Detected climate change signals in atmospheric circulation: mechanisms, puzzles and opportunities

Tiffany A Shaw[1], Julie M. Arblaster[2], Thomas Birner[3], Amy Hawes Butler[4], Daniela I.V. Domeisen[5], Chaim I Garfinkel[6], Hella Garny[7], Kevin M Grise[8], and Alexey Yurievich Karpechko[9]

[1]University of Chicago
[2]Monash University
[3]Ludwig-Maximilians University
[4]NOAA Chemical Sciences Laboratory
[5]ETH Zurich
[6]Hebrew University of Jerusalem
[7]Deutsches Zentrum für Luft- und Raumfahrt
[8]University of Virginia
[9]Finnish Meteorological Institute

October 09, 2024

Emerging climate change signals in atmospheric circulation

T. A. Shaw[1], J. M. Arblaster[2], T. Birner[3,8], A. H. Butler[4], D. I.V. Domeisen[5.6], C.I. Garfinkel[7], H. Garny[8], K. M. Grise[9], and A. Yu. Karpechko[10]

[1]The University of Chicago, Chicago, IL, USA.

[2]ARC Centre of Excellence for the Weather of the 21st Century, Monash University, Victoria, Australia.

[3]Ludwig-Maximilians-University Munich, Munich, Germany.

[4]National Oceanic and Atmospheric Administration, Chemical Sciences Laboratory, Boulder, CO, USA.

[5]Université de Lausanne, Lausanne, Switzerland

[6]ETH Zurich, Zurich, Switzerland

[7]Fredy & Nadine Herrmann Institute of Earth Sciences, The Hebrew University of Jerusalem, Israel.

[8]Deutsches Zentrum für Luft- und Raumfahrt, Oberpfaffenhofen, Germany,

[9]University of Virginia, Charlottesville, VA, USA

[10]Finnish Meteorological Institute, Helsinki, Finland.

Corresponding author: Tiffany Shaw (tas1@uchicago.edu)

Key Points:

- Long term trends in atmospheric circulation are emerging across different regions and seasons with some attributed to human activities.
- Many circulation signals have been linked to dynamical mechanisms involving thermodynamic changes, although discrepancies remain.
- Emerging signals in combination with new tools promise considerable progress in understanding the dynamical response in the coming decade.

Abstract

The circulation response to climate change shapes regional climate and extremes. Over the last decade an increasing number of atmospheric circulation signals have been documented, with some attributed to human activities. The circulation signals represent an exciting opportunity for improving our understanding of dynamical mechanisms, testing our theories and reducing uncertainties. The signals have also presented puzzles that represent an opportunity for better understanding the circulation response to climate change, its contribution to climate extremes, interactions with moisture, and connection to thermodynamic discrepancies. The next decade is likely to be a golden age for dynamics with many advances possible.

Plain Language Summary

Regional climate change signals in atmospheric circulation (wind and pressure) have been documented in many regions. Some of the signals are expected and have been attributed to human activities whereas others are not. The next decade represents an exciting time to better understand the dynamical mechanisms underlying these signals and their relationship to thermodynamic signals with the goal of improving regional climate prediction.

1 Introduction

The emergence and attribution of thermodynamic signals in response to anthropogenic climate change is well appreciated. Indeed global-mean warming over land and ocean, amplified warming in the tropical upper troposphere, rising of the tropopause, cooling of the stratosphere, regional land warming, and Arctic amplification of surface warming have all been attributed to human activities (IPCC 2021). Thermodynamically driven changes in regional hot extremes, heavy precipitation and drought have also been confidently attributed to human activities in some regions (IPCC 2021, Fig. SPM.3). This progress on thermodynamic signals has been achieved through a multi-pronged approach: detection of observed signals, attribution to human activities, and understanding of the underlying mechanisms using climate model simulations that exhibit fidelity in the signal and mechanisms.

Atmospheric circulation is well-known to affect regional climate through changes in fluid-dynamic variables, including atmospheric wind, pressure. These changes can subsequently influence moisture, clouds and radiation. Many generations of climate models have predicted robust circulation responses to climate change by the end of the century, including an upward shift and acceleration of the subtropical jet streams, weakening and expansion of the Hadley circulation, poleward shifts of the eddy-driven jet streams, strengthening of the storm tracks in the Southern Hemisphere and seasonally varying storm track responses in the Northern Hemisphere. In general, circulation signals are more uncertain as compared to thermodynamic ones, especially at the regional scale, due to large internal variability and the lack of sufficiently strong constraints on atmospheric dynamics (Shepherd, 2014). Furthermore, competing influences on dynamics in a changing climate, e.g. Arctic versus tropical warming, cloud shortwave versus longwave responses, aerosol cooling versus greenhouse gas warming, etc also can lead to a weak net dynamical response (Perlwitz 2012, Shaw et al., 2016). Hence dynamic variables are considered to have a lower signal-to-noise ratio, which has cascading impacts on hydrological cycle signals (Elbaum et al 2022).

Over the last decade an increasing number of atmospheric circulation signals, here defined as statistically significant linear trends over the satellite era or longer, have been documented in the literature. These signals are part of a growing number of regional climate change signals, some of which exhibit discrepancies with model predictions (Shaw et al. 2024). Here we focus specifically on atmospheric circulation signals that have been documented in the literature since recent assessments (IPCC 2021, Shepherd 2014). We specifically highlight signals that have emerged and been attributed to human activities; discuss progress on understanding dynamical mechanisms underlying the signals; and describe remaining puzzles, including the role of internal variability versus the forced response versus observational uncertainty, model-observation discrepancies and the impact of mean state biases. We discuss the importance of linking statistical analysis and understanding of dynamic and thermodynamic signals. In particular, some thermodynamic signals exhibit discrepancies with model predictions, e.g. the “pattern effect” of SST trends, and are potentially linked to the atmospheric circulation, e.g. via thermodynamic gradients and cloud radiative effects. Finally, we highlight how circulation signals, along with existing and emerging tools, represent an exciting opportunity for making progress in the next few decades on understanding the dynamical mechanisms behind the circulation response to climate change.

2 Circulation signals

The number of atmospheric circulation signals reported in the literature across different regions, hemispheres, and seasons has grown significantly in recent years (Table 1). Some are zonal-mean signals (8 out of 20**)** but many are regional **(**12 out of 20). For example, increased sea-level pressure near South-West Western Australia is associated with recent drying trends in this region (Fig. 1a,c,e; Hope et al 2006). Furthermore, many Southern Hemisphere signals are zonally symmetric, leading to similar impacts across longitudinal regions (Kang et al. 2024).

In some cases the signals have been detected and attributed to human activities (see below and Table 1). In other cases the role of internal variability and/or reanalysis biases still needs to be assessed. In many cases the sign of the signal is consistent with model predictions, however in some cases there is a discrepancy between observations and models. In still other cases, expected regional signals, like reduced precipitation in the Central and Western Mediterranean associated with higher sea-level pressure, will take more time to emerge (Fig. 1b,d,f) (Seager et al. 2024).

One of the earliest examples of an atmospheric circulation signal being formally attributed to human activities involved ozone depletion (Gillett et al., 2013). The circulation signals include an increase in the strength of the winds in the southern hemisphere stratosphere, an associated delay of the spring-time breakdown of the stratospheric polar vortex, and a poleward shift of the eddy-driven tropospheric jet stream (Fig. 2) and southern Hadley cell edge in austral summer (Thompson et al. 2011, Lee & Feldstein 2013, WMO 2018). Since the 2000s, ozone recovery, which opposes the influence of greenhouse gas increases on the circulation, has been associated with reduced SH circulation trends (Banerjee et al 2020, Zambri et al. 2021), though these are sensitive to end points (Fig. 2).

In recent years several more atmospheric circulation signals have been attributed to human activities (Table 1), including greenhouse gas emissions, but also with ozone depletion or aerosol emissions either in isolation or in combination (e.g. Gillett et al., 2016). In the Northern Hemisphere the combination of anthropogenic greenhouse gas and aerosol emissions have weakened the summertime circulation as measured by the zonal-mean storm tracks (eddy kinetic energy, Chemke & Coumou 2024), zonal-mean jet stream, and regional surface cyclone activity (mean sea level pressure, Kang et al. 2024b). Improved estimates of anthropogenic aerosol forcing were important for the improved Northern Hemisphere summertime storm track signals in CMIP5 versus CMIP6 (Chemke & Coumou 2024). The weakening of the East Asian summertime jet stream has been attributed exclusively to anthropogenic aerosol emissions (Dong et al. 2022).

The weakening of the annual-mean Northern Hemisphere Hadley cell has also been attributed to anthropogenic greenhouse gas and aerosol emissions (Chemke & Yuval 2023, Lionello et al. 2024). The poleward shift of the Southern Hemisphere Hadley cell edge has been attributed to ozone depletion and anthropogenic greenhouse gas emissions (Grise et al. 2019; Lionello et al. 2024).

3 Progress in understanding mechanisms

Many dynamical mechanisms have been proposed to explain atmospheric circulation responses to anthropogenic forcing that have been robustly predicted by generations of climate models (Thompson et al. 2011, Vallis et al. 2015, Hoskins & Woollings 2015, Shaw, 2019, Wills et al. 2019). Here we highlight progress on understanding mechanisms underlying the response to ozone depletion, greenhouse gas and aerosol forcing as they relate to the circulation signals listed in Table 1.

3.1 Ozone depletion

Ozone depletion reduces the shortwave absorption of ultraviolet radiation, cooling the lower stratosphere. This cooling induces an increase of the meridional temperature gradient and a strengthening of the stratospheric zonal wind consistent with thermal wind balance. Imposing a cooling of the lower stratosphere in idealized model simulations leads to a poleward shift of the tropospheric eddy-driven jet (Polvani & Kushner 2002, Kushner & Polvani 2004, Butler et al. 2010). However, the tropospheric response to stratospheric forcing is sensitive to the state of the troposphere (Chan & Plumb, 2009; Garfinkel et al. 2013). A mechanism proposed to explain the poleward shift of the eddy-driven jet stream in the lower atmosphere links the change in stratospheric winds to a modification of the eastward propagation of tropospheric eddies thereby affecting the momentum flux (Chen & Held 2007). At this time, there is still not a complete mechanistic understanding that connects the ozone hole to the shift of the jet stream and Hadley cell edge (Thompson et al. 2011, Kidston et al. 2015). This lack of understanding may in part be due to the complex dynamical interactions that are found to be crucial for a downward impact (Kidston et al. 2015).

3.2 Greenhouse gas forcing

Greenhouse gas increases lead to tropical upper tropospheric warming consistent with moist adiabatic adjustment (Manabe & Wetherald, 1975, Held 1993). This response increases the meridional temperature gradient near the tropopause, strengthening the subtropical jet and shear via thermal wind balance (Allen & Sherwood, 2008; Lee et al., 2019). This direct impact of the tropics on the atmospheric circulation is confirmed by a CO_2 increase only in the tropics in model simulations (Shaw & Tan 2018, Shaw, 2019).

The shift of the jet stream and Hadley cell in response to greenhouse gas increases have been argued to be connected to this tropical warming response (Lorenz & DeWeaver, 2007; Lu et al. 2007; Lu et al., 2014; Butler et al., 2010). However the poleward shift of the midlatitude near-surface jet and Hadley cell edge and the strengthening of the subtropical jet happen on distinct timescales (compare red and blue lines in Fig. 3), suggesting the shifts are driven by different mechanisms (Chemke & Polvani, 2019, 2021; Menzel et al., 2019). Recent studies suggest midlatitude processes including local moisture gradient, latent heat release, vertical temperature gradient (static stability), and cloud changes are more important than tropical changes (Shaw & Voigt, 2016, Voigt & Shaw, 2016, Chemke & Polvani 2019, 2021, Garfinkel et al., 2024; Lachmy, 2022; Tamarin-Brodsky & Kaspi, 2017; Tan & Shaw, 2020; Voigt et al., 2021). The importance of moisture and clouds has been revealed by advancing theory to incorporate moisture (e.g., Tamarin-Brodsky & Kaspi, 2017; Shaw et al. 2018, Lachmy, 2022) and simulations across the model hierarchy (Garfinkel et al 2024; Ghosh et al., 2024, Tan & Shaw 2020, Ceppi & Hartmann, 2016; Voigt & Shaw, 2015).

The signal of Northern Hemisphere summertime circulation weakening has been linked to a weakening of the near-surface temperature gradient due to Arctic amplification (Coumou et al. 2015), however recent work shows the contribution of Arctic sea ice loss and Arctic amplification to the circulation signal is negligible (Blackport et al, 2019; Blackport & Screen, 2021; Kang et al. 2023). Instead the weakening signal is related to high latitude warming over land (not ocean or sea ice) induced by greenhouse gas and aerosol forcing (Dong et al., 2022; Chemke & Coumou 2024, Kang et al., 2024b).

The strengthening of the Southern Hemisphere wintertime storm tracks, which occurs robustly across all longitudes, has been connected to several mechanisms: An increase in mean available potential energy due to increased latitudinal temperature gradients aloft (O'Gorman 2010); increased surface flux trends that reflect equatorward ocean energy transport and Southern Ocean cooling (Shaw et al. 2022); and changes in the vertical structure of the jet stream (Chemke et al. 2022).

