

Surface sampling of wetlands in a tropical dry forest after a large fire: Wind direction dominates the transport and deposition of fire proxies

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Abstract

A major fire raged through the dry tropical forests of Bandipur Tiger Reserve, a 'biodiversity hotspot' in southern India, in February 2019. The fire occurred in patches, burning >10,000 acres of forest, becoming one of the largest forest fires in India in recent times. Very few studies have been able to capture the dynamics of fire proxies from an active surface fire, especially in the dry tropics. We opportunistically sampled two wetlands roughly sandwiched between the largest and the second-largest burnt patches a week after the fire was extinguished. We collected surface samples from each wetland and looked at popular fire proxies – macrocharcoal, microcharcoal, microcharcoal/pollen (C/P) ratio and the abundance and distribution of Polycyclic Aromatic Hydrocarbons (PAHs). Macrocharcoal counts were low (mean ~5), while the C/P ratio was ~1. Low Molecular Weight (LMW) PAH molecules Phenanthrene, Anthracene, Fluoranthene and Pyrene were found in both sites, while High Molecular Weight (HMW) PAHs were only found in one. None of the proxies is particularly indicative of the large surface fire that occurred ~15 km away. Analysis of wind speed and direction from weather station data and forward and backward HYSPLIT model trajectories tell us that both wetlands were not downstream of the smoke plume. There was also no recorded precipitation between the fire event and the sampling date. Through this opportunistic study of fire proxies, we show that wind direction and wet scavenging are essential factors determining the transport and deposition of fire proxies in this environment. Hence reconstruction of fire histories should be done using multi-site data since the absence of fire proxies does not equate to the absence of fire.

1 Introduction

2
3 Approximately 420 million years ago (late Silurian Period) the emergence of land vascular plants
4 (Beerling, 2007) and the rise in the photosynthetic atmospheric oxygen (>13%), made fire ignition
5 possible on Earth (Bowman et al., 2009; Glasspool et al., 2004; Scott, 2000a; Scott and Glasspool,
6 2006). By Late Tertiary, increase in oxygen (Berner, 2006) and burnable biomass, fire became a
7 global phenomenon capable of changing and shaping every terrestrial ecosystem across the globe
8 (Herring, 2013; Jia et al., 2003; Pausas and Keeley, 2009). In addition to being a global terrestrial
9 modifier, fire also played a critical role in human evolution as well. Cooking meat aided in the
10 development of brain size, bipedalism, and utilization of stone tools (Dunbar, 2009). The knowledge
11 of ignition and management of fire came out much later, out of necessity, due to the temperature
12 drop from the warm Pliocene to Pleistocene ice ages (deMenocal, 2004; Glikson, 2013). From 790-
13 690 kyr, humans started using fire more extensively (deMenocal, 2004) and subsequently, they
14 started to make appreciable effects on the forest ecology (Archibald et al., 2012; Bowman et al.,
15 2011; D. Burton, 2009; Marlon et al., 2009). With these increased anthropogenic pressures and the
16 changes induced due to climate change and global warming on fire-prone ecosystems, several
17 concerns have been raised about their fate and resilience (Bowman, 2015; Flannigan et al., 2009;
18 Foley et al., 2013; Lawson et al., 2013; Schumacher and Bugmann, 2006). The information that is
19 required to answer such concerns is achieved through the study of the long-term trends of vegetation
20 and fire histories.

21
22 In the last 50 years, several palaeoecological investigation techniques have been developed to aid
23 this process. This includes the investigation through charcoal particles (Clark, 1988; Leys et al.,
24 2013; Scott, 2000b; Tinner and Hu, 2003; Vachula, 2021; Whitlock and Larsen, 2002) Polycyclic
25 Aromatic Hydrocarbons (PAHs) (Denis et al., 2012; Edwards, 1983; Kong et al., 2021; Musa
26 Bandowe et al., 2014; Vachula et al., 2022; Yunker et al., 2002) in archeological material and natural
27 paleoenvironmental archives, magnetic susceptibility of soil (Gedye et al., 2000; Rummery, 1983),
28 fire scars in dendrochronology (Stephens et al., 2003), and also from historical documents. Of all
29 these methods, charcoal-based reconstruction is the most widely used due to several advantages.
30 Charcoal is inert (Scott, 2010), which results in its being well-preserved in a wide variety of
31 environmental conditions. Charcoal has the potential to predict various characteristics of fire like

intensity (Duffin et al., 2008), severity (Whitlock and Larsen, 2002), extent, and frequency (Leys et al., 2013). Charcoal morphologies have also been used to trace the type of vegetation burnt (Enache and Cumming, 2006; Frank-DePue et al., 2022; Jensen et al., 2007; Mustaphi and Pisaric, 2014). The latter has developed an elaborate classification of sedimentary charcoal particles to identify fuel sources. Particles were categorized into seven major morphological classes (A to G) based on the overall shape, subdivided into 27 subclasses based on dominant surface textures or major features. With this classification scheme, they introduced an identification flow chart to identify fuel sources. Charcoal has also been combined with other proxies such as pollen (MacDonald et al., 1991), phytolith (Gu et al., 2007), molecular fire proxies (Argiriadis et al., 2018; Conedera et al., 2009; Denis et al., 2012; Musa Bandowe et al., 2014; Vachula et al., 2022) waxes (Lerch et al., 2022) and, sterols (Argiriadis et al., 2018). Since several other proxies can be simultaneously analyzed and creatively combined with charcoal, it provides robust ways of examining linkages between forest fires, vegetation, and anthropogenic activities.

