



## 25 **ABSTRACT**

26 The recent development of the TOUGH3 code allows for a faster and more reliable fluid flow  
27 simulator. At the same time, new versions of FLAC3D are released periodically, allowing for  
28 new features and faster execution. In this paper, we present the first implementation of the  
29 coupling between TOUGH3 and FLAC3Dv6/7, maintaining parallel computing capabilities  
30 for the coupled fluid flow and geomechanical codes. We compare the newly developed  
31 version with analytical solutions and with the previous approach, and provide some  
32 performance analysis on different meshes and varying the number of running processors.  
33 Finally, we present two case studies related to fault reactivation during CO<sub>2</sub> sequestration and  
34 nuclear waste disposal. The use of parallel computing allows for meshes with a larger number  
35 of elements, and hence more detailed understanding of thermo-hydro-mechanical processes  
36 occurring at depth.

37  
38

39 **Keywords:** THM processes; coupled simulator; parallel computing; TOUGH3; FLAC3D

40

## 41 **1. INTRODUCTION**

42 The current development of georesources exploitation strongly relies on numerical simulation  
43 of the processes occurring at depth. Understanding of the coupled thermo-hydro-mechanical  
44 processes is essential to assess properly the changes in system conditions as well as to study  
45 the risks associated with the underground exploitation (e.g. loss of circulation; caprock  
46 failure; induced seismicity). Model developments and their applications constitute a huge  
47 step towards understanding coupled processes. Several numerical simulators are already  
48 available in the literature for the study of coupled processes at various levels of complexity.  
49 Some models allow for all Thermo-Hydro-Mechanical-Chemical (THMC) couplings.

50 Examples are: TOUGHREACT-FLAC3D (Taron and Elsworth, 2009; Rutqvist et al., 2014;  
51 Rutqvist, 2017), OpenGeoSys (Kolditz et al., 2012), Dumux (Flemisch et al., 2011),  
52 COMSOL Multiphysics (COMSOL, 2020). In addition, other simulators have been applied to  
53 study partial processes coupling, THM or THC: e.g. Sierra Mechanics (Newell et al., 2017),  
54 3DEC (Itasca, 2016), CODE-BRIGHT (Olivella et al., 1996), CSMP++ (Yapparova et al.,  
55 2017; Salimzadeh et al., 2018), PFLOTRAN (Lichtner et al., 2017).

56 In numerical modelling, the governing equations (conservation laws of mass, momentum and  
57 energy) are solved considering the relationship among processes (e.g. coupling of two or  
58 more processes), and completed with constitutive laws, initial conditions and boundary  
59 conditions. One factor determining the computational effort is the number of simulated  
60 THMC coupled processes. Another factor is the numerical scheme. In the literature, the term  
61 *monolithic* refers to a scheme in which the physical equations for multiple processes are  
62 solved simultaneously, which may be computationally expensive. More loose couplings exist,  
63 such as one-way (i.e. a given process influences another, but not vice versa) or two-way  
64 sequential (i.e. the different processes are considered in sequence). Such schemes are less  
65 computationally intensive and may refer to the same simulator, but often the integration of  
66 different codes is used to take advantage of specialized codes and to increase the types of  
67 simulated processes (Kulik et al., 2012; Rutqvist, 2017).

68 Given the complexity of the coupled processes, verification of the numerical approach is  
69 often an issue. Analytical solutions are only available for very simplified processes (e.g. only  
70 for fully saturated medium), and observations from lab and *in situ* experiments involve  
71 significant uncertainties. Benchmarking activities involving code-to-code comparison and  
72 validation against analytical solutions and experimental data are often in play for developing  
73 numerical models (Blanco-Martín et al., 2015; Garitte et al., 2017; Rutenberg et al., 2018;  
74 Birkholzer et al., 2019). The TOUGH family of codes are commonly applied to model the

75 coupling of fluid flow and heat transport in geological media (Pruess et al., 2012; Jung et al.,  
76 2017), and have been extended to consider coupling to geomechanical processes. In  
77 particular, several TOUGH-based geomechanical codes have been developed to solve THM  
78 problems (Rutqvist, 2017), among which TOUGH-FLAC is the most widely used, with  
79 recent applications featuring inverse modeling (Blanco-Martín et al., 2016; Rinaldi et al.,  
80 2017) and finite strain deformation (Blanco-Martín et al., 2017). Since its initial development  
81 in the late 1990s (Rutqvist et al., 2002), TOUGH-FLAC has been applied to study  
82 geomechanical aspects of CO<sub>2</sub> sequestration, nuclear waste disposal, enhanced geothermal  
83 systems, underground gas storage and compressed air energy storage, gas production from  
84 hydrate-bearing formations, induced seismicity, as well as for the implementation and the  
85 study of constitutive equations (Rutqvist, 2017 and references therein).

86 The most common version of TOUGH-FLAC accounts for the two-way sequential coupling  
87 of TOUGH2 (Pruess et al., 2012) for the simulation of non-isothermal, multi-phase and  
88 multi-component fluid flow with upgraded versions of FLAC3D (e.g. ver.4/5, Itasca, 2011)  
89 for solving the mechanical equilibrium. The equations for fluid flow and geomechanics are  
90 solved sequentially, and the approach is unconditionally stable, using the fixed-stress split  
91 sequential scheme: the flow sub-problem is solved first with a fixed total stress field, which is  
92 then modified in the subsequent geomechanics sub-problem by using modified variables from  
93 the flow step (Kim et al., 2011).

94 Despite the wide use of TOUGH-FLAC coupled simulations in the literature, applications are  
95 often limited to a relatively small computational domain, with a number of elements usually  
96 smaller than 50,000. In this work, we moved one step forward by coupling for the first time  
97 the newly developed TOUGH3 (Jung et al., 2017) with versions 6.0 and 7.0 of FLAC3D  
98 (Itasca, 2017), hence implementing a parallelized version of the well-known coupled  
99 simulator. The coupled simulator integrates all the new functionalities of TOUGH3,

100 including the use of PETSc solvers, together with the improved solver performance in  
101 FLAC3Dv6/7 as well as the possibility of using Python scripting compared to the FISH  
102 programming embedded in previous versions of FLAC3D. After verifying the correctness of  
103 the approach comparing the simulation results with an analytical solution and with the  
104 previous version, we evaluate the performance of the newly developed approach. Finally, we  
105 present results of two case studies, aimed at understanding the potential for fault reactivation  
106 during CO<sub>2</sub> sequestration and the evolution of stress and strain during nuclear waste disposal  
107 in a deep geological repository.

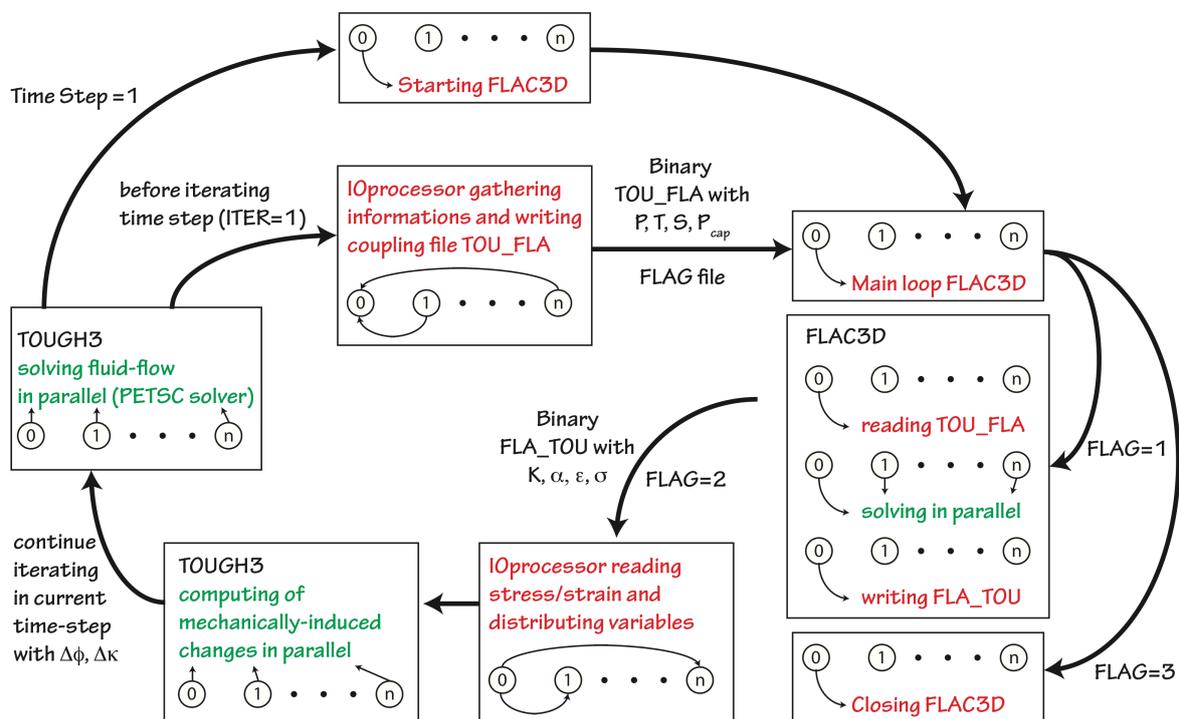
## 108 **2. COUPLING APPROACH**

109 The main modifications of the approach compared to the previous versions are:

- 110 1- Use of TOUGH3 with parallel computing and use of FLAC3Dv6/7, both allowing for  
111 faster calculation.
- 112 2- Use of binary instead of ASCII exchange files. This is particularly important when  
113 dealing with computational domains containing a large number of elements.
- 114 3- Use of a flag system to avoid saving and restoring the FLAC3D model state to a  
115 binary file at each time step. In the new approach, the FLAC3D model state is only  
116 loaded at the beginning of the simulation, then paused during TOUGH3 execution,  
117 thereby avoiding overhead caused by restoring/saving the mechanical state at each  
118 time step.
- 119 4- Use of Python routines to read/write coupling files in FLAC3D. Through the use of  
120 NumPy (Harris et al., 2020), this can produce up to 34× faster execution (10× in  
121 average) in variable allocation compared to the previously used FISH scripting  
122 (Itasca, 2017).

