Near-Automated Estimate of City Nitrogen Oxides Emissions Applied to South and Southeast Asia

Gongda Lu¹, Eloise Ann Marais¹, Karn Vohra¹, Rebekah P Horner¹, Dandan Zhang², Randall V Martin³, and Sarath K Guttikunda⁴

¹University College London ²Washington University in St Louis ³Washington University in St. Louis ⁴Urban Emissions

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Abstract

Cities in South and Southeast Asia are developing rapidly without routine, up-to-date knowledge of air pollutant precursor emissions. This data deficit can potentially be addressed for nitrogen oxides (NO_x) by deriving city NO_x emissions from satellite observations of nitrogen dioxide (NO_2) sampled under windy conditions. NO_2 plumes of isolated cities are aligned along a consistent wind-rotated direction and a best-fit Gaussian is applied to estimate emissions. This approach currently relies on non-standardized selection of the area to sample around the city centre and Gaussian fits often fail or yield nonphysical parameters. Here, we automate this approach by defining many (54) sampling areas that we test with TROPOspheric Monitoring Instrument (TROPOMI) NO_2 observations for 2019 over 19 cities in South and Southeast Asia. Our approach is efficient, adaptable to many cities, standardizes and eliminates sensitivity of the Gaussian fit to sampling area choice, and increases success of deriving annual emissions from 40-60% with one sampling area to 100% (all 19 cities) with 54. The annual emissions we estimate range from 16 ± 5 mol s⁻¹ for Yangon (Myanmar) and Bangalore (India) to 125 ± 41 mol s⁻¹ for Dhaka (Bangladesh). With the enhanced success of our approach, we find evidence from comparison of our top-down emissions to past studies and to inventory estimates that the wind rotation and EMG fit approach may be biased, as it does not adequately account for spatial and seasonal variability in NO_x photochemistry. Further methodological development is needed to enhance its accuracy and to exploit it to derive sub-annual emissions.

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- 3 Gongda Lu¹, Eloise A. Marais¹, Karn Vohra¹, Rebekah P. Horner¹, Dandan Zhang²,
- 4 Randall V. Martin², Sarath Guttikunda^{3,4}
- ⁵ ¹Department of Geography, University College London, Gower Street, London, UK.
- ⁶ ²Department of Energy, Environmental, and Chemical Engineering, Washington University
- 7 in St. Louis, St. Louis, MO, USA.
- 8 ³Transportation Research and Injury Prevention (TRIP) Center, Indian Institute of
- 9 Technology, New Delhi, 110016, India.
- ⁴Urban Emissions, New Delhi, 110016, India.
- 11
- 12 Corresponding author: Eloise A. Marais (<u>e.marais@ucl.ac.uk</u>)
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14 Key points:

- A refined approach to estimate nitrogen oxides emissions for isolated cities using
 wind fields and satellite nitrogen dioxide data.
- Many sampling areas defined for each city, increasing success of deriving emissions
 from 40-60% for one area to 100% for 54 areas.
- Applied to 19 cities in South and Southeast Asia to estimate annual emissions of 23 to
 181 kilotonnes in 2019.
- 21

22 Abstract

23 Cities in South and Southeast Asia are developing rapidly without routine, up-to-date 24 knowledge of air pollutant precursor emissions. This data deficit can potentially be addressed for nitrogen oxides (NO_x) by deriving city NO_x emissions from satellite observations of 25 nitrogen dioxide (NO₂) sampled under windy conditions. NO₂ plumes of isolated cities are 26 27 aligned along a consistent wind-rotated direction and a best-fit Gaussian is applied to estimate emissions. This approach currently relies on non-standardized selection of the area to sample 28 around the city centre and Gaussian fits often fail or yield non-physical parameters. Here, we 29 30 automate this approach by defining many (54) sampling areas that we test with TROPOspheric Monitoring Instrument (TROPOMI) NO2 observations for 2019 over 19 cities in South and 31 Southeast Asia. Our approach is efficient, adaptable to many cities, standardizes and eliminates 32 33 sensitivity of the Gaussian fit to sampling area choice, and increases success of deriving annual emissions from 40-60% with one sampling area to 100% (all 19 cities) with 54. The annual 34 emissions we estimate range from 16±5 mol s⁻¹ for Yangon (Myanmar) and Bangalore (India) 35 36 to 125±41 mol s⁻¹ for Dhaka (Bangladesh). With the enhanced success of our approach, we find evidence from comparison of our top-down emissions to past studies and to inventory 37 estimates that the wind rotation and EMG fit approach may be biased, as it does not adequately 38 account for spatial and seasonal variability in NO_x photochemistry. Further methodological 39

- 40 development is needed to enhance its accuracy and to exploit it to derive sub-annual emissions.
- 41

42 Plain Language Summary

43 Cities are a large source of nitrogen oxides (NO_x) that go on to form many types of air pollutants of harm to human health. City NO_x emissions estimated with observations from space-based 44 45 instruments are vital in regions that lack access to up-to-date, locally developed inventories. Success of obtaining satellite-derived emissions hinges on user selection of a sampling area 46 around each city centre. Here we present an automated, efficient method that uses many (54) 47 sampling areas. When tested on 19 cities in South and Southeast Asia, annual NO_x emissions 48 49 are obtained for all 19 cities compared to about half the selected cities when using a single sampling area. With this updated approach, we estimate total NO_x emissions in 2019 that range 50 from 23 kilotonnes for Yangon and Bangalore to almost 10-times more (181 kilotonnes) for 51 52 Dhaka. The greater success of our updated approach also helps us identify that the accuracy of emissions derivation from satellite observations should be further improved by accounting for 53 the influence of spatial and seasonal variability in NO_x photochemistry. 54

55

56 1 Introduction

Nitrogen oxides (NO_x \equiv NO₂ + NO) react to form particulate nitrate and tropospheric 57 ozone and deposit to sensitive habitats (Luo et al., 2019; Sillman, 1999), thus degrading air 58 59 quality, altering climate, and adversely affecting human health and the environment (Grulke & Heath, 2020; Lelieveld et al., 2015; Yue et al., 2017; Marais et al., 2023). Controls targeting 60 anthropogenic sources of NO_x have been extensively implemented in cities in Europe, the US 61 and China (Curier et al., 2014; de Foy et al., 2016; Silvern et al., 2019). In cities in other parts 62 of the world, particularly South and Southeast Asia, NO_x is increasing rapidly due to fast 63 economic development and limited or absent air quality policies (Vohra et al., 2021; 2022). 64 65 Vohra et al. (2022) used 14 years of satellite observations of NO₂ from the Ozone Monitoring Instrument (OMI) to infer increases of ~1-14 % a⁻¹ in surface NO₂ pollution in almost all rapidly 66 developing large cities in South and Southeast Asia. Only in Jakarta did NO₂ decline due to 67 emission controls applied to vehicles (Vohra et al., 2022). Population projections suggest that, 68

69 by 2100, one-fifth of the world's most populous cities will be in Southeast Asia (Hoornweg & 70 Pope, 2017), necessitating reliable and up-to-date NO_x emissions estimates for assessing the 71 impact of this growth on urban air quality and for informing air quality policies.

72 Bottom-up inventories provide estimates of anthropogenic NO_x emissions, but publicly available versions for South and Southeast Asia do not adequately represent contemporary 73 local conditions, as these are derived using outdated activity data, are resource-intensive to 74 produce so lag the present day, are at spatial resolutions that are coarser than many cities in the 75 region, and data needed to compile the inventories do not exist for many countries (Kurokawa 76 77 & Ohara, 2020). The two most used bottom-up inventories for these regions are the Regional Emission inventory in Asia (REAS) (Kurokawa & Ohara, 2020) and the inventory known as 78 MIX, a mosaic of REAS and other regional inventories (Li et al., 2017). REAS and MIX are 79 80 at ~25 km resolution, MIX only covers 2 years of data, and the most recent years are 2015 for REAS and 2010 for MIX. Still, REAS and inventories used to create MIX are routinely 81 incorporated in global inventories such as the Community Emissions Data System (CEDS_{GBD-} 82 83 MAPS) (McDuffie et al., 2020), and Hemispheric Transport of Air Pollution (HTAP) (Crippa et 84 al., 2023).

85 Independent and contemporary estimates of city NO_x emissions can be derived with satellite observations of tropospheric NO₂ vertical column densities (VCDs) without the need 86 for resource-intensive computer models. A method first proposed by Beirle et al. (2011) 87 involves selecting isolated cities and treating these as large point sources of NO_x. In this 88 approach, individual satellite pixels within a target domain centred on a city centre were split 89 90 into eight major wind directions to resolve the city plume in each direction. A mathematical 91 function was then fit to the plume to account for its Gaussian shape and exponential decay of NO₂. This fit, referred to as an Exponential Modified Gaussian (EMG), yields parameters that 92 are then used to estimate NO_x emissions. It also yields an effective lifetime of NO_x for the city 93 plume that is dominated by dispersion for the windy conditions sampled. As dispersion 94 95 dominates, the derived lifetime is much shorter than the chemical lifetime of NO_x that includes conversion to nitric acid (HNO₃) or organic nitrates (de Foy et al., 2014; Laughner & Cohen, 96 97 2019) and, to a lesser extent, dry deposition of NO₂ (Zhang et al., 2012). Beirle et al. (2011) 98 used OMI observations of NO₂ to derive NO_x emissions for eight global megacities. The Beirle 99 et al. (2011) approach required many (four) years of OMI data to achieve distinct plumes in each wind direction. 100

Valin et al. (2013) expanded on the approach developed by Beirle et al. (2011) by 101 demonstrating that all satellite data can instead be aligned along a single upwind-downwind 102 direction relative to the city centre. This approach reduced the number of observations needed 103 to distribute the data by wind direction and so extended application to a greater number of 104 geographically isolated cities over shorter sampling periods. Wind rotation of OMI 105 observations and the EMG fit have since been used to calculate city NO_x emissions 106 predominantly in the US (de Foy et al., 2014; Goldberg et al., 2019a; Lu et al., 2015) and for 107 select cities worldwide (Goldberg et al., 2021). Following the 2017 launch of the higher spatial 108 109 resolution TROPOspheric Monitoring Instrument (TROPOMI), the wind rotation, EMG fit, and related approaches have been extended to smaller isolated cities and shorter sampling 110 periods than was possible with OMI. Applications include cities in western Europe (Lorente et 111 al., 2019; Pope et al., 2022), China (Wu et al., 2021), the US (Goldberg et al., 2019b), and 112 worldwide (Lange et al., 2022), as well as investigating changes in NO_x emissions due to 113 COVID-19 lockdown measures in the New York Metropolitan Area (Tzortziou et al., 2022) 114 and for select cities in India, Argentina, and Spain (Lange et al., 2022). So far, the wind rotation 115 and EMG fit has only been applied to 5-13 cities in South and Southeast Asia as part of global 116 117 studies (Goldberg et al., 2021; Lange et al., 2022).

Even though there has been substantial development and use of the EMG fit, it still requires that a user define a sampling area around the city that effectively captures the wind rotated plume. The area selected varies with city size and plume length (Lu et al., 2015; Goldberg et al., 2019a; Lange et al., 2022). This approach often yields no or poor EMG fits and non-physical best-fit parameters (Laughner & Cohen, 2019), decreasing the likelihood of deriving top-down emissions. Selecting appropriate city-specific areas for the wide-ranging city sizes in South and Southeast Asia is also time consuming and not standardized.

Here we develop a near-automated and efficient EMG fitting routine for deriving annual city NO_x emissions, demonstrate the utility of this automation by applying it to TROPOMI NO_2 observations over isolated cities in South and Southeast Asia with wideranging city sizes, compare our top-down emissions to past studies and a global bottom-up inventory, and exploit the greater success of our updated sampling to identify opportunities to further develop the EMG fit approach.

