

Characterization of Relativistic Electron Precipitation Events Observed by the CALET Experiment Using Self-Organizing-Maps

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Key Points:

- Relativistic Electron Precipitation (REP) is observed by the CALET experiment from the International Space Station
- The Self-Organizing-Map technique is used for automatic detection and classification of rapidly varying REP intervals
- The Self-Organizing-Maps distinguish between different REP populations

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Abstract

Relativistic electron precipitation (REP) is a relatively high-latitude phenomenon where high-energy electrons trapped in the outer radiation belt are lost into the Earth's atmosphere. REP events observed at low Earth orbit show varying temporal profiles and global distributions. While the precipitation origin has been attributed to specific wave modes or scattering sources, the sorting of REP events by type or driver remains an unsolved challenge. In this study, we analyze the temporal profile of relativistic electron precipitation events observed by the CALorimetric Electron Telescope (CALET) experiment on board the International Space Station. We use an unsupervised machine learning technique called Self-Organizing-Maps (SOM) to automatically detect and then classify relativistic electron events observed by the two scintillator layers at the top of the apparatus, sensitive to electrons with energies > 1.5 MeV and > 3.4 MeV, respectively. We calculate the power spectral density (PSD) of the count rates observed by both sensors and use them as an input for the SOM. The SOM technique groups the PSDs by their similarity, resulting in a classification of relativistic electron events by the periodicity of the observed precipitation. We investigate the L-shell and magnetic local time distribution of the resulting classification, and energy spectral index associated with the observations. Clear precipitation patterns are observed and compared to past precipitation categorization attempts as well as known distributions of various scattering mechanisms. The classification reveals features through the sorting of the variability of the rapid precipitation, allowing the identification of different precipitation populations with varying properties.

Plain Language Summary

Fast electrons are normally trapped by the Earth's magnetic field. However, they often get released in bursts and impact the upper layers of the atmosphere near the poles. The underlying processes are still not well understood and debated. In this study we use an unsupervised artificial intelligence technique called Self-Organizing-Maps (SOM) to automatically detect and classify the observations made by a charged particle detector onboard the International Space Station (ISS). The SOM categorizes the bursts based on their variability and group together observations by their similarity. We compare the categorization with the spatial location of the electron bursts. Clear patterns are observed and compared with past categorizations attempts.

1 Introduction

Relativistic Electron Precipitation (REP) refers to electrons with energies greater than hundreds of keV and initially trapped in the outer Van Allen radiation belt that fall into the upper atmosphere due to pitch angle scattering in the loss cone (Shprits et al., 2006; Loto’Aniu et al., 2006; Millan & Thorne, 2007). This phenomenon represents a source of radiation capable of generating atmospheric heating as well as posing a long term health risk for airline pilots and in both, short and long term for astronauts, especially during extravehicular activities (RA et al., 1995; Dachev, 2018; Ueno et al., 2020; Xu et al., 2021). Currently, the most widely accepted mechanism for REP is pitch angle scattering associated with wave-particle interaction or current sheet scattering (CSS) (Summers & Thorne, 2003; W. Li & Hudson, 2019). The former process occurs as result of the resonance of magnetospheric waves with parallel velocity of counter-streaming energetic electrons (Lorentzen et al., 2001; Millan & Thorne, 2007; Blum, Halford, et al., 2015; Blum & Breneman, 2020). Meanwhile, the latter arises from the violation of the first adiabatic invariant when the Earth’s magnetic field curvature radius is comparable to the gyroradius of the electrons. It mainly occurs near the equatorial region of the current sheet, hence the name current sheet scattering (Sergeev & Tsyganenko, 1982; Sergeev et al., 1983; Capannolo et al., 2022). Since both mechanisms can generate large losses of relativistic electrons, they are important for maintaining the equilibrium of the outer Van Allen belt, and efforts continue to be made to obtain direct observations of both scattering mechanisms in the radiation belts and precipitation into the upper atmosphere.

Several direct REP measurements have been conducted by spacecraft and balloons during the last four decades. The Heavy Ion Large Telescope (HILT) experiment from the Solar, Anomalous, and Magnetospheric Particle Explorer (SAMPEX) observed that REP events usually have a latitudinal extension of 2-3°, and can persist for several hours (Blake et al., 1996). SAMPEX observations also showed the existence of 10-30 seconds time-scale precipitation bands mostly observed in the dusk-midnight sector and of more rapid variations (~ 100 ms) known as microbursts predominantly observed in the dawn-noon sector (Nakamura et al., 1995; Blake et al., 1996; Bortnik et al., 2006; Blum, Li, & Denton, 2015; Crew et al., 2016; Shumko et al., 2018). These REP events have been categorized based on their location in L-shell and MLT coordinate as well as with their correlation with proton precipitation, and lower energetic electrons. Yahnin et al. (2016)

86 identified a total of three groups of REP events. The first group corresponds to electrons
 87 from the isotropic zone near the trapped limit for electrons. This type of precipitation
 88 always occurs in the nightside and is likely result of CSS. They also observed a second
 89 and third group from electrons deeper in the trapped zone which suggest they are the
 90 result from the interaction with waves. The second group corresponds to relativistic elec-
 91 trons observed simultaneously with lower energetic electrons ($> 30\text{keV}$). These events
 92 are observed at all MLTs, with a maximum at the pre-midnight sector, and they are more
 93 likely to be related to electrostatic waves near the upper-hybrid-frequency, and plasma-
 94 spheric hiss. The third group corresponds to REP events correlated with energetic pro-
 95 tons observations, suggesting an interaction with EMIC waves, mostly observed in the
 96 dusk and pre-midnight sectors.

97 Blum et al. (2013) and K. Zhang et al. (2017) used the Colorado Student Space
 98 Weather Experiment (CSSWE) cubesat and Balloon Array for Radiation-belt Relativis-
 99 tic Electron Losses (BARREL) to study a total of three different precipitation bands events
 100 during 18-19 January 2013. Both studies estimated a net loss of the 0.58-1.63 MeV elec-
 101 trons close to 5% of the total electron content, showing the significance of precipitation
 102 bands as nearly 15-20 events could deplete the outer belt. Similarly, Shekhar et al. (2020)
 103 used NOAA/POES satellites and BARREL to quantify the relativistic electron loss for
 104 11 events on January 17, 2013. They estimate a net loss of 5% of the electrons with en-
 105 ergies above 700 keV.

