# European soil NOx emissions derived from satellite NO2 observations

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#### Abstract

We introduce an innovative method to distinguish soil nitrogen oxides (NOx=NO+NO2) emissions from satellite-based total NOx emissions using its seasonal characteristics. To evaluate the approach, we compare the deviation between the tropospheric NO2 concentration observed by satellite and two atmospheric composition model simulations driven by the newly estimated soil NOx emissions and the Copernicus Atmosphere Monitoring Service (CAMS) inventory. The estimated average soil NOx emissions in Europe are 2.5 kg N ha-1 yr-1 in 2019, and the annual soil NOx emissions is approximately 2.5 times larger than that of the CAMS inventory. Our method can easily be extended to other regions at middle or high latitudes with similar seasonal characteristics of soil emissions. The soil emissions are subtracted from the total NOx emissions using known CO2/NOx factors from bottom-up inventories.

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## European soil NOx emissions derived from satellite NO2 observations

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## 10 Key Points:

- An innovative method is introduced to derive soil NO<sub>x</sub> emissions in Europe from satellite NO<sub>2</sub> observations.
   The resulting soil NO<sub>x</sub> emissions are at least two times larger than widely used bottom up
- The resulting soil NO<sub>x</sub> emissions are at least two times larger than widely used bottom-up soil NO<sub>x</sub> emission estimates.
- This satellite observation-based method provides a valuable independent estimate of the soil NO<sub>x</sub> emissions.
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#### 18 Abstract

We introduce an innovative method to distinguish soil nitrogen oxides ( $NO_x = NO + NO_2$ ) 19 emissions from satellite-based total NO<sub>x</sub> emissions using its seasonal characteristics. To evaluate 20 the approach, we compare the deviation between the tropospheric NO<sub>2</sub> concentration observed 21 by satellite and two atmospheric composition model simulations driven by the newly estimated 22 23 soil NO<sub>x</sub> emissions and the Copernicus Atmosphere Monitoring Service (CAMS) inventory. The estimated average soil NO<sub>x</sub> emissions in Europe are 2.5 kg N ha<sup>-1</sup> yr<sup>-1</sup> in 2019, and the annual 24 soil NO<sub>x</sub> emissions is approximately 2.5 times larger than that of the CAMS inventory. Our 25 method can easily be extended to other regions at middle or high latitudes with similar seasonal 26 characteristics of soil emissions. The soil emissions are subtracted from the total  $NO_x$  emissions 27 yielding realistic anthropogenic  $NO_x$  emissions. We further show this also yields realistic 28 29 anthropogenic  $CO_2$  emissions using known  $CO_2/NO_x$  factors from bottom-up inventories.

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#### 31 Plain Language Summary

32 Soil nitrogen oxide emissions ( $NO_x = NO + NO_2$ ) are an important source of air pollution, accounting for about 15% of global NO<sub>x</sub> emissions. Unfortunately, soil emissions are not always 33 accurately described by current bottom-up inventories. Accurate quantification is beneficial for 34 clarifying the contribution of biogenic sources to air quality and developing more targeted air 35 quality measures. We present an innovative method for estimating soil NO<sub>x</sub> emissions from 36 satellite-based total NO<sub>x</sub> emissions. The newly estimated annual emissions in Europe are about 37 38 2.5 times higher than reported in previous studies. The method is evaluated by comparing the deviation between the simulated and satellite observed tropospheric NO<sub>2</sub> concentrations. This 39 method can also be extended to other regions around the world with similar seasonal 40 characteristics of soil NO<sub>x</sub> emissions. Anthropogenic NO<sub>x</sub> emissions are determined by 41 subtracting the soil  $NO_x$  emissions from total  $NO_x$  emissions. We further show these 42 anthropogenic  $NO_x$  emissions can be converted into realistic  $CO_2$  emissions by using known 43 CO<sub>2</sub>/NO<sub>x</sub> emission factors. 44