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132 **2 Materials and Methods**

133 **2.1 TROPOMI NO₂ and City Selection**

We use Level 2 TROPOMI NO₂ tropospheric column VCDs for 2019 from the 134 Sentinel-5P Products Algorithm Laboratory (S5P-PAL) portal (https://data-portal.s5p-135 pal.com/; last acquired 30 January 2022). These data have been retrieved with a consistent 136 137 algorithm (version 02.03.01) and corrected for a low bias in NO₂ over polluted scenes (Eskes et al., 2021). TROPOMI achieves daily global coverage with a swath width of 2600 km, an 138 equator crossing time of 13:30 local solar time (LST), and a nadir pixel resolution that increased 139 on 5 August 2019 from 7 km \times 3.5 km to 5.5 km \times 3.5 km. We use cloud-free, high-quality 140 data identified with a quality flag ≥ 0.75 (van Geffen et al., 2021). 141

To identify isolated cities appropriate for top-down estimate of NO_x emissions, we first 142 oversample TROPOMI NO₂ to obtain high-resolution gridded annual means $(0.05^{\circ} \times 0.05^{\circ})$; 143 ~6 km latitude \times ~5 km longitude) by weighting areas of overlap between the satellite pixels 144 and cells on a fixed latitude-longitude grid using tessellation (Sun et al., 2018). We use the 145 resultant gridded TROPOMI NO₂ shown in Figure 1 to manually select 19 cities that are 146 isolated hotspots. The 19 selected cities are Karachi, Islamabad, and Lahore in Pakistan; Kabul 147 in Afghanistan; Ahmedabad, Mumbai, Delhi, Bangalore, Chennai, and Kolkata in India; 148 Colombo in Sri Lanka; Dhaka in Bangladesh; Yangon in Myanmar; Bangkok in Thailand; 149 Kuala Lumpur in Malaysia; the sovereign city Singapore; Ho Chi Minh City in Vietnam; 150 Jakarta in Indonesia; and Manila in the Philippines. Other hotspots in Figure 1 are either not 151 cities, such as the coal-fired power plants concentrated in eastern India, or are not isolated, such 152 as Hanoi, Haiphong and Nam Dinh in northern Vietnam. 153

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Figure 1. Annual mean TROPOMI tropospheric NO2 VCDs over South and Southeast Asia in 156 2019. Maps show South (left) and Southeast (right) Asia TROPOMI NO2 oversampled to 0.05° 157 158 \times 0.05°. The 19 selected cities, numbered from east to west, are Karachi (1), Islamabad (5), and Lahore (6) in Pakistan; Kabul (2) in Afghanistan; Ahmedabad (3), Mumbai (4), Delhi (7), 159 Bangalore (8), Chennai (10), and Kolkata (11) in India; Colombo (9) in Sri Lanka; Dhaka (12) 160 161 in Bangladesh; Yangon (13) in Myanmar; Bangkok (14) in Thailand; Kuala Lumpur (15) in Malaysia; the sovereign city Singapore (16); Ho Chi Minh City (17) in Vietnam; Jakarta (18) 162 in Indonesia; and Manila in the Philippines (19). 163

165 **2.2 Wind Rotation and EMG Fit**

Figure 2 illustrates the major steps involved in the wind rotation and EMG fit to derive 166 annual NO_x emissions for Singapore. The wind fields we use to calculate wind direction and 167 speed to retain TROPOMI NO₂ observations under windy conditions are the fifth generation 168 European (ERA5) 169 ReAnalysis 3D hourly u and v wind components (https://cds.climate.copernicus.eu/cdsapp#!/home; last acquired 18 March 2022) provided at 170 $0.25^{\circ} \times 0.25^{\circ}$ resolution. At each TROPOMI NO₂ pixel, we compute collocated mean ERA5 171 wind speeds and directions 30 min around 13:30 LST, the TROPOMI overpass time, in the 172 lowest 5 layers (\geq 900 hPa) to capture dispersion of mixed-layer near-surface NO₂ plumes. 173 Within a $4^{\circ} \times 4^{\circ}$ domain around each city centre, we isolate TROPOMI pixels with coincident 174 wind speeds $> 2 \text{ m s}^{-1}$, the threshold typically used for windy conditions (Beirle et al., 2011; 175 Pope et al., 2022). We rotate each TROPOMI NO₂ pixel by the angle of its wind direction, 176 preserving the distance of the pixel from the city centre. This aligns all pixels along the same 177 "upwind-downwind" direction that in our work is from north to south (Figure 2(a)). After wind 178 rotating all pixels in a year (as in Figure 2), we grid pixels onto a uniform $0.05^{\circ} \times 0.05^{\circ}$ grid 179 using simple point-in-box averaging (Figure 2(a)) and fill empty grid cells (grey squares in 180 Figure 2(a)) using nearest-neighbour interpolation to reduce low biases in the steps that follow. 181

182 Next, the 2D map in Figure 2(b) is converted to 1D line densities by summing all grid 183 cells in the across-wind (east-to-west) direction in 0.05° upwind-downwind (north-to-south) 184 increments. In the standard approach, a single area smaller than the $4^{\circ} \times 4^{\circ}$ domain is used,

defined by the distance upwind, downwind, and across-wind of the city centre. Instead of using 185 a single area, we define multiple areas that encompass the range of sizes typically used in past 186 studies (Goldberg et al., 2021; Lange et al., 2022; Laughner & Cohen, 2019). These, defined 187 as distances from the city centre, are 0.5°, 0.75°, and 1° upwind, 0.5°, 0.75°, 1.0°, 1.25°, 1.5°, 188 1.75°, 2.0° downwind, and 0.5°, 0.75°, and 1.0° across-wind, with the requirement that the 189 distance downwind of the city centre is \geq the distance upwind to capture the extent of the city 190 plume. This yields 54 areas and associated line densities. The sizes of the smallest and largest 191 192 areas sampled and the across-wind 0.05° increments summed to obtain line densities in the 193 smallest area sampled are shown in Figure 2(b).

194 The EMG model we use to fit to the observed 1D line densities is the Laughner & 195 Cohen (2019) formulation:

196
$$F(x|a, x_0, \mu_x, \sigma_x, B) = \frac{a}{2x_0} \exp\left(\frac{\mu_x}{x_0} + \frac{\sigma_x^2}{2x_0^2} - \frac{x}{x_0}\right) \operatorname{erfc}\left(-\frac{1}{\sqrt{2}}\left[\frac{x-\mu_x}{\sigma_x} - \frac{\sigma_x}{x_0}\right]\right) + B$$
(1),

where *x* is the distance of each line density upwind and downwind of the city centre (Figure 2(c)) and *a*, x_0 , μ_x , σ_x and *B* are best-fit parameters. Of these, *a* is total NO₂ in the plume (in moles), x_0 is the *e*-folding distance or length scale of NO₂ decay (in km), μ_x is the location of the apparent source relative to the city centre (in km) or the peak of the Gaussian fit that in Figure 2(c) is located ~20 km downwind or south of the city centre, σ_x is the Gaussian smoothing length scale (in km) that is ~2.355 × the Full Width at Half Maximum (FWHM), and *B* is background NO₂ (in moles m⁻¹).

We use initial guesses for the best-fit parameters in Equation (1) that are similar to those 204 205 from Laughner & Cohen (2019), but our fitting procedure differs. Laughner & Cohen (2019) used a non-linear interior point minimization algorithm (the *fmincon* function in MATLAB) to 206 optimize model parameters with 10 iterations per line density. Instead, we perform the fit with 207 the scipy.optimize.curve fit module from SciPy Python package version 1.7.3 and iterate on 208 209 the fit until the difference in fitting parameters between the current and previous iteration is negligible (< 0.001%) for at most 10 iterations. Fit convergence is usually achieved after 3 210 211 iterations. Only good-quality fits are retained, identified with goodness-of-fits $(R^2) > 0.8$, as in Laughner & Cohen (2019). We further screen for physically implausible best-fit parameters 212 using criteria similar to Laughner & Cohen (2019): a is positive, x_0 is at least 1.6 km 213 (approximately 1/e of the grid resolution), μ_x is within the sampling area, the emission width 214 is less than the *e*-folding distance ($\sigma_x < x_0$), background NO₂ is positive and less than the 215 maximum line density value, and the e-folding distance occurs between the plume centre and 216 the edge of the sampling area. We introduce an additional requirement to ensure that x_0 is within 217 the sampling area ($x_0 <$ length of sampling area downwind of the city centre). 218

The Singapore example in Figure 2 is an ideal city, as all 54 EMG fits are successful. Figure 2(c) shows that the observed line densities are most sensitive to the across-wind length, as this determines the amount of NO_2 summed to yield each line density. We will demonstrate in Section 3 that for many of the cities in Figure 1 a large number of EMG fits fail to meet the conditions for success, necessitating as many as 54 fits.

The successful EMG fits are used to calculate effective NO_x lifetimes (τ_{NO_x} ; reported in h) and midday NO_x emissions (E_{NO_x} ; in moles s⁻¹):

$$226 \qquad \tau_{NO_x} = \frac{x_0}{\omega} \tag{2}$$

227
$$E_{NO_x} = \gamma \times \frac{a}{\tau_{NO_x}}$$
(3),

- where ω is the sampling area mean wind speed (in m s⁻¹) and γ is the unitless molar ratio of [NO_x]/[NO₂] to convert moles NO₂ to moles NO_x. The up to 54 individual estimates of τ_{NO_x} and E_{NO_x} are averaged to obtain values for each city.
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Figure 2. Illustration of major steps in the wind rotation and EMG fit to derive annual NO_x 234 emissions for Singapore. The main steps in each panel are wind rotate and grid windy scene 235 TROPOMI NO₂ pixels to $0.05^{\circ} \times 0.05^{\circ}$ (a), fill data gaps (b), and fit the EMG function (Eq. 236 (1)) (solid lines) to observed line densities (filled circles) (c). In (b), black rectangles show the 237 extent of the largest and smallest sampling areas and dashed lines in the smallest area show the 238 0.05° increments used to calculate the line densities in (c). All 54 successful EMG fits, 18 lines 239 for each of the three across-wind lengths, are shown in (c). Values in (c) give the mean and 240 standard deviation of the city NO_x emissions (Eq. (3)), effective NO_x lifetime (Eq. (2)), and 241 sampling area ERA5 wind speed. The goodness-of-fit (\mathbb{R}^2) is ≥ 0.99 for all fits in (c). 242

We use the same $[NO_x]/[NO_2] = 1.32$ value as Beirle et al. (2011) and subsequent studies to represent rapid cycling between NO and NO₂. Liu et al. (2022) determined with synthetic experiments that city NO_x emissions are relatively unaffected by variability in $[NO_x]/[NO_2]$, but that study was for US cities. Surface measurements aid in determining

suitability of $[NO_x]/[NO_2] = 1.32$, but these are limited to cities in India and have data quality 248 issues (Vohra et al., 2021). Instead, we use the GEOS-Chem model to assess suitability of the 249 1.32 value. We simulate the model in 2019 and sample the lowest model layer around the 250 TROPOMI overpass time. We use output from a coarse and finer resolution version of GEOS-251 Chem to also test sensitivity of this ratio to model resolution, especially given many of these 252 cities are coastal (Figure 1). We use the classical configuration of the model that operates on a 253 254 single computational node, called GEOS-Chem Classic (GCClassic), and the highperformance model configuration (GCHP) that is a parallelized across multiple computational 255 nodes to enable finer resolution global simulations (Eastham et al., 2018). GCClassic is version 256 13.3.4 (https://doi.org/10.5281/zenodo.5764874) run on a fixed $2^{\circ} \times 2.5^{\circ}$ global grid and 257 GCHP is version 13.4.1 (https://doi.org/10.5281/zenodo.6564711) run on a C360 global grid 258 (~25 km \times ~31 km). GCClassic and GCHP use the same vertical grid and chemical mechanism. 259 For GCClassic, grid squares that overlap with each city are sampled, whereas for GCHP, we 260 261 use city sampling extents determined from a combination of administrative and geographic boundary shapefiles and Google Maps (Figure S1). Midday sampling is at 12:00 to 15:00 LST 262 from GCClassic and 13:00 to 14:00 LST from GCHP. At midday, NO_x is in photochemical 263 steady state, so the relative abundance of NO and NO₂ is insensitive to the extent of the 264 sampling window around midday (Potts et al., 2021). 265

We calculate uncertainties in the NO_x emissions by adding individual errors in 266 quadrature. These include best-fit parameters x_0 and a, sampling area mean wind speed ω , the 267 TROPOMI NO₂ observations, and $[NO_x]/[NO_2]$. We use the relative standard deviation from 268 all successful EMG fits to calculate city-specific errors in x_0 and a. For ω , we consider errors 269 due to the choice of spatial and temporal sampling and the threshold used for windy conditions. 270 We use the Beirle et al. (2011) estimated 10% error in temporal sampling choice and 5% error 271 due to vertical sampling choice. We conduct our own tests of the sensitivity to threshold and 272 spatial sampling choice. For $[NO_x]/[NO_2]$ we assess whether the 10% error attributed to this 273 274 variable by Beirle et al. (2011) is appropriate by quantifying the percent deviation of GCClassic and GCHP [NO_x]/[NO₂] from 1.32. Beirle et al. (2011) applied a 30% error to OMI that is also 275 appropriate for TROPOMI. Even though uncertainties in TROPOMI slant columns (NO₂ along 276 the viewing path) are much less than those from OMI (van Geffen et al., 2020), the air mass 277 factor used to convert slant columns to VCDs remains the largest contributor to errors in NO₂ 278 279 VCDs and is similar for OMI and TROPOMI (van Geffen et al., 2021).

