometer. *International Journal of Mass Spectrometry* 263(2–3): 152–170:  
581 doi:10.1016/j.ijms.2007.01.010.
- 582 Edwards NT (1983) Polycyclic Aromatic Hydrocarbons (PAH's) in the Terrestrial Environment—  
583 A Review. *Journal of Environmental Quality* 12(4): 427–441:  
584 doi:10.2134/jeq1983.00472425001200040001x.
- 585 Enache MD and Cumming BF (2006) Tracking recorded fires using charcoal morphology from  
586 the sedimentary sequence of Prosser Lake, British Columbia (Canada). *Quaternary Research*  
587 65(02): 282–292: doi:10.1016/j.yqres.2005.09.003.
- 588 Feurdean A (2021) Experimental production of charcoal morphologies to discriminate fuel  
589 source and fire type: an example from Siberian taiga. *Biogeosciences* 18(12): 3805–3821:  
590 doi:10.5194/bg-18-3805-2021.
- 591 Flannigan M, Stocks B, Turetsky M and Wotton M (2009) Impacts of climate change on fire  
592 activity and fire management in the circumboreal forest. *Global Change Biology* 15(3): 549–560:  
593 doi:10.1111/j.1365-2486.2008.01660.x.
- 594 Foley SF, Gronenborn D, Andreae MO, Kadereit JW, Esper J, Scholz D, et al. (2013) The  
595 Palaeoanthropocene – The beginnings of anthropogenic environmental change. *Anthropocene*  
596 3: 83–88: doi:10.1016/j.ancene.2013.11.002.
- 597 Frank-DePue L, Vachula RS, Balascio NL, Cahoon K and Kaste JM (2022) Trends in  
598 sedimentary charcoal shapes correspond with broad-scale land-use changes: insights gained  
599 from a 300-year lake sediment record from eastern Virginia, USA. *Journal of Paleolimnology*.  
600 Available at: <https://link.springer.com/10.1007/s10933-022-00260-x>: doi:10.1007/s10933-022-  
601 00260-x.
- 602 Gedye SJ, Jones RT, Tinner W, Ammann B and Oldfield F (2000) The use of mineral  
603 magnetism in the reconstruction of fire history: a case study from Lago di Origgio, Swiss Alps.  
604 *Palaeogeography, Palaeoclimatology, Palaeoecology* 164(1–4): 101–110: doi:10.1016/S0031-  
605 0182(00)00178-4.

- 606 Glasspool IJ, Edwards D and Axe L (2004) Charcoal in the Silurian as evidence for the earliest  
607 wildfire. *Geology* 32(5): 381: doi:10.1130/G20363.1.
- 608 Glikson A (2013) Fire and human evolution: The deep-time blueprints of the Anthropocene.  
609 *Anthropocene* 3: 89–92: doi:10.1016/j.ancene.2014.02.002.
- 610 Gu Y, Pearsall DM, Xie S and Yu J (2007) Vegetation and fire history of a Chinese site in  
611 southern tropical Xishuangbanna derived from phytolith and charcoal records from Holocene  
612 sediments. *Journal of Biogeography* 0(0): 070821084123002-??? doi:10.1111/j.1365-  
613 2699.2007.01763.x.
- 614 Gunnar E (1960) The acetolysis method-a revised description. *Sven Bot Tidskr* 516–564.
- 615 Herring JR (2013) Charcoal Fluxes into Sediments of the North Pacific Ocean: The Cenozoic  
616 Record of Burning. In: Sundquist ET and Broecker WS (eds) *Geophysical Monograph Series*.  
617 Washington, D. C.: American Geophysical Union, 419–442. Available at:  
618 <http://doi.wiley.com/10.1029/GM032p0419>: doi:10.1029/GM032p0419.
- 619 Higuera P, Peters M, Brubaker L and Gavin D (2007) Understanding the origin and analysis of  
620 sediment-charcoal records with a simulation model. *Quaternary Science Reviews* 26(13–14):  
621 1790–1809: doi:10.1016/j.quascirev.2007.03.010.
- 622 Iyer V, Kawale S, Kotak V and Panchal M (2021) Mapping of forest fire area in Bandipur  
623 National Park, India using Sentinel 2B and Landsat 8 satellite data. .
- 624 Jensen K, Lynch EA, Calcote R and Hotchkiss SC (2007) Interpretation of charcoal  
625 morphotypes in sediments from Ferry Lake, Wisconsin, USA: do different plant fuel sources  
626 produce distinctive charcoal morphotypes? *The Holocene* 17(7): 907–915:  
627 doi:10.1177/0959683607082405.
