Strong linkage between observed daily precipitation extremes and anthropogenic emissions across the contiguous United States

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Abstract

The results of probabilistic event attribution studies depend on the choice of the extreme value statistics used in the analysis, particularly with the arbitrariness in the selection of appropriate thresholds to define extremes. We bypass this issue by using the Extended Generalized Pareto Distribution (ExtGPD), which jointly models low precipitation with a generalized Pareto distribution and extremes with a different Pareto tail, to conduct daily precipitation attribution across the contiguous United States (CONUS). We apply the ExtGPD to 12 general circulation models from the Coupled Model Intercomparison Project Phase 6 and compare counterfactual scenarios with and without anthropogenic emissions. Observed precipitation by the Climate Prediction Centre is used for evaluating the GCMs. We find that greenhouse gases rather than natural variability can explain the observed magnitude of extreme daily precipitation, especially in the temperate regions. Our results highlight an unambiguous linkage of anthropogenic emissions to daily precipitation extremes across CONUS.

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2	anthropogenic emissions across the contiguous United States
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In this study, we investigate how human-induced emissions affect daily rainfall extremes across
the United States. The attribution of an extreme event to human-induced emissions depends on
the selected extreme event statistics, with setting a threshold to define what counts as an extreme
event remaining a major challenge. To overcome this, we used the Extended Generalized Pareto

28 Distribution (ExtGPD) that jointly models both low and heavy rainfall events without defining a

threshold, providing a more complete picture of the full distribution including extremes. We

30 fitted the ExtGPD to 12 general circulation models and compared scenarios with and without

human-induced emissions. Our findings suggest that human emissions are responsible for the

32 observed intensity of daily rainfall extremes across the United States, especially in regions with

temperate climates, and that these extremes would have been smaller without greenhouse gases.

34 Highlights

- We apply Extended Generalized Pareto Distribution for probabilistic event attribution as
 it bypasses issues with threshold specification.
- Anthropogenic emissions are responsible for the observed magnitude of extreme daily
 precipitation across the United States.
- The study underscores the urgent need for mitigation, revealing a clear link between
 anthropogenic activities and extreme precipitation.

41 **1. Introduction**

Climate attribution studies that examine the role of anthropogenic climate change in altering the 42 probability of observed weather extremes have proliferated since Allen (2003), who proposed a 43 simple probabilistic framework for detecting the role of anthropogenic forcing in the occurrence 44 of an observed extreme. The availability of large ensemble climate model simulations and the 45 Detection and Attribution Model Intercomparison Project (DAMIP) gave further impetus to 46 climate attribution studies. The goal of these studies is to quantify the contribution of historic 47 emissions and natural forcing in altering the risk of observed extremes and to project their 48 changes (e.g., Gillett et al., 2016). This is generally accomplished using frequentist probabilistic 49 methods (Oldenborgh (2021) and references therein). Most of these studies are event-specific 50 attribution of observed extraordinary weather extremes in the recent years, while a few of them 51 evaluate the role of anthropogenic emissions in altering the observed changes in extremes based 52 on fingerprint techniques (e.g., Risser et al., 2022, Kirchmeier-Young & Zhang, 2020). 53 54 Previous studies have reported an observed intensification of precipitation extremes in the central

and eastern parts of the contiguous United States (CONUS) and no detectable change is reported

56 in the western United States (Easterling et al., 2017; Guo et al., 2019; Trenberth, 2018).

57 However, the existing event attribution studies have reported a low to medium confidence in human attribution to observed precipitation extremes across the CONUS (Seneviratne et al., 58 59 2021). For instance, event attribution studies found that the 3-day rainfall that caused the Louisiana floods of 2016 had become 40% more likely since 1900 (Van Der Wiel et al., 2017) 60 and extreme precipitation associated with Hurricane Harvey in August 2017 in Houston was 61 intensified due to global warming (e.g., Wang et al., 2018; Wehner & Sampson, 2021; Zhao et 62 al., 2018). Similarly, external forcing reportedly caused the intensification and increased 63 frequency of 1-day and 5-day annual maximum precipitation across North America based on 64 optimal fingerprinting techniques (Kirchmeier-Young et al., 2020). On the other hand, an 65 unequivocal role of human forcing was not detected in the 2013 Colorado heavy rainfall events 66 (Hoerling et al., 2014; Pall et al., 2017). Most of the existing attribution studies are based on 67 specific observed precipitation extremes spanning a few days in duration and impacting a 68 specific region of interest. Moreover, the use of different approaches in the extreme event 69 definition and statistical modelling makes a direct comparison of these studies difficult. 70

