

1 **Drought-busting ‘miracles’ in the Colorado River Basin may become less frequent and less**
2 **powerful under climate warming**
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34 **Abstract:** Drought-busting events in the Colorado River Basin, such as the “Miracle May” of
35 2015 that greatly alleviated an unprecedented water shortage, have been observed for more than
36 a century. But while such events are much prayed for in times of drought, they have not been
37 well researched or even characterized. In this study, conducted in collaboration with water
38 managers from across the basin, we propose a definition for “miracle events” that reflects real-
39 world, actionable relevance. The resulting characterization offers a framework by which to
40 quantify the frequency and strength of extreme dry-to-wet springtime transitions. While limited
41 by uncertainties in model simulations and the myriad hydrological futures these simulations seek
42 to project, and thus requiring cautious interpretation, this study finds that such transitions may
43 become less frequent and less powerful under climate warming.

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45

46 **Keywords:** Miracle Spring, Colorado River Basin, drought index, climate models, stakeholders,
47 water managers

48

49 **1. Introduction**

50

51 In early 2015, the Colorado River Basin (CRB) faced an unprecedented water shortage. In the
52 “Four Corners” states of Utah, Arizona, New Mexico and Colorado, which comprise the majority
53 of the basin’s territory, three-quarters of the total landmass were in “abnormally dry” conditions,
54 which had persisted—or became worse—since 2012 without a single month of reprieve. From a
55 near-capacity high of 1,216 feet above sea level in 1998, Lake Mead, the largest reservoir in the
56 United States, had fallen to a historic low, plummeting more than 140 feet. Officials from the
57 U.S. Bureau of Reclamation estimated that Colorado River Compact allocations were deficient in
58 the order of 1.2 million acre-feet per year. And water managers across the region were imposing
59 historic curtailments. But in spring of 2015 a nearly threefold increase in the basin’s average
60 precipitation flipped the river system from severe shortage to flooding status in a matter of
61 weeks, and water managers rejoiced at the “Miracle May,” which boosted the basin’s water
62 storage and alleviated the immediate concerns about shortages.

63

64 This wasn’t the first “miracle” event, so dubbed because the anomalously strong rainfall had not
65 been predicted during springtime water planning. The term appears to have been coined in 1991,
66 when California was experiencing its worst drought since the Dust Bowl before a “Miracle
67 March” brought record-breaking spring snowfall that tripled the mountain snowpack. Other
68 winter-to-spring, dry-to-wet transitions that replenished the water supply at critical moments of
69 need are reflected in the historical record, including the prodigious precipitation events in 1915,
70 1945 and 1976.

71

72 Facing repeated, prolonged drought events across the West in recent decades (Sullivan et al.,
73 2019; Udall and Overpeck, 2017) alongside growing demand for water that continues to cut
74 away at the CRB’s finite supply (Rajagopalan et al., 2009; Woodhouse et al., 2021), water
75 managers may be allured by the opportunity to know whether climate models can project such
76 seemingly abrupt and extreme dry-to-wet springtime transitions, particularly when faced with the
77 prospect of a warming world. Under future warming scenarios, the contribution of Upper Basin
78 snowmelt, which comprises about 92% of the runoff for the entire CRB, could fall by one-third
79 (Lukas and Payton, 2020; Li et al., 2017). A warmer climate is also likely to produce a low-to-no
80 snow future in the CRB, which consequently impacts streamflow and water supply (Siirila-
81 Woodburn et al., 2021; Woodhouse et al., 2021; Lukas and Payton, 2020). As these futures begin
82 to emerge, phenomena that at one time seemed like unforeseeable and wondrous events may be
83 an important part of the scientific process of water management and planning—if, that is, the
84 frequency and intensity of these events can actually be predicted.

85

86 Thus far, however, springtime miracle precipitation events in the CRB have not been
87 quantitatively defined nor universally characterized. Thus, to understand the drivers of critical
88 dry-to-wet springtime transitions, how representable they are in climate models, and how they
89 might evolve in response to global climate change, a scientific definition that reflects real-world
90 management relevance must first be developed.

91

92 Since a common definition for extreme dry-to-wet transition events does not exist, and no
93 specific criteria have previously been developed, this research tests different metrics with a
94 variety of drought indices to create what we have dubbed the “Miracle Index.” Utilizing multiple

95 datasets to evaluate historical dry-to-wet transitions, and the extent to which a warmer climate
96 may impact similar situations in the future, the “Miracle Index” quantifies events that begin with
97 at least four consecutive anomalously dry months followed by at least three consecutive
98 anomalously wet months, hinged at a “month zero,” as shown in Figure 1.
99

100 As this index is not just intended for conceptual but also practical use, regional water managers
101 have been involved from the point of conception, and the research questions that drove this study
102 were based on their input during a series of meetings aimed at understanding how miracle events
103 might be relevant to water management (see Section 2.1). These questions include: How can
104 extreme dry-to-wet transitions factor into water management? How often do these events occur?
105 Can climate models predict these miracle events? Would past dry-to-wet transitions have
106 happened without a warmer climate? How will miracle events change in the future? Will
107 projected changes impact water management?
108

109 **2. Data and Methodology**

110

111 This study introduces a definition for extreme springtime dry-to-wet transitions (hereafter
112 “miracle events”) using common precipitation and drought indices, with an emphasis on creating
113 a metric that is relevant and useful to water managers in the CRB. To this end, we incorporate
114 the past experiences and future needs of water managers, based on a series of stakeholder
115 discussions, into our analysis of various drought indices, which have been chosen to better
116 understand how miracle events might be differentially characterized, with the intention of
117 arriving at an index that balances scientific value and utility.
118

119 **2.1. Stakeholders Discussions and Survey Data**

120

121 Co-production is a process that involves iterative and continual engagement between scientists
122 and decision-makers (in this case water managers) to collaboratively develop actionable science
123 (Jagannathan et al 2020, Kirchhoff et al. 2013). For this research, we used co-production
124 principles and approaches to develop a research plan, finalize research methods, and engage in
125 an iterative process based on feedback about the management relevance of key scientific results.
126 We engaged water managers from five different state and federal agencies located in Colorado
127 and Utah, and seven water management professionals from these agencies, who are also co-
128 authors of this paper, participated in iterative engagements with project scientists throughout the
129 course of the research.
130

131 Upon approval from Lawrence Berkeley National Laboratory’s Human Subjects Committee
132 Institutional Review Board for key engagements, three focus group discussions and three surveys
133 were conducted, interspersed by informal conversations and discussions over the course of two
134 years. The first focus group discussion aimed to co-produce the questions, approach and
135 potential outputs of the research. Subsequent focus groups were conducted to elicit iterative
136 feedback on interim results and approach. The surveys were intended to gather written feedback
137 on further directions for the work. Through these discussions and surveys, the managers and
138 scientists co-developed several components of the research, such as the key metrics of interest,
139 the spatial domain that is most decision-relevant, the uncertainty range, and the interpretation of
140 climate model projections. Notably, this feedback resulted in a broadening of the scope of the

141 research from only looking at specific events (such as the “Miracle May” of 2015 and similar
142 “miracle months”) to looking at the transitory periods that are centered on such events.