280 **2.3 Bottom-up Anthropogenic Emissions**

We compare our top-down estimates to anthropogenic NO_x emissions from the widely 281 used bottom-up HTAP inventory version 3 (HTAP v3) (Crippa et al., 2023). HTAP v3 has 282 high enough spatial resolution $(0.1^{\circ} \times 0.1^{\circ})$ to resolve cities selected in Figure 1. The most 283 recent year is 2018, achieved by extending emissions from the regional REAS inventory ending 284 in 2015 to the year 2018 with trends from the Emissions Database for Global Atmospheric 285 Research (EDGAR) inventory. The same sampling boundaries as GCHP are used (Section 2.2; 286 Figure S1). The HTAP v3 NO_x emissions include contributions from aviation, transport (road, 287 rail, pipeline, inland waters), shipping, energy, industry, and residential sectors. 288

Cities targeted can be influenced by non-anthropogenic NO_x sources, such as open burning of biomass (Marvin et al., 2021) and natural sources such as soils (Weng et al., 2020) and lightning (Miyazaki et al., 2014). We assess suitability of comparing our top-down emissions to anthropogenic bottom-up emissions only by determining the percent contribution of anthropogenic emissions to total NO_x emissions. To do this, we simulate total NO_x emissions with the Harmonized Emissions Component (HEMCO) standalone model version 3.0.0 (<u>https://zenodo.org/records/4984639</u>; last accessed 20 March 2022) (Lin et al., 2021) and 296 sample the same spatial extent as GCHP and HTAP_v3 (Figure S1). HEMCO is run at a spatial 297 resolution of $0.25^{\circ} \times 0.3125^{\circ}$ (~ 28 km latitude × ~ 33 km longitude). HEMCO calculates open 298 biomass burning emissions using the Global Fire Emissions Database with small fires 299 (GFED4s) inventory (Randerson et al., 2017) and reads in and processes lightning and soil NO_x 300 from offline emissions at the same resolution as HEMCO (Murray et al., 2012; Weng et al., 301 2020).

Bottom-up emissions from HTAP v3 are 24-h means, whereas top-down estimates 302 derived using TROPOMI are representative of midday emissions. Goldberg et al. (2021) 303 304 multiplied satellite-derived midday NO_x emissions by 0.77 to convert midday top-down NO_x emissions to 24-h means for comparison to bottom-up inventories. This value was inferred 305 from bottom-up emissions estimates for the Netherlands, so may not be suitable for the selected 306 307 cities in South and Southeast Asia. The hourly scaling factors used by HEMCO for the chosen cities range from 0.70 to 1.16. These are for the year 2000 and are extrapolations of values for 308 conditions in Europe, so may not be suitable for the year and cities targeted in this study. Given 309 310 this, we do not scale top-down emissions and instead discuss whether differences in averaging times contribute to discrepancies between top-down and bottom-up emissions estimates. 311

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313 **3 Results and Discussion**

314 **3.1 Wind Rotation and EMG Fit Metrics**

Isolating windy condition (> 2 m s^{-1}) satellite pixels removes 8-34% of all 2019 quality-315 and cloud-screened TROPOMI NO2 pixels for most cities in Figure 1. Cities with greater data 316 317 loss are Lahore (43% data loss), Kabul (58%) and Islamabad (63%). No spatial data gap filling (Section 2.2, Figure 2) is needed within the areas sampled, due to the high sampling frequency 318 of TROPOMI. If only a single domain size is selected, annual EMG fits meet all criteria for 319 success for 7 to 12 of the 19 cities in Figure 1, depending on the sampling area chosen. Using 320 321 our extended method, we successfully derive annual NO_x emissions for all 19 cities, due to the enhanced probability of obtaining at least one successful EMG fit. 322

Figure shows the number of successful EMG fits (orange bars) range from 3 (Kabul) to 323 all 54 (Singapore). Singapore, Dhaka, Jakarta, Karachi, Manila, and Mumbai are least impacted 324 325 by the choice of sampling area. The 6 cities in Figure 3 with < 20 fits are most likely to fail if only a single sampling area is used. For all retained EMG fits, differences between observed 326 and fitted NO₂ line densities, the fit residuals, are negligible. The most common causes for a 327 failed EMG fit rank as: background NO₂ (B in Equation (1)) > maximum NO₂ line density (36%) 328 of all fits conducted), $R^2 \le 0.8$ (24%), emission width > *e*-folding distance (19%), total plume 329 NO₂ (a in Equation (1)) < 0 (13%), and e-folding distance > the downwind length of the 330 sampling area (12%). Multiple causes can co-occur in a single fit, so cumulative percentages 331 332 exceed 100%.

We also test sensitivity of top-down NO_x emissions to the choice of wind speed 333 threshold and horizontal sampling extent to attribute an error to these. For this, we apply a 334 stricter wind speed threshold of 3 m s⁻¹ and test the difference in NO_x emissions if instead of 335 filtering for windy conditions using pixel-mean wind fields, we calculate a sampling-area mean 336 wind speed to filter for windy conditions as in Goldberg et al. (2019a). We apply these 337 conditions to a mid-sized sampling area of 0.75° upwind, 1.5° downwind, and $\pm 0.75^{\circ}$ across-338 wind. Variability in NO_x emissions for cities with successful EMG fits for all 4 wind sampling 339 340 conditions is at most 10% (Figure S2). Given these results, we attribute a 10% error to the choice of horizontal sampling and to the wind speed threshold. 341

GCClassic (coarse resolution) annual mean [NO_x]/[NO₂] for the target cities ranges 342 343 from 1.25 (Dhaka) to 1.41 (Kabul). The range in ratios from GCHP (finer resolution) is wider at 1.24 (Ahmedabad) to 1.64 (Kolkata). The difference in ratios between the coarse and fine 344 resolution models is typically $\pm 10\%$, except for a few cities with ratios from the fine resolution 345 model that exceed the coarse resolution model by 14% for Singapore, 16% for Lahore, 23% 346 for Dhaka, and 23% for Kolkata. This is because the fine resolution model better resolves the 347 city plume that includes a greater proportion of NO_x as NO from fresh emission sources. As 348 the difference between the model city ratios and the 1.32 value is $\pm 10\%$ for most cities, we use 349 the same 10% error for $[NO_x]/[NO_2]$ as Beirle et al. (2011). 350

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Figure 3. Successful EMG fits and top-down NO_x emissions for the cities targeted in this study. Bars are emissions (green) and the corresponding number of successful fits (orange). Black error lines are NO_x emission standard deviations for all successful fits. The orange dashed line at 54 indicates the maximum possible EMG fits. Emissions multiplied by ~1.45 yields emissions in Gg NO₂ a⁻¹.

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359 **3.2 Top-Down NO_x Emissions**

Green bars in Figure 3 show the mean annual top-down NO_x emissions for all cities (values are in Table S1). These range from $\sim 16 \text{ mol s}^{-1}$ for Bangalore and Yangon to $\sim 125 \text{ mol}$ s⁻¹ for Dhaka. The range in the total mass of NO_x emitted for these cities, assuming the midday

emission rate is reasonably representative of the 24-h emission rate, is 23-181 Gg NO_x as NO₂. 363 Emissions for most cities are $< 50 \text{ mol s}^{-1}$ ($<73 \text{ Gg NO}_x$ as NO₂ a⁻¹). Cities with emissions 364 between 50-100 mol s⁻¹ (73-145 Gg NO_x as NO₂ a⁻¹) include Karachi, Delhi, and Jakarta and > 365 100 mol s⁻¹ (> 145 Gg NO_x as NO₂ a⁻¹) include Bangkok, Singapore, and Dhaka. Emission 366 rates for Bangkok, Dhaka and Singapore are comparable to the range of top-down emissions 367 estimated for large, polluted cities in China using the EMG approach (Wu et al., 2021). The 368 369 effective lifetimes for the cities in Figure 1 (shown in Figure S3) range from 1.2 h for Colombo to 6.3 h for Kuala Lumpur. Variability in effective lifetimes depends most strongly on the 370 downwind extent of the plume. The Pearson's correlation coefficient, R, between city mean 371 372 effective lifetimes and x_0 values is 0.90.

373 For the target cities, the relative standard deviations of annual NO_x emissions (black 374 error lines in Figure 3) range from just 1% for Bangalore to 27% for Kuala Lumpur. This is far less than the equivalent Gaussian fit uncertainty of 10-50% estimated by Beirle et al. (2011) 375 for a single sampling area. The relatively large variability in Kuala Lumpur NO_x emissions is 376 because the smaller EMG sampling areas do not fully encompass the elongated wind rotated 377 city NO₂ plume, causing a low bias in NO_x emissions for the smaller areas sampled. The effect 378 of this is dampened by the almost 30 successful fits used to obtain mean NO_x emissions for this 379 city. The relative standard deviations of the NO_x lifetimes (Figure S3) range from 3% for 380 Bangalore to 37% for Chennai. The relative standard deviations of other parameters are $\sim 6\%$ 381 for wind speeds (Figure S4), 4% (Bangalore) to 38% (Chennai) for x_0 , and 4% (Kabul and 382 Bangalore) to 37% (Bangkok) for a. 383

The overall uncertainty in annual NO_x emissions we obtain by adding all error contributions in quadrature ranges from 32% for Bangalore and Yangon to 55% for Bangkok. Values for all cities are in Table S1. The TROPOMI NO₂ VCDs make the largest contribution to the overall uncertainty. The higher-end of our uncertainty estimates is similar to the typical ~50% uncertainty reported in past studies (Beirle et al., 2011; Verstraeten et al., 2018; Goldberg et al., 2021). We use our overall uncertainties in the comparison of our top-down emissions to values from the literature and from HTAP in the sections that follows.

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392 **3.3 Comparison to Top-Down Estimates from Past Studies**

To assess our approach, we compare in Figure 4 our annual NO_x emissions to values 393 from past studies that used similar sampling time periods and a single sampling area. These 394 include multivear (2017-2019) mean emissions from Goldberg et al. (2021) obtained using the 395 OMI sensor and emissions from Lange et al. (2022) obtained with select days of TROPOMI 396 data from 2018 to 2020. Goldberg et al. (2021) estimated emissions for 10 of the 19 cities in 397 our study. These we read from their Figure S10 for Karachi, Figure S11 for 4 cities in India, 398 and Figure S13 for 5 cities in Southeast Asia and divide by the 0.77 midday to 24-h scaling 399 factor used in that study. Emissions are reported by Lange et al. (2022) for 5 of the 19 cities in 400 our study. Based on the regression statistics in Figure 4, our emissions are typically ~26% more 401 than estimates from these past studies. Exceptions are Mumbai, Ahmedabad, and Chennai that 402 in our study are 16-29% less than Goldberg et al. (2021). Lange et al. (2022) used an earlier 403 version of the TROPOMI data product that has a known low bias in NO₂ VCDs over very 404 polluted scenes (van Geffen et al., 2022). Differences in TROPOMI data products are the likely 405 cause for our higher Delhi (by 27%) and Singapore (by 18%) emissions. Relatively small error 406 estimates from Lange et al. (2022) are because they only propagate error contributions from 407 the wind speed data and the EMG fit. 408

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Figure 4. Comparison of our and past top-down NO_x emissions. Symbols compare our emissions to those from Goldberg et al. (2021) (red) and Lange et al. (2022) (blue). Error bars are overall uncertainties for our study (Section 2.2, Table S1), the same 53% uncertainty applied to all cities by Goldberg et al. (2021) and the city-specific uncertainties for Lange et al. (2022). Lines are the Theil regression fit (solid black) and 1:1 relationship (dashed grey). Inset text gives the regression statistics and Pearson's correlation coefficient (R). Arrows and inset text for Dhaka give the error values that extend beyond the plotting range.