The definition of extremes is an important criterion in determining the apparent role of 71 72 anthropogenic climate change (e.g. Philip et al., 2020). Kirchmeier-Young et al. (2019) suggested that longer spatial and temporal scales increase the signal-to-noise ratio of extreme weather 73 74 events, making it easier to attribute longer- than shorter-scale events to human influences. The statistical approaches used for determining the changes in the frequency of extremes also 75 influence the attribution statements (Naveau et al., 2020). The block maxima or peak-over-76 77 threshold (POT) approaches are usually employed to select the extreme events that are used for 78 the subsequent statistical modelling. While a block maxima approach reduces the sample size of observed extremes to one per year, the POT method is constrained by the arbitrary choice of the 79 threshold (e.g., Nerantzaki et al., 2023). The threshold selection in a POT method represents a 80 tradeoff between ensuring an adequate sample size for statistical testing and adhering to the 81 assumptions of extreme value distributions. The different approaches in event definition and 82 arbitrary choice of thresholds make it difficult to interpret the existing studies. This necessitates 83 the use of statistical approaches that are not dependent on threshold specification for climate 84

85 attribution studies.

To address this issue, we employ the extended Generalized Pareto distribution (ExtGPD: Naveau 86 et al., 2016), which jointly models low precipitation with a generalized Pareto distribution and 87 heavy rainfall with a different Pareto tail, for building the statistical framework for climate 88 attribution. The ExtGPD helps in modelling the entire range of data without the need for 89 specifying a threshold, thereby increasing the sample size needed for fitting the statistical model. 90 Here, we use the ExtGPD for providing a comprehensive statement on the role of anthropogenic 91 forcing on altering the risk of daily precipitation extremes by sampling the entire precipitation 92 time series across CONUS. We use counterfactual simulations from 12 general circulation 93 models (GCMs) that participated in the DAMIP of the 6th Coupled Model Intercomparison 94 Project (CMIP6) for generating the attribution statements. We organize the study into three 95 sections: (1) evaluation of the performance of the GCMs in capturing daily precipitation 96 extremes; (2) attribution of human contribution to daily precipitation extremes by comparing 97 historical simulations with counterfactual scenarios; and (3) estimation of the sensitivity of 98 99 anthropogenic forcing to the magnitude of extremes and climatic regions.

100 **<u>2. Data and Methods</u>**

101 **2.1 Data**

We used daily precipitation by the Climate Prediction Centre (CPC) available at 0.25° resolution 102 over CONUS. For the evaluation and attribution study, we used 12 CMIP6 GCMs (Table S1) 103 simulations from hist-nat, hist-GHG and historical scenarios. The hist-nat simulations are based 104 on natural forcing (solar irradiation and volcanic aerosols) and exclude anthropogenic forcing 105 (anthropogenic aerosols and emissions). The hist-GHG simulations are forced by well-mixed 106 greenhouse gas emissions and exclude both natural and anthropogenic aerosols. The historical 107 108 scenarios of GCMs are based on anthropogenic (GHG and aerosols) and natural forcing (volcanic and solar). The historical simulations of the GCMs are forced from 1850-2014 (Gillett 109 110 et al., 2016). We used 34-year simulations from 1981-2014 to focus on the most recent decades 111 and to mitigate potential issues related to the presence of non-stationarities. We regridded the CPC daily precipitation and climate model simulations to 1-degree resolution using bilinear 112 interpolation to ensure consistency among the datasets. We excluded daily gridded precipitation 113 below 0.1 mm of both CPC and CMIP6 datasets for the analysis to remove drizzle bias in the 114 GCMs (e.g., Dai et al., 2006; DeMott et al., 2007). 115

116 **2.2. Methods**

117 **2.2.1.** Extreme value modelling

118 The generalized Pareto distribution (GPD) family is widely used for modelling extreme

119 precipitation exceeding a particular threshold as it is appropriate for modelling heavy tail

distributions (e.g., Coles, 2001) (equation 1); however, the threshold estimation is challenging. A

