



## Abstract

Operational numerical weather predictions (NWP) have been improved considerably over recent decades; however, they still occasionally generate large forecast errors referred to as “forecast busts”. This study investigates forecast busts over the Arctic between 2008 and 2019 using operational forecasts from five NWP centers. Forecasts with an anomaly correlation coefficient below the climatological 10<sup>th</sup> percentile, and a root-mean-square error above the 90<sup>th</sup> percentile, are regarded as “busts”. The percentage of forecast busts decreased from 2008 (7% to 13%) to 2012, and was between 2% and 6% for the period 2012–2019. The seasonal cycle of the forecast busts shows peaks in May and July–September. The forecast bust occurred more frequently when the initial pattern was the Greenland Blocking (GB) or Arctic Cyclone (AC) pattern rather than one of the other patterns. These results help users to be careful when they use the forecasts initialized on GB and AC patterns.

## Plain Language Summary

Recently, human activity in the Arctic region, such as trans-Arctic shipping, has increased due to the reduction in Arctic sea ice. Accurate weather forecasts will become increasingly important as the level of human activity in the Arctic continues to increase. However, despite the gradual improvement in NWP models over recent decades, operational forecasting still generates the occasional poor forecast, referred to as a “forecast bust.” This study investigates Arctic forecast busts using data from five operational forecasts.

Between 7% and 13% of Arctic forecasts were classified as “busts” in 2008, but the percentage decreased significantly from 2008 to 2012. Between 2% and 6% of forecasts were busts over the period 2012–2019. Regarding the seasonal cycle of busts in 2008–2019, Arctic forecast busts were most frequent in the May and July–September periods (~6% to 7%), but less frequent between December and March (~4%). The forecast busts occurred most frequently when the initial pattern used was the Greenland Blocking (GB) or Arctic Cyclone (AC) pattern. Therefore, users should access the forecast uncertainty using ensemble forecasts and compare the forecasts generated by different NWP centers when the initial weather pattern is GB or AC.

## 1 Introduction

Improvements in both our understanding of dynamical and physical processes, as well as in computational efficiency, have allowed numerical weather predictions (NWP) to improve significantly over recent decades (Bauer et al. 2019). The leading NWP centers across the globe now routinely provide high-resolution deterministic and low-resolution ensemble forecasts on medium-range timescales. However, NWP occasionally generate very poor forecasts despite the huge improvements in forecast skill. These occasional poor forecasts are referred to as “forecast busts” (Rodwell et al., 2013).

Forecast busts across Europe have been investigated in many previous studies. Rodwell et al. (2013) showed that the composite field from the verifying analysis of forecast busts across Europe shows blocking over Scandinavia. They also showed that the composite of the initial analysis for these busts showed the Rockies trough accompanied by high convective available potential energy (CAPE) over North America. Lillo and Parsons (2017) used empirical orthogonal function analysis to study the same forecast busts as Rodwell et al. (2013). They showed that the forecast busts occurred during large-scale pattern transition caused by amplification of Rossby waves. Grams et al. (2018) also showed the importance of moist

processes associated with the warm conveyor belt for the European forecast busts. Magnusson (2017) studied the error sources associated with three European forecast busts using ensemble sensitivity analysis and a relaxation experiment, and found that the error sources originated from the tropical Pacific, North America, and North Atlantic.

Recently, human activity in the Arctic region, such as trans-Arctic shipping, has increased due to the reduction in Arctic sea ice (Eguiluz et al., 2016; Melia et al., 2016). Accurate weather forecasts are becoming increasingly important as human activity continues to increase in the Arctic. Although the forecast skill over the Arctic associated with operational predictions has been increasing for the past 10 years (Jung and Matsueda, 2016), operational predictions occasionally generate very poor forecasts (9 and 10 in July in Fig. 1), as in the case of the European forecast busts. Yamagami et al. (2018a, b, 2019) showed that operational ensemble forecasts generally predict the position of extraordinary Arctic cyclones 2.5–4.5 days before they reach their mature stage. This suggests that such extraordinary Arctic cyclones could be one of the possible events that lead to the occurrence of Arctic forecast busts. Our forecast skill with respect to the Arctic atmosphere has a large influence on our ability to accurately forecast Arctic sea ice (Nakanowatari et al., 2018). In addition, the Arctic forecast skill influences the midlatitude forecast skill on medium- to extended-range timescales (Jung et al., 2014). Day et al. (2019) showed that the influence of the Arctic forecast on the mid-latitude forecast skill increased during periods affected by Scandinavian blocking. These previous studies indicate that Arctic forecast busts would significantly influence the forecasts of other climate systems and other regions.

In this study, we investigated the characteristics of forecast busts over the Arctic by using operational forecasts from major NWP centers around the world. We also investigated the relationship between forecast busts and weather patterns over the Arctic in summer.

## 2 Data and Methods

### 2.1 Forecast data

The operational forecast data used in this study are available from the TIGGE database (Swinbank et al., 2016) managed by the European Centre for Medium-range Weather Forecasts (ECMWF). We used ensemble forecast data from five NWP centers: the Canadian Meteorological Centre (CMC), ECMWF, the Japan Meteorological Agency (JMA), the US National Centers for Environmental Prediction (NCEP), and the UK Met Office (UKMO). These five NWP centers show higher performance than the other NWP centers available at the TIGGE database in the Northern Hemisphere (Matsueda and Tanaka, 2008; Swinbank et al., 2016) and over the Arctic (Jung and Matsueda, 2016). Ensemble forecasts initialized at 1200 UTC on every day from 1 January 2008 to 31 December 2019 were used in this study. Note that there are some missing data for each NWP centers (In particular, there are long missing periods in 2017 and 2018 for CMC and 2014 for UKMO). The forecast data had a grid spacing of  $2.5^\circ$  and a temporal resolution of 1 day.

