Sheltering of sea ice ridges in the ice-ocean drag force: implications from laboratory experiments

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Abstract

The increasing movement and deformation of Arctic sea ice cover results in pronounced drag sheltering effects behind sea ice pressure ridges. This needs to be accounted for in the parameterization of the form drag of ridges, thereby posing a challenge to evaluate the ice–ocean dynamic feedback. Laboratory experiments were conducted in a water tank to explore the sheltering effect between adjacent ridges of various geometries. The form drag forces on the keel models were measured, and the particle image velocimetry (PIV) system was employed to capture the flow fields surrounding the models to explain the variations in the drag force. The key sheltering parameters were the ratio between keel spacing and keel depth L/H, flow velocity u, and keel slope angle α . The results showed that the drag force F1 on the upstream keel was close to the value of the single keel case, while the drag force F2 on the downstream keel was lower, for L/H [?] 10 even opposite to the flow direction. Having changed from negative to positive, the sheltering coefficient $\Gamma = F1/F2$ increased with increasing L/H. Γ decreased remarkably with steepening α and was independent of u. To fully incorporate the effects of the L/H and α , we propose a new sheltering function fitted with the experimental results: $\Gamma = [1-1.56\exp(sL/H)]^*1.20\alpha$ -0.08, $s=0.001\alpha$ -0.15. This function is compared with the previous sheltering functions and the actual ice conditions in the Arctic Ocean, pointing the way to obtain the final sheltering functions applicable to sea ice dynamics models.

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1	Sheltering of sea ice ridges in the ice–ocean drag force: implications
2	from laboratory experiments
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9	Key points
10	• The drag force on the downstream sea ice ridge is reduced
11	significantly and even becomes negative due to the sheltering of the
12	upstream ridge.
13	• The sheltering effect weakens with increasing dimensionless
14	spacing, strengthens with increasing keel angle but isn't related to flow
15	speed.
16	• A new sheltering function has been proposed, providing a
17	reference for the parameterisation of ice-ocean drag coefficient.

18 Abstract

19	The increasing movement and deformation of Arctic sea ice cover results in
20	pronounced drag sheltering effects behind sea ice pressure ridges. This needs to be
21	accounted for in the parameterisation of the form drag of ridges, thereby posing a
22	challenge to evaluate the ice-ocean dynamic feedback. Laboratory experiments were
23	conducted in a water tank to explore the sheltering effect between adjacent ridges of
24	various geometries. The form drag forces on the keel models were measured, and the
25	particle image velocimetry (PIV) system was employed to capture the flow fields
26	surrounding the models to explain the variations in the drag force. The key sheltering
27	parameters were the ratio between keel spacing and keel depth L/H , flow velocity u ,
28	and keel slope angle α . The results showed that the drag force F_1 on the upstream keel
29	was close to the value of the single keel case, while the drag force F_2 on the
30	downstream keel was lower, for $L/H \le 10$ even opposite to the flow direction. Having
31	changed from negative to positive, the sheltering coefficient $\Gamma = F_1/F_2$ increased with
32	increasing L/H. Γ decreased remarkably with steepening α and was independent of u .
33	To fully incorporate the effects of the L/H and α , we propose a new sheltering
34	function fitted with the experimental results: $\Gamma = [1-1.56 \exp(sL/H)] \times 1.20 \alpha^{-0.08}$,
35	$s = 0.001\alpha$ -0.15. This function is compared with the previous sheltering functions and

the actual ice conditions in the Arctic Ocean, pointing the way to obtain the finalsheltering functions applicable to sea ice dynamics models.

38

Plain Language Summary

39 The wake generated by ridge keels in drifting sea ice causes critical sheltering on 40 the downstream ice-ocean interaction, as the keels occupy a significant fraction of the 41 under-ice mixed layer. The flow pattern in the under-ice boundary layer is difficult to 42 measure in situ. Therefore, the sheltering effect was investigated in laboratory 43 experiments, including drag force measurements and flow field mapping with 44 particles. The results revealed that the sheltering effect is sensitive to the keel 45 dimensionless spacing and slope angle but independent of the incoming flow speed. 46 Based on finite depth experiments, a new sheltering function was fitted for the 47 sheltering effect.

48 **1 Introduction**

The decline in Arctic sea ice thickness during recent decades has intensified sea ice movement (Kwok et al., 2018; Itkin et al., 2017; Stroeve et al., 2014), leading to the emergence of more deformed ice due to the low strength of thin ice (Itkin et al., 2017; Rampal et al., 2009; Albedyll et al., 2021; Herman, 2011; Leppäranta, 2011). It

53	is anticipated to have a significant impact on sea ice feedback processes, such as ice
54	drift patterns and momentum transfer (Castellani et al., 2014; Armitage et al., 2020;
55	Boutin et al., 2019; Thielen K., 2021; Tschudi et al., 2020). The drag coefficients are
56	critical parameters in these processes, and their modelling has evolved from
57	semiempirical constants to parameterisation schemes associated with sea ice
58	morphology (McPhee, 2008; Leppäranta and Omstedt, 1990; Lu et al., 2011; Brenner
59	et al., 2021; Tsamados et al., 2014). The wind and ocean drag forces are partitioned
60	for the skin friction drag, floe edge form drag, and ridge form drag (Arya, 1975;
61	Garbrech et al., 2002), where the ridge term becomes more important with increasing
62	proportion of deformed sea ice (Tsamados et al., 2014; Brenner et al., 2021; Castellani
63	et al., 2018).
C A	

A remarkable sheltering effect exists between adjacent ridges due to the wake (Garbrech et al., 1999) that reduces the form drag of downstream ridges. To account for this effect, a sheltering function has been used in drag parameterisation (Lu et al., 2011; Lüpkes et al., 2012; Tasmados et al., 2014). Two forms of sheltering function, viz. exponential and power laws, have been applied: $\Gamma_1 = 1 - \exp(-0.18L/H)$ and $\Gamma_2 = [1-(L/H)^{-0.5}]^2$, where *L* is the keel spacing and *H* is the keel depth.