Mechanisms explaining regional signals are related to stationary wave changes. The strengthening summertime Northern Hemisphere stationary wave signal has been connected to a teleconnection from the tropical Pacific (Sun et al. 2022) and soil moisture deficits (Teng et al. 2022). A related signal is the increase in extratropical heatwaves in summertime (e.g., Russo & Domeisen, 2023, Domeisen et al. 2023), which have been suggested to be related to increased "waviness" of the jet stream and the increased occurrence of so-called resonance events (Kornhuber et al., 2017; Mann et al., 2018), often associated with double jets (Rousi et al., 2022). However the quantitative mechanism underlying this link has not been established. Instead, anthropogenic aerosol forcing has been argued to be important for regional heat wave signals (Schumacher et al. 2024).

During wintertime the strengthening high over the Mediterranean has been connected to the large-scale upper-tropospheric circulation and land-sea contrast response, and specifically to a less rapid warming of the Mediterranean sea than of the surrounding land (Tuel & Eltahir 2020). The large-scale tropospheric circulation response consists of an eastward shift of wintertime stationary waves associated with strengthened eastward subtropical upper-level jet (Simpson et al., 2016; Wills et al 2019). This eastward shift is associated with uncertainty in regional climate change in e.g., Western North America (Simpson et al., 2016). Finally, the pattern of sea surface temperature warming can modify regional circulation and subtropical precipitation responses to greenhouse gas forcing (Zappa et al 2020).

3.3 Aerosol forcing

The mechanism proposed to explain the regional circulation signals in response to aerosol forcing involves the aerosol direct effect (aerosol-radiation interactions). Regions with reductions in aerosol optical depth, e.g. Eurasia and Eastern North America, show increases in clear-sky surface shortwave radiation (unmasking effect) whereas regions with increases in aerosol optical depth, e.g. South and East Asia, show a decrease in clear-sky surface shortwave radiation. The surface radiation signals weaken the meridional surface temperature gradient from the tropics to the extratropics, which following thermal wind balance weakens the summertime Eurasian jet. The shortwave radiation signals are coupled via the longitudinal circulation to the downstream ocean leading to a weakening of the storm tracks (Kang et al. 2024b).

Other studies have proposed additional mechanisms linked to the indirect influence of aerosols on clouds. For example, sulfate aerosols may brighten clouds which reflect more radiation to space, leading to a change in radiative balance that promotes poleward heat transport by the atmosphere and ocean (Needham & Randall, 2023).

4 Puzzles

4.1 Model-observation discrepancies

The lengthening observational record has provided some “puzzles” where there are apparent discrepancies between observed and modeled signals (Shaw et al. 2024). There are several well-known thermodynamic discrepancies, including opposite signed SST trends in observations and models in the tropical Pacific (Lee et al., 2022; Seager et al. 2022; Wills et al., 2022) and Southern Ocean (Wills et al., 2022; Kang et al., 2023).

In addition, important circulation discrepancies have been identified. In particular, the Walker circulation trend is toward a strengthening in observations but a weakening in models (Chung et al., 2019). Also, there is a strengthening of the Northern Hemisphere Hadley cell in reanalysis data but a weakening in models, though there is evidence that the reanalysis trends are artificial (Chemke & Polvani 2019b).

Similar to thermodynamic discrepancies, there are also cases where models capture the signal but it is underestimated as compared to reanalysis trends even after accounting for internal variability: increased Southern Hemisphere storminess trends (Chemke et al., 2022; Shaw et al., 2022) and North Atlantic lower-tropospheric jet strength trend (Blackport & Fyfe 2022, compare model distributions in colors to black line representing reanalysis in Fig. 4). In other cases the models overestimate the trends (strengthening of the upper-tropospheric jet stream; Woollings et al., 2023).

The relationship between thermodynamic and dynamic discrepancies is an active area of research. Recent papers show SST trend discrepancies impact Southern Hemisphere storminess and midlatitude jet trends (Yang et al., 2021; Kang et al., 2024), and heatwave trends over Europe are underestimated in models due to a discrepancy in the dynamical contribution (compare black dots representing models to colored lines representing observations in Fig. 5), although the details of this circulation trend discrepancy are not well understood and remain to be investigated (Vautard et al., 2023).

An important limitation of atmospheric circulation signals that needs to be taken into account when comparing model and observed signals is that atmospheric circulation signals rely heavily on reanalysis products. Such datasets can exhibit drifts and jumps due to changes in the underlying data sources (SPARC, 2022). In the Southern Hemisphere there is considerable spread in circulation signals across these products (Martineau et al. 2024, Kang et al. 2024). In the Northern Hemisphere, diabatic heating biases in reanalysis products have been shown to impact Hadley cell signals (Chemke & Polvani 2019). Surface pressure observations have been used to resolve the discrepancy in Hadley cell signals (Chemke & Yuval 2023).

4.2 **Disentangling forced response from internal variability**

One of the major challenges in comparing observed and model circulation signals is the confounding factors of internal variability, which can mask or exacerbate forced trends in the climate system, and observational uncertainty. For example, recent work for the Brewer-Dobson circulation trends shows that observational uncertainty can be large enough to account for the discrepancy with simulated Brewer-Dobson circulation trends in the middle stratosphere (Garny et al, submitted to RoG).

One way to separate the forced response from internal variability is using single forcing simulations. For example, if the signal is present in response to greenhouse gas or aerosol forcing only, and observational and model uncertainty is low, then it is likely a forced response. If the signal is present in the experiments without anthropogenic forcing (e.g. the preindustrial control experiment), then one cannot rule out the role of internal variability. Another way to quantify the role of internal variability is using large ensemble simulations with identical external forcing and slightly different initial conditions (Deser et al., 2020; Maher et al., 2021). The two approaches are combined in single-forcing large ensembles, which have been used to reconcile some discrepancies (by accounting for internal variability), such as the poleward expansion of the Hadley cell edge documented in the late 2000s (Grise et al., 2019) or cold winters over subpolar Eurasia from 1998 to 2012 (Garfinkel et al 2017; Outten et al 2022). However, given the relatively large magnitude of internal variability at regional scales (particularly in the extratropics during wintertime) and potential model errors, acknowledging a range of plausible future circulation trends ("storylines") is necessary for impacts planning (Zappa & Shepherd, 2017; Mindlin et al., 2020; Schmidt & Grise, 2021; Williams et al., 2024).

While large ensembles can help disentangle the signal from the noise, recent work has highlighted a signal-to-noise issue in coupled models suggesting that models may not properly represent the magnitude of forced signals relative to internal variability. This "signal-to-noise paradox" manifests most clearly when the ensemble-mean signal correlates better with observations of the real world than with individual members of the initialized model forecast ensemble (Weisheimer et al. 2024).

4.3 **Role of mean state biases/spread for future change**

The spread in model climatologies has been used to constrain thermodynamic climate change signals, e.g. the snow-ice albedo feedback (Hall and Qu 2006), through emergent constraints. Emergent constraints are statistical relationships between a model's representation of a particular physical process in the current climate and its future projection. Emergent constraints are most robust when they are supported by a plausible physical mechanism.

Several emergent constraints have been proposed for circulation signals (Simpson et al., 2021): for example, the Southern Hemisphere eddy-driven jet position (Kidston & Gerber, 2010), and the wintertime stationary wave response over the North Pacific (Simpson et al., 2016). In both cases, a mechanism was proposed to explain the emergent constraint: fluctuation dissipation theorem for jet position, and jet stream strength affecting stationary wavelength. Unfortunately, some dynamical emergent constraints are not robust across CMIP versions (Wu et al., 2019; Curtis et al., 2020; Karpechko et al. 2024). Furthermore, the Southern Hemisphere jet position constraint, which is only robust in wintertime (Simpson & Polvani, 2016), appears to be an artifact of the zonal mean (Breul et al., 2023).

Mean state model biases can have important implications for the forced response. For example, even if a model accurately simulates the observed circulation response to climate change (e.g., a poleward shift of the eddy-driven jet stream), if the circulation feature does not have the correct location or magnitude in the present-day climate, then the model's projected future climate change signal may be biased in terms of location and/or magnitude (Maraun et al., 2017; Grise, 2022). Systematically addressing this issue globally is challenging and requires a detailed understanding of the circulation features for all relevant regions.

5 Opportunities for progress

Understanding the emerging circulation signals and unraveling the puzzles they present provide exciting opportunities for making progress in understanding the dynamical response to climate change. Some opportunities for future research are:

5.1 **Investigate signals across the seasonal cycle**

Almost all of the dynamical signals in Table 1 are for the winter and summer seasons. Investigating signals in other seasons such as autumn and spring as well as seasonal transitions is important. During these seasons some signals may be stronger (Watt-Meyer et al., 2019) because there potentially exist fewer competing thermodynamic signals.

It is also unclear how climate change affects the seasonal cycle of dynamical features beyond the monsoons, which exhibit a well-documented delay in response to climate change (e.g., Seth et al., 2013) and the stratospheric polar vortex, which is projected to form earlier and decay later in the future (Ayarzaguena et al., 2020; Rao and Garfinkel 2021). Quantifying and understanding the seasonality of dynamical changes has important implications for impacts such as severe weather, ecosystems, forest fires, and agriculture.

5.2 **Move beyond the longitudinal and time mean**

Almost all of the dynamical signals in Table 1 reflect the time-mean. Circulation extremes have received only limited attention beyond blocking. Yet, recent work suggests the signal of climate change may be larger in the tails of the circulation distribution (Shaw & Miyawaki, 2024). It is also important to understand how circulation trends affect trends in other variables such as heat waves (Vautard et al., 2023).

Along similar lines, for a wide range of extremes and processes, there is much work to be done to understand how the dynamical response to climate change varies across different regions. For example, insights have been gained into recent trends by defining the Hadley Cell for different regional sectors (Nguyen et al., 2018; Staten et al., 2019; Hoskins et al., 2020; Gillett et al., 2021). The well-known model-observation discrepancy in tropical SST trends (Wills et al. 2022, Seager et al. 2022) represents an opportunity for understanding how tropical climate change affects regional circulation trends and this should be investigated further. Ultimately, teleconnections bridging different regions will change due to mean state changes under climate change and more work is needed to understand how.

5.3 **Use signals to test mechanisms and model fidelity**

Now that circulation signals are emerging, the dynamical mechanisms underlying the circulation trends can be compared to theoretical expectations and model predictions. Applying the numerous theoretical frameworks that have been proposed to explain dynamical responses to climate change (Vallis et al. 2015, Hoskins & Woollings 2015, Shaw, 2019, Wills et al. 2019) offers great potential for progress. Large ensemble, single forcing simulations (Smith et al. 2022) can also be leveraged to attribute observed circulation changes, to investigate whether internal variability involves dynamical mechanisms that are distinct from the forced response to anthropogenic climate change, to clarify the relative importance of different anthropogenic forcings, to showcase examples where models lack fidelity, to isolate and potentially correct signal-to-noise biases (section 4.2), and to directly examine how climate forcings affect the tails of the distribution (e.g. section 5.2).

5.4 Leverage the power of existing and emerging tools

Existing tools such as idealized models (Schemm & Röthlisberger, 2024; Jiménez-Esteve & Domeisen, 2022; Jiménez-Esteve et al, 2022), model hierarchies (Maher et al, 2019), mechanism denial experiments targeted toward understanding circulation signals and nudging (Hitchcock et al. 2022) are all powerful for understanding mechanisms and unraveling the relationship between circulation signals and other trends, or to understand the role of mean-state biases in the atmospheric circulation (e.g. Friesen et al., 2022). The impacts of known thermodynamic biases, e.g. SST trend biases, can be understood and quantified through targeted model experiments, e.g. using pacemaker simulations with coupled models (Kang et al. 2024).

Several new tools have emerged in the last decade that can be leveraged for making progress. Subseasonal to seasonal (S2S) forecasting models has emerged as a more widespread tool, with large ensembles of S2S forecasts that can be leveraged for understanding dynamical mechanisms and model-observation discrepancies. By pooling different ensemble members and different initializations for a given target forecast, and by assuming that atmospheric initial conditions are lost within the first month, tens of thousands of potential realizations of climate can be created (e.g. Kelder et al., 2020; Kolstad et al., 2022). This method could be exploited to improve mechanistic understanding of data-limited dynamical processes such as teleconnections. S2S ensemble forecasts can additionally be used to diagnose common model biases that also exist on climate timescales (L'Heureux et al., 2022; Garfinkel et al 2022; Lawrence et al, 2022; Beverley et al., 2023; Randall & Emanuel, 2024).

The use of AI/ML methods has exploded in the last few years. Physics-informed and explainable AI has the potential to advance our understanding of circulation signals (Connolly et al. 2023). In particular, these methods may be able to "learn" the source of discrepancies between models and observations, and structural uncertainties across different models.