Discrepancies between Goldberg et al. (2021) and our emissions are not as 419 straightforward to diagnose, as Goldberg et al. (2021) use NO₂ VCDs from a different sensor 420 (OMI) and apply a systematic 37% increase to NO_x emissions to correct for a low bias in OMI 421 attributed to the coarse resolution a priori used in the NO₂ VCDs retrieval. Sampling area 422 choice may also be a factor. For example, the smallest of our 54 areas yields NO_x emissions of 423 102 mol s⁻¹ for Singapore that is 10 mol s⁻¹ less than the mean of all EMG fits. Goldberg et al. 424 425 (2021) used year-round OMI data for all cities except Delhi and Karachi. As these cities are north of 25°N, only May-September observations were used by Goldberg et al. (2021). We find 426 that Delhi and Karachi mean May-September TROPOMI NO2 VCDs in 2019 averaged within 427 the $4^{\circ} \times 4^{\circ}$ domain selected for each city (Figure 2(a)-(b)) are 11-12% less than those in 428 429 October-April, due to the shorter photochemical lifetime of NO_x in the warmer months. Open biomass burning emissions also influence seasonality in the TROPOMI NO₂ VCDs, but the 430 431 EMG fit accounts for this by distinguishing background NO_2 (*B* in Equation (1)) from NO_2 in 432 the city plume (*a* in Equation (1)).

We find that if we apply the EMG fit to individual months for Delhi and Karachi, all 54 EMG fits fail for Delhi in July-August and yield spurious results in September due to large data loss resulting from persistent clouds during the monsoon season. All 12 months are retained for Karachi, Singapore and Manila. November-April mean values of *a* are 21% more than in May-October for Karachi, 9% more for Singapore, and 39% more for Manila. This suggests that using NO₂ VCDs for a portion of the year may yield systematic biases in 439 emissions that may not reflect seasonality in the underlying activities affecting the emissions. Larger wintertime than summertime emissions have also been reported in the global study of 440 Lange et al. (2022). They quantified summer-to-winter emission ratios of ~0.5 for Colombo 441 and Delhi. The top-down emissions calculation (Equation (3)) does not fully account for 442 seasonality in photochemistry. The derived effective NO_x lifetimes used to calculate NO_x 443 emissions (Equation (2)) are mostly influenced by dispersion. As a result, the effective lifetimes 444 445 are much shorter than the expected chemical lifetimes of NO_x (de Foy et al., 2014). In the synthetic experiment scenarios tested by de Foy et al. (2014), the EMG fit applied to wind 446 rotated data yielded an effective lifetime of 4 h for a 12-h chemical lifetime scenario. According 447 to Shah et al. (2020), the chemical lifetime of NO_x for central-eastern China centred at ~35°N. 448 449 the northerly portion of our domain, ranges from ~6 h in summer to ~24 h in winter. None of the monthly effective lifetimes for our target cities reproduces this seasonality and the longest 450 lifetime is 13.3±3.7 h for Yangon in November. The implication is that the size of absolute 451 452 emissions derived with sub-annual satellite data may be biased, but should have negligible effect if used to quantify relative trends, as in Goldberg et al. (2021) and Laughner & Cohen 453 (2019), for example. 454

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456 **3.4 Comparison to Bottom-up Emissions**

Figure 5 compares annual top-down and bottom-up NO_x emissions. According to our 457 HEMCO simulations, anthropogenic sources account for most (>87%) annual NO_x emissions. 458 The relative differences between our top-down estimates and the bottom-up inventory are 459 within 50% for Mumbai (1%), Bangkok (2%), Chennai (9%), Ahmedabad (11%), Kolkata 460 (21%), Singapore (21%), Bangalore (32%), Manila (35%), and Kuala Lumpur (46%). A 50-461 100% difference occurs for Ho Chi Minh City (53%), Jakarta (54%), Delhi (64%), and 462 463 Colombo (91%). Even greater relative differences occur for Karachi (2.1 times), Islamabad (2.1 times), Lahore (2.4 times), Yangon (3.3 times), Dhaka (6.9 times), and Kabul (11-fold). 464 The largest absolute discrepancies are for Dhaka and Jakarta. Bottom-up emissions are 107 465 mol s⁻¹ less than the top-down values for Dhaka and 78 mol s⁻¹ more for Jakarta. On a mass 466 basis, this is equivalent to a 155 Gg NO_x as NO₂ underestimate for Dhaka and a 113 Gg NO_x 467 as NO₂ overestimate for Jakarta. 468

The different years used (2018 for HTAP, 2019 for TROPOMI) should at most account 469 for a 14% difference in emissions, based on the size of annual trends inferred by Vohra et al. 470 471 (2022) using long-term observations of OMI NO₂ VCDs over large and fast-growing cities in South and Southeast Asia. Vohra et al. (2022) identified that emission inventories do not 472 capture the steep decline in NO_x emissions in Jakarta attributed to national policies targeting 473 vehicles. In addition to misrepresenting annual changes in underlying activities, the emission 474 factors are mostly informed by studies in China and Japan (Kurokawa & Ohara, 2020). The 475 bottom-up and top-down emissions differences for many cities also exceed the $\pm 30\%$ 476 difference that results from the choice of bottom-up emissions grid sampling and the $\pm 30\%$ 477 478 difference from the timing of the top-down (midday) and bottom-up (24-h) estimates inferred by Goldberg et al. (2021). 479

480 Apparent in Figure 5 is a latitudinal pattern in the discrepancies. Top-down emissions 481 are greater than bottom-up emissions for cities to the north and vice versa for cities to the south, 482 so that in general top-down emissions exceed bottom-up emissions in South Asia and vice 483 versa in Southeast Asia. NO_x chemical loss varies with latitude, due to variability in the amount 484 of sunlight available to form hydroxyl and peroxy radicals required to form HNO₃ and organic 485 nitrates, the main daytime chemical loss pathway for NO_x . This latitudinal pattern is likely 486 because the EMG fit also does not fully account for spatial variability in NO_x photochemistry, 487 imparting a bias in the top-down emissions. The size of this bias will depend on the relative 488 contribution of NO_x chemical loss to total loss in the wind rotated plume.

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Figure 5. Comparison of annual top-down and bottom-up NO_x emissions for target cities. Data are coloured by city centre latitude and split into top-down NO_x emissions < 40 mol s⁻¹ (a) and \geq 40 mol s⁻¹ (b). Error bars are the overall uncertainty in top-down emissions estimates. Grey lines indicate 1:1 agreement (solid) and ±50% difference (dashed). The bottom-up emissions sampling extent of each city is in Figure S1. Data used to generate the figure are in Table S1.

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498 4 Conclusions

City nitrogen oxides (NO_x) emissions can be derived with a now well-established 499 approach using satellite observations of nitrogen dioxide (NO₂), wind rotation and a Gaussian 500 501 fit to the city plume. Issues with this approach are that the choice of sampling area around the city centre is not standardized and so is prone to subjective area selection and the Gaussian fit 502 often fails or yields non-physical best-fit parameters. Here we address these issues by applying 503 54 sampling areas to isolated cities. We test our method with TROPOspheric Monitoring 504 Instrument (TROPOMI) NO₂ observations for 2019 over 19 large, isolated cities in South and 505 Southeast Asia that lack contemporary, publicly available bottom-up emissions estimates. 506

507 Annual NO_x emissions, obtained for all 19 cities, are < 73 Gg NO_x as NO₂ a⁻¹ for most cities, between 73-145 Gg NO_x as NO₂ a^{-1} for Karachi, Delhi, and Jakarta and > 145 Gg NO_x 508 as NO₂ a⁻¹ for Bangkok, Dhaka, and Singapore. The overall uncertainty in the annual emissions 509 is 30-60%. Our emissions estimates are in general \sim 27% more than past studies that use a single 510 sampling area, due to differences in satellite data products and months targeted. The latter we 511 suggest may lead to biases, as the top-down emissions estimate does not properly account for 512 seasonality in photochemical loss of NO_x. Relative differences between our top-down estimates 513 and a widely used bottom-up inventory are < 50% for 9 of the 19 cities, within 50-100% for 514 Ho Chi Minh City, Jakarta, Delhi, and Colombo, and much greater for Karachi (2.1 times), 515 Islamabad (2.1 times), Lahore (2.4 times), Yangon (3.3 times), Dhaka (6.9 times), and Kabul 516 (11-fold). There is a latitudinal dependence of the size of these discrepancies that we suggest 517

is because the top-down approach also does not properly account for spatial variability in the chemical lifetime of NO_x .

520 The increased success of deriving NO_x emissions with our updated approach enables 521 us to identify that further development is needed to account for time and space variability in 522 the chemical lifetime of NO_x to fully exploit the top-down approach to interrogate seasonality 523 in emissions, to validate bottom-up emissions, to exploit hourly observations from 524 geostationary instruments, and to inform air quality regulation.

525 Data and Software Availability

526 The TROPOMI tropospheric columns for 2019 are publicly available from the S5P-PAL Data 527 Portal (https://data-portal.s5p-pal.com/). GEOS-Chem source codes are preserved on Zenodo 528 by The International GEOS-Chem User Community (2021) for GCClassic version 13.3.4 and 520 htt The International GEOS Chem User Community (2022) for GCClassic version 12.4.1

529 by The International GEOS-Chem User Community (2022) for GCHP version 13.4.1.

530 Author Contributions

GL developed the methodology, GL and EAM processed, analysed and interpreted the data.
GL and EAM prepared the manuscript. KV assisted in data collection and analysis. RPH and
DZ conducted the GEOS-Chem simulations (RPH: GCClassic; DZ: GCHP). RVM contributed

- to the methodology. SG contributed to interpretation of the results. All co-authors provided
- 535 editorial input.

536 **Conflicts of interest**

537 The authors declare there are no conflicts of interest.

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Near-Automated Estimate of City Nitrogen Oxides Emissions Applied to South and Southeast Asia

- 3 Gongda Lu¹, Eloise A. Marais¹, Karn Vohra¹, Rebekah P. Horner¹, Dandan Zhang²,
- 4 Randall V. Martin², Sarath Guttikunda^{3,4}
- ⁵ ¹Department of Geography, University College London, Gower Street, London, UK.
- ⁶ ²Department of Energy, Environmental, and Chemical Engineering, Washington University
- 7 in St. Louis, St. Louis, MO, USA.
- 8 ³Transportation Research and Injury Prevention (TRIP) Center, Indian Institute of
- 9 Technology, New Delhi, 110016, India.
- ⁴Urban Emissions, New Delhi, 110016, India.
- 11
- 12 Corresponding author: Eloise A. Marais (<u>e.marais@ucl.ac.uk</u>)
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14 Key points:

- A refined approach to estimate nitrogen oxides emissions for isolated cities using
 wind fields and satellite nitrogen dioxide data.
- Many sampling areas defined for each city, increasing success of deriving emissions
 from 40-60% for one area to 100% for 54 areas.
- Applied to 19 cities in South and Southeast Asia to estimate annual emissions of 23 to
 181 kilotonnes in 2019.
- 21