70 The exponential form was originally derived from experiments on terrestrial

71	forests of varying density and spacing for the decay of downwind speed (Nägeli,
72	1964). Hanssen-Bauer and Gjessing (1988) applied this function for the first time to
73	calculate drag attenuation in a field of ice floes. It was also utilised in the optimised
74	drag parametrisation considering probability functions of a surface parameter
75	developed by Mai et al. (1996). Lüpkes et al. (2012) adjusted the attenuation
76	parameter from -0.18 to -0.5 based on tunnel observations by Lopez et al. (2005), and
77	they further transformed this into a function of sea ice fraction by assuming a
78	relationship between the spacing and sea ice fraction. However, this simplified
79	scheme works less well at either high or low ice concentration.
80	Regarding the power law form, Tennekes and Lumley (1972) deduced that the
81	maximum velocity loss behind a cylinder in a turbulent wake is a power function of
82	the distance in a two-dimensional boundary-free plane shear flow. Based on this,
83	Steele et al. (1989) provided a power-law decay function for sheltering between ice
84	floes in a coupled sea ice-ocean model and analysed this sheltering with respect to
85	floes size. Lu et al. (2011) then employed this function in the parameterisation of sea
86	ice-ocean drag coefficient.
07	

87 Tsamados et al. (2014) used the exponential sheltering function as they 88 considered it to be simultaneously valid for the air-ice and ice-ocean interfaces rather

89	than the power law. Brenner et al. (2021) compared the ice-ocean drag coefficient
90	calculated from the force-balance using observational data with the parameterisation
91	of Lu et al. (2011) and Tsamados et al. (2014). They showed that the difference of
92	sheltering functions is an important factor affecting the prediction accuracy of the
93	drag parameterisation, opening new ideas for improving the parameterisation schemes.
94	Both exponential and power-law functions have been recognised and cited to date,
95	providing valuable references for subsequent research. However, significant
96	differences exist between land and oceanic boundary layers regarding highly
97	permeable forest and variable sea ice structures (Hanssen-Bauer & Gjessing, 1988).
98	The power function was derived for a cylinder in boundary-free plane flow, while sea
99	ice drifts on the ocean surface, and its accuracy has not been verified by laboratory
100	experiments or numerical simulations. Neither method considers the highly variable
101	structural geometry of ridges. Thus, their applicability for sea ice is worth discussing.
102	Although the wake for structures in tandem configuration has been widely studied in
103	hydrodynamics (Kurtulmuş, 2022; Liu & Chen, 2002; Du et al., 2019; Luo et al., 1999;
104	Rui et al., 2012; Pinarbasi et al., 2015; Kumar & Sen, 2021; Kim et al., 2008;
105	Sohankar, 2012; Lin et al., 2002), no adequate wake functions exist to our knowledge
106	since the studies have mainly focused on small spacings (Du et al., 2019).

107	It is quite laborious to conduct in situ measurements for the sheltering around
108	ridges, especially due to the harsh conditions beneath sea ice, as opposed to the
109	situation with the atmospheric boundary layer above ice surface (Prinsenberg &
110	Peterson, 2002; Castellani et al., 2014; Garbrech et al., 1999). The depth of ridge
111	keels is 4-5 times the sail height (Timco & Burden, 1997; Kharitonov & Borodkin,
112	2020; Strub-Klein & Sudom, 2012), and the keel depths are a fraction of 1/5 to 1/2 of
113	the mixed layer depth while the corresponding fraction is usually less than 1/100 in
114	the atmospheric boundary layer (McPhee, 2002). Therefore, the sheltering between
115	keels deserves more attention. A combination of hydrodynamic laboratory
116	experiments and numerical modelling seems to be the best way to address this issue.
117	The former provides the physical truth of flow characteristics at ridges, and the latter
118	extends the experimental results to real ocean conditions (Mortikov, 2016; Pite et al.,
119	1995; Zu et al., 2021).
120	The goal of the present research was to investigate the sheltering effect between
121	sea ice ridges through a series of water tank experiments, thus providing references
122	for further numerical modelling and keel drag parameterisation schemes. This study is

123 a follow-up to the water tank experiments of Zu et al. (2022) and Wang et al. (2022).

124 Zu et al. (2022) examined the parameterization of the drag coefficient for a single

125	ridge, while Wang et al. (2022) addressed the feasibility of an experimental system to
126	study the sheltering between ridges using a fixed keel shape. Here this experimente
127	has been improved from the preliminary tests and covered more comprehensive
128	factors. This paper is organised as follows: the theory of parameterization on ice-
129	ocean drag coefficient is briefly summarized in Section 2 as a background. And in
130	Section 3, the experimental setup is introduced, including the force and flow field
131	measurement system. The dependence of the drag force and sheltering effect on the
132	flow velocity, dimensionless spacing, and keel slope angle are investigated in Section
133	4. Section 5 discusses the sheltering in finite water depth, the impact of sheltering
134	function on ice-ocean drag coefficient and sea ice motion, and limitations in the
135	laboratory experiment. Finally, conclusions are drawn in Section 6.

136 **2** Theory of ice-ocean drag coefficient parameterization

137 Ice-ocean shear stress τ_{io} is commonly described by the quadratic drag law:

138
$$\tau_{io} = \rho C_{io}(\boldsymbol{u}_i - \boldsymbol{u}_o) ||\boldsymbol{u}_i - \boldsymbol{u}_o|$$
(1)

139 where ρ is seawater density, C_{io} is the ice-ocean drag coefficient, \boldsymbol{u}_i is the drift velocity 140 of ice and \boldsymbol{u}_o is ocean current velocity at the reference depth. Based on the drag 141 partition theory, the total ice-ocean drag force is given as the sum of the form drags of

142the floe edge and ridge keel, and the skin friction drag (Arya, 1975). Accordingly, the143total ice-ocean drag coefficient C_{io} in sea ice geometry-based parameterisation144schemes can be calculated as follows (Lu et al.,2011; Tsamados et al., 2014):145 $C_{io} = C_r + C_f + C_s$ (2)

146
$$C_{\rm r} = \frac{1}{2} c_{\rm r} A \frac{H}{L} \Gamma\left(\frac{L}{H}\right) P_{\rm r0}$$
(2a)

147
$$C_{\rm f} = \frac{1}{2} c_{\rm f} A \frac{d}{L_{\rm f}} \Gamma\left(\frac{d}{L_{\rm f}}\right) P_{\rm f0} \tag{2b}$$

148
$$C_{\rm s} = c_{\rm s} A (1-m \frac{H}{L}), \frac{H}{L} \le \frac{1}{m}$$
(2c)

149 where C_r , C_f and C_s represent ridge keel drag coefficient, floe edge drag coefficient, 150and skin drag coefficient, respectively, c_r , c_f and c_s are the corresponding local drag 151 coefficients, A is the sea ice concentration, H is keel depth, L is keel spacing, and $d/L_{\rm f}$ 152is the ratio of the floe draft to floe length. The ratio H/L is the ridging intensity, and AH/L equals the sum of keel depths per unit length. In Eq. (2a, b) P_{r0} and P_{f0} are 153 154 boundary layer integration functions, with some differences due to the choice of u_0 . 155 Tsamados et al. (2014) set u_0 equal to the velocity associated with the inner boundary 156 layer which obeys the logarithmic distribution law, $P_{r0} = P_{f0} = \left[\ln(h_{ref}/z_0) / \ln(z_{ref}/z_0) \right]^2$, 157 where z_0 is the roughness length and the selection of h_{ref} and z_{ref} is referred to 158 Tsamados et al. (2014). $P_{r0} = P_{f0} = 1$ when u_0 was selected as the undisturbed far-field 159 flow velocity by Lu et al. (2011). These two parameterisations of Eq. 2a using

160	different boundary layer functions can be converted, with a transformation factor of
161	0.76 proposed by Zu et al. (2022). In Eq. (2c), m is skin drag attenuation parameter. In
162	Eq. (2a, b), $\Gamma = \Gamma(L/H)$ is the empirical sheltering functions mentioned above. The
163	same sheltering function is employed between adjacent ice floes (Eq. 2b) and between
164	adjacent ice ridges (Eq. 2a). We only focus on the sheltering function in ridge keel
165	drag in this study.