121 large threshold could reduce the sample size and leads to higher uncertainty in the parameter

estimations, while a smaller threshold does not satisfy the approximations of the GPD leading to

model errors (e.g., Rivoire et al., 2021). Naveau et al. (2016) proposed a transition function to

the GPD that provides a smooth connection between the upper tail and the main body of the

125 function. This approach, called the extended GPD (ExtGPD), avoids the need for threshold

specification and helps in sampling the entire timeseries for modelling extremes. This approach

has been applied and tested in various contexts (e.g. Haruna et al., 2023; Gamet & Jalbert, 2022;

128 Rivoire et al., 2021; Legrand et al., 2023).

129 The probability distribution of the GPD when the shape ξ (upper tail parameter) is larger than 130 zero (i.e., heavy tail behavior) is:

131
$$H(x_{\sigma}) = 1 - \left(1 + \frac{\xi x}{\sigma}\right)^{-1/\xi} \text{ for } \xi > 0$$
(1)

132 where, σ is the scale parameter of the GPD distribution and x is the random variable.

133 The ExtGPD is a transformation of the GPD, such that:

134
$$G(v) = v^k; \text{ where, } v = H(\chi/\sigma)$$
(2)

135 and k > 0, is the lower tail parameter.

136 Naveau et al. (2016) proposed four different statistical formulations for the extended GPD

137 model. We use the ExtGPD with three parameters (i.e., scale, lower tail parameter and the shape

138 or upper tail parameter) owing to its simplicity and convergence of the statistical framework

139 (equation 2). We estimate the parameters using the probability weighted moments method as it

140 converges better than the maximum likelihood estimation (Naveau et al., 2016). We use the

- following initial values for the parameters: shape $\xi = 0.2$; scale, $\sigma = 1$; lower tail parameter, k =
- 142 0.5. We fit the ExtGPD to CPC and GCM daily precipitation values above 0.1 mm for each 1-

143 degree grid and estimate daily precipitation extremes of various return periods. We evaluate the

- 144 performance of the GCMs in simulating observed precipitation based on the ability of the
- 145 historical GCM simulations to capture CPC precipitation extremes: for each grid, we select only
- 146 those models whose *n*-year return level of precipitation in historical simulations falls within the
- 147 95% confidence interval of the *n*-year return level of the CPC observations. Thus, we identify a
- subset of good-performing GCMs for each grid, which we use in the subsequent event
- 149 attribution.

150 **2.2.2.** Climate attribution of extreme precipitation.

151 We use the 34-year counterfactual scenarios of hist-nat and hist-GHG (1981-2014) from the

subset of GCMs to make a probabilistic attribution statement. We estimated the ratio of change

in return level of extreme precipitation of the counterfactual scenarios to that of the historical

simulations to quantify the impact of anthropogenic emissions (equation 3).

155
$$Attribution Ratio = \frac{P_{Counterfactual} - P_{historical}}{P_{historical}},$$
(3)

156 where *counterfactual* stands for both hist–nat and hist–GHG.

A ratio greater (smaller) than zero implies that the counterfactual scenario increases (decreases) the *n*-year return level precipitation compared to the historical forcing. Further, we evaluate the sensitivity of the attribution ratio to different return periods and climate zones. The CONUS mainly falls within three climatic zones: arid, temperate and cold as per Köppen-Geiger climate classification (Beck et al., 2018). We separately assess the attribution ratio for each of the three climate zones. The analysis for the ExtGPD is performed in R using the mev and gmm libraries (Belzile, 2015; Chaussé, 2010).

164 **3. Results**

165 **3.1. Evaluation of CMIP6 GCMs**

166 We validated the performance of the GCMs by comparing their ability to capture the observed

167 extremes. The spatial pattern of extremes is well captured by CESM2, CanESM5, E3SM-2-0,

168 FGOALs-g3 and MRI-ESM2-0 (Figure 1b). The largest values in precipitation extremes based

169 on CPC (Figure 1a) tend to be concentrated along the Gulf Coast and to decrease as we move

inland northward, and on the Sierra Navada in the west; most GCMs can capture this gradient,

