# Improving signal-to-noise ratios of ambient noise cross-correlation functions using local attributes

Bin He<sup>1</sup>, Hejun Zhu<sup>1</sup>, and David E Lumley<sup>1</sup>

<sup>1</sup>University of Texas at Dallas

December 1, 2023

#### Abstract

For seismographic stations with short acquisition duration, the signal-to-noise ratios (SNRs) of ambient noise cross-correlation functions (CCFs) are typically low, preventing us from accurately measuring surface wave dispersion curves or waveform characteristics. In addition, with low-quality CCFs, it is difficult to monitor temporal variations of subsurface physical states or extract relatively weak signals such as body waves. In this study, we propose to use local attributes to improve the SNRs of ambient noise CCFs, which allows us to enhance the quality of CCFs for stations with limited acquisition duration. Two local attributes: local cross-correlation and local similarity, are used in this study. The local cross-correlation allows us to extend the dimensionality of daily CCFs with computational costs similar to global cross-correlation. Taking advantage of this extended dimensionality, the local similarity is then used to measure non-stationary similarity between the extended daily CCFs with a reference stacking trace, which enables us to design better stacking weights to enhance coherent features and attenuate incoherent background noises. Ambient noise recorded by several broadband stations from the USArray in North Texas and Oklahoma, the Superior Province Rifting EarthScope Experiment in Minnesota and Wisconsin and a high-frequency nodal array deployed in the San Bernardino basin are used to demonstrate the performance of the proposed approach for improving the SNR of CCFs.































# Improving signal-to-noise ratios of ambient noise cross-correlation functions using local attributes

Bin He<sup>1</sup>, Hejun Zhu<sup>1,2</sup>, David Lumley<sup>1,2</sup>

<sup>1</sup>Department of Geosciences, The University of Texas at Dallas, Richardson, TX 75080, USA
 <sup>2</sup>Department of Physics, the University of Texas at Dallas, Richardson, TX 75080, USA

# 6 Key Points:

1

2

3

8

# Signal-to-noise ratio of ambient noise cross-correlation functions can be improved by using local cross-correlation and local similarity.

- The local cross-correlation function is efficiently implemented by solving the heat
   equation.
- Applications to broadband and high-frequency nodal arrays validate the effective ness of the proposed method.

Corresponding author: Hejun Zhu, hejun.zhu@utdallas.edu

#### 13 Abstract

For seismographic stations with short acquisition duration, the signal-to-noise ra-14 tios (SNRs) of ambient noise cross-correlation functions (CCFs) are typically low, pre-15 venting us from accurately measuring surface wave dispersion curves or waveform char-16 acteristics. In addition, with low-quality CCFs, it is difficult to monitor temporal vari-17 ations of subsurface physical states or extract relatively weak signals such as body waves. 18 In this study, we propose to use local attributes to improve the SNRs of ambient noise 19 CCFs, which allows us to enhance the quality of CCFs for stations with limited acqui-20 sition duration. Two local attributes: local cross-correlation and local similarity, are used 21 in this study. The local cross-correlation allows us to extend the dimensionality of daily 22 CCFs with computational costs similar to global cross-correlation. Taking advantage of 23 this extended dimensionality, the local similarity is then used to measure non-stationary 24 similarity between the extended daily CCFs with a reference stacking trace, which en-25 ables us to design better stacking weights to enhance coherent features and attenuate 26 incoherent background noises. Ambient noise recorded by several broadband stations from 27 the USArray in North Texas and Oklahoma, the Superior Province Rifting EarthScope 28 Experiment in Minnesota and Wisconsin and a high-frequency nodal array deployed in 29 the San Bernardino basin are used to demonstrate the performance of the proposed ap-30 proach for improving the SNR of CCFs. 31

32

## Plain Language Summary

Seismic ambient noise has been widely used for imaging and monitoring subsur-33 face structures by using cross-correlation functions (CCFs) of continuous recordings be-34 tween station pairs. Typically, we have to stack a lot of CCFs to enhance signals (e.g., 35 surface and body waves) of CCFs. However, for temporal monitoring purposes or those 36 deployed arrays with short acquisition duration, it is impossible to have a lot of stack-37 ing. The lack of enough stacking could result in low signal-to-noise ratios (SNRs) of the 38 CCFs. We proposed a new approach to improve the SNRs of CCFs by using two local 39 attributes: local cross-correlation and local similarity. We first extend the CCFs to a higher 40 dimension by using local cross-correlation, and then the local similarity is used to de-41 fine a better weighting factor for the final stacking. We use several field data examples 42 to prove the effectiveness of the proposed approach. 43

### 44 1 Introduction

With the assumption of homogeneously distributed noise sources, both theoreti-45 cal and experimental studies have demonstrated that empirical Green's functions can 46 be retrieved by cross-correlating continuous ambient noise records between two seismo-47 graphic stations (Aki, 1957; Claerbout, 1968; Buckingham et al., 1992; Lobkis & Weaver, 48 2001; Weaver & Lobkis, 2001; Shapiro & Campillo, 2004; Wapenaar, 2004; Nakata et al., 49 2019). Both surface (Campillo & Paul, 2003; Shapiro & Campillo, 2004; Sabra et al., 2005) 50 and body wave signals (Draganov et al., 2009; Zhan et al., 2010; Poli et al., 2012; Lin 51 et al., 2013; Nakata et al., 2015) have been successively extracted from ambient noise record-52 ings. With empirical Green's functions between each pair of stations in a seismic array, 53 we are able to measure surface wave phase or group dispersion curves, and then perform 54 surface wave tomography to estimate seismic properties in the subsurface. Compared 55 with earthquake tomography, ambient noise tomography allows us to image tectonically 56 inactive regions and achieve better cross-path coverages. To date, ambient noise cross-57 correlation functions (CCFs) have been successively used to investigate velocity as well 58 as anisotropic structures within the crust and uppermost mantle (Shapiro et al., 2005; 59 Y. Yang et al., 2007; Lin et al., 2008; Moschetti et al., 2010; Yao et al., 2010; Huang et 60 al., 2010; Lin et al., 2011; Shen & Ritzwoller, 2016). Recently, ambient noise CCFs are 61 combined with waveform inversion to better constrain subsurface velocity structures (Gao 62 & Shen, 2014; M. Chen et al., 2014; Lee et al., 2014; Y. Liu et al., 2017; Sager et al., 2018; 63 K. Wang et al., 2018; Zhu, 2018; Sager et al., 2020; Fan et al., 2022; Maguire et al., 2022). 64 Taking advantage of continuous records, ambient noise CCFs can also be utilized to mon-65 itor temporal evolution of physical states in the subsurface (Brenguier et al., 2008; Nakata 66 & Snieder, 2011; Hadziioannou et al., 2011; Mainsant et al., 2012; De Plaen et al., 2016; 67 Q.-Y. Wang & Yao, 2020; Le Breton et al., 2021; Mao et al., 2022). 68

Previous studies have demonstrated that we are able to improve the signal-to-noise 69 ratios (SNRs) of ambient noise CCFs as well as reducing the effect of source direction-70 ality by increasing the stacking duration of signals (Bensen et al., 2007). To obtain high-71 quality CCFs, typically we have to use long continuous records and sometimes have to 72 stack data with several years of acquisition. However, for some temporary experiments 73 e.g., Brenguier et al. (2008); Issa et al. (2017); G. Liu et al. (2018); Dougherty et al. (2019); 74 G. Chen et al. (2023); Wu et al. (2023), the SNRs of ambient noise CCFs are typically 75 low, preventing us to measure robust surface wave dispersion curves or waveform char-76

-3-

acteristics. For these cases, we have to design effective approaches to improve the SNRs 77 of ambient noise CCFs. In addition, the SNRs are crucial for extracting weak coherent 78 signals, such as body waves (Snieder, 2004; Nakata et al., 2015, 2016). There are sev-79 eral studies towards improving the SNR of ambient noise CCFs. For instance, Baig et 80 al. (2009) designed better-stacking weights in the time-frequency domain with discrete 81 orthogonal S transform, which allowed them to measure robust Rayleigh and Love ar-82 rivals for stations with long offsets. G. Li et al. (2018) further compared the performances 83 of time and frequency domain inverse S transforms for this stacking procedure. Schimmel 84 et al. (2011) used instantaneous phase coherence to avoid strong amplitude arrivals, such 85 as earthquake signals, and enhance coherent features in noise records. Seats et al. (2012) 86 and Clarke et al. (2011) utilized overlapped moving windows (Welch's method) to im-87 prove the convergence of CCFs towards stable Green's functions. In addition, wavelet 88 and curvelet transforms have also been applied to denoise CCFs by Stehly et al. (2011) 89 and Mao et al. (2022). Furthermore, Weaver and Yoritomo (2018) proposed several schemes 90 to choose optimal weights for stacking so that the effective incident intensity distribu-91 tion is closer to isotropic. Xie et al. (2020) used the root-mean-square ratio to remove 92 those CCFs with low SNR. A systematic evaluation and comparison of the performance 93 of several stacking methods is discussed by X. Yang et al. (2023). 94

In this study, we propose to use local attributes (Rickett & Lumley, 2001; Hale, 2006; 95 Fomel, 2007a) to denoise CCFs when we have short acquisition durations. In compar-96 ison with global attributes, the local attributes enable us to extract non-stationary char-97 acteristics in seismic data. There are a variety of local attributes, including local sim-98 ilarity (Fomel, 2007a), local cross-correlation (Hale, 2006), local frequency (Fomel, 2007a) 99 and local skewness (Fomel & van der Baan, 2014), etc. In this study, we use local cross-100 correlation (Hale, 2006) to extend the dimensionality of daily CCFs, and then use stack-101 ing weights measured by local similarity to improve the stacking quality of CCFs. The 102 local cross-correlation can be considered as a natural extension of the Welch's method 103 (Seats et al., 2012) but with longer overlap windows and much lower computational costs. 104 Its overall computational cost is similar to the conventional global cross-correlation. The 105 local similarity has been used to measure non-stationary similarity between time-lapse 106 images (Fomel & Jin, 2009), improve stacking quality of normal moveout data (G. Liu 107 et al., 2009) as well as angle-domain common-imaging-gathers (Lin et al., 2011). All these 108

-4-

studies demonstrate the capability of using local similarity to improve the stacking qual-ity.

We first review local cross-correlation and local similarity. Next, they are used to improve the SNRs of ambient noise CCFs with one-day and one-month acquisition durations. Continuous noise records from several broadband USArray stations in North Texas and Oklahoma, the Superior Province Rifting EarthScope Experiment in Minnesota and a high-frequency nodal array in the San Bernardino basin are used to demonstrate the performance of the proposed approach.

#### 117 2 Method

118

## 2.1 Local cross-correlation

The global cross-correlation function  $c(\tau)$  between two signals f(t) and g(t) can be defined (Nakata & Snieder, 2011; Harris et al., 2020) as follows:

$$c(\tau) = \int_{-\infty}^{\infty} f(t+\tau)g(t)dt \quad .$$
(1)

The result of the global cross-correlation is a function of time lag  $\tau$ , which cannot cap-119 ture non-stationary time shifts between two input signals and therefore makes it diffi-120 cult to separate signals and noises. In order to fulfill this goal, one way is to perform time-121 windowed cross-correlation, which has much higher computational costs in comparison 122 to Equation 1. In addition, there might be leakage and edge problems due to the selected 123 window functions. To solve these problems, Hale (2006) designed an efficient algorithm 124 to compute the local CCF and measure time-varying correlations between two signals. 125 It was first used to measure non-stationary warpings between time-lapse migration im-126 ages. Later, the local cross-correlation was applied to measure non-stationary travel time 127 differences between two seismograms, which can be utilized to constrain subsurface ve-128 locity structures through full-waveform inversion (Díaz & Sava, 2015). 129

The idea of local cross-correlation is similar to the moving window cross-correlation for capturing transient time shifts, but with much higher computational efficiency. In addition, it helps us to separate signal and noise in a higher dimension. Taking advantage of several properties of Gaussian windows, such as the product of two Gaussian windows is still a Gaussian function, the computational cost of local cross-correlation can be reduced as similar to that of global cross-correlation. For instance, the Gaussian-windowed

-5-

versions of signals f(t) and g(t) can be represented as

$$\hat{f}(t,t_0) = f(t)\omega(t_0 - t)$$
 ,  
 $\hat{g}(t,t_0) = g(t)\omega(t_0 - t)$  . (2)

where  $\omega(t_0-t)$  is the Gaussian window function located at  $t_0$ , i.e.,  $\omega(t_0-t) = \frac{1}{\sqrt{2\pi\sigma}}e^{-(t_0-t)^2/2\sigma^2}$ .

Here  $\sigma$  is the standard deviation, which controls the width of the Gaussian window. With

these windowed signals, we are able to compute the cross-correlation function at each

 $_{140}$  local time step  $t_0$ . Without optimization, the computational cost of this moving window

cross-correlation is  $O(N_l N_w N_s)$ . Here,  $N_l$ ,  $N_w$  and  $N_s$  are the numbers of samples for

time lags, Gaussian window, and input signals, respectively.

<sup>143</sup> The windowed cross-correlation can be written as

$$c(t_0,\tau) = \int_{-\infty}^{\infty} f(t+\tau,t_0+\tau)g(t,t_0)dt$$
  
$$= \int_{-\infty}^{\infty} f(t+\tau)\omega(t_0-t)g(t)\omega(t_0-t)dt$$
  
$$= \int_{-\infty}^{\infty} f(t+\tau)g(t)W(t,t_0)dt \quad , \qquad (3)$$

144 where

$$W(t,t_0) = \omega(t_0 - t)\omega(t_0 - t) = \frac{1}{2\pi\sigma^2}e^{-(t_0 - t)^2/\sigma^2} \quad .$$
(4)

 $W(t, t_0)$  is again a Gaussian function with a smaller width  $\sigma/\sqrt{2}$  compared to  $w(t_0 - t)$ . Based on this property, we can first compute the pre-stacked cross-correlation between two signals, i.e.,  $h(t, \tau) = f(t+\tau)g(t)$ . It is then filtered by a Gaussian function  $W(t, t_0)$ as

$$c(t_0,\tau) = \int_{-\infty}^{\infty} h(t,\tau) W(t,t_0) dt \quad .$$
(5)

Combining with some efficient recursive Gaussian filter algorithms (Deriche, 1993; Alvarez & Mazorra, 1994; Van Vliet et al., 1998), the computational cost of implementing Equation 3 can be similar as the global cross-correlation shown in Equation 1, i.e.,  $O(N_l N_s)$ . This makes it attractive for processing large-scale datasets, such as continuous ambient noise records.

154

## 2.2 Recursive Gaussian convolution

How to efficiently compute the convolution of a signal or an image with a Gaussian function is a common problem in data processing. There are several ways to perform this filtering procedure. For instance, by taking advantage of the Fourier transform, Gaussian convolution can be computed efficiently as elemental-wise multiplication in the

<sup>159</sup> frequency domain. Another efficient way to implement the Gaussian filter is to consider

<sup>160</sup> the result as the solution of the following heat equation:

$$\frac{\partial u}{\partial y} - \frac{\partial^2 u}{\partial x^2} = 0 ,$$

$$u(x,0) = f(x) .$$
(6)

Here, if we let  $y = \sigma^2/2$ , then the final solution of the above heat equations u(x, y) is

equivalent to the convolution of signal f(x) with a Gaussian function  $\omega(x) = \frac{1}{\sqrt{2\pi\sigma}} e^{-x^2/2\sigma^2}$ .

<sup>163</sup> This heat equation can be solved by using finite-difference schemes. Alvarez and Mazorra

<sup>164</sup> (1994) used this property to derive an efficient procedure to compute the Gaussian con-

volution. For instance, they discretized the simulation domain with

$$u_i^j = u(i\Delta x, j\Delta y) \quad , \tag{7}$$

Here  $\Delta x$  and  $\Delta y$  are the grid spacings along the spatial and time directions, *i* and *j* are the associate indices. Then, the finite-difference solution of the above heat equation can be written as

$$\frac{u_i^{j+1} - u_i^j}{\Delta t} - \frac{u_{i+1}^{j+1} + u_{i-1}^{j+1} - 2u_i^{j+1}}{\Delta x^2} = 0 \quad .$$
(8)

169 Let  $\lambda = \Delta y / \Delta x^2$ , then

$$(1+2\lambda)u_i^{j+1} - \lambda u_{i-1}^{j+1} - \lambda u_{i+1}^{j+1} = u_i^j \quad .$$
(9)

<sup>170</sup> We can solve this implicit finite-difference equation using the following filter

$$H(z) = \frac{\nu}{\lambda} \frac{1}{(1 - \nu z^{-1})(1 - \nu z)} \quad , \tag{10}$$

171 with

$$\nu = \frac{1 + 2\lambda - \sqrt{(1 + 4\lambda)}}{2\lambda} \quad . \tag{11}$$

<sup>172</sup> The above filtering procedure can be efficiently solved by using the following three steps,

marching from  $u_i^j$  to  $u_i^{j+1}$  as

$$u_{i}^{j'} = u_{i}^{j} + \nu u_{i-1}^{j'} ,$$
  

$$u_{i}^{j''} = u_{i}^{j'} + \nu u_{i+1}^{j''} ,$$
  

$$u_{i}^{j+1} = \frac{\nu}{\lambda} u_{i}^{j''} .$$
(12)

- <sup>174</sup> The first and second steps in the above procedure are the applications of causal and acausal
- <sup>175</sup> filters, respectively. The third step is just a simple multiplication. Thanks to the uncon-
- ditional stability of this implicit solution (Alvarez & Mazorra, 1994), only a few itera-
- tions (four iterations are used in this study to propagate along the time dimension y)

enable us to converge to the solution of the Gaussian filter. Therefore, the time grid can be defined by  $\Delta y = \sigma^2/(2N_y)$ , where  $N_y$  is the total number of iterations. Combining this recursive Gaussian filter with Equation 5, we are able to efficiently compute the local cross-correlation function. For example, for a certain  $\tau$ , the source term for Equation 6 is the pre-stacked cross-correlation  $h(t, \tau)$ . Then solution of the corresponding heat equation is the non-stationary CCFs  $u(t, \tau)$  at this particular  $\tau$ .