143

144 2.2. Observational data

145

146 To develop a definition of historical miracle events, we explored datasets including precipitation,
147 Palmer Drought Severity Index (PDSI), Palmer Hydrological Drought Index (PHDI), and
148 Standardized Precipitation Index (SPI, Guenang and Kamga, 2014) to analyze how these indices
149 change due to miracle precipitation in the Upper Basin and Lower Basin (UB and LB,
150 respectively). The observational data used in this study were downloaded from the U.S. National
151 Climatic Data Center for the Colorado Basin (Vose et al., 2014,
152 <https://www1.ncdc.noaa.gov/pub/data/cirs/climdiv/>) and assessed in different timescales (e.g. 1-,
153 3-, 6-, 12-, 24- month) from 1895 to 2019. Here, SPI3 means a running three month mean of SPI
154 representing the end month, and so forth, as we considered short-term, seasonal, and long-term
155 SPI analyses to see the change in drought in the CRB at different timescales. In this regard, SPI3
156 depicts semi-persistent moisture conditions and provides a seasonal estimation of precipitation.
157 Note that evaporative effect is not considered here due to the difficulty in deriving the
158 Standardized Precipitation-Evapotranspiration Index (SPEI) from model outputs.

159

160 While we initially analyzed both the UB and LB, based on stakeholder feedback we decided to
161 focus on the upper basin, which is the totality of the river network north of Lee’s Ferry in
162 northern Arizona, because the vast majority of natural flow of the Colorado River comes from
163 that segment of the CRB (Lukas and Payton, 2020). This decision was reviewed, and affirmed,
164 after it became clear that miracle events are significantly more common in the UB (as discussed
165 below).

166

167 2.3 Climate modeling data

168

169 Based on feedback from water managers, we used both Coupled Model Intercomparison Project
170 Phases 5 and 6 (CMIP5 and CMIP6) of IPCC global simulation data to investigate modeling
171 relevance and analyze projected miracle events. One limitation to the use of both phases of these
172 climate projections is that the CMIP5 data used here have been bias-corrected and statistically
173 downscaled, with a horizontal resolution of 0.125 degree (produced by the Bureau of
174 Reclamation CMIP5-BCSD; Reclamation, 2013; Maurer et al., 2007), while the CMIP6 data
175 were obtained from <http://climexp.knmi.nl/> with the default 1.25-degree horizontal resolution
176 and without statistical downscaling or bias correction (CMIP6 downscaling datasets are not yet
177 available). Thus, comparisons between the two sets of climate projections are limited. However,
178 the stakeholders preferred to have all possible information sources to consider when it comes to
179 future projections.

180

181 Each model of CMIP5 used in this study was driven by historical forcings (observations of
182 aerosols, greenhouse gasses, and solar irradiance) from 1950 to 2005, with the follow-on period
183 of 2006-2100 driven by forcings from Representative Concentration Pathway 8.5 (RCP8.5)
184 (Freychet et al., 2015). We used historical precipitation from an ensemble of 81 simulations from
185 30 global climate models of CMIP5 and an ensemble of 69 simulations from 28 models for
186 CMIP6. The CMIP6 includes shared socio-economic pathway (SSP) projections intended for

187 exploration of future climate conditions under various scenarios of population change, economic
188 growth, urbanization, technological development, and other factors that may influence emissions
189 and other climate-impacting physical processes (Mishra et al. 2020; Zhai et al. 2020). The
190 current study focuses on the SSP585 scenario, which assumes high radiative forcing (8.5 W m^{-2})
191 by the end of the century. SPI at different timescales was calculated to evaluate the dry, wet and
192 miracle events in both climate models. Historical (1850–2014) and future projected (2015–2100)
193 precipitation under different scenarios were used to derive SPI.

194

195 2.4 Regional pseudo-global warming simulation

196

197 The Pacific Northwest National Laboratory (PNNL) developed high-resolution pseudo global
198 warming (PGW) simulations using the Weather Research and Forecasting (WRF, Skamarock et
199 al., 2008) model. These simulations (hereafter PNNL-WRF) cover the Western U.S. with a grid
200 spacing of 6-km. The PNNL-WRF output includes a control simulation of 30 years driven by
201 boundary conditions from the North American Regional Reanalysis (NARR) during 1981 to
202 2020 (Chen et al. 2018). For future projections, five PGW simulations were produced by PNNL-
203 WRF by adding the NARR boundary conditions to the mean monthly perturbations derived from
204 the climate change signals simulated by five CMIP5 models (CanESM2, CESM1-CAM5,
205 GFDL-ESM2M, HadGEM2-ES, and MPI-ESM-MR) for the period of 2041-2070 following a
206 high-end scenario (RCP8.5) relative to the historical simulations for 1981-2010 from the same
207 five CMIP5 models. The PGW simulations provide a direct comparison for how historical
208 climate events may behave under warmer conditions.

209

210 3. Characterizing springtime miracle events

211

212 Different datasets were used to evaluate and define the extreme dry-to-wet transitions in the
213 CRB. “Month 0” is defined as a month following at least four consecutive winter months of
214 drought conditions and preceding at least three consecutive wet months that reverse the drought
215 situation through the impending summer. For standardized indices like PDSI and SPI, this
216 definition means a persistent sign reversal happening in month 0 as illustrated in Figure 1.
217 Snowpack metrics are not being considered here because the “miracle season” coincides with the
218 snowmelt season, such that even if the increased precipitation added snow on the mountains, it
219 would not last long. (Rainfall in May could accelerate snowmelt; this is discussed later.) Since
220 different drought indices may yield different results, we analyzed many indices and, in
221 collaboration with stakeholders, have chosen to display the results from three: PDSI, SPI3 and
222 SPI6.

223

224 The miracle events that are described by each of these indices do reveal some divergence, as
225 shown in Figure 2. Despite these differences, however, there is a 50% overlap between the
226 miracle events identified using the three indices. In the UB, both PDSI and SPI3 produced 11
227 miracle events (Figures 2a and 2b) while SPI6 (Figure 2c) produced 13 miracle events. Of note,
228 the LB produced fewer than five miracle events regardless of the index used (result not shown).