### 2.2 Forecast skill and threshold of forecast bust

To detect the forecast busts, we used the uncentered anomaly correlation coefficient (ACC) and root-mean-square error (RMSE) of the latitude-weighted geopotential height at 500 hPa ( $Z500$ ) over the Arctic ( $\geq 65^\circ\text{N}$ ) as follows (Wilks 2019):

$$\text{ACC} = \frac{\sum_{i=1}^N (Z500_f - Z500_c)(Z500_a - Z500_c)}{\sqrt{\sum_{i=1}^N (Z500_f - Z500_c)^2} \sqrt{\sum_{i=1}^N (Z500_a - Z500_c)^2}},$$

$$\text{RMSE} = \sqrt{\frac{1}{N} \sum_{i=1}^N (Z500_f - Z500_a)^2},$$

where  $Z500_f$ ,  $Z500_a$ , and  $Z500_c$  are the predicted, analyzed, and climatological  $Z500$ , respectively, and  $N$  is the total number of grid points over the Arctic. We used the own-control analysis (an initial field of the control forecast) from each NWP center to calculate the ACC and RMSE in a bias-free manner. The climatological  $Z500$  was calculated using the ECMWF Reanalysis 5 (ERA5) data (Hersbach et al., 2020).

Rodwell et al. (2013) defined the threshold for forecast busts over Europe as a forecast with an ACC of less than 0.4 and an RMSE greater than 60 m at a lead time of 144 hours. However, the forecast skill differs among the NWP centers and in different regions. To obtain a subjective threshold for the Arctic forecast busts, we calculated the probability density functions (PDFs) of the ACC and RMSE in each month using control forecasts from each NWP center at a lead time of 144 hours over our analysis period from 2008–2019. Then, the climatological 10<sup>th</sup> percentile value of the ACC and 90<sup>th</sup> percentile value of the RMSE were retrieved from each PDF. When the control forecasts showed an ACC of less than the 10<sup>th</sup> percentile value of ACC and an RMSE greater than the 90<sup>th</sup> percentile value of RMSE for each month at a lead time of 144 hours, the forecasts were regarded as forecast busts.

### 3 Results

#### 3.1 ACC and RMSE distributions

The distribution of the ACC for Z500 in the Arctic at a lead time of 144 hours shows that 50% of the forecasts (25<sup>th</sup> to 75<sup>th</sup> percentile values, colored box in Fig. 2a) between 2008 and 2019 had ACC values of 0.6–0.9 in all months and for all NWP centers, except for CMC in May, June, and October, JMA in June and July, and NCEP in June. The ACC was typically highest in February (the average ACC was 0.73–0.81, white circle) and lowest in June (the average was 0.63–0.71). The 10<sup>th</sup>, 25<sup>th</sup>, 75<sup>th</sup>, and 90<sup>th</sup> percentiles, as well as the median ACC values also showed a similar seasonal cycle to the average. All of the NWP centers showed the largest standard deviation of ACC in May or June, except for UKMO whose standard deviation was the largest in October. CMC, ECMWF, JMA, and NCEP showed the second largest standard deviation in October. The standard deviation of ACC was the smallest in March for CMC, November for ECMWF and JMA, and December for NCEP.

The RMSE at lead times of 144 hours also shows a similar seasonal cycle to the ACC (Fig. 2b). The RMSE was highest in January (the average RMSE was 81.7–95.7 m) and lowest in August (the average was 66.6–74.2 m). The seasonal cycles of the 10<sup>th</sup>, 25<sup>th</sup>, 75<sup>th</sup>, and 90<sup>th</sup> percentiles, as well as the median values, were similar to that of the average RMSE for all NWP centers. The standard deviation of the RMSE was highest in December or January and lowest in June or August.

Over all, ECMWF showed the highest skill in all months, and CMC or JMA showed the lowest skill among the five NWP centers. The standard deviations of ACC and RMSE were smallest for ECMWF, indicating that the quality of the ECMWF forecast is more stable compared with the other centers.

As mentioned above, the bust threshold was the 10<sup>th</sup> percentile value of the ACC and 90<sup>th</sup> percentile value of the RMSE in each month for the individual NWP centers (Table S1). In July 2016, the ACCs for the ECMWF, NCEP and UKMO forecasts initialized on 10 were lower than the 10<sup>th</sup> percentile value (dotted lines in Fig. 1a), and at the same time, the RMSEs for these forecasts were higher than the 90<sup>th</sup> percentile value (Fig. 1b). On the other hand, the ACC for the JMA forecast initialized on 10 was lower than the 10<sup>th</sup> percentile value, but the RMSE for its forecasts was lower than the 90<sup>th</sup> percentile values. Thus, we regarded the ECMWF, NCEP, and UKMO forecasts as busts, but the CMC and JMA forecasts were not.

