Moisture migration on the shear zone controls landslide failure models

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Abstract: The shear behavior of loess is of paramount importance to 6 understanding of mechanisms triggering landslides. In this research, using a ring 7 shear apparatus, a series of ring-shear tests was performed on the Lanzhou 8 9 loess to examine the residual shear behavior of the loess with different water 10 contents. The results show that the increase of water content results in the reduction of void ratio of the samples, when the water content continuously 11 12 increase to a special threshold, the void ratio of the samples increase lightly then reach a stable state. Correspondingly, the cohesion increase and the friction 13 angle decrease with increasing water content until the water content threshold, 14 15 and thereafter the cohesion decreased quickly to zero and the friction angle stabilized at a residual value. Futhermore, the localization of shear deformation 16 result in difference between shear zone and soli layers below and above the 17 shear zone in void ratio, causing the two different types of water migration on 18 shear zone. Before the water content threshold, the water is sucked towards 19 shear zone from soil layers above and below the shear zone, and then the water 20 is expelled shear zone into soil layers above and below the shear zone. The 21

22	water migration can influence shear strength of soil which depends on the type
23	of water migration on shear zone. These results reveal that the volumetric
24	change of soil not only has great influence on shear behavior during wetting but
25	also is related to the water migration during shearing.

Keywords: Loess landslide, shear zone, volumetric response, water
migration, residual shear strength, ring shear tests

30 **1 Introduction**

Landslides occurred in loess setting cause serious casualties and much 31 destruction almost every year. Loess landslides are often triggered by rainfall 32 and irrigation (Gibbs and Holland, 1960; Lutenegger and Hallberg, 1988; 33 Derbyshire et al., 1991; Derbyshire et al., 1994; Dijkstra et al., 1994; Derbyshire, 34 2001; Ueno et al., 2005), but sometimes earthquake is also an important trigger 35 36 to them (Lutenegger, 1981; Ishihara et al., 1990; Zhang and Wang, 2007). In 37 these conditions, the initiation and movement of loess landslides were generally implicated in reduction of shear strength of soil accompanied by gradual or rapid 38 39 increase of water content.

There has been considerable research into examining the effect of water 40 content on shear strength of loess, and much of this effort has recognized that 41 42 the impacts of water content are of paramount importance to understanding the 43 mechanisms of loess landslides (Gibbs and Holland, 1960; Lutenegger, 1981; Ishihara et al., 1990; Derbyshire et al., 1994; Dijkstra et al., 1994; Zhang and 44 Wang, 2007; Zhang et al., 2009). Most of the previous studies were performed 45 46 on undisturbed loess samples in triaxial tests, mainly focusing on the natural and saturated samples (Gibbs and Holland, 1960; Audric and Bouquier, 1976; 47 RenéJacques, 1988; Tan, 1988; Abduljauwad and Al-Gassous, 1991; Phien-wej 48 et al., 1992; Zhang et al., 2009). Additionally, there have been few attempts to 49

examine shear strength behavior of remolded loess samples in ring shear tests 50 (Derbyshire et al., 1994; Dijkstra et al., 1994; Zhang and Wang, 2007). 51 Derbyshire et al. (1994) and Dijkstra et al. (1994) conducted a series of tests at 52 different water contents on remolded loess samples taken from Chinese 53 Lanzhou region, using a modified Bromhead ring shear apparatus. The results 54 showed that the apparent cohesion increased and the internal friction angle 55 56 decreased with increasing water content untill a certain threshold, after which 57 the apparent cohesion decreased rapidly and the internal friction angle stabilized at a residual value theoretically. In practice, however, the leakage of water and 58 59 fine particle during shearing caused an increase in the internal friction angle due to the deficiencies of the apparatus (Derbyshire et al., 1994). Derbyshire et al. 60 (1994) also pointed out that the leakage also affected the cohesion of soil. 61 Hence, we can find that change in water content during shearing has impact on 62 63 strength behavior of the loess. Furthermore, field investigation to four loess landslides have shown that the water content on shear zone is higher than that 64 of soil mass above and below the shear zone (Long et al., 2007). The authors 65 66 hypothesized that the shear zones were related to pedological soil acting upon a relatively impermeable layer, causing higher water content on shear zone. To 67 the phenomena existed, we need more detailed study to clarify the effect of 68 water content change on strength behavior of a soil during wetting and shearing. 69

However, the strength behaviors of soil under different water contents dry to saturated conditions have been studied in less detail. Furthermore, there is no research concerned with water change within samples with various water contents during shearing and its relation with strength behavior of soil.