- 171 even though some of them fail to do so. For instance, ACCESS-CM2, ACCESS-ESM1-5, BCC-CSM2-MR, GFDL-ESM4 and MIROC6 overestimate extremes in the central part of the study 172 173 region. We observed the GCMs to perform well in simulating the extreme rainfall patterns in the U.S. Northeast, West Coast, the Great Plains of the North and South, and southeastern regions, 174 which receive relatively high extreme precipitation (Figure 1). We also examined the 175 performance of GCMs in capturing higher return level precipitation (Figures S1-S3) and found 176 similar spatial patterns to those in the 100-year return period. Overall, most GCMs capture the 177 spatial pattern of extreme precipitation across CONUS, even though a few of them overestimate 178 extremes, limiting their use in subsequent modelling. 179
- 180

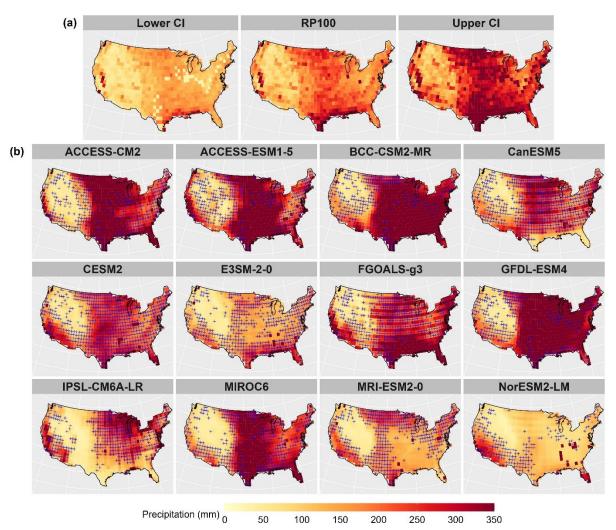


Figure 1. Evaluation of GCMs. Panels (a) show the 100-year daily precipitation (mm) and their upper and lower confidence interval (95% CI) based on CPC observations from 1981-2014.

Panels (b) show the 100-year daily precipitation extremes for the 12 GCMs from 1981-2014. The

blue crosses highlight the grids that fall within the 95% CI of the CPC observations.

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187 To mitigate issues with the biases in the estimates of extremes in certain GCMs, for the 188 attribution we considered only those models whose *n*-year precipitation falls within the 95% confidence interval (CI) of the observations for each grid (Figure 1b). This approach helps in 189 190 removing the GCMs that significantly underestimate or overestimate the observed precipitation 191 pattern, helping to reduce the uncertainty in climate attribution statements. More than five GCMs satisfy the above condition in most grids across CONUS, with the highest number of GCMs 192 available in the western and central United States, as well as the northeastern regions (Figure 193 194 S4). At least three GCMs with satisfactory performance are available in more than 87% grids (730 grids of the total 831 grids) across CONUS, thereby ensuring the robustness of the 195 estimates. In short, over five GCM simulations of *n*-year return level precipitation are consistent 196 with the corresponding CPC simulation of precipitation extremes across most of the CONUS and 197 we use them for climate attribution (Figure S4). Moreover, the whole time series of precipitation 198 (>0.1 mm) are sampled by fitting the ExtGPD distribution. Therefore, the comparison of 199 extremes in CMIP6 GCMs and CPC observations provides more confidence in the GCM 200

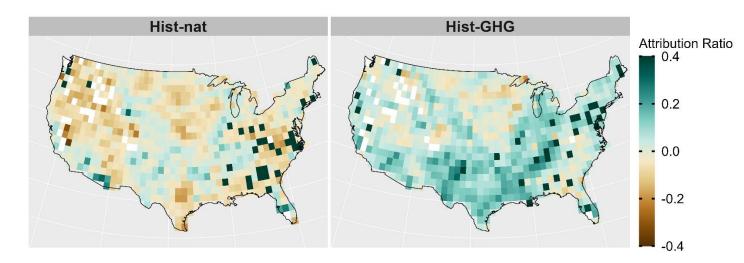
201 precipitation simulations.