#### 3.2 Frequency of forecast busts

The proportion of forecasts that were busts over the Arctic was highest in 2008 for all NWP centers except NCEP (Fig. 3a). In 2008, about 13% of forecasts were busts for CMC and ECMWF, 10% for JMA and UKMO, and 7.5% for NCEP. The proportion of forecast busts decreased significantly from 2008 to 2012 for all NWP centers, falling to between 3% and 5% (i.e., ca. 10–18 days) in 2012. The decrease in forecast busts indicates the improvements in the

forecast systems (e.g., model resolution, assimilation systems, and boundary conditions). Rodwell et al. (2013) showed that the number of European busts had a local maximum in 2008, suggesting that the frequency of less predictable patterns for the operational ECMWF model was higher in 2008 than in the other years. As with the European flow patterns, the frequency of less predictable Arctic flow patterns might be higher in 2008 than in the other years. Although after 2012 the proportion of busts remained below 6% for all NWP centers, except for JMA and NCEP in 2015, the year for the local maximum differed among the NWP centers. ECMWF showed the highest percentage in 2013 and decreased gradually after 2014. JMA and NCEP showed the highest percentage in 2015, and JMA (NCEP) showed a higher percentage in 2017 (2014) as well. The percentage of busts for UKMO was a higher in 2014 and 2016. In contrast, all NWP centers recorded their lowest proportion of forecast busts in 2018, except for CMC. This lowest percentage of busts for 2018 might indicate the higher flow-dependent predictability over the Arctic than that for the other years. Another possible explanation of the low proportion in 2018 is the contribution of the Special Observing Periods Northern Hemisphere 1 (SOP-NH1: 1 February to 31 March) and 2 (SOP-NH2: 1 July to 30 September) that formed part of the Year of Polar Prediction (YOPP; Jung et al., 2016). About 2000 and 3000 extra radiosonde observations were conducted in the Arctic region during the SOP-NH1 and SOP-NH2 periods, respectively. These extra observations were assimilated into operational forecasts. Thus, the analysis uncertainties may have been reduced during these periods compared with other periods, resulting in the lowest proportion of busts.

The seasonal cycle of the proportion of forecast busts has two peaks (Fig. 3b). One is in May, and the other is in mid- to late-summer. Although all NWP centers show the peak in May clearly, the later peak differs among the NWP centers (July for ECMWF; August for CMC, JMA, and UKMO; and September for NCEP). At these peaks, the proportion of forecast busts was approximately 6% to 7% (ca. 21–26 days). The proportion of forecast busts was lowest in winter, with a value of around 4% (ca. 14 days). As a large number of the Arctic forecast busts occurred in summer, we focus on these summertime busts in the next subsection.

### 3.3 Frequency of summer forecast busts and its relationship to Arctic weather patterns

As with the annual bust proportion, the number of forecast busts in summer generally decreased from 2008 to 2019 for all NWP centers except JMA (Fig. 4). More than 10 busts occurred over the period 2008–2010 for all NWP centers. In particular, 15 ECMWF forecasts were busts in 2008 (Fig. 4b). After 2011, the number of forecast busts was at most six during summer for NCEP and UKMO (Fig. 4d and 4e), indicating that the summer busts have a similar interannual variability to the annual busts for these two centers (Fig. 3a). For the CMC and ECMWF, the number of busts remained relatively high until 2013, but the number decreased significantly after 2013 (Fig. 4a and 4b). In contrast to these NWP centers, the number of forecast busts for JMA was large, even in 2016 and 2017, indicating that the summer forecast busts contribute to the higher proportion of annual busts for JMA in 2016 and 2017 (Fig. 3a).

To investigate the atmospheric situation over the Arctic associated with these busts, we classified the Arctic atmospheric circulation into five weather patterns based on the *k*-means clustering method for twenty non-normalized principal components of Z500 anomaly over the

Arctic area, as used by Matsueda and Kyouda (2016) and Matsueda and Palmer (2018). The five weather patterns are called as the Arctic Dipole (AD), Greenland Blocking (GB), Arctic Cyclone (AC), Beaufort High (BH), and Summer North Atlantic Oscillation (SNAO; Fig. S1, Text S1). The classification revealed that forecasts initialized on any of these weather patterns can bust (Fig. 4a–e). However, the dominant initial weather pattern for busts differed among the years and NWP centers.

Over the period 2008–2013, a large number of forecasts initialized on the GB pattern were busts for all NWP centers. In particular, between 40% and 80% of the forecast busts initialized on the GB pattern. The number of forecast busts initialized on the GB pattern decreased until 2013 for all NWP centers. For JMA, NCEP, and UKMO at a lead time of 144 hours, the predicted BH pattern (left number at top-right corner in Fig. 4h–j) was a smaller number than the analyzed BH pattern in ERA5 (Fig. 4m–o). These results imply that the transition from initial GB pattern (Fig. S2b) to the BH pattern (Fig. S2d) after 144 hours is less frequent in the forecast than in analysis. In contrast, the predicted AD pattern (Fig. 4f–i) was a larger number than the analyzed AD pattern (Fig. 4k–n) for CMC, ECMWF, JMA, and NCEP. The transition from initial GB pattern to the AD pattern (Fig. S2a) would be more frequent in forecasts than in analysis. These suggest that the westward propagation of high pressure over Greenland is difficult to predict for the NWP models.

