

Numerical investigation of the refractive properties of near-horizontal shore platforms and their effects on harmonic and stationary wave patterns

Raphael Krier-Mariani¹, Stephenson Wayne¹, Sarah Wakes², and Mark E. Dickson³

¹University of Otago

²Department of Mathematics and Statistics, University of Otago

³University of Auckland

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Abstract

Near-horizontal shore platforms display highly irregular plan shapes, but little is known about the way in which these irregularities influence the significant wave height (H_s) on the platforms and the frequency components of the nearshore wave field. We use a nonlinear Boussinesq wave model to conduct harmonic and bispectral mode decomposition analyses, studying the control of concave and convex platform edges over wind (WW: 0.125 - 0.33 Hz), swell (SW: 0.05 - 0.125 Hz) and infragravity (IG: 0.008 - 0.05 Hz) frequencies. For breaking and non-breaking waves, increasing the platform edge concavity intensified wave divergence and subsequent attenuation of SW and IG across the outer platforms, reducing by up to 25%. Increasing the platform edge convexity intensified focusing and amplification of SW and WW over the outer platforms, increasing by up to 18% and 55% for breaking and non-breaking waves. In the presence of breaking, IG amplification depended on the generation of wave divergence across the inner platform, a condition determined by a critical convex curvature threshold ($K=1.8$) balancing wave focusing from refraction and defocusing from breaking. We find that convex curvature can determine the relative dominance of WW, SW and IG across platforms. Alongshore, coherent wave interactions governed IG stationary patterns defined by a node near the platform centreline and two antinodes on either side of concave edges. A node was generated at the platform centreline, and two antinodes were observed on either side of the convex edges for $K>1.8$, with the opposite pattern observed for $K<1.8$.

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Numerical investigation of the refractive properties of near-horizontal shore platforms and their effects on harmonic and stationary wave patterns

R. Krier-Mariani¹, W. Stephenson¹, S. Wakes², M. Dickson³

¹ School of Geography, University of Otago, New Zealand

² Department of Mathematics and Statistics, University of Otago, New Zealand

³ School of Environment, University of Auckland, New Zealand

Corresponding author: Raphael, Krier-Mariani (raphael.m.krier@gmail.com; raphael.krier@postgrad.otago.ac.nz)

28 **Key Points:**

- 29
- Through refraction, concave and convex near-horizontal shore platforms can separate the frequency components of the wavefield.
- 30
- Refraction patterns controlled by platform edge convexity affect the dominance of wind, swell and infragravity waves across platforms.
- 31
- Coherent amplification from the intersection of refracted infragravity waves controls the nodal state of alongshore stationary waves.
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35 **Abstract**

36 Near-horizontal shore platforms display highly irregular plan shapes, but little is
37 known about the way in which these irregularities influence the significant wave height (\widehat{H}_s)
38 on the platforms and the frequency components of the nearshore wavefield. We use a non-
39 linear Boussinesq wave model to conduct harmonic and bispectral mode decomposition
40 analyses, studying the control of concave and convex platform edges over wind (WW: 0.125
41 - 0.33 Hz), swell (SW: 0.05 - 0.125 Hz) and infragravity (IG: 0.008 - 0.05 Hz) frequencies. For
42 breaking and non-breaking waves, increasing the platform edge concavity intensified wave
43 divergence and subsequent attenuation of SW and IG across the outer platforms, reducing \widehat{H}_s
44 by up to 25%. Increasing the platform edge convexity intensified focusing and amplification
45 of SW and WW over the outer platforms, increasing \widehat{H}_s by up to 18% and 55% for breaking
46 and non-breaking waves. In the presence of breaking, IG amplification depended on the
47 generation of wave divergence across the inner platform, a condition determined by a critical
48 convex curvature threshold ($|\mathcal{K}|=1.8$) balancing wave focusing from refraction and
49 defocusing from breaking. We find that convex curvature can determine the relative
50 dominance of WW, SW and IG across platforms. Alongshore, coherent wave interactions
51 governed IG stationary patterns defined by a node near the platform centreline and two
52 antinodes on either side of concave edges. A node was generated at the platform centreline,
53 and two antinodes were observed on either side of the convex edges for $|\mathcal{K}|>1.8$, with the
54 opposite pattern observed for $|\mathcal{K}|<1.8$.

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56 Plain Language Summary

57 Near-horizontal shore platforms fronting coastal cliffs act as wave energy buffers,
58 regulating wave-induced erosion in rock coast environments. Genuine research endeavours
59 have permitted establishing the link between near-horizontal platform morphology and wave
60 transformation across-shore. However, the effects of alongshore variations in near-horizontal
61 platform morphology on the properties of nearshore wavefields remain sparsely
62 documented. As ocean waves share akin refractive properties to light rays, it can be assumed
63 that, similarly to optical lenses, shore platforms can separate waves according to their
64 frequency depending on their geometry. Subsequently, the convergence and divergence of
65 refracted wave trains of similar phases and frequencies could affect the properties of the
66 nearshore wavefield. The present research investigates this phenomenon over concave and
67 convex edge platforms and its impact on the nearshore wavefield characteristics. Our results
68 show that wave refraction over near-horizontal platforms with concave and convex edges
69 affects the relative dominance of short, medium and long-period waves across shore and
70 results in alongshore stationary wave patterns near the shoreline with nodal states varying in
71 relation to platform edge geometry. Such patterns likely result in alongshore variations in
72 wave erosion and the generation of wave-generated currents shaping rock coasts in the
73 platform.

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78 1 Introduction

79 Near-horizontal shore platforms, defined by a low gradient ($\tan\beta < 0.0175$) and a
80 steep seaward edge, are prevailing coastal landforms in rock coast environments (Sunamura,
81 1992; Trenhaile, 1999). These landforms have an essential role in wave transformation
82 processes, regulating wave erosive forces at the shoreline (Stephenson and Kirk, 2000;
83 Matsumoto et al., 2016a,b). Thus, an accurate description of the geomorphic control exerted
84 by shore platforms on nearshore wave transformation patterns is necessary for improving
85 rock coast geomorphological models.

86 Studies have investigated the control of near-horizontal shore platform morphology on
87 the cross-shore evolution of the wavefield (e.g. Beetham and Kench, 2011; Marshall and
88 Stephenson, 2011; Ogawa et al., 2016). Wave breaking induced by the sharp depth transition
89 at the seaward edge of a platform results in the dissipation of incident swell waves (SW: 0.05
90 Hz < f < 0.125 Hz) and the generation of low-frequency infragravity waves (IG: f < 0.05 Hz)
91 over the platform (Poate et al., 2020). Across the platform surface, IG gradually amplify due
92 to shoaling and energy is transferred from high to lower frequencies, becoming the dominant
93 frequency component over the inner platform (Beetham and Kench, 2011; Marshall and
94 Stephenson, 2011; Ogawa et al., 2011). Wind waves (WW: 0.125 < f < 0.33 Hz) can propagate
95 onto platforms from offshore and, in some cases, be locally generated over the outer platform
96 to become the dominant frequency component in this area (Marshall and Stephenson, 2011;
97 Ogawa et al., 2011). These observations were summarised in the conceptual model of Ogawa
98 et al. (2011), indicating that it is common for the outer platform, platform centre, and inner
99 platform to be dominated by WW, SW and IG, respectively. Ogawa et al. (2011) suggested
100 that these zones shift across-shore with tidal elevation and showed that the relative
101 submergence of shore platforms (depth at the seaward edge/incident wave height) is a critical
102 factor controlling the relative dominance of SW and IG. Collectively, understanding the
103 behaviour of each frequency band of the wavefield helps to depict the variation of significant
104 wave height (H_s) across platforms affecting erosion of the platform and cliff (Trenhaile,
105 2000). However, the impact of shore platform morphology on two-dimensional wave
106 transformation processes and effect on the frequency bands composing the wavefield have
107 been overlooked.

108 Few field studies have considered the impact of the platform morphology of near-
109 horizontal platforms on two-dimensional wave transformation patterns (Krier-Mariani et al.
110 2022, 2023). Krier-Mariani et al. (2023) showed that directional patterns controlled by
111 irregularities in platform morphology generated localised areas of wave ray convergence and
112 divergence as well as alongshore variations in standing IG patterns, influencing the wave
113 energy distribution over the platform surfaces. Based on these observations, Krier-Mariani et
114 al. (2023) introduced a conceptual model in which concave and convex platform edge
115 geometries would control wave ray convergence and divergence patterns over the platform
116 surface, subsequently affecting the IG energy levels and SW decay rates. However, the

117 influence of platform edge geometry on two-dimensional wave patterns could not be clearly
118 isolated from field observations.

119 In the absence of detailed field studies on the effects of platform edge geometry on
120 wave transformation characteristics, the literature on morphologically analogous submerged
121 flat structures is useful. Depending on their geometry, submerged flats can separate the
122 frequency components of the wavefield, refracting and reorganising the wave crests of
123 incident waves according to their frequency (Jarry et al., 2011; Griffiths and Porter, 2012; Li
124 et al., 2020). This phenomenon can result in complex refraction patterns specific to each
125 frequency component of the wavefield, leading to the generation of caustic rays (clusters of
126 caustic points generated by wave ray intersection) over submerged surfaces (e.g. Mandlier
127 and Kench, 2012). Patterns of wave ray convergence and divergence induced by refraction
128 over submerged flat structures significantly impact the wavefield characteristics. Wave ray
129 convergence results in a localised enhancement of wave height (e.g. Ito and Tanimoto, 1972;
130 Berkhoff et al., 1982), skewness and kurtosis (Janssen and Herbers, 2009; Jarry et al., 2011;
131 Lawrence et al., 2022) while wave ray divergence has the opposite effects.

132 Although relatively few studies have considered the impact of submerged flat
133 geometries on the cross-shore evolution of harmonic and subharmonic components of the
134 wavefield, harmonic components amplification has been observed in areas of wave
135 convergence (e.g. Lynett and Liu, 2004; Gouin et al., 2017). According to Li et al. (2020), this
136 phenomenon could be attributed to the non-linear effects of convergence on wave height
137 amplification. As the geometry of submerged flats influences the cross-shore pattern of wave
138 convergence (intensity and location) of each harmonic, it likely also influences the cross-shore
139 patterns of wave harmonics amplification, intrinsically affecting the dominance of different
140 wave frequencies across platforms. This hypothesis as yet to be verified.

141 It has proven difficult to establish causality between patterns of wave ray intersection,
142 increased nonlinearity and alongshore wave height amplification for random wavefields,
143 notably due to the limitation of wave ray tracking techniques to evaluate complex wave ray
144 crossing patterns in dense constellations of caustics (Ito and Tanimoto, 1972). Another way
145 of approaching this problem involves considering the impact of coherent wave interaction
146 patterns on the amplification of dominant frequency components of the wavefield. Coherent

147 wave interaction refers to the non-linear process occurring at the intersection of waves with
148 similar frequency, waveform and phase. It has been identified as a fundamental non-linear
149 wave amplification process in optics (e.g. Young, 1802), quantum mechanics (e.g., Weiland
150 and Wihelmsson, 1977; Falk, 1979; Inouye et al., 1999; Kozuma et al., 1999) and geoscience
151 (e.g. Harid et al., 2014). There have been few investigations of this process in coastal wave
152 studies, but Dalrymple (1975) demonstrated that this process could result in the formation of
153 alongshore stationary wave patterns in random wavefields and the subsequent formation of
154 nearshore currents. More recently, Tamura et al. (2020) showed that, similar to light
155 refraction through a prism, ocean wave refraction over a submarine canyon could separate
156 waves of a random wavefield according to their frequency and phase, favouring coherent
157 wave interactions. Based on this theoretical grounding, it is hypothesised that by controlling
158 the refraction patterns of individual frequency components of the wavefield, submerged flat
159 (e.g. shore platforms) geometry affects coherent wave amplification over submerged flat
160 surfaces, leading to the generation of alongshore stationary wave patterns for SW and IG.

161 The impact of shore platform geometry on the behaviour of wave harmonics and
162 stationary wave patterns remains to be evaluated in detail on near-horizontal platform
163 surfaces. However, such a task was proven to be difficult during field observations due to the
164 variable nature of nearshore wavefields and the morphological complexity of shore platforms
165 (e.g. Krier-Mariani et al. 2022, 2023). Therefore, this study adopts a numerical modelling
166 approach to address the question: How do mesoscale variations in platform edge geometry
167 affect the behaviour of wave harmonics and the subsequent wave height distribution across
168 and along platform surfaces?

169 **2 Method**

170 **2.1 Model set up**

171 The phase-resolving Boussinesq wave model FUNWAVE_TVD V3.6 (Shi et al., 2012)
172 was used to investigate two-dimensional wave transformation over shore platforms. This
173 model treats wave transformation in the time domain and provides a robust representation
174 of non-linear processes, refraction and diffraction while retaining information on the wave

175 phase (Sheremet et al., 2011; Buckley et al., 2015, Buckley et al., 2018; Thomas and
176 Dwarakish, 2015).

177

178 **2.1.1 Domain**

179 Idealised three-dimensional near-horizontal platform morphologies were incorporated
180 into a 1274 m (x-axis) to 300 m (y-axis) domain (Fig. 1a). A 0.35 m deep, 250 m wide shallow
181 planar surface was included at the landward extremity to absorb wave energy and minimise
182 resonance. The platforms were defined by a constant gradient of 0.35 degrees, a width of 300
183 m (at the centreline, $y = 150$ m) and a 3m high seaward cliff of 45 degrees. The nearshore
184 bathymetry profile was composed of a 480 m subtidal ramp (at the centreline) with a gradient
185 of 0.35 degrees followed by an 8 m deep and 635 m wide flat.

186 Planform geometry was represented using three generic edge geometries defined as
187 straight, concave and convex. The degree of curvature of the concave ($\mathcal{K} < 0$) and convex
188 ($\mathcal{K} > 0$) edge geometries was derived from the parametric ellipse equation. The semi-major
189 axis (a , along the x-axis) was kept constant (120 m) to avoid modifying the cross-shore profile
190 along the centreline, and various degrees of edge curvature were obtained from 2 m
191 increments along the semi-minor axis (b , along the y-axis) between 50 to 100 m, resulting in
192 26 cases with edge curvatures ($|\mathcal{K}| = |a/b|$) ranging from 1.2 to 2.4 (Fig. 1b-f). The
193 bathymetry was smooth to reduce noise generated by sharp edges and interpolated to a 2 m
194 grid adopted to ensure model stability following a series of sensitivity analyses, providing a
195 realistic representation of model resolution used in previous research in nearshore areas (e.g.
196 Su et al., 2021).

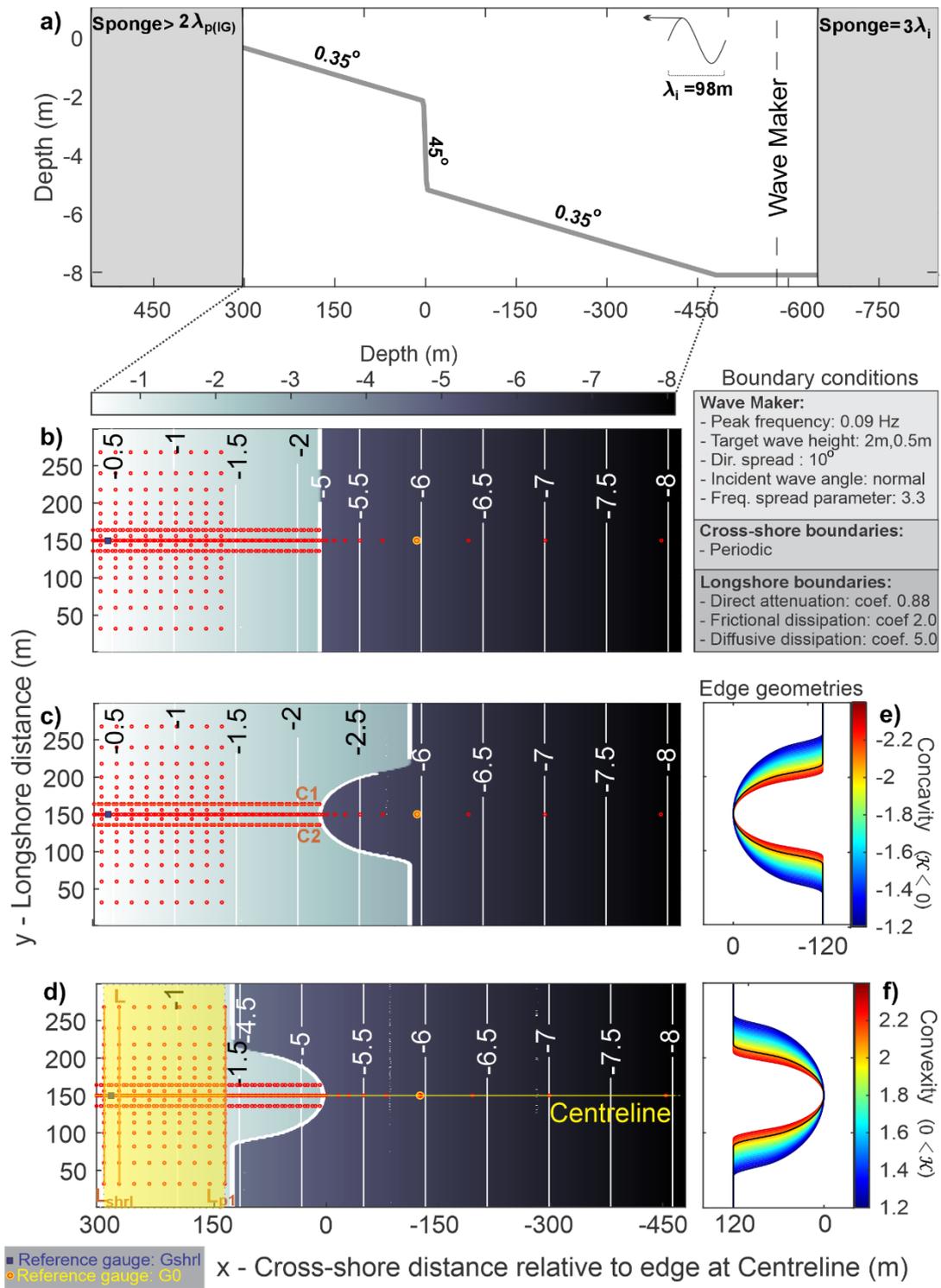


Figure 1: Boussinesq wave model configuration showing the bathymetry profile along the centreline ($y = 150\text{m}$) (a), the model domain for the straight, concave and convex platforms (b-d), and the range of platform edge curvatures considered (e,f). Specifications of the boundary conditions are annotated in the figure. The red dots mark the location of the virtual gauges used for analysis. The yellow shaded area (between L_p and L_{srl}) represents the inner platform section considered for alongshore analysis.

197 **2.1.2 Wave conditions**

198 The model was forced by irregular waves with a directional spread of 10 degrees. An
199 internal wavemaker (Wei et al., 1999) was located on the deep flat at the bottom of the
200 subtidal ramp, five wavelengths (λ_i) away from the platform edge to avoid distortion of the
201 initial wave crests. Irregular waves were generated using a JONSWAP wave spectra
202 (Hasselmann et al., 1973) with a fixed peak enhancement factor of 3.3, a peak frequency (f_p)
203 of 0.09 Hz and direction of 0° (shore-normal) to simplify the visualisation of the refraction
204 effects induced by different platform edge geometries.

205 Two sets of simulations were generated to investigate the transformation of: (1) waves
206 propagating across the platform surface without breaking ($H_s= 0.5$ m), as such waves can
207 release large amounts of erosive energy when they break against cliffs (Thompson et al.,
208 2019; Thompson et al., 2022); and (2) wave breaking at the seaward edge ($H_s= 2$ m) decaying
209 across the platform, which are typically used to define variation of wave erosive force across
210 platforms in geomorphological models (e.g. Trenhaile, 2000; Matsumoto et al., 2016a,b).
211 These two sets of simulations combined with the range of platform concave and convex
212 curvatures resulted in 106 simulations (including straight edge reference cases). The default
213 breaking index of FUNWAVE-TVD ($\gamma_b=0.80$) was used to represent wave breaking, providing
214 a close representation of the breaking conditions for steep submerged slopes (Blenkinsopp
215 and Chaplin, 2008). The effects of bottom friction were not considered (i.e. the frictional
216 dissipation coefficient was set to $C_d= 0.002$, representing a smooth surface).

217

218 **2.1.3 Boundary conditions**

219 The domain boundary conditions were defined to minimise reflection. Periodic
220 boundaries (Chen et al., 2003) were applied to the northern and southern extremities of the
221 domain, allowing waves to propagate out of the domain. Following Shi et al. (2016), sponge
222 layers employing a direct damping coefficient as well as dissipation by friction and diffusion
223 were used to reduce noise and dampen wave energy at the eastern and western sides of the
224 domain (Fig. 1a). The width of the sponge layer at the shallow western side of the domain
225 was chosen to correspond to twice the peak wavelength of the IG at this location (estimated

226 during trial runs using the virtual gauge G_{shrl} at the shoreline, Fig. 1), to avoid reflection and
227 the subsequent generation of standing IG waves.

228

229 **2.1.4 Model Validation**

230 Due to the lack of two-dimensional field measurements in similar near-horizontal
231 shore platform settings, no direct validation of our model simulations was carried out.
232 However, a number of studies validated FUNWAVE-TVD against field observations over coral
233 reefs, proving the model's ability to represent wave transformation over smooth submerged
234 flats with sharp seaward edges (e.g. Mendonca et al., 2008; Su et al., 2015; Zhang et al., 2019).
235 As the present study explores wave processes such as refraction and non-linear energy
236 transfer fairly well represented by the model (Griffiths and Porter, 2012; Su et al., 2015) and
237 does not investigate subsequent processes such as wave-driven circulation and sediment
238 transport, it is deemed unnecessary to validate the model with experimental data at this stage
239 (similar inference were made in da Silva et al., 2023).

240 **2.2 Measurements and analysis**

241 To determine the impact of planform geometries on wave transformation across the
242 platforms, the spectral evolutions of waves propagating across concave and convex platforms
243 (affected by two-dimensional transformation processes) were compared to the spectral
244 evolution of waves propagating across the straight-edge platform (only affected by on-
245 dimensional transformation processes). This approach permitted the identification of spectral
246 anomalies representing the energy variations for specific harmonics induced by refraction.
247 Positive and negative anomalies indicate harmonic amplifications and attenuation,
248 respectively. Combined, the harmonic anomalies result in anomalies of significant wave
249 height across platforms ($\Delta\widehat{H}_s$). Following Baldock et al. (2020), the cross-shore patterns of
250 $\Delta\widehat{H}_s$ were then compared to the directional patterns along the platform centrelines to identify
251 the effects of refraction patterns controlled by platform edge geometry on significant wave
252 height distribution across platforms.

253 In the alongshore, the effects of coherent wave interaction induced by refraction over
254 concave and concave platforms on the generation of stationary wave patterns were

255 considered. For this purpose, the bispectrum (Hasselmann et al., 1963) provides a convenient
256 representation of the wavefield as it holds information on the wave phase, frequency and
257 power necessary to detect phase coupling. The bispectrum, defined from the third moment
258 of the free surface elevation time series, also represents a measure of skewness, which
259 increases in areas of wave ray intersection (Janssen and Herbers, 2009; Jarry et al., 2011;
260 Lawrence et al., 2022). Following Kim and Powers (1979), who investigated the impact of
261 coherent interactions of random electromagnetic waves on plasma density fluctuation using
262 bispectral properties, the frequency, phase and power information yielded by the bispectrum
263 were used to identify patterns of coherent wave interactions over the inner platforms. A
264 modal decomposition method based on bispectral properties (Appendix 2), the Bispectral
265 Mode Decomposition or BMD (Schmidt, 2020), was employed to identify the modal state of
266 coherent structures for self-interacting harmonic components within the SW and IG
267 frequency bands. The areas of coherent wave interactions were then compared to the wave
268 height distribution of SW and IG over platforms of various geometries to identify patterns of
269 coherent wave amplification.

270

271 **2.2.1 Wave measurements**

272 Wave records were obtained from virtual gauges recording surface elevation (η) as
273 well as u and v velocity components at 2 Hz (Fig. 1b-d). In the cross-shore direction, the gauge
274 spacing along the centreline increased seaward from the platform edge (increment based on
275 geometric series starting with a spacing of 4 m with an increment factor of 1.5). On the
276 platforms, the gauge spacing was irregular but not exceeding 6 m along the centreline,
277 transects C_1 and C_2 . The distance between the gauges composing the alongshore transects
278 (between L_0 and L_{shrl}) increased on either side of the centreline from 6 to 30 m (with an
279 increment factor of 1.25). Statistical analyses of the wavefield properties were based on an
280 observation window of 2048 seconds, starting 230 seconds after the start of the simulations,
281 marking the time at which SW reached the landward extremity of the domain and IG were
282 generated.

283

284 2.2.2 Definition of wave height

285 The significant wave height (H_s) was defined from the spectra moment (e.g. Thornton
286 and Guza, 1983):

$$H_s = 4 \sqrt{\int_{f_{min}}^{f_{max}} S(f) \cdot df} \quad (1)$$

287

288 The wave spectra estimates $S(f)$ were generated using the Welch (1967) method with
289 segment lengths of 512 samples, 50% overlap and a Hanning window resulting in 20 Degrees
290 of Freedom (Priestley, 1981). To provide a more detailed representation of the wavefield, the
291 gravity and infragravity waves were further divided into two frequency bands, encapsulating
292 the dominant harmonics observed within the WW, SW, and IG (high and low) frequency
293 ranges across the domain (Table 1). The wave height associated with each of these frequency
294 bands was determined using:

295

$$H_{f_{np}} = 4 \sqrt{\int_{f_{low}}^{f_{high}} S(f) \cdot df} \quad (2)$$

296 where n denotes the rank of the harmonic, f_{low} and f_{high} represents the lower and higher
297 frequencies of the power spectral density peak associated with this harmonic, Table 1. The
298 reference incident wave height (H_0) was defined from measurements taken at the gauge G_0
299 located at the top of the subtidal ramp (Fig. 1) and was used to normalise the wave height on
300 the platform surface ($\widehat{H}_s(x) = H_s(x)/H_0$, $\widehat{H}_{f_{np}}(s) = H_{f_{np}}(x)/H_0$). For simplicity,
301 normalised wave heights are hereafter referred to as wave heights.

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309 **Table 1** Frequency band analysis parameters

Conventional Frequency class	Frequency subclass	Corresponding Harmonic	Frequency (f_{np})	Frequency range ($f_{low} - f_{high}$)
Gravity waves	Wind waves (WW)	Second harmonic	f_{2p}	0.15 – 0.20 Hz
	Swell waves (SW)	Principal harmonic	f_p	0.06 – 0.12 Hz
Infragravity waves	Infragravity High (IG _H)	Second subharmonic	$f_{1/2p}$	0.04 – 0.05 Hz
	Infragravity Low* (IG _L)	Fifth subharmonic	$f_{1/5p}$	0.008 – 0.03 Hz

310 *Note that the typical cutoff frequency for the lower portion of the IG frequency band is 0.005 Hz (e.g. Pequignet et al., 2014;
 311 Gawehn et al., 2016). However, the chosen cutoff frequency of 0.008 Hz is more appropriate to describe the low IG in the
 312 simulated wavefield as it corresponds to a trough in the power spectra estimate across the entire domain, which provides a
 313 better physical representation of the low IG.
 314

315 2.2.3 Definition of peak direction

316 The angle α between the peak direction of waves propagating on either side of the
 317 centreline (along the cross-shore transects C1 and C2, Fig. 1) was used to investigate the
 318 evolution of wave convergence and divergence along the platform centrelines. The peak
 319 direction of waves over the platform was estimated from the directional wave spectra
 320 $S(f, \theta) = S(f)G(\theta|f)$ calculated from the free surface elevation (η) and velocity
 321 components (u and v) time series by applying the Extension of the Maximum Entropy
 322 Principle (EMEP) method (Hashimoto et al., 1994). To this effect, segments of 512 samples
 323 were used to estimate the frequency spectra ($S(f)$) and 200 iterations to define the
 324 approximation of the spreading function ($G(\theta|f)$) resulting in 76 frequency bins and
 325 directional bins of 5°.
 326

327 2.2.4 Identification of coherent wave interaction patterns

328 The BMD was applied to the free surface elevation time series recorded by the two-
 329 dimensional virtual gauge array between L_p and L_{shrl} , marking the boundaries of the spatial
 330 domain ξ (Fig. 1). The welch periodograms employed in the BMD were computed using
 331 segments of 512 samples, 50% overlap and a Hanning window resulting in 20 Degrees of
 332 Freedom. Patterns of coherent wave interactions were identified from coherent self-

333 interaction maps ($\psi_{k,k}$) which are defined by the product of cross-frequency fields $\phi_{k \circ k}$ and
 334 the bispectral modes ϕ_{k+k} obtained from the BMD:

$$\psi_{k,k}(\xi, f_k, f_k) = |\phi_{k \circ k} \circ \phi_{k+k}| \quad (3)$$

335
 336 where the frequency k considered were f_p and $f_{1/5p}$, representing the dominant harmonics
 337 in the SW and IG frequency bands. The cross-frequency fields $\phi_{k \circ k}$ are maps of phase
 338 alignment for these frequencies, while bispectral modes ϕ_{k+k} represent the amplitude of
 339 oscillations of the sea surface at frequency $2k$. Conventionally, the largest values of the
 340 normalised coherent self-interaction maps $\widehat{\psi}_{k,k}$ indicate areas where phase coupling has the
 341 strongest effect on wave amplitude for the sum frequency $2k$. The interaction maps for
 342 straight wave crests with parallel wave rays are expected to be homogeneous alongshore. In
 343 contrast, for cases where wave crests are bent and wave rays intersect, interaction maps will
 344 be non-homogenous alongshore and display maxima in areas of wave ray intersection. In the
 345 presence of coherent wave amplification, maxima in coherent self-interaction maps
 346 correspond to areas of wave height amplification at frequency k .

347

348 **3 Results**

349 **3.1 Impact of planform geometry on across-shore wave transformation**

350 **3.1.1 Non-breaking waves ($H_0 = 0.5$ m)**

351 The spatial evolution of the spectral properties of non-breaking waves propagating
 352 across the domain was examined for the three types of platform geometries (Fig. 2), for which
 353 the power spectra density was concentrated around four distinctive frequency components
 354 (Fig. 2a): the second and the principal harmonics (f_{2p} and f_p) within the WW and SW
 355 frequency bands; and the second and fifth subharmonics ($f_{1/2p}$ and $f_{1/5p}$) within the IG_H and
 356 IG_L frequency bands.

357 The spectral anomalies observed over the concave platforms indicated an attenuation
 358 of the principal harmonic (Fig. 2b-e). This phenomenon intensified with increasing degrees of
 359 curvature (with minimum spectral anomalies at peak frequency reducing from -0.09 m² Hz⁻¹

360 at $|\mathcal{K}|= 1.2$ to $-0.41\text{m}^2 \text{ Hz}^{-1}$ at $|\mathcal{K}|= 2.4$). In contrast, an amplification of the second and
361 principal harmonics was observed across convex platforms (Fig. 2g-j), intensifying with
362 increasing edge curvature (with maximum spectral anomalies at peak frequency increasing
363 from $0.52 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 1.2$ to $1.37 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 2.4$).

364 The variation of spectral characteristics of each harmonic over the concave and convex
365 platforms can be expressed in terms of mean wave height anomalies ($\overline{\Delta H_{f_{np}}}$). The most
366 significant impacts of platform curvature on mean wave height anomalies were observed
367 within the WW and SW frequency bands. The mean wave height anomalies associated with
368 the second and the principal harmonics displayed a very strong linear dependency ($R^2 > 0.9$)
369 to the degree of platform edge curvature (Fig. 2f,k). The increase of curvature from $|\mathcal{K}|= 1.2$
370 to 2.4 promoted the attenuation of harmonics within the WW and SW frequency bands across
371 concave platforms and the amplification of these waves across convex platforms. The
372 attenuation of the second and principal harmonics across concave platforms of high curvature
373 $|\mathcal{K}|= 2.4$ corresponded to 9% and 15% of H_0 . Across convex platforms of high curvature, $|\mathcal{K}|=$
374 2.4, the amplification of the second and principal harmonics reached up to 11% and 29% of
375 H_0 . The mean wave height anomalies for the subharmonic in the IG_H and IG_L frequency bands
376 were negligible for nonbreaking waves.

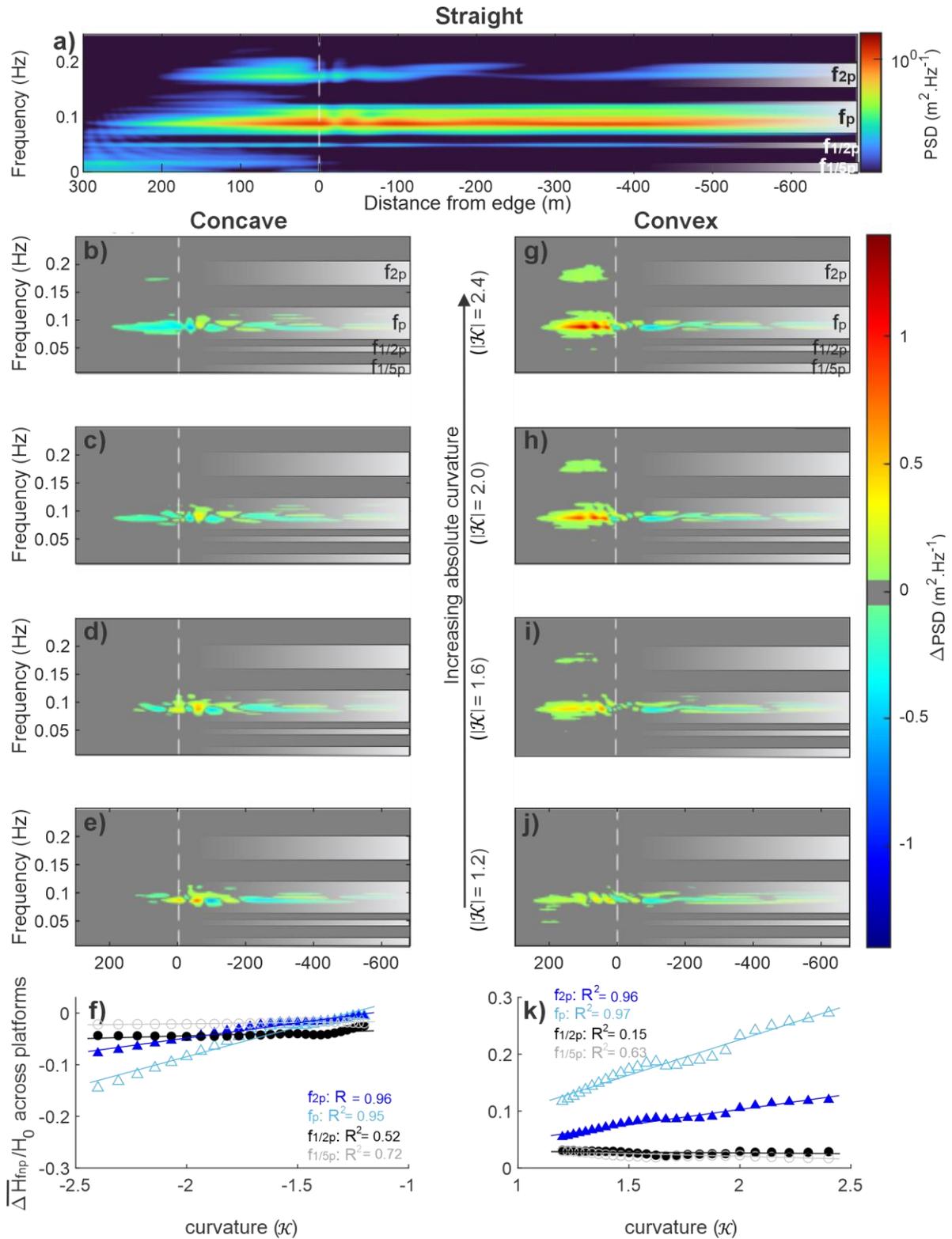


Figure 2: Impact of platform edge curvature on the harmonic components of the wavefield for non-breaking waves showing: the spectral anomalies in relation to the straight edge platform (a) for concave platform geometries (b-e) and convex platform geometries (g-j); and the impact of curvature on the mean wave height of each harmonic across the concave (f) and convex (k) platforms.

377

378

379 $\Delta\widehat{H}_s$ across the platform can be assumed to be impacted by refraction patterns
380 controlled by platform geometry. To explore this process, directional patterns and $\Delta\widehat{H}_s$ across
381 the centrelines of each platform were compared for various edge curvatures. The cross-shore
382 patterns of $\Delta\widehat{H}_s$ presented in Fig. 3a,b were modulated by the platform edge curvature, with
383 magnitude increasing with curvature for both types of platform geometries. As a result, for
384 high degrees of curvature ($|\mathcal{K}| = 2.4$), the negative anomalies across the concave platforms
385 indicated a maximum of 25% attenuation in significant wave height (Fig. 3a), while the
386 positive anomalies across the convex platforms (Fig. 3b) indicated a 55% amplification of
387 significant wave height. The location of the largest $\Delta\widehat{H}_s$ shifted across platforms in relation
388 to curvature. For concave platforms, the largest negative $\Delta\widehat{H}_s$ over the outer platform shifted
389 landward with decreasing curvature from $|\mathcal{K}| = 2.4$ to 1.9. Similarly, the largest positive $\Delta\widehat{H}_s$
390 across the convex platforms shifted landward, reaching the inner platform for $|\mathcal{K}| < 1.6$. For
391 low degrees of concave curvatures ($|\mathcal{K}| < 1.9$), corresponding to curvatures for which
392 amplification of wave energy seaward of the platform edge was observed (Fig. 2b-e), wave
393 transformation patterns across the platform centreline were affected by the preconditioning
394 of incident waves occurring off the platform edge. Therefore, the description of the following
395 results focuses on concave edge curvatures, $|\mathcal{K}| > 1.9$.

396 Similarly to the cross-shore evolution of $\Delta\widehat{H}_s$, the peak magnitude of wave ray
397 divergence observed across concave platforms and convergence across the convex platforms
398 decreased and shifted landward from the mid-platform ($x \approx 150$ m) to the outer platform
399 with decreasing curvature (Fig. 3c,d). A Spearman rank correlation (Fig. 3e,f) revealed that the
400 dependency of cross-shore $\Delta\widehat{H}_s$ on the directional patterns observed over the concave
401 platforms was only relevant (moderate to strong, $\rho_s > 0.4$) for platform edge curvatures
402 exceeding 1.9. In contrast, a strong relationship ($\rho_s > 0.6$) as observed between wave height
403 anomaly and directional patterns over convex platforms for the majority of platform edge
404 curvatures, indicating that $\Delta\widehat{H}_s$ across the convex platforms were predominantly controlled
405 by the wave convergence and divergence across the centreline.

406

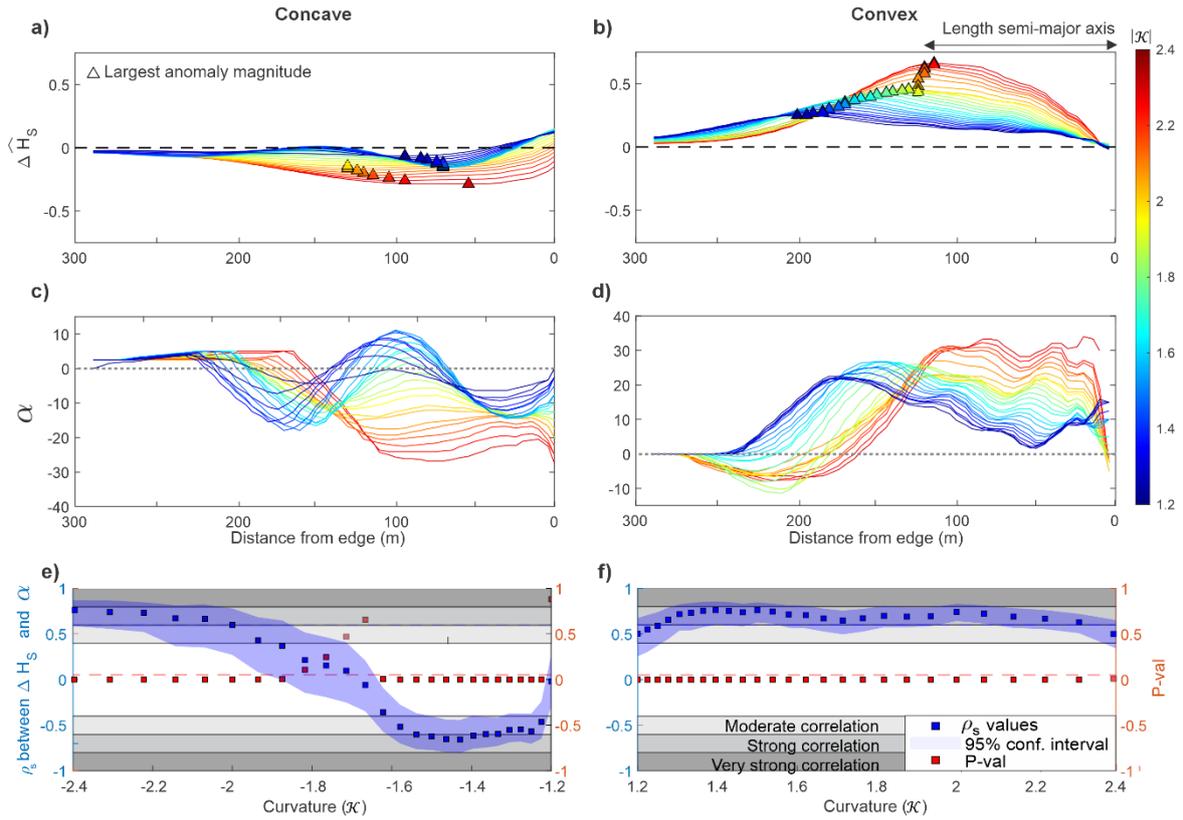


Figure 3: Relationship between directional patterns and significant wave height anomalies of non-breaking waves across concave (left) and convex platforms (right) at the centreline ($y = 150$ m) for different degrees of curvature showing: the significant wave height anomalies (a,b) and the cross-shore directional patterns (c,d). The impact of directional pattern on wave anomaly pattern was assessed using a spearman correlation between the two parameters (e,f).

407

408 3.1.2 Broken waves ($H_0 = 2.0$ m)

409 The spectral evolution of broken waves across concave platforms displayed a complex
 410 pattern of spectral anomalies (Fig. 4b-e), with a clear amplification of the principal harmonic
 411 corresponding to SW over the outer platform for degrees of curvature below 1.9 (at f_p ,
 412 positive anomalies reached $1.5 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K} = 2.0$ and 2.4 , 0 to 150 m from the edge) and a
 413 clear attenuation of this harmonic for degrees of curvature exceeding 1.9 (at f_p , negative
 414 spectral anomaly reached $-1.2 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K} = 1.2$ and 1.6 , 0 to 150 m from the edge). These
 415 differences were related to the amplification of the principal harmonic for edge curvatures
 416 below 1.9, displaying anomalies reaching up to $\sim 12 \text{ m}^2 \text{ Hz}^{-1}$ in the vicinity of the concave edge
 417 sections (0 to -120 m from the edge) before reaching the platform surface (Fig. 4d,e). Over
 418 convex platforms, the principal harmonic presented the largest amplification (Fig. 4g-j), which

419 intensified over the outer platform with increasing curvature (positive anomaly at f_p reached
420 $2.1 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 1.2$ and $4.1 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 2.4$, 0 to 150 m from edge). In contrast, the
421 amplification of subharmonics within the IG_H and IG_L frequency bands toward the shoreline
422 observed along the platform centreline was stronger for low convex edge curvatures than for
423 high convex edge curvatures (positive anomaly at $f_{1/5p}$ reached $0.7 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 1.2$ and 0.5
424 $\text{m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 2.4$, 150 to 300 m from edge).

425 Relationships between edge curvature and mean wave height anomalies across both
426 platform types were observed (Fig. 4f,k). For anomalies in the WW and SW frequencies, the
427 mean wave height anomalies of the second and the principal harmonics presented a strong
428 linear dependence on the degree of edge curvature of concave and convex edges ($R^2 > 0.90$).
429 In the IG_H frequency band, mean wave height anomalies associated with the second
430 subharmonic were linearly dependent on the curvature across concave platforms ($R^2 = 0.67$).
431 The mean wave height anomalies associated with the fifth subharmonic in the IG_L frequency
432 band decreased linearly ($R^2 = 0.94$) with curvature over the convex platforms.

433 Variations in edge curvature affected the relative importance of WW, SW, IG_H and IG_L
434 anomalies across the platforms. For concave platforms, the increase of concave edge
435 curvature promoted attenuation of all frequency bands, but particularly for WW and SW. For
436 the harmonic components within the WW and SW frequency bands, the mean wave height
437 attenuation across concave platforms was negligible for low curvature ($\overline{\Delta H_{f_{2p}}}$ and $\overline{\Delta H_{f_p}}$ and
438 representing less than 1% of H_0 at $|\mathcal{K}|=1.2$) but intensified for high degrees of curvature
439 ($\overline{\Delta H_{f_{2p}}}$ and $\overline{\Delta H_{f_p}}$ representing less than 3% and 6% of H_0 at $|\mathcal{K}|=2.4$). Across the convex
440 platforms of low curvature ($1.2 < |\mathcal{K}| < 1.75$), the largest amplification of mean wave height
441 was observed for the principal harmonic ($\overline{\Delta H_{f_p}}$ representing 4% to 7% of H_0), followed by the
442 fifth subharmonic ($\overline{\Delta H_{f_{1/5p}}}$ representing 3% to 3.5% of H_0). The amplification of the fifth
443 subharmonic became less important with increasing curvature, while the mean wave height
444 of the second harmonic was amplified. For convex curvatures exceeding 1.75, the principal
445 harmonic displayed the largest amplification ($\overline{\Delta H_{f_p}}$ representing 7% to 10% of H_0), followed
446 by the second harmonic ($\overline{\Delta H_{f_{2p}}}$ representing 3% to 4% of H_0). Thus, the reduction of convex
447 edge curvature promoted the amplification of IG_L , while the increase of convex edge
448 curvature promoted the amplification of WW and SW.

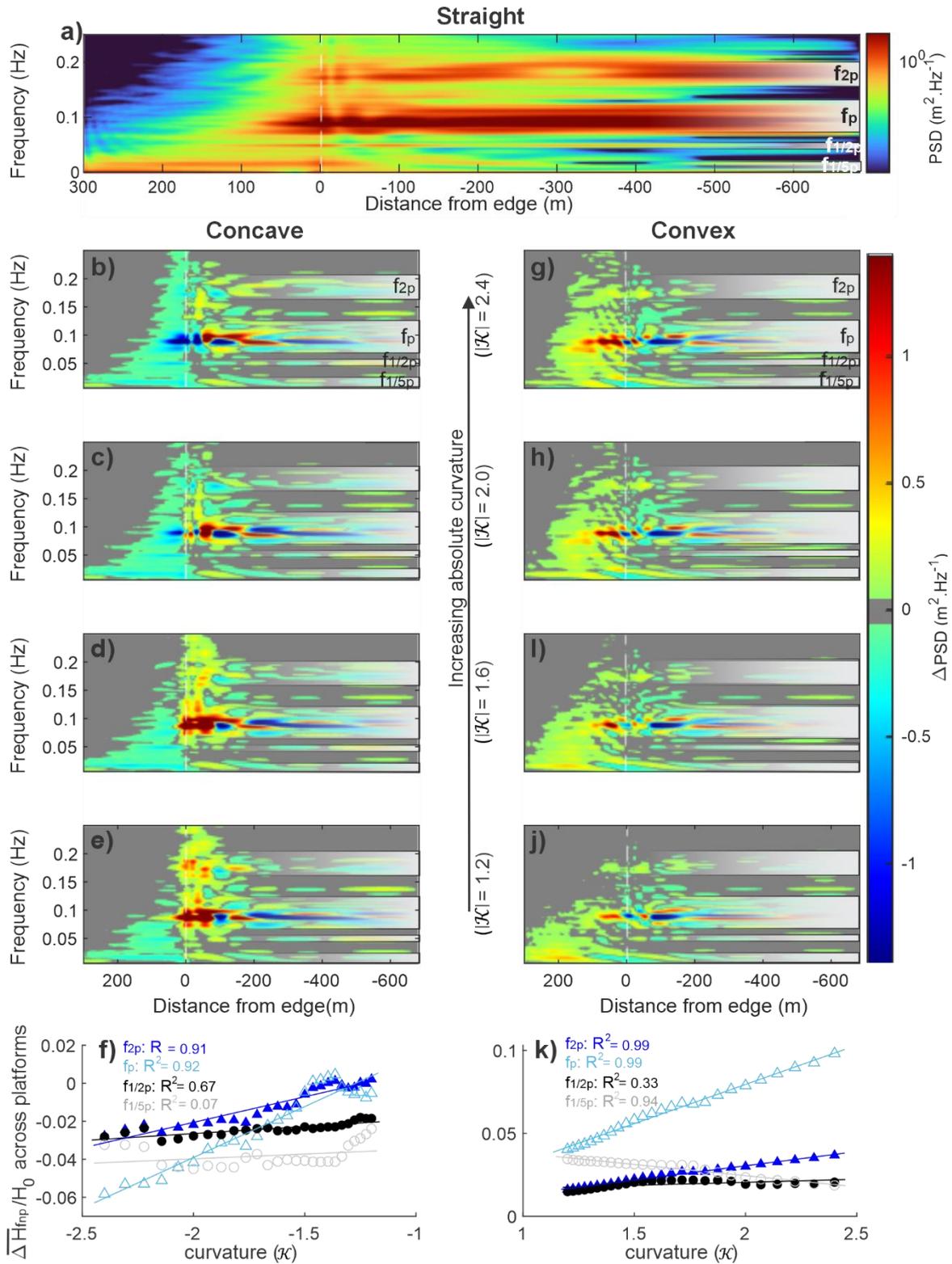


Figure 4: Impact of platform edge curvature on the harmonic components of the wavefield for broken waves showing: the spectral anomalies in relation to the straight edge platform (a) for concave platform geometries (b-e) and convex platform geometries (g-j); and the impact of curvature on the mean wave height of each harmonic across the concave (f) and convex (k) platforms.

449

450 The relationship between $\Delta \widehat{H}_s$ (Fig. 5a,b) and directional patterns (Fig. 5c,d) observed
451 across the concave and convex platform centrelines was more complex for broken than non-
452 breaking waves. The main difference with the non-breaking waves resided in the seaward
453 shift of the maximum divergence (Fig. 5c) and convergence (Fig.5d) locations over the outer
454 concave and convex platforms, respectively. This shift was particularly pronounced for convex
455 shore platforms with low degrees of curvature ($|\mathcal{K}| < 1.8$), for which the peaks of convergence
456 observed mid-platform ($x \approx 175$ m, Fig. 3d) were attenuated.

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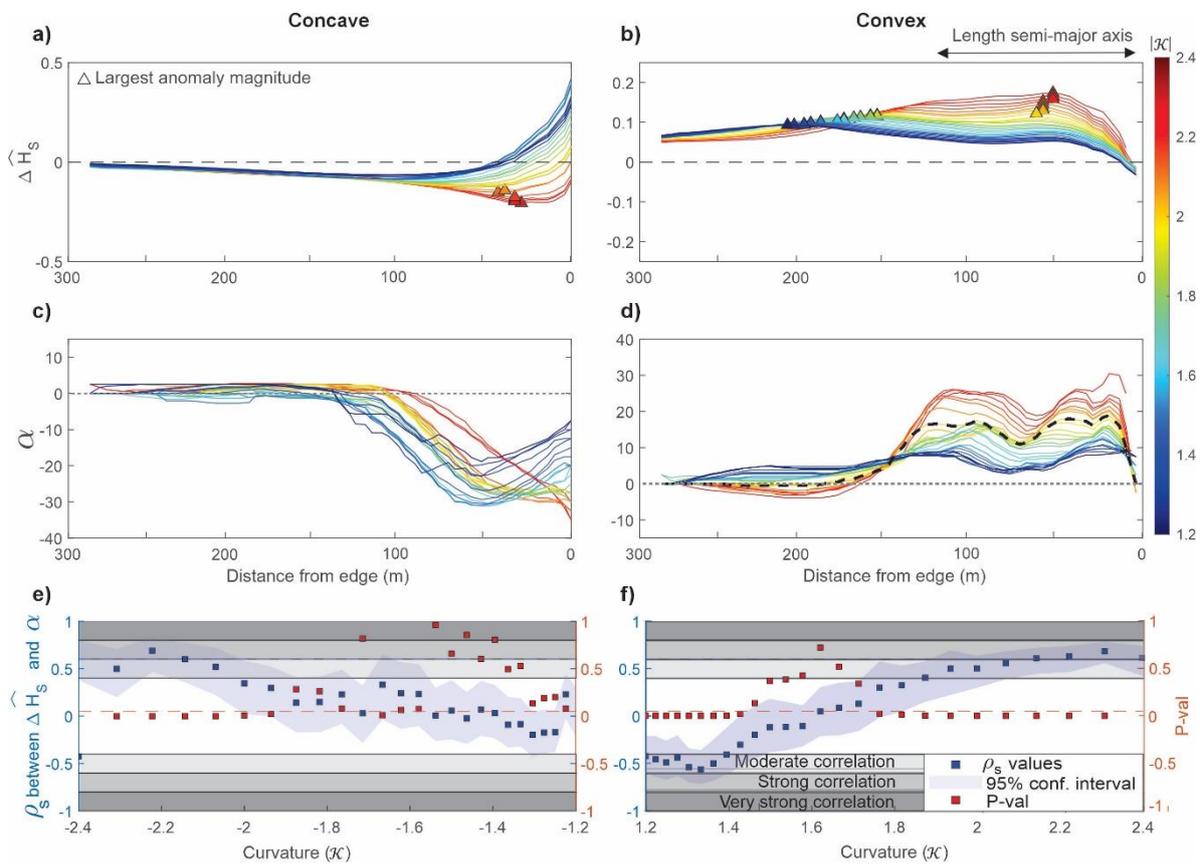


Figure 5: Relationship between directional patterns and significant wave height anomalies of broken waves across concave (left) and convex platforms (right) at the centreline ($y = 150$ m) showing: the significant wave height anomalies (a,b), the cross-shore directional patterns (c,d) and spearman correlation between these two parameters (e,f).

459

460

461

462 For concave platforms, the seaward shift of maximum divergence zones coincided
463 with a seaward shift of the location of the largest negative anomalies (representing a 25%
464 attenuation in \widehat{H}_s , Fig. 5a). As a result, the relationship between $\Delta\widehat{H}_s$ and directional patterns
465 of broken waves remained moderate to strong ($0.4 < \rho_s < 0.6$) for concave edge curvatures
466 exceeding 1.9 (Fig. 5e), indicating that for large degrees of edge curvature, the $\Delta\widehat{H}_s$ observed
467 across the concave platform depended on the directional patterns along the centreline. For
468 convex platforms, the seaward shift of the maximum convergence locations (Fig. 5d)
469 coincided with a seaward shift of the largest positive anomalies for curvatures over 1.8
470 (representing an 18% amplification in \widehat{H}_s , Fig. 5b). However, for curvatures lower than 1.8,
471 the maximum anomalies shifted landward. Thus, the correlations between $\Delta\widehat{H}_s$ and
472 directional patterns across the centreline were moderate to strong ($0.4 < \rho_s < 0.6$) for convex
473 edge curvatures exceeding 1.8, and weak ($\rho_s < 0.4$) for curvatures dropping below 1.8 (Fig.
474 5f). This phenomenon can be explained by analysing the relative influence of each harmonic
475 component on $\Delta\widehat{H}_s$ observed across the platforms (Fig. 6).

476 For convex curvatures exceeding 1.8, the decrease of wave convergence over the
477 outer platform and wave ray divergence over the inner platforms (Fig. 5c) coincided with a
478 reduction of wave height anomalies for all harmonic components over the inner platform (Fig.
479 6). This reduction was particularly important for the fifth subharmonic, $\Delta\widehat{H}_{f_{1/5p}}$, representing
480 5% of the observed amplification of significant wave height at $x = 190$ m against 10% at $x =$
481 130 m for $|\mathcal{K}|=2.4$. In contrast, convex edge curvature below 1.8 inhibited the formation of a
482 divergence zone, ensuring the sustainability of wave ray convergence across the entire
483 platform. Under these conditions, the wave height anomalies within the WW and SW
484 frequency bands were sustained across the entire platform, and anomalies within the IG_L
485 frequency band were amplified over the inner platform ($\Delta\widehat{H}_{f_{1/5p}}$ representing 15% of the
486 observed amplification of normalised significant wave height at $x = 190$ m for $|\mathcal{K}|=1.2$) to
487 become the dominant type of anomaly at this location. Thus, $\Delta\widehat{H}_s$ became predominantly
488 controlled by the behaviour of IG_L as curvature decreased ($1.4 < \mathcal{K} < 1.8$). For very low degrees
489 of curvature ($\mathcal{K} < 1.4$), the amplification of IG_L was of such importance that $\Delta\widehat{H}_s$ were

490 amplified over the inner platform despite the decrease in wave convergence, resulting in a
 491 negative correlation ($-0.6 < \rho_s < -0.4$) between $\Delta\widehat{H}_s$ and directional patterns (Fig. 5c).

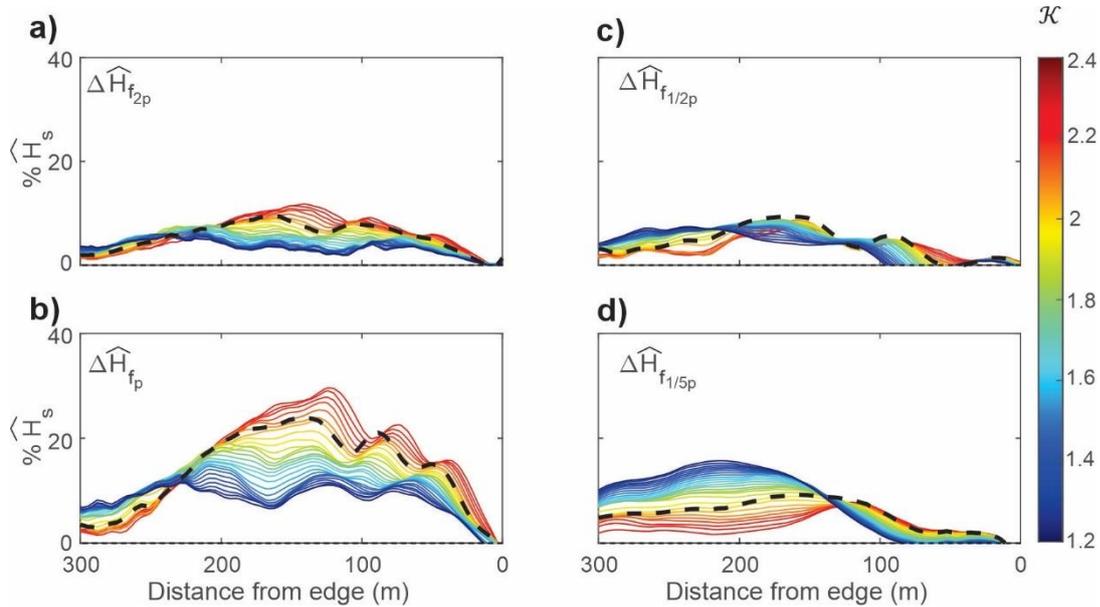


Figure 6: Percentage of significant wave height variations across the centreline ($y = 150$ m) convex platforms associated with anomalies of the second higher harmonics (a), the principal harmonic (b), the second subharmonic (c) and the fifth subharmonic (d) components ($H_s = H_{sStraight} + \Delta H_s$). The dashed line represents a curvature of 1.8, marking the threshold for the formation of a divergence zone over the inner platforms.