202 **3.2.** Climate attribution

203 We estimate the attribution ratio by taking the ratio of the change in return level in the counterfactual scenario (hist-nat and hist-GHG) to that of the historical simulations (equation 3). 204 We estimate the multi-model mean attribution ratio of the subset of GCMs that performed well 205 compared to CPC for each grid based on 100-year precipitation (Figure 2 and Figures S6-S7). 206 207 Most regions across the country exhibit an attribution ratio below 0 in the hist-nat scenario and above 0 in the hist-GHG scenario (Figure 2). The results show a counterfactual scenario of 208 natural-only forcing would have made the 100-year daily precipitation event smaller than the 209 observations (Figure 2). Likewise, a well-mixed greenhouse gas emission scenario that excludes 210 both anthropogenic and natural aerosols would have enhanced the 100-year daily precipitation 211 extremes during the observational period (Figure 2). We observed consistent patterns in the 212 attribution ratio at higher return levels of daily precipitation (Figure S8). Overall, we identify the 213

- significant role of anthropogenic emissions in exacerbating the daily precipitation extremes
- across the CONUS.

216



217

Figure 2. Multi-model mean attribution ratio of hist-nat and hist-GHG simulations estimated
using the subset of GCMs that performed well with respect to CPC for the 100-year daily
precipitation. For each grid, the mean is estimated based on the GCMs that capture the observed
extremes (Figure S4). The white color corresponds to grids for which all GCMs failed to capture
the observed extremes within the 95% confidence interval.

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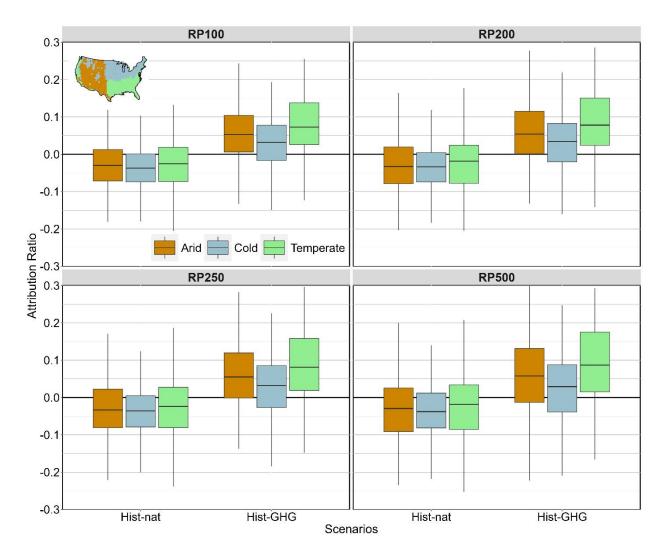
Five or more GCMs that captures the historical extremes are considered for computing the multi-224 model mean attribution ratio in most grids (Figure S4). We compared the sign of the attribution 225 ratio of the subset of GCMs in each grid to assess the robustness of the estimates of different 226 GCMs. The majority of the GCMs shows a negative sign of change in the hist-nat scenario and a 227 positive sign of change in the hist-GHG scenario (Figures S5-S7). We found that a high 228 229 percentage of GCMs agree in the sign of change at regions with high strength in attribution (Figure 2 and Figure S5). The Northwest, U.S. East Coast, Midwest and southern regions show a 230 high negative attribution ratio and have a high percentage of GCMs agreeing on the sign of 231 change (negative attribution ratio) in the hist-nat scenario (Figure 2 and Figure S5). Similarly, the 232 233 Southern Plains, the east-central United States, U.S. Northeast and West Coast show high positive attribution ratio and have a higher percentage of GCMs agreeing on the positive sign of 234 change (Figures 2 and Figure S5). Thus, the grids with higher magnitude of the multi-model 235 236 mean attribution ratio have more GCMs that agree on the sign of change, making the result

237 robust across the GCMs. The higher magnitude of extreme precipitation in a GHG-only scenario (hist-GHG) compared to the historical simulations highlights the modulating effect of natural and 238 239 anthropogenic aerosols in reducing precipitation intensities. Aerosols, particularly sulphate aerosols, have a net negative impact on radiative forcing due to their high reflectance and 240 indirect effect on cloud formation, leading to a reduction in precipitation (e.g., Allan et al., 2020; 241 Risser et al., 2022). The positive attribution ratio in the hist-GHG also coincides with regions 242 with high sulphate aerosols emissions across CONUS, including the northeastern, the east-243 central and southwestern United States and the U.S. West Coast (Risser et al., 2022). The 244 increased risk of extreme daily precipitation in the hist-GHG scenario can be attributed to not 245 accounting for the role of aerosols. A lower daily precipitation magnitude in the natural-forcing 246 scenario emphasizes the role of GHG emissions in exacerbating the precipitation extremes. We 247 have not directly compared the precipitation magnitudes in the two counterfactual scenarios of 248 well-mixed GHG and natural only forcing. However, it is evident that GHG forcing alone would 249 have enhanced the magnitude of observed extremes compared to natural forcing (Figure 2 and 250 Figures S6-S7). 251