The number of forecast busts initialized on the GB pattern decreased significantly over the period 2014–2019. Since the frequency of analyzed GB pattern in summer over the period 2014–2019 (22.3%) was almost similar to that over the period 2008–2013 (23.7%), this reduction indicates that the westward propagation of high pressure could be predicted correctly after 2013 due to improvements in NWP systems. Although CMC and ECMWF show no dominant initial weather pattern associated with the busts over the period 2014–2019 (Fig. 4a and 4b), the dominant initial weather pattern for JMA, NCEP, and UKMO was AC (Fig. 4c–e). In particular, almost all of the JMA forecast busts in 2012, 2016, and 2017 were initialized on the AC pattern. The summers of 2012 and 2016 saw the development of extreme Arctic cyclones (Simmonds and Rudeva, 2012; Yamagami et al. 2017). Besides, the AC pattern was dominant in the summer of 2017 (51 days/91 days), and an extraordinary AC was detected on 10 in August 2017 using the threshold in Yamagami et al. (2018). For the extraordinary ACs, CMC and ECMWF showed a higher prediction skill for the central pressure than JMA, NCEP, and UKMO (Yamagami et al., 2019). In contrast, CMC showed the lowest prediction skill in the central position among the five NWP centers. These results suggest that busts associated with extraordinary ACs would have occurred due to the error for the AC deepening. However, during AC pattern, some busts were associated with the extraordinary ACs, the others were associated with ordinary ACs. These results indicate that the JMA, NCEP, and UKMO models have difficulties predicting the wandering, persistence, and decay of the ACs.

There are no dominant predicted and analyzed weather patterns at lead times of 144 hours (Fig. 4f–o). Unlike the Scandinavian blocking for forecast busts in Europe (Rodwell et al., 2013), the NWP models do not have a specific weather pattern over the Arctic in verifying analysis.

### 3.2 A descriptive heading about methods

## 4 Summary and Conclusions

Despite the gradual improvement in the NWP models over recent decades, operational forecasts continue to occasionally generate very poor forecasts, which are referred to as “forecast busts.” This study investigated the characteristics of the Arctic forecast busts using operational forecasts from five leading NWP centers. To define the threshold for Arctic forecast busts, we assessed the forecast skill of the operational forecasts from each month over the period 2008–2019. The ACC (RMSE) over the Arctic was highest in February (January) and lowest in June (August) for all NWP centers. We used the 10<sup>th</sup> percentile of the ACC and 90<sup>th</sup> percentile of the RMSE from each NWP center for each month as the subjective threshold for forecast busts over the Arctic.

Considering the proportion of forecasts in each year that were busts, 7% to 13% were busts in 2008, but the proportion of busts then decreased significantly from 2008 to 2012 for all NWP centers. The proportion of forecast busts was between 2% and 6% for all NWP centers from 2013 to 2019, but the year of local maximum differed among the NWP centers. The monthly variability of forecast busts showed that the proportion of busts increased in the May and July–September periods (~7%), but decreased in December–March (~4%).

To investigate the relationship between forecast busts and atmospheric circulation in summer, we classified the Arctic atmospheric circulation into five patterns. The five atmospheric patterns were Arctic Dipole (AD), Greenland Blocking (GB), Arctic Cyclone (AC), Beaufort High (BH), and Summer NAO (SNAO). The dominant initial weather pattern associated with forecast busts was GB between 2008 and 2013 for all NWP centers. For the JMA, NCEP, and UKMO forecast busts, the AC pattern also shows a higher proportion. Although the forecast busts initialized on the GB pattern decreased after 2013 for all NWP centers, the forecast busts initialized on the AC pattern were still dominant for JMA, NCEP, and UKMO. In contrast, the CMC and ECMWF forecast busts did not show a specific initial weather pattern in recent years. The Arctic forecast busts were not associated with specific weather patterns at a lead time of 144 hours, suggesting that the summer busts presumably occurred associated with the difference in the position of synoptic systems (e.g., difference in direction of ACs’ wandering).

The European forecast busts for the ECMWF Integrated Forecasting System occurred during Scandinavian blocking episodes (Rodwell et al., 2013). The source of the forecast errors associated with the European forecast busts were over North America and the Pacific equator (Magnusson 2017). Over the Arctic, the forecast busts were associated with the initial GB and AC patterns. For the ECMWF bust initialized on 10 July 2016 (Fig. 1), the initial weather pattern was the GB pattern and persisted up to a lead time of 96 hours. The GB pattern changed to the AD pattern at a lead time of 120 hours, and the AD pattern persisted up to a lead time of 144 hours. The comparison between higher- and lower-skill five members showed the large positive and negative differences across the polar vortex at a lead time of 144 hours (Fig. S2g), and its source was the initial difference around the polar vortex (Fig. S2a). Besides, the spread of the control analysis among the five NWP centers in summer (Fig. S3a) was large over the Pacific side of the Arctic Ocean and Greenland, as with in winter (Bauer et al., 2015). The analysis spread classified by the weather patterns was large around the polar vortex for all patterns and over Greenland for the GB and SNAO patterns. These areas are one of the possible sources of the initial errors for the Arctic forecast busts. The observations over the Arctic region have



potential impacts on improvements of forecasts over the Arctic (Yamazaki et al. 2015; Lawrence et al. 2019). This study also supports the impact of the increase in Arctic observation conducted by YOPP SOP-NH1 and 2 on the operational global forecasts. Therefore, the additional observations on the larger spread area for each pattern could reduce the Arctic forecast busts.

This study suggests that users should access the forecast uncertainty using ensemble forecasts and the differences in forecasts among the NWP centers when forecasts are initialized on the GB and AC patterns, especially on AC pattern in recent years. For the European forecast busts, Grams et al. (2018) showed that both moist and dry processes associated with the warm conveyor belt contribute to the large errors. Parsons et al. (2019) also showed that the mesoscale convective systems over North America led to the European forecast busts. In addition, Day et al. (2019) showed that the deterioration of the Arctic forecast reduces forecast skill in the mid-latitudes during Greenland blocking episodes. Further studies of the detailed processes associated with error growth in Arctic forecast busts and the impact of Arctic forecast busts on mid-latitude forecast skill will be needed in the future.