- 628 Jia G, Peng P, Zhao Q and Jian Z (2003) Changes in terrestrial ecosystem since 30 Ma in East  
629 Asia: Stable isotope evidence from black carbon in the South China Sea. *Geology* 31(12): 1093:  
630 doi:10.1130/G19992.1.
- 631 Kong S-R, Yamamoto M, Shaari H, Hayashi R, Seki O, Mohd Tahir N, et al. (2021) The  
632 significance of pyrogenic polycyclic aromatic hydrocarbons in Borneo peat core for the  
633 reconstruction of fire history. *PLOS ONE* 16(9): e0256853: doi:10.1371/journal.pone.0256853.
- 634 Kulkarni C, Finsinger W, Anand P, Nogué S and Bhagwat SA (2021) Synergistic impacts of  
635 anthropogenic fires and aridity on plant diversity in the Western Ghats: Implications for  
636 management of ancient social-ecological systems. *Journal of Environmental Management* 283:  
637 111957: doi:10.1016/j.jenvman.2021.111957.
- 638 Laflamme RE and Hites RA (1978) The global distribution of polycyclic aromatic hydrocarbons  
639 in recent sediments. *Geochimica et Cosmochimica Acta* 42(3): 289–303: doi:10.1016/0016-  
640 7037(78)90182-5.
- 641 Lawson IT, Tzedakis PC, Roucoux KH and Galanidou N (2013) The anthropogenic influence on  
642 wildfire regimes: charcoal records from the Holocene and Last Interglacial at Ioannina, Greece.  
643 *Journal of Biogeography* 40(12): 2324–2334: doi:10.1111/jbi.12164.

- 644 Lerch M, Unkelbach J, Schneider F, Zech M and Klinge M (2022) Holocene vegetation  
645 reconstruction in the forest–steppe of Mongolia based on leaf waxes and macro-charcoals in  
646 soils. *E&G Quaternary Science Journal* 71(1): 91–110: doi:10.5194/egqsj-71-91-2022.
- 647 Leys B, Brewer SC, McConaghy S, Mueller J and McLauchlan KK (2015) Fire history  
648 reconstruction in grassland ecosystems: amount of charcoal reflects local area burned.  
649 *Environmental Research Letters* 10(11): 114009: doi:10.1088/1748-9326/10/11/114009.
- 650 Leys B, Carcaillet C, Dezileau L, Ali AA and Bradshaw RHW (2013) A comparison of charcoal  
651 measurements for reconstruction of Mediterranean paleo-fire frequency in the mountains of  
652 Corsica. *Quaternary Research* 79(3): 337–349: doi:10.1016/j.yqres.2013.01.003.
- 653 Lu H, Zhu L and Zhu N (2009) Polycyclic aromatic hydrocarbon emission from straw burning  
654 and the influence of combustion parameters. *Atmospheric Environment* 43(4): 978–983:  
655 doi:10.1016/j.atmosenv.2008.10.022.
- 656 Lynch JA, Clark JS and Stocks BJ (2004) Charcoal production, dispersal, and deposition from  
657 the Fort Providence experimental fire: interpreting fire regimes from charcoal records in boreal  
658 forests. *Canadian Journal of Forest Research* 34(8): 1642–1656: doi:10.1139/x04-071.
- 659 MacDonald GM, Larsen CPS, Szeicz JM and Moser KA (1991) The reconstruction of boreal  
660 forest fire history from lake sediments: A comparison of charcoal, pollen, sedimentological, and  
661 geochemical indices. *Quaternary Science Reviews* 10(1): 53–71: doi:10.1016/0277-  
662 3791(91)90030-X.
- 663 Mackay D and Shiu WY (1977) Aqueous solubility of polynuclear aromatic hydrocarbons.  
664 *Journal of Chemical & Engineering Data* 22(4): 399–402: doi:10.1021/je60075a012.
- 665 Marlon JR, Bartlein PJ, Carcaillet C, Gavin DG, Harrison SP, Higuera PE, et al. (2009) Erratum:  
666 Climate and human influences on global biomass burning over the past two millennia. *Nature*  
667 *Geoscience* 2(4): 307–307: doi:10.1038/ngeo468.
- 668 McGrath TE, Chan WG and Hajaligol MR (2003) Low temperature mechanism for the formation  
669 of polycyclic aromatic hydrocarbons from the pyrolysis of cellulose. *Journal of Analytical and*  
670 *Applied Pyrolysis* 66(1–2): 51–70: doi:10.1016/S0165-2370(02)00105-5.