Figure 1 shows a simple example to illustrate the advantages of local cross-correlation 184 over global cross-correlation. There are non-stationary time shifts for three events at 1, 185 3 and 5 s between two input signals. The global cross-correlation is shown in Figure 1b. 186 Although it captures the correct time shifts around -0.4 s, 0.1 s and 0.3 s, it does not 187 carry any useful information about the locations of these time shifts. The local cross-188 correlation is presented in Figure 1c, with three energy spots located at the right loca-189 tions with correct time lags. This local cross-correlation function enables us to extend 190 the dimensionality of the input signals as well as the cross-correlation function, which 191 can be utilized to design a better strategy to improve stacking quality as discussed be-192 low. 193

<sup>194</sup> 2.3 Local similarity

Local similarity is a local attribute that allows us to measure non-stationary similarity between two time series or images (Fomel & Jin, 2009; G. Liu et al., 2009, 2011). First, the global similarity between two signals f(t) and g(t) can be defined as follows

$$s = \frac{\int f(t)g(t)dt}{\sqrt{\int f^2(t)dt}\sqrt{\int g^2(t)dt}} \quad .$$
(13)

In comparison with the global cross-correlation in Equation 1, the global similarity can be considered as a zero-lag cross-correlation function normalized by the energies of each signal. Therefore, it is insensitive to the absolute amplitudes of input signals. Its magnitude ranges from -1 to 1. Similar to local cross-correlation, local similarity was designed to measure non-stationary similarity between time-lapse images (Rickett & Lumley, 2001; Fomel & Jin, 2009). Once taking the square of Equation 13 on both sides, we can write its discretized version as

$$s^{2} = \frac{(\mathbf{f}^{T}\mathbf{g})(\mathbf{g}^{T}\mathbf{f})}{(\mathbf{f}^{T}\mathbf{f})(\mathbf{g}^{T}\mathbf{g})} \quad .$$
(14)

Fomel (2007a) considered the above equation as the solution to the following two least-

206 squares problems:

$$\mathbf{s}_{1} = \arg \min_{\mathbf{s}_{1}} |\mathbf{F}\mathbf{s}_{1} - \mathbf{g}| ,$$
  
$$\mathbf{s}_{2} = \arg \min_{\mathbf{s}_{2}} |\mathbf{G}\mathbf{s}_{2} - \mathbf{f}| .$$
(15)

Here,  $\mathbf{F}$  and  $\mathbf{G}$  are the diagonal operators constructed from the elements of input vec-

tors f and g, respectively. Shaping regularization (Fomel, 2007b) and a conjugate gra-

dient method can be used to solve the above least-squares problems. Then, the local sim-

ilarity can be computed as elemental-wise multiplication between  $s_1$  and  $s_2$  as

$$\mathbf{s} = \sqrt{\mathbf{s}_1 \mathbf{s}_2} \quad . \tag{16}$$

As pointed out in Fomel (2007a), the normalized local correlation can be considered as the first iteration from local similarity. Compared to a scalar value for the global similarity in Equation 13, the local similarity is a vector with the same length as the input signals. It can also be utilized to measure the similarity between two 2D or 3D images with high efficiency.

216

### 2.4 Stacking strategy with local similarity

From local cross-correlation, we obtain a function  $c(t, \tau)$ . The simplest way to calculate the CCF  $c_1(\tau)$  is to use the following linear stacking

$$c_1(\tau) = \frac{1}{N} \sum_t c(t,\tau)$$
 (17)

If we use the above expression to compute noise CCF, its result is similar to the direct implementation of Equation 1. In this study, we choose  $c_1(\tau)$  as a reference trace and then compute the local similarity  $s(t, \tau)$  between  $c_1(\tau)$  and  $c(t, \tau)$ . Once with the local similarity, we can use the following strategy to improve the stacking quality

$$c_2(\tau) = \frac{1}{N} \sum_t \alpha(t,\tau) c(t,\tau) \quad , \tag{18}$$

223 with

$$\alpha(t,\tau) = \begin{cases} s(t,\tau) - \alpha_0 & , & \text{if } \alpha(t,\tau) > \alpha_0 \\ 0 & , & \text{if } \alpha(t,\tau) \le \alpha_0 \end{cases}$$
(19)

Here  $\alpha_0$  is a pre-defined parameter to decide the stacking level. This stacking procedure has been applied to enhance coherent signals for normal moveout corrected data (G. Liu et al., 2009) and angle-domain common-image-gathers (G. Liu et al., 2011). In this study, we use it to improve the stacking quality of noise CCFs. In practice,  $\alpha(t, \tau)$  is a soft thresh-

- <sup>228</sup> old solved in the seislet domain (Daubechies et al., 2004; G. Liu et al., 2009). Compared
- with the similarity shown in Figure 1d, the clipped similarity ( $\alpha(t,\tau)$  in Eq. 19) in Fig-
- <sup>230</sup> ure 1f shows zero background values away from signal regions, while keeping almost 1.0
- weights around the signal regions. This helps to suppress noises away from the signal
- regions. Because this synthetic test does not involve noises, the weighted prestack local
- correlations ( $c_2(\tau)$  in Eq. 18) shown in Figure 1e are basically identical to the prestack
- local correlations shown in Figure 1c. To validate the effectiveness of the proposed method
- to improve the SNR, we give its definition used in this study as

$$SNR = \sqrt{\frac{\frac{1}{N_s} \int_{s_1}^{s_2} f^2(t) dt}{\frac{1}{N_n} \int_{n_1}^{n_2} f^2(t) dt}} \quad .$$
(20)

Where  $[s_1, s_2]$  and  $[n_1, n_2]$  represent signal and noise windows,  $N_s$  and  $N_n$  denote the number of time samples for signal and noise, respectively.

238 **2.5 Workflow** 

In summary, the proposed method based on two local attributes can be implemented 239 through the following six steps: 240 1. Download and preprocess continuous ambient noise data for each station, and cut 241 them into daily recordings. 242 2. Calculate the pre-stacked cross-correlation function  $h(t,\tau) = f(t+\tau)g(t)$  as shown 243 in Equation 3 between daily recordings from station pairs. 244 3. Stack all daily global stacking CCFs and extend it along the time axis to obtain 245 two-dimensional data as a reference. 246 4. Solve the heat equation in Equation 6 to obtain a daily local CCF. Note here, the 247 pre-stacked CCF is the source term f(x) in Equation 6, and the local CCF is also 248 two-dimensional data. 249 5. Calculate the daily local similarity between the reference CCF (step 3) and the 250 daily local CCF(step 4). 251 6. Stack all daily local CCFs weighted by the local similarity according to Equation 252 18 for each station pair. 253

#### <sup>254</sup> 3 Numerical examples

We download continuous ambient noise data, which are then processed by using the NoisePy package (Jiang & Denolle, 2020) following standard procedures from Bensen et al. (2007). It includes deconvolution of instrument responses, bandpass filter (2–10 s), time domain normalization and spectral whitening. The preprocessed data are used to retrieve the CCFs and demonstrate the advantages of the proposed method.

260

#### 3.1 One-day records

Figures 2a and b are vertical component displacements for the USArray stations 261 TA.234A and TA.Z36A recorded on October 10th, 2010. The inter-station between this 262 station pair is 213 km. As shown in Figure 3a, the SNR of the global cross-correlation 263 is relatively low because we only use daily records. Only causal Rayleigh waves can be 264 clearly observed in this CCF. Both amplitude and time-frequency spectra of the global 265 cross-correlation function are presented in Figures 3b and c, respectively. Besides the dom-266 inant Rayleigh wave arrival, there are many background noises, especially from 0.2 to 267 0.5 Hz. The local cross-correlation between Figures 2a and b is shown in Figure 2c. With 268 the local cross-correlation, we are able to extend the 1D CCF to a 2D image, which can 269 be used to measure local similarity and design a better strategy to improve stacking qual-270 ity. Although there are many incoherent background noises in Figure 2c, it includes a 271 vertical coherent feature around time lags between 60 to 90 s. If we use a simple linear 272 stacking for Figure 2c along the time axis (Equation 17), we will obtain the same sig-273 nal as the global cross-correlation shown in Figure 3a. Next, we compute the local sim-274 ilarity between Figure 2c and the raw stacked CCF shown in Figure 3a. In the local sim-275 ilarity result (Figure 2d), the coherent signals between 60 to 90 s stand out in compar-276 ison to the background noise level. With the stacking weights based on the local sim-277 ilarity (Equation 18), the new CCF and its amplitude/time-frequency spectra are pre-278 sented in Figures 3d–f. Compared with the global CCF in Figures 3a–c, the incoherent 279 noises are significantly reduced without changing the dominant Rayleigh wave arrival. 280 In the meanwhile, the overall amplitude spectrum does not change before and after ap-281 plying the stacking weights (Figures 3b and e). Now, it is much easier for us to measure 282 the phase or group dispersion curves of Rayleigh waves for the newly stacked signal (Fig-283 ure 3d). 284

#### 3.2 One-month records

Figures 4a and b present ambient noise CCFs for daily records (October 2010) by 286 using global cross-correlation and local attributes, respectively. With these local attributes, 287 incoherent background noises are attenuated while surface wave signals are preserved and 288 much easier for us to measure. We observe alternative causal and acausal surface wave 289 signals, which might come from the changes in ambient noise source distributions. Since 290 there is data redundancy for Figure 4b, instead of directly stacking it over the time axis, 291 we calculate local similarity again between the raw stacked CCF (Figure 5a) with Fig-292 ure 4b, which is shown in Figure 4c. Again, coherent signals, such as surface wave pack-293 ages, stand out in comparison with background noises. Figures 5a and d compare the 294 raw stacked CCF and the new signal with local attributes. We observe significant im-295 provement of the new signal in terms of SNR. Based on the time-frequency analysis shown 296 in Figures 5c and f, the dispersive characteristics of Rayleigh waves are preserved while 297 background noises from 0.1 to 0.3 Hz are attenuated. Similar to the previous daily ex-298 ample, the overall amplitude spectra do not change too much between these two CCFs. 299

Next, we compare the convergence of the raw global cross-correlation and the new 300 result based on local attributes. Figure 6a shows the convergence of the raw daily global 301 CCFs over one month. With the increasing stacking duration, the SNR is improved and 302 coherent Rayleigh wave signals stand out gradually. However, the convergence and im-303 provement of SNR are relatively low (Figure 7b). Here we use a one-month stacked CCF 304 as the reference trace to compute the similarity between two signals (Equation 13). In 305 this test, the signal window is [-100, -50] s and [50, 100] s, while the noise window ranges 306 from [-200, -150] s and [150, 200] s. Improvements for the local attribute stacking CCF 307 are shown in Figure 6b. In addition, another causal signal appears around 50 s after 10 308 days of stacking, which could be body waves or higher mode surface waves with an ap-309 parent velocity about 4.3 km/s. But those incoherent events (e.g., 20 s) are relatively 310 well suppressed. We also speculate that the improvement of SNR with our method is sig-311 nificant with only several days of stacking (Figure 7b). By stacking with more than 10 312 days, both the correlation coefficient and SNR gradually become stable (Figure 7). It 313 is intriguing that the first 5 days CCFs show strong acausal signals (e.g., -80 s in Fig-314 ures 6a-b) and our improved SNR shows a bump around 5 days (Figure 7b). This might 315 indicate the noise sources were mainly from station Z36A to station 234A before the first 316 5 days. Then more complex noise sources from stations 234A to station Z36A appeared 317

-12-

and resulted in stronger causal signals (e.g. 50-80 s) and complex noises in CCFs. As a result, the SNR goes lower from 5-12 day stacking and gradually turns high with longer stacking.

We further test the proposed approach for 20 USArray stations deployed in North 321 Texas and Oklahoma (Figure 8). One-month data (October 2010) are used for both raw 322 stacking and the new approach. Comparisons of these two results can be found in Fig-323 ure 9. We observe the SNRs of the new approach are much higher in comparison with 324 the classical stacking approach. Except for the dominant Rayleigh waves, we also no-325 tice that there might be additional earlier arriving weaker events emerging from 500 km 326 distance (highlighted by dark blue arrows). For example, the early arrival at about 640 327 km has an apparent velocity of 3.5 km/s, which is higher than the group velocity ( $\sim 2.6$ 328 km/s) of the dominant Rayleigh waves. It could be a candidate for head/diving body 329 waves or higher mode surface waves contaminated with noises. Even they are weak sig-330 nals, our method can retain them as long as they are coherent. 331

Next, we test the proposed approach for several stations from three dense arrays 332 and surrounding stations deployed in Minnesota and Wisconsin around the Midconti-333 nent Rift (Figure 10; Wolin et al., 2015). Figure 11 shows the comparisons of one-month 334 stacked CCFs based on the conventional approach and the proposed procedure. It is ob-335 vious the SNRs are improved quite a lot for most traces shown in Figure 11c with the 336 proposed method. The overall average of the SNR from the traditional stacking method 337 (4.8) is much smaller than the one (35.4) from the proposed method. Although the noise 338 level between 6-15 s period is lower, the SNRs are improved from 6.7 to 53.8 by using 339 the proposed stacking method as shown in Figure 12. It is interesting that SNRs for data 340 in 15-30 s period band (Figure 13) are smaller compared to those in the periods of 6-15 341 s, and the proposed method helps us to improve the SNRs as expected. Panels d-f show 342 the improved SNRs from two-month stacking compared with the monthly stacking (pan-343 els a-c) of Figures 11-13. We speculate that with a two-month stacking duration, the im-344 provements of SNRs using conventional linear stacking are larger than the proposed method, 345 but the SNRs from one-month stacking of the proposed method are still higher than those 346 from two-month stacking with the conventional method. This is important for monitor-347 ing time-lapse changes of near-surface velocity changes with temporary arrays which usu-348 ally have quite short acquisition durations (Nakata et al., 2016; Issa et al., 2017; Mor-349 dret et al., 2020; Zhang et al., 2022). 350

-13-

351

## 3.3 Applications to a nodal array

Between 2017 and 2019, 10 linear dense Distribution of Basin Amplification Seis-352 mic Investigation (BASIN) nodal arrays (SG1–SG4, and SB1–SB6) were deployed in the 353 San Gabriel and San Bernardino basins for Fine characterization of basin shapes and depths 354 (Y. Li et al., 2022; X. Wang et al., 2021). We apply our method to the SB1 array to val-355 idate its performance. The SB1 array (Figure. 14a) consists of 239 Fairfield ZLand nodes 356 with standard 5 Hz 3-component geophones with spatial sampling of  $\sim 250$  m. It was de-357 ployed for approximately one month. The basin depth beneath the SB1 array is about 358 0-3.0 km (Y. Li et al., 2022), therefore, it is essential to have more measurements at lower 359 period bands, such as 1.0-5.0 s. The data downloading and preprocessing are similar to 360 previous tests except the continuous noises were down-sampled with a sampling frequency 361 of 4 Hz and then bandpass filtered between 1-20 s. 362

Taking the first station as the master station, the corresponding CCFs are arranged 363 according to their offset and displayed in Figure 14. As expected, with 5 days of stack-364 ing (Figure 14b), the CCFs from traditional linear stacking have strong noises for all sta-365 tion pairs. The 30-day stacking clearly improves the data quality so that we can observe 366 the dominant acausal Rayleigh waves. Compared with CCFs from linear stacking, the 367 proposed method helps us to enhance coherent signals even with only 5 days of stack-368 ing. The 30-day stacking further improves the data quality. Next, we measure phase dis-369 persion curves for the stacked CCFs, which are the input for surface wave tomography 370 (Yao et al., 2006; Fang et al., 2015). There are three criteria to make sure the measured 371 dispersion curves are stable. (1) The SNR is larger than 5.0; (2) The inter-station dis-372 tance is larger than 1.5 wavelength at corresponding periods (Bensen et al., 2007; G. Chen 373 et al., 2023). (3) The picked phase velocity is within  $\pm 12\%$  of a 3D reference phase ve-374 locity model. Here the 3D reference phase velocity model is constructed from a local to-375 mographic model, CVM-S 4.26 (Lee et al., 2014), by conducting a forward modeling based 376 on the fast-marching method (Rawlinson & Sambridge, 2004; Fang et al., 2015). Thanks 377 to the improvement of SNR from our method, more high-quality dispersion curves pass 378 the selection criteria with 5-day stacking compared with those measured from linear stack-379 ing as illustrated in Figure 15. More importantly, the number of measurements for 1-380 5 s and 10-15 s period bands are also increased. These are further improved for 30-day 381 stacked CCFs using the proposed method. 382

-14-

### **4 Discussions**

Commonly used ambient noise processing procedures (Bensen et al., 2007) require stacking over long, continuous records to enhance coherent signals, also require the assumption of evenly distributed ambient sources in order to obtain good estimates of Green's functions. However, both conditions impose constraints on the application of temporary arrays with short acquisition duration.