106 Thorne and Kennel (1971) suggested that Electromagnetic Ion Cyclotron (EMIC)
 107 waves can generate REP in the $E > 1$ MeV range, which would imply simultaneous ob-
 108 servation of REP and increases in proton precipitation in the anisotropic proton zone
 109 where protons are unstable to wave growth. This correlation was observed by Søråas et
 110 al. (2005) using the Polar Operational Environmental Satellites (POES) by matching the
 111 proton flux increases observed by the P1 (52 keV differential proton flux) and relativis-
 112 tic electron increases observed by P6 (> 800 keV when used for electrons) channels. Sandanger
 113 et al. (2007, 2009) used the same channels to show that the proton and electron enhance-
 114 ments are consistent with scattering into the loss cone by EMIC waves. Carson et al. (2013)
 115 analyzed EMIC-driven REP using 12 years of POES observations and found that the ma-
 116 jority of events occur in the pre-midnight and midnight sectors around $L \sim 5$. Other space-
 117 craft such as the FIREBIRD-II cubesats observed electron precipitation in the 200-300
 118 keV range while in conjunction with EMIC waves detected by the the Van Allen Probes,

119 suggesting that EMIC waves can efficiently scatter a wide energy range of electrons (Capannolo
 120 et al., 2021). REP has been also observed by the Balloon Array for Radiation belt Rel-
 121 ativistic Electron Losses (BARREL) in conjunction to EMIC waves measured by a Geo-
 122 stationery Operational Environmental Satellite (GOES) spacecraft at dusk (Blum, Hal-
 123 ford, et al., 2015), and at pre-midnight by Van Allen Probes (J. Zhang et al., 2016). How-
 124 ever, EMIC-driven events account for only a portion of all the REP occurring in the mid-
 125 night sector, as CSS also plays an important role scattering relativistic electrons in the
 126 current sheet (Smith et al., 2016; Shekhar et al., 2017; Capannolo et al., 2022).

127 The periodicities observed by low altitude orbit spacecraft can help to detect REP
 128 events and also to distinguish between radiation belt crossings, precipitation bands, or
 129 microbursts. They can be examined with spectrograms to investigate the time evolution
 130 of the REP (Nakamura et al., 1995). Kataoka et al. (2016) used four-month data from
 131 the CALorimetric Electron Telescope (CALET) on the International Space Station (ISS)
 132 to show that 5-20s (50 - 200 mHz) periodicities are frequently present during REP events.
 133 These periodicities have been regularly observed (Mursula et al., 2001; Jacobs, 2012),
 134 and they have been associated with nonlinear wave growth of EMIC-triggered emissions
 135 as proposed by several numerical simulations (Omura & Zhao, 2012; Shoji & Omura, 2013;
 136 Kubota et al., 2015).

137 The use of periodicity analysis is an alternative to other methods currently used
 138 for the identification of REP events. In general, automatic algorithms are more efficient
 139 than methods based on visual inspection of data, and are less sensitive to biases in the
 140 analysis of large amount of data (Bortnik et al., 2007). However, they are susceptible
 141 to noise-to-signal ratio problems inducing false positive cases if the detection threshold
 142 is reduced with the intention of identifying small amplitude events (Guralnik & Srivas-
 143 tava, 1999). Currently, microburst-detection algorithms have shown to be effective, but
 144 have not been equally efficient for the detection of precipitation bands (O'brien et al.,
 145 2003; Blum, Li, & Denton, 2015). We present here a novel method for automatic detec-
 146 tion and analysis of REP.

147 **1.1 The CALET Experiment**

148 The CALET experiment was designed to observe high-energy cosmic rays and has
 149 been operational since October 2015. The instrument is attached to the Japanese Mod-

150 ule “Kibo” at the ISS and has the scientific objective to study high-energy phenomena
 151 of the Universe (Torii & Collaboration, 2007; Torii, 2016; Asaoka et al., 2018; Torii et
 152 al., 2019). The two scintillator arrays constituting the charge detector (CHDX and CHDY)
 153 at the top of the apparatus used to identify the incident cosmic rays charge are also sen-
 154 sitive to electrons with energies > 1.5 MeV and > 3.4 MeV, respectively (Bruno et al.,
 155 2022). This makes the CHDX, CHDY pair suitable for the detection of hard spectra REP
 156 events (Kataoka et al., 2016). This capability is particularly useful since CALET is one
 157 of the few instruments available at this energy range for conjugate MeV electrons stud-
 158 ies during the Van Allen probes era (2012-2019). Its data have already been used for the
 159 study of the relation between some REP events and magnetospheric waves (Kataoka et
 160 al., 2020; Bruno et al., 2022).

161 The REP events observed by CALET are identified by isolated increases in count
 162 rates measured by the CHDX/CHDY detectors. Figure 1a shows several hours of data
 163 from November 10, 2015 where the peaks correspond to relativistic electrons. Figure 1b
 164 shows an example of a REP observation. They are characterized by rapid variations that
 165 can last from a few seconds to several minutes. In some cases, both types of profiles (smooth
 166 and rapid profiles) are present at the same time (see Figure 1c). Automated detection
 167 algorithms for these types of events can be more complex to design as they would also
 168 require a previous knowledge about the existence of each type of signature and their char-
 169 acteristics. Another class of events consists in smooth profiles mostly associated with pro-
 170 tons detected in the South-Atlantic-Anomaly region and, similarly, electrons in the in-
 171 nermost part of the outer radiation belt ($L \sim 3$) (Kataoka et al., 2016, 2020; Bruno et
 172 al., 2022). Such events are identified as a gradual increase-then-decrease of the count rates
 173 with a timescale typically of 5-10 minutes. Figures 1e to g consist of the continuous wavelet
 174 transform (Aguiar-Conraria & Soares, 2014) of the observations showing the contrast-
 175 ing variability of the CHDX channel for smooth and rapid relativistic electrons profiles,
 176 respectively. It is important to mention that since REP events can last several hours and
 177 extend in latitude and longitude, the same REP event can be detected during consec-
 178 utive orbits (Nakamura et al., 1995; Blake et al., 1996; Blum et al., 2013; Bruno et al.,
 179 2022).