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## 46 **1 Introduction**

Nitrogen oxides ( $NO_x = NO + NO_2$ ) are important pollutants and their subsequent oxidation 47 48 products have detrimental impacts on human health and crop production (Skalska et al., 2010). Soil NO<sub>x</sub> emissions are the largest contributor to the NO<sub>x</sub> budget besides combustion sources, 49 contributing up to ~15% of global NO<sub>x</sub> emissions (Hudman et al., 2012; Vinken et al., 2014; 50 51 Weng et al., 2020). The relative contribution of soil  $NO_x$  to total  $NO_x$  emissions is gradually increasing due to steadily declining anthropogenic NO<sub>x</sub> emissions as a result of successful 52 emission reduction strategies in, e.g., China (van der A et al., 2017; Lu et al., 2021), the USA 53 (Zhang et al., 2003; Silvern et al., 2019), and Europe (Rafaj et al., 2015; Skiba et al., 2020). 54 Furthermore, soil NO<sub>x</sub> emissions play a non-negligible role in rural air pollution especially 55 during summer time while fossil fuel combustion emissions are relatively constant over the year 56 57 (Fortems-Cheiney et al., 2021; Wang et al., 2022). The precise quantification of soil NO<sub>x</sub> emissions is therefore essential for assessing emission control strategies and a better 58 understanding of air quality. 59

Two microbial processes, nitrification and denitrification, are the main sources of soil  $NO_x$  and 60 they occur in agricultural and natural ecosystems (Hall et al., 1996; Pilegaard, 2013). Key factors 61 that regulate NO<sub>x</sub> emissions from soil are: temperature, soil moisture and texture, soil pH, 62 nutrient availability, ecosystem types, agricultural management and ambient atmospheric NO<sub>x</sub> 63 concentration (Hall et al., 1996; Butterbach-Bahl et al., 2013; Medinets et al., 2015). Chamber 64 studies and field measurements are commonly employed to investigate the response of soil NO<sub>x</sub> 65 emissions to rewetting of dry soils (Garcia-Montiel et al., 2003; Hickman et al., 2021), fertilizer-66 induced change (Liu et al., 2017; Song et al., 2020; Hui et al., 2023) and atmospheric deposition 67 (Hall and Matson, 1999; Venterea et al., 2003; Koehler et al., 2009; Eickenscheidt and Brumme, 68 2012). Global and regional soil NO<sub>x</sub> emissions are generally estimated by three different model-69 based methods: simple scaling (Davidson and Kingerlee, 1997), empirical models (Yienger and 70 Levy II, 1995; Yan et al., 2005; Weng et al., 2020; Simpson and Darras, 2021) and process-71 oriented models (Butterbach-Bahl et al., 2009; Molina-Herrera et al., 2017). However, these 72 models in general disagree about the soil  $NO_x$  quantities and their spatial patterns. 73

74 Satellite-based observations provide an alternative method to derive soil  $NO_x$  emissions. Bertram et al. (2005) and Zörner et al. (2016) found that SCIAMACHY (Scanning Imaging Absorption 75 spectroMeter for Atmospheric CHartographY) observations captured the brief, high-intensity soil 76 NO<sub>x</sub> pulses in response to fertilizer application or rainfall events in agricultural regions and semi-77 arid ecosystems. Other studies constrained soil  $NO_x$  emissions top-down using retrieved  $NO_2$ 78 vertical column densities (VCDs) from the Ozone Monitoring Instrument (OMI) for East China 79 80 (Lin, 2012) and globally (Vinken et al., 2014). Huber et al. (2020) used the unprecedented spatiotemporal resolution of the TROPOMI NO<sub>2</sub> product to quantify soil-driven contributions of 81 cropland to regional NO<sub>x</sub> emissions by a box model on daily to seasonal scales for the U.S. 82 Southern Mississippi River Valley. Furthermore, other studies estimate NO<sub>x</sub> emissions by 83 analyzing the relationship between observed NO<sub>2</sub> concentrations and NO<sub>x</sub> emissions with 84 inversion techniques that consider the transport process of NO<sub>x</sub> (Mijling and van der A, 2012; 85 86 Miyazaki et al., 2012). However, such methods estimate only total NO<sub>x</sub> emissions, encompassing both natural and anthropogenic sources. 87

88 In this study, we introduce a new method for estimating soil NO<sub>x</sub> emissions in individual grid cells based on its seasonal variations. This method is a post-processing of the total  $NO_x$ 89 emissions derived by the inverse algorithm DECSO (Daily Emission estimation Constrained by 90 Satellite Observations, Mijling and van der A, 2012; Ding et al., 2017a) applied to NO2 91 observations over Europe by TROPOspheric Monitoring Instrument (TROPOMI) on Sentinel 5 92 Precursor (S5-P) satellite. We evaluate the performance of our method by comparing the 93 94 deviation of the tropospheric NO<sub>2</sub> concentrations between atmospheric chemistry model simulations and observations by TROPOMI. Finally, we explore the potential to use the 95 difference between total satellite-derived NO<sub>x</sub> emissions and soil NO<sub>x</sub> emissions for indirectly 96 97 estimating fossil-fuel CO<sub>2</sub> emissions.