22 Abstract

23 Cities in South and Southeast Asia are developing rapidly without routine, up-to-date 24 knowledge of air pollutant precursor emissions. This data deficit can potentially be addressed for nitrogen oxides (NO_x) by deriving city NO_x emissions from satellite observations of 25 nitrogen dioxide (NO₂) sampled under windy conditions. NO₂ plumes of isolated cities are 26 27 aligned along a consistent wind-rotated direction and a best-fit Gaussian is applied to estimate emissions. This approach currently relies on non-standardized selection of the area to sample 28 around the city centre and Gaussian fits often fail or yield non-physical parameters. Here, we 29 30 automate this approach by defining many (54) sampling areas that we test with TROPOspheric Monitoring Instrument (TROPOMI) NO2 observations for 2019 over 19 cities in South and 31 Southeast Asia. Our approach is efficient, adaptable to many cities, standardizes and eliminates 32 33 sensitivity of the Gaussian fit to sampling area choice, and increases success of deriving annual emissions from 40-60% with one sampling area to 100% (all 19 cities) with 54. The annual 34 emissions we estimate range from 16±5 mol s⁻¹ for Yangon (Myanmar) and Bangalore (India) 35 36 to 125±41 mol s⁻¹ for Dhaka (Bangladesh). With the enhanced success of our approach, we find evidence from comparison of our top-down emissions to past studies and to inventory 37 estimates that the wind rotation and EMG fit approach may be biased, as it does not adequately 38 account for spatial and seasonal variability in NO_x photochemistry. Further methodological 39

- 40 development is needed to enhance its accuracy and to exploit it to derive sub-annual emissions.
- 41

42 Plain Language Summary

43 Cities are a large source of nitrogen oxides (NO_x) that go on to form many types of air pollutants of harm to human health. City NO_x emissions estimated with observations from space-based 44 45 instruments are vital in regions that lack access to up-to-date, locally developed inventories. Success of obtaining satellite-derived emissions hinges on user selection of a sampling area 46 around each city centre. Here we present an automated, efficient method that uses many (54) 47 sampling areas. When tested on 19 cities in South and Southeast Asia, annual NO_x emissions 48 49 are obtained for all 19 cities compared to about half the selected cities when using a single sampling area. With this updated approach, we estimate total NO_x emissions in 2019 that range 50 from 23 kilotonnes for Yangon and Bangalore to almost 10-times more (181 kilotonnes) for 51 52 Dhaka. The greater success of our updated approach also helps us identify that the accuracy of emissions derivation from satellite observations should be further improved by accounting for 53 the influence of spatial and seasonal variability in NO_x photochemistry. 54

55

56 1 Introduction

Nitrogen oxides (NO_x \equiv NO₂ + NO) react to form particulate nitrate and tropospheric 57 ozone and deposit to sensitive habitats (Luo et al., 2019; Sillman, 1999), thus degrading air 58 59 quality, altering climate, and adversely affecting human health and the environment (Grulke & Heath, 2020; Lelieveld et al., 2015; Yue et al., 2017; Marais et al., 2023). Controls targeting 60 anthropogenic sources of NO_x have been extensively implemented in cities in Europe, the US 61 and China (Curier et al., 2014; de Foy et al., 2016; Silvern et al., 2019). In cities in other parts 62 of the world, particularly South and Southeast Asia, NO_x is increasing rapidly due to fast 63 economic development and limited or absent air quality policies (Vohra et al., 2021; 2022). 64 65 Vohra et al. (2022) used 14 years of satellite observations of NO₂ from the Ozone Monitoring Instrument (OMI) to infer increases of ~1-14 % a⁻¹ in surface NO₂ pollution in almost all rapidly 66 developing large cities in South and Southeast Asia. Only in Jakarta did NO₂ decline due to 67 emission controls applied to vehicles (Vohra et al., 2022). Population projections suggest that, 68

69 by 2100, one-fifth of the world's most populous cities will be in Southeast Asia (Hoornweg & 70 Pope, 2017), necessitating reliable and up-to-date NO_x emissions estimates for assessing the 71 impact of this growth on urban air quality and for informing air quality policies.

72 Bottom-up inventories provide estimates of anthropogenic NO_x emissions, but publicly available versions for South and Southeast Asia do not adequately represent contemporary 73 local conditions, as these are derived using outdated activity data, are resource-intensive to 74 produce so lag the present day, are at spatial resolutions that are coarser than many cities in the 75 region, and data needed to compile the inventories do not exist for many countries (Kurokawa 76 77 & Ohara, 2020). The two most used bottom-up inventories for these regions are the Regional Emission inventory in Asia (REAS) (Kurokawa & Ohara, 2020) and the inventory known as 78 MIX, a mosaic of REAS and other regional inventories (Li et al., 2017). REAS and MIX are 79 80 at ~25 km resolution, MIX only covers 2 years of data, and the most recent years are 2015 for REAS and 2010 for MIX. Still, REAS and inventories used to create MIX are routinely 81 incorporated in global inventories such as the Community Emissions Data System (CEDS_{GBD-} 82 83 MAPS) (McDuffie et al., 2020), and Hemispheric Transport of Air Pollution (HTAP) (Crippa et 84 al., 2023).

85 Independent and contemporary estimates of city NO_x emissions can be derived with satellite observations of tropospheric NO₂ vertical column densities (VCDs) without the need 86 for resource-intensive computer models. A method first proposed by Beirle et al. (2011) 87 involves selecting isolated cities and treating these as large point sources of NO_x. In this 88 approach, individual satellite pixels within a target domain centred on a city centre were split 89 90 into eight major wind directions to resolve the city plume in each direction. A mathematical 91 function was then fit to the plume to account for its Gaussian shape and exponential decay of NO₂. This fit, referred to as an Exponential Modified Gaussian (EMG), yields parameters that 92 are then used to estimate NO_x emissions. It also yields an effective lifetime of NO_x for the city 93 plume that is dominated by dispersion for the windy conditions sampled. As dispersion 94 95 dominates, the derived lifetime is much shorter than the chemical lifetime of NO_x that includes conversion to nitric acid (HNO₃) or organic nitrates (de Foy et al., 2014; Laughner & Cohen, 96 97 2019) and, to a lesser extent, dry deposition of NO₂ (Zhang et al., 2012). Beirle et al. (2011) 98 used OMI observations of NO₂ to derive NO_x emissions for eight global megacities. The Beirle 99 et al. (2011) approach required many (four) years of OMI data to achieve distinct plumes in each wind direction. 100

Valin et al. (2013) expanded on the approach developed by Beirle et al. (2011) by 101 demonstrating that all satellite data can instead be aligned along a single upwind-downwind 102 direction relative to the city centre. This approach reduced the number of observations needed 103 to distribute the data by wind direction and so extended application to a greater number of 104 geographically isolated cities over shorter sampling periods. Wind rotation of OMI 105 observations and the EMG fit have since been used to calculate city NO_x emissions 106 predominantly in the US (de Foy et al., 2014; Goldberg et al., 2019a; Lu et al., 2015) and for 107 select cities worldwide (Goldberg et al., 2021). Following the 2017 launch of the higher spatial 108 109 resolution TROPOspheric Monitoring Instrument (TROPOMI), the wind rotation, EMG fit, and related approaches have been extended to smaller isolated cities and shorter sampling 110 periods than was possible with OMI. Applications include cities in western Europe (Lorente et 111 al., 2019; Pope et al., 2022), China (Wu et al., 2021), the US (Goldberg et al., 2019b), and 112 worldwide (Lange et al., 2022), as well as investigating changes in NO_x emissions due to 113 COVID-19 lockdown measures in the New York Metropolitan Area (Tzortziou et al., 2022) 114 and for select cities in India, Argentina, and Spain (Lange et al., 2022). So far, the wind rotation 115 and EMG fit has only been applied to 5-13 cities in South and Southeast Asia as part of global 116 117 studies (Goldberg et al., 2021; Lange et al., 2022).

Even though there has been substantial development and use of the EMG fit, it still requires that a user define a sampling area around the city that effectively captures the wind rotated plume. The area selected varies with city size and plume length (Lu et al., 2015; Goldberg et al., 2019a; Lange et al., 2022). This approach often yields no or poor EMG fits and non-physical best-fit parameters (Laughner & Cohen, 2019), decreasing the likelihood of deriving top-down emissions. Selecting appropriate city-specific areas for the wide-ranging city sizes in South and Southeast Asia is also time consuming and not standardized.

Here we develop a near-automated and efficient EMG fitting routine for deriving annual city NO_x emissions, demonstrate the utility of this automation by applying it to TROPOMI NO_2 observations over isolated cities in South and Southeast Asia with wideranging city sizes, compare our top-down emissions to past studies and a global bottom-up inventory, and exploit the greater success of our updated sampling to identify opportunities to further develop the EMG fit approach.

131

132 **2 Materials and Methods**

133 **2.1 TROPOMI NO₂ and City Selection**

We use Level 2 TROPOMI NO₂ tropospheric column VCDs for 2019 from the 134 Sentinel-5P Products Algorithm Laboratory (S5P-PAL) portal (https://data-portal.s5p-135 pal.com/; last acquired 30 January 2022). These data have been retrieved with a consistent 136 137 algorithm (version 02.03.01) and corrected for a low bias in NO₂ over polluted scenes (Eskes et al., 2021). TROPOMI achieves daily global coverage with a swath width of 2600 km, an 138 equator crossing time of 13:30 local solar time (LST), and a nadir pixel resolution that increased 139 on 5 August 2019 from 7 km \times 3.5 km to 5.5 km \times 3.5 km. We use cloud-free, high-quality 140 data identified with a quality flag ≥ 0.75 (van Geffen et al., 2021). 141

To identify isolated cities appropriate for top-down estimate of NO_x emissions, we first 142 oversample TROPOMI NO₂ to obtain high-resolution gridded annual means $(0.05^{\circ} \times 0.05^{\circ})$; 143 ~6 km latitude \times ~5 km longitude) by weighting areas of overlap between the satellite pixels 144 and cells on a fixed latitude-longitude grid using tessellation (Sun et al., 2018). We use the 145 resultant gridded TROPOMI NO₂ shown in Figure 1 to manually select 19 cities that are 146 isolated hotspots. The 19 selected cities are Karachi, Islamabad, and Lahore in Pakistan; Kabul 147 in Afghanistan; Ahmedabad, Mumbai, Delhi, Bangalore, Chennai, and Kolkata in India; 148 Colombo in Sri Lanka; Dhaka in Bangladesh; Yangon in Myanmar; Bangkok in Thailand; 149 Kuala Lumpur in Malaysia; the sovereign city Singapore; Ho Chi Minh City in Vietnam; 150 Jakarta in Indonesia; and Manila in the Philippines. Other hotspots in Figure 1 are either not 151 cities, such as the coal-fired power plants concentrated in eastern India, or are not isolated, such 152 as Hanoi, Haiphong and Nam Dinh in northern Vietnam. 153

154





Figure 1. Annual mean TROPOMI tropospheric NO2 VCDs over South and Southeast Asia in 156 2019. Maps show South (left) and Southeast (right) Asia TROPOMI NO2 oversampled to 0.05° 157 158 \times 0.05°. The 19 selected cities, numbered from east to west, are Karachi (1), Islamabad (5), and Lahore (6) in Pakistan; Kabul (2) in Afghanistan; Ahmedabad (3), Mumbai (4), Delhi (7), 159 Bangalore (8), Chennai (10), and Kolkata (11) in India; Colombo (9) in Sri Lanka; Dhaka (12) 160 161 in Bangladesh; Yangon (13) in Myanmar; Bangkok (14) in Thailand; Kuala Lumpur (15) in Malaysia; the sovereign city Singapore (16); Ho Chi Minh City (17) in Vietnam; Jakarta (18) 162 in Indonesia; and Manila in the Philippines (19). 163

165 **2.2 Wind Rotation and EMG Fit**

Figure 2 illustrates the major steps involved in the wind rotation and EMG fit to derive 166 annual NO_x emissions for Singapore. The wind fields we use to calculate wind direction and 167 speed to retain TROPOMI NO₂ observations under windy conditions are the fifth generation 168 European (ERA5) 169 ReAnalysis 3D hourly u and v wind components (https://cds.climate.copernicus.eu/cdsapp#!/home; last acquired 18 March 2022) provided at 170 $0.25^{\circ} \times 0.25^{\circ}$ resolution. At each TROPOMI NO₂ pixel, we compute collocated mean ERA5 171 wind speeds and directions 30 min around 13:30 LST, the TROPOMI overpass time, in the 172 lowest 5 layers (\geq 900 hPa) to capture dispersion of mixed-layer near-surface NO₂ plumes. 173 Within a $4^{\circ} \times 4^{\circ}$ domain around each city centre, we isolate TROPOMI pixels with coincident 174 wind speeds $> 2 \text{ m s}^{-1}$, the threshold typically used for windy conditions (Beirle et al., 2011; 175 Pope et al., 2022). We rotate each TROPOMI NO₂ pixel by the angle of its wind direction, 176 preserving the distance of the pixel from the city centre. This aligns all pixels along the same 177 "upwind-downwind" direction that in our work is from north to south (Figure 2(a)). After wind 178 rotating all pixels in a year (as in Figure 2), we grid pixels onto a uniform $0.05^{\circ} \times 0.05^{\circ}$ grid 179 using simple point-in-box averaging (Figure 2(a)) and fill empty grid cells (grey squares in 180 Figure 2(a)) using nearest-neighbour interpolation to reduce low biases in the steps that follow. 181