74 In this paper, we examine loess samples taken from a landslide, occurring on Jiuzhoutai area of Lanzhou City, China. The approach of our research is to 75 76 investigate experimentally, using a ring shear apparatus, the strength behavior 77 of the loess samples suffered large shear displacement at different normal 78 stresses by varying water content. The main purpose of our research is to study 79 the effect of water content on residual strength behavior of the loess. We focus on analyzing variation of residual shear strength parameters (cohesion and 80 friction angle) of soil with different water contents, and pay special attention to 81 the water content change at different layers within samples after shearing and its 82 83 relation with the residual shear strength parameters. The results of our research showed that water migration on shear zone can afford a new understanding to 84 mechanisms triggering loess landslides, especially in unsaturated settings. 85

86

87 2 Material and methods

88 2.1 Testing sample

We took loess samples from a landslide, occurring on Jouzhoutai area of Lanzhou city, Gansu Province, China (Fig. 1). Many landslides occurred without direct triggers such as rainfall or earthquake, and this landslide was a typical case. Despite the modest volume involved, this kind of landslide has caused many fatalities and much destruction, due to their high velocity along steep slope and the absence of incipient movement evidence.

The loess is an aeolian silt with a mean particle diameter of 0.018, coefficient of uniformity of 5.50 and coefficient of gradation of 1.05. The grain size distribution of the loess samples is presented in Fig. 2. The sample approximately is consisted of about 92% silt, 7% clay and 1% sand. Some basic physical properties of the samples are listed in Table 1. It has low plastic limit, low in-situ bulk density and greatly natural void ratio.

101

102 2.2 Sample preparation

The samples were prepared in terms of different initial water contents. Distilled water was first added to oven dried samples to reach the desired water contents and the samples were stirred evenly. Then, the samples were sealed by thin plastic film and stored for 24 hours in an air-conditioned room such that the sample has uniform distribution of moisture. After that, the prepared samples were placed in the shear box using moist tamping method (Finno et al., 1997).

This method has the advantage that high void ratio specimen can be easily
achieved (Finno et al., 1997). The samples were placed in three layers, and then
each layer was damped to obtain the designed densities.

112

113 2.3 Testing apparatus

114 The ring shear apparatus has been widely used in examining the residual 115 shear strength of soils for the analysis of slope stability (Bishop et al., 1971; 116 Bromhead, 1979; Sassa et al., 2004; Wang and Sassa, 2009; Wang et al., 2010). 117 The ring shear apparatus employed in the present research is the fifth version 118 (DPRI-5), which was developed by Disaster Prevention Research Institute (DPRI), Kyoto University (Sassa et al., 2004). The apparatus has a shear box 119 120 with 120 mm inner diameter, 180 mm outer diameter, 115 mm height. The apparatus allows shear tests on soils under drained or completely undrainded 121 122 condition. The apparatus has also a special structure that prevents soil or water 123 leakage during long shear displacement.

The schematic of the ring-shear apparatus is shown in Fig 3. The overview of the apparatus is shown in Fig. 3a. The shear mode of a sample in the torque-controlled ring-shear apparatus was shown conceptually in Fig. 3b. The sample in the ring-shear box is laterally confined between pairs of doughnut-shaped upper and lower confining rings. During the test, the sample is

loaded normally through an annular loading platen connected to a load piston. The lower half of the shear box rotates in both directions, driven by a servomotor through a transmission system, while the upper part is kept steady by means of two retaining torque arms. Fig. 3c illustrates enlarged diagram of half of the cross section the undrained ring-shear box and the pore-water pressure measurement system. The detailed information on ring shear tests can be found in relevant literatures (Wang and Sassa, 2002; Sassa et al., 2003).

136

137 2.4 Testing program and procedure

138 To measure the cohesion and friction angle at different water contents, 7 tests (T₁-T₇) were performed for unsaturated loess samples of different water 139 contents at four normal stresses (50, 100, 150, 250 kpa) by multistage testing 140 procedure (Stark and Eid, 1994). Another test (T₈) was performed for saturated 141 142 loess sample at normal stress of 250 kpa by single-stage testing procedure (Tiwari and Marui, 2004). Choosing such the two testing procedures is to ensure 143 least variation of the samples and to obtain identical properties related to these 144 145 samples. The testing procedures will be described in detail in the following 146 section.

147 To examine whether shear displacement during multistage shearing has 148 impact on shear resistances of soil, 3 ancillary tests (T₉-T₁₁) were performed by

individual shear tests for unsaturated samples with the same water content at a
given normal stress under different shear displacements (10, 30, 40 cm). The
tests are similar to the single-stage testing for saturated samples. Another 3
ancillary tests (T₁₂-T₁₄) were conducted by single-stage testing procedure at
different normal stresses for saturated samples. These tests aim at achieving
void ratio change before and after shearing for a complete analysis.