To improve the SNR of CCFs stacked with a short duration, Xie et al. (2020) pro-389 posed a root-mean-square-ratio selective (RMSRS) stacking procedure to remove those 390 CCFs that negatively contribute to the SNR of the final stacked CCF. It is realized by 391 comparing the root-mean-square ratio of signals and noises for each CCF and the stacked 392 CCF. Therefore, its effectiveness depends on an accurate definition of signal/noise win-393 dows. Here, we compare the conventional linear stacking, proposed procedure and RM-394 SRS stacking for the station pair TA.A12A-TA.A18A at 10-35 s period band. The sig-395 nal window (Figure 17c) is defined with the reference time  $t_{ref}$  and maximum period of 396 interest  $T_{max}$  as  $[t_{ref} - 2T_{max}, t_{ref} + 2T_{max}]$ . Here, the reference time  $t_{ref} = d/v_{ref}$  is 397 defined by the inter-station distance d (425 km) and a reference group velocity  $v_{ref}$  (3.0 398 km/s). The noise window is defined from 0 to the signal windows and  $4T_{max}$  out of the 399 signal windows (Xie et al., 2020). Compared with linear stacking, the RMSRS stacking 400 successfully suppresses the noises within the defined noise windows and therefore improves 401 the SNR as highlighted in Figures 17d and f. Similarly, our approach also successfully 402 attenuates those incoherent noises and significantly improves the SNR, which is about 403 two times the other two methods. However, the coherent signal at about 75 seconds with 404 an apparent velocity of 5.6 km/s is retained by our approach due to the similarity be-405 tween local CCFs and the stacked CCFs. The main reason that the RMSRS-based method 406 helps to suppress this signal is because they are selected as noises. Such selection seems 407 to be challenging to deal with low-SNR CCFs at shorter period bands (3-16 s) obtained 408 from the high-frequency nodal array (Figure 14). As illustrated in Figures 14d, the im-409 provement of selective stacking is limited compared with linear stacking. On the con-410 trary, our approach significantly improves the stacking quality and helps us to obtain more 411 high-quality dispersion curves. We note here that, for fair comparisons, the RMSRS-based 412 method is implemented on the raw stacked CCFs. For better performance, Xie et al. (2020) 413 suggested using RMSRS stacking at several narrow period bands so that it could define 414 a better selection window, which is out of the scope of this study. 415

-15-

Although the proposed approach helps us to improve the SNR of CCFs with only 416 several days of stacking, the non-causality and asymmetry, which mainly arise from the 417 uneven distribution of ambient noise sources, remain challenging. For tomography pur-418 poses, theoretical works indicate that phase velocities can be estimated from the empir-419 ical Green's functions, which are obtained by taking the negative time derivative of the 420 symmetric cross-correlation under the assumption of a spatially homogeneous ambient-421 noise source distribution (Lobkis & Weaver, 2001; Sabra et al., 2005; Snieder, 2004; Yao 422 et al., 2006; Lin et al., 2008). As suggested by Yao et al. (2006), inhomogeneous source 423 distribution may contribute to 1-3 percent inconsistency between phase velocity mea-424 surements and the traditional earthquake-based two-station method between periods of 425 20–30 s. Therefore, the "symmetric" CCF is usually taken by the average of the cross-426 correlation at positive and negative correlation lag times (Yao et al., 2006; Lin et al., 2008). 427 However, how much such averaged CCFs affect tomography results remains unknown. 428 On the other hand, several studies suggested measuring full-waveform differences of CCFs 429 and source location simultaneously (Tromp et al., 2010; Sager et al., 2018; Datta et al., 430 2023), which naturally mitigates the uncertainty caused by the source distributions. In 431 both cases, it is important for us to obtain high SNR CCFs, especially for short-duration 432 nodal arrays. In addition, because the high-quality phase velocity measurements obtained 433 from 5-day stacking (Figure 15c) by the proposed approach is comparable with those ob-434 tained from the 30-day linear stacking, the surface wave-based monitoring (Durand et 435 al., 2011; Brenguier et al., 2020) seems to be possible, albeit coda waves are mostly used 436 (Mao et al., 2019, 2020; Luo et al., 2023). Another potential application for our approach 437 could be weak coherent signal extraction, such as body waves (Zhan et al., 2010; Poli 438 et al., 2012; Nakata et al., 2015, 2016; Mao et al., 2020). Body waves extracted from am-439 bient noise CCFs have been proven to be capable of improving the imaging resolution 440 compared to surface wave tomography (Nakata et al., 2015). As illustrated in Figures 441 6 and 14, those coherent signals with an apparent velocity larger than 4.0 km/s might 442 be good candidates for body waves, although we are not able to rule out the possibil-443 ity of higher-order surface waves. 444

The last factor we need to consider is the computational and memory cost for our approach. Taking station pair TA.A12A and TA.A18A as an example, we have 30 days of recordings and 86400-time samples for each day. The cross-correlation time lag varies from -480 s to 480 s for every 1 s. We calculate the global cross-correlations every 60 min-

-16-

utes with 75% overlaps, which yields 91 subset CCFs for each day. They are then nor-449 malized and stacked to obtain the dayily CCF. Finally, 30 days of global cross-correlation 450 and linear stacking take 47.9 s and 0.02 s, respectively, while the local cross-correlation, 451 local similarity and weighted stacking take 196.7 s, 12,150 s and 78.3 s. We note here, 452 that the local cross-correlation solved with our method actually contains 86,400 CCFs 453 for each day. It helps us to extend the dimension of CCFs dramatically (534 times the 454 number of global CCFs), but with only four times the computational cost compared to 455 global stacking. Whereas, such a high dimension, in turn, greatly decreases the efficiency 456 of our approach. To mitigate this problem, the local CCFs are downsampled 60 times 457 (from 1.0 s to 60 s) by taking the average CCFs for every 60 samples, and then the com-458 putational cost for local similarity and weighted stacking is decreased to 167.1 and 10.0 459 s, respectively. The final stacked CCFs by using these local CCFs are almost the same 460 as the original local CCFs. Overall, the computational cost for the proposed approach 461 is 7.8 times the computational cost compared to the hourly global stacking, but with 15.8 462 times more CCFs. In addition, increasing the cross-correlation time window (e.g., from 463 one hour to three hours) does not affect the efficiency of our approach, but will increase 464 the computational time for global cross-correlation. Finally, we compare the computa-465 tion cost using the nodal array. For each station, we have 345,600 time samples per day 466 with a 4 Hz sampling frequency. The cross-correlation time lag varies from -120 s to 120 467 s for every 0.25 s. The global and local CCFs are calculated the same as in the previ-468 ous example. To save computational costs, we downsample the local CCFs 60 times (from 469 0.25 s to 15 s). Then all station-pairs are distributed to 72 CPU cores for parallel com-470 putations. Our approach takes 5792.0 s and the traditional linear stacking takes 792.0 471 s. The final stacked CCFs are compared in section 3.3. In conclusion, our approach sig-472 nificantly helps us to improve the SNR of CCFs, but with about 7.5 times the compu-473 tation cost compared to traditional linear stacking. Such extra computational costs are 474 bearable compared to the following computational costs tomography (Zhu, 2018; Wu et 475 al., 2023; G. Chen et al., 2023). 476

#### 477 **5** Conclusion

In this study, by taking advantage of local attributes, we present a new approach
to increase the SNRs of ambient noise CCFs. Two local attributes are used in this study:
local cross-correlation and local similarity. The local cross-correlation is employed to ex-

-17-

tend the dimensionality of daily CCFs, and the local similarity is used to design better-481 stacking weights to enhance coherent signals and attenuate incoherent background noises. 482 Applications to ambient noise records from several broadband stations and a high-frequency 483 nodal array demonstrate the performance of the proposed approach. With higher SNRs, 484 we are able to extract more high-quality dispersion curves, which are important for sur-485 face wave tomography. In addition, 5-day stacking by our approach can produce CCFs 486 comparable to 30-day linear stacking in terms of SNRs, demonstrating its potential ap-487 plications for time-lapse monitoring. In addition, extracting coherent weak signals, such 488 as body waves, could be another application of the proposed approach. 489

#### 490 Open Research

All seismic data used in this study can be obtained from the IRIS Data Management Center (https://ds.iris.edu/ds) under the network codes TA and XI. We use the Noisepy (https://noise-python.readthedocs.io/en/latest/) for parallel data downloading and preprocessing (Jiang & Denolle, 2020). The open software Madagascar (Fomel et al., 2013) download from (http://www.ahay.org) is used to calculate local similarity and plot figures. We also use PyGMT (Wessel et al., 2019) downloaded from (https://www.pygmt.org/latest/) for plotting figures.

498

#### Author Contributions

The authors confirm their contribution to the paper as follows: study conception and design: Hejun Zhu; data collection: Hejun Zhu and Bin He; analysis and interpretation of results: Hejun Zhu, Bin He and David Lumley; draft manuscript preparation: Hejun Zhu, Bin He and David Lumley. All authors reviewed the results and approved the final version of the manuscript.

#### 504 Acknowledgments

This paper is contribution no. \*\*\* from the Department of Geosciences at the University of Texas at Dallas. The numerical results are computed through the Optane nodes on the UTD Seismology Group HPC clusters. This research is partially supported by the sponsors of the UT Dallas 3D + 4D Seismic Full Waveform Inversion research consortium.

### 510 References

- Aki, K. (1957). Space and time spectra of stationary stochastic waves, with special
   reference to microtremors. Bulletin of the Earthquake Research Institute, 35,
   415–456.
- Alvarez, L., & Mazorra, L. (1994). Signal and image restoration using shock fil ters and anisotropic diffusion. SIAM journal on numerical analysis, 31(2),
   590–605.
- Baig, A. M., Campillo, M., & Brenguier, F. (2009). Denoising seismic noise cross
   correlations. Journal of Geophysical Research: Solid Earth, 114 (B8).
- <sup>519</sup> Bensen, G., Ritzwoller, M., Barmin, M., Levshin, A. L., Lin, F., Moschetti, M.,
- Yang, Y. (2007). Processing seismic ambient noise data to obtain reli able broad-band surface wave dispersion measurements. *Geophysical journal international*, 169(3), 1239–1260.
- <sup>523</sup> Brenguier, F., Courbis, R., Mordret, A., Campman, X., Boué, P., Chmiel, M., ...
- others (2020). Noise-based ballistic wave passive seismic monitoring. part 1: body waves. *Geophysical Journal International*, 221(1), 683–691.
- Brenguier, F., Shapiro, N. M., Campillo, M., Ferrazzini, V., Duputel, Z., Coutant,
   O., & Nercessian, A. (2008). Towards forecasting volcanic eruptions using
   seismic noise. *Nature Geoscience*, 1(2), 126–130.
- Buckingham, M. J., Berknout, B. V., & Glegg, S. A. (1992). Imaging the ocean with
   ambient noise. *Nature*, 356 (6367), 327–329.
- Campillo, M., & Paul, A. (2003). Long-range correlations in the diffuse seismic coda.
   Science, 299(5606), 547–549.
- <sup>533</sup> Chen, G., Chen, J., Tape, C., Wu, H., & Tong, P. (2023). Double-difference ad <sup>534</sup> joint tomography of the crust and uppermost mantle beneath alaska. Journal
   <sup>535</sup> of Geophysical Research: Solid Earth, 128(1), e2022JB025168.
- <sup>536</sup> Chen, M., Huang, H., Yao, H., van der Hilst, R., & Niu, F. (2014). Low wave speed
  <sup>537</sup> zones in the crust beneath se tibet revealed by ambient noise adjoint tomogra<sup>538</sup> phy. *Geophysical Research Letters*, 41(2), 334–340.
- <sup>539</sup> Claerbout, J. F. (1968). Synthesis of a layered medium from its acoustic transmis <sup>540</sup> sion response. *Geophysics*, 33(2), 264–269.
- <sup>541</sup> Clarke, D., Zaccarelli, L., Shapiro, N., & Brenguier, F. (2011). Assessment of res <sup>542</sup> olution and accuracy of the moving window cross spectral technique for mon-

543	itoring crustal temporal variations using ambient seismic noise. Geophysical
544	$Journal \ International, \ 186(2), \ 867-882.$
545	Datta, A., Shekar, B., & Kumar, P. L. (2023). Acoustic full waveform inversion for
546	2-d ambient noise source imaging. Geophysical Journal International, $234(3)$ ,
547	1628 - 1639.
548	Daubechies, I., Defrise, M., & De Mol, C. (2004). An iterative thresholding algo-
549	rithm for linear inverse problems with a sparsity constraint. <i>Communications</i>
550	on Pure and Applied Mathematics: A Journal Issued by the Courant Institute
551	of Mathematical Sciences, 57(11), 1413–1457.
552	De Plaen, R. S., Lecocq, T., Caudron, C., Ferrazzini, V., & Francis, O. (2016).
553	Single-station monitoring of volcanoes using seismic ambient noise. $Geophysical$
554	Research Letters, $43(16)$ , 8511–8518.
555	Deriche, R. (1993). Recursively implementating the gaussian and its derivatives (Un-
556	published doctoral dissertation). INRIA.
557	Díaz, E., & Sava, P. (2015). Data domain wavefield tomography using local correla-
558	tion functions. In 2015 seg annual meeting.
559	Dougherty, S. L., Cochran, E. S., & Harrington, R. M. (2019). The large-n seismic
560	survey in oklahoma (lasso) experiment. Seismological Research Letters, $90(5)$ ,
561	2051 - 2057.
562	Draganov, D., Campman, X., Thorbecke, J., Verdel, A., & Wapenaar, K. (2009). Re-
563	flection images from ambient seismic noise. Geophysics, $74(5)$ , A63–A67.
564	Durand, S., Montagner, J., Roux, P., Brenguier, F., Nadeau, R., & Ricard, Y.
565	(2011). Passive monitoring of anisotropy change associated with the park-
566	field 2004 earthquake. Geophysical Research Letters, $38(13)$ .
567	Fan, X., Guo, Z., Zhao, Y., & Chen, QF. (2022). Crust and uppermost man-
568	tle magma plumbing system beneath changbaishan intraplate volcano,
569	china/north korea, revealed by ambient noise adjoint tomography. $Geophysical$
570	Research Letters, $49(12)$ , e2022GL098308.
571	Fang, H., Yao, H., Zhang, H., Huang, YC., & van der Hilst, R. D. (2015). Direct
572	inversion of surface wave dispersion for three-dimensional shallow crustal struc-
573	ture based on ray tracing: methodology and application. Geophysical Journal
574	International, $201(3)$ , $1251-1263$ .
575	Fomel, S. (2007a). Local seismic attributes. <i>Geophysics</i> , 72(3), A29–A33.

- Fomel, S. (2007b). Shaping regularization in geophysical-estimation problems. *Geo- physics*, 72(2), R29–R36.
- Fomel, S., & Jin, L. (2009). Time-lapse image registration using the local similarity attribute. *Geophysics*, 74(2), A7–A11.
- Fomel, S., Sava, P., Vlad, I., Liu, Y., & Bashkardin, V. (2013). Madagascar: Open source software project for multidimensional data analysis and reproducible
   computational experiments. *Journal of Open Research Software*, 1(1).
- Fomel, S., & van der Baan, M. (2014). Local skewness attribute as a seismic phase
   detector. *Interpretation*, 2(1), SA49–SA56.
- Gao, H., & Shen, Y. (2014). Upper mantle structure of the cascades from full-wave
   ambient noise tomography: Evidence for 3d mantle upwelling in the back-arc.
   *Earth and Planetary Science Letters*, 390, 222–233.
- Hadziioannou, C., Larose, E., Baig, A., Roux, P., & Campillo, M. (2011). Improving
   temporal resolution in ambient noise monitoring of seismic wave speed. Jour nal of Geophysical Research: Solid Earth, 116(B7).
- Hale, D. (2006). Fast local cross-correlations of images. In Seg technical program expanded abstracts 2006 (pp. 3160–3164). Society of Exploration Geophysicists.
- Harris, C. R., Millman, K. J., van der Walt, S. J., Gommers, R., Virtanen, P., Cour napeau, D., ... Oliphant, T. E. (2020, September). Array programming with
   NumPy. Nature, 585(7825), 357–362. Retrieved from https://doi.org/
   10.1038/s41586-020-2649-2 doi: 10.1038/s41586-020-2649-2
- Huang, H., Yao, H., & van der Hilst, R. D. (2010). Radial anisotropy in the crust of
   se tibet and sw china from ambient noise interferometry. *Geophysical Research Letters*, 37(21).
- Issa, N. A., Lumley, D., & Pevzner, R. (2017). Passive seismic imaging at reservoir
   depths using ambient seismic noise recorded at the otway co2 geological stor age research facility. *Geophysical Journal International*, 209(3), 1622–1628.
- Jiang, C., & Denolle, M. A. (2020). Noisepy: A new high-performance python tool for ambient-noise seismology. *Seismological Research Letters*, 91(3), 1853– 1866.
- Le Breton, M., Bontemps, N., Guillemot, A., Baillet, L., & Larose, É. (2021). Landslide monitoring using seismic ambient noise correlation: challenges and applications. *Earth-Science Reviews*, 216, 103518.