180 The data used in this study have a continuous coverage from October 2015 to Oc-
 181 tober 2021. The data has quasi-periodic sampling time resolution, with an average pe-
 182 riod of 1 second. The ISS (therefore CALET) is located at low Earth orbit (LEO) at 370-

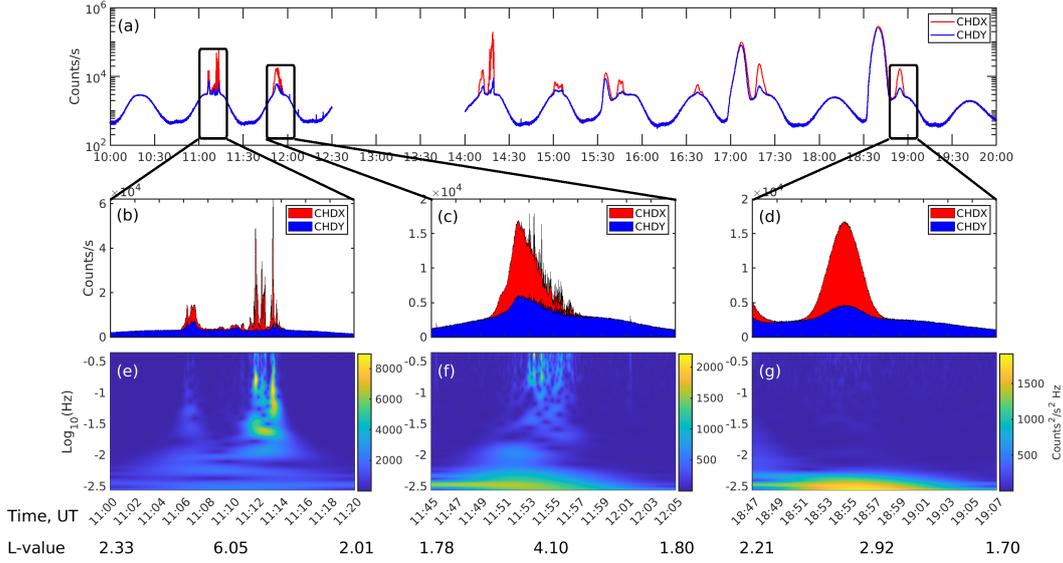


Figure 1. (a) Ten hours of CALET CHDX (red; $E > 1.5$ MeV) and CHDY (blue; $E > 3.4$ MeV) data in counts/s. The sharp spikes represent sudden increases of relativistic electrons. (b) Example of a rapid electron precipitation event. (c) Example of a combined smooth and rapid profiles of relativistic electrons. (d) Example of a smooth profile of relativistic electrons. (e) Continuous wavelet transform of the rapid electron precipitation event shown in plot b. (f) Continuous wavelet transform of combined event shown in plot c. (g) Continuous wavelet transform of a smooth profile of relativistic electrons event shown in d. Data gap is present between 12:30 and 14:00.

183 460 km of altitude and has an inclination of 51.6° . ~~As a result~~In consequence, the ISS
 184 visits L=4-7 regions several times a day at a similar magnetic local time (MLT) enabling
 185 periodic sampling of the outer radiation belt. The ISS exhibits a precession time of 60
 186 days. This implies that the MLT at which the CALET probes the high L-shell also fol-
 187 lows the same 60-day periodicity.

188 2 Methodology

189 A self-organizing-map (SOM) is an unsupervised machine learning technique used
 190 to define an ordered mapping, as a projection from a set of given data items onto a reg-
 191 ular, usually two-dimensional grid of nodes. A data item will be mapped into the most
 192 similar node, based on the smallest distance from the node in some metric (Kohonen,
 193 1982, 1990, 2013). The SOM technique has been widely used for unsupervised cluster-

194 ing of different kinds of data set in biology, chemistry, sociology, and economics (Akman
 195 et al., 2019; Mele & Crowley, 2008; Yang & Chou, 2003; Collan et al., 2007), but lately
 196 also for identification of magnetospheric regions (Innocenti et al., 2021) and categoriza-
 197 tion of plasma waves (Vech & Malaspina, 2021).

198 The SOM is a competitive learning neural network model. The neural network con-
 199 sists of a grid of nodes initially built from randomly selected samples in the data set. This
 200 means that identical results can theoretically only be guaranteed when the same seed
 201 is used during the pseudo-random selection of samples. However, converging results will
 202 generate mirrored, rotated, or identical maps. Since the distribution of the clusters with
 203 respect to each other in mirrored or rotated maps is the same, the interpretation of the
 204 map remains unchanged in these cases.

205 The learning process is based on an iterative search of the *best-matching-unit* (BMU)
 206 for each one of the samples in the data set. The BMU is the most similar unit (or node)
 207 to each value of the data set during each iteration. The similarity between the nodes and
 208 the data can be evaluated with multiple metrics; the most popular one, used in this study
 209 is the Euclidean distance $\left(d(q_i, p_i) = \sqrt{\sum (q_i - p_i)^2}\right)$ where q and p represent the cur-
 210 rent sample and current unit, respectively. During each iteration, the BMU and the nodes
 211 surrounding it are updated to become more similar to the latest input sample evaluated.
 212 The updates are made based in the learning rate $(\eta = \eta_0 e^{-t\lambda})$ that controls how much
 213 weight the last sample has on the update of the BMU. The radius of influence $(\sigma = \sigma_0 e^{-t\beta})$
 214 determines the influence of the input vector in the surrounding clusters where t corre-
 215 spond to the current iteration and λ and β are the respective decay rates for the learn-
 216 ing rate and the radius of influence, respectively. For both steps we used $\eta_0 = 0.1$, $\sigma_0 =$
 217 $\sqrt{2}$ and $\beta = 0.1$. The behavior of the SOM to these free parameters is standard to any
 218 SOM, they are initially defined defined by the size of the map and the similarity between
 219 the observations and later adjusted for better performance. Different parameters will de-
 220 termine how fast (i.e. after how many iterations) the map converges to a stable solution
 221 or if it does not converge at all. We tested multiple combinations of parameters and se-
 222 lected the above because they result in the map converging to the same result even when
 223 different seeds are used for the random selection of the initial map, which is evidence of
 224 a converging solution. In addition, as it will be shown below, we observed only a small
 225 number of incorrectly classified observations using these parameters.