166 **3 Methods**

167 **3.1 Experimental setup**

168 The laboratory experiment was conducted in a water tank which was 11.5 m long, 169 0.45 m wide, and 0.50 m deep (Fig. 1a). The basic experimental setup was originated 170 by Wang et al. (2022) but has been improved here as summarized below. Similarity 171treatment was adopted for the major parameters based on dimensional analysis. The 172 dependent quantities were the drag force F, keel spacing L, keel depth H, water depth 173 D, flow velocity u, keel slope angle α , fluid density ρ , and kinematic viscosity 174 coefficient of fluid v. H, ρ , and u were selected as the fundamental dimensional 175quantities, and the derived dimensionless quantities were (Table 1): $\pi_1 = c_r$ is the 176 local drag coefficient, π_2 is the dimensionless spacing, π_3 is Reynolds number (Re),

177 and π_4 is dimensionless depth. It is necessary to ensure that the ranges of these 178 dimensionless quantities in the water tank experiment are consistent with the natural 179 conditions.

Dimensionless quantities	π_1	<i>π</i> ₂	π ₃	π_4	π_5
Expression	$\frac{F}{ ho u^2 H}$	$\frac{L}{H}$	$\frac{uH}{v}$	$\frac{D}{H}$	α

Table 1 Similarity basis of the model experiment

180

181

182 The keel model was set up as a solid isosceles triangle, which has been used in 183 many laboratory experiments (e.g., Pite et al., 1995; Zu et al., 2021) and geophysical 184 applications (Roberts et al., 2019; Lemieux et al., 2015). The model material was 185 plexiglass, as used by Waters and Bruno (1995), Pite et al. (1995), and Zu et al. (2021). 186 The model shape corresponds to observed sea ice ridge keels (Timco & Burden, 1997), 187 but the macroporosity of 11-20% between keel blocks (Kharitonov & Borodkin, 2020; 188 Strub-Klein & Sudom, 2012) was ignored. A low macroporosity may occur in nature 189 near the waterline due to ridge consolidation (Leppäranta et al., 1995; Roberts et al., 190 2019). The keel slope angle of first-year ice is mostly in the range of $20^{\circ}-50^{\circ}$, with a 191 mean value of 29.6° (Strub-Klein & Sudom, 2012; Timco & Burden, 1997; Davis &

192 Wadhams, 1995) and a maximum of 87° (Kharitonov & Borodkin, 2020). The keel 193 depth (minus background ice thickness) mostly varies from 2 to 12 m, with an average 194 of 5 m and maximum exceeding 27 m, (Metzger et al., 2021, Kharitonov & Borodkin, 195 2020; Wadhams et al., 2011). 196 The parameter settings for the experiment are given in Table 2. All four keel 197 depths H were measured with the fixed tank depth D = 0.45 m, thus D/H = 3.75-22.5, 198 whereas the variable water depth experiments were conducted with $\alpha = 45^{\circ}$, H = 0.12199 m, and D = 0.30 - 0.50 m. For the keel model H = 0.02 m, the lowest flow velocity is u = 0.2 m/s since the drag force is too small (down to 10^{-2} N) to be accurate for u = 0.1200 201 m/s. The model for $\alpha = 90^{\circ}$ was a 5 mm thick rectangular slice. 202 The dimensionless spacing L/H is a classical parameter in wake analysis (Du et 203 al., 2019), and it is equal to the inverse ridging intensity, which is a common spatial 204 statistic of ridging (see Leppäranta, 2011). It was ensured that the keel models of 205 different depths corresponded to fixed L/H. The min and max L/H achievable for 206 different keel sizes depend on the width of the model and the length of the tank: $L_{min} =$ 207 $2(H+0.02)/\tan(\alpha)$, where 0.02 m was freeboard, and $L_{max} = 6$ m in the stable uniform 208 flow section, accounting for the length of the tank and the locations of the inlet and 209 outlet. The characteristic length in the Reynolds number was chosen as the wetted

210	perimeter length of the keel model, $Re = [2uH/\sin(\alpha)]/v$, $v = 1.003 \times 10^{-6} \text{ m}^2/\text{s}$. The
211	range of Re was 8×10^3 to 2.47×10^5 . The laboratory experiments were considered
212	similar to the natural marine environment when the flow was turbulent or $Re > 10^4$
213	(Zu et al., 2021). The cases of $Re < 10^4$ were at the margin and may occur in nature at
214	near stationary ice conditions.
215	

216Table 2 Experimental cases for the drag force measurement. The asterisk shows217the cases the flow field was measured by the PIV system.

Drag force measurement			
<i>u</i> (m/s)	0.1, 0.2, 0.25*, 0.3*, 0.35		
<i>H</i> (m)	0.02*,0.04*,0.08*,0.12*		
α (°)	20*,30,45*,60*,90		
L/H ^{Note1}	2.5*, 5*, 7.5*, 10*, 20*, 30, 50,70,100		
$D(m)^{Note2}$	0.50,0.45*,0.40,0.35,0.30		

218 **3.2 Drag force measurements**

219	The drag force measurement device is shown in Fig. 1b. The selected speed was
220	obtained by adjusting the frequency of the flow generator machine, which is linearly
221	related to the speed. A flow stabiliser was added to the inlet of the tank to ensure a
222	uniform inflow. Three brackets were placed on the side rails of the tank, and they

223	were used to connect the acoustic Doppler velocimeter (ADV), the upstream and
224	downstream keel models, abbreviated as keel-1 and keel-2, respectively.
225	Real-time monitoring and recording of the incoming flow velocity was
226	conducted using the ADV. The force sensors were fixed to the keel models vertically
227	and connected to the screw on the bracket integrally, enabling height adjustment.
228	Two-component force sensors were chosen to measure the form drag force in the flow
229	direction, with a range of 10 N. To reduce errors in the data, the spacing of the keels
230	was adjusted by only moving keel-2. In addition, a round rod (6 mm diameter) was
231	fixed to keel-1 bracket as a nonload item (Fig. 1b). Its original goal was to record the
232	vibration of the flow generator and the brackets above the tank that helped to remove
233	the disturbances on the measurements of keel drag force. This was similar with the
234	experiments used to record the noise of cylinders with a large aspect ratio in a wind
235	field (Reza et al., 2023). We verified that the effect of disturbances such as flow
236	generators was extremely small (<0.9%) and could be ignored. Altogether, 730 groups
237	of force measurement data were obtained (Table 2).