609	Lee, EJ., Chen, P., Jordan, T. H., Maechling, P. B., Denolle, M. A., & Beroza,
610	G. C. (2014). Full-3-d tomography for crustal structure in southern california
611	based on the scattering-integral and the adjoint-wavefield methods. $Journal \ of$
612	Geophysical Research: Solid Earth, 119(8), 6421–6451.
613	Li, G., Niu, F., Yang, Y., & Xie, J. (2018). An investigation of time-frequency
614	domain phase-weighted stacking and its application to phase-velocity extrac-
615	tion from ambient noise's empirical green's functions. Geophysical Journal
616	International, $212(2)$ , $1143-1156$ .
617	Li, Y., Villa, V., Clayton, R., & Persaud, P. (2022). Shear wave velocities in the san
618	gabriel and san bernardino basins, california. Authorea Preprints.
619	Lin, FC., Moschetti, M. P., & Ritzwoller, M. H. (2008). Surface wave tomogra-
620	phy of the western united states from ambient seismic noise: Rayleigh and
621	love wave phase velocity maps. $Geophysical Journal International, 173(1),$
622	281-298.
623	Lin, FC., Ritzwoller, M. H., Yang, Y., Moschetti, M. P., & Fouch, M. J. (2011).
624	Complex and variable crustal and uppermost mantle seismic anisotropy in the
625	western united states. Nature Geoscience, $4(1)$ , 55–61.
626	Lin, FC., Tsai, V. C., Schmandt, B., Duputel, Z., & Zhan, Z. (2013). Extracting
627	seismic core phases with array interferometry. Geophysical Research Letters,
628	40(6), 1049-1053.
629	Liu, G., Fomel, S., & Chen, X. (2011). Stacking angle-domain common-image gath-
630	ers for normalization of illumination. Geophysical Prospecting, $59(2)$ , 244–255.
631	Liu, G., Fomel, S., Jin, L., & Chen, X. (2009). Stacking seismic data using local cor-
632	relation. Geophysics, $74(3)$ , V43–V48.
633	Liu, G., Persaud, P., & Clayton, R. W. (2018). Structure of the northern los angeles
634	basins revealed in teleseismic receiver functions from short-term nodal seismic
635	arrays. Seismological Research Letters, $89(5)$ , 1680–1689.
636	Liu, Y., Niu, F., Chen, M., & Yang, W. (2017). 3-d crustal and uppermost man-
637	tle structure beneath ne china revealed by ambient noise adjoint tomography.
638	Earth and Planetary Science Letters, 461, 20–29.

Lobkis, O. I., & Weaver, R. L. (2001). On the emergence of the green's function
in the correlations of a diffuse field. *The Journal of the Acoustical Society of America*, 110(6), 3011–3017.

-22-
642	Luo, B., Zhang, S., & Zhu, H. (2023). Monitoring seasonal fluctuation and long-term
643	trends for the greenland ice sheet using seismic noise auto-correlations. Geo-
644	physical Research Letters, $50(7)$ , e2022GL102146.
645	Maguire, R., Schmandt, B., Li, J., Jiang, C., Li, G., Wilgus, J., & Chen, M. (2022).
646	Magma accumulation at depths of prior rhyolite storage beneath yellowstone
647	caldera. Science, 378(6623), 1001–1004.
648	Mainsant, G., Larose, E., Brönnimann, C., Jongmans, D., Michoud, C., & Jaboyed-
649	off, M. (2012). Ambient seismic noise monitoring of a clay landslide: Toward
650	failure prediction. Journal of Geophysical Research: Earth Surface, 117(F1).
651	Mao, S., Campillo, M., van Der Hilst, R. D., Brenguier, F., Stehly, L., & Hillers,
652	G. (2019). High temporal resolution monitoring of small variations in crustal
653	strain by dense seismic arrays. Geophysical Research Letters, $46(1)$ , 128–137.

- Mao, S., Lecointre, A., van der Hilst, R. D., & Campillo, M. (2022). Space-time
   monitoring of groundwater fluctuations with passive seismic interferometry.
   *Nature communications*, 13(1), 1–9.
- Mao, S., Mordret, A., Campillo, M., Fang, H., & van der Hilst, R. D. (2020). On the
   measurement of seismic traveltime changes in the time-frequency domain with
   wavelet cross-spectrum analysis. *Geophysical Journal International*, 221(1),
   550–568.
- Mordret, A., Courbis, R., Brenguier, F., Chmiel, M., Garambois, S., Mao, S., ... others (2020). Noise-based ballistic wave passive seismic monitoring-part 2: surface waves. *Geophysical Journal International*, 221(1), 692–705.
- Moschetti, M., Ritzwoller, M., Lin, F., & Yang, Y. (2010). Seismic evidence for
   widespread western-us deep-crustal deformation caused by extension. Nature,
   464 (7290), 885–889.
- Nakata, N., Boué, P., Brenguier, F., Roux, P., Ferrazzini, V., & Campillo, M. (2016).
  Body and surface wave reconstruction from seismic noise correlations between
  arrays at piton de la fournaise volcano. *Geophysical Research Letters*, 43(3),
  1047–1054.
- Nakata, N., Chang, J. P., Lawrence, J. F., & Boué, P. (2015). Body wave extraction
   and tomography at long beach, california, with ambient-noise interferometry.
   *Journal of Geophysical Research: Solid Earth*, 120(2), 1159–1173.
- <sup>674</sup> Nakata, N., Gualtieri, L., & Fichtner, A. (2019). Seismic ambient noise. Cambridge

-23-

675	University Press.
676	Nakata, N., & Snieder, R. (2011). Near-surface weakening in japan after the 2011
677	tohoku-oki earthquake. Geophysical Research Letters, $38(17)$ .
678	Poli, P., Pedersen, H., & Campillo, M. (2012). Emergence of body waves from cross-
679	correlation of short period seismic noise. Geophysical Journal International,
680	188(2), 549-558.
681	Rawlinson, N., & Sambridge, M. (2004). Wave front evolution in strongly heteroge-
682	neous layered media using the fast marching method. Geophysical Journal In-
683	ternational, 156(3), 631-647.
684	Rickett, J., & Lumley, D. (2001). Cross-equalization data processing for time-lapse
685	seismic reservoir monitoring: A case study from the gulf of mexico. Geo-
686	$physics, \ 66(4), \ 1015-1025.$
687	Sabra, K. G., Gerstoft, P., Roux, P., Kuperman, W., & Fehler, M. C. (2005). Ex-
688	tracting time-domain green's function estimates from ambient seismic noise.
689	Geophysical research letters, $32(3)$ .
690	Sager, K., Boehm, C., Ermert, L., Krischer, L., & Fichtner, A. (2020). Global-scale
691	full-waveform ambient noise inversion. Journal of Geophysical Research: Solid
692	Earth, 125(4), e2019JB018644.
693	Sager, K., Ermert, L., Boehm, C., & Fichtner, A. (2018). Towards full waveform am-
694	bient noise inversion. Geophysical Journal International, $212(1)$ , 566–590.
695	Schimmel, M., Stutzmann, E., & Gallart, J. (2011). Using instantaneous phase
696	coherence for signal extraction from ambient noise data at a local to a global
697	scale. Geophysical Journal International, $184(1)$ , $494-506$ .
698	Seats, K. J., Lawrence, J. F., & Prieto, G. A. (2012). Improved ambient noise cor-
699	relation functions using welch s method. Geophysical Journal International,
700	188(2), 513-523.
701	Shapiro, N. M., & Campillo, M. (2004). Emergence of broadband rayleigh waves
702	from correlations of the ambient seismic noise. $Geophysical Research Letters$ ,
703	31(7).
704	Shapiro, N. M., Campillo, M., Stehly, L., & Ritzwoller, M. H. (2005). High-
705	resolution surface-wave tomography from ambient seismic noise. Science,
706	307(5715),1615-1618.
707	Shen, W., & Ritzwoller, M. H. (2016). Crustal and uppermost mantle structure be-

708	neath the united states. Journal of Geophysical Research: Solid Earth, 121(6),
709	4306-4342.
710	Snieder, R. (2004). Extracting the green's function from the correlation of coda
711	waves: A derivation based on stationary phase. Physical review $E$ , $69(4)$ ,
712	046610.
713	Stehly, L., Cupillard, P., & Romanowicz, B. (2011). Towards improving ambient
714	noise tomography using curvelet denoising filters and sem simulations of seis-
715	mic ambient noise simultaneously. Comptes Rendus Geoscience, 343(8-9),
716	591-599.
717	Tromp, J., Luo, Y., Hanasoge, S., & Peter, D. (2010). Noise cross-correlation sensi-
718	tivity kernels. Geophysical Journal International, 183(2), 791–819.
719	Van Vliet, L. J., Young, I. T., & Verbeek, P. W. (1998). Recursive gaussian deriva-
720	tive filters. In Proceedings. fourteenth international conference on pattern
721	recognition (cat. no. 98ex170) (Vol. 1, pp. 509–514).
722	Wang, K., Yang, Y., Basini, P., Tong, P., Tape, C., & Liu, Q. (2018). Refined
723	crustal and uppermost mantle structure of southern california by ambient noise
724	adjoint tomography. Geophysical Journal International, 215(2), 844–863.
725	Wang, QY., & Yao, H. (2020). Monitoring of velocity changes based on seismic
726	ambient noise: A brief review and perspective. Earth and Planetary Physics,
727	4(5), 532-542.
728	Wang, X., Zhan, Z., Zhong, M., Persaud, P., & Clayton, R. W. (2021). Urban basin
729	structure imaging based on dense arrays and bayesian array-based coherent
730	receiver functions. Journal of Geophysical Research: Solid Earth, 126(9),
731	e2021JB022279.
732	Wapenaar, K. (2004). Retrieving the elastodynamic green's function of an arbitrary $\left( \frac{1}{2} \right)$
733	inhomogeneous medium by cross correlation. Physical review letters, $93(25)$ ,
734	254301.
735	Weaver, R. L., & Lobkis, O. I. (2001). Ultrasonics without a source: Thermal
736	fluctuation correlations at mhz frequencies. Physical Review Letters, 87(13),
737	134301.
738	Weaver, R. L., & Yoritomo, J. Y. (2018). Temporally weighting a time varying noise
739	field to improve green function retrieval. The Journal of the Acoustical Society
740	of America, 143(6), 3706–3719.

-25-

- Wessel, P., Luis, J., Uieda, L., Scharroo, R., Wobbe, F., Smith, W. H., & Tian, D.
  (2019). The generic mapping tools version 6. *Geochemistry, Geophysics, Geosystems, 20*(11), 5556–5564.
  Wolin, E., van der Lee, S., Bollmann, T. A., Wiens, D. A., Revenaugh, J., Dar-
- byshire, F. A., ... Wysession, M. E. (2015). Seasonal and diurnal variations in
  long-period noise at spree stations: The influence of soil characteristics on shallow stations' performance. Bulletin of the Seismological Society of America,
  105(5), 2433–2452.
- Woollard, G. (1965). The bouguer gravity anomaly map of the united states. Eos,
   Transactions American Geophysical Union, 46(1), 197–202.
- Wu, S.-M., Huang, H.-H., Lin, F.-C., Farrell, J., & Schmandt, B. (2023). Extreme
   seismic anisotropy indicates shallow accumulation of magmatic sills beneath
   yellowstone caldera. *Earth and Planetary Science Letters*, 616, 118244.
- Xie, J., Yang, Y., & Luo, Y. (2020). Improving cross-correlations of ambient noise
   using an rms-ratio selection stacking method. *Geophysical Journal Interna- tional*, 222(2), 989–1002.
- Yang, X., Bryan, J., Okubo, K., Jiang, C., Clements, T., & Denolle, M. A. (2023).
   Optimal stacking of noise cross-correlation functions. *Geophysical Journal International*, 232(3), 1600–1618.
- Yang, Y., Ritzwoller, M. H., Levshin, A. L., & Shapiro, N. M. (2007). Ambient noise
   rayleigh wave tomography across europe. *Geophysical Journal International*,
   168(1), 259–274.
- Yao, H., van Der Hilst, R. D., & De Hoop, M. V. (2006). Surface-wave array tomography in se tibet from ambient seismic noise and two-station analysis—i. phase
  velocity maps. *Geophysical Journal International*, 166(2), 732–744.
- Yao, H., Van Der Hilst, R. D., & Montagner, J.-P. (2010). Heterogeneity and
   anisotropy of the lithosphere of se tibet from surface wave array tomography.
   *Journal of Geophysical Research: Solid Earth*, 115(B12).
- Zhan, Z., Ni, S., Helmberger, D. V., & Clayton, R. W. (2010). Retrieval of moho reflected shear wave arrivals from ambient seismic noise. *Geophysical Journal International*, 182(1), 408–420.
- Zhang, Z., Nakata, N., Karplus, M., Kaip, G., & Yi, J. (2022). Shallow ice-sheet
  composite structure revealed by seismic imaging near the west antarctic ice



Figure 1. Comparison of local and global cross-correlations. Panel(a) shows two signals with non-stationary time shifts. Three events with time shifts of -0.4 s, 0.1 s and 0.3 s are used. Panels (b) and (c) present global and local cross-correlations, respectively. Here, the 1D global cross-correlations are extended along the time dimension for better comparisons with local crosscorrelations.  $\sigma$  is set to 0.2 s so that it is small enough to capture the non-stationary property of these two signals. Panels (d) and (f) show the similarity (s in Equation 16) between local and global cross-correlations before and after applying a threshold ( $\alpha(t, \tau)$  in Equation 19), respectively. Panel e shows the prestack local cross-correlations weighted by local similarity ( $c_2$  in Equation 18)



<sup>778</sup> Journal International, 214(1), 716–730.



Figure 2. Local cross-correlation and local similarity for one-day CCFs between stations TA.234A and TA.Z36A on October 10, 2010. Panels (a) and (b) show one-day records for stations TA.234A and TA.Z36A, respectively. Panel (c) presents local cross-correlation between these two signals. Panel (d) shows the local similarity between panel (c) and the linearly stacked signal shown in Figure 3a. The causal Rayleigh waves between 60-80 s can be clearly observed in panel (d).



**Figure 3.** Comparisons of one-day CCFs from simple stacking and stacking with local attributes. Panel (a) shows the CCF with simple linear stacking (Equation 1). Panels (b) and (c) are the amplitude and time-frequency spectra of panel (a), respectively. Panel (d) is the CCF stacked with local similarity. Panels (e) and (f) are the amplitude and time-frequency spectra of panel (d).



Figure 4. Comparisons of daily CCFs within one month (October 2010) for simple stacking (a) and stacking with local attribute (b). Panel (c) shows the local similarity between panel (b) and the linear stacked result shown in Figure 5a.



Figure 5. The same setting as Figure 3 except for monthly stacked results on October 2010.



Figure 6. Comparisons of convergence for conventional (a) and the proposed procedure (b). The causal and acasual signals could be clearly identified around  $\pm 80$  s with an apparent group velocity of 2.7 km/s. After 10 days of stacking, we observe a causal coherent signal appear around 50 s with an apparent velocity of 4.3 km/s, possibly representing body waves.



Figure 7. Improvements of correlation coefficients (a) and SNRs (b) for one-month stacking results. Black stars and open red circles are the results of the conventional stacking procedure and the proposed approach. The reference trace for calculating the correlation coefficients is the 30-day raw stack.



Figure 8. Distributions of the USArray stations in north Texas and Oklahoma used for comparisons in Figure 9.



Figure 9. Comparisons of one-month stacked CCFs averaged from negative and positive lags from conventional approach (a) and the proposed procedure (b) for station pairs shown in Figure 8.

## -34-



Figure 10. 35 seismic stations used for comparisons shown in Figures 11-13. The SM, SN and SS stations are parts of the Superior Province Rifting Earthscope Experiment (Wolin et al., 2015). Other stations come from the USArray Transportable Array. The background color is the Bouguer gravity anomaly (Woollard, 1965), where the linear feature with positive (blue) values highlights the extension of the Mid-continent Rift.



Figure 11. Comparisons of one-month stacked CCFs averaged from negative and positive lags (without bandpass filter) from conventional approach (a) and proposed procedure (b). The improvement of averaged SNR for each trace is shown in panel (c) with black (conventional) and red (proposed) dots. Panels d-f are the same as panels a-c except for two-month stacking. The short magenta and blue solid lines in panel (d) represent the signal and noise windows used to calculate the SNRs shown in panels c and f.



Figure 12. The same setting as Figure 11, but for the stacked data bandpass filtered between 6-15 s.



Figure 13. Same as Figure 11 but from the stacked data bandpass filtered between 15-30 s.



Figure 14. (a) Station distribution of one dense array deployed in the San Bernardino basin. The red triangle denotes the master station while the other black ones are stations used to calculate the CCFs. The red dots represent faults from the U.S. Geological Surveys. Shot gathers of 5 days (b) linear, (e) proposed local and (d) root-mean-square ratio based selective (RMSRS) (Xie et al., 2020) stacking. The blue and dark green lines in panels c and e are acausal and causal arrival times with a group velocity of 1.5 km/s and 3.5 km/s, respectively. They are used to highlight the potential ranges of Rayleigh wave arrivals. The blue and dark green lines in panels d and g are used to define the signal windows for root-mean-square ratio calculation. Panels (e-g) display the corresponding 30-day stacking results.



Figure 15. Comparison of selected phase dispersion curves from 5-day raw (a), proposed (b) and RMSRS (c) stacking CCFs displayed in Figures 14b-d. Panels (b), (d) and (f) show the number of measurements for every 0.5 s from 1.0 s to 15.0 s.



Figure 16. Similar to Figure 15 but from 30-day stacking CCFs displayed in Figures 14e-g



**Figure 17.** Comparisons of convergence for conventional (a), the proposed procedure (b) and root-mean-square-ratio selective (RMSRS) (Xie et al., 2020) stacking of the TA.A12A-TA.A18A station pair. Panels (d-e) compare 15-day stacked CCFs of the three stacking methods at 10-20, 15-30 and 20-35 s period bands. The blue numbers are the SNR for the corresponding trace. To calculate the SNR, we choose 57-221 s as the signal window and 221-357 s as the noise window. The magenta short lines (57-221 s) are used to define the signal windows for RMSRS calculation. The green dashed rectangles are used to highlight the improvement of the RMSS stacking method. Both RMSRS stacking and our proposed approaches help us to improve the SNR at different period bands.

