Evaluating the Sustainability of Groundwater Resources: A Framework Incorporating the Ecological Groundwater Depth and Reliability, Resilience, and Vulnerability Indexes

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Abstract

Groundwater resource sustainability faces significant challenges due to groundwater overdraft and waterlogging. Here we propose a novel framework for evaluating the sustainability of groundwater resources. The framework incorporates a dynamic calculation of the ecological groundwater depth (EGWD) at the grid scale, considering multiple protective targets. To quantitatively evaluate the groundwater sustainability, we utilize reliability, resilience, and vulnerability, to measure the frequency, duration, and extent of unsatisfactory conditions. We apply this framework to the lower part of Tao'er River Basin in China. During the non-growth period and growth period, the upper thresholds of the EGWD range from 1.16 to 2.05 meters and 1.16 to 4.05 meters, respectively. The lower thresholds range from 6.28 to 33.54 meters and 4.87 to 30.72 meters, respectively. Future climate change improves reliability performances in regions with deep groundwater depths. Although the precipitation infiltration increases in future scenarios, prolonged duration and enhanced intensity of extreme climate events lead to decreased resilience and vulnerability performances under climate change. The proportion of areas with resilience values less than 1/12 expands to 2⁻³ times that of the historical scenario. Furthermore, we observe that more areas face the dual challenges of groundwater depletion and waterlogging under future climate change, particularly in high-emission scenarios. This study enhances understanding of groundwater resource sustainability by considering the spatial-temporal distribution of the EGWD, climate change impacts, and the identification of key regions for management. The insights can inform the development of effective strategies for sustainable groundwater resource management.

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11	Key Points:
12	• A framework is proposed to evaluate the sustainability of groundwater resources under the
13	impacts of climate change and/or human activities.
14	• The framework considers multiple targets to calculate the ecological groundwater depth with
15	spatial heterogeneity and temporal dynamics.
16	• The framework incorporates reliability, resilience, and vulnerability to evaluate the
17	sustainable performance of groundwater resources.
18	

19 Abstract

20 Groundwater resource sustainability faces significant challenges due to groundwater overdraft and 21 waterlogging. Here we propose a novel framework for evaluating the sustainability of 22 groundwater resources. The framework incorporates a dynamic calculation of the ecological 23 groundwater depth (EGWD) at the grid scale, considering multiple protective targets. To 24 quantitatively evaluate the groundwater sustainability, we utilize reliability, resilience, and 25 vulnerability, to measure the frequency, duration, and extent of unsatisfactory conditions. We 26 apply this framework to the lower part of Tao'er River Basin in China. During the non-growth 27 period and growth period, the upper thresholds of the EGWD range from 1.16 to 2.05 meters and 1.16 to 4.05 meters, respectively. The lower thresholds range from 6.28 to 33.54 meters and 4.87 28 29 to 30.72 meters, respectively. Future climate change improves reliability performances in regions 30 with deep groundwater depths. Although the precipitation infiltration increases in future scenarios, 31 prolonged duration and enhanced intensity of extreme climate events lead to decreased resilience 32 and vulnerability performances under climate change. The proportion of areas with resilience 33 values less than 1/12 expands to $2\sim3$ times that of the historical scenario. Furthermore, we observe 34 that more areas face the dual challenges of groundwater depletion and waterlogging under future 35 climate change, particularly in high-emission scenarios. This study enhances understanding of 36 groundwater resource sustainability by considering the spatial-temporal distribution of the 37 EGWD, climate change impacts, and the identification of key regions for management. The 38 insights can inform the development of effective strategies for sustainable groundwater resource 39 management.

Key words: groundwater resource sustainability; ecological groundwater depth; reliabilityresilience-vulnerability; climate change; groundwater resource management

42 **1 Introduction**

Groundwater, as the largest accessible freshwater resource, plays an increasingly prominent
role in providing freshwater supply and accounts for approximately one-third of global freshwater
extractions (Doll et al., 2012; Famiglietti, 2014; Gorelick & Zheng, 2015). Proper management of

46 groundwater resources is essential for ensuring long-term water supply for both human and 47 ecosystem needs, particularly within the ambit of compounded climatic shifts and anthropogenic 48 interventions (Mays, 2013). However, in regions with limited access to surface water, escalating 49 demands often result in groundwater overdraft. Managing groundwater resources in such areas 50 becomes challenging, as they must address a wide range of issues associated with global change, 51 including damages on groundwater-dependent ecosystems, groundwater depletion, land 52 subsidence, seawater intrusion, and deterioration of groundwater quality (Dangar et al., 2021; 53 Duran-Llacer et al., 2022; Gejl et al., 2019; Kumar et al., 2022; Ye et al., 2016). Furthermore, 54 regions characterized by shallow groundwater depths, caused by excessive irrigation and poor 55 drainage, experience soil salinization and waterlogging (Metternicht and Zinck, 2003; Moreira et al., 2015; Singh, 2014; Wang et al., 2021). This phenomenon significantly affects the eco-56 57 environment and agricultural safety in these areas. It is evident that the sustainable management of 58 groundwater resources faces significant threats under the combined effects of future climate 59 change and human activities (Malekinezhad & Banadkooki, 2018; Scanlon et al., 2023).

60 The concept of groundwater sustainability is initially proposed by Alley et al. (1999) and has 61 since undergone continuous refinement and development. Groundwater sustainability is defined as 62 the capacity of groundwater systems to consistently provide a consistent supply of sufficient and 63 high-quality water for both ecosystems and human society (Ruan and Wu, 2022). Prior research 64 has presented two approaches to evaluate groundwater sustainability. The first approach focuses 65 on estimating groundwater storage using satellite remote sensing to calculate sustainability 66 indexes, such as GRACE-groundwater drought index, for the evaluation of groundwater resource 67 sustainability (Thomas, 2019; Thomas, Famiglietti, Landerer, et al., 2017; Thomas, Caineta & 68 Nanteza, 2017). The second approach aims to establish a comprehensive evaluation framework 69 that considers various aspects, including resources, economy, and ecology (Bui et al., 2018; 70 Karimi et al., 2022; Samani et al., 2021). It evaluates the sustainability of groundwater resources 71 and ascertain the degree of sustainability using various indexes such as groundwater sustainability 72 infrastructure index (Pandey et al., 2011) and groundwater sustainability index (Singh & Bhakar, 73 2021). The above two kinds of evaluation approach are both based on the groundwater storage. 74

Groundwater depth, as the most tangible variable of groundwater resources, provides a more

75 direct insight into the interplay between underground aquifers and ecological contexts. Unlike 76 groundwater storge, groundwater depth facilitates a quantitative correlation with factors such as 77 soil salinization, vegetation growth, and land subsidence, thereby enabling efficient evaluation of 78 groundwater resource sustainability (Yang et al., 2019; Zhang et al., 2022). Ecological 79 groundwater depth (EGWD), comprising upper and lower thresholds (Zhang, 1981; Zhang et al., 80 2003), serves as a criterion to determine whether groundwater depths remain within an appropriate 81 range. Various targets, such as soil salinization (Singh, 2014), vegetation degradation (Huang et 82 al., 2019; Kath, 2018), land subsidence (Ye et al., 2016), and aquifer depletion (Konikow and 83 Kendy, 2005), are employed to establish the EGWD.