226 Here we implement the SOM technique to classify the observations from CALET
 227 and analyze the precipitation patterns found. This process is performed in two steps:
 228 (1) detection of rapid electron precipitation observation events; (2) and classification of
 229 rapid precipitation observation events. The first step uses the Power-Spectral-Density
 230 (PSD) calculated from 10 minute windows of data as input for the SOM while the sec-
 231 ond step uses an interval-integrated-PSD. The details of the implementation of the SOM
 232 are explained in the following section [and a diagram of the methodology implemented](#)
 233 [can be found in the Supporting Information \(Figure S2\).](#)

234 The CALET data set is collected at a nearly constant rate of 1-second. The data
 235 is re-sampled to 1-second as uniform sampling is required for spectral analysis used in
 236 this study. The re-sampling helps to reduce aliasing and contributes to removing small
 237 data gaps. Windows with gaps larger than 3 seconds are discarded as they would intro-
 238 duce a spurious response during the application of the [Fast-Fourier-Transform](#) (FFT).
 239 The re-sampled data is subdivided into 10 minute windows starting from October 2015
 240 until October 2021. This choice is based on the ISS orbital constraints, since REP events
 241 can be observed only for a few minutes during each pass. The SOM technique is applied
 242 two times for similar, but different objectives. In the first place, the SOM is applied with
 243 the objective of distinguishing rapid precipitation from the smooth profile intervals, and
 244 background noise. During this step, the SOM is implemented using the PSD of each one
 245 of the windows as input; the PSDs are calculated from the count rates of the CHD chan-
 246 nels to capture the intrinsic variability of the observed relativistic electrons. In this step,
 247 the Euclidean distance is computed using the current event spectral power at each fre-
 248 quency (q_i) and the current node spectral power at each frequency (p_i). The PSD are
 249 calculated using the [Fast-Fourier-Transform](#) FFT with 100 ~~FFT~~ points, in order to compute
 250 the PSD while keeping a fast computational time. The number of FFT points should
 251 be increased if data with higher sampling resolution is used. The output is a map of clus-
 252 ters where every cluster contains a subset of PSD with shared similarities in overall power
 253 and power distribution in frequency. Exclusively focusing on REP events, we chose clus-
 254 ters with zero smooth profiles or background noise, effectively eliminating these elements
 255 from the analyzed sample. We are then left with a “cleaned” data set of only rapid pre-
 256 cipitation observations for further study.

257 During the second step only rapid precipitation observations are considered. How-
 258 ever, it is possible that more than one rapid precipitation event occurs during a 10 minute

259 interval. In order to separate multiple REP events in the same 10 minute interval, we
 260 apply a 60-second long moving window that computes the variance along the interval,
 261 and normalize such variance by dividing by its norm. Based on visual inspection we es-
 262 tablished a tolerance of 0.1 [counts/s]^2 to detect the beginning and ending of the rapid
 263 precipitation event. There are cases when two consecutive rapid precipitation observa-
 264 tions occur in periods less than 10 minutes: this usually happens when the ISS crosses
 265 a region where relativistic precipitation is observed by CALET during the ascending and
 266 descending orbital passes. When these observations are separated, two time series of dif-
 267 ferent lengths are generated. To obtain two same-length time series consistent with the
 268 rest of the data, the edges of the series are filled with generated background noise sim-
 269 ilar to that seen by CALET when only background particles are observed.

270 The lower limit on the PSDs are set to 100 mHz to remove the effects associated
 271 with the rapid movement of the ISS. The upper limit of the PSDs are set to 500 mHz
 272 since aliasing due to 1 second sampling rate should equally affect the detection of all very
 273 rapid precipitation observed. Since the SOM technique is most efficient with a low num-
 274 ber of variables, we created an equivalent representation of the PSDs with a lower num-
 275 ber of variables by dividing each PSD in 10 bins with 50% overlap and integrating the
 276 PSD each bin to obtain a simplified PSD profile. This procedure allows us to simplify
 277 the PSD and forces the SOM to classify by overall power of the PSD and power distri-
 278 bution in frequency.

279 Once all the events are individualized and standardized, we apply the k -means tech-
 280 nique to calculate the number of clusters (k -value) that minimize total variances between
 281 all the events contained in each cluster. Finally, the SOM technique is applied to the interval-
 282 integrated-PSD of the rapid precipitation observations with the objective of classifying
 283 different features of the precipitation in order to identify different types of rapid precip-
 284 itation events.

285 The output is a grid of clusters (or map) where each cluster consists of precipita-
 286 tion events with similar PSD characteristics. We examine the properties of the precip-
 287 itation events in each cluster to explore their dependence on various variables and bet-
 288 ter determine the physical meaning behind the SOM's categorizations.