182 Next, the 2D map in Figure 2(b) is converted to 1D line densities by summing all grid 183 cells in the across-wind (east-to-west) direction in 0.05° upwind-downwind (north-to-south) 184 increments. In the standard approach, a single area smaller than the $4^{\circ} \times 4^{\circ}$ domain is used,

defined by the distance upwind, downwind, and across-wind of the city centre. Instead of using 185 a single area, we define multiple areas that encompass the range of sizes typically used in past 186 studies (Goldberg et al., 2021; Lange et al., 2022; Laughner & Cohen, 2019). These, defined 187 as distances from the city centre, are 0.5°, 0.75°, and 1° upwind, 0.5°, 0.75°, 1.0°, 1.25°, 1.5°, 188 1.75°, 2.0° downwind, and 0.5°, 0.75°, and 1.0° across-wind, with the requirement that the 189 distance downwind of the city centre is \geq the distance upwind to capture the extent of the city 190 plume. This yields 54 areas and associated line densities. The sizes of the smallest and largest 191 192 areas sampled and the across-wind 0.05° increments summed to obtain line densities in the 193 smallest area sampled are shown in Figure 2(b).

194 The EMG model we use to fit to the observed 1D line densities is the Laughner & 195 Cohen (2019) formulation:

196
$$F(x|a, x_0, \mu_x, \sigma_x, B) = \frac{a}{2x_0} \exp\left(\frac{\mu_x}{x_0} + \frac{\sigma_x^2}{2x_0^2} - \frac{x}{x_0}\right) \operatorname{erfc}\left(-\frac{1}{\sqrt{2}}\left[\frac{x-\mu_x}{\sigma_x} - \frac{\sigma_x}{x_0}\right]\right) + B$$
(1),

where *x* is the distance of each line density upwind and downwind of the city centre (Figure 2(c)) and *a*, x_0 , μ_x , σ_x and *B* are best-fit parameters. Of these, *a* is total NO₂ in the plume (in moles), x_0 is the *e*-folding distance or length scale of NO₂ decay (in km), μ_x is the location of the apparent source relative to the city centre (in km) or the peak of the Gaussian fit that in Figure 2(c) is located ~20 km downwind or south of the city centre, σ_x is the Gaussian smoothing length scale (in km) that is ~2.355 × the Full Width at Half Maximum (FWHM), and *B* is background NO₂ (in moles m⁻¹).

We use initial guesses for the best-fit parameters in Equation (1) that are similar to those 204 205 from Laughner & Cohen (2019), but our fitting procedure differs. Laughner & Cohen (2019) used a non-linear interior point minimization algorithm (the *fmincon* function in MATLAB) to 206 optimize model parameters with 10 iterations per line density. Instead, we perform the fit with 207 the scipy.optimize.curve fit module from SciPy Python package version 1.7.3 and iterate on 208 209 the fit until the difference in fitting parameters between the current and previous iteration is negligible (< 0.001%) for at most 10 iterations. Fit convergence is usually achieved after 3 210 211 iterations. Only good-quality fits are retained, identified with goodness-of-fits $(R^2) > 0.8$, as in Laughner & Cohen (2019). We further screen for physically implausible best-fit parameters 212 using criteria similar to Laughner & Cohen (2019): a is positive, x_0 is at least 1.6 km 213 (approximately 1/e of the grid resolution), μ_x is within the sampling area, the emission width 214 is less than the *e*-folding distance ($\sigma_x < x_0$), background NO₂ is positive and less than the 215 maximum line density value, and the e-folding distance occurs between the plume centre and 216 the edge of the sampling area. We introduce an additional requirement to ensure that x_0 is within 217 the sampling area ($x_0 <$ length of sampling area downwind of the city centre). 218

The Singapore example in Figure 2 is an ideal city, as all 54 EMG fits are successful. Figure 2(c) shows that the observed line densities are most sensitive to the across-wind length, as this determines the amount of NO_2 summed to yield each line density. We will demonstrate in Section 3 that for many of the cities in Figure 1 a large number of EMG fits fail to meet the conditions for success, necessitating as many as 54 fits.

The successful EMG fits are used to calculate effective NO_x lifetimes (τ_{NO_x} ; reported in h) and midday NO_x emissions (E_{NO_x} ; in moles s⁻¹):

$$226 \qquad \tau_{NO_x} = \frac{x_0}{\omega} \tag{2}$$

227
$$E_{NO_x} = \gamma \times \frac{a}{\tau_{NO_x}}$$
(3),

- where ω is the sampling area mean wind speed (in m s⁻¹) and γ is the unitless molar ratio of [NO_x]/[NO₂] to convert moles NO₂ to moles NO_x. The up to 54 individual estimates of τ_{NO_x} and E_{NO_x} are averaged to obtain values for each city.
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- 232





Figure 2. Illustration of major steps in the wind rotation and EMG fit to derive annual NO_x 234 emissions for Singapore. The main steps in each panel are wind rotate and grid windy scene 235 TROPOMI NO₂ pixels to $0.05^{\circ} \times 0.05^{\circ}$ (a), fill data gaps (b), and fit the EMG function (Eq. 236 (1)) (solid lines) to observed line densities (filled circles) (c). In (b), black rectangles show the 237 extent of the largest and smallest sampling areas and dashed lines in the smallest area show the 238 0.05° increments used to calculate the line densities in (c). All 54 successful EMG fits, 18 lines 239 for each of the three across-wind lengths, are shown in (c). Values in (c) give the mean and 240 standard deviation of the city NO_x emissions (Eq. (3)), effective NO_x lifetime (Eq. (2)), and 241 sampling area ERA5 wind speed. The goodness-of-fit (\mathbb{R}^2) is ≥ 0.99 for all fits in (c). 242

We use the same $[NO_x]/[NO_2] = 1.32$ value as Beirle et al. (2011) and subsequent studies to represent rapid cycling between NO and NO₂. Liu et al. (2022) determined with synthetic experiments that city NO_x emissions are relatively unaffected by variability in $[NO_x]/[NO_2]$, but that study was for US cities. Surface measurements aid in determining

suitability of $[NO_x]/[NO_2] = 1.32$, but these are limited to cities in India and have data quality 248 issues (Vohra et al., 2021). Instead, we use the GEOS-Chem model to assess suitability of the 249 1.32 value. We simulate the model in 2019 and sample the lowest model layer around the 250 TROPOMI overpass time. We use output from a coarse and finer resolution version of GEOS-251 Chem to also test sensitivity of this ratio to model resolution, especially given many of these 252 cities are coastal (Figure 1). We use the classical configuration of the model that operates on a 253 254 single computational node, called GEOS-Chem Classic (GCClassic), and the highperformance model configuration (GCHP) that is a parallelized across multiple computational 255 nodes to enable finer resolution global simulations (Eastham et al., 2018). GCClassic is version 256 13.3.4 (https://doi.org/10.5281/zenodo.5764874) run on a fixed $2^{\circ} \times 2.5^{\circ}$ global grid and 257 GCHP is version 13.4.1 (https://doi.org/10.5281/zenodo.6564711) run on a C360 global grid 258 (~25 km \times ~31 km). GCClassic and GCHP use the same vertical grid and chemical mechanism. 259 For GCClassic, grid squares that overlap with each city are sampled, whereas for GCHP, we 260 261 use city sampling extents determined from a combination of administrative and geographic boundary shapefiles and Google Maps (Figure S1). Midday sampling is at 12:00 to 15:00 LST 262 from GCClassic and 13:00 to 14:00 LST from GCHP. At midday, NO_x is in photochemical 263 steady state, so the relative abundance of NO and NO₂ is insensitive to the extent of the 264 sampling window around midday (Potts et al., 2021). 265

We calculate uncertainties in the NO_x emissions by adding individual errors in 266 quadrature. These include best-fit parameters x_0 and a, sampling area mean wind speed ω , the 267 TROPOMI NO₂ observations, and $[NO_x]/[NO_2]$. We use the relative standard deviation from 268 all successful EMG fits to calculate city-specific errors in x_0 and a. For ω , we consider errors 269 due to the choice of spatial and temporal sampling and the threshold used for windy conditions. 270 We use the Beirle et al. (2011) estimated 10% error in temporal sampling choice and 5% error 271 due to vertical sampling choice. We conduct our own tests of the sensitivity to threshold and 272 spatial sampling choice. For $[NO_x]/[NO_2]$ we assess whether the 10% error attributed to this 273 274 variable by Beirle et al. (2011) is appropriate by quantifying the percent deviation of GCClassic and GCHP [NO_x]/[NO₂] from 1.32. Beirle et al. (2011) applied a 30% error to OMI that is also 275 appropriate for TROPOMI. Even though uncertainties in TROPOMI slant columns (NO₂ along 276 the viewing path) are much less than those from OMI (van Geffen et al., 2020), the air mass 277 factor used to convert slant columns to VCDs remains the largest contributor to errors in NO₂ 278 279 VCDs and is similar for OMI and TROPOMI (van Geffen et al., 2021).

280 **2.3 Bottom-up Anthropogenic Emissions**

We compare our top-down estimates to anthropogenic NO_x emissions from the widely 281 used bottom-up HTAP inventory version 3 (HTAP v3) (Crippa et al., 2023). HTAP v3 has 282 high enough spatial resolution $(0.1^{\circ} \times 0.1^{\circ})$ to resolve cities selected in Figure 1. The most 283 recent year is 2018, achieved by extending emissions from the regional REAS inventory ending 284 in 2015 to the year 2018 with trends from the Emissions Database for Global Atmospheric 285 Research (EDGAR) inventory. The same sampling boundaries as GCHP are used (Section 2.2; 286 Figure S1). The HTAP v3 NO_x emissions include contributions from aviation, transport (road, 287 rail, pipeline, inland waters), shipping, energy, industry, and residential sectors. 288

Cities targeted can be influenced by non-anthropogenic NO_x sources, such as open burning of biomass (Marvin et al., 2021) and natural sources such as soils (Weng et al., 2020) and lightning (Miyazaki et al., 2014). We assess suitability of comparing our top-down emissions to anthropogenic bottom-up emissions only by determining the percent contribution of anthropogenic emissions to total NO_x emissions. To do this, we simulate total NO_x emissions with the Harmonized Emissions Component (HEMCO) standalone model version 3.0.0 (<u>https://zenodo.org/records/4984639</u>; last accessed 20 March 2022) (Lin et al., 2021) and 296 sample the same spatial extent as GCHP and HTAP_v3 (Figure S1). HEMCO is run at a spatial 297 resolution of $0.25^{\circ} \times 0.3125^{\circ}$ (~ 28 km latitude × ~ 33 km longitude). HEMCO calculates open 298 biomass burning emissions using the Global Fire Emissions Database with small fires 299 (GFED4s) inventory (Randerson et al., 2017) and reads in and processes lightning and soil NO_x 300 from offline emissions at the same resolution as HEMCO (Murray et al., 2012; Weng et al., 301 2020).