- 671 Mooney SD and Tinner W (2011) The analysis of charcoal in peat and organic sediments. 19.
- 672 Musa Bandowe BA, Srinivasan P, Seelge M, Sirocko F and Wilcke W (2014) A 2600-year  
673 record of past polycyclic aromatic hydrocarbons (PAHs) deposition at Holzmaar (Eifel,  
674 Germany). *Palaeogeography, Palaeoclimatology, Palaeoecology* 401: 111–121:  
675 doi:10.1016/j.palaeo.2014.02.021.
- 676 Mustaphi CJC and Pisaric MFJ (2014) A classification for macroscopic charcoal morphologies  
677 found in Holocene lacustrine sediments. *Progress in Physical Geography: Earth and*  
678 *Environment* 38(6): 734–754: doi:10.1177/0309133314548886.
- 679 Nogué S, Tovar C, Bhagwat SA, Finsinger W and Willis KJ (2018) Exploring the Ecological  
680 History of a Tropical Agroforestry Landscape Using Fossil Pollen and Charcoal Analysis from  
681 Four Sites in Western Ghats, India. *Ecosystems* 21(1): 45–55: doi:10.1007/s10021-017-0132-1.

- 682 Pausas JG and Keeley JE (2009) A Burning Story: The Role of Fire in the History of Life.  
683 *BioScience* 59(7): 593–601: doi:10.1525/bio.2009.59.7.10.
- 684 Pereboom EM, Vachula RS, Huang Y and Russell J (2020) The morphology of experimentally  
685 produced charcoal distinguishes fuel types in the Arctic tundra. *The Holocene* 30(7): 1091–  
686 1096: doi:10.1177/0959683620908629.
- 687 Pisaric MFJ (2002) [No title found]. *Journal of Paleolimnology* 28(3): 349–354:  
688 doi:10.1023/A:1021630017078.
- 689 Ranjbar Jafarabadi A, Riyahi Bakhtiari A, Aliabadian M and Shadmehri Toosi A (2017) Spatial  
690 distribution and composition of aliphatic hydrocarbons, polycyclic aromatic hydrocarbons and  
691 hopanes in superficial sediments of the coral reefs of the Persian Gulf, Iran. *Environmental*  
692 *Pollution* 224: 195–223: doi:10.1016/j.envpol.2017.01.080.
- 693 Rummary TA (1983) The use of magnetic measurements in interpreting the fire histories of lake  
694 drainage basins. *Hydrobiologia* 103(1): 53–58: doi:10.1007/BF00028427.
- 695 Schumacher S and Bugmann H (2006) The relative importance of climatic effects, wildfires and  
696 management for future forest landscape dynamics in the Swiss Alps: FUTURE FOREST  
697 DYNAMICS IN THE SWISS ALPS. *Global Change Biology* 12(8): 1435–1450:  
698 doi:10.1111/j.1365-2486.2006.01188.x.
- 699 Scott AC (2000a) The Pre-Quaternary history of fire. *Palaeogeography, Palaeoclimatology,*  
700 *Palaeoecology* 164(1–4): 281–329: doi:10.1016/S0031-0182(00)00192-9.
- 701 Scott AC (2000b) The Pre-Quaternary history of fire. *Palaeogeography, Palaeoclimatology,*  
702 *Palaeoecology* 164(1–4): 281–329: doi:10.1016/S0031-0182(00)00192-9.
- 703 Scott AC (2010) Charcoal recognition, taphonomy and uses in palaeoenvironmental analysis.  
704 *Palaeogeography, Palaeoclimatology, Palaeoecology* 291(1–2): 11–39:  
705 doi:10.1016/j.palaeo.2009.12.012.
- 706 Scott AC and Glasspool IJ (2006) The diversification of Paleozoic fire systems and fluctuations  
707 in atmospheric oxygen concentration. *Proceedings of the National Academy of Sciences*  
708 103(29): 10861–10865: doi:10.1073/pnas.0604090103.
- 709 Stephens SL, Skinner CN and Gill SJ (2003) Dendrochronology-based fire history of Jeffrey  
710 pine - mixed conifer forests in the Sierra San Pedro Martir, Mexico. *Canadian Journal of Forest*  
711 *Research* 33(6): 1090–1101: doi:10.1139/x03-031.
- 712 Stevenson J and Haberle S (2005) PALAEOWORKS TECHNICAL PAPERS 5. .