252 We observed distinct regional patterns in the attribution ratio. The Southern Plains, easterncentral and north-eastern United States, which receive intense precipitation, exhibit a higher 253 254 sensitivity of precipitation extremes to anthropogenic forcing (right panel in Figures 1a and 2). On the contrary, the western United States, including the U.S. Northwest and the Northern 255 Plains, show a higher reduction in extremes under natural-only forcing (left panel in Figure 2). 256 257 This is consistent with the spatial pattern of observed extreme precipitation across CONUS 258 (Figure 1a). An intensification of heavy precipitation is observed since 1979 in the central and eastern United States, attributed to the increased frequency of mesoscale convective systems that 259 cause heavy precipitation during the warm season (Easterling et al., 2017). Similarly, a linear 260 261 increase in precipitation extremes is reported in the Midwest, U.S. East Coast and the Great Plains excluding the northwest regions (Dong et al., 2021), with a low confidence in the increase 262 of extreme precipitation in the western regions (Seneviratne et al., 2021). Overall, we found the 263 Southern Plains and the northeast regions, which receive intense daily precipitation and an 264 observed increase in extremes, are highly sensitive to anthropogenic GHG emissions, more so 265 than some of the drier northwestern regions. 266

267 **3.3.** Role of climate patterns and return period on the attribution ratio

We observed regional variability in the attribution ratio under the two counterfactual scenarios. 268 To understand the role of climatic regions on the attribution ratio, we estimated the multi-model 269 mean attribution ratio for the three major Köppen-Geiger climate classes over CONUS, namely 270 271 arid, cold and temperate regions (Beck et al., 2018). The temperate regions exhibit the highest increase in extreme precipitation under hist-GHG simulations, followed by the arid regions, with 272 the areas belonging to the cold region showing the lowest signal (Figure 3). The vast number of 273 event attribution studies conducted in the southern and central United States highlights the role 274 275 of anthropogenic emissions in increasing the probability of observed precipitation extremes in 276 the temperate eastern United States (Van Der Wiel et al., 2017; Wang et al., 2018; Zhao et al., 2018), consistent with our findings. In the hist-nat scenario, the temperate and arid regions show 277 278 a relatively low decrease in precipitation extremes compared to the cold regions, which show the highest decrease (Figure 3). Moreover, the cold regions have the smallest interquartile spread in 279 280 the attribution ratio in the hist-nat scenario. However, the difference among the climatic regions is much less in hist-nat compared to hist-GHG scenario. As observed earlier, the arid and cold 281 282 regions, which received low precipitation extremes during the study period, have a low attribution ratio compared to the temperate regions in the hist-GHG scenario, which receive 283 284 intense daily precipitation. This implies that GHG emissions exacerbate precipitation extremes in wetter regions compared to drier ones. 285



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Figure 3. Attribution ratio for the three main Köppen-Geiger climate classes across CONUS. The spatial extent of the three main climate classes (arid, cold, and temperate) is shown in the inset figure. The box plots depict the attribution ratio for the three climate classes in a hist-nat and hist-GHG scenario for 100, 200, 250 and 500-year daily precipitation. In the boxplots, the dark horizontal line represents the median, the box represents the interquartile range, and the whiskers correspond to the 5th and 95th percentile values.

293

We also compared the sensitivity of the attribution ratio to the magnitude of the extremes. We

- observed a slight increase in the risk of extreme precipitation of higher return periods (RP=200,
- 296 250, 500) due to anthropogenic emissions (Figure 3 and Figures S8-S9). There is a nominal
- 297 increase (decrease) in the multi-model mean attribution ratio in the hist-GHG (hist-nat) scenario
- with an increase in precipitation return period (Figure 3 and Figure S8). However, there is a
- substantial increase in the interquartile range as the return period increases (Figure 3 and Figure

S9). The temperate climatic regions exhibit a relatively higher increase in the attribution ratio in the hist-GHG scenario compared to the hist-nat (Figure 3). Overall, we find that the likelihood of higher magnitude precipitation extremes would have increased in a counterfactual scenario of anthropogenic emissions compared to the historical scenario. Previous studies outline that the probability of rare extremes increases at a higher rate in a warming world (Myhre et al., 2019), in agreement with our results, even though at a more muted level.