## Acknowledgments

The author thanks to the ECMWF for providing ERA5 (<https://doi.org/10.24381/cds.bd0915c6>) and TIGGE (<https://apps.ecmwf.int/datasets/data/tigge/>) datasets. This study was supported by the Japan Society for the Promotion of Science (JSPS) Grant-in-Aid for Research Activity Start-up, Number 19K23454.

## References

- Bauer, P., Thorpe, A., & Brunet, G. (2015). The quiet revolution of numerical weather prediction. *Nature*, 525, 47–55. <https://doi.org/10.1038/nature14956>
- Bauer, P., Magnusson, L., Thépaut, J.-N., & Hamill, T. M. (2016). Aspects of ECMWF model performance in polar areas. *Quarterly Journal of the Royal Meteorological Society*, 142, 583–596. <https://doi.org/10.1002/qj.2449>
- Day, J. J., Sandu, I., Magnusson, L., Rodwell, M. J., Lawrence, H., Bormann, N., & Jung, T. (2019). Increased Arctic influence on the midlatitude flow during Scandinavian Blocking episodes. *Quarterly Journal of the Royal Meteorological Society*, 145, 3846–3862. <https://doi.org/10.1002/qj.3673>
- Eguíluz, V. M., Fernández-Gracia, J., Irigoien, X., & Duarte, C. M. (2016). A quantitative assessment of Arctic shipping in 2010–2014. *Nature Scientific Reports*, 6, 30682. <https://doi.org/doi:10.1038/srep30682>
- Grams, C. M., Magnusson, L., & Madonna, E. (2018). An atmospheric dynamics perspective on the amplification and propagation of forecast error in numerical weather prediction models: A case study. *Quarterly Journal of the Royal Meteorological Society*, 144, 2577–2591. <https://doi.org/10.1002/qj.3353>
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., et al. (2020). The ERA5 Global Reanalysis. *Quarterly Journal of the Royal Meteorological Society*, 146, 1999–2049. <https://doi.org/10.1002/qj.3803>
- Jung, T., & Matsueda, M. (2016). Verification of global numerical weather forecasting systems in polar regions using TIGGE data. *Quarterly Journal of the Royal Meteorological Society*, 142, 574–582. <https://doi.org/10.1002/qj.2437>

- Jung, T., Kasper, M. A., Semmler, T., & Serrar, S. (2014). Arctic influence on subseasonal midlatitude prediction. *Geophysical Research Letters*, 41, 3676–3680. <https://doi.org/10.1002/2014GL059961>
- Jung, T., Gordon, N. D., Bauer, P., Bromwich, D. H., Chevallier, M., Day, J. J., et al. (2016). Advancing polar prediction capabilities on daily to seasonal time scales. *Bulletin of the American Meteorological Society*, 97, 1631–1647. <https://doi.org/10.1175/BAMS-D-14-00246.1>
- Lawrence, H., Bormann, N., Sandu, I., Day, J., Farnan, J., & Bauer, P. (2019). Use and impact of Arctic observations in the ECMWF Numerical Weather Prediction system. *Quarterly Journal of the Royal Meteorological Society*, 145, 3432–3454. <https://doi.org/10.1002/qj.3628>
- Lillo, S. P., & Parsons, D. B. (2017). Investigating the dynamics of error growth in ECMWF medium-range forecast busts. *Quarterly Journal of the Royal Meteorological Society*, 143, 1211–1226. <https://doi.org/10.1002/qj.2938>
- Magnusson, L. (2017). Diagnostic methods for understanding the origin of forecast errors. *Quarterly Journal of the Royal Meteorological Society*, 143, 2129–2142. <https://doi.org/10.1002/qj.3072>
- Matsueda, M., & Kyouda, M. (2016). Wintertime East Asian flow patterns and their predictability on medium-range timescales. *SOLA*, 12, 121–126. <https://doi.org/10.2151/sola.2016-027>
- Matsueda, M., & Palmer, T. N. (2018). Estimates of flow-dependent predictability of wintertime Euro-Atlantic weather regimes in medium-range forecasts. *Quarterly Journal of the Royal Meteorological Society*, 144, 1012–1027. <https://doi.org/10.1002/qj.3265>
- Matsueda, M., & Tanaka, H. L. (2008). Can MCGE outperform the ECMWF ensemble? *SOLA*, 4, 77–80. <https://doi.org/10.2151/sola.2008-020>
- Melia, N., Haines, K., & Hawkins, E. (2016). Sea ice decline and 21st century trans-Arctic shipping routes. *Geophysical Research Letters*, 43, 9720–9728. <https://doi.org/10.1002/2016GL069315>
- Nakanowatari, T., Inoue, J., Sato, K., Bertino, L., Xie, J., Matsueda, M., et al. (2018). Medium-range predictability of early summer sea ice thickness distribution in the East Siberian Sea based on the TOPAZ4 ice-ocean data assimilation system. *The Cryosphere*, 12, 2005–2020. <https://doi.org/10.5194/tc-12-2005-2018>
- Parsons, D. B., Lillo, S. P., Rattray, C. P., Bechtold, P., Rodwell, M. J., & Bruce, C. M. (2019). The role of continental mesoscale convective systems in forecast busts within global weather prediction systems. *Atmosphere*, 10, 1–32. <https://doi.org/10.3390/atmos10110681>
- Rodwell, M. J., Magnusson, L., Bauer, P., Bechtold, P., Bonavita, M., Cardinali, C., et al. (2013). Characteristics of occasional poor medium-range weather forecasts for Europe. *Bulletin of the American Meteorological Society*, 94, 1393–1405. <https://doi.org/10.1175/BAMS-D-12-00099.1>
- Simmonds, I., & Rudeva, I. (2012). The great Arctic cyclone of August 2012. *Geophysical Research Letters*, 39, L23709. <https://doi.org/10.1029/2012GL054259>
- Swinbank, R., Kyouda, M., Buchanan, P., Froude, L., Hamill, T.M., Hewson, T.D., et al. (2016). The TIGGE Project and Its Achievements. *Bulletin of the American Meteorological Society*, 97, 49–67. <https://doi.org/10.1175/BAMS-D-13-00191.1>
- Wilks, D. S. (2019). Statistical methods in the atmospheric sciences (4th edition). Oxford: Elsevier Academic Press.