Finally, high resolution global models going down to kilometer scale resolution present an exciting opportunity for understanding how large- and small-scale dynamics interact. In order to answer outstanding questions, carefully designed mechanistic model experiments across the model hierarchy are still crucial, which should be informed by results from new high-resolution (or large ensemble) model experiments. High resolution models also have the potential to reveal where model-observation discrepancies are the result of not properly representing small-scale dynamics in both the atmosphere and ocean.

A new era of climate change research is upon us, one where atmospheric circulation signals are emerging, attribution is becoming possible and puzzles and discrepancies are accumulating. There is an opportunity to embrace these signals and the puzzles they present, including cases where there is a lack of consensus, and use the signals as an opportunity to further advance our understanding of the climate system and improve predictions of regional climate change.

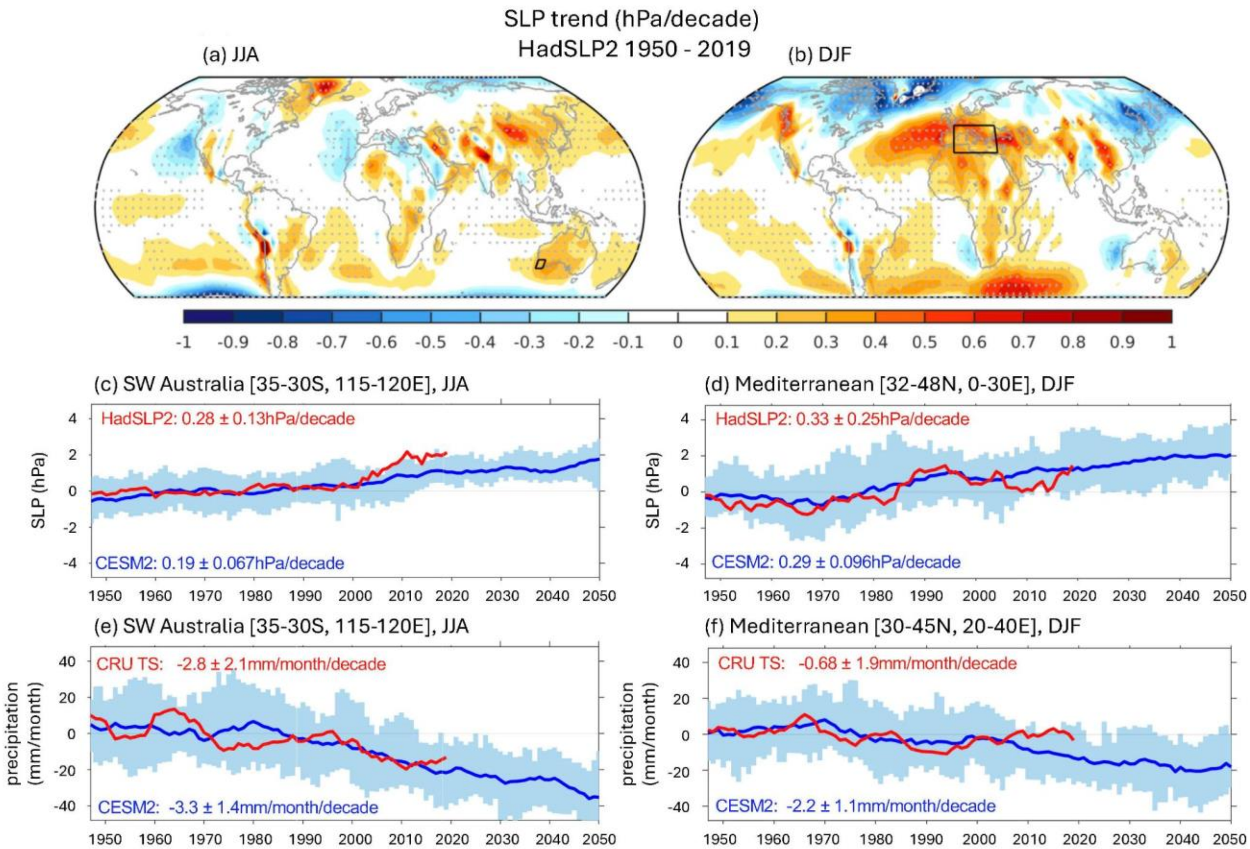


Figure 1: Regional circulation signals for JJA (left) and DJF (right). (a,b) Spatial structure of SLP trends [hPa/decade] from 1950-2019 in observations with stippling indicating statistically significant linear trends at the 0.05 level. Time series of (c) SLP [hPa] and (e) precipitation [mm/month] anomalies in observations (red line, HadSLPv2 for SLP, and CRU TS v4.07 for precipitation) over South-West Australia (black box in a) during JJA. (d,f) DJF SLP and precipitation over Mediterranean regions (black box in b) defined in Tuel and Eltahir (2020). Mean (blue line) and range (blue shading) of the 15-member historical-GHG only simulation in CESM2 of SLP and precipitation (Simpson et al 2023). All time series have been smoothed with a 10-year running mean.

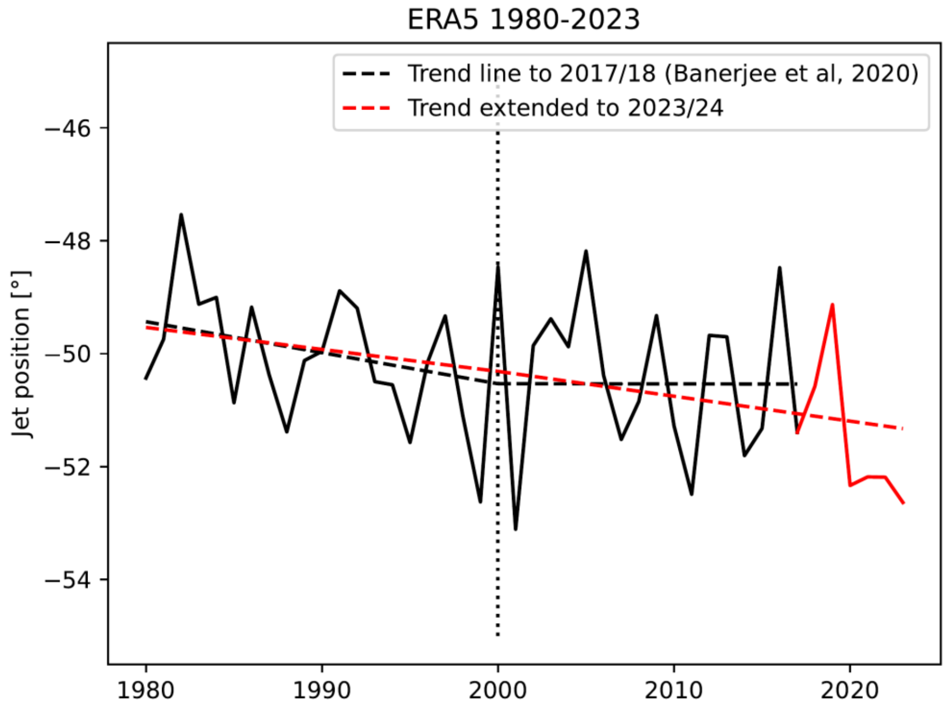


Figure 2: SH mid-latitude jet stream position response to ozone depletion. Jet position in DJF from ERA5, reproducing Banerjee et al. 2020, for years 1980/81-2017/18 (black lines), and extended time series to 2023/24 (red lines). Trends are fitted by continuous piecewise linear regression (following Banerjee et al), and trend values are -0.5°/dec for the ozone depletion period (1980/81 to 2000/01), and 0.0°/dec for 2000/01-2017/18. For the extended time series, trend values are -0.4°/dec for both ozone depletion and recovery periods, emphasizing the sensitivity of trend estimates from short records to end points.

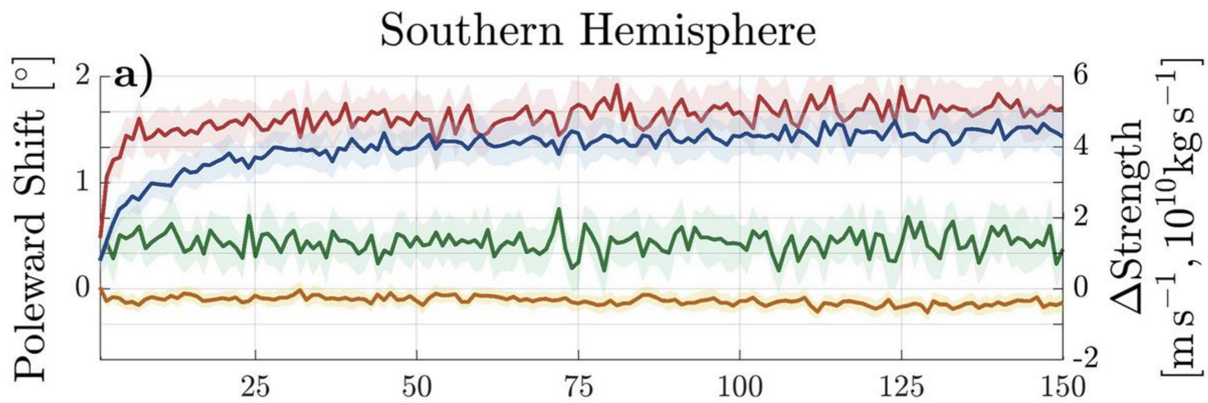


Figure 3: Time series of southern hemispheric response in model years to quadrupling atmospheric CO_2 for (a) the Hadley cell (HC) edge (red) and strength (orange) and the subtropical jet (STJ) location (green) and strength (blue). For each plot, shading represents the 95% confidence interval of model spread. Taken from Menzel et al. (2019).

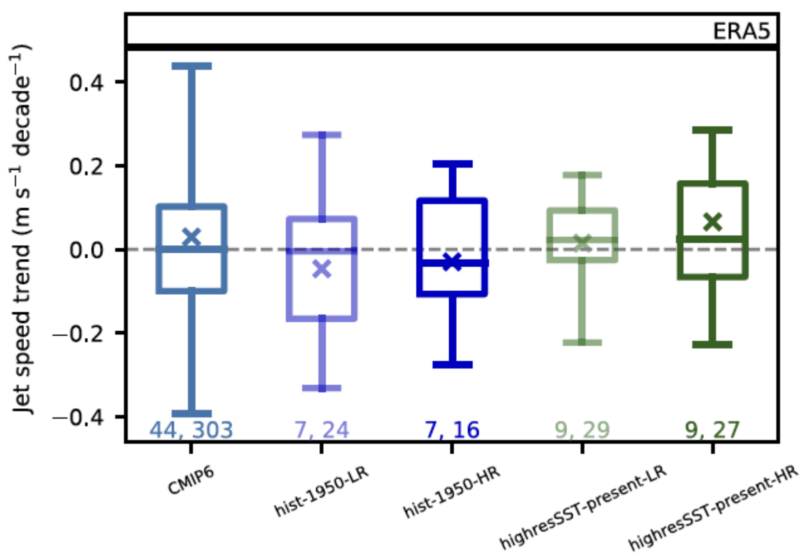


Figure 4: Trends in North Atlantic lower-tropospheric (700 hPa) jet stream strength from 1951-2014 in reanalysis data (ERA5) and across coupled (CMIP6) climate model ensemble, and low (LR) and high (HR) resolution HighResMIP climate model ensemble. The box represents upper and lower quartile ranges, and the whiskers represent the minimum and maximum from all ensemble members. The lines in the boxes indicate the median from all ensembles, and the crosses represent the multimodel mean. The two numbers at the bottom

indicate the total number of models (left) and total number of ensemble members (right) from each experiment. Taken from Blackport & Fyfe (2022).

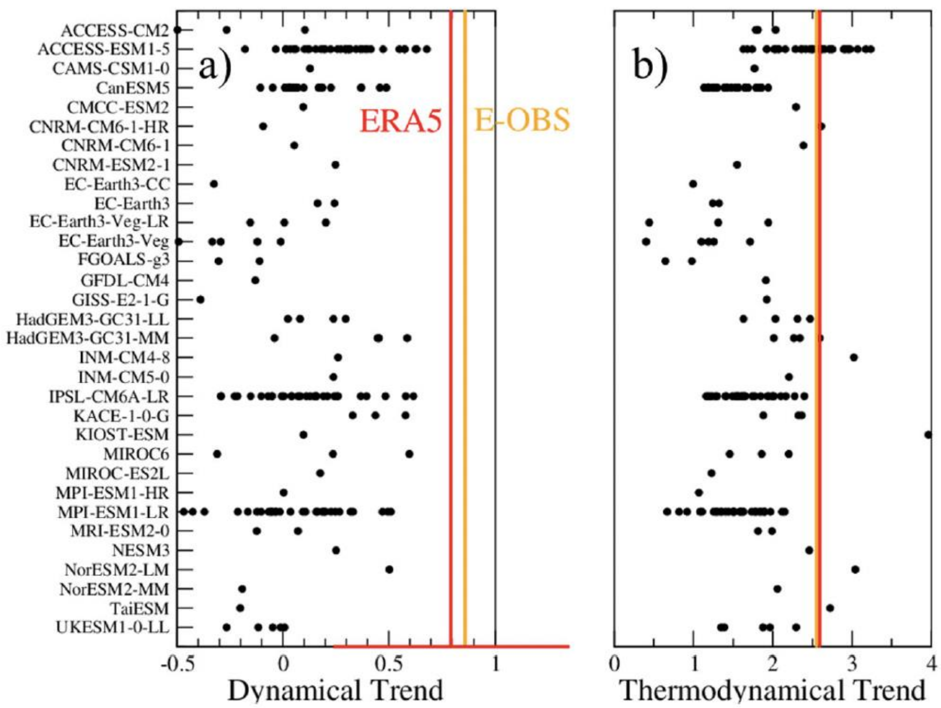


Figure 5: Dynamical (a) and thermodynamical (b) contributions to the summer TXx (summer maximum of maximal daily temperature) trends from ERA5 ECMWF Reanalysis (red line), E-OBS observation (orange line), and the 170 CMIP6 model simulations (names in ordinate) that were available (black dots) averaged over Western Europe. The thermodynamical contributions are simply calculated as residual by subtracting the dynamical trend from the total trend. For reference, the red bar at the bottom of (a) represents the 95% confidence interval of the estimate of the ERA5 TXx dynamical trend, estimated with a Gaussian assumption, i. e. the interval is calculated as plus or minus 2* the standard deviation (STD) of the error estimate on the trend coefficient. This confidence range describes the uncertainty related to the internal variability. This shows that this confidence range, calculated with the single realization of the observation, is consistent with the uncertainty range calculated from simulation members (respective standard deviations for observed trend and simulated trends of 0.28 and 0.25). Taken from Vautaurd et al. (2023).