492

493

494 3.2 Effects of platform edge geometry on alongshore wave height patterns

495 3.2.1 Non-breaking waves

496 Coherent self-interaction maps were plotted to investigate the impact of platform edge
 497 geometry on alongshore wave height variation over the inner platform for non-breaking
 498 waves (Fig. 7). Maxima in these maps ($\widehat{\psi} \approx 1$) correspond to areas of strongest coherent
 499 interaction for the dominant frequency components within the SW and IG_L frequency bands
 500 (f_p and $f_{1/5p}$). Over the concave platforms, zones of coherent self-interaction for the principal
 501 harmonic (f_p) shifted alongshore from the platform centrelines to become concentrated near
 502 the northern and southern extremities of the platform as the edge curvature increased (Fig.
 503 7a-d). These alongshore variations were predominantly observed between $x = 130$ and 175
 504 m, where divergence along the centreline was the strongest (Fig. 3c). In contrast, coherent

505 self-interaction maps for the fifth subharmonic ($f_{1/5p}$) were more homogenous alongshore
506 (Fig. 7e-h), except near the shoreline, where coherent self-interactions were predominantly
507 observed at the platform centreline. Over the convex platform, coherent self-interactions of
508 the principal harmonic (Fig. 7i-l) were concentrated toward the platform centreline for edge
509 curvatures between $|\mathcal{K}|=1.2$ to 1.6 (Fig. 7k,l), but as edge curvature increased, coherent wave
510 interaction for this harmonic predominantly occurred at the northern and southern
511 extremities of the platforms (Fig. 7i,j). For the fifth subharmonic, coherent self-interactions
512 were focussed near the platform centreline for low edge curvature and spread alongshore
513 toward the shoreline $|\mathcal{K}|=1.2$ (Fig. 7p). As curvature increased, the areas of fifth subharmonic
514 coherent self-interactions near the shoreline split into two peaks on either side of the
515 platform centreline (Fig. 7m,n). This phenomenon was observed at curvatures for which a
516 mild divergence was observed over the inner platform (Fig. 3d).

517 The wave height distribution over the inner sections of concave and convex platforms
518 is shown in Fig. 8 and 9. The spatial distribution of the significant wave height (\widehat{H}_s) presented
519 the strongest similitudes ($R^2 > 0.9$) with the wave height patterns of the principal harmonic
520 (\widehat{H}_{fp}) regardless of the platform geometry and curvature. This indicates a strong control of
521 SW on the patterns of significant wave height variations over the inner platform. In contrast,
522 the correlation between the wave height patterns of the fifth subharmonic and the significant
523 wave height patterns over the inner platforms of concave and convex geometries was weak
524 ($R^2 < 0.4$), indicating that IG_L had little impact on the variations of significant wave height at
525 this location.

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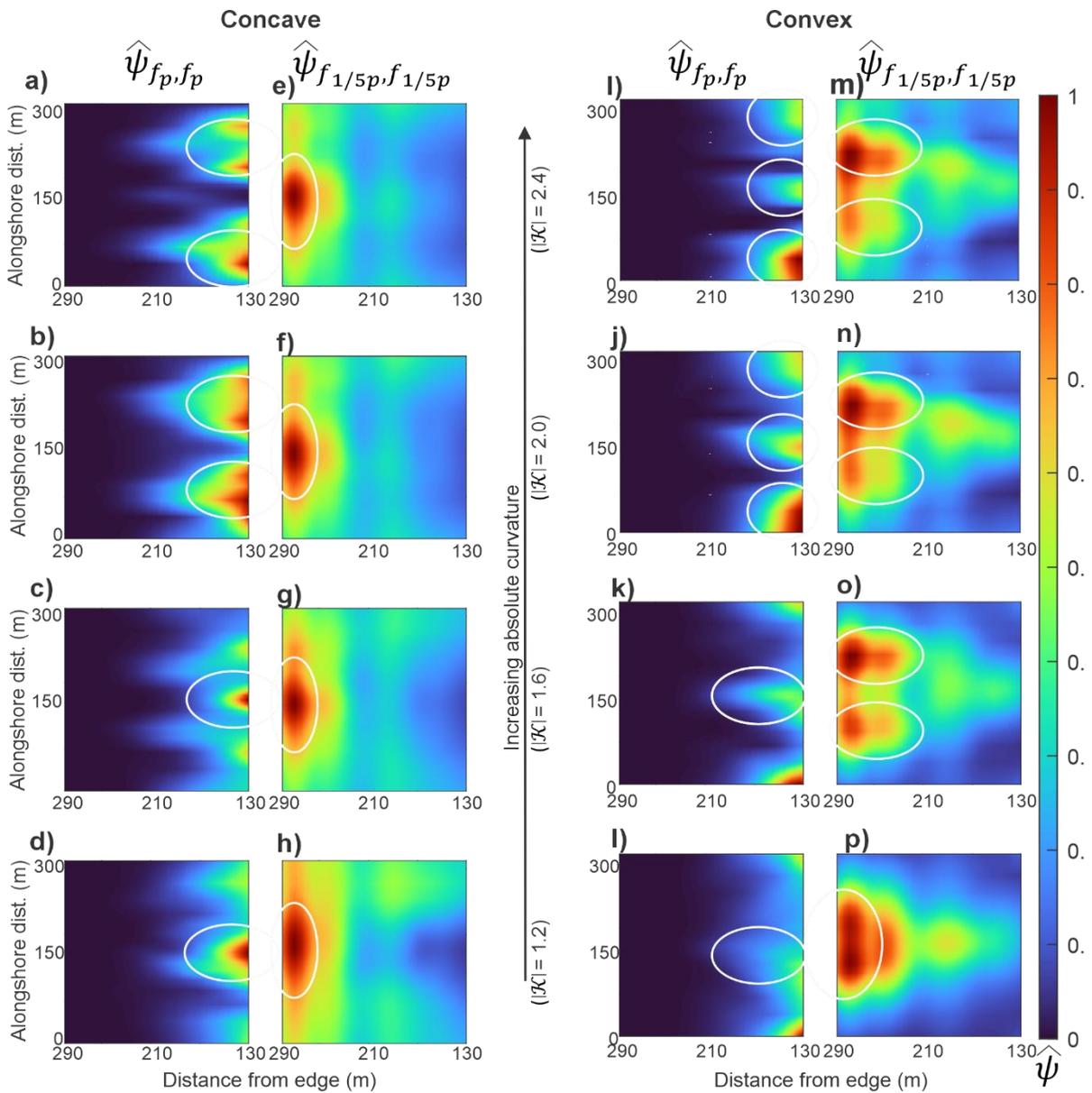


Figure 7: Coherent self-interaction maps defined from the bispectral modal state of self-interacting components for the principal harmonic (f_p) and the fifth subharmonic ($f_{1/5p}$) of non-breaking waves over the inner platform (Fig. 1) at different concave (a-h) and convex (i-p) edge curvatures. The centreline is located at $y = 150$ m. Values of $\hat{\psi}$ of 1 indicate areas of the largest coherent wave interactions. The white ellipses highlight the zones of strong coherent wave interactions.

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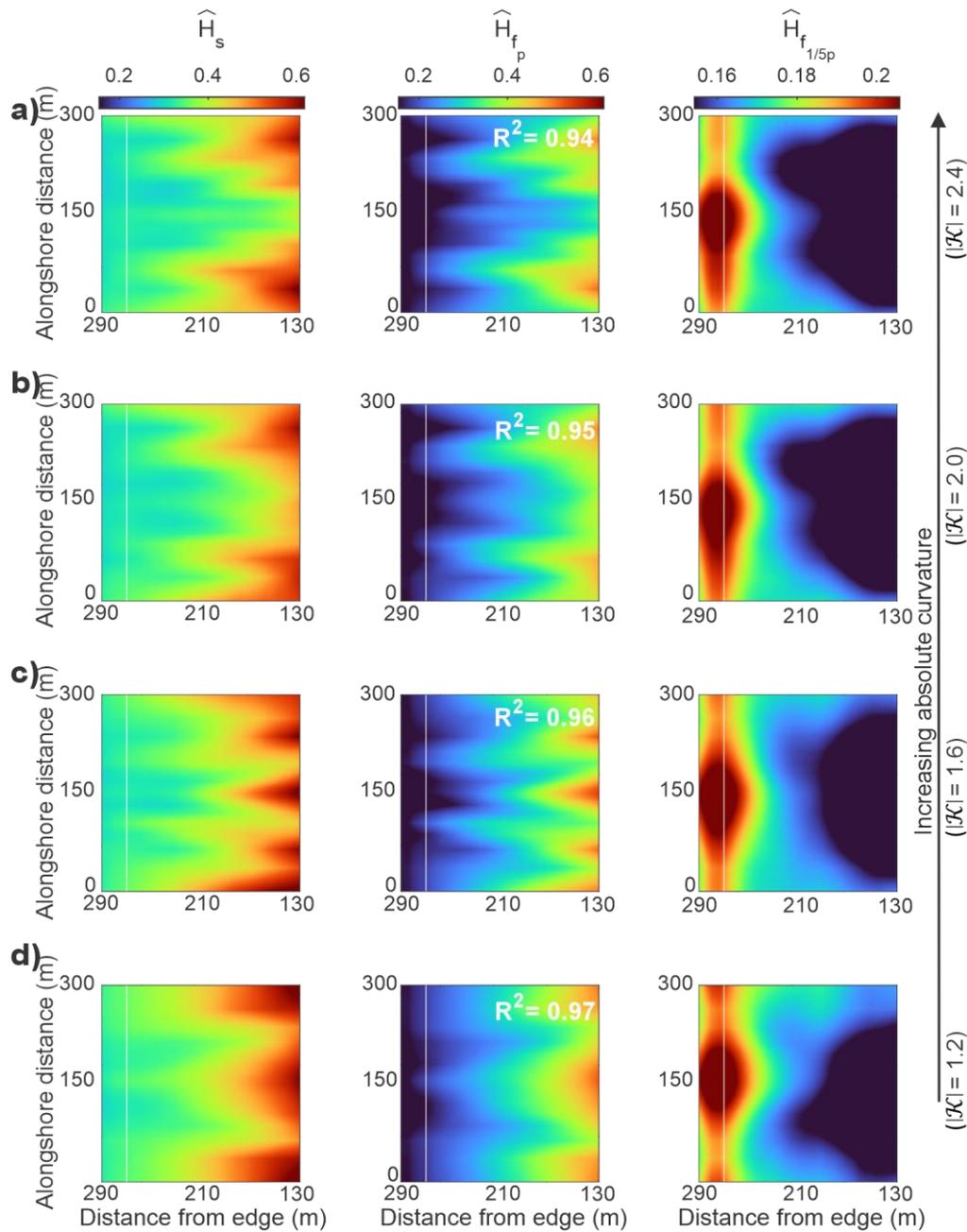


Figure 8: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of non-breaking waves over the inner platform (Fig. 1) for various concave (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline (Fig. 1). The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature (only $R^2 \geq 0.4$ is shown, representing moderate to very strong correlations)

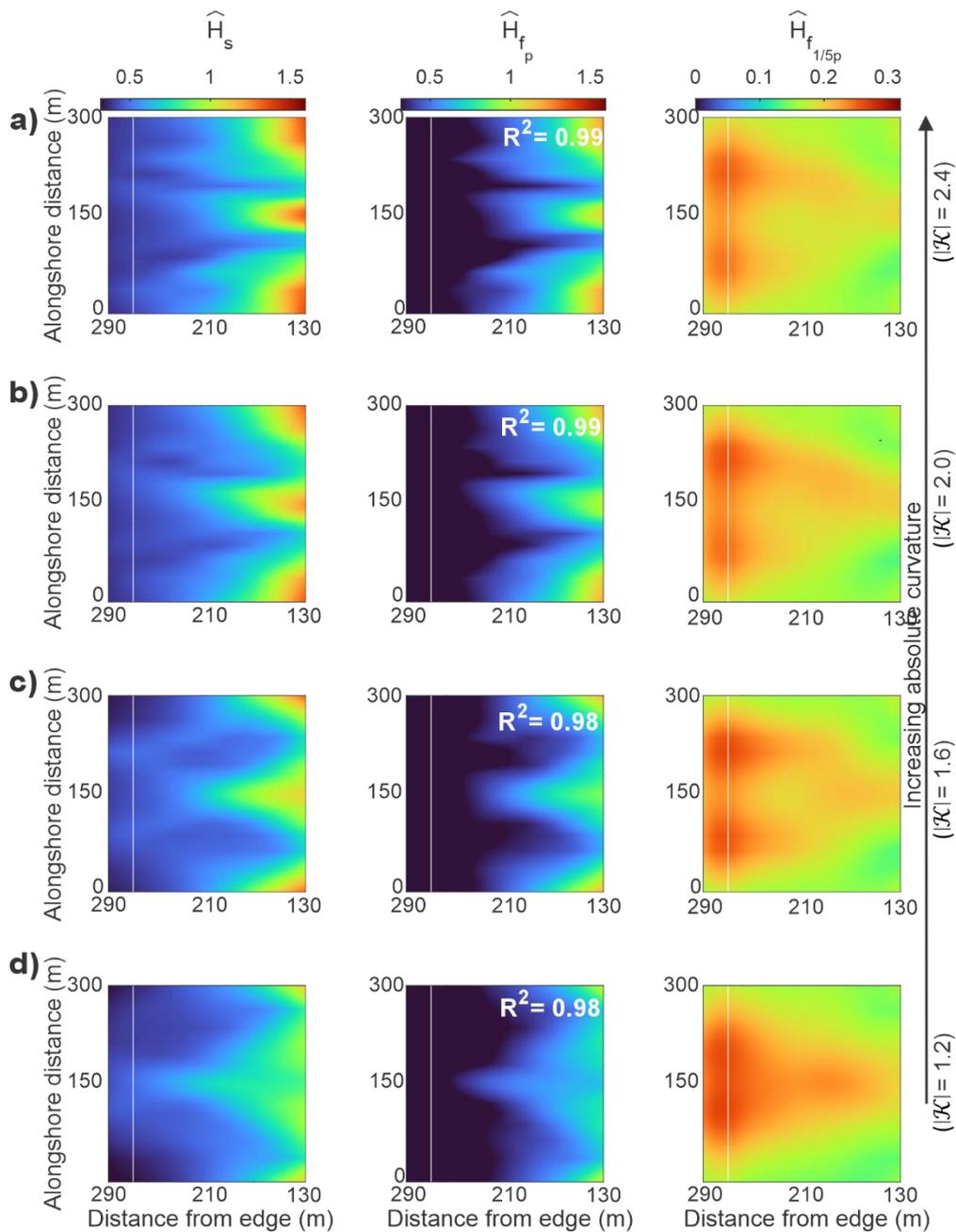


Figure 9: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of non-breaking waves over the inner platform (Fig. 1) for various convex (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline (Fig. 1). The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature (only $R^2 \geq 0.4$ is shown, representing moderate to very strong correlations)

541

542 A strong relationship, $R^2 > 0.8$, was observed between modal coherent self-interaction
 543 patterns and wave height patterns of the principal harmonic and fifth subharmonics over the
 544 inner sections of concave (Fig. 10a) and convex platforms (Fig. 10b). This observation indicates
 545 that the alongshore variations of the principal harmonic (SW) were predominantly controlled
 546 by coherent wave interaction, which in turn drove the alongshore variations in significant
 547 wave height over the inner section of both concave and convex platforms. The resulting
 548 stationary patterns in significant wave height along the shoreline were characterised by a
 549 decrease of significant wave height toward the centreline of concave platforms (Fig. 11a),
 550 which became more pronounced with increasing curvature (maximum alongshore difference
 551 in $\widehat{H}_s=0.05$ at $|\mathcal{K}|=1.2$, increasing to 0.06 at $|\mathcal{K}|=2.4$, Fig. 10a). Over convex platforms,
 552 stationary patterns for normalised significant wave height were characterised by an increase
 553 of significant wave height toward the platform centreline at low degrees of curvature (Fig.
 554 11b), resulting in an alongshore difference in $\widehat{H}_s \approx 0.15$ for $|\mathcal{K}| < 1.8$ near the shoreline. A
 555 progressive amplification of the lobes on either side of the centreline was observed as
 556 curvature increased, resulting in a more homogenous alongshore distribution of significant
 557 wave height for high degrees of curvature ($\widehat{H}_s \approx 0.06$ for $|\mathcal{K}| > 2$, Fig. 11b).

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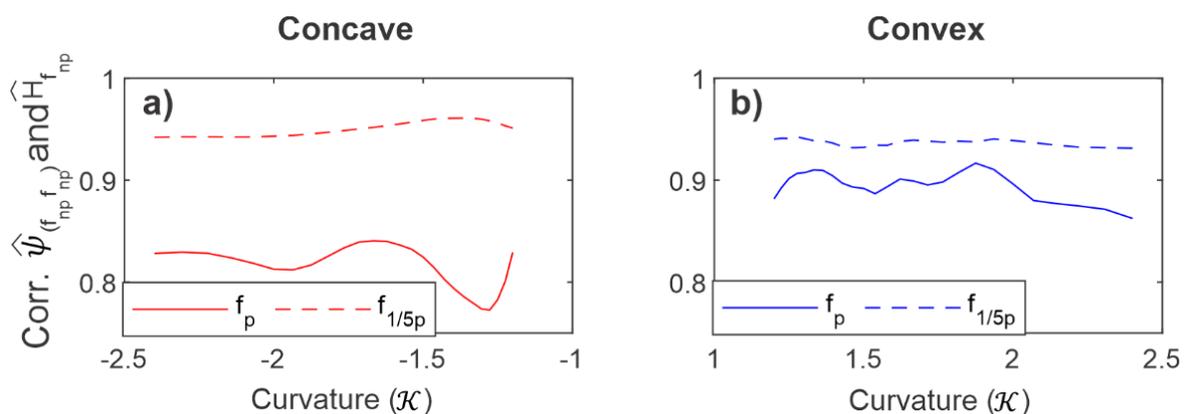


Figure 10: Correlation between interaction maps (Fig. 7) and wave height patterns (Fig. 8, 9) for the principal harmonic (f_p) and the fifth subharmonic ($f_{1/5p}$) of non-breaking waves over the inner platform of concave (a) and convex (b) edges.

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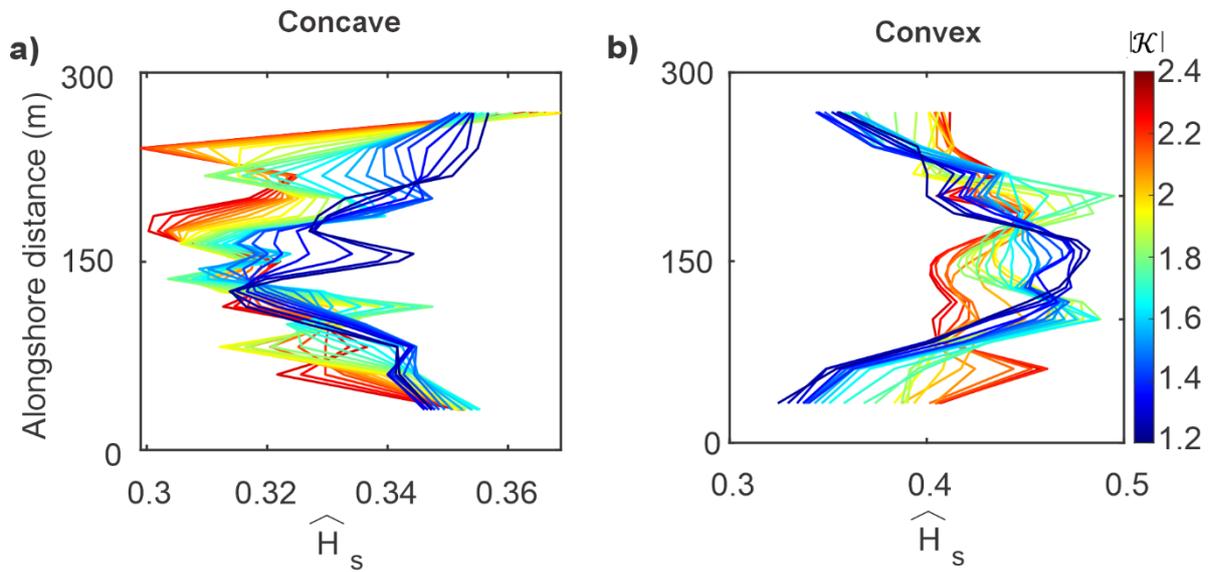


Figure 11: Alongshore variations in significant wave height patterns 20 m from the shoreline for non-breaking waves (transect L, Fig. 1) in relation to concave (a) and convex (d) edge curvatures

562 3.2.2 Broken waves

563 Coherent self-interaction patterns of the principal harmonic and fifth subharmonic of
 564 broken waves displayed alongshore variabilities over both concave and convex platforms (Fig.
 565 12). Over the concave platforms, the coherent self-interaction zone of the principal harmonic
 566 was concentrated toward the centreline for low degrees of curvature ($|\mathcal{K}|=1.2$), spreading
 567 alongshore as the degree of curvature increased (Fig. 12a-d). Zones of coherent self-
 568 interactions for the fifth subharmonic were predominantly observed on the northern and
 569 southern extremities of the platforms and became more distinct as the edge curvature
 570 increased (Fig. 12e-h). Over the convex platforms, coherent self-interactions of the principal
 571 harmonic were the strongest on the northern and southern extremities of the platforms at
 572 $x \approx 190$ m. For the fifth subharmonic (Fig. 12m-p), areas of coherent self-interaction were
 573 concentrated along the platform centrelines for low degrees of curvature ($|\mathcal{K}|=1.2$ and 1.6),
 574 but spread either side of the platform centrelines for high degrees of curvature ($|\mathcal{K}|=2.0$ and
 575 2.4). The differences in coherent self-interaction patterns between low and high degrees of
 576 curvature were characterised by a mild divergence over the inner section of convex platforms
 577 for curvatures greater than 1.8 (Fig. 7d).

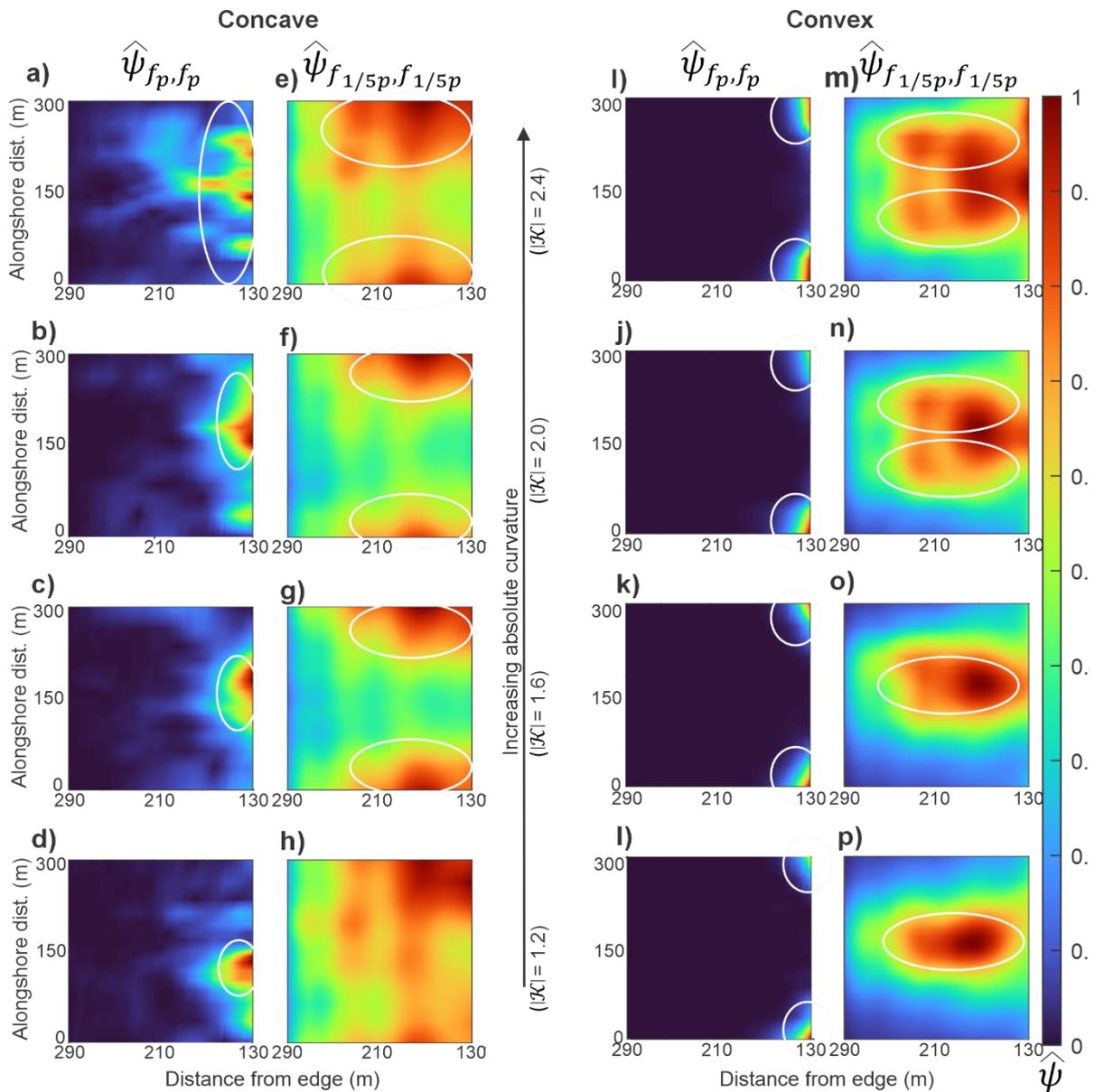


Figure 12: Coherent self-interaction maps defined from the bispectral modal state of self-interacting components for the principal harmonics (f_p) and the fifth subharmonic ($f_{1/5p}$) of broken waves over the inner platform (Fig. 1) at different concave (a-h) and convex (i-p) edge curvatures. The centreline is located at $y = 150$ m. Values of $\hat{\psi}$ of 1 indicate areas of strong coherent wave interactions. The white ellipses highlight the zones of strong coherent wave interactions.

582

583 For broken waves, the influence of IG_L on significant wave height distribution over the
584 inner sections of concave and convex platforms was greater than for non-breaking waves (Fig.
585 13, Fig. 14). Over the inner section of the concave platforms (Fig. 13), the fifth subharmonic
586 had greater wave height than the principal harmonic. Thus, the wave height patterns of the
587 fifth subharmonic had a greater impact on the significant wave height patterns ($0.85 < R^2 <$
588 0.91 for $1.2 < |\mathcal{K}| < 2.4$) than the principal harmonic ($0.74 < R^2 < 0.86$ for $1.2 < |\mathcal{K}| < 2.4$) in this
589 region. The wave height of the principal harmonic and fifth subharmonic decreased from the
590 northern and southern extremities of the platforms to the platform centrelines. The
591 combined effect of these patterns was a net alongshore decrease of significant wave height
592 toward the platform centrelines. Over the inner section of convex platforms (Fig. 14), the
593 principal harmonic displayed the greatest wave height (maximum $\widehat{H}_{fp} \approx 0.5$) on the northern
594 and southern sides of the platform between $x \approx 130-190$ m. The wave height of the fifth
595 subharmonic was relatively smaller, reaching a maximum at the platform centreline
596 (maximum $\widehat{H}_{fp} \approx 0.22-0.27$), regardless of the curvature. As a result, the wave height
597 distribution of the principal harmonic exerted a strong control on the significant wave height
598 pattern over the inner platforms ($0.9 < R^2 < 0.95$) in comparison to the control exerted by the
599 fifth subharmonic ($0.4 < R^2 < 0.58$). However, the wave height of the principal harmonic
600 significantly decreased past $x \approx 190$ m, becoming comparable to the wave height of the fifth
601 subharmonic. Thus, alongshore variations in significant wave height were controlled by
602 alongshore patterns of both principal harmonic and fifth subharmonic for $x \geq 190$ m. For low
603 degrees of edge curvature ($|\mathcal{K}|=1.2$), the maximum wave height of the fifth subharmonic was
604 observed at the platforms' centreline and evolved with increasing curvature to form two
605 maxima on either side of the centreline for large degrees of edge curvature ($|\mathcal{K}|= 2.2$ and
606 2.4). This evolution was clearly observed in the significant wave height pattern between $x =$
607 190 and 300 m, underlining the influence of the fifth subharmonic (IG_L) on the alongshore
608 variation of significant wave height at the shoreline.