- Yamagami, A., Matsueda, M., & Tanaka, H. L. (2017). Extreme Arctic Cyclone in August 2016. *Atmospheric Science Letters*, 18, 307–314, <https://doi.org/10.1002/asl.757>
- Yamagami, A., Matsueda, M., & Tanaka, H. L. (2018). Predictability of the 2012 great Arctic cyclone on medium-range timescales. *Polar Science*, 15, 13–23. <https://doi.org/10.1016/j.polar.2018.01.002>
- Yamagami, A., Matsueda, M., & Tanaka, H. L. (2018b). Medium-range forecast skill for extraordinary Arctic cyclones in summer of 2008–2016. *Geophysical Research Letters*, 45, 4429–4437. <https://doi.org/10.1029/2018GL077278>
- Yamagami, A., Matsueda, M., & Tanaka, H. L. (2019). Skill of medium-range reforecast for summertime extraordinary Arctic Cyclones in 1986–2016. *Polar Science*, 20, 107–116. <https://doi.org/10.1016/j.polar.2019.02.003>
- Yamazaki, A., J. Inoue, K. Dethloff, M. Maturilli, and G. König-Langlo (2015). Impact of radiosonde observations on forecasting summertime Arctic cyclone formation. *Journal of Geophysical Research: Atmospheres*, 120, <https://doi.org/10.1002/2014JD022925>

**Figure 1.** (a) ACC and (b) RMSE of geopotential height at 500 hPa over the Arctic area ( $\geq 65^\circ\text{N}$ ) at a lead time of 144 hours for the control forecasts from CMC (yellow), ECMWF (blue), JMA (red), NCEP (green), and UKMO (purple) initialized from 1 to 21 July 2016. The colored broken lines denote the climatological 10<sup>th</sup> percentile ACC and 90<sup>th</sup> percentile RMSE values for each NWP center.

**Figure 2.** (a) ACC and (b) RMSE distributions of geopotential height at 500 hPa at a lead time of 144 hours over the Arctic area ( $\geq 65^\circ\text{N}$ ) for the CMC (yellow), ECMWF (blue), JMA (red), NCEP (green), and UKMO (purple) control forecasts initialized for each month. Box limits indicate the 50<sup>th</sup> percentile of the scores, and the horizontal bar in the box shows the median value. The vertical lines extending from the box show the range of scores from 10<sup>th</sup> to 90<sup>th</sup> percentile. The white circles show the average values.

**Figure 3.** Percentage of forecast busts over the Arctic area in (a) each year and (b) each month for the CMC (yellow), ECMWF (blue), JMA (red), NCEP (green), and UKMO (purple) control forecasts.

**Figure 4.** Frequency of the Arctic Dipole (AD, purple), Greenland Blocking (GB, green), Arctic Cyclone (AC, red), Beaufort High (BH, blue), and Summer NAO (SNAO, orange) weather patterns for busts of control forecasts at (a–e) the initial time and (f–j) a lead time of 144 hours for (a, f) CMC, (b, g) ECMWF, (c, h) JMA, (d, i) NCEP, and (e, j) UKMO in summer (June–August) over the period 2008–2019. (k–o) Frequency of analyzed weather patterns at a lead time of 144 hours for each NWP center calculated using the ERA5 data. The numbers of busts for total (black) and each regimes (colored) over the period 2008–2013 (left) and 2014–2019 (right) are given at the top-right corner of each frame.

Figure 1.

# Forecast skill

(control run, Arctic, 2016, 12UTC, FT+144hr)

