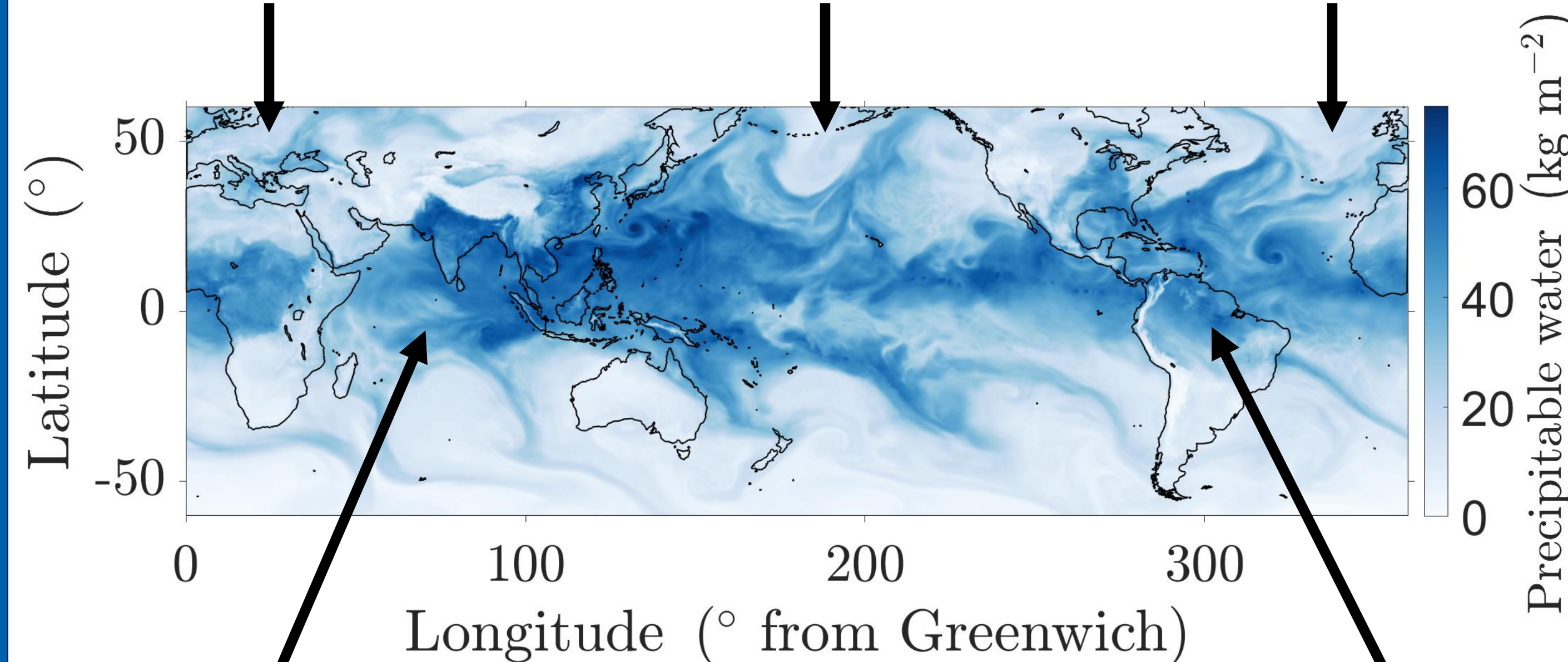


Motivation

- To first order, the size of extra-tropical storms is set by the Rossby radius of deformation, proportional to the atmospheric stratification divided by the Coriolis parameter.