84 However, previous studies have primarily focused on calculating the EGWD predicated upon 85 individual protection target in a region thus neglecting the coupling impacts of multiple protection 86 targets. Different areas have diverse protection targets, the requirements of which on groundwater depth also vary in different seasons. It may lead to incomplete consideration of the spatial 87 88 heterogeneity and temporal dynamics of the EGWD. Furthermore, while various indicators and 89 criteria have been proposed to evaluate groundwater sustainability based on groundwater storage 90 and depths, these measures exhibit a gap in providing a comprehensive evaluation of groundwater 91 resource sustainability that incorporates the frequency, duration, and extent of unsatisfactory states 92 within the groundwater system.

93 To fill the above research gaps, this paper proposes a novel framework for evaluating the 94 sustainability of groundwater resources. Initially, we calculate the dynamic EGWD at the grid 95 scale, incorporating a multitude of protection targets as criteria to ascertain suitable ranges of 96 groundwater depths. Subsequently, an integrated surface water and groundwater model 97 incorporating diverse scenarios of interests to decision makers is developed to simulate variations 98 of groundwater depths. Finally, we deploy a suit of evaluation indexes encompassing reliability, 99 resilience, and vulnerability (RRV) to quantitatively measure the frequency, duration, and extent 100 of groundwater depths surpassing threshold ranges of the EGWD. These indexes, entrenched 101 within the domain of water supply operation of reservoirs (Cai et al., 2002; Hashimoto et al., 102 1982; McMahon et al., 2006; Sandoval-Solis et al., 2011; Zhang et al., 2017), offer a holistic 103 evaluation of groundwater resource sustainability by drawing an analogy between fluctuations in

104 groundwater depth and the operational dynamics of reservoir systems. We apply this framework to 105 the lower reach of the Tao'er River Basin (LTRB) in Northeast China. By evaluating the 106 sustainability of groundwater resources under future climate change, the framework identifies key 107 areas within the LTRB necessitating targeted groundwater resource stewardship, thus engendering 108 consequential insights pertinent to the realm of sustainable management strategies.

109 The subsequent sections of this paper are structured as follows. Section 2 presents the 110 detailed framework for evaluating groundwater resource sustainability, while Section 3 provides 111 an overview of the LTRB's fundamental characteristics and the application of the evaluation 112 framework within the basin. Section 4 discusses the spatial-temporal distribution of the EGWD, 113 the RRV performances of the groundwater system under climate change, groundwater resource 114 sustainability issues, and the identification of key areas for future management. Finally, Section 5 115 encapsulates the synthesis of findings generated by this inquiry, the concomitant implications, the 116 inherent limitations of this study, and prospective avenues for future work.

117 2 The Framework for Evaluating the Sustainability of Groundwater

118 **Resource**

119 **2.1 Overview**

120 This study presents a novel conceptual framework for evaluating the sustainability of 121 groundwater resources, consisting of three essential components (Figure 1): (1) the calculation of 122 the EGWD, (2) the development of an integrated hydrological model incorporating diverse 123 scenarios, and (3) the calculation of RRV indexes. The EGWD is regarded as the criterion to judge 124 whether the groundwater depth remains in a reasonable range. The integrated hydrological model 125 can accurately simulate long-term trajectories of groundwater depths in scenarios of interest to 126 decision makers. For quantitatively evaluating the groundwater sustainability, the EGWD and 127 simulation results of groundwater depths are used to calculate RRV indexes measuring the 128 frequency, duration, and extent of unsatisfactory conditions. For a more intricate understanding of 129 each component within the framework, please refer to sections 2.2~2.4 below.





132 **2.2 Ecological Groundwater Depth**

133 The EGWD refers to the groundwater depth that can satisfies the requirements of eco-134 environment without causing any deterioration, encompassing both upper and lower thresholds 135 (Zhang, 1981; Zhang et al., 2003). In general, shallow groundwater depths may result in soil 136 salinization due to the strong soil evaporation with the moisture evaporating to the air and salts 137 accumulating in the soil (Singh, 2021). Conversely, deep groundwater depths can give rise to 138 various eco-geological environmental issues, such as vegetation degradation (Xu & Su, 2019), 139 land subsidence (Ye et al., 2016), groundwater depletion (Konikow & Kendy, 2005), groundwater 140 quality deterioration (Gejl et al., 2019), seawater intrusion (Tomaszkiewicz et al., 2014). The 141 EGWD influenced by multiple protection targets (Figure 2), is dominated by different protection 142 targets in distinct areas and time periods. To comprehensively evaluate the EGWD, this study 143 synergistically integrates multiple protection targets, culminating in the computation of grid-scale 144 EGWD across different temporal periods.





146 Figure 2 Schematic diagram of the ecological groundwater depth

145

147 2.2.1 Upper Thresholds of the Ecological Groundwater Depth

The upper threshold of the EGWD is primarily oriented towards the mitigation of soil salinization and the facilitation of optimal growth of non-aquatic vegetation species (Jia et al., 2015; Zhai et al., 2021). During the non-growth period spanning from November to April of the subsequent calendar year, vegetations undergoes phases of desiccation, senescence, mortality, or dormancy, leading to a diminished requirement for vegetation growth proliferation. During this phase, the groundwater depth necessary to prevent soil salinization is determined by the capillary

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154 rise height. Conversely, encompassing the growth period spanning from May to October, the 155 imperative groundwater depth for preempting soil salinization phenomena becomes a function of 156 the summation between the capillary rise height and the depth of vegetation root systems. When 157 the groundwater depth falls below the capillary rise height, groundwater recharges the soil through 158 the capillarity, leading to high soil moisture and strong soil evaporation. The moisture 159 transpiration process results in aqueous volatilization into the atmosphere, with concomitant 160 accumulation of salts within the soil matrix, thereby instigating plausible events of soil 161 salinization occurrences. Additionally, instances of shallow groundwater depths can exert an 162 elevating influence upon soil moisture content and concurrently impose limitations upon effective soil aeration within the root zone. It thus becomes incumbent that the capillary rise height remains 163 164 situated below the substratum spanned by the vegetation root depth during the growth period. In 165 hydrological contexts dominated by perennial or seasonal surface water bodies, typified by aquatic 166 ecosystems, the risk of soil salinization is minimal. Therefore, the upper threshold of the EGWD is 167 not concerned within such aquatic settings (Lu, 2020).

168 The determination of the upper threshold of the EGWD involves two crucial parameters: the 169 capillary rise height and the depth of vegetation root systems. The capillary rise height is subject 170 to the physicochemical properties of the soil, and can be attained through scrutiny of controlled 171 laboratory experimentation (Dong et al., 2008), the field investigation (Pan et al., 2018), and the 172 computational application of empirical formulations (Useviciute & Baltrenaite-Gediene, 2022). 173 The vegetation root depth exhibits variations across disparate ecological settings, an outcome of 174 the diversity in vegetation species and the differences of root systems spanning ecosystems 175 encompassing paddy fields, drylands, grasslands, and woodlands. The maximum root depth of 176 different vegetations species within a specific ecosystem is selected as the vegetation root depth 177 for that ecosystem. For each vegetation species, the depth at which roots account for more than 178 80% of the total root length is considered as the root depth of that particular vegetation. 179 Information on vegetation species and their respective root depths in various ecosystems can be 180 obtained through field investigations or by referring to relevant literature and books (Lu, 2020; 181 Zhao, 2012).