Acknowledgments

The authors acknowledge the participants of the WCRP APARC DynVar/SNAP Workshop that took place 9-13 October 2023 in Munich, Germany. TAS acknowledges support from NSF (AGS-2300037) and NOAA (NA23OAR4310597). Support from the Swiss National Science Foundation through project PP00P2_198896 to DD is gratefully acknowledged. This project has received funding from the European Research Council (ERC) under the European Union's Horizon 2020 research and innovation programme (Grant agreement No. 847456). CIG acknowledges the support of the Israel Science Foundation (grant agreement 1727/21). KMG acknowledges support from NSF (AGS-2330009). JMA acknowledges support from the ARC Centre of Excellence for the Weather of the 21st Century (CE230100012) and partial support from the Regional and Global Model Analysis component of the Earth and Environmental System Modelling Program of the U.S. Department of Energy's Office of Biological and Environmental Research via National Science Foundation IA 1947282. AYK acknowledges support from the European Union's Horizon 2020 research and innovation framework programme under Grant agreement 774101003590 (PolarRES).

Open Research

No data was generated.

Conflict of Interest

The authors declare no conflicts of interest relevant to this study.

References

Allen, R. J., & Sherwood, S. C. (2008). Warming maximum in the tropical upper troposphere deduced from thermal winds. *Nature Geoscience*, *1*, 399.

Allen, R. J., Sherwood, S. C., Norris, J. R., & Zender, C. S. (2012). Recent Northern Hemisphere tropical expansion primarily driven by black carbon and tropospheric ozone. *Nature*, *485*(7398), 350–354. https://doi.org/10.1038/nature11097

Ayarzagüena, B., Charlton-Perez, A. J., Butler, A. H., Hitchcock, P., Simpson, I. R., Polvani, L. M., et al. (2020). Uncertainty in the response of sudden stratospheric warmings and stratosphere-troposphere coupling to quadrupled CO2 concentrations in CMIP6 models. *Journal of Geophysical Research: Atmospheres*, 125, e2019JD032345. https://doi.org/10.1029/2019JD032345

Banerjee, A., Fyfe, J. C., Polvani, L. M., Waugh, D., & Chang, K.-L. (2020). A pause in Southern Hemisphere circulation trends due to the Montreal Protocol. *Nature, 579*(7800), 544–548. https://doi.org/10.1038/s41586-020-2120-4

Beverley, J. D., Newman, M., & Hoell, A. (2023). Rapid development of systematic ENSO-related seasonal forecast errors. *Geophysical Research Letters*, 50, e2022GL102249. https://doi.org/10.1029/2022GL102249

Blackport, R., Screen, J. A., van der Wiel, K., & Bintanja, R. (2019). Minimal influence of reduced Arctic sea ice on coincident cold winters in mid-latitudes. Nature climate change, 9(9), 697-704.

Blackport, R., & Screen, J. A. (2021). Observed statistical connections overestimate the causal effects of Arctic sea ice changes on midlatitude winter climate. Journal of Climate, 34(8), 3021-3038.

Blackport, R., & Fyfe, J.C. (2022). Climate models fail to capture strengthening wintertime North Atlantic jet and impacts on Europe. *Science Advances*, *8*, eabn3112. https://doi.org/10.1126/sciadv.abn3112

Borowiak, A., A. King, T. Lane, The link between the Madden-Julian Oscillation and rainfall trends in northwest Australia, Geophysical Research Letters, 50 (2023), 10.1029/2022GL101799e2022GL101799

Breul, P., Ceppi, P., & Shepherd, T. G. (2023). Revisiting the wintertime emergent constraint of the southern hemispheric midlatitude jet response to global warming. *Weather and Climate Dynamics, 4,* 39–47. https://doi.org/10.5194/wcd-4-39-2023, 2023.

Butler, A. H., Thompson, D. W. J., & Heikes, R. (2010). The Steady-State Atmospheric Circulation Response to Climate Change–like Thermal Forcings in a Simple General Circulation Model. *Journal of Climate*, *23*(13), 3474–3496. https://doi.org/10.1175/2010JCLI3228.1

Ceppi, P., & Hartmann, D. L. (2016). Clouds and the Atmospheric Circulation Response to Warming. *Journal of Climate*, *29*(2), 783–799. https://doi.org/10.1175/JCLI-D-15-0394.1

Chan, C. J., & Plumb, R. A. (2009). The response to stratospheric forcing and its dependence on the state of the troposphere. Journal of the Atmospheric Sciences, 66(7), 2107-2115.

Chang, E. K. M., Ma, C.-G., Zheng, C. & Yau, A.M.W. (2016). Observed and projected decrease in Northern Hemisphere extratropical cyclone activity in summer and its impacts on maximum temperature. *Geophysical Research Letters*, *43*, 2200–2208. https://doi.org/10.1002/2016GL068172

Chemke, R., & Polvani, L. M. (2019b). Opposite tropical circulation trends in climate models and in reanalyses. *Nature Geoscience*, *12*(7), 528–532. https://doi.org/10.1038/s41561-019-0383-x

Chemke, R., & Polvani, L. M. (2019). Exploiting the Abrupt 4 x CO2 Scenario to Elucidate Tropical Expansion Mechanisms. *J. Climate*, *32*, 859–875.

Chemke, R., & Polvani, L. M. (2021). Elucidating the Mechanisms Responsible for Hadley Cell Weakening Under 4 × CO 2 Forcing. *Geohys. Res. Lett*, 10.1029/2020GL090348

Chemke, R., & Yuval, J. (2023). Human-induced weakening of the Northern Hemisphere tropical circulation. *Nature 617*, 529–532. https://doi.org/10.1038/s41586-023-05903-1

Chemke, R., & Coumou, D. (2024). Human influence on the recent weakening of storm tracks in boreal summer. *npj Climate and Atmos. Sci.* 10.1038/s41612-024-00640-2

Chemke, R., Polvani, L. M., Kay, J. E., & Orbe, C. (2021). Quantifying the role of ocean coupling in Arctic amplification and sea-ice loss over the 21st century. *Npj Climate and Atmospheric Science*, *4*(1), 1–9. https://doi.org/10.1038/s41612-021-00204-8

Chemke, R., Ming, Y. & Yuval, J. (2022). The intensification of winter mid-latitude storm tracks in the Southern Hemisphere. *Nature Climate Change, 12*, 553–557. https://doi.org/10.1038/s41558-022-01368-8

Chung, E.-S., Timmermann, A., Soden, B. J., Ha, K.-J., Shi, L., & John, V. O. (2019). Reconciling opposing Walker circulation trends in observations and model projections. *Nature Climate Change*, *9*(5), 405–412. https://doi.org/10.1038/s41558-019-0446-4

Connolly, C., E. A. Barnes, P. Hassanzadeh, M. Pritchard (2023) Using neural networks to learn the jet stream forced response from natural variability, Artificial Intelligence for the Earth Systems 2 (2), e220094

Coumou, D., Di Capua, G., Vavrus, S., Wang, L., & Wang, S. (2018). The influence of Arctic amplification on mid-latitude summer circulation. *Nature Communications*, *9*(1), 2959. https://doi.org/10.1038/s41467-018-05256-8

Coumou, D., Lehmann, J., & Beckmann, J. (2015). The weakening summer circulation in the Northern Hemisphere mid-latitudes. *Science*, *348*(6232), 324–327. https://doi.org/10.1126/science.1261768

Cox, T., Donohoe, A., Armour, K. C., Frierson, D. M. W. and G. H. Roe. (2024). Trends in Atmospheric Heat Transport Since 1980. J. Climate, doi: 10.1175/JCLI-D-23-0385.1.

Curtis, P. E., Ceppi, P., & Zappa, G. (2020). Role of the mean state for the Southern Hemispheric jet stream response to CO2 forcing in CMIP6 models. *Environmental Research Letters*, *15*, 064011. https://doi.org/10.1088/1748-9326/ab8331

Deser, C., Lehner, F., Rodgers, K.B., Ault, T., Delworth, T.L., DiNezio, P.N., et al. (2020). Insights from Earth system model initial-condition large ensembles and future prospects. *Nature Climate Change*, *10*, 277–286. https://doi.org/10.1038/s41558-020-0731-2

Domeisen, D. I. V., Eltahir, E. A. B., Fischer, E. M., Knutti, R., Perkins-Kirkpatrick, S. E., Schär, C., Seneviratne, S. I., Weisheimer, A., & Wernli, H. (2023). Prediction and projection of heatwaves. *Nature Reviews Earth & Environment*, *4*(1), 36–50. https://doi.org/10.1038/s43017-022-00371-z

Dong, B., Sutton, R. T., Shaffrey, L., & Harvey, B. (2022). Recent decadal weakening of the summer Eurasian westerly jet attributable to anthropogenic aerosol emissions. *Nature Communications*, *13*(1), 1148. https://doi.org/10.1038/s41467-022-28816-5

Elbaum, E., Garfinkel, C. I., Adam, O., Morin, E., Rostkier-Edelstein, D., & Dayan, U. (2022). Uncertainty in projected changes in precipitation minus evaporation: Dominant role of dynamic circulation changes and weak role for thermodynamic changes. *Geophysical Research Letters*, 49, e2022GL097725. https://doi.org/10.1029/2022GL097725

Eyring, V., et al., 2021: Human Influence on the Climate System. In Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, pp. 423–552, doi:10.1017/9781009157896.005.

Franzke, C. L. E., & Harnik, N. (2023). Long-Term Trends of the Atmospheric Circulation and Moist Static Energy Budget in the JRA-55 Reanalysis. *Journal of Climate, 36*(9), 2959-2984. https://doi.org/10.1175/JCLI-D-21-0724.1

Garfinkel, C. I., Son, S. W., Song, K., Aquila, V. and Oman, L. D. (2013) The Effect of Tropospheric Jet Latitude on Coupling between the Stratospheric Polar Vortex and the Troposphere. J. Climate, 26, 2077–2095.

Garfinkel, C. I., Son, S. W., Song, K., Aquila, V. and Oman, L. D. (2017) Stratospheric variability contributed to and sustained the recent hiatus in Eurasian winter warming. Geophysical research letters 44, no. 1: 374-382.

Garfinkel, C. I., Schwartz, C., Domeisen, D. I. V., Son, S.-W., Butler, A. H., & White, I. P. (2018). Extratropical atmospheric predictability from the quasi-biennial oscillation in subseasonal forecast models. *Journal of Geophysical Research: Atmospheres*, 123, 7855–7866. https://doi.org/10.1029/2018JD028724

Garfinkel, C. I., Chen, W., Li, Y., Schwartz, C., Yadav, P., & Domeisen, D. (2022). The Winter North Pacific Teleconnection in Response to ENSO and the MJO in Operational Subseasonal Forecasting Models Is Too Weak. *Journal of Climate*, *35*(24), 8013-8030. https://doi.org/10.1175/JCLI-D-22-0179.1

Garfinkel, C. I., Keller, B., Lachmy, O., White, I., Gerber, E. P., Jucker, M., & Adam, O. (2024). Impact of parameterized convection on the storm track and near-surface jet response to global warming: Implications for mechanisms of the future poleward shift. *Journal of Climate*, *37*, 2541–2564. https://doi.org/10.1175/JCLI-D-23-0105.1

Garny, H. , F. Ploeger , M. Abalos, H. Bönisch, A. Castillos, T. von Clarmann, M. Diallo , A. En-gels , J. C. Laube , M. Linz, J. Neu , A. Podglajen , E. Ray, L. Rivoire, L. N. Saunders , G. Stiller, F. Voet, K. A. Walker: Age of stratospheric air: Progress on processes, observations and long-term trends, *submitted to Reviews of Geophysics.*

Gertler, C.G., & O'Gorman, P.A. (2019). Changing available energy for extratropical cyclones and associated convection in Northern Hemisphere summer. *Proceedings of the National Academy of Sciences*, *116*, 4105-4110. https://doi.org/10.1073/pnas.1812312116

Ghosh, S., Lachmy, O., & Kaspi, Y. (2024). The role of diabatic heating in the midlatitude atmospheric circulation response to climate change. *Journal of Climate*, *1*(aop). https://doi.org/10.1175/JCLI-D-23-0345.1

Gillett, N. P., J. C. Fyfe, and D. E. Parker (2013), Attribution of observed sea level pressure trends to greenhouse gas, aerosol, and ozone changes, Geophys. Res. Lett., 40, 2302–2306, doi:10.1002/grl.50500.