However, these reconstruction techniques with charcoal are not without problems. The lack of standardization in extraction procedures, identification, and quantification of macrocharcoal (Conedera et al., 2009) is of great concern. The uncertainties in relationships between particle size, burnt area, and dispersal distance (Clark, 1988; Higuera et al., 2007; Marlon et al., 2009; Vachula, 2021; Vachula et al., 2018; Vachula and Richter, 2018) as well as the use of charcoal morphology to trace fuel sources are yet to be resolved (Enache and Cumming, 2006; Frank-DePue et al., 2022; Mustaphi and Pisaric, 2014). There are also variability associated with the production rate of charcoal depending on the vegetation burned (Feurdean, 2021; Pereboom et al., 2020; Vachula et al., 2022). Additionally, factors like wind speed and direction, morphology and size of the airborne particles, geographic location of sampling and source sites are also likely to affect the distribution, deposition, and retrieval of these proxies.

Another recently developing fire indicator is the polycyclic aromatic hydrocarbons (PAHs). They are produced either by the combustion of fossil fuels or by the burning of organic materials during a forest fire (Edwards, 1983). The source of the PAH can be distinguished by certain ratios of specific PAHs (Denis et al., 2012; Vachula et al., 2022; Yunker et al., 2002). There are several advantages associated with this proxy. The PAH produced can vary depending on the combustion temperature

(Blumer, 1976; Laflamme and Hites, 1978) and they have a broader temperature range than charcoal (Conedera et al., 2009; Lu et al., 2009). Moreover, PAH assemblages also vary with fuel type, fuel moisture and oxygen availability (Burns et al., 1997; Lu et al., 2009; Yang et al., 2007). Thus, PAH analysis can give a lot of information about the fire event and the fuel sources, similar to charcoal. A few studies have in fact used PAH as indicators of fire along with charcoal (Denis et al., 2012; Kong et al., 2021; Vachula et al., 2022; Yunker et al., 2002) but there are problems associated with this as well. The significant ones being lack of understanding of the factors that control their production, spatial distribution, and deposition mechanism. The PAH in the sediment can also undergo microbial degradation, especially the low molecular ones, leading to its depletion in the sediment (Yan et al., 2022, 2021). Further, the low molecular ones are also subjected to degradation during transport (Vachula et al., 2022).

Thus, certain aspects of these proxies (charcoal and PAH) are not very well understood and thus require a modern fire-proxy calibration to see how these proxies are captured and preserved in environmental archives. It also gives an idea on how they behave during after the fire event and also about the transport mechanism. Studies have been conducted on various fire-prone ecosystems like temperate, boreal, subalpine forests, grasslands, and savannahs across the world, making use of atmospheric and lake sediment traps, lake sediment cores, and lake surface sampling (Aleman et al., 2013; Clark et al., 1998; Leys et al., 2015; Lynch et al., 2004; Pisaric, 2002; Tinner et al., 2006). But none have been reported so far from the frequently burning tropical dry forests, especially from India, and this would be the first study reporting it.

In this study, we look at the modern fire-proxy relationship in the dry tropical forests of southern India, through an opportunistic sampling of two wetlands in the Mudumalai Tiger Reserve (MTR), after a fire broke out on 22 February 2019. The fire occurred in patches, in the forests of Bandipur, Wayanad, and Mudumalai Tiger reserves of the Nilgiris Biosphere Reserve in the Western Ghats. According to a newspaper report, this was purported to be a man-made disaster caused by miscreants deliberately setting elephant dung on fire (Chennabasaveshwar, 2019). The initial fires were brought under control within hours, but the dominant easterly winds of > 9km/hr speed (Kargudi weather station data), aided its spread, burning an area of more than 10,920 acres in just three days. The fire was brought under control by 26 February by the collective effort of forest officials, local

people, volunteers from the public and the Indian Air Force. A satellite study reported that the land temperature during the fire reached up to 140°C (Iyer et al., 2021). Even a week after the main fire event, complete extinguishment of the residual fire was not achieved.

Methods

Sample location

The samples were collected from two wetlands, Ombetta (110°34'36.73"N, 76°03'29.24"E) and Imbrella (11°03'36.73"N, 76°03'29.24"E), one week after the fire, on 7 March 2019 (Fig 1). The fire occurred in patches with the largest area burnt in the Bandipur Tiger reserve (BTR), while a second, smaller patch burnt in the adjacent Mudumalai Tiger reserve (MTR). Ombetta is ~5 km north of the MTR patch and 15 km south of the BTR patch. Imbrella is located ~7 km north of the MTR patch and 10 km south of the BTR fire patch. These two wetlands are separated by a distance of about three kilometers.

Sample Collection

A total of 20 samples, 10 from each wetland were collected by walking along the edges. The ten samples comprise of five surface samples collected by scooping 2 cm of the surface and five bottle core samples collected by pressing a 5 cm tall plastic bottle into the sediment and collecting the contents. The collected samples were stored in a freezer in the field base at -20° C. They were transported in frozen state to the lab at the Indian Institute of Science, Bengaluru, where they were freeze-dried (Labconco bulk drier), crushed gently using pestle and mortar and stored in room temperature until further analysis.