Figure 1.



Figure 2.



Figure 3.



0.1

0.4

Figure 4.







Figure 5.



Figure 6.





Figure 7.



Figure 8.


Figure 9.

# Raw stack



# Stack with local attributes



Figure 10.





Figure 11.

## Raw stack

Stack with local attributes



avg\_snr=45.0093

	Α
400	to many and a state of the first of the state of the stat
200	An Mar and Mar and a second and
	mMMmmmmmmmmmmmmmmmmmmmmmmmmmmmmmmmmmmmm







Figure 12.

## Raw stack

Stack with local attributes













Figure 13.

### Raw stack

Stack with local attributes



Figure 14.











Figure 15.

## Raw stack



Figure 16.

# Raw stack



Figure 17.



### Improving signal-to-noise ratios of ambient noise cross-correlation functions using local attributes

Bin He<sup>1</sup>, Hejun Zhu<sup>1,2</sup>, David Lumley<sup>1,2</sup>

<sup>1</sup>Department of Geosciences, The University of Texas at Dallas, Richardson, TX 75080, USA
 <sup>2</sup>Department of Physics, the University of Texas at Dallas, Richardson, TX 75080, USA

### 6 Key Points:

1

2

3

8

### Signal-to-noise ratio of ambient noise cross-correlation functions can be improved by using local cross-correlation and local similarity.

- The local cross-correlation function is efficiently implemented by solving the heat
   equation.
- Applications to broadband and high-frequency nodal arrays validate the effective ness of the proposed method.

Corresponding author: Hejun Zhu, hejun.zhu@utdallas.edu

### 13 Abstract

For seismographic stations with short acquisition duration, the signal-to-noise ra-14 tios (SNRs) of ambient noise cross-correlation functions (CCFs) are typically low, pre-15 venting us from accurately measuring surface wave dispersion curves or waveform char-16 acteristics. In addition, with low-quality CCFs, it is difficult to monitor temporal vari-17 ations of subsurface physical states or extract relatively weak signals such as body waves. 18 In this study, we propose to use local attributes to improve the SNRs of ambient noise 19 CCFs, which allows us to enhance the quality of CCFs for stations with limited acqui-20 sition duration. Two local attributes: local cross-correlation and local similarity, are used 21 in this study. The local cross-correlation allows us to extend the dimensionality of daily 22 CCFs with computational costs similar to global cross-correlation. Taking advantage of 23 this extended dimensionality, the local similarity is then used to measure non-stationary 24 similarity between the extended daily CCFs with a reference stacking trace, which en-25 ables us to design better stacking weights to enhance coherent features and attenuate 26 incoherent background noises. Ambient noise recorded by several broadband stations from 27 the USArray in North Texas and Oklahoma, the Superior Province Rifting EarthScope 28 Experiment in Minnesota and Wisconsin and a high-frequency nodal array deployed in 29 the San Bernardino basin are used to demonstrate the performance of the proposed ap-30 proach for improving the SNR of CCFs. 31

32

### Plain Language Summary

Seismic ambient noise has been widely used for imaging and monitoring subsur-33 face structures by using cross-correlation functions (CCFs) of continuous recordings be-34 tween station pairs. Typically, we have to stack a lot of CCFs to enhance signals (e.g., 35 surface and body waves) of CCFs. However, for temporal monitoring purposes or those 36 deployed arrays with short acquisition duration, it is impossible to have a lot of stack-37 ing. The lack of enough stacking could result in low signal-to-noise ratios (SNRs) of the 38 CCFs. We proposed a new approach to improve the SNRs of CCFs by using two local 39 attributes: local cross-correlation and local similarity. We first extend the CCFs to a higher 40 dimension by using local cross-correlation, and then the local similarity is used to de-41 fine a better weighting factor for the final stacking. We use several field data examples 42 to prove the effectiveness of the proposed approach. 43

### 44 1 Introduction

With the assumption of homogeneously distributed noise sources, both theoreti-45 cal and experimental studies have demonstrated that empirical Green's functions can 46 be retrieved by cross-correlating continuous ambient noise records between two seismo-47 graphic stations (Aki, 1957; Claerbout, 1968; Buckingham et al., 1992; Lobkis & Weaver, 48 2001; Weaver & Lobkis, 2001; Shapiro & Campillo, 2004; Wapenaar, 2004; Nakata et al., 49 2019). Both surface (Campillo & Paul, 2003; Shapiro & Campillo, 2004; Sabra et al., 2005) 50 and body wave signals (Draganov et al., 2009; Zhan et al., 2010; Poli et al., 2012; Lin 51 et al., 2013; Nakata et al., 2015) have been successively extracted from ambient noise record-52 ings. With empirical Green's functions between each pair of stations in a seismic array, 53 we are able to measure surface wave phase or group dispersion curves, and then perform 54 surface wave tomography to estimate seismic properties in the subsurface. Compared 55 with earthquake tomography, ambient noise tomography allows us to image tectonically 56 inactive regions and achieve better cross-path coverages. To date, ambient noise cross-57 correlation functions (CCFs) have been successively used to investigate velocity as well 58 as anisotropic structures within the crust and uppermost mantle (Shapiro et al., 2005; 59 Y. Yang et al., 2007; Lin et al., 2008; Moschetti et al., 2010; Yao et al., 2010; Huang et 60 al., 2010; Lin et al., 2011; Shen & Ritzwoller, 2016). Recently, ambient noise CCFs are 61 combined with waveform inversion to better constrain subsurface velocity structures (Gao 62 & Shen, 2014; M. Chen et al., 2014; Lee et al., 2014; Y. Liu et al., 2017; Sager et al., 2018; 63 K. Wang et al., 2018; Zhu, 2018; Sager et al., 2020; Fan et al., 2022; Maguire et al., 2022). 64 Taking advantage of continuous records, ambient noise CCFs can also be utilized to mon-65 itor temporal evolution of physical states in the subsurface (Brenguier et al., 2008; Nakata 66 & Snieder, 2011; Hadziioannou et al., 2011; Mainsant et al., 2012; De Plaen et al., 2016; 67 Q.-Y. Wang & Yao, 2020; Le Breton et al., 2021; Mao et al., 2022). 68

Previous studies have demonstrated that we are able to improve the signal-to-noise 69 ratios (SNRs) of ambient noise CCFs as well as reducing the effect of source direction-70 ality by increasing the stacking duration of signals (Bensen et al., 2007). To obtain high-71 quality CCFs, typically we have to use long continuous records and sometimes have to 72 stack data with several years of acquisition. However, for some temporary experiments 73 e.g., Brenguier et al. (2008); Issa et al. (2017); G. Liu et al. (2018); Dougherty et al. (2019); 74 G. Chen et al. (2023); Wu et al. (2023), the SNRs of ambient noise CCFs are typically 75 low, preventing us to measure robust surface wave dispersion curves or waveform char-76

-3-

acteristics. For these cases, we have to design effective approaches to improve the SNRs 77 of ambient noise CCFs. In addition, the SNRs are crucial for extracting weak coherent 78 signals, such as body waves (Snieder, 2004; Nakata et al., 2015, 2016). There are sev-79 eral studies towards improving the SNR of ambient noise CCFs. For instance, Baig et 80 al. (2009) designed better-stacking weights in the time-frequency domain with discrete 81 orthogonal S transform, which allowed them to measure robust Rayleigh and Love ar-82 rivals for stations with long offsets. G. Li et al. (2018) further compared the performances 83 of time and frequency domain inverse S transforms for this stacking procedure. Schimmel 84 et al. (2011) used instantaneous phase coherence to avoid strong amplitude arrivals, such 85 as earthquake signals, and enhance coherent features in noise records. Seats et al. (2012) 86 and Clarke et al. (2011) utilized overlapped moving windows (Welch's method) to im-87 prove the convergence of CCFs towards stable Green's functions. In addition, wavelet 88 and curvelet transforms have also been applied to denoise CCFs by Stehly et al. (2011) 89 and Mao et al. (2022). Furthermore, Weaver and Yoritomo (2018) proposed several schemes 90 to choose optimal weights for stacking so that the effective incident intensity distribu-91 tion is closer to isotropic. Xie et al. (2020) used the root-mean-square ratio to remove 92 those CCFs with low SNR. A systematic evaluation and comparison of the performance 93 of several stacking methods is discussed by X. Yang et al. (2023). 94

In this study, we propose to use local attributes (Rickett & Lumley, 2001; Hale, 2006; 95 Fomel, 2007a) to denoise CCFs when we have short acquisition durations. In compar-96 ison with global attributes, the local attributes enable us to extract non-stationary char-97 acteristics in seismic data. There are a variety of local attributes, including local sim-98 ilarity (Fomel, 2007a), local cross-correlation (Hale, 2006), local frequency (Fomel, 2007a) 99 and local skewness (Fomel & van der Baan, 2014), etc. In this study, we use local cross-100 correlation (Hale, 2006) to extend the dimensionality of daily CCFs, and then use stack-101 ing weights measured by local similarity to improve the stacking quality of CCFs. The 102 local cross-correlation can be considered as a natural extension of the Welch's method 103 (Seats et al., 2012) but with longer overlap windows and much lower computational costs. 104 Its overall computational cost is similar to the conventional global cross-correlation. The 105 local similarity has been used to measure non-stationary similarity between time-lapse 106 images (Fomel & Jin, 2009), improve stacking quality of normal moveout data (G. Liu 107 et al., 2009) as well as angle-domain common-imaging-gathers (Lin et al., 2011). All these 108

-4-

studies demonstrate the capability of using local similarity to improve the stacking qual-ity.

We first review local cross-correlation and local similarity. Next, they are used to improve the SNRs of ambient noise CCFs with one-day and one-month acquisition durations. Continuous noise records from several broadband USArray stations in North Texas and Oklahoma, the Superior Province Rifting EarthScope Experiment in Minnesota and a high-frequency nodal array in the San Bernardino basin are used to demonstrate the performance of the proposed approach.

### 117 2 Method

118

### 2.1 Local cross-correlation

The global cross-correlation function  $c(\tau)$  between two signals f(t) and g(t) can be defined (Nakata & Snieder, 2011; Harris et al., 2020) as follows:

$$c(\tau) = \int_{-\infty}^{\infty} f(t+\tau)g(t)dt \quad .$$
(1)

The result of the global cross-correlation is a function of time lag  $\tau$ , which cannot cap-119 ture non-stationary time shifts between two input signals and therefore makes it diffi-120 cult to separate signals and noises. In order to fulfill this goal, one way is to perform time-121 windowed cross-correlation, which has much higher computational costs in comparison 122 to Equation 1. In addition, there might be leakage and edge problems due to the selected 123 window functions. To solve these problems, Hale (2006) designed an efficient algorithm 124 to compute the local CCF and measure time-varying correlations between two signals. 125 It was first used to measure non-stationary warpings between time-lapse migration im-126 ages. Later, the local cross-correlation was applied to measure non-stationary travel time 127 differences between two seismograms, which can be utilized to constrain subsurface ve-128 locity structures through full-waveform inversion (Díaz & Sava, 2015). 129

The idea of local cross-correlation is similar to the moving window cross-correlation for capturing transient time shifts, but with much higher computational efficiency. In addition, it helps us to separate signal and noise in a higher dimension. Taking advantage of several properties of Gaussian windows, such as the product of two Gaussian windows is still a Gaussian function, the computational cost of local cross-correlation can be reduced as similar to that of global cross-correlation. For instance, the Gaussian-windowed

-5-

versions of signals f(t) and g(t) can be represented as

$$\hat{f}(t,t_0) = f(t)\omega(t_0 - t)$$
 ,  
 $\hat{g}(t,t_0) = g(t)\omega(t_0 - t)$  . (2)

where  $\omega(t_0-t)$  is the Gaussian window function located at  $t_0$ , i.e.,  $\omega(t_0-t) = \frac{1}{\sqrt{2\pi\sigma}}e^{-(t_0-t)^2/2\sigma^2}$ .

Here  $\sigma$  is the standard deviation, which controls the width of the Gaussian window. With

these windowed signals, we are able to compute the cross-correlation function at each

 $_{140}$  local time step  $t_0$ . Without optimization, the computational cost of this moving window

cross-correlation is  $O(N_l N_w N_s)$ . Here,  $N_l$ ,  $N_w$  and  $N_s$  are the numbers of samples for

time lags, Gaussian window, and input signals, respectively.

<sup>143</sup> The windowed cross-correlation can be written as

$$c(t_0,\tau) = \int_{-\infty}^{\infty} f(t+\tau,t_0+\tau)g(t,t_0)dt$$
  
$$= \int_{-\infty}^{\infty} f(t+\tau)\omega(t_0-t)g(t)\omega(t_0-t)dt$$
  
$$= \int_{-\infty}^{\infty} f(t+\tau)g(t)W(t,t_0)dt \quad , \qquad (3)$$

144 where

$$W(t,t_0) = \omega(t_0 - t)\omega(t_0 - t) = \frac{1}{2\pi\sigma^2}e^{-(t_0 - t)^2/\sigma^2} \quad .$$
(4)

 $W(t, t_0)$  is again a Gaussian function with a smaller width  $\sigma/\sqrt{2}$  compared to  $w(t_0 - t)$ . Based on this property, we can first compute the pre-stacked cross-correlation between two signals, i.e.,  $h(t, \tau) = f(t+\tau)g(t)$ . It is then filtered by a Gaussian function  $W(t, t_0)$ as

$$c(t_0,\tau) = \int_{-\infty}^{\infty} h(t,\tau) W(t,t_0) dt \quad .$$
(5)

Combining with some efficient recursive Gaussian filter algorithms (Deriche, 1993; Alvarez & Mazorra, 1994; Van Vliet et al., 1998), the computational cost of implementing Equation 3 can be similar as the global cross-correlation shown in Equation 1, i.e.,  $O(N_l N_s)$ . This makes it attractive for processing large-scale datasets, such as continuous ambient noise records.

154

### 2.2 Recursive Gaussian convolution

How to efficiently compute the convolution of a signal or an image with a Gaussian function is a common problem in data processing. There are several ways to perform this filtering procedure. For instance, by taking advantage of the Fourier transform, Gaussian convolution can be computed efficiently as elemental-wise multiplication in the

<sup>159</sup> frequency domain. Another efficient way to implement the Gaussian filter is to consider

the result as the solution of the following heat equation:

$$\frac{\partial u}{\partial y} - \frac{\partial^2 u}{\partial x^2} = 0 ,$$

$$u(x,0) = f(x) .$$
(6)

Here, if we let  $y = \sigma^2/2$ , then the final solution of the above heat equations u(x, y) is

equivalent to the convolution of signal f(x) with a Gaussian function  $\omega(x) = \frac{1}{\sqrt{2\pi\sigma}} e^{-x^2/2\sigma^2}$ .

<sup>163</sup> This heat equation can be solved by using finite-difference schemes. Alvarez and Mazorra

<sup>164</sup> (1994) used this property to derive an efficient procedure to compute the Gaussian con-

volution. For instance, they discretized the simulation domain with

$$u_i^j = u(i\Delta x, j\Delta y) \quad , \tag{7}$$

Here  $\Delta x$  and  $\Delta y$  are the grid spacings along the spatial and time directions, *i* and *j* are the associate indices. Then, the finite-difference solution of the above heat equation can be written as

$$\frac{u_i^{j+1} - u_i^j}{\Delta t} - \frac{u_{i+1}^{j+1} + u_{i-1}^{j+1} - 2u_i^{j+1}}{\Delta x^2} = 0 \quad .$$
(8)

169 Let  $\lambda = \Delta y / \Delta x^2$ , then

$$(1+2\lambda)u_i^{j+1} - \lambda u_{i-1}^{j+1} - \lambda u_{i+1}^{j+1} = u_i^j \quad .$$
(9)

<sup>170</sup> We can solve this implicit finite-difference equation using the following filter

$$H(z) = \frac{\nu}{\lambda} \frac{1}{(1 - \nu z^{-1})(1 - \nu z)} \quad , \tag{10}$$

171 with

$$\nu = \frac{1 + 2\lambda - \sqrt{(1 + 4\lambda)}}{2\lambda} \quad . \tag{11}$$

<sup>172</sup> The above filtering procedure can be efficiently solved by using the following three steps,

marching from  $u_i^j$  to  $u_i^{j+1}$  as

$$u_{i}^{j'} = u_{i}^{j} + \nu u_{i-1}^{j'} ,$$
  

$$u_{i}^{j''} = u_{i}^{j'} + \nu u_{i+1}^{j''} ,$$
  

$$u_{i}^{j+1} = \frac{\nu}{\lambda} u_{i}^{j''} .$$
(12)

- <sup>174</sup> The first and second steps in the above procedure are the applications of causal and acausal
- <sup>175</sup> filters, respectively. The third step is just a simple multiplication. Thanks to the uncon-
- ditional stability of this implicit solution (Alvarez & Mazorra, 1994), only a few itera-
- tions (four iterations are used in this study to propagate along the time dimension y)

enable us to converge to the solution of the Gaussian filter. Therefore, the time grid can be defined by  $\Delta y = \sigma^2/(2N_y)$ , where  $N_y$  is the total number of iterations. Combining this recursive Gaussian filter with Equation 5, we are able to efficiently compute the local cross-correlation function. For example, for a certain  $\tau$ , the source term for Equation 6 is the pre-stacked cross-correlation  $h(t, \tau)$ . Then solution of the corresponding heat equation is the non-stationary CCFs  $u(t, \tau)$  at this particular  $\tau$ .

Figure 1 shows a simple example to illustrate the advantages of local cross-correlation 184 over global cross-correlation. There are non-stationary time shifts for three events at 1, 185 3 and 5 s between two input signals. The global cross-correlation is shown in Figure 1b. 186 Although it captures the correct time shifts around -0.4 s, 0.1 s and 0.3 s, it does not 187 carry any useful information about the locations of these time shifts. The local cross-188 correlation is presented in Figure 1c, with three energy spots located at the right loca-189 tions with correct time lags. This local cross-correlation function enables us to extend 190 the dimensionality of the input signals as well as the cross-correlation function, which 191 can be utilized to design a better strategy to improve stacking quality as discussed be-192 low. 193

<sup>194</sup> 2.3 Local similarity

Local similarity is a local attribute that allows us to measure non-stationary similarity between two time series or images (Fomel & Jin, 2009; G. Liu et al., 2009, 2011). First, the global similarity between two signals f(t) and g(t) can be defined as follows

$$s = \frac{\int f(t)g(t)dt}{\sqrt{\int f^2(t)dt}\sqrt{\int g^2(t)dt}} \quad .$$
(13)

In comparison with the global cross-correlation in Equation 1, the global similarity can be considered as a zero-lag cross-correlation function normalized by the energies of each signal. Therefore, it is insensitive to the absolute amplitudes of input signals. Its magnitude ranges from -1 to 1. Similar to local cross-correlation, local similarity was designed to measure non-stationary similarity between time-lapse images (Rickett & Lumley, 2001; Fomel & Jin, 2009). Once taking the square of Equation 13 on both sides, we can write its discretized version as

$$s^{2} = \frac{(\mathbf{f}^{T}\mathbf{g})(\mathbf{g}^{T}\mathbf{f})}{(\mathbf{f}^{T}\mathbf{f})(\mathbf{g}^{T}\mathbf{g})} \quad .$$
(14)

Fomel (2007a) considered the above equation as the solution to the following two least-

206 squares problems:

$$\mathbf{s}_{1} = \arg \min_{\mathbf{s}_{1}} |\mathbf{F}\mathbf{s}_{1} - \mathbf{g}| ,$$
  
$$\mathbf{s}_{2} = \arg \min_{\mathbf{s}_{2}} |\mathbf{G}\mathbf{s}_{2} - \mathbf{f}| .$$
(15)

Here,  $\mathbf{F}$  and  $\mathbf{G}$  are the diagonal operators constructed from the elements of input vec-

tors f and g, respectively. Shaping regularization (Fomel, 2007b) and a conjugate gra-

dient method can be used to solve the above least-squares problems. Then, the local sim-

ilarity can be computed as elemental-wise multiplication between  $s_1$  and  $s_2$  as

$$\mathbf{s} = \sqrt{\mathbf{s}_1 \mathbf{s}_2} \quad . \tag{16}$$

As pointed out in Fomel (2007a), the normalized local correlation can be considered as the first iteration from local similarity. Compared to a scalar value for the global similarity in Equation 13, the local similarity is a vector with the same length as the input signals. It can also be utilized to measure the similarity between two 2D or 3D images with high efficiency.

216

### 2.4 Stacking strategy with local similarity

From local cross-correlation, we obtain a function  $c(t, \tau)$ . The simplest way to calculate the CCF  $c_1(\tau)$  is to use the following linear stacking

$$c_1(\tau) = \frac{1}{N} \sum_t c(t,\tau)$$
 (17)

If we use the above expression to compute noise CCF, its result is similar to the direct implementation of Equation 1. In this study, we choose  $c_1(\tau)$  as a reference trace and then compute the local similarity  $s(t, \tau)$  between  $c_1(\tau)$  and  $c(t, \tau)$ . Once with the local similarity, we can use the following strategy to improve the stacking quality

$$c_2(\tau) = \frac{1}{N} \sum_t \alpha(t,\tau) c(t,\tau) \quad , \tag{18}$$

223 with

$$\alpha(t,\tau) = \begin{cases} s(t,\tau) - \alpha_0 &, & \text{if } \alpha(t,\tau) > \alpha_0 \\ 0 &, & \text{if } \alpha(t,\tau) \le \alpha_0 \end{cases}$$
(19)

Here  $\alpha_0$  is a pre-defined parameter to decide the stacking level. This stacking procedure has been applied to enhance coherent signals for normal moveout corrected data (G. Liu et al., 2009) and angle-domain common-image-gathers (G. Liu et al., 2011). In this study, we use it to improve the stacking quality of noise CCFs. In practice,  $\alpha(t, \tau)$  is a soft thresh-

- <sup>228</sup> old solved in the seislet domain (Daubechies et al., 2004; G. Liu et al., 2009). Compared
- with the similarity shown in Figure 1d, the clipped similarity ( $\alpha(t,\tau)$  in Eq. 19) in Fig-
- <sup>230</sup> ure 1f shows zero background values away from signal regions, while keeping almost 1.0
- weights around the signal regions. This helps to suppress noises away from the signal
- regions. Because this synthetic test does not involve noises, the weighted prestack local
- correlations ( $c_2(\tau)$  in Eq. 18) shown in Figure 1e are basically identical to the prestack
- local correlations shown in Figure 1c. To validate the effectiveness of the proposed method
- to improve the SNR, we give its definition used in this study as

$$SNR = \sqrt{\frac{\frac{1}{N_s} \int_{s_1}^{s_2} f^2(t) dt}{\frac{1}{N_n} \int_{n_1}^{n_2} f^2(t) dt}} \quad .$$
(20)

Where  $[s_1, s_2]$  and  $[n_1, n_2]$  represent signal and noise windows,  $N_s$  and  $N_n$  denote the number of time samples for signal and noise, respectively.

238 **2.5 Workflow** 

In summary, the proposed method based on two local attributes can be implemented 239 through the following six steps: 240 1. Download and preprocess continuous ambient noise data for each station, and cut 241 them into daily recordings. 242 2. Calculate the pre-stacked cross-correlation function  $h(t,\tau) = f(t+\tau)g(t)$  as shown 243 in Equation 3 between daily recordings from station pairs. 244 3. Stack all daily global stacking CCFs and extend it along the time axis to obtain 245 two-dimensional data as a reference. 246 4. Solve the heat equation in Equation 6 to obtain a daily local CCF. Note here, the 247 pre-stacked CCF is the source term f(x) in Equation 6, and the local CCF is also 248 two-dimensional data. 249 5. Calculate the daily local similarity between the reference CCF (step 3) and the 250 daily local CCF(step 4). 251 6. Stack all daily local CCFs weighted by the local similarity according to Equation 252 18 for each station pair. 253

### <sup>254</sup> 3 Numerical examples

We download continuous ambient noise data, which are then processed by using the NoisePy package (Jiang & Denolle, 2020) following standard procedures from Bensen et al. (2007). It includes deconvolution of instrument responses, bandpass filter (2–10 s), time domain normalization and spectral whitening. The preprocessed data are used to retrieve the CCFs and demonstrate the advantages of the proposed method.

260

### 3.1 One-day records

Figures 2a and b are vertical component displacements for the USArray stations 261 TA.234A and TA.Z36A recorded on October 10th, 2010. The inter-station between this 262 station pair is 213 km. As shown in Figure 3a, the SNR of the global cross-correlation 263 is relatively low because we only use daily records. Only causal Rayleigh waves can be 264 clearly observed in this CCF. Both amplitude and time-frequency spectra of the global 265 cross-correlation function are presented in Figures 3b and c, respectively. Besides the dom-266 inant Rayleigh wave arrival, there are many background noises, especially from 0.2 to 267 0.5 Hz. The local cross-correlation between Figures 2a and b is shown in Figure 2c. With 268 the local cross-correlation, we are able to extend the 1D CCF to a 2D image, which can 269 be used to measure local similarity and design a better strategy to improve stacking qual-270 ity. Although there are many incoherent background noises in Figure 2c, it includes a 271 vertical coherent feature around time lags between 60 to 90 s. If we use a simple linear 272 stacking for Figure 2c along the time axis (Equation 17), we will obtain the same sig-273 nal as the global cross-correlation shown in Figure 3a. Next, we compute the local sim-274 ilarity between Figure 2c and the raw stacked CCF shown in Figure 3a. In the local sim-275 ilarity result (Figure 2d), the coherent signals between 60 to 90 s stand out in compar-276 ison to the background noise level. With the stacking weights based on the local sim-277 ilarity (Equation 18), the new CCF and its amplitude/time-frequency spectra are pre-278 sented in Figures 3d–f. Compared with the global CCF in Figures 3a–c, the incoherent 279 noises are significantly reduced without changing the dominant Rayleigh wave arrival. 280 In the meanwhile, the overall amplitude spectrum does not change before and after ap-281 plying the stacking weights (Figures 3b and e). Now, it is much easier for us to measure 282 the phase or group dispersion curves of Rayleigh waves for the newly stacked signal (Fig-283 ure 3d). 284

### 3.2 One-month records

Figures 4a and b present ambient noise CCFs for daily records (October 2010) by 286 using global cross-correlation and local attributes, respectively. With these local attributes, 287 incoherent background noises are attenuated while surface wave signals are preserved and 288 much easier for us to measure. We observe alternative causal and acausal surface wave 289 signals, which might come from the changes in ambient noise source distributions. Since 290 there is data redundancy for Figure 4b, instead of directly stacking it over the time axis, 291 we calculate local similarity again between the raw stacked CCF (Figure 5a) with Fig-292 ure 4b, which is shown in Figure 4c. Again, coherent signals, such as surface wave pack-293 ages, stand out in comparison with background noises. Figures 5a and d compare the 294 raw stacked CCF and the new signal with local attributes. We observe significant im-295 provement of the new signal in terms of SNR. Based on the time-frequency analysis shown 296 in Figures 5c and f, the dispersive characteristics of Rayleigh waves are preserved while 297 background noises from 0.1 to 0.3 Hz are attenuated. Similar to the previous daily ex-298 ample, the overall amplitude spectra do not change too much between these two CCFs. 299

Next, we compare the convergence of the raw global cross-correlation and the new 300 result based on local attributes. Figure 6a shows the convergence of the raw daily global 301 CCFs over one month. With the increasing stacking duration, the SNR is improved and 302 coherent Rayleigh wave signals stand out gradually. However, the convergence and im-303 provement of SNR are relatively low (Figure 7b). Here we use a one-month stacked CCF 304 as the reference trace to compute the similarity between two signals (Equation 13). In 305 this test, the signal window is [-100, -50] s and [50, 100] s, while the noise window ranges 306 from [-200, -150] s and [150, 200] s. Improvements for the local attribute stacking CCF 307 are shown in Figure 6b. In addition, another causal signal appears around 50 s after 10 308 days of stacking, which could be body waves or higher mode surface waves with an ap-309 parent velocity about 4.3 km/s. But those incoherent events (e.g., 20 s) are relatively 310 well suppressed. We also speculate that the improvement of SNR with our method is sig-311 nificant with only several days of stacking (Figure 7b). By stacking with more than 10 312 days, both the correlation coefficient and SNR gradually become stable (Figure 7). It 313 is intriguing that the first 5 days CCFs show strong acausal signals (e.g., -80 s in Fig-314 ures 6a-b) and our improved SNR shows a bump around 5 days (Figure 7b). This might 315 indicate the noise sources were mainly from station Z36A to station 234A before the first 316 5 days. Then more complex noise sources from stations 234A to station Z36A appeared 317

-12-

and resulted in stronger causal signals (e.g. 50-80 s) and complex noises in CCFs. As
a result, the SNR goes lower from 5-12 day stacking and gradually turns high with longer
stacking.

We further test the proposed approach for 20 USArray stations deployed in North 321 Texas and Oklahoma (Figure 8). One-month data (October 2010) are used for both raw 322 stacking and the new approach. Comparisons of these two results can be found in Fig-323 ure 9. We observe the SNRs of the new approach are much higher in comparison with 324 the classical stacking approach. Except for the dominant Rayleigh waves, we also no-325 tice that there might be additional earlier arriving weaker events emerging from 500 km 326 distance (highlighted by dark blue arrows). For example, the early arrival at about 640 327 km has an apparent velocity of 3.5 km/s, which is higher than the group velocity ( $\sim 2.6$ 328 km/s) of the dominant Rayleigh waves. It could be a candidate for head/diving body 329 waves or higher mode surface waves contaminated with noises. Even they are weak sig-330 nals, our method can retain them as long as they are coherent. 331

Next, we test the proposed approach for several stations from three dense arrays 332 and surrounding stations deployed in Minnesota and Wisconsin around the Midconti-333 nent Rift (Figure 10; Wolin et al., 2015). Figure 11 shows the comparisons of one-month 334 stacked CCFs based on the conventional approach and the proposed procedure. It is ob-335 vious the SNRs are improved quite a lot for most traces shown in Figure 11c with the 336 proposed method. The overall average of the SNR from the traditional stacking method 337 (4.8) is much smaller than the one (35.4) from the proposed method. Although the noise 338 level between 6-15 s period is lower, the SNRs are improved from 6.7 to 53.8 by using 339 the proposed stacking method as shown in Figure 12. It is interesting that SNRs for data 340 in 15-30 s period band (Figure 13) are smaller compared to those in the periods of 6-15 341 s, and the proposed method helps us to improve the SNRs as expected. Panels d-f show 342 the improved SNRs from two-month stacking compared with the monthly stacking (pan-343 els a-c) of Figures 11-13. We speculate that with a two-month stacking duration, the im-344 provements of SNRs using conventional linear stacking are larger than the proposed method, 345 but the SNRs from one-month stacking of the proposed method are still higher than those 346 from two-month stacking with the conventional method. This is important for monitor-347 ing time-lapse changes of near-surface velocity changes with temporary arrays which usu-348 ally have quite short acquisition durations (Nakata et al., 2016; Issa et al., 2017; Mor-349 dret et al., 2020; Zhang et al., 2022). 350

-13-

351

### 3.3 Applications to a nodal array

Between 2017 and 2019, 10 linear dense Distribution of Basin Amplification Seis-352 mic Investigation (BASIN) nodal arrays (SG1–SG4, and SB1–SB6) were deployed in the 353 San Gabriel and San Bernardino basins for Fine characterization of basin shapes and depths 354 (Y. Li et al., 2022; X. Wang et al., 2021). We apply our method to the SB1 array to val-355 idate its performance. The SB1 array (Figure. 14a) consists of 239 Fairfield ZLand nodes 356 with standard 5 Hz 3-component geophones with spatial sampling of  $\sim 250$  m. It was de-357 ployed for approximately one month. The basin depth beneath the SB1 array is about 358 0-3.0 km (Y. Li et al., 2022), therefore, it is essential to have more measurements at lower 359 period bands, such as 1.0-5.0 s. The data downloading and preprocessing are similar to 360 previous tests except the continuous noises were down-sampled with a sampling frequency 361 of 4 Hz and then bandpass filtered between 1-20 s. 362

Taking the first station as the master station, the corresponding CCFs are arranged 363 according to their offset and displayed in Figure 14. As expected, with 5 days of stack-364 ing (Figure 14b), the CCFs from traditional linear stacking have strong noises for all sta-365 tion pairs. The 30-day stacking clearly improves the data quality so that we can observe 366 the dominant acausal Rayleigh waves. Compared with CCFs from linear stacking, the 367 proposed method helps us to enhance coherent signals even with only 5 days of stack-368 ing. The 30-day stacking further improves the data quality. Next, we measure phase dis-369 persion curves for the stacked CCFs, which are the input for surface wave tomography 370 (Yao et al., 2006; Fang et al., 2015). There are three criteria to make sure the measured 371 dispersion curves are stable. (1) The SNR is larger than 5.0; (2) The inter-station dis-372 tance is larger than 1.5 wavelength at corresponding periods (Bensen et al., 2007; G. Chen 373 et al., 2023). (3) The picked phase velocity is within  $\pm 12\%$  of a 3D reference phase ve-374 locity model. Here the 3D reference phase velocity model is constructed from a local to-375 mographic model, CVM-S 4.26 (Lee et al., 2014), by conducting a forward modeling based 376 on the fast-marching method (Rawlinson & Sambridge, 2004; Fang et al., 2015). Thanks 377 to the improvement of SNR from our method, more high-quality dispersion curves pass 378 the selection criteria with 5-day stacking compared with those measured from linear stack-379 ing as illustrated in Figure 15. More importantly, the number of measurements for 1-380 5 s and 10-15 s period bands are also increased. These are further improved for 30-day 381 stacked CCFs using the proposed method. 382

-14-

### **4 Discussions**

Commonly used ambient noise processing procedures (Bensen et al., 2007) require stacking over long, continuous records to enhance coherent signals, also require the assumption of evenly distributed ambient sources in order to obtain good estimates of Green's functions. However, both conditions impose constraints on the application of temporary arrays with short acquisition duration.

To improve the SNR of CCFs stacked with a short duration, Xie et al. (2020) pro-389 posed a root-mean-square-ratio selective (RMSRS) stacking procedure to remove those 390 CCFs that negatively contribute to the SNR of the final stacked CCF. It is realized by 391 comparing the root-mean-square ratio of signals and noises for each CCF and the stacked 392 CCF. Therefore, its effectiveness depends on an accurate definition of signal/noise win-393 dows. Here, we compare the conventional linear stacking, proposed procedure and RM-394 SRS stacking for the station pair TA.A12A-TA.A18A at 10-35 s period band. The sig-395 nal window (Figure 17c) is defined with the reference time  $t_{ref}$  and maximum period of 396 interest  $T_{max}$  as  $[t_{ref} - 2T_{max}, t_{ref} + 2T_{max}]$ . Here, the reference time  $t_{ref} = d/v_{ref}$  is 397 defined by the inter-station distance d (425 km) and a reference group velocity  $v_{ref}$  (3.0 398 km/s). The noise window is defined from 0 to the signal windows and  $4T_{max}$  out of the 399 signal windows (Xie et al., 2020). Compared with linear stacking, the RMSRS stacking 400 successfully suppresses the noises within the defined noise windows and therefore improves 401 the SNR as highlighted in Figures 17d and f. Similarly, our approach also successfully 402 attenuates those incoherent noises and significantly improves the SNR, which is about 403 two times the other two methods. However, the coherent signal at about 75 seconds with 404 an apparent velocity of 5.6 km/s is retained by our approach due to the similarity be-405 tween local CCFs and the stacked CCFs. The main reason that the RMSRS-based method 406 helps to suppress this signal is because they are selected as noises. Such selection seems 407 to be challenging to deal with low-SNR CCFs at shorter period bands (3-16 s) obtained 408 from the high-frequency nodal array (Figure 14). As illustrated in Figures 14d, the im-409 provement of selective stacking is limited compared with linear stacking. On the con-410 trary, our approach significantly improves the stacking quality and helps us to obtain more 411 high-quality dispersion curves. We note here that, for fair comparisons, the RMSRS-based 412 method is implemented on the raw stacked CCFs. For better performance, Xie et al. (2020) 413 suggested using RMSRS stacking at several narrow period bands so that it could define 414 a better selection window, which is out of the scope of this study. 415

-15-

Although the proposed approach helps us to improve the SNR of CCFs with only 416 several days of stacking, the non-causality and asymmetry, which mainly arise from the 417 uneven distribution of ambient noise sources, remain challenging. For tomography pur-418 poses, theoretical works indicate that phase velocities can be estimated from the empir-419 ical Green's functions, which are obtained by taking the negative time derivative of the 420 symmetric cross-correlation under the assumption of a spatially homogeneous ambient-421 noise source distribution (Lobkis & Weaver, 2001; Sabra et al., 2005; Snieder, 2004; Yao 422 et al., 2006; Lin et al., 2008). As suggested by Yao et al. (2006), inhomogeneous source 423 distribution may contribute to 1-3 percent inconsistency between phase velocity mea-424 surements and the traditional earthquake-based two-station method between periods of 425 20–30 s. Therefore, the "symmetric" CCF is usually taken by the average of the cross-426 correlation at positive and negative correlation lag times (Yao et al., 2006; Lin et al., 2008). 427 However, how much such averaged CCFs affect tomography results remains unknown. 428 On the other hand, several studies suggested measuring full-waveform differences of CCFs 429 and source location simultaneously (Tromp et al., 2010; Sager et al., 2018; Datta et al., 430 2023), which naturally mitigates the uncertainty caused by the source distributions. In 431 both cases, it is important for us to obtain high SNR CCFs, especially for short-duration 432 nodal arrays. In addition, because the high-quality phase velocity measurements obtained 433 from 5-day stacking (Figure 15c) by the proposed approach is comparable with those ob-434 tained from the 30-day linear stacking, the surface wave-based monitoring (Durand et 435 al., 2011; Brenguier et al., 2020) seems to be possible, albeit coda waves are mostly used 436 (Mao et al., 2019, 2020; Luo et al., 2023). Another potential application for our approach 437 could be weak coherent signal extraction, such as body waves (Zhan et al., 2010; Poli 438 et al., 2012; Nakata et al., 2015, 2016; Mao et al., 2020). Body waves extracted from am-439 bient noise CCFs have been proven to be capable of improving the imaging resolution 440 compared to surface wave tomography (Nakata et al., 2015). As illustrated in Figures 441 6 and 14, those coherent signals with an apparent velocity larger than 4.0 km/s might 442 be good candidates for body waves, although we are not able to rule out the possibil-443 ity of higher-order surface waves. 444

The last factor we need to consider is the computational and memory cost for our approach. Taking station pair TA.A12A and TA.A18A as an example, we have 30 days of recordings and 86400-time samples for each day. The cross-correlation time lag varies from -480 s to 480 s for every 1 s. We calculate the global cross-correlations every 60 min-

-16-

utes with 75% overlaps, which yields 91 subset CCFs for each day. They are then nor-449 malized and stacked to obtain the dayily CCF. Finally, 30 days of global cross-correlation 450 and linear stacking take 47.9 s and 0.02 s, respectively, while the local cross-correlation, 451 local similarity and weighted stacking take 196.7 s, 12,150 s and 78.3 s. We note here, 452 that the local cross-correlation solved with our method actually contains 86,400 CCFs 453 for each day. It helps us to extend the dimension of CCFs dramatically (534 times the 454 number of global CCFs), but with only four times the computational cost compared to 455 global stacking. Whereas, such a high dimension, in turn, greatly decreases the efficiency 456 of our approach. To mitigate this problem, the local CCFs are downsampled 60 times 457 (from 1.0 s to 60 s) by taking the average CCFs for every 60 samples, and then the com-458 putational cost for local similarity and weighted stacking is decreased to 167.1 and 10.0 459 s, respectively. The final stacked CCFs by using these local CCFs are almost the same 460 as the original local CCFs. Overall, the computational cost for the proposed approach 461 is 7.8 times the computational cost compared to the hourly global stacking, but with 15.8 462 times more CCFs. In addition, increasing the cross-correlation time window (e.g., from 463 one hour to three hours) does not affect the efficiency of our approach, but will increase 464 the computational time for global cross-correlation. Finally, we compare the computa-465 tion cost using the nodal array. For each station, we have 345,600 time samples per day 466 with a 4 Hz sampling frequency. The cross-correlation time lag varies from -120 s to 120 467 s for every 0.25 s. The global and local CCFs are calculated the same as in the previ-468 ous example. To save computational costs, we downsample the local CCFs 60 times (from 469 0.25 s to 15 s). Then all station-pairs are distributed to 72 CPU cores for parallel com-470 putations. Our approach takes 5792.0 s and the traditional linear stacking takes 792.0 471 s. The final stacked CCFs are compared in section 3.3. In conclusion, our approach sig-472 nificantly helps us to improve the SNR of CCFs, but with about 7.5 times the compu-473 tation cost compared to traditional linear stacking. Such extra computational costs are 474 bearable compared to the following computational costs tomography (Zhu, 2018; Wu et 475 al., 2023; G. Chen et al., 2023). 476

### 477 **5** Conclusion

In this study, by taking advantage of local attributes, we present a new approach
to increase the SNRs of ambient noise CCFs. Two local attributes are used in this study:
local cross-correlation and local similarity. The local cross-correlation is employed to ex-

-17-

tend the dimensionality of daily CCFs, and the local similarity is used to design better-481 stacking weights to enhance coherent signals and attenuate incoherent background noises. 482 Applications to ambient noise records from several broadband stations and a high-frequency 483 nodal array demonstrate the performance of the proposed approach. With higher SNRs, 484 we are able to extract more high-quality dispersion curves, which are important for sur-485 face wave tomography. In addition, 5-day stacking by our approach can produce CCFs 486 comparable to 30-day linear stacking in terms of SNRs, demonstrating its potential ap-487 plications for time-lapse monitoring. In addition, extracting coherent weak signals, such 488 as body waves, could be another application of the proposed approach. 489

### 490 Open Research

All seismic data used in this study can be obtained from the IRIS Data Management Center (https://ds.iris.edu/ds) under the network codes TA and XI. We use the Noisepy (https://noise-python.readthedocs.io/en/latest/) for parallel data downloading and preprocessing (Jiang & Denolle, 2020). The open software Madagascar (Fomel et al., 2013) download from (http://www.ahay.org) is used to calculate local similarity and plot figures. We also use PyGMT (Wessel et al., 2019) downloaded from (https://www.pygmt.org/latest/) for plotting figures.

498

### Author Contributions

The authors confirm their contribution to the paper as follows: study conception and design: Hejun Zhu; data collection: Hejun Zhu and Bin He; analysis and interpretation of results: Hejun Zhu, Bin He and David Lumley; draft manuscript preparation: Hejun Zhu, Bin He and David Lumley. All authors reviewed the results and approved the final version of the manuscript.

### 504 Acknowledgments

This paper is contribution no. \*\*\* from the Department of Geosciences at the University of Texas at Dallas. The numerical results are computed through the Optane nodes on the UTD Seismology Group HPC clusters. This research is partially supported by the sponsors of the UT Dallas 3D + 4D Seismic Full Waveform Inversion research consortium.
## 510 References

- Aki, K. (1957). Space and time spectra of stationary stochastic waves, with special
   reference to microtremors. Bulletin of the Earthquake Research Institute, 35,
   415–456.
- Alvarez, L., & Mazorra, L. (1994). Signal and image restoration using shock fil ters and anisotropic diffusion. SIAM journal on numerical analysis, 31(2),
   590–605.
- Baig, A. M., Campillo, M., & Brenguier, F. (2009). Denoising seismic noise cross
   correlations. Journal of Geophysical Research: Solid Earth, 114 (B8).
- <sup>519</sup> Bensen, G., Ritzwoller, M., Barmin, M., Levshin, A. L., Lin, F., Moschetti, M.,
- Yang, Y. (2007). Processing seismic ambient noise data to obtain reli able broad-band surface wave dispersion measurements. *Geophysical journal international*, 169(3), 1239–1260.
- <sup>523</sup> Brenguier, F., Courbis, R., Mordret, A., Campman, X., Boué, P., Chmiel, M., ...
- others (2020). Noise-based ballistic wave passive seismic monitoring. part 1: body waves. *Geophysical Journal International*, 221(1), 683–691.
- Brenguier, F., Shapiro, N. M., Campillo, M., Ferrazzini, V., Duputel, Z., Coutant,
   O., & Nercessian, A. (2008). Towards forecasting volcanic eruptions using
   seismic noise. *Nature Geoscience*, 1(2), 126–130.
- Buckingham, M. J., Berknout, B. V., & Glegg, S. A. (1992). Imaging the ocean with
   ambient noise. *Nature*, 356 (6367), 327–329.
- Campillo, M., & Paul, A. (2003). Long-range correlations in the diffuse seismic coda.
   Science, 299(5606), 547–549.
- <sup>533</sup> Chen, G., Chen, J., Tape, C., Wu, H., & Tong, P. (2023). Double-difference ad <sup>534</sup> joint tomography of the crust and uppermost mantle beneath alaska. Journal
   <sup>535</sup> of Geophysical Research: Solid Earth, 128(1), e2022JB025168.
- <sup>536</sup> Chen, M., Huang, H., Yao, H., van der Hilst, R., & Niu, F. (2014). Low wave speed
   <sup>537</sup> zones in the crust beneath se tibet revealed by ambient noise adjoint tomogra <sup>538</sup> phy. *Geophysical Research Letters*, 41(2), 334–340.
- <sup>539</sup> Claerbout, J. F. (1968). Synthesis of a layered medium from its acoustic transmis <sup>540</sup> sion response. *Geophysics*, 33(2), 264–269.
- <sup>541</sup> Clarke, D., Zaccarelli, L., Shapiro, N., & Brenguier, F. (2011). Assessment of res <sup>542</sup> olution and accuracy of the moving window cross spectral technique for mon-

543	itoring crustal temporal variations using ambient seismic noise. Geophysical
544	$Journal \ International, \ 186(2), \ 867-882.$
545	Datta, A., Shekar, B., & Kumar, P. L. (2023). Acoustic full waveform inversion for
546	2-d ambient noise source imaging. Geophysical Journal International, $234(3)$ ,
547	1628 - 1639.
548	Daubechies, I., Defrise, M., & De Mol, C. (2004). An iterative thresholding algo-
549	rithm for linear inverse problems with a sparsity constraint. <i>Communications</i>
550	on Pure and Applied Mathematics: A Journal Issued by the Courant Institute
551	of Mathematical Sciences, 57(11), 1413–1457.
552	De Plaen, R. S., Lecocq, T., Caudron, C., Ferrazzini, V., & Francis, O. (2016).
553	Single-station monitoring of volcanoes using seismic ambient noise. $Geophysical$
554	Research Letters, $43(16)$ , 8511–8518.
555	Deriche, R. (1993). Recursively implementating the gaussian and its derivatives (Un-
556	published doctoral dissertation). INRIA.
557	Díaz, E., & Sava, P. (2015). Data domain wavefield tomography using local correla-
558	tion functions. In 2015 seg annual meeting.
559	Dougherty, S. L., Cochran, E. S., & Harrington, R. M. (2019). The large-n seismic
560	survey in oklahoma (lasso) experiment. Seismological Research Letters, $90(5)$ ,
561	2051 - 2057.
562	Draganov, D., Campman, X., Thorbecke, J., Verdel, A., & Wapenaar, K. (2009). Re-
563	flection images from ambient seismic noise. Geophysics, $74(5)$ , A63–A67.
564	Durand, S., Montagner, J., Roux, P., Brenguier, F., Nadeau, R., & Ricard, Y.
565	(2011). Passive monitoring of anisotropy change associated with the park-
566	field 2004 earthquake. Geophysical Research Letters, $38(13)$ .
567	Fan, X., Guo, Z., Zhao, Y., & Chen, QF. (2022). Crust and uppermost man-
568	tle magma plumbing system beneath changbaishan intraplate volcano,
569	china/north korea, revealed by ambient noise adjoint tomography. $Geophysical$
570	Research Letters, $49(12)$ , e2022GL098308.
571	Fang, H., Yao, H., Zhang, H., Huang, YC., & van der Hilst, R. D. (2015). Direct
572	inversion of surface wave dispersion for three-dimensional shallow crustal struc-
573	ture based on ray tracing: methodology and application. Geophysical Journal
574	International, $201(3)$ , $1251-1263$ .
575	Fomel, S. (2007a). Local seismic attributes. <i>Geophysics</i> , 72(3), A29–A33.

- Fomel, S. (2007b). Shaping regularization in geophysical-estimation problems. *Geo- physics*, 72(2), R29–R36.
- Fomel, S., & Jin, L. (2009). Time-lapse image registration using the local similarity attribute. *Geophysics*, 74(2), A7–A11.
- Fomel, S., Sava, P., Vlad, I., Liu, Y., & Bashkardin, V. (2013). Madagascar: Open source software project for multidimensional data analysis and reproducible
   computational experiments. *Journal of Open Research Software*, 1(1).
- Fomel, S., & van der Baan, M. (2014). Local skewness attribute as a seismic phase
   detector. *Interpretation*, 2(1), SA49–SA56.
- Gao, H., & Shen, Y. (2014). Upper mantle structure of the cascades from full-wave
   ambient noise tomography: Evidence for 3d mantle upwelling in the back-arc.
   *Earth and Planetary Science Letters*, 390, 222–233.
- Hadziioannou, C., Larose, E., Baig, A., Roux, P., & Campillo, M. (2011). Improving
   temporal resolution in ambient noise monitoring of seismic wave speed. Jour nal of Geophysical Research: Solid Earth, 116(B7).
- Hale, D. (2006). Fast local cross-correlations of images. In Seg technical program expanded abstracts 2006 (pp. 3160–3164). Society of Exploration Geophysicists.
- Harris, C. R., Millman, K. J., van der Walt, S. J., Gommers, R., Virtanen, P., Cour napeau, D., ... Oliphant, T. E. (2020, September). Array programming with
   NumPy. Nature, 585(7825), 357–362. Retrieved from https://doi.org/
   10.1038/s41586-020-2649-2 doi: 10.1038/s41586-020-2649-2
- Huang, H., Yao, H., & van der Hilst, R. D. (2010). Radial anisotropy in the crust of
   se tibet and sw china from ambient noise interferometry. *Geophysical Research Letters*, 37(21).
- Issa, N. A., Lumley, D., & Pevzner, R. (2017). Passive seismic imaging at reservoir
   depths using ambient seismic noise recorded at the otway co2 geological stor age research facility. *Geophysical Journal International*, 209(3), 1622–1628.
- Jiang, C., & Denolle, M. A. (2020). Noisepy: A new high-performance python tool for ambient-noise seismology. *Seismological Research Letters*, 91(3), 1853– 1866.
- Le Breton, M., Bontemps, N., Guillemot, A., Baillet, L., & Larose, É. (2021). Landslide monitoring using seismic ambient noise correlation: challenges and applications. *Earth-Science Reviews*, 216, 103518.

609	Lee, EJ., Chen, P., Jordan, T. H., Maechling, P. B., Denolle, M. A., & Beroza,
610	G. C. (2014). Full-3-d tomography for crustal structure in southern california
611	based on the scattering-integral and the adjoint-wavefield methods. $Journal \ of$
612	Geophysical Research: Solid Earth, 119(8), 6421–6451.
613	Li, G., Niu, F., Yang, Y., & Xie, J. (2018). An investigation of time-frequency
614	domain phase-weighted stacking and its application to phase-velocity extrac-
615	tion from ambient noise's empirical green's functions. Geophysical Journal
616	International, $212(2)$ , $1143-1156$ .
617	Li, Y., Villa, V., Clayton, R., & Persaud, P. (2022). Shear wave velocities in the san
618	gabriel and san bernardino basins, california. Authorea Preprints.
619	Lin, FC., Moschetti, M. P., & Ritzwoller, M. H. (2008). Surface wave tomogra-
620	phy of the western united states from ambient seismic noise: Rayleigh and
621	love wave phase velocity maps. $Geophysical Journal International, 173(1),$
622	281-298.
623	Lin, FC., Ritzwoller, M. H., Yang, Y., Moschetti, M. P., & Fouch, M. J. (2011).
624	Complex and variable crustal and uppermost mantle seismic anisotropy in the
625	western united states. Nature Geoscience, $4(1)$ , 55–61.
626	Lin, FC., Tsai, V. C., Schmandt, B., Duputel, Z., & Zhan, Z. (2013). Extracting
627	seismic core phases with array interferometry. Geophysical Research Letters,
628	40(6), 1049-1053.
629	Liu, G., Fomel, S., & Chen, X. (2011). Stacking angle-domain common-image gath-
630	ers for normalization of illumination. Geophysical Prospecting, $59(2)$ , 244–255.
631	Liu, G., Fomel, S., Jin, L., & Chen, X. (2009). Stacking seismic data using local cor-
632	relation. Geophysics, $74(3)$ , V43–V48.
633	Liu, G., Persaud, P., & Clayton, R. W. (2018). Structure of the northern los angeles
634	basins revealed in teleseismic receiver functions from short-term nodal seismic
635	arrays. Seismological Research Letters, $89(5)$ , 1680–1689.
636	Liu, Y., Niu, F., Chen, M., & Yang, W. (2017). 3-d crustal and uppermost man-
637	tle structure beneath ne china revealed by ambient noise adjoint tomography.
638	Earth and Planetary Science Letters, 461, 20–29.

Lobkis, O. I., & Weaver, R. L. (2001). On the emergence of the green's function
in the correlations of a diffuse field. *The Journal of the Acoustical Society of America*, 110(6), 3011–3017.

-22-

642	Luo, B., Zhang, S., & Zhu, H. (2023). Monitoring seasonal fluctuation and long-term
643	trends for the greenland ice sheet using seismic noise auto-correlations. Geo-
644	physical Research Letters, $50(7)$ , e2022GL102146.
645	Maguire, R., Schmandt, B., Li, J., Jiang, C., Li, G., Wilgus, J., & Chen, M. (2022).
646	Magma accumulation at depths of prior rhyolite storage beneath yellowstone
647	caldera. Science, 378(6623), 1001–1004.
648	Mainsant, G., Larose, E., Brönnimann, C., Jongmans, D., Michoud, C., & Jaboyed-
649	off, M. (2012). Ambient seismic noise monitoring of a clay landslide: Toward
650	failure prediction. Journal of Geophysical Research: Earth Surface, 117(F1).
651	Mao, S., Campillo, M., van Der Hilst, R. D., Brenguier, F., Stehly, L., & Hillers,
652	G. (2019). High temporal resolution monitoring of small variations in crustal
653	strain by dense seismic arrays. Geophysical Research Letters, $46(1)$ , 128–137.

- Mao, S., Lecointre, A., van der Hilst, R. D., & Campillo, M. (2022). Space-time
   monitoring of groundwater fluctuations with passive seismic interferometry.
   *Nature communications*, 13(1), 1–9.
- Mao, S., Mordret, A., Campillo, M., Fang, H., & van der Hilst, R. D. (2020). On the
   measurement of seismic traveltime changes in the time-frequency domain with
   wavelet cross-spectrum analysis. *Geophysical Journal International*, 221(1),
   550–568.
- Mordret, A., Courbis, R., Brenguier, F., Chmiel, M., Garambois, S., Mao, S., ... others (2020). Noise-based ballistic wave passive seismic monitoring-part 2: surface waves. *Geophysical Journal International*, 221(1), 692–705.
- Moschetti, M., Ritzwoller, M., Lin, F., & Yang, Y. (2010). Seismic evidence for
   widespread western-us deep-crustal deformation caused by extension. Nature,
   464 (7290), 885–889.
- Nakata, N., Boué, P., Brenguier, F., Roux, P., Ferrazzini, V., & Campillo, M. (2016).
  Body and surface wave reconstruction from seismic noise correlations between
  arrays at piton de la fournaise volcano. *Geophysical Research Letters*, 43(3),
  1047–1054.
- Nakata, N., Chang, J. P., Lawrence, J. F., & Boué, P. (2015). Body wave extraction
   and tomography at long beach, california, with ambient-noise interferometry.
   *Journal of Geophysical Research: Solid Earth*, 120(2), 1159–1173.
- <sup>674</sup> Nakata, N., Gualtieri, L., & Fichtner, A. (2019). Seismic ambient noise. Cambridge

-23-

675	University Press.
676	Nakata, N., & Snieder, R. (2011). Near-surface weakening in japan after the 2011
677	tohoku-oki earthquake. Geophysical Research Letters, $38(17)$ .
678	Poli, P., Pedersen, H., & Campillo, M. (2012). Emergence of body waves from cross-
679	correlation of short period seismic noise. Geophysical Journal International,
680	188(2), 549-558.
681	Rawlinson, N., & Sambridge, M. (2004). Wave front evolution in strongly heteroge-
682	neous layered media using the fast marching method. Geophysical Journal In-
683	ternational, 156(3), 631-647.
684	Rickett, J., & Lumley, D. (2001). Cross-equalization data processing for time-lapse
685	seismic reservoir monitoring: A case study from the gulf of mexico. Geo-
686	$physics, \ 66(4), \ 1015-1025.$
687	Sabra, K. G., Gerstoft, P., Roux, P., Kuperman, W., & Fehler, M. C. (2005). Ex-
688	tracting time-domain green's function estimates from ambient seismic noise.
689	Geophysical research letters, $32(3)$ .
690	Sager, K., Boehm, C., Ermert, L., Krischer, L., & Fichtner, A. (2020). Global-scale
691	full-waveform ambient noise inversion. Journal of Geophysical Research: Solid
692	Earth, 125(4), e2019JB018644.
693	Sager, K., Ermert, L., Boehm, C., & Fichtner, A. (2018). Towards full waveform am-
694	bient noise inversion. Geophysical Journal International, $212(1)$ , 566–590.
695	Schimmel, M., Stutzmann, E., & Gallart, J. (2011). Using instantaneous phase
696	coherence for signal extraction from ambient noise data at a local to a global
697	scale. Geophysical Journal International, $184(1)$ , $494-506$ .
698	Seats, K. J., Lawrence, J. F., & Prieto, G. A. (2012). Improved ambient noise cor-
699	relation functions using welch s method. Geophysical Journal International,
700	188(2), 513-523.
701	Shapiro, N. M., & Campillo, M. (2004). Emergence of broadband rayleigh waves
702	from correlations of the ambient seismic noise. $Geophysical Research Letters$ ,
703	31(7).
704	Shapiro, N. M., Campillo, M., Stehly, L., & Ritzwoller, M. H. (2005). High-
705	resolution surface-wave tomography from ambient seismic noise. Science,
706	307(5715),1615-1618.
707	Shen, W., & Ritzwoller, M. H. (2016). Crustal and uppermost mantle structure be-

708	neath the united states. Journal of Geophysical Research: Solid Earth, 121(6),
709	4306-4342.
710	Snieder, R. (2004). Extracting the green's function from the correlation of coda
711	waves: A derivation based on stationary phase. Physical review $E$ , $69(4)$ ,
712	046610.
713	Stehly, L., Cupillard, P., & Romanowicz, B. (2011). Towards improving ambient
714	noise tomography using curvelet denoising filters and sem simulations of seis-
715	mic ambient noise simultaneously. Comptes Rendus Geoscience, 343(8-9),
716	591-599.
717	Tromp, J., Luo, Y., Hanasoge, S., & Peter, D. (2010). Noise cross-correlation sensi-
718	tivity kernels. Geophysical Journal International, 183(2), 791–819.
719	Van Vliet, L. J., Young, I. T., & Verbeek, P. W. (1998). Recursive gaussian deriva-
720	tive filters. In Proceedings. fourteenth international conference on pattern
721	recognition (cat. no. 98ex170) (Vol. 1, pp. 509–514).
722	Wang, K., Yang, Y., Basini, P., Tong, P., Tape, C., & Liu, Q. (2018). Refined
723	crustal and uppermost mantle structure of southern california by ambient noise
724	adjoint tomography. Geophysical Journal International, 215(2), 844–863.
725	Wang, QY., & Yao, H. (2020). Monitoring of velocity changes based on seismic
726	ambient noise: A brief review and perspective. Earth and Planetary Physics,
727	4(5), 532-542.
728	Wang, X., Zhan, Z., Zhong, M., Persaud, P., & Clayton, R. W. (2021). Urban basin
729	structure imaging based on dense arrays and bayesian array-based coherent
730	receiver functions. Journal of Geophysical Research: Solid Earth, 126(9),
731	e2021JB022279.
732	Wapenaar, K. $(2004)$ . Retrieving the elastodynamic green's function of an arbitrary
733	inhomogeneous medium by cross correlation. Physical review letters, $93(25)$ ,
734	254301.
735	Weaver, R. L., & Lobkis, O. I. (2001). Ultrasonics without a source: Thermal
736	fluctuation correlations at mhz frequencies. Physical Review Letters, 87(13),
737	134301.
738	Weaver, R. L., & Yoritomo, J. Y. (2018). Temporally weighting a time varying noise
739	field to improve green function retrieval. The Journal of the Acoustical Society
740	of America, 143(6), 3706–3719.

-25-

- Wessel, P., Luis, J., Uieda, L., Scharroo, R., Wobbe, F., Smith, W. H., & Tian, D.
  (2019). The generic mapping tools version 6. *Geochemistry, Geophysics, Geosystems, 20*(11), 5556–5564.
  Wolin, E., van der Lee, S., Bollmann, T. A., Wiens, D. A., Revenaugh, J., Dar-
- byshire, F. A., ... Wysession, M. E. (2015). Seasonal and diurnal variations in
  long-period noise at spree stations: The influence of soil characteristics on shallow stations' performance. Bulletin of the Seismological Society of America,
  105(5), 2433–2452.
- Woollard, G. (1965). The bouguer gravity anomaly map of the united states. Eos,
   Transactions American Geophysical Union, 46(1), 197–202.
- Wu, S.-M., Huang, H.-H., Lin, F.-C., Farrell, J., & Schmandt, B. (2023). Extreme
   seismic anisotropy indicates shallow accumulation of magmatic sills beneath
   yellowstone caldera. *Earth and Planetary Science Letters*, 616, 118244.
- Xie, J., Yang, Y., & Luo, Y. (2020). Improving cross-correlations of ambient noise
   using an rms-ratio selection stacking method. *Geophysical Journal Interna- tional*, 222(2), 989–1002.
- Yang, X., Bryan, J., Okubo, K., Jiang, C., Clements, T., & Denolle, M. A. (2023).
   Optimal stacking of noise cross-correlation functions. *Geophysical Journal International*, 232(3), 1600–1618.
- Yang, Y., Ritzwoller, M. H., Levshin, A. L., & Shapiro, N. M. (2007). Ambient noise
   rayleigh wave tomography across europe. *Geophysical Journal International*,
   168(1), 259–274.
- Yao, H., van Der Hilst, R. D., & De Hoop, M. V. (2006). Surface-wave array tomography in se tibet from ambient seismic noise and two-station analysis—i. phase
  velocity maps. *Geophysical Journal International*, 166(2), 732–744.
- Yao, H., Van Der Hilst, R. D., & Montagner, J.-P. (2010). Heterogeneity and
   anisotropy of the lithosphere of se tibet from surface wave array tomography.
   *Journal of Geophysical Research: Solid Earth*, 115(B12).
- Zhan, Z., Ni, S., Helmberger, D. V., & Clayton, R. W. (2010). Retrieval of moho reflected shear wave arrivals from ambient seismic noise. *Geophysical Journal International*, 182(1), 408–420.
- Zhang, Z., Nakata, N., Karplus, M., Kaip, G., & Yi, J. (2022). Shallow ice-sheet
  composite structure revealed by seismic imaging near the west antarctic ice



Figure 1. Comparison of local and global cross-correlations. Panel(a) shows two signals with non-stationary time shifts. Three events with time shifts of -0.4 s, 0.1 s and 0.3 s are used. Panels (b) and (c) present global and local cross-correlations, respectively. Here, the 1D global cross-correlations are extended along the time dimension for better comparisons with local crosscorrelations.  $\sigma$  is set to 0.2 s so that it is small enough to capture the non-stationary property of these two signals. Panels (d) and (f) show the similarity (s in Equation 16) between local and global cross-correlations before and after applying a threshold ( $\alpha(t, \tau)$  in Equation 19), respectively. Panel e shows the prestack local cross-correlations weighted by local similarity ( $c_2$  in Equation 18)



<sup>778</sup> Journal International, 214(1), 716–730.



Figure 2. Local cross-correlation and local similarity for one-day CCFs between stations TA.234A and TA.Z36A on October 10, 2010. Panels (a) and (b) show one-day records for stations TA.234A and TA.Z36A, respectively. Panel (c) presents local cross-correlation between these two signals. Panel (d) shows the local similarity between panel (c) and the linearly stacked signal shown in Figure 3a. The causal Rayleigh waves between 60-80 s can be clearly observed in panel (d).



**Figure 3.** Comparisons of one-day CCFs from simple stacking and stacking with local attributes. Panel (a) shows the CCF with simple linear stacking (Equation 1). Panels (b) and (c) are the amplitude and time-frequency spectra of panel (a), respectively. Panel (d) is the CCF stacked with local similarity. Panels (e) and (f) are the amplitude and time-frequency spectra of panel (d).



Figure 4. Comparisons of daily CCFs within one month (October 2010) for simple stacking (a) and stacking with local attribute (b). Panel (c) shows the local similarity between panel (b) and the linear stacked result shown in Figure 5a.



Figure 5. The same setting as Figure 3 except for monthly stacked results on October 2010.



Figure 6. Comparisons of convergence for conventional (a) and the proposed procedure (b). The causal and acasual signals could be clearly identified around  $\pm 80$  s with an apparent group velocity of 2.7 km/s. After 10 days of stacking, we observe a causal coherent signal appear around 50 s with an apparent velocity of 4.3 km/s, possibly representing body waves.



Figure 7. Improvements of correlation coefficients (a) and SNRs (b) for one-month stacking results. Black stars and open red circles are the results of the conventional stacking procedure and the proposed approach. The reference trace for calculating the correlation coefficients is the 30-day raw stack.



Figure 8. Distributions of the USArray stations in north Texas and Oklahoma used for comparisons in Figure 9.



Figure 9. Comparisons of one-month stacked CCFs averaged from negative and positive lags from conventional approach (a) and the proposed procedure (b) for station pairs shown in Figure 8.

## -34-



Figure 10. 35 seismic stations used for comparisons shown in Figures 11-13. The SM, SN and SS stations are parts of the Superior Province Rifting Earthscope Experiment (Wolin et al., 2015). Other stations come from the USArray Transportable Array. The background color is the Bouguer gravity anomaly (Woollard, 1965), where the linear feature with positive (blue) values highlights the extension of the Mid-continent Rift.



Figure 11. Comparisons of one-month stacked CCFs averaged from negative and positive lags (without bandpass filter) from conventional approach (a) and proposed procedure (b). The improvement of averaged SNR for each trace is shown in panel (c) with black (conventional) and red (proposed) dots. Panels d-f are the same as panels a-c except for two-month stacking. The short magenta and blue solid lines in panel (d) represent the signal and noise windows used to calculate the SNRs shown in panels c and f.



Figure 12. The same setting as Figure 11, but for the stacked data bandpass filtered between 6-15 s.



Figure 13. Same as Figure 11 but from the stacked data bandpass filtered between 15-30 s.



Figure 14. (a) Station distribution of one dense array deployed in the San Bernardino basin. The red triangle denotes the master station while the other black ones are stations used to calculate the CCFs. The red dots represent faults from the U.S. Geological Surveys. Shot gathers of 5 days (b) linear, (e) proposed local and (d) root-mean-square ratio based selective (RMSRS) (Xie et al., 2020) stacking. The blue and dark green lines in panels c and e are acausal and causal arrival times with a group velocity of 1.5 km/s and 3.5 km/s, respectively. They are used to highlight the potential ranges of Rayleigh wave arrivals. The blue and dark green lines in panels d and g are used to define the signal windows for root-mean-square ratio calculation. Panels (e-g) display the corresponding 30-day stacking results.



Figure 15. Comparison of selected phase dispersion curves from 5-day raw (a), proposed (b) and RMSRS (c) stacking CCFs displayed in Figures 14b-d. Panels (b), (d) and (f) show the number of measurements for every 0.5 s from 1.0 s to 15.0 s.



Figure 16. Similar to Figure 15 but from 30-day stacking CCFs displayed in Figures 14e-g



**Figure 17.** Comparisons of convergence for conventional (a), the proposed procedure (b) and root-mean-square-ratio selective (RMSRS) (Xie et al., 2020) stacking of the TA.A12A-TA.A18A station pair. Panels (d-e) compare 15-day stacked CCFs of the three stacking methods at 10-20, 15-30 and 20-35 s period bands. The blue numbers are the SNR for the corresponding trace. To calculate the SNR, we choose 57-221 s as the signal window and 221-357 s as the noise window. The magenta short lines (57-221 s) are used to define the signal windows for RMSRS calculation. The green dashed rectangles are used to highlight the improvement of the RMSS stacking method. Both RMSRS stacking and our proposed approaches help us to improve the SNR at different period bands.