To perform multistage testing for unsaturated samples, each sample was consolidated at initial normal stress of 50 kpa, and then was sheared at a shear rate of 0.1 mm/s up to steady state at 10 cm shear displacement under drained condition. Then the procedure was repeated on the same sample for the normal stresses of 100, 150 and 250 kpa to obtain cohesion and angle friction of the sample.

161 To perform single-stage testing for saturated samples, each sample was first 162 saturated with the help of carbon dioxide and de-aired water. The degree of 163 saturation was checked by using the B_D parameter proposed by Sassa (1985), in 164 this study, the samples was fully saturated ($B_D = 1.0$). The saturated sample was 165 consolidated at normal stress of 250 kpa, and was sheared at shear rate of 0.1 mm/s up to steady state, i.e., until pore pressure and shear resistance remained 166 constant. To obtain the friction angle and cohesion of the saturated sample, the 167 residual failure line (R.F.L.) was measured after the undrained shear test. After 168

169 the undrained shearing testing was completed, the upper drainage line was switched to a drainage condition so that the generated porewater pressure could 170 171 dissipate, while the lower part of the ring shear apparatus was kept rotating at a small constant speed (0.1 mm/s). The stress shifted from Point SSP to Point 172 RS1, where the generated pore-water pressure dissipated completely. 173 Thereafter, the loaded normal stress was reduced to a small value at a very slow 174 175 rate (0.05 kPa/s) after shearing to ensure the drained condition, while the shear 176 resistance was measured. The stress moved from RS1 to RS2. The line 177 between RS1 and RS2 shows the residual failure line (R.F.L.) of this sample. 178 The measured shear resistance can be regarded as the ultimate shear strength for this sample at the applied normal stress (Wang et al., 2007). 179

To analyze void ratio after consolidation (before shearing) and after shearing, sample height during shearing was recorded concurrently with progress of shear deformation for all of the tests. Additionally, the initial water content was calibrated before each sample was placed into shear box. After shearing was completed, the shear box was opened, and samples were taken from three different layers (i.e., below, within and above the shear zone), and then their water content were measured.

187

188 **3 Results**

Because of the imposed sample preparation and testing procedure, the only difference is water content to the unsaturated samples (T₁-T₇), apart from the tests (T₉-T₁₁) being used to examine the effect shear displacement on shear resistance. To the saturated samples (T₈, T₁₂-T₁₄), the testing procedure is different from that of unsaturated samples. Hence, we present one example of the unsaturated samples and of the saturated samples, respectively. Furthermore, the results of the tests (T₉-T₁₁) were presented completely.

196

197 3.1 Effect of water content on shear strength

To exemplify the shear strength behavior at different water contents, results of unsaturated sample (T₄) and saturated sample (T₈) are presented in Fig. 4 and 5, respectively.

Fig. 4 shows that the results of multistage shear testing on the unsaturated sample with an initial water content of 12.98% (T₄). Fig. 4a shows that variation of shear resistance with progress of shear displacement at four normal stresses. Fig. 4b shows that the variation of sample height with progress of shear displacement. In Fig. 4 it can be seen that the shear resistance maintain a steady state and the sample height decreased gradually with increasing shear displacement.

208 Fig. 5 shows that the results of single-stage testing on saturated sample (T_8) . Fig. 5a shows the variation of normal stress, shear resistance, and pore 209 210 pressure with progress of shear displacement. It is found that some pore pressure was built up before peak shear strength, while after failure, pore 211 pressure showed a sharp increase, and the shear resistance decreased 212 213 gradually with progress of shear displacement and finally reached a constant, 214 i.e., constant volume, constant normal effective stress and constant velocity 215 (Poulos, 1981). Fig. 5b shows the effective stress path and residual failure line. 216 The measured shear resistance was regarded as the ultimate shear strength for the saturated sample, as suggested by Wang et al. (2007). 217

Fig.6 summarizes the variation of residual shear strength with normal stress (i.e., failure envelope) on the tests (T₁-T₈), ranging from 2.97% to 25.03% in water content on shear zone. The slope and the intercept of the failure envelope denoted residual friction angle (φ) and cohesion (*C*) for each sample, respectively.

223

3.2 Effect of shear displacement on shear strength

To examine the effect of testing procedure on shear strength behavior, the tree tests (T₉-T₁₁) were conducted at almost the same initial water content (15.59% \pm 0.03) and the same normal stress of 250 kpa. Fig. 7 illustrates the

228 results of variation of stress ratio with different progress of shear displacements (10, 30, 40 cm). As shown in Fig. 7, there is very light increase in stress ratio with 229 230 progress of shear displacement, but the effect of the increase was negligible to define cohesion and friction angle of sample in our tests. The results show that 231 multistage testing can produce results similar to sing-stage testing, and can 232 obtain identical properties of sample after test. This finding is consistent with 233 234 those observed in clay tests performed by other authors (Stark and Eid, 1994; 235 Tiwari and Marui, 2004).