123 One of the limitations of FLAC3D is that it runs only on Windows-based machines. Then, to  
124 make use of the MPI parallelization provided in TOUGH3, we compiled the code in the

125 Cygwin environment or Windows Subsystem for Linux (Windows 10), which also allows  
 126 calls to Windows executables (i.e. FLAC3D batch). All the simulations presented here were  
 127 carried out on a workstation with two 6-core CPUs Intel Xeon E5-2643 (24 threads) at 3.40  
 128 GHz and equipped with 64 GB RAM.  
 129 This hardware configuration allows using up to 12 processors. Each of the two codes,  
 130 evaluated separately, provides a speed increase up to  $6\times$  for a quite large computational  
 131 domain (800,000 elements) on this machine when employing the maximum number of cores  
 132 (supplementary material, Fig. S1).  
 133 The general coupling approach between any TOUGH-code and FLAC3D is based on file  
 134 exchange to share variables and/or properties. In MPI codes, and in particular for TOUGH3,  
 135 a processor is designed as “IOProcessor” and takes care of all the input/output functionalities  
 136 of the code. Figure 1 describes the coupling approach for TOUGH3 and FLAC3D.



137  
 138 *Figure 1.* Coupling approach between TOUGH3 and FLAC3D for each time step. Green parts are executed in  
 139 parallel, while red parts are executed in serial.  $P$ ,  $T$ ,  $S$ ,  $P_{cap}$  are pore pressure, temperature, saturation, and  
 140 capillary pressure, respectively.  $K$ ,  $\alpha$ ,  $\epsilon$ ,  $\sigma$ , refer to bulk modulus, Biot’s coefficient, strain, and stress, while  $\Delta\phi$   
 141 and  $\Delta\kappa$  stand for porosity and permeability changes.

142 As in previous versions of TOUGH-FLAC, TOUGH3 is the main code, which runs the  
143 simulation through time and modifies the time step according to the stability of the fluid flow  
144 sub-problem. At each time step, FLAC3D computes the mechanical equilibrium to a  
145 predefined convergence threshold. As a rule of thumb, the approach is quite stable for poro-  
146 elastic problems, but it may require an appropriate choice of time stepping when including  
147 time-dependent deformation (e.g. creep).

148 For each time step, before starting the iterations for solving the fluid flow sub-problem,  
149 TOUGH3 invokes a subroutine to gather the arrays from all the MPI processes. Such arrays  
150 (pressure, temperature, saturation and capillary pressure) are written to file TOU\_FLA by the  
151 IOProcessor, together with a flag (1). The IOProcessor performs this entire stage in serial,  
152 while all the other  $n$  MPI processes are idle. During this process, FLAC3D is idle waiting for  
153 the flag to change. Then, FLAC3D (i) reads the TOU\_FLA file (in serial), (ii) solves for  
154 mechanical equilibrium (in parallel), and (iii) writes the FLA\_TOU file to transfer data to  
155 TOUGH3 (in serial), and modifies the flag (2). At this stage, the subroutine invoked  
156 previously by TOUGH3 is waiting for FLAC3D to finish execution, then the IOProcessor  
157 serially reads the flag and the FLA\_TOU file and distributes the variables/properties (bulk  
158 modulus, Biot's coefficient, strain, and stress) to all  $n$  MPI processes.

159 Finally, the parallel computing can restart with the calculation of mechanically-induced  
160 changes of flow properties and with continuation of the flow iterations to finish the current  
161 time step. When TOUGH3 is at the last time step, it will issue a flag (3) that FLAC3D will  
162 interpret and save the final state in a binary file. The mechanical state as well as the flow  
163 variables can also be saved at predefined times during execution. The use of Python routines  
164 within FLAC3D allows for easier handling of arrays (e.g. mapping of a given variable or  
165 extra post-processing computation), and also for handling and personalizing the entire output  
166 functions, making redundant the use of classical TOUGH3 output functions.

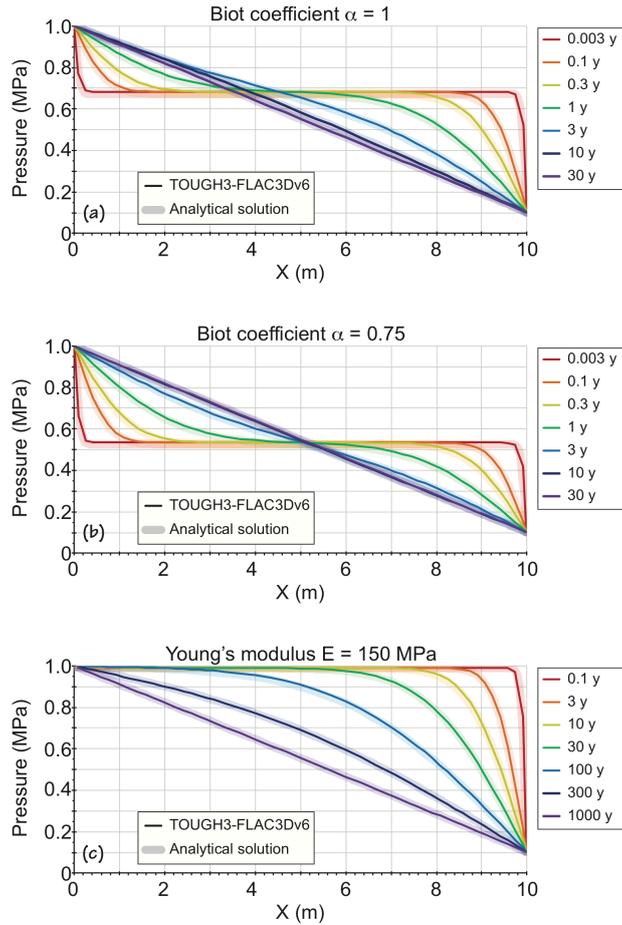
### 167 **3. VERIFICATION AND PERFORMANCE**

#### 168 **3.1 Analytical solution**

169 An analytical solution can be derived for a Terzaghi-like problem (Lux et al., 2015). The  
170 problem here is part of the BenVaSim initiative to verify and benchmark several numerical  
171 codes (Rutenberg et al., 2018). While being a simplified numerical exercise, the model setup  
172 physically resembles a dam construction in a flooded drift with a pore pressure gradient  
173 allowing water flowing through the host rock. As the current approach is based on previous  
174 versions, more verifications can be found elsewhere (Blanco-Martín et al., 2017).

175 The model is one dimensional, fully saturated 10 m-long domain with displacement  
176 completely fixed in  $y$ - and  $z$ - directions and at  $x = 10$  m. The initial pressure and total stress  
177 in the model are set to 0.1013 MPa. Pore pressure is 1 MPa at left boundary ( $x = 0$  m), and  
178 0.1013 MPa at the right boundary ( $x = 10$  m). A total stress of 1 MPa is applied at time  $t=0^+$   
179 at  $x = 0$  m. The base case scenario accounts for a porous and low permeable material  
180 (porosity  $\phi = 0.15$ ; permeability  $\kappa = 10^{-20}$  m<sup>2</sup>) with stiff and deformable matrix (Young's  
181 modulus  $E = 8$  GPa; Biot's coefficient  $\alpha = 1$ , and Poisson's ratio  $\nu = 0$  to allow for 1D  
182 problem). Variations to the base case account for compressible grains (Biot's coefficient  $\alpha =$   
183 0.75) and for very soft material (Young's modulus  $E = 150$  MPa).

184 Figure 2 shows the profiles of the pore pressure at different times comparing the numerical  
185 (solid line) and the analytical (shaded area) solutions. Results show that for different times,  
186 the developed approach is able to match the analytical solution, with only minor differences  
187 mostly related to space discretization (50 elements for the 10 m long domain). The initial  
188 pressurization of the system (undrained response at 0.003 y) is somewhat larger than the  
189 analytical solution due to coarse time discretization, while the steady state is well matched (at  
190 30 y for the base case scenario). A more compressible grain will result in less undrained



191

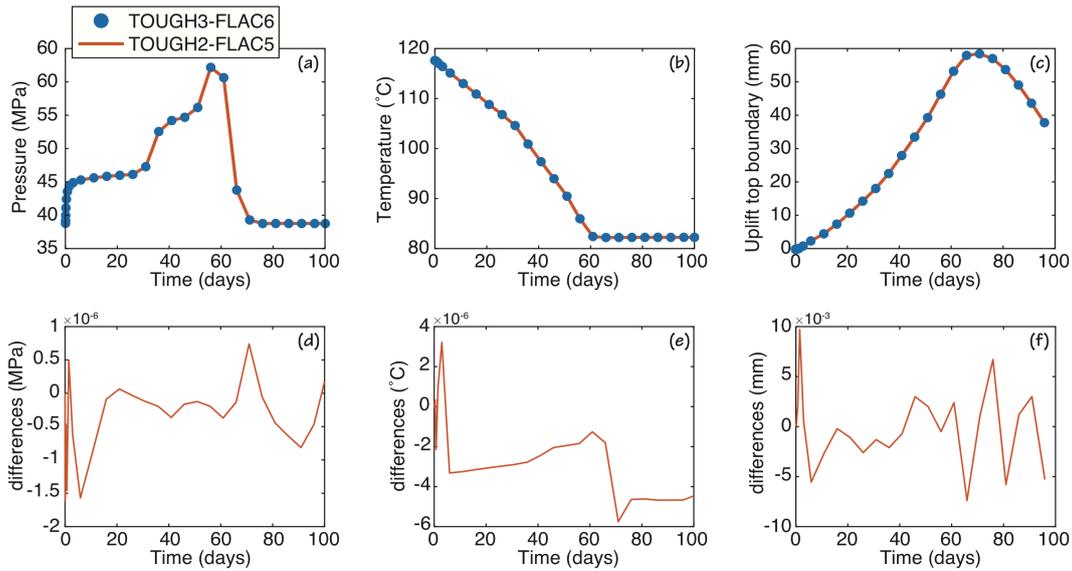
192 *Figure 2.* Verification of the coupling approach with an analytical solution in a low permeable environment. (a)  
 193 Base case scenario; (b) Compressible grains (Biot coefficient  $\alpha = 0.75$ ); (c) Soft matrix (Young's modulus  $E =$   
 194 150 MPa).

195 pressurization (Fig. 2b), while the soft matrix will result in much larger deformation,  
 196 allowing for larger porosity decrease in the matrix, and larger pressure increase during the  
 197 undrained response (Fig. 2c). As it can be seen, the numerical approach is in good agreement  
 198 with the analytical solution also for more critical scenarios.

### 199 **3.2 Comparison with previous simulator**

200 In order to verify the validity of TOUGH3-FLAC3Dv6/7, we compared results with the  
 201 previous TOUGH2-FLAC3Dv5 (Blanco-Martín et al., 2017). We checked the results for  
 202 variables such as injection pressure and temperature as well as the uplift of the top boundary.