229

230 Based on all the miracle events identified by PDSI, SPI3 and SPI6, we constructed a composite,
231 shown in Figure 3. A notable caveat to this discussion is that there is ongoing debate as to which
232 indices best characterize different types of drought, such as meteorological drought vs.

233 hydrological drought (Slette et al., 2019), and both PDSI and SPI are meteorological drought
234 indices.

235
236 To validate the depiction of these events and examine the long-term behavior of drought with
237 respect to a miracle event, we also plotted a key hydrological index, the PHDI, as well as SPI1
238 (one month) and SPI24 (two years). While PDSI considers longer-term dryness that will affect
239 streamflow, PHDI offers a mechanism by which to assess how much precipitation is needed to
240 end a drought. The SPI1 histogram reveals the subseasonal behavior of precipitation with respect
241 to month 0, and the result in Fig. 3 suggests a prolonged precipitation change that is of seasonal
242 or longer term. This result indicates that the hydroclimatic change associated with an extreme
243 dry-to-wet transition is a seasonal anomaly that may be part of seasonal-to-interannual or even
244 longer variations. This inference is supported by the PHDI, which shows a more delayed
245 response, suggesting replenished water storage following month 0.

247 3.1 Relevant climatic anomalies

248
249 The CRB has experienced some persistent climate anomalies in recent decades, and we also
250 looked at how these anomalies have changed. In response to stakeholder interest, we analyzed
251 the frequency of “wet-wet” anomalies from a wet winter to a wet spring, based on December-
252 February and March-May seasons, using the SPI3 drought index for the CRB. “Dry-dry”
253 anomalies are also defined similarly. These persistent anomalies (“wet-wet” and “dry-dry”)
254 account for most of the historical years, reflecting the lower-frequency nature of hydroclimatic
255 variation in this region (Wang et al. 2018). Figure 4 shows the historical distribution of the
256 tendency in SPI3 based on the count of each type of anomaly within a 15-year bin. Both the UB
257 and LB have experienced an increasing trend in the frequency of dry-dry anomalies, meaning
258 that both winter and spring have become drier in recent years, echoing the worsened drought
259 situation as reported in the literature. The increased dry-dry anomalies and decreased wet-wet
260 anomalies are present in the lower CRB (Figure 4b); however only the trends in dry-dry
261 anomalies are statistically significant ($p < .01$, Figure 4).

262
263 By considering all possible combinations of dry-wet, wet-wet, dry-dry, and wet-dry situations
264 shown in Figures 4c and 4d, we found that the UB has more extreme miracle events than the
265 lower basin. This difference is conspicuous in the pie-charts (Figure 4c and 4d) and it is worth
266 noting that this form of graphic depiction was well received by the water managers as an
267 informative piece of climate analysis. We further looked at the historical precipitation and
268 drought strength (PDSI) over the two basins and found that annual precipitation has been
269 decreasing in both basins (Figure S1). Both basins show a similar inter-annual variability of
270 precipitation and drought. Both basins also show a significant decrease in precipitation and,
271 subsequently, PDSI. Note that the decreasing rate of PDSI is more profound than that of
272 precipitation in both basins, reflecting the known compound effect of the declining snowpack
273 and drying soil on precipitation deficit in the CRB (BOR 2013). Figure S2 shows that the wet-
274 wet situation has decreased while the dry-dry anomalies have become more frequent in recent
275 years, consistent with SPI3 (Figure 4).

277 3.2 Climate projections

278

279 We analyzed the CMIP5/6 data to assess the future of extreme dry-to-wet transitions using SPI3
280 criteria. SPI was found to be preferable to PDSI for model-based analyses as the latter has not
281 been found to be accurately reflected in CMIP projections (Yang, et al., 2020). (The choice of
282 SPI3 is further discussed in Section 4.) As shown in Figure 5, the models produce extreme dry-
283 to-wet transitions in numbers that are similar to observations, with around one event every ten
284 years. The uncertainty among the individual models/ensembles is quite large, as shown by the
285 error bars for both CMIP5 and CMIP6, and the frequency of miracle events is also consistently
286 higher in CMIP5 than CMIP6. This difference could result from the fact that the CMIP5 used in
287 this analysis is a downscaled and bias-corrected dataset and can capture the UB's dry-to-wet
288 spring transitions better than the coarser-CMIP6 data without bias correction.

289
290 Both models, however, produce a similarly decreasing frequency of extreme dry-to-wet
291 springtime transitions in a warming climate, with CMIP5 projecting a ~50% decrease and
292 CMIP6 projecting a ~15% decrease from the most recent decade (2009-2019) to the final decade
293 of this century. Notably, both models project relatively steady miracle events through the mid-
294 century, roughly in line with past observations, at which point a gradual decline in these extreme
295 transitions is suggested (Figure 5). The main reason for the decreasing miracle numbers may be
296 twofold. First, while the models project increased annual precipitation as the climate warms, the
297 same models project decreasing May-June precipitation; this may result from the enhanced and
298 poleward shift of the North Pacific Subtropical High during April-June, diverging moisture away
299 from the southwestern US (Song et al. 2018a). Second, CMIP5 projects that the frequency of
300 drought (dry-dry) will mildly decrease in the future (Fig. S3a) while the occurrence of
301 anomalously wet seasons will increase (wet-wet; Fig. S3c). However, CMIP6 suggests a
302 significant decreasing trend in drought frequency (Fig. S3b) in conjunction with a ~50% increase
303 in wet-wet (Fig. S3d). In other words, less drought and more precipitation during the winter-to-
304 spring transition may reduce the possibility of reaching the miracle event criteria, especially in
305 CMIP6, but the reduction in May-June precipitation also can diminish the likelihood of miracle
306 events. This may appear to be contradictory to what some researchers have projected will be a
307 low-to-no snow future in the western United States (Siirila-Woodburn et al. 2021), and the
308 discrepancy speaks to the large uncertainty that still plagues the current generation of climate
309 models. While there are still many questions to answer about the capability of these models to
310 accurately capture the forces that would lead to changes in miracle events, it is intriguing that
311 both sets of models project a declining frequency of miracle events after mid-century in the
312 upper CRB as an integrated effect of the changing characteristics of wet and dry seasons. This
313 has led us to perform the next analysis.

