

Abstract

In the last decade, much work has been done to better understand methane (CH_4) emissions from the oil and gas (O&G) industry in the United States. Ethane (C_2H_6), a gas that is co-emitted with thermogenic sources of CH_4 , is emitted in the US almost entirely by the O&G sector. In this study, we perform an inverse analysis on 300 hours of atmospheric boundary layer C_2H_6 measurements to estimate C_2H_6 emissions from the US O&G sector. Measurements were collected from 2017-2019 as part of the Atmospheric Carbon and Transport (ACT) America aircraft campaign and encompass much of the central and eastern United States. We find that for the fall, winter, and spring campaigns, C_2H_6 data consistently exceeds values that would be expected based on EPA O&G leak rate estimates. C_2H_6 observations from the summer 2019 dataset show significantly lower C_2H_6 enhancements in the southcentral region that cannot be reconciled with data from the other three seasons, either due to complex meteorological conditions or a temporal shift in the emissions. Converting the fall, winter, and spring season posterior C_2H_6 emissions estimate to an inventory of O&G CH_4 emissions, we estimate that O&G CH_4 emissions are larger than EPA inventory values by more than 50%. Uncertainties in the gas composition data limit the effectiveness of using C_2H_6 as a proxy for O&G CH_4 emissions. These limits could be resolved retroactively by increasing the availability of industry-collected gas composition data.

Plain Language Summary

Methane is a potent greenhouse gas responsible for a quarter of the warming the climate has experienced thus far. The oil and gas sector is a significant source of methane through leaks in its infrastructure. Recent studies of individual basins have found emissions from oil and gas in the US to be greater than inventory estimates, but difficulties arise with source attribution in broader scale studies due to the numerous potential sources of methane. This study quantifies methane emissions from oil and gas by looking at atmospheric ethane, a gas whose emissions stem almost entirely from oil and gas in the US. Hundreds of hours of ethane observations were collected via aircraft over the course of 4 seasons between 2017-2019. These observations are compared with model-projected ethane values based on our current knowledge of ethane emissions, and those emissions are adjusted to best match the observed data. We find ethane emissions are grossly underestimated in the US. Because ethane is co-emitted with oil and gas methane sources, this underestimation of ethane reflects similar underestimations in oil and gas methane emissions. We conclude that US inventories are underestimating methane emissions from oil and gas by more than 50%.

1 Introduction

Methane (CH_4) is an important greenhouse gas with 28-35 times the warming potential of carbon dioxide over a 100 year period (Myhre et al., 2013). Global CH_4 concentrations in the atmosphere have nearly tripled since pre-industrial times, mainly driven by anthropogenic activity and are responsible for a fourth of the increased radiative forcing on the planet (Myhre et al., 2013). Although CH_4 concentrations stabilized for a brief period in the early 2000s, global concentrations began increasing again by 2007, with an increasing growth rate continuing through present date (Nisbet et al., 2019).

Efforts to understand causes for increasing global CH_4 trends are hampered by difficulties related to source attribution (Saunio et al., 2020). CH_4 has numerous anthropogenic sources, including animal agriculture, fossil fuel extraction, and waste management. In addition to anthropogenic emitters, there are natural sources of CH_4 emissions that play a large role in the global CH_4 budget. Of particular importance to the global CH_4 budget are emissions from anaerobic respiration in wetlands, which create unique challenges to CH_4 source attribution. Unlike most anthropogenic sources, whose emis-

71 sions can be spatially mapped out using inventory techniques and are relatively consis-
72 tent in magnitude on an annual timeframe, wetland emissions are more difficult to de-
73 scribe, with an uncertain spatial pattern and large seasonal variability in emissions based
74 on soil temperature and moisture (Yvon-Durocher et al., 2014). Many prior studies have
75 examined seasonal emissions and magnitudes of local wetland sources using flux tower
76 measurements (Grant & Roulet, 2002; Matthes et al., 2014), but extrapolating these re-
77 sults to a continental or global scale is challenging.

78 Difficulties with source attribution extend into the United States, where both an-
79 thropogenic and natural sources play a large role in the country’s CH_4 budget. In the
80 US, greenhouse gas emissions from anthropogenic sources are quantified through a bottom-
81 up inventory created by the United States Environmental Protection Agency (EPA). The
82 largest sources of anthropogenic CH_4 emissions projected by this inventory are from leaks
83 in oil and gas (O&G) infrastructure, enteric fermentation and manure management re-
84 lated to livestock, and anaerobic respiration occurring in landfills (US Environmental Pro-
85 tection Agency, 2020). These sources were responsible for 83% of US anthropogenic CH_4
86 emissions in the 2018 inventory estimate. Natural sources in the US are dominated by
87 wetlands and are not tracked by the EPA’s inventory. WetCHARTs, a global wetland
88 emissions ensemble, is generally used as the prior for CH_4 emissions from wetlands in
89 the US (Bloom et al., 2017; Sheng et al., 2017; Maasakkers et al., 2016). WetCHARTs
90 ensemble members estimates the country’s wetland emissions to vary from values that
91 are insignificant to totals rivaling those from anthropogenic emissions, with seasonal de-
92 pendence and spatial variability between ensemble members, resulting in large uncer-
93 tainties in the overall US CH_4 budget.

94 Contrary to recent global increases in the atmospheric growth rate of CH_4 , the EPA’s
95 inventory estimates of anthropogenic CH_4 emission in the US show a 10% decrease in
96 the last decade, from 28 Tg in 2008 to 25 Tg in 2018. This decrease is driven primar-
97 ily by projected decreases in emissions from the energy sector, despite a 50% increase
98 in gas production and a >200% increase in oil production during the 10 year period (US
99 Energy Information Administration, 2020b). Over the last several years, various atmo-
100 spheric studies monitoring emissions from O&G from individual wellpads (Rella et al.,
101 2015; Robertson et al., 2017; Caulton et al., 2019), basins (Karion et al., 2015; Barkley
102 et al., 2017; Peischl et al., 2018), and entire regions (Barkley et al., 2019b) have consis-
103 tently found emission rates larger than the EPA inventory, raising concerns of a broad
104 underestimation of leaks from the O&G sector (Alvarez et al., 2018). However, large-
105 scale CH_4 inversion studies involving the US have not been as conclusive, with differing
106 opinions, as to the accuracy of inventory emissions from O&G, animal agriculture, and
107 wetlands (Sheng et al., 2018; Yu et al., 2020; Maasakkers et al., 2019). The enormous
108 spread of uncertainty regarding the magnitude, spatial distribution, and seasonality of
109 CH_4 emissions from wetlands, as well as the numerous other potential sources of CH_4 ,
110 poses a challenge to large scale studies attempting source attribution of detected CH_4
111 signals.

112 One common method to disaggregate certain sources is to measure both CH_4 and
113 ethane (C_2H_6) concentrations (McKain et al., 2015; Barkley et al., 2019a). Sources of
114 C_2H_6 coincide with thermogenic CH_4 emitters (O&G extraction, biomass burning), but
115 not biogenic sources (wetlands, animal agriculture, landfills). Thus, a correlation (or lack
116 of) between C_2H_6 and CH_4 can provide information on the sector responsible for observed
117 CH_4 enhancements. While global C_2H_6 emissions are split between the energy sector and
118 biomass burning (Xiao et al., 2008), the National Emissions Inventory (NEI) 2011 C_2H_6
119 inventory attributes 95% of C_2H_6 emissions in the US to the energy sector, making it
120 a simple tracer for identifying O&G emissions in the region ((US Environmental Pro-
121 tection Agency, 2014)). This technique has been used in various smaller-scale aircraft
122 mass balance studies across individual O&G basins to verify that the measured enhance-
123 ments are associated with O&G activity (Peischl et al., 2018; Roscioli et al., 2015). Sim-

124 ilarly, C_2H_6 concentrations from a network of flask measurements have been used to infer
125 trends in emissions from O&G emissions on broader scales (Xiao et al., 2008; Tzompa-
126 Sosa et al., 2017).

