Resonances in fluid-filled cracks of complex geometry and application to very long period (VLP) seismic signals at Mayotte submarine volcano

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Abstract

Fluid-filled cracks sustain a slow guided wave (Krauklis wave or crack wave) whose resonant frequencies are widely used for interpreting long period (LP) and very long period (VLP) seismic signals at active volcanoes. Significant efforts have been made to model this process using analytical developments along an infinite crack or numerical methods on simple crack geometries. In this work, we develop an efficient hybrid numerical method for computing resonant frequencies of complex-shaped fluid-filled cracks and networks of cracks and apply it to explain the ratio of spectral peaks in the VLP signals from the Fani Maoré submarine volcano that formed in Mayotte in 2018. By coupling triangular boundary elements and the finite volume method, we successfully handle complex geometries and achieve computational efficiency by discretizing solely the crack surfaces. The resonant frequencies are directly determined through eigenvalue analysis. After proper verification, we systematically analyze the resonant frequencies of rectangular and elliptical cracks, quantifying the effect of aspect ratio and crack stiffness ratio. We then discuss theoretically the contribution of fluid viscosity and seismic radiation to energy dissipation. Finally, we obtain a crack geometry that successfully explains the characteristic ratio between the first two modes of the VLP seismic signals from the Fani Maoré submarine volcano in Mayotte. Our work not only reveals rich eigenmodes in complex-shaped cracks but also contributes to illuminating the subsurface plumbing system of active volcanoes. The developed model is readily applicable to crack wave resonances in other geological settings, such as glacier hydrology and hydrocarbon reservoirs.

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Resonances in fluid-filled cracks of complex geometry and application to very long period (VLP) seismic signals at Mayotte submarine volcano

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Key Points:

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13	• Hybrid method combining BEM and FVM efficiently computes resonant frequen-
14	cies of complex-shaped fluid-filled cracks
15	• Elliptical crack shares similar modes with rectangular crack but a crack network
16	produces more complex resonances
17	- A dumbbell-shaped crack explains ratio of first two modes (~2.5) in the VLP seis-
18	mic signal at Mayotte submarine volcano

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19 Abstract

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40 1 Introduction

Slow guided waves that propagate along fluid-filled cracks, named crack waves or 41 Krauklis waves, can be used for inferring the geometries of subsurface cracks and the fluid 42 properties in a wide range of geological settings (Krauklis, 1962; Ferrazzini & Aki, 1987; 43 Paillet & White, 1982; B. Chouet, 1986; Korneev, 2008; Tang & Cheng, 1989; Lipovsky 44 & Dunham, 2015). In volcanology, crack wave resonances along magma-filled sills and 45 dikes have been used for interpreting long period (LP, 0.5-2 s) and very long period (VLP, 46 2 to 100 s) seismic signals at many volcanos, including Mount Redoubt (B. A. Chouet 47 et al., 1994), Aso (Kawakatsu et al., 2000; Niu & Song, 2020), Galeras (Cruz & Chouet, 48 1997), Asama (Fujita & Ida, 2003), Kusatsi-Shirane (Kumagai et al., 2003; Nakano & 49 Kumagai, 2005), Etna (Lokmer et al., 2008), and Erebus (Aster, 2019). Crack waves (and 50

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their induced tube waves in wellbores) are used for diagnosing the fracture geometries 51 in unconventional hydrocarbon reservoirs (Henry et al., 2002; Tary et al., 2014; Lipovsky 52 & Dunham, 2015; Liang et al., 2017). The resonating or humming signals in glaciers have 53 also been attributed to crack waves (Métaxian et al., 2003; Stuart et al., 2005; Gräff et 54 al., 2019; McQuillan & Karlstrom, 2021). Natural cracks in the subsurface are complex 55 in shape and usually form an inter-connected network. Therefore, efficient methods for 56 computing resonant modes of single cracks and networks of cracks are necessary for in-57 terpreting frequencies measured in the field. 58

Since its first discovery by Krauklis (1962), crack waves have been studied analyt-59 ically (Aki et al., 1977; Ferrazzini & Aki, 1987; Korneev, 2008; Lipovsky & Dunham, 2015), 60 experimentally (Tang & Cheng, 1988; Nakagawa et al., 2016; Cao et al., 2021), and nu-61 merically by various methods (e.g., B. Chouet, 1986; Yamamoto & Kawakatsu, 2008; 62 Frehner & Schmalholz, 2010; O'Reilly et al., 2017; Liang et al., 2020; Shauer et al., 2021; 63 Jin et al., 2022). Analytically derived dispersion relations are useful for understanding 64 the propagation behavior but are meant for an infinitely long crack and do not account 65 for the restriction of the finite crack tip. The finite difference method (FDM) is normally 66 based on cartesian grids in 2D (Fehler & Aki, 1978) or 3D (B. Chouet, 1986; Liang et 67 al., 2020) and limited to a tabular crack shape. Maeda and Kumagai (2013) and Maeda 68 and Kumagai (2017) performed a large number of numerical simulations on rectangu-69 lar cracks using a FDM simulator developed by B. Chouet (1986). With that, they ob-70 tained a set of empirical fitting formulas for resonant frequencies given the crack aspect 71 ratio α and stiffness ratio $C_L = K_f L/(Gw_0)$, where K_f is the fluid bulk modulus, G 72 the solid shear modulus, L the crack length and w_0 the crack aperture. However, such 73 relations only apply to longitudinal or transverse modes on rectangular cracks (Maeda 74 & Kumagai, 2013, 2017). Notably, O'Reilly et al. (2017) simulated a non-planar fluid-75 filled crack using FDM on a curvilinear grid and adopted a lubrication-type approxima-76 tion in the fluid (Lipovsky & Dunham, 2015), neglecting fluid acoustics in the crack width 77 direction while resolving the narrow viscous boundary layer close to the crack wall. This 78 treatment removes the time step restriction introduced by extremely fine mesh size in 79 the crack width direction and accelerates the computation. However, their work was lim-80 ited to 2D geometries. The finite element method (FEM) is more flexible for handling 81 complex crack geometries and has been used to study crack waves in 2D (Frehner & Schmal-82 holz, 2010; Frehner, 2013) and 3D (Shauer et al., 2021). Particularly, Shauer et al. (2021) 83

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produced the first simulation of an elliptical fluid-filled crack using the generalized finite 84 element method (GFEM). However, similar to FDM, FEM needs to discretize the vol-85 ume, which results in a large number of elements and high computational cost. On the 86 other hand, the boundary element method (BEM) reduces the simulation space from a 87 domain to boundary surfaces, drastically decreasing the number of degrees of freedom, 88 and has been used to study waves in fluid-filled cracks (Yamamoto & Kawakatsu, 2008; 89 Pointer et al., 1998; Jin et al., 2022) and other inclusions (Zheng et al., 2016; Sun et al., 90 2020). However, previous BEM simulations are either in two dimensions or focus on the 91 wave diffraction instead of analyzing the resonant frequencies. Currently, the study of 92 resonant frequencies of complex-shaped fluid-filled cracks and crack networks in three 93 dimensions remain unknown. 94

In this work, we propose an efficient hybrid numerical method to simulate crack 95 wave resonance in complex-shaped cracks or crack networks filled with an inviscid fluid, 96 by coupling the boundary element method (BEM) for the solid response and the finite 97 volume method (FVM) for acoustics in the fluid. By using triangular elements in both 98 BEM and FVM on the crack surfaces, we successfully handle complex crack shapes and 99 intersections. We restrict our attention to the low frequency limit where the crack wave 100 is much slower than the solid body waves, such that the solid response can be approx-101 imated as quasi-static (Korneev, 2008; Lipovsky & Dunham, 2015; Liang et al., 2020). 102 An eigenvalue analysis is performed to extract the resonant modes directly in the fre-103 quency domain, circumventing errors from time discretization and spectral analysis of 104 the time domain simulation data. We first verify our method by comparing results with 105 analytical solutions in the rigid wall limit and with numerical solutions from existing meth-106 ods for both a rectangular (B. Chouet, 1986; Maeda & Kumagai, 2017) and elliptical cracks 107 (Shauer et al., 2021). An example is then provided to demonstrate the simulation ca-108 pability for intersecting cracks. The effect of crack aspect ratio and stiffness ratio on res-109 onant frequencies (longitudinal, transverse, and mixed modes) is systematically inves-110 tigated for both rectangular and elliptical cracks. Although our current model does not 111 include viscous or radiation loss, we provide some theoretical discussion on these effects 112 under simple assumptions (boundary layer limit and quasi-dynamic approximation). Fi-113 nally, we present a crack shape compatible with the first two spectral peaks of VLP seis-114 mic signals from the Fani Maoré, Mayotte submarine volcano and discuss the potential 115 of the methodology for future applications in volcanology and other geological settings. 116

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Figure 1. Schematics of an arbitrarily-shaped fluid-filled crack, its spatial discretization (with unknown variables placed in the element centroids, red dots), and a zoom-in view at an intersection between two cracks.

117 2 Methods

In this section, we present the governing equations, discretization, and eigenmode analysis for computing the resonant frequencies.

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2.1 Governing equations

We consider the oscillation of inviscid fluids in complex-shaped thin cracks embed-121 ded in a 3D homogeneous linear elastic solid (Figure 1). The initial opening of the crack 122 is w_0 , which is assumed to be a constant and much smaller than the wavelength λ . We 123 adopt a similar lubrication approximation as B. Chouet (1986), Yamamoto and Kawakatsu 124 (2008) and O'Reilly et al. (2017), and treat the fluid pressure and velocities as uniform 125 in the crack thickness direction, reducing the crack from a 3D body to a 2D surface S. 126 Following O'Reilly et al. (2017), we consider a small crack curvature so that its effect 127 on the fluid momentum balance is negligible. Thus, the mass and momentum balance 128 of the fluid on the crack surface are written as 129

$$\frac{1}{w_0}\frac{\partial w}{\partial t} + \frac{1}{K_f}\frac{\partial p}{\partial t} + \frac{\partial v_\xi}{\partial \xi} + \frac{\partial v_\eta}{\partial \eta} = 0, \tag{1}$$

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$$\rho_f \frac{\partial v_\xi}{\partial t} + \frac{\partial p}{\partial \xi} = 0, \tag{2}$$

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$$\rho_f \frac{\partial v_\eta}{\partial t} + \frac{\partial p}{\partial \eta} = 0, \tag{3}$$

where ρ_f and K_f are fluid density and bulk modulus, w is the crack opening perturbation, p is the fluid pressure perturbation, t is time, and ξ and η are two locally perpendicular coordinates tangent to the crack surface, v_{ξ} and v_{η} are the fluid velocities in the ξ and η directions, respectively. Eliminating v_{ξ} and v_{η} in (1) using (2)-(3), we have

$$\rho_f \left(\frac{1}{w_0} \frac{\partial^2 w}{\partial t^2} + \frac{1}{K_f} \frac{\partial^2 p}{\partial t^2} \right) - \Delta p = 0, \tag{4}$$

where $\Delta = \frac{\partial^2}{\xi^2} + \frac{\partial^2}{\eta^2}$ is the tangential Laplace operator along the crack surface. The coupling between fluid and solid is encapsulated in the relation between the crack opening perturbation w and pressure perturbation p, which must balance the solid normal stress perturbation σ_n on the crack wall (assumed positive in compression). Since we focus on the low frequency limit, the solid response is approximately quasi-static (Korneev, 2008; Lipovsky & Dunham, 2015; Liang et al., 2020), and p for a linear elastic solid can be expressed as (Segall, 2010):

$$p(x) = \int_{S} K(x,\xi) w(\xi) \, dA, \tag{5}$$

where $K(x,\xi)$ is the Green's function that relates a unit open dislocation impulse at ξ to the normal stress change at x. The expressions of K in an elastic whole space and half space are available analytically for a uniform dislocation on both rectangular elements (Okada, 1985, 1992) and triangular elements (Nikkhoo & Walter, 2015).

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2.2 Discretization

We discretize the crack surfaces into N_e triangular elements. The unknown average pressures $\bar{\mathbf{p}}$ and openings $\bar{\mathbf{w}}$, placed at element centroids (as shown in Figure 1), are related by

$$\bar{\mathbf{p}} = \mathbf{K}\bar{\mathbf{w}},\tag{6}$$

where **K** is a N_e by N_e matrix and K(i, j) denotes the fluid pressure (or solid normal stress) change at the centroid of the *i*-th element caused by a unit open dislocation on the *j*-th element. We use the full space Green's function in this study but one can also use the half space solution.

We then discretize the tangential Laplacian operator by a finite volume scheme with a two-point flux (TPF) approximation following Karimi-Fard et al. (2004), which has been widely used for diffusive flows through a discrete fracture network in hydrocarbon reservoirs (e.g., Li & Lee, 2008; Moinfar et al., 2013; Xu et al., 2017; Berre et al., 2019). This scheme is only first-order accurate and is thus rarely used in wave propagation problems due to the strong numerical diffusion in time domain simulations (e.g., Durran, 2013). However, it is a sufficient scheme for our problem as we focus on resolving only the spa-

tial distribution of eigenmodes in the frequency domain and the low order of accuracy

can be remedied by using more elements. Here, we briefly present the key derivation steps

and the readers are referred to Karimi-Fard et al. (2004) for a detailed description.

We consider an arbitrary planar triangular element i with a surface S_i and boundary edges l_{ij} , where j is the index of the neighboring elements. Each i and j pair forms a hydraulic connection. When multiple cracks intersect, multiple connections share the same edge. We integrate equation (4) over each element i's surface, leading to:

$$\rho_f A_i \left[\frac{1}{w_0} \frac{\partial^2 \bar{w}_i}{\partial t^2} + \frac{1}{K_f} \frac{\partial^2 \bar{p}_i}{\partial t^2} \right] = \int_{S_i} \Delta p ds, \tag{7}$$

169 where

$$\bar{p}_i = \frac{1}{A_i} \int_{S_i} p ds,\tag{8}$$

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$$\bar{w}_i = \frac{1}{A_i} \int_{S_i} w ds,\tag{9}$$

are the average pressure and opening of element i, respectively. Applying the divergence

theorem to the right hand side of equation (7), we have:

$$\int_{S_i} \Delta p ds = \int_{S_i} \vec{\nabla} \cdot \vec{\nabla} p ds = \int_l \frac{\partial p}{\partial n} dl = -\sum_{j=1}^{n_c} D_{i \to j} Q_{i \to j}, \tag{10}$$

where $\partial p/\partial n$ is the pressure gradient normal to the boundary edges, n_c is the total number of connections in contact with element i, $Q_{i \to j}$ is the flux going out from element i to element j. Since $Q_{i \to j} = -Q_{j \to i}$, we only store $Q_{i \to j}$ for each (i, j) pair and its positive flux direction is pre-defined by an indicator function $I_{i \to j} = -I_{j \to i} = 1$. $D_{i \to j}$ is the discrete divergence operator and $D_{i \to j} = I_{i \to j} = 1$.

The assumption of the TPF scheme is to approximate the flux term in the following form (equation (7) in Karimi-Fard et al. (2004)):

$$Q_{i \to j} = I_{i \to j} T_{ij} (p_i - p_j), \tag{11}$$

where p_i and p_j are pressures defined at the centroids of the two neighboring elements.

 T_{ij} is the scalar transmissibility and is expressed as

$$T_{ij} = \frac{\alpha_i \alpha_j}{\sum_{k=1}^{n_c} \alpha_k},\tag{12}$$

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$$\alpha_k = \frac{l_{ij}}{d_k} \vec{n}_k \cdot \vec{f}_k, \tag{13}$$

where l_{ij} is the length of the connecting edge, d_k and \vec{f}_k are the length and unit directional vector from midpoint of the edge to the centroid of element k, \vec{n}_k is a unit normal vector perpendicular to the edge and pointing towards element k, as shown in Figure 1. Fluxes on the crack boundaries are set to zero. Combining equations (10) and (11), we have:

$$\int_{S_i} \Delta p ds = -\sum_{j=1}^{n_c} D_{i \to j} I_{i \to j} T_{ij} \left(p_i - p_j \right). \tag{14}$$

It is apparent that changing the positive flux direction from $i \to j$ to $j \to i$ flips the sign of both $D_{i\to j}$ and $I_{i\to j}$ and thus results in the same Laplacian term. Substituting equation (14) into equation (7) and rewriting in the matrix form, we have the spatially discretized equation without external forcing:

$$\rho_f\left(\frac{1}{w_0}\mathbf{K}^{-1} + \frac{1}{K_f}\right)\frac{\partial^2 \bar{\mathbf{p}}}{\partial t^2} = -\mathbf{A}^{-1}\mathbf{D}\mathbf{Q} = -\mathbf{A}^{-1}\mathbf{D}\mathbf{T}\bar{\mathbf{p}},\tag{15}$$

where **A** is a diagonal matrix of size N_e by N_e denoting the area of each element, $\mathbf{Q} = \mathbf{T}\mathbf{\bar{p}}$ is the flux vector whose size is the total number of connections N_c , **T** is the transmissibility matrix (including the indicator function) of size N_c by N_e that maps the vector $\mathbf{\bar{p}}$ to \mathbf{Q} , and \mathbf{D} is the divergence matrix of size N_e by N_c that maps \mathbf{Q} to the net flux out of each element. The structure of matrices \mathbf{D} and \mathbf{T} for a system of three intersecting crack elements are described in Appendix A.

We further introduce the following dimensionless quantities:

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$$\mathbf{K}^* = \mathbf{K}/(G/L), \mathbf{A}^* = \mathbf{A}/L^2, w = w^*/w_0, t^* = t/(L/c_l), \bar{\mathbf{p}}^* = \bar{\mathbf{p}}/(\rho_f c_l^2),$$
(16)

where G is the solid shear modulus, L is a representative length of the crack and $c_l = \sqrt{Gw_0/(\rho_f L)}$ is a representative crack wave speed. Different non-dimensionalization strategies exist, such as the one by B. Chouet (1986) which normalizes wave speeds by the solid compressional wave speed c_p . We choose c_l instead, because in the long wavelength limit, where compliance of the crack dominates, this choice conveniently results in a fundamental frequency of the order of unity. The nondimensionalised equation is

$$\left(\frac{1}{C_L}\mathbf{I} + (\mathbf{K}^*)^{-1}\right)\frac{\partial^2 \bar{\mathbf{p}}^*}{\partial t^{*2}} = -\mathbf{A}^{*-1}\mathbf{D}\mathbf{T}\bar{\mathbf{p}}^*,\tag{17}$$

where $C_L = K_f L/Gw_0$ is the key dimensionless parameter, named crack stiffness ra-

tio by B. Chouet (1986). The crack wave limit is achieved with $C_L \gg 1$, where the crack

 $_{207}$ is much more compliant than the fluid. C_L can be related to the representative crack

wave speed c_l by $C_L = c_f^2/c_l^2$, where c_f is the fluid acoustic wave speed. The crack topol-

 $_{209}$ ogy (for instance, the aspect ratio α for a rectangular or elliptical crack) and solid Pois-

son's ratio ν_s are encapsulated into the dimensionless stiffness matrix \mathbf{K}^* . The solid Pois-

son's ratio is set to 0.25 throughout this manuscript, unless otherwise mentioned.

2.3 Eigenmode analysis

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We directly obtain the resonant frequencies through eigenmode analysis in the frequency domain. The spatially discretized dimensionless equation is written as

$$\frac{\partial^2 \bar{\mathbf{p}}^*}{\partial t^{*2}} = -\mathbf{B} \bar{\mathbf{p}}^*,\tag{18}$$

215 where

$$\mathbf{B} = \left(\frac{1}{C_L}\mathbf{I} + (\mathbf{K}^*)^{-1}\right)^{-1} \mathbf{A}^{*-1} \mathbf{D} \mathbf{T}.$$
(19)

²¹⁶ The nondimensionalised Fourier transform is defined as

$$\hat{u}(\omega^*) = \int_{-\infty}^{+\infty} u(t^*) e^{i\omega^* t^*} dt^*,$$
(20)

217 where

$$\omega^* = \omega/\left(c_l/L\right),\tag{21}$$

218 is the dimensionless angular frequency. The dimensionless frequency is

$$f^* = \omega^* / (2\pi) = f / (c_l / L).$$
 (22)

Taking the Fourier transform of equation (18), we have:

$$(\omega^*)^2 \hat{\mathbf{p}} = \mathbf{B} \hat{\mathbf{p}},\tag{23}$$

where $(\omega^*)^2$ and \hat{p} are the eigenvalues and eigenvectors of the real matrix **B**. Since we deal with inviscid fluids, we only seek real positive eigenvalues, which correspond to undamped oscillatory modes. The resulting eigenvectors determine the spatial distribution of the pressure on the crack surface. Solving the resonant frequencies in dimensionless form is advantageous, because one can easily scale the solution to other parameters, such as crack length, crack width and solid stiffness, given the same dimensionless parameters, C_L , ν_s and crack topology.

²²⁷ **3** Verification and examples

In this section, we first verify our implementation by comparing our results to analytical solutions in the rigid solid limit and numerical solutions from existing studies.

- ²³⁰ We then present an example of simple intersecting crack geometry to demonstrate the
- utility of our method.

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Mode	Theoretical value	Numerical value	Error $(\%)$
1	0.5000	0.5004	0.074
2	0.8333	0.8325	0.105
3	0.9718	0.9706	0.127
4	1.0000	1.0008	0.078
5	1.3017	1.3018	0.009
6	1.5000	1.5010	0.066
7	1.6667	1.6647	0.118
8	1.7159	1.7172	0.072
9	1.7401	1.7370	0.175
10	1.9437	1.9413	0.122
11	2.0000	2.0014	0.068
12	2.1667	2.1693	0.123
13	2.2423	2.2409	0.060
14	2.5000	2.4954	0.184
15	2.5000	2.5010	0.042
16	2.5495	2.5426	0.270

 Table 1. The error between the theoretical and numerical resonant frequencies for the first 16 modes

3.1 Comparison with analytical solutions in a rigid solid

We compute the \mathbf{K} matrix using the subroutines developed by Nikkhoo and Wal-233 ter (2015), which have been extensively used by other studies. The bulk part that needs 234 to be validated is the FVM discretization of the Laplacian term. For that, we set solid 235 rigidity to infinity and compare the numerical results to the analytical solution of the 236 resonant frequencies of linear acoustic waves in a 2D rectangular domain with zero-flux 237 boundaries (Rona, 2007). The solution is in a dimensionless form with a rectangular do-238 main of size 1 by 0.5 and a wave speed of 1. The comparison results for the first 16 modes 239 are tabulated in Table 1. The excellent agreement between our numerical results and the 240

analytical solutions, with relative differences smaller than 0.2%, verifies our FVM dis-

²⁴² cretization of the Laplacian term.

243

3.2 Comparison to numerical solutions by existing studies

We compare solutions by our method (BEM+FVM) to those by B. Chouet (1986), 244 Maeda and Kumagai (2017) and Shauer et al. (2021). With B. Chouet (1986) and Maeda 245 and Kumagai (2017), we compare resonant frequencies of longitudinal modes for a rect-246 angular crack for various values of C_L (5, 15, 25, 50, 75, 100). With the GFEM by Shauer 247 et al. (2021), we compare solutions of multiple modes on both rectangular and ellipti-248 cal cracks. The eigenmodes can be straightforwardly classified as longitudinal (variation 249 only along the major crack axis), transverse (variation only along the minor crack axis), 250 and mixed modes for a rectangular crack, but less so for an elliptical crack. Since the 251 method by Shauer et al. (2021) discretizes the problem in time and, therefore, does not 252 readily provide resonant frequencies, we ran their code to excite the fluid oscillation on 253 the crack with $C_L = 100$ by a point injection source and then extract the resonant fre-254 quencies from the spectral peaks of the pressure records at a few receiving points. We 255 use a Gaussian time function for the injection source $f(t) = \exp\left(-(t-t_c)^2/T^2\right)$ with 256 $t_c = 0.5, T = 0.1$, to ensure a smooth start and a sufficiently wide spectrum to cover 257 enough eigenmodes. Note that if either excitation or receiving points are placed on the 258 nodal line, the eigenmode can not be excited or recorded. Therefore, not all eigenmodes 259 are excited in the time domain simulation and we also only compare selective modes with 260 Shauer et al. (2021), which is sufficient for verification purposes. The detailed geome-261 tries and simulation data are presented in Appendix B. Notably, the code of Shauer et 262 al. (2021) has the capability of both considering (fully dynamic, FD) or neglecting the 263 solid inertia (quasi-static, QS), allowing to investigate the impact of the solid inertia on 264 crack wave resonant frequencies. 265

Tables 2 and 3 show the comparison of dimensionless resonant frequencies of seletive eigenmodes from the GFEM program by Shauer et al. (2021) with those by our method for a rectangular and elliptical crack, respectively, with an aspect ratio of 0.5, major axis length of 1, and C_L of 100. The relative difference between our results and those from Shauer et al. (2021) are near 2% or less, with or without solid inertia. This close agreement not only demonstrates the validity of our approach but also reassures that the quasi-static solid response is a very good approximation when computing the

	Rosona	nt froque	meios dot	octod at	rocoiving	points by CEEM	BEM+EVM	Error FD	Error QS
Mode	nesona	nt neque	incles det	ecteu at	receiving	, points by GFEM		(%)	(%)
	(-0.5	(5, 0)	(-0.2,	0.25)		(0, 0.25)			
	FD	QS	FD	QS	FD	QS			
1	1.236	1.236	1.236	1.236			1.210	2.15	2.15
2	2.727	2.691					2.662	2.44	1.09
3			2.890	2.873	2.818	2.782	2.835	1.94	1.34
4			3.453	3.418			3.373	2.37	1.33
5	4.454	4.400	4.436	4.382			4.385	1.57	0.34
6					4.526	4.491	4.466	1.34	0.56
7			6.035	5.964			5.980	0.92	0.27
8	6.399	6.273			6.344	6.236	6.330	1.09	0.90

 Table 2.
 Resonant frequencies by BEM+FVM and GFEM with or without inertia (FD or QS),

 rectangular crack

The bold values are used for error calculation. We use a of CL=100 and aspect ratio of 0.5 Mode eigenfunctions are shown in Figure 2.

 Table 3.
 Resonant frequencies by BEM+FVM and GFEM without inertia (QS), elliptical crack

M.J.	The resor	ant frequencies can be	detected at detection points		E (07)
Mode	(-0.5, 0)	(0, 0.25)	(0, 0)	BEM+FVM	Error (%)
1	1.527			1.518	0.59
2	3.027		3.090	3.050	0.75
3		3.290		3.241	1.51
5	4.890			4.816	1.54
7	6.890	6.853	6.944	6.771	1.76
8		7.308	7.235	7.107	1.80

The bold values are used for error calculation. We use a of CL=100 and aspect ratio of 0.5 Mode eigenfunctions are shown in Figure 3.

crack wave resonant frequencies, at least for a C_L of 100. A similar conclusion has also 273 been reached by Shauer et al. (2021). Since we assume a quasi-static solid response, it 274 is reasonable that our results have a better agreement to those by GFEM without in-275 ertia. 276



Figure 2. Dimensionless frequencies and eigenfunctions of the first 16 resonant modes (numbered in an ascending order in frequencies) of a rectangular crack with $C_L=100$ and aspect ratio of 0.5 calculated by BEM+FVM. The errors of selective resonant frequencies between the BEM+FVM and GFEM without inertia are shown in Table 2. The white color indicates the nodal lines.

The pressure eigenfunctions of the first 16 resonant modes are displayed in Figure 2 for a rectangular crack and Figure 3 for an elliptical crack, showing a rich spectrum of spatial variations including longitudinal, transverse, and mixed modes. Different modes can produce different near and far field radiation patterns, that may be detectable in real 280 seismic data (e.g., Liang et al., 2020).

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The dimensionless frequencies of the first 9 longitudinal modes for rectangular cracks 282 by various methods with different crack stiffness ratios are shown in Figure 4. The re-283 sults of Shauer et al. (2021) are only computed for a C_L of 100. For ease of comparison, 284 we convert dimensionless frequencies f^* in our studies to those in B. Chouet (1986) f_C^* , 285



Figure 3. Same as Figure 2 but for an elliptical crack.

which are related by $f_C^* = f^* c_l / c_P$. Overall, our results match well with those by Shauer 286 et al. (2021) (relative error < 3%) and also qualitatively well with those by B. Chouet 287 (1986) and Maeda and Kumagai (2017). However, there are quantitative discrepancies 288 between our results and those by B. Chouet (1986) (relative error 8.83-23.43%) and Maeda 289 and Kumagai (2017) (relative error 2.72-16.63%, see the supporting information for tab-290 ulated errors). Particularly, both B. Chouet (1986) and Maeda and Kumagai (2017) sys-291 tematically give lower frequencies than those by our method and Shauer et al. (2021). 292 We suspect these discrepancies are likely due to differences in spatial and temporal sam-293 pling, or domain sizes used in the FDM code in B. Chouet (1986) and Maeda and Ku-294 magai (2017). Particularly, a truncated domain in the FDM results in a more compli-295 ant solid response (Korneev et al., 2014), which in turn results in a lower crack wave speed 296 and resonant frequencies. Our method uses boundary elements and thus an infinite do-297 main is directly satisfied. For this reason, when comparing results with Shauer et al. (2021), 298 we deliberately used a very large domain (10 times the length of the crack) in the GFEM 299 code to minimize its boundary effect using an unstructured grid, coarsening in regions 300 far from the crack. 301



Figure 4. Dimensionless frequencies f_C^* of longitudinal modes for rectangular cracks with different C_L (5, 15, 25, 50, 75, 100) by various methods, and a zoom-in view of the case $C_L = 100$ on the right panel. Results by B. Chouet (1986) and Maeda and Kumagai (2017) are slightly shifted in the horizontal axis to avoid overcrowding the figure.

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3.3 An example of intersecting cracks

We now apply our method to one example of intersecting cracks, one full ellipse 303 with a half-elliptical branch, and obtain the first 16 eigenmodes, shown in Figure 5. In-304 teractions between multiple cracks result in more complex resonant modes than in sin-305 gle cracks (shown in Figures 3 and 4). For example, the fundamental mode now involves 306 fluid exchange between the major crack and the branch, and has a lower frequency than 307 the fundamental mode of the major crack (the second mode in this case). When nodal 308 lines coincide with the intersecting edge, resonances can be isolated on the major crack, 309 such as modes 2, 7, 8, 13 and 16. Temporal manifestation of these modes requires a more 310 peculiar condition: the excitation must not be located in the branch. One can certainly 311 add more complexities in the crack network, such as asymmetries, non-planarity or more 312 intricate coupling, and expect to encounter richer eigenmodes. However, such modeling 313 only becomes meaningful when more compelling observations exist and require. We will 314 demonstrate later how a particular crack shape can explain the ratio of the first two spec-315 tral peaks in the VLP seismic data at the Fani Maoré, Mayotte submarine volcano. Ex-316

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- cept for that, we decide to leave the analysis of eigenmodes of a more complex crack net-
- ³¹⁸ work for future investigation.



Figure 5. The first 16 eigenmodes of a simple two-intersecting-cracks geometry: a half ellipse intersecting a full ellipse (aspect ratio 0.5) along its minor axis. The major axis length of the full elliptical crack is chosen as L for the non-dimensionalisation.

³¹⁹ 4 Effect of aspect ratio α and crack stiffness ratio C_L

In this section, we present the effect of α and C_L on the resonant frequencies of rectangular and elliptical cracks, with major and minor axes in the x- and y-directions, respectively. Maeda and Kumagai (2017) presented a similar analysis for rectangular cracks, but only on longitudinal and transverse modes. Here, we include the mixed modes and the results for elliptical cracks. We fixed $C_L = 100$ when varying α (from 0.05 to 1.00 with an increment of 0.05) and fix $\alpha = 0.5$ when varying C_L (from 5 to 100 with an increment of 5). The frequencies of the first 16 eigenmodes are tabulated in the Support-

ing information. Here, we select 9 representative modes and visualize them in Figures 327 6-9. For rectangular cracks, we associate to each mode a pair of numbers (i, j) that de-328 note the number of half wavelengths in the x- and y-directions. For instance, the fun-329 damental mode (1,0) is a longitudinal mode with one half wavelength pressure variation 330 in the x-direction and quasi-uniform in the y-direction. Such numbering becomes less 331 obvious for elliptical cracks, especially when the aspect ratio approaches 1, for which the 332 eigenfunctions are better characterized by radial and circumferential variations. Nonethe-333 less, for the ease of comparing results with rectangular cracks, we still number the rep-334 resentative modes in Figures 7 and 9 approximately into longitudinal, transverse, and 335 mixed modes. 336



Figure 6. Dimensionless resonant frequencies of representative modes of rectangular cracks as a function of the aspect ratio α . C_L is fixed to 100. The eigenfunctions displayed are for an aspect ratio of 0.55. Certain high order mixed and transverse modes rank outside of the first 16 eigenmodes that we store, which causes the apparent absence of data at low aspect ratios.

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4.1 Effect of aspect ratio

The variation of resonant frequencies with aspect ratio is shown in Figure 6 and 7 for rectangular and elliptical cracks, respectively. For both cases, decreasing the aspect ratio increases the crack stiffness from the transverse direction and results in higher resonant frequencies for all the modes. This effect is relatively mild for longitudinal modes but rather steep for transverse and mixed modes. For instance, when α of a rectangu-



Figure 7. Same as Figure 6 but for elliptical cracks. Note that the mode numbers (i, j) are not strictly valid for an elliptical crack but are useful for our interpretation (more explanation in the main text).

lar crack decreases from 1.0 to 0.1, the frequency of the fundamental mode (1,0) increases by ~ 1.8 fold while the frequency of the mixed mode (1,1) increases by ~ 14.6 fold. As a result, the first few resonant modes are predominantly longitudinal for both rectangular and elliptical cracks at low aspect ratios (below 0.2). For a similar mode, the resonant frequency of an elliptical crack is consistently higher than that of the rectangular crack. This is expected as the elliptical crack is narrower in the transverse direction and thus stiffer than a rectangular crack of the same length and aspect ratio.

Another clear feature, for both rectangular and elliptical cracks, is that frequen-350 cies of modes with same wavelengths in the transverse direction converge as α decreases. 351 For instance, frequencies of mixed modes (1, 1) and (2, 1) converge to the values of trans-352 verse mode (0, 1). Similar convergence also exists for modes (0, 2) and (1, 2). This is 353 expected because the crack wave speed, in the limit of low aspect ratio, is primarily con-354 trolled by the short wavelength in the transverse direction. As α increases, the frequen-355 cies of different modes become more intermingled and mode degeneration occurs, where 356 modes with distinct eigenfunctions share the same frequency. It is well known that mode 357 degeneration occurs at $\alpha = 1$ due to the geometric symmetry of a square or circle. What 358 we show here is that mode degeneration also occurs at intermediate aspect ratios. For 359

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Figure 8. Resonant frequencies of representative modes of rectangular cracks as a function of the crack stiffness C_L . The aspect ratio α is set as 0.5



Figure 9. Same as Figure 8 but for elliptical cracks

instance, modes (3, 0) and (0, 1) for both a rectangular and elliptical crack share similar frequencies when $\alpha \approx 0.35$.

362 4.2 Effect of crack stiffness ratio

Since the normalization constant $(c_f/\sqrt{C_L})/L$ for frequency changes with C_L , we visualize the actual resonant frequency f, instead of f^* . We use $c_f = 1$ m/s and L =1 m to scale f^* to f. C_L is the key dimensionless parameter that controls the crack wave propagation: the higher the value of C_L , the lower the phase velocity (e.g., B. Chouet,

³⁶⁷ 1986; Maeda & Kumagai, 2017). As a result, the resonant frequencies of all modes for

both rectangular (Figure 8) and elliptical cracks (Figure 9) decrease continuously as C_L

increases. Again, for a similar mode, the resonant frequencies of an elliptical crack is con-

sistently higher than those of a rectangular crack given the same axial lengths.

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5 Energy dissipation

Since we currently focus on computing crack resonant frequencies in complex crack geometries, we assume both an inviscid fluid and a quasi-static solid and we do not consider energy dissipation, from either fluid viscosity or seismic radiation. When damping exists, the resonant frequency becomes complex and the rate of decay is quantified by the quality factor

$$Q = \frac{\operatorname{Re}\left(f\right)}{2\operatorname{Im}\left(f\right)},\tag{24}$$

which is also the number of cycles for an oscillation's energy to fall off to $e^{-2\pi} \sim 0.2\%$ 377 of its original value. The effect of viscous damping has been investigated analytically with 378 fully dynamic (Korneev, 2008) and quasi-static solid response (Lipovsky & Dunham, 2015) 379 on an infinite crack. However, the applicability of the Q formula on a finite crack has 380 not yet been tested. In addition, the convoluted derivation in Korneev (2008) makes it 381 difficult to quantify the relative contribution of different dissipative sources to the to-382 tal energy loss. On the other hand, numerical studies on rectangular cracks (e.g., Ku-383 magai & Chouet, 2000) have investigated the Q caused by seismic radiation but adopted 384 a simplistic treatment of the fluid viscosity, either an inviscid or fully-developed flow. In 385 this section, we offer a semi-analytical discussion of energy dissipation under a few as-386 sumptions and attempt to address two questions: (1) does the formula of Q developed 387 by Lipovsky and Dunham (2015) for an infinite crack also apply to a finite crack? (2)388 which of the two sources of energy dissipation, fluid viscosity and seismic radiation, is 389 more significant? 390

391

5.1 The applicability of Q formula from dispersion to a finite crack

We consider a viscous fluid with kinematic viscosity μ . To focus on the effect of the finite geometry, we compare analytical solutions by Lipovsky and Dunham (2015) to numerical solutions by Liang et al. (2020) for a rectangular crack, both of which as³⁹⁵ sume a quasi-static solid response. For simplicity, we focus on the boundary layer limit

$$\zeta = w_0 / \sqrt{4\nu/\omega} \gg 1, \tag{25}$$

where the crack aperture w_0 is much larger than the thickness of the viscous boundary layer $\sqrt{4v/\omega}$. In this limit, Q is high and ω can be well approximately by the inviscid solution. The analytical formula of Q for crack waves with real wavenumber is given by equation (80) in Lipovsky and Dunham (2015) and, after neglecting the small imaginary part of phase velocity when $\zeta \gg 1$, we have:

$$Q = \sqrt{2}\zeta.$$
 (26)

The hypothesis is that this expression for Q also holds, at least approximately for a fi-401 nite rectangular crack, regardless of its geometric shape, as long as w_0 , μ and ω are known. 402 We perform numerical simulations using the program by Liang et al. (2020), who em-403 ployed a finite difference method on a stretched grid to deal with the narrow viscous bound-404 ary layer. We set L = 100 m, $K_f = 1$ Pa, G = 1 Pa, $w_0 = 1$ m, which results in a C_L 405 of 100, and solve for the inviscid resonant angular frequencies ω of rectangular cracks 406 of two aspect ratios, 0.5 and 1.0. We then adjust μ so that ζ takes the values of 10, 20, 407 40, 60, 80, 100 and 200. We consider the first two modes of the crack with aspect ratio 408 of 0.5 and the fundamental mode of the square crack to represent different mode types 409 and crack shapes. The Q values of viscous cases are obtained using the methodology by 410 Liang et al. (2020) and the comparison to equation (26) is shown in Figure 10. 411

As shown in Figure 10, the prediction by the analytical formula in Lipovsky and 412 Dunham (2015) matches well the numerical solutions. The agreement gets better at large 413 ζ , where the assumption of boundary layer limit becomes more accurate. The differences 414 between the numerical and analytical solutions are less than 5% at Q > 40, while the 415 difference at Q = 10 is ~ 14%. Another encouraging finding is that aspect ratios and 416 mode numbers of rectangular cracks have a negligible impact on the value of Q as long 417 as ζ is the same. We thus postulate that the Q formula is likely to hold also for other 418 crack shapes or even a crack network. We further propose that one may first approxi-419 mate the resonant frequency ω of complex shaped cracks using the inviscid solution ef-420 ficiently determined by our method, and then directly estimate Q using the analytical 421 formula. However, future numerical studies considering both complex crack geometry 422 and fluid viscosity are necessary to rigorously test this hypothesis. 423

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Figure 10. Quality factor of various resonant modes of rectangular cracks as a function of the boundary layer thickness ratio ζ .

5.2 The competition between radiation and viscous damping

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Instead of considering the fully dynamic solid response (e.g., Korneev, 2008), we 425 assume a quasi-dynamic solid response (e.g., Rice, 1993; Geubelle & Rice, 1995), which 426 allows to explicit extract the instantaneous long wavelength emission perpendicular to 427 the crack surface, the radiation damping (RD) term. We also consider an infinite crack 428 in two dimensions for the ease of theoretical treatment following Lipovsky and Dunham 429 (2015). By neglecting the wave mediated stresses and the seismic diffraction at the fi-430 nite crack tips, the radiation we consider is an underestimate, but it is still useful for un-431 derstanding the relative importance of various dissipation sources. Since resonances tend 432 to be overdamped in the fully developed flow limit $\zeta \gg 40$ (Korneev, 2008; Lipovsky 433 & Dunham, 2015), we continue to focus on the boundary layer limit $\zeta \ll 40$. We ex-434 plicitly identify the radiation and viscous damping terms in the governing equation and 435 then compute the ratio of their magnitudes. 436

The width-averaged crack wave equation considering viscous wall traction is obtained by combining the mass and momentum balance equations in Lipovsky and Dunham (2015),

$$\frac{\rho_f}{K_f}\frac{\partial^2 p}{\partial t^2} + \frac{\rho_f}{w_0}\frac{\partial^2 w}{\partial t^2} - \frac{\partial^2 p}{\partial x^2} - \frac{2}{w_0}\frac{\partial \tau}{\partial x} = 0,$$
(27)

where τ is the wall shear traction. We introduce the double Fourier transform of an ar-

441 bitrary function F(x,t) as

$$\hat{F}(k,\omega) = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} F(x,t) e^{-i(kx-\omega t)} dt dx.$$
(28)

 A_{442} Applying it to equation (27) leading to

$$-\omega^2 \rho_f \left(\frac{\hat{p}}{K_f} + \frac{\hat{w}}{w_0}\right) + k^2 \hat{p} - ik \frac{2}{w_0} \hat{\tau} = 0.$$
⁽²⁹⁾

⁴⁴³ Using equation (38) in Lipovsky and Dunham (2015) and neglecting horizontal wall mo-

tion, the wall shear traction is related to fluid pressure by

$$\hat{\tau} = -ikw_0\Omega\hat{p}/2,\tag{30}$$

445 where

$$\Omega = \left(\sqrt{i}/\zeta\right) \tanh\left(\zeta/\sqrt{i}\right),\tag{31}$$

and tends to \sqrt{i}/ζ in the boundary layer limit. Therefore, the viscous damping (VD) term in the equation is

$$VD = k^2 \Omega \hat{p}. \tag{32}$$

Applying the quasi-dynamic solid response (Geubelle & Rice, 1995), the fluid pressure and crack opening are related by

$$\hat{p} = \frac{Gk\hat{w}}{2\left(1 - v_s\right)} - i\omega\eta_R\hat{w},\tag{33}$$

where the two terms on the right hand side are the quasi-static response and radiation 450 damping (RD), respectively, and $\eta_R = \rho_s c_p/2$ is the radiation damping coefficient. The 451 ratio between the QS and RD terms is approximately $c_s^2/(c_p c)$ (dropping terms involv-452 ing Poisson's ratio), where $c = \omega/k$ is the crack wave phase velocity. In the low-frequency 453 limit, which we are interested in, the crack wave speed is much smaller than the speeds 454 of the solid body waves, $c \ll c_s \sim c_p$, and thus the RD term is much smaller than the 455 QS term. Substituting equation (33) into (29) and approximating \hat{w}/\hat{p} using the QS part, 456 we obtain the RD term in equation (29) as 457

$$RD = i\omega^3 \frac{\rho_f n_R}{K_f} \hat{w} \approx i\omega^3 \frac{\rho_f n_R}{K_f} \frac{2(1-v_s)}{Gk} \hat{p}.$$
(34)

458 The ratio between RD and VD is

$$\frac{RD}{VD} = ic^3 \frac{\rho_f n_R}{K_f} \frac{2(1-\upsilon_s)}{G\Omega} = i \frac{c_p c^3}{c_f^2 c_s^2} \frac{(1-\upsilon_s)}{\Omega}.$$
(35)

459 In the boundary layer limit, the magnitude of this ratio becomes

$$\left|\frac{RD}{VD}\right| = \frac{c_p c^3}{c_f^2 c_s^2} \zeta \left(1 - v_s\right). \tag{36}$$

- 460 When the overall damping is small, the crack wave phase velocity as a function of wave-
- length λ is well approximated by the inviscid dispersion relation:

$$c = \sqrt{\frac{2\pi G w_0}{\lambda \rho_f \left(1 - \upsilon_s\right)}} \sim c_f / sqrt C_\lambda,\tag{37}$$

where $C_{\lambda} = K_f \lambda / G w_0$ is a crack stiffness ratio similar to C_L but replacing L by λ . Finally, we obtain RD/VD, which scales as

$$\left|\frac{RD}{VD}\right| \sim \frac{c_p c_f}{c_s^2} \frac{\zeta}{C_\lambda^{3/2}} = \frac{c_p c_s}{c_f^2} \zeta \left(\frac{w_0 \rho_s}{\lambda \rho_f}\right)^{3/2},\tag{38}$$

after dropping small constants such as 2, π and μ_s . |RD/VD| is governed by three di-464 mensionless parameters: $\frac{c_p c_f}{c_s^2}$, C_{λ} and ζ . The first parameter one is controlled by the body 465 wave speeds of the solid and fluid and is not related to the crack geometry. For a typ-466 ical crustal rock and liquid fluid, for instance with $c_f = 1500$ m/s, $c_p = 4500$ m/s, and 467 $c_s=2500~{\rm m/s},\,\frac{c_pc_f}{c_\circ^2}$ is near unity. However, exsolved gases in liquid fluid, common in 468 shallow volcanic or geothermal environments (e.g., Kumagai & Chouet, 1999, 2001), can 469 significantly decrease the sound speed of the mixture, resulting in a much smaller $\frac{c_p c_f}{c^2}$. 470 The trade-off between C_{λ} and ζ in controlling |RD/VD| is displayed in Figure 11. In 471 the regime of high ζ and low C_{λ} , seismic radiation dominates over viscous damping, while 472 in the regime of low ζ and high C_{λ} vice versa. Note that increasing λ or decreasing fre-473 quency ω while fixing other parameters increases C_{λ} and simultaneously decreases ζ , both 474 of which lead to a lower percentage of damping in radiation. 475

6 Application to VLP seismic signals during the Mayotte volcano-seismic crisis

Since 10 May 2018, an unprecedented submarine volcano-seismic crisis occurred 478 30 km east of Mayotte Island (France), featuring a lithosphere-scale dyke intrusion and 479 drainage ($\sim 5 \text{ km}^3$) of deep magma reservoirs and producing exceptionally deep seismic-480 ity and substantial surface deformation (Cesca, Letort, et al., 2020; Feuillet et al., 2021; 481 Saurel et al., 2021; Mittal et al., 2022; Mercury et al., 2022; Retailleau et al., 2022). By 482 mid June of 2018, sustained long duration and highly oscillatory VLP seismic signals (see 483 an example in Figure 12a) have been observed and persist since, which are associated 484 with resonances of magma-filled cracks excited by nearby volcano-tectonic (VT) events 485



Figure 11. |RD/VD| as a function of the crack stiffness ratio C_{λ} and the boundary layer thickness ratio ζ . Parameters used are $c_f = 1500$ m/s, $c_p = 4500$ m/s, and $c_s = 2500$ m/s.

or possible piston collapse movements (Cesca, Letort, et al., 2020; Feuillet et al., 2021). 486 The stack of spectra of multiple VLP events reveals multiple resonant modes, among which 487 the fundamental mode with period ~ 15.5 s is present in all events, but not all higher modes 488 are manifested in each event, probably due to differences in the excitation. The funda-489 mental frequency can be readily explained by the crack model upon choosing a proper 490 crack length and aperture (Cesca, Letort, et al., 2020). However, as shown in Figure 12b, 491 the uneven spacing between resonant modes implies additional complexity in the source. 492 Particularly, the ratio between the first higher mode and the fundamental mode is $f_2/f_1 \approx$ 493 2.5. As shown in Figures 6 and 7, this value can not be explained by a simple rectan-494 gular or elliptical crack. Here, we show this observation can be explained by a dumbbell-495 shaped crack (Figure 12c). This crack shape is compatible with the f_2/f_1 data, but might 496 still differ from the real crack geometry in Mayotte as we have not made a systematic 497 attempt to also match the frequencies of other higher modes. However, this example is 498 sufficient to demonstrate the potential application of the developed method. One pro-499 found question is perhaps whether one can reconstruct the topology of the crack given 500 the information of all the resonant frequencies. Mark Kac also asked a similar question 501 "Can one hear the shape of a drum?" (Kac, 1966). Unfortunately, the answer is nega-502 tive: there exist multiple isospectral geometries that share the same resonant frequen-503

cies, as mathematically proven by Gordon et al. (1992). However, these isospectral geometries are rare even though they do exist and one can still decipher the shape of the
resonator given additional constraints of the vibration pattern, which in practice requires
dense geophysical observation particularly in the near field. A formal inversion procedure would need to be developed in the future to find the optimal crack geometry or topology of interconnected crack networks that best explains all the observed resonant frequencies and other geophysical constraints.



Figure 12. (a) Normalized vertical acceleration waveform of an representative VLP event (on 11 November 2018, bandpass filtered to 0.02-0.1 Hz) at the nearest broadband seismic station YTMZ on land, during the volcano-seismic crisis near Mayotte. (b) Stacked spectrum of 21 strong VLP signals compiled by Cesca, Letort, et al. (2020), highlighting multiple unevenly spaced resonant modes (dashed lines). Particularly, the frequency ratio between the first two modes $f_2/f_1 \approx 2.5$. The blue ticks indicate the integer multiples of the fundamental frequency. (c) Eigenmodes of a possible crack shape that satisfies $f_2/f_1 \approx 2.5$.

511 7 Summary

We have developed a hybrid method that couples the boundary element and finite 512 volume method to efficiently compute the resonant modes of fluid-filled cracks with com-513 plex geometry. Particularly, the BEM reduces three dimensional cracks to 2D surfaces, 514 substantially decreasing the number of degrees of freedom. By performing eigenmode 515 analysis in the frequency domain, we avoid errors from both the time discretization and 516 spectral analysis of the time domain data. We solve the problem in dimensionless form 517 so that the results can be conveniently scaled to other crack sizes. After proper verifi-518 cation, we apply our method to an example of a crack network, revealing distinct res-519 onant frequencies and vibration patterns, which may be utilized to infer more accurately 520 crack shapes from seismic data. 521

We then systematically analyze the influence of crack aspect ratio and crack stiff-522 ness on the resonant frequencies for both rectangular and elliptical cracks, which are com-523 mon models for interpreting real data. In general, rectangular and elliptical cracks share 524 similar eigenmode types and frequencies, while the elliptical crack has slightly higher res-525 onance frequencies due to the reduced length of the minor axis. At a high aspect ratio, 526 the frequencies of various mode types (longitudinal, transverse and mixed) are intermin-527 gled and mode degeneration occurs. Reducing the aspect ratio increases the frequencies 528 of all the modes, but more intensely for transverse and mixed modes than for longitu-529 dinal modes. In addition, at low aspect ratio, frequencies of modes (transverse or mixed) 530 with the same wavelengths in the transverse direction converge and differentiating them 531 requires additional knowledge of their vibration patterns. On the other hand, increas-532 ing C_L results in a decrease in resonant frequencies for all modes, regardless of the crack 533 geometry, which is primarily due to the decrease in crack wave propagation speed. 534

The major part of this work does not consider fluid viscosity or seismic radiation, 535 and thus cannot be used to directly compute the quality factor Q. However, by making 536 a few assumptions, we offer additional theoretical discussion on the energy dissipation. 537 First, by comparing numerical to analytical solutions, we confirm that the simple for-538 mula $Q = \sqrt{2}\zeta$ derived by Lipovsky and Dunham (2015) is a rather good approxima-539 tion for a rectangular crack when the thickness of the viscous boundary layer is much 540 smaller than the crack width, regardless of crack aspect ratio or vibrational mode. This 541 is an encouraging finding that suggests one may first obtain the inviscid resonant fre-542

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quencies using our method and then apply analytical formula to compute Q. Note that 543 this formula still does not consider seismic radiation. We then derived the relative ra-544 tio of the radiation damping to viscous damping, assuming a quasi-dynamic solid response 545 on an infinite crack. We show that this ratio is primarily controlled by three dimension-546 less parameters: $c_p c_f / c_s^2$, C_λ and ζ . Particularly, in the limit of high ζ and low C_λ , seis-547 mic radiation dominates over viscous damping while the opposite is true in the limit of 548 low ζ and high C_{λ} . Note that the seismic radiation considered here is a lower bound as 549 we neglected the wave-mediated stresses and the seismic radiation at the finite crack tip. 550 However, our theoretical development still offers a valuable insight into the partition of 551 damping in crack waves. 552

Finally, we obtain one possible crack shape, a "dumbbell", that successfully explains 553 the ratio of frequencies of the first two modes in the VLP seismic data during the 2018 554 Fani Maoré, Mayotte submarine volcanic eruption. This shape is one possibility and may 555 be updated when additional higher modes and geophysical constraints are integrated into 556 the analysis. In addition, the method developed here can be directly applied to other 557 scenarios, such as unconventional oil and gas fields and glacier hydraulics. Future work 558 requires a rigorous treatment of fluid viscosity, elastodynamics, and coupling to other 559 geometries such as conduits and equidimensional chambers. 560

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9 Data Availability Statement

The source code and the input files associated with the simulation cases are included in the Zenodo data respository at Liang et al. (2023). The VLP catalog of the Mayotte crisis is provided by Cesca, Heimann, et al. (2020) and is freely available online.

571 **References**

572	Aki, K., Fehler, M., & Das, S. (1977). Source mechanism of volcanic tremor: Fluid-
573	driven crack models and their application to the 1963 Kīlauea eruption. $Jour$
574	nal of Volcanology and Geothermal Research, $2(3)$, 259–287. doi: 10.1016/0377
575	-0273(77)90003-8

- Aster, R. C. (2019). Interrogating a surging glacier with seismic interferometry. Geo *physical Research Letters*, 46(14), 8162–8165. doi: 10.1029/2019GL084286
- Berre, I., Doster, F., & Keilegavlen, E. (2019). Flow in fractured porous media:
 A review of conceptual models and discretization approaches. Transport in
 Porous Media, 130(1), 215–236. doi: 10.1007/s11242-018-1171-6
- Cao, H., Medici, E., & Askari, R. (2021). Physical modeling of fluid-filled fractures
 using the dynamic photoelasticity technique. *Geophysics*, 86(1), T33–T43. doi:
 10.1190/geo2020-0037.1
- ⁵⁸⁴ Cesca, S., Heimann, S., Letort, J., Razafindrakoto, H., Dahm, T., & Cotton, F.
- Seismic catalogues of the 2018–2019 volcano-seismic crisis offshore
 mayotte, comoro islands. v. 1.0 (october 2019). available at gfz data ser-
- ssr
 vices.
 Retrieved from https://doi.org/10.5880/GFZ.2.1.2019.004
 doi:

 ssr
 10.5880/GFZ.2.1.2019.004
 doi:
- Cesca, S., Letort, J., Razafindrakoto, H. N., Heimann, S., Rivalta, E., Isken, M. P.,
 ... others (2020). Drainage of a deep magma reservoir near mayotte inferred
 from seismicity and deformation. *Nature Geoscience*, 13(1), 87–93. doi:
 10.1038/s41561-019-0505-5
- ⁵⁹³ Chouet, B. (1986). Dynamics of a fluid-driven crack in three dimensions by the finite
 difference method. Journal of Geophysical Research: Solid Earth, 91 (B14),
 ⁵⁹⁵ 13967–13992. doi: 10.1029/JB091iB14p13967
- ⁵⁹⁶ Chouet, B. A., Page, R. A., Stephens, C. D., Lahr, J. C., & Power, J. A. (1994).
 ⁵⁹⁷ Precursory swarms of long-period events at redoubt volcano (1989–1990),
 ⁵⁹⁸ alaska: their origin and use as a forecasting tool. Journal of Volcanology and
- ⁵⁹⁹ Geothermal Research, 62(1-4), 95–135. doi: 10.1016/0377-0273(94)90030-2
- Cruz, F. G., & Chouet, B. A. (1997). Long-period events, the most characteristic
 seismicity accompanying the emplacement and extrusion of a lava dome in
 galeras volcano, colombia, in 1991. Journal of Volcanology and Geothermal
 Research, 77(1-4), 121–158. doi: 10.1016/S0377-0273(96)00091-1

604	Durran, D. R. (2013). Numerical methods for wave equations in geophysical fluid dy-
605	namics (Vol. 32). Springer Science & Business Media.
606	Fehler, M., & Aki, K. (1978). Numerical study of diffraction of plane elas-
607	tic waves by a finite crack with application to location of a magma lens.
608	$Bulletin \ of \ the \ Seismological \ Society \ of \ America, \ 68(3), \ 573-598. $ doi:
609	10.1785/BSSA0680030573
610	Ferrazzini, V., & Aki, K. (1987). Slow waves trapped in a fluid-filled infinite crack:
611	Implication for volcanic tremor. Journal of Geophysical Research: Solid Earth,
612	92(B9), 9215-9223.doi: 10.1029/JB092iB09p09215
613	Feuillet, N., Jorry, S., Crawford, W. C., Deplus, C., Thinon, I., Jacques, E.,
614	\dots others (2021). Birth of a large volcanic edifice offshore mayotte via
615	lithosphere-scale dyke intrusion. Nature Geoscience, $14(10)$, 787–795. doi:
616	10.1038/s41561-021-00809-x
617	Frehner, M. (2013). Krauklis wave initiation in fluid-filled fractures by a passing
618	body wave. In Poromechanics v: Proceedings of the fifth biot conference on
619	poromechanics (pp. 92–100). doi: 10.1061/9780784412992.011
620	Frehner, M., & Schmalholz, S. M. (2010). Finite-element simulations of stoneley
621	guided-wave reflection and scattering at the tips of fluid-filled fractures. Geo -
622	physics, 75(2), T23-T36. doi: 10.1190/1.3340361
623	Fujita, E., & Ida, Y. (2003). Geometrical effects and low-attenuation reso-
624	nance of volcanic fluid inclusions for the source mechanism of long-period
625	earthquakes. Journal of Geophysical Research: Solid Earth, $108(B2)$. doi:
626	10.1029/2002JB001806
627	Geubelle, P. H., & Rice, J. R. (1995). A spectral method for three-dimensional elas-
628	todynamic fracture problems. Journal of the Mechanics and Physics of Solids,
629	43(11), 1791-1824.doi: 10.1016/0022-5096(95)00043-I
630	Gordon, C., Webb, D. L., & Wolpert, S. (1992). One cannot hear the shape of a
631	drum. Bulletin of the American Mathematical Society, $27(1)$, 134–138. doi: 10
632	.1090/S0273-0979-1992-00289-6
633	Gräff, D., Walter, F., & Lipovsky, B. P. (2019). Crack wave resonances within the
634	basal water layer. Annals of Glaciology, $60(79)$, 158–166. doi: 10.1017/aog
635	.2019.8
636	Henry, F., Fokkema, J., & De Pater, C. (2002). Experiments on stoneley wave prop-

-30-

637	agation in a borehole intersected by a finite horizontal fracture. In $64th~eage$
638	conference & exhibition (pp. cp–5). doi: 10.3997/2214-4609-pdb.5.P143
639	Jin, Y., Zheng, Y., Huang, L., & Ehlig-Economides, C. (2022). Characterizing
640	hydraulic fractures using the transient pressure surge effect. In $Spe/aapg/seg$
641	unconventional resources technology conference (p. D021S028R002). doi:
642	10.15530/urtec-2022-3718981
643	Kac, M. (1966) . Can one hear the shape of a drum? The american mathematical
644	monthly, 73 (4P2), 1–23. doi: 10.1080/00029890.1966.11970915
645	Karimi-Fard, M., Durlofsky, L. J., & Aziz, K. (2004). An efficient discrete-fracture
646	model applicable for general-purpose reservoir simulators. SPE journal, $9(02)$,
647	227–236. doi: 10.2118/88812-PA
648	Kawakatsu, H., Kaneshima, S., Matsubayashi, H., Ohminato, T., Sudo, Y., Tsutsui,
649	T., Legrand, D. (2000) . Aso94: Aso seismic observation with broadband
650	instruments. Journal of Volcanology and Geothermal Research, 101(1-2),
651	129–154. doi: 10.1016/S0377-0273(00)00166-9
652	Korneev, V. (2008) . Slow waves in fractures filled with viscous fluid. Geophysics,
653	73(1), N1–N7. doi: 10.1190/1.2802174
654	Korneev, V., Danilovskaya, L., Nakagawa, S., & Moridis, G. (2014). Krauklis wave
655	in a trilayer. $Geophysics$, 79(4), L33–L39. doi: 10.1190/geo2013-0216.1
656	Krauklis, P. V. (1962). On some low-frequency oscillations of a fluid layer in an elas-
657	tic medium. Prikl. Mat. Mekh., $26(6)$, 1111–1115. doi: 10.1016/0021-8928(63)
658	90084-4
659	Kumagai, H., & Chouet, B. A. (1999) . The complex frequencies of long-period seis-
660	mic events as probes of fluid composition beneath volcanoes. Geophysical Jour-
661	nal International, $138(2)$, F7–F12. doi: 10.1046/j.1365-246X.1999.00911.x
662	Kumagai, H., & Chouet, B. A. (2000). Acoustic properties of a crack containing
663	magmatic or hydrothermal fluids. Journal of Geophysical Research: Solid
664	Earth, 105(B11), 25493–25512. doi: 10.1029/2000JB900273
665	Kumagai, H., & Chouet, B. A. (2001). The dependence of acoustic properties of
666	a crack on the resonance mode and geometry. Geophysical research letters,
667	28(17), 3325-3328. doi: 10.1029/2001GL013025
668	Kumagai, H., Miyakawa, K., Negishi, H., Inoue, H., Obara, K., & Suetsugu, D.
669	(2003). Magmatic dike resonances inferred from very-long-period seismic

-31-

670	signals. Science, 299(5615), 2058–2061. Retrieved from http://science
671	.sciencemag.org/content/299/5615/2058 doi: 10.1126/science.1081195
672	Li, L., & Lee, S. H. (2008). Efficient field-scale simulation of black oil in a nat-
673	urally fractured reservoir through discrete fracture networks and homoge-
674	nized media. SPE Reservoir evaluation & engineering, $11(04)$, 750–758. doi:
675	10.2118/103901-PA
676	Liang, C., Karlstrom, L., & Dunham, E. M. (2020). Magma oscillations in a
677	conduit-reservoir system, application to very long period (vlp) seismicity at
678	basaltic volcanoes: 1. theory. Journal of Geophysical Research: Solid Earth,
679	125(1), e2019JB017437. doi: 10.1029/2019JB017437
680	Liang, C., O'Reilly, O., Dunham, E. M., & Moos, D. (2017). Hydraulic fracture di-
681	agnostics from krauklis-wave resonance and tube-wave reflections. <i>Geophysics</i> ,
682	82(3), D171-D186.doi: 10.1190/geo2016-0480.1
683	Liang, C., Peng, J., Ampuero, JP., Shauer, N., & Dai, K. (2023, August). Dataset
684	for "Resonances in fluid-filled cracks of complex geometry and application
685	to very long period (VLP) seismic signals at Mayotte submarine volcano".
686	Zenodo. Retrieved from https://doi.org/10.5281/zenodo.8275079 doi:
687	10.5281/zenodo.8275079
688	Lipovsky, B. P., & Dunham, E. M. (2015). Vibrational modes of hydraulic frac-
689	tures: Inference of fracture geometry from resonant frequencies and attenua-
690	tion. Journal of Geophysical Research: Solid Earth, 120(2), 1080–1107. doi:
691	10.1002/2014JB011286
692	Lokmer, I., Saccorotti, G., Di Lieto, B., & Bean, C. J. (2008). Temporal evolution
693	of long-period seismicity at etna volcano, italy, and its relationships with the
694	2004–2005 eruption. Earth and Planetary Science Letters, 266(1-2), 205–220.
695	doi: 10.1016/j.epsl.2007.11.017
696	Maeda, Y., & Kumagai, H. (2013). An analytical formula for the longitudinal res-
697	onance frequencies of a fluid-filled crack. Geophysical Research Letters, $40(19)$,
698	5108–5112. doi: $10.1002/grl.51002$
699	Maeda, Y., & Kumagai, H. (2017). A generalized equation for the resonance fre-
700	quencies of a fluid-filled crack. Geophysical Journal International, $209(1)$, 192–
701	201. doi: 10.1093/gji/ggx019
702	McQuillan, M., & Karlstrom, L. (2021). Fluid resonance in elastic-walled englacial

703	transport networks. Journal of Glaciology, $67(266),999{-}1012.$ doi: 10.1017/jog
704	.2021.48
705	Mercury, N., Lemoine, A., Doubre, C., Bertil, D., van Der Woerd, J., Hoste-
706	Colomer, R., & Battaglia, J. (2022). Onset of a submarine eruption east of
707	mayotte, comoros archipelago: the first ten months seismicity of the seismo-
708	volcanic sequence (2018–2019). Comptes Rendus. Géoscience, 354 (S2), 105–
709	136. doi: 10.5802/crgeos.191
710	Métaxian, JP., Araujo, S., Mora, M., & Lesage, P. (2003). Seismicity related to the
711	glacier of cotopaxi volcano, ecuador. Geophysical Research Letters, $30(9)$. doi:
712	10.1029/2002GL016773
713	Mittal, T., Jordan, J. S., Retailleau, L., Beauducel, F., & Peltier, A. (2022). May-
714	otte 2018 eruption likely sourced from a magmatic mush. Earth and Planetary
715	Science Letters, 590, 117566. doi: 10.1016/j.epsl.2022.117566
716	Moinfar, A., Varavei, A., Sepehrnoori, K., & Johns, R. T. (2013, 07). Develop-
717	ment of an Efficient Embedded Discrete Fracture Model for 3D Compositional
718	Reservoir Simulation in Fractured Reservoirs. SPE Journal, $19(02)$, 289-303.
719	doi: 10.2118/154246-PA
720	Nakagawa, S., Nakashima, S., & Korneev, V. A. (2016). Laboratory measurements
721	of guided-wave propagation within a fluid-saturated fracture. Geophysical
722	Prospecting, $64(1)$, 143–156. doi: 10.1111/1365-2478.12223
723	Nakano, M., & Kumagai, H. (2005). Response of a hydrothermal system to
724	magmatic heat inferred from temporal variations in the complex frequen-
725	cies of long-period events at kusatsu-shirane volcano, japan. Journal of
726	volcanology and geothermal research, 147(3-4), 233–244. doi: 10.1016/
727	j.jvolgeores.2005.04.003
728	Nikkhoo, M., & Walter, T. R. (2015). Triangular dislocation: an analytical, artefact-
729	free solution. Geophysical Journal International, 201(2), 1119–1141. doi: 10
730	.1093/gji/ggv035
731	Niu, J., & Song, TR. A. (2020). Real-time and in-situ assessment of conduit
732	permeability through diverse long-period tremors beneath as volcano,
733	japan. Journal of Volcanology and Geothermal Research, 401, 106964. doi:
734	10.1016/j.jvolgeores.2020.106964
735	Okada, Y. (1985). Surface deformation due to shear and tensile faults in a half-

736	space. Bulletin of the seismological society of America, $75(4)$, 1135–1154.
737	Okada, Y. (1992). Internal deformation due to shear and tensile faults in a half-
738	space. Bulletin of the Seismological Society of America, $82(2)$, 1018–1040.
739	O'Reilly, O., Dunham, E. M., & Nordström, J. (2017). Simulation of wave propa-
740	gation along fluid-filled cracks using high-order summation-by-parts operators
741	and implicit-explicit time stepping. SIAM Journal on Scientific Computing,
742	<i>39</i> (4), B675–B702. Retrieved from https://doi.org/10.1137/16M1097511
743	doi: 10.1137/16M1097511
744	Paillet, F. L., & White, J. E. (1982, 08). Acoustic modes of propagation in
745	the borehole and their relationship to rock properties. $Geophysics, 47(8),$
746	1215-1228. Retrieved from https://doi.org/10.1190/1.1441384 doi:
747	10.1190/1.1441384
748	Pointer, T., Liu, E., & Hudson, J. A. (1998). Numerical modelling of seismic
749	waves scattered by hydrofractures: application of the indirect boundary el-
750	ement method. $Geophysical Journal International, 135(1), 289-303.$ doi:
751	10.1046/j.1365-246X.1998.00644.x
752	Retailleau, L., Saurel, JM., Laporte, M., Lavayssière, A., Ferrazzini, V., Zhu,
753	W., \ldots others (2022). Automatic detection for a comprehensive view of
754	mayotte seismicity. Comptes Rendus. Géoscience, $354(S2)$, $153-170$. doi:
755	$10.5802/\mathrm{crgeos.}133$
756	Rice, J. R. (1993). Spatio-temporal complexity of slip on a fault. Journal of Geo-
757	physical Research: Solid Earth, $98(B6)$, $9885-9907$. doi: $10.1029/93$ JB00191
758	Rona, A. (2007). The acoustic resonance of rectangular and cylindrical cavities.
759	Journal of Algorithms & Computational Technology, $1(3)$, $329-356$. doi:
760	10.1260/174830107782424110
761	Saurel, JM., Jacques, E., Aiken, C., Lemoine, A., Retailleau, L., Lavayssière, A.,
762	\ldots others (2021). May otte seismic crisis: building knowledge in near real-time
763	by combining land and ocean-bottom seismometers, first results. $Geophysical$
764	Journal International, 228(2), 1281–1293. doi: 10.1093/gji/ggab392
765	Segall, P. (2010). Earthquake and volcano deformation. Princeton University Press.
766	doi: $10.1515/9781400833856$
767	Shauer, N., Desmond, K. W., Gordon, P. A., Liu, F., & Duarte, C. A. (2021). A
768	three-dimensional generalized finite element method for the simulation of wave

-34-

769	propagation in fluid-filled fractures. Computer Methods in Applied Mechanics
770	and Engineering, 386, 114136. doi: 10.1016/j.cma.2021.114136
771	Stuart, G., Murray, T., Brisbourne, A., Styles, P., & Toon, S. (2005). Seismic emis-
772	sions from a surging glacier: Bakaninbreen, svalbard. Annals of Glaciology, 42,
773	151–157. doi: 10.3189/172756405781812538
774	Sun, F., Gong, Y., & Dong, C. (2020). A novel fast direct solver for 3d elas-
775	tic inclusion problems with the isogeometric boundary element method.
776	Journal of Computational and Applied Mathematics, 377, 112904. doi:
777	10.1016/j.cam.2020.112904
778	Tang, X., & Cheng, C. (1988). Wave propagation in a fluid-filled fracture—an exper-
779	imental study. Geophysical Research Letters, $15(13)$, 1463–1466. doi: 10.1029/
780	GL015i013p01463
781	Tang, X., & Cheng, C. (1989). A dynamic model for fluid flow in open borehole frac-
782	tures. Journal of Geophysical Research: Solid Earth, $94(B6)$, 7567–7576. doi:
	10,1020/IB004;B06p07567
783	10.1029/3D0341D00p07307
783 784	Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of
783 784 785	Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments.
783 784 785 786	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi:
783 784 785 786 787	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904
783 784 785 786 787 788	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-
783 784 785 786 787 788 788	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir sim-
783 784 785 786 787 788 788 789 790	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi:
783 784 785 786 787 788 789 790 791	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA
783 784 785 786 787 788 789 790 791 792	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the
783 784 785 786 787 788 789 790 791 792 793	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>,
783 784 785 786 787 788 789 790 791 792 793 794	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>, 174(3), 1174–1186. doi: 10.1111/j.1365-246X.2008.03871.x
783 784 785 786 787 788 789 790 791 792 793 794 795	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>, 174(3), 1174–1186. doi: 10.1111/j.1365-246X.2008.03871.x Zheng, Y., Malallah, A. H., Fehler, M. C., & Hu, H. (2016). 2d full-waveform model-
 783 784 785 786 787 788 789 790 791 792 793 794 795 796 	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>, 174 (3), 1174–1186. doi: 10.1111/j.1365-246X.2008.03871.x Zheng, Y., Malallah, A. H., Fehler, M. C., & Hu, H. (2016). 2d full-waveform modeling of seismic waves in layered karstic media. <i>Geophysics</i>, 81(2), T25–T34. doi:

798

Appendix A Matrices D and T for a simple crack intersection

In this section, we show step by step how to construct matrices **D** and **T** for a simple crack intersection shown in Figure A1. The element number and positive flux direc-
- tion of each active connection as labeled. The boundary edges have zero flux and they
- $_{802}$ do not contribute to **D** and **T**. Thus, we have five elements and five active connections
- numbered as $\{2 \rightarrow 1, 3 \rightarrow 2, 4 \rightarrow 2, 5 \rightarrow 2, 5 \rightarrow 4\}$, where $i \rightarrow j$ defines the positive flux direction. The size of both **D** and **T** are 5 by 5.
 - T_{54} T_{52} T_{32} (3) $(4) \leftarrow 2$ T_{42} T_{21} (1)

Figure A1. Geometry of a simple crack intersection. The element number and the positive flow direction of each active connection (non-zero flux) are indicated by the circled number and arrow, respectively. The scalar transmisibilities are labled near each connection.

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Let's first consider the matrix \mathbf{D} , which sums the flux from active connections to 805 obtain the net out-flux from each element. We consider the first row of \mathbf{D} as an exam-806 ple, relevant for element 1. The only connection that contributes to the net out-flux of 807 element 1 is connection 1 with the positive direction of $2 \rightarrow 1$, the opposite to the out-808 flux direction. Thus, D(1,1) = -1 and other entries of the first row are zeros. How-809 ever, for element 2, the positive flux of connection 1 aligns with the outflux direction, 810 which leads to D(2,1) = 1. Similarly, other entries of matrix **D** can be determined and 811 the matrix \mathbf{D} is: 812

$$\mathbf{D} = \begin{bmatrix} -1 & 0 & 0 & 0 & 0 \\ 1 & -1 & -1 & -1 & 0 \\ 0 & 1 & 0 & 0 & 0 \\ 0 & 0 & 1 & 0 & -1 \\ 0 & 0 & 0 & 1 & 1 \end{bmatrix}.$$
 (A1)

We now proceed to construct the matrix \mathbf{T} , which computes the flux on each active connection from the pressure on each cell. Note that we only store the flux in the positive direction. For instance, the flux on the first connection is $Q_{2\rightarrow 1} = T_{21} (p_2 - p_1)$, which means $T(1,2) = -T(1,1) = T_{21}$. Similarly, other entries of the matrix \mathbf{T} can be com-

 $_{817}$ puted and the full expression of T is:

$$\mathbf{T} = \begin{bmatrix} -T_{21} & T_{21} & 0 & 0 & 0\\ 0 & -T_{32} & T_{32} & 0 & 0\\ 0 & -T_{42} & 0 & T_{42} & 0\\ 0 & -T_{52} & 0 & 0 & T_{52}\\ 0 & 0 & 0 & -T_{54} & T_{54} \end{bmatrix}.$$
 (A2)

Appendix B Resonant frequencies from time domain results by GFEM

In this section, we explain the procedure to obtain selective resonant frequencies 819 from the time domain simulation results using the GFEM code developed by Shauer et 820 al. (2021). As shown in Figure B1, we apply injection sources with a gaussian source time 821 function on the certain position on the crack (red stars), obtain the pressure time series 822 (duration of 50 s) on three receiving points (blue triangles), and then extract the res-823 onant frequencies at spectral peaks. For the rectangular crack, we place one source at 824 the upperleft corner, which manages to excite all the first eight modes, and three receivers 825 (R1, R2, and R3) at (-0.5, 0), (-0.20, 0.25), and (0, 0.25), respectively. Different receivers 826 sample different eigenmodes. For instance, receiver R1 samples modes 1, 2, 5, and 8 as 827 shown in Figure B1-c. The modes sampled by R2 and R3 are shown in Table 2. We make 828 this choice to selectively sample closely-spaced modes, for instance mode 2 and 3, at dif-829 ferent receivers to avoid ambiguity. 830

For the elliptical crack, we place two sources at the leftmost and uppermost ends, and three receivers at (-0.5, 0), (0, 0.25), (0, 0) respectively. Due to the excitation and monitoring geometry, we focus only sampling the longitudinal and transverse modes, which are clearly seperated peaks in the spectrum. The eigenmodes sampled by different receivers are shown in Figure B1-f and Table 3.



Figure B1. (a, d) The source and receiver positions. (b, e) Pressure time series at three receivers. (c, f) The normalized spectral amplitude of data at receiver R1. The vertial black dashed lines are the resonant frequencies (with mode number labelled) computed by BEM+FVM method.

Resonances in fluid-filled cracks of complex geometry and application to very long period (VLP) seismic signals at Mayotte submarine volcano

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Key Points:

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13	• Hybrid method combining BEM and FVM efficiently computes resonant frequen-
14	cies of complex-shaped fluid-filled cracks
15	• Elliptical crack shares similar modes with rectangular crack but a crack network
16	produces more complex resonances
17	- A dumbbell-shaped crack explains ratio of first two modes (~2.5) in the VLP seis-
18	mic signal at Mayotte submarine volcano

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19 Abstract

Fluid-filled cracks sustain a slow guided wave (Krauklis wave or crack wave) whose res-20 onant frequencies are widely used for interpreting long period (LP) and very long pe-21 riod (VLP) seismic signals at active volcanoes. Significant efforts have been made to model 22 this process using analytical developments along an infinite crack or numerical methods 23 on simple crack geometries. In this work, we develop an efficient hybrid numerical method 24 for computing resonant frequencies of complex-shaped fluid-filled cracks and networks 25 of cracks and apply it to explain the ratio of spectral peaks in the VLP signals from the 26 Fani Maoré submarine volcano that formed in Mayotte in 2018. By coupling triangu-27 lar boundary elements and the finite volume method, we successfully handle complex ge-28 ometries and achieve computational efficiency by discretizing solely the crack surfaces. 29 The resonant frequencies are directly determined through eigenvalue analysis. After proper 30 verification, we systematically analyze the resonant frequencies of rectangular and ellip-31 tical cracks, quantifying the effect of aspect ratio and crack stiffness ratio. We then dis-32 cuss theoretically the contribution of fluid viscosity and seismic radiation to energy dis-33 sipation. Finally, we obtain a crack geometry that successfully explains the character-34 istic ratio between the first two modes of the VLP seismic signals from the Fani Maoré 35 submarine volcano in Mayotte. Our work not only reveals rich eigenmodes in complex-36 shaped cracks but also contributes to illuminating the subsurface plumbing system of 37 active volcanoes. The developed model is readily applicable to crack wave resonances 38 in other geological settings, such as glacier hydrology and hydrocarbon reservoirs. 39

40 1 Introduction

Slow guided waves that propagate along fluid-filled cracks, named crack waves or 41 Krauklis waves, can be used for inferring the geometries of subsurface cracks and the fluid 42 properties in a wide range of geological settings (Krauklis, 1962; Ferrazzini & Aki, 1987; 43 Paillet & White, 1982; B. Chouet, 1986; Korneev, 2008; Tang & Cheng, 1989; Lipovsky 44 & Dunham, 2015). In volcanology, crack wave resonances along magma-filled sills and 45 dikes have been used for interpreting long period (LP, 0.5-2 s) and very long period (VLP, 46 2 to 100 s) seismic signals at many volcanos, including Mount Redoubt (B. A. Chouet 47 et al., 1994), Aso (Kawakatsu et al., 2000; Niu & Song, 2020), Galeras (Cruz & Chouet, 48 1997), Asama (Fujita & Ida, 2003), Kusatsi-Shirane (Kumagai et al., 2003; Nakano & 49 Kumagai, 2005), Etna (Lokmer et al., 2008), and Erebus (Aster, 2019). Crack waves (and 50

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their induced tube waves in wellbores) are used for diagnosing the fracture geometries 51 in unconventional hydrocarbon reservoirs (Henry et al., 2002; Tary et al., 2014; Lipovsky 52 & Dunham, 2015; Liang et al., 2017). The resonating or humming signals in glaciers have 53 also been attributed to crack waves (Métaxian et al., 2003; Stuart et al., 2005; Gräff et 54 al., 2019; McQuillan & Karlstrom, 2021). Natural cracks in the subsurface are complex 55 in shape and usually form an inter-connected network. Therefore, efficient methods for 56 computing resonant modes of single cracks and networks of cracks are necessary for in-57 terpreting frequencies measured in the field. 58

Since its first discovery by Krauklis (1962), crack waves have been studied analyt-59 ically (Aki et al., 1977; Ferrazzini & Aki, 1987; Korneev, 2008; Lipovsky & Dunham, 2015), 60 experimentally (Tang & Cheng, 1988; Nakagawa et al., 2016; Cao et al., 2021), and nu-61 merically by various methods (e.g., B. Chouet, 1986; Yamamoto & Kawakatsu, 2008; 62 Frehner & Schmalholz, 2010; O'Reilly et al., 2017; Liang et al., 2020; Shauer et al., 2021; 63 Jin et al., 2022). Analytically derived dispersion relations are useful for understanding 64 the propagation behavior but are meant for an infinitely long crack and do not account 65 for the restriction of the finite crack tip. The finite difference method (FDM) is normally 66 based on cartesian grids in 2D (Fehler & Aki, 1978) or 3D (B. Chouet, 1986; Liang et 67 al., 2020) and limited to a tabular crack shape. Maeda and Kumagai (2013) and Maeda 68 and Kumagai (2017) performed a large number of numerical simulations on rectangu-69 lar cracks using a FDM simulator developed by B. Chouet (1986). With that, they ob-70 tained a set of empirical fitting formulas for resonant frequencies given the crack aspect 71 ratio α and stiffness ratio $C_L = K_f L/(Gw_0)$, where K_f is the fluid bulk modulus, G 72 the solid shear modulus, L the crack length and w_0 the crack aperture. However, such 73 relations only apply to longitudinal or transverse modes on rectangular cracks (Maeda 74 & Kumagai, 2013, 2017). Notably, O'Reilly et al. (2017) simulated a non-planar fluid-75 filled crack using FDM on a curvilinear grid and adopted a lubrication-type approxima-76 tion in the fluid (Lipovsky & Dunham, 2015), neglecting fluid acoustics in the crack width 77 direction while resolving the narrow viscous boundary layer close to the crack wall. This 78 treatment removes the time step restriction introduced by extremely fine mesh size in 79 the crack width direction and accelerates the computation. However, their work was lim-80 ited to 2D geometries. The finite element method (FEM) is more flexible for handling 81 complex crack geometries and has been used to study crack waves in 2D (Frehner & Schmal-82 holz, 2010; Frehner, 2013) and 3D (Shauer et al., 2021). Particularly, Shauer et al. (2021) 83

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produced the first simulation of an elliptical fluid-filled crack using the generalized finite 84 element method (GFEM). However, similar to FDM, FEM needs to discretize the vol-85 ume, which results in a large number of elements and high computational cost. On the 86 other hand, the boundary element method (BEM) reduces the simulation space from a 87 domain to boundary surfaces, drastically decreasing the number of degrees of freedom, 88 and has been used to study waves in fluid-filled cracks (Yamamoto & Kawakatsu, 2008; 89 Pointer et al., 1998; Jin et al., 2022) and other inclusions (Zheng et al., 2016; Sun et al., 90 2020). However, previous BEM simulations are either in two dimensions or focus on the 91 wave diffraction instead of analyzing the resonant frequencies. Currently, the study of 92 resonant frequencies of complex-shaped fluid-filled cracks and crack networks in three 93 dimensions remain unknown. 94

In this work, we propose an efficient hybrid numerical method to simulate crack 95 wave resonance in complex-shaped cracks or crack networks filled with an inviscid fluid, 96 by coupling the boundary element method (BEM) for the solid response and the finite 97 volume method (FVM) for acoustics in the fluid. By using triangular elements in both 98 BEM and FVM on the crack surfaces, we successfully handle complex crack shapes and 99 intersections. We restrict our attention to the low frequency limit where the crack wave 100 is much slower than the solid body waves, such that the solid response can be approx-101 imated as quasi-static (Korneev, 2008; Lipovsky & Dunham, 2015; Liang et al., 2020). 102 An eigenvalue analysis is performed to extract the resonant modes directly in the fre-103 quency domain, circumventing errors from time discretization and spectral analysis of 104 the time domain simulation data. We first verify our method by comparing results with 105 analytical solutions in the rigid wall limit and with numerical solutions from existing meth-106 ods for both a rectangular (B. Chouet, 1986; Maeda & Kumagai, 2017) and elliptical cracks 107 (Shauer et al., 2021). An example is then provided to demonstrate the simulation ca-108 pability for intersecting cracks. The effect of crack aspect ratio and stiffness ratio on res-109 onant frequencies (longitudinal, transverse, and mixed modes) is systematically inves-110 tigated for both rectangular and elliptical cracks. Although our current model does not 111 include viscous or radiation loss, we provide some theoretical discussion on these effects 112 under simple assumptions (boundary layer limit and quasi-dynamic approximation). Fi-113 nally, we present a crack shape compatible with the first two spectral peaks of VLP seis-114 mic signals from the Fani Maoré, Mayotte submarine volcano and discuss the potential 115 of the methodology for future applications in volcanology and other geological settings. 116

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Figure 1. Schematics of an arbitrarily-shaped fluid-filled crack, its spatial discretization (with unknown variables placed in the element centroids, red dots), and a zoom-in view at an intersection between two cracks.

117 2 Methods

In this section, we present the governing equations, discretization, and eigenmode analysis for computing the resonant frequencies.

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2.1 Governing equations

We consider the oscillation of inviscid fluids in complex-shaped thin cracks embed-121 ded in a 3D homogeneous linear elastic solid (Figure 1). The initial opening of the crack 122 is w_0 , which is assumed to be a constant and much smaller than the wavelength λ . We 123 adopt a similar lubrication approximation as B. Chouet (1986), Yamamoto and Kawakatsu 124 (2008) and O'Reilly et al. (2017), and treat the fluid pressure and velocities as uniform 125 in the crack thickness direction, reducing the crack from a 3D body to a 2D surface S. 126 Following O'Reilly et al. (2017), we consider a small crack curvature so that its effect 127 on the fluid momentum balance is negligible. Thus, the mass and momentum balance 128 of the fluid on the crack surface are written as 129

$$\frac{1}{w_0}\frac{\partial w}{\partial t} + \frac{1}{K_f}\frac{\partial p}{\partial t} + \frac{\partial v_\xi}{\partial \xi} + \frac{\partial v_\eta}{\partial \eta} = 0, \tag{1}$$

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$$\rho_f \frac{\partial v_\xi}{\partial t} + \frac{\partial p}{\partial \xi} = 0, \tag{2}$$

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$$\rho_f \frac{\partial v_\eta}{\partial t} + \frac{\partial p}{\partial \eta} = 0, \tag{3}$$

where ρ_f and K_f are fluid density and bulk modulus, w is the crack opening perturbation, p is the fluid pressure perturbation, t is time, and ξ and η are two locally perpendicular coordinates tangent to the crack surface, v_{ξ} and v_{η} are the fluid velocities in the ξ and η directions, respectively. Eliminating v_{ξ} and v_{η} in (1) using (2)-(3), we have

$$\rho_f \left(\frac{1}{w_0} \frac{\partial^2 w}{\partial t^2} + \frac{1}{K_f} \frac{\partial^2 p}{\partial t^2} \right) - \Delta p = 0, \tag{4}$$

where $\Delta = \frac{\partial^2}{\xi^2} + \frac{\partial^2}{\eta^2}$ is the tangential Laplace operator along the crack surface. The coupling between fluid and solid is encapsulated in the relation between the crack opening perturbation w and pressure perturbation p, which must balance the solid normal stress perturbation σ_n on the crack wall (assumed positive in compression). Since we focus on the low frequency limit, the solid response is approximately quasi-static (Korneev, 2008; Lipovsky & Dunham, 2015; Liang et al., 2020), and p for a linear elastic solid can be expressed as (Segall, 2010):

$$p(x) = \int_{S} K(x,\xi) w(\xi) \, dA, \tag{5}$$

where $K(x,\xi)$ is the Green's function that relates a unit open dislocation impulse at ξ to the normal stress change at x. The expressions of K in an elastic whole space and half space are available analytically for a uniform dislocation on both rectangular elements (Okada, 1985, 1992) and triangular elements (Nikkhoo & Walter, 2015).

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2.2 Discretization

We discretize the crack surfaces into N_e triangular elements. The unknown average pressures $\bar{\mathbf{p}}$ and openings $\bar{\mathbf{w}}$, placed at element centroids (as shown in Figure 1), are related by

$$\bar{\mathbf{p}} = \mathbf{K}\bar{\mathbf{w}},\tag{6}$$

where **K** is a N_e by N_e matrix and K(i, j) denotes the fluid pressure (or solid normal stress) change at the centroid of the *i*-th element caused by a unit open dislocation on the *j*-th element. We use the full space Green's function in this study but one can also use the half space solution.

We then discretize the tangential Laplacian operator by a finite volume scheme with a two-point flux (TPF) approximation following Karimi-Fard et al. (2004), which has been widely used for diffusive flows through a discrete fracture network in hydrocarbon reservoirs (e.g., Li & Lee, 2008; Moinfar et al., 2013; Xu et al., 2017; Berre et al., 2019). This scheme is only first-order accurate and is thus rarely used in wave propagation problems due to the strong numerical diffusion in time domain simulations (e.g., Durran, 2013). However, it is a sufficient scheme for our problem as we focus on resolving only the spa-

tial distribution of eigenmodes in the frequency domain and the low order of accuracy

can be remedied by using more elements. Here, we briefly present the key derivation steps

and the readers are referred to Karimi-Fard et al. (2004) for a detailed description.

We consider an arbitrary planar triangular element i with a surface S_i and boundary edges l_{ij} , where j is the index of the neighboring elements. Each i and j pair forms a hydraulic connection. When multiple cracks intersect, multiple connections share the same edge. We integrate equation (4) over each element i's surface, leading to:

$$\rho_f A_i \left[\frac{1}{w_0} \frac{\partial^2 \bar{w}_i}{\partial t^2} + \frac{1}{K_f} \frac{\partial^2 \bar{p}_i}{\partial t^2} \right] = \int_{S_i} \Delta p ds, \tag{7}$$

169 where

$$\bar{p}_i = \frac{1}{A_i} \int_{S_i} p ds,\tag{8}$$

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$$\bar{w}_i = \frac{1}{A_i} \int_{S_i} w ds,\tag{9}$$

are the average pressure and opening of element i, respectively. Applying the divergence

theorem to the right hand side of equation (7), we have:

$$\int_{S_i} \Delta p ds = \int_{S_i} \vec{\nabla} \cdot \vec{\nabla} p ds = \int_l \frac{\partial p}{\partial n} dl = -\sum_{j=1}^{n_c} D_{i \to j} Q_{i \to j}, \tag{10}$$

where $\partial p/\partial n$ is the pressure gradient normal to the boundary edges, n_c is the total number of connections in contact with element i, $Q_{i \to j}$ is the flux going out from element i to element j. Since $Q_{i \to j} = -Q_{j \to i}$, we only store $Q_{i \to j}$ for each (i, j) pair and its positive flux direction is pre-defined by an indicator function $I_{i \to j} = -I_{j \to i} = 1$. $D_{i \to j}$ is the discrete divergence operator and $D_{i \to j} = I_{i \to j} = 1$.

The assumption of the TPF scheme is to approximate the flux term in the following form (equation (7) in Karimi-Fard et al. (2004)):

$$Q_{i \to j} = I_{i \to j} T_{ij} (p_i - p_j), \tag{11}$$

where p_i and p_j are pressures defined at the centroids of the two neighboring elements.

 T_{ij} is the scalar transmissibility and is expressed as

$$T_{ij} = \frac{\alpha_i \alpha_j}{\sum_{k=1}^{n_c} \alpha_k},\tag{12}$$

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$$\alpha_k = \frac{l_{ij}}{d_k} \vec{n}_k \cdot \vec{f}_k, \tag{13}$$

where l_{ij} is the length of the connecting edge, d_k and \vec{f}_k are the length and unit directional vector from midpoint of the edge to the centroid of element k, \vec{n}_k is a unit normal vector perpendicular to the edge and pointing towards element k, as shown in Figure 1. Fluxes on the crack boundaries are set to zero. Combining equations (10) and (11), we have:

$$\int_{S_i} \Delta p ds = -\sum_{j=1}^{n_c} D_{i \to j} I_{i \to j} T_{ij} \left(p_i - p_j \right). \tag{14}$$

It is apparent that changing the positive flux direction from $i \to j$ to $j \to i$ flips the sign of both $D_{i\to j}$ and $I_{i\to j}$ and thus results in the same Laplacian term. Substituting equation (14) into equation (7) and rewriting in the matrix form, we have the spatially discretized equation without external forcing:

$$\rho_f\left(\frac{1}{w_0}\mathbf{K}^{-1} + \frac{1}{K_f}\right)\frac{\partial^2 \bar{\mathbf{p}}}{\partial t^2} = -\mathbf{A}^{-1}\mathbf{D}\mathbf{Q} = -\mathbf{A}^{-1}\mathbf{D}\mathbf{T}\bar{\mathbf{p}},\tag{15}$$

where **A** is a diagonal matrix of size N_e by N_e denoting the area of each element, $\mathbf{Q} = \mathbf{T}\mathbf{\bar{p}}$ is the flux vector whose size is the total number of connections N_c , **T** is the transmissibility matrix (including the indicator function) of size N_c by N_e that maps the vector $\mathbf{\bar{p}}$ to \mathbf{Q} , and \mathbf{D} is the divergence matrix of size N_e by N_c that maps \mathbf{Q} to the net flux out of each element. The structure of matrices \mathbf{D} and \mathbf{T} for a system of three intersecting crack elements are described in Appendix A.

We further introduce the following dimensionless quantities:

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$$\mathbf{K}^* = \mathbf{K}/(G/L), \mathbf{A}^* = \mathbf{A}/L^2, w = w^*/w_0, t^* = t/(L/c_l), \bar{\mathbf{p}}^* = \bar{\mathbf{p}}/(\rho_f c_l^2),$$
(16)

where G is the solid shear modulus, L is a representative length of the crack and $c_l = \sqrt{Gw_0/(\rho_f L)}$ is a representative crack wave speed. Different non-dimensionalization strategies exist, such as the one by B. Chouet (1986) which normalizes wave speeds by the solid compressional wave speed c_p . We choose c_l instead, because in the long wavelength limit, where compliance of the crack dominates, this choice conveniently results in a fundamental frequency of the order of unity. The nondimensionalised equation is

$$\left(\frac{1}{C_L}\mathbf{I} + (\mathbf{K}^*)^{-1}\right)\frac{\partial^2 \bar{\mathbf{p}}^*}{\partial t^{*2}} = -\mathbf{A}^{*-1}\mathbf{D}\mathbf{T}\bar{\mathbf{p}}^*,\tag{17}$$

where $C_L = K_f L/Gw_0$ is the key dimensionless parameter, named crack stiffness ra-

tio by B. Chouet (1986). The crack wave limit is achieved with $C_L \gg 1$, where the crack

 $_{207}$ is much more compliant than the fluid. C_L can be related to the representative crack

wave speed c_l by $C_L = c_f^2/c_l^2$, where c_f is the fluid acoustic wave speed. The crack topol-

 $_{209}$ ogy (for instance, the aspect ratio α for a rectangular or elliptical crack) and solid Pois-

son's ratio ν_s are encapsulated into the dimensionless stiffness matrix \mathbf{K}^* . The solid Pois-

son's ratio is set to 0.25 throughout this manuscript, unless otherwise mentioned.

2.3 Eigenmode analysis

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We directly obtain the resonant frequencies through eigenmode analysis in the frequency domain. The spatially discretized dimensionless equation is written as

$$\frac{\partial^2 \bar{\mathbf{p}}^*}{\partial t^{*2}} = -\mathbf{B} \bar{\mathbf{p}}^*,\tag{18}$$

215 where

$$\mathbf{B} = \left(\frac{1}{C_L}\mathbf{I} + (\mathbf{K}^*)^{-1}\right)^{-1} \mathbf{A}^{*-1} \mathbf{D} \mathbf{T}.$$
(19)

²¹⁶ The nondimensionalised Fourier transform is defined as

$$\hat{u}(\omega^*) = \int_{-\infty}^{+\infty} u(t^*) e^{i\omega^* t^*} dt^*,$$
(20)

217 where

$$\omega^* = \omega/\left(c_l/L\right),\tag{21}$$

218 is the dimensionless angular frequency. The dimensionless frequency is

$$f^* = \omega^* / (2\pi) = f / (c_l / L).$$
 (22)

Taking the Fourier transform of equation (18), we have:

$$(\omega^*)^2 \hat{\mathbf{p}} = \mathbf{B} \hat{\mathbf{p}},\tag{23}$$

where $(\omega^*)^2$ and \hat{p} are the eigenvalues and eigenvectors of the real matrix **B**. Since we deal with inviscid fluids, we only seek real positive eigenvalues, which correspond to undamped oscillatory modes. The resulting eigenvectors determine the spatial distribution of the pressure on the crack surface. Solving the resonant frequencies in dimensionless form is advantageous, because one can easily scale the solution to other parameters, such as crack length, crack width and solid stiffness, given the same dimensionless parameters, C_L , ν_s and crack topology.

²²⁷ **3** Verification and examples

In this section, we first verify our implementation by comparing our results to analytical solutions in the rigid solid limit and numerical solutions from existing studies.

- ²³⁰ We then present an example of simple intersecting crack geometry to demonstrate the
- utility of our method.

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Mode	Theoretical value	Numerical value	Error $(\%)$
1	0.5000	0.5004	0.074
2	0.8333	0.8325	0.105
3	0.9718	0.9706	0.127
4	1.0000	1.0008	0.078
5	1.3017	1.3018	0.009
6	1.5000	1.5010	0.066
7	1.6667	1.6647	0.118
8	1.7159	1.7172	0.072
9	1.7401	1.7370	0.175
10	1.9437	1.9413	0.122
11	2.0000	2.0014	0.068
12	2.1667	2.1693	0.123
13	2.2423	2.2409	0.060
14	2.5000	2.4954	0.184
15	2.5000	2.5010	0.042
16	2.5495	2.5426	0.270

 Table 1. The error between the theoretical and numerical resonant frequencies for the first 16 modes

3.1 Comparison with analytical solutions in a rigid solid

We compute the \mathbf{K} matrix using the subroutines developed by Nikkhoo and Wal-233 ter (2015), which have been extensively used by other studies. The bulk part that needs 234 to be validated is the FVM discretization of the Laplacian term. For that, we set solid 235 rigidity to infinity and compare the numerical results to the analytical solution of the 236 resonant frequencies of linear acoustic waves in a 2D rectangular domain with zero-flux 237 boundaries (Rona, 2007). The solution is in a dimensionless form with a rectangular do-238 main of size 1 by 0.5 and a wave speed of 1. The comparison results for the first 16 modes 239 are tabulated in Table 1. The excellent agreement between our numerical results and the 240

analytical solutions, with relative differences smaller than 0.2%, verifies our FVM dis-

²⁴² cretization of the Laplacian term.

243

3.2 Comparison to numerical solutions by existing studies

We compare solutions by our method (BEM+FVM) to those by B. Chouet (1986), 244 Maeda and Kumagai (2017) and Shauer et al. (2021). With B. Chouet (1986) and Maeda 245 and Kumagai (2017), we compare resonant frequencies of longitudinal modes for a rect-246 angular crack for various values of C_L (5, 15, 25, 50, 75, 100). With the GFEM by Shauer 247 et al. (2021), we compare solutions of multiple modes on both rectangular and ellipti-248 cal cracks. The eigenmodes can be straightforwardly classified as longitudinal (variation 249 only along the major crack axis), transverse (variation only along the minor crack axis), 250 and mixed modes for a rectangular crack, but less so for an elliptical crack. Since the 251 method by Shauer et al. (2021) discretizes the problem in time and, therefore, does not 252 readily provide resonant frequencies, we ran their code to excite the fluid oscillation on 253 the crack with $C_L = 100$ by a point injection source and then extract the resonant fre-254 quencies from the spectral peaks of the pressure records at a few receiving points. We 255 use a Gaussian time function for the injection source $f(t) = \exp\left(-(t-t_c)^2/T^2\right)$ with 256 $t_c = 0.5, T = 0.1$, to ensure a smooth start and a sufficiently wide spectrum to cover 257 enough eigenmodes. Note that if either excitation or receiving points are placed on the 258 nodal line, the eigenmode can not be excited or recorded. Therefore, not all eigenmodes 259 are excited in the time domain simulation and we also only compare selective modes with 260 Shauer et al. (2021), which is sufficient for verification purposes. The detailed geome-261 tries and simulation data are presented in Appendix B. Notably, the code of Shauer et 262 al. (2021) has the capability of both considering (fully dynamic, FD) or neglecting the 263 solid inertia (quasi-static, QS), allowing to investigate the impact of the solid inertia on 264 crack wave resonant frequencies. 265

Tables 2 and 3 show the comparison of dimensionless resonant frequencies of seletive eigenmodes from the GFEM program by Shauer et al. (2021) with those by our method for a rectangular and elliptical crack, respectively, with an aspect ratio of 0.5, major axis length of 1, and C_L of 100. The relative difference between our results and those from Shauer et al. (2021) are near 2% or less, with or without solid inertia. This close agreement not only demonstrates the validity of our approach but also reassures that the quasi-static solid response is a very good approximation when computing the

	Recomment frequencies detected at receiving points by CEEM			BEM+EVM	Error FD	Error QS			
Mode	nesona	Resonant nequencies detected at receiving points by GTEM				DEM+P VM	(%)	(%)	
	(-0.5, 0)		(-0.2, 0.25)		(0, 0.25)				
	FD	QS	FD	QS	FD	QS			
1	1.236	1.236	1.236	1.236			1.210	2.15	2.15
2	2.727	2.691					2.662	2.44	1.09
3			2.890	2.873	2.818	2.782	2.835	1.94	1.34
4			3.453	3.418			3.373	2.37	1.33
5	4.454	4.400	4.436	4.382			4.385	1.57	0.34
6					4.526	4.491	4.466	1.34	0.56
7			6.035	5.964			5.980	0.92	0.27
8	6.399	6.273			6.344	6.236	6.330	1.09	0.90

 Table 2.
 Resonant frequencies by BEM+FVM and GFEM with or without inertia (FD or QS),

 rectangular crack

The bold values are used for error calculation. We use a of CL=100 and aspect ratio of 0.5 Mode eigenfunctions are shown in Figure 2.

 Table 3.
 Resonant frequencies by BEM+FVM and GFEM without inertia (QS), elliptical crack

M.J.	The resonant frequencies can be detected at detection points				E (07)	
Mode	(-0.5, 0)	(0, 0.25)	(0, 0)	BEM+FVM	Error (%)	
1	1.527			1.518	0.59	
2	3.027		3.090	3.050	0.75	
3		3.290		3.241	1.51	
5	4.890			4.816	1.54	
7	6.890	6.853	6.944	6.771	1.76	
8		7.308	7.235	7.107	1.80	

The bold values are used for error calculation. We use a of CL=100 and aspect ratio of 0.5 Mode eigenfunctions are shown in Figure 3.

crack wave resonant frequencies, at least for a C_L of 100. A similar conclusion has also 273 been reached by Shauer et al. (2021). Since we assume a quasi-static solid response, it 274 is reasonable that our results have a better agreement to those by GFEM without in-275 ertia. 276



Figure 2. Dimensionless frequencies and eigenfunctions of the first 16 resonant modes (numbered in an ascending order in frequencies) of a rectangular crack with $C_L=100$ and aspect ratio of 0.5 calculated by BEM+FVM. The errors of selective resonant frequencies between the BEM+FVM and GFEM without inertia are shown in Table 2. The white color indicates the nodal lines.

The pressure eigenfunctions of the first 16 resonant modes are displayed in Figure 2 for a rectangular crack and Figure 3 for an elliptical crack, showing a rich spectrum of spatial variations including longitudinal, transverse, and mixed modes. Different modes can produce different near and far field radiation patterns, that may be detectable in real 280 seismic data (e.g., Liang et al., 2020).

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The dimensionless frequencies of the first 9 longitudinal modes for rectangular cracks 282 by various methods with different crack stiffness ratios are shown in Figure 4. The re-283 sults of Shauer et al. (2021) are only computed for a C_L of 100. For ease of comparison, 284 we convert dimensionless frequencies f^* in our studies to those in B. Chouet (1986) f_C^* , 285



Figure 3. Same as Figure 2 but for an elliptical crack.

which are related by $f_C^* = f^* c_l / c_P$. Overall, our results match well with those by Shauer 286 et al. (2021) (relative error < 3%) and also qualitatively well with those by B. Chouet 287 (1986) and Maeda and Kumagai (2017). However, there are quantitative discrepancies 288 between our results and those by B. Chouet (1986) (relative error 8.83-23.43%) and Maeda 289 and Kumagai (2017) (relative error 2.72-16.63%, see the supporting information for tab-290 ulated errors). Particularly, both B. Chouet (1986) and Maeda and Kumagai (2017) sys-291 tematically give lower frequencies than those by our method and Shauer et al. (2021). 292 We suspect these discrepancies are likely due to differences in spatial and temporal sam-293 pling, or domain sizes used in the FDM code in B. Chouet (1986) and Maeda and Ku-294 magai (2017). Particularly, a truncated domain in the FDM results in a more compli-295 ant solid response (Korneev et al., 2014), which in turn results in a lower crack wave speed 296 and resonant frequencies. Our method uses boundary elements and thus an infinite do-297 main is directly satisfied. For this reason, when comparing results with Shauer et al. (2021), 298 we deliberately used a very large domain (10 times the length of the crack) in the GFEM 299 code to minimize its boundary effect using an unstructured grid, coarsening in regions 300 far from the crack. 301



Figure 4. Dimensionless frequencies f_C^* of longitudinal modes for rectangular cracks with different C_L (5, 15, 25, 50, 75, 100) by various methods, and a zoom-in view of the case $C_L = 100$ on the right panel. Results by B. Chouet (1986) and Maeda and Kumagai (2017) are slightly shifted in the horizontal axis to avoid overcrowding the figure.

302 3

3.3 An example of intersecting cracks

We now apply our method to one example of intersecting cracks, one full ellipse 303 with a half-elliptical branch, and obtain the first 16 eigenmodes, shown in Figure 5. In-304 teractions between multiple cracks result in more complex resonant modes than in sin-305 gle cracks (shown in Figures 3 and 4). For example, the fundamental mode now involves 306 fluid exchange between the major crack and the branch, and has a lower frequency than 307 the fundamental mode of the major crack (the second mode in this case). When nodal 308 lines coincide with the intersecting edge, resonances can be isolated on the major crack, 309 such as modes 2, 7, 8, 13 and 16. Temporal manifestation of these modes requires a more 310 peculiar condition: the excitation must not be located in the branch. One can certainly 311 add more complexities in the crack network, such as asymmetries, non-planarity or more 312 intricate coupling, and expect to encounter richer eigenmodes. However, such modeling 313 only becomes meaningful when more compelling observations exist and require. We will 314 demonstrate later how a particular crack shape can explain the ratio of the first two spec-315 tral peaks in the VLP seismic data at the Fani Maoré, Mayotte submarine volcano. Ex-316

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- cept for that, we decide to leave the analysis of eigenmodes of a more complex crack net-
- ³¹⁸ work for future investigation.



Figure 5. The first 16 eigenmodes of a simple two-intersecting-cracks geometry: a half ellipse intersecting a full ellipse (aspect ratio 0.5) along its minor axis. The major axis length of the full elliptical crack is chosen as L for the non-dimensionalisation.

³¹⁹ 4 Effect of aspect ratio α and crack stiffness ratio C_L

In this section, we present the effect of α and C_L on the resonant frequencies of rectangular and elliptical cracks, with major and minor axes in the x- and y-directions, respectively. Maeda and Kumagai (2017) presented a similar analysis for rectangular cracks, but only on longitudinal and transverse modes. Here, we include the mixed modes and the results for elliptical cracks. We fixed $C_L = 100$ when varying α (from 0.05 to 1.00 with an increment of 0.05) and fix $\alpha = 0.5$ when varying C_L (from 5 to 100 with an increment of 5). The frequencies of the first 16 eigenmodes are tabulated in the Support-

ing information. Here, we select 9 representative modes and visualize them in Figures 327 6-9. For rectangular cracks, we associate to each mode a pair of numbers (i, j) that de-328 note the number of half wavelengths in the x- and y-directions. For instance, the fun-329 damental mode (1,0) is a longitudinal mode with one half wavelength pressure variation 330 in the x-direction and quasi-uniform in the y-direction. Such numbering becomes less 331 obvious for elliptical cracks, especially when the aspect ratio approaches 1, for which the 332 eigenfunctions are better characterized by radial and circumferential variations. Nonethe-333 less, for the ease of comparing results with rectangular cracks, we still number the rep-334 resentative modes in Figures 7 and 9 approximately into longitudinal, transverse, and 335 mixed modes. 336



Figure 6. Dimensionless resonant frequencies of representative modes of rectangular cracks as a function of the aspect ratio α . C_L is fixed to 100. The eigenfunctions displayed are for an aspect ratio of 0.55. Certain high order mixed and transverse modes rank outside of the first 16 eigenmodes that we store, which causes the apparent absence of data at low aspect ratios.

337

4.1 Effect of aspect ratio

The variation of resonant frequencies with aspect ratio is shown in Figure 6 and 7 for rectangular and elliptical cracks, respectively. For both cases, decreasing the aspect ratio increases the crack stiffness from the transverse direction and results in higher resonant frequencies for all the modes. This effect is relatively mild for longitudinal modes but rather steep for transverse and mixed modes. For instance, when α of a rectangu-



Figure 7. Same as Figure 6 but for elliptical cracks. Note that the mode numbers (i, j) are not strictly valid for an elliptical crack but are useful for our interpretation (more explanation in the main text).

lar crack decreases from 1.0 to 0.1, the frequency of the fundamental mode (1,0) increases by ~ 1.8 fold while the frequency of the mixed mode (1,1) increases by ~ 14.6 fold. As a result, the first few resonant modes are predominantly longitudinal for both rectangular and elliptical cracks at low aspect ratios (below 0.2). For a similar mode, the resonant frequency of an elliptical crack is consistently higher than that of the rectangular crack. This is expected as the elliptical crack is narrower in the transverse direction and thus stiffer than a rectangular crack of the same length and aspect ratio.

Another clear feature, for both rectangular and elliptical cracks, is that frequen-350 cies of modes with same wavelengths in the transverse direction converge as α decreases. 351 For instance, frequencies of mixed modes (1, 1) and (2, 1) converge to the values of trans-352 verse mode (0, 1). Similar convergence also exists for modes (0, 2) and (1, 2). This is 353 expected because the crack wave speed, in the limit of low aspect ratio, is primarily con-354 trolled by the short wavelength in the transverse direction. As α increases, the frequen-355 cies of different modes become more intermingled and mode degeneration occurs, where 356 modes with distinct eigenfunctions share the same frequency. It is well known that mode 357 degeneration occurs at $\alpha = 1$ due to the geometric symmetry of a square or circle. What 358 we show here is that mode degeneration also occurs at intermediate aspect ratios. For 359

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Figure 8. Resonant frequencies of representative modes of rectangular cracks as a function of the crack stiffness C_L . The aspect ratio α is set as 0.5



Figure 9. Same as Figure 8 but for elliptical cracks

instance, modes (3, 0) and (0, 1) for both a rectangular and elliptical crack share similar frequencies when $\alpha \approx 0.35$.

362 4.2 Effect of crack stiffness ratio

Since the normalization constant $(c_f/\sqrt{C_L})/L$ for frequency changes with C_L , we visualize the actual resonant frequency f, instead of f^* . We use $c_f = 1$ m/s and L =1 m to scale f^* to f. C_L is the key dimensionless parameter that controls the crack wave propagation: the higher the value of C_L , the lower the phase velocity (e.g., B. Chouet,

³⁶⁷ 1986; Maeda & Kumagai, 2017). As a result, the resonant frequencies of all modes for

both rectangular (Figure 8) and elliptical cracks (Figure 9) decrease continuously as C_L

increases. Again, for a similar mode, the resonant frequencies of an elliptical crack is con-

sistently higher than those of a rectangular crack given the same axial lengths.

371

5 Energy dissipation

Since we currently focus on computing crack resonant frequencies in complex crack geometries, we assume both an inviscid fluid and a quasi-static solid and we do not consider energy dissipation, from either fluid viscosity or seismic radiation. When damping exists, the resonant frequency becomes complex and the rate of decay is quantified by the quality factor

$$Q = \frac{\operatorname{Re}\left(f\right)}{2\operatorname{Im}\left(f\right)},\tag{24}$$

which is also the number of cycles for an oscillation's energy to fall off to $e^{-2\pi} \sim 0.2\%$ 377 of its original value. The effect of viscous damping has been investigated analytically with 378 fully dynamic (Korneev, 2008) and quasi-static solid response (Lipovsky & Dunham, 2015) 379 on an infinite crack. However, the applicability of the Q formula on a finite crack has 380 not yet been tested. In addition, the convoluted derivation in Korneev (2008) makes it 381 difficult to quantify the relative contribution of different dissipative sources to the to-382 tal energy loss. On the other hand, numerical studies on rectangular cracks (e.g., Ku-383 magai & Chouet, 2000) have investigated the Q caused by seismic radiation but adopted 384 a simplistic treatment of the fluid viscosity, either an inviscid or fully-developed flow. In 385 this section, we offer a semi-analytical discussion of energy dissipation under a few as-386 sumptions and attempt to address two questions: (1) does the formula of Q developed 387 by Lipovsky and Dunham (2015) for an infinite crack also apply to a finite crack? (2)388 which of the two sources of energy dissipation, fluid viscosity and seismic radiation, is 389 more significant? 390

391

5.1 The applicability of Q formula from dispersion to a finite crack

We consider a viscous fluid with kinematic viscosity μ . To focus on the effect of the finite geometry, we compare analytical solutions by Lipovsky and Dunham (2015) to numerical solutions by Liang et al. (2020) for a rectangular crack, both of which as³⁹⁵ sume a quasi-static solid response. For simplicity, we focus on the boundary layer limit

$$\zeta = w_0 / \sqrt{4\nu/\omega} \gg 1, \tag{25}$$

where the crack aperture w_0 is much larger than the thickness of the viscous boundary layer $\sqrt{4v/\omega}$. In this limit, Q is high and ω can be well approximately by the inviscid solution. The analytical formula of Q for crack waves with real wavenumber is given by equation (80) in Lipovsky and Dunham (2015) and, after neglecting the small imaginary part of phase velocity when $\zeta \gg 1$, we have:

$$Q = \sqrt{2}\zeta.$$
 (26)

The hypothesis is that this expression for Q also holds, at least approximately for a fi-401 nite rectangular crack, regardless of its geometric shape, as long as w_0 , μ and ω are known. 402 We perform numerical simulations using the program by Liang et al. (2020), who em-403 ployed a finite difference method on a stretched grid to deal with the narrow viscous bound-404 ary layer. We set L = 100 m, $K_f = 1$ Pa, G = 1 Pa, $w_0 = 1$ m, which results in a C_L 405 of 100, and solve for the inviscid resonant angular frequencies ω of rectangular cracks 406 of two aspect ratios, 0.5 and 1.0. We then adjust μ so that ζ takes the values of 10, 20, 407 40, 60, 80, 100 and 200. We consider the first two modes of the crack with aspect ratio 408 of 0.5 and the fundamental mode of the square crack to represent different mode types 409 and crack shapes. The Q values of viscous cases are obtained using the methodology by 410 Liang et al. (2020) and the comparison to equation (26) is shown in Figure 10. 411

As shown in Figure 10, the prediction by the analytical formula in Lipovsky and 412 Dunham (2015) matches well the numerical solutions. The agreement gets better at large 413 ζ , where the assumption of boundary layer limit becomes more accurate. The differences 414 between the numerical and analytical solutions are less than 5% at Q > 40, while the 415 difference at Q = 10 is ~ 14%. Another encouraging finding is that aspect ratios and 416 mode numbers of rectangular cracks have a negligible impact on the value of Q as long 417 as ζ is the same. We thus postulate that the Q formula is likely to hold also for other 418 crack shapes or even a crack network. We further propose that one may first approxi-419 mate the resonant frequency ω of complex shaped cracks using the inviscid solution ef-420 ficiently determined by our method, and then directly estimate Q using the analytical 421 formula. However, future numerical studies considering both complex crack geometry 422 and fluid viscosity are necessary to rigorously test this hypothesis. 423

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Figure 10. Quality factor of various resonant modes of rectangular cracks as a function of the boundary layer thickness ratio ζ .

5.2 The competition between radiation and viscous damping

424

Instead of considering the fully dynamic solid response (e.g., Korneev, 2008), we 425 assume a quasi-dynamic solid response (e.g., Rice, 1993; Geubelle & Rice, 1995), which 426 allows to explicit extract the instantaneous long wavelength emission perpendicular to 427 the crack surface, the radiation damping (RD) term. We also consider an infinite crack 428 in two dimensions for the ease of theoretical treatment following Lipovsky and Dunham 429 (2015). By neglecting the wave mediated stresses and the seismic diffraction at the fi-430 nite crack tips, the radiation we consider is an underestimate, but it is still useful for un-431 derstanding the relative importance of various dissipation sources. Since resonances tend 432 to be overdamped in the fully developed flow limit $\zeta \gg 40$ (Korneev, 2008; Lipovsky 433 & Dunham, 2015), we continue to focus on the boundary layer limit $\zeta \ll 40$. We ex-434 plicitly identify the radiation and viscous damping terms in the governing equation and 435 then compute the ratio of their magnitudes. 436

The width-averaged crack wave equation considering viscous wall traction is obtained by combining the mass and momentum balance equations in Lipovsky and Dunham (2015),

$$\frac{\rho_f}{K_f}\frac{\partial^2 p}{\partial t^2} + \frac{\rho_f}{w_0}\frac{\partial^2 w}{\partial t^2} - \frac{\partial^2 p}{\partial x^2} - \frac{2}{w_0}\frac{\partial \tau}{\partial x} = 0,$$
(27)

where τ is the wall shear traction. We introduce the double Fourier transform of an ar-

441 bitrary function F(x,t) as

$$\hat{F}(k,\omega) = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} F(x,t) e^{-i(kx-\omega t)} dt dx.$$
(28)

 A_{442} Applying it to equation (27) leading to

$$-\omega^2 \rho_f \left(\frac{\hat{p}}{K_f} + \frac{\hat{w}}{w_0}\right) + k^2 \hat{p} - ik \frac{2}{w_0} \hat{\tau} = 0.$$
⁽²⁹⁾

⁴⁴³ Using equation (38) in Lipovsky and Dunham (2015) and neglecting horizontal wall mo-

tion, the wall shear traction is related to fluid pressure by

$$\hat{\tau} = -ikw_0\Omega\hat{p}/2,\tag{30}$$

445 where

$$\Omega = \left(\sqrt{i}/\zeta\right) \tanh\left(\zeta/\sqrt{i}\right),\tag{31}$$

and tends to \sqrt{i}/ζ in the boundary layer limit. Therefore, the viscous damping (VD) term in the equation is

$$VD = k^2 \Omega \hat{p}. \tag{32}$$

Applying the quasi-dynamic solid response (Geubelle & Rice, 1995), the fluid pressure and crack opening are related by

$$\hat{p} = \frac{Gk\hat{w}}{2\left(1 - v_s\right)} - i\omega\eta_R\hat{w},\tag{33}$$

where the two terms on the right hand side are the quasi-static response and radiation 450 damping (RD), respectively, and $\eta_R = \rho_s c_p/2$ is the radiation damping coefficient. The 451 ratio between the QS and RD terms is approximately $c_s^2/(c_p c)$ (dropping terms involv-452 ing Poisson's ratio), where $c = \omega/k$ is the crack wave phase velocity. In the low-frequency 453 limit, which we are interested in, the crack wave speed is much smaller than the speeds 454 of the solid body waves, $c \ll c_s \sim c_p$, and thus the RD term is much smaller than the 455 QS term. Substituting equation (33) into (29) and approximating \hat{w}/\hat{p} using the QS part, 456 we obtain the RD term in equation (29) as 457

$$RD = i\omega^3 \frac{\rho_f n_R}{K_f} \hat{w} \approx i\omega^3 \frac{\rho_f n_R}{K_f} \frac{2(1-v_s)}{Gk} \hat{p}.$$
(34)

458 The ratio between RD and VD is

$$\frac{RD}{VD} = ic^3 \frac{\rho_f n_R}{K_f} \frac{2(1-\upsilon_s)}{G\Omega} = i \frac{c_p c^3}{c_f^2 c_s^2} \frac{(1-\upsilon_s)}{\Omega}.$$
(35)

459 In the boundary layer limit, the magnitude of this ratio becomes

$$\left|\frac{RD}{VD}\right| = \frac{c_p c^3}{c_f^2 c_s^2} \zeta \left(1 - v_s\right). \tag{36}$$

- 460 When the overall damping is small, the crack wave phase velocity as a function of wave-
- length λ is well approximated by the inviscid dispersion relation:

$$c = \sqrt{\frac{2\pi G w_0}{\lambda \rho_f \left(1 - \upsilon_s\right)}} \sim c_f / sqrt C_\lambda,\tag{37}$$

where $C_{\lambda} = K_f \lambda / G w_0$ is a crack stiffness ratio similar to C_L but replacing L by λ . Finally, we obtain RD/VD, which scales as

$$\left|\frac{RD}{VD}\right| \sim \frac{c_p c_f}{c_s^2} \frac{\zeta}{C_\lambda^{3/2}} = \frac{c_p c_s}{c_f^2} \zeta \left(\frac{w_0 \rho_s}{\lambda \rho_f}\right)^{3/2},\tag{38}$$

after dropping small constants such as 2, π and μ_s . |RD/VD| is governed by three di-464 mensionless parameters: $\frac{c_p c_f}{c_s^2}$, C_{λ} and ζ . The first parameter one is controlled by the body 465 wave speeds of the solid and fluid and is not related to the crack geometry. For a typ-466 ical crustal rock and liquid fluid, for instance with $c_f = 1500$ m/s, $c_p = 4500$ m/s, and 467 $c_s=2500~{\rm m/s},\,\frac{c_pc_f}{c_\circ^2}$ is near unity. However, exsolved gases in liquid fluid, common in 468 shallow volcanic or geothermal environments (e.g., Kumagai & Chouet, 1999, 2001), can 469 significantly decrease the sound speed of the mixture, resulting in a much smaller $\frac{c_p c_f}{c^2}$. 470 The trade-off between C_{λ} and ζ in controlling |RD/VD| is displayed in Figure 11. In 471 the regime of high ζ and low C_{λ} , seismic radiation dominates over viscous damping, while 472 in the regime of low ζ and high C_{λ} vice versa. Note that increasing λ or decreasing fre-473 quency ω while fixing other parameters increases C_{λ} and simultaneously decreases ζ , both 474 of which lead to a lower percentage of damping in radiation. 475

6 Application to VLP seismic signals during the Mayotte volcano-seismic crisis

Since 10 May 2018, an unprecedented submarine volcano-seismic crisis occurred 478 30 km east of Mayotte Island (France), featuring a lithosphere-scale dyke intrusion and 479 drainage ($\sim 5 \text{ km}^3$) of deep magma reservoirs and producing exceptionally deep seismic-480 ity and substantial surface deformation (Cesca, Letort, et al., 2020; Feuillet et al., 2021; 481 Saurel et al., 2021; Mittal et al., 2022; Mercury et al., 2022; Retailleau et al., 2022). By 482 mid June of 2018, sustained long duration and highly oscillatory VLP seismic signals (see 483 an example in Figure 12a) have been observed and persist since, which are associated 484 with resonances of magma-filled cracks excited by nearby volcano-tectonic (VT) events 485



Figure 11. |RD/VD| as a function of the crack stiffness ratio C_{λ} and the boundary layer thickness ratio ζ . Parameters used are $c_f = 1500$ m/s, $c_p = 4500$ m/s, and $c_s = 2500$ m/s.

or possible piston collapse movements (Cesca, Letort, et al., 2020; Feuillet et al., 2021). 486 The stack of spectra of multiple VLP events reveals multiple resonant modes, among which 487 the fundamental mode with period ~ 15.5 s is present in all events, but not all higher modes 488 are manifested in each event, probably due to differences in the excitation. The funda-489 mental frequency can be readily explained by the crack model upon choosing a proper 490 crack length and aperture (Cesca, Letort, et al., 2020). However, as shown in Figure 12b, 491 the uneven spacing between resonant modes implies additional complexity in the source. 492 Particularly, the ratio between the first higher mode and the fundamental mode is $f_2/f_1 \approx$ 493 2.5. As shown in Figures 6 and 7, this value can not be explained by a simple rectan-494 gular or elliptical crack. Here, we show this observation can be explained by a dumbbell-495 shaped crack (Figure 12c). This crack shape is compatible with the f_2/f_1 data, but might 496 still differ from the real crack geometry in Mayotte as we have not made a systematic 497 attempt to also match the frequencies of other higher modes. However, this example is 498 sufficient to demonstrate the potential application of the developed method. One pro-499 found question is perhaps whether one can reconstruct the topology of the crack given 500 the information of all the resonant frequencies. Mark Kac also asked a similar question 501 "Can one hear the shape of a drum?" (Kac, 1966). Unfortunately, the answer is nega-502 tive: there exist multiple isospectral geometries that share the same resonant frequen-503

cies, as mathematically proven by Gordon et al. (1992). However, these isospectral geometries are rare even though they do exist and one can still decipher the shape of the
resonator given additional constraints of the vibration pattern, which in practice requires
dense geophysical observation particularly in the near field. A formal inversion procedure would need to be developed in the future to find the optimal crack geometry or topology of interconnected crack networks that best explains all the observed resonant frequencies and other geophysical constraints.



Figure 12. (a) Normalized vertical acceleration waveform of an representative VLP event (on 11 November 2018, bandpass filtered to 0.02-0.1 Hz) at the nearest broadband seismic station YTMZ on land, during the volcano-seismic crisis near Mayotte. (b) Stacked spectrum of 21 strong VLP signals compiled by Cesca, Letort, et al. (2020), highlighting multiple unevenly spaced resonant modes (dashed lines). Particularly, the frequency ratio between the first two modes $f_2/f_1 \approx 2.5$. The blue ticks indicate the integer multiples of the fundamental frequency. (c) Eigenmodes of a possible crack shape that satisfies $f_2/f_1 \approx 2.5$.

511 7 Summary

We have developed a hybrid method that couples the boundary element and finite 512 volume method to efficiently compute the resonant modes of fluid-filled cracks with com-513 plex geometry. Particularly, the BEM reduces three dimensional cracks to 2D surfaces, 514 substantially decreasing the number of degrees of freedom. By performing eigenmode 515 analysis in the frequency domain, we avoid errors from both the time discretization and 516 spectral analysis of the time domain data. We solve the problem in dimensionless form 517 so that the results can be conveniently scaled to other crack sizes. After proper verifi-518 cation, we apply our method to an example of a crack network, revealing distinct res-519 onant frequencies and vibration patterns, which may be utilized to infer more accurately 520 crack shapes from seismic data. 521

We then systematically analyze the influence of crack aspect ratio and crack stiff-522 ness on the resonant frequencies for both rectangular and elliptical cracks, which are com-523 mon models for interpreting real data. In general, rectangular and elliptical cracks share 524 similar eigenmode types and frequencies, while the elliptical crack has slightly higher res-525 onance frequencies due to the reduced length of the minor axis. At a high aspect ratio, 526 the frequencies of various mode types (longitudinal, transverse and mixed) are intermin-527 gled and mode degeneration occurs. Reducing the aspect ratio increases the frequencies 528 of all the modes, but more intensely for transverse and mixed modes than for longitu-529 dinal modes. In addition, at low aspect ratio, frequencies of modes (transverse or mixed) 530 with the same wavelengths in the transverse direction converge and differentiating them 531 requires additional knowledge of their vibration patterns. On the other hand, increas-532 ing C_L results in a decrease in resonant frequencies for all modes, regardless of the crack 533 geometry, which is primarily due to the decrease in crack wave propagation speed. 534

The major part of this work does not consider fluid viscosity or seismic radiation, 535 and thus cannot be used to directly compute the quality factor Q. However, by making 536 a few assumptions, we offer additional theoretical discussion on the energy dissipation. 537 First, by comparing numerical to analytical solutions, we confirm that the simple for-538 mula $Q = \sqrt{2}\zeta$ derived by Lipovsky and Dunham (2015) is a rather good approxima-539 tion for a rectangular crack when the thickness of the viscous boundary layer is much 540 smaller than the crack width, regardless of crack aspect ratio or vibrational mode. This 541 is an encouraging finding that suggests one may first obtain the inviscid resonant fre-542

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quencies using our method and then apply analytical formula to compute Q. Note that 543 this formula still does not consider seismic radiation. We then derived the relative ra-544 tio of the radiation damping to viscous damping, assuming a quasi-dynamic solid response 545 on an infinite crack. We show that this ratio is primarily controlled by three dimension-546 less parameters: $c_p c_f / c_s^2$, C_λ and ζ . Particularly, in the limit of high ζ and low C_λ , seis-547 mic radiation dominates over viscous damping while the opposite is true in the limit of 548 low ζ and high C_{λ} . Note that the seismic radiation considered here is a lower bound as 549 we neglected the wave-mediated stresses and the seismic radiation at the finite crack tip. 550 However, our theoretical development still offers a valuable insight into the partition of 551 damping in crack waves. 552

Finally, we obtain one possible crack shape, a "dumbbell", that successfully explains 553 the ratio of frequencies of the first two modes in the VLP seismic data during the 2018 554 Fani Maoré, Mayotte submarine volcanic eruption. This shape is one possibility and may 555 be updated when additional higher modes and geophysical constraints are integrated into 556 the analysis. In addition, the method developed here can be directly applied to other 557 scenarios, such as unconventional oil and gas fields and glacier hydraulics. Future work 558 requires a rigorous treatment of fluid viscosity, elastodynamics, and coupling to other 559 geometries such as conduits and equidimensional chambers. 560

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9 Data Availability Statement

The source code and the input files associated with the simulation cases are included in the Zenodo data respository at Liang et al. (2023). The VLP catalog of the Mayotte crisis is provided by Cesca, Heimann, et al. (2020) and is freely available online.

571 **References**

572	Aki, K., Fehler, M., & Das, S. (1977). Source mechanism of volcanic tremor: Fluid-
573	driven crack models and their application to the 1963 Kīlauea eruption. $Jour$
574	nal of Volcanology and Geothermal Research, $2(3)$, 259–287. doi: 10.1016/0377
575	-0273(77)90003-8

- Aster, R. C. (2019). Interrogating a surging glacier with seismic interferometry. Geo *physical Research Letters*, 46(14), 8162–8165. doi: 10.1029/2019GL084286
- Berre, I., Doster, F., & Keilegavlen, E. (2019). Flow in fractured porous media:
 A review of conceptual models and discretization approaches. Transport in
 Porous Media, 130(1), 215–236. doi: 10.1007/s11242-018-1171-6
- Cao, H., Medici, E., & Askari, R. (2021). Physical modeling of fluid-filled fractures
 using the dynamic photoelasticity technique. *Geophysics*, 86(1), T33–T43. doi:
 10.1190/geo2020-0037.1
- ⁵⁸⁴ Cesca, S., Heimann, S., Letort, J., Razafindrakoto, H., Dahm, T., & Cotton, F.
- Seismic catalogues of the 2018–2019 volcano-seismic crisis offshore
 mayotte, comoro islands. v. 1.0 (october 2019). available at gfz data ser-
- ssr
 vices.
 Retrieved from https://doi.org/10.5880/GFZ.2.1.2019.004
 doi:

 ssr
 10.5880/GFZ.2.1.2019.004
 doi:
- Cesca, S., Letort, J., Razafindrakoto, H. N., Heimann, S., Rivalta, E., Isken, M. P.,
 ... others (2020). Drainage of a deep magma reservoir near mayotte inferred
 from seismicity and deformation. *Nature Geoscience*, 13(1), 87–93. doi:
 10.1038/s41561-019-0505-5
- ⁵⁹³ Chouet, B. (1986). Dynamics of a fluid-driven crack in three dimensions by the finite
 difference method. Journal of Geophysical Research: Solid Earth, 91 (B14),
 ⁵⁹⁵ 13967–13992. doi: 10.1029/JB091iB14p13967
- ⁵⁹⁶ Chouet, B. A., Page, R. A., Stephens, C. D., Lahr, J. C., & Power, J. A. (1994).
 ⁵⁹⁷ Precursory swarms of long-period events at redoubt volcano (1989–1990),
 ⁵⁹⁸ alaska: their origin and use as a forecasting tool. Journal of Volcanology and
- ⁵⁹⁹ Geothermal Research, 62(1-4), 95–135. doi: 10.1016/0377-0273(94)90030-2
- Cruz, F. G., & Chouet, B. A. (1997). Long-period events, the most characteristic
 seismicity accompanying the emplacement and extrusion of a lava dome in
 galeras volcano, colombia, in 1991. Journal of Volcanology and Geothermal
 Research, 77(1-4), 121–158. doi: 10.1016/S0377-0273(96)00091-1

604	Durran, D. R. (2013). Numerical methods for wave equations in geophysical fluid dy-
605	namics (Vol. 32). Springer Science & Business Media.
606	Fehler, M., & Aki, K. (1978). Numerical study of diffraction of plane elas-
607	tic waves by a finite crack with application to location of a magma lens.
608	$Bulletin \ of \ the \ Seismological \ Society \ of \ America, \ 68(3), \ 573-598. $ doi:
609	10.1785/BSSA0680030573
610	Ferrazzini, V., & Aki, K. (1987). Slow waves trapped in a fluid-filled infinite crack:
611	Implication for volcanic tremor. Journal of Geophysical Research: Solid Earth,
612	92(B9), 9215-9223.doi: 10.1029/JB092iB09p09215
613	Feuillet, N., Jorry, S., Crawford, W. C., Deplus, C., Thinon, I., Jacques, E.,
614	\dots others (2021). Birth of a large volcanic edifice offshore mayotte via
615	lithosphere-scale dyke intrusion. Nature Geoscience, $14(10)$, 787–795. doi:
616	10.1038/s41561-021-00809-x
617	Frehner, M. (2013). Krauklis wave initiation in fluid-filled fractures by a passing
618	body wave. In Poromechanics v: Proceedings of the fifth biot conference on
619	poromechanics (pp. 92–100). doi: 10.1061/9780784412992.011
620	Frehner, M., & Schmalholz, S. M. (2010). Finite-element simulations of stoneley
621	guided-wave reflection and scattering at the tips of fluid-filled fractures. Geo -
622	physics, 75(2), T23-T36. doi: 10.1190/1.3340361
623	Fujita, E., & Ida, Y. (2003). Geometrical effects and low-attenuation reso-
624	nance of volcanic fluid inclusions for the source mechanism of long-period
625	earthquakes. Journal of Geophysical Research: Solid Earth, $108(B2)$. doi:
626	10.1029/2002JB001806
627	Geubelle, P. H., & Rice, J. R. (1995). A spectral method for three-dimensional elas-
628	todynamic fracture problems. Journal of the Mechanics and Physics of Solids,
629	43(11), 1791-1824.doi: 10.1016/0022-5096(95)00043-I
630	Gordon, C., Webb, D. L., & Wolpert, S. (1992). One cannot hear the shape of a
631	drum. Bulletin of the American Mathematical Society, $27(1)$, 134–138. doi: 10
632	.1090/S0273-0979-1992-00289-6
633	Gräff, D., Walter, F., & Lipovsky, B. P. (2019). Crack wave resonances within the
634	basal water layer. Annals of Glaciology, $60(79)$, 158–166. doi: 10.1017/aog
635	.2019.8
636	Henry, F., Fokkema, J., & De Pater, C. (2002). Experiments on stoneley wave prop-

-30-

637	agation in a borehole intersected by a finite horizontal fracture. In $64th~eage$
638	conference & exhibition (pp. cp–5). doi: 10.3997/2214-4609-pdb.5.P143
639	Jin, Y., Zheng, Y., Huang, L., & Ehlig-Economides, C. (2022). Characterizing
640	hydraulic fractures using the transient pressure surge effect. In $Spe/aapg/seg$
641	unconventional resources technology conference (p. D021S028R002). doi:
642	10.15530/urtec-2022-3718981
643	Kac, M. (1966) . Can one hear the shape of a drum? The american mathematical
644	monthly, 73 (4P2), 1–23. doi: 10.1080/00029890.1966.11970915
645	Karimi-Fard, M., Durlofsky, L. J., & Aziz, K. (2004). An efficient discrete-fracture
646	model applicable for general-purpose reservoir simulators. SPE journal, $9(02)$,
647	227–236. doi: 10.2118/88812-PA
648	Kawakatsu, H., Kaneshima, S., Matsubayashi, H., Ohminato, T., Sudo, Y., Tsutsui,
649	T., Legrand, D. (2000) . Aso94: Aso seismic observation with broadband
650	instruments. Journal of Volcanology and Geothermal Research, 101(1-2),
651	129–154. doi: 10.1016/S0377-0273(00)00166-9
652	Korneev, V. (2008) . Slow waves in fractures filled with viscous fluid. Geophysics,
653	73(1), N1–N7. doi: 10.1190/1.2802174
654	Korneev, V., Danilovskaya, L., Nakagawa, S., & Moridis, G. (2014). Krauklis wave
655	in a trilayer. $Geophysics$, 79(4), L33–L39. doi: 10.1190/geo2013-0216.1
656	Krauklis, P. V. (1962). On some low-frequency oscillations of a fluid layer in an elas-
657	tic medium. Prikl. Mat. Mekh., $26(6)$, 1111–1115. doi: 10.1016/0021-8928(63)
658	90084-4
659	Kumagai, H., & Chouet, B. A. (1999) . The complex frequencies of long-period seis-
660	mic events as probes of fluid composition beneath volcanoes. Geophysical Jour-
661	nal International, $138(2)$, F7–F12. doi: 10.1046/j.1365-246X.1999.00911.x
662	Kumagai, H., & Chouet, B. A. (2000). Acoustic properties of a crack containing
663	magmatic or hydrothermal fluids. Journal of Geophysical Research: Solid
664	Earth, 105(B11), 25493–25512. doi: 10.1029/2000JB900273
665	Kumagai, H., & Chouet, B. A. