

## 742 Appendix A GPS Analysis

743 The hypothesis of episodic slow slip has been postulated by employing solely hor-  
 744 izontal GPS recordings. Here, we use all three components of GPS recording (two hor-  
 745 izontal and one vertical) to demonstrate the 3D deformation of the continental crust over  
 746 time and how their magnitudes relate to tremor location.

### 747 A1 Reliability of Vertical GPS Measurements

748 Uncertainty in vertical GPS measurements is approximately 3 times that of hor-  
 749 izontal measurements. More importantly, we recognize that seasonal variations in sur-  
 750 face mass variations can have substantial impact on vertical GPS measurements Blewitt  
 751 et al. (2001); Dong et al. (2002); Bettinelli et al. (2008).

752 Here, however, we ignore the effect of seasonal changes on vertical GPS measure-  
 753 ments because it is extremely challenging to decouple the effect of seasonal surficial mass  
 754 changes from displacement due to tectonic deformation. This task become especially chal-  
 755 lenging in Cascadia where the episodic deformation cycle spans 13–14 months which is  
 756 close to seasonal cycles (12 months).

757 Nonetheless, we observe that

- 758 • vertical GPS measurements are large and in many cases an order of magnitude  
 759 larger than horizontal displacements,
- 760 • there is a close correspondence between sudden changes in horizontal displacements  
 761 (horizontal GPS reversals) and rapid vertical GPS measurements on numerous oc-  
 762 casions, and
- 763 • vertical cyclic displacement patterns (Figure A3 and A4) show close spatial cor-  
 764 respondence with spatial tremor patterns in Cascadia and Alaska.

765 Given the above observations, one might argue that vertical displacements observed con-  
 766 tain significant imprints of tectonic deformation.

### 767 A2 Data Processing

768 Prior to hodogram analysis, GPS data are detrended using a 1001 point median  
 769 filter to eliminate long-term trends, and thereafter filtered using a 11-point median fil-  
 770 ter to suppress short-term noise bursts. GPS stations with significant noise that could  
 771 not be corrected from using the above filtering operations were not used in the analy-  
 772 sis.

773 Computation of the net vertical and horizontal GPS displacements was done by  
 774 fitting ellipsoids to the hodograms. Projection of the major axis of the ellipse on the ver-  
 775 tical direction and the horizontal plane yield the net vertical and horizontal displacements,  
 776 respectively.

### 777 A3 Displacements due to Buckling and Collapse

778 Figure A1 shows the expected temporal evolution of the vertical (blue) and hor-  
 779 izontal (red) displacements of four locations A, B, C, and D, (phase  $T_0$ , Figure 2) on the  
 780 surface of a continental plate through a buckling and collapsing cycle. The magnitude  
 781 of the horizontal displacement is expected to decrease monotonically from the corner of  
 782 the accretionary wedge (location A) landward as depicted by the decreasing range of the  
 783 horizontal displacement moving from A through D. The vertical displacement, however,  
 784 is small at location A, attains a maximum at location C, and tapers off to a small value  
 785 further landwards (location D).

786 An efficient technique to analyze and quantify such multi-component data is to gener-  
787 ate hodograms which are a display of the motion of a point as a function of time. Fig-  
788 ure A1 shows the hodograms for each of the four locations A, B, C, and D on the right.  
789 The path followed by a particle during the buckling phase is different from that followed  
790 during the collapse phase, thereby resulting in hysteresis of the particle motion. Note  
791 that such hysteresis demonstrates a non-linear particle motion (Figure A1) as opposed  
792 to an expected linear motion for the case of slow slip. Moreover, it is clear from the hodograms  
793 that the horizontal displacement decreases monotonically from the corner of the accre-  
794 tionary wedge (location A) landward, while the vertical displacement attains a maximum  
795 right above the narrow tremor zone.

796 Figure A2 shows an example of a hodogram obtained from GPS data. This data  
797 comes from the Albert Head GPS site on Vancouver Island in Victoria, British Columbia  
798 – the data for which was originally employed by Rogers and Dragert (2003) to hypoth-  
799 esize the process of slow slip. Note the hysteresis and the prominent vertical displace-  
800 ment observed at this site which is quite similar to the pattern expected for surface lo-  
801 cation C (Figure A1) right above a tremor belt. Other studies (Wech et al., 2009; Wells  
802 et al., 2017) indeed map significant tremor activity beneath this GPS site.