238

Fig. 1 (a) A schematic drawing of the study tank. W1 and W2 represent the windows

of the PIV shots, and the shaded area in the middle is the overlap. (b) A photograph

- showing the model test device.
- 242 **3.3 Flow field measurement**

Particle image velocimetry (PIV) is an instantaneous, multipoint, noncontact hydrodynamic measurement method that can record velocity distributions over many spatial points in the same transect, providing a rich spatial structure of the flow field (Westerweel et al., 2013). The PIV system was used in some groups in the experiments (Table 2) with the shooting positions (windows 1 and 2) shown in Fig. 1(a). The origin of the coordinate system was located at the intersection of the midline of keel-1 and the free surface, and the positive axis directions are shown by the arrowsin the fig.1(a).

251	The PIV installation diagram is depicted in Fig. 2(a). The steps of the PIV
252	method were as follows: First, to eliminate the excess light in the environment,
253	darkroom conditions were created, and the test area was shielded. The tracer particles
254	added to the fluid were hollow glass beads with a diameter of 50 μm and a density of
255	1.05 g/cm ³ , and with excellent light reflectivity and high followability. A continuous
256	laser with a wavelength of 532 nm was used as the light source to illuminate the area
257	of interest. A black coating was added to the model to reduce its surface reflection and
258	the edge effect. Finally, the Photron FastCAM SA5 ultrahigh-speed camera was used
259	to take continuous images at a fixed position in the given time interval, with a
260	resolution of 1024 \times 1024 pixels (Fig. 2b). It can be adjusted in both the horizontal
261	and vertical directions, ensuring that the camera is perpendicular to the laser surface
262	when capturing.
263	The Window 1 and Window 2 were both of approximately 45×45 cm size. The
264	overlapping areas of the windows were averaged and then spliced together. The
265	captured frame rate was 250FPS and the storage time for the cases was 60s. An

266 interrogation window of 32×32 pixels was used to capture a sort of high

- 267 displacement, and then 16×16 pixels were taken in the final pass to give a refined
- 268 vector resolution (Xue et al., 2023). The maximum pixel displacement in the PIV
- image was 5.47 pixels at u = 0.25 m/s. Considering the systematic error of 2D PIV, the
- 270 window deformation, and the error from speed control, the relative uncertainty was 4-
- 271 5% (Singha et al., 2009; Jin et al., 2014), equivalent to 0.22-0.27 pixels.



- Fig. 2 (a) PIV experimental installation diagram; (b) Camera arrangement
- 274 diagram
- 275
- **4 Results**

4.1 Drag force on ice ridges

Fig. 3 shows the drag forces on the upstream and downstream keels vs. L/H for α

279 = 45° and H = 0.08 m. The other cases are not shown because of the similar patterns. 280 Fig. 3(a) shows that L/H influences F_1 only within L/H < 10, where F_1 increased by 5% 281 due to the presence of keel-2. Thus, the upstream keel drag is the same as in the case 282 of an individual keel which has been investigated thoroughly by Wu (2016) and 283 provides a reference for the decay of the downstream drag F_2 . Fig. 3(b) reveals that F_2 284 was reduced as a result of sheltering by the upstream keel, and this attenuation 285 decreased with increasing L/H as F_2 gradually approached F_1 as described in more

detail in Section 3.4.



Fig. 3 (a) The drag force F_1 and (b) the drag force F_2 vs. the L/H for $\alpha = 45^{\circ}$ and H =

289 0.08 m.

290

291 The drag force $F_2 = F_2(u^2)$ for $\alpha = 45^\circ$ and H = 0.08 m is shown in Fig. 4. The 292 dependence was linear, consistent with the classical drag force formula (Tennekes & Lumley, 1972; Arya, 1975; Wamser & Martinson, 1993). F_2 appeared negative at L/H ≤ 5 , when keel-2 was located in the recirculation zone formed by flow separation behind keel-1. This reverse drag force was also linear with u^2 . The presence of different ranges of recirculation zones has been demonstrated in submerged wake experiments for a conical island (Oruo et al., 2017) and in tank experiments with rectangular columns placed in series (Pinarbasi et al., 2015).



299

300 Fig. 4 The drag forces F_2 vs. the velocity squared u^2 when $\alpha = 45^\circ$ and H = 0.08 m

301

In the mechanism of drag variation, the main turbulent structures and the development in the wake field around keels are essential. The time-averaged streamlines and dimensionless velocity \bar{u}/u_0 for $\alpha = 45^\circ$, $u_0 = 0.25$ m/s, and L/H = 5 (H= 0.08 m) are plotted in Fig. S1, corresponding to $F_1 = 1.94$ N and $F_2 = -0.06$ N. The characteristic structures in the wake mainly included recirculation zones, vortex 307 shedding, separated shear layers and an acceleration zone (Ouro et al., 2017),
308 analysed in Text S1. The development of turbulent structure leads to changes in the
309 pressure difference between the leading and lee edges of the keel-2, which is the
310 essence of the drag force.

311 **4.2 Influence of flow velocity on the sheltering effect**

312 The ratio of the drag forces $\Gamma = F_2/F_1$ is defined as the sheltering coefficient, 313 which characterises the strength of the sheltering between the keels. Fig. 5 shows this 314 coefficient as a function of the squared flow velocity for $\alpha = 45^{\circ}$ and H = 0.12 m. As 315 u^2 increased, Γ remained stable within 4.2%. According to the dimensional analysis in 316 Table 1, the flow velocity *u* impacted the sheltering effect by changing only *R*e, which 317 is crucial in the laminar flow regime but barely affects the turbulent regime (Zu et al., 318 2021). Since the ocean currents under sea ice are mostly turbulent (McPhee, 2008; 319 Kharitonov & Borodkin, 2020), Γ is independent of the *u*. Based on the quadratic drag laws, $\Gamma = F_2/F_1 \sim u_2^2/u_1^2 = \text{constant}$ at a fixed L/H, where u_1 and u_2 are the flow 320 321 velocities at keel-1 and keel-2, respectively. Thus, u_2/u_1 depends only on L/H for the 322 fixed $\alpha = 45^{\circ}$, implying that the velocity deficit caused by keel was not connected 323 with the incoming flow. Li and Sherman (2015), in a study focused on the

- 324 aerodynamics and morphodynamics of the sheltering effect behind also confirmed
- 325 that the ratio of the disturbed velocity to the incoming velocity at a certain elevation
- 326 and distance was nearly an independent factor.

1.2 0.8	•	•	\$	•	*	<i>L/H</i> ■ 2.5 ♦ 5 ▲ 7.5 ▼ 10
⊾ _{0.4}	▼	▼ ▲	▼ ▲	▼ ▲	▼ ▲	 20 20 30 ★ 50
0.0	Ŷ	¢	♦	♦	¢	
-0.4 0.	00	 0.0	$\frac{1}{100}$	0.10 n^2/s^2)	<u> </u>	0.15

327

328 Fig. 5 The sheltering coefficient Γ vs. the squared velocity u^2 for $\alpha = 45^{\circ}$ and H = 0.12

m

329

330

4.3 Influence of the ridge spacing on the sheltering effect



3.38 1. The sheltering coefficient was negative and changed sign at approximately L/H = 5, 3.39 a similar critical value of L/d=5 was found for the drag on the downstream cylinder 340 (diameter *d*) in simulations with two cylinders in tandem by Li et al. (2018). An 341 examination of the flow pattern was needed to interpret the phenomenon. For L/H > 5, 342 sheltering gradually became weaker.