306 4. Discussion and Conclusions

We used the ExtGPD to assess the performance of 12 DAMIP GCMs in capturing the daily 307 precipitation extremes. Unlike the GPD, the ExtGPD samples the entire time series and removes 308 309 the arbitrariness in threshold selection in modelling extremes. Hence, the comparison of CPC 310 observations and GCMs using the ExtGPD provides a robust estimation of the climate model performance, and we found that most GCMs capture the spatial pattern of observed extremes. 311 Our probabilistic attribution study shows an unambiguous connection of human emissions to 312 daily precipitation extremes across CONUS. We report higher confidence in the attribution of 313 daily precipitation extremes to well-mixed GHG emissions in the Northeast, Southern Plains, and 314 U.S. West Coast, which receive high extreme precipitation, consistent with previous studies (e.g., 315 Easterling et al., 2017, Dong et al., 2016). The arid regions in the west and the colder Midwest 316 show a higher sensitivity to a counterfactual scenario of natural forcing, implying a reduction in 317 318 extreme precipitation magnitude compared to the observed extremes. The spatial distribution of 319 the attribution ratio identified in the study aligns with the observed pattern of precipitation extremes, suggesting a tendency for increased extreme magnitudes in regions experiencing 320 higher levels of extremes. 321

322 One limitation of this study is the lower spatial resolution of the DAMIP GCMs (~120 km),

323 which leads to an inherent disadvantage in the accurate simulation of the regional precipitation

324 patterns. However, the attribution ratio defined in the study is based on the counterfactual

scenarios of the same climate models forced with different external forcings. Therefore,

assuming that the response of the GCMs to the anthropogenic and natural forcing would be the

327 same across resolutions, then our attribution statements would not be impacted by their coarse

resolution. This assumption should be verified in future studies using a suite of models with

329 different resolutions. To strengthen our results and lend more credence to their robustness, we

only considered a subset of GCMs that performed well in reproducing the observed statistical
 properties of rainfall extremes, and used the ExtGPD, which samples the entire time series with
 no need for thresholds.

Anthropogenic warming has already increased the global mean temperature by 1°C from the pre-333 334 industrial level and the current emission trajectory is expected to exceed 1.5°C by the middle of the 21st century (Masson-Delmotte et al., 2018). An increase in global mean temperature 335 increases the water holding capacity of the atmosphere according to the Clausius Clapeyron 336 relationship, which in turn increases the frequency of precipitation extremes (Allan & Soden, 337 2008; Fowler et al., 2021; Papalexiou & Montanari, 2019; Westra et al., 2014). Our results 338 339 highlight the human influence on observed daily precipitation extremes across CONUS and emphasize the need for active reduction in human emissions to mitigate the intensification of 340 341 precipitation extremes in future.

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Open Research: The CPC and CMIP6 multi-model precipitation data used in the study are
 freely downloadable from https://psl.noaa.gov/data/gridded/data.unified.daily.conus.html and

356 <u>https://esgf-data.dkrz.de/search/esgf-dkrz/</u>, respectively. The CMIP6 models used in the study are

357 listed in Table S1. The analysis in this study is performed in R, an open-source computational

software. The codes used for the analysis and data visualization are made available at

359 <u>http://www.hydroshare.org/resource/d372565acce441bebdd49a0cf0307ff3</u>.

360 **Declaration of Competing Interest:** The authors declare no competing personal relationships or 361 financial interests that could have appeared to influence the work reported in this paper.

362 Author Contribution: Nanditha J. S.: Methodology, Software, Formal analysis, Investigation,

363 Data curation, Writing – original draft, Writing – review & editing, Visualization. Gabriele

- 364 Villarini: Conceptualization, Methodology, Investigation, Writing review & editing,
- 365 Supervision, Project administration, Funding acquisition. Hanbeen Kim: Software, Writing –
- 366 review & editing, Visualization. Philippe Naveau: Investigation, Writing review & editing

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