### 98 2 Materials and Methods

#### 99 2.1 NO<sub>x</sub> emissions from DECSO

100 NO<sub>x</sub> emissions are derived by the state-of-the-art inverse algorithm DECSO (Daily Emission 101 estimation Constrained by Satellite Observations, Mijling and van der A, 2012; Ding et al., 102 2017a). DECSO is specifically developed for daily updates of emissions of short-lived

atmospheric constituents using satellite observations. The algorithm solves the sensitivity of 103 104 concentrations to emissions using a single forward run of the chemical transport model CHIMERE v2020 (Menut et al., 2021) and a simplified 2D trajectory analysis. An extended 105 Kalman filter is used for assimilation of the observed column concentrations in the inversion step. 106 DECSO is able to provide total emissions from biogenic (originating from soil for NO<sub>x</sub>) and 107 anthropogenic sources for short-lived chemical species and it can detect new emission sources 108 that may be missing in bottom-up inventories. It has been validated (Ding et al., 2017b) and 109 successfully applied to different regions using OMI and TROPOMI observations (Ding et al., 110 2015; Ding et al., 2018; Ding et al., 2020; van der A et al., 2020; Ding et al., 2022). In this study, 111 monthly NO<sub>x</sub> emissions in 2019 over Europe (10°W-30°E, 35-55°N) are derived from 112 TROPOMI NO<sub>2</sub> observations using DECSO on a spatial resolution of  $0.2^{\circ} \times 0.2^{\circ}$ . These total 113 emissions are used as input to isolate soil NO<sub>x</sub> emissions in a post-processing step, which is 114 explained below. 115

- 116 2.2 Soil NO<sub>x</sub> emissions estimates
- 117 Several studies have shown that soil NO<sub>x</sub> emissions are significantly influenced by land use type
- 118 (Valente and Thornton, 1993; Verchot et al., 1999; Yan et al., 2005). The soil emissions in our
- study area originate from four main land use types: forest, croplands, shrub and grassland (Figure
- 120 S1). Here we merged shrub and grassland into one category (called "other biogenic") considering
- 121 their limited occurrence in the study area (Table S1c).
- We use the following five steps (see flow chart in Figure S2) to separate soil  $NO_x$  emissions
- 123 from total  $NO_x$  emissions:

(1) We select pixels dominated by the biogenic sector using the proportion of each land use type.
The minimum thresholds of the three land use ratios (forest, crop, and other biogenic sources)
are set to 0.5 for individual grid cells to make sure the cell is dominated by one of the biogenic
source sector types. For these pixels, the fraction of urban coverage is required to be less than
0.02 to eliminate the interference of anthropogenic emissions as much as possible. The selected
pixels are referred to as biogenic pixels.

- 130 (2) To exclude the remaining anthropogenic emissions in the selected grid cells, we subtract 131 CAMS anthropogenic  $NO_x$  emissions (version 5.3, called CAMS-ant) from the DECSO total 132  $NO_x$  emissions. Note that this is only done for the selected biogenic pixels. If negative values 133 occur after subtraction, they are set to zero. A sensitivity analysis with respect to this step is 134 described in Section 3.1.
- (3) In order to better reflect the spatial heterogeneity of soil emissions, we divide the researcharea equally into 5 subregions in the latitude direction by 2 subregions in the longitude direction.
- 137 In each of these 10 subregions, the average monthly emissions of the selected pixels are fitted
- with a Gaussian function  $f(x) = A e^{\frac{-(x-B)^2}{2c^2}}$  over one year. We chose a Gaussian function as soil NOx emissions in Europe vary slowly with season with typically a winter minimum and summer
- 140 maximum. The fitting parameters A, B, and C are obtained for pixels dominated by each of the
- 141 land use types separately (see step 1). A represents the maximum soil  $NO_x$  emissions in a year, B
- represents the month when the maximum soil emissions occur, and C determines the width of the Gaussian curve and thus the length of the season, which also affects the amount of winter soil
- Gaussian curve and thus the length of the season, which also affects the amount of winter so NO<sub>x</sub> emissions. Examples of the Gaussian fitting can be found in Figure S3.