236

237 3.3 Effect of water content and shearing on void ratio

Fig. 8 shows the variation in void ratio with water content of the samples 238 (T₁-T₈) at four different normal stresses. Fig. 8a plots void ratio after 239 240 consolidation (before shearing) against the initial water content before sample 241 was placed into shear box. Fig. 8b plots void ratio after shearing against the water content on shear zone after shearing was finished. The void ratio of each 242 sample was calculated using the variation of sample height after consolidation 243 244 and shearing (as shown in Fig. 4b) by the formula ($e=G_s/p_d - 1$), in which G_s is specific gravity of the sample, pd was dry density of the sample at a given normal 245 246 stress.

In Fig. 8, it can be seen that the void ratio of sample decrease with water content until certain water content threshold, either before or after shearing, and thereafter it begins to increase and reaches a stable state near saturation. It is interesting to note that the threshold is about the plastic limit of the samples. This indicated that the inconsistent change in void ratio is related to the transition of soil state from plastic to liquid.

By comparing Fig.8a to b, it can be found that the void ratio of the samples was higher before shearing (after consolidation) than after shearing until a certain water content threshold (18.47%), the difference decreasing as the samples are close to saturated condition. It shows that the contraction of the samples occurred during shearing.

The comparison also showed that the water content after shearing increased with the reduction of void ratio and is higher than that before test until the water content threshold (18.47%), after which the water content after test decreased with increasing void ratio and is smaller than that of before shearing. This result revealed that the change in water content was accompanied by change in void ratio during shearing.

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265 3.4 Shearing-induced change within sample

266 Change of the samples within shear box, here concerning shear zone and water content, were observed and measured for each test. Fig. 9 shows 267 example of shear zone formed in our tests and a sketch of water content test. 268 Observation of shear deformation of the samples showed that an annular shear 269 270 zone with a lens-shaped cross section was formed in each test, as shown in Fig. 9a and b. The observation showed that the shear deformation in ring shear tests 271 272 was localized in the shear zone. This phenomenon is consistent with that 273 observed in sand ring-shear tests performed by Wang and Sassa (2002) and 274 Moore and Iverson (2002).

Fig. 9c presents a sketch of water content test for each test. The samples were taken from layers below, within and above the shear zone, respectively. Because the shear zone is thin (as shown in Fig. 9a and b), we collected totally the shear zone, and divided into three parts. The results showed that the water content within sample changed along shear zone during shearing for all of the tests, causing the development of pore pressure gradient in each sample.

281

282 4 Discussions

283 4.1 Volumetric response

In our tests the increase of water content resulted in the samples to contraction until the water content threshold (18.47%), whereas the continued

286 increase of water content resulted in the samples to swell compared to the sample before the water content threshold (as shown in Fig. 8). This process 287 was explained because the soil is of metastable structure (Barden et al., 1973; 288 Fredlund and Morgenstern, 1976). Nevertherless, it was well documented that 289 290 this process has great impact on the changes in physical and mechanical properties of the loess due to volumetric change after wetting (Feda, 1966; 291 292 Šajgalik, 1990; Derbyshire et al., 1994; Dijkstra, 2001; Kruse et al., 2007). It has 293 also found that the change in water content on shear zone was accompanied by 294 the volumetric change during shearing (Fig. 8). To this kind of phenomenon, 295 Terzaghi et al (1996) have pointed out that it is related to change in density of soil during shearing, and this can modify shear resistance. Clearly, volumetric 296 response of soil during wetting and shearing is of great importance to 297 298 understand variations of shear strength parameters and water migration in our 299 tests.

300

301 4.2 Cohesion and friction angle

The effect of water content on cohesion (*C*) and friction angle (φ) is illustrated in Fig. 10, where *C* and φ are plotted against the water content on shear zone for the tests (T₁-T₈). As shown, with an increase in water content, *C* increased and φ decreased until the water content threshold (18.47%), and thereafter, *C*

decreased quickly to zero and φ increased lightly and then stabilized at a 306 residual value. This phenomenon is similar to that obtained in ring shear tests on 307 308 the Lanzhou loess performed by Derbyshire et al. (1994) and Dijkstra et al. (1994). They explained that this phenomenon is related to the bonds between 309 particles influenced by thickness change of water molecules surrounding 310 particles when the water content increases. However, the nature of bands 311 312 between particles is extremely complex (Barden et al., 1973; Terzaghi et al., 313 1996; Craig, 2004). Here, this phenomenon shown in Fig. 10 was explained in 314 terms of the volumetric change during wetting as follows.