343

344 Fig. 6 The sheltering coefficient Γ vs. dimensionless spacing L/H when $\alpha = 45^{\circ}$. The

345 blue curve represents the fit for $\Gamma = \Gamma(L/H)$. The dashed line at L/H=30 separated the

growth and asymptotic parts of the curve.



Fig. 7. Contours of time-average streamlines (a, b), turbulent kinetic energy

349 $1/2(\overline{u'^2}+\overline{w'^2})$ (c, d), and dimensionless Reynolds shear stress $(\overline{u'w'}/u_0^2)$ (e, f). L/H=5

350 (b, d, f) and L/H=10 (H=0.04 m) (a, c, e). $\alpha = 45^{\circ}$, $u_0=0.25$ m/s, and u' and w' are the

351 velocity fluctuations along the flow and vertical directions, respectively.

352



359 turbulence structure and development mechanism of the wake as the fundamental360 reason for the drag variation.

361	When the two keels are very close (Fig. 7b, d, f), the shear layer separated by
362	keel-1 will adhere to the surface of keel-2 where the vortex wraps around, so they can
363	be treated as one bluff body (Du et al., 2019). This flow pattern is denoted as a single
364	slender-body regime (Sohankar, 2012), which exposed the leading face of keel-2 to a
365	strong suction and resulted in a thrust (negative F_2) rather than a drag force with $\Gamma < 0$
366	for $L/H=5$. The peak values of TKE and Reynolds stress occurred underneath keel-2
367	along the transition from the recirculation region to the far wake, corresponding to the
368	region where shear-layer vortices lost coherence due to interactions with the incoming
369	flow and backflow. The continuous vortex structure was gradually shed. The Reynolds
370	stress was consistently higher up to $x/H=14$, which meant a long section of intensive
371	turbulent momentum exchange. Thereafter, the stress gradually dissipated downwards
372	in the direction of the flow.
373	Figs. 7a, c, e show the reattached flow regime (Sohankar, 2012), characterised by
374	synchronised vortex shedding from each cylinder. The separated shear layer at keel-1
375	found sufficient space to develop and roll into the distance between two keels, with

376 the reattachment positioned in front of keel-2 (x/H = 6-8). Hence, secondary

377	separation occurred at the tip of keel-2 which presents a co-shedding vortex flow
378	regime (Kurtulmuş, 2022; Sohankar, 2012). The vortex length became distant with
379	the first flow separation reattaching at $x/H = 7$, while the second reattachment point
380	was shifted to $x/H = 14$. The locations of the peak values of TKE and Reynolds stress
381	were advanced to $x/H = 2-4$ between keels, accompanied by the enhancement of the
382	TKE and the weakening of the Reynolds stress which was elevated at the intersection
383	of eddies and incoming flows behind keel-2 due to the induced secondary separation,
384	but it was weaker. The reduction in the turbulence intensity in the wake also indicated
385	a decrease in the sheltering capacity of the wake.
386	In the range of $L/H \ge 30$, Γ increased with L/H towards the asymptote, and the
387	wake field was gradually restored to an undisturbed state. The third mode, appearing
388	in this far-wake region, manifested as vortices falling off independently of the two
389	structures. In the aerodynamic experiments by Sakamoto et al. (1987) on two square
390	prisms in tandem, this wake mode occurred when $L/H > 27$.

4.4 Influence of the keel slope angle on the sheltering effect

- 392 Fig. 8 shows Γ against α , with the error bars representing the standard deviation.
- 393 Since negative values of Γ occurred, the logarithmic coordinate transformation was

394	performed for Γ +0.5. With increasing slope, the sheltering coefficient tended to
395	decrease. Linear fits in the plot were excellent ($R^2=0.98$), implying a power function
396	relationship Γ +0.5 = 3.02* $\alpha^{-0.31}$ (20° $\leq \alpha \leq 90^{\circ}$). The decline in Γ was strong at α =
397	20° - 30° , which is the common range of the slope Arctic sea-ice ridges (Timco &
398	Burden, 1997; Strub-Klein & Sudom, 2012). Furthermore, the decrease of Γ caused
399	by the increasing α revealed different characteristics in the regimes $L/H < 20$ and L/H
400	\geq 20, represented in red and blue colours, respectively. The wake can be divided into
401	near and far wake zones (Vermeer et al., 2003; Williamson, 1996), and the influence
402	of the keel shape on the wake field was mainly concentrated in the near-wake region,
403	while the geometry was no longer important in the far-wake region. For instance,
404	when α increased from 20° to 90°, Γ decreased from 0.65 to 0.22 at $L/H = 10$ and
405	from 0.93 to 0.78 at $L/H = 50$.



407 Fig. 8 The sheltering coefficient Γ vs. slope angle α . The dashed red curves represent

408 the near-wake region, the dashed dark blue curves represent the far-wake region, and

409 the solid blue line represents the fitted curve for
$$L/H = 10$$
.

410

411 More insight into the flow characteristics in the near-wake region could be 412 obtained by extracting the PIV flow field results at different slopes. Fig. 9 illustrates 413 the time-average streamlines, TKE, and Reynolds shear stress for $\alpha = 20^{\circ}$ and $\alpha = 60^{\circ}$ 414 when L/H=10, corresponding to $\Gamma = 0.65$ and 0.31. The result can be analysed in 415 conjunction with the flow field at $\alpha = 45^{\circ}$ and L/H=10 in Fig. 7(a, c, e). The wake 416 flow extent at $\alpha = 20^{\circ}$ had a small longitudinal range of z/H = 1-1.5, and the 417 secondary separation was weak near keel-2, with no significant vortex production 418 behind (Fig. 9a). The attachment point was at x/H = 5-6. The flow separation at keel-1 419 with $\alpha = 60^{\circ}$ was more intense (Fig. 9b), and the reattachment point was shifted to x/H420 = 7-8, implying a deferred shedding process of the vortex. Steeper angles lead to 421 increased pressure gradients around the keel, which affect the flow separation process 422 and the vortices that roll up behind the keel, resulting in different sized vortex 423 structures, vortex shedding patterns, and dissipation (Wang et al., 2017). 424 Overall, the wake generated with $\alpha = 60^{\circ}$ produced stronger turbulence intensity

425 and more active flow separation processes compared to $\alpha = 20^{\circ}$ and $\alpha = 45^{\circ}$, leading

426	to an increase in sheltering strength and extent. The larger range of TKE and
427	Reynolds stress peak values was maintained up to $x/H = 7-8$ and $z/H = 3-4$. The
428	increase in the Reynolds stress enhanced the momentum transfer across the shear
429	layer, thereby delaying wake recovery. The first triggered peak value occurred at x/H
430	= 2-4, developing and dissipating along the flow direction. The Reynolds stress
431	peaked again but at a significantly lower level than in the first one due to the
432	secondary disturbance of keel-2 that was also related to α . The Reynolds stresses
433	behind keel-2 were higher with $\alpha = 60^{\circ}$ than with 45° and 20° due to the larger
434	velocity gradient of the fluid flowing through the keel. The greater shear stress
435	intensifies the level of turbulence in the wake, accompanied by a significant increase
436	in TKE.