Bottom-up emissions from HTAP v3 are 24-h means, whereas top-down estimates 302 derived using TROPOMI are representative of midday emissions. Goldberg et al. (2021) 303 304 multiplied satellite-derived midday NO_x emissions by 0.77 to convert midday top-down NO_x emissions to 24-h means for comparison to bottom-up inventories. This value was inferred 305 from bottom-up emissions estimates for the Netherlands, so may not be suitable for the selected 306 307 cities in South and Southeast Asia. The hourly scaling factors used by HEMCO for the chosen cities range from 0.70 to 1.16. These are for the year 2000 and are extrapolations of values for 308 conditions in Europe, so may not be suitable for the year and cities targeted in this study. Given 309 310 this, we do not scale top-down emissions and instead discuss whether differences in averaging times contribute to discrepancies between top-down and bottom-up emissions estimates. 311

312

313 **3 Results and Discussion**

314 **3.1 Wind Rotation and EMG Fit Metrics**

Isolating windy condition (> 2 m s^{-1}) satellite pixels removes 8-34% of all 2019 quality-315 and cloud-screened TROPOMI NO2 pixels for most cities in Figure 1. Cities with greater data 316 317 loss are Lahore (43% data loss), Kabul (58%) and Islamabad (63%). No spatial data gap filling (Section 2.2, Figure 2) is needed within the areas sampled, due to the high sampling frequency 318 of TROPOMI. If only a single domain size is selected, annual EMG fits meet all criteria for 319 success for 7 to 12 of the 19 cities in Figure 1, depending on the sampling area chosen. Using 320 321 our extended method, we successfully derive annual NO_x emissions for all 19 cities, due to the enhanced probability of obtaining at least one successful EMG fit. 322

Figure shows the number of successful EMG fits (orange bars) range from 3 (Kabul) to 323 all 54 (Singapore). Singapore, Dhaka, Jakarta, Karachi, Manila, and Mumbai are least impacted 324 325 by the choice of sampling area. The 6 cities in Figure 3 with < 20 fits are most likely to fail if only a single sampling area is used. For all retained EMG fits, differences between observed 326 and fitted NO₂ line densities, the fit residuals, are negligible. The most common causes for a 327 failed EMG fit rank as: background NO₂ (B in Equation (1)) > maximum NO₂ line density (36%) 328 of all fits conducted), $R^2 \le 0.8$ (24%), emission width > *e*-folding distance (19%), total plume 329 NO₂ (a in Equation (1)) < 0 (13%), and e-folding distance > the downwind length of the 330 sampling area (12%). Multiple causes can co-occur in a single fit, so cumulative percentages 331 332 exceed 100%.

We also test sensitivity of top-down NO_x emissions to the choice of wind speed 333 threshold and horizontal sampling extent to attribute an error to these. For this, we apply a 334 stricter wind speed threshold of 3 m s⁻¹ and test the difference in NO_x emissions if instead of 335 filtering for windy conditions using pixel-mean wind fields, we calculate a sampling-area mean 336 wind speed to filter for windy conditions as in Goldberg et al. (2019a). We apply these 337 conditions to a mid-sized sampling area of 0.75° upwind, 1.5° downwind, and $\pm 0.75^{\circ}$ across-338 wind. Variability in NO_x emissions for cities with successful EMG fits for all 4 wind sampling 339 340 conditions is at most 10% (Figure S2). Given these results, we attribute a 10% error to the choice of horizontal sampling and to the wind speed threshold. 341

GCClassic (coarse resolution) annual mean [NO_x]/[NO₂] for the target cities ranges 342 343 from 1.25 (Dhaka) to 1.41 (Kabul). The range in ratios from GCHP (finer resolution) is wider at 1.24 (Ahmedabad) to 1.64 (Kolkata). The difference in ratios between the coarse and fine 344 resolution models is typically $\pm 10\%$, except for a few cities with ratios from the fine resolution 345 model that exceed the coarse resolution model by 14% for Singapore, 16% for Lahore, 23% 346 for Dhaka, and 23% for Kolkata. This is because the fine resolution model better resolves the 347 city plume that includes a greater proportion of NO_x as NO from fresh emission sources. As 348 the difference between the model city ratios and the 1.32 value is $\pm 10\%$ for most cities, we use 349 the same 10% error for $[NO_x]/[NO_2]$ as Beirle et al. (2011). 350

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Figure 3. Successful EMG fits and top-down NO_x emissions for the cities targeted in this study. Bars are emissions (green) and the corresponding number of successful fits (orange). Black error lines are NO_x emission standard deviations for all successful fits. The orange dashed line at 54 indicates the maximum possible EMG fits. Emissions multiplied by ~1.45 yields emissions in Gg NO₂ a⁻¹.

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359 **3.2 Top-Down NO_x Emissions**

Green bars in Figure 3 show the mean annual top-down NO_x emissions for all cities (values are in Table S1). These range from $\sim 16 \text{ mol s}^{-1}$ for Bangalore and Yangon to $\sim 125 \text{ mol}$ s⁻¹ for Dhaka. The range in the total mass of NO_x emitted for these cities, assuming the midday

emission rate is reasonably representative of the 24-h emission rate, is 23-181 Gg NO_x as NO₂. 363 Emissions for most cities are $< 50 \text{ mol s}^{-1}$ ($<73 \text{ Gg NO}_x$ as NO₂ a⁻¹). Cities with emissions 364 between 50-100 mol s⁻¹ (73-145 Gg NO_x as NO₂ a⁻¹) include Karachi, Delhi, and Jakarta and > 365 100 mol s⁻¹ (> 145 Gg NO_x as NO₂ a⁻¹) include Bangkok, Singapore, and Dhaka. Emission 366 rates for Bangkok, Dhaka and Singapore are comparable to the range of top-down emissions 367 estimated for large, polluted cities in China using the EMG approach (Wu et al., 2021). The 368 369 effective lifetimes for the cities in Figure 1 (shown in Figure S3) range from 1.2 h for Colombo to 6.3 h for Kuala Lumpur. Variability in effective lifetimes depends most strongly on the 370 downwind extent of the plume. The Pearson's correlation coefficient, R, between city mean 371 372 effective lifetimes and x_0 values is 0.90.

373 For the target cities, the relative standard deviations of annual NO_x emissions (black 374 error lines in Figure 3) range from just 1% for Bangalore to 27% for Kuala Lumpur. This is far less than the equivalent Gaussian fit uncertainty of 10-50% estimated by Beirle et al. (2011) 375 for a single sampling area. The relatively large variability in Kuala Lumpur NO_x emissions is 376 because the smaller EMG sampling areas do not fully encompass the elongated wind rotated 377 city NO₂ plume, causing a low bias in NO_x emissions for the smaller areas sampled. The effect 378 of this is dampened by the almost 30 successful fits used to obtain mean NO_x emissions for this 379 city. The relative standard deviations of the NO_x lifetimes (Figure S3) range from 3% for 380 Bangalore to 37% for Chennai. The relative standard deviations of other parameters are $\sim 6\%$ 381 for wind speeds (Figure S4), 4% (Bangalore) to 38% (Chennai) for x_0 , and 4% (Kabul and 382 Bangalore) to 37% (Bangkok) for a. 383

The overall uncertainty in annual NO_x emissions we obtain by adding all error contributions in quadrature ranges from 32% for Bangalore and Yangon to 55% for Bangkok. Values for all cities are in Table S1. The TROPOMI NO₂ VCDs make the largest contribution to the overall uncertainty. The higher-end of our uncertainty estimates is similar to the typical ~50% uncertainty reported in past studies (Beirle et al., 2011; Verstraeten et al., 2018; Goldberg et al., 2021). We use our overall uncertainties in the comparison of our top-down emissions to values from the literature and from HTAP in the sections that follows.

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392 **3.3 Comparison to Top-Down Estimates from Past Studies**

To assess our approach, we compare in Figure 4 our annual NO_x emissions to values 393 from past studies that used similar sampling time periods and a single sampling area. These 394 include multivear (2017-2019) mean emissions from Goldberg et al. (2021) obtained using the 395 OMI sensor and emissions from Lange et al. (2022) obtained with select days of TROPOMI 396 data from 2018 to 2020. Goldberg et al. (2021) estimated emissions for 10 of the 19 cities in 397 our study. These we read from their Figure S10 for Karachi, Figure S11 for 4 cities in India, 398 and Figure S13 for 5 cities in Southeast Asia and divide by the 0.77 midday to 24-h scaling 399 factor used in that study. Emissions are reported by Lange et al. (2022) for 5 of the 19 cities in 400 our study. Based on the regression statistics in Figure 4, our emissions are typically ~26% more 401 than estimates from these past studies. Exceptions are Mumbai, Ahmedabad, and Chennai that 402 in our study are 16-29% less than Goldberg et al. (2021). Lange et al. (2022) used an earlier 403 version of the TROPOMI data product that has a known low bias in NO₂ VCDs over very 404 polluted scenes (van Geffen et al., 2022). Differences in TROPOMI data products are the likely 405 cause for our higher Delhi (by 27%) and Singapore (by 18%) emissions. Relatively small error 406 estimates from Lange et al. (2022) are because they only propagate error contributions from 407 the wind speed data and the EMG fit. 408

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Figure 4. Comparison of our and past top-down NO_x emissions. Symbols compare our emissions to those from Goldberg et al. (2021) (red) and Lange et al. (2022) (blue). Error bars are overall uncertainties for our study (Section 2.2, Table S1), the same 53% uncertainty applied to all cities by Goldberg et al. (2021) and the city-specific uncertainties for Lange et al. (2022). Lines are the Theil regression fit (solid black) and 1:1 relationship (dashed grey). Inset text gives the regression statistics and Pearson's correlation coefficient (R). Arrows and inset text for Dhaka give the error values that extend beyond the plotting range.

Discrepancies between Goldberg et al. (2021) and our emissions are not as 419 straightforward to diagnose, as Goldberg et al. (2021) use NO₂ VCDs from a different sensor 420 (OMI) and apply a systematic 37% increase to NO_x emissions to correct for a low bias in OMI 421 attributed to the coarse resolution a priori used in the NO₂ VCDs retrieval. Sampling area 422 choice may also be a factor. For example, the smallest of our 54 areas yields NO_x emissions of 423 102 mol s⁻¹ for Singapore that is 10 mol s⁻¹ less than the mean of all EMG fits. Goldberg et al. 424 425 (2021) used year-round OMI data for all cities except Delhi and Karachi. As these cities are north of 25°N, only May-September observations were used by Goldberg et al. (2021). We find 426 that Delhi and Karachi mean May-September TROPOMI NO2 VCDs in 2019 averaged within 427 the $4^{\circ} \times 4^{\circ}$ domain selected for each city (Figure 2(a)-(b)) are 11-12% less than those in 428 429 October-April, due to the shorter photochemical lifetime of NO_x in the warmer months. Open biomass burning emissions also influence seasonality in the TROPOMI NO₂ VCDs, but the 430 431 EMG fit accounts for this by distinguishing background NO_2 (*B* in Equation (1)) from NO_2 in 432 the city plume (*a* in Equation (1)).