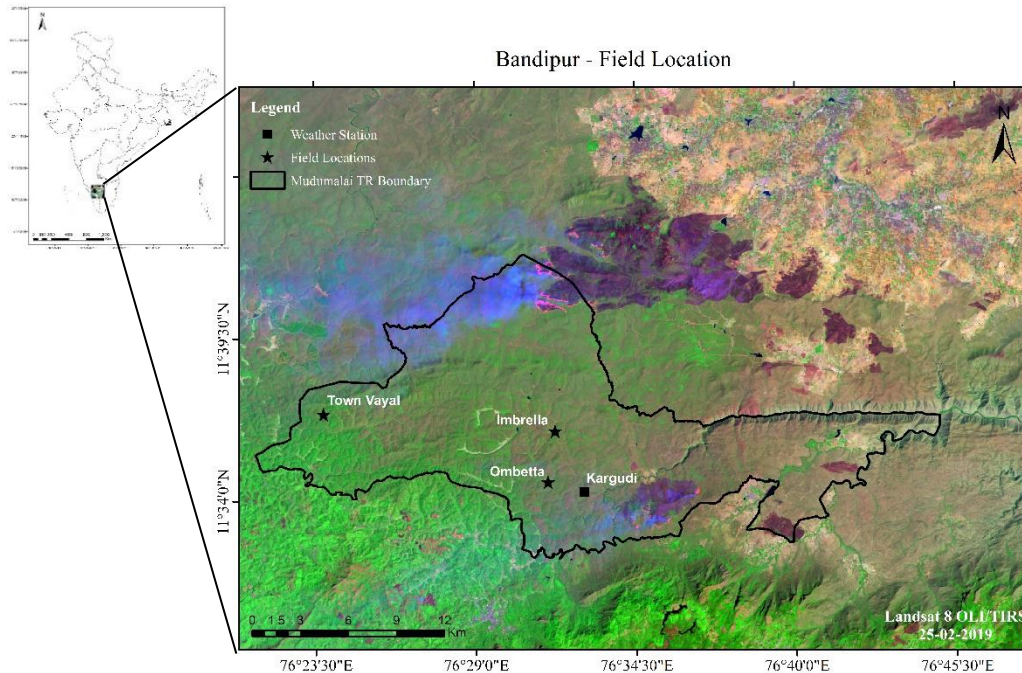


Fig 1: Landsat Image of Mudumalai Tiger Reserve (MTR) during the final day of the fire. The black line demarcates MTR boundaries. Imbrella and Ombetta (black stars) are the sampling wetlands, black square is the Kargudi weather station from which the wind data was collected. Two dark violet patches on the top of MTR and southeast of Kargudi are the major fire patches. The fire started in east and spread towards the west. You can see smoke arising moving west.

Macrocharcoal analysis

One gram of sample was taken and kept in 10ml of 10% sodium hexametaphosphate ($\text{NaPO}_3)_6$ solution overnight to deflocculate the sediments. We found Sodium hexametaphosphate was better in dispersing the clay than KOH (Andreola et al., 2004). This was followed by Sodium hypochlorite (NaOCl) treatment to bleach the organics (Stevenson and Haberle, 2005). The samples were then sieved through a 125 μm mesh and residue $>125 \mu\text{m}$ was collected for identification. Macrocharcoal particles are black, opaque and usually planar and were counted under a stereomicroscope (LeicaS4E) at 10X magnification. For each sample, three extractions were done and the charcoal particles counted, its mean, median, maximum and minimum values are given in fig 2.

Microcharcoal and Pollen counts

Microcharcoal extraction was done using the same method used for pollen (Bennett, 1990), the standard procedure followed at the French Institute of Pondicherry (FIP). In this method, samples are treated with a dispersing agent and coarse particles ($>150\text{ }\mu\text{m}$) and clay particles are removed. Then it is treated with 10% hydrochloric acid (HCl) to remove carbonates and (48%) hydrofluoric Acid (HF) to remove silicates. This is followed by acetolysis (9:1 mixture of acetic anhydride $(\text{CH}_3\text{CO})_2\text{O}$ and sulfuric acid (H_2SO_4) to remove polysaccharides (Gunnar, 1960). The sample is then centrifuged and the supernatant is removed. The extracts were mounted on glass slides using glycerin. Both microcharcoal and pollen grains were counted (minimum of 1000 pollen) under a microscope (Olympus CX43). Pollen counting was done using the Thanikaimoni reference collection at (FIP) (Anupama et al., 2014;).

Polycyclic Aromatic Hydrocarbons (PAHs)

Sediment samples, in powdered form ($\sim 6\text{ g}$), were combined homogeneously with unsaturated silica. This blend was subsequently packed into stainless steel enclosures and subjected to lipid extraction in a Buchi Speed Extractor E-914. The solvent mixture used was dichloromethane (DCM) and methanol in a ratio of 93:7. The extractor parameters were set to conduct two cycles at a temperature of $100\text{ }^\circ\text{C}$ and a pressure of 70 bar, facilitating the maximal extraction of organic material. Post extraction, the substance obtained was concentrated carefully at $30\text{ }^\circ\text{C}$ using a Buchi P-6 Multivapor. The total lipid extract (TLE) dissolved in 2 ml of hexane, was then transferred into a glass column filled with 2 cm of glass wool and 17 cm of silica gel. 20 ml of hexane was used to isolate the saturated hydrocarbon component from the TLE. The aromatic fraction was eluted using 80 ml of a solution of hexane-DCM in the ratio of 4:1. This fraction was then completely condensed via dry nitrogen gas and dissolved in a final volume of 0.5 ml DCM for further analysis (Ajay et al., 2021; Behera et al., 2022)