203 We account for a 3D computational domain ( $10 \text{ km} \times 10 \text{ km} \times 4 \text{ km}$ ) with homogeneous  
204 properties (permeability  $\kappa = 5 \cdot 10^{-15} \text{ m}^2$ , porosity  $\phi = 0.1$ , density  $\rho = 2550 \text{ kg/m}^3$ ). The model  
205 ranges from a depth of -2 km to -6 km, and top and bottom boundaries, as well as the  
206 boundaries at  $x = y = 10 \text{ km}$ , are open to fluid flow. The boundaries at  $x = y = 0 \text{ km}$  are closed  
207 and allow for symmetry. Mechanically, we assume a poro-elastic material (Young's modulus  
208  $E = 10 \text{ GPa}$ , Poisson's ratio  $\nu = 0.25$ , Biot's coefficient  $\alpha = 1$ ), with the top and side  
209 boundaries ( $x = y = 10 \text{ km}$ ) at fixed stress conditions, with rollers for all the other boundaries.  
210 We assume initial hydrostatic gradient for pore pressure, geothermal gradient for temperature  
211 ( $30 \text{ }^\circ\text{C/km}$ ), and lithostatic gradient for stresses. We simulate 60 days of cold-water injection  
212 ( $T = 10 \text{ }^\circ\text{C}$ ) in a saturated medium with variable rate (30 days at  $30 \text{ kg/s}$ , 20 days at  $60 \text{ kg/s}$ ,  
213 and 10 days at  $90 \text{ kg/s}$ ), followed by 40 days of shut-in period for a total simulation time of  
214 100 days. The injection region is at a depth of 4 km and extends over a region  $50 \text{ m} \times 50 \text{ m} \times$   
215  $50 \text{ m}$ . For both simulations, we use a mesh with  $\sim 46,500$  elements. Porosity changes depend  
216 on the bulk modulus and on the total volumetric strain, when larger than  $10^{-4}$ . The FLAC3D  
217 mechanical ratio between the maximum unbalanced force magnitude and the average applied  
218 force magnitude is set to  $10^{-7}$ . The TOUGH3 convergence criterion is set to  $10^{-5}$ .  
219 As shown in Figure 3, the two approaches are in extremely good agreement, with differences  
220 in pressure in the order of some Pa, minor differences in temperature, and differences in  
221 uplift in the order of some microns.  
222

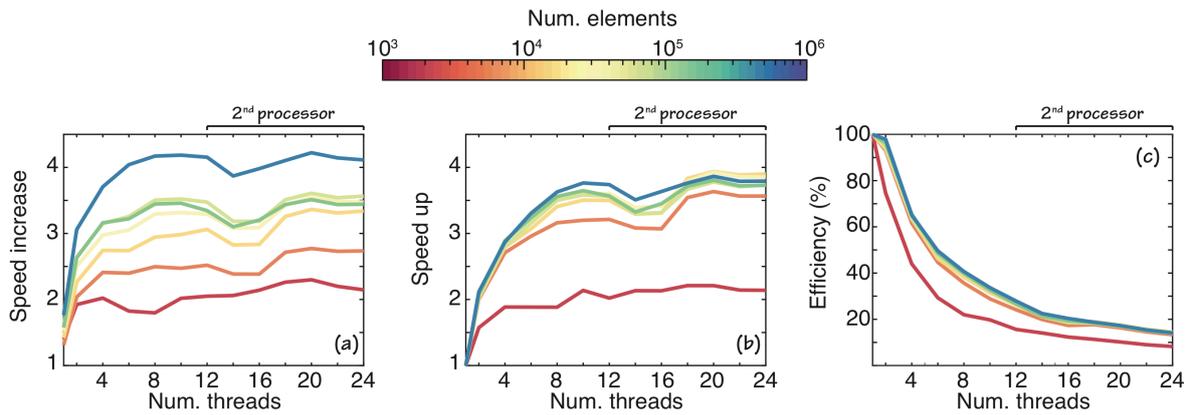


223

224 *Figure 3.* Comparison between TOUGH2-FLAC3Dv5 (red line) and TOUGH3-FLAC3Dv6 (blue dots) for  
 225 pressure near the injection point (a), temperature near the injection point (b), and uplift of the top boundary (c)  
 226 for a mesh with 46,464 elements. (d-f): differences between the two approaches.

### 227 **3.3. Performance Tests**

228 Figure 4a-b show the performance of TOUGH3-FLAC3Dv6 compared to the previous  
 229 version and to serial execution as a function of number of threads and for various sizes of the  
 230 computational domain. We varied the number of threads in FLAC3D and the number of cores  
 231 in TOUGH3 accordingly (each core allows for two threads). For each computational domain,  
 232 we have created similar initial conditions (via steady state simulation), to avoid biases on the  
 233 final simulation, which is the same as what was described in Section 3.2. Some differences  
 234 may arise for small meshes in the final displacement, which depends on the exact position of  
 235 the monitoring point, which is set the nearest to the position  $x=300$  m,  $y=300$  m, and  $z=-2000$   
 236 m (supplementary material, Fig. S2). For a single core and thread, the code is 1.5x faster  
 237 compared to the previous version only owing to better I/O handling and use of Python in  
 238 FLAC3D. Due to parallelization overhead (communication and domain decomposition), the  
 239 speed up clearly depends on the size of the mesh, with an increase of up to 2x faster for a  
 240 coarse mesh with 4,000 elements and up to 4x faster for a relatively fine



241

242 *Figure 4.* (a) Speed increase for TOUGH3-FLAC3Dv6 compared to TOUGH2-FLAC3Dv5 as a function of the  
 243 number of threads for different mesh sizes (colormap). (b) Speed up of the current approach (execution time for  
 244 a single thread compared to multi-thread) as a function of number of threads for different mesh sizes (colormap)  
 245 (c) Efficiency of TOUGH3-FLAC3Dv6 for different mesh sizes (colormap)

246 mesh with 800,000 elements. Interestingly, for all the cases the performances do not improve  
 247 after reaching 12 threads (6 cores). This could be linked to the configuration of the  
 248 workstation used for testing, that accounts for two 6-core processors. In fact, FLAC3D's  
 249 developers suggest the use of single processor workstation<sup>1</sup>. Therefore, the scaling for  
 250 multiple cores is not effective past 12 threads, and the efficiency (speed up/number of cores)  
 251 of the coupled approach strongly drops below 50% (Figure 4c). Overall, the speed up is not  
 252 dramatically smaller when compared to what we obtain on the same workstation for using  
 253 TOUGH3 or FLAC3D independently from each other: a speed increase of 6× (supplementary  
 254 material, Fig. S1).

255

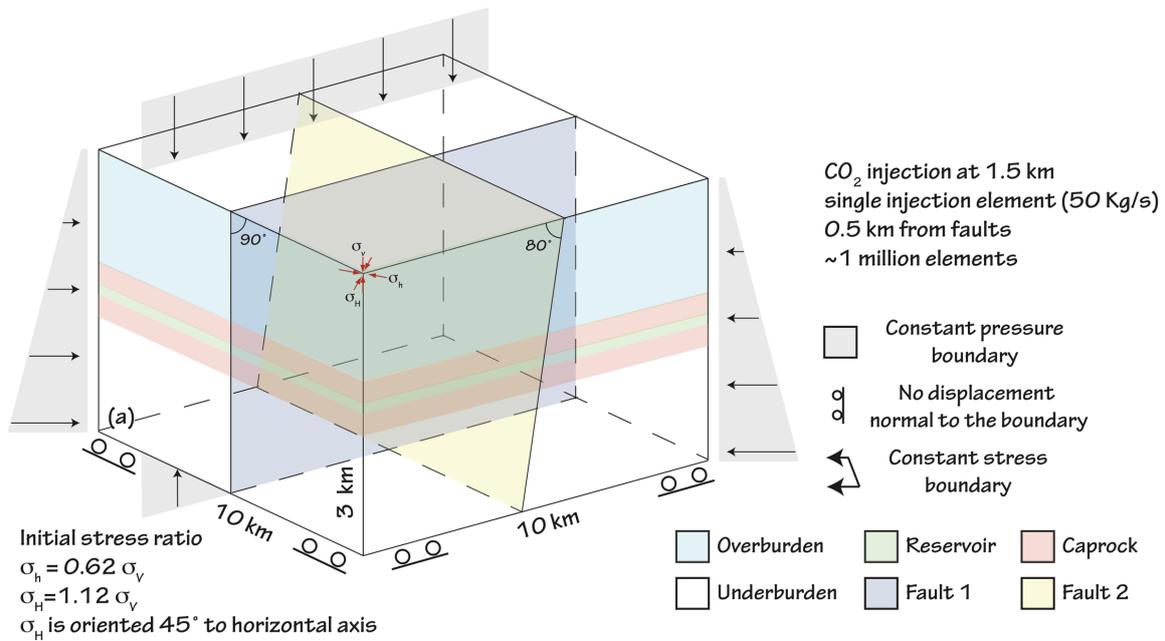
---

<sup>1</sup> <https://www.itascainternational.com/software/faqs/does-itasca-software-support-multiple-processors>

256 **4. CASE STUDIES**

257 **4.1 Fault reactivation during CO<sub>2</sub> sequestration**

258 The model presented here closely follows previous works addressing the same topic  
 259 (Rutqvist et al., 2016). A three dimensional model was already proposed by Rinaldi et al.  
 260 (2015), who addressed the effect of the well orientation on induced seismicity and CO<sub>2</sub>  
 261 leakage through the fault. Here, thanks to the faster solver, we introduce a further  
 262 complication in the model, which is the presence of a multiple fault system (Fig. 5). A similar  
 263 model was also employed for studying the natural seismicity occurring at Matsushiro, Japan  
 264 (Cappa et al., 2009), but here the much larger number of elements allows for better details.  
 265 We simulate a 3D computational domain 10 km × 10 km × 3 km with about 1 million  
 266 elements. The two fault zones strike N90° and N180° while dipping 90° and 80°,  
 267 respectively, and intersect at the center of the computational domain, assuming the north is  
 268 oriented in the y-direction. (Fig. 5). Injection occurs in a 100 m thick reservoir, bounded by



269

270 *Figure 5. Computational domain for the study of fault reactivation during CO<sub>2</sub> injection.*

271 two 150 m thick caprock formations, at a distance of 500 m from each fault (at a single point  
272  $x = 4500$  m,  $y = 4500$  m,  $z = -1500$  m), and with a constant rate of 50 kg/s (1.6 Mt/y) for a  
273 total of 3 years. Initial conditions account for hydrostatic pressure and geothermal gradient,  
274 and the simulation is considered isothermal (i.e., temperature is only needed to calculate the  
275 fluid viscosity and density). Initial stress follows a strike-slip regime, with both maximum  
276 and minimum principal stresses horizontal. The maximum horizontal stress ( $\sigma_H$ ) is oriented  
277 N45° with a stress ratio  $\sigma_H/\sigma_V = 1.12$ , while the minimum horizontal stress ( $\sigma_h$ ) has a ratio of  
278  $\sigma_h/\sigma_V = 0.62$ , with the vertical one ( $\sigma_V$ ) being the lithostatic stress. Boundaries are all open to  
279 fluid flow with constant lithostatic stress and hydrostatic pore pressure, except for the bottom  
280 where the displacement normal to the boundary is null. The system is initially fully saturated  
281 with brine, with retention curves for capillary pressure and relative permeability following  
282 van Genuchten (1980).

