Ionospheric response over Thailand from the 15 January 2022 eruption of the Hunga Tonga-Hunga Ha'apai volcano

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Abstract

This study reports on the upper atmospheric response over Thailand to the Hunga Tonga volcano's eruption on January 15th, 2022. The eruption occurred during the geomagnetic storm recovery phase, providing a rare comparison between effects from outside (geomagnetic storm) and inside atmosphere (volcanic eruption). About nine hours later, we observed post-eruption fluctuations in the ionosphere total electron content (TEC). TEC was recorded in Thailand approximately ten times remarkably higher than typical levels from this large perturbation. The initial impact reached Thailand with speed of ~275 m/s. Detrended TEC (dTEC) revealed mixed wave packets at various intervals. Two significant traveling ionospheric disturbance (TID) waves reached Thailand at 13 UT and 14 UT, respectively. Equatorial plasma bubbles (EPBs) were observed between 12-15 UT and 17-18 UT over Thailand. Our findings could provide insight into how communication signals over Thailand are affected by both disturbances, particularly in the case of widespread volcanic eruptions.



Solar Wind and Geomagnetic Parameters







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Key Points:

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18	• Hunga Tonga-Hunga Ha'apai eruption caused ionospheric disturbances in Thai-
19	land 9 hours after the eruption with a wavefront speed of ${\sim}275~{\rm m/s}$
20	• Two significant traveling ionospheric disturbance shock waves reached Thailand
21	at 13 UT and 14 UT $$
22	• Equatorial plasma bubbles triggered by the eruption spread horizontally for 100

Equatorial plasma bubbles triggered by the eruption spread horizontally for 1000 km and vertically for 700 km were spotted in Thailand 23

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24 Abstract

This study reports on the upper atmospheric response over Thailand to the Hunga 25 Tonga volcano's eruption on January 15th, 2022. The eruption occurred during the ge-26 omagnetic storm recovery phase, providing a rare comparison between effects from out-27 side (geomagnetic storm) and inside atmosphere (volcanic eruption). About nine hours 28 later, we observed post-eruption fluctuations in the ionosphere total electron content (TEC). 29 TEC was recorded in Thailand approximately ten times remarkably higher than typi-30 cal levels from this large perturbation. The initial impact reached Thailand with speed 31 32 of ~ 275 m/s. Detrended TEC (dTEC) revealed mixed wave packets at various intervals. Two significant traveling ionospheric disturbance (TID) waves reached Thailand at 13 33 UT and 14 UT, respectively. Equatorial plasma bubbles (EPBs) were observed between 34 12-15 UT and 17-18 UT over Thailand. Our findings could provide insight into how com-35 munication signals over Thailand are affected by both disturbances, particularly in the 36 case of widespread volcanic eruptions. 37

³⁸ Plain Language Summary

The Hunga Tonga eruption had a significant impact on the upper atmosphere glob-39 ally. This study investigates the eruption's impact on Thailand's ionosphere by compar-40 ing it to geomagnetic storm-induced changes preceding the volcanic eruption. Studying 41 how waves moved through the atmosphere and changes in observable upper atmospheric 42 phenomena, we found that substantial perturbations in total electron content, or TEC, 43 are about ten times higher than usual. The first perturbation arrived over Thailand nine 44 hours after the eruption, indicating a speed of approximately 275 meters per second. Mixed 45 wave patterns also occurred at specific intervals throughout the day. Since the first wave 46 packet appeared after the eruption, two large waves reached Thailand about four and 47 five hours later, respectively. Additionally, ionospheric bubbles formed over Thailand that 48 night. Overall, this analysis shows how the eruption's impact interacted with other at-49 mospheric dynamics, highlighting the complexity of regional-scale changes in the iono-50 sphere. 51

52 1 Introduction

The opportunity to investigate the consequences of both external (geomagnetic storm) 53 and internal (geological activity) drivers, the case of the Hunga Tonga-Hunga Ha'apai 54 (HTHH) volcanic eruption on January 15th, 2022, is unique. The HTHH volcano located 55 at 20.54°S geographic latitude and 175.38°W geographic longitude in the South Pacific 56 is particularly intriguing to be observed for the co-volcanic ionospheric disturbances (CVIDs) 57 following the moderate geomagnetic storm (G2, Kp 6-) event influenced by a coronal mass 58 ejection (CME) on January 14th, 2022. In addition, a minor geomagnetic storm (G1, 59 Kp 5-) was produced at the end of the UT time on day 15th from the impact of a high-60 speed solar wind stream originating from a coronal hole (Aa et al., 2022). Generally, the 61 energy from energetic particles during the geomagnetic storm event produces high-latitude 62 heating, which can produce TIDs. The propagation of these waves is typically studied 63 by deriving from TIDs, the detrended change in TEC (dTEC). 64

The eruption took place at ~4:14 UT on January 15th, 2022, during the recovery phase of the January 14th geomagnetic storm. Previous studies in the Asian sector have reported TIDs (Pradipta et al., 2023; Rakesh et al., 2022; Li et al., 2023; Sun et al., 2022; Saito, 2022; Hong et al., 2022; Zhang et al., 2022; Astafyeva et al., 2022; Aa et al., 2022; Muafiry et al., 2022). The HTHH eruption and its aftermath (Tarumi & Yoshizawa, 2023) generated acoustic and gravity waves creating the Lamb waves over a wide range frequency (Astafyeva et al., 2022; Zhang et al., 2022) that spread globally for several days after the



Solar Wind and Geomagnetic Parameters

Figure 1. (a) VTEC and average VTEC during quiet phase (gray line) from KMI6 (Thailand station) shows the typical day-to-night variation found at this location, (b) Kp index, and (c) SYM/H index throughout January 13th-17th, 2022. The onset of the geomagnetic storm, occurring around 10:53 UT on January 14th, is marked by the red dashed line. The beginning of the recovery phase, indicated by the minimum SYM/H time around 22:30 UT on January 14th, is represented by the green dashed line. The volcanic eruption, occurring at approximately 4:14 UT on January 15th, is denoted by the blue dashed line. There is a sudden increase in VTEC after the SYM/H index reaches its minimum value. Two significant VTEC wavefronts arrive in Thailand. The initial and subsequent peak arrival times in Thailand are labeled with an orange arrow labeled (1) and a purple arrow labeled (2), respectively. Sky blue and green arrows indicate two troughs before the initial and after the subsequent peaks, respectively.

explosion. CVIDs often represent quasi-periodic variations of ionospheric electron den sity or TEC with periods of 12–30 minutes (Shults et al., 2016).

The Earth's ionospheric disturbances in low-latitude regions are known to be caused 74 by sudden composition changes originating from external and internal sources. This work, 75 focuses on the response over Thailand, indicating a growing interest in the field of iono-76 spheric research in Thailand and filling in gaps in observations. These dynamic phenom-77 ena are crucial to understanding ionospheric behavior. An important ionospheric focus 78 has been characterizing ionospheric gradients, waves, and plasma bubbles. The causes 79 of the ionospheric gradients are varied and include traveling ionosphere disturbances (TIDs) 80 and ionospheric irregularities (Liu et al., 2018). Ionospheric irregularities over Thailand 81 are influenced by several factors, including solar activity, geomagnetic conditions, and 82 local atmospheric dynamics, contributing to TEC variability and irregularities near the 83 equatorial region. Moreover, ionospheric irregularities in Thailand primarily occur at night, with the appearance of spread F and pre-midnight scintillation inhibited by magnetic 85 activity (Charoenkalunyuta et al., 2012). These irregularities can significantly affect ra-86 dio wave propagation, satellite communications, and navigation systems. 87

⁸⁸ 2 Data and Method

This study investigates ionospheric irregularities using regional TEC data from the global navigation satellite system (GNSS) TEC data from the global Madrigal database and Thailand sites in Southeast Asia, which is near the equatorial region. The analysis period is particularly interesting because the observations reported here occurred during a geomagnetic disturbance and eruption of the Hunga Tonga volcano.

For regional data, the 1-second TEC and the 5-minute rate of TEC change index 94 (ROTI) were obtained at Bangkok (KMI6: 13.77°, 100.53°) and Chumphon (CPN1: 10.76°, 95 99.37°) stations referenced to 350 km altitudes over Thailand (Bumrungkit et al., 2018). 96 The slant measurements of TEC (STEC), derived from GNSS satellite ranging data ob-97 servations, are converted to vertical TEC (VTEC) using the 350 km assumed pierce point 98 location. The time variations of the VTEC from January 13th to 17th, 2022, were binned 99 at 20-minute temporal resolution across a range of $5^{\circ} \times 5^{\circ}$ in latitude and longitude of 100 the Thailand region (black line in Figure 1a) in comparison with the average VTEC dur-101 ing the quiet phase from January 10th-12th, 2022 (gray line in Figure 1a). The distance-102 time of VTEC and ROTI in Thailand were calculated (Figures 3a and 3b). The phase 103 speed is determined by analyzing the maximum VTEC within 20-minute and 200-kilometer 104 intervals. 105

To process the data, we applied the Savitzky-Golay moving average method for calculating the dTEC (Savitzky & Golay, 1964; Astafyeva et al., 2022; Zhang et al., 2022; Sun et al., 2022; Li et al., 2023; Aa et al., 2022; Hong et al., 2022; Saito, 2022). The dTEC data were extracted from particular lines of sights (LOSs) corresponding to satellite paths from Hunga Tonga to Thailand during 7 UT to 21 UT (Figure 4). Subsequently, disturbed wave periods were identified using Lomb-Scargle periodogram analysis (Lomb, 1976; Scargle, 1982).

From more than 5000 stations, 5-minute VTEC data referenced to the 350-km pierce point altitude were obtained from the worldwide Madrigal GNSS database (http:-//millstonehill.haystack.mit.edu/). The global distance-time VTEC variations were calculated for ten latitude-longitude-distance bins from Tonga to Thailand along the great-arc circle (Figure 2). The propagation speed was based on the maximum VTEC within the mean 20-minute observation, using 52 minutes after the eruption until the first peak arrived in Thailand at around 12:30 UT.

Very High Frequency (VHF) radar (transmitting at 39.65 MHz) data were obtained from the Prachomklao VHF radar station at the King Mongkut's Institute of Technology Ladkrabang (KMITL) Chumphon campus, provided by the National Institute of Information and Communications Technology (NICT) and KMITL (Thanakulketsarat et al., 2023).

3 Results and Discussion

Figure 1 depicts temporal VTEC variations in the VTEC from the KMI6 station 126 (located at 10° latitude and 100° longitude), alongside Kp and SYM/H indices for the 127 period spanning from January 13th to 17th, 2022. This time duration covers varying con-128 ditions from quiet time through a geomagnetic storm event and up to a volcanic erup-129 tion (illustrated by the dashed blue line) during the geomagnetic storm recovery phase. 130 The precise onset time (depicted by the dashed red line) of the moderate geomagnetic 131 storm event on January 14th was difficult to determine due to a slight increase in the 132 SYM/H index. We define the onset time as approximately 10:53 UT, based on a minor 133 peak in the SYM/H index, reaching ~ 17 nT, coinciding with the beginning of the rise 134 in the Kp index. The initial and main phases of the geomagnetic storm each lasted around 135 5-6 hours. The recovery phase commenced with the first minimum in the SYM/H in-136



Figure 2. (a) VTEC at 12:30 UT from the worldwide Madrigal database. (b) The variations in VTEC across global distances, from Tonga to Thailand, at the time of the volcanic eruption around 4:14 UT, are represented by a red dashed line. The propagation speed along this distance-time VTEC trajectory from Tonga to Thailand is 313 m/s, as indicated by the solid magenta line in (a).

dex, ~-100 nT, reaching a peak Kp index at around 22:30 UT on January 14th (indicated by the dashed green line).

Throughout the main phase of the geomagnetic storm, VTEC variations in Thai-139 land remained relatively stable. However, following two occurrences of minimum SYM/H 140 index values, coinciding with the highest Kp index readings between 21-23 UT on Jan-141 uary 14th, significant increases in VTEC were observed during the daytime on January 142 15th, continuing through the time of the volcanic eruption. This suggests that the ob-143 served VTEC variations were primarily influenced by the geomagnetic storm rather than 144 the volcanic eruption, given the propagation time. In this case, the observed VTEC in-145 creases at the KMI6 station indicate the occurrence of the ionospheric positive storm con-146 current with the first minimum SYM/H index rather than during the main phase from 147 11-16 UT on January 14th. 148

Figure 2b illustrates the propagation at a speed of approximately 331 m/s since the HTHH eruption. The sharp increase in VTEC observed on the distance-time graph for the KMI6 and CPN1 stations indicates a front speed of ~275 m/s (Figure 3a). This speed is lower than that derived from the global propagated direction. The VTEC observed farther away from HTHH, at approximately 10,000-km range (represented by the solid olive line in Figure 3b), corresponds to the VTEC observed by KMI6 in Thailand.

A trough of VTEC from the VTEC depletion (sky blue arrow in Figure 1a) occurred 155 around 1.5 hours ($\sim 11 \text{ UT}$) before the initial rise in VTEC observed in Thailand at around 156 12:30 UT (orange arrow (1) in Figure 1a). This first VTEC peak took ~ 9 hours to ar-157 rive in Thailand after the volcanic eruption, consistent with findings from previous stud-158 ies in nearby regions (Sun et al., 2022; Hong et al., 2022; Saito, 2022). A second VTEC 159 peak occurred at $\sim 13:45$ UT (purple arrow (2) in Figure 1a), followed by another trough 160 of TEC at $\sim 15:30$ UT (green arrow in Figure 1a). On January 16th, during the daytime, 161 VTEC increased slightly earlier than the average observed during quiet conditions, likely 162



Figure 3. Distance-time variations from Tonga to Thailand were analyzed using data from KMI6 and CPN1 stations, showing: (a) VTEC with the propagation speed (magenta solid line) of sudden VTEC changed in Thailand, occurring around 275 m/s at approximately 9 hours after the HTHH volcanic eruption (red dashed line), and (b) ROTI displaying the expansion of EPBs horizontally over 1000-km range. (c) Spectra of VHF radar varying with altitudes measured by CPN1 station at (1) 13 UT, (2) 14 UT, (3) 15 UT, (4) 16 UT, (5) 17 UT, and (6) 18 UT. The EPB at 14 UT and 15 UT reaches altitudes up to 700 km, fades at 16 UT, and reappears again at 17 UT and 18 UT.



Figure 4. The dTEC along satellite's lines of sight in oriented directions from Hunga Tonga to Thailand. Mixed wave packets from 7 UT to 21 UT, occurring at various intervals: 8.5-10 UT (green double-head arrow), 12-15 UT (gray double-head arrow), 16-18 UT (sky blue double-head arrow), and 18.5-20 UT (pink double-head arrow). Between 12-15 UT, disturbances intensified around 12 UT, about 8 hours after the HTHH eruption. Fluctuations in dTEC peaked around 13 UT (orange arrow) and again at 14 UT (purple arrow), resembling TID shock waves. Intensity decreased after 15 UT.

due to residual enhanced VTEC levels. Subsequently, a decline in VTEC was observed, indicating diminishing volcanic impact, followed by a well-defined recovery phase.

The dTEC from 7 UT to 21 UT (illustrated in Figure 4) reveals mixed wave pack-165 ets extracted along the satellite's lines of sight oriented from Hunga Tonga to Thailand 166 at various intervals: 8.5-10 UT (green double-head arrow), 12-15 UT (gray double-head 167 arrow), 16-18 UT (sky blue double-head arrow), and 18.5-20 UT (pink double-head ar-168 row). During the 12-15 UT interval, these disturbances began intensifying around 12 UT, 169 approximately 8 hours after the eruption of HTHH. The sudden fluctuations in dTEC, 170 ranging between 1-3 TECU around 13 UT (referred to the orange arrow (1) in Figure 171 4), demonstrated a significant increase, resembling a TID shock wave. This coincides with 172 the first peak of VTEC arrival in Thailand, occurring at approximately 12:30 UT (marked 173 174 by the orange arrow (1) in Figure 1a), before gradually diminishing.

Around 14 UT, another significant increase in dTEC occurred, indicating the ar-175 rival of another significant TID shock wave (referred to as purple (2') in Figure 4), cor-176 responding to the sudden fluctuations of the second peak of VTEC at $\sim 13:45$ UT (pur-177 ple arrow (2) in Figure 1a). These fluctuations gradually diminish in intensity from 15 178 UT onwards, showing only small variations thereafter. This second disturbance decayed 179 more rapidly, resembling the rate of decay observed in the first wave packet. Other small 180 ionospheric conditions remained significantly disturbed for several hours, with background 181 disturbances persisting until returning to normal storm-time conditions over Thailand. 182

At a distance of 10,000-km range, we noted three occurrences where the dTEC envelopes correspond with Zhang et al. (2022) during the same time intervals. Specifically, wave packet (1') corresponds to the dTEC arrival time (\sim 13 UT) for wave travel along the great circle route at a velocity of 350 m/s, while wave packet (2') corresponds to the dTEC arrival time (~14 UT) for wave travel along the great circle route at a velocity of 300 m/s. Additionally, the dTEC observed in the interval 8.5-10 UT, 16-18 UT, and 18.5-20 UT could be associated with observing other medium-scale TID (MSTID) with the 20-minute wave period. However, this TID dissipated quickly due to its rapid velocity. The faster wave would have already reached the 10,000 km mark because it diminishes more rapidly.

During 12-15 UT on January 15th, a high level of variation in the distance-time 193 profile of ROTI (KMI6) (Figure 3b) was observed over a horizontal distance of more than 194 1000 km. The changes in ROTI demonstrated fluctuations, indicating the presence of 195 equatorial plasma bubbles (EPBs) over Thailand, consistent with the observations of EPB 196 "C" reported by Sun et al. (2022). The hourly VHF radar (CPN1) (Figure 3c) also presents 197 the formation dynamics of EPBs. Discrepancies between the VHF radar data and ROTI 198 were observed. It could be due to differences in observation locations or the propaga-199 tion characteristics of EPBs. Spatial expansion at very high altitudes, up to approximately 200 700 km, was noted at 14 UT (Figure 3c-2) and 15 UT (Figure 3c-3) based on spectral 201 data analysis. Consequently, the EPBs disappeared but reassembled around 16 UT (Fig-202 ure 3c-4) at different locations, suggesting the formation of new EPBs during the recov-203 ery phase between 17-18 UT (Figures 3c-5 and 3c-6). 204

205 4 Summary

The data collected from Thai stations highlighted substantial deviations from the 206 usual ionospheric conditions attributed to moderate and minor geomagnetic storm drivers 207 and a volcanic eruption. Disturbances observed on January 15th, characterized by a sig-208 nificant increase in VTEC, were typical of geomagnetic storms. However, unusual hor-209 izontal wave-like perturbations in dTEC, commencing around 9 hours after the eruption, 210 were related to volcanic activity. Two significant TID shock waves reached Thailand at 211 13 UT (350 m/s) and 14 UT (300 m/s), respectively, lasting about 3 hours. The observed 212 VTEC depletion on January 16th confirmed the impact of geomagnetic storm activity 213 during the recovery phase in low-latitude regions. On this day, distinct EPBs were ob-214 served on a large scale, spanning over 1000 km horizontally and up to 700 km in alti-215 tude. These EPBs appear to have been produced by the HTHH eruption. In the future, 216 the formations of EPBs resulting from the volcanic eruption, particularly compared to 217 those arising during the main phase of geomagnetic storm events, should be further in-218 vestigated. Additional data from diverse stations, incorporating information or relevant 219 events or aftermath effects of the explosion, is essential to comprehensively analyze these 220 phenomena. Moreover, this event holds significance for understanding disruptions caused 221 by external and internal energy sources, with potential implications for future commu-222 nication signal impacts. 223

224 Open Research

The solar wind parameters and the geomagnetic indices are obtained by NASA/GSFC's 225 Space Physics Data Facility's OMNIWeb service (https://omniweb.gsfc.nasa.gov/), and 226 the World Data Center for Geomagnetism, Kyoto (https://wdc.kugi.kyoto-u.ac.jp/). The 227 SuperMAG collaborators (https://supermag.jhuapl.edu/info/?page=acknowledgement) 228 provide AE index data. The Kp index is available from GFZ German Research Centre 229 for Geosciences (https://kp.gfz-potsdam.de/en/data). The KMI6 and CPN1 GNSS data 230 are provided by the Thai GNSS and Space Weather Information Data Center hosted at 231 King Mongkut's Institute of Technology Ladkrabang (KMITL), Thailand (http://iono-232 gnss.kmitl.ac.th). The worldwide GNSS data are available from the CEDAR Madrigal 233 database (http:// cedar.openmadrigal.org/). The VHF data are contributed from the 234 Southeast Asia Low-latitude IOnosphere Network (SEALION) under the National In-235

stitute of Information and Communications Technology (NICT) service (https://aer-nc web.nict.go.jp/sealion/index.html).

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Figure 1.

Solar Wind and Geomagnetic Parameters



Figure 2.

(a) Madrigal GNSS TEC 2022-01-15 (12:30 UT) (b) Distance-time VTEC variations from Tonga to Thailand 2022-01-15







Figure 3.



Figure 4.

dTEC (KMI6) 2022-01-15



Ionospheric response over Thailand from the 15 January 2022 eruption of the Hunga Tonga-Hunga Ha'apai volcano

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Key Points:

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18	• Hunga Tonga-Hunga Ha'apai eruption caused ionospheric disturbances in Thai-
19	land 9 hours after the eruption with a wavefront speed of ${\sim}275~{\rm m/s}$
20	• Two significant traveling ionospheric disturbance shock waves reached Thailand
21	at 13 UT and 14 UT $$
22	• Equatorial plasma bubbles triggered by the eruption spread horizontally for 100

Equatorial plasma bubbles triggered by the eruption spread horizontally for 1000 km and vertically for 700 km were spotted in Thailand 23

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24 Abstract

This study reports on the upper atmospheric response over Thailand to the Hunga 25 Tonga volcano's eruption on January 15th, 2022. The eruption occurred during the ge-26 omagnetic storm recovery phase, providing a rare comparison between effects from out-27 side (geomagnetic storm) and inside atmosphere (volcanic eruption). About nine hours 28 later, we observed post-eruption fluctuations in the ionosphere total electron content (TEC). 29 TEC was recorded in Thailand approximately ten times remarkably higher than typi-30 cal levels from this large perturbation. The initial impact reached Thailand with speed 31 32 of ~ 275 m/s. Detrended TEC (dTEC) revealed mixed wave packets at various intervals. Two significant traveling ionospheric disturbance (TID) waves reached Thailand at 13 33 UT and 14 UT, respectively. Equatorial plasma bubbles (EPBs) were observed between 34 12-15 UT and 17-18 UT over Thailand. Our findings could provide insight into how com-35 munication signals over Thailand are affected by both disturbances, particularly in the 36 case of widespread volcanic eruptions. 37

³⁸ Plain Language Summary

The Hunga Tonga eruption had a significant impact on the upper atmosphere glob-39 ally. This study investigates the eruption's impact on Thailand's ionosphere by compar-40 ing it to geomagnetic storm-induced changes preceding the volcanic eruption. Studying 41 how waves moved through the atmosphere and changes in observable upper atmospheric 42 phenomena, we found that substantial perturbations in total electron content, or TEC, 43 are about ten times higher than usual. The first perturbation arrived over Thailand nine 44 hours after the eruption, indicating a speed of approximately 275 meters per second. Mixed 45 wave patterns also occurred at specific intervals throughout the day. Since the first wave 46 packet appeared after the eruption, two large waves reached Thailand about four and 47 five hours later, respectively. Additionally, ionospheric bubbles formed over Thailand that 48 night. Overall, this analysis shows how the eruption's impact interacted with other at-49 mospheric dynamics, highlighting the complexity of regional-scale changes in the iono-50 sphere. 51

52 1 Introduction

The opportunity to investigate the consequences of both external (geomagnetic storm) 53 and internal (geological activity) drivers, the case of the Hunga Tonga-Hunga Ha'apai 54 (HTHH) volcanic eruption on January 15th, 2022, is unique. The HTHH volcano located 55 at 20.54°S geographic latitude and 175.38°W geographic longitude in the South Pacific 56 is particularly intriguing to be observed for the co-volcanic ionospheric disturbances (CVIDs) 57 following the moderate geomagnetic storm (G2, Kp 6-) event influenced by a coronal mass 58 ejection (CME) on January 14th, 2022. In addition, a minor geomagnetic storm (G1, 59 Kp 5-) was produced at the end of the UT time on day 15th from the impact of a high-60 speed solar wind stream originating from a coronal hole (Aa et al., 2022). Generally, the 61 energy from energetic particles during the geomagnetic storm event produces high-latitude 62 heating, which can produce TIDs. The propagation of these waves is typically studied 63 by deriving from TIDs, the detrended change in TEC (dTEC). 64

The eruption took place at ~4:14 UT on January 15th, 2022, during the recovery phase of the January 14th geomagnetic storm. Previous studies in the Asian sector have reported TIDs (Pradipta et al., 2023; Rakesh et al., 2022; Li et al., 2023; Sun et al., 2022; Saito, 2022; Hong et al., 2022; Zhang et al., 2022; Astafyeva et al., 2022; Aa et al., 2022; Muafiry et al., 2022). The HTHH eruption and its aftermath (Tarumi & Yoshizawa, 2023) generated acoustic and gravity waves creating the Lamb waves over a wide range frequency (Astafyeva et al., 2022; Zhang et al., 2022) that spread globally for several days after the



Solar Wind and Geomagnetic Parameters

Figure 1. (a) VTEC and average VTEC during quiet phase (gray line) from KMI6 (Thailand station) shows the typical day-to-night variation found at this location, (b) Kp index, and (c) SYM/H index throughout January 13th-17th, 2022. The onset of the geomagnetic storm, occurring around 10:53 UT on January 14th, is marked by the red dashed line. The beginning of the recovery phase, indicated by the minimum SYM/H time around 22:30 UT on January 14th, is represented by the green dashed line. The volcanic eruption, occurring at approximately 4:14 UT on January 15th, is denoted by the blue dashed line. There is a sudden increase in VTEC after the SYM/H index reaches its minimum value. Two significant VTEC wavefronts arrive in Thailand. The initial and subsequent peak arrival times in Thailand are labeled with an orange arrow labeled (1) and a purple arrow labeled (2), respectively. Sky blue and green arrows indicate two troughs before the initial and after the subsequent peaks, respectively.

explosion. CVIDs often represent quasi-periodic variations of ionospheric electron den sity or TEC with periods of 12–30 minutes (Shults et al., 2016).

The Earth's ionospheric disturbances in low-latitude regions are known to be caused 74 by sudden composition changes originating from external and internal sources. This work, 75 focuses on the response over Thailand, indicating a growing interest in the field of iono-76 spheric research in Thailand and filling in gaps in observations. These dynamic phenom-77 ena are crucial to understanding ionospheric behavior. An important ionospheric focus 78 has been characterizing ionospheric gradients, waves, and plasma bubbles. The causes 79 of the ionospheric gradients are varied and include traveling ionosphere disturbances (TIDs) 80 and ionospheric irregularities (Liu et al., 2018). Ionospheric irregularities over Thailand 81 are influenced by several factors, including solar activity, geomagnetic conditions, and 82 local atmospheric dynamics, contributing to TEC variability and irregularities near the 83 equatorial region. Moreover, ionospheric irregularities in Thailand primarily occur at night, with the appearance of spread F and pre-midnight scintillation inhibited by magnetic 85 activity (Charoenkalunyuta et al., 2012). These irregularities can significantly affect ra-86 dio wave propagation, satellite communications, and navigation systems. 87

⁸⁸ 2 Data and Method

This study investigates ionospheric irregularities using regional TEC data from the global navigation satellite system (GNSS) TEC data from the global Madrigal database and Thailand sites in Southeast Asia, which is near the equatorial region. The analysis period is particularly interesting because the observations reported here occurred during a geomagnetic disturbance and eruption of the Hunga Tonga volcano.

For regional data, the 1-second TEC and the 5-minute rate of TEC change index 94 (ROTI) were obtained at Bangkok (KMI6: 13.77°, 100.53°) and Chumphon (CPN1: 10.76°, 95 99.37°) stations referenced to 350 km altitudes over Thailand (Bumrungkit et al., 2018). 96 The slant measurements of TEC (STEC), derived from GNSS satellite ranging data ob-97 servations, are converted to vertical TEC (VTEC) using the 350 km assumed pierce point 98 location. The time variations of the VTEC from January 13th to 17th, 2022, were binned 99 at 20-minute temporal resolution across a range of $5^{\circ} \times 5^{\circ}$ in latitude and longitude of 100 the Thailand region (black line in Figure 1a) in comparison with the average VTEC dur-101 ing the quiet phase from January 10th-12th, 2022 (gray line in Figure 1a). The distance-102 time of VTEC and ROTI in Thailand were calculated (Figures 3a and 3b). The phase 103 speed is determined by analyzing the maximum VTEC within 20-minute and 200-kilometer 104 intervals. 105

To process the data, we applied the Savitzky-Golay moving average method for calculating the dTEC (Savitzky & Golay, 1964; Astafyeva et al., 2022; Zhang et al., 2022; Sun et al., 2022; Li et al., 2023; Aa et al., 2022; Hong et al., 2022; Saito, 2022). The dTEC data were extracted from particular lines of sights (LOSs) corresponding to satellite paths from Hunga Tonga to Thailand during 7 UT to 21 UT (Figure 4). Subsequently, disturbed wave periods were identified using Lomb-Scargle periodogram analysis (Lomb, 1976; Scargle, 1982).

From more than 5000 stations, 5-minute VTEC data referenced to the 350-km pierce point altitude were obtained from the worldwide Madrigal GNSS database (http:-//millstonehill.haystack.mit.edu/). The global distance-time VTEC variations were calculated for ten latitude-longitude-distance bins from Tonga to Thailand along the great-arc circle (Figure 2). The propagation speed was based on the maximum VTEC within the mean 20-minute observation, using 52 minutes after the eruption until the first peak arrived in Thailand at around 12:30 UT.

Very High Frequency (VHF) radar (transmitting at 39.65 MHz) data were obtained from the Prachomklao VHF radar station at the King Mongkut's Institute of Technology Ladkrabang (KMITL) Chumphon campus, provided by the National Institute of Information and Communications Technology (NICT) and KMITL (Thanakulketsarat et al., 2023).

3 Results and Discussion

Figure 1 depicts temporal VTEC variations in the VTEC from the KMI6 station 126 (located at 10° latitude and 100° longitude), alongside Kp and SYM/H indices for the 127 period spanning from January 13th to 17th, 2022. This time duration covers varying con-128 ditions from quiet time through a geomagnetic storm event and up to a volcanic erup-129 tion (illustrated by the dashed blue line) during the geomagnetic storm recovery phase. 130 The precise onset time (depicted by the dashed red line) of the moderate geomagnetic 131 storm event on January 14th was difficult to determine due to a slight increase in the 132 SYM/H index. We define the onset time as approximately 10:53 UT, based on a minor 133 peak in the SYM/H index, reaching ~ 17 nT, coinciding with the beginning of the rise 134 in the Kp index. The initial and main phases of the geomagnetic storm each lasted around 135 5-6 hours. The recovery phase commenced with the first minimum in the SYM/H in-136



Figure 2. (a) VTEC at 12:30 UT from the worldwide Madrigal database. (b) The variations in VTEC across global distances, from Tonga to Thailand, at the time of the volcanic eruption around 4:14 UT, are represented by a red dashed line. The propagation speed along this distance-time VTEC trajectory from Tonga to Thailand is 313 m/s, as indicated by the solid magenta line in (a).

dex, ~-100 nT, reaching a peak Kp index at around 22:30 UT on January 14th (indicated by the dashed green line).

Throughout the main phase of the geomagnetic storm, VTEC variations in Thai-139 land remained relatively stable. However, following two occurrences of minimum SYM/H 140 index values, coinciding with the highest Kp index readings between 21-23 UT on Jan-141 uary 14th, significant increases in VTEC were observed during the daytime on January 142 15th, continuing through the time of the volcanic eruption. This suggests that the ob-143 served VTEC variations were primarily influenced by the geomagnetic storm rather than 144 the volcanic eruption, given the propagation time. In this case, the observed VTEC in-145 creases at the KMI6 station indicate the occurrence of the ionospheric positive storm con-146 current with the first minimum SYM/H index rather than during the main phase from 147 11-16 UT on January 14th. 148

Figure 2b illustrates the propagation at a speed of approximately 331 m/s since the HTHH eruption. The sharp increase in VTEC observed on the distance-time graph for the KMI6 and CPN1 stations indicates a front speed of ~275 m/s (Figure 3a). This speed is lower than that derived from the global propagated direction. The VTEC observed farther away from HTHH, at approximately 10,000-km range (represented by the solid olive line in Figure 3b), corresponds to the VTEC observed by KMI6 in Thailand.

A trough of VTEC from the VTEC depletion (sky blue arrow in Figure 1a) occurred 155 around 1.5 hours ($\sim 11 \text{ UT}$) before the initial rise in VTEC observed in Thailand at around 156 12:30 UT (orange arrow (1) in Figure 1a). This first VTEC peak took ~ 9 hours to ar-157 rive in Thailand after the volcanic eruption, consistent with findings from previous stud-158 ies in nearby regions (Sun et al., 2022; Hong et al., 2022; Saito, 2022). A second VTEC 159 peak occurred at $\sim 13:45$ UT (purple arrow (2) in Figure 1a), followed by another trough 160 of TEC at $\sim 15:30$ UT (green arrow in Figure 1a). On January 16th, during the daytime, 161 VTEC increased slightly earlier than the average observed during quiet conditions, likely 162



Figure 3. Distance-time variations from Tonga to Thailand were analyzed using data from KMI6 and CPN1 stations, showing: (a) VTEC with the propagation speed (magenta solid line) of sudden VTEC changed in Thailand, occurring around 275 m/s at approximately 9 hours after the HTHH volcanic eruption (red dashed line), and (b) ROTI displaying the expansion of EPBs horizontally over 1000-km range. (c) Spectra of VHF radar varying with altitudes measured by CPN1 station at (1) 13 UT, (2) 14 UT, (3) 15 UT, (4) 16 UT, (5) 17 UT, and (6) 18 UT. The EPB at 14 UT and 15 UT reaches altitudes up to 700 km, fades at 16 UT, and reappears again at 17 UT and 18 UT.



Figure 4. The dTEC along satellite's lines of sight in oriented directions from Hunga Tonga to Thailand. Mixed wave packets from 7 UT to 21 UT, occurring at various intervals: 8.5-10 UT (green double-head arrow), 12-15 UT (gray double-head arrow), 16-18 UT (sky blue double-head arrow), and 18.5-20 UT (pink double-head arrow). Between 12-15 UT, disturbances intensified around 12 UT, about 8 hours after the HTHH eruption. Fluctuations in dTEC peaked around 13 UT (orange arrow) and again at 14 UT (purple arrow), resembling TID shock waves. Intensity decreased after 15 UT.

due to residual enhanced VTEC levels. Subsequently, a decline in VTEC was observed, indicating diminishing volcanic impact, followed by a well-defined recovery phase.

The dTEC from 7 UT to 21 UT (illustrated in Figure 4) reveals mixed wave pack-165 ets extracted along the satellite's lines of sight oriented from Hunga Tonga to Thailand 166 at various intervals: 8.5-10 UT (green double-head arrow), 12-15 UT (gray double-head 167 arrow), 16-18 UT (sky blue double-head arrow), and 18.5-20 UT (pink double-head ar-168 row). During the 12-15 UT interval, these disturbances began intensifying around 12 UT, 169 approximately 8 hours after the eruption of HTHH. The sudden fluctuations in dTEC, 170 ranging between 1-3 TECU around 13 UT (referred to the orange arrow (1) in Figure 171 4), demonstrated a significant increase, resembling a TID shock wave. This coincides with 172 the first peak of VTEC arrival in Thailand, occurring at approximately 12:30 UT (marked 173 174 by the orange arrow (1) in Figure 1a), before gradually diminishing.

Around 14 UT, another significant increase in dTEC occurred, indicating the ar-175 rival of another significant TID shock wave (referred to as purple (2') in Figure 4), cor-176 responding to the sudden fluctuations of the second peak of VTEC at $\sim 13:45$ UT (pur-177 ple arrow (2) in Figure 1a). These fluctuations gradually diminish in intensity from 15 178 UT onwards, showing only small variations thereafter. This second disturbance decayed 179 more rapidly, resembling the rate of decay observed in the first wave packet. Other small 180 ionospheric conditions remained significantly disturbed for several hours, with background 181 disturbances persisting until returning to normal storm-time conditions over Thailand. 182

At a distance of 10,000-km range, we noted three occurrences where the dTEC envelopes correspond with Zhang et al. (2022) during the same time intervals. Specifically, wave packet (1') corresponds to the dTEC arrival time (\sim 13 UT) for wave travel along the great circle route at a velocity of 350 m/s, while wave packet (2') corresponds to the dTEC arrival time (~14 UT) for wave travel along the great circle route at a velocity of 300 m/s. Additionally, the dTEC observed in the interval 8.5-10 UT, 16-18 UT, and 18.5-20 UT could be associated with observing other medium-scale TID (MSTID) with the 20-minute wave period. However, this TID dissipated quickly due to its rapid velocity. The faster wave would have already reached the 10,000 km mark because it diminishes more rapidly.

During 12-15 UT on January 15th, a high level of variation in the distance-time 193 profile of ROTI (KMI6) (Figure 3b) was observed over a horizontal distance of more than 194 1000 km. The changes in ROTI demonstrated fluctuations, indicating the presence of 195 equatorial plasma bubbles (EPBs) over Thailand, consistent with the observations of EPB 196 "C" reported by Sun et al. (2022). The hourly VHF radar (CPN1) (Figure 3c) also presents 197 the formation dynamics of EPBs. Discrepancies between the VHF radar data and ROTI 198 were observed. It could be due to differences in observation locations or the propaga-199 tion characteristics of EPBs. Spatial expansion at very high altitudes, up to approximately 200 700 km, was noted at 14 UT (Figure 3c-2) and 15 UT (Figure 3c-3) based on spectral 201 data analysis. Consequently, the EPBs disappeared but reassembled around 16 UT (Fig-202 ure 3c-4) at different locations, suggesting the formation of new EPBs during the recov-203 ery phase between 17-18 UT (Figures 3c-5 and 3c-6). 204

205 4 Summary

The data collected from Thai stations highlighted substantial deviations from the 206 usual ionospheric conditions attributed to moderate and minor geomagnetic storm drivers 207 and a volcanic eruption. Disturbances observed on January 15th, characterized by a sig-208 nificant increase in VTEC, were typical of geomagnetic storms. However, unusual hor-209 izontal wave-like perturbations in dTEC, commencing around 9 hours after the eruption, 210 were related to volcanic activity. Two significant TID shock waves reached Thailand at 211 13 UT (350 m/s) and 14 UT (300 m/s), respectively, lasting about 3 hours. The observed 212 VTEC depletion on January 16th confirmed the impact of geomagnetic storm activity 213 during the recovery phase in low-latitude regions. On this day, distinct EPBs were ob-214 served on a large scale, spanning over 1000 km horizontally and up to 700 km in alti-215 tude. These EPBs appear to have been produced by the HTHH eruption. In the future, 216 the formations of EPBs resulting from the volcanic eruption, particularly compared to 217 those arising during the main phase of geomagnetic storm events, should be further in-218 vestigated. Additional data from diverse stations, incorporating information or relevant 219 events or aftermath effects of the explosion, is essential to comprehensively analyze these 220 phenomena. Moreover, this event holds significance for understanding disruptions caused 221 by external and internal energy sources, with potential implications for future commu-222 nication signal impacts. 223

224 Open Research

The solar wind parameters and the geomagnetic indices are obtained by NASA/GSFC's 225 Space Physics Data Facility's OMNIWeb service (https://omniweb.gsfc.nasa.gov/), and 226 the World Data Center for Geomagnetism, Kyoto (https://wdc.kugi.kyoto-u.ac.jp/). The 227 SuperMAG collaborators (https://supermag.jhuapl.edu/info/?page=acknowledgement) 228 provide AE index data. The Kp index is available from GFZ German Research Centre 229 for Geosciences (https://kp.gfz-potsdam.de/en/data). The KMI6 and CPN1 GNSS data 230 are provided by the Thai GNSS and Space Weather Information Data Center hosted at 231 King Mongkut's Institute of Technology Ladkrabang (KMITL), Thailand (http://iono-232 gnss.kmitl.ac.th). The worldwide GNSS data are available from the CEDAR Madrigal 233 database (http:// cedar.openmadrigal.org/). The VHF data are contributed from the 234 Southeast Asia Low-latitude IOnosphere Network (SEALION) under the National In-235

stitute of Information and Communications Technology (NICT) service (https://aer-nc web.nict.go.jp/sealion/index.html).

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