Gillett, N. P., Shiogama, H., Funke, B., Hegerl, G., Knutti, R., Matthes, K., Santer, B. D., Stone, D., & Tebaldi, C. (2016). The Detection and Attribution Model Intercomparison Project (DAMIP v1.0) contribution to CMIP6. *Geoscientific Model Development*, *9*(10), 3685–3697. https://doi.org/10.5194/gmd-9-3685-2016

Gillett, Z. E., H. H. Hendon, J. M. Arblaster, and E.-P. Lim. (2021). Tropical and extratropical influences on the variability of the Southern Hemisphere wintertime subtropical jet. *Journal of Climate*, **34**, 4009–4022, https://doi.org/10.1175/JCLI-D-20-0460.1.

Grise, K. M. (2022). Atmospheric circulation constraints on 21st century seasonal precipitation storylines for the southwestern United States. *Geophysical Research Letters*, 49, e2022GL099443. https://doi.org/10.1029/2022GL099443

Grise, K. M., Davis, S.M., Simpson, I.R., Waugh, D.W., Fu, Q., Allen, R.J., et al. (2019). Recent tropical expansion: Natural variability or forced response? *Journal of Climate*, *32*, 1551–1571. https://doi.org/10.1175/JCLI-D-18-0444.1

Hall, A., & Qu, X. (2006). Using the current seasonal cycle to constrain snow albedo feedback in future climate change. *Geophysical Research Letters*, *33*, L03502. https://doi.org/10.1029/2005GL025127.

Hanna, E., Fettweis, X., & Hall, R. J. (2018). Brief communication: Recent changes in summer Greenland blocking captured by none of the CMIP5 models. *The Cryosphere*, *12*, 3287–3292. https://doi.org/10.5194/tc-12-3287-2018

Heidemann, H., Cowan, T., Henley, B.J., Ribbe, J., Freund, M. & Power, S. (2023) Variability and long-term change in Australian monsoon rainfall: a review. WIREs Climate Change, 14, e823.

Hitchcock, P. and Coauthors, 2022: Stratospheric Nudging And Predictable Surface Impacts (SNAPSI): a protocol for investigating the role of stratospheric polar vortex disturbances in subseasonal to seasonal forecasts, Geo. Mod. Dev, 15, 5073–5092

Held, I. M. (1993), Large-Scale dynamics and climate change, Bull. Amer. Met. Soc., 74, doi:10.1175/1520-0477.

Hope, P. K., Drosdowsky, W., and Nicholls, N. (2006): Shifts in the synoptic systems influencing southwest Western Australia, *Clim. Dynam.*, 26, 751–764, https://doi.org/10.1007/s00382-006-0115-y.

Hoskins, B., Woollings, T. (2015) Persistent Extratropical Regimes and Climate Extremes. *Curr Clim Change Rep* **1**, 115–124. https://doi.org/10.1007/s40641-015-0020-8

Hoskins, B.J., Yang, G.-Y., & Fonseca, R.M. (2020). The detailed dynamics of the June–August Hadley Cell. *Quarterly Journal of the Royal Meteorological Society*, *146*, 557–575. https://doi.org/10.1002/qj.3702

Huang, C. S. Y., & Nakamura, N. (2016). Local Finite-Amplitude Wave Activity as a Diagnostic of Anomalous Weather Events. *Journal of the Atmospheric Sciences*, *73*(1), 211-229. https://doi.org/10.1175/JAS-D-15-0194.1

IPCC (2021). *Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change* [Masson-Delmotte, V., P. Zhai, A. Pirani, S.L. Connors, C. Péan, S. Berger, N. Caud, Y. Chen, L. Goldfarb, M.I. Gomis, M. Huang, K. Leitzell, E. Lonnoy, J.B.R. Matthews, T.K. Maycock, T. Waterfield, O. Yelekçi, R. Yu, and B. Zhou (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, In press, doi:10.1017/9781009157896.

Kang, J. M., Shaw, T. A., & Sun, L. (2023). Arctic Sea Ice Loss Weakens Northern Hemisphere Summertime Storminess but Not Until the Late 21st Century. *Geophysical Research Letters*, *50*(9), e2022GL102301. https://doi.org/10.1029/2022GL102301

Kang, J. M., Shaw, T. A., Kang, S., Simpson, I. R. & Yu, Y. (2024). Revisiting the reanalysis-model discrepancy in Southern Hemisphere winter storm track trends. 10.22541/essoar.171224128.81410474/v1

Kang, J. M., Shaw, T. A. and Sun, L. (2024b). Anthropogenic aerosols have significantly weakened the regional summertime circulation in the Northern Hemisphere during the satellite era. 10.22541/essoar.171901442.29170078/v1

Kang, S.M., Yu, Y., Deser, C., Zhang, X., Kang, I.-S., Lee, S.-S., Rodgers, K.B., & Ceppi, P. (2023). Global impacts of recent Southern Ocean cooling. *Proceedings of the National Academy of Sciences, 120*, e2300881120. https://doi.org/10.1073/pnas.230088112

Karpechko, A. Yu., Wu, Z., Simpson, I. R., Kretschmer, M., Afargan-Gerstman, H., Butler, A. H., Domeisen, D. I.V., Garny, H., Lawrence, Z, Manzini, E., & Sigmond, M.: Northern Hemisphere Stratosphere-Troposphere Circulation Change in CMIP6 1 Models. Part 2: Mechanisms and Sources of the Spread (2024), *Journal of Geophysical Research: Atmospheres,* doi: 10.1029/2024JD040823

Kelder, T., Müller, M., Slater, L.J., Marjoribanks, T.I., Wilby, R.L., Prudhomme, C., Bohlinger, P., Ferranti, L., & Nipen, T. (2020). Using UNSEEN trends to detect decadal changes in 100-year precipitation extremes. *npj Climate and Atmospheric Science*, *3*, 47. https://doi.org/10.1038/s41612-020-00149-4

Kidston, J., & Gerber, E.P. (2010). Intermodel variability of the poleward shift of the austral jet stream in the

CMIP3 integrations linked to biases in 20th century climatology. *Geophysical Research Letters*, *37*, L09708. https://doi.org/10.1029/2010GL042873.

Kidston, J., Scaife, A., Hardiman, S. *et al.* Stratospheric influence on tropospheric jet streams, storm tracks and surface weather. *Nature Geosci* **8**, 433–440 (2015). https://doi.org/10.1038/ngeo2424

Kornhuber, K., Petoukhov, V., Karoly, D., Petri, S., Rahmstorf, S., & Coumou, D. (2017). Summertime Planetary Wave Resonance in the Northern and Southern Hemispheres. *Journal of Climate*, *30*(16), 6133–6150. https://doi.org/10.1175/JCLI-D-16-0703.1

Kolstad, E. W., Lee, S. H., Butler, A. H., Domeisen, D. I. V., & Wulff, C. O. (2022). Diverse surface signatures of stratospheric polar vortex anomalies. *Journal of Geophysical Research: Atmospheres*, *127*, e2022JD037422. https://doi.org/10.1029/2022JD037422

Knutson, T. R. & Ploshay, J. Sea level pressure trends: model-based assessment of detection, attribution, and consistency with CMIP5 historical simulations. J. Clim. 34, 327–346 (2021).

Lachmy, O. (2022). The relation between the latitudinal shifts of midlatitude diabatic heating, eddy heat flux and the eddy-driven jet in CMIP6 models. *Journal of Geophysical Research: Atmospheres*, *127*, e2022JD036556. https://doi.org/10.1029/2022JD036556

Lawrence, Z. D., Abalos, M., Ayarzagüena, B., Barriopedro, D., Butler, A. H., Calvo, N., ... & Wu, R. W. Y. (2022). Quantifying stratospheric biases and identifying their potential sources in subseasonal forecast systems. Weather and Climate Dynamics, Weather Clim. Dynam., 3, 977–1001

Lee, S., & Feldstein, S.B. (2013). Detecting Ozone- and Greenhouse Gas–Driven Wind Trends with Observational Data. *Science, 339*,563-567. https://doi.org/10.1126/science.1225154

Lee, S., L'Heureux, M., Wittenberg, A.T., Seager, R., O'Gorman, P.A., & Johnson, N.C. (2022). On the future zonal contrasts of equatorial Pacific climate: Perspectives from Observations, Simulations, and Theories. *npj Climate and Atmospheric Science*, *5*, 82. https://doi.org/10.1038/s41612-022-00301-2

Lee, S. H., Williams, P. D., & Frame, T. H. A. (2019). Increased shear in the North Atlantic upper-level jet stream over the past four decades. *Nature*, *572*(7771), 639–642. https://doi.org/10.1038/s41586-019-1465-z

L'Heureux, M.L., Tippett, M.K., & Wang, W. (2022). Prediction Challenges From Errors in Tropical Pacific Sea Surface Temperature Trends. *Frontiers in Climate*, *4*, 837483. https://doi.org/10.3389/fclim.2022.837483

Lionello, P., D'Agostino, R., Ferreira, D., Nguyen, H., & Singh, M. S. (2024). The Hadley circulation in a changing climate. *Ann NY Acad Sci.*, 1534, 69–93. https://doi.org/10.1111/nyas.15114

Lorenz, D. J., & DeWeaver, E. T. (2007). Tropopause height and zonal wind response to global warming in the IPCC scenario integrations. *Journal of Geophysical Research: Atmospheres*, *112*(D10). https://doi.org/10.1029/2006JD008087

Lorenz, D. J., & Hartmann, D. L. (2001). Eddy–Zonal Flow Feedback in the Southern Hemisphere. *Journal of the Atmospheric Sciences*, *58*(21), 3312-3327. https://doi.org/10.1175/1520-0469(2001)058<3312:EZFFIT>2.0.CO;2

Lu, J., Leung, L. R., Yang, Q., Chen, G., Collins, W. D., Li, F., Hou, Z. J., & Feng, X. (2014). The robust dynamical contribution to precipitation extremes in idealized warming simulations across model resolutions. *Geophysical Research Letters*, *41*, 2971–2978. https://doi.org/10.1002/2014GL059532

Maher, P., Gerber, E. P., Medeiros, B., Merlis, T. M., Sherwood, S., Sheshadri, A., ... & Zurita-Gotor, P. (2019). Model hierarchies for understanding atmospheric circulation. Reviews of Geophysics, 57(2), 250-280.

Maher, N., Milinski, S., & Ludwig, R. (2021). Large ensemble climate model simulations: introduction, overview, and future prospects for utilising multiple types of large ensemble. *Earth System Dynamics*, *12*, 401–418. https://doi.org/10.5194/esd-12-401-2021

Manabe, S., & Wetherald, R. T. (1975). The Effects of Doubling the CO2 Concentration on the climate of a General Circulation Model. *Journal of the Atmospheric Sciences*, *32*(1), 3–15. https://doi.org/10.1175/1520-0469(1975)032<0003:TEODTC>2.0.CO;2

Mann, M. E., Rahmstorf, S., Kornhuber, K., Steinman, B. A., Miller, S. K., Petri, S., & Coumou, D. (2018). Projected changes in persistent extreme summer weather events: The role of quasi-resonant amplification. *Science Advances*, *4*(10). https://doi.org/10.1126/sciadv.aat3272

Maraun, D., Shepherd, T. G., Widmann, M., Zappa, G., Walton, D., Gutierrez, J. M., et al. (2017). Towards process-informed bias correction of climate change simulations. *Nature Climate Change*, *7*(11), 764–773. https://doi.org/10.1038/nclimate3418

Martineau, P., Behera, S. K., Nonaka, M., Nakamura, H. & Kosaka, Y. (2024), ‘Seasonally dependent increases in subweekly temperature variability over Southern Hemisphere landmasses detected in multiple reanalyses’, Weather and Climate Dynamics 5(1), 1–15.