The fractionated aliquots were subjected to analytical testing using gas chromatography-mass spectrometry (GC-MS; Agilent 7890B/5977 MSD), adhering to the procedure detailed in (Behera et al., 2022). A non-polar capillary column (HP5-MS, with dimensions of $30\text{ m} \times 250\text{ }\mu\text{m} \times 0.25$

μm) was utilized for this analysis, with helium as carrier gas. The injection method used was splitless, with the inlet temperature set at 320 °C and a maximum sample volume of 1 μL. The GC oven's programming commenced with a base temperature of 40 °C, held steady for 2 minutes, followed by a progressive increase to a final temperature of 320°C at a gradient of 4°C per minute. The rate of flow for the helium gas within the column was sustained at 1.4 cm²/sec. As for the mass spectrometer detector (MSD) conditions, an EI ionization source of 70 eV was used, with a mass range spanning from 45–600 atomic mass units (amu). The multiplier voltage was set to 2341 V, with the ion-source temperature maintained at 230 °C.

The polycyclic aromatic hydrocarbons (PAHs) were recognized by aligning the distinct mass spectra with existing literature and the National Institute of Standards and Technology (NIST) database. In order to evaluate the quantity of PAHs present within the samples, an external calibration curve was generated. This was based on analytical standards with a range of concentrations (n=8), incorporating the Sigma CRM47930 PAHs mix with concentrations of 50, 100, 200, 300, 400, 600, 800, and 1200 ng/mL. Several indices related to PAHs were determined to aid in the quantification and understanding of the varied origins contributing to the organic content in the system.

Wind Trajectory Analysis

To track the wind direction, we used the Hybrid Single-Particle Lagrangian Integrated Trajectory model (HYSPLIT) (Draxler and Hess, 1998). This computer model is widely used to forecast wildfire smoke trajectories as it can impact air quality and human health. It also has archival data of the movement of air parcels. This archival information was combined with two functions of the model – 1. Forward trajectory that gives the direction the air parcel moved from a given point of interest and 2. Backward trajectory that gives the direction from which air parcels came to the point of interest. In our case, we used forward trajectories for the fire locations to see the direction of potential charcoal transport and backward trajectories on sampling locations to see the potential sources of charcoal particles deposited. The model was run on all days starting from 22nd February to the date of sample collection. It was run for 24 hours each day with new trajectories starting every 2 hours. Since this was a large fire, we considered wind parcels at

different altitudes ranging from 100 meters to 1 kilometer from the ground level. We also collected wind speed and wind direction data from Kargudi weather station, about 2 km and 4 km southeast of Ombetta and Imbrella, respectively. The Kargudi weather Station is close to the fire patch in MTR (~5 km) (fig 1).

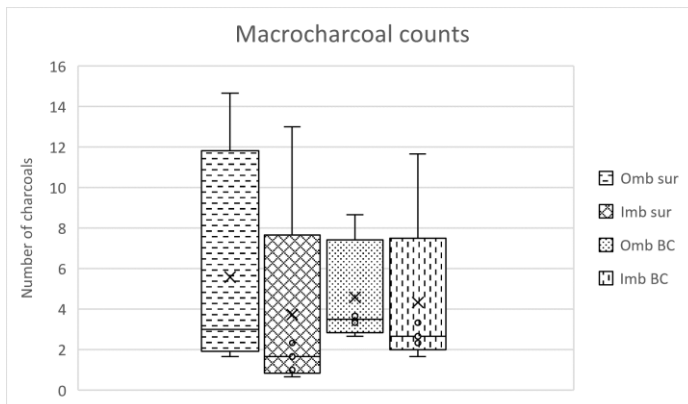


Fig 2: Box plot of macrocharcoal counts per gram of soil. Three extracts were done for each sample. The mean, median, maximum and minimum values of charcoal obtained are given in this plot. For all the samples, average number of charcoals lies between 4 and 6.

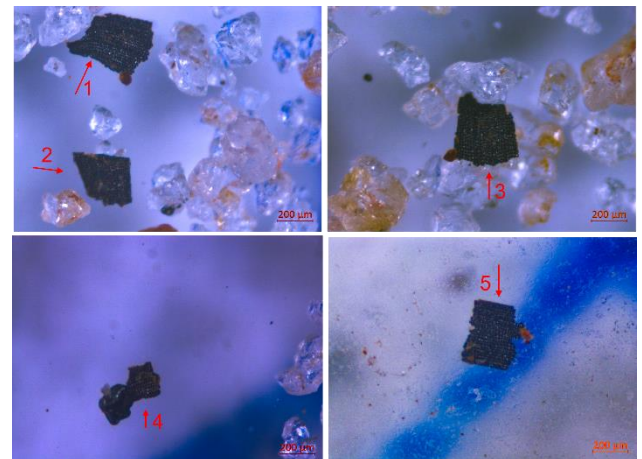


Fig 3 macrocharcoal particles (red arrows) found from the wetland samples. Only macrocharcoal of this type was found from the samples. This according to Mustaphi and Pisaric, 2014's classification scheme is of wood origin. The W/L ratio of these charcoals also show that they are of wood origin. The transparent crystals surrounding the charcoals are quartz. The W/L ratio of 1, 2, 3, 4 and 5 are 0.95, 0.80,