315 The friction angle should represent only the contribution of physical bonding to shear resistance for uncemented soil (Terzaghi et al., 1996). This means that 316 the friction angle of uncemented soil is only dependent on the shear resistance 317 318 of soil, which in turn depends on effective normal stress (Terzaghi et al., 1996; 319 Craig, 2004). Therefore, the friction angel of uncemented soil should be a function of shear resistance or effective normal stress. In our tests the samples 320 are not related to cementation because the tests were conducted on completely 321 322 remolded samples (Derbyshire et al., 1994; Dijkstra et al., 1994). It is has well 323 known that the increase of water content and the reduction of void ratio will result in the increase of the pore water pressure, causing the reduction of effective 324 normal stress (Barden et al., 1973; Terzaghi et al., 1996; Craig, 2004). Therefore, 325

in our tests the friction angle decreased gradually with decreasing effective normal stress when the water content is continuously close to the threshold (18.47%), after which the friction angel retained a stable value as the sample reached constant, i.e. constant volume and constant effective normal stress. The process is consistent with change in residual shear resistance with increasing water content, as shown in Fig. 11.

332 The cohesion of soil is attributed to negative pore water pressure within void 333 space very small in size between particles (Craig, 2004). The negative pore 334 water pressure is also referred to as soil suction that mainly depends on pore 335 size and water content (Fredlurid and Rahurcjo, 1993; Vanapalli et al., 1996; Craig, 2004). It was well documented that the suction will increase with 336 decreasing void ratio and water content (Fredlurid and Rahurcjo, 1993; Craig, 337 2004). In our tests, this appears to have a paradox that the change in suction 338 339 with the increase of water content and the reduction of void ratio. However, the previous research has shown that the suction of silt with metastable structure 340 increase with decreasing void ratio, even though the water content was always 341 342 an increasing variable before a critical value (Matyas and Radhakrishna, 1968; Fredlund and Morgenstern, 1976). Furthermore, it was well documented that the 343 Lanzhou loess is typically metastable structure silt (Derbyshire et al., 1994; 344 Dijkstra et al., 1994; Derbyshire, 2001; Dijkstra, 2001; Kruse et al., 2007). 345

346 Therefore, we concluded that the change in cohesion of the samples is related to 347 suction of soil, although the suction was not measured in our tests. Before the water content threshold (18.47%), the cohesion increased due to the increase of 348 suction, because a reduction in void ratio resulted in a reduction in pore size 349 350 between particles, causing the increase of suction. After the water content threshold (18.47%), metastable structure of soil was destroyed when the water 351 352 content increase continuously, and then an increase in void ratio caused an 353 increase in pore size between particles, leading to the very limited suction and 354 high pore-water pressure, and as a result, the cohesion deceased rapidly to 355 zero.

356

357 4.3 Water migration

358 The water content change within samples for the tests (T₁-T₈) was illustrated 359 Fig. 12 for a better comparison, although the phenomenon of water migration occurred in all of the tests. In Fig. 12, we found two different types of water 360 migration on shear zone. The first one is that water migrate towards the shear 361 362 zone from the soil layers below and above the shear zone, as the water content is smaller than 18.47%, and thereafter, the second one occurs, is that water 363 migrate outwards the shear zone into the soil layers below and above the shear 364 zone. The comparison of void ratio and water content before and after shearing 365

(as shown in Fig. 8) has shown that the water migration, which is related to
 volumetric change during shearing, depends on the pre-shear water content.

In our tests, the volumetric change during shearing may be localized on the 368 369 shear zone, or is at least very much stronger on shear zone than on soil layers 370 above and below the shear zone, because the shear deformation was localized on the shear zone in ring shear tests (Fig. 9a and b). In other words, there is a 371 372 distinct difference between shear zone and others position within samples due to 373 the localization of shear deformation. The kind of difference has been observed 374 from special attention to shear zone in ring shear tests (Wang and Sassa, 2002; 375 Wafid Agung et al., 2004; Fukuoka et al., 2006; Wang and Sassa, 2009). The 376 following explanation to water migration will be focus on the difference.

To the first type, the void ratio of shear zone is lower than that of the soil layers above and below the shear zone due to localization of shear deformation. As a result the suction is higher on shear zone than on soil layers above and below the shear zone, because the suction increase with decreasing void ratio at a given condition (Fredlurid and Rahurcjo, 1993). Therefore, the water migrated towards the shear zone during shearing as a result of pore-volume change along shear zone.

To the second type, the suction of soil is very limited due to high saturation (Fredlurid and Rahurcjo, 1993; Vanapalli et al., 1999), although the void ratio is

still lower on shear zone than in the soil layers above and below the shear zone.
However, Terzaghi et al. (1996) have suggested that in saturated soils density
change during shearing is achieved by expelling or by taking in water. Therefore,
the reduction of water content on shear zone is related to the reduction of void
ratio along shear zone. The results in our tests showed that this kind of
phenomena also occur near to saturated soils.