Menzel, M. E., Waugh, D., & Grise, K. (2019). Disconnect Between Hadley Cell and Subtropical Jet Variability and Response to Increased CO2. *Geophysical Research Letters*, *46*(12), 7045–7053. https://doi.org/10.1029/2019GL083345

Mindlin, J., Shepherd, T.G., Vera, C.S., Osman, M., Zappa, G., Lee, R.W., & Hodges, K.I. (2020). Storyline description of Southern Hemisphere midlatitude circulation and precipitation response to greenhouse gas forcing. *Climate Dynamics, 54*, 4399–4421. https://doi.org/10.1007/s00382-020-05234-1

Needham, M. R., & Randall, D. A. (2023). Anomalous Northward Energy Transport due to Anthropogenic Aerosols during the Twentieth Century. *Journal of Climate*, *36*(19), 6713–6728. https://doi.org/10.1175/JCLI-D-22-0798.1

Nguyen, H., Hendon, H. H., Lim, E. P., Boschat, G., Maloney, E., & Timbal, B. (2018). Variability of the extent of the Hadley circulation in the southern hemisphere: a regional perspective. *Climate Dynamics*, *50*(1-2), 129-142. https://doi.org/10.1007/s00382-017-3592-2

O'Gorman, P. A. (2010). Understanding the varied response of the extratropical storm tracks to climate change. *Proceedings of the National Academy of Sciences*, *107*(45), 19176–19180. https://doi.org/10.1073/pnas.1011547107

Outten, S., Li C., King, M. P., Suo, L., Siew, P. Y., Davy, R., Dunn-Sigouin, E. et al. (2022) Reconciling conflicting evidence for the cause of the observed early 21st century Eurasian cooling. Weather and Climate Dynamics Discussions 2022: 1-32.

Perlwitz, J. Tug of war on the jet stream. *Nature Clim. Change* 1, 29–31 (2012).

Randall, D. A., & Emanuel, K. (2024). The Weather–Climate Schism. *Bulletin of the American Meteorological Society*, *105*(1), E300-E305. https://doi.org/10.1175/BAMS-D-23-0124.1

Rao, J., & Garfinkel, C. I. (2021). Projected changes of stratospheric final warmings in the Northern and Southern Hemispheres by CMIP5/6 models. *Climate Dynamics*. https://doi.org/10.1007/s00382-021-05647-6

Rousi, E., Kornhuber, K., Beobide-Arsuaga, G., Luo, F., & Coumou, D. (2022). Accelerated western European heatwave trends linked to more-persistent double jets over Eurasia. *Nature Communications*, *13*(1), 3851. https://doi.org/10.1038/s41467-022-31432-y

Russo, E., & Domeisen, D. I. V. (2023). Increasing Intensity of Extreme Heatwaves: The Crucial Role of Metrics. *Geophysical Research Letters*, *50*(14), e2023GL103540. https://doi.org/10.1029/2023GL103540

Schemm, S., & Röthlisberger, M. (2024). Aquaplanet simulations with winter and summer hemispheres: model setup and circulation response to warming. *Weather and Climate Dynamics*, *5*, 43–63. https://doi.org/10.5194/wcd-5-43-2024

Schmidt, D. F., & Grise, K. M. (2021). Drivers of Twenty-First-Century U.S. Winter Precipitation Trends in CMIP6 Models: A Storyline-Based Approach. *Journal of Climate, 34*(16), 6875-6889. https://doi.org/10.1175/JCLI-D-21-0080.1

Schumacher, D.L., Singh, J., Hauser, M. *et al.* (2024) Exacerbated summer European warming not captured by climate models neglecting long-term aerosol changes. *Commun Earth Environ* **5**, 182. https://doi.org/10.1038/s43247-024-01332-8

Seager, R., Henderson, N., & Cane, M. (2022). Persistent Discrepancies between Observed and Modeled Trends in the Tropical Pacific Ocean. *Journal of Climate*, *35*(14), 4571-4584. https://doi.org/10.1175/JCLI-D-21-0648.1

Seager, R., Wu, Y., Cherchi, A., Simpson, I. R., Osborn, T. J., Kushnir, Y., Lukovic, J., Liu, H., & Nakamura, J. (2024). Recent and near-term future changes in impacts-relevant seasonal hydroclimate in the world's Mediterranean climate regions. International Journal of Climatology, 1–29. https://doi.org/10.1002/joc.8551

Seth, A., Rauscher, S. A., Biasutti, M., Giannini, A., Camargo, S. J., & Rojas, M. (2013). CMIP5 Projected Changes in the Annual Cycle of Precipitation in Monsoon Regions. *Journal of Climate, 26*(19), 7328-7351. https://doi.org/10.1175/JCLI-D-12-00726.1

Shaw, T. A. (2019). Mechanisms of Future Predicted Changes in the Zonal Mean Mid-Latitude Circulation. *Current Climate Change Reports, 5*(4), 345–357. https://doi.org/10.1007/s40641-019-00145-8

Shaw, T. A., P. Arias, M. Collins and Coauthors (2024), Regional Climate Change: consensus, discrepancies, and ways forward, Frontiers in Climate, 10.3389/fclim.2024.1391634.

Shaw, T. A., Baldwin, M., Barnes, E. A., Caballero, R., Garfinkel, C. I., Hwang, Y.-T., Li, C., O’Gorman, P. A., Riviere, G., Simpson, I. R., & Voigt, A. (2016). Storm track processes and the opposing influences of climate change. *Nature Geoscience*, *9*(9), 656–664.

Shaw, T. A., Barpanda, P., & Donohoe, A. (2018). A moist static energy framework for zonal-mean storm track intensity. *Journal of the Atmospheric Sciences*, *75*, 1979–1994. https://doi.org/10.1175/JAS-D-17-0183.1

Shaw, T. A., Miyawaki, O., & Donohoe, A. (2022). Stormier Southern Hemisphere induced by topography and ocean circulation. *Proceedings of the National Academy of Sciences*, *119*(50), e2123512119. https://doi.org/10.1073/pnas.2123512119

Shaw, T.A., & Miyawaki, O. (2024). Fast upper-level jet stream winds get faster under climate change. *Nature Climate Change*, *14*, 61–67. https://doi.org/10.1038/s41558-023-01884-1

Shaw, T. A., & Tan, Z. (2018). Testing Latitudinally Dependent Explanations of the Circulation Response to Increased CO2 Using Aquaplanet Models. *Geophysical Research Letters*, *45*, 9861–9869. https://doi.org/10.1029/2018GL078974

Shaw, T. A., Barpanda, P. and A. Donohoe. (2018). A Moist Static Energy Framework for Zonal-Mean Storm-Track Intensity. *J. Atmos. Sci.*, *75*, 1979–1994. https://doi.org/10.1175/JAS-D-17-0183.1

Shaw, T. A., & Voigt, A. (2016). What can moist thermodynamics tell us about circulation shifts in response to uniform warming? *Geophysical Research Letters*, *43*(9), 4566–4575. https://doi.org/10.1002/2016GL068712

Shepherd, T. G. (2014). Atmospheric circulation as a source of uncertainty in climate change projections. *Nature Geoscience*, *7*(10), 703–708. https://doi.org/10.1038/ngeo2253

Shrestha, S., & Soden, B. J. (2023). Anthropogenic weakening of the atmospheric circulation during the satellite era. *Geophysical Research Letters*, 50, e2023GL104784. https://doi.org/10.1029/2023GL104784

Simpson, I. R., & Polvani, L.M. (2016). Revisiting the relationship between jet position, forced response, and annular mode variability in the southern midlatitudes. *Geophysical Research Letters*, *43*, 2896–2903. https://doi.org/10.1002/2016GL067989.

Simpson, I., Seager, R., Ting, M. & Shaw, T.A. (2016). Causes of change in Northern Hemisphere winter meridional winds and regional hydroclimate. *Nature Climate Change*, *6*, 65–70. https://doi.org/10.1038/nclimate2783

Simpson, I. R., McKinnon, K. A., Davenport, F. V., Tingley, M., Lehner, F., Al Fahad, A., & Chen, D. (2021). Emergent Constraints on the Large-Scale Atmospheric Circulation and Regional Hydroclimate: Do They Still Work in CMIP6 and How Much Can They Actually Constrain the Future?. *Journal of Climate, 34*(15), 6355-6377. https://doi.org/10.1175/JCLI-D-21-0055.1

Simpson, I. R., Rosenbloom, N., Danabasoglu, G., Deser, C., Yeager, S. G., McCluskey, C. S., ... & Rodgers, K. B.

(2023). The CESM2 single-forcing large ensemble and comparison to CESM1: Implications for experimental design. Journal of Climate, 36(17), 5687-5711.

Smith, D. et al. (2022). Attribution of multi-annual to decadal changes in the climate system: The Large Ensemble Single Forcing Model Intercomparison Project (LESFMIP), Frontiers Climate, 10.3389/fclim.2022.955414

SPARC, 2022: SPARC Reanalysis Intercomparison Project (S-RIP) Final Report. Masatomo Fujiwara, Gloria L. Manney, Lesley J. Gray, and Jonathon S. Wright (Eds.), SPARC Report No. 10, WCRP-6/2021, doi: 10.17874/800dee57d13, available at www.sparc-climate.org/publications/sparc-reports.

Staten, P. W., Grise, K. M., Davis, S. M., Karnauskas, K., & Davis, N. (2019). Regional widening of tropical overturning: Forced change, natural variability, and recent trends. *Journal of Geophysical Research: Atmospheres*, 124, 6104–6119. https://doi.org/10.1029/2018JD030100

Sun, X., Ding, Q., Wang, S.-Y. S., Topál, D., Li, Q., Castro, C., Teng, H., Luo, R., & Ding, Y. (2022). Enhanced jet stream waviness induced by suppressed tropical Pacific convection during boreal summer. *Nature Communications*, *13*(1), 1288. https://doi.org/10.1038/s41467-022-28911-7

Tamarin-Brodsky, T., & Kaspi, Y. (2017). Enhanced poleward propagation of storms under climate change. *Nature Geoscience*, *10*(12), 908–913. https://doi.org/10.1038/s41561-017-0001-8

Tan, Z., & Shaw, T. A. (2020). Quantifying the Impact of Wind and Surface Humidity-Induced Surface Heat Exchange on the Circulation Shift in Response to Increased CO2. *Geophysical Research Letters*, *47*(18), e2020GL088053. https://doi.org/10.1029/2020GL088053

Teng, H., Leung, R., Branstator, G., Lu, J., & Ding, Q. (2022). Warming Pattern over the Northern Hemisphere Midlatitudes in Boreal Summer 1979–2020. *Journal of Climate*, *35*(11), 3479–3494. https://doi.org/10.1175/JCLI-D-21-0437.1

Thompson, D. W. J., Solomon, S., Kushner, P. J., England, M. H., Grise, K. M., & Karoly, D. J. (2011). Signatures of the Antarctic ozone hole in Southern Hemisphere surface climate change. *Nature Geoscience*, *4*(11), 741–749. https://doi.org/10.1038/ngeo1296

Tuel, A., & Eltahir, E. A. B. (2020). Why Is the Mediterranean a Climate Change Hot Spot?. *Journal of Climate, 33*(14), 5829-5843. https://doi.org/10.1175/JCLI-D-19-0910.1

Vallis, G. K., P. Zurita-Gotor, C. Cairns, J. Kidston, (2015) Response of the large-scale structure of the atmosphere to global warming, Q. J. Roy. Met. Soc, https://doi.org/10.1002/qj.2456.

Vautard, R., Cattiaux, J., Happé, T., Singh, J., Bonnet, R., Cassou, C., Coumou, D., D'Andrea, F., Faranda, D., Fischer, E., Ribes, A., Sippel, S., & Yiou, P. (2023). Heat extremes in Western Europe increasing faster than simulated due to atmospheric circulation trends. *Nature Communications*, *14*(1), Article 1. https://doi.org/10.1038/s41467-023-42143-3

Voigt, A., Albern, N., Ceppi, P., Grise, K., Li, Y., & Medeiros, B. (2021). Clouds, radiation, and atmospheric circulation in the present-day climate and under climate change. *WIREs Climate Change*, *12*, e694. https://doi.org/10.1002/wcc.694

Voigt, A., & Shaw, T. A. (2015). Circulation response to warming shaped by radiative changes of clouds and water vapour. *Nature Geoscience*, *8*(2), 102–106. https://doi.org/10.1038/ngeo2345

Watt-Meyer, O., Frierson, D. M. W., & Fu, Q. (2019). Hemispheric Asymmetry of Tropical Expansion Under CO2 Forcing. *Geophysical Research Letters*, *46*, 9231–9240. https://doi.org/10.1029/2019GL083695

Weisheimer, A., Baker, L. H., Bröcker, J., Garfinkel, C. I., Hardiman, S. C., Hodson, D. L., Palmer, T. N., Robson, J.

I., Scaife, A. A., Screen, J. A., Shepherd, T. G., Smith, D. M., & Sutton, R. T. (2024). The Signal-to-Noise Paradox in Climate Forecasts: Revisiting our Understanding and Identifying Future Priorities. *Bulletin of the American Meteorological Society*, https://doi.org/10.1175/BAMS-D-24-0019.1

Williams, R. S., Marshall, G. J., Levine, X., Graff, L. S., Handorf, D., Johnston, N. M., Karpechko, A. Y., Orr, A., Van de Berg, W. J., Wijngaard, R. R., & Mooney, P. A. (2024). Future Antarctic Climate: Storylines of Midlatitude Jet Strengthening and Shift Emergent from CMIP6. *Journal of Climate*, *37*(7), 2157-2178. https://doi.org/10.1175/JCLI-D-23-0122.1

Wills, R. C. J., Dong, Y., Proistosecu, C., Armour, K. C., & Battisti, D. S. (2022). Systematic Climate Model Biases in the Large-Scale Patterns of Recent Sea-Surface Temperature and Sea-Level Pressure Change. *Geophysical Research Letters*, *49*(17), e2022GL100011. https://doi.org/10.1029/2022GL100011

Wills, R.C.J., White, R.H. & Levine, X.J. (2019) Northern Hemisphere Stationary Waves in a Changing Climate. Curr Clim Change Rep **5**, 372–389. https://doi.org/10.1007/s40641-019-00147-6

WMO (World Meteorological Organization), Scientific Assessment of Ozone Depletion: 2018, Global Ozone Research and Monitoring Project–Report No. 58, 588 pp., Geneva, Switzerland, 2018.