Results

Macro and Microcharcoal

Macrocharcoal counts of both sets of samples, surface and bottle core, were disproportionately low compared to the range and extent of the forest fire (see fig 2). The maximum number of macrocharcoal counted was 15 from one Ombetta surface sample, while the average number of particles was ~5 for all samples from both the wetlands. All the charcoal particles observed have

a textured broad shape indicative of wood burning (Mustaphi and Pisaric, 2014) (fig 3). The width to length ratio of all these particles are >0.5 which again indicates a woody origin (Aleman et al., 2013) (fig 3). Macrocharcoal counting of the top 5 cm of Town Vayal (TV), another wetland from a very low fire frequency area of Mudumalai was also done. This top layer sample had 13 charcoals per gram of sample. This sample also had morphologically similar type of macrocharcoal in them. Microcharcoal to pollen (C/P) ratio was 1.2 and 1.1 for Ombetta and Imbrella respectively. Length to width ratio of the microcharcoals gave values ranging from 1.2 (leaf) to >2.13 (wood) to >3.62 (grass) (Cui et al., 2009) (fig 4). Thus, vegetation of all types is represented in the pollen slides while only wood charcoals are represented in the macrocharcoal.

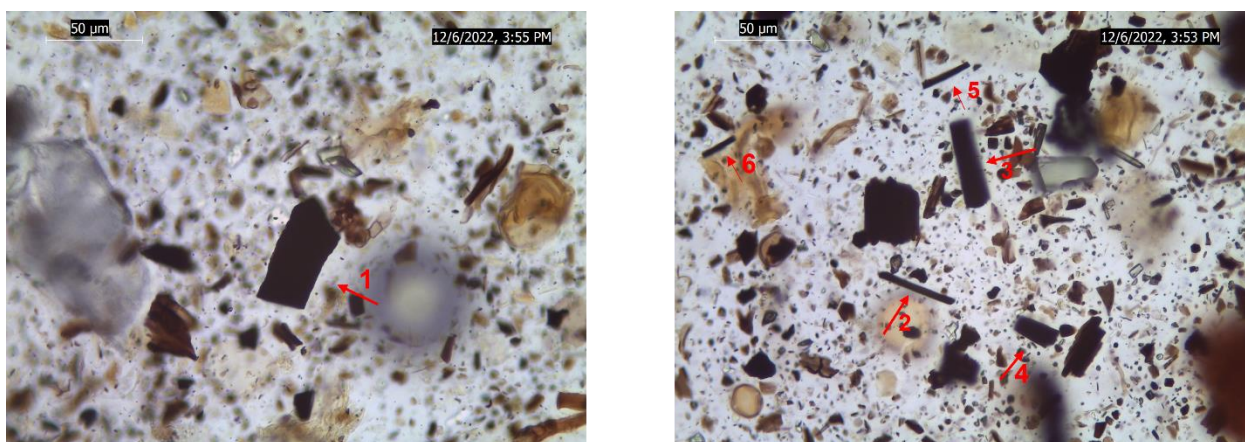


Fig 4: microcharcoal images of the wetland samples. The red arrows show the microcharcoal. Unlike macrocharcoal, microcharcoal of all sizes are found in the sample. The L/W values of 1,2,3,4,5, and 6 are 1.97, 12.3, 3.90, 2.19, 1.22, and 6.55

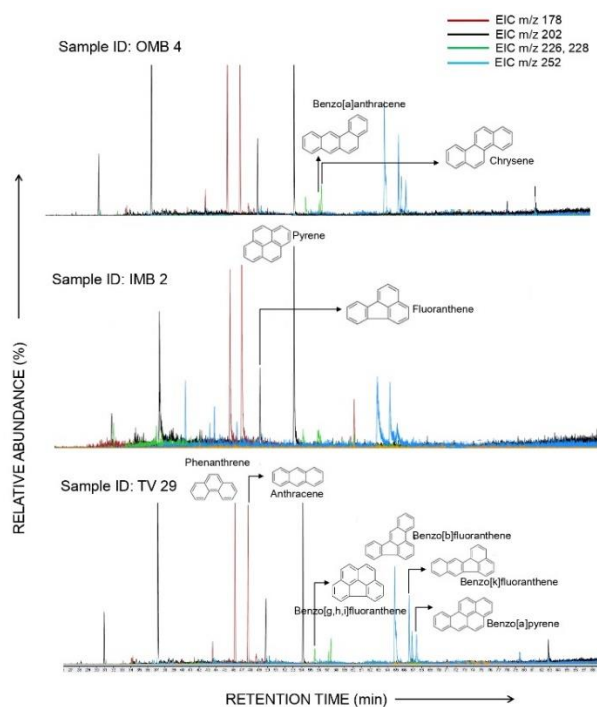


Fig 5: Chromatograms of the 10 PAHs obtained. These samples are representative of all the samples analysed. OMB and IMB are the wetland samples and TV29 is a surface sample from a low fire-frequency area of MTR. The PAHs found include phenanthrene, anthracene, fluoranthene, pyrene, benzo[g,h,i]fluoranthene, benz[a]anthracene, chrysene, benzo[b]fluoranthene, and benzo[a]pyrene, shown in the figures with representative peak positions.

PAH Analysis

Out of the 32 PAH compounds reported in literature (Dzepina et al., 2007), 10 compounds were detected from the nine surface samples of the wetlands (fig 5). This includes four Low Molecular Weight (LMW) ones, phenanthrene, anthracene, fluoranthene and pyrene and six High Molecular Weight (HMW) ones, benzo[g,h,i]fluoranthene, benz[a]anthracene, Chrysene, benzo[b]fluoranthene, and benzo[a]pyrene. All the Imbrella samples are missing benzo[g,h,i]fluoranthene and benz[a]anthracene except for IMB 4 and IMB 5 respectively. LMW/total PAH value shows that Ombetta has low LMW/ Σ PAHs (mean ~ 0.37) while Imbrella has comparatively high values (mean ~ 0.74). However, Imbrella is generally low in PAHs (mean of total PAHs ~ 0.76 ng/g) compared to Ombetta (mean of total PAHs ~ 3.5 ng/g). Imbrella also lacks many of the higher end compounds. The Σ LMW PAHs of Imbrella is also low compared to Ombetta. Ombetta is rich in HMW PAHs (avg ~ 0.63). Furthermore, Town Vayal surface sample, which is from a low fire-frequency area, also yielded the same 10 compounds and is rich in LMW compounds ($>80\%$ of total PAH detected). The Σ PAHs for town vayal is comparable to Ombetta (2.94 ng/g) but is generally rich in LMW PAHs than HMW unlike Ombetta.

Ratios of different PAHs were used to determine the source of PAHs. Yunker et al., 2002 show that samples with anthracene / (anthracene (Ant)+ phenanthrene (Ph)) ratio >0.1 are pyrogenic in origin. This value for Ombetta, Imbrella and Town Vayal lies between 0.75 - 0.81, 0.38-0.61 and 0.69 respectively. Similarly, for the Fluoranthene (Fla) / (fluoranthene + pyrene (Py)) ratio >0.5 indicate pyrogenic source. This value lies between 0.53-0.61, 0.4-0.63, and 0.58 for Ombetta, Imbrella and town vayal respectively. One sample of Imbrella (IMB 1) shows petrogenic signature, while rest of the samples including the town vayal are pyrogenic in origin (Fig 6a). Ratios of Fla/Py vs Ph/Ant (Fig 6b) and Fla/ Py vs LMW/HMW (Fig 6c) (Ranjbar Jafarabadi et al., 2017) also suggest pyrogenic source for all the samples except for town vayal which has a high LMW/HMW value (Fig 6c).

Table 1 PAH analysis summary

Site	Σ PAH	Σ LMW	Σ HMW	$a/(a+p)$	$f/(f+p)$	LMW/HMW	f/p
IMB	0.764637	100277	41398	0.5614	0.56	0.863	1.34
OMB	3.5893	242044	432627	0.7802	0.58	0.592	1.39
TV	2.9468	427586	125852	0.6925	0.58	3.397	1.40

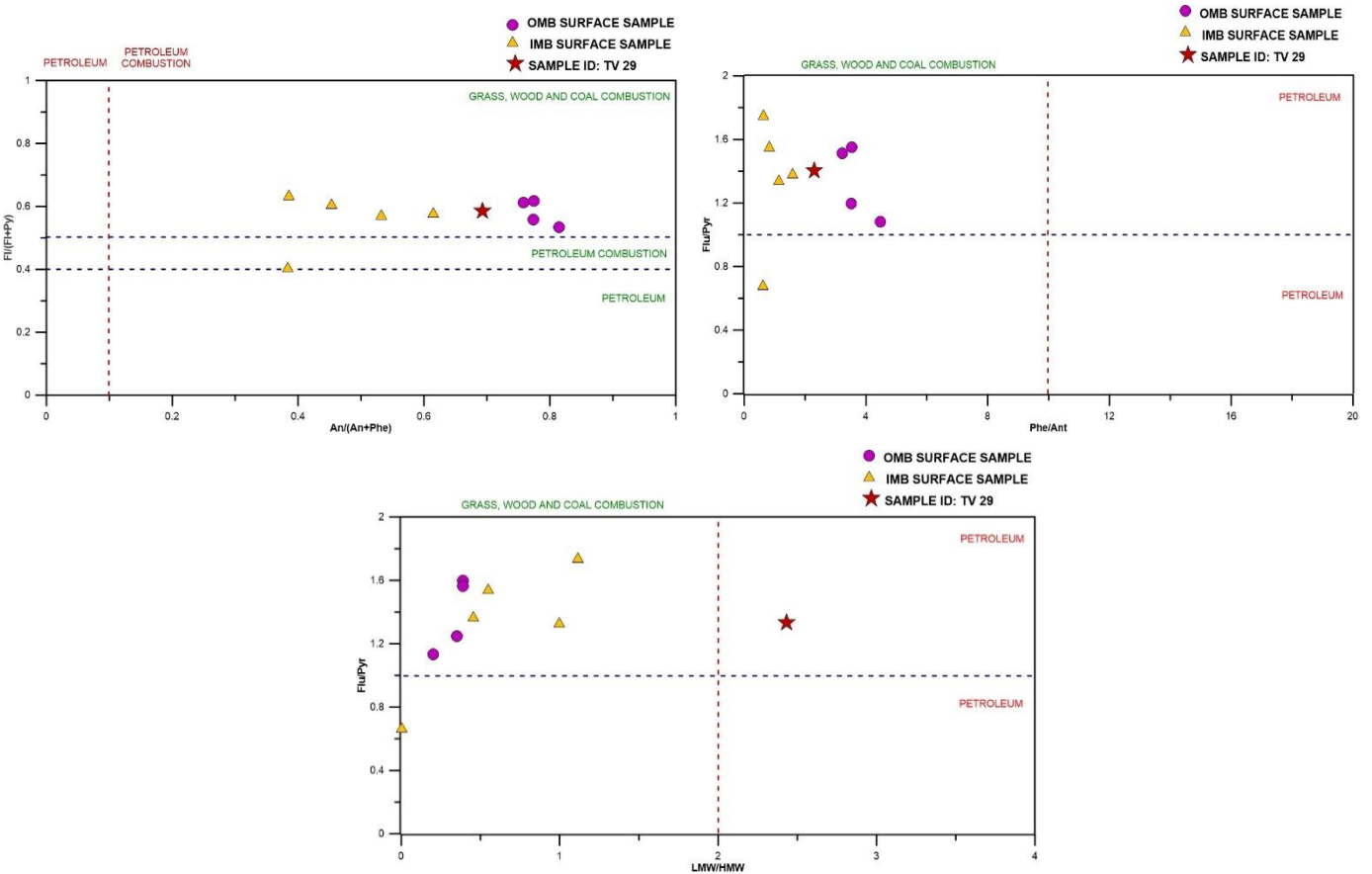


Fig 6: PAH source determination plots

Wind Trajectory Analysis

From the Kargudi weather station data, we created a wind rose diagram to observe the general wind pattern during and after the fire events (fig 7). During the fire event (22 – 26 February 2019), the weather station recorded a wind speed ranging from 0.3 to 18.3 km/hr. The wind rose diagram (fig 7) showed that winds with speed greater than 9.3 km/hr dominated these days. More than 60% of the winds were from the Northeast, East or Southeast and thus in general, easterly in direction. This seems to have aided the spread of the fire towards the west. The westerly winds start to get stronger towards the end of the fire event. After the fire event, from 27 February 2019, wind speed ranged between 0.3 to 12.3 km/hr, more than 65% of the wind was westerly.

HYSPLIT model was used to look at wind trajectories every hour of the day at different heights from the surface. The forward trajectory models from the major fire locations were created. The results show that wind from the larger Bandipur fire patch did not pass through any of the sample locations both during and after the fire event except on 6 March, 2019, one day before the sample collection. Forward trajectory models on the smaller MTR fire patch show that the wind was passing through the sample locations both during and after the fire days. This implies that whatever fire proxies were captured by the wetlands would be dominated by the MTR fire source. A backward trajectory of the wind parcels at the sample locations was also analyzed that confirmed the conclusions from the forward trajectory models (REF Supplementary material (III)).



Fig 7: Wind rose diagram created from the Kargudi weather station data. The left side image is after fire (27th feb to 7 of march, 2019). Right side image is from 22nd feb to 26th feb, 2019, when the fire was prominent. During the fire days the wind was mostly westerly. More than 40% of them are high speed wind (15.3-18.3 km/hr). Easterly winds start to develop towards the end of fire. After the fire event, the wind is mostly from the east with speed ranging 12.3-6.3 km/hr. The change in wind direction has in fact helped in curbing the fire.

Discussion

Macrocharcoal and Microcharcoal Count

Macrocharcoal counts from both wetlands are low (mean ~5) for all the samples. Whitlock and Larsen (2002), in their study, show that macrocharcoal count >50 was indicative of fire events. However, we must note that tropical forests are not included in this study and hence this arbitrary value may not be applicable to our study from tropical India. Since ours is the first study to try and evaluate fire-proxy relationships in the dry tropical forests of India, we do not yet have baseline values for a fire peak representation in surface sediments in this landscape. Hence we chose to compare our results to investigations on sediment cores from the larger region of the Western Ghats (Bhagwat et al., 2012, 2014; Kulkarni et al., 2021; Nogué et al., 2018) and a surface sample (top 5cm) from Town vayal. These investigations on fire-vegetation-human inter-relationships have been done in the Kodagu district of Karnataka, central Western Ghats in anthropogenically modified landscapes, albeit in tropical wet evergreen forests and very low burning part of the Mudumalai Tiger reserve, respectively. On one of the sediment cores, the authors report an increase of fire in two time periods: ~1800-1400 yr BP and 400-0 yr BP, was represented by 61 charcoal particles per cm³ of sample (Bhagwat et al., 2012; Kulkarni et al., 2021) respectively. This shows that >50 charcoal count for fire indication may hold for tropical India as well. However further study has to be done to confirm this. The surface of these sediment core had about 18 charcoals per cm³ of sample. This low count is not very surprising as the region is very unlikely to burn as it is a plantation area. This strongly implies that the surface of our wetlands is not representative of the large surface fire. The understanding that macrocharcoal as a result of their larger size and limited mobility would only be able to travel small distances and represent a local fire event aligns well with our findings as both the wetlands are 5-7 kms away from either of the fire patches (Clark, 1988; Mooney and Tinner, 2011).

Macrocharcoal: Morphology

The Bandipur fire was one of the largest recent forest fires and burnt large swathes of the forest, including canopy trees and understory vegetation – both grass and shrub. Yet, only macrocharcoal, possibly of woody plants origin, were found (irregular, opaque with textured surface), while charcoal from the other vegetation types were not found. However, microcharcoal indicative of diverse vegetation types were found in the samples (based on the large range of sizes and shapes (Fig 4). Macrocharcoal of similar morphology is also found from the low burning, 800-year, old Town Vayal sample (unpublished data) as well from its surface. Thus, the presence of this macrocharcoal in both low burning location and >7km from the fire edge (OMB and IMB) shows that it can travel larger distances. It could also be distinctive of the tropical dry deciduous forest.