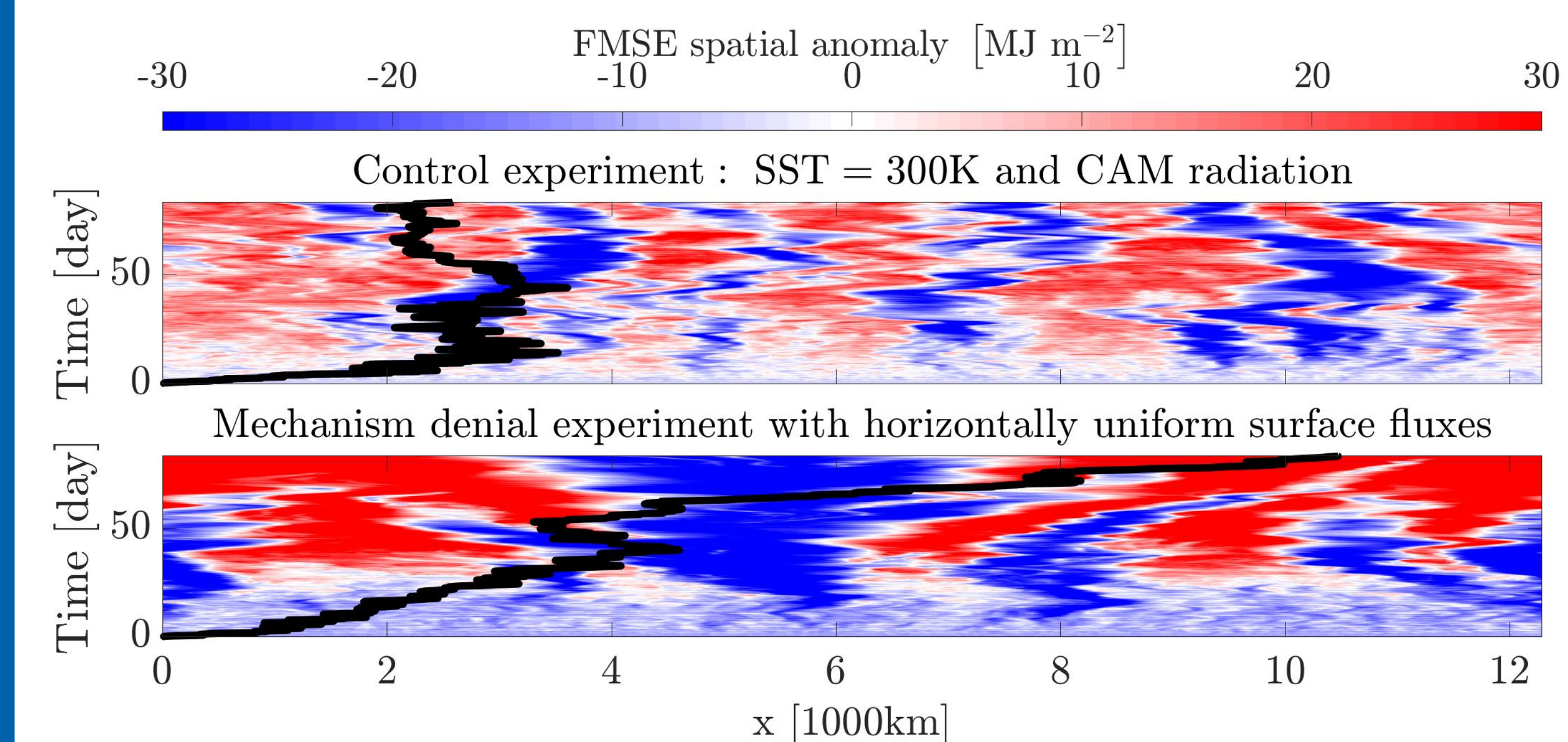
- However, what sets the scale L at which convective activity organizes near the Equator, where the Coriolis parameter is small, remains an open question.
- Previous theories [1-4] predict the order of magnitude of L and some of its dependencies (surface temperature, boundary layer properties) by assuming a dominant mechanism.
- Here, we take an alternative approach by (1) formulating a budget for L and diagnosing contributions to its evolution from different processes in (2-3) 3D cloud-permitting sim. with interactive rad., surface fluxes & large-scale dyn. and (4) reanalysis data [5], satellite observations [6] & global cloud-permitting sim. [7-8].

How do **radiation**, **surface enthalpy fluxes** and **advection** contribute to the emergence and evolution of a dominant **size** for **convective aggregation**?