438	Fig. 9 Contours of time-average streamline (a, b), turbulent kinetic energy
439	$1/2\left(\overline{u'^2}+\overline{w'^2}\right)$ (c, d), and Reynolds shear stress $\overline{u'w'}/u_0^2$ (e, f) for $\alpha = 20^\circ$ (a, c, e) and
440	$\alpha = 60^{\circ}$ (b, d, f) when $L/H = 10$, corresponding to $\Gamma = 0.65$ and 0.31, respectively.
441	

442 5 Discussion

443 **5.1 Sheltering functions**

444 The sheltering function was affected not only by the dimensionless spacing L/H445 but also by the slope angle α , but it was independent of the flow velocity u. In 446 Sections 4.2-4.4, the curve $\Gamma = \Gamma(L/H)$ at $\alpha = 45^{\circ}$ was better fitted by the exponential 447 function (Fig. 6), similar with the previously presented sheltering function. The 448 relationship between Γ and α was given by a power law (Fig. 8). To obtain a new 449 expression of the sheltering function based on the present experimental data, the two 450 functional forms above are combined. The least squares method was used for the 451 optimal solutions of the pending parameters. The resulting sheltering function $\Gamma = \Gamma$ 452 $(L/H, \alpha)$ is:

453
$$\Gamma = [1-1.56 \exp(sL/H)] \times 1.2a^{-0.08}$$
, $s = 0.001a - 0.15$ (3)

454 The range of application is $2 \le L/H \le 110$ and $10^\circ \le \alpha \le 90^\circ$. Formerly s was a

455	constant $s = -0.18$ (Hanssen-Bauer & Gjessing, 1988) and $s = -0.5$ (Lüpkes et al.,
456	2012), but here it was found that s increased with α . To account for this, s was taken
457	as a linear function of α . The exponential factor reflects the influence of both
458	arguments, while the power function factor depends only on the slope angle. First, the
459	exponent <i>sL/H</i> determines the sign of the sheltering coefficient. At $sL/H = -0.44$, we
460	have $\Gamma = 0$ and $s = -0.14$ when $\alpha = 10^{\circ}$ and -0.06 when $\alpha = 90^{\circ}$. The minimum and
461	maximum values of sL/H were -16.8 and -0.12, respectively; $\Gamma > 0$ when -16.8 $< sL/H$
462	< -0.44, while $\Gamma < 0$ when -0.44< $sL/H <$ -0.12. In the limiting cases, when $L \rightarrow 0$,
463	$\Gamma \rightarrow -0.672 \alpha^{-0.08}$, and when $L \rightarrow$ infinity, $\Gamma \rightarrow 1.20 \alpha^{-0.08}$. Ninety percent of the increase
464	in Γ with increasing L was concentrated at $L/H \le 30$. In this range, the reduction in Γ
465	could reach 69% due to the steepening of the keel slope. Fig. 10 shows the fitted
466	curves (Eq. 3) for $\alpha = 20^{\circ}$, 45° and 60° within $2 \le L/H \le 110$, with $R^2 = 0.97$, RMSE
467	=0.07.



469	Fig. 10 Comparison of Eq. (3) with the experimental data. The adjusted R^2 is
470	0.97, and the root mean square error is 0.0697. The solid points are the experimental
471	results, and the error bars represent the standard deviations.
472	
473	Fig. 11 compares our formula (Eq. 3) with the two previously reported equations.
474	The pink shaded area presents Eq. (3), extended through $\alpha = 10^{\circ}-90^{\circ}$. The average
475	slope angles in different papers varies, but they are generally about 30° which was
476	chosen here as the representative angle (Timco & Burden, 1997; Strub-Klein &
477	Sudom, 2012). Eq. (3) was then $\Gamma_{30} = 0.92-1.44 \exp(-0.12L/H)$. The factor -0.12 in
478	the exponent is larger than the fixed constant -0.18 (Hanssen-Bauer & Gjessing, 1988)
479	and -0.5 (Lüpkes et al., 2012). The minimum value of this factor in Eq. (3) was -0.14,
480	and thus Eq. (3) represents stronger sheltering than the exponential function Γ_1 but
481	weaker than the power function in Γ_2 . Compared to Γ_{30} , the exponential function was
482	15.2% higher, while the power function was 17.1% lower within $L/H=2.5-100$. Eq. (3)
483	includes the reverse drag force at small spacings and the effect of the keel slope.





Fig. 11 Comparison between Eq. (3) (the pink shaded for $\alpha = 10^{\circ} - 90^{\circ}$) and the previously reported exponential (green dashes) and power (blue dotted lines)

487 functions. $\alpha = 30^{\circ}$ represents the typical slope angle of keels.

488

489 The sheltering coefficient in Eq. (3) can reach down to -0.44, while in the other 490 two functions, the minimum values were 0.30 and 0.08 when L/H = 2. Such small 491 spacings are uncommon in a real sea ice cover, but ridges sometimes form in closely 492 packed clusters. Valenti et al. (2021) proposed from observations that L follows the 493 lognormal distribution presented by Wadhams and Davy (1986), with the third 494 quartile values Q3 = 292 m and 277 m in the Beaufort and Chukchi Seas, respectively, 495 and the average keel depth is approximately 5-8.5 m. As a result, L/H was mostly in 496 the range of less than 60, accompanied by Γ from -0.14 to 0.93 for $\alpha = 30^{\circ}$ (Eq. 3). 497 Thus, the form drag from ice ridges was significantly weakened due to the sheltering.

498	This effect will be strengthened as the frequency of ice ridges increases during the
499	winter (Steele et al., 1989). Besides, the occurrence of ridge keels tends to be
500	consecutive and overlapping in the spatial underice profiles from the moored IPS and
501	ADCP data of Valenti et al. (2021), resulting in the sheltering of the wake being
502	superimposed and enhanced.
503	The inclination angle of obstacles has been validated in hydrodynamics and sea
504	ice dynamics as a key parameter of the drag force of two square pillars (Du et al.,
505	2019; Yen et al., 2008) and ridge keels (Zu et al., 2021). In particular, the geometry of
506	ridged ice varies greatly for the depth and cross-sectional shape, e.g., a maximum
507	depth of 27 m and a maximum slope angle of 87.5° were observed by Kharitonov &
508	Borodkin (2020). The differences in α depend on many factors, such as wind speed
509	and current direction, which drive sea ice deformation. Observations show keel slopes
510	in the Barents Sea of 8°-29° and in the Shokalsky Strait of 11.2°-87.5° (Sand et al.,
511	2015; Kharitonov & Borodkin, 2020). Neglecting the changes in the slope angle, i.e.,
512	with the fixed angle of 30° in Eq. (3), the form drag on ice ridges in the Barents Sea
513	will be underestimated by 17.7% within $L/H = 10-40$, while it may be overestimated
514	by as much as 43.5% in the Shokalsky Strait.