127 Advancements in technology have resulted in the increased availability of instru-
128 mentation capable of measuring C_2H_6 concentrations precisely at high temporal reso-
129 lution (Weibring et al., 2020; Kostinek et al., 2019; Yacovitch et al., 2014), expanding
130 our capabilities of solving for C_2H_6 emissions. The Atmospheric Carbon and Transport
131 - America (ACT-America) mission was a 5-season aircraft campaign across the central
132 and eastern US designed to examine various trace gases and their transport in the at-
133 mosphere. During four of the seasonal campaigns, continuous C_2H_6 measurements were
134 collected, producing more than 500 hours of C_2H_6 data capable of capturing various plume
135 structures related to O&G activity on a regional scale. In this study, we examine the char-
136 acteristics of this unprecedented dataset and use it to quantify C_2H_6 emissions compared
137 to bottom-up inventory estimates. These C_2H_6 emissions are then combined with US gas
138 composition data to infer CH_4 leak rates from major US O&G basins.

139 2 Methods

140 2.1 ACT-America Aircraft Campaign and Observational Dataset

141 The ACT-America campaign was a NASA Earth Venture suborbital aircraft mis-
142 sion with flight activities spanning summer 2016 to summer 2019. During this time, 5
143 individual campaigns covering all 4 seasons (summer twice) were conducted using two
144 aircraft collecting meteorological, greenhouse gas, and trace gas data within fair weather
145 and frontal weather patterns. For all 5 campaigns, continuous CH_4 data were collected
146 on both aircraft using a commercial PICARRO G2401-m instrument adapted with a cus-
147 tom inlet system for drying and conditioning the sample air (DiGangi et al., 2018). Dur-
148 ing the Winter 2017, Fall 2017, Spring 2018, and Summer 2019 campaigns, continuous
149 C_2H_6 data were collected on the B200 aircraft using the CAMS-2 (Compact Airborne
150 Multi-Species Spectrometer) instrument (Weibring et al., 2020). The CAMS-2 C_2H_6 mea-
151 surements when averaged over time and linearly regressed versus NOAA portable flask
152 packages collected during the flights yielded slopes in the 0.990 to 1.031 range across sea-
153 sonal campaigns, with r^2 values between the two measurements of 0.996 (Baier et al.,
154 2020; Weibring et al., 2020), providing high confidence in the accuracy of its measure-
155 ments. Additionally, during the Fall 2017 and Summer 2019 campaigns, the C130 air-
156 craft was equipped with a quantum and interband cascade laser spectrometer (QCLS)
157 capable of continuous in situ C_2H_6 measurements (Kostinek et al., 2019). The QCLS in-
158 strument performed in-flight two-point calibrations every three to ten minutes, ensur-
159 ing accurate measurements throughout the flights. Altogether, the entire C_2H_6 dataset
160 contains more than 500 hours of continuous airborne C_2H_6 observations, of which 300
161 hours were within the atmospheric boundary layer (ABL), making it an ideal dataset
162 to study atmospheric C_2H_6 plumes and structures throughout the central and eastern
163 US.

164 From the observational dataset, large C_2H_6 plumes were consistently observed down-
165 wind of Texas/Oklahoma/Louisiana in the southcentral US and along the western Ap-
166 palachians in the northeast, corresponding with O&G activity in each of these regions
167 (Figure 1). C_2H_6 plumes associated with southcentral O&G sources could be observed
168 in the midwestern flights as far north as Minnesota when consistent southerly winds were
169 present. C_2H_6 mixing ratios within the plumes were largest during the fall and winter
170 campaign, likely related to the low boundary layer mixing depth during those seasons.
171 During the Summer 2019 campaign, large C_2H_6 plumes were detected in the western Ap-
172 palachian similar to other seasons, but signals in the southcentral were substantially smaller
173 in a way that boundary layer depth alone cannot explain. Additionally, boundary layer

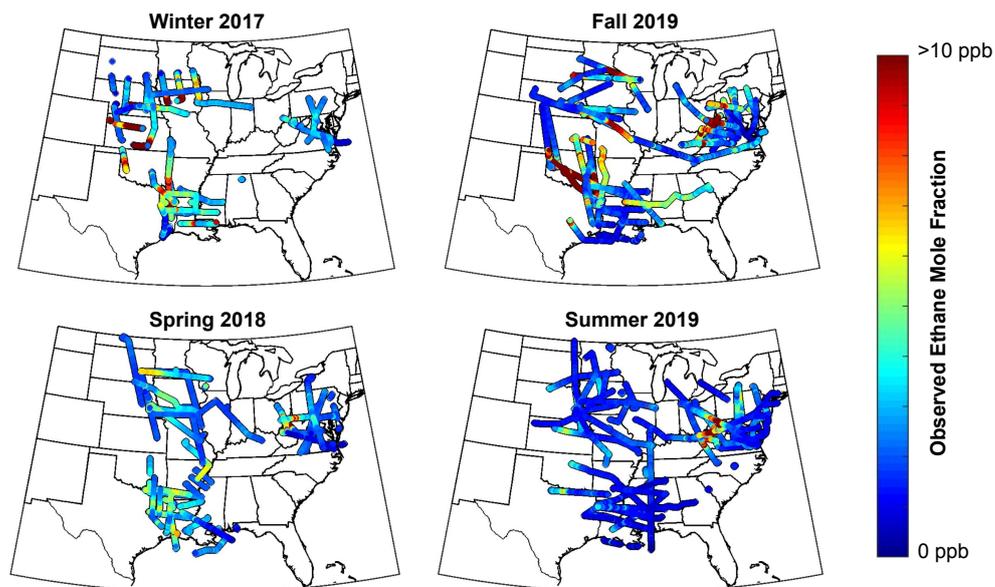


Figure 1. Continuous boundary layer C_2H_6 measurements collected by the B200 and C130 aircraft during each seasonal campaign. Only data within the atmospheric boundary layer (≤ 1000 m AGL) are shown.

174 C_2H_6 concentrations in the summer were often similar in value to their free tropospheric
 175 counterparts, a trait not observed during other seasons (Figure S2).

176 Influence functions for observations were created using the Lagrangian particle dis-
 177 persion model FLEXPART-WRF (Brioude et al., 2013) to provide information on the
 178 area captured by the ACT-America C_2H_6 dataset. The model was run at 27 km reso-
 179 lution over the North American domain and was meteorologically driven using WRF-
 180 Chem simulations developed as part of the ACT-America campaign (Feng et al., 2019).
 181 Every 30 seconds of flight time, 5,000 particles were released from the aircraft location
 182 and traced back 10 days in time. Further details on the development of the influence func-
 183 tions can be found in Cui et al. (2015).

184 Averaging the influence functions provides information on which O&G basins have
 185 substantial coverage by the ACT flight campaign (Figure 2). Much of the midwestern
 186 and southcentral US is captured across all four seasons, encompassing many of the ma-
 187 jor O&G basins including Haynesville, Fayetteville, Anadarko, Eagle Ford, Woodford,
 188 and the Barnett. A sharp decrease in the influence map can be observed west of central
 189 Texas. For this reason, we do not include the Permian basin in western Texas/New Mex-
 190 ico in our analysis of the southcentral US and consider it and other western basins (Bakken,
 191 Denver-Julesburg, four-corners region, California) to be outside the scope of this study.
 192 In the northeast, many flight tracks were designed to fly downwind of the Marcellus and
 193 Utica gas basins in the western Appalachian regions of West Virginia and western Penn-
 194 sylvania, and thus this area has significant coverage in the influence maps. An excep-
 195 tion to this occurs during the winter campaign, where many of the downwind measure-
 196 ments were performed by the C130 aircraft which did not have a C_2H_6 instrument at
 197 the time, resulting in only 3 flights with direct C_2H_6 measurements downwind of the re-
 198 gion (see Figure S3 for a seasonal breakdown of the influence functions).

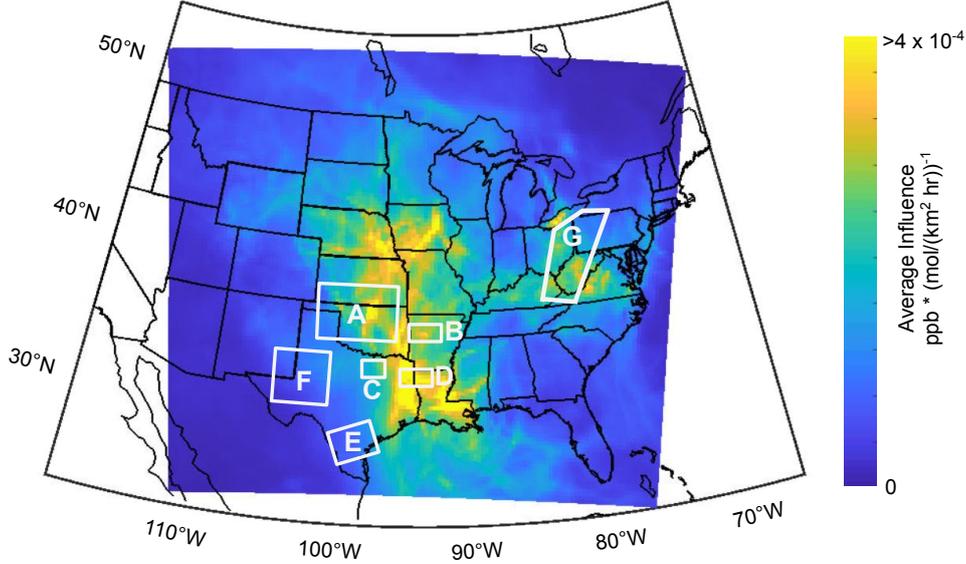


Figure 2. Averaged influence function from the boundary layer observations used in the C_2H_6 inversion. Brighter colors indicate areas whose surface interactions were captured more frequently by the boundary layer observations. O&G basins pertinent to this study are highlighted in boxes and are labeled as follows: (A.) Anadarko. (B.) Fayetteville. (C.) Barnett. (D.) Haynesville. (E.) Eagle Ford. (F.) Permian. (G.) Appalachia.

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2.2 Inversion Framework

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Observational C_2H_6 data collected from the ACT-America campaign is used in this study to perform an inversion and learn about C_2H_6 emissions from O&G sources in the flight domain. The basic inversion framework used is similar to previous work (Lauvaux et al., 2012; Sheng et al., 2018) and is expressed by cost function,

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$$J(\mathbf{x}) = \frac{1}{2}(\mathbf{y} - \mathbf{H}\mathbf{x})^T \mathbf{R}^{-1}(\mathbf{y} - \mathbf{H}\mathbf{x}) + \frac{1}{2}(\mathbf{x} - \mathbf{x}_0)^T \mathbf{B}^{-1}(\mathbf{x} - \mathbf{x}_0) \quad (1)$$

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In this equation, we solve for a posterior emissions grid \mathbf{x} that minimizes the cost function J using influence functions (\mathbf{H}) that translate the flux field to a modelled enhancement ($\mathbf{H}\mathbf{x}$). In the cost function, two terms control the solution. The first term is a cost related to the mismatch between the posterior modelled enhancements versus observed concentrations (\mathbf{y}), with greater discrepancies resulting in a larger cost term. Here, \mathbf{R} is the observation error covariance matrix, and weights the first term based on the confidence in the observations and model transport. The second term in the cost function equation is a cost related to the change between the posterior flux (\mathbf{x}) and the prior flux (\mathbf{x}_0), with larger changes resulting in a greater cost. Here, \mathbf{B} is the flux error covariance matrix, and weights the second term based on the confidence in the prior flux field. Minimizing the cost function with respect to \mathbf{x} yields

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$$\mathbf{x} = \mathbf{x}_0 + \mathbf{B}\mathbf{H}^T(\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R})^{-1}(\mathbf{y} - \mathbf{H}\mathbf{x}_0) \quad (2)$$

217 and solving for \mathbf{x} yields the posterior flux field. For this study, observations from
 218 each season are grouped together and an posterior flux map is solved for each of the sea-
 219 sonal flight campaigns.

220 A best guess C_2H_6 emissions map was created to serve as a prior for the inversion
 221 (Figure 4). To create this prior, CH_4 emissions from the O&G sector of the EPA Grid-
 222 ded 2012 CH_4 Emissions Inventory (Maasakkers et al., 2016) were multiplied by expected
 223 molar $\text{C}_2\text{H}_6/\text{CH}_4$ ratios of each basin (Table S2), resulting in an C_2H_6 emissions map.
 224 For Eagle Ford, Haynesville, Fayetteville, Barnett, Permian, Denver-Julesburg, and the
 225 Bakken O&G basins, atmospheric measurements from NOAA aircraft studies were avail-
 226 able to derive observed basin-wide $\text{C}_2\text{H}_6/\text{CH}_4$ ratios (Peischl et al., 2015, 2018; Tzompa-
 227 Sosa et al., 2017). For the Anadarko and Appalachian region where representative at-
 228 mospheric measurements were not available, data from the United States Geological Sur-
 229 vey were used to create a spatial map of $\text{C}_2\text{H}_6/\text{CH}_4$ ratios to apply to these regions (US
 230 Geological Survey, 2018; Kitanidis, 1997)(See supplemental section S1 for additional info).
 231 For all emissions related to transmission, storage, and distribution, an $\text{C}_2\text{H}_6/\text{CH}_4$ ratio
 232 of 0.027 was applied (Plant et al., 2019). For prior C_2H_6 emissions in Canada and Mex-
 233 ico, we use CH_4 emissions from O&G facilities provided in Sheng et al. (2017) and ap-
 234 ply the mean $\text{C}_2\text{H}_6/\text{CH}_4$ ratio from the USGS database of 0.085 to convert it to C_2H_6 emis-
 235 sions. These sources outside the US have minimal representation in our influence func-
 236 tions and have no impact on the overall solution. Although biomass burning and bio-
 237 fuels are also significant sources of C_2H_6 on a global scale, these sources are estimated
 238 to be only a few percent of the total C_2H_6 emissions in the US (Tzompa-Sosa et al., 2017).
 239 Overall, this C_2H_6 emissions map represents our best guess as to representing C_2H_6 emis-
 240 sions based on the EPA’s bottom-up O&G CH_4 emissions inventory, and will be referred
 241 to as the "Default" map henceforth.

242 Two additional C_2H_6 emission maps were created to test the sensitivity of the in-
 243 version to different priors. The first alternative map was created by taking the EPA CH_4
 244 emissions map used in the creation of the Default inventory and applying a flat $\text{C}_2\text{H}_6/\text{CH}_4$
 245 ratio of 0.085. This ratio preserves the total C_2H_6 emissions from the "Default" map but
 246 redistributes them in a way that removes knowledge of the unique gas compositions of
 247 different basins, and is referred to henceforth as the "Flat Rate" inventory. The second
 248 alternative map is based on ethane emissions provided by the US National Emissions In-
 249 ventory 2011 inventory (NEI2011) C_2H_6 , an C_2H_6 map available in GEOS-Chem and used
 250 in Tzompa-Sosa et al. (2017). In addition to having a different spatial distribution com-
 251 pared to the Default inventory, total emissions in the NEI2011 inventory are roughly half
 252 the Default total. C_2H_6 emissions from the Emissions Database for Global Atmospheric
 253 Research (EDGAR) v4.3.2 were originally considered as well for this analysis, but to-
 254 tal emissions were 8 times lower than our Default inventory and were decided to be too
 255 inaccurate to serve as a useful prior (Huang et al., 2017).

256 Continuous boundary layer C_2H_6 measurements from the ACT-America campaign
 257 were used as observational input for this study, with their associated FLEXPART-derived
 258 influence functions used for \mathbf{H} in the inversion (section 2.1). Observations greater than
 259 1000 meters above ground level were not included in the analysis to remove any data above
 260 the ABL. For this study, we restrict the domain of our inversion to the area within the
 261 4 corners (23.7°N, 110.72°W), (23.0°N, 77.5°W), (49.9°N, -67.3°W), (51.1°N, -119.0°W,
 262 domain shown as the colored region in Figure 2). Because the influence functions only
 263 provide a local enhancement inside the study domain, for each flight date a background
 264 value is determined to represent the C_2H_6 mixing ratios entering the domain. This value
 265 is chosen by taking the 5th percentile of the observed boundary layer C_2H_6 concentra-
 266 tions on a given flight and subtracting it from the observations, producing an observed
 267 C_2H_6 enhancement. The 5th percentile of model-projected enhancements along the flight
 268 track is then added onto the observed enhancements in order to align the modelled and
 269 observed background values. This final step is necessary in rare scenarios where mod-

270 elled O&G enhancements are influencing the entire aircraft transect, thus impacting the
 271 observed background concentrations (see Barkley et al. (2019b) for more details). In cal-
 272 culating the modelled C_2H_6 enhancements for this study, we treat C_2H_6 as an inert gas
 273 rather than a reactive one due to its long average lifetime (weeks to months) relative to
 274 the length of time the local plumes travel from the source to the aircraft (hours to days).
 275 For more information on the possible influence of C_2H_6 loss rates, see supplemental sec-
 276 tion S1.

277 To run an inversion, values must be assigned to the \mathbf{R} and \mathbf{B} matrices related to
 278 the uncertainty in the observation/transport and prior flux fields respectively. For the
 279 observational/transport uncertainty matrix \mathbf{R} , we use a method similar to the residual
 280 error method discussed in Sheng et al. (2017). For each flight, modelled enhancements
 281 are first scaled by a constant to have the same overall enhancement as the observed en-
 282 hancements. This step is performed to remove any existing overall bias that may exist
 283 in the prior inventory for the calculation of \mathbf{R} . After removing this bias, the residual er-
 284 rors are calculated between the model and observation and the standard deviation of this
 285 error is used to represent the \mathbf{R} value along its diagonal for a given flight, with no value
 286 assigned for the off-diagonals. Values for the diagonal elements of \mathbf{R} varied across flights,
 287 but seasonal averages for the standard deviation of the error ranged from 0.7 ppb in the
 288 spring to 1.8 ppb in the fall. This process results in flights with large observational and
 289 transport uncertainty on days with large enhancements and poor correlations between
 290 the observed and modelled values, thus giving these flights less weight in the inversion
 291 solution. Similarly, flights where observed and modelled plume structures align have a
 292 smaller \mathbf{R} value assigned and thus are given greater weight in the overall solution. This
 293 method for classifying transport uncertainty is particularly effective for an C_2H_6 inver-
 294 sion study, since the locations of the sources (i.e. O&G infrastructure) are known with
 295 high confidence such that misaligned plumes would most likely be caused by errors in
 296 the transport rather than problems with the spatial mapping of the flux.

297 For the flux uncertainty matrix \mathbf{B} , there lacks a clear answer on what the uncer-
 298 tainty of the prior fluxes should be. Since the primary source of C_2H_6 in the US is from
 299 O&G production and processing, the location of C_2H_6 emitters in the US should be ac-
 300 curate. However, leak rates from O&G activity have been shown in various studies to
 301 be beyond the uncertainty bounds of the EPA 2012 Gridded CH_4 Inventory (Alvarez et
 302 al., 2018; Maasackers et al., 2016). For this reason, a value for the flux uncertainty \mathbf{B}
 303 is selected based on mathematical constraints of the inversion rather than an understand-
 304 ing of the prior flux map. Specifically, all flux grids are assigned an error as a percent-
 305 age of their prior, where the percent uncertainty is selected using a chi-squared metric,
 306 defined as

$$307 \quad \chi^2 = \frac{1}{m} (\mathbf{y} - \mathbf{H}\mathbf{x}_0)^T (\mathbf{H}\mathbf{B}\mathbf{H}' + \mathbf{R})^{-1} (\mathbf{y} - \mathbf{H}\mathbf{x}_0) \quad (3)$$

308 where m is the number of observations and χ^2 is the chi-squared metric used to as-
 309 sess whether the inversion errors satisfy a Gaussian distribution. Here, we select a per-
 310 cent error for the flux uncertainty in \mathbf{B} that brings χ^2 close to 1. For the Default, Flat
 311 Rate, and NEI2011 inventory, the assigned flux errors along the diagonal of \mathbf{B} were 50,
 312 50, and 80 percent of the total grid emissions respectively. Additionally, a correlation
 313 length with an e-folding decay length of 50 km is assigned to the off-diagonal elements
 314 in \mathbf{B} . This added correlation provides consistent shifts in the emissions within a basin,
 315 while allowing changes across basins to behave uniquely. Implications of the \mathbf{R} and \mathbf{B}
 316 matrices selected for this study are explored further in the sensitivity analysis (Table S4).

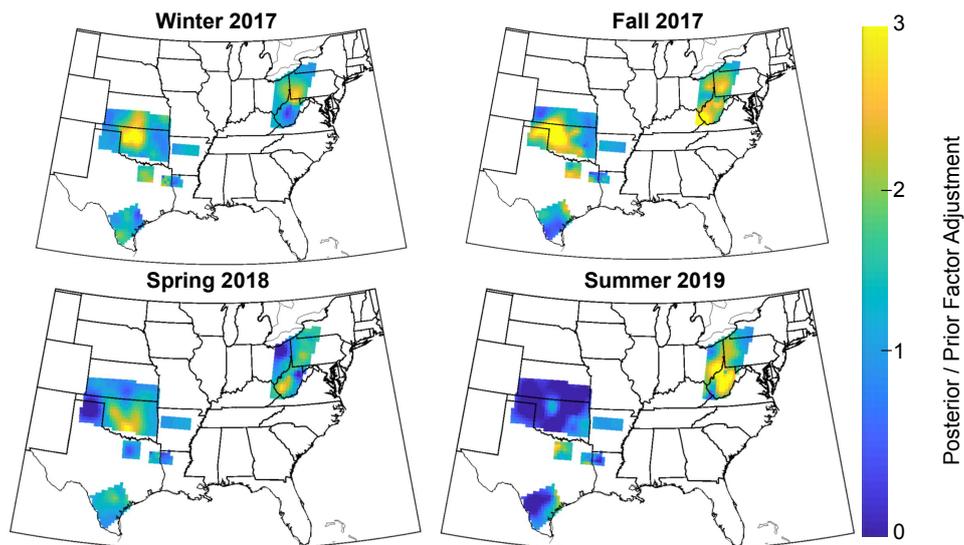


Figure 3. Fractional changes between the Default posterior and prior C_2H_6 flux map created by an inversion conducted for each individual season. Only basins within the scope of this study are shown on the map.