2. Long-channel simulations

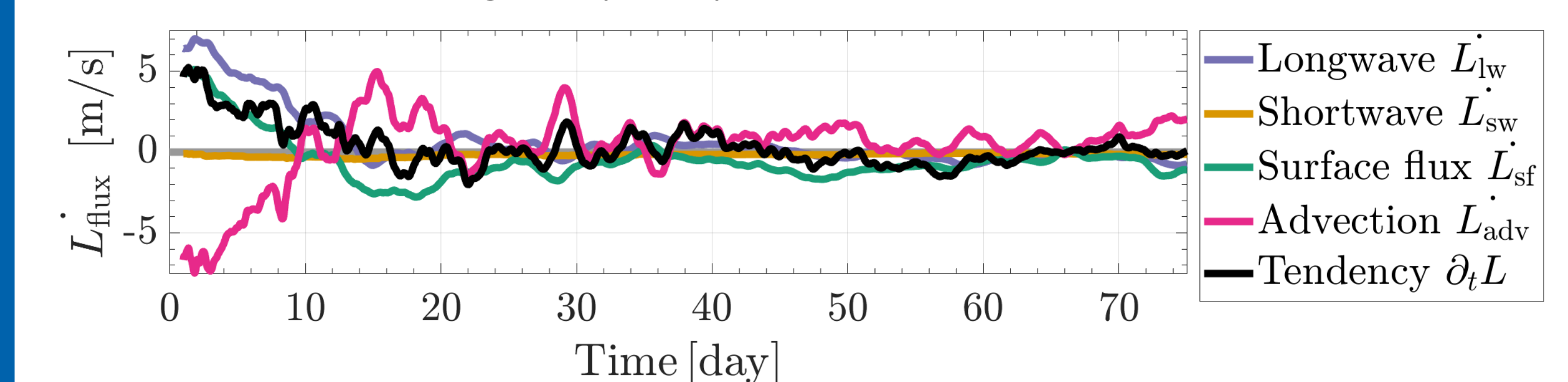
Measuring the convective-aggregation length scale (L)

Method: We simulate radiative-convective equilibrium in a long-channel domain [2] using the cloud-permitting model SAM [7] and measure the scale of West-East anomalies in time:



Understanding the evolution of L in time

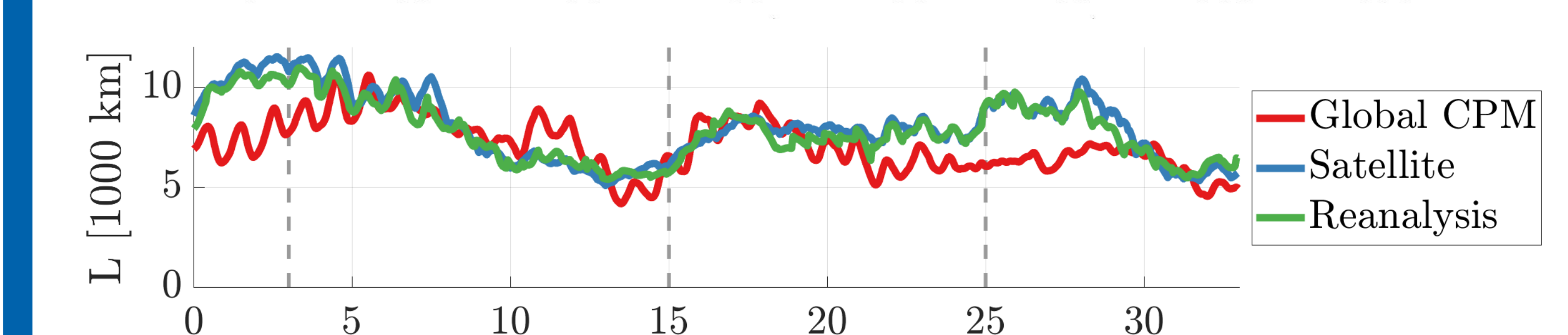
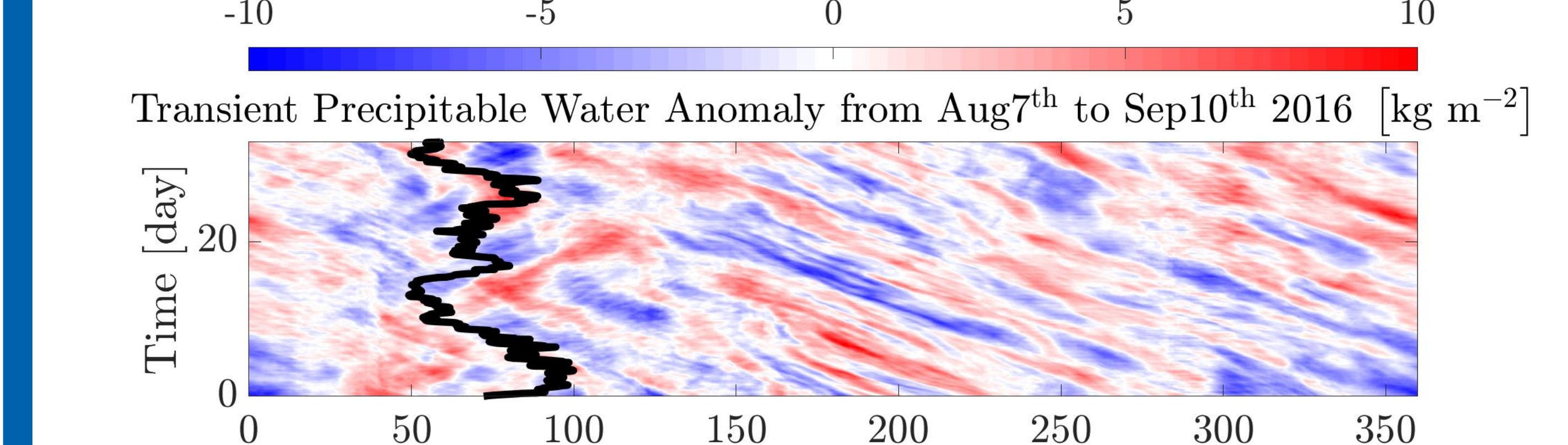
Method: We use the L budget to quantify the contribution of each flux to L 's evolution:



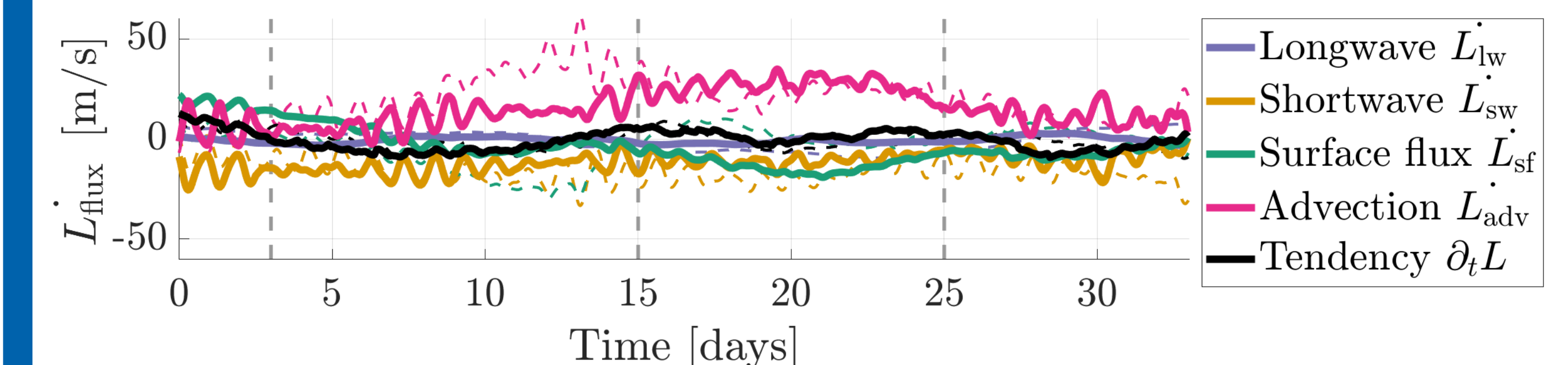
4. Observed variability in Tropics

Measuring the scale of transient precipitable water anomalies (L)

Method: We average the observed [5] transients of precipitable water from South to North in the Tropics (e.g. 15S-15N) and focus on the scale of West-East anomalies in August 2016:



We compare the evolution and budget of L across observations [5-6] and global CPM [7-8]:



1. Theory

Budget for column frozen moist static energy (H)

Definition: H , a proxy for convective activity, is defined as the sum of column internal energy, potential energy and latent heat.

$$H \text{ [J m}^{-2}] \stackrel{\text{def}}{=} \int_0^{p_s} \frac{dp}{g} (c_p T + gz + L_{\text{vap}} q_{\text{vap}} - L_{\text{fus}} q_{\text{ice}})$$

Budget: H is conserved under vertical convective mixing and altered by the net energy flux through the col. boundaries: Longwave, shortwave, surf. fluxes & horizontal advection.

$$\frac{\partial H}{\partial t} \text{ [W m}^{-2}] = \dot{H}_{\text{lw}} + \dot{H}_{\text{sw}} + \dot{H}_{\text{sf}} + \dot{H}_{\text{adv}} = \sum_{\text{flux}=\text{lw,sw,sf,adv}} \dot{H}_{\text{flux}}$$

Budget for spatial spectrum of moist static energy (φ)

Definition: φ , a measure of the scale-by-scale variance of H , is defined as the modulus of the spatial Fourier transform \hat{H} of H .

$$\varphi \text{ [J}^2 \text{ m}^{-2}] \stackrel{\text{def}}{=} \hat{H}^* \hat{H}$$

Budget: φ is altered by the scale-by-scale coherence between H and energy fluxes: at each scale, variance is reinforced by positive coherences & destroyed by neg. coherences.

$$\frac{\partial \varphi}{\partial t} \text{ [J}^2 \text{ m}^{-2} \text{ s}^{-1}] = \sum_{\text{flux}=\text{lw,sw,sf,adv}} 2\text{Re} \left(\hat{H}^* \hat{H}_{\text{flux}} \right) = \sum_{\text{flux}=\text{lw,sw,sf,adv}} \dot{\varphi}_{\text{flux}}$$

Budget for convective-aggregation length scale (L)

Definition: L , the distance between pos. & neg. anomalies of H , is formally defined as the spectral mean of the wavelength, weighted by the spatial power spectrum φ of H .

$$L \text{ [m]} \stackrel{\text{def}}{=} \frac{1}{\langle \varphi \rangle} \left\langle \frac{2\pi\sqrt{n}}{\|k\|} \varphi \right\rangle = \frac{\langle \lambda \varphi \rangle}{\langle \varphi \rangle}$$

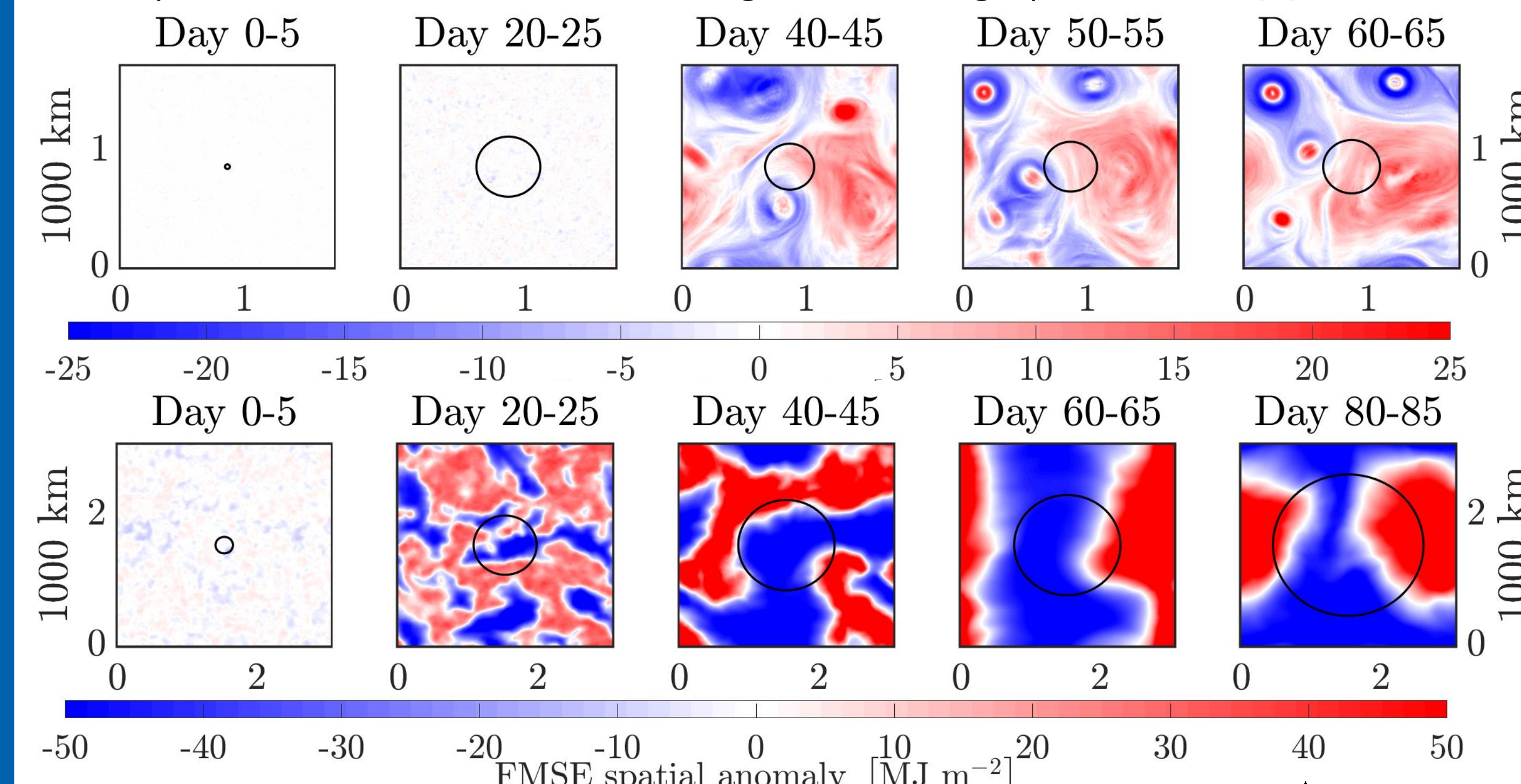
Budget: L is altered by the expansion tendency of each energy flux, given by the product of the aggregation rate and a length scale factor which vanishes if fluxes operate at the scale L .