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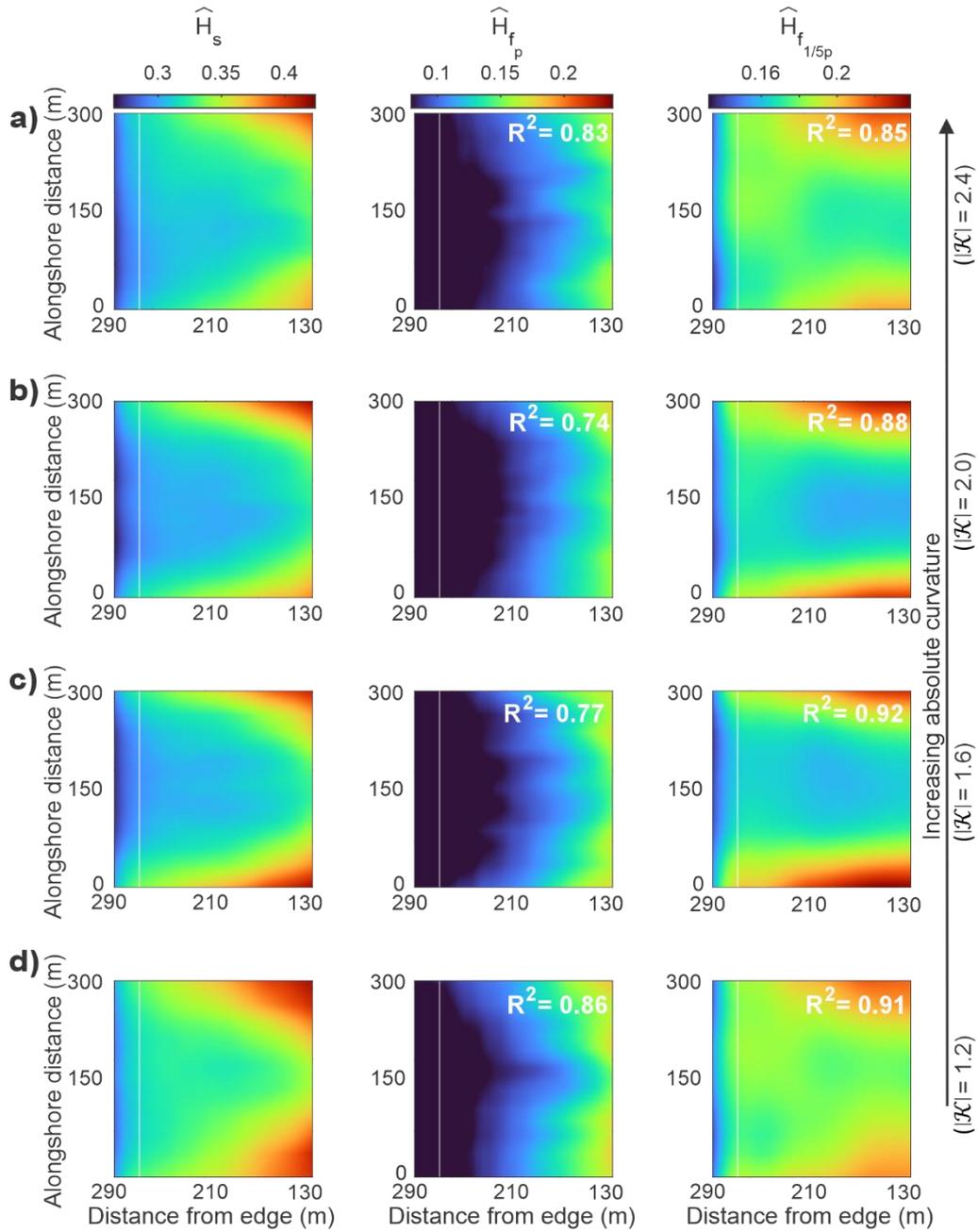


Figure 13: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of broken waves over the inner platform (Fig. 1) for various concave (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline. The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature.

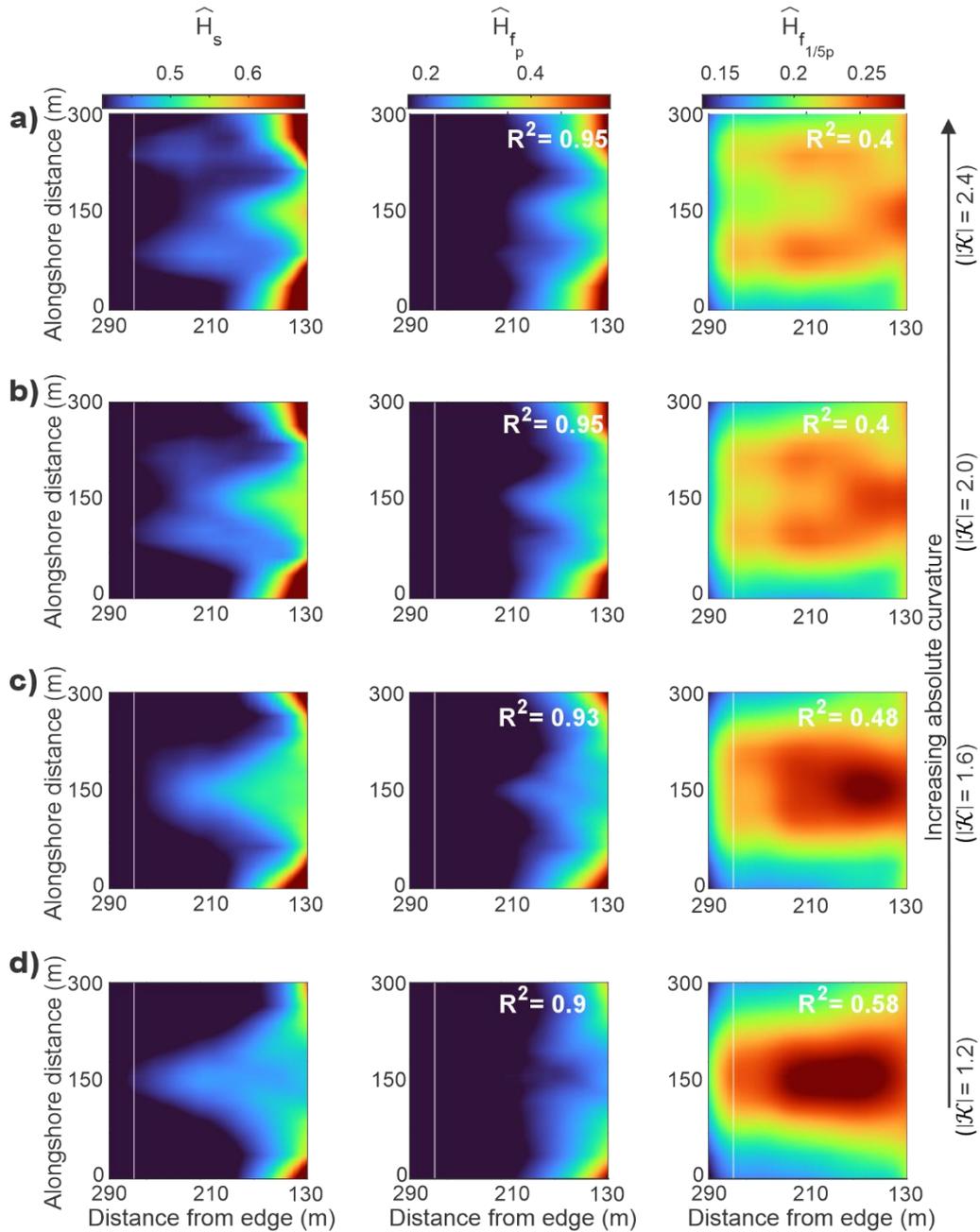


Figure 14: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of broken waves over the inner platform (Fig. 1) for various convex (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline. The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature.

613

614

615 Strong ($R^2 > 0.8$) and moderate to strong ($0.5 < R^2 < 0.7$) relationships were observed
616 over the inner sections of concave (Fig. 15a) and convex platforms (Fig. 15b) between modal
617 coherent self-interaction patterns and wave height patterns for the fifth subharmonics and
618 principal harmonics, respectively. The implication is that coherent wave amplification
619 influenced the longshore patterns of wave height for the principal harmonic and fifth
620 subharmonic over the inner platform, although this process had a smaller impact on the
621 principal harmonic. Thus, coherent wave amplification at IG frequencies was the principal
622 process controlling alongshore variations of significant wave height along the shoreline. The
623 resulting stationary patterns in significant wave height along the shoreline were marked by a
624 decrease of significant wave height toward the centreline of concave platforms, which
625 became more pronounced with increasing curvature (maximum alongshore difference in
626 $\widehat{H}_s=0.04$ at $|\mathcal{K}|=1.2$, increasing to 0.02 at $|\mathcal{K}|=2.4$, Fig. 16a). For convex platforms, an increase
627 of significant wave height toward the platform centreline was observed at low degrees of
628 curvature, resulting in maximum alongshore variations of significant wave height $\widehat{H}_s \approx 0.08$
629 for $|\mathcal{K}|<1.8$. A progressive amplification of the lobes on either side of the centreline
630 generated two wave height maxima for high degrees of curvature, for which maximum
631 alongshore variations of significant wave height $\widehat{H}_s \approx 0.06$ for $|\mathcal{K}|>2$ (Fig. 16b).

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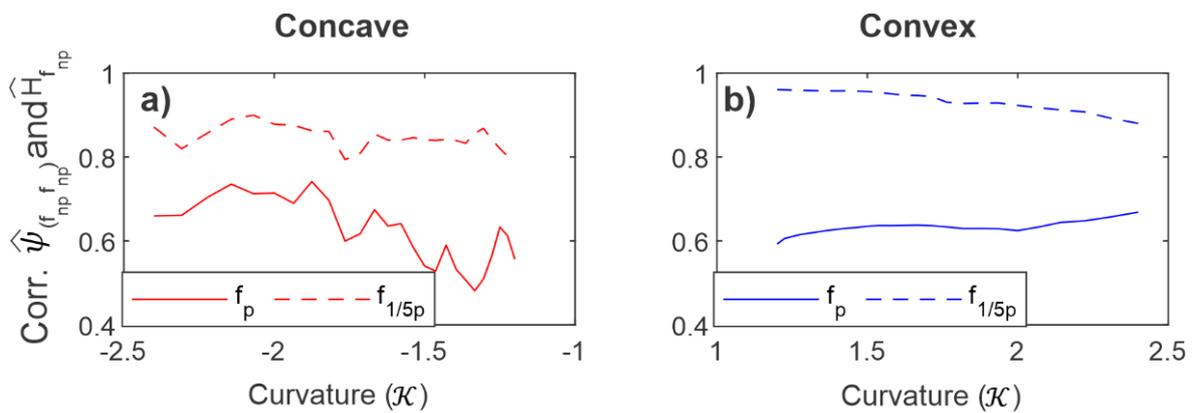


Figure 15: Correlation between interaction maps and wave height patterns for the principal harmonic (f_p) and the fifth subharmonic ($f_{1/5p}$) of broken waves over the inner platform of concave (a) and convex (b) edges.

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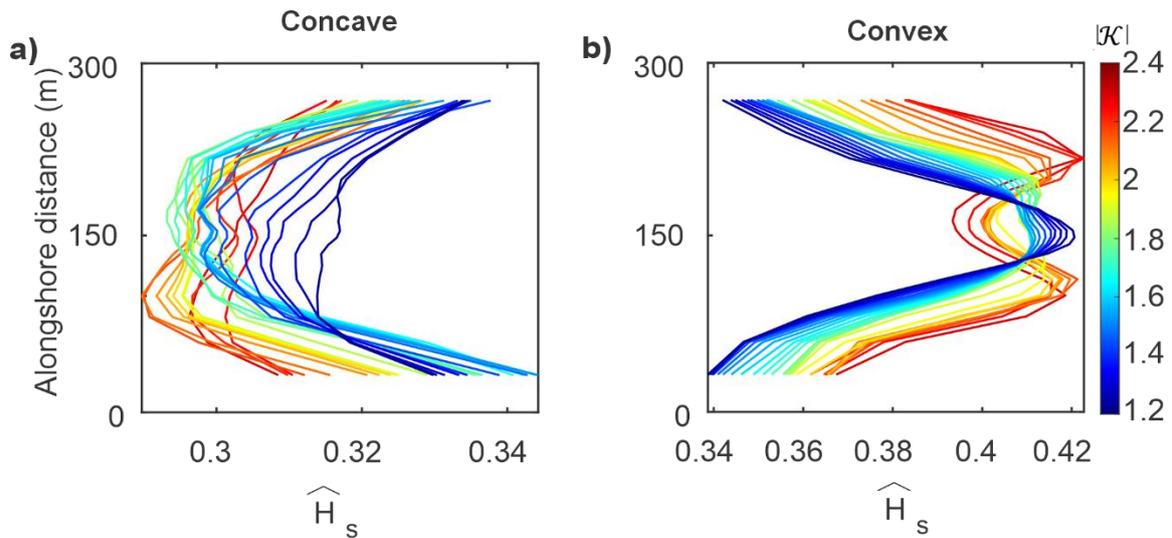


Figure 16: Alongshore variations in normalised significant wave height patterns 20 m from the shoreline (transect L) for broken waves in relation to concave (a) and convex (d) curvature.

636

637 4 Discussion

638 4.1 Impact of platform edge geometry on wave transformation across shore platforms

639 Modelling studies generally investigate the impact of refraction on wave energy
 640 distribution over a fixed curvature (e.g. Berkhoff et al., 1982; Li et al., 2020), whereas in this
 641 work, we varied the degree of edge curvature and analysed its effect on the transformation
 642 of harmonic components. Increasing concave edge curvature decreased the wave energy of
 643 harmonic components in WW, SW, IG_H and IG_L frequency bands (Fig. 2b-e, 4b-e), accounting
 644 for up to 25% reduction in \widehat{H}_s (Fig. 3a, 5a). In contrast, increasing convex edge curvature
 645 amplified both the second and principal harmonics in the WW and SW frequency bands over
 646 the outer platforms, increasing \widehat{H}_s by up to 55% and 18% for non-breaking and broken waves
 647 (Fig. 3b, 5b) while minimizing the amplification of the fifth subharmonic within the IG_L
 648 frequency band over the inner platforms (Fig. 2g-j, 4g-j). Thus, it is clear that morphological
 649 variability in platform edge curvature influences significant wave height on shore platforms,
 650 and this has potential implications for backwear and downwear erosion processes on rock
 651 coasts (e.g. Trenhaile, 1987, Matsumoto et al., 2016a,b).

652 Amplification of higher harmonics associated with wave refraction has previously been
653 associated with wave focusing (Gouin et al., 2017), but the impact of edge curvature on
654 amplification has not been considered. Our modelling results show that increasing convex
655 edge curvature enhances wave focussing over the outer platform (Fig. 3, 5), promoting the
656 generation of higher harmonic from non-linear triadic (sum) interactions (Janssen and
657 Herbers, 2009; Jarry et al., 2011; Lawrence et al., 2022). Shore platform studies have linked
658 the generation and dominance of high-frequency waves over the outer section of near-
659 horizontal platforms to locally generated wind waves (e.g. Ogawa et al., 2011; 2016). Though
660 this process cannot be ruled out, the nearshore wind speed required for locally produced WW
661 energy is substantial (Hasselmann et al., 1973), and the generation of higher harmonics from
662 non-linear triadic interaction caused by wave refraction appears to be a more plausible
663 physical interpretation for high-frequency wave generation on the outer sections of shore
664 platforms.

665 Research has demonstrated that wave amplification in the IG frequency band over near-
666 horizontal platforms is influenced by the ratio of water depth at the cliff toe to the platform
667 width, and relative submergence (Beetham and Kench, 2011; Ogawa et al., 2015). We show
668 that edge curvature exerts an additional morphological control on IG amplification across
669 convex platforms by affecting the balance between focusing intensity from refraction and
670 defocusing effects from wave breaking controlled by convex edges (Fig. 4, 5). Although a
671 decrease in convex edge curvature should theoretically result in a landward shift of the focal
672 point over submerged flats (e.g. Mandlier and Kench, 2012), a seaward shift of the focal point
673 was observed in this study for broken waves (Fig. 5d). This phenomenon can be attributed to
674 the defocussing effects resulting from the enhancement of radiation stress and wave-
675 generated current by wave breaking (Yoon et al., 2004; Choi et al., 2009). A critical curvature
676 was found for which the intensity of wave focusing by wave refraction is not strong enough
677 to overcome the defocusing effects of wave breaking, identified here as $|\mathcal{K}|=1.8$ (Fig. 5d).
678 When the critical curvature is exceeded, wave rays intersect across the platform centreline,
679 in which cases, IG amplification is minimised by wave ray divergence over the inner platform
680 (Fig. 6d). In contrast, for convex edge curvatures lower than the critical curvature, wave rays
681 do not intersect across the platform centreline, sustaining wave convergence across the
682 entire platform. In this case, IG amplification is promoted over the inner platform,

683 representing up to 15% of the increase in significant wave height at this position (Fig. 6d).
684 Thus, the present research identifies convergence as a key mechanism acting on the growth
685 of IG across shore platform, working in conjunction with other processes such as to
686 breakpoint forcing (Poate et al., 2020) and energy transfer from higher frequencies and
687 shoaling (Beetham and Kench, 2011).

688 In a conceptual model, Ogawa et al. (2011) described spatial zones on a shore platform
689 that are dominated by different wave types and pointed out that the spatial characteristics
690 change according to the tidal stage. Here, we show that shifts in dominant wave types across
691 shore platforms are also controlled by convex edge curvature (certainly at high tide). High
692 convex curvature amplifies harmonics within the WW and SW frequency bands over the outer
693 platform and inhibits the amplification of IG over the inner platform (Fig. 4g). Low degrees of
694 convex curvature have the opposite effect (Fig. 4j), resulting in the seaward shift of the zones
695 dominated by WW and SW frequency bands. The influence of refraction patterns generated
696 by convex edge geometries on the collective behaviour of harmonics affected the significant
697 wave height patterns across convex platforms. Baldock et al. (2020) hypothesised that the
698 significant wave height across convex platforms is defined by a specific balance between
699 cross-shore energy loss from dissipation and energy gain from oblique refracted SW, resulting
700 in a correlation between significant wave height anomalies and refraction patterns. The
701 strong relationship ($\rho_s > 0.6$) observed between directional patterns and significant wave
702 height anomalies of non-breaking waves (Fig. 3f) suggest that, for this wave state, energy is
703 effectively gained across the platform from refracted SW. However, for broken waves, the
704 correlation between directional patterns and significant wave height anomalies ($\Delta\widehat{H}_s$)
705 decreased with curvature to become weak ($\rho_s < 0.4$) below the critical curvature threshold
706 (Fig. 5f). These differences are attributed to the influence of IG growth on significant wave
707 height over the inner platform due to the amplification of IG from post-breaking energy
708 transfer from high to low frequency (Poate et al., 2020), and low sustained convergence
709 across the platform (Fig. 5b, 6d). These results underline the importance of considering the
710 refraction patterns of both SW and IG when investigating significant wave height patterns
711 across convex submerged flats.

712

713 4.2 Impact of platform edge geometry on along shore wave transformation

714 Coherent wave amplification is identified here as a crucial process affecting SW and IG_L
715 height distribution in the inner section of near-horizontal platforms. For non-breaking waves,
716 the dominant frequencies within both SW and IG_L frequency bands presented strong
717 correlations between patterns of coherent wave interaction and wave height distribution
718 over concave and convex platforms (Fig. 10). However, this correlation decreased for the
719 principal harmonic in the SW frequency bands in the presence of wave breaking (Fig. 15),
720 perhaps due to the combination of defocussing (Yoon et al., 2004) and dissipation effects
721 (Farrell et al., 2009) associated with wave breaking.

722 The present observations validate the hypothesis of Winter et al. (2017) on the
723 formation of alongshore stationary IG patterns from wave refraction over convex platforms.
724 However, our results suggest that such patterns are generated by coherent wave interaction
725 following the generation of caustic rays in the IG_L frequency band (Fig. 7m-p, 12m-p) rather
726 than alongshore standing waves, as Winter et al. (2017) suggested. In fact, the latter would
727 require interacting IG to propagate alongshore in opposite directions. The directional analysis
728 presented here precludes such a possibility ($\alpha \approx 0^\circ$ near the shoreline, Fig. 3, 5). As the
729 present paper demonstrates the coherent wave interaction plays crucial role in alongshore IG
730 wave patterns, a coherent wave class should be added to the resonant, progressive-
731 dissipative, standing and progressive-growing low-frequency wave classes previously
732 identified over submerged flats (Gawehn et al., 2016).

733 The combined modes of SW and IG coherent wave amplification exerted a crucial
734 control on the stationary patterns of significant wave heights over the inner platform. For
735 non-breaking waves, significant wave height variations over the inner platform are
736 predominantly controlled by coherent wave amplification of SW (Fig. 7, 8). In contrast, for
737 breaking waves, the distribution of significant wave height over the inner platform became
738 controlled by coherent wave amplification occurring within both SW and IG_L frequency
739 components as IG_L became a prominent wave type in this region (Ogawa et al. 2011) (Fig. 13,
740 14). The present results support the conceptual model presented by Krier-Mariani et al.
741 (2022), suggesting that patterns of wave ray intersection on either side of concave edge
742 sections result in stationary SW and IG amplification patterns. The control exerted by the

743 critical curvature on the alongshore distribution of SW and IG over convex platforms can be
744 attributed to the generation of a terminal point (marking the transition from wave ray
745 convergence to divergence) at the platform centreline for convex curvatures exceeding the
746 critical curvature. In such cases, caustic rays are formed on either side of the centreline
747 (Mandlier and Kench, 2012), promoting coherent wave interactions in these regions (Fig. 12a,
748 14a).

749 It follows that the control exerted by platform curvature on coherent wave
750 amplification plays an essential role in the nodal state of significant wave height along the
751 shoreline (Fig. 11, 16). For concave platforms, a node near the centreline and antinodes on
752 the northern and southern extremities of the platform were observed. For convex edge
753 curvature under the critical curvature threshold, an antinode was observed along the
754 platform centreline where waves converged, while for curvature exceeding the critical
755 curvature threshold, two antinodes were observed on either side of the platform centreline.
756 While such patterns could wrongly be associated with edge waves, the present results
757 support the observations of Dalrymple (1975), who first associated nodal and anti-nodal
758 points in alongshore wave height patterns with coherent wave interaction.

759 It has previously been established that by controlling the nodal state of significant wave
760 height along the shoreline of open coasts, coherent wave amplification could lead to the
761 formation of rip currents (Dalrymple, 1975; Wei and Dalrymple, 2017). This mechanism is
762 expected to impact circulation patterns over near-horizontal platforms equally. da Silva et al.
763 (2023) identified alongshore pressure gradient as the dominant driver of circulation patterns
764 in the lee of submerged flats, resulting in two or four-cell circulation systems. A two-cell
765 system is typically characterised by an alongshore diverging flow from the lee of the
766 submerged flat edge to the shoreline, while a four-cell system is characterised by an
767 alongshore diverging flow at the lee of the submerged flat and a converging flow at the
768 shoreline. Considering the present results, it can be hypothesised that stationary wave
769 patterns and the subsequent alongshore pressure gradient generated by coherent wave
770 amplification drive the formation of circulation cells over convex platforms. Theoretically, the
771 formation of two antinodes over the inner platform would result in a four-cell circulation
772 system (Fig. 14a,b, Fig. 16b), while an antinode across the entire platform would result in a

773 two-cell circulation system (Fig. 14c,d, Fig.16b). These differences depend on whether or not
774 the submerged flat geometry allows for the formation of a terminal point. Previous studies
775 established that this condition was predominantly controlled by the distance between the
776 seaward edge of submerged flats and the shoreline (da Silva et al., 2022, 2023; Ranasinghe et
777 al., 2006, 2010), while we show that the degree of edge curvature is equally important (Fig.
778 11, 16).

779 **5 Conclusions**

780 This study employed an exploratory numerical modelling approach to investigate the
781 impact of concave and convex platform edge geometries on the behaviour of wave harmonics
782 and the subsequent wave height distribution patterns over near-horizontal shore platforms.
783 Harmonic analyses show that refraction patterns controlled by concave and convex platform
784 edge curvatures result in wave height variation for the principal and second higher harmonics
785 over the outer platform, and for the subharmonics over the inner platform. Wave divergence
786 across concave edge platforms decreased the height of harmonics within both SW and IG
787 frequency bands, resulting in the attenuation of significant wave height for high degrees of
788 curvature. Over the outer section of convex platforms, increasing curvature intensified wave
789 focusing and amplified the principal and second harmonics within the SW frequency band. A
790 critical curvature value of 1.8 demarcates the formation of a wave ray divergence zone over
791 the inner platform, conditioned by the balance between wave focusing from wave refraction
792 and wave defocusing from wave breaking. Below this threshold, wave convergence amplified
793 IG over the inner platform, but over this threshold, wave divergence reduced the
794 amplification of IG over the inner platform. Through these mechanisms, it is apparent that
795 edge curvature can influence both the relative dominance of SW and IG frequencies, and the
796 pattern of significant wave height transformation across near-horizontal platforms. Using a
797 high-order spectral decomposition method, this study further demonstrated that coherent
798 wave amplification influences stationary IG and SW patterns over the inner platform,
799 affecting the alongshore distribution of significant wave height. We found that platform
800 geometry controls the nodal state of the stationary patterns along the shoreline, possibly
801 resulting in alongshore variation of wave erosive force and the generation of wave-generated
802 currents shaping rock coasts.

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809

810 **Open research**

811 The numerical model input files and post processed data for the model simulations
812 (Using FUNWAVE 3.6) over idealised shore platform geometries are available at *Public release*
813 *planned after review.*

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Numerical investigation of the refractive properties of near-horizontal shore platforms and their effects on harmonic and stationary wave patterns

R. Krier-Mariani¹, W. Stephenson¹, S. Wakes², M. Dickson³

¹ School of Geography, University of Otago, New Zealand

² Department of Mathematics and Statistics, University of Otago, New Zealand

³ School of Environment, University of Auckland, New Zealand

Corresponding author: Raphael, Krier-Mariani (raphael.m.krier@gmail.com; raphael.krier@postgrad.otago.ac.nz)

28 **Key Points:**

- 29
- Through refraction, concave and convex near-horizontal shore platforms can separate the frequency components of the wavefield.
- 30
- Refraction patterns controlled by platform edge convexity affect the dominance of wind, swell and infragravity waves across platforms.
- 31
- Coherent amplification from the intersection of refracted infragravity waves controls the nodal state of alongshore stationary waves.
- 32
- 33
- 34

35 **Abstract**

36 Near-horizontal shore platforms display highly irregular plan shapes, but little is
37 known about the way in which these irregularities influence the significant wave height (\widehat{H}_s)
38 on the platforms and the frequency components of the nearshore wavefield. We use a non-
39 linear Boussinesq wave model to conduct harmonic and bispectral mode decomposition
40 analyses, studying the control of concave and convex platform edges over wind (WW: 0.125
41 - 0.33 Hz), swell (SW: 0.05 - 0.125 Hz) and infragravity (IG: 0.008 - 0.05 Hz) frequencies. For
42 breaking and non-breaking waves, increasing the platform edge concavity intensified wave
43 divergence and subsequent attenuation of SW and IG across the outer platforms, reducing \widehat{H}_s
44 by up to 25%. Increasing the platform edge convexity intensified focusing and amplification
45 of SW and WW over the outer platforms, increasing \widehat{H}_s by up to 18% and 55% for breaking
46 and non-breaking waves. In the presence of breaking, IG amplification depended on the
47 generation of wave divergence across the inner platform, a condition determined by a critical
48 convex curvature threshold ($|\mathcal{K}|=1.8$) balancing wave focusing from refraction and
49 defocusing from breaking. We find that convex curvature can determine the relative
50 dominance of WW, SW and IG across platforms. Alongshore, coherent wave interactions
51 governed IG stationary patterns defined by a node near the platform centreline and two
52 antinodes on either side of concave edges. A node was generated at the platform centreline,
53 and two antinodes were observed on either side of the convex edges for $|\mathcal{K}|>1.8$, with the
54 opposite pattern observed for $|\mathcal{K}|<1.8$.

55

56 Plain Language Summary

57 Near-horizontal shore platforms fronting coastal cliffs act as wave energy buffers,
58 regulating wave-induced erosion in rock coast environments. Genuine research endeavours
59 have permitted establishing the link between near-horizontal platform morphology and wave
60 transformation across-shore. However, the effects of alongshore variations in near-horizontal
61 platform morphology on the properties of nearshore wavefields remain sparsely
62 documented. As ocean waves share akin refractive properties to light rays, it can be assumed
63 that, similarly to optical lenses, shore platforms can separate waves according to their
64 frequency depending on their geometry. Subsequently, the convergence and divergence of
65 refracted wave trains of similar phases and frequencies could affect the properties of the
66 nearshore wavefield. The present research investigates this phenomenon over concave and
67 convex edge platforms and its impact on the nearshore wavefield characteristics. Our results
68 show that wave refraction over near-horizontal platforms with concave and convex edges
69 affects the relative dominance of short, medium and long-period waves across shore and
70 results in alongshore stationary wave patterns near the shoreline with nodal states varying in
71 relation to platform edge geometry. Such patterns likely result in alongshore variations in
72 wave erosion and the generation of wave-generated currents shaping rock coasts in the
73 platform.

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78 1 Introduction

79 Near-horizontal shore platforms, defined by a low gradient ($\tan\beta < 0.0175$) and a
80 steep seaward edge, are prevailing coastal landforms in rock coast environments (Sunamura,
81 1992; Trenhaile, 1999). These landforms have an essential role in wave transformation
82 processes, regulating wave erosive forces at the shoreline (Stephenson and Kirk, 2000;
83 Matsumoto et al., 2016a,b). Thus, an accurate description of the geomorphic control exerted
84 by shore platforms on nearshore wave transformation patterns is necessary for improving
85 rock coast geomorphological models.

86 Studies have investigated the control of near-horizontal shore platform morphology on
87 the cross-shore evolution of the wavefield (e.g. Beetham and Kench, 2011; Marshall and
88 Stephenson, 2011; Ogawa et al., 2016). Wave breaking induced by the sharp depth transition
89 at the seaward edge of a platform results in the dissipation of incident swell waves (SW: 0.05
90 Hz < f < 0.125 Hz) and the generation of low-frequency infragravity waves (IG: f < 0.05 Hz)
91 over the platform (Poate et al., 2020). Across the platform surface, IG gradually amplify due
92 to shoaling and energy is transferred from high to lower frequencies, becoming the dominant
93 frequency component over the inner platform (Beetham and Kench, 2011; Marshall and
94 Stephenson, 2011; Ogawa et al., 2011). Wind waves (WW: 0.125 < f < 0.33 Hz) can propagate
95 onto platforms from offshore and, in some cases, be locally generated over the outer platform
96 to become the dominant frequency component in this area (Marshall and Stephenson, 2011;
97 Ogawa et al., 2011). These observations were summarised in the conceptual model of Ogawa
98 et al. (2011), indicating that it is common for the outer platform, platform centre, and inner
99 platform to be dominated by WW, SW and IG, respectively. Ogawa et al. (2011) suggested
100 that these zones shift across-shore with tidal elevation and showed that the relative
101 submergence of shore platforms (depth at the seaward edge/incident wave height) is a critical
102 factor controlling the relative dominance of SW and IG. Collectively, understanding the
103 behaviour of each frequency band of the wavefield helps to depict the variation of significant
104 wave height (H_s) across platforms affecting erosion of the platform and cliff (Trenhaile,
105 2000). However, the impact of shore platform morphology on two-dimensional wave
106 transformation processes and effect on the frequency bands composing the wavefield have
107 been overlooked.