(4) Since the parameters obtained in step 3 represent soil emissions with a specific land use ratio larger than 0.5 (set in step 1) but still with mixed land use types, we use the solution of formula S2 to obtain the typical parameters of pure pixels, *i.e.* the land use ratio of one of the three types, either forest, crop, or other biogenic sources, equals 1. In this way, we obtain 30 sets of parameters (A, B, and C) representing soil emissions for three land-use types and 10 subregions separately. To smooth the transitions between subregions, we perform a two-dimensional interpolation to obtain the parameters for each land-use type and for each grid cell separately.

(5) We assume that the land use ratio directly determines the proportion of soil  $NO_x$  emissions. The monthly soil emissions per grid cell is calculated by multiplying the ratio of the three land use types by the three Gaussian functions of the corresponding soil emission types, and adding them together.

(6) If the soil emission calculated at a certain grid cell is larger than the total emission of DECSO
 in a certain month, the soil emission of this month is set to be equal to this total emission of
 DECSO. In this way the total of the derived DECSO emissions remain conserved. The end
 product will be called DECSO-soil from here.

Figure S4 shows the three key parameters A, B, and C that depict the seasonal characteristics of 160 soil NO<sub>x</sub> emissions for the three different land use types, with significant zonal and meridional 161 differentiation. The value of parameter A, representing the maximum soil NOx emissions during 162 the year, for forests and croplands are generally similar (Figure S4 a-c). The month of the 163 maximum soil emissions (parameter B) occurs a bit later in forest areas (July - August) than in 164 croplands areas (June - July) (see Figure S4 d-f). The parameter C represents the width of the 165 Gaussian fit and this also affects winter soil NO<sub>x</sub> emissions. For all three land use types, 166 parameter C shows a clear decreasing trend with increasing latitude (Figure S4 g-i). This is 167 because the higher the latitude, the lower the winter temperature, and the lower the microbial 168 activity, resulting in a shorter active season. 169

170 2.3 Emission inventories and land use dataset

In this study, three emission inventories are used for comparison with our estimates. They are the 171 172 CAMS soil emissions inventory (CAMS-GLOB-SOIL version 2.4, henceforth called CAMSsoil), the Harvard-NASA Emissions Component (HEMCO) soil emissions inventory (version 173 2021, called HEMCO-soil) and the National Long-range Transboundary Air Pollution (LRTAP) 174 NO<sub>x</sub> emissions (called LRTAP-NO<sub>x</sub>). CAMS-soil provides gridded global monthly soil NO 175 176 emissions as total values and for separate source sectors at spatial resolution of  $0.5^{\circ} \times 0.5^{\circ}$ . It is based on empirical formulas and process parameter models (Simpson and Darras, 2021). 177 HEMCO-soil provides global hourly soil NO<sub>x</sub> emissions at a horizontal resolution of  $0.25^{\circ}$  lat. × 178 0.3125° lon. (Weng et al., 2020), (Keller et al., 2014). LRTAP-NO<sub>x</sub> provide country level yearly 179 NO<sub>x</sub> emissions for agriculture and other sectors and is provided by the European Environment 180 Agency. Global monthly bottom-up anthropogenic NO<sub>x</sub> (version 5.3, called CAMS-ant) and CO<sub>2</sub> 181 182 emissions (version 4.2, called CAMS-CO<sub>2</sub>) inventories are both obtained from the Copernicus Atmosphere Monitoring Service (CAMS) at a 0.1°×0.1° horizontal resolution (Soulie et al., 183 2023). All emission data are for 2019 and are regridded to the same domain and resolution of 184 DECSO  $(0.2^{\circ} \times 0.2^{\circ})$ . The land use data Land Cover are obtained from the Copernicus Global 185 Land Service (version 3.0.1, Buchhorn et al., 2020). The original 23 land use classes of the Land 186 Cover database were first grouped into 8 new main classes, comprising ocean, urban, cropland, 187

grassland, bare land, inland water, forest, and shrub defined in Table S1. The land use ratio for

each class was calculated by re-gridding the original 100m resolution Land Cover product to the DECSO grid of  $0.2^{\circ}$ 

190 DECSO grid of  $0.2^{\circ}$ .

191 2.4 Evaluation of derived soil emissions by comparing modelled concentrations to satellite