283

284 *Table 1.* Hydraulic and mechanical properties for the different domains in the CO<sub>2</sub> injection case study  
285

	<b>Reservoir</b>	<b>Caprock</b>	<b>Overburden</b>	<b>Underburden</b>	<b>Damage zone</b>	<b>Fault core 1</b>	<b>Fault core 2</b>
Permeability (m <sup>2</sup> )	10 <sup>-13</sup>	10 <sup>-19</sup>	10 <sup>-14</sup>	10 <sup>-18</sup>	10 <sup>-15</sup>	10 <sup>-17</sup>	10 <sup>-17</sup>
Porosity	0.1	0.1	0.1	0.1	0.1	0.1	0.1
Rock density (kg/m <sup>3</sup> )	2260	2260	2260	2260	2260	2260	2260
Residual CO <sub>2</sub> saturation (-)	0.05	0.05	0.05	0.05	0.05	0.05	0.05
Residual liquid saturation (-)	0.3	0.3	0.3	0.3	0.3	0.3	0.3
van Genuchten (1980), $p_0$ (kPa)	19.9	621	19.9	621	19.9	19.9	19.9
van Genuchten (1980), $m$ (-)	0.457	0.457	0.457	0.457	0.457	0.457	0.457
Young's modulus (GPa)	10	10	10	10	10	5	5
Poisson's ratio	0.25	0.25	0.25	0.25	0.25	0.25	0.25
Peak/residual friction angle (°)	-	-	-	-	-	31/29	31/29
Dilation angle (°)	-	-	-	-	-	10	10

286

287 For the sake of simplicity, we do not include permeability changes at this stage, as we focus  
288 on the reactivation of the faults. Permeability changes are instead more relevant when  
289 studying CO<sub>2</sub> leakage (Rinaldi et al., 2014). We account for the full hydro-mechanical  
290 coupling by modeling porosity changes as function of the volumetric strain and pore  
291 pressure.

292 We assume for all rock formations elastic rheology, except the core of both faults, which  
293 follows a strain-softening ubiquitous-joint model with frictional law depending on the  
294 accumulated plastic strain (Cappa and Rutqvist, 2011). Both elastic and hydraulic properties  
295 for the different domains are listed in Table 1. The simulation execution time with the given  
296 setup is comparable to a similar case in 2D and single fault for the previous version of the  
297 simulator (some hours).

298 Pressure evolution and CO<sub>2</sub> plume are shown in Figure 6. Results show that the  
299 pressurization of the reservoir is quite fast with changes up to 4 MPa near the injection point.  
300 Both faults start pressurizing right after injection starts, and less than 5 MPa are needed to  
301 reactivate both faults, with reactivation time depending on the fault strength (or actually on  
302 the fault orientation with respect to the state of stress). The CO<sub>2</sub> plume is still confined close  
303 to the injection, extending up to 200 m when both faults are reactivated.

304 Figure 7 shows how the rupture starts occurring on Fault 1 (the vertical blue fault, Fig. 5),  
305 and it is followed several days after by reactivation on Fault 2 (the dipping yellow fault, Fig.  
306 5). This is consistent with the state of stress, according to which Fault 1 is favorably oriented  
307 for shear activation. Given the frictional law in the ubiquitous joint model, the friction angle  
308 drops in the ruptured area (i.e. the one where plastic strain accumulates) from the peak value  
309 (31°) to the residual (29°). Reactivation on Fault 1 occurs after only 30 days of injection with  
310 maximum slip of 0.6 cm (Fig. 7a-b), and it is followed by the reactivation on Fault 2 after 70  
311 days of injection with a maximum slip of 0.2 cm, which involves only a small minor patch on

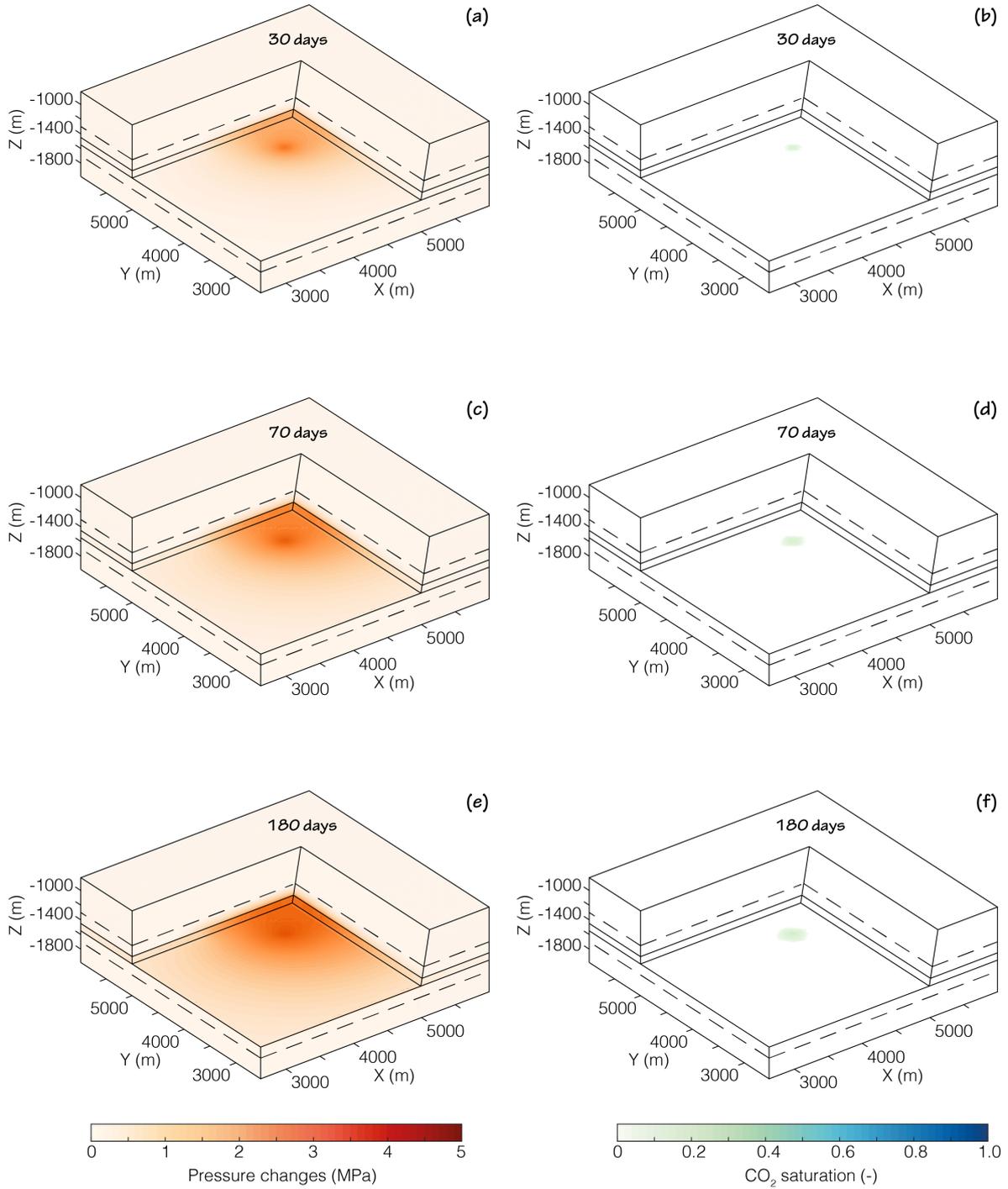
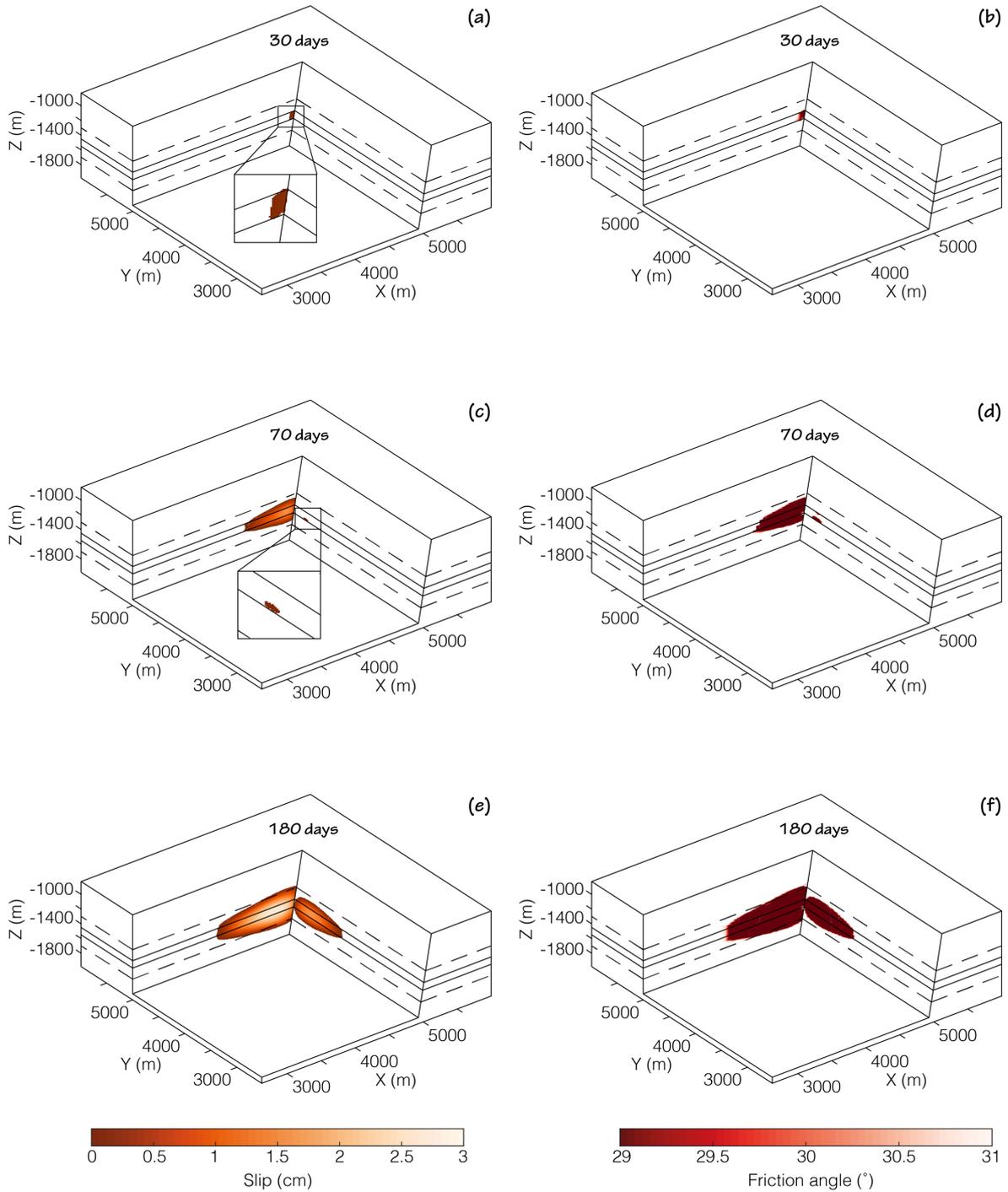