108 Few field studies have considered the impact of the platform morphology of near-
109 horizontal platforms on two-dimensional wave transformation patterns (Krier-Mariani et al.
110 2022, 2023). Krier-Mariani et al. (2023) showed that directional patterns controlled by
111 irregularities in platform morphology generated localised areas of wave ray convergence and
112 divergence as well as alongshore variations in standing IG patterns, influencing the wave
113 energy distribution over the platform surfaces. Based on these observations, Krier-Mariani et
114 al. (2023) introduced a conceptual model in which concave and convex platform edge
115 geometries would control wave ray convergence and divergence patterns over the platform
116 surface, subsequently affecting the IG energy levels and SW decay rates. However, the

117 influence of platform edge geometry on two-dimensional wave patterns could not be clearly
118 isolated from field observations.

119 In the absence of detailed field studies on the effects of platform edge geometry on
120 wave transformation characteristics, the literature on morphologically analogous submerged
121 flat structures is useful. Depending on their geometry, submerged flats can separate the
122 frequency components of the wavefield, refracting and reorganising the wave crests of
123 incident waves according to their frequency (Jarry et al., 2011; Griffiths and Porter, 2012; Li
124 et al., 2020). This phenomenon can result in complex refraction patterns specific to each
125 frequency component of the wavefield, leading to the generation of caustic rays (clusters of
126 caustic points generated by wave ray intersection) over submerged surfaces (e.g. Mandlier
127 and Kench, 2012). Patterns of wave ray convergence and divergence induced by refraction
128 over submerged flat structures significantly impact the wavefield characteristics. Wave ray
129 convergence results in a localised enhancement of wave height (e.g. Ito and Tanimoto, 1972;
130 Berkhoff et al., 1982), skewness and kurtosis (Janssen and Herbers, 2009; Jarry et al., 2011;
131 Lawrence et al., 2022) while wave ray divergence has the opposite effects.

132 Although relatively few studies have considered the impact of submerged flat
133 geometries on the cross-shore evolution of harmonic and subharmonic components of the
134 wavefield, harmonic components amplification has been observed in areas of wave
135 convergence (e.g. Lynett and Liu, 2004; Gouin et al., 2017). According to Li et al. (2020), this
136 phenomenon could be attributed to the non-linear effects of convergence on wave height
137 amplification. As the geometry of submerged flats influences the cross-shore pattern of wave
138 convergence (intensity and location) of each harmonic, it likely also influences the cross-shore
139 patterns of wave harmonics amplification, intrinsically affecting the dominance of different
140 wave frequencies across platforms. This hypothesis as yet to be verified.

141 It has proven difficult to establish causality between patterns of wave ray intersection,
142 increased nonlinearity and alongshore wave height amplification for random wavefields,
143 notably due to the limitation of wave ray tracking techniques to evaluate complex wave ray
144 crossing patterns in dense constellations of caustics (Ito and Tanimoto, 1972). Another way
145 of approaching this problem involves considering the impact of coherent wave interaction
146 patterns on the amplification of dominant frequency components of the wavefield. Coherent

147 wave interaction refers to the non-linear process occurring at the intersection of waves with
148 similar frequency, waveform and phase. It has been identified as a fundamental non-linear
149 wave amplification process in optics (e.g. Young, 1802), quantum mechanics (e.g., Weiland
150 and Wihelmsson, 1977; Falk, 1979; Inouye et al., 1999; Kozuma et al., 1999) and geoscience
151 (e.g. Harid et al., 2014). There have been few investigations of this process in coastal wave
152 studies, but Dalrymple (1975) demonstrated that this process could result in the formation of
153 alongshore stationary wave patterns in random wavefields and the subsequent formation of
154 nearshore currents. More recently, Tamura et al. (2020) showed that, similar to light
155 refraction through a prism, ocean wave refraction over a submarine canyon could separate
156 waves of a random wavefield according to their frequency and phase, favouring coherent
157 wave interactions. Based on this theoretical grounding, it is hypothesised that by controlling
158 the refraction patterns of individual frequency components of the wavefield, submerged flat
159 (e.g. shore platforms) geometry affects coherent wave amplification over submerged flat
160 surfaces, leading to the generation of alongshore stationary wave patterns for SW and IG.

161 The impact of shore platform geometry on the behaviour of wave harmonics and
162 stationary wave patterns remains to be evaluated in detail on near-horizontal platform
163 surfaces. However, such a task was proven to be difficult during field observations due to the
164 variable nature of nearshore wavefields and the morphological complexity of shore platforms
165 (e.g. Krier-Mariani et al. 2022, 2023). Therefore, this study adopts a numerical modelling
166 approach to address the question: How do mesoscale variations in platform edge geometry
167 affect the behaviour of wave harmonics and the subsequent wave height distribution across
168 and along platform surfaces?

169 **2 Method**

170 **2.1 Model set up**

171 The phase-resolving Boussinesq wave model FUNWAVE_TVD V3.6 (Shi et al., 2012)
172 was used to investigate two-dimensional wave transformation over shore platforms. This
173 model treats wave transformation in the time domain and provides a robust representation
174 of non-linear processes, refraction and diffraction while retaining information on the wave

175 phase (Sheremet et al., 2011; Buckley et al., 2015, Buckley et al., 2018; Thomas and
176 Dwarakish, 2015).

177

178 **2.1.1 Domain**

179 Idealised three-dimensional near-horizontal platform morphologies were incorporated
180 into a 1274 m (x-axis) to 300 m (y-axis) domain (Fig. 1a). A 0.35 m deep, 250 m wide shallow
181 planar surface was included at the landward extremity to absorb wave energy and minimise
182 resonance. The platforms were defined by a constant gradient of 0.35 degrees, a width of 300
183 m (at the centreline, $y = 150$ m) and a 3m high seaward cliff of 45 degrees. The nearshore
184 bathymetry profile was composed of a 480 m subtidal ramp (at the centreline) with a gradient
185 of 0.35 degrees followed by an 8 m deep and 635 m wide flat.

186 Planform geometry was represented using three generic edge geometries defined as
187 straight, concave and convex. The degree of curvature of the concave ($\mathcal{K} < 0$) and convex
188 ($\mathcal{K} > 0$) edge geometries was derived from the parametric ellipse equation. The semi-major
189 axis (a , along the x-axis) was kept constant (120 m) to avoid modifying the cross-shore profile
190 along the centreline, and various degrees of edge curvature were obtained from 2 m
191 increments along the semi-minor axis (b , along the y-axis) between 50 to 100 m, resulting in
192 26 cases with edge curvatures ($|\mathcal{K}| = |a/b|$) ranging from 1.2 to 2.4 (Fig. 1b-f). The
193 bathymetry was smooth to reduce noise generated by sharp edges and interpolated to a 2 m
194 grid adopted to ensure model stability following a series of sensitivity analyses, providing a
195 realistic representation of model resolution used in previous research in nearshore areas (e.g.
196 Su et al., 2021).

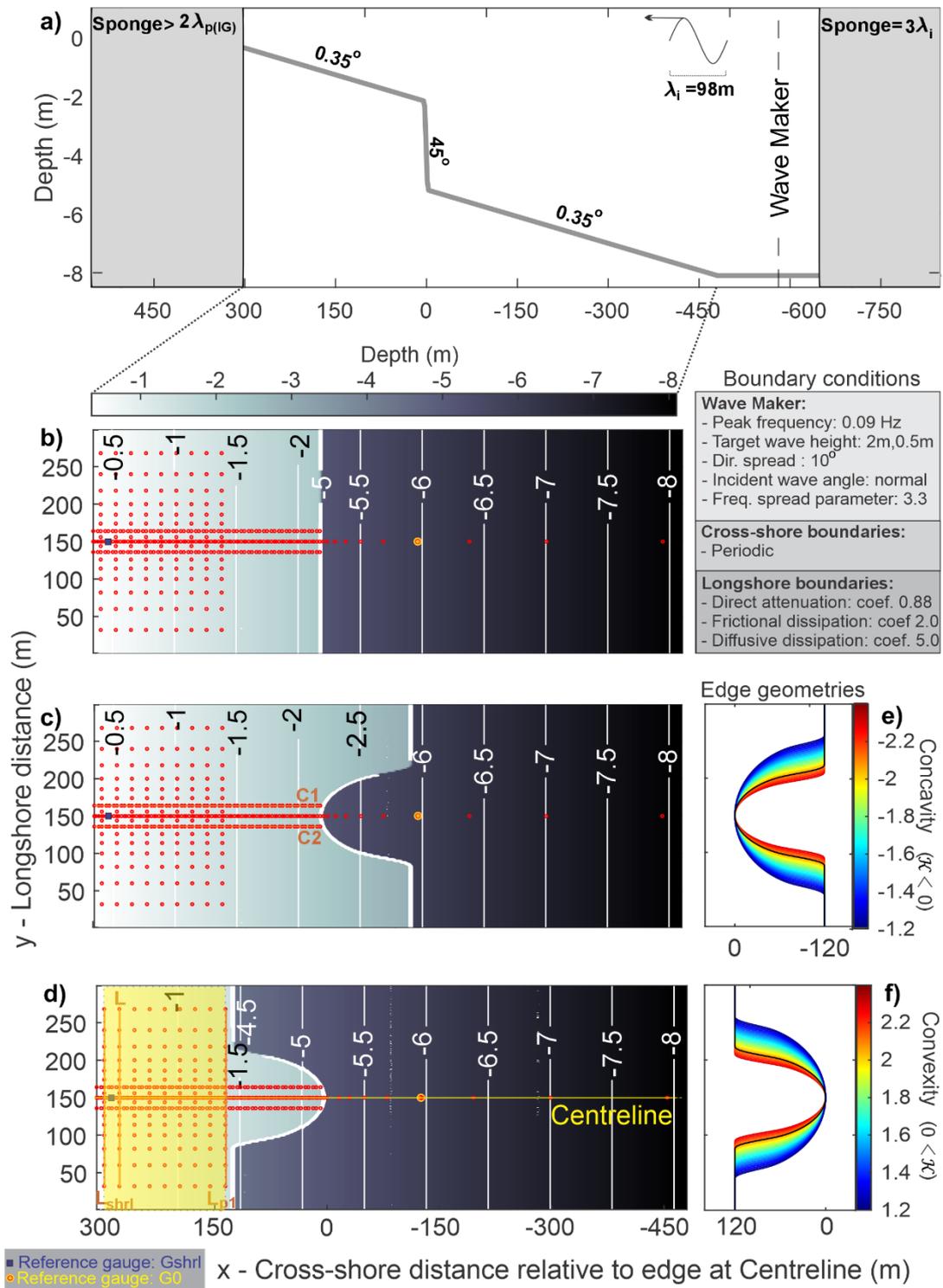


Figure 1: Boussinesq wave model configuration showing the bathymetry profile along the centreline ($y = 150\text{m}$) (a), the model domain for the straight, concave and convex platforms (b-d), and the range of platform edge curvatures considered (e,f). Specifications of the boundary conditions are annotated in the figure. The red dots mark the location of the virtual gauges used for analysis. The yellow shaded area (between L_p and L_{srl}) represents the inner platform section considered for alongshore analysis.

197 **2.1.2 Wave conditions**

198 The model was forced by irregular waves with a directional spread of 10 degrees. An
199 internal wavemaker (Wei et al., 1999) was located on the deep flat at the bottom of the
200 subtidal ramp, five wavelengths (λ_i) away from the platform edge to avoid distortion of the
201 initial wave crests. Irregular waves were generated using a JONSWAP wave spectra
202 (Hasselmann et al., 1973) with a fixed peak enhancement factor of 3.3, a peak frequency (f_p)
203 of 0.09 Hz and direction of 0° (shore-normal) to simplify the visualisation of the refraction
204 effects induced by different platform edge geometries.

205 Two sets of simulations were generated to investigate the transformation of: (1) waves
206 propagating across the platform surface without breaking ($H_s= 0.5$ m), as such waves can
207 release large amounts of erosive energy when they break against cliffs (Thompson et al.,
208 2019; Thompson et al., 2022); and (2) wave breaking at the seaward edge ($H_s= 2$ m) decaying
209 across the platform, which are typically used to define variation of wave erosive force across
210 platforms in geomorphological models (e.g. Trenhaile, 2000; Matsumoto et al., 2016a,b).
211 These two sets of simulations combined with the range of platform concave and convex
212 curvatures resulted in 106 simulations (including straight edge reference cases). The default
213 breaking index of FUNWAVE-TVD ($\gamma_b=0.80$) was used to represent wave breaking, providing
214 a close representation of the breaking conditions for steep submerged slopes (Blenkinsopp
215 and Chaplin, 2008). The effects of bottom friction were not considered (i.e. the frictional
216 dissipation coefficient was set to $C_d= 0.002$, representing a smooth surface).

217

218 **2.1.3 Boundary conditions**

219 The domain boundary conditions were defined to minimise reflection. Periodic
220 boundaries (Chen et al., 2003) were applied to the northern and southern extremities of the
221 domain, allowing waves to propagate out of the domain. Following Shi et al. (2016), sponge
222 layers employing a direct damping coefficient as well as dissipation by friction and diffusion
223 were used to reduce noise and dampen wave energy at the eastern and western sides of the
224 domain (Fig. 1a). The width of the sponge layer at the shallow western side of the domain
225 was chosen to correspond to twice the peak wavelength of the IG at this location (estimated

226 during trial runs using the virtual gauge G_{shrl} at the shoreline, Fig. 1), to avoid reflection and
227 the subsequent generation of standing IG waves.

228

229 **2.1.4 Model Validation**

230 Due to the lack of two-dimensional field measurements in similar near-horizontal
231 shore platform settings, no direct validation of our model simulations was carried out.
232 However, a number of studies validated FUNWAVE-TVD against field observations over coral
233 reefs, proving the model's ability to represent wave transformation over smooth submerged
234 flats with sharp seaward edges (e.g. Mendonca et al., 2008; Su et al., 2015; Zhang et al., 2019).
235 As the present study explores wave processes such as refraction and non-linear energy
236 transfer fairly well represented by the model (Griffiths and Porter, 2012; Su et al., 2015) and
237 does not investigate subsequent processes such as wave-driven circulation and sediment
238 transport, it is deemed unnecessary to validate the model with experimental data at this stage
239 (similar inference were made in da Silva et al., 2023).

240 **2.2 Measurements and analysis**

241 To determine the impact of planform geometries on wave transformation across the
242 platforms, the spectral evolutions of waves propagating across concave and convex platforms
243 (affected by two-dimensional transformation processes) were compared to the spectral
244 evolution of waves propagating across the straight-edge platform (only affected by on-
245 dimensional transformation processes). This approach permitted the identification of spectral
246 anomalies representing the energy variations for specific harmonics induced by refraction.
247 Positive and negative anomalies indicate harmonic amplifications and attenuation,
248 respectively. Combined, the harmonic anomalies result in anomalies of significant wave
249 height across platforms ($\Delta\widehat{H}_s$). Following Baldock et al. (2020), the cross-shore patterns of
250 $\Delta\widehat{H}_s$ were then compared to the directional patterns along the platform centrelines to identify
251 the effects of refraction patterns controlled by platform edge geometry on significant wave
252 height distribution across platforms.

253 In the alongshore, the effects of coherent wave interaction induced by refraction over
254 concave and concave platforms on the generation of stationary wave patterns were

255 considered. For this purpose, the bispectrum (Hasselmann et al., 1963) provides a convenient
256 representation of the wavefield as it holds information on the wave phase, frequency and
257 power necessary to detect phase coupling. The bispectrum, defined from the third moment
258 of the free surface elevation time series, also represents a measure of skewness, which
259 increases in areas of wave ray intersection (Janssen and Herbers, 2009; Jarry et al., 2011;
260 Lawrence et al., 2022). Following Kim and Powers (1979), who investigated the impact of
261 coherent interactions of random electromagnetic waves on plasma density fluctuation using
262 bispectral properties, the frequency, phase and power information yielded by the bispectrum
263 were used to identify patterns of coherent wave interactions over the inner platforms. A
264 modal decomposition method based on bispectral properties (Appendix 2), the Bispectral
265 Mode Decomposition or BMD (Schmidt, 2020), was employed to identify the modal state of
266 coherent structures for self-interacting harmonic components within the SW and IG
267 frequency bands. The areas of coherent wave interactions were then compared to the wave
268 height distribution of SW and IG over platforms of various geometries to identify patterns of
269 coherent wave amplification.

270

271 **2.2.1 Wave measurements**

272 Wave records were obtained from virtual gauges recording surface elevation (η) as
273 well as u and v velocity components at 2 Hz (Fig. 1b-d). In the cross-shore direction, the gauge
274 spacing along the centreline increased seaward from the platform edge (increment based on
275 geometric series starting with a spacing of 4 m with an increment factor of 1.5). On the
276 platforms, the gauge spacing was irregular but not exceeding 6 m along the centreline,
277 transects C_1 and C_2 . The distance between the gauges composing the alongshore transects
278 (between L_0 and L_{shrl}) increased on either side of the centreline from 6 to 30 m (with an
279 increment factor of 1.25). Statistical analyses of the wavefield properties were based on an
280 observation window of 2048 seconds, starting 230 seconds after the start of the simulations,
281 marking the time at which SW reached the landward extremity of the domain and IG were
282 generated.

283

284 2.2.2 Definition of wave height

285 The significant wave height (H_s) was defined from the spectra moment (e.g. Thornton
286 and Guza, 1983):

$$H_s = 4 \sqrt{\int_{f_{min}}^{f_{max}} S(f) \cdot df} \quad (1)$$

287

288 The wave spectra estimates $S(f)$ were generated using the Welch (1967) method with
289 segment lengths of 512 samples, 50% overlap and a Hanning window resulting in 20 Degrees
290 of Freedom (Priestley, 1981). To provide a more detailed representation of the wavefield, the
291 gravity and infragravity waves were further divided into two frequency bands, encapsulating
292 the dominant harmonics observed within the WW, SW, and IG (high and low) frequency
293 ranges across the domain (Table 1). The wave height associated with each of these frequency
294 bands was determined using:

295

$$H_{f_{np}} = 4 \sqrt{\int_{f_{low}}^{f_{high}} S(f) \cdot df} \quad (2)$$

296 where n denotes the rank of the harmonic, f_{low} and f_{high} represents the lower and higher
297 frequencies of the power spectral density peak associated with this harmonic, Table 1. The
298 reference incident wave height (H_0) was defined from measurements taken at the gauge G_0
299 located at the top of the subtidal ramp (Fig. 1) and was used to normalise the wave height on
300 the platform surface ($\widehat{H}_s(x) = H_s(x)/H_0$, $\widehat{H}_{f_{np}}(s) = H_{f_{np}}(x)/H_0$). For simplicity,
301 normalised wave heights are hereafter referred to as wave heights.

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309 **Table 1** Frequency band analysis parameters

Conventional Frequency class	Frequency subclass	Corresponding Harmonic	Frequency (f_{np})	Frequency range ($f_{low} - f_{high}$)
Gravity waves	Wind waves (WW)	Second harmonic	f_{2p}	0.15 – 0.20 Hz
	Swell waves (SW)	Principal harmonic	f_p	0.06 – 0.12 Hz
Infragravity waves	Infragravity High (IG _H)	Second subharmonic	$f_{1/2p}$	0.04 – 0.05 Hz
	Infragravity Low* (IG _L)	Fifth subharmonic	$f_{1/5p}$	0.008 – 0.03 Hz

310 *Note that the typical cutoff frequency for the lower portion of the IG frequency band is 0.005 Hz (e.g. Pequignet et al., 2014;
 311 Gawehn et al., 2016). However, the chosen cutoff frequency of 0.008 Hz is more appropriate to describe the low IG in the
 312 simulated wavefield as it corresponds to a trough in the power spectra estimate across the entire domain, which provides a
 313 better physical representation of the low IG.

314

315 2.2.3 Definition of peak direction

316 The angle α between the peak direction of waves propagating on either side of the
 317 centreline (along the cross-shore transects C1 and C2, Fig. 1) was used to investigate the
 318 evolution of wave convergence and divergence along the platform centrelines. The peak
 319 direction of waves over the platform was estimated from the directional wave spectra
 320 $S(f, \theta) = S(f)G(\theta|f)$ calculated from the free surface elevation (η) and velocity
 321 components (u and v) time series by applying the Extension of the Maximum Entropy
 322 Principle (EMEP) method (Hashimoto et al., 1994). To this effect, segments of 512 samples
 323 were used to estimate the frequency spectra ($S(f)$) and 200 iterations to define the
 324 approximation of the spreading function ($G(\theta|f)$) resulting in 76 frequency bins and
 325 directional bins of 5°.

326

327 2.2.4 Identification of coherent wave interaction patterns

328 The BMD was applied to the free surface elevation time series recorded by the two-
 329 dimensional virtual gauge array between L_p and L_{shrl} , marking the boundaries of the spatial
 330 domain ξ (Fig. 1). The welch periodograms employed in the BMD were computed using
 331 segments of 512 samples, 50% overlap and a Hanning window resulting in 20 Degrees of
 332 Freedom. Patterns of coherent wave interactions were identified from coherent self-

333 interaction maps ($\psi_{k,k}$) which are defined by the product of cross-frequency fields $\phi_{k \circ k}$ and
 334 the bispectral modes ϕ_{k+k} obtained from the BMD:

$$\psi_{k,k}(\xi, f_k, f_k) = |\phi_{k \circ k} \circ \phi_{k+k}| \quad (3)$$

335
 336 where the frequency k considered were f_p and $f_{1/5p}$, representing the dominant harmonics
 337 in the SW and IG frequency bands. The cross-frequency fields $\phi_{k \circ k}$ are maps of phase
 338 alignment for these frequencies, while bispectral modes ϕ_{k+k} represent the amplitude of
 339 oscillations of the sea surface at frequency $2k$. Conventionally, the largest values of the
 340 normalised coherent self-interaction maps $\widehat{\psi}_{k,k}$ indicate areas where phase coupling has the
 341 strongest effect on wave amplitude for the sum frequency $2k$. The interaction maps for
 342 straight wave crests with parallel wave rays are expected to be homogeneous alongshore. In
 343 contrast, for cases where wave crests are bent and wave rays intersect, interaction maps will
 344 be non-homogenous alongshore and display maxima in areas of wave ray intersection. In the
 345 presence of coherent wave amplification, maxima in coherent self-interaction maps
 346 correspond to areas of wave height amplification at frequency k .

347

348 **3 Results**

349 **3.1 Impact of planform geometry on across-shore wave transformation**

350 **3.1.1 Non-breaking waves ($H_0 = 0.5$ m)**

351 The spatial evolution of the spectral properties of non-breaking waves propagating
 352 across the domain was examined for the three types of platform geometries (Fig. 2), for which
 353 the power spectra density was concentrated around four distinctive frequency components
 354 (Fig. 2a): the second and the principal harmonics (f_{2p} and f_p) within the WW and SW
 355 frequency bands; and the second and fifth subharmonics ($f_{1/2p}$ and $f_{1/5p}$) within the IG_H and
 356 IG_L frequency bands.

357 The spectral anomalies observed over the concave platforms indicated an attenuation
 358 of the principal harmonic (Fig. 2b-e). This phenomenon intensified with increasing degrees of
 359 curvature (with minimum spectral anomalies at peak frequency reducing from -0.09 m² Hz⁻¹

360 at $|\mathcal{K}|= 1.2$ to $-0.41\text{m}^2 \text{ Hz}^{-1}$ at $|\mathcal{K}|= 2.4$). In contrast, an amplification of the second and
361 principal harmonics was observed across convex platforms (Fig. 2g-j), intensifying with
362 increasing edge curvature (with maximum spectral anomalies at peak frequency increasing
363 from $0.52 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 1.2$ to $1.37 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 2.4$).

364 The variation of spectral characteristics of each harmonic over the concave and convex
365 platforms can be expressed in terms of mean wave height anomalies ($\overline{\Delta H_{f_{np}}}$). The most
366 significant impacts of platform curvature on mean wave height anomalies were observed
367 within the WW and SW frequency bands. The mean wave height anomalies associated with
368 the second and the principal harmonics displayed a very strong linear dependency ($R^2 > 0.9$)
369 to the degree of platform edge curvature (Fig. 2f,k). The increase of curvature from $|\mathcal{K}|= 1.2$
370 to 2.4 promoted the attenuation of harmonics within the WW and SW frequency bands across
371 concave platforms and the amplification of these waves across convex platforms. The
372 attenuation of the second and principal harmonics across concave platforms of high curvature
373 $|\mathcal{K}|= 2.4$ corresponded to 9% and 15% of H_0 . Across convex platforms of high curvature, $|\mathcal{K}|=$
374 2.4, the amplification of the second and principal harmonics reached up to 11% and 29% of
375 H_0 . The mean wave height anomalies for the subharmonic in the IG_H and IG_L frequency bands
376 were negligible for nonbreaking waves.

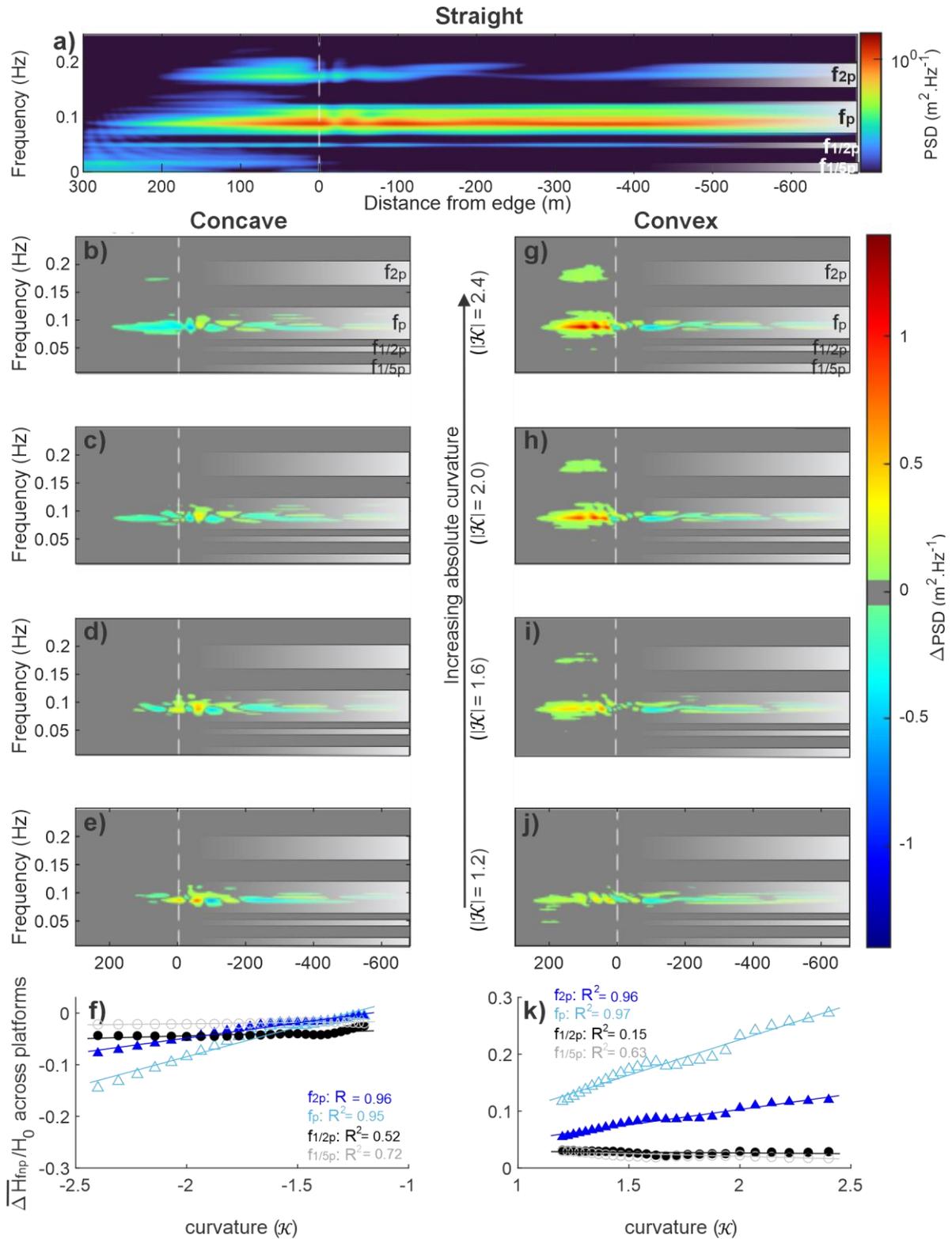


Figure 2: Impact of platform edge curvature on the harmonic components of the wavefield for non-breaking waves showing: the spectral anomalies in relation to the straight edge platform (a) for concave platform geometries (b-e) and convex platform geometries (g-j); and the impact of curvature on the mean wave height of each harmonic across the concave (f) and convex (k) platforms.

377

378

379 $\Delta\widehat{H}_s$ across the platform can be assumed to be impacted by refraction patterns
380 controlled by platform geometry. To explore this process, directional patterns and $\Delta\widehat{H}_s$ across
381 the centrelines of each platform were compared for various edge curvatures. The cross-shore
382 patterns of $\Delta\widehat{H}_s$ presented in Fig. 3a,b were modulated by the platform edge curvature, with
383 magnitude increasing with curvature for both types of platform geometries. As a result, for
384 high degrees of curvature ($|\mathcal{K}| = 2.4$), the negative anomalies across the concave platforms
385 indicated a maximum of 25% attenuation in significant wave height (Fig. 3a), while the
386 positive anomalies across the convex platforms (Fig. 3b) indicated a 55% amplification of
387 significant wave height. The location of the largest $\Delta\widehat{H}_s$ shifted across platforms in relation
388 to curvature. For concave platforms, the largest negative $\Delta\widehat{H}_s$ over the outer platform shifted
389 landward with decreasing curvature from $|\mathcal{K}| = 2.4$ to 1.9. Similarly, the largest positive $\Delta\widehat{H}_s$
390 across the convex platforms shifted landward, reaching the inner platform for $|\mathcal{K}| < 1.6$. For
391 low degrees of concave curvatures ($|\mathcal{K}| < 1.9$), corresponding to curvatures for which
392 amplification of wave energy seaward of the platform edge was observed (Fig. 2b-e), wave
393 transformation patterns across the platform centreline were affected by the preconditioning
394 of incident waves occurring off the platform edge. Therefore, the description of the following
395 results focuses on concave edge curvatures, $|\mathcal{K}| > 1.9$.

396 Similarly to the cross-shore evolution of $\Delta\widehat{H}_s$, the peak magnitude of wave ray
397 divergence observed across concave platforms and convergence across the convex platforms
398 decreased and shifted landward from the mid-platform ($x \approx 150$ m) to the outer platform
399 with decreasing curvature (Fig. 3c,d). A Spearman rank correlation (Fig. 3e,f) revealed that the
400 dependency of cross-shore $\Delta\widehat{H}_s$ on the directional patterns observed over the concave
401 platforms was only relevant (moderate to strong, $\rho_s > 0.4$) for platform edge curvatures
402 exceeding 1.9. In contrast, a strong relationship ($\rho_s > 0.6$) as observed between wave height
403 anomaly and directional patterns over convex platforms for the majority of platform edge
404 curvatures, indicating that $\Delta\widehat{H}_s$ across the convex platforms were predominantly controlled
405 by the wave convergence and divergence across the centreline.

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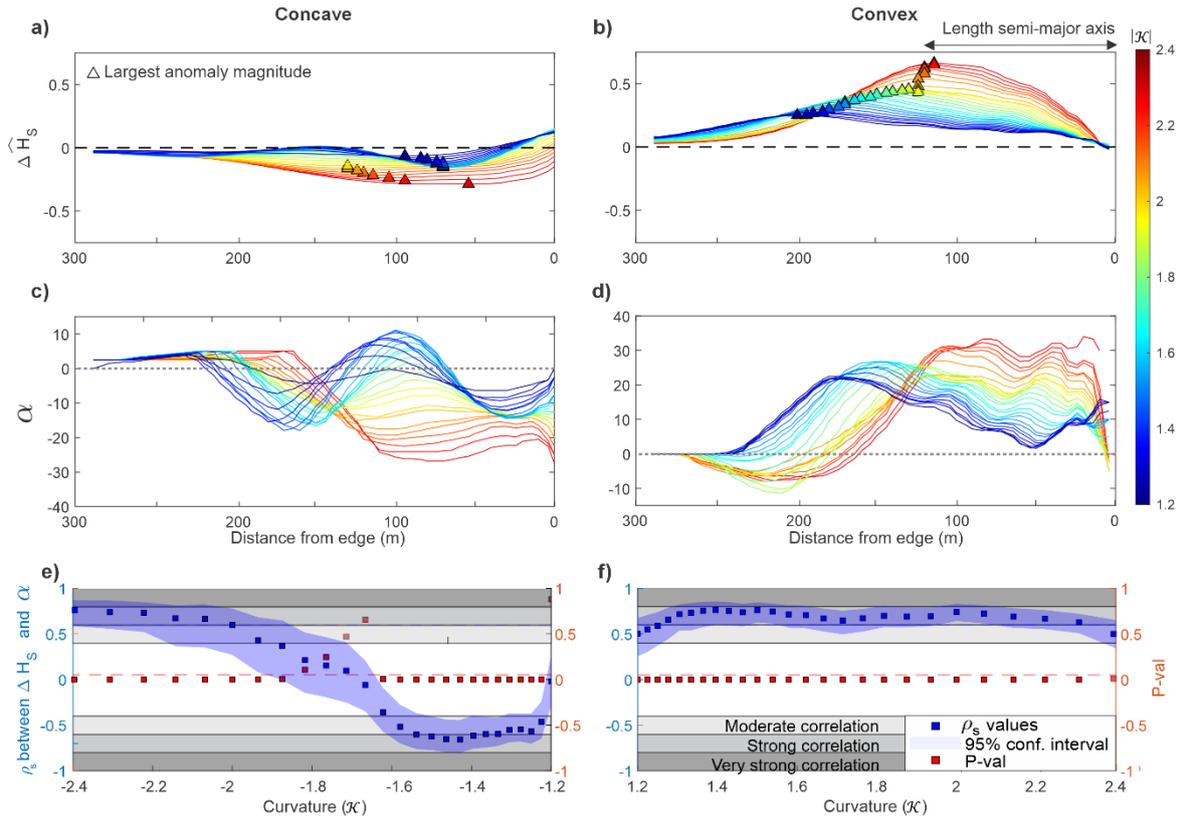


Figure 3: Relationship between directional patterns and significant wave height anomalies of non-breaking waves across concave (left) and convex platforms (right) at the centreline ($y = 150$ m) for different degrees of curvature showing: the significant wave height anomalies (a,b) and the cross-shore directional patterns (c,d). The impact of directional pattern on wave anomaly pattern was assessed using a spearman correlation between the two parameters (e,f).

407

408 3.1.2 Broken waves ($H_0 = 2.0$ m)

409 The spectral evolution of broken waves across concave platforms displayed a complex
 410 pattern of spectral anomalies (Fig. 4b-e), with a clear amplification of the principal harmonic
 411 corresponding to SW over the outer platform for degrees of curvature below 1.9 (at f_p ,
 412 positive anomalies reached $1.5 \text{ m}^2 \text{ Hz}^{-1}$ at $\kappa = 2.0$ and 2.4 , 0 to 150 m from the edge) and a
 413 clear attenuation of this harmonic for degrees of curvature exceeding 1.9 (at f_p , negative
 414 spectral anomaly reached $-1.2 \text{ m}^2 \text{ Hz}^{-1}$ at $\kappa = 1.2$ and 1.6 , 0 to 150 m from the edge). These
 415 differences were related to the amplification of the principal harmonic for edge curvatures
 416 below 1.9, displaying anomalies reaching up to $\sim 12 \text{ m}^2 \text{ Hz}^{-1}$ in the vicinity of the concave edge
 417 sections (0 to -120 m from the edge) before reaching the platform surface (Fig. 4d,e). Over
 418 convex platforms, the principal harmonic presented the largest amplification (Fig. 4g-j), which

419 intensified over the outer platform with increasing curvature (positive anomaly at f_p reached
420 $2.1 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 1.2$ and $4.1 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 2.4$, 0 to 150 m from edge). In contrast, the
421 amplification of subharmonics within the IG_H and IG_L frequency bands toward the shoreline
422 observed along the platform centreline was stronger for low convex edge curvatures than for
423 high convex edge curvatures (positive anomaly at $f_{1/5p}$ reached $0.7 \text{ m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 1.2$ and 0.5
424 $\text{m}^2 \text{ Hz}^{-1}$ at $\mathcal{K}= 2.4$, 150 to 300 m from edge).

425 Relationships between edge curvature and mean wave height anomalies across both
426 platform types were observed (Fig. 4f,k). For anomalies in the WW and SW frequencies, the
427 mean wave height anomalies of the second and the principal harmonics presented a strong
428 linear dependence on the degree of edge curvature of concave and convex edges ($R^2 > 0.90$).
429 In the IG_H frequency band, mean wave height anomalies associated with the second
430 subharmonic were linearly dependent on the curvature across concave platforms ($R^2 = 0.67$).
431 The mean wave height anomalies associated with the fifth subharmonic in the IG_L frequency
432 band decreased linearly ($R^2 = 0.94$) with curvature over the convex platforms.

433 Variations in edge curvature affected the relative importance of WW, SW, IG_H and IG_L
434 anomalies across the platforms. For concave platforms, the increase of concave edge
435 curvature promoted attenuation of all frequency bands, but particularly for WW and SW. For
436 the harmonic components within the WW and SW frequency bands, the mean wave height
437 attenuation across concave platforms was negligible for low curvature ($\overline{\Delta H_{f_{2p}}}$ and $\overline{\Delta H_{f_p}}$ and
438 representing less than 1% of H_0 at $|\mathcal{K}|=1.2$) but intensified for high degrees of curvature
439 ($\overline{\Delta H_{f_{2p}}}$ and $\overline{\Delta H_{f_p}}$ representing less than 3% and 6% of H_0 at $|\mathcal{K}|=2.4$). Across the convex
440 platforms of low curvature ($1.2 < |\mathcal{K}| < 1.75$), the largest amplification of mean wave height
441 was observed for the principal harmonic ($\overline{\Delta H_{f_p}}$ representing 4% to 7% of H_0), followed by the
442 fifth subharmonic ($\overline{\Delta H_{f_{1/5p}}}$ representing 3% to 3.5% of H_0). The amplification of the fifth
443 subharmonic became less important with increasing curvature, while the mean wave height
444 of the second harmonic was amplified. For convex curvatures exceeding 1.75, the principal
445 harmonic displayed the largest amplification ($\overline{\Delta H_{f_p}}$ representing 7% to 10% of H_0), followed
446 by the second harmonic ($\overline{\Delta H_{f_{2p}}}$ representing 3% to 4% of H_0). Thus, the reduction of convex
447 edge curvature promoted the amplification of IG_L , while the increase of convex edge
448 curvature promoted the amplification of WW and SW.

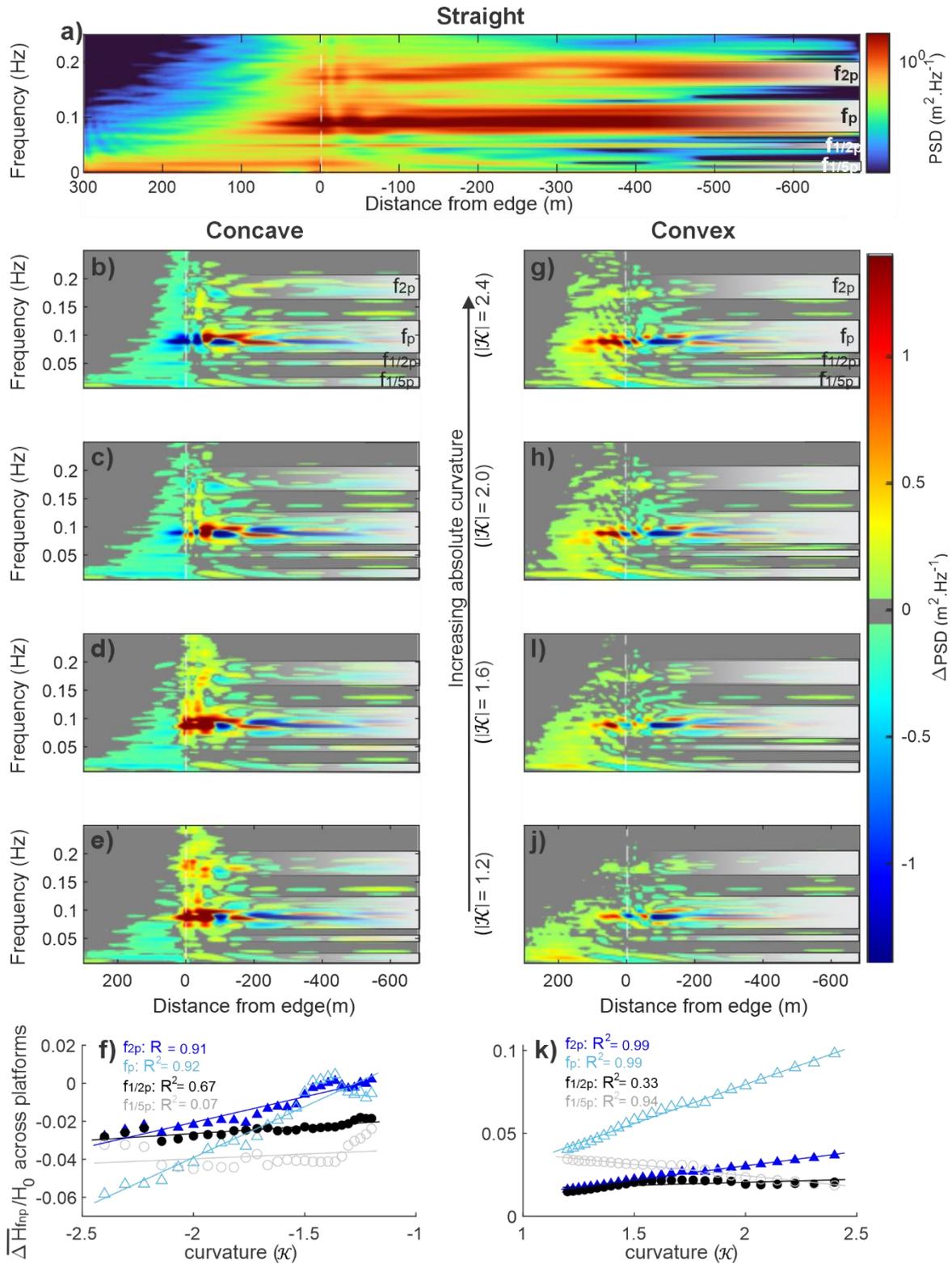


Figure 4: Impact of platform edge curvature on the harmonic components of the wavefield for broken waves showing: the spectral anomalies in relation to the straight edge platform (a) for concave platform geometries (b-e) and convex platform geometries (g-j); and the impact of curvature on the mean wave height of each harmonic across the concave (f) and convex (k) platforms.

449

450 The relationship between $\Delta \widehat{H}_s$ (Fig. 5a,b) and directional patterns (Fig. 5c,d) observed
451 across the concave and convex platform centrelines was more complex for broken than non-
452 breaking waves. The main difference with the non-breaking waves resided in the seaward
453 shift of the maximum divergence (Fig. 5c) and convergence (Fig.5d) locations over the outer
454 concave and convex platforms, respectively. This shift was particularly pronounced for convex
455 shore platforms with low degrees of curvature ($|\mathcal{K}| < 1.8$), for which the peaks of convergence
456 observed mid-platform ($x \approx 175$ m, Fig. 3d) were attenuated.

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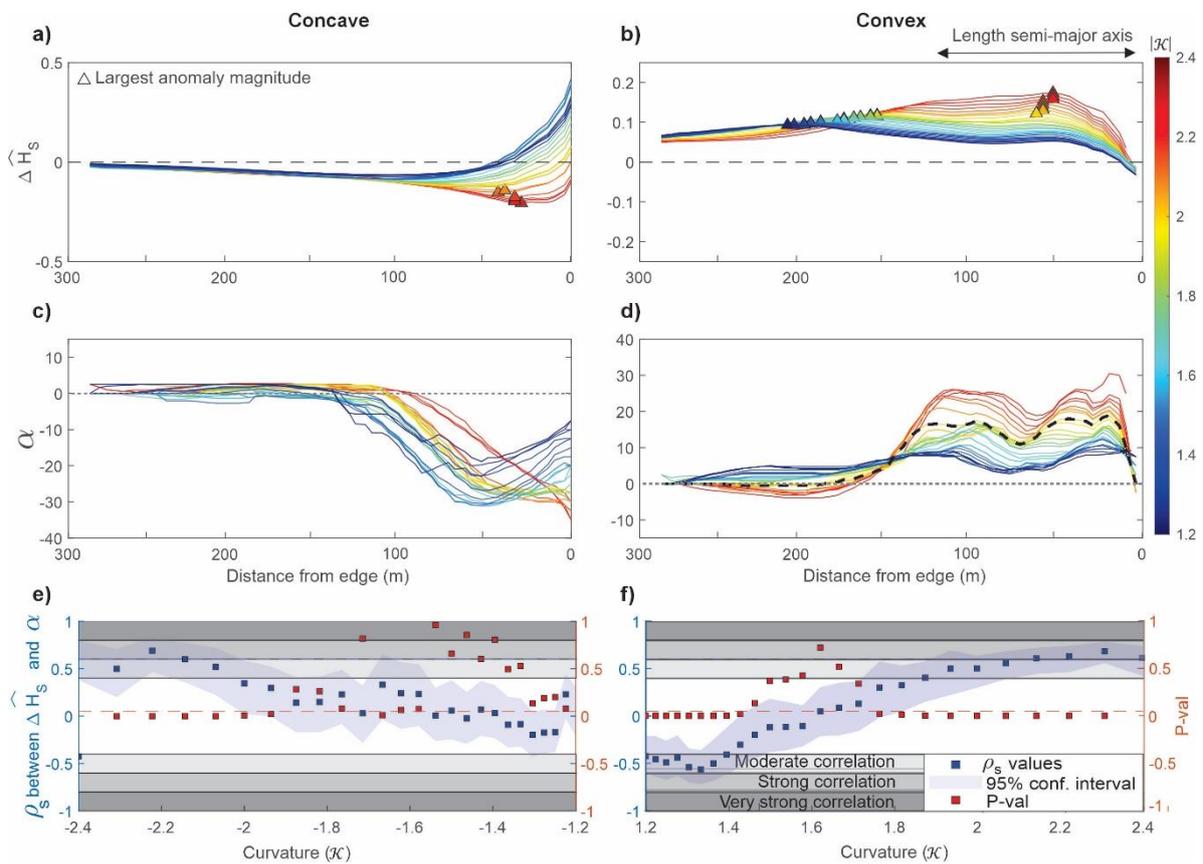


Figure 5: Relationship between directional patterns and significant wave height anomalies of broken waves across concave (left) and convex platforms (right) at the centreline ($y = 150$ m) showing: the significant wave height anomalies (a,b), the cross-shore directional patterns (c,d) and spearman correlation between these two parameters (e,f).

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461

462 For concave platforms, the seaward shift of maximum divergence zones coincided
463 with a seaward shift of the location of the largest negative anomalies (representing a 25%
464 attenuation in \widehat{H}_s , Fig. 5a). As a result, the relationship between $\Delta\widehat{H}_s$ and directional patterns
465 of broken waves remained moderate to strong ($0.4 < \rho_s < 0.6$) for concave edge curvatures
466 exceeding 1.9 (Fig. 5e), indicating that for large degrees of edge curvature, the $\Delta\widehat{H}_s$ observed
467 across the concave platform depended on the directional patterns along the centreline. For
468 convex platforms, the seaward shift of the maximum convergence locations (Fig. 5d)
469 coincided with a seaward shift of the largest positive anomalies for curvatures over 1.8
470 (representing an 18% amplification in \widehat{H}_s , Fig. 5b). However, for curvatures lower than 1.8,
471 the maximum anomalies shifted landward. Thus, the correlations between $\Delta\widehat{H}_s$ and
472 directional patterns across the centreline were moderate to strong ($0.4 < \rho_s < 0.6$) for convex
473 edge curvatures exceeding 1.8, and weak ($\rho_s < 0.4$) for curvatures dropping below 1.8 (Fig.
474 5f). This phenomenon can be explained by analysing the relative influence of each harmonic
475 component on $\Delta\widehat{H}_s$ observed across the platforms (Fig. 6).

476 For convex curvatures exceeding 1.8, the decrease of wave convergence over the
477 outer platform and wave ray divergence over the inner platforms (Fig. 5c) coincided with a
478 reduction of wave height anomalies for all harmonic components over the inner platform (Fig.
479 6). This reduction was particularly important for the fifth subharmonic, $\Delta\widehat{H}_{f_{1/5p}}$, representing
480 5% of the observed amplification of significant wave height at $x = 190$ m against 10% at $x =$
481 130 m for $|\mathcal{K}|=2.4$. In contrast, convex edge curvature below 1.8 inhibited the formation of a
482 divergence zone, ensuring the sustainability of wave ray convergence across the entire
483 platform. Under these conditions, the wave height anomalies within the WW and SW
484 frequency bands were sustained across the entire platform, and anomalies within the IG_L
485 frequency band were amplified over the inner platform ($\Delta\widehat{H}_{f_{1/5p}}$ representing 15% of the
486 observed amplification of normalised significant wave height at $x = 190$ m for $|\mathcal{K}|=1.2$) to
487 become the dominant type of anomaly at this location. Thus, $\Delta\widehat{H}_s$ became predominantly
488 controlled by the behaviour of IG_L as curvature decreased ($1.4 < \mathcal{K} < 1.8$). For very low degrees
489 of curvature ($\mathcal{K} < 1.4$), the amplification of IG_L was of such importance that $\Delta\widehat{H}_s$ were

490 amplified over the inner platform despite the decrease in wave convergence, resulting in a
 491 negative correlation ($-0.6 < \rho_s < -0.4$) between $\Delta\widehat{H}_s$ and directional patterns (Fig. 5c).

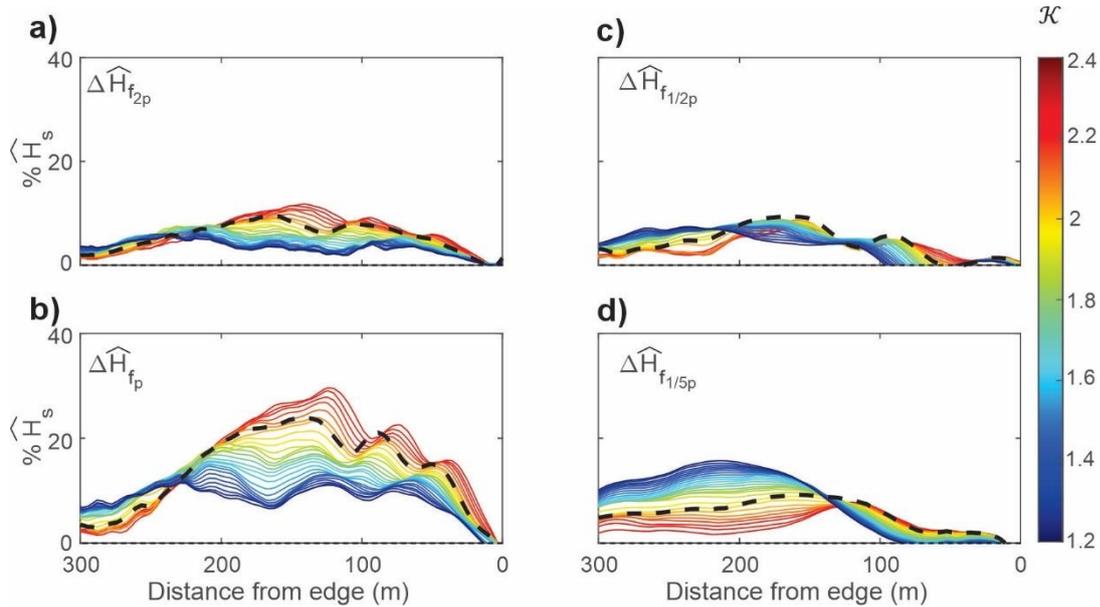


Figure 6: Percentage of significant wave height variations across the centreline ($y = 150$ m) convex platforms associated with anomalies of the second higher harmonics (a), the principal harmonic (b), the second subharmonic (c) and the fifth subharmonic (d) components ($H_s = H_{sStraight} + \Delta H_s$). The dashed line represents a curvature of 1.8, marking the threshold for the formation of a divergence zone over the inner platforms.

492

493

494 3.2 Effects of platform edge geometry on alongshore wave height patterns

495 3.2.1 Non-breaking waves

496 Coherent self-interaction maps were plotted to investigate the impact of platform edge
 497 geometry on alongshore wave height variation over the inner platform for non-breaking
 498 waves (Fig. 7). Maxima in these maps ($\widehat{\psi} \approx 1$) correspond to areas of strongest coherent
 499 interaction for the dominant frequency components within the SW and IG_L frequency bands
 500 (f_p and $f_{1/5p}$). Over the concave platforms, zones of coherent self-interaction for the principal
 501 harmonic (f_p) shifted alongshore from the platform centrelines to become concentrated near
 502 the northern and southern extremities of the platform as the edge curvature increased (Fig.
 503 7a-d). These alongshore variations were predominantly observed between $x = 130$ and 175
 504 m, where divergence along the centreline was the strongest (Fig. 3c). In contrast, coherent

505 self-interaction maps for the fifth subharmonic ($f_{1/5p}$) were more homogenous alongshore
506 (Fig. 7e-h), except near the shoreline, where coherent self-interactions were predominantly
507 observed at the platform centreline. Over the convex platform, coherent self-interactions of
508 the principal harmonic (Fig. 7i-l) were concentrated toward the platform centreline for edge
509 curvatures between $|\mathcal{K}|=1.2$ to 1.6 (Fig. 7k,l), but as edge curvature increased, coherent wave
510 interaction for this harmonic predominantly occurred at the northern and southern
511 extremities of the platforms (Fig. 7i,j). For the fifth subharmonic, coherent self-interactions
512 were focussed near the platform centreline for low edge curvature and spread alongshore
513 toward the shoreline $|\mathcal{K}|=1.2$ (Fig. 7p). As curvature increased, the areas of fifth subharmonic
514 coherent self-interactions near the shoreline split into two peaks on either side of the
515 platform centreline (Fig. 7m,n). This phenomenon was observed at curvatures for which a
516 mild divergence was observed over the inner platform (Fig. 3d).

517 The wave height distribution over the inner sections of concave and convex platforms
518 is shown in Fig. 8 and 9. The spatial distribution of the significant wave height (\widehat{H}_s) presented
519 the strongest similitudes ($R^2 > 0.9$) with the wave height patterns of the principal harmonic
520 (\widehat{H}_{fp}) regardless of the platform geometry and curvature. This indicates a strong control of
521 SW on the patterns of significant wave height variations over the inner platform. In contrast,
522 the correlation between the wave height patterns of the fifth subharmonic and the significant
523 wave height patterns over the inner platforms of concave and convex geometries was weak
524 ($R^2 < 0.4$), indicating that IG_L had little impact on the variations of significant wave height at
525 this location.

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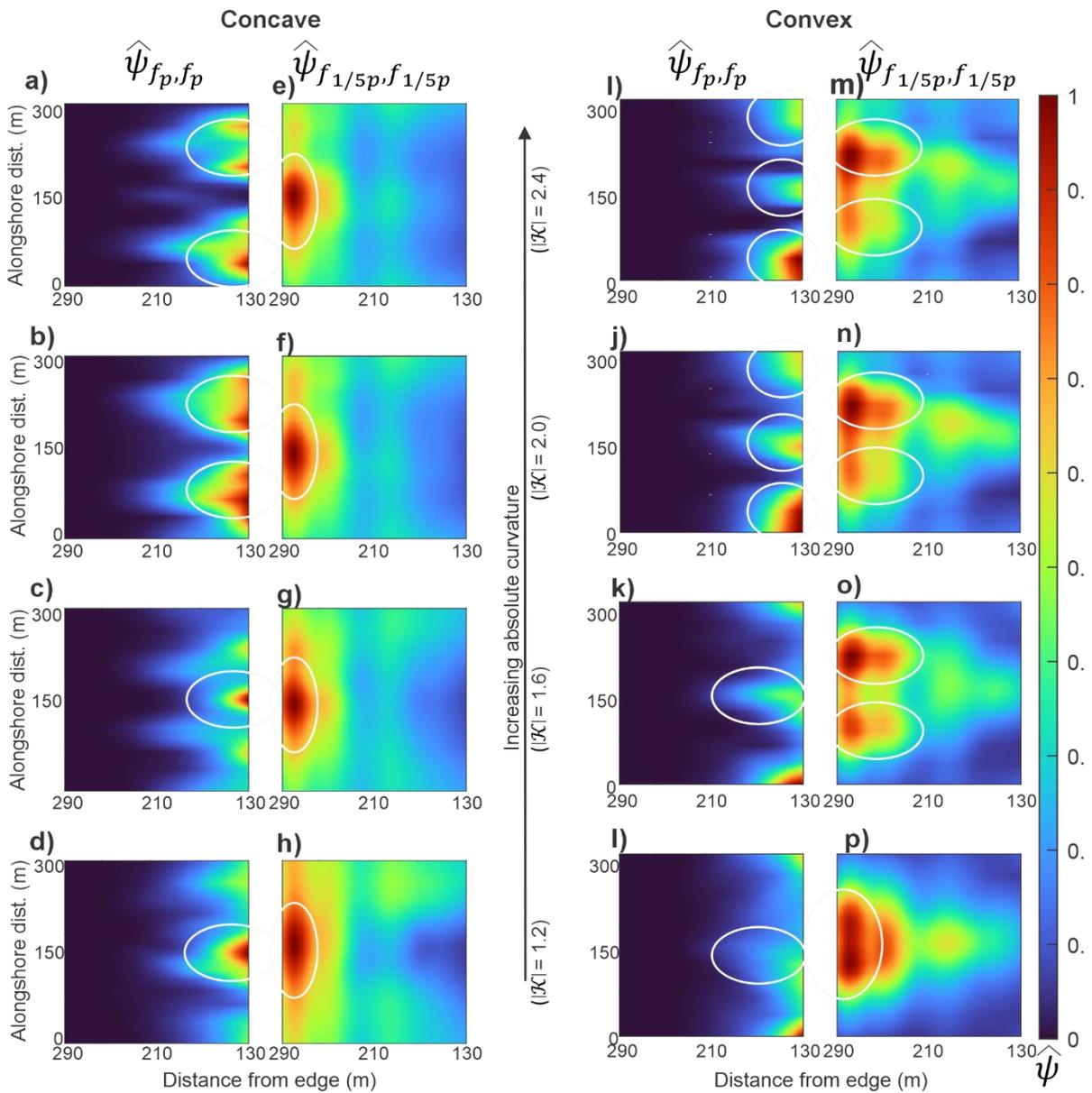


Figure 7: Coherent self-interaction maps defined from the bispectral modal state of self-interacting components for the principal harmonic (f_p) and the fifth subharmonic ($f_{1/5p}$) of non-breaking waves over the inner platform (Fig. 1) at different concave (a-h) and convex (i-p) edge curvatures. The centreline is located at $y = 150$ m. Values of $\hat{\psi}$ of 1 indicate areas of the largest coherent wave interactions. The white ellipses highlight the zones of strong coherent wave interactions.