192 observations

We conduct two comparative experiments to simulate tropospheric NO<sub>2</sub> columns, which use 193 either CAMS soil emissions or DECSO soil emissions. We evaluate the performance of the 194 newly estimated soil emissions in this study by comparing the Root Mean Square Error (RMSE) 195 196 between the simulated tropospheric NO<sub>2</sub> concentration and the TROPOMI observed tropospheric NO<sub>2</sub> concentration of these two comparative experiments). The tropospheric NO<sub>2</sub> columns were 197 198 simulated by an extended version of ECMWF's Integrated Forecasting System (IFS) called "IFS-COMPO" (Flemming et al., 2015; Huijnen et al., 2019). IFS-COMPO is part of the global 199 200 component of the Copernicus Atmosphere Monitoring Service (CAMS) and has been employed to supply global analyses and forecasts of atmospheric composition in an operational mode 201 202 starting from 2014. The version of IFS-COMPO employed here is based on IFS CY48R1 (ECMWF, 2023), but with only tropospheric chemistry activated. Its default anthropogenic 203 emissions, based on CAMS-GLOB-ANT v5.3 (Soulie et al., 2023) are adopted. The model is 204 driven by our newly estimated soil NO<sub>x</sub> emissions, and CAMS soil NO<sub>x</sub> emissions (version 2.4, 205 Simpson and Darras, 2021) for reference. IFS-COMPO was run for the year 2019 at a horizontal 206 resolution of approximately 40 km with 137 vertical layers and 900s time steps and with a one-207 208 month spin up period. When we compare TROPOMI NO<sub>2</sub> observations with the IFS-COMPO simulation, only observations with a quality flag above 0.75 are used to avoid retrievals for 209 ground pixels covered with snow, ice or high cloud radiance fraction, as well as problematic 210 retrievals. The model outputs are interpolated to the local overpass time of TROPOMI and the 211 averaging kernel is applied to the modelled NO<sub>2</sub> profile. The collocated observation-model pairs 212 are re-gridded to a regular latitude-longitude grid with a 0.25° resolution using an area-weighted 213 averaging considering the area of the TROPOMI-pixel if the coverage of the grid cell is above 214 50% (Douros et al., 2023). The only difference between the two comparative model experiments 215 is the input of soil NO<sub>x</sub> emissions. 216

### 217 **3 Results**

218 3.1 Comparison of Soil NOx emissions with CAMS

Figure 1 shows the spatial distribution of calculated soil NO<sub>x</sub> emissions for each sector (forests 219 and croplands sectors) in the study area during summer (May-August). The yearly averaged soil 220 NO<sub>x</sub> emissions for the entire domain from forests, croplands, and other biological sources are 2.6, 221 2.6 and 2.0 kg N ha<sup>-1</sup> yr<sup>-1</sup> respectively (in May-August shown in Figure 1 they are on average 3.7, 222 3.6 and 2.9 kg N ha<sup>-1</sup> yr<sup>-1</sup>), which fall within the estimated range of forest emissions (0.35 to 15.9 223 kg N ha<sup>-1</sup> yr<sup>-1</sup> in Saxony of Germany; Molina-Herrera et al. 2017) and are of the same order of 224 magnitude for croplands as estimated by Yan et al. (1.08 kg N ha<sup>-1</sup> yr<sup>-1</sup> globally; 2005). Regions 225 with high CAMS-soil emissions, such as the Castile-León plain in Spain and the Po River plain 226 227 in Italy, display strong similarities with the spatial distribution of DECSO-soil NO<sub>x</sub> emissions of the croplands sector (Figure 1 c-d). Furthermore, the CAMS soil NO<sub>x</sub> emission inventory has 228 very low emissions in forest areas resulting in lower emission estimates in the northwestern 229 Iberian Peninsula, the forest areas of Romania and the south-central France (Figure 1 a-d). Note 230

the high correlation ( $R^2 = 0.53$  in Figure S6) between the DECSO forest emissions (Figure 1a)



and the difference map shown in Figure 1b.

233

Figure 1. The spatial distribution of the derived soil NO<sub>x</sub> emissions during summer represented by the average emissions from May to August in 2019 from (a) forest and (c) croplands. (d) shows the CAMS-soil NO<sub>x</sub> emissions in Europe during summer. The difference between CAMSsoil and DECSO-soil is shown in (b). The soil emissions calculated from DECSO total emissions are regridded to the resolution of CAMS-soil, which is  $0.5^{\circ} \times 0.5^{\circ}$ .