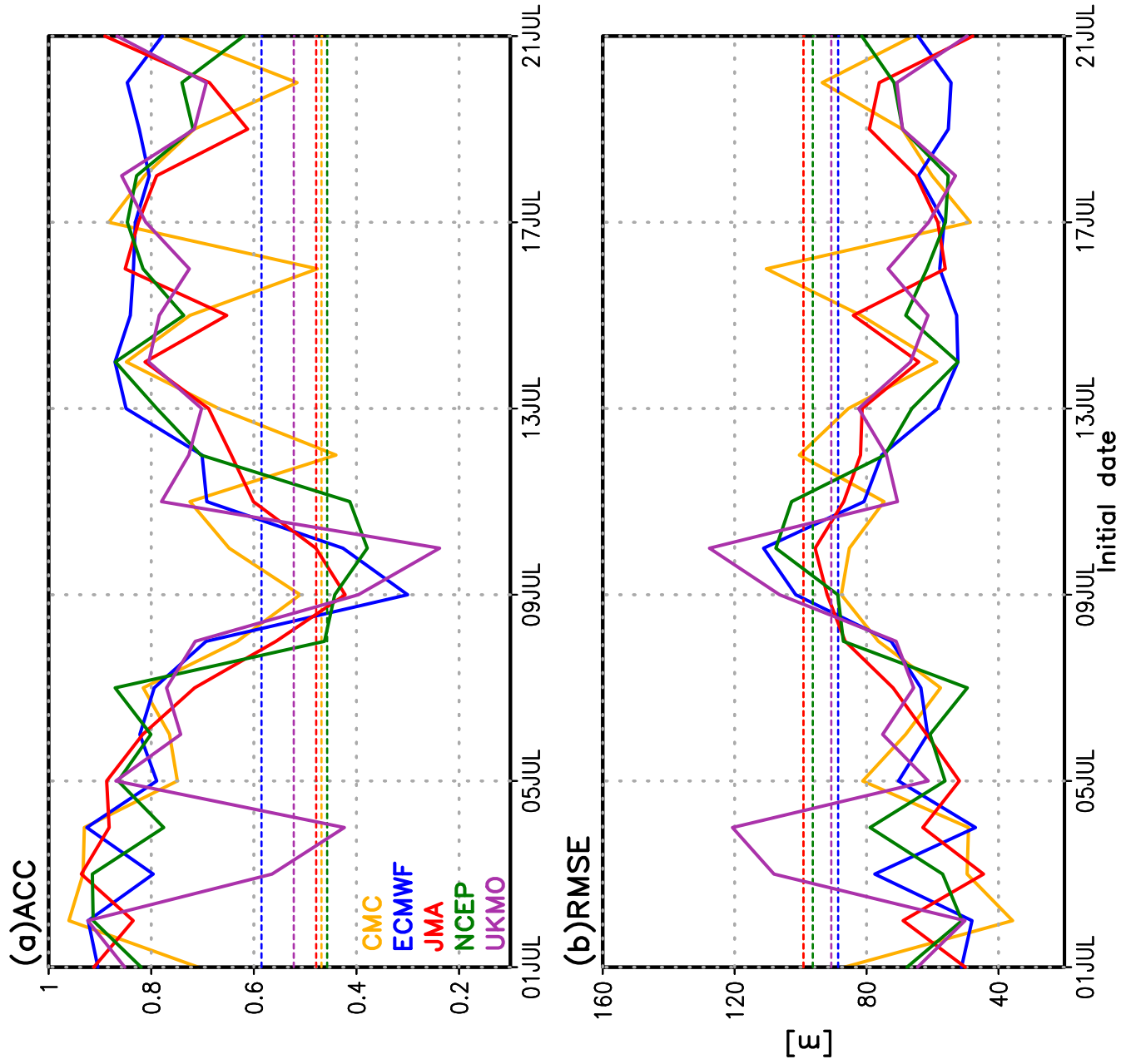


Figure 2.

Score threshold for forecast bust over Arctic  
(FT+144 hr)

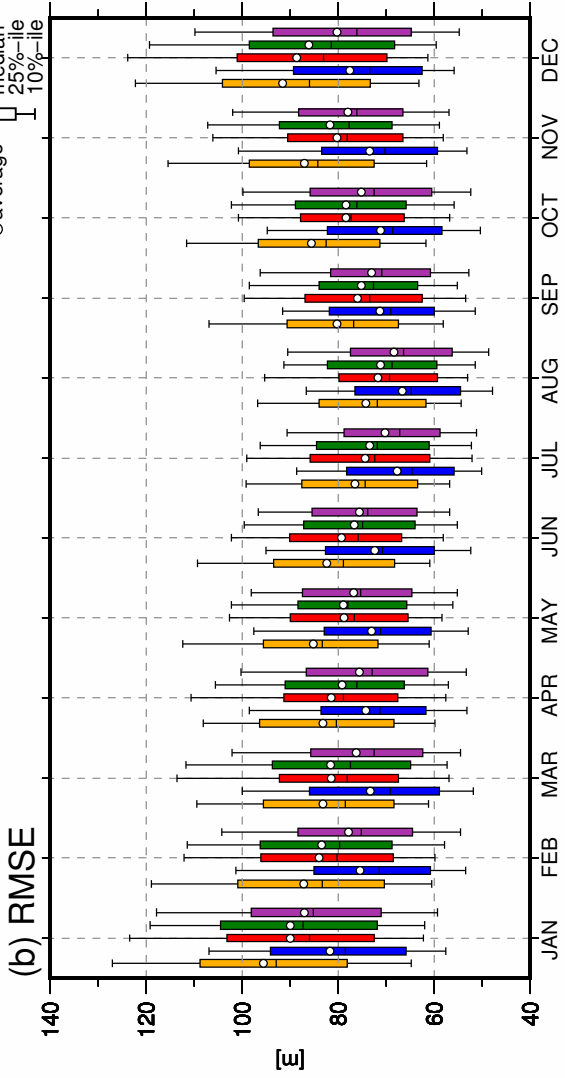
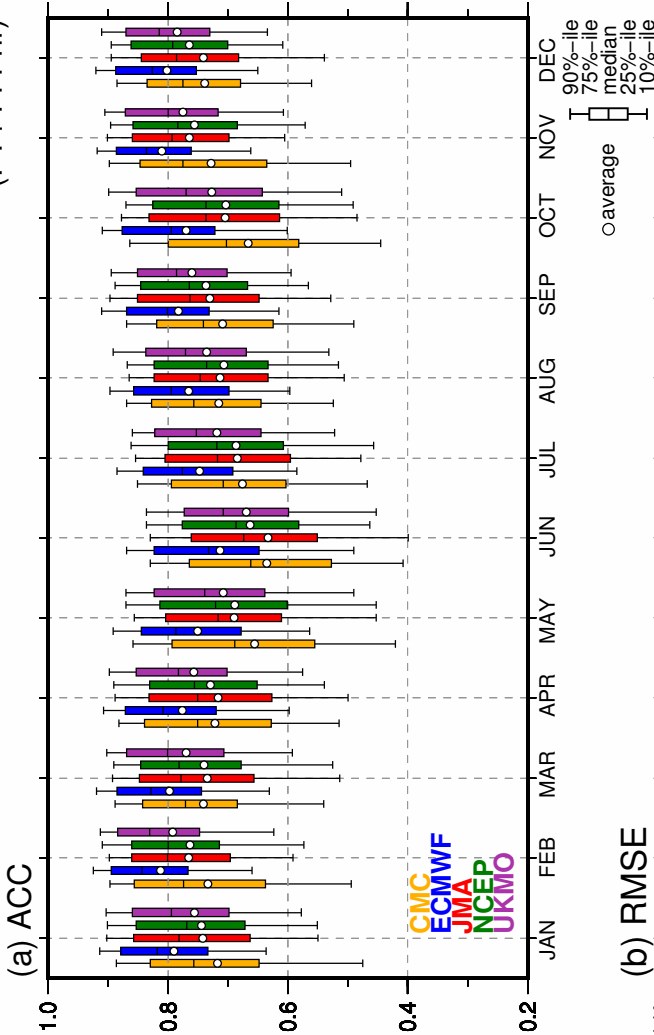