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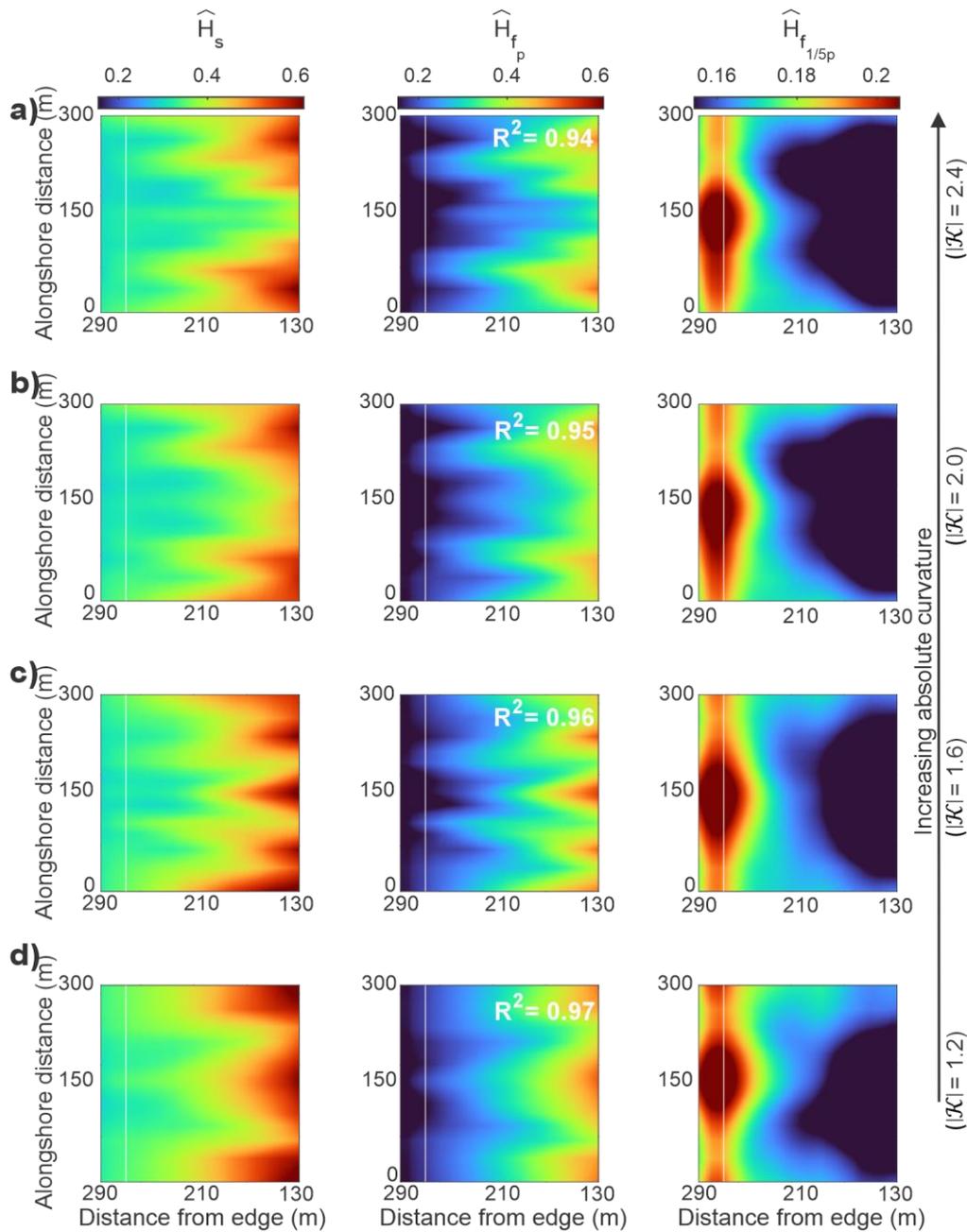


Figure 8: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of non-breaking waves over the inner platform (Fig. 1) for various concave (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline (Fig. 1). The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature (only $R^2 \geq 0.4$ is shown, representing moderate to very strong correlations)

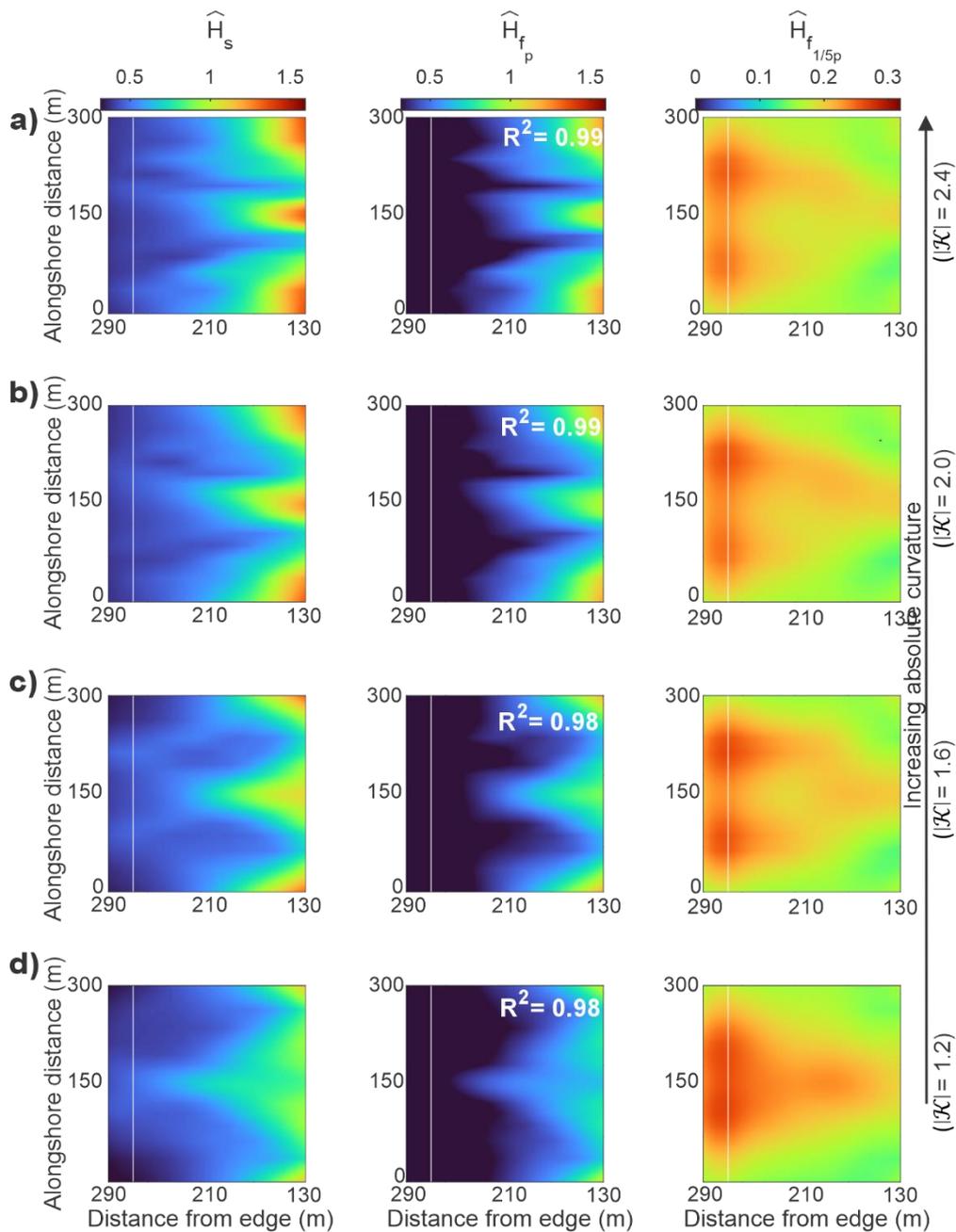


Figure 9: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of non-breaking waves over the inner platform (Fig. 1) for various convex (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline (Fig. 1). The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature (only $R^2 \geq 0.4$ is shown, representing moderate to very strong correlations)

541

542 A strong relationship, $R^2 > 0.8$, was observed between modal coherent self-interaction
 543 patterns and wave height patterns of the principal harmonic and fifth subharmonics over the
 544 inner sections of concave (Fig. 10a) and convex platforms (Fig. 10b). This observation indicates
 545 that the alongshore variations of the principal harmonic (SW) were predominantly controlled
 546 by coherent wave interaction, which in turn drove the alongshore variations in significant
 547 wave height over the inner section of both concave and convex platforms. The resulting
 548 stationary patterns in significant wave height along the shoreline were characterised by a
 549 decrease of significant wave height toward the centreline of concave platforms (Fig. 11a),
 550 which became more pronounced with increasing curvature (maximum alongshore difference
 551 in $\widehat{H}_s=0.05$ at $|\mathcal{K}|=1.2$, increasing to 0.06 at $|\mathcal{K}|=2.4$, Fig. 10a). Over convex platforms,
 552 stationary patterns for normalised significant wave height were characterised by an increase
 553 of significant wave height toward the platform centreline at low degrees of curvature (Fig.
 554 11b), resulting in an alongshore difference in $\widehat{H}_s \approx 0.15$ for $|\mathcal{K}| < 1.8$ near the shoreline. A
 555 progressive amplification of the lobes on either side of the centreline was observed as
 556 curvature increased, resulting in a more homogenous alongshore distribution of significant
 557 wave height for high degrees of curvature ($\widehat{H}_s \approx 0.06$ for $|\mathcal{K}| > 2$, Fig. 11b).

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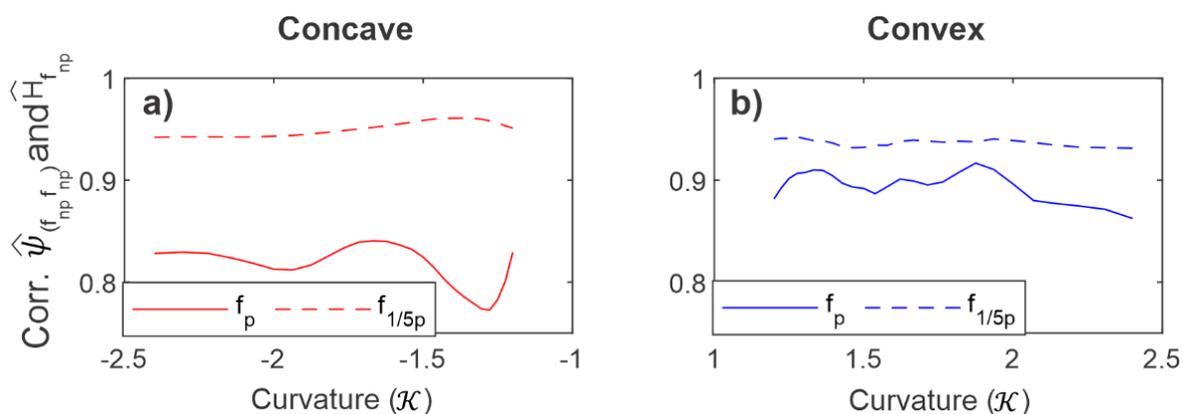


Figure 10: Correlation between interaction maps (Fig. 7) and wave height patterns (Fig. 8, 9) for the principal harmonic (f_p) and the fifth subharmonic ($f_{1/5p}$) of non-breaking waves over the inner platform of concave (a) and convex (b) edges.

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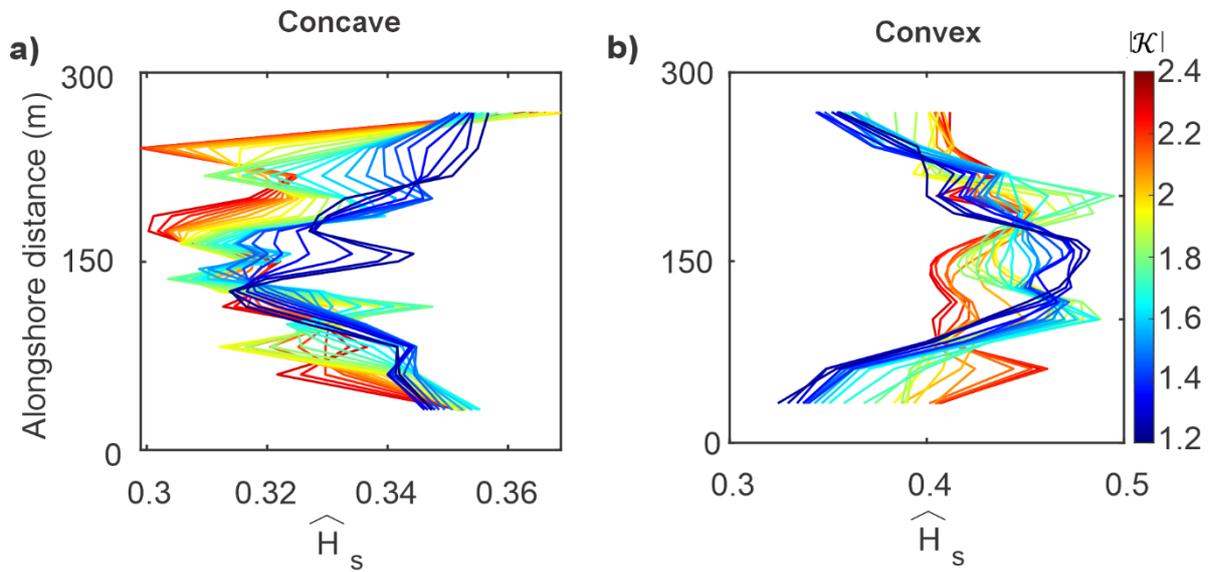


Figure 11: Alongshore variations in significant wave height patterns 20 m from the shoreline for non-breaking waves (transect L, Fig. 1) in relation to concave (a) and convex (d) edge curvatures

562 3.2.2 Broken waves

563 Coherent self-interaction patterns of the principal harmonic and fifth subharmonic of
 564 broken waves displayed alongshore variabilities over both concave and convex platforms (Fig.
 565 12). Over the concave platforms, the coherent self-interaction zone of the principal harmonic
 566 was concentrated toward the centreline for low degrees of curvature ($|\mathcal{K}|=1.2$), spreading
 567 alongshore as the degree of curvature increased (Fig. 12a-d). Zones of coherent self-
 568 interactions for the fifth subharmonic were predominantly observed on the northern and
 569 southern extremities of the platforms and became more distinct as the edge curvature
 570 increased (Fig. 12e-h). Over the convex platforms, coherent self-interactions of the principal
 571 harmonic were the strongest on the northern and southern extremities of the platforms at
 572 $x \approx 190$ m. For the fifth subharmonic (Fig. 12m-p), areas of coherent self-interaction were
 573 concentrated along the platform centrelines for low degrees of curvature ($|\mathcal{K}|=1.2$ and 1.6),
 574 but spread either side of the platform centrelines for high degrees of curvature ($|\mathcal{K}|=2.0$ and
 575 2.4). The differences in coherent self-interaction patterns between low and high degrees of
 576 curvature were characterised by a mild divergence over the inner section of convex platforms
 577 for curvatures greater than 1.8 (Fig. 7d).

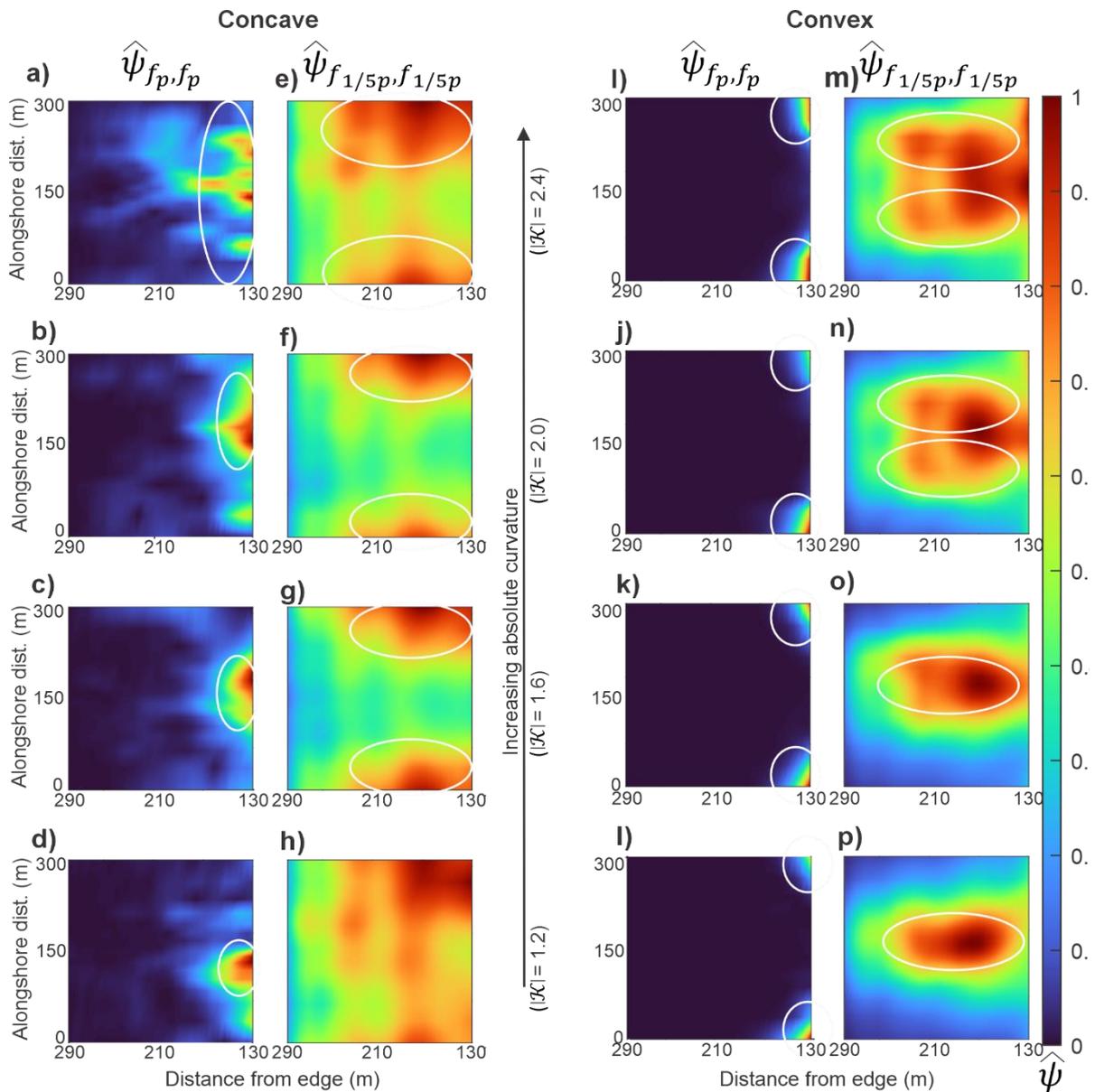


Figure 12: Coherent self-interaction maps defined from the bispectral modal state of self-interacting components for the principal harmonics (f_p) and the fifth subharmonic ($f_{1/5p}$) of broken waves over the inner platform (Fig. 1) at different concave (a-h) and convex (i-p) edge curvatures. The centreline is located at $y = 150$ m. Values of $\hat{\psi}$ of 1 indicate areas of strong coherent wave interactions. The white ellipses highlight the zones of strong coherent wave interactions.

582

583 For broken waves, the influence of IG_L on significant wave height distribution over the
584 inner sections of concave and convex platforms was greater than for non-breaking waves (Fig.
585 13, Fig. 14). Over the inner section of the concave platforms (Fig. 13), the fifth subharmonic
586 had greater wave height than the principal harmonic. Thus, the wave height patterns of the
587 fifth subharmonic had a greater impact on the significant wave height patterns ($0.85 < R^2 <$
588 0.91 for $1.2 < |\mathcal{K}| < 2.4$) than the principal harmonic ($0.74 < R^2 < 0.86$ for $1.2 < |\mathcal{K}| < 2.4$) in this
589 region. The wave height of the principal harmonic and fifth subharmonic decreased from the
590 northern and southern extremities of the platforms to the platform centrelines. The
591 combined effect of these patterns was a net alongshore decrease of significant wave height
592 toward the platform centrelines. Over the inner section of convex platforms (Fig. 14), the
593 principal harmonic displayed the greatest wave height (maximum $\widehat{H}_{fp} \approx 0.5$) on the northern
594 and southern sides of the platform between $x \approx 130-190$ m. The wave height of the fifth
595 subharmonic was relatively smaller, reaching a maximum at the platform centreline
596 (maximum $\widehat{H}_{fp} \approx 0.22-0.27$), regardless of the curvature. As a result, the wave height
597 distribution of the principal harmonic exerted a strong control on the significant wave height
598 pattern over the inner platforms ($0.9 < R^2 < 0.95$) in comparison to the control exerted by the
599 fifth subharmonic ($0.4 < R^2 < 0.58$). However, the wave height of the principal harmonic
600 significantly decreased past $x \approx 190$ m, becoming comparable to the wave height of the fifth
601 subharmonic. Thus, alongshore variations in significant wave height were controlled by
602 alongshore patterns of both principal harmonic and fifth subharmonic for $x \geq 190$ m. For low
603 degrees of edge curvature ($|\mathcal{K}|=1.2$), the maximum wave height of the fifth subharmonic was
604 observed at the platforms' centreline and evolved with increasing curvature to form two
605 maxima on either side of the centreline for large degrees of edge curvature ($|\mathcal{K}|= 2.2$ and
606 2.4). This evolution was clearly observed in the significant wave height pattern between $x =$
607 190 and 300 m, underlining the influence of the fifth subharmonic (IG_L) on the alongshore
608 variation of significant wave height at the shoreline.

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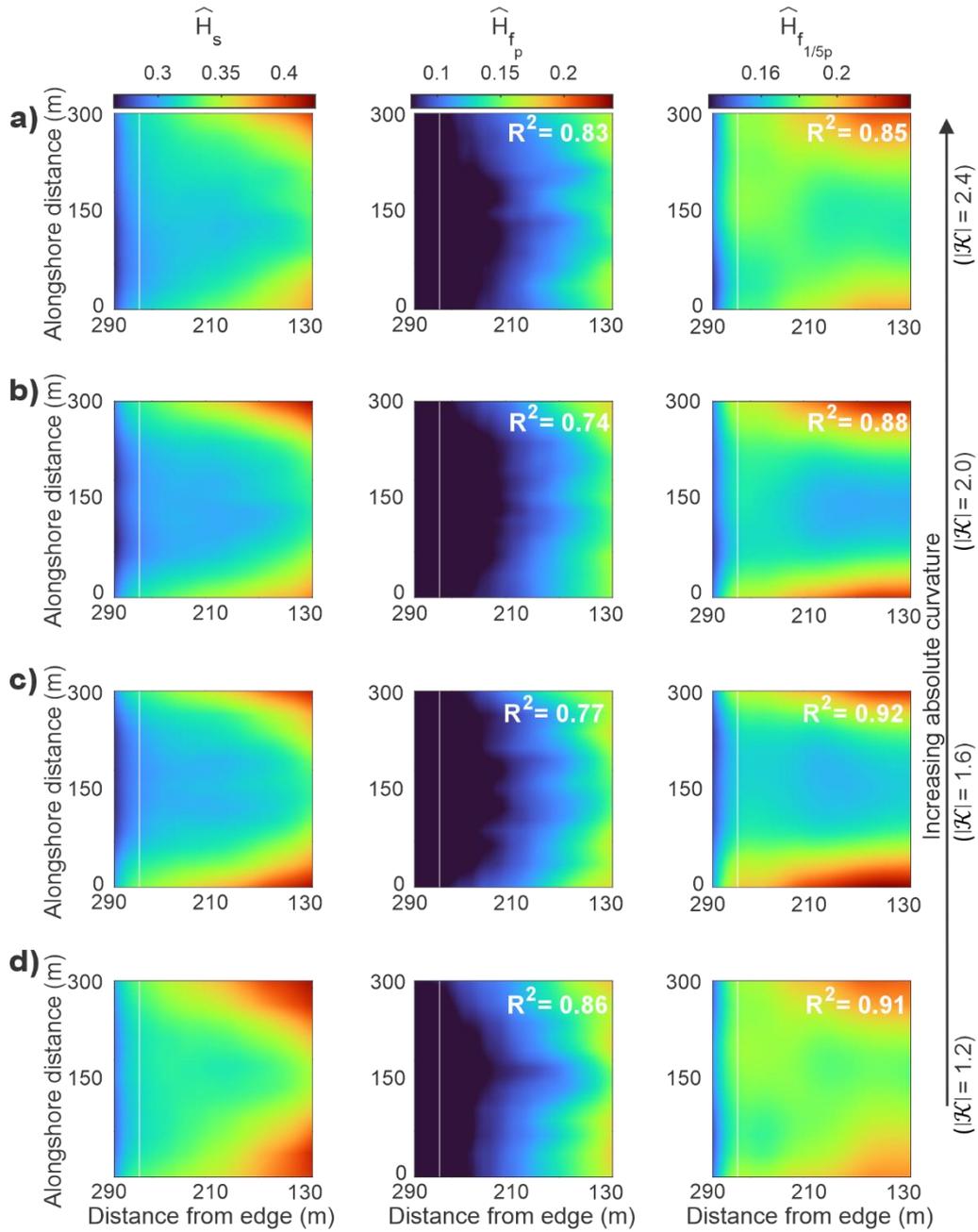


Figure 13: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of broken waves over the inner platform (Fig. 1) for various concave (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline. The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature.

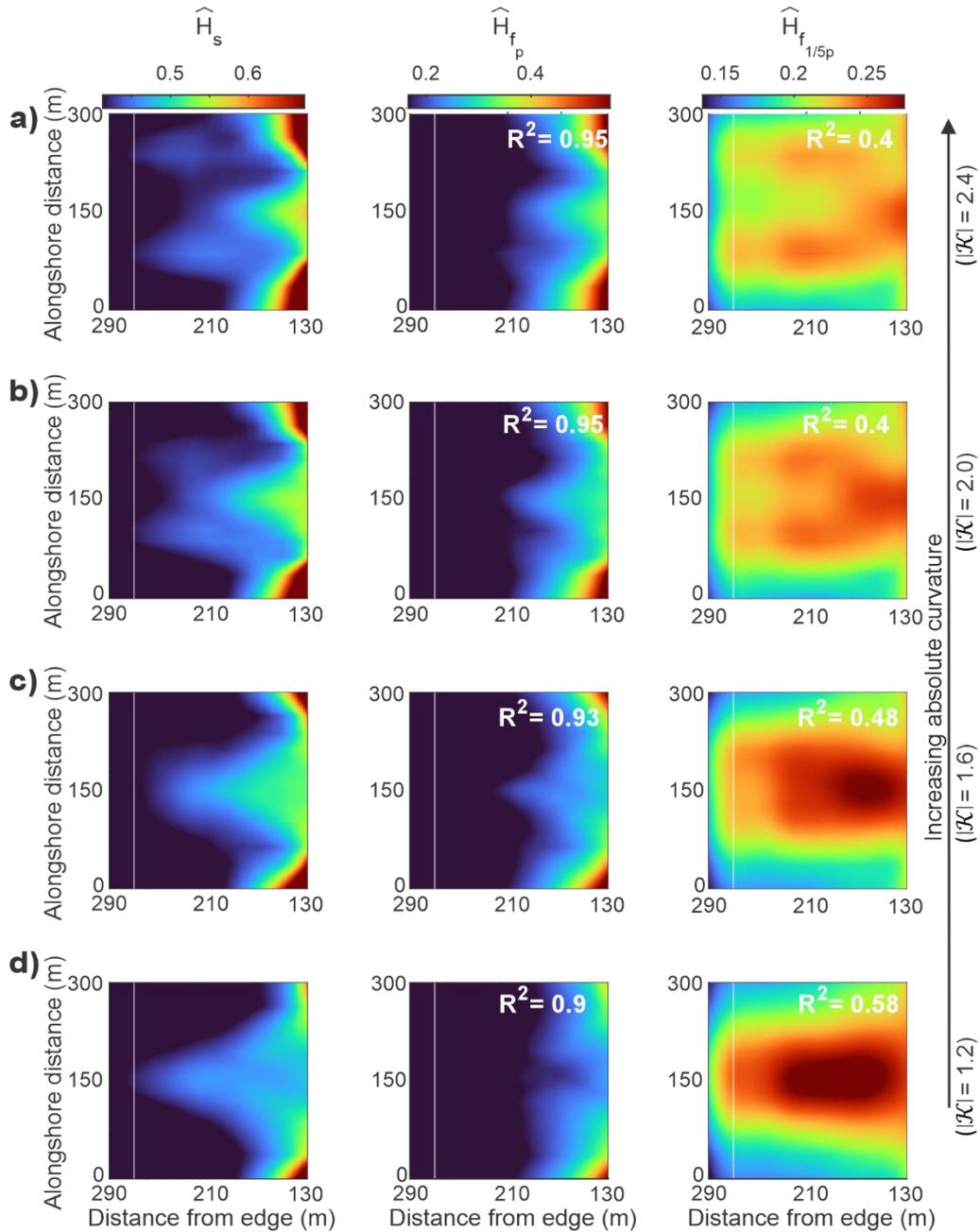


Figure 14: Wave height distribution for the entire frequency range (\widehat{H}_s), the principal harmonic (\widehat{H}_{f_p}) and the fifth subharmonic ($\widehat{H}_{f_{1/5p}}$) of broken waves over the inner platform (Fig. 1) for various convex (a-d) edge curvatures. The white line represents the alongshore transect L, 20 m from the shoreline. The centreline is located at $y = 150$ m. The R^2 values indicate the correlation between wave height patterns of the principal harmonic and fifth subharmonic with the significant wave height pattern for the same degree of curvature.