We compared the sum of all DECSO soil to sum of all CAMS soil emissions in our study 239 domain. Our derived total annual soil NO<sub>x</sub> emissions are 1.1 Tg N yr<sup>-1</sup>, which is more than 2.5 240 times larger than the total of CAMS-soil (0.4 Tg N yr<sup>-1</sup>) and about 2.3 times higher than 241 HEMCO (0.5 Tg N yr<sup>-1</sup>) (Figure S7). The average soil NO<sub>x</sub> emissions in the study area are 2.5 kg 242 N ha<sup>-1</sup> yr<sup>-1</sup> in 2019. Figure 2a shows that the obtained typical monthly time profile of soil NO<sub>x</sub> 243 emissions is similar to that of CAMS. The spatial distribution and the amount of the DECSO 244 cropland emissions are comparable to the CAMS soil emissions. CAMS-soil and LRTAP NO<sub>x</sub> 245 emissions from agriculture sector are also consistent for national total numbers (Figure 2b). 246 Furthermore, we found that the discrepancy with CAMS is more significant in countries with a 247 large proportion of forest area, such as the Spain (138 Gg N yr<sup>-1</sup> for DECSO-soil and 48 Gg N yr<sup>-1</sup> 248 <sup>1</sup> for CAMS-soil) and France (130 Gg N yr<sup>-1</sup> for DECSO-soil and 64 Gg N yr<sup>-1</sup> for CAMS-soil). 249 And the deviation is smaller in countries with a large proportion of non-forest area (Figure 2b), 250 such as the Netherlands (about 8 Gg N yr<sup>-1</sup> for both DECSO-soil and CAMS-soil) and Belgium 251 (7 Gg N yr<sup>-1</sup> for DECSO-soil and 5 Gg N yr<sup>-1</sup> for CAMS-soil). Figures 2c and S8 show that after 252 excluding soil emissions, the difference between anthropogenic NO<sub>x</sub> emissions derived with 253 DECSO based on satellite observations and CAMS anthropogenic emissions becomes noticeably 254

## smaller (DECSO anthropogenic NO<sub>x</sub> is 4.9 kg N ha<sup>-1</sup> yr<sup>-1</sup> and CAMS-anthropogenic is 4.8 kg N ha<sup>-1</sup> yr<sup>-1</sup>).



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Figure 2. (a) Monthly comparison of derived soil  $NO_x$  emissions for three land use types with CAMS-soil. The estimated upper limit and the lower limit of emissions as described in below are shown by the dashed line. (b) National soil  $NO_x$  emissions from DECSO-soil and CAMS-soil. (c) The monthly proportion of anthropogenic and soil NOx emissions of DECSO and CAMS. (d) The spatial distributions of DECSO-soil emissions in 2019 during summer (May to August).

263 3.2 Uncertainty analysis

The biggest uncertainty in our method is caused by the correction for anthropogenic emissions in 264 the selected biogenic grid cells (step 2 in Section 2.2). Therefore, we estimated the upper and 265 lower limit of the calculated soil emissions, by performing a sensitivity test. We first assume all 266 selected biogenic grid cells are without remaining anthropogenic emissions, resulting in an upper 267 limit of the derived soil emissions. On the other hand, the lower limit of emissions is obtained by 268 assuming that the emissions of the selected biogenic grid cells are completely anthropogenic in 269 wintertime as biogenic activity is at a minimum in Europe during winter. Thus we replaced the 270 271 anthropogenic emissions of CAMS (used in step 2 of Section 2.2) by the average of the DECSO total emissions in January and December. This results in an upper limit of about 33% higher 272 emissions and a lower limit that is about 14% lower than the calculated DECSO-soil emissions 273 (Figure S5). 274

275 The derived soil  $NO_x$  emissions are sensitive to uncertainties in the derived DECSO emissions.

The DECSO emissions have a precision of about 30% for monthly emissions in a single grid cell.