392 It has been reported that the evident reduction of shear strength on 393 overconsolidated clay samples was attributed to the light increase of water 394 content on shear zone in undrained triaxial tests (Bishop, 1961; Skempton and 395 La Rochelle, 1965), whereas other studies have shown that the reduction of water content on shear zone can increase the shear strength on saturated 396 sensitive clay samples in undrained triaxial tests (Taylor, 1951; Crawford, 1961). 397 The above results, along with our test results showed that the water migration on 398 399 shear zone is different for different types of soil, even though the sample were all 400 located in saturated condition. However, it is clear that the water migration has great impact on shear strength of soil, and that the effect is dependent on the 401 402 type of water migration on shear zone. Furthermore, the magnitude of the effect of water migration on shear strength may have difference to different soils. The 403 effect of water migration on shear strength can be concluded as follows, 404 405 although in our tests concurrent changes in water migration and mobilized shear

strength could not be measured. If the water migrate towards shear zone, shear strength on shear zone may be decease due to the reduction of suction or the increase of pore pressure. If the water migrate outwards shear zone, the shear strength may be increase due to the reduction of water content on shear zone.

410

411 5 Conclusions

A series of tests were conducted on loess to examine their shear behavior of using a ring shear apparatus. Based on test results presented in this paper, the following conclusions can be drawn.

(1) The increase of water content caused the reduction of void ratio of the samples until the water content threshold of 18.47%, and thereafter, the continued increase of water content caused the light increase of void ratio then reached a stable state. Furthermore, contraction during shearing occurred in all of the tests due to the reduction of void ratio.

(2) The cohesion increased and the friction angle decreased with increasing
water content until the water content threshold of 18.47%, after which the
cohesion decreased quickly to zero and the friction angle increased lightly and
then stabilized at a residual value.

(3) The localization of shear zone was accompanied by shear deformation forall of the tests. The two types of water migration on shear zone occurred in our

tests. The water was sucked towards shear zone from soil layers above and
below the shear zone when the water content is smaller than 18.47%, whereas
thereafter the water was expelled shear zone into soil layers above and below
the shear zone.

(4) The effect of water migration on shear strength of soil depends on the
type of water migration on shear zone. The water migration may play a negative
role in decreasing shear strength when water migrate towards shear zone,
whereas it may play a positive role in increasing shear strength when water
migrate outwards shear zone.

(5) There was a same water content threshold for changes in void ratio,
strength parameters and water migration. The threshold is close to the plastic
limit of the samples. All of the changes revealed the dependence of the shear
behavior and the water migration on soil volumetric response during wetting and
shearing.

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446 **Reference**:

447 Abduljauwad, S.N., Al-Gassous, K.A., 1991. Soil deformation and shear strength characteristics of the

448 Sana'a soil, Yemen Arab Republic. Engineering Geology 31, 291-314.

- 449 Audric, T., Bouquier, L., 1976. Collapsing behaviour of some loess soils from Normandy. Quarterly
- 450 Journal of Engineering Geology and Hydrogeology 9, 265-277.
- 451 Barden, L., McGown, A., Collins, K., 1973. The collapse mechanism in partly saturated soil. Engineering 452 Geology 7, 49-60.
- Bishop, A.W., 1961. Discussion on soil properties and their measurement. Proc. 5th Int. Conf. Soil Mech.,
 Paris, pp. 89-93.
- 455 Bishop, A.W., Green, G.E., Garga, V.K., Andresen, A., Brown, J.D., 1971. A new ring shear apparatus and
- 456 its application to the measurement of residual strength. Géotechnique 21, 273-328.
- 457 Bromhead, E.N., 1979. A simple ring shear apparatus. Ground Engineering 12, 40-44.
- 458 Craig, R.F., 2004. Craig's Soil Mechanics, 7th Edition. Taylor & Francis Group.
- 459 Crawford, C.B., 1961. The influence of strain on shearing resistance of sensitive clay. ASTM 61,460 1250-1265.
- 461 Derbyshire, E., 2001. Geological hazards in loess terrain, with particular reference to the loess regions of
 462 China. Earth-Science Reviews 54, 231-260.
- Derbyshire, E., Dijkstra, T.A., Smalley, I.J., Li, Y., 1994. Failure mechanisms in loess and the effects of
 moisture content changes on remoulded strength. Quaternary International 24, 5-15.
- Derbyshire, E., Wang, J., Jin, Z., Billard, A., Egels, Y., Kasser, M., Jones, D.K.C., Muxart, T., Owen, L., 1991.
 Landslides in the Gansu loess of China. Catena Supplement 20, 119-145.
- 467 Dijkstra, T.A., 2001. Geotechnical thresholds in the Lanzhou loess of China. Quaternary International468 76-77, 21-28.
- 469Dijkstra, T.A., Rogers, C.D.F., Smalley, I.J., Derbyshire, E., Li, Y.J., Meng, X.M., 1994. The loess of470north-central China: Geotechnical properties and their relation to slope stability. Engineering Geology 36,
- 471 153-171.
- Feda, J., 1966. Structural stability of subsident loess soil from Praha-Dejvice. Engineering Geology 1,201-219.
- Finno, R.J., Harris, W.W., Mooney, M.A., Viggiani, G., 1997. Shear bands in plane strain compression of
 loose sand. Geotechnique 47, 149-165.
- 476 Fredlund, D.G., Morgenstern, N.R., 1976. Constitutive relations for volume change in unsaturated soils.
- 477 Canadian Geotechnical Journal 13, 261–276.
- 478 Fredlurid, D.G., Rahurcjo, H., 1993. Soil Mechanics for Unsaturated Soils. John Wiley & Sons, Inc.
- 479 Fukuoka, H., Sassa, K., Wang, G., Sasaki, R., 2006. Observation of shear zone development in ring-shear
 480 apparatus with a transparent shear box. Landslides 3, 239-251.
- 481 Gibbs, H.J., Holland, W.Y., 1960. Petrographic and engineering properties of loess. U.S. Bureau of
- 482 Reclamation, Denver. Engineering Monographs 28, 1-37.
- 483 Ishihara, K., Okusa, S., Oyagi, N., Ischuk, A., 1990. Liquefaction- induced flowslide in the collapsible loess
- 484 deposit in Soviet Tajik. Soils and Foundations 30, 73-89.

- 485 Kruse, G.A.M., Dijkstra, T.A., Schokking, F., 2007. Effects of soil structure on soil behaviour: Illustrated
- 486 with loess, glacially loaded clay and simulated flaser bedding examples. Engineering Geology 91, 34-45.
- 487 Long, J., Li, T., Lei, X., Yang, S., 2007. Study on physical properties of soil in sliding zone of loess landslide.

488 Chinese Journal of Geotechnical Engineering 29, 289-293. (in Chinese).

- 489 Lutenegger, A.J., 1981. Stability of loess in light of the inactive particle theory. Nature 291, 360-360.
- 490 Lutenegger, A.J., Hallberg, G.R., 1988. Stability of loess. Engineering Geology 25, 247-261.
- 491 Matyas, E.L., Radhakrishna, H.S., 1968. Volume change characteristics of partially saturated soil.
 492 Géotechnique 18, 432-448.
- 493 Moore, P.L., Iverson, N.R., 2002. Slow episodic shear of granular materials regulated by dilatant 494 strengthening. Geology 30, 843-846.
- Phien-wej, N., Pientong, T., Balasubramaniam, A.S., 1992. Collapse and strength characteristics of loessin Thailand. Engineering Geology 32, 59-72.
- 497 Poulos, S.J., 1981. The steady state of deformation. Journal of the Geotechnical Engineering Division,498 ASCE 107, 553-562.
- RenéJacques, B., 1988. Some specific problems of wetted loessial soils in civil engineering. EngineeringGeology 25, 303-324.
- 501 Šajgalik, J., 1990. Sagging of loesses and its problems. Quaternary International 7-8, 63-70.
- 502 Sassa, K., 1985. The mechanism of debris flows. In Proceedings of the 11th International Conference on 503 Soil Mechanics and Foundation Engineering, San Francisco, Calif. Vol. 3, pp. 1173–1176.
- 504 Sassa, K., Fukuoka, H., Wang, G., Ishikawa, N., 2004. Undrained dynamic-loading ring-shear apparatus 505 and its application to landslide dynamics. Landslides 1, 7-19.
- 506 Sassa, K., Wang, G., Fukuoka, H., 2003. Performing undrained shear tests on saturated sands in a new 507 intelligent type of ring shear apparatus. Geotechnical Testing Journal 26, 257-265.
- 508 Skempton, A.W., La Rochelle, P., 1965. The Bradwell slip: A short-term failure in London Clay. 509 Géotechnique 15, 221-242.
- 510 Stark, T.D., Eid, H.T., 1994. Drained Residual Strength of Cohesive Soils. Journal of geotechnical 511 engineering 120, 856-871.
- 512 Tan, T.K., 1988. Fundamental properties of loess from Northwestern China. Engineering Geology 25,513 103-122.
- Taylor, D.W., 1951. Research on shearing resistance of clay: Sections on water migration studies. U.S.
 Army Engineers Waterways Experiment Station Report 36-42, 9-12.
- 516 Terzaghi, K., Peck, R.B., Mesri, G., 1996. Soil Mechanics in Engineering Practice, 3rd Edition. John Wiley &
 517 Sons, Inc.
- 518 Tiwari, B., Marui, H., 2004. Objective Oriented Multistage Ring Shear Test for Shear Strength of 519 Landslide Soil. Journal of Geotechnical and Geoenvironmental Engineering 130, 217-222.
- 520 Ueno, M., Nishimura, T., Kato, M., Nakamura, H., Zeng, S., 2005. Variation of shearing characteristics of
 521 loess soil after irrigation. Northwestern Seismological Journal 27, 128-134.
- 522 Vanapalli, S.K., Fredlund, D.G., Pufahl, D.E., 1996. The Relationship Between the Soil-Water
- 523 Characteristic Curve and the Unsaturated Shear Strength of a Compacted Glacial Till. Geotechnical
- 524 Testing Journal 19, 259-268.