Polycyclic Aromatic Hydrocarbons (PAHs)

All of our wetland samples show pyrogenic source for the PAHs (fig 6). TV is more proximate to human occupation; it is likely that the signatures of PAHs there are of a mixed pyro-petrogenic origin. The higher LMW/HMW value for TV sample would be indicative of that. However, since the sample was collected in what is well within Tiger Reserve boundaries, we expect that most of the PAHs would be of pyrogenic origin for the purposes of this study.

The individual PAH compounds can give an idea of the spatial distribution of the fire (Vachula et al., 2022). According to Vachula et al., 2022, naphthalene, anthracene, acenaphthylene, and fluorine record local fire histories (within 40km) whereas phenanthrene, benzo[g,h,i] perylene, chrysene, benzo[k]fluoranthene, benzo[b]fluoranthene record both local and regional fires. Our samples are rich in compounds that can record both local and regional fires and thus indicate that the big near fire event was not well captured by the wetland.

LMW/ Σ PAH is found to show a positive correlation with accumulation rate of sedimentary macrocharcoal of all size classes (Vachula et al., 2022). They thus postulate that the adsorption of PAHs (LMWs) to charcoal can be an important mode of transport and deposition for them.

However, (Kong et al., 2021) in their study of the peat samples from Malaysia observed a positive correlation of the HMW PAHs with charcoal abundance. The LMW was in fact decreasing with increasing charcoal abundance. In our case, both the wetlands and Town Vayal core top have low charcoal counts. While Town Vayal shows a high LMW/ Σ PAH, Ombetta is rich in HMW PAHs. This difference in concentration of LMW and HMW in low charcoal scenario highlights the possibility for more research in linking either of these PAHs with charcoal abundance. Based on our results neither of the PAHs show any relationship with macrocharcoal. This indicates that the increase in concentration is independent of adsorption to charcoal as a transportation mechanism.

HMW PAHs are produced more in a high temperature scenario (McGrath et al., 2003). Forest fires, depending on the fuel that is burned and the climate, can be highly variable in terms of temperature of burning. Even though both Ombetta and Imbrella are close to each other (~3km), there is a huge distinction between them in terms of the PAH that were captured (LMW vs HMW). Imbrella is closer to the bigger fire patch. However, the Σ PAHs that got captured by the wetland is low compared to Ombetta. This shows that the signals from the bigger BTR patch were not captured by the wetland. HYSPLIT model also confirms that none of the wind parcel from the BTR passed through our sample locations. The high value of HMWs in Ombetta (Ombetta is closer to MTR patch) indicates that the signal might be from the MTR patch. HYSPLIT model results show that some of the winds did go through our sample location (Ombetta).

High percentages of LMW PAHs compared to HMW PAHs in the wet forest sediments of Town vayal shows that they can travel farther and get deposited. Their presence in low amounts even in Imbrella shows that they can get deposited in less transport-favorable environments as well. The LMW PAHs have low gas-particle partition coefficient (Alam et al., 2014; Xie et al., 2014) and so they are mostly associated with gas phase(Wang et al., 2016). This property of the LMW PAHs makes sense with their long range of transport. The opposite is true for HMW PAHs. Thus, high percentages of HMW PAHs in Ombetta and absence/ low concentration in both Imbrella and Town Vayal shows that they are more local fire indicators and are particle or transport medium dependent. However, more modern-day calibration has tinbe done on this regard before using them as reliable fire proxies.

According to Wang et al., 2016, PAH flux deposition of LMW PAHs (2 + 3 rings) are high with wet condition. This is because the LMWs are more water soluble (Mackay and Shiu, 1977) and tend to be associated more with gaseous phase. The HMW PAHs (4+5+6 rings) on the other hand, favor dry deposition. The PAH flux deposition was also found to be high on a calm day (Wang et al., 2016). In our case, from the fire start day to the day of sample collection, the climate was dry and windy. So, most of the LMW PAHs might have been carried away by the wind as they are more associated with the gaseous phase. Thus, the dry deposition mode should have been more favored in such conditions. However, the absence of HMW PAHs in Imbrella shows that wind does determine their transport and deposition just as is the case for macrocharcoal.

Conclusion

From our opportunistic sampling of wetland surfaces following a large fire event, we report the following important findings. First, we demonstrate the importance of wind direction in proxy transport and deposition. We see that all fire proxies considered in both wetlands – macrocharcoal, microcharcoal/pollen ratio and PAHs is disproportionately low and not representative of one of the largest forest fire events in the recent past in India. This is ascribed to the lack of favorable wind for deposition of these proxies. Second, the distance travelled by macrocharcoal (blocky texture) in our study site. For the smaller burnt area whose downwind fire proxies we seem to have captured through our sampling of these two wetland surfaces, charcoal particles of size 120-200 μ has been transported up to 7 km which to our knowledge is reported for the first time in India. Third, we report ambiguity in the use of PAH as fire proxy. HMW PAHs seem to capture local and high intensity fire well while LMW PAHs seem to be regional in nature. We also report ambiguity with respect to transportation and preservation of PAHs and their relationship with wet and dry deposition events. More modern-day fire calibration is required to have a more nuanced understanding of the factors that affect fire proxy transport and deposition and to reconstruct past fire events with more confidence.

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