277 However, for this analysis on average soil emissions, the DECSO emissions are averaged over

pixels over the whole region and thus strongly reduced compared to single grid cells. Therefore,

- the error of the anthropogenic emission correction mentioned above is dominating, and we estimate the uncertainty on the average soil emissions to be about 30%.
- 281 3.3 Assessment of the DECSO soil emissions using IFS-COMPO simulations

Figure 3 shows the change of RMSE ( $\Delta$ RMSE%) between the TROPOMI observations and the 282 simulated tropospheric NO<sub>2</sub> concentration in the IFS-COMPO model driven by the DECSO-soil 283 and CAMS-soil emissions. The smaller the deviation, the higher the reliability of the soil 284 emissions compared to TROPOMI. Figure 3a-d shows the spatial distribution and seasonal 285 variation of  $\Delta$ RMSE% calculated by formula S3. A negative  $\Delta$ RMSE% represents that the model 286 287 simulation deviation driven by DECSO-soil is smaller than that driven by CAMS-soil, meaning that the DECSO-soil are more consistent with TROPOMI observations than that of CAMS-soil. 288 289 While we use the same TROPOMI NO<sub>2</sub> observations as employed in the DECSO optimization procedure, the atmospheric composition modeling framework is fully independent to DECSO. 290 We found that simulations driven by DECSO soil emissions performed significantly better than 291 using CAMS soil over most of Eastern Europe, North Africa, and Spain (blue area in Figure 3), 292 293 especially in spring and autumn (Figure S9), when the percentual emissions changes with respect to CAMS-Soil are largest. The spatial distribution of changes in  $\Delta RMSE\%$  in areas dominated 294 by rural area, forest, and croplands area is shown in Figure S10-S12. Overall, the simulated 295 RMSE% of DECSO soil is lower than that of CAMS soil, about 6% lower in spring and 2% 296 lower in autumn (Figure S9). In general, the newly calculated soil emissions significantly reduce 297 the error of the simulated and observed tropospheric NO<sub>2</sub> concentrations, which shows the 298 299 consistency of the DECSO-soil. The negative  $\Delta RMSE\%$  over forest shows that soil NOx emissions over forest are underestimated by CAMS. 300



301

**Figure 3.** The deviation of observed and simulated tropospheric NO<sub>2</sub> concentrations driven by DECSO-soil and CAMS-soil (a-d) represented by  $\Delta$ RMSE%. The average of  $\Delta$ RMSE% in (a) spring, (b) summer, (c) autumn and (d) winter calculated by formula S3. RMSE refers to the average difference between the simulated tropospheric NO2 concentration and the observed tropospheric NO2 concentration. Subtracting RMSE of experiment 2 from that of experiment 1 yields  $\Delta$ RMSE. Dividing  $\Delta$ RMSE by the average of the simulated tropospheric NO2 concentration results of the two experiments results in  $\Delta$ RMSE%. A negative  $\Delta$ RMSE% shown

- in blue means that the DECSO-soil are more consistent with TROPOMI observations than that ofCAMS-soil.
- 311 3.4 Indirect estimates of anthropogenic CO2 emissions

Since anthropogenic  $NO_x$  and  $CO_2$  emissions are usually released simultaneously, several studies 312 have used the NO<sub>x</sub> emissions retrieved from satellite observations to infer the anthropogenic CO<sub>2</sub> 313 emissions of countries or regions by a top-down method (de Laat and van der A, 2019; Zheng et 314 al., 2020; Li et al., 2023; Miyazaki and Bowman, 2023). However, these studies did not consider 315 the fact that the NO<sub>x</sub> emissions retrieved based on satellite observations include non-316 317 anthropogenic soil  $NO_x$  emissions. After subtracting soil  $NO_x$  emissions from the total  $NO_x$  of DECSO, we can calculate the co-emitted CO<sub>2</sub> emissions by multiplying DECSO anthropogenic 318 319 NO<sub>x</sub> emissions with the NO<sub>x</sub>/CO<sub>2</sub> emission factors obtained from CAMS inventory. The spatial pattern of CO<sub>2</sub> emissions based on DECSO has a high overall consistency with the bottom-up 320 CAMS emission inventory (Figure 4). The annual CO<sub>2</sub> emissions derived from DECSO (called 321 DECSO-CO2) in the study area in 2019 is 3.7 Gt, which is comparable with the 3.2 Gt of the 322 323 CAMS inventory (called CAMS-CO2). Overall, this reflects the potential of using DECSO to indirectly infer fossil-fuel CO<sub>2</sub> emissions, especially for regions where CO<sub>2</sub> emissions are less 324 well-known than in Europe. 325



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Figure 4. The spatial distributions of (a) estimated annual CO<sub>2</sub> emissions using DECSO, and (b) bottom-up CO<sub>2</sub> emission inventory CAMS.