613

614

615 Strong ($R^2 > 0.8$) and moderate to strong ($0.5 < R^2 < 0.7$) relationships were observed
616 over the inner sections of concave (Fig. 15a) and convex platforms (Fig. 15b) between modal
617 coherent self-interaction patterns and wave height patterns for the fifth subharmonics and
618 principal harmonics, respectively. The implication is that coherent wave amplification
619 influenced the longshore patterns of wave height for the principal harmonic and fifth
620 subharmonic over the inner platform, although this process had a smaller impact on the
621 principal harmonic. Thus, coherent wave amplification at IG frequencies was the principal
622 process controlling alongshore variations of significant wave height along the shoreline. The
623 resulting stationary patterns in significant wave height along the shoreline were marked by a
624 decrease of significant wave height toward the centreline of concave platforms, which
625 became more pronounced with increasing curvature (maximum alongshore difference in
626 $\widehat{H}_s=0.04$ at $|\mathcal{K}|=1.2$, increasing to 0.02 at $|\mathcal{K}|=2.4$, Fig. 16a). For convex platforms, an increase
627 of significant wave height toward the platform centreline was observed at low degrees of
628 curvature, resulting in maximum alongshore variations of significant wave height $\widehat{H}_s \approx 0.08$
629 for $|\mathcal{K}|<1.8$. A progressive amplification of the lobes on either side of the centreline
630 generated two wave height maxima for high degrees of curvature, for which maximum
631 alongshore variations of significant wave height $\widehat{H}_s \approx 0.06$ for $|\mathcal{K}|>2$ (Fig. 16b).

632

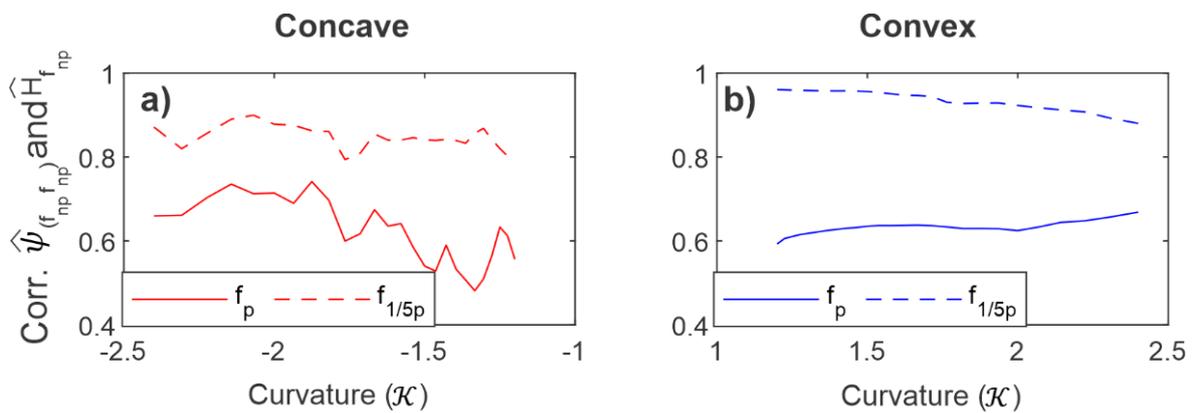


Figure 15: Correlation between interaction maps and wave height patterns for the principal harmonic (f_p) and the fifth subharmonic ($f_{1/5p}$) of broken waves over the inner platform of concave (a) and convex (b) edges.

633

634

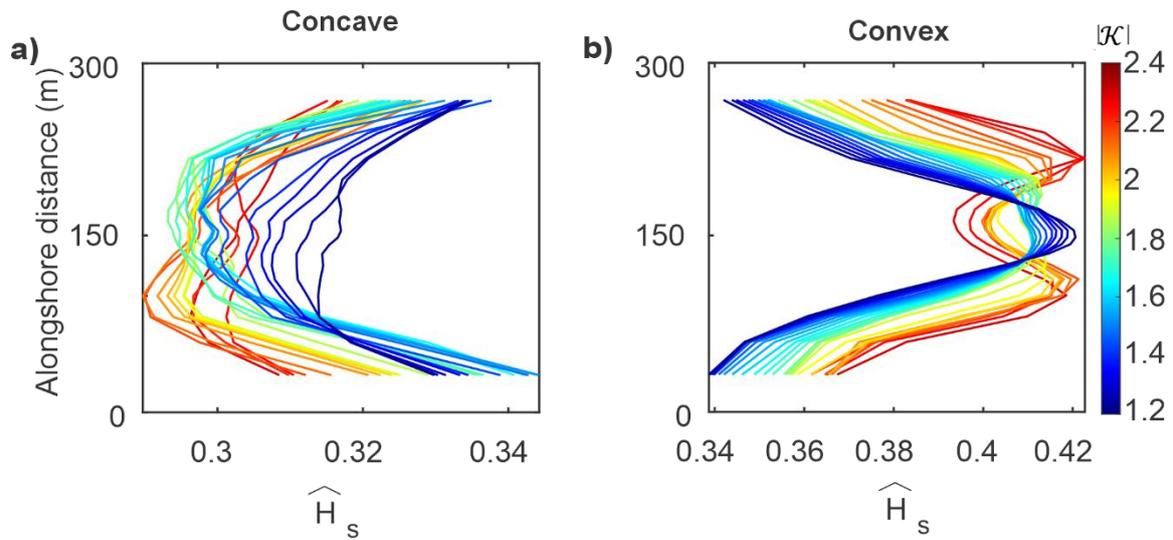


Figure 16: Alongshore variations in normalised significant wave height patterns 20 m from the shoreline (transect L) for broken waves in relation to concave (a) and convex (d) curvature.

636

637 4 Discussion

638 4.1 Impact of platform edge geometry on wave transformation across shore platforms

639 Modelling studies generally investigate the impact of refraction on wave energy
 640 distribution over a fixed curvature (e.g. Berkhoff et al., 1982; Li et al., 2020), whereas in this
 641 work, we varied the degree of edge curvature and analysed its effect on the transformation
 642 of harmonic components. Increasing concave edge curvature decreased the wave energy of
 643 harmonic components in WW, SW, IG_H and IG_L frequency bands (Fig. 2b-e, 4b-e), accounting
 644 for up to 25% reduction in \widehat{H}_s (Fig. 3a, 5a). In contrast, increasing convex edge curvature
 645 amplified both the second and principal harmonics in the WW and SW frequency bands over
 646 the outer platforms, increasing \widehat{H}_s by up to 55% and 18% for non-breaking and broken waves
 647 (Fig. 3b, 5b) while minimizing the amplification of the fifth subharmonic within the IG_L
 648 frequency band over the inner platforms (Fig. 2g-j, 4g-j). Thus, it is clear that morphological
 649 variability in platform edge curvature influences significant wave height on shore platforms,
 650 and this has potential implications for backwear and downwear erosion processes on rock
 651 coasts (e.g. Trenhaile, 1987, Matsumoto et al., 2016a,b).

652 Amplification of higher harmonics associated with wave refraction has previously been
653 associated with wave focusing (Gouin et al., 2017), but the impact of edge curvature on
654 amplification has not been considered. Our modelling results show that increasing convex
655 edge curvature enhances wave focussing over the outer platform (Fig. 3, 5), promoting the
656 generation of higher harmonic from non-linear triadic (sum) interactions (Janssen and
657 Herbers, 2009; Jarry et al., 2011; Lawrence et al., 2022). Shore platform studies have linked
658 the generation and dominance of high-frequency waves over the outer section of near-
659 horizontal platforms to locally generated wind waves (e.g. Ogawa et al., 2011; 2016). Though
660 this process cannot be ruled out, the nearshore wind speed required for locally produced WW
661 energy is substantial (Hasselmann et al., 1973), and the generation of higher harmonics from
662 non-linear triadic interaction caused by wave refraction appears to be a more plausible
663 physical interpretation for high-frequency wave generation on the outer sections of shore
664 platforms.

665 Research has demonstrated that wave amplification in the IG frequency band over near-
666 horizontal platforms is influenced by the ratio of water depth at the cliff toe to the platform
667 width, and relative submergence (Beetham and Kench, 2011; Ogawa et al., 2015). We show
668 that edge curvature exerts an additional morphological control on IG amplification across
669 convex platforms by affecting the balance between focusing intensity from refraction and
670 defocusing effects from wave breaking controlled by convex edges (Fig. 4, 5). Although a
671 decrease in convex edge curvature should theoretically result in a landward shift of the focal
672 point over submerged flats (e.g. Mandlier and Kench, 2012), a seaward shift of the focal point
673 was observed in this study for broken waves (Fig. 5d). This phenomenon can be attributed to
674 the defocussing effects resulting from the enhancement of radiation stress and wave-
675 generated current by wave breaking (Yoon et al., 2004; Choi et al., 2009). A critical curvature
676 was found for which the intensity of wave focusing by wave refraction is not strong enough
677 to overcome the defocusing effects of wave breaking, identified here as $|\mathcal{K}|=1.8$ (Fig. 5d).
678 When the critical curvature is exceeded, wave rays intersect across the platform centreline,
679 in which cases, IG amplification is minimised by wave ray divergence over the inner platform
680 (Fig. 6d). In contrast, for convex edge curvatures lower than the critical curvature, wave rays
681 do not intersect across the platform centreline, sustaining wave convergence across the
682 entire platform. In this case, IG amplification is promoted over the inner platform,

683 representing up to 15% of the increase in significant wave height at this position (Fig. 6d).
684 Thus, the present research identifies convergence as a key mechanism acting on the growth
685 of IG across shore platform, working in conjunction with other processes such as to
686 breakpoint forcing (Poate et al., 2020) and energy transfer from higher frequencies and
687 shoaling (Beetham and Kench, 2011).

688 In a conceptual model, Ogawa et al. (2011) described spatial zones on a shore platform
689 that are dominated by different wave types and pointed out that the spatial characteristics
690 change according to the tidal stage. Here, we show that shifts in dominant wave types across
691 shore platforms are also controlled by convex edge curvature (certainly at high tide). High
692 convex curvature amplifies harmonics within the WW and SW frequency bands over the outer
693 platform and inhibits the amplification of IG over the inner platform (Fig. 4g). Low degrees of
694 convex curvature have the opposite effect (Fig. 4j), resulting in the seaward shift of the zones
695 dominated by WW and SW frequency bands. The influence of refraction patterns generated
696 by convex edge geometries on the collective behaviour of harmonics affected the significant
697 wave height patterns across convex platforms. Baldock et al. (2020) hypothesised that the
698 significant wave height across convex platforms is defined by a specific balance between
699 cross-shore energy loss from dissipation and energy gain from oblique refracted SW, resulting
700 in a correlation between significant wave height anomalies and refraction patterns. The
701 strong relationship ($\rho_s > 0.6$) observed between directional patterns and significant wave
702 height anomalies of non-breaking waves (Fig. 3f) suggest that, for this wave state, energy is
703 effectively gained across the platform from refracted SW. However, for broken waves, the
704 correlation between directional patterns and significant wave height anomalies ($\Delta\widehat{H}_s$)
705 decreased with curvature to become weak ($\rho_s < 0.4$) below the critical curvature threshold
706 (Fig. 5f). These differences are attributed to the influence of IG growth on significant wave
707 height over the inner platform due to the amplification of IG from post-breaking energy
708 transfer from high to low frequency (Poate et al., 2020), and low sustained convergence
709 across the platform (Fig. 5b, 6d). These results underline the importance of considering the
710 refraction patterns of both SW and IG when investigating significant wave height patterns
711 across convex submerged flats.

712

713 4.2 Impact of platform edge geometry on along shore wave transformation

714 Coherent wave amplification is identified here as a crucial process affecting SW and IG_L
715 height distribution in the inner section of near-horizontal platforms. For non-breaking waves,
716 the dominant frequencies within both SW and IG_L frequency bands presented strong
717 correlations between patterns of coherent wave interaction and wave height distribution
718 over concave and convex platforms (Fig. 10). However, this correlation decreased for the
719 principal harmonic in the SW frequency bands in the presence of wave breaking (Fig. 15),
720 perhaps due to the combination of defocussing (Yoon et al., 2004) and dissipation effects
721 (Farrell et al., 2009) associated with wave breaking.

722 The present observations validate the hypothesis of Winter et al. (2017) on the
723 formation of alongshore stationary IG patterns from wave refraction over convex platforms.
724 However, our results suggest that such patterns are generated by coherent wave interaction
725 following the generation of caustic rays in the IG_L frequency band (Fig. 7m-p, 12m-p) rather
726 than alongshore standing waves, as Winter et al. (2017) suggested. In fact, the latter would
727 require interacting IG to propagate alongshore in opposite directions. The directional analysis
728 presented here precludes such a possibility ($\alpha \approx 0^\circ$ near the shoreline, Fig. 3, 5). As the
729 present paper demonstrates the coherent wave interaction plays crucial role in alongshore IG
730 wave patterns, a coherent wave class should be added to the resonant, progressive-
731 dissipative, standing and progressive-growing low-frequency wave classes previously
732 identified over submerged flats (Gawehn et al., 2016).

733 The combined modes of SW and IG coherent wave amplification exerted a crucial
734 control on the stationary patterns of significant wave heights over the inner platform. For
735 non-breaking waves, significant wave height variations over the inner platform are
736 predominantly controlled by coherent wave amplification of SW (Fig. 7, 8). In contrast, for
737 breaking waves, the distribution of significant wave height over the inner platform became
738 controlled by coherent wave amplification occurring within both SW and IG_L frequency
739 components as IG_L became a prominent wave type in this region (Ogawa et al. 2011) (Fig. 13,
740 14). The present results support the conceptual model presented by Krier-Mariani et al.
741 (2022), suggesting that patterns of wave ray intersection on either side of concave edge
742 sections result in stationary SW and IG amplification patterns. The control exerted by the

743 critical curvature on the alongshore distribution of SW and IG over convex platforms can be
744 attributed to the generation of a terminal point (marking the transition from wave ray
745 convergence to divergence) at the platform centreline for convex curvatures exceeding the
746 critical curvature. In such cases, caustic rays are formed on either side of the centreline
747 (Mandlier and Kench, 2012), promoting coherent wave interactions in these regions (Fig. 12a,
748 14a).

749 It follows that the control exerted by platform curvature on coherent wave
750 amplification plays an essential role in the nodal state of significant wave height along the
751 shoreline (Fig. 11, 16). For concave platforms, a node near the centreline and antinodes on
752 the northern and southern extremities of the platform were observed. For convex edge
753 curvature under the critical curvature threshold, an antinode was observed along the
754 platform centreline where waves converged, while for curvature exceeding the critical
755 curvature threshold, two antinodes were observed on either side of the platform centreline.
756 While such patterns could wrongly be associated with edge waves, the present results
757 support the observations of Dalrymple (1975), who first associated nodal and anti-nodal
758 points in alongshore wave height patterns with coherent wave interaction.

759 It has previously been established that by controlling the nodal state of significant wave
760 height along the shoreline of open coasts, coherent wave amplification could lead to the
761 formation of rip currents (Dalrymple, 1975; Wei and Dalrymple, 2017). This mechanism is
762 expected to impact circulation patterns over near-horizontal platforms equally. da Silva et al.
763 (2023) identified alongshore pressure gradient as the dominant driver of circulation patterns
764 in the lee of submerged flats, resulting in two or four-cell circulation systems. A two-cell
765 system is typically characterised by an alongshore diverging flow from the lee of the
766 submerged flat edge to the shoreline, while a four-cell system is characterised by an
767 alongshore diverging flow at the lee of the submerged flat and a converging flow at the
768 shoreline. Considering the present results, it can be hypothesised that stationary wave
769 patterns and the subsequent alongshore pressure gradient generated by coherent wave
770 amplification drive the formation of circulation cells over convex platforms. Theoretically, the
771 formation of two antinodes over the inner platform would result in a four-cell circulation
772 system (Fig. 14a,b, Fig. 16b), while an antinode across the entire platform would result in a

773 two-cell circulation system (Fig. 14c,d, Fig.16b). These differences depend on whether or not
774 the submerged flat geometry allows for the formation of a terminal point. Previous studies
775 established that this condition was predominantly controlled by the distance between the
776 seaward edge of submerged flats and the shoreline (da Silva et al., 2022, 2023; Ranasinghe et
777 al., 2006, 2010), while we show that the degree of edge curvature is equally important (Fig.
778 11, 16).

779 **5 Conclusions**

780 This study employed an exploratory numerical modelling approach to investigate the
781 impact of concave and convex platform edge geometries on the behaviour of wave harmonics
782 and the subsequent wave height distribution patterns over near-horizontal shore platforms.
783 Harmonic analyses show that refraction patterns controlled by concave and convex platform
784 edge curvatures result in wave height variation for the principal and second higher harmonics
785 over the outer platform, and for the subharmonics over the inner platform. Wave divergence
786 across concave edge platforms decreased the height of harmonics within both SW and IG
787 frequency bands, resulting in the attenuation of significant wave height for high degrees of
788 curvature. Over the outer section of convex platforms, increasing curvature intensified wave
789 focusing and amplified the principal and second harmonics within the SW frequency band. A
790 critical curvature value of 1.8 demarcates the formation of a wave ray divergence zone over
791 the inner platform, conditioned by the balance between wave focusing from wave refraction
792 and wave defocusing from wave breaking. Below this threshold, wave convergence amplified
793 IG over the inner platform, but over this threshold, wave divergence reduced the
794 amplification of IG over the inner platform. Through these mechanisms, it is apparent that
795 edge curvature can influence both the relative dominance of SW and IG frequencies, and the
796 pattern of significant wave height transformation across near-horizontal platforms. Using a
797 high-order spectral decomposition method, this study further demonstrated that coherent
798 wave amplification influences stationary IG and SW patterns over the inner platform,
799 affecting the alongshore distribution of significant wave height. We found that platform
800 geometry controls the nodal state of the stationary patterns along the shoreline, possibly
801 resulting in alongshore variation of wave erosive force and the generation of wave-generated
802 currents shaping rock coasts.

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809

810 **Open research**

811 The numerical model input files and post processed data for the model simulations
812 (Using FUNWAVE 3.6) over idealised shore platform geometries are available at *Public release*
813 *planned after review.*

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Appendix 1

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Description of the fully nonlinear Boussinesq wave model FUNWAVE-TVD

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The numerical phase resolving model FUNWAVE_TVD is based on the conservative form of the fully nonlinear Boussinesq equations formulated by Shi et al. (2012). Following Tonelli and Petti (2009) wave breaking forces the model to switch from Boussinesq equations, where dispersive and nonlinear effects are of a similar order of magnitude, to the nonlinear shallow water equation, where nonlinearity dominates. This model employs a Total Variation Diminishing (TVD) spatial discretisation scheme to solve the fully non-linear Boussinesq equation (combining finite-volume for nonlinear terms and finite-difference for dispersive terms) and incorporates a time-dependent reference level (Kennedy et al. 2001) moving with the instantaneous free surface to calculate the velocity potential. The combination of the shock-capturing TVD scheme and moving reference provides robust performance in simulating breaking waves and optimising nonlinear behaviour. Furthermore, the model uses an adaptative time stepping defined from a third-order Strong Stability-Preserving (SSP) Runge–Kutta scheme (Gottlieb et al., 2001) to increase model stability. The conservative form of the fully nonlinear Boussinesq equations in FUNWAVE_TVD employs a modification of the leading order pressure term in the momentum equation using a modified surface gradient term such as:

$$\eta_t + \nabla \cdot \mathbf{M} = 0 \quad (\text{A1.1})$$

$$\begin{aligned} M_t + \nabla \cdot \left[\frac{MM}{H_{tot}} \right] + \nabla \left[\frac{1}{2} g(\eta^2 + 2h\eta) \right] \\ = H_{tot} \{ \bar{u}_{2,t} + u_\alpha \cdot \nabla \bar{u}_2 + \bar{u}_2 \cdot \nabla u_\alpha - V'_{1,t} - V_1'' - V_2 - V_3 - R \} \\ + g\eta \nabla h \end{aligned} \quad (\text{A1.2})$$

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23 where ∇ denotes the horizontal partial derivative $((\partial/\partial x), (\partial/\partial y))$, η is the free surface
 24 elevation, h is the water depth, $H_{tot} = h + \eta$ is the total local water depth and g is the
 25 gravitational acceleration, the terms $\nabla \left[\frac{1}{2}g(\eta^2 + 2h\eta) \right]$ and $g\eta\nabla h$ are components of the
 26 surface gradient. The horizontal volume flux is expressed as:

$$M = H_{tot}\{u_\alpha + \bar{u}_2\} \quad (\text{A1.3})$$

27
 28 where u_α is the horizontal velocity at the reference level $z_\alpha = \zeta h + \beta\eta$ (from Kennedy et al.
 29 (2001)) with $\zeta = -0.53$ and $\beta = 0.47$. While u_2 is the depth dependant correction at $O(\mu^2)$ (with
 30 μ representing the ratio of depth over wave length) that is expressed as:

$$u_2(z) = (z_\alpha - z)\nabla A + \frac{1}{2}(z_\alpha^2 - z^2)\nabla B \quad (\text{A1.4})$$

31
 32 with $\nabla A = \nabla \cdot (hu_\alpha)$ and $\nabla B = \nabla \cdot u_\alpha$. The depth-averaged contribution to the horizontal
 33 velocity field is given by:

$$\bar{u}_2 = \frac{1}{H_{tot}} \int_{-h}^{\eta} u_2(z) dz = \left[\frac{z_\alpha^2}{2} - \frac{1}{6}(h^2 - h\eta + \eta^2) \right] \nabla B + \left[z_\alpha + \frac{1}{2}(h - \eta) \right] \nabla A \quad (\text{A1.5})$$

34
 35 V_1 and V_2 represent the dispersive terms of the Boussinesq equation defined as:

$$V_1 = \left\{ \frac{z_\alpha^2}{2} \nabla B + z_\alpha \nabla A \right\}_t - \nabla \left[\frac{\eta^2}{2} B_t + \eta A_t \right] \quad (\text{A1.6})$$

$$V_2 = \nabla \left\{ (z_\alpha - \eta)(U_\alpha \cdot \nabla) A + \frac{1}{2}(z_\alpha^2 - \eta^2)(U_\alpha \cdot \nabla) B + \frac{1}{2}[A + \eta B]^2 \right\} \quad (\text{A1.7})$$

36
 37
 38 with V_3 representing the second order $(O(\mu^2))$ effect of the vertical velocity, which is
 39 expressed as:

$$V_3 = \omega_0 i^z \times \bar{u}_2 + \omega_2 i^z \times u_\alpha \quad (\text{A1.8})$$

44 Where with i^z the unit vector in the vertical direction and:

45

$$\omega_0 = (\nabla \times u_\alpha) \cdot i^z = v_{\alpha,x} - u_{\alpha,y} \quad (\text{A1.9})$$

$$\omega_2 = (\nabla \times \bar{u}_2) \cdot i^z = z_{\alpha,x}(A_y + z_\alpha B_y) - z_{\alpha,y}(A_x + z_\alpha B_x) \quad (\text{A1.10})$$

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47 R in Eq. A3.2 represents the combination of diffusive (R_s) and dissipative (R_f) terms (Chen et

48 al., 1999) induced by sub-grid lateral turbulent mixing and bottom friction, $R = R_s + R_f$, with

49 R_f , expressed as:

$$R_f = \frac{C_d}{h + \eta} u_\alpha |u_\alpha| \quad (\text{A1.11})$$

50 where C_d is the bottom friction coefficient.

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Appendix 2

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Description of the Bispectra Mode Decomposition

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Advantages and limitations of orthogonal decomposition methods

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Development and advantages of high-order spectral decomposition methods

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Investigating the generation of stationary patterns from coherent wave amplification requires a decomposition method capable of holding information on both spectral and phase characteristics of the wavefield. Such information can be provided by high-order statistical analyses such as bispectrum, defined from the third moment of the data field (Hasselmann et al., 1963). The bispectrum presents attractive properties to identify coherent wave amplification. It is not only capable of detecting quadratic phase coupling for specific sets of frequencies but also represents a measure of skewness, which is expected to increase in areas of wave ray intersection (e.g. Janssen and Herbers, 2009). Despite these advantages, the bispectrum is only applicable to one-dimensional spatial domains. To overcome this limitation, Schmidt (2020) recently introduced the Bispectra Mode Decomposition (BMD), which consists of maximising the expansion coefficients of a spatial integral measure of the bispectrum. Thus, the BMD can be regarded as a decomposition method based on the same principle as the spectral EOF but applied to higher-order spectral analysis.

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95 Description of the Bispectral Mode decomposition method

96 In the BMD approach, the time series of two-dimensional sea surface elevation
 97 observations defined in the time domain and cartesian coordinate system ($q(\xi, t) \in \mathbb{C}^{M \times N_t}$)
 98 are first redefined in the frequency domain using Welch's method (Welch, 1967) such as:

$$\hat{q}(\xi, f_k) = \sum_{j=0}^{N_{FFT}-1} q(\xi, t_{j+1}) e^{-i2\pi jk/N_{FFT}} \quad (A2.1)$$

with $k = 0, \dots, N_{FFT} - 1$

99 where $q(\xi, t_j) \in \mathbb{C}^M$ represents the two-dimensional sea surface observations in the spatial
 100 domain ξ defined by a number of points $M = N_x, N_y, N_z$ at a sample time t_j with $j = 0, \dots, N_t$.
 101 N_{FFT} represents the number of samples in one of the N_{blk} segments used to calculate the
 102 Fourier transform. Two-dimensional observations are, therefore, redefined in the space-
 103 frequency domain $\hat{q}(\xi, f_k) \in \mathbb{C}^{M \times N_{blk}}$.

104 The product of the Fourier coefficients used to define the bispectrum for frequencies k
 105 and l is obtained from the Hadamard product of the matrices $\hat{q}(\xi, f_k) \equiv \hat{q}_k$ and $\hat{q}(\xi, f_l) \equiv \hat{q}_l$
 106 such as:

$$\hat{q}_{k \circ l} = \hat{q}_k \circ \hat{q}_l \quad (A2.2)$$

107 The spatial integral measure of the bispectrum is therefore expressed as:

$$b(f_k, f_l) = E \left[\int_{\Omega} \hat{q}_k^* \circ \hat{q}_l^* \circ \hat{q}_{k+l} d\xi \right] = E [\hat{q}_{k \circ l}^H \hat{q}_{k+l}] \quad (A2.3)$$

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109 where $E[.]$ is the expectation operator, $(.)^*$ and $(.)^H$ denote the complex conjugate and
 110 transpose, respectively. Assuming that the observed fluid is incompressible, the form of the
 111 triadic interaction in the Navier-Stokes is used in the BMD to establish a causal relationship
 112 between the product of the two interacting frequency components represented by the term
 113 $\hat{q}_{k \circ l}$ in Eq. A2.3, generating the third frequency component represented by the term \hat{q}_{k+l} .
 114 Therefore, the interacting and resulting frequency components are linked by a shared
 115 expansion coefficient, a_{ij} , in the modal decomposition and defined by the linear expansions:

$$\phi_{k \circ l}^{[i]}(\xi, f_k, f_l) = \sum_{j=1}^{N_{blk}} a_{ij}(f_{k+l}) \hat{q}_{k \circ l}^{[j]} \quad (\text{A2.4})$$

$$\phi_{k+l}^{[i]}(\xi, f_{k+l}) = \sum_{j=1}^{N_{blk}} a_{ij}(f_{k+l}) \hat{q}_{k+l}^{[j]} \quad (\text{A2.5})$$

116 The cross-frequency fields $\phi_{k \circ l}$ are maps of phase alignment between two frequency
 117 components, while bispectral modes ϕ_{k+l} are linear combinations of Fourier modes related
 118 to the amplitude of oscillations of the sea surface at frequency $k + l$. Consequently, the
 119 modal decomposition in the BMD is defined from the spectral properties of each segment
 120 obtained from the Welch method rather than from the raw two-dimensional time series of
 121 observations conventionally used in the EOF analysis. Eq. A2.4 and A2.4 can be, therefore,
 122 regarded as the product of expansion coefficients and data matrices such as:

$$\phi_{k \circ l}^{[i]} = \hat{Q}_{k \circ l} a_i \quad (\text{A2.6})$$

$$\phi_{k+l}^{[i]} = \hat{Q}_{k+l} a_i \quad (\text{A2.7})$$

123 Where $\hat{Q}_{k \circ l}$ and $\hat{Q}_{k+l} \in \mathbb{C}^{M \times N_{blk}}$ and $a_i = [a_{i1}(f_{k+l}), \dots, a_{iN_{blk}}(f_{k+l})]^T$ represents the i -th
 124 vector of expansion coefficients for the (k, l) frequency doublets, with $(\cdot)^T$ denoting the
 125 transpose. To optimally represent the sea surface characteristics in terms of integral
 126 bispectral density, the set of expansion coefficients a_1 maximising the value of $b(f_k, f_l)$ in Eq.
 127 A2.3 is defined from the numerical radius of the complex product matrix B representing the
 128 bispectral density matrix:

$$B = \hat{Q}_{k \circ l}^H \hat{Q}_{k+l} \quad (\text{A2.8})$$

129 To seek the expansion coefficients corresponding to the largest eigenvalue λ_{max} . This method
 130 allows defining an optimal approximation of the eigenvalue characterising the integral
 131 bispectral density of the wavefield for each pair of frequency components, referred to as the
 132 mode bispectrum $\lambda_1(f_k, f_l)$. That is, the integral bispectral density is best represented by the
 133 first mode of the BMD, with other modes having a minimal impact. The peak magnitude of
 134 the optimal complex eigenvalue $|\lambda_1(f_k, f_l)|$ for the set of frequencies f_k and f_l is analogue
 135 to the peak magnitude found in the bispectrum. Therefore, the BMD defines the modal states
 136 of the wavefield in regard to interactions between frequency components, which allows the

137 extraction of spatial structures of phase coupling and resulting triadic interactions in two
138 dimensions.

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