8

2.2.2 Lower Thresholds of the Ecological Groundwater Depth

183 The lower threshold of the EGWD is subject to the influence of multiple protection targets 184 (Figure 2). Prolonged and intense groundwater exploitation can lead to a decline in regional 185 groundwater level, thereby resulting in insufficient water supply to vegetation roots and 186 consequent degradation of vegetation integrity (Meng et al., 2019; Zhang et al., 2011). This 187 decline in groundwater level also transfers the load originally carried by groundwater to the soil 188 particles, which increases the risk of land subsidence (Ye et al., 2016). Additionally, high-intensity 189 groundwater exploitation may lead to the aquifer dewatering and groundwater depletion (Bierkens 190 & Wada, 2019), significantly impairing groundwater resource sustainability. Apart from these 191 overarching protection targets, specific areas may have unique protection needs. For example, in 192 groundwater wellfield protection zones, the decline of the groundwater level facilitates the 193 infiltration of the water with poor quality into the aquifer, leading to the groundwater quality 194 deterioration (Yoneda et al., 2001). In coastal areas, low groundwater level easily causes seawater 195 intrusion, resulting in saline groundwater (Tomaszkiewicz et al., 2014). Therefore, the calculation 196 of the lower threshold of the EGWD necessitates a comprehensive analysis of specific protection 197 requirements in different areas and time periods, considering appropriate protection targets.

198

(1) Groundwater depth thresholds for preventing vegetation degradation

199 Vegetation undergoes processes such as desiccation, senescence, mortality, or dormancy 200 during the non-growth period, thereby making the target of preventing vegetation degradation 201 insignificant. In agricultural areas, the groundwater depth threshold for mitigating vegetation 202 degradation is ignored due to implementation of irrigation practices during periods of water 203 deficiency. However, in grasslands, woodlands, and wetlands, this threshold is taken into account.

204 Previous studies have proposed two ideas for determining the groundwater depth threshold 205 for preventing vegetation degradation. The first posits that groundwater will no longer support the 206 vegetation growth when the groundwater depth exceeds the extreme evaporation depth. Hence, the 207 extreme evaporation depth is regarded as the threshold (H₁) for preventing vegetation degradation 208 (Fan et al., 2004; Pan et al., 2018). The extreme evaporation depth relies on soil properties and can 209 be obtained through numerical simulations (Shah et al., 2007), lysimeter measurements (Ma et al., 210 2019), empirical formula calculations (Fu et al., 2008), and dynamic data correlation methods **211** (Zhao, 2012).

212 The second idea involves analyzing the relationship between the actual growth state of 213 vegetation and groundwater depth to determine the threshold (H₂) for preventing vegetation 214 degradation (Huang et al., 2019; Jia et al., 2015; Kath, 2018). To encapsulate the growth status of 215 vegetation accurately, precise vegetation indexes are embraced, such as the normalized vegetation 216 index (NDVI) (Xu and Su, 2019) and the enhanced vegetation index (EVI) (Elmore et al., 2006). 217 H₂ is determined by mapping the relationships between these vegetation indexes and groundwater 218 depths while examining the variations in vegetation indexes as groundwater depths increase. 219 Taking NDVI as an example, there are two relationships between NDVI and groundwater depths 220 with the increase of groundwater depths (Dang et al., 2019). First, NDVI initially increases and 221 then decreases with increasing groundwater depth. The groundwater depth corresponding to the 222 maximum NDVI is identified as H₂. Second, NDVI initially increases and then stabilizes with 223 increasing groundwater depth. The groundwater depth corresponding to the stable NDVI is 224 identified as H2. Given the diverse vegetation types and their distinct responses to groundwater 225 depths in different ecosystems (e.g., grassland, woodland, wetland), the relationships between 226 NDVI of different ecosystems and groundwater depths are mapped to determine H₂.

227 This study takes a comprehensive approach by incorporating both soil properties and the 228 actual growth states of vegetation to determine the groundwater depth threshold for preventing 229 vegetation degradation. It introduces a synthetic method to calculate this threshold. During the 230 growth period, if the ecological index of a specific grid is lower than the maximum ecological 231 index observed during the non-growth period of that same grid, the vegetation growth state is 232 deemed poor. In such cases, the groundwater depth threshold for preventing vegetation 233 degradation is determined as the smaller value between H₁ and H₂. Conversely, if the ecological 234 index during the growth period exceeds the maximum ecological index during the non-growth 235 period, the vegetation growth state is deemed satisfactory. Consequently, the groundwater depth 236 threshold for preventing vegetation degradation is determined as larger value between H_1 and H_2 .

237

(2) Groundwater depth thresholds for preventing land subsidence

238 Land subsidence is influenced by various factors including groundwater depth, strata239 lithologic structure, hydrogeological conditions. In order to calculate the groundwater depth

240 threshold for preventing land subsidence, several methods are employed, all of which require the 241 determination of an acceptable land subsidence depth. One approach involves analyzing long-term 242 land subsidence observations, whether via land subsidence observation system or the remote 243 sensing satellite to derive the statistical relationship between the groundwater depth decline and 244 land subsidence, thereby determining the groundwater depth threshold for preventing land 245 subsidence (Lu et al., 2022; Peng et al., 2022). Another method integrates groundwater dynamics 246 with soil deformation modeling, facilitating the simulation of the stratum deformation caused by 247 groundwater level declines, ultimately providing the groundwater depth threshold for preventing 248 land subsidence (Ma & Luo, 2015). Furthermore, the layer-wise summation method (Equation 1) 249 leverages geological and hydrogeological data to calculate the groundwater depth threshold for 250 preventing land subsidence (Gong, 1996).

$$D = D_0 + S / \left(\mu_s \gamma \sum_{i=1}^n m_{vi} H_i \right)$$
(1)

where *D* is the groundwater depth threshold for preventing land subsidence, m; D_0 is the groundwater depth before the groundwater level decreases, m; *S* is the total land subsidence caused by the decline of groundwater level; μ_s is the land subsidence experience coefficient; γ is the bulk density of water, kN/kg; m_{vi} is the volume compression coefficient of layer i; H_i is the thickness of layer i, m; *n* is the total number of layers.

257 (3) Groundwater depth thresholds for preventing Quaternary aquifer depletion

258 The Quaternary aquifer depletion is closely associated with groundwater exploitation. Many 259 studies have highlighted the risk of aquifer depletion when the groundwater level experiences 260 continuous decline or groundwater exploitation exceeds the allowable exploitable yield (Myriam 261 et al., 2018; Sophocleous, 1997; Zhao et al., 2019). However, a consensus has not been reached 262 regarding the specific threshold at which the groundwater level decline signifies aquifer depletion. 263 This study assumes that the Quaternary aquifer is depleted when the groundwater depth exceeds 264 two-thirds of the aquifer's thickness, which is defined as the groundwater depth threshold for 265 preventing Quaternary aquifer depletion (Lu, 2020). Through borehole data analysis, it is possible 266 to ascertain the thickness of Quaternary aquifer, thereby enabling the calculation of the 267 groundwater depth threshold for preventing Quaternary aquifer depletion.