- 713 Tinner W, Hofstetter S, Zeugin F, Conedera M, Wohlgemuth T, Zimmermann L, et al. (2006)  
714 Long-distance transport of macroscopic charcoal by an intensive crown fire in the Swiss Alps -  
715 implications for fire history reconstruction. *The Holocene* 16(2): 287–292:  
716 doi:10.1191/0959683606hl925rr.
- 717 Tinner W and Hu FS (2003) Size parameters, size-class distribution and area-number  
718 relationship of microscopic charcoal: relevance for fire reconstruction. *The Holocene* 13(4):  
719 499–505: doi:10.1191/0959683603hl615rp.

- 720 Tissot C, Chikini H and T.S N ( (n.d.)) *Pollen of wet evergreen forests of the Western Ghats,*  
721 *India.* .
- 722 Vachula RS (2021) A meta-analytical approach to understanding the charcoal source area  
723 problem. *Palaeogeography, Palaeoclimatology, Palaeoecology* 562: 110111:  
724 doi:10.1016/j.palaeo.2020.110111.
- 725 Vachula RS, Karp AT, Denis EH, Balascio NL, Canuel EA and Huang Y (2022) Spatially  
726 calibrating polycyclic aromatic hydrocarbons (PAHs) as proxies of area burned by vegetation  
727 fires: Insights from comparisons of historical data and sedimentary PAH fluxes.  
728 *Palaeogeography, Palaeoclimatology, Palaeoecology* 596: 110995:  
729 doi:10.1016/j.palaeo.2022.110995.
- 730 Vachula RS and Richter N (2018) Informing sedimentary charcoal-based fire reconstructions  
731 with a kinematic transport model. *The Holocene* 28(1): 173–178:  
732 doi:10.1177/0959683617715624.
- 733 Vachula RS, Russell JM, Huang Y and Richter N (2018) Assessing the spatial fidelity of  
734 sedimentary charcoal size fractions as fire history proxies with a high-resolution sediment record  
735 and historical data. *Palaeogeography, Palaeoclimatology, Palaeoecology* 508: 166–175:  
736 doi:10.1016/j.palaeo.2018.07.032.
- 737 Wang Q, Liu M, Li Y, Liu Y, Li S and Ge R (2016) Dry and wet deposition of polycyclic aromatic  
738 hydrocarbons and comparison with typical media in urban system of Shanghai, China.  
739 *Atmospheric Environment* 144: 175–181: doi:10.1016/j.atmosenv.2016.08.079.
- 740 Whitlock C and Larsen C (2002) Charcoal as a Fire Proxy. In: Smol JP, Birks HJB, Last WM,  
741 Bradley RS and Alverson K (eds) *Tracking Environmental Change Using Lake Sediments.*  
742 Dordrecht: Springer Netherlands, 75–97. Available at: [http://link.springer.com/10.1007/0-306-47668-1\\_5](http://link.springer.com/10.1007/0-306-47668-1_5):  
743 doi:10.1007/0-306-47668-1\_5.
- 744 Xie M, Hannigan MP and Barsanti KC (2014) Gas/particle partitioning of n-alkanes, PAHs and  
745 oxygenated PAHs in urban Denver. *Atmospheric Environment* 95: 355–362:  
746 doi:10.1016/j.atmosenv.2014.06.056.
- 747 Yan H, Yan Z, Wang L, Hao Z and Huang J (2022) Toward understanding submersed  
748 macrophyte *Vallisneria spiralis*-microbe partnerships to improve remediation potential for PAH-  
749 contaminated sediment. *Journal of Hazardous Materials* 425: 127767:  
750 doi:10.1016/j.jhazmat.2021.127767.
- 751 Yan Z, Song N, Wang C and Jiang H (2021) Functional potential and assembly of microbes  
752 from sediments in a lake bay and adjoining river ecosystem for polycyclic aromatic hydrocarbon  
753 biodegradation. *Environmental Microbiology* 23(2): 628–640: doi:10.1111/1462-2920.15104.
- 754 Yang C-R, Lin T-C and Chang F-H (2007) Particle size distribution and PAH concentrations of  
755 incense smoke in a combustion chamber. *Environmental Pollution* 145(2): 606–615:  
756 doi:10.1016/j.envpol.2005.10.036.
- 757 Yunker MB, Macdonald RW, Vingarzan R, Mitchell RH, Goyette D and Sylvestre S (2002) PAHs  
758 in the Fraser River basin: a critical appraisal of PAH ratios as indicators of PAH source and  
759 composition. *Organic Geochemistry* 33(4): 489–515: doi:10.1016/S0146-6380(02)00002-5.