3 Results and Discussion

3.1 C_2H_6 Inversion Results

For each season, posterior C_2H_6 flux maps were successful in reducing both the error and bias and increasing the correlation between the observed and modelled signal for that season (Table 1, Figures S4-S9). Notably, across all 4 seasons there was an overall low bias between the modelled prior and observed C_2H_6 enhancements, resulting in seasonal posterior maps that generally increased C_2H_6 emissions in order to compensate. Regionally, observed C_2H_6 plumes were underestimated the largest in the western Appalachian region across all 4 seasons (Figure 3). In this region, posterior C_2H_6 emissions were increased by 50 to 150% more than the prior to correct for the underestimation. In the southcentral US, a similar low bias was observed in the winter, fall, and spring campaign. This led to posterior solutions for these three seasons that show a systematic increase to the total posterior flux of 50%. This increase is mostly uniform across the individual O&G basins, with a notable exception for the Haynesville basin whose posterior is consistently within 11% of the Default prior across all seasons.

Observations from the southcentral and midwestern US during the Summer 19 campaign behave like outliers compared to the rest of the dataset. Of the 87 flights used in the inversion analysis, only 12 had an overall negative bias when comparing the observed C_2H_6 to the model prior (observed enhancement less than modelled enhancement), 8 of which occurred in the southcentral and midwestern portion of the Summer 19 campaign. These low observations have a profound effect on the posterior emissions in the southcentral US for the Summer 19 campaign, with total emissions from this region that are less than a third of posterior emission maps from the other three seasonal campaigns. To demonstrate how unrepresentative the Summer 19 results are compared to the rest of the dataset, we take the posterior inventory derived for each season and apply it to the entire 4 season dataset (Table 2). In doing this, we find that the winter, fall, and spring posteriors all produce similar statistical improvements to the overall dataset compared

| | Default Inventory Winter 2017 Performance | | Default Inventory Fall 2017 Performance | | Default Inventory Spring 2018 Performance | | Default Inventory Summer 2019 Performance | |
|--|---|-----------|---|-----------|---|-----------|---|-----------|
| | Prior | Posterior | Prior | Posterior | Prior | Posterior | Prior | Posterior |
| Southcentral Total C ₂ H ₆ (mol s ⁻¹) | 488 | 709 | 488 | 777 | 488 | 601 | 488 | 199 |
| Western Appalachia Total C ₂ H ₆ (mol s ⁻¹) | 125 | 184 | 125 | 294 | 125 | 185 | 125 | 269 |
| Mean Absolute Error (ppb) | 1.39 | 0.99 | 1.65 | 1.16 | 0.68 | 0.54 | 0.85 | 0.59 |
| Mean Bias (ppb) | 1.20 | 0.61 | 1.50 | 0.76 | 0.55 | 0.37 | 0.25 | 0.20 |
| y,Hx Correlation | 0.85 | 0.90 | 0.75 | 0.83 | 0.61 | 0.72 | 0.44 | 0.74 |

Table 1. Table describing the performance of each of the seasonal posteriors relative to the observations from that season.

344 to the prior. However, when applying the summer C₂H₆ posterior to the 4 season dataset,
 345 not only does it perform substantially worse than each of the other seasonal posteriors,
 346 it also performs worse than the prior, with increases to the absolute error, bias, and a
 347 substantial decrease in the model-obs correlation coefficient from 0.71 to 0.50.

348 One possible explanation for the discrepancy in summer could be related to errors
 349 in the model transport simulation. The mean absolute error in the modelled boundary
 350 layer wind direction compared to flight observations were on average 30 degrees per flight
 351 in the summer campaign versus 15 in each of the other three seasons, likely related to
 352 the slower wind speeds observed during summer (Table S3). These directional transport
 353 errors can create misalignments in observed versus modelled plumes, leading to an over-
 354 all reduction in the posterior emissions relative to the truth. Compounding on these wind
 355 issues is the possibility of increased convective activity in summer. While boundary layer
 356 statistics are tracked in the model and compared to observations, it is more difficult to
 357 assess how much of the signal is lost out the top of the boundary layer due to summer-
 358 time convective processes. Another, non-transport related possibility is that there could
 359 be a non-trivial loss of C₂H₆ due to an increase in the OH chemical sink during the sum-
 360 mer months. In this study, the impacts of C₂H₆ loss were not considered, as C₂H₆ has
 361 an average lifetime of two months and most plumes were captured within 48 hours of
 362 release from the source (Burkholder et al., 2015). However, in conditions with excessive
 363 heat and large OH concentrations that can occur in the summer, the lifetime of C₂H₆
 364 can be reduced to as little as four days in the most extreme conditions (see supplement-
 365 al section S2 for more details). Even so, a lifetime of four days would only have minor
 366 impacts on local plumes, and the extreme conditions leading to high C₂H₆ loss would
 367 only exist for a short period in the afternoon hours. Furthermore, summer C₂H₆ obser-
 368 vations from the northeast were elevated to levels similar to those observed during the
 369 other seasons, seemingly unaffected by summertime conditions. For these reasons, it is
 370 unlikely that a chemical loss could be responsible for the low values observed in the south-
 371 central US. One final possibility is that the low C₂H₆ enhancements observed in the Sum-
 372 mer 19 campaign are due to a real and significant temporal change in C₂H₆ emissions
 373 in the southcentral US during this period, particularly in the Anadarko basin. However,
 374 the near-zero emission rate solution provided by the inversion posterior in the Anadarko
 375 lacks any sort of real-world explanation for such a large shift compared to previous sea-

| | Default Inventory Prior | Default Inventory Winter 2017 Posterior | Default Inventory Fall 2017 Posterior | Default Inventory Spring 2018 Posterior | Default Inventory Summer 2019 Posterior |
|---|-------------------------|---|---------------------------------------|---|---|
| Southcentral Total C₂H₆ (mol s⁻¹) | 488 | 709 | 777 | 601 | 199 |
| Western Appalachia Total C₂H₆ (mol s⁻¹) | 125 | 184 | 294 | 185 | 269 |
| Four Season Mean Absolute Error (ppb) | 1.15 | 1.02 | 0.96 | 1.04 | 1.18 |
| Four Season Mean Bias (ppb) | 0.86 | 0.21 | 0.23 | 0.58 | 0.94 |
| Four Season y,Hx Correlation | 0.71 | 0.73 | 0.78 | 0.71 | 0.50 |

Table 2. Table describing the performance of each of the individual seasonal posterior fluxes when each is applied relative to all four seasons of observations. Green areas highlight statistical improvement compared to the prior, whereas red boxes show degradation.