268 (4) The calculation of lower thresholds of the EGWD

By performing calculations to determine the groundwater depth thresholds for all protection targets, the minimum value among these thresholds for each grid is considered as the lower threshold of the EGWD.

272 **2.3** Hydrological Model and Scenario Design

273 Accurately simulating long-term trajectories of groundwater depths is crucial for the 274 sustainable evaluation of groundwater resources, especially considering the influence of climate 275 change and/or human activities. Integrated surface water and groundwater models, such as 276 HydroGeoSphere (Brunner & Simmons, 2012), SWAT-MODFLOW (Kim et al., 2008), GSFLOW 277 (Markstrom et al., 2008), HEIFLOW (Tian, Zheng, Wu, et al., 2015; Tian, Zheng, Zheng, et al., 278 2015), prove to be effective tools for simulating hydrological cycle processes, including variations 279 of the groundwater depth, groundwater flow movements, and interactions between surface water 280 and groundwater. Importantly, the simulations should involve the potential effects of climate 281 change and/or human activities on the groundwater depth. The Coupled Model Intercomparison 282 Project Phase 6 (CMIP6), initiated by the World Climate Research Programme, provides access to 283 meteorological data of multiple general circulation models (GCMs) and shared socioeconomic 284 pathway (SSPs) developed by several research institutions around the world. Human activities 285 significantly affect groundwater depth through land-use pattern changes and the implementation of 286 water resources management measures. Examples include the conversion from drylands to 287 paddies, the conversion from paddies to drylands, the alteration of the groundwater supply ratio, 288 and exploitation scheduling. Incorporating these factors into simulations provides valuable 289 insights for groundwater resource management.

290

2.4 Reliability, Resilience, and Vulnerability of Groundwater System

The sustainability evaluation of groundwater resources requires appropriate indexes to describe the characteristics of the groundwater depth exceeding threshold ranges of the EGWD. To determine the performance of the groundwater system, we adopt the RRV metrics, commonly used to evaluate water scarcity in water supply systems (Cai et al., 2002; Hashimoto et al., 1982; McMahon et al., 2006; Sandoval-Solis et al., 2011; Zhang et al., 2017). RRV is used to describe
the frequency, duration, and extent of the system being in an unsatisfactory state. In this study,
RRV indexes are used to evaluate the sustainability performance of groundwater system by
simulating the variations of groundwater depths under different scenarios.

Reliability represents the probability of a system being in a satisfactory state (Hashimoto et al., 1982). Specifically, reliability is defined as the ratio of the periods during which groundwater depths remain within threshold ranges of the EGWD to the total simulation periods. A higher reliability value indicates a more sustainable groundwater system.

$$\alpha = \frac{1}{n} \sum_{i=1}^{n} Z_i$$
(2)

304 where α is the reliability; *n* is the total number of simulation periods; *t* is the current period; 305 Z_t is the status of the current period. Z_t equals 1 with the groundwater system being in a 306 satisfactory state, otherwise Z_t equals 0.

Resilience represents the duration for a system to recover to the satisfactory state from an unsatisfactory state (Hashimoto et al., 1982). In our study, resilience is defined as the reciprocal of the longest duration of groundwater depths exceeding threshold ranges of the EGWD during the simulation period (Moy et al., 1986). A larger resilience value indicates a more resilient system, as it signifies a shorter duration for groundwater depths to return to threshold ranges of the EGWD and attain a satisfactory state.

313

$$\beta = \frac{1}{MAX(A_i)} \tag{3}$$

314 where β is the resilience; A_i is the duration of *i*th consecutive unsatisfactory state. When the 315 groundwater depth is in the threshold range of the EGWD in all simulation periods, β equals 1.

316 Vulnerability represents the extent to which a system is deviates from a satisfactory state 317 (Hashimoto et al., 1982). In our study, vulnerability is specifically defined as the maximum value 318 of groundwater depth exceeding threshold ranges of the EGWD throughout the simulation period. 319 To ensure comparability and account for variations in the EGWD across different areas and time 320 periods, vulnerability is nondimensionalized and normalized following the approach suggested by 321 Loucks (1997) and McMahon et al. (2006). A smaller normalized vulnerability value indicates a 322 more sustainable groundwater system, signifying a lower extent of groundwater depths exceeding 323 threshold ranges of the EGWD.

324
$$\gamma = \frac{\max(GWD_t - L_m) / L_m - \min(\max(GWD_t - L_m) / L_m)}{\max(\max(GWD_t - L_m) / L_m) - \min(\max(GWD_t - L_m) / L_m)}$$
(4)

325 where γ is the nondimensionalized and normalization vulnerability. *GWD_t* is the groundwater 326 depth in period t; *m* is the 12 months of the year; L_m is the monthly threshold of the EGWD.

327 **3** Case Study

328 The lower reach of Tao'er River Basin (LTRB) in northeast China, with an area of 5,900 km², 329 is selected as the study area (Figure 3). The average annual precipitation is 386.6 mm, the average 330 annual evaporation is 1,091.4 mm, and the average annual temperature is 5.8 °C. 331 Geomorphologically, the area primarily comprises of piedmont inclined plains and low plains. The 332 piedmont inclined plain is distributed in the Tao'er River alluvial fan, while the low plain with the 333 flat terrain is distributed in the eastward of alluvial fan. The lithology of the Quaternary pore 334 phreatic aquifers in the river valley, alluvial fan, and low plain is predominantly composed of 335 Holocene alluvial sand gravel, Pleistocene alluvial sand gravel, Pleistocene alluvial sand gravel, 336 alluvial and lacustrine fine sand, and silky fine sand (Li et al., 2021).

337 The fragile eco-environment in the LTRB is highly susceptible to climate change and human 338 activities. From 1990 to 2000, the farmland area increases from 2,972 to 10,872 km², with an 339 increase of about 20%, among which the paddy area increases from 139 to 871 km², with an 340 increase of over five times. The irrigation accounts for more than 80% of the total water 341 consumption, leading to a substantial surge in water demand due to the increased agricultural land. 342 The irrigation is dominated by groundwater, which accounts for more than 80% of the irrigation. 343 Tens of thousands of pumping wells drilled for irrigation leads to continuous decline in regional 344 groundwater levels. Previous reports (SRWRC, 2017; WRMC & WCSDI, 2015) have highlighted 345 that the actual groundwater exploitation within the LTRB far exceeds the permissible exploitable 346 yield, leading to groundwater depletion. This is further evidenced by the decline in groundwater 347 levels, which are more than 2 m below the streambed elevation, indicating that the stream has 348 been disconnected from groundwater in the LTRB (Ma et al., 2022; Wang et al., 2014). The vast 349 stream leakage to groundwater, coupled with the impact of a series of droughts, results in frequent 350 streamflow drying up events, which poses a significant threat to the health of river ecosystem. 351 Furthermore, soil salinization has been observed in the southwest LTRB due to shallow

groundwater depths and strong groundwater evaporation rates. Consequently, there is an urgent
need to investigate whether groundwater depths exceed threshold ranges of the EGWD for the
sustainable groundwater exploitation within the LTRB.