3 Results

3.1 First Step Results

The objective of the first step is to detect rapid precipitation observations without the need of extensive visual inspection of the data, or an algorithm that requires a detailed knowledge about the characteristics of the data. The SOM is able to not only identify isolated rapid precipitation intervals, but also events where smooth profiles and rapid precipitation occur simultaneously. In order to validate the SOM technique with CALET data, we visually inspected all the clusters to verify the observations were correctly classified. During the time period covered in this study, the SOM identified 1448 rapid precipitation events, 21301 intervals were classified as smooth profiles of relativistic electrons and the rest (275241) identified as background noise. We visually inspected all 1448 events classified by the SOM as rapid precipitation and found 87 events (6.0%) to be false positives (events incorrectly classified as rapid). We also performed a survey over half of the events that were classified as smooth profile events to quantify the number of false negatives. From a visual survey of 11545 events that were identified as smooth profile events, we found 27 false negatives or 0.23% of the events. The number of windows classified as noise is too large to be evaluated by visual inspection, so we randomly selected 5000 time windows classified by the SOM as noise for visual identification. Of this sample we found 9 false negatives, or 0.18% of the events. We performed a z-test to calculate a confidence interval and found the total number of false negatives in the background noise to be 495 ± 172 with a 95% of confidence.

As demonstrated by Figure 2a, the geographic distribution of smooth-profile events concentrates in the South-Atlantic-Anomaly region. Another component is present in the southern hemisphere around $L \sim 3$, corresponding to trapped and quasi-trapped (drift-loss-cone) electrons in the inner boundary of the outer radiation belt, where trapped electrons correspond to electrons that can stably drift around Earth unless perturbed, and quasi-trapped correspond to electrons that will bounce several times before precipitation occurs (Selesnick et al., 2003; Tu et al., 2010; Pham et al., 2017; K. Zhang et al., 2017). Meanwhile, the rapid profiles (Figure 2c) are typically detected at higher magnetic latitudes, mapping to the footprint of the outer radiation belt ($L \sim 4-6$).

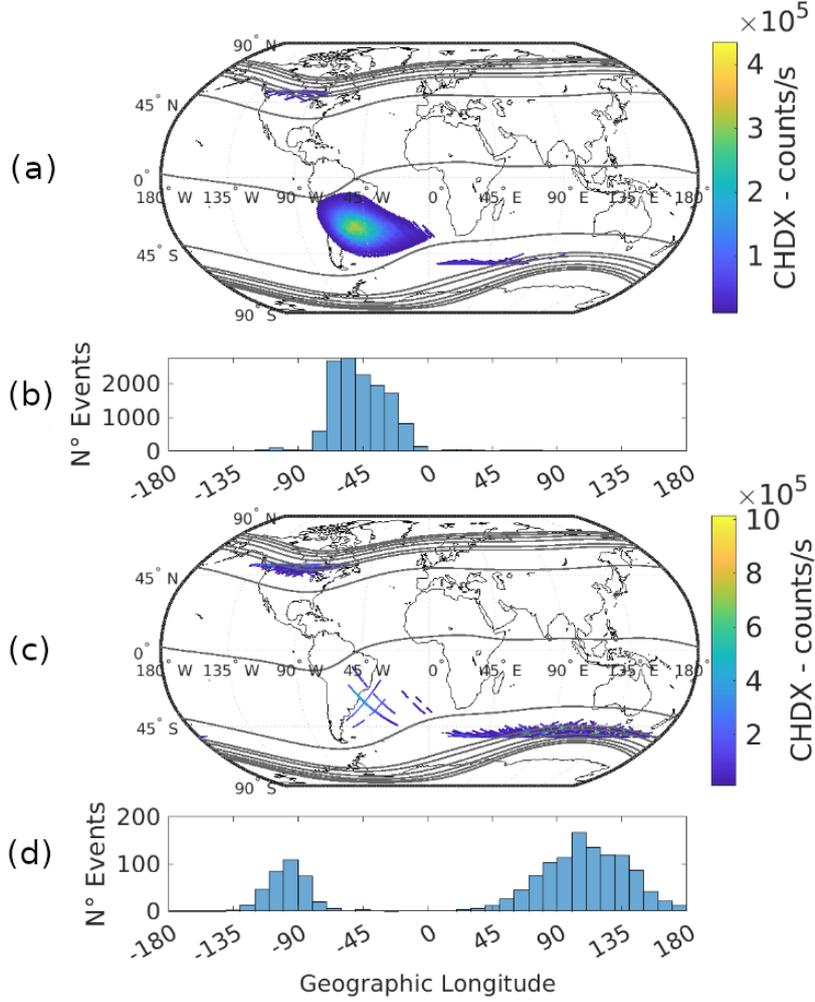


Figure 2. (a) Positions of the first 1000 smooth profiles as an example. (b) Geographic longitude histogram of all smooth profile events. (c) Positions of all REP events between October 2015 to October 2021. The color indicates the counts/s observed by CALET. L-shell curves from L=1 to L=8 in dark gray. (d) Geographic longitude histogram of all rapid events. Note that the latitude of CALET observations is constrained by the inclination of the ISS orbit (51.6°).

319 **3.2 Second Step Results**

320 The objective of the second step is to analyze the maps generated in step 2, con-
 321 sisting of only REP events to uncover patterns associated with the magnitude and vari-
 322 ability of the REP events observed. The number of clusters is determined using the k -
 323 means (k being the number of clusters) technique (Likas et al., 2003). k -means acts as
 324 a classifier that minimizes the within-cluster variances given a predetermined number

325 of clusters. We found that the optimal number of clusters is 15 and to simplify the anal-
 326 ysis the rapid precipitation observation events are classified into 16 clusters to keep the
 327 map square. With the objective to study the precipitation L-MLT patterns and the as-
 328 sociated variability, we evaluate the clusters in terms of the frequency interval that car-
 329 ries the maximum power in the PSD, the energy spectra index, and the distance to the
 330 plasmopause of the observations.

331 The median frequency at which the PSD peaks considering all events is 183 mHz
 332 ($T=5.5$ seconds). Figure 3 shows the PSD of the classified clusters. Clusters 1-4 and 6-
 333 8, and 11 have most of their power above 183 mHz, while clusters 5, 9-10, and 12-16 are
 334 dominated by lower frequency signatures. Since each cluster is filled with individual PSDs
 335 corresponding to precipitation events, for each of them we calculate a median curve of
 336 the PSDs using the median value at each frequency. We also estimate the 25% and 75%
 337 curves to observe the distribution of the variability of the events at each frequency.