We find that if we apply the EMG fit to individual months for Delhi and Karachi, all 54 EMG fits fail for Delhi in July-August and yield spurious results in September due to large data loss resulting from persistent clouds during the monsoon season. All 12 months are retained for Karachi, Singapore and Manila. November-April mean values of *a* are 21% more than in May-October for Karachi, 9% more for Singapore, and 39% more for Manila. This suggests that using NO₂ VCDs for a portion of the year may yield systematic biases in 439 emissions that may not reflect seasonality in the underlying activities affecting the emissions. Larger wintertime than summertime emissions have also been reported in the global study of 440 Lange et al. (2022). They quantified summer-to-winter emission ratios of ~0.5 for Colombo 441 and Delhi. The top-down emissions calculation (Equation (3)) does not fully account for 442 seasonality in photochemistry. The derived effective NO_x lifetimes used to calculate NO_x 443 emissions (Equation (2)) are mostly influenced by dispersion. As a result, the effective lifetimes 444 445 are much shorter than the expected chemical lifetimes of NO_x (de Foy et al., 2014). In the synthetic experiment scenarios tested by de Foy et al. (2014), the EMG fit applied to wind 446 rotated data yielded an effective lifetime of 4 h for a 12-h chemical lifetime scenario. According 447 to Shah et al. (2020), the chemical lifetime of NO_x for central-eastern China centred at ~35°N. 448 449 the northerly portion of our domain, ranges from ~6 h in summer to ~24 h in winter. None of the monthly effective lifetimes for our target cities reproduces this seasonality and the longest 450 lifetime is 13.3±3.7 h for Yangon in November. The implication is that the size of absolute 451 452 emissions derived with sub-annual satellite data may be biased, but should have negligible effect if used to quantify relative trends, as in Goldberg et al. (2021) and Laughner & Cohen 453 (2019), for example. 454

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456 **3.4 Comparison to Bottom-up Emissions**

Figure 5 compares annual top-down and bottom-up NO_x emissions. According to our 457 HEMCO simulations, anthropogenic sources account for most (>87%) annual NO_x emissions. 458 The relative differences between our top-down estimates and the bottom-up inventory are 459 within 50% for Mumbai (1%), Bangkok (2%), Chennai (9%), Ahmedabad (11%), Kolkata 460 (21%), Singapore (21%), Bangalore (32%), Manila (35%), and Kuala Lumpur (46%). A 50-461 100% difference occurs for Ho Chi Minh City (53%), Jakarta (54%), Delhi (64%), and 462 463 Colombo (91%). Even greater relative differences occur for Karachi (2.1 times), Islamabad (2.1 times), Lahore (2.4 times), Yangon (3.3 times), Dhaka (6.9 times), and Kabul (11-fold). 464 The largest absolute discrepancies are for Dhaka and Jakarta. Bottom-up emissions are 107 465 mol s⁻¹ less than the top-down values for Dhaka and 78 mol s⁻¹ more for Jakarta. On a mass 466 basis, this is equivalent to a 155 Gg NO_x as NO₂ underestimate for Dhaka and a 113 Gg NO_x 467 as NO₂ overestimate for Jakarta. 468

The different years used (2018 for HTAP, 2019 for TROPOMI) should at most account 469 for a 14% difference in emissions, based on the size of annual trends inferred by Vohra et al. 470 471 (2022) using long-term observations of OMI NO₂ VCDs over large and fast-growing cities in South and Southeast Asia. Vohra et al. (2022) identified that emission inventories do not 472 capture the steep decline in NO_x emissions in Jakarta attributed to national policies targeting 473 vehicles. In addition to misrepresenting annual changes in underlying activities, the emission 474 factors are mostly informed by studies in China and Japan (Kurokawa & Ohara, 2020). The 475 bottom-up and top-down emissions differences for many cities also exceed the $\pm 30\%$ 476 difference that results from the choice of bottom-up emissions grid sampling and the $\pm 30\%$ 477 478 difference from the timing of the top-down (midday) and bottom-up (24-h) estimates inferred by Goldberg et al. (2021). 479

480 Apparent in Figure 5 is a latitudinal pattern in the discrepancies. Top-down emissions 481 are greater than bottom-up emissions for cities to the north and vice versa for cities to the south, 482 so that in general top-down emissions exceed bottom-up emissions in South Asia and vice 483 versa in Southeast Asia. NO_x chemical loss varies with latitude, due to variability in the amount 484 of sunlight available to form hydroxyl and peroxy radicals required to form HNO₃ and organic 485 nitrates, the main daytime chemical loss pathway for NO_x . This latitudinal pattern is likely 486 because the EMG fit also does not fully account for spatial variability in NO_x photochemistry, 487 imparting a bias in the top-down emissions. The size of this bias will depend on the relative 488 contribution of NO_x chemical loss to total loss in the wind rotated plume.

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Figure 5. Comparison of annual top-down and bottom-up NO_x emissions for target cities. Data are coloured by city centre latitude and split into top-down NO_x emissions < 40 mol s⁻¹ (a) and \geq 40 mol s⁻¹ (b). Error bars are the overall uncertainty in top-down emissions estimates. Grey lines indicate 1:1 agreement (solid) and ±50% difference (dashed). The bottom-up emissions sampling extent of each city is in Figure S1. Data used to generate the figure are in Table S1.

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498 4 Conclusions

City nitrogen oxides (NO_x) emissions can be derived with a now well-established 499 approach using satellite observations of nitrogen dioxide (NO₂), wind rotation and a Gaussian 500 501 fit to the city plume. Issues with this approach are that the choice of sampling area around the city centre is not standardized and so is prone to subjective area selection and the Gaussian fit 502 often fails or yields non-physical best-fit parameters. Here we address these issues by applying 503 54 sampling areas to isolated cities. We test our method with TROPOspheric Monitoring 504 Instrument (TROPOMI) NO₂ observations for 2019 over 19 large, isolated cities in South and 505 Southeast Asia that lack contemporary, publicly available bottom-up emissions estimates. 506

507 Annual NO_x emissions, obtained for all 19 cities, are < 73 Gg NO_x as NO₂ a⁻¹ for most cities, between 73-145 Gg NO_x as NO₂ a^{-1} for Karachi, Delhi, and Jakarta and > 145 Gg NO_x 508 as NO₂ a⁻¹ for Bangkok, Dhaka, and Singapore. The overall uncertainty in the annual emissions 509 is 30-60%. Our emissions estimates are in general \sim 27% more than past studies that use a single 510 sampling area, due to differences in satellite data products and months targeted. The latter we 511 suggest may lead to biases, as the top-down emissions estimate does not properly account for 512 seasonality in photochemical loss of NO_x. Relative differences between our top-down estimates 513 and a widely used bottom-up inventory are < 50% for 9 of the 19 cities, within 50-100% for 514 Ho Chi Minh City, Jakarta, Delhi, and Colombo, and much greater for Karachi (2.1 times), 515 Islamabad (2.1 times), Lahore (2.4 times), Yangon (3.3 times), Dhaka (6.9 times), and Kabul 516 (11-fold). There is a latitudinal dependence of the size of these discrepancies that we suggest 517

is because the top-down approach also does not properly account for spatial variability in the chemical lifetime of NO_x .

520 The increased success of deriving NO_x emissions with our updated approach enables 521 us to identify that further development is needed to account for time and space variability in 522 the chemical lifetime of NO_x to fully exploit the top-down approach to interrogate seasonality 523 in emissions, to validate bottom-up emissions, to exploit hourly observations from 524 geostationary instruments, and to inform air quality regulation.

525 Data and Software Availability

526 The TROPOMI tropospheric columns for 2019 are publicly available from the S5P-PAL Data 527 Portal (https://data-portal.s5p-pal.com/). GEOS-Chem source codes are preserved on Zenodo 528 by The International GEOS-Chem User Community (2021) for GCClassic version 13.3.4 and 520 htt The International GEOS Chem User Community (2022) for GCClassic version 12.4.1

529 by The International GEOS-Chem User Community (2022) for GCHP version 13.4.1.

530 Author Contributions

GL developed the methodology, GL and EAM processed, analysed and interpreted the data.
GL and EAM prepared the manuscript. KV assisted in data collection and analysis. RPH and
DZ conducted the GEOS-Chem simulations (RPH: GCClassic; DZ: GCHP). RVM contributed

- to the methodology. SG contributed to interpretation of the results. All co-authors provided
- 535 editorial input.

536 **Conflicts of interest**

537 The authors declare there are no conflicts of interest.

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	Readure PUBLICATIONS
1	
2	JGR Atmospheres
3	Supporting Information for
4 5	Near-Automated Estimate of City Nitrogen Oxides Emissions Applied to South and Southeast Asia
6 7	Gongda Lu ¹ , Eloise A. Marais ¹ , Karn Vohra ¹ , Rebekah P. Horner ¹ , Dandan Zhang ² , Randall V. Martin ² , Sarath Guttikunda ^{3,4}
8 9 10 11 12 13 14	¹ Department of Geography, University College London, Gower Street, London, UK. ² Department of Energy, Environmental, and Chemical Engineering, Washington University in St. Louis, St. Louis, MO, USA. ³ Transportation Research and Injury Prevention (TRIP) Center, Indian Institute of Technology, New Delhi, 110016, India. ⁴ Urban Emissions, New Delhi, 110016, India.
15 16	Corresponding author: Eloise A. Marais (<u>e.marais@ucl.ac.uk</u>)
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19 20	Figures S1 to S4

City (Country) ^a	Top-down NO _x emissions [mol s ⁻¹] ^b	NO _x lifetimes [h] ^{b,c}	Wind speeds [m s ⁻¹] ^d	Bottom-up NO _x emissions [mol s ⁻¹] ^e
1. Karachi (Pakistan)	52.9 ± 18.7	3.1 ± 0.6	5.6 ± 0.1	24.7
2. Kabul (Afghanistan)	18.8 ± 6.1	1.5 ± 0.3	2.8 ± 0	1.8
3. Ahmedabad (India)	21.7 ± 9.2	3.8 ± 0.8	4.3 ± 0.1	19.5
4. Mumbai (India)	45.6 ± 16.5	2.9 ± 0.6	4.2 ± 0.1	45.3
5. Islamabad (Pakistan)	21.7 ± 8.4	2.2 ± 0.5	3.0 ± 0	10.6
6. Lahore (Pakistan)	33.4 ± 11.9	3.0 ± 0.5	3.4 ± 0	14.0
7. Delhi (India)	89.0 ± 31.9	2.5 ± 0.5	4.3 ± 0	54.4
8. Bangalore (India)	15.5 ± 5.0	3.5 ± 0.5	3.7 ± 0	22.9
9. Colombo (Sri Lanka)	20.7 ± 7.6	1.2 ± 0.3	5.7 ± 0.2	10.9
10. Chennai (India)	25.3 ± 13.6	4.9 ± 2.0	5.2 ± 0	27.9
11. Kolkata (India)	42.5 ± 15.8	2.7 ± 0.5	4.1 ± 0	35.1
12. Dhaka (Bangladesh)	124.8 ± 41.1	2.6 ± 0.4	3.8 ± 0	18.2
13. Yangon (Myanmar)	16.1 ± 5.2	2.1 ± 0.3	3.6 ± 0	4.9
14. Bangkok (Thailand)	102.3 ± 55.9	2.5 ± 0.7	4.4 ± 0	104.4
15. Kuala Lumpur (Malaysia)	41.4 ± 22.0	6.3 ± 1.9	3.7 ± 0.1	76.3
16. Singapore	112.1 ± 37.7	2.4 ± 0.4	5.1 ± 0.1	141.1
17. Ho Chi Minh City (Vietnam)	25.2 ± 11.3	4.9 ± 1.4	4.9 ± 0.3	16.4
18. Jakarta (Indonesia)	65.8 ± 32.5	3.3 ± 1.1	4.2 ± 0.2	144.2
19. Manila (Philippines)	40.5 ± 17.4	3.3 ± 0.7	6.1 ± 0.2	62.5

21 Table S1: Annual top-down NOx emissions and effective lifetimes, sampling area mean wind 22 speeds, and bottom-up NO_x emissions for cities in South and Southeast Asia

^a Numbered according to labels in Figure 1.
 ^b Errors in emissions and lifetimes calculated by adding individual errors in quadrature (see Section 2.2 for details).
 ^c Effective lifetime, as loss is dominated by dispersion.
 ^d Calculated using ERA5 reanalysis midday hourly wind fields (see Section 2.2 for details).
 ^e HTAP version 3 anthropogenic emissions inventory 24-h emission rates (see Section 2.3 for details).





Figure S1. Sampling areas of bottom-up emissions for target cities in South and Southeast Asia. Hatching identifies the sampling extent for each city. City and sampling boundaries are determined using the Database of Global Administrative Areas (GADM) (https://gadm.org/; last accessed 17 March 2023) and Google Maps for all cities and the Humanitarian Data Exchange (https://data.humdata.org/; last accessed 17 March 2023) to map Laguna de Bay bordering Manila. Background maps are from © Google Maps, 2023.



Figure S2. Sensitivity of annual top-down NO_x emissions to wind speed selection. Wind speeds

tested are individual pixels with speeds > 2 m s⁻¹ (red) and > 3 m s⁻¹ (yellow), and sampling area (1.5° downwind, 0.75° upwind, $\pm 0.75^{\circ}$ across-wind) mean speeds > 2 m s⁻¹ (green) and > 3 m s⁻¹

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(blue). Only cities with successful EMG fits for all wind speed selections are shown.



41 Figure S3. Annual effective NOx lifetimes from all successful EMG fits for target cities in South 42 43 and Southeast Asia. Red bars are the means of NO2 lifetimes and black error lines are the standard

deviations from all successful EMG fits.



Figure S4. Annual mean wind speeds for target cities in South and Southeast Asia. Yellow bars

45 46 47 are the sampling area mean wind speeds and black error lines are the standard deviations from all

successful annual EMG fits.

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