314 3.3 Pseudo-global warming scenario

315
316 Using PNNL-WRF PGW simulations, we assessed the effect of climate warming on miracle
317 events. As shown in Figure 6, the control simulation produces a frequency of dry-to-wet
318 transitions that is similar to observations, with one to two events per decade. Applying the
319 RCP8.5 level of warming to the historical period decreases the frequency of miracle events
320 (Figure 6), a result that is consistent with our CMIP projection analysis.

321
322
323 We further diagnosed the precipitation change in the PGW simulation and found that most of the
324 months during the year would receive more precipitation than the control run, except for May

325 (Figure S4), again consistent with the CMIP projections. The PGW simulation lends further
326 support to the implication of fewer miracle events under continuously warming climate
327 conditions. The spring (May/June) drying consistently found in CMIP5/6 and the PGW
328 simulations in the UB has been revealed by Gao et al. (2015) based on regional climate
329 simulations and by Song et al. (2018a) from CMIP5 simulations. They attributed the robust
330 spring drying in the southwestern U.S. to increased moisture divergence by the North Pacific
331 Subtropical High (NPSH), which is enhanced by future warming. Song et al. (2018a) found that
332 under global warming, the NPSH is enhanced due to larger land-sea thermal contrast, but more
333 importantly the seasonal march of the subtropical high (and monsoon onset) is delayed and this
334 directly enhances the NPSH in spring more than summer (Song et al., 2018b). Therefore, the
335 seasonally dependent change in the NPSH accentuates the spring drying that contributes to the
336 reduced likelihood of extreme dry-to-wet transition in the UB.

337
338 The change in SPI3 between the control and PGW simulations (Figure S5) shows that the
339 strength of dry-to-wet transition is reduced in a warmer climate, compounded by weaker SPI3
340 before and during month 0. The strength of an extreme dry-to-wet transition can be assessed by
341 averaging SPI3 from the four months before month 0 (dry index) and subtracting the mean SPI3
342 of the three wet months in and after month 0 (wet index), and the difference between these
343 indices implies a miracle event's strength. With additional warming, miracle strength declines
344 37% over the historical climate control run (Table 1). Thus, a warmer climate may not only
345 reduce the frequency of extreme springtime dry-to-wet transitions but the strength of such
346 miracle events, too.

347 348 3.4 Stakeholder engagement

349 Stakeholders were essential to formulating the management-relevant research questions upon
350 which this study was centered. At the onset of this research, the water managers indicated a
351 desire to better understand miracle events, as this knowledge could inform reservoir operations,
352 flood control planning, irrigation operations, and municipal use, as well as other water users in
353 the basin. The key stakeholder questions relevant to these desires were: What constitutes a
354 miracle event? What are the metrics of such an event? Can miracle events be predicted? Do
355 different modeling approaches yield different answers? Will the likelihood of miracle events
356 change in the future?

357
358
359 In addition to a specific focus on miracle events, the water managers also suggested that a
360 broader understanding of dry-to-wet winter-to-spring transitions would be additionally relevant
361 for decision-making. Their questions related to a more general understanding of drought-busting
362 spring weather included: To what extent are dry winters and wet springs correlated? What do
363 observations and climate models suggest about how often dry winters are followed by wet
364 springs? How might these conditions change in the future? What are the dynamical drivers at
365 play?

366
367 In the initial engagements, water managers identified several miracle events that were of
368 relevance to their regions, with a particular interest in the miracle events that occurred in 1992
369 and 2015. They also suggested that the UB and LB should be considered separately and informed
370 the decision to focus on the UB, given its dominant role in the Colorado River water supply. In

371 our ongoing discussions, the managers also pointed out that spatial scales play a significant role
372 in defining miracle events, as conditions in one region may be very different from events in
373 another. For example, the managers suggested that while the May 2015 event was significant for
374 most localities within the CRB, other regions, such as the area managed by Denver Water, were
375 not as greatly affected.

376

377 Stakeholder input was also key to our decision to examine miracle events using several indices,
378 as their work necessarily incorporates multiple perspectives. The managers expressed a general
379 preference for the indices that showed the “Miracle May” of 2015, as this basin-wide event was
380 immediately and intuitively familiar. This event was reflected in both the PDSI and SPI3
381 analyses and was one of the factors that helped inform our decision to examine the elements of
382 SPI3 in the CMIP models. Managers also indicated a desire to better understand how and why
383 results differ between CMIP versions, as they have recently begun updating their plans to reflect
384 the relatively recent release of CMIP6 data.

385

386 The managers expressed a desire to gain a broader understanding of the patterns of wet-wet, wet-
387 dry, dry-dry, and dry-wet winter-to-spring transitions, and the pie-charts showing these results
388 (Figure 4) were consistently flagged as being most useful for drought planning, reservoir
389 releases, inter- and intra-state water storage management, and demand management
390 programming. The managers noted the utility of knowledge such as the fact that wet winters are
391 generally more likely to be followed by wet springs, and dry winters are generally more likely to
392 be followed by dry springs. These stakeholders further expressed that it was additionally helpful
393 to understand how these patterns might hold up in the future, a question that was answered in
394 part by all three sets of climate projections indicating a decreasing trend of miracle events as the
395 climate warms. As one manager noted, “if we know the miracle spring signal is decreasing, we
396 know that a miracle spring is less likely to save us.”

397

398 **4. Discussion and Summary**

399

400 It may not be possible to perfectly define the conditions that constitute a “miracle.” This is
401 ultimately a phenomenon to which water managers and climate researchers alike might apply
402 Justice Potter Stewart’s famous standard of “I know it when I see it.” But as “miracle” has
403 become an increasingly common colloquial description of rapidly drought-ending precipitation
404 events, we believe there is great value in a good-faith attempt to define the kinds of conditions
405 that are embraced within that shorthand description.

406

407 The principal challenge in this endeavor is selecting the right metrics to define a miracle event.
408 As previously noted, as even the definition of “drought” is subject to broad interpretation, it
409 makes sense that the variables constituting a “drought-busting” event might be similarly rife
410 grounds for debate. Indeed, stakeholders suggested many different ways that one might arrive at
411 a “miracle index,” including the application of data related to snow water equivalent
412 calculations, streamflows, and reservoir levels. As these variables were difficult to derive in
413 climate projections, and as stakeholder consensus that a simple-to-assess and relatively easy-to-
414 model equation would offer the broadest potential actionability, SPI3 emerged as the key index
415 upon which to base our analysis, with PDSI and SPI6 playing supporting roles to offer a holistic

416 picture of past and future miracle events. Future work should consider SPEI and other
417 hydrological drought indices.