- Vanapalli, S.K., Fredlund, D.G., Pufahl, D.E., 1999. The influence of soil structure and stress history on the
 soil-water characteristics of a compacted till. Géotechnique 49, 143-159.
- 527 Wafid Agung, M., Sassa, K., Fukuoka, H., Wang, G., 2004. Evolution of Shear-Zone Structure in Undrained
- 528 Ring-Shear Tests. Landslides 1, 101-112.
- Wang, G., Sassa, K., 2002. Post-failure mobility of saturated sands in undrained load-controlled ring
 shear tests. Canadian Geotechnical Journal 39, 821-837.
- Wang, G., Sassa, K., 2009. Seismic loading impacts on excess pore-water pressure maintain landslide
 triggered flowslides. Earth Surface Processes and Landforms 34, 232-241.
- Wang, G., Sassa, K., Fukuoka, H., Tada, T., 2007. Experimental Study on the Shearing Behavior of
 Saturated Silty Soils Based on Ring-Shear Tests. Journal of Geotechnical and Geoenvironmental
 Engineering 133, 319-333.
- 536 Wang, G., Suemine, A., Schulz, W.H., 2010. Shear-rate-dependent strength control on the dynamics of
- rainfall-triggered landslides, Tokushima Prefecture, Japan. Earth Surface Processes and Landforms 35,407-416.
- 539 Zhang, D., Wang, G., 2007. Study of the 1920 Haiyuan earthquake-induced landslides in loess (China).
- 540 Engineering Geology 94, 76-88.
- 541 Zhang, D., Wang, G., Luo, C., Chen, J., Zhou, Y., 2009. A rapid loess flowslide triggered by irrigation in
- 542 China. Landslides 6, 55-60.
- 543
- 544

Table captions:

Table 1. Basic physical properties of loess used in this study

Figure captions:

Fig.1 Location of study area and sampling site

Fig. 2 Grain size distributions curve of loess used in this study

Fig. 3 Ring-shear apparatus DPRI-Ver.5. (a) Overview of the ring shear apparatus with a transparent shear box; (b) Sample in ring-shear apparatus; (c) Cross section through center of undrained shear box of ring-shear apparatus.

Fig. 4 Example of Drained shear test on the unsaturated sample. (a) Shear resistance against shear displacement at different normal stresses. (b) Sample height against shear displacement at different normal stresses.

Fig. 5 Undrained shear test on the saturated sample (T_8). (a) Variation of normal stress, pore pressure, and shear resistance with shear displacement. (b)

Residual failure line and effective stress path.

Fig. 6 Residual shear strength against normal stress at different water contents one shear zone for the test (T_1-T_8) .

Fig. 7 Stress ratio against shear displacement on samples with the same water content at a given normal stress.

Fig. 8 Void ratio against water content at different normal stresses. (a) Void ration against initial water content after consolidation (before shearing). (b) Void

ration against final water content on shear zone after shearing at each shear steps.

Fig. 9 Example of shear zones on the sample T4. (a) Formation of the shear zone after shearing; (b) Change in basic properties of the shear zone (c) Sketch map of water content test at different soil layers.

Fig. 10 Cohesion and residual friction angle at different water contents.

Fig. 11 Shear resistance against water content at different normal stresses.

Fig. 12 Water content of samples at different soil layers.

Property (definition and method)	Loess
Specific gravity (Gs)	2.71
Initial moist bulk density (g/cm ³)	1.50
Initial water content (%)	7.40
Initial void ratio	0.92
Liquid limit (%)	27.74
Plastic limit (%)	17.68
Plasticity index (%)	10.06

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Fig. 12 Water content of samples at different soil layers.

Note: I: initial water content before test; U, L: water content of soil layer above and below the shear zone; SZ: water content of soil layer on the shear zone; A: average water content of sample after test