515 5.2 Views on sea ice dynamics

516	The most significant advance between the new sheltering function Γ and the
517	previous versions Γ_1 and Γ_2 is the includes of the keel angle, inevitably inducing a
518	change in C_{io} in Eq. (2), which further changes the contribution of oceanic drag force
519	τ_{io} in the sea ice momentum equation. In the free-drifting conditions, sea ice motion
520	depends mainly on the wind and current drag forces, providing a good scenario to
521	check the impact of the new sheltering function. Fig 12. shows the variation of C_{io} (a-
522	c), wind factor (the ratio of sea ice drift velocity to wind velocity) $ u_i / u_a $ (d-f) and
523	the deflection angle θ (g-i) calculated using three different sheltering functions (a, d,
524	g: Γ_1 ; b, e, h : Γ_2 ; c, f, i: Γ in Eq.(3)) with keel geometric parameters H/L and α when
525	A = 0.8. The drag coefficients are calculated using Eq. (2) with reference to the
526	parametrisation scheme of Lu et al. (2011), and the calculation of the wind factor and
527	deflection angle follows the momentum equation of free-drifting sea ice by
528	Leppäranta (2011), with the formulas and parameter values shown in Text S2.
529	Combining data from Valenti et al. (2021) on the distribution of H and L , and
530	observations of H/L from Berner et al. (2021), it can be found that most cases fell
531	within $H/L \ge 0.01$. However, earlier observations show that is predominantly in the
532	range of 0.001-0.01 (Davis & Wadhams, 1995; Granberg & Leppäranta, 1999). As a
533	result, the range of H/L is set to 0.001 to 0.1, covering most keels.

534	Based on Fig. 12(a-c), C_{io} increases as H/L increase but the impact of α is not
535	obvious when $0.001 \le H/L \le 0.01$. Differences between the three schemes are clearly
536	visible when $H/L \ge 0.01$. C_{i0} is higher with Γ_1 and lower with Γ_2 , and in between with
537	Γ . Γ affects C_{io} by changing the contribution of C_r in Eq.(2a). When $H/L \ge 0.01$, C_{io}
538	using Γ varies nonlinearly with the coupling of H/L and α in Fig. 12c, instead of
539	monotonically decreasing as in Fig. 12a and b. L/H and α actions on the wake
540	structure have been illustrated in the PIV results above. The max C_{io} in Fig. 12c only
541	appeared in the region of $0.05 \le H/L \le 0.1$ and $\alpha = 10^{\circ}-30^{\circ}$.
542	$ u_i / u_a $ is a key indicator of wind stress transfer, determining sea ice drift. Based
543	on the general solution of the free drift equation, $ u_i / u_a $ is a function of Nansen
544	number (Na) and Rossby number (R) where $Na \propto (C_{ai}/C_{io})^{0.5}$ and $R \propto (C_{ai}C_{io})^{-0.5}$. For
545	$ R < 0.2$, $ u_i / u_a \approx Na \propto (C_{ai}/C_{io})^{0.5}$ (Leppäranta. 2011). So, $ u_i / u_a $ is in positive
546	correlation with Γ^{-1} . $ u_i / u_a $ are mostly in 0.012-0.024, consistent with the
547	observation of Cole et al. (2017) in the Canada Basin. When $H/L \ge 0.01$, $ u_i / u_a $ also
548	varies nonlinearly due to the sheltering effect. Interestingly, in the range of $0.06 \le H/L$
549	≤ 0.1 , u_i / u_a reaches its maximum 0.024 when α varied from 10 to 90°, implying the
550	amplification of stress transfer from the atmosphere.




560	θ is a function of $(C_{ai}C_{io})^{-0.5}$, then θ is positively correlated with Γ^{-1} . The variation
561	pattern of θ is independent of the chosen values of wind angle and oceanic Ekman
562	angle for $ R < 0.2$, as shown in Text S2. θ presented in Fig13. i is higher than that in
563	Fig12. g while lower than Fig12. h. With the increasing of H/L from 0.01 to 0.1, θ
564	decreases about 36% when $\alpha = 30^{\circ}$. With the increasing of α from 10 to 90° when
565	$H/L=0.05$ θ increases about 24%

566 **5.3 Laboratory experiments vs. polar oceans - limitations and uncertainties**

567 This paper has been devoted to laboratory experiments on the sheltering effect 568 between ice ridges and a new sheltering function has been proposed. The 569 experimental results provide solid references for parametrisation in sea ice dynamics, 570 but differences in the boundary conditions between the water tank and the Arctic 571 Ocean must be clarified because they partly limit the direct application to polar oceans. 572 Firstly, the upper boundary in the tank was a free surface, while the polar ocean 573 surface is covered with ice. This mainly affects whether the incoming reference 574 velocity is chosen to be a logarithmically-law distributed flow velocity in the inner 575 boundary layer Lüpkes et al., 2012) or the far-field velocity independent of the 576 thickness of the boundary layer (Shaw et al., 2008; Österlund et al., 2000). The inner

577	logarithmic layer is approximately 2 m deep in winter (Kharitonov & Borodkin,
578	2020) and less in summer (Cole et al., 2017; Gallaher et al., 2016), as estimated by
579	an empirical formula related to the Ekman depth (Shaw et al., 2008). Therefore, the
580	logarithmic layer is of concern primarily for small ice ridges in winter.
581	The conversion factors of the two reference velocities for the single keel form
582	drag were derived by Zu et al. (2021) as 1.1-1.3, decreasing with increasing ridge
583	height. However, we are concerned with the effect of the reference velocity on
584	sheltering characterised by F_1/F_2 . Section 4.2 showed that the wake generated by a
585	fixed shape keel is independent of u in the turbulent conditions, consistent with the
586	attenuation after passing through the shelterbelt for different wind speeds (Ozawa et
587	al., 2007) and the review of sheltering behind sand fences (Li & Sherman, 2015).
588	Therefore, it is reasonable to assume that the errors introduced by differences in the
589	upper boundary are small and do not significantly impact the form or accuracy of the
590	sheltering function.
591	Secondly, compared to the natural conditions in the Arctic Ocean where $D >> H$,
592	the shallow water in the experimental tank has led to increased sheltering. To better
593	understand this, we performed variable depth experiments with D varying from 0.30–

594 0.50 when $\alpha = 45^{\circ}$, H = 0.12m. The resulting Γ vs. L/H is given in Fig. 13. The dark