Woollings, T., Drouard, M., O'Reilly, C.H., Sexton, D.M.H., & McSweeney, C. (2023). Trends in the atmospheric jet streams are emerging in observations and could be linked to tropical warming. *Communications Earth & Environment, 4*, 125. https://doi.org/10.1038/s43247-023-00792-8

Wu, Y., Simpson, I. R., & Seager, R. (2019). Intermodel spread in the Northern Hemisphere stratospheric polar vortex response to climate change in the CMIP5 models. *Geophysical Research Letters*, 46, 13290–13298. https://doi.org/10.1029/2019GL085545

Yang D., Arblaster J. M., Meehl G. A. and England M. H. (2021). The role of coupled feedbacks in the decadal variability of the Southern hemisphere Eddy-driven jet, *J. Geophys. Res*. 126 e2021JD035023. https://doi.org/10.1029/2021JD035023

Yeager, S.G., Chang, P., Danabasoglu, G., Rosenbloom, N., Zhang, Q., Castruccio, F.S., Gopal, A., Rencurrel, M.C., & Simpson, I.R. (2023). Reduced Southern Ocean warming enhances global skill and signal-to-noise in an eddy-resolving decadal prediction system. *npj Climate and Atmospheric Science, 6*, 107. https://doi.org/10.1038/s41612-023-00434-y

Zambri, B., Solomon, S., Thompson, D. W. J., & Fu, Q. (2021). Emergence of Southern Hemisphere stratospheric circulation changes in response to ozone recovery. *Nature Geoscience*. https://doi.org/10.1038/s41561-021-00803-3

Zaplotnik, Ž., Pikovnik, M., & Boljka, L. (2022). Recent Hadley Circulation Strengthening: A Trend or Multidecadal Variability?. *Journal of Climate, 35*(13), 4157-4176. https://doi.org/10.1175/JCLI-D-21-0204.1

Zappa, G., & Shepherd, T. G. (2017). Storylines of Atmospheric Circulation Change for European Regional Climate Impact Assessment. *Journal of Climate*, *30*(16), 6561-6577. https://doi.org/10.1175/JCLI-D-16-0807.1

Zappa, G., P. Ceppi, and T. G. Shepherd. (2020) Time-evolving sea-surface warming patterns modulate the climate change response of subtropical precipitation over land. Proceedings of the National Academy of Sciences 117, no. 9: 4539-4545.

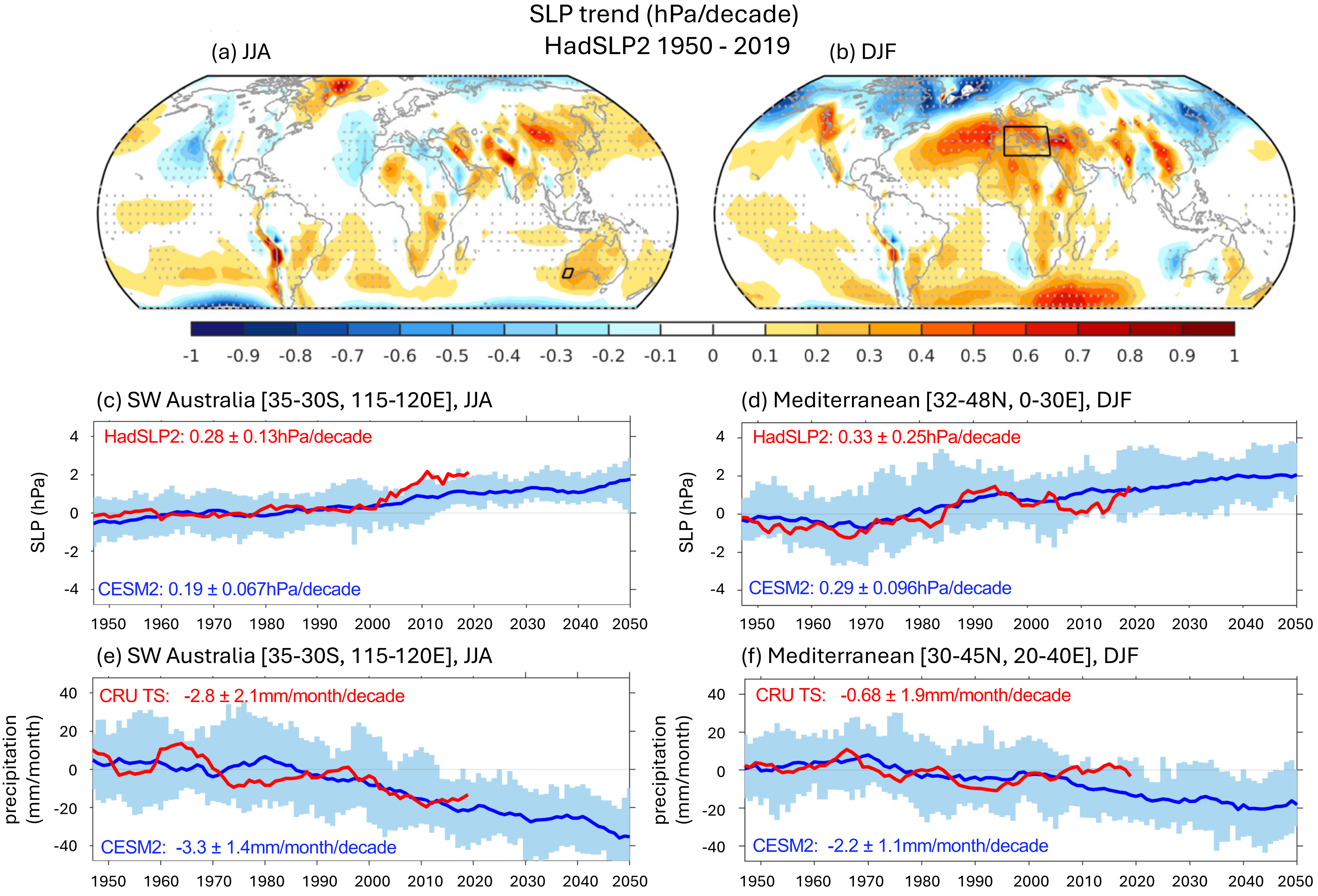
Zhao, X., & Allen, R.J. (2019). Strengthening of the Walker Circulation in recent decades and the role of natural sea surface temperature variability. *Environmental Research Communications*, *1*, 021003. https://doi.org/10.1088/2515-7620/ab0dab.

Table 1. Emerging Circulation signals

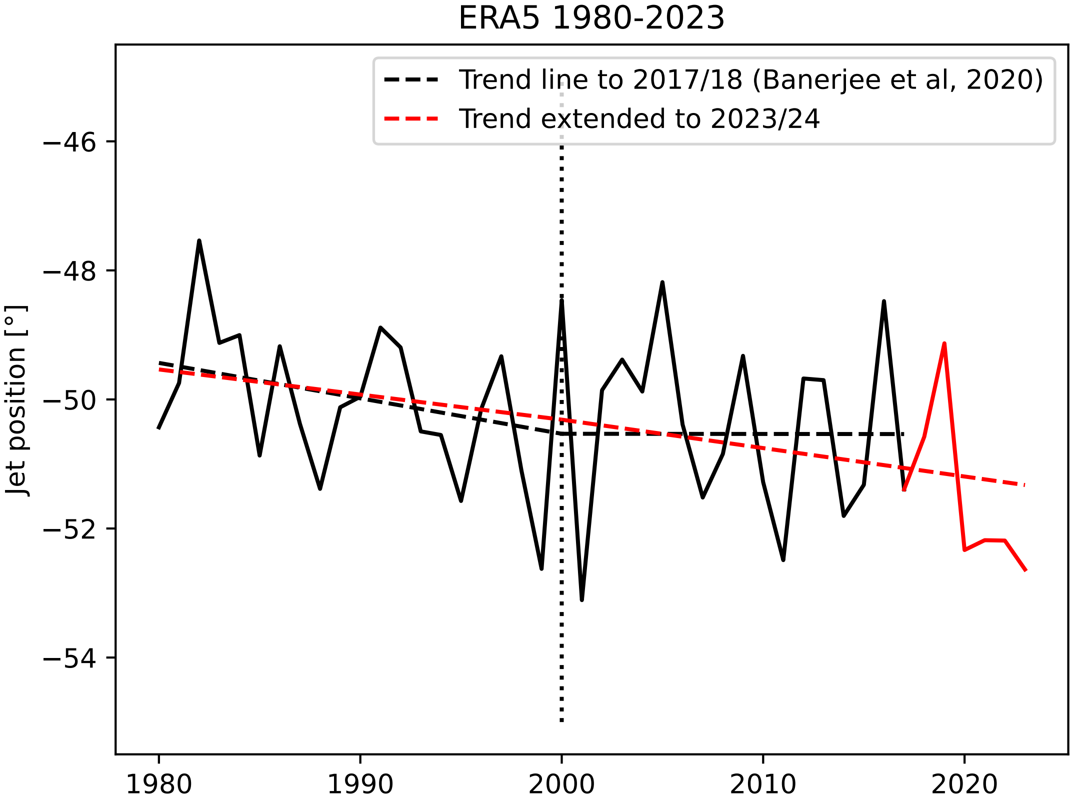
Atmospheric circulation signals (statistically significant long term trends) that have been reported in the literature. Following IPCC terminology signals are labeled detected if the likelihood of occurrence by chance due to internal variability is small and attributed if the causal human driver (greenhouse gas, aerosol, ozone forcing, etc.) has been determined.

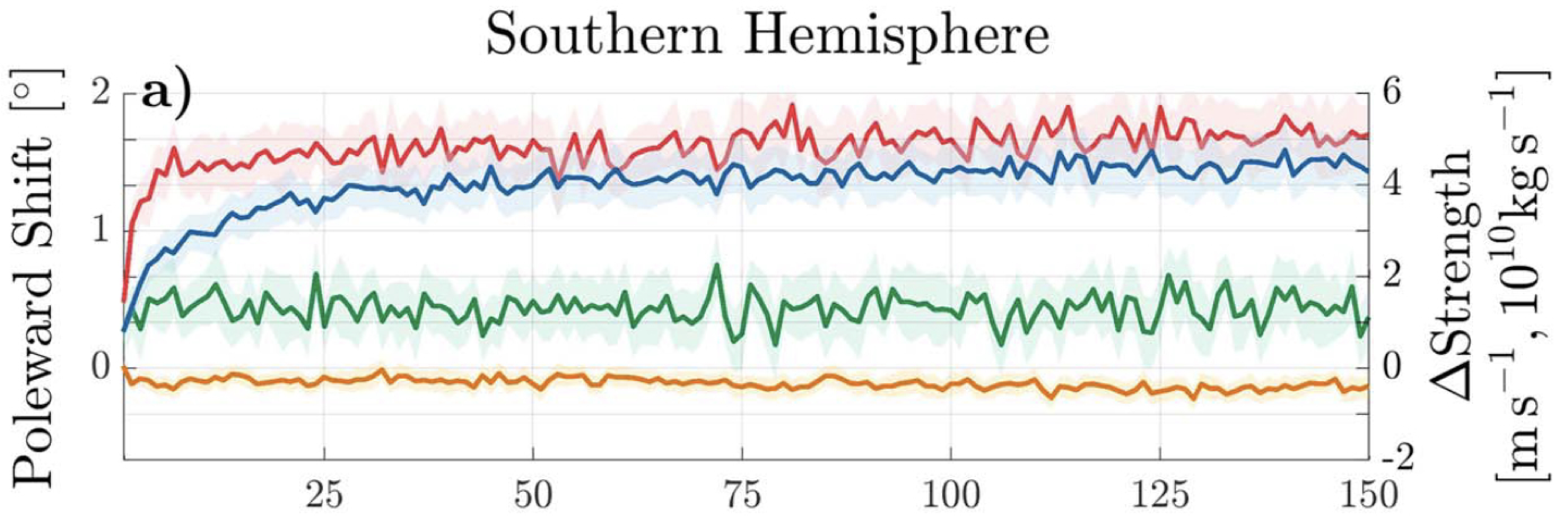
Signal	**Region**	**Season**	**Reference**	**Detected**	**Attributed**
Increased wind shear (zonal wind)	North Atlantic	Annual mean	Lee et al. (2019)		
Upper-troposphere jet strengthening (zonal wind)	Zonal-mean	DJF	Woollings et al. (2023), Franzke & Harnik (2023)		
Lower-troposphere jet strengthening (zonal wind, mean sea level pressure)	North Atlantic	DJF	Blackport & Fyfe (2022), Wills et al. (2022)		
Lower-troposphere jet poleward shift (zonal wind)	Zonal-mean	DJF	Lee & Feldstein (2013), Woollings et al. (2023)	x	x
Mid-troposphere jet weakening (zonal wind)	N. Hemisphere Zonal-mean	JJA	Coumou et al. (2015), Kang et al. (2024b)	x	x
Upper-troposphere jet weakening (zonal wind)	Eurasia	JJA	Dong et al. (2022)	x	x
Storm track weakening (eddy kinetic energy)	N. Hemisphere Zonal-mean	JJA	Coumou et al. (2015, Chang et al. (2016), Gertler & O'Gorman (2019), Kang et al. (2023), Cox et al. (2024), Chemke & Coumou (2024)	x	x
Extratropical cyclone activity weakening (mean sea level pressure)	North Atlantic, North Pacific	JJA	Kang et al. (2024b)	x	x
Increased blocking (500 hPa geopotential height)	Greenland	JJA	Hanna et al. (2018)		
Storm track strengthening (eddy kinetic energy)	S. Hemisphere Zonal-mean	JJA	Chemke et al. (2022)	x	
Storm track strengthening (eddy kinetic energy)	S. Hemisphere Zonal-mean	Annual mean	Shaw et al. (2022), Cox et al. (2024)		