418
419 Thus defined, there is tremendous potential in further exploration of miracle events, most notably
420 in the realms of spatiality and consequentiality. Just like miracles of the unquantifiable sort, after
421 all, a miracle event may not have equal benefits across an entire basin, or even sub-regions of
422 that basin. An extreme dry-to-wet transition during spring may mitigate a given winter drought
423 *in summa*, but the impact of that transition may result in vastly differing effects on agriculture
424 and ecology from place to place, most notably when it comes to fires and floods. This is perhaps
425 best reflected in the “Miracle May” event of 2015, which saved the CRB from a severe water
426 shortage but also resulted in massive flooding in Texas (Wang et al. 2015).

427
428 Given that each miracle event affects different watersheds in different ways, every such event
429 will be felt and examined in different ways. This may complicate the path to a holistic analysis.
430 The implication of the presented research, however, offers a starting point for additional
431 examinations of this phenomenon to understand its climatological characteristics and provides
432 valuable clues about the future of extreme dry-to-wet springtime transitions. It may be true, as is
433 often said, that “miracles never cease,” but our analysis suggests that, in a warming world,
434 miracle events in the CRB may soon begin to slow down and weaken.

435 436 **Acknowledgement:**

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438 Biological and Environmental Research program under Award Number DE-SC0016605.

439 440 **Open Research**

441 PDSI and SPI data was used in this study can be freely accessed from
442 <https://www1.ncdc.noaa.gov/pub/data/cirs/climdiv/>, CMIP5 downscaled and bias corrected data
443 are freely available at [https://gdo-
444 dcp.ucllnl.org/downscaled_cmip_projections/dcpInterface.html#Projections:%20Complete%20A
445 rchives](https://gdo-dcp.ucllnl.org/downscaled_cmip_projections/dcpInterface.html#Projections:%20Complete%20Archives). CMIP6 data are generated at
446 https://climexp.knmi.nl/selectfield_cmip6.cgi?id=someone@somewhere. The regional climate
447 model (PNNL-WRF) data are available upon request by email to Lai-Yung Ruby Leung
448 (ruby.leung@pnnl.gov).

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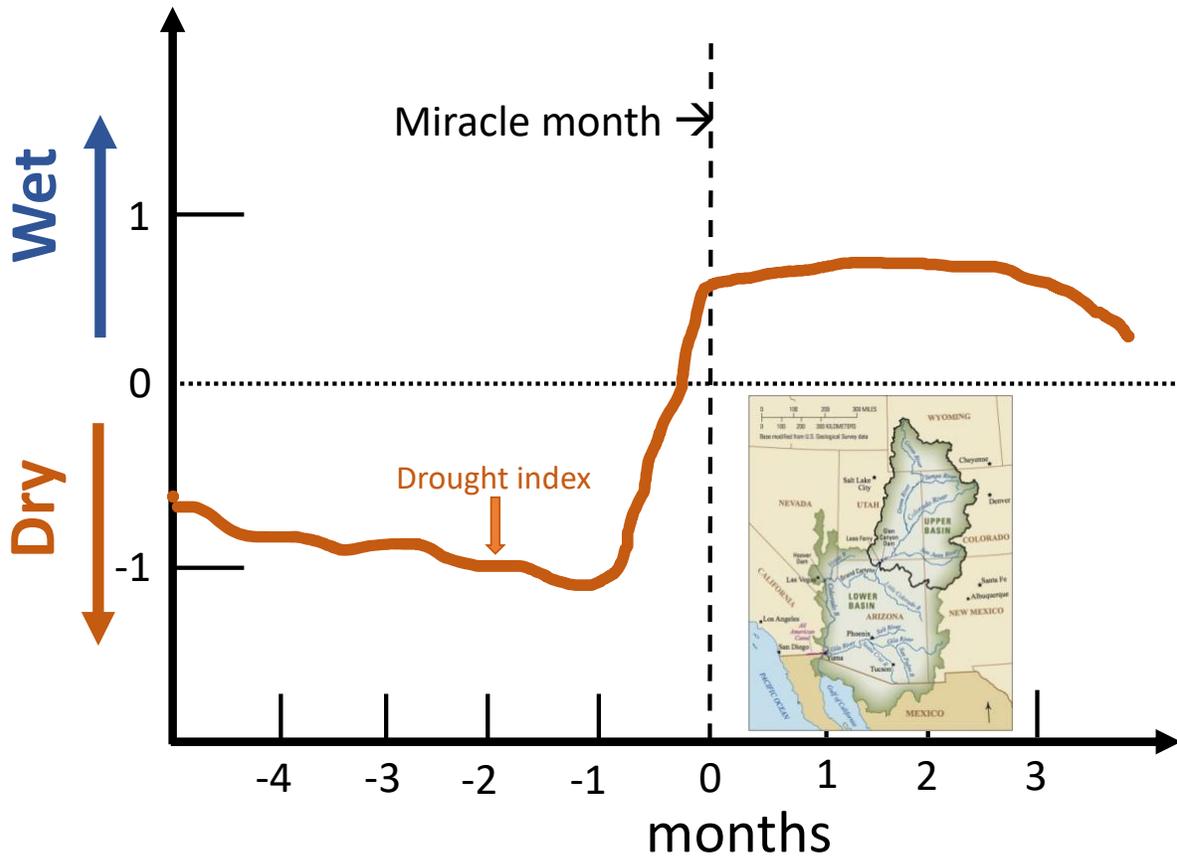
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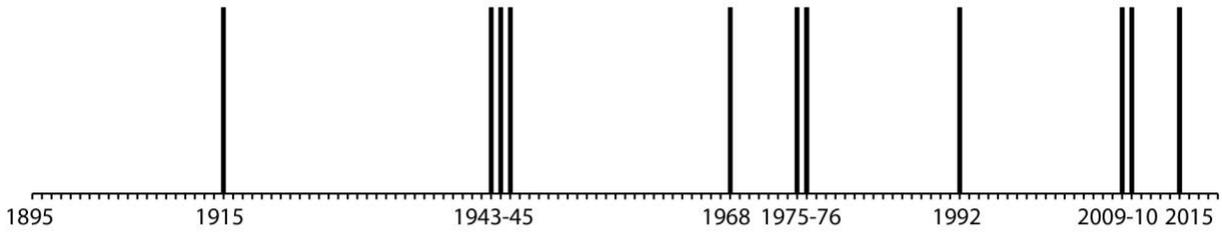
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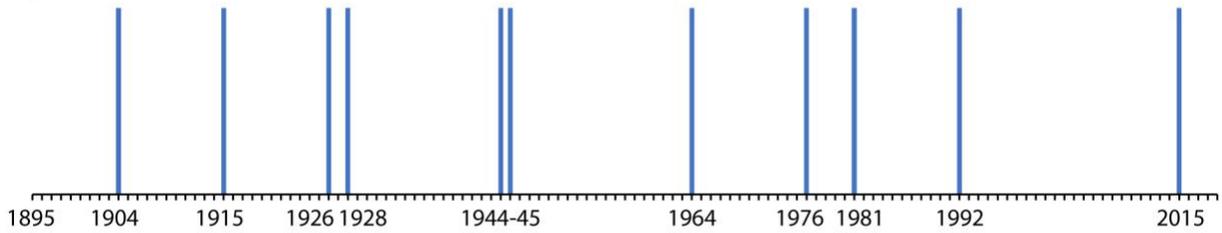