355

Figure 3 The study domain. (a) The location of the Tao'er River Basin; (b) The location of the LTRB; (c) The land-use in the LTRB. Reasons for selecting the LTRB as the study domain are as follows. First, groundwater overdraft resulting in various eco-environmental issues, mainly occurs in the LTRB, which is a crucial region to investigate the sustainability of groundwater resources. Second, there are adequate boreholes and groundwater level observations for establishing groundwater model in the LTRB. Third, the groundwater model is more suitable for the simulation

362 of the movements of the loose rock pore groundwater, which is mainly distributed in the LTRB.

363 The methods presented in Section 2.2 are used for calculating the EGWD in the LTRB. The 364 upper threshold of the EGWD is determined by considering capillary rise heights specific to 365 different soils and vegetation root depths in various ecosystems. On the other hand, the lower 366 threshold of the EGWD takes into account three protection targets: vegetation degradation, land 367 subsidence, and Quaternary aquifer depletion. The reasons are as follows. First, high-intensity 368 groundwater extraction in the LTRB has led to a regional decline in groundwater levels, causing 369 insufficient water supply for vegetation and subsequent degradation. Moreover, the decline of 370 groundwater level transfers pressure to soil particles, posing a risk of land subsidence. Lastly, the 371 Quaternary aquifer, primarily exploited for groundwater, serves as a crucial storage reservoir, with 372 good water storage conditions, necessitating the prevention of depletion due to overexploitation. 373 Detailed descriptions of data and methods for calculating the EGWD are presented in Table 1.

Table 1 The descriptions of data and methods for calculating the EGWD in the LTRB.

Types of the EGWD	Protection targets	Descriptions
Upper thresholds	Soil salinization	Capillary rise heights of different soils
	and the growth	(Supplementary Figure S1) are obtained from
	of non-aquatic	experimental results in Zhao (2012), and empirical
	vegetation	values in CGS (2012). Vegetation root depths of
		paddy, dryland, grassland and woodland ecosystems
		are obtained referring to the field investigation results
		in Lu (2020) and Zhao (2012) and the empirical
		values in Chen (2001) and Yu (2003).
Lower thresholds	Vegetation	Extreme evaporation depths of different soils are
(the shallowest of	degradation (D_1)	obtained referring to calculation results in Fan (2004)
$D_1, D_2, and D_3)$		and Zhao (2012) and empirical values in CGS (2012).
		NDVI is selected as the ecological indicator of
		vegetation growth state. Monthly NDVI from 2001 to
		2016 is obtained from Resources and Environmental
		Science and Data Center, Institute of Geographic
		Academy of Sciences (https://www.reade.or)
	T 1 1 1	Academy of Sciences (<u>nups://www.resdc.cn</u>).
	Land subsidence	The layer-wise summation method described in
	(D_2)	Section 2.2.2 is used to calculate the groundwater
		depth threshold for preventing land subsidence. S is 20
		mm (Lu, 2020). μ_s and m_{vi} is 0.4 and 0.025 MPa ⁻ ,
		respectively, referring to Chang (1990).
	Quaternary	The thickness of Quaternary aquifer is obtained
	aquiter depletion	according to the borehole data, two-thirds of which is
	(D_3)	taken as the groundwater depth threshold for
		preventing Quaternary aquifer depletion (Lu, 2020).

376 An integrated SWAT-MODFLOW model is developed for the simulation of groundwater 377 depths. The SWAT model encompasses the whole Tao'er River Basin and is driven by daily 378 meteorological data, including precipitation, maximum and minimum temperature, wind speed, 379 and solar radiation data at four weather stations (Figure 3b), covering the period from 1990 to 380 2016. The period is divided into the warm-up period (1990 - 1992), calibration period (2005 -381 2016), and validation period (1993 - 2004). The MODFLOW model, focusing on the LTRB, 382 employs uniform 1 km \times 1 km grids to discretize the Quaternary pore aquifer. The modelling 383 domain, containing 5903 active grids, is generalized as a layer of unequal thickness, isotropic two-384 dimensional unsteady flow aquifer according to the borehole data. The streambed leakage, precipitation infiltration, and phreatic evaporation are obtained from SWAT model, and then 385 386 transmitted to Visual MODFLOW through River Module, Recharge Module, and 387 Evapotranspiration Module for groundwater simulation. The MODLFOW model is run with 10-388 day stress period from 2001 to 2016. The first eight years (2001 - 2008) are considered as the 389 calibration period, and the remaining eight years (2009 - 2016) are used for validation. Further 390 details regarding the setup, calibration, and validation of the integrated SWAT-MODFLOW can be 391 found in our previous study (Wang et al. 2023). Monthly simulation results of integrated SWAT-392 MODFLOW model are used to calculate RRV.

393 For scenario design, we consider the impact of climate change on groundwater depths and 394 assume that historical human activities will not change in the future. This is due to the fact that 395 there were no major changes in land-use and groundwater supply and exploitation practices in the 396 LTRB over the past two decades. Meteorological data, including daily precipitation, daily 397 maximum and minimum temperature, from 5 GCMs and 3 SSPs, are obtained from CMIP6. 398 Combined with historical meteorological data, a total number of 16 scenarios are designed (Table 399 2). Future meteorological data exhibit lower spatial resolution and deviations compared to 400 historical data, necessitating spatial downscaling and deviation correction. The historical daily 401 meteorological data from four weather stations (Figure 3b) serve as a reference for the correction 402 process. First, the inverse distance weight method (Chaplot et al., 2006) is used to downscale the 403 future gridded meteorological data to the four weather stations. Second, the quantile mapping 404 method (Wilcke et al., 2013) is applied to modify the biased meteorological data. All scenarios'

- 405 meteorological data are fed into the integrated SWAT-MODLFOW model to simulate groundwater
- 406 depths, and simulation results from 5 GCMs of the same emission scenario are as a set to calculate

407 RRV.

- 408
- **409** Table 2 The climate scenarios characterized by various Shared Socioeconomic Pathways (SSPs)

Scenario	SSPs	GCMs	Time Range (Resolution)	Spatial Resolution
S0	History	-	1969-2016 (daily)	4 stations
S1	SSP126	BCC-CSM2-MR	2015-2064 (daily)	1.125°×1.125°
S2		GFDL-ESM4		1.250°×1.000°
S3		IPSL-CM6A-LR		2.500°×1.268°
S4		MPI-EMS1-2-LR		1.875°×1.875°
S5		MRI-ESM2-0		1.125°×1.125°
S6	SSP245	BCC-CSM2-MR		1.125°×1.125°
S7		GFDL-ESM4		1.250°×1.000°
S8		IPSL-CM6A-LR		2.500°×1.268°
S9		MPI-EMS1-2-LR		1.875°×1.875°
S10		MRI-ESM2-0		1.125°×1.125°
S11	SSP585	BCC-CSM2-MR		1.125°×1.125°
S12		GFDL-ESM4		1.250°×1.000°
S13		IPSL-CM6A-LR		2.500°×1.268°
S14		MPI-EMS1-2-LR		1.875°×1.875°
S15		MRI-ESM2-0		1.125°×1.125°

410 and General Circulation Models (GCMs).

411 **4 Results**

412 **4.1 Ecological Groundwater Depth in the LTRB**

Upper and lower thresholds of the EGWD demonstrate notable spatial heterogeneity and temporal variability (Figure 4). Upper thresholds of the EGWD range from 1.16 to 2.05 meters during the non-growth period, with a majority of grids displaying values below 2 meters (Figure 4a). Upper thresholds of the EGWD range from 1.16 to 4.05 meters during the growth period, with most grids maintaining values below 3 meters (Figure 4b). Grids with larger upper thresholds of the EGWD are mainly covered with woodlands, where the vegetation roots are more developed and rooted deeper. On the other hand, lower thresholds of the EGWD range from 6.28 to 33.54

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420 meters during the non-growth period (Figure 4c) and from 4.87 to 30.72 meters during the growth 421 period (Figure 4d), with the majority of grids exhibiting values below 20 meters. Lower thresholds 422 of the EGWD in the middle and border grids of the LTRB are deeper than 15 meters. In contrast, 423 grids in the south, western river valley, and wetlands have shallow lower thresholds of the EGWD, 424 which is less than 10 meters. These spatial and temporal dynamics in EGWD thresholds 425 underscore the influence of local vegetation characteristics and hydrogeological conditions on 426 groundwater depth variations.

427 During the non-growth period, the mitigation of soil salinization is the primary protection 428 target for upper thresholds of the EGWD, which is controlled by both the mitigation of soil 429 salinization and the facilitation of optimal growth of non-aquatic vegetation species during the 430 growth period. Lower thresholds of the EGWD are spatially controlled by different protection 431 targets (Figure 4e and 4f). In southwest grids, the prevention of Quaternary aquifer depletion, due 432 to its thin thickness, emerges as the prominent target during the non-growth period. Conversely, in 433 other grids, the prevention of land subsidence remains the primary target. During the growth 434 period, the preservation of vegetation becomes a dominant target in grasslands, wetlands, and 435 forests, while the protection targets in remaining grids align with those observed during the non-436 growth period.



437

Figure 4 Ecological groundwater depths and their dominant protection targets in the LTRB. (a)(b) Upper thresholds of the EGWD during the non-growth period and growth period, respectively.
(c)-(d) Lower thresholds of the EGWD during the non-growth period and growth period,
respectively. (e)-(f) Dominant protection targets that control lower thresholds of the EGWD in
different grids during the non-growth period and growth period, respectively.

4.2 Reliability, Resilience, and Vulnerability of Groundwater System

There are more areas where groundwater depths exceed threshold ranges of the EGWD in future climate scenarios, as indicated by the increased presence of green areas compared to the historical scenario (Figure 5). In the historical, SSP126, SSP245, and SSP585 scenarios, areas 447 with reliability values less than 1.0 accounts for 26%, 41%, 44%, and 50%, respectively. 448 Compared to the precipitation in the historical scenario, the projected precipitation in the SSP126, SSP245, and SSP585 scenarios shows an increase of 29%, 22%, and 29%, respectively. 449 450 Meanwhile, the evapotranspiration is increased by 31%, 25%, and 29% (Supplementary Figure 451 S2). Combined with the effects of rainfall and evaporation, the average precipitation infiltration 452 into groundwater is also increased by 24%, 14%, and 30% during the simulation period, 453 respectively (Supplementary Figure S2). This enhanced infiltration raises the possibility of 454 groundwater depths exceeding upper thresholds of the EGWD in areas with shallow groundwater 455 depths, such as the river valley between the Tao'er River and Jiaoliu River and the southwestern 456 LTRB. Specifically, we find that 20%, 11%, and 19% of the whole basin experience such a 457 transition in the SSP126, SSP245, and SSP585 scenarios, respectively. On the other hand, due to 458 the increased precipitation infiltration into groundwater, the proportion of the LTRB area with low 459 reliability values (<0.8) is projected to be 11%, 15%, and 11% for the SSP126, SSP245, and 460 SSP585 scenarios, respectively, all of which are lower than 16% observed in the historical 461 scenario. These areas with increased reliability in the future climate scenarios are mainly 462 concentrated in the central LTRB.



464 Figure 5 Reliability performances of groundwater system in the LTRB. Panels (a) to (d) present
465 the spatial distribution of reliability values in the history, SSP126, SSP245, and SSP585 scenarios,
466 respectively. Panel (e) presents the proportions of reliability values within different intervals.

463

Compared to the historical scenario, future climate scenarios exhibit poor resilience performances in the LTRB (Figure 6). The proportion of the LTRB area with resilience values less than 1/12 accounts for 8%, 17%, 22% and 26% in the historical, SSP126, SSP245, and SSP585 scenarios, respectively. This indicates a significant expansion of areas where groundwater depths exceed threshold ranges of the EGWD and fails to recover within one year. Unlike reliability that reflects the average performance of system during the simulation horizon, resilience reflects the

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474 system performance during the extreme dry or wet events. In this study, for example, we define 475 the extreme dry or wet events as periods when the precipitation infiltration into groundwater falls 476 below or exceeds 25% of the average precipitation infiltration during the simulation stages. 477 Extreme dry events occurring in future climate scenarios exhibit longer durations compared to the 478 historical scenario (Supplementary Figure S3). Specifically, the historical scenario exhibits 479 extreme dry and wet periods less than 20 months and 10 months, respectively, while future 480 scenarios experience extreme dry and wet periods more than 30 months and 20 months, 481 respectively. In extreme dry periods, the reduction in precipitation infiltration into groundwater 482 makes the groundwater replenishment insufficient to compensate for the groundwater exploitation, 483 which results in prolonging the duration of groundwater depths exceeding threshold ranges of the 484 EGWD in the central and southern LTRB. Furthermore, extreme wet periods contribute to the 485 decreased resilience of groundwater depths exceeding upper thresholds of the EGWD in the river 486 valley between Tao'er River and Jiaoliu River and southwestern LTRB.



Figure 6 Resilience performances of groundwater system in the LTRB. Panels (a) to (d) present the spatial distributions of resilience values in the history, SSP126, SSP245, and SSP585 scenarios, respectively. Panel (e) presents the proportions resilience values within different intervals. The legend number, such as 1/12, indicates consecutive 12-month durations during which the groudwater depth exceeds the upper or lower thresholds of the EGWD, with similar implications conveyed by the numbers 1/3, 1/6, and 1/9.

487

In comparison to the historical scenario, future climate scenarios exhibit poorer vulnerability
performances in the LTRB (Figure 7). Areas with vulnerability values below 0.2 accounts for
95%, 89%, 90% and 88% of the LTRB in the historical, SSP126, SSP245, and SSP585 scenarios.

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498 The distribution of areas with high vulnerability has undergone a shift between the historical and 499 future climate scenarios. In the historical scenario, the river valley between Tao'er River and 500 Jiaoliu River exhibits high vulnerability values, whereas this issue is alleviated in future climate 501 scenarios. The southern and eastern LTRB encounter scattered areas with high vulnerability values 502 in the historical scenario, but grids with high vulnerability in these areas increase and cluster in 503 future climate scenarios. Vulnerability, similar to resilience, represents the performance during an 504 extreme dry or wet period. As mentioned above, although there is an overall increase in 505 precipitation infiltration in future climate scenarios, the duration of extreme drought events 506 becomes longer. It indicates that the cumulative drought intensity of single extreme drought event 507 in future climate scenarios exceeding that in the historical scenario (Supplementary Figure S3). 508 Enhanced extreme drought events lead to deeper groundwater depths in future climate scenarios. 509 For example, in the southern LTRB, future climate scenarios depict the deepest groundwater depth 510 approximately 6 meters below the lower threshold of the EGWD, which is about 2 m in the 511 historical scenario. Analogously, the prolonged duration of extreme wet conditions in future 512 climate scenarios increases the intensity of groundwater recharge, rising groundwater levels. For 513 example, in some grids of eastern LTRB, the shallowest groundwater depth exceeds the upper 514 threshold of the EGWD by about 1 m during extreme wet events in future climate scenarios, while 515 it remains in the threshold range of the EGWD in the historical scenario.



517 Figure 7 Vulnerability performances of groundwater system in the LTRB. Panels (a) to (d) present
518 the spatial distribution of vulnerability values in the history, SSP126, SSP245, and SSP585
519 scenarios, respectively. Panel (e) presents the proportions of vulnerability values within different
520 intervals.

516

4.3 Sustainability of Groundwater Resources in Some Key Subbasins

522 There is an expansion of areas experiencing groundwater resource unsustainability in future 523 climate scenarios, primarily concentrated in regions I, II, and III (Figure 8). The criterion 524 employed in this study to determine the sustainability of groundwater resources is based on the 525 comprehensive performance of reliability, resilience and vulnerability metrics. Specifically, when the reliability index falls below 0.8, resilience is less than 1/12, or vulnerability surpasses 0.2, it is suggested that the sustainability of groundwater resources becomes compromised. The expansion of areas experiencing groundwater resource unsustainability can be predominantly attributed to poor performances of resilience and vulnerability during extreme dry and wet periods (Figure 6 and 7).

531 Distinct challenges regarding the sustainability of groundwater resources are observed across 532 various regions. Some areas face the specific issue of groundwater depths surpassing either the 533 upper or lower threshold of EGWD, while others encounter both situations intermittently during 534 different periods (Figure 8). In regions I and II, the unsustainability of groundwater resources can 535 be chiefly attributed to groundwater depths falling below lower thresholds of the EGWD. The 536 extent of groundwater resource unsustainability in the SSP126 scenario closely resembles that of 537 the historical scenario. Whereas the SSP245 and SSP585 scenarios exhibit an expanding scope of 538 this issue. In these scenarios, the duration of extreme dry periods exceeds 40 months and even 539 reaches 65 months in the SSP585 scenario, compared to only 20 months in the historical scenario. 540 The cumulative drought intensity of single extreme drought event in future climate scenarios 541 surpasses that of the historical scenario (Supplementary Figure S3), resulting in poorer resilience 542 and vulnerability performances. Moreover, in region II, the prevention of Quaternary aquifer 543 depletion, due to its thin thickness, emerges as the prominent control target (Figure 4e and 4f). The 544 groundwater depth is prone to decline rapidly when the groundwater exploitation occurs. 545 Therefore, except for the impact of extreme drought, poor water-abundance is another reason for 546 groundwater depths falling below lower thresholds of the EGWD.

547 The unsustainability of groundwater resources in regions III primarily stems from 548 groundwater depths exceeding upper thresholds of the EGWD. The primary reason for the issue is 549 the increased precipitation in future climate scenarios. Shallow groundwater depths, which are 550 even less than 3 m in some areas (Supplementary Figure S4), also contribute to groundwater 551 depths exceeding upper thresholds in this region. Soil salinization caused by shallow groundwater 552 depths has been observed in region III. Furthermore, in the river valley of region III, the extensive 553 leakage from the stream that replenishes groundwater serves as a significant factor in groundwater 554 depths surpassing upper thresholds.



- 566 exceed upper thresholds of the EGWD occasionally in the medium and long term (2037~2064)
- 567 due to the climate transitioning towards wetter conditions. Groundwater depths in region B exceed

upper thresholds of the EGWD during the initial simulation years, then rapidly decrease to threshold ranges of the EGWD. Although region B does not experience groundwater depths falling below lower thresholds of the EGWD in scenario S15, the situation appears in scenarios S5 (the MRI-ESM2-0 and SSP126 scenario) and S6 (the BCC-CSM2-MR and SSP245 scenario) (Supplementary Figures S9b and S10b). Notably, in grids near the river in region B, there is a high probability of groundwater depths surpassing upper thresholds of the EGWD due to the leakage from the stream in the future (Supplementary Figure S19).

575 Climate change induces fluctuations of the intensity of precipitation infiltration recharge into 576 groundwater during different periods in the future. The intensity of precipitation infiltration 577 recharge into groundwater is initially weak in the near future and subsequently increase. This 578 causes these areas encountering dual issues of groundwater depletion and waterlogging during 579 different periods. This phenomenon is even more pronounced in the scenario of high emissions. 580 Moreover, simulations of groundwater depth trajectories have certain variations in the same 581 emission scenario, which are caused by uncertainties in future predictions of precipitation and 582 temperature of different GCMs (Figure 9 and Supplementary Figures S5 – S18).



Figure 9 Projected trajectories of the groundwater depth in scenario S0 and scenario S15 in a grid
at (a) region A, (b) region B, (c) region C, and (d) region D. The coordinates in the figure are the
locations of the grids at region A, region B, region C, and region D. The results of other scenarios
are displayed in Supplementary Figures S5 – S18.

588 **5 Summary and Discussion**

589 **5.1 Summary**

590 This paper develops a novel framework for evaluating the sustainability of groundwater 591 resources. Assuming that the groundwater depth fluctuations resemble the operation of reservoirs, 592 the framework utilizes the grid-scale dynamic EGWD considering multiple protective targets as 593 the criterion to determine whether the groundwater depth remains within an appropriate range. To 594 conduct a comprehensive evaluation of groundwater resource sustainability, the framework 595 incorporates reliability, resilience, and vulnerability indexes to quantify the frequency, duration, 596 and extent of groundwater depths exceeding threshold ranges of the EGWD. These quantifications 597 are based on simulation results of an integrated surface water and groundwater model under 598 multiple scenarios. By employing this framework, we can comprehensively assess risks associated 599 with groundwater overexploitation and potential for groundwater resource utilization in different 600 areas and periods. Moreover, it facilitates the identification of key areas that require focused 601 attention in the future groundwater resources management efforts.

602 The proposed framework is applied to the LTRB located in Northeast China for evaluating 603 the sustainability of groundwater resources under future climate change. The gird-scale dynamic 604 EGWD considering 4 protective targets is calculated, and the dominated protected target is 605 determined. During the non-growth period and growth period, upper thresholds of the EGWD 606 exhibit a range of 1.16~2.05 meters and 1.16~4.05 meters, respectively, while lower thresholds of 607 the EGWD show a range of 6.28~33.54 meters and 4.87~30.72 meters, respectively. Long-term 608 trajectories of groundwater depths are simulated using an integrated SWAT-MODFLOW model 609 under 16 historical and future climate scenarios. The results show that increased precipitation in 610 future climate scenarios improves the reliability performance in areas with deep groundwater 611 depths. However, it also increases the possibility of groundwater depths exceeding upper thresholds of the EGWD in areas with shallow groundwater depths. Notably, changes in future 612 613 climate conditions lead to longer duration and greater intensity of extreme dry and wet events 614 compared to the historical scenario. In the historical scenario, extreme dry and wet events last less 615 than 20 months and 10 months, respectively. While in future climate scenarios, these extreme

616 events extend beyond 30 months for dry periods and 20 months for wet periods. Consequently, the 617 cumulative intensity of single extreme climate event in future climate scenarios exceeds that of the 618 historical scenario. It causes a significant expansion of areas experiencing poor performances of 619 resilience and vulnerability indicators in future climate scenarios. These findings indicate that 620 even with an increase in average precipitation in the future, extreme climate conditions give rise to 621 different issues of groundwater depletion and waterlogging in different regions. Moreover, climate 622 change induces fluctuations of the intensity of precipitation infiltration recharge into groundwater 623 during different periods in the future, which causes dual challenges of groundwater depletion and 624 waterlogging during different periods in some regions. This phenomenon is even more 625 pronounced in the scenario of high emissions. Overall, these results highlight the significant 626 challenges faced in managing groundwater resources in the future, emphasizing the importance 627 and urgency of effective groundwater resource management in the LTRB.

628

5.2 Implications for regional groundwater resource management

There are significant challenges to the sustainability of groundwater resource management in the LTRB under future climate change. The proposed framework in this paper enables the evaluation of overexploitation risks and exploitation potential of groundwater resources in different areas and periods under the impacts of climate change, and facilitates the identification of key areas for future groundwater resource management. Our results provide some implications to facilitate informed decision-making and enable effective measures to be implemented in managing groundwater resources in the future.

636 There is a significant need for the groundwater management sector to enhance its capacity to 637 address extreme climate conditions. Although the reliability performance of groundwater system 638 improves due to the increased precipitation in future climate scenarios, extreme climate events 639 lead to poorer performances of resilience and vulnerability compared to the historical scenario 640 (Figures $5 \sim 7$). In the context of impending climate changes, extreme climate conditions impose 641 more stringent demands on groundwater resource management (Candela et al., 2012; Le Brocque 642 et al., 2018), which is vital for ensuring regional security of water supply, food and ecology (Liu et 643 al., 2022; Pagán et al., 2016; Yan et al., 2014).

644 Optimizing the crop planting structure can be an effective measure to achieve spatially 645 sustainable groundwater exploitation (Luo et al., 2022). In the LTRB, where the unsustainability 646 of groundwater resources is observed due to groundwater depths exceeding threshold ranges of the 647 EGWD (Figure 8), the predominant land use is cultivated land, with groundwater irrigation 648 accounting for over 80% of the total water supply. This indicates that the crop planting structure 649 plays a significant role in determining the spatial patterns of groundwater exploitation. 650 Implementing a conversion from paddies to drylands in areas with high-intensity groundwater 651 exploitation and a tendency for groundwater depths to fall below the lower threshold of the 652 EGWD, such as region I, can increase the groundwater level. Implementing a conversion from 653 drylands to paddies in areas with low-intensity groundwater exploitation and shallow groundwater 654 depths, where groundwater depths are likely to exceed upper thresholds of the EGWD, such as 655 region III, can decrease the groundwater level. This highlights the importance of considering the 656 spatial optimization of crop planting structure as a strategy to achieve sustainable groundwater 657 exploitation. By implementing appropriate land use changes, it is possible to mitigate the 658 challenges associated with groundwater unsustainability and promote the long-term viability of 659 groundwater resources.

660 Taking a long-term perspective, it is crucial to manage groundwater exploitation in a manner 661 that makes full use of the regulation and storage capacity of the aquifer. If the historical 662 groundwater exploitation scheduling continues to be implemented in future climate scenarios, 663 there are areas experiencing dual issues of groundwater depletion and waterlogging during 664 different periods (Figures 8 and 9). To address these challenges, it is necessary to implement 665 different measures during different periods in the future. In the near future, groundwater resource 666 management measures should be implemented to mitigate the risk of groundwater depths falling 667 below the lower thresholds of the EGWD. In the medium term, measures should be taken to 668 prevent the risk of groundwater depths exceeding the upper thresholds of the EGWD. 669 Additionally, continuous monitoring of groundwater dynamics is essential to maintain 670 groundwater depths within the threshold ranges of the EGWD in the long run. While it will make 671 the management of groundwater resources more challenges, it is important to recognize that the 672 aquifer, that is, groundwater reservoir, possesses strong interannual regulation capabilities (Dai et al., 2005). Therefore, it is necessary to manage groundwater exploitation from a long-term
perspective, leveraging the aquifer's regulation and storage abilities to achieve rational
groundwater exploitation over time.

676 **5.3 Limitations and future work**

677 A comprehensive exploration that considers the coupling impacts of both climate change and 678 human activities on groundwater resources is crucial for a holistic understanding of groundwater 679 system. The evaluation of groundwater resource sustainability in this study only focuses on the 680 impact of future climate change. Although it can assist in formulating management measures for 681 groundwater resources in the future, the growing prominence of human activities and their effects 682 on groundwater resources are not considered in this study. By incorporating the effects of specific 683 human activities, such as planting patterns (Wang, et al., 2022), irrigation system layouts 684 (Shandany et al., 2018), and irrigation efficiency (Zhang et al., 2017), alongside climate change, a 685 more accurate evaluation of groundwater resource sustainability can be achieved. Understanding 686 the interplay between these factors and groundwater resources can establish a more robust and 687 adaptive approach to groundwater resource management and ensure the long-term sustainability of 688 this vital water source.

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692 Data availability

All the data used in this study are previously published and can be accessed. The monthly normalized vegetation index (NDVI) data can be obtained from Resources and Environmental Science and Data Center, Institute of Geographic Sciences and Natural Resources Research, Chinese Academy of Sciences (<u>https://www.resdc.cn</u>). Public data sets used to drive the integrated SWAT-MODFLOW model simulations are available via Geospatial Data Cloud site, Computer

698 Network Information Center, Chinese Academy of Sciences (http://www.gscloud.cn) (DEM), 699 Resources and Environment Science and Data Center, Chinese Academy of Sciences 700 (http://www.resdc.cn) (land-use map), National Cryosphere Desert Data Center (Lu & Liu, 2019) 701 (soil map), National Meteorological Information Center, China Meteorological Administration 702 (http://data.cma.cn) (precipitation, maximum and minimum temperature, relative humidity, wind 703 speed), Global Land Evaporation Amsterdam Model (Martens et al., 2017; Miralles et al., 2011) 704 (remote sensing evapotranspiration), and Coupled Model Intercomparison Project Phase 6 705 (CMIP6) (https://esgf-node.llnl.gov/search/cmip6) (future precipitation, maximum and minimum 706 temperature). Other data used to develop the integrated SWAT-MODFLOW model, including 707 river network, boreholes, surface water and groundwater extracts, streamflow observations, and 708 groundwater level observations, are available from the corresponding authors upon reasonable 709 request.

710 Author contributions

Bo Xu conceived the research. Mingjun Wang performed the research. Mingjun Wang, Bo
Xu, and Chi Zhang analyzed the results and drafted the manuscript. Yong Peng and Yu Li polished
the manuscript. Xinqiang Du constructed the ecological groundwater depth. Chi Zhang and Bing
Yu provided funding support.

715 Competing interests

716 The authors declare no competing interests.

717 Additional information

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