Figure 3.



# Ratio of forecast bust over Arctic

(90 %-ile, 144 hr)

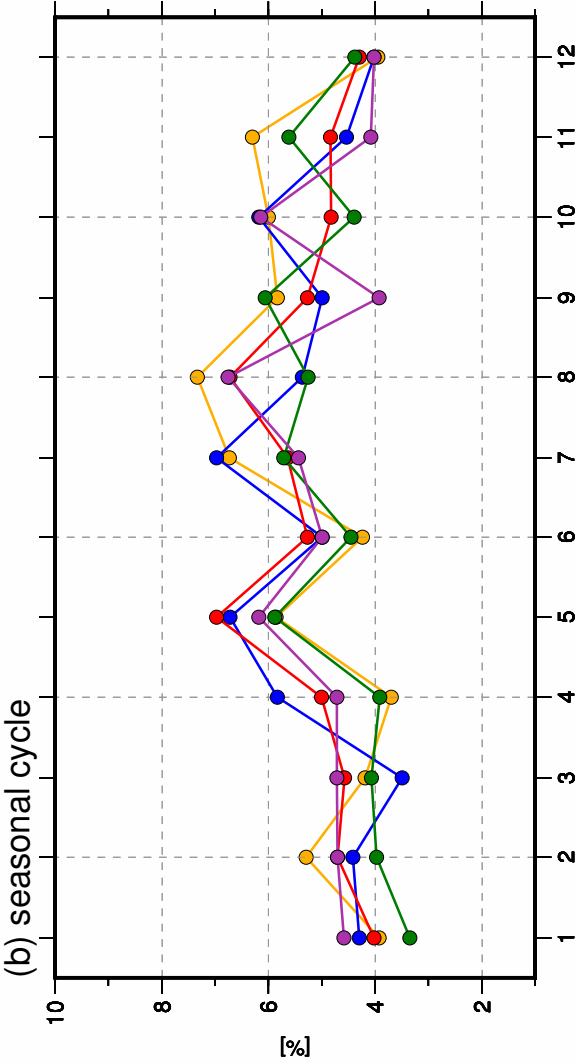
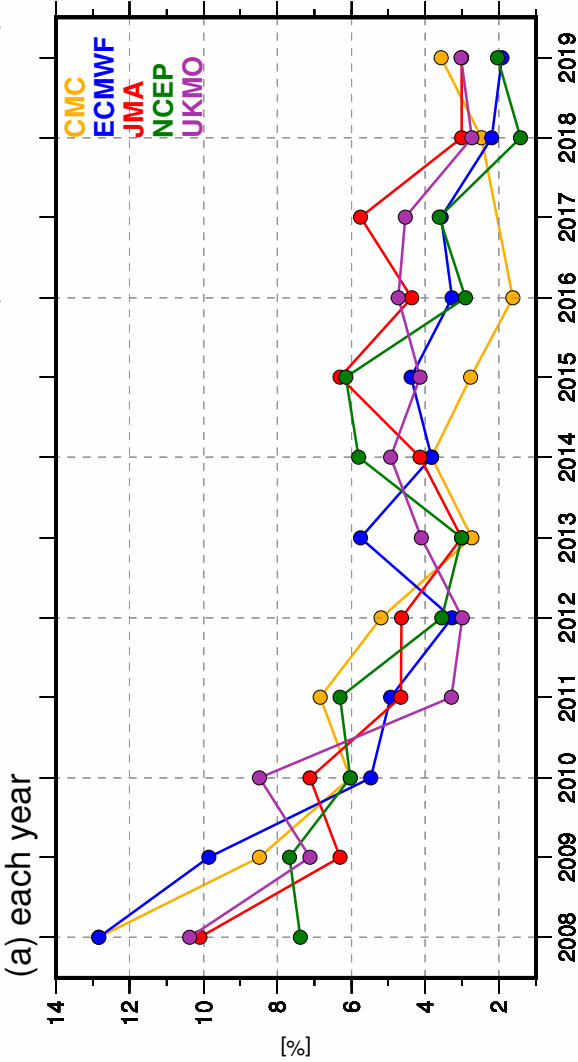


Figure 4.

## Frequency of Arctic weather patterns for busts (JJA)