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535 Figure 1: Schematic of miracle spring events with Colorado River Basins map (inserted). 'Month
536 0' represents the miracle month while negative months are dry and positive months are wet
537 months.
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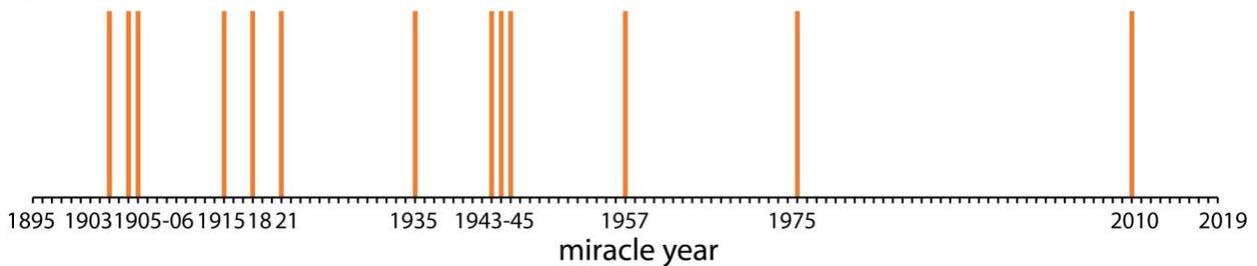
a) PDSI