Figure 6. Distribution of pressure changes (a,c,e) and CO<sub>2</sub> saturation (b,d,f) after 30, 70, and 180 days of injection operations.



316

317

*Figure 7.* Distribution of plastic slip (a,b,c) and friction angle (d,e,f) at 30, 70, and 180 days of injection.

318

the fault plane (Fig. 7c-d). Worth to note that while the injection continues, the rupture

319

continues extending on the two faults, reaching a maximum extent after 180 days of

320

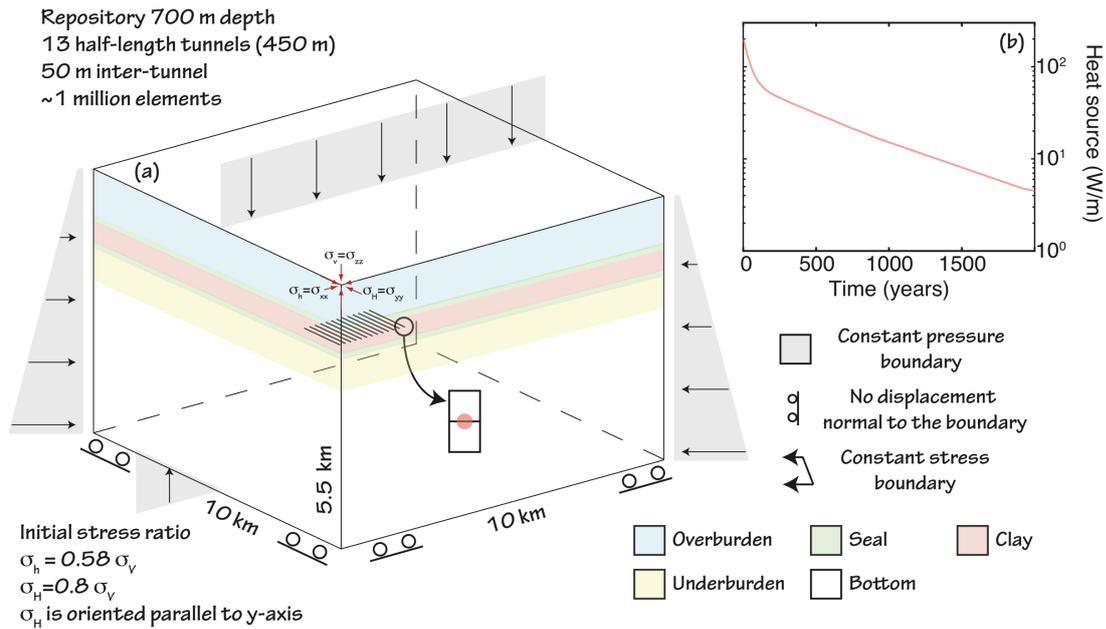
injection, which is much shorter than the total injection time of 3 years. The rupture area is

321 larger on Fault 1 given the more optimal orientation for shear rupture with respect to the  
322 stress field.

#### 323 **4.2 Potential for fault reactivation during geological nuclear waste disposal**

324 The model presented in this section aims at understanding the stress and strain changes  
325 occurring at depth in a nuclear waste geological repository. We simulate the heat generated  
326 by several, parallel nuclear waste emplacement tunnels located in argillaceous clay host rock,  
327 i.e. following the Swiss concept for geological nuclear waste disposal (NAGRA, 2016). The  
328 use of a refined mesh allows for more details. Figure 8a shows the three-dimensional  
329 computational domain, with dimensions  $10 \text{ km} \times 10 \text{ km} \times 5.5 \text{ km}$  and starting at the ground  
330 surface. We simulate conditions during 2000 years after nuclear waste disposal in a  
331 repository located in a clay layer with anisotropic permeability at a depth of 700 m, and  
332 embedded within two seal formations as well as under- and overburden. Thanks to symmetry,  
333 we simulate only a quarter of the domain, and simulate 13 half-length tunnels. Each tunnel  
334 has a length of 450 m and the tunnels are 50 m apart. Each element of the tunnel is a heat  
335 source variable in time (Fig. 8b), and is connected to two elements of the main computational  
336 domain (Fig. 8a,). The boundaries at  $x = 0 \text{ m}$  and  $y = 0 \text{ m}$  are closed to fluid flow and have  
337 null displacement normal to the boundary to simulate symmetry. The other boundaries are  
338 open to fluid flow and have constant lithostatic stress and hydrostatic pore pressure, with the  
339 exception of the bottom boundary where the normal displacement is blocked, and the pore  
340 pressure is set to  $\sim 54 \text{ MPa}$ . Initial conditions follow hydrostatic and geothermal ( $30 \text{ C}^\circ/\text{km}$ )  
341 gradients, while we impose normal stress conditions with the lithostatic vertical

342



343

344 *Figure 8.* Computational domain for the study of potential fault reactivation during geological nuclear waste  
 345 disposal.

346 stress ( $\sigma_v$ ) being the maximum principal stress and with  $\sigma_{xx} = \sigma_h$  as minimum principal stress.

347 We impose stress ratios  $\sigma_h / \sigma_v = 0.58$  and  $\sigma_H / \sigma_v = 0.8$ . We simulate full thermo-hydro-

348 mechanical coupling by assuming that the porosity depends on strain, pressure, and

349 temperature. We neglect at the current stage any permeability variation as a function of the

350 stress field. For simplicity, all the layers follow an elastic rheology, and we varied only the

351 hydraulic properties (Table 2). We use here single phase conditions, assuming that the

352 tunnels are already fully saturated at hydrostatic pressure conditions after emplacement, i.e.

353 ignoring some of the short-term re-saturation process that may take tens of years (Rutqvist et

354 al., 2014) as well as potential gas generation (Fall et al., 2014; Shaw, 2015). The full 2000

355 years are simulated in a bit more than 2 hours.

356 Figure 9 shows the temperature and pore pressure distribution at different stages of evolution.

357 Thermal effects are slower and only few °C changes are observed 10 years after emplacement

358 in the near repository region (Fig. 9a), but the temperature changes are large enough to enable

359 so-called thermal pressurization, a phenomenon known to occur when heating pore fluids in

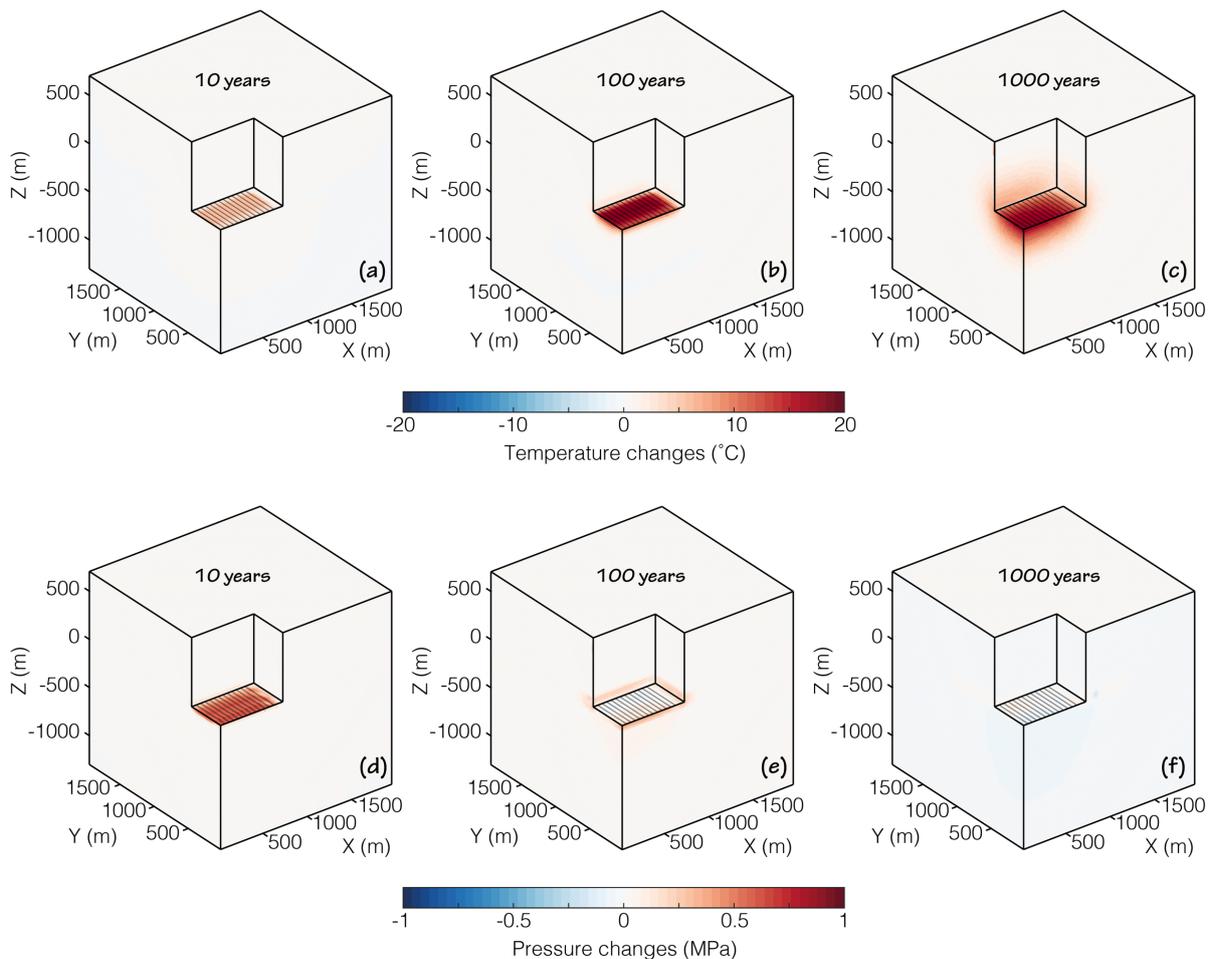
360  
361  
362

Table 2. Hydraulic and mechanical properties for the different domains in the deep geological repository case study. Note that for the clay we assume anisotropic permeability.

	Soil	Overburden	Seal	Clay	Underburden	Bottom
Permeability ( $k_h/k_v$ ) ( $m^2$ )	$10^{-14}$	$10^{-15}$	$10^{-17}$	$3 \cdot 10^{-19}/6 \cdot 10^{-20}$	$10^{-17}$	$10^{-18}$
Porosity	0.1	0.1	0.05	0.074	0.01	0.01
Rock density ( $kg/m^3$ )	2430	2430	2430	2430	2430	2430
Thermal conductivity ( $W/m \text{ } ^\circ C$ )	3.2	3.2	3.2	3.2	3.2	3.2
Specific Heat ( $J/kg \text{ } ^\circ C$ )	920	920	920	920	920	920
Young's modulus (GPa)	8	8	8	8	8	8
Poisson's ratio	0.27	0.27	0.27	0.27	0.27	0.27
Linear thermal expansion coefficient ( $^\circ C^{-1}$ )	$2 \cdot 10^{-5}$	$2 \cdot 10^{-5}$	$2 \cdot 10^{-5}$	$2 \cdot 10^{-5}$	$2 \cdot 10^{-5}$	$2 \cdot 10^{-5}$

363

364



365

366  
367

Figure 9. Distribution of changes in temperature (a,b,c) and pressure (d,e,f) in a deep geological repository for the storage of nuclear waste 10, 100, and 1000 years after emplacement.

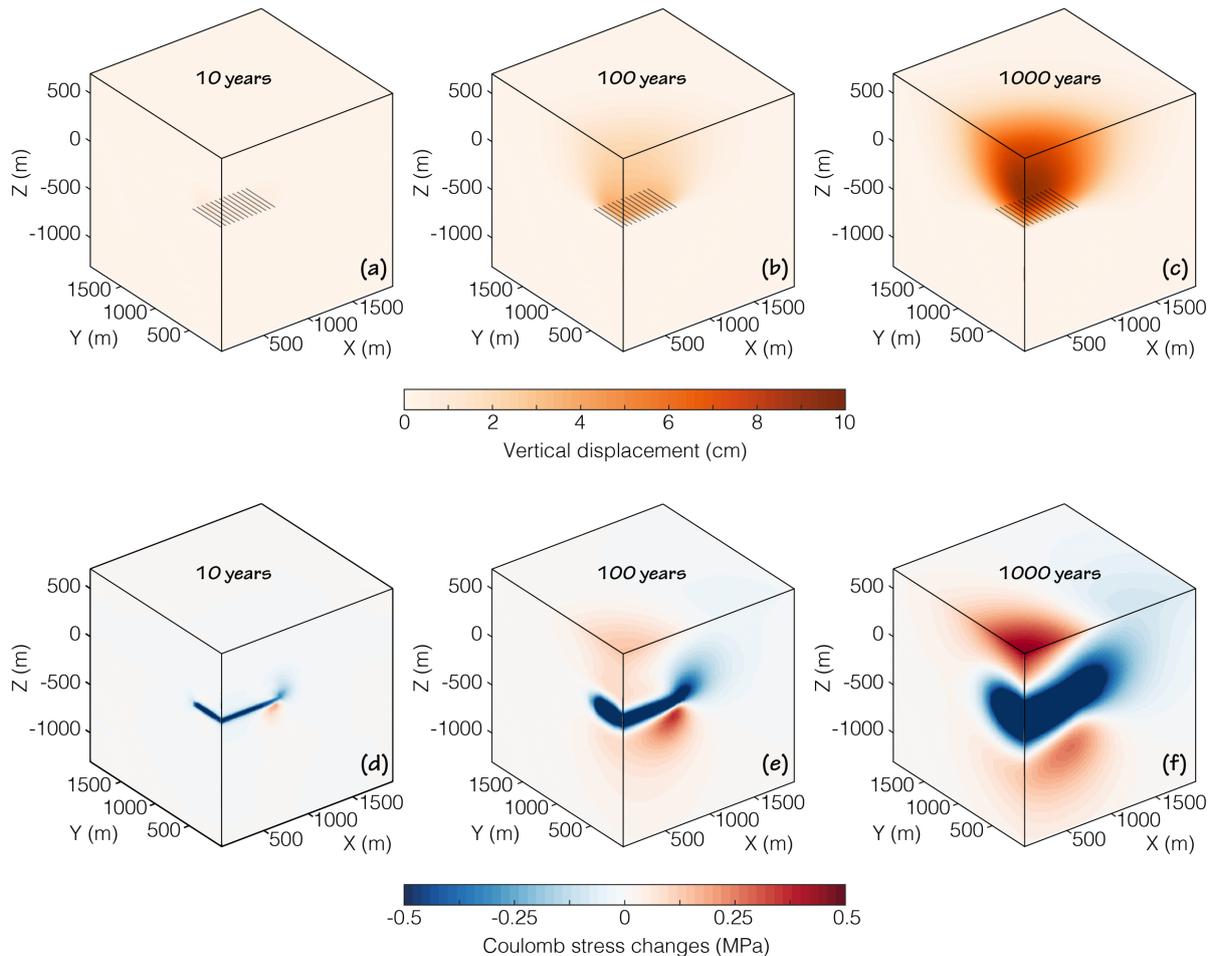
368 low permeability rocks, such as shale (Gens et al., 2007; Ghabezloo and Sulem, 2009; Zhang  
369 et al., 2017). In this case, the thermal pressurization causes a relatively small pore pressure  
370 increase of about 1 MPa (Fig. 9d). The domain remains fully saturated.

371 At later stages, temperature changes start distributing in the domain up to several hundreds of  
372 meters outside the repository after 1000 years (Fig. 9c). At the same time, the pressure  
373 changes are diffused in the low permeable clay formation, and only minor changes are  
374 observed after 1000 years (Fig. 9f).

375 The temperature changes due to the heat generated by the nuclear waste can be responsible of  
376 quite large deformation at the ground surface, up to several cm uplift after 1000 years (Fig.  
377 10a, b, and c). We also evaluated the potential for fault reactivation. Starting from the  
378 changes in the full stress tensor, we evaluate the Coulomb Stress change as  $\Delta CFS = \Delta\tau +$   
379  $\mu(\Delta\sigma_n + \Delta p)$ , where  $\Delta\tau$  is the change in shear stress,  $\Delta\sigma_n$  is the change in normal stress and  
380  $\Delta p$  is the change in pore pressure, with the convention of stresses negative for compression.

381 Shear and normal stresses are calculated for faults striking parallel to the tunnels and with  $80^\circ$   
382 dip angle toward the repository (i.e. strike N180°). Figures 10d-f show how the repository  
383 itself is undergoing stress shadow (negative Coulomb stress changes), meaning that failure is  
384 hindered. At early times, the thermal pressurization is causing only more compression and  
385 stabilizing faults in the near repository region (Fig. 10d). Failure of the considered fault  
386 orientation is, however, favored at greater depth, where more seismogenic faults could be  
387 present. This is linked to the shear transfer caused by temperature changes, and it is then  
388 particularly relevant when the thermal effect starts distributing outside the clay formation  
389 (e.g. at 100 or 1000 years – Fig. 10e-f). These results are valid for steeply dipping faults and  
390 are well in agreement with a recent 2D study on fault reactivation during disposal of nuclear  
391 waste at depth (Urpi et al., 2019).

392



393

394 *Figure 10.* Distribution of vertical displacement (a,b,c) and Coulomb stress changes (d,e,f) in a deep geological  
 395 repository for the storage of nuclear waste 10, 100, and 1000 years after emplacement. Coulomb stress is  
 396 calculated for faults dipping  $80^\circ$  and striking  $N180^\circ$ , assuming the north is oriented in the  $y$ -direction.

### 397 **5. CONCLUSION AND FUTURE OUTLOOK**

398 We coupled for the first time the recently released TOUGH3 with FLAC3Dv6/7. The results  
 399 are well in agreement with analytical solutions and with the previous version of the simulator  
 400 (TOUGH2-FLAC3Dv5). The coupled code can now use all the new functionalities of  
 401 TOUGH3, including, among others, the flexibility of a Fortran 90 code and the use of PETSc.  
 402 Together with the claimed improved solver performance, FLAC3Dv6/7 includes an internal  
 403 link with Python scripting, which provides an average  $10\times$  faster I/O and variables  
 404 assignment performance compared to the FISH programming embedded in FLAC3Dv5.

405 The removal of some bottlenecks, thanks to the use of binary files rather than ASCII, and the  
406 removal of save/restore operations for FLAC3D, largely helped in improving the  
407 performances of the coupled approach. The current coupled code allows up to 4-fold increase  
408 in execution speed for a two 6-core processors workstation compared to the previous version  
409 for a mesh with about 800,000 elements.

410 The possibility to run problems with very large number of elements in the computational  
411 mesh will enable a more detailed description of the thermo-hydro-mechanical processes  
412 occurring at depth. We have provided in this work two examples: one relates to fault  
413 reactivation during CO<sub>2</sub> sequestration, and the second one relates to nuclear waste disposal.  
414 For the first test case, we were able to simulate reactivation of intersecting faults during  
415 injection operations. The example highlights the time of induced fault reactivation for each  
416 fault. In the second test case, we demonstrate the use of the approach to simulate multiple  
417 emplacement tunnels at high details. Albeit simplified, the example shows the evolution of  
418 stress and strain in a deep geological nuclear waste repository, including the potential for  
419 fault reactivation.

420 The current coupled code, however, does not provide yet the improved performances that  
421 could be desired for High-Performance Computing in clusters. Certainly, a drawback with the  
422 use of TOUGH-FLAC is the limitation to run exclusively on Windows-based machines.

423 However, Itasca plans on releasing a MPI version of their code in the future (Rutqvist, 2017),  
424 and the current version of TOUGH3-FLAC3D constitutes a strong base for future coupling.

425

## 426 **ACKNOWLEDGMENTS**

427 The research was funded by a research agreement between ENSI and SED and by a Swiss  
428 National Science Foundation (SNSF) Ambizione Energy research grant (PZENP2\_160555).  
429 Funding for LBNL was provided by the Spent Fuel and Waste Disposition Campaign, Office  
430 of Nuclear Energy of the U.S. DOE, under Contract Number DE-AC02-05CH11231. The

431 authors are grateful to Michael Ruthenberg and the BenVaSim community for useful  
432 discussions.

433

434

## 435 **REFERENCES**

- 436 Birkholzer, J.T., Tsang, C.-F., Bond, A.E., Hudson, J.A., Jing, L., Stephansson, O., 2019. 25  
437 years of DECOVALEX - Scientific advances and lessons learned from an international  
438 research collaboration in coupled subsurface processes. *Int. J. Rock Mech. Min. Sci.* 122,  
439 103995. doi:10.1016/j.ijrmms.2019.03.015
- 440 Blanco-Martín, L., Rutqvist, J., Birkholzer, J.T., 2017. Extension of TOUGH-FLAC to the  
441 finite strain framework. *Comput. Geosci.* 108, 64–71. doi:10.1016/j.cageo.2016.10.015
- 442 Blanco-Martín, L., Rutqvist, J., Doughty, C., Zhang, Y., Finsterle, S., Oldenburg, C.M.,  
443 2016. Coupled geomechanics and flow modeling of thermally induced compaction in  
444 heavy oil diatomite reservoirs under cyclic steaming. *J. Pet. Sci. Eng.* 147, 474–484.  
445 doi:10.1016/j.petrol.2016.09.002
- 446 Blanco-Martín, L., Wolters, R., Rutqvist, J., Lux, K.-H., Birkholzer, J.T., 2015. Comparison  
447 of two simulators to investigate thermal–hydraulic–mechanical processes related to  
448 nuclear waste isolation in saliferous formations. *Comput. Geotech.* 66, 219–229.  
449 doi:10.1016/j.compgeo.2015.01.021
- 450 Cappa, F., Rutqvist, J., 2011. Impact of CO<sub>2</sub> geological sequestration on the nucleation of  
451 earthquakes. *Geophys. Res. Lett.* 38, L17313. doi:10.1029/2011GL048487
- 452 Cappa, F., Rutqvist, J., Yamamoto, K., 2009. Modeling crustal deformation and rupture  
453 processes related to upwelling of deep CO<sub>2</sub>-rich fluids during the 1965–1967  
454 Matsushiro earthquake swarm in Japan. *J. Geophys. Res.* 114, B10304.  
455 doi:10.1029/2009JB006398
- 456 COMSOL, 2020. COMSOL Multiphysics® v. 5.6. www.comsol.com. COMSOL AB,  
457 Stockholm, Sweden.
- 458 Fall, M., Nasir, O., Nguyen, T.S., 2014. A coupled hydro-mechanical model for simulation of  
459 gas migration in host sedimentary rocks for nuclear waste repositories. *Eng. Geol.* 176,  
460 24–44. doi:10.1016/j.enggeo.2014.04.003
- 461 Flemisch, B., Darcis, M., Erbertseder, K., Faigle, B., Lauser, A., Mosthaf, K., Müthing, S.,  
462 Nuske, P., Tatomir, A., Wolff, M., Helmig, R., 2011. DuMux: DUNE for multi-  
463 {phase,component,scale,physics,...} flow and transport in porous media. *Adv. Water*  
464 *Resour.* 34, 1102–1112. doi:10.1016/j.advwatres.2011.03.007
- 465 Garitte, B., Nguyen, T.S., Barnichon, J.D., Graupner, B.J., Lee, C., Maekawa, K., Manepally,  
466 C., Ofoegbu, G., Dasgupta, B., Fedors, R., Pan, P.Z., Feng, X.T., Rutqvist, J., Chen, F.,  
467 Birkholzer, J., Wang, Q., Kolditz, O., Shao, H., 2017. Modelling the Mont Terri HE-D  
468 experiment for the Thermal–Hydraulic–Mechanical response of a bedded argillaceous  
469 formation to heating. *Environ. Earth Sci.* 76, 345. doi:10.1007/s12665-017-6662-1
- 470 Gens, A., Vaunat, J., Garitte, B., Wileveau, Y., 2007. In situ behaviour of a stiff layered clay  
471 subject to thermal loading: observations and interpretation. *Géotechnique* 57, 207–228.  
472 doi:10.1680/geot.2007.57.2.207
- 473 Ghabezloo, S., Sulem, J., 2009. Stress dependent thermal pressurization of a fluid-saturated  
474 rock. *Rock Mech. Rock Eng.* 42, 1–24. doi:10.1007/s00603-008-0165-z
- 475 Harris, C.R., Millman, K.J., van der Walt, S.J., Gommers, R., Virtanen, P., Cournapeau, D.,  
476 Wieser, E., Taylor, J., Berg, S., Smith, N.J., Kern, R., Picus, M., Hoyer, S., van  
477 Kerkwijk, M.H., Brett, M., Haldane, A., del Río, J.F., Wiebe, M., Peterson, P., Gérard-

478 Marchant, P., Sheppard, K., Reddy, T., Weckesser, W., Abbasi, H., Gohlke, C.,  
479 Oliphant, T.E., 2020. Array programming with NumPy. *Nature* 585, 357–362.  
480 doi:10.1038/s41586-020-2649-2

481 Itasca, 2011. FLAC3d v5.0, Fast Lagrangian Analysis of Continua in 3 Dimensions, User's  
482 Guide.

483 Itasca, 2016. 3DEC v5.2, Distinct-Element Modeling of Jointed and Blocky Material in 3D.

484 Itasca, 2017. FLAC3d V6.0, Fast Lagrangian Analysis of Continua in 3 Dimensions, User's  
485 Guide.

486 Jung, Y., Pau, G.S.H., Finsterle, S., Pollyea, R.M., 2017. TOUGH3: A new efficient version  
487 of the TOUGH suite of multiphase flow and transport simulators. *Comput. Geosci.* 108,  
488 2–7. doi:10.1016/j.cageo.2016.09.009

489 Kim, J., Tchelepi, H.A., Juanes, R., 2011. Stability and convergence of sequential methods  
490 for coupled flow and geomechanics: Fixed-stress and fixed-strain splits. *Comput.*  
491 *Methods Appl. Mech. Eng.* 200, 1591–1606. doi:10.1016/j.cma.2010.12.022

492 Kolditz, O., Bauer, S., Bilke, L., Böttcher, N., Delfs, J.O., Fischer, T., Görke, U.J.,  
493 Kalbacher, T., Kosakowski, G., McDermott, C.I., Park, C.H., Radu, F., Rink, K., Shao,  
494 H., Shao, H.B., Sun, F., Sun, Y.Y., Singh, A.K., Taron, J., Walther, M., Wang, W.,  
495 Watanabe, N., Wu, Y., Xie, M., Xu, W., Zehner, B., 2012. OpenGeoSys: an open-source  
496 initiative for numerical simulation of thermo-hydro-mechanical/chemical (THM/C)  
497 processes in porous media. *Environ. Earth Sci.* 67, 589–599. doi:10.1007/s12665-012-  
498 1546-x

499 Kulik, D.A., Wagner, T., Dmytrieva, S. V., Kosakowski, G., Hingerl, F.F., Chudnenko, K.  
500 V., Berner, U.R., 2012. GEM-Selektor geochemical modeling package: revised  
501 algorithm and GEMS3K numerical kernel for coupled simulation codes. *Comput.*  
502 *Geosci.* doi:10.1007/s10596-012-9310-6

503 Lichtner, P.C., Hammond, G.E., Lu, C., Karra, S., Bisht, G., Andre, B., Mills, R.T., Kumar,  
504 J., Frederik, J.M., 2017. PLFOTRAN User Manual.

505 Lux, K.-H., Rutenberg, M., Seeska, R., Feierabend, J., Düsterloh, U., 2015. Kopplung der  
506 Softwarecodes FLAC3D und TOUGH2 in Verbindung mit in situ-, laborativen und  
507 numerischen Untersuchungen zum thermisch-hydraulisch-mechanisch gekoppelten  
508 Verhalten von Tongestein unter Endlagerbedingungen.

509 NAGRA, 2016. Arbeitsbericht NAB 16-43: Geomechanische Unterlagen.

510 Newell, P., Martinez, M.J., Eichhubl, P., 2017. Impact of layer thickness and well orientation  
511 on caprock integrity for geologic carbon storage. *J. Pet. Sci. Eng.* 155, 100–108.  
512 doi:10.1016/j.petrol.2016.07.032

513 Olivella, S., Gens, A., Carrera, J., Alonso, E.E., 1996. Numerical formulation for a simulator  
514 (CODE\_BRIGHT) for the coupled analysis of saline media. *Eng. Comput.* 13, 87–112.  
515 doi:10.1108/02644409610151575

516 Pruess, K., Oldenburg, C., Moridic, G., 2012. TOUGH2 User's Guide Version 2.1, Report  
517 LBNL-43134, Lawrence Berkeley National Laboratory, Berkeley, CA, USA.

518 Rinaldi, A.P., Rutqvist, J., Cappa, F., 2014. Geomechanical effects on CO<sub>2</sub> leakage through  
519 fault zones during large-scale underground injection. *Int. J. Greenh. Gas Control* 20,  
520 117–131. doi:10.1016/j.ijggc.2013.11.001

521 Rinaldi, A.P., Rutqvist, J., Finsterle, S., Liu, H.-H., 2017. Inverse modeling of ground surface  
522 uplift and pressure with iTOUGH-PEST and TOUGH-FLAC: The case of CO<sub>2</sub> injection  
523 at In Salah, Algeria. *Comput. Geosci.* 108, 98–109. doi:10.1016/j.cageo.2016.10.009

524 Rinaldi, A.P., Vilarrasa, V., Rutqvist, J., Cappa, F., 2015. Fault reactivation during CO<sub>2</sub>  
525 sequestration: Effects of well orientation on seismicity and leakage. *Greenh. Gases Sci.*  
526 *Technol.* 5, 645–656. doi:10.1002/ghg.1511

527 Rutenberg, M., Feierabend, J., Lux, K., Maßmann, J., Sentís, M.L., Graupner, B.J.,

528 Hansmann, J., Czaikowski, O., Wieczorek, K., Friedenber, L., Hotzel, S., Kock, I.,  
 529 Rutqvist, J., Hu, M., Rinaldi, A.P., 2018. BENVASIM—A BENCHMARKING OF  
 530 SIMULATORS FOR MODELING TH<sup>2</sup>M PROCESSES IN THE CONTEXT OF  
 531 RADIOACTIVE WASTE DISPOSAL, in: Proceedings of the TOUGH Symposium  
 532 2018. Berkeley, CA, USA.

533 Rutqvist, J., Wu, Y.-S., Tsang, C.-F., Bodvarsson, G., 2002. A modeling approach for  
 534 analysis of coupled multiphase fluid flow, heat transfer, and deformation in fractured  
 535 porous rock. *Int. J. Rock Mech. Min. Sci.* 39, 429–442. doi:10.1016/S1365-  
 536 1609(02)00022-9

537 Rutqvist, J., 2017. An overview of TOUGH-based geomechanics models. *Comput. Geosci.*  
 538 108, 56–63. doi:10.1016/j.cageo.2016.09.007

539 Rutqvist, J., Rinaldi, A.P., Cappa, F.F., Jeanne, P., Mazzoldi, A., Urpi, L., Guglielmi, Y.,  
 540 Vilarrasa, V., 2016. Fault activation and induced seismicity in geological carbon storage  
 541 – Lessons learned from recent modeling studies. *J. Rock Mech. Geotech. Eng.* 8, 789–  
 542 804. doi:10.1016/j.jrmge.2016.09.001

543 Rutqvist, J., Zheng, L., Chen, F., Liu, H.-H., Birkholzer, J., 2014. Modeling of Coupled  
 544 Thermo-Hydro-Mechanical Processes with Links to Geochemistry Associated with  
 545 Bentonite-Backfilled Repository Tunnels in Clay Formations. *Rock Mech. Rock Eng.* 47,  
 546 167–186. doi:10.1007/s00603-013-0375-x

547 Salimzadeh, S., Paluszny, A., Zimmerman, R.W., 2018. Effect of cold CO<sub>2</sub> injection on  
 548 fracture apertures and growth. *Int. J. Greenh. Gas Control* 74, 130–141.  
 549 doi:10.1016/j.ijggc.2018.04.013

550 Shaw, R.P., 2015. Gas Generation and Migration in Deep Geological Radioactive Waste  
 551 Repositories, Geological Society, London, Special Publications. Geological Society of  
 552 London. doi:10.1144/SP415

553 Taron, J., Elsworth, D., 2009. Thermal–hydrologic–mechanical–chemical processes in the  
 554 evolution of engineered geothermal reservoirs. *Int. J. Rock Mech. Min. Sci.* 46, 855–864.  
 555 doi:10.1016/j.ijrmms.2009.01.007

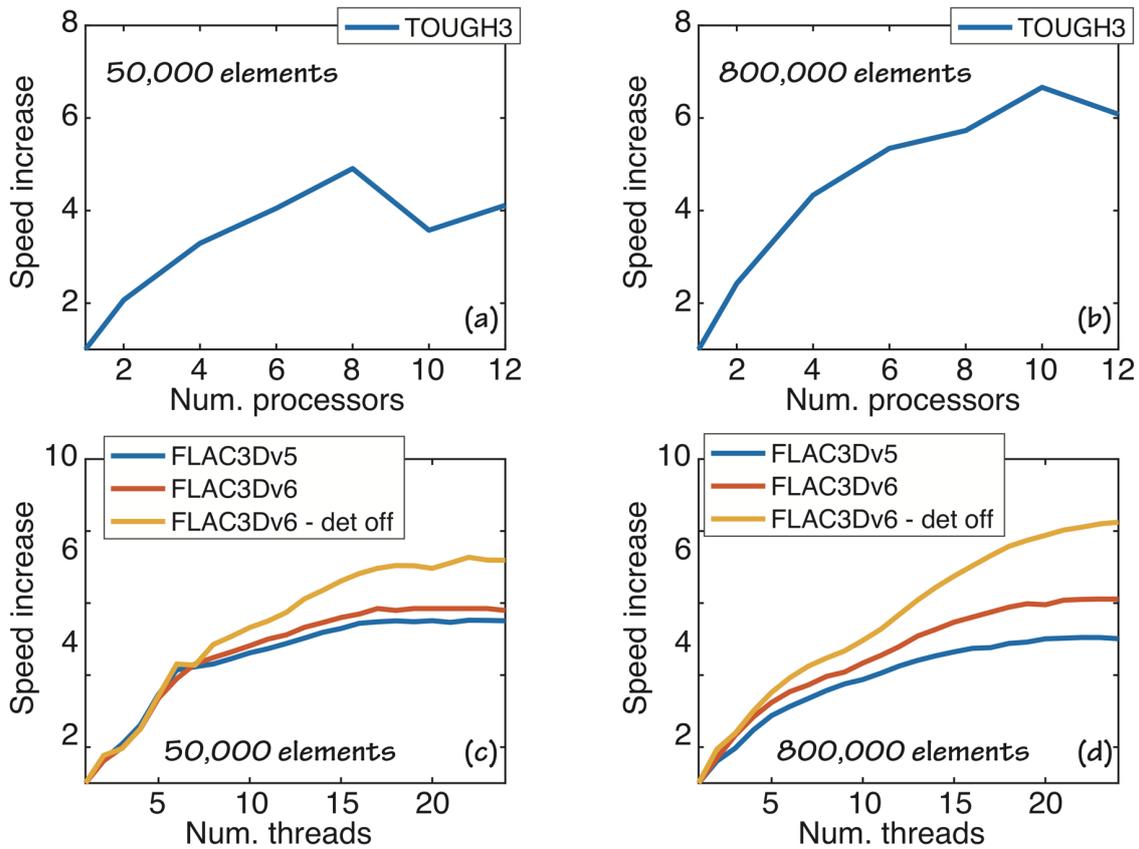
556 Urpi, L., Rinaldi, A.P., Rutqvist, J., Wiemer, S., 2019. Fault Stability Perturbation by  
 557 Thermal Pressurization and Stress Transfer Around a Deep Geological Repository in a  
 558 Clay Formation. *J. Geophys. Res. Solid Earth* 124, 8506–8518.  
 559 doi:10.1029/2019JB017694

560 van Genuchten, M.T., 1980. A Closed-form Equation for Predicting the Hydraulic  
 561 Conductivity of Unsaturated Soils. *Soil Sci. Soc. Am. J.*  
 562 doi:10.2136/sssaj1980.03615995004400050002x

563 Yapparova, A., Gabellone, T., Whitaker, F., Kulik, D.A., Matthäi, S.K., 2017. Reactive  
 564 Transport Modelling of Dolomitisation Using the New CSMP++GEM Coupled Code:  
 565 Governing Equations, Solution Method and Benchmarking Results. *Transp. Porous  
 566 Media* 117, 385–413. doi:10.1007/s11242-017-0839-7

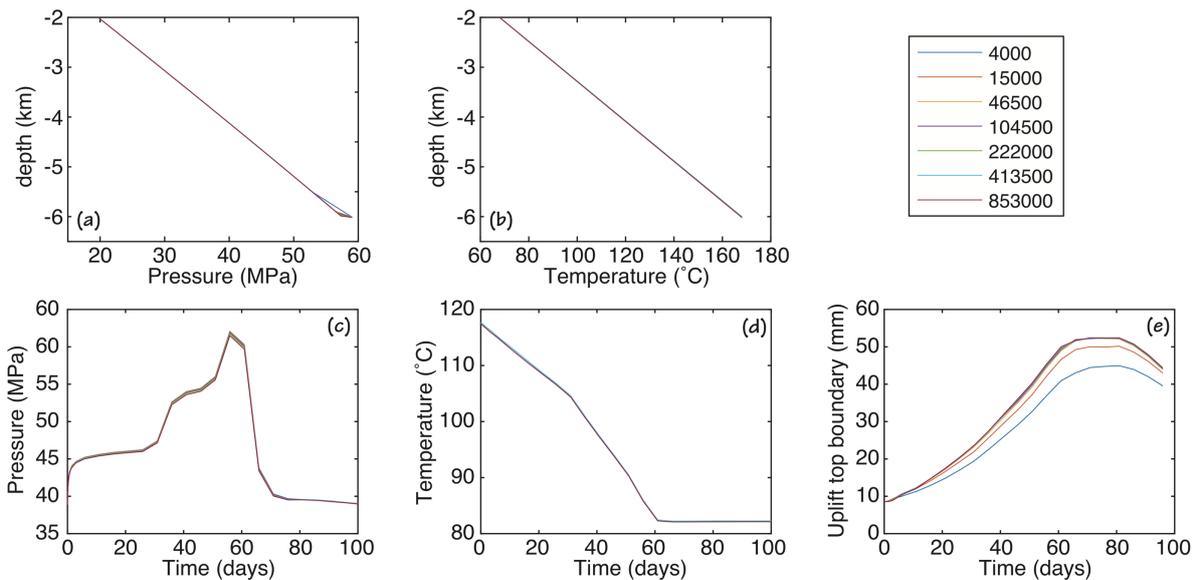
567 Zhang, C.-L., Conil, N., Armand, G., 2017. Thermal effects on clay rocks for deep disposal  
 568 of high-level radioactive waste. *J. Rock Mech. Geotech. Eng.* 9, 463–478.  
 569 doi:10.1016/j.jrmge.2016.08.006

570  
 571



573

574 *Figure S1.* (a,b) Evaluation of speed increase for TOUGH3. (c,d) Evaluation of the speed increase in FLAC3D.



575

576 *Figure S2.* Simulations for different computational domain sizes: (a) Initial pressure; (b) Initial temperature; (c)  
577 Pressure evolution; (d) Temperature evolution; (e) Displacement of top boundary.