$$\frac{\partial L}{\partial t} \text{ [m s}^{-1}] = \sum_{\text{flux}=\text{lw,sw,sf,adv}} \frac{\langle \dot{\varphi}_{\text{flux}} \rangle}{\langle \varphi \rangle} \times \left(\frac{\langle \lambda \dot{\varphi}_{\text{flux}} \rangle}{\langle \dot{\varphi}_{\text{flux}} \rangle} - L \right) = \sum_{\text{flux}=\text{lw,sw,sf,adv}} \dot{L}_{\text{flux}}$$

3. Two-dimensional clusters

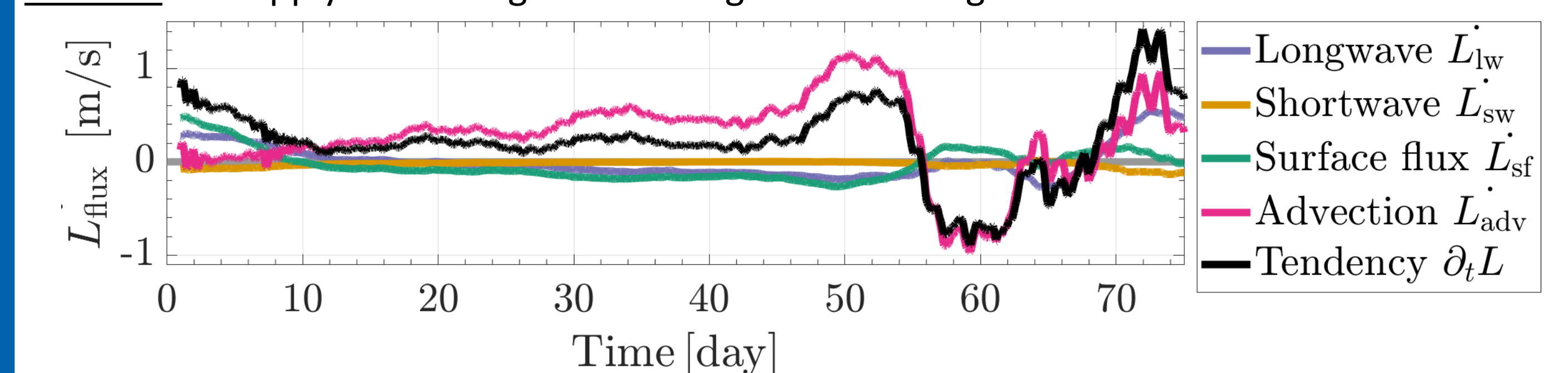
Measuring the size of convective clusters (L)

Method: We simulate radiative-convective equilibrium in a rotating square domain of Coriolis parameter $3 \times 10^{-4} \text{ s}^{-1}$ and in a large non-rotating square domain [7]:



Understanding the evolution of L in time

Method: We apply the L budget to the large non-rotating simulation



Conclusions

Summary

- We have developed a budget that relates the evolution of the convective-aggregation length scale to the vertically-integrated energy fluxes (radiative, surface, and advective).
- Longwave cooling** expands the convective-aggregation scale until convection aggregates.
- Shortwave heating** shrinks the convective-aggregation length scale.
- Surface fluxes** expand L when it is small & become main shrinking term when L is large.
- Advection of energy** expands L because it homogenizes small-scale convective anomalies.
- We can generalize the definition and budget of L to two dimensions ($n=2$) and apply it to non-rotating convective clusters as well as tropical cyclones.
- In the real Tropics, the budget can be used to understand the West-East scale of the transient zonal anomalies of precipitable water.
- We can measure the effect of radiative and convective biases on the spatial organization of water vapor in the Tropics by comparing the L budget across models and observations.

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