Hadley cell shift (mass stream function)	S. Hemisphere Zonal-mean	Annual mean	Grise et al. (2019), Lionello et al. (2024)	x	x
Hadley cell intensity (mass stream function)	N. Hemisphere Zonal-mean	Annual mean	Chemke & Yuval (2023), Lionello et al. (2024)	x	x
Walker circulation strengthening (mean sea level pressure, surface winds)	Pacific	Annual mean	Chung et al. (2019), Zhao and Allen (2019)	x	
Weakening of upward vertical motion (500 hPa vertical motion)	Global	Annual mean	Shrestha & Soden (2023)	x	x
Increasing sea level pressure	South west Western Australia	JJA	Hope et al (2006); Knutson & Ploshay (2021); Figure 1c	x	
Increasing stationary wave amplitude (200 hPa geopotential height)	N. Hemisphere (200 hPa geopotential height)	JJA	Teng et al. (2022), Sun et al. (2022)		
Strengthening stationary waves (sea level pressure)	Mediterranean	DJF	Tuel & Eltahir (2020); Figure 1d	x	
Strengthening summer Monsoon	N. Hemisphere	JJA	Eyring et al. (2021)		
	Australian	DJF	Borowiak et al (2023); Heidemann et al (2023)	x	

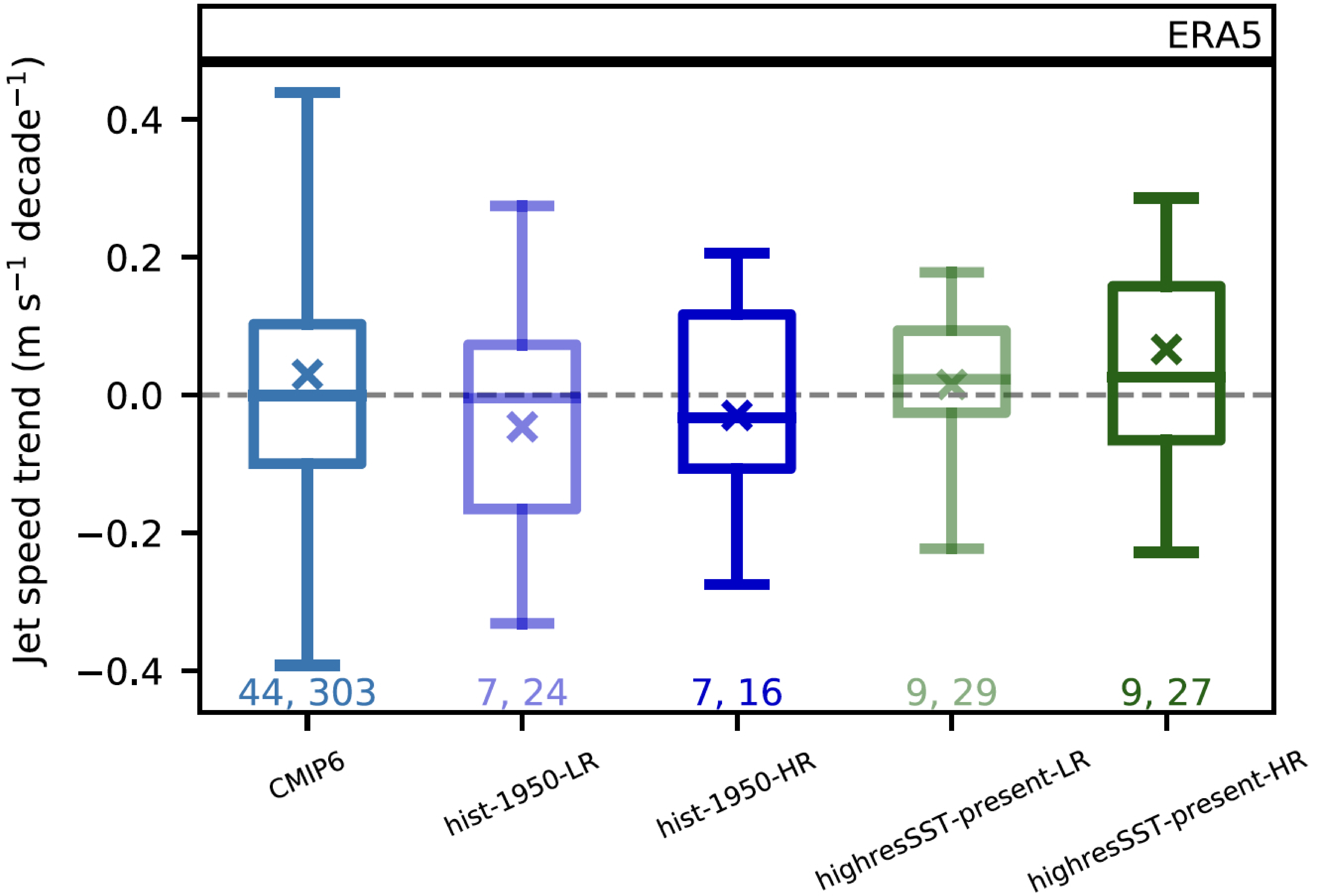


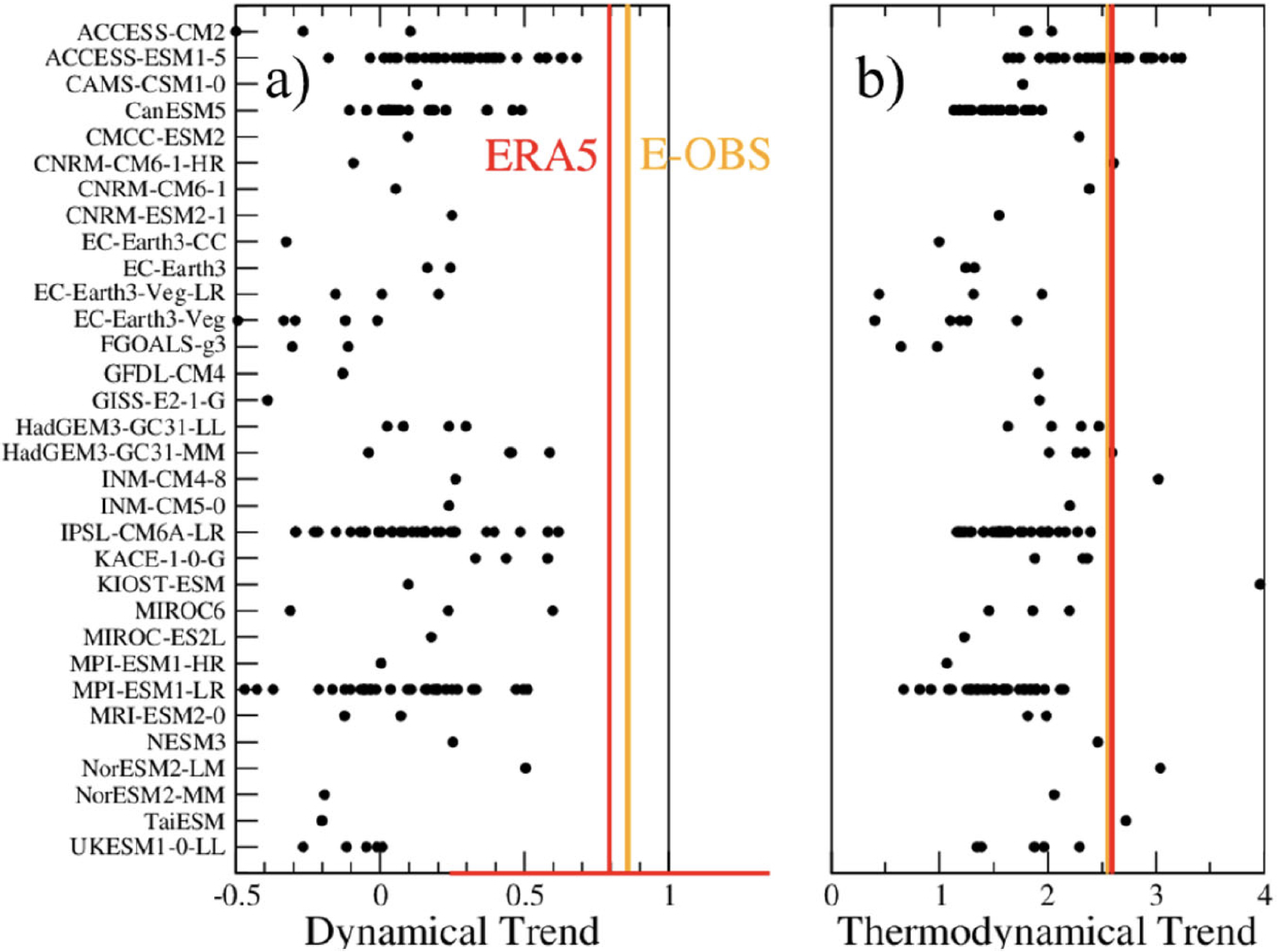
SLP trend (hPa/decade)
HadSLP2 1950 - 2019
(a) JJA
(b) DJF
-1 -0.9 -0.8 -0.7 -0.6 -0.5 -0.4 -0.3 -0.2 -0.1 0 0.1 0.2 0.3 0.4 0.5 0.6 0.7 0.8 0.9 1
(c) SW Australia [35-30S, 115-120E], JJA
HadSLP2: 0.28 ± 0.13hPa/decade
CESM2: 0.19 ± 0.067hPa/decade
SLP (hPa)
(d) Mediterranean [32-48N, 0-30E], DJF
HadSLP2: 0.33 ± 0.25hPa/decade
CESM2: 0.29 ± 0.096hPa/decade
(e) SW Australia [35-30S, 115-120E], JJA
CRU TS: -2.8 ± 2.1mm/month/decade
CESM2: -3.3 ± 1.4mm/month/decade
precipitation (mm/month)
(f) Mediterranean [30-45N, 20-40E], DJF
CRU TS: -0.68 ± 1.9mm/month/decade
CESM2: -2.2 ± 1.1mm/month/decade
1950 1960 1970 1980 1990 2000 2010 2020 2030 2040 2050


ERA5 1980-2023
Trend line to 2017/18 (Banerjee et al, 2020)
Trend extended to 2023/24
Jet position [°]
−46
−48
−50
−52
−54
1980
1990
2000
2010
2020



Southern Hemisphere
a)
Poleward Shift [°]
ΔStrength [m s−1, 10^10 kg s−1]
2
1
0
6
4
2
0
-2
25
50
75
100
125
150


ERA5
Jet speed trend (m s^{-1} decade^{-1})
0.4
0.2
0.0
−0.2
−0.4
44, 303
7, 24
7, 16
9, 29
9, 27
CMIP6
hist-1950-LR
hist-1950-HR
highresSST-present-LR
highresSST-present-HR



ACCESS-CM2
ACCESS-ESM1-5
CAMS-CSM1-0
CanESM5
CMCC-ESM2
CNRM-CM6-1-HR
CNRM-CM6-1
CNRM-ESM2-1
EC-Earth3-CC
EC-Earth3
EC-Earth3-Veg-LR
EC-Earth3-Veg
FGOALS-g3
GFDL-CM4
GISS-E2-1-G
HadGEM3-GC31-LL
HadGEM3-GC31-MM
INM-CM4-8
INM-CM5-0
IPSL-CM6A-LR
KACE-1-0-G
KIOST-ESM
MIROC6
MIROC-ES2L
MPI-ESM1-HR
MPI-ESM1-LR
MRI-ESM2-0
NESM3
NorESM2-LM
NorESM2-MM
TaiESM
UKESM1-0-LL
a)
b)
ERA5
E-OBS
-0.5
0
0.5
1
0
1
2
3
4
Dynamical Trend
Thermodynamical Trend