338 For each cluster four representative values are calculated for the events in the re-
 339 spective cluster: The median of the frequency at the PSD maximum amplitude of each
 340 PSD; the median of the maximum amplitude of each PSD; the median of the maximum
 341 spectral hardness; and the median of the distance to the plasmopause. The clusters are
 342 then sorted using each one of these values. We compare the group of clusters that show
 343 the maximum dissimilarity to enhance the characteristics that could be useful for anal-
 344 ysis. We achieve this by comparing the clusters below the 25 and above 75 percentile,
 345 respectively of the four computed values that represent one characteristic of the clusters.

346 When comparing the representative frequency at the PSD maximum amplitude,
 347 the two groups show differences in their MLT and L-shell distributions. Since the dis-
 348 tributions were close to a Gaussian, we performed the significance Z-test with a Z-value=5.86.
 349 Similarly, we use Monte Carlo test to compute the probability that such distribution dif-
 350 ference can be due to randomness. The median difference between both distributions is
 351 larger than in 96.2% of random distributions computed. Both tests are performed with
 352 a 95% of confidence to estimate that the discrepancy between both distributions is sta-
 353 tistically significant.

354 Figure 4 shows the distribution in [the mapmaps](#) of different characteristics of the
 355 clustered events. They demonstrate how other characteristics associated to the events
 356 distribute when the SOM organizes the events by their PSD. Figure 4a shows the result-

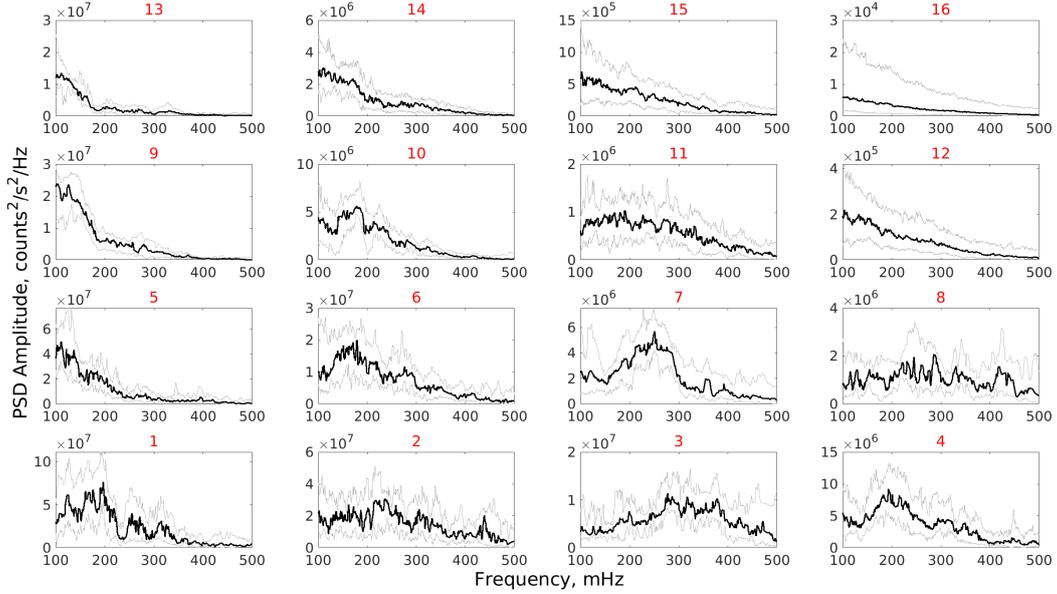


Figure 3. PSD of all clusters. The black curves represent the median of all the events of each cluster. Gray curves correspond to the 25 and 75 percentiles at each frequency of the PSD of the events in each cluster. Red numbers indicate the number of the respective cluster.

357 ing map where the colors indicate the median frequency in each cluster at which power
 358 spectral density peaks. The clusters where the dominant frequency is above 183 mHz
 359 contain events where high-frequency periodicity electron precipitation is dominant. Fig-
 360 ure 4b displays the percentage of REP weighed by the total number of passes through
 361 every L-MLT grid cell, demonstrating that low-periodicity ([below the 25 percentile](#))
 362 events are dominant at pre-midnight and between L=5-6. Figures 4c shows that high-
 363 periodicity ([above the 75 percentile](#)) events occur at local times, but are more likely
 364 to occur in the midnight sector between L=5-7.

365 Figure 4d displays the median value of the highest amplitude in the PSD for each
 366 cluster. We use again the clusters where the median frequency is below the 25th or above
 367 the 75th percentile to separate them into two groups. Figure 4e shows that small am-
 368 plitude events occur predominantly at midnight. In the midnight sector they are observed
 369 at L=5-7. Finally, Figure 4f demonstrates that rapid precipitation with larger amplitudes
 370 is dominant in the pre-midnight sector and between L=5-6.

371 We also evaluate the event energy spectra in each cluster. We use the ratio between
 372 the count rates measured by the two sensors to calculate an energy spectral index (CHDX/CHDY).

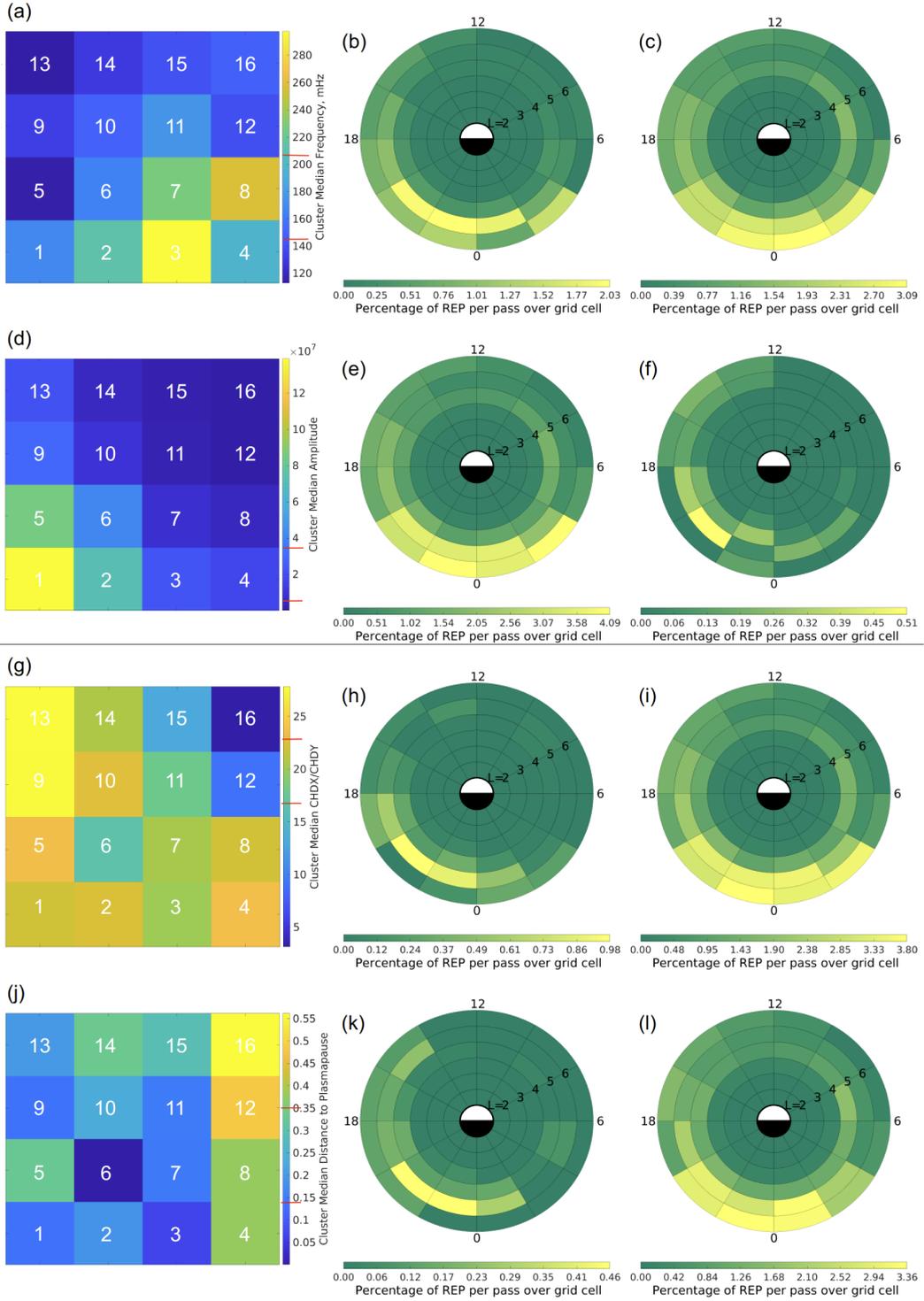


Figure 4. Left column: Maps of clusters. Middle column: Bivariate histogram of clusters under the 25 percentile. Right column: Bivariate histogram of clusters above the 75 percentile. (a-c) Median of the dominant periodicities. (d-f) Median of the PSD amplitude. (g-i) Median of the energy spectral index. (j-l) Median of the distance to the plasmopause. L-shell histograms can be found in the Supporting Information.

373 The energy spectral index was calculated using the maximum CHDX/CHDY ratio dur-
 374 ing the event. Since CHDX and CHDY detect electrons with energies above 1.5 and 3.4
 375 MeV, respectively, larger values of the spectral index correspond to a softer spectrum
 376 associated with the precipitation. Figure 4g shows the same kind of map as Figures 4a
 377 and 4d, but with the color code denoting the median energy spectral index of each clus-
 378 ter. Figure 4h demonstrates that events from clusters with softer energy spectral index
 379 (CHDX/CHDY above 75th percentile) are concentrated in the pre-midnight sector and
 380 at $L = 5-6$. In contrast, Figure 4i shows that events from clusters with a harder energy
 381 spectral index (CHDX/CHDY below 25th percentile) are common at all MLT, but pre-
 382 dominantly observed in the midnight sector at $L=5-7$.

383 Lastly, we calculate the location of the events with respect to the plasmopause. Moldwin
 384 et al. (2002) MLT-dependant empirical model has been used in numerous other stud-
 385 ies to analyze the spatial distribution of waves (Carson et al., 2013; D. Wang et al., 2015;
 386 Saikin et al., 2016), and to investigate the location of the outer belt with respect to the
 387 plasmopause (X. Li et al., 2006) among other studies. We use this model to calculate
 388 the location of the plasmopause and estimate its distance to the REP detection location
 389 (ΔL) in order to see if different precipitation types exhibit any correlation by their dis-
 390 tance to the plasmopause. Figure 4j shows the median ΔL of the events in each cluster.
 391 Figure 4k shows that the clusters with a median distance to the plasmopause below the
 392 25th percentile are more common near the pre-midnight sector $L=5-6$. In contrast, the
 393 clusters with a median distance above the 75th percentile are more frequent in the pre-
 394 midnight and midnight sectors at $L=5-7$.

395 4 Discussion

396 The results presented in the previous sections suggest that the SOM is an efficient
 397 tool for separating different types of REP observations time series by classifying their
 398 PSD. It effectively distinguishes rapid precipitation events from smooth profiles, and back-
 399 ground noise, eliminating the need for extensive visual inspection or the use of standard
 400 automated algorithms that are often sensitive to signal-to-noise ratio detection thresh-
 401 olds. The SOM is also capable of classifying rapid precipitation events by the period-
 402 icities and the power of the PSD. We use the median, in addition to 25 and 75 percentiles
 403 values of the dominant frequency, peak PSD amplitude, energy spectral index, and dis-
 404 tance to plasmopause of the rapid precipitation as reference to separate the precipita-

405 tion into different populations. The results show that precipitation with different char-
 406 acteristics can generate similar precipitation patterns, demonstrating the classification
 407 of rapid precipitation is a multidimensional problem. They also reveal features through
 408 sorting based on PSD alone, allowing the identification of different precipitation com-
 409 ponents with varying properties.

410 Microbursts and precipitation bands are example of two types of REP with the pe-
 411 riodicity of the electron fluxes among the observational characteristics that distinguish
 412 them. While whistler mode chorus waves are the primary mechanism believed to drive
 413 microbursts, electrostatic and EMIC waves are believed to drive precipitation bands (Thorne
 414 & Kennel, 1971; Blum, Li, & Denton, 2015). However, the observation of precipitation
 415 bands at conjugated locations and consecutive orbits (Blake et al., 1996), suggests that
 416 their characteristic signature is related to spatial rather than temporal characteristics
 417 (Lorentzen et al., 2001; Bortnik et al., 2006; Blum, Li, & Denton, 2015).

418 Carson et al. (2013); Z. Wang et al. (2014); Gasque et al. (2021) used an algorithm
 419 applied to POES P1 (52 keV differential proton flux) and P6 (>800 keV when used for
 420 electrons) channels to detect EMIC-driven REP events. These authors found that EMIC-
 421 driven REP are predominantly detected in the dusk-midnight sector around $L \sim 5$. The
 422 CSS mechanism also occurs in the midnight sector and it is sometimes even more effi-
 423 cient than wave-driven REP. While previous studies have associated REP near midnight
 424 to EMIC waves, it has been speculated that softer REP events are driven by CSS while
 425 harder precipitation events are more likely to be driven by EMIC waves (Smith et al.,
 426 2016; Shekhar et al., 2018; Capannolo et al., 2021). Capannolo et al. (2022) performed
 427 a conservative classification between EMIC-driven and CSS-driven REP events to en-
 428 sure events were truly driven by one mechanism alone and found that near 40% of the
 429 classified events were CSS-driven.

430 The results of this analysis show similarities with the findings of aforementioned
 431 studies. For instance, Figures 4h and i show that REP events can be separated by their
 432 relative spectral hardness into at least two populations that overlap near midnight. Fig-
 433 ure 4h shows that softer precipitation events mainly occur in the pre-midnight sector be-
 434 tween $L=4-5$. Figure 4i shows that hard precipitation is observed at all MLT, but they
 435 are mainly localized in the midnight sector. Some events are seen in the morning sec-
 436 tor where microbursts are commonly observed. However, the microburst variability (~ 100

ms) cannot be resolved by the 1 second time resolution of the CALET data set used in this work, making it hard to investigate the origin of these hard precipitation events.

The classification by variability also separates the REP observations into two populations. Kataoka et al. (2016) also studied periodicities observed by CALET in the dusk and pre-midnight sector finding similar periodicities that have been associated to non-linear wave growth of EMIC-triggered emissions by several numerical simulations and observational studies. The REP events with low variability are more frequent in the pre-midnight sector where EMIC-driven precipitation is more commonly observed. On the other hand, REP events with higher variability are more frequently observed in the mid-night sector near where CSS-driven precipitation is more frequent. The morphology of the REP patterns and their similarity to EMIC-driven REP and CSS-driven REP patterns reported by Yahnin et al. (2016) suggests a potential connection between the variability observed and the driver of the precipitation. Low variability REP are more common in the same region where EMIC-driven REP are often observed. Meanwhile, high variability REP are observed where CSS-driven REP are most commonly observed.

In all cases, an L-MLT pattern with a predominant occurrence of events in the pre-midnight sector has been generated, similar to the one observed for EMIC-driven REP. However, the REP classification cannot be directly associated with drivers without conjugated observations. Typically, EMIC-driven REP and CSS-driven REP have been difficult to distinguish from each other as they often can occur simultaneously, and currently only a small portion of the observations can be truly classified as EMIC or CSS-driven precipitation (Capannolo et al., 2022).

The results of this work demonstrate that there is information about REP hidden in the variability of the observations that can be used for future studies to distinguish and analyze their drivers.

5 Conclusions

In this study, we have described a new use of an unsupervised machine learning technique to classify time series of relativistic electron precipitation. We have tested the capabilities of SOM for the analysis of the rapid relativistic electron precipitation observed by CALET in the 2015-2021 period. The SOM has automatically detected rapid electron precipitation intervals and classified them by the main characteristics of their PSD.

468 It has been shown that the SOM technique is a robust method for event detection. The
469 SOM is capable of detecting rapid electron precipitation events of any magnitude, even
470 if they are superposed with smooth profiles. The SOM has also been implemented to clas-
471 sify the rapid precipitation observations. The output of the maps suggests that the SOM
472 technique can categorize rapid precipitation into different types of precipitation with dif-
473 ferent properties.

474 The energy spectral index and distance to the plasmopause exhibit a similar L-MLT
475 pattern than the one obtained for periodicities and amplitude, but when different clus-
476 ters are grouped. This is most likely due to the existence of multiple high and low pe-
477 riodicities for rapid precipitation, such as microbursts and precipitation bands that have
478 different L-MLT distributions. This also reveals the complexities of REP analysis as mul-
479 tiple precipitation types, with different characteristics, may have similar patterns in L-
480 MLT. It also shows that unsupervised machine learning is a useful tool for disentangling
481 this multidimensional problem.

482 We have demonstrated that this technique has the potential for the identification
483 of electron precipitation in LEO observations, and to distinguish different types of pre-
484 cipitation. As next steps, we plan to use it in conjugated studies between CALET and
485 the Van Allen Probes that would help to determine the specific common characteristics
486 of rapid precipitation observations in each one of the clusters obtained from the SOM.

487 **6 Data Availability Statement**

488 The CALET data used in this study are publicly available (`data.darts.isas.jaxa`
489 `.jp/pub/calet/cal-v1.1/CHD/level1.1/obs/`) in ASCII format in the Data ARchives
490 and Transmission System (DARTS) of the Japan Aerospace Exploration Agency (JAXA).
491 The catalog of REP observation events is attached in a text file. The Supporting Infor-
492 mation contains a description of the catalog and necessary considerations when using
493 the catalog.

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