#### 803 **A4 Horizontal and Vertical Displacements in Cascadia and Alaska**

804 We generate hodograms for all the GPS measurements at sites in the Cascadia sub-  
805 duction zone and in Alaska and thereafter compute the vertical displacement, horizon-  
806 tal displacement, and their ratio. These attributes for Cascadia and Alaska are shown  
807 in Figures A3 and A4, respectively. Note that in both cases, the horizontal displacement  
808 decreases monotonically from the margin landwards; while the vertical displacement in-  
809 creases as one moves landwards from the margin, attains a maximum, and decreases there-  
810 after. The belt of maximum vertical displacements along the Cascadia margin has a close  
811 correspondence to the tremor maps generated by Wech et al. (2009); Wells et al. (2017).  
812 Similarly, the maximum vertical displacements in Alaska encompass the tremor activ-  
813 ity mapped by Y. Ohta et al. (2006) and Peterson and Christensen (2009) (in addition  
814 to showing locations where additional tremor activity could be expected).

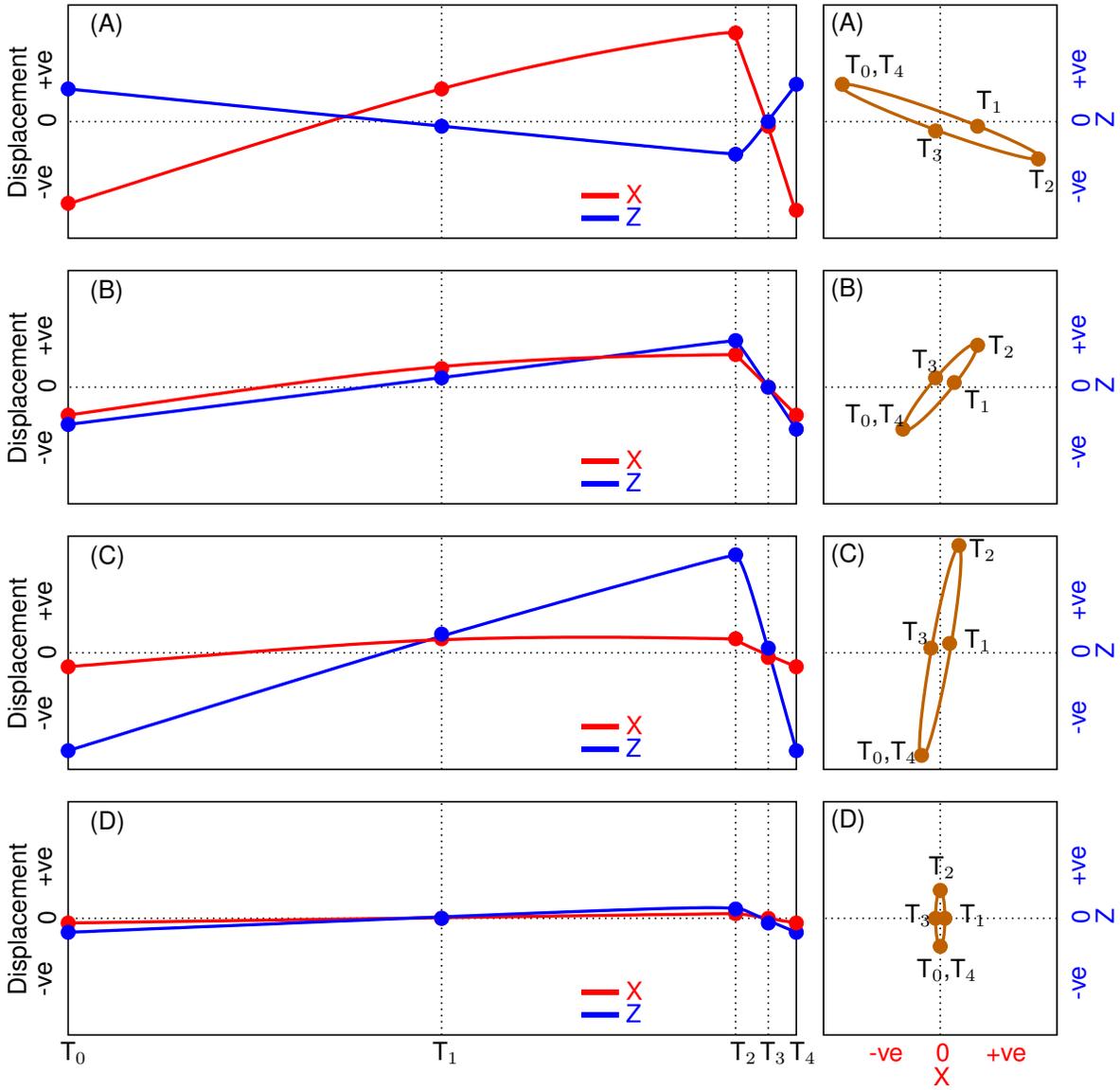
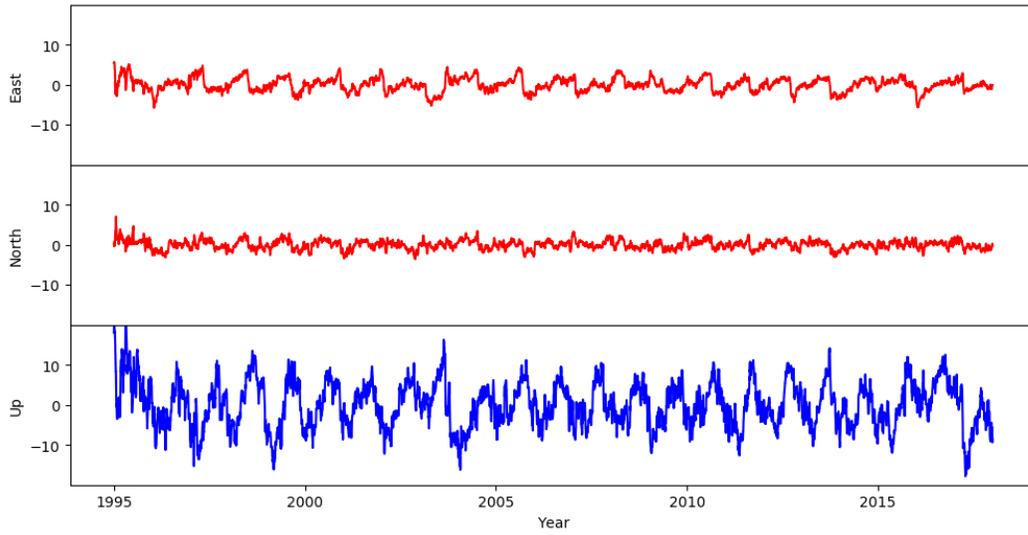
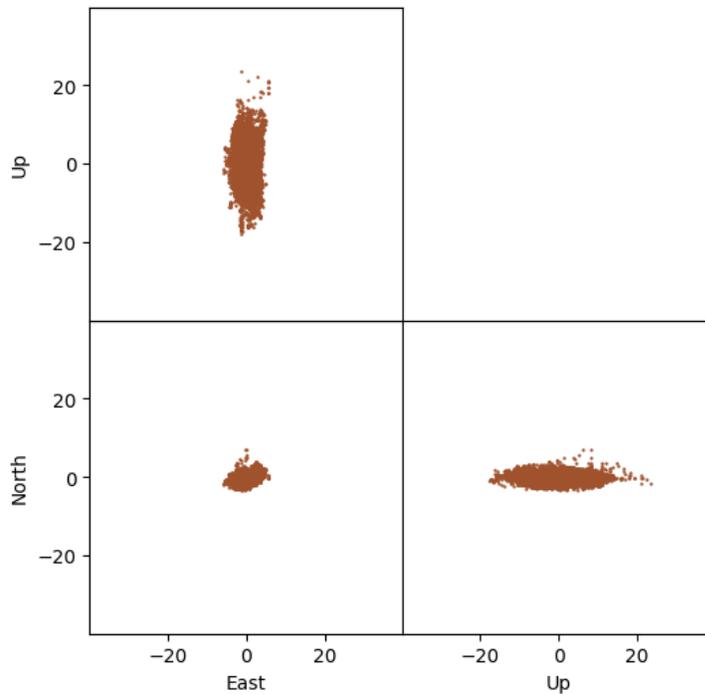


Figure A1: **Time-dependent detrended displacements (left column) and corresponding hodograms (right column) of points A through D (Figure 2) during a single cycle of Episodic Buckling and Collapse.** Horizontal displacement X is shown in red and vertical displacement Z in blue. The different phases of the subduction cycle are also denoted.