b) SPI03

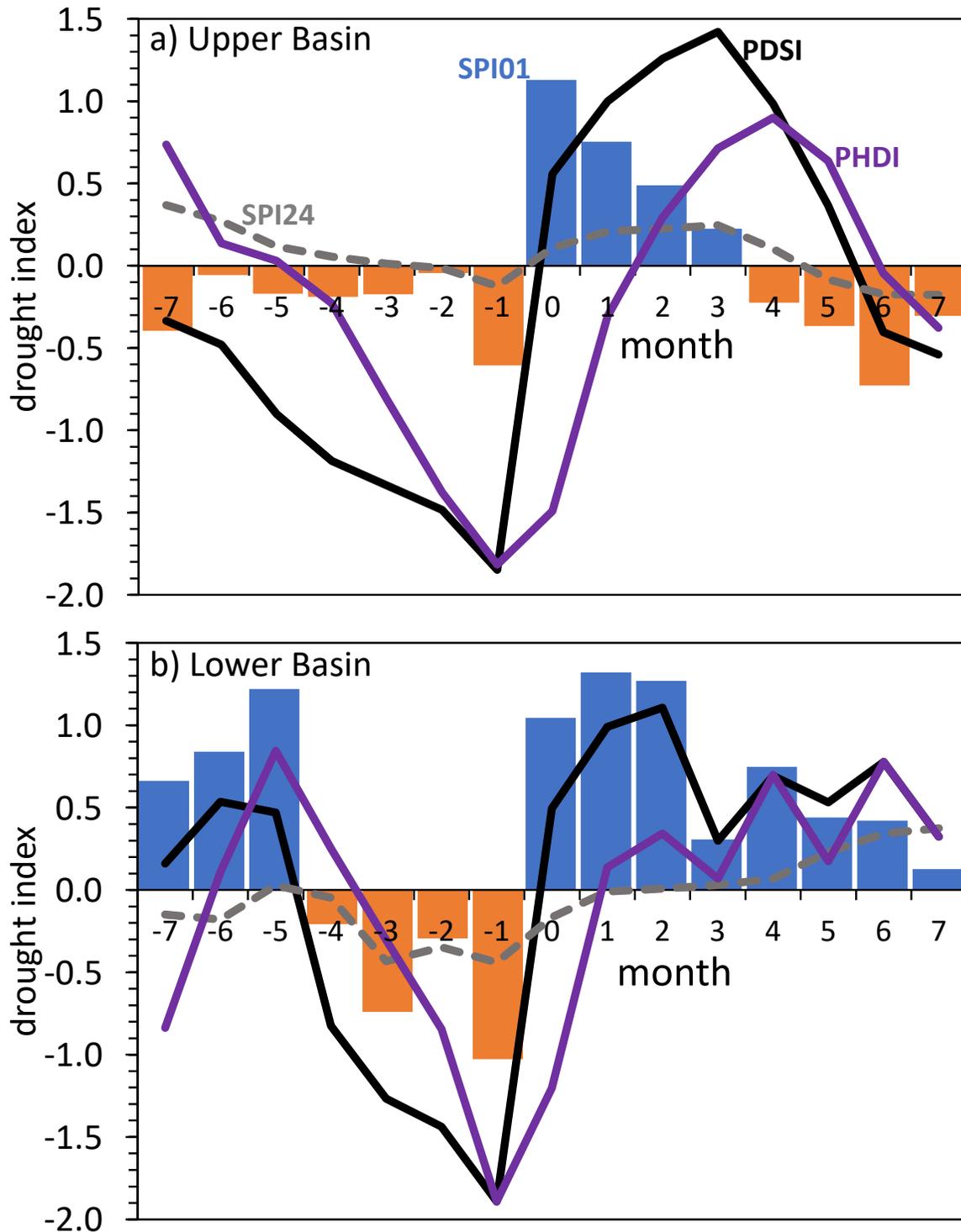


c) SPI06



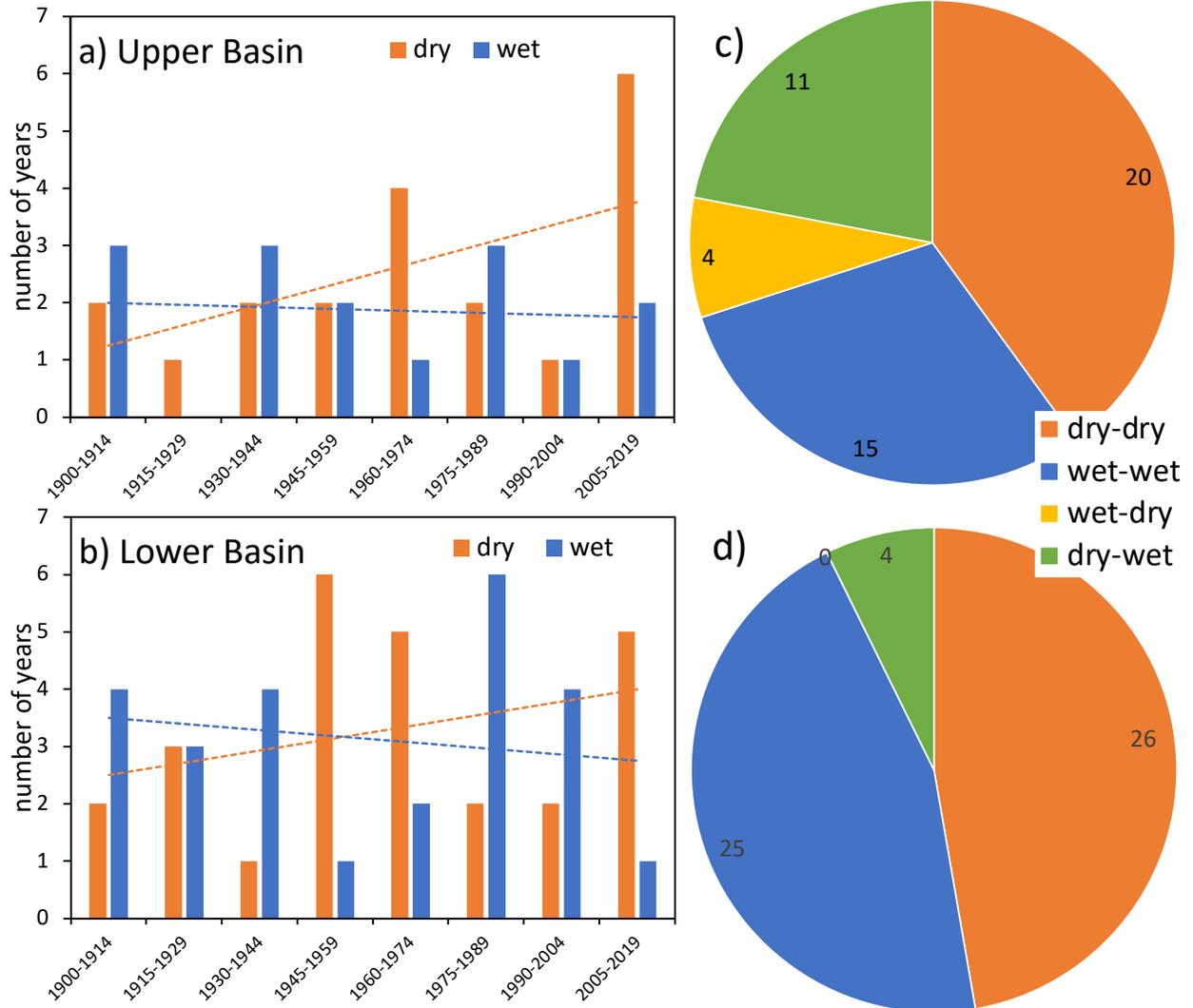
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Figure 2: Historical miracle events calculated from different drought indices considering four consecutive negative drought index followed by three consecutive positive drought index. Upper panel (a) shows the miracle years based on PDSI (upper panel), middle panel (b) shows the miracle years based on SPI3, and lower panel (c) based on SPI6.



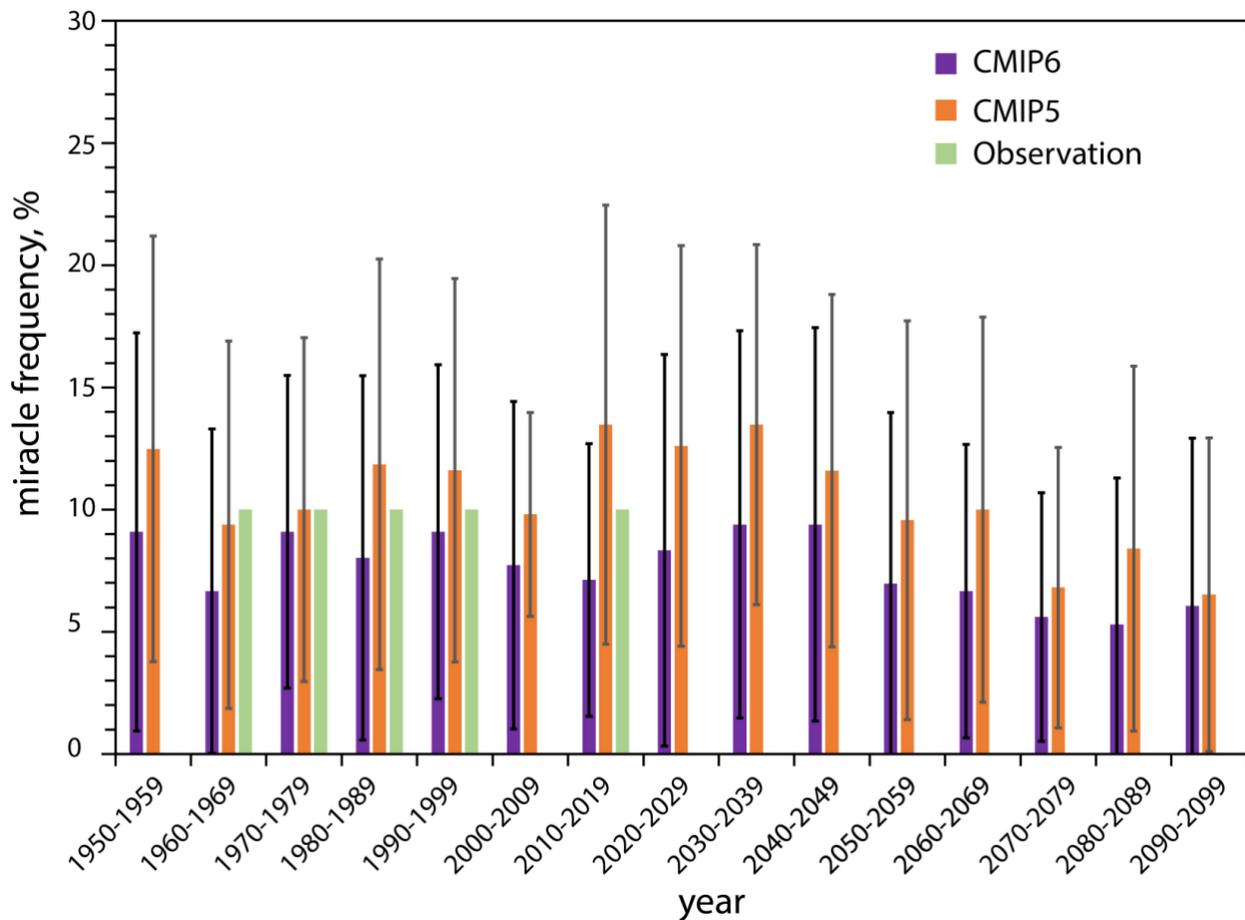
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556 Figure 3: Average drought indices during Miracle years that is calculated based on PDSI for
 557 Upper basin a) and lower basin b). Bar plot shows the SPI1, dashed line shows the SPI24, black
 558 line shows the PDSI and purple line shows the PHDI. Miracle month is centered at 0 when
 559 drought index (PDSI) changed from negative to positive value.