595	red curve represents the fit for $D = 0.30$ m ($R^2 = 0.99$), while the light red curve
596	represents the fit for $D = 0.50$ m ($R^2 = 0.98$). The shallow depth may restrict the
597	ability of wakes to expand vertically, inhibiting the spread of eddies along the depth
598	direction and resulting in accelerated wake recovery (Ouro & Nishino, 2021; Singha
599	et al., 2009; Akilli, & Rockwell, 2002). As a result of this feature, Γ gradually
600	increased with decreasing D. D/H was not inserted in the form of the sheltering
601	function but rather illustrated by the determination of its coefficients. No attempt has
602	been made here to fit depth D into the sheltering function for this stage, because Γ is
603	expected to be depth-independent beyond the critical depth (Zu et al., 2021).
604	However, this does not deny the above evaluations of sea ice dynamic parameters in
605	section 5.2. Because Fig. 12(a-c) was recalculated by using approximately 50% lower
606	coefficients of the exponential function in Eq. (3), but the results revealed that the
607	variation trend changed little and the maximum value of C_{io} increased only below
608	15%. The given portion 50% is roughly determined through linearly extending the
609	form of $\Gamma_{D=0.3}$ and $\Gamma_{D=0.5}$ in Fig.13 to the critical depth in Zu et al. (2021).





Fig. 13 The sheltering coefficient $\Gamma = \Gamma(L/H)$ when the tank depth *D* is increased from 0.30 m to 0.50 m at $\alpha = 45^{\circ}$ and H=0.12 m (*D*/*H*=2.5-4.17). The dark red/light red curve represents the fit of Γ for D = 0.30 m/D = 0.50 m; $R^2 = 0.99$ and 0.98. The fits for the other *D* are similar, and not shown here.

616 Thirdly, the sheltering strength of the wake was overestimated in the far-wake 617 region (L/H > 30). Asymptotically, the sheltering function approached a limit less than 618 1, within 0.82-0.97 in Fig. 11. This was somewhat different from our intuition and 619 was caused by the nature of laboratory experiments where the system is forced by the 620 inflow boundary condition. The flow section in the tank was 11 m long inevitably 621 experienced a loss of flow velocity and the floating free surface which may result in Γ 622 < 1. Since the momentum was fed into the system from the boundary, passing a keel 623 resulted in a loss of momentum due to the form drag. Consequently, the drag force

624	was lower than expected at keel-2, with Γ <1. In nature, momentum is fed into the
625	system by distributed forcing over the whole ice cover, and a corresponding
626	momentum loss does not occur. Thus, when the ridge spacing tends to infinity, the
627	downstream keel drag force is not affected by upstream sheltering, corresponding to
628	Γ=1.

5 Conclusion

630	Laboratory experiments were carried out to investigate the form drag of adjacent
631	sea ice ridges and the sheltering effect of varying keel depth H , dimensionless spacing
632	L/H , keel slope angle α , and flow velocity u . The conclusions are drawn as below.
633	The drag force F_1 on the upstream keel exhibited stability as L/H increased and
634	was close to the case of an individual keel; just within $L/H < 10$, F_1 increased by 5%
635	due to the presence of downstream keel. The drag force F_2 on the downstream keel
636	became larger with increasing L/H and gradually approached an asymptote. At $L/H \le$
637	10, F_2 was opposite to the flow direction, as the downstream keel was located within
638	the recirculation zone created by the upstream one. F_2 was proportional to the square
639	of <i>u</i> , consistent with the classical drag force formula.
640	The sheltering coefficient $\Gamma = F_2/F_1$ characterises the sheltering strength. Γ is

641	related to L/H and α but is independent of u . A new formula was proposed by fitting
642	the experimental results: $\Gamma = [1-1.56 \exp(sL/H)] \times 1.20a^{-0.08}$, $s = 0.001a - 0.15$; s was
643	changed from a constant to a linear function of α . Overall, the function can be treated
644	as the product of an exponential function and a power function. On the one hand,
645	there is an exponential relationship between Γ and L/H . Γ increased as the L/H
646	increases and underwent the transition from negative to positive at $L/H = 5$. The time-
647	averaged streamlines, turbulent kinetic energies and Reynolds stresses for $L/H = 5$ and
648	L/H = 10, corresponding to the single body regime and the reattached flow regime,
649	were calculated using the PIV results to validate the effect of L/H variations on wake
650	turbulence characteristics. On the other hand, there is a power relationship between Γ
651	and α which is characterised a decreasing trend as α increases. By comparing the flow
652	flied results for $\alpha = 20^{\circ}$, 45° and 60°, the wake expanded longitudinally as the angle
653	increased, accompanied by larger recirculation zone, stronger TKE, and greater
654	Reynolds stress. However, the influence of α on Γ was also coupled with L/H . α was
655	mostly impactful in the near-wake region.
656	A sea ice free-drift dynamics case was calculated using three different sheltering
657	function. The differences in results of C_{io} , wind factor $ u_i / u_a $ and the deflection angle

658 θ are more pronounced at $H/L \ge 0.01$, corresponding to more dense ridged ice

42

659	nowadays in observations than before (Berner et al. 2021; Davis & Wadhams, 1995).
660	The values calculated using the new sheltering function are larger than the power
661	function, smaller than the exponential function. The coupling of H/L and α leads to a
662	non-linear variation of C_{io} , $ u_i / u_a $ and θ at $H/L \ge 0.01$, however, the remaining two
663	functions could not to capture this feature. In the range of $0.05 \le H/L \le 0.1$, the
664	sheltering strength is enhanced and $ u_i / u_a $ increases from 0.012 to 0.024 when α
665	increased from 10 to 90°. It is contrast to a constant wind factor for different ridge
666	slope as using previous sheltering functions.
667	As a solid step towards the final objective of parameterization, the new function
668	has a clear physical meaning that incorporates both the impact of the dimensionless

spacing and slope angle on the sheltering, which is an improvement of the previous

results by Steele et al. (1989). The outcome has clarified the major parameters

670

671 affecting sheltering as well as the turbulent structure and characteristics of the flow

672 field around double keels. Because of the limitations from water tank experiments,

Eq. (3) cannot be further optimised and then directly applied to the polar ocean.
However, it can be straightforwardly improved and extended to depth-independent
situations using numerical simulations in the next step. Possible directions of future

676 research include, but are not limited to, (i) considering the probability distribution

677	function of geometric parameters and spacings of ice ridges, (ii) considering the
678	porosity of ice ridges, and (iii) probing the mechanisms of variation in shelter effects
679	in a stratified two-layer flow. It is anticipated that a universal relationship can be
680	obtained that is applicable in sea ice dynamics modelling.
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687 Data Availability Statement

688 The laboratory experiment data include drag force measurements and PIV flow field

- 689 measurements shown in this manuscript can be accessed via the website
- 690 (https://doi.org/10.5281/zenodo.10065506)

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Figure 1.



Figure 2.





Figure 3.



Figure 4.



Figure 5.


Figure 6.



Figure 7.



Figure 8.



Figure 9.



Figure 10.



Figure 11.



Figure 12.



Figure 13.