376 sons. Production data shows the Anadarko basin was at its peak O&G production rates
 377 during the period (US Energy Information Administration, 2020a). Regardless of the rea-
 378 son, the low C₂H₆ observations measured in the southcentral during the Summer 19 months
 379 are not representative of data from the winter, fall, and spring campaigns, which show
 380 strong consistency in the location and expected magnitudes of regional C₂H₆ enhance-
 381 ments. For this work, we choose to discard summer data in evaluating our best estimate
 382 of US C₂H₆ emissions, but consider it a source of uncertainty and intrigue for future re-
 383 search.

384 To create our best guess regarding C₂H₆ emissions across the southcentral and east-
 385 ern US, we weight the Default posterior flux maps from winter, fall, and spring with equal
 386 confidence and average them together (Figure 4). This averaged posterior solution pro-
 387 vides consistent improvement compared to the prior, reducing the absolute error between
 388 the model and observations in 84% of flights, providing confidence that this averaged so-
 389 lution serves as an improvement to a large majority of the observational data for the three
 390 seasons it represents. C₂H₆ emissions from the 3 season posterior are almost universally
 391 increased compared to the prior, with a 43% increase overall in the southcentral US and
 392 a 76% increase in the Appalachia. With these large increases in the emissions the mean
 393 obs-model bias is reduced, decreasing from 1.14 ppb to 0.70 ppb. The inability to elim-
 394 inate this bias is a natural result of errors in the transport preventing the inversion from
 395 solving for misaligned observed plumes, and thus the emission increases in the 3 season
 396 posterior are likely still an underestimation of the true emissions.

397 To better understand the sensitivity of our 3 season posterior to the prior, the in-
 398 version was rerun using the two alternative prior maps discussed in section 2.2 and a 3
 399 season posterior was created from each set of priors (Figures 4 and 5) For every basin,
 400 posteriors from the three inventories converged towards a similar solution compared to
 401 their priors. An example of this is the Haynesville basin, where the two alternative prior
 402 inventories have a factor of 4 spread between their emissions, but their posteriors con-

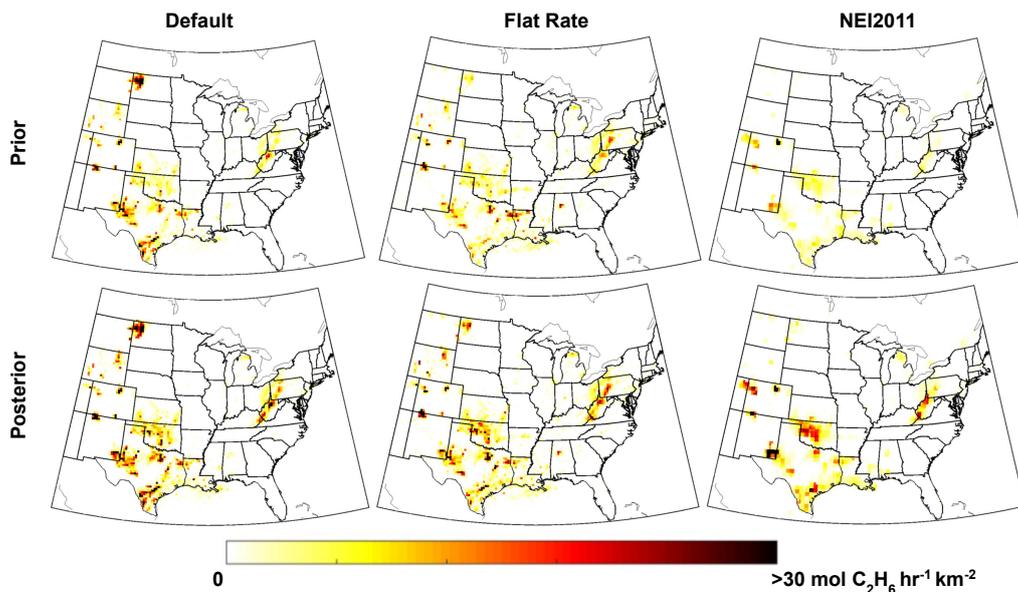


Figure 4. Prior C₂H₆ inventories used in this study and their respective 3 season mean posteriors. "Default" represents the best guess prior from this study based on multiplying the O&G sector of the Gridded EPA 2012 CH₄ Emissions Inventory by projected C₂H₆/CH₄ ratios of individual basins. "Flat Rate" multiplies the same CH₄ inventory by a flat rate C₂H₆/CH₄ ratio of 0.085, producing a similar total as "Default" with a different spatial representation. "NEI2011" comes from the NEI 2011 C₂H₆ inventory.

403 converge to within 20% of the Default posterior solution. Ultimately, the 2 alternative pos-
 404 teriors produced solutions for the entire southcentral US that were within 1 mol/s of each
 405 other, despite starting 183 mol/s apart. Their similar solutions for the southcentral re-
 406 gion are 12% lower than the posterior calculated using the Default inventory. This dif-
 407 ference is driven primarily by the Eagle Ford basin, which had the largest absolute spreads
 408 between prior inventories and whose western portion lies on the edge of the region of in-
 409 fluence captured by the ACT-America flights (Figure 2), making it more difficult for the
 410 inversion to constrain using the available dataset. In the northeast, there was a similar
 411 consensus among the three inventories of a substantial increase in emissions required to
 412 recreate the large plumes observed there. While the emissions in the NEI2011 posterior
 413 do not increase to the same levels as the other two priors (171 vs 221 and 239 mol/s),
 414 the NEI2011 prior in the northeast was missing nearly all of the emissions from conven-
 415 tional gas activity in western Pennsylvania, resulting in a posterior map that still ap-
 416 pears to be underestimating the sources in that region even after substantial increases
 417 given the large positive bias between the observations and modelled posterior. Overall
 418 the NEI2011 prior, whose total C₂H₆ emissions are 53% less than the Default prior,
 419 leads to a total posterior C₂H₆ emissions estimate that is only 14% lower than the Default pos-
 420 terior.

421 In addition to testing the sensitivity of the inversion posterior to different priors,
 422 we also test its sensitivity to numerous other conditions, including adjustments to the
 423 magnitude of the prior, changes to the selection of the background C₂H₆ term, elimina-
 424 tion of the length correlations applied to the prior flux uncertainty, adjustment of the
 425 observation matrix error term to a constant (giving equal uncertainty to all observations),
 426 and solving for a posterior solution using all three seasons of data simultaneously rather

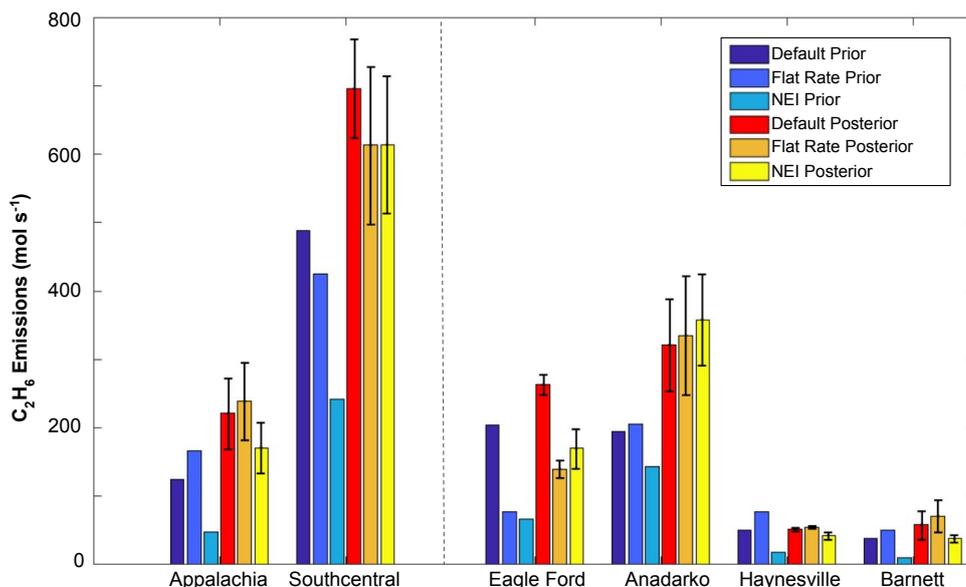


Figure 5. Priors and their respective 3 season posterior C_2H_6 emissions for major basins in this study. "Southcentral" is the sum of Eagle Ford, Anadarko, Haynesville, and Barnett. Error bars show the standard deviation of the individual season solutions for each basin.

427 than averaging 3 unique posterior maps (Table S4). All resulting posteriors produce sim-
 428 ilar results compared to the Default posterior, providing further confidence in the solu-
 429 tion. In particular, using our Default prior multiplied by a factor of 3 converged to a sim-
 430 ilar solution as our Default posterior, reducing concerns that using a prior with a low
 431 bias could be resulting in an large underestimation of the solution. One source of uncer-
 432 tainty that is difficult to address in our sensitivity tests are errors in the transport
 433 used to create the influence functions. Average flight wind speed biases were under 1.3
 434 m/s and wind direction absolute errors were close to 15 degrees for the winter, fall, and
 435 spring campaigns when compared to ACT-America boundary layer aircraft data (Ta-
 436 ble S3). However, modelled boundary layer heights were 15 to 22 percent lower on av-
 437 erage seasonally compared to observations. A negative bias in the modelled boundary
 438 layer would result in an equally proportionally higher value in the influence functions
 439 and model-projected enhancements. Applying a correction factor to the influence func-
 440 tions relative to the mean bias for each season and rerunning the inversion produces an
 441 overall posterior that is 22% higher than the Default posterior. Consistent with our past
 442 work that has applied corrections to simulated errors in ABL depth and winds (Barkley
 443 et al., 2017, 2019a, 2019b), we consider the posterior created using the ABL depth cor-
 444 rection to be an equally plausible solution in our best estimate of the C_2H_6 emissions,
 445 as the boundary layer bias is a source of error with a known and somewhat correctable
 446 bias on the solution. The boundary layer bias, along with a negative enhancement bias
 447 still present in the model versus observational comparison of the posterior C_2H_6 enhance-
 448 ments, are both potential reasons to suspect the Default posterior on its own may still
 449 be underestimating overall C_2H_6 emissions despite the large increase in the emissions re-
 450 lative to the prior.

451 3.2 Interpretation of CH_4 Emissions from the O&G Sector

452 The Default prior inventory developed in this study is created by multiplying the
 453 EPA 2012 Gridded CH_4 Emissions Inventory for O&G sources by the suspected C_2H_6/CH_4

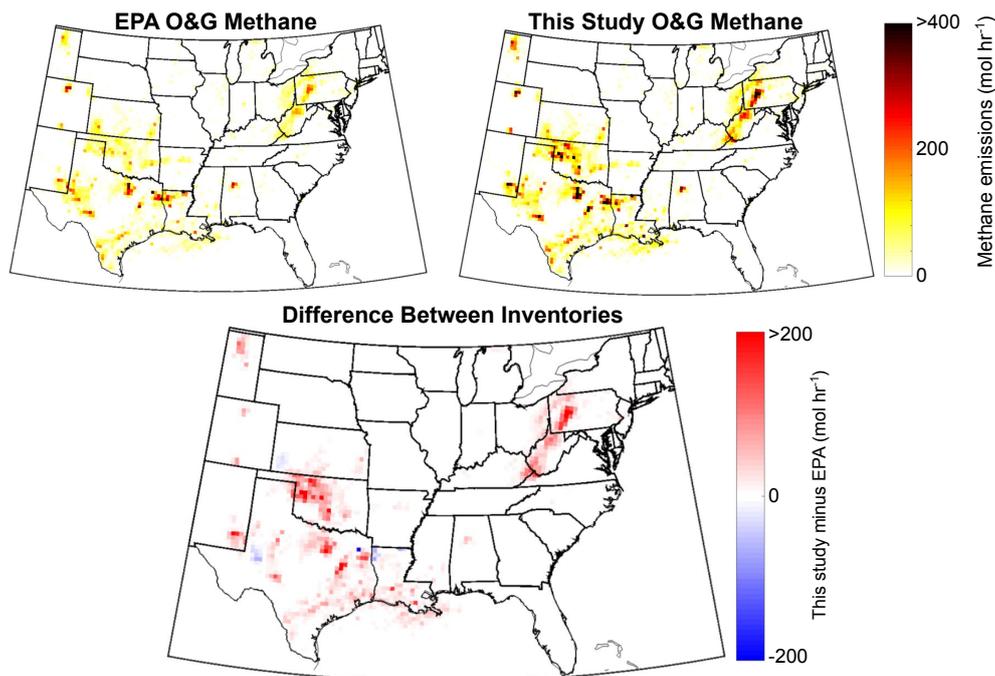


Figure 6. Top-left: Oil and gas CH₄ emissions from the EPA 2012 Gridded CH₄ Inventory. Top-right: Oil and gas CH₄ emissions estimated from the C₂H₆ posterior in this study. Bottom: Difference between the two inventories.

454 ratio of each grid's emissions. If the assumed ratios are correct, and emissions primar-
 455 ily occur in situations where the gas content is unaltered (i.e. processes unrelated to gas
 456 separation), then changes between the posterior and prior Default C₂H₆ inventory should
 457 proportionally reflect changes in the EPA 2012 CH₄ inventory for O&G. Thus, we can
 458 use our Default posterior C₂H₆ inventory to create our best interpretation of O&G CH₄
 459 emissions in the southcentral and eastern US.

460 Using the converted C₂H₆ posterior as a proxy for O&G CH₄ emissions, our inven-
 461 tory projects that O&G CH₄ emissions are almost universally larger than the 2012 EPA
 462 Inventory estimates (Figure 6). In the southcentral US, we estimate emissions to be 48%
 463 higher than inventory estimates (77% using the ABL-adjusted posterior). Of this increase,
 464 two-thirds of it is driven by increases in the Anadarko basin, which had the largest pro-
 465 portional change in the C₂H₆ posterior and is the largest source of CH₄ emissions in the
 466 region. Of the four remaining southcentral basins captured in this study, the Haynesville
 467 basin is the only basin that did not see a significant increase in its emissions relative to
 468 the prior. As noted in section 2.1, our solution does not include potential changes from
 469 the Permian basin which is outside of the region of influence captured by ACT-America
 470 campaign. O&G production in the Permian has increased by more than a factor of 3 since
 471 2012 and some of the largest CH₄ signals from the basin have been measured using satel-
 472 lite observations in the area (Zhang et al., 2020). As such, it is likely that emissions in
 473 the Permian follow a similar pattern of underestimation observed for other southcentral
 474 basins in our study.

475 In the Appalachians, we estimate CH₄ emissions to be 77% higher than EPA 2012
 476 inventory estimates (105% using the ABL-adjusted posterior), the largest difference of
 477 any basin in this study. Part of the discrepancy between inventory results and the pos-
 478 terior may be related to the increased presence of unconventional natural gas activity

479 in the Marcellus shale. Between 2012 and 2018 Pennsylvania and West Virginia under-
480 went some of the largest gas production growth in the US, with annual production tripling
481 during the period (US Energy Information Administration, 2020a), a change that would
482 not be captured in the EPA 2012 inventory. However, Pennsylvania state inventories,
483 which provide annual inventory estimates of unconventional natural gas activity in the
484 state using methodologies similar to the EPA, show CH₄ emissions from unconventional
485 activity only increased by 20% during this period, and that these unconventional wells
486 only represent a portion of O&G emissions in the region (Omara et al., 2016), with much
487 of the emissions coming from pre-existing conventional activity. Thus, it is unlikely that
488 changes in unconventional activity between 2012 and the time of this study would be re-
489 sponsible for the 77% increase in regional emission rates found from the inversion results
490 relative to the EPA 2012 Gridded Inventory, and that the discrepancy would still be present
491 in an updated EPA inventory.

492 The CH₄ inventory estimates for individual basins from this study generally align
493 with mass balance studies of corresponding basins. In the Haynesville basin we calcu-
494 late an O&G emission rate of 43 Mg/hr, compared to 42 Mg/hr from Peischl et al. (2018)
495 and 76 Mg/hr from Cui et al. (2017), which includes non-O&G values in its total as well.
496 In the Barnett, we calculate emissions to be 57 Mg/hr, larger than 46 Mg/hr found in
497 Peischl et al. (2018) but close to the 60 Mg/hr found in Karion et al. (2015). In Eagle
498 Ford, both the western and eastern basin in this study had a combined emission rate of
499 68 Mg/hr versus 83 Mg/hr in Peischl et al. (2018). Additionally, the large values observed
500 in the Appalachian match findings that show emissions from unconventional O&G in-
501 frastructure in the Marcellus are greatly underestimated by EPA inventory values (Barkley
502 et al., 2019a; Caulton et al., 2019).

503 The interpretation of O&G CH₄ emissions using C₂H₆ observations has a unique
504 advantage compared to more traditional methodologies that rely on CH₄ measurements
505 due to the simplicity of C₂H₆ sources. In the US where CH₄ emissions have near equal
506 contributions from fossil fuels, agriculture, and wetlands, each of which have their own
507 uncertainties, C₂H₆ emissions are dominated almost entirely by the O&G sector. Fur-
508 thermore, there is high confidence in the spatial mapping of O&G sources in the US due
509 to extensive documentation of the various components associated with O&G extraction,
510 simplifying interpretation of atmospheric C₂H₆ data. As an example, in the ACT-America
511 campaign, the majority of aircraft flight tracks were hundreds to thousands of kilome-
512 ters away from O&G basins. Despite this, the model prior was able to consistently track
513 C₂H₆ enhancements from these sources in the winter, fall, and spring, with correlations
514 between the model vs observed boundary layer C₂H₆ enhancements of 0.85, 0.75, and
515 0.61 respectively. The high skill in tracking enhancements from a single sector with well-
516 defined locations creates a scenario where a stable posterior solution can be generated
517 through various inverse methodologies (Table S4).

518 Despite high confidence in the C₂H₆ posterior, the conversion of this posterior to
519 O&G CH₄ emissions is entirely dependent on the quality and availability of information
520 related to the C₂H₆/CH₄ ratio for each basin. During the time observations from this
521 study were collected, there were numerous recent flights available from a separate study
522 that captured the C₂H₆/CH₄ ratios of various basins (Peischl et al., 2018), providing con-
523 fidence in our C₂H₆ to CH₄ conversions for those locations. However, basins that are ge-
524 ographically broad, such as the Anadarko and Marcellus, are more difficult to charac-
525 terize a ratio for using atmospheric data. Furthermore, the average gas composition of
526 a basin can change over time, making ratios found in older studies less applicable (Lan
527 et al., 2019). Publicizing upstream gas composition data collected by major O&G com-
528 panies would be one immediate solution and effectively eliminate any uncertainty intro-
529 duced in an C₂H₆ to CH₄ inventory conversion.

530 One other source of uncertainty related to the C₂H₆ to CH₄ inventory conversion
531 presented in this study is the presence of coal mines in the Appalachia that overlap with

532 regional gas production. These mines are a dominant source of CH_4 emissions in the re-
533 gion, but little information is available regarding C_2H_6 emissions associated with them.
534 For this study, we treat C_2H_6 emissions from these sources to be negligible based mea-
535 sured values from a 1973 geological survey that found little to no C_2H_6 in many of the
536 mines relevant to this study (Kim, 1973). However, ratios as high as 0.08 have been ob-
537 served in western Kentucky (Strapoć et al., 2007), raising the possibility that some por-
538 tion of the C_2H_6 observed in this study in the Appalachia could be attributable to un-
539 derground mines. Air from major coal ventilation shafts is sampled 4 times each year
540 to quantify the amount of CH_4 present for safety precautions (US Environmental Pro-
541 tection Agency, 2017). Measuring C_2H_6 mixing ratios in these samples would provide
542 additional information to expand our understanding of C_2H_6 emissions in the northeast-
543 ern US.

544 4 Conclusion

545 Using the largest collection of airborne boundary layer C_2H_6 data to date, an in-
546 version was performed to estimate C_2H_6 and CH_4 emissions from various O&G basins
547 across the eastern US. From this dataset, we estimate that a large portion of C_2H_6 emis-
548 sions cannot be explained using O&G emission data from the EPA 2012 Gridded CH_4
549 Inventory and existing C_2H_6 to CH_4 emissions ratio data. We conclude that the EPA
550 CH_4 emissions inventory significantly underestimates CH_4 emissions from O&G sources.
551 From this dataset, our results yield CH_4 emissions estimates similar in magnitude to na-
552 tional estimates in Alvarez et al. (2018). In particular, this study finds the largest dis-
553 crepancies occurring in the Anadarko and western Appalachian O&G plays, though all
554 basins other than Haynesville showed an increase in expected C_2H_6 emissions relative
555 to the prior. We also find C_2H_6 emissions from O&G that are much higher than exist-
556 ing C_2H_6 inventories, a fact that should be taken into account in any future efforts to
557 investigate O&G CH_4 emissions using atmospheric C_2H_6 observations, as well as stud-
558 ies relying on existing C_2H_6 inventories to account for certain chemical reactions.

559 This study reveals the potential to use broad-scale continuous C_2H_6 data to con-
560 strain CH_4 emissions from the O&G sector on a nationwide scale. Our confidence in the
561 spatial distribution of emissions associated with the O&G sector and the dominant role
562 of O&G in US C_2H_6 emissions allows for modelling and interpretation of observed sig-
563 nals without large concerns for source attribution. Furthermore, misalignment in observed
564 versus modelled plumes may serve as a useful diagnostic of model transport errors due
565 to the well-documented spatial knowledge of C_2H_6 emitters in the US. One current weak-
566 ness with relying on C_2H_6 observations to understand CH_4 emissions from O&G is the
567 requirement of knowing the C_2H_6 to CH_4 ratio of gas composition and basin emissions
568 in order to convert C_2H_6 emissions into CH_4 emissions. Gas composition data is collected
569 by individual companies but not shared publicly. Should some form of this information
570 be made available for the public, it would greatly enhance the confidence of CH_4 emis-
571 sion estimates of the O&G industry using C_2H_6 and other trace gases.

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