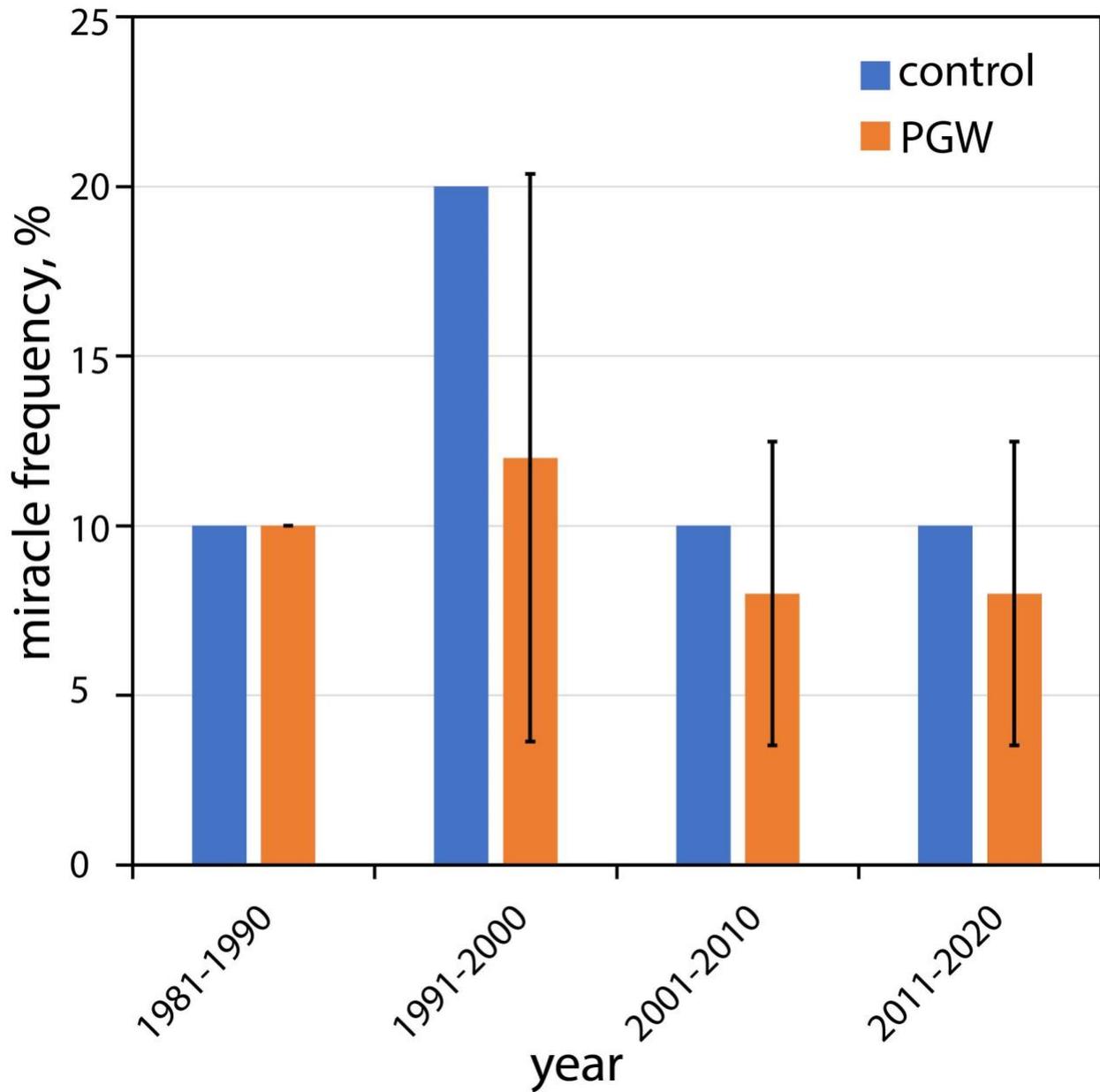
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Figure 4: Average dry and wet frequency based on winter and spring SPI3 value considering the months from December to May for a) upper and b) lower basins. The right panel c) and d) show the frequency pie-chart of four different transitions from winter (Dec-Feb) to spring (Mar-May) for upper and lower basins, respectively.



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Figure 5: Miracle spring precipitation frequency (per 10 years) based on observational (green bars), CMIP5 (orange bars) and CMIP6 (purple bars). Historical period is considered from 1950 while future projections are considered from 2006 for CMIP5 and from 2015 for CMIP6. Error bars in CMIPs data are based on ± 0.8 standard deviation. SPI3 is used to calculate the miracle spring precipitation. There were no miracle events during 2000-2009 in the observation.



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Figure 6: Miracle spring precipitation frequency from regional climate model simulation for the historical control run (blue bars) and pseudo global warming (PGW run, orange bar) that utilized climate change signals from 5 global climate models. Error bars in PGW frequency based on ± 0.8 standard deviation.

591 Table 1: Dry- and wet-indices from average miracle years based on regional climate model
592 simulations (PNNL WRF) from control and PGW runs from (1980-2020). Dry index is the total
593 SPI3 value from four dry months before miracle and wet index is the total SPI3 from following
594 three wet months. Miracle strength is the difference between dry- and wet-indices and change in
595 miracle strength is calculated as control miracle strength/PGW miracle strength *100%.
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	control	PGW
dry index	-3.20	-3.36
wet index	2.28	2.10
miracle strength	-0.92	-1.25
change in miracle strength, %	37	

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