#### 329 4 Conclusions

We have developed a method for estimating soil NO<sub>x</sub> emissions based on their seasonal 330 characteristics, which we derive from the non-urban regions in our study domain, in our case 331 Europe. The method starts from satellite-based total NOx emissions derived with the DECSO 332 emission inversion system. The estimated soil NO<sub>x</sub> emissions based on DECSO is 2.5 kg N ha<sup>-1</sup> 333  $yr^{-1}$  for Europe in 2019. We found that the existing widely used soil NO<sub>x</sub> emission inventories 334 CAMS and HEMCO (based on empirical and statistical models) report lower soil NO<sub>x</sub> emissions 335 by about 2.5 times. To assess the reliability of the derived DECSO soil  $NO_x$  emissions, we tested 336 them using IFS-COMPO simulations. The model-simulated tropospheric NO<sub>2</sub> concentrations 337 338 driven by DECSO soil  $NO_x$  are closer to the  $NO_2$  concentrations observed by TROPOMI than the simulation driven by CAMS soil emissions. The improvement was especially observed in 339 spring, with a RMSE% reduction of 6%. When checking the spatial distribution (Fig.2), it seems 340 that the discrepancy originates mainly from the forests, where the DECSO derived soil emissions 341

are much higher than those in the CAMS inventory. Possibly the soil NO<sub>x</sub> emissions from forests

in Europe are currently underestimated. Not many studies are yet performed to European forest
 emissions, but Molina-Herrera et al. (2017) concluded that for the state of State of Saxony,
 Germany both agricultural and forest area are significant sources of soil NO<sub>x</sub>.

The seasonal characteristic of DECSO-soil is consistent with the European regional soil  $NO_x$ 346 emissions calculated by Simpson and Darras (2021) based on empirical formulas and process 347 parameter models (see Figure S13b). Regions with similar seasonal patterns of soil NO<sub>x</sub> 348 emissions as the European region are found at mid-latitudes including North America, North 349 Africa, East Asia, Russia (Figure S13 from Simpson and Darras, 2021) making these regions 350 suitable for deriving soil NOx emissions from satellite with the same approach. For mid-latitude 351 regions in the southern hemisphere such as Australia, this method can also be used by shifting 352 the peak parameter to wintertime. 353

354 Our method exploits observations from satellites for a better understanding of the amount and spatiotemporal variation of soil NO<sub>x</sub> emissions. The method, starting from DECSO total 355 emissions, is computationally fast and regionally consistent. After isolating the contribution of 356 357 soil NO<sub>x</sub>, the remainder can be attributed to anthropogenic emissions and the total amount and spatial patterns of anthropogenic CO<sub>2</sub> emissions can be indirectly estimated. The results for 358 Europe are consistent with the bottom-up CO<sub>2</sub> inventory, which demonstrate the potential for 359 DECSO to expand its application to other regions in the world with less information on CO<sub>2</sub> 360 emissions. 361

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371

#### 372 **Open Research**

373	TROPOMI data	a is availa	able at: <u>https:/</u>	/www.tro	pomi.eu/data-products/nitrogen-dioxide. CA	MS soil
374	NOx emissions	are availa	able at: https://p	permalin	c.aeris-data.fr/CAMS-GLOB-SOIL. HEMCO	soil NOx
375	emissions	are	available	at:	https://figshare.com/articles/dataset/Glob	al_high-
376	resolution_emis	ssions_of_	_soil_NOx_sea_	_salt_ae	osols_and_biogenic_VOCs/9962216/4.	National
377	Long-range Tra	ansbounda	ary Air Pollutio	n (LRTA	P) NOx emissions are obtained from the E	uropean

- 378 Environment Agency (https://www.eea.europa.eu/data-and-maps/dashboards/air-pollutant-
- 379 emissions-data-viewer-5). CAMS anthropogenic NOx and CO2 emissions are obtained from ECCAD
- 380 (https://permalink.aeris-data.fr/CAMS-GLOB-ANT). The land use data Land Cover as input data for
- 381 our method are downloaded from the Copernicus Global Land Service
- 382 (https://land.copernicus.eu/global/products/lc).
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