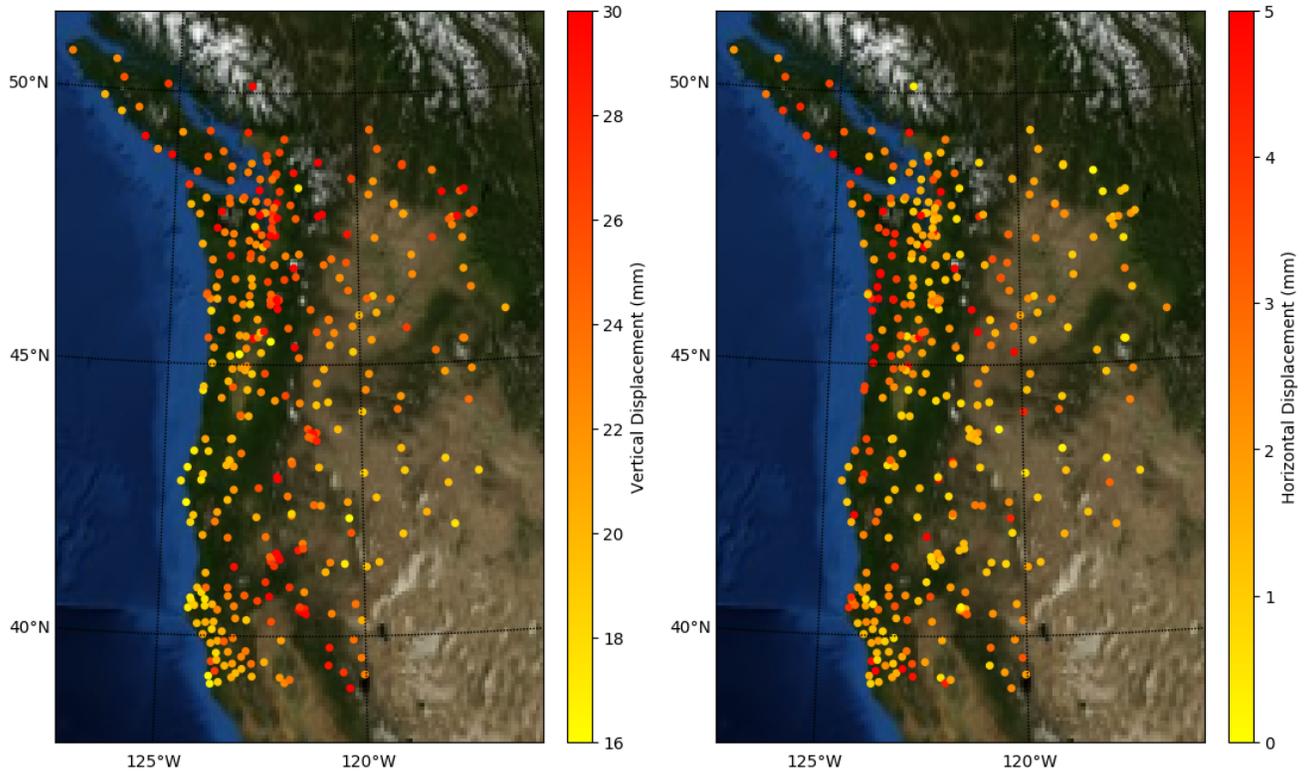


(a) GPS



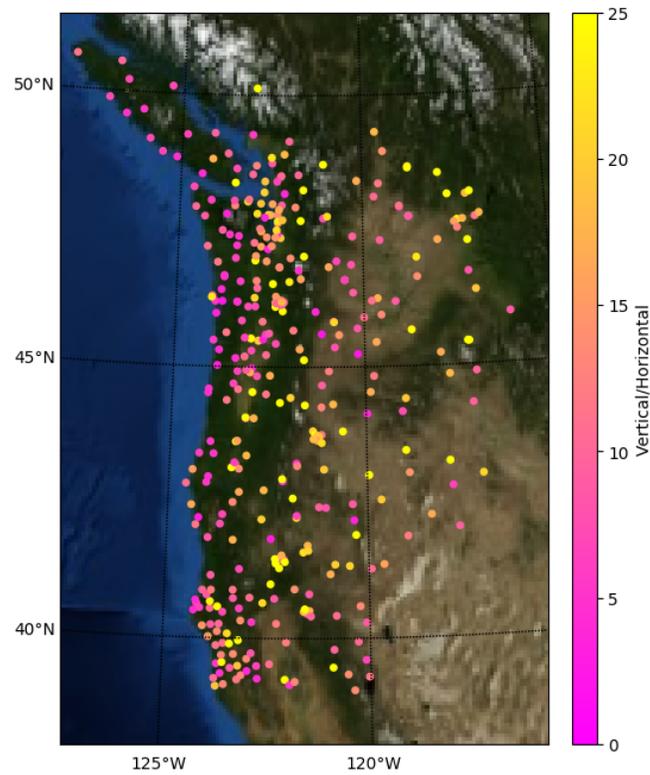
(b) Hodogram

Figure A2: East, North, and vertical components of GPS data and corresponding hodogram from the Albert Head GPS site on Vancouver Island in Victoria, British Columbia and corresponding hodogram on the right. All data have been detrended and filtered. The hodogram is displayed in the form of projections on the three orthogonal planes.



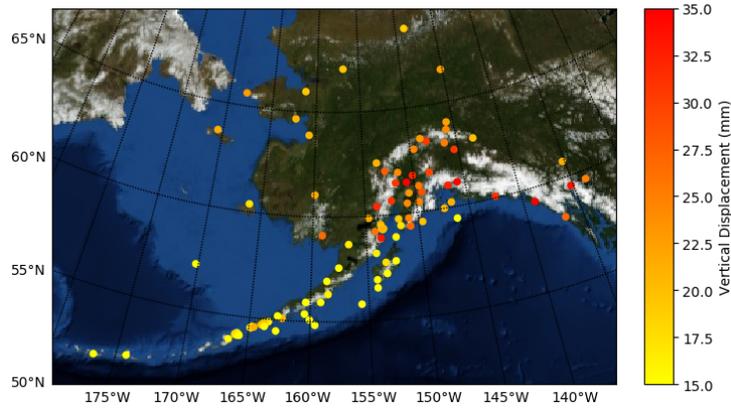
(a) Vertical displacement

(b) Horizontal displacement

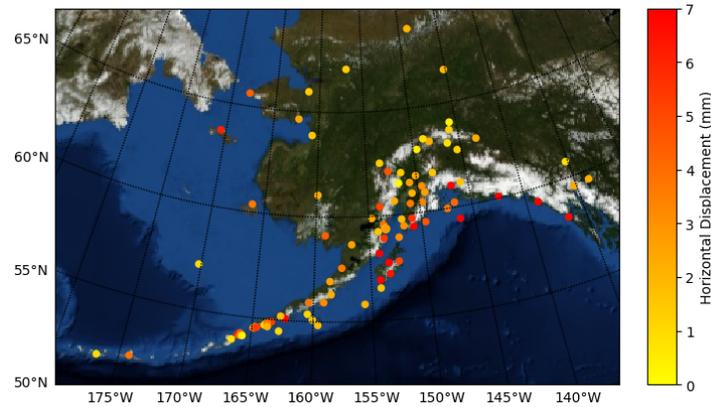


(c) Vertical-Horizontal Ratio

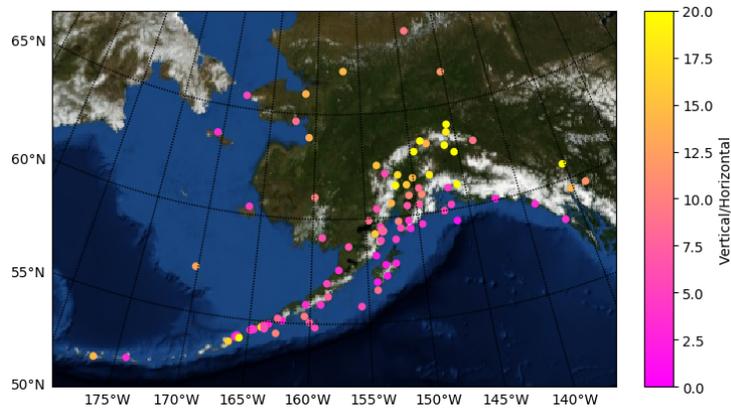
Figure A3: Measures of surface deformation in Cascadia subduction zone. **a**, Net vertical displacement and **b**, net horizontal displacement computed from GPS measurements, and **c**, their ratio. All color scales have been truncated to expose the patterns.



(a) Vertical displacement



(b) Horizontal displacement



(c) Vertical-Horizontal Ratio

Figure A4: **Measures of surface deformation in Alaska.** **a**, Net vertical displacement and **b**, net horizontal displacement computed from GPS measurements, and **c**, their ratio. All color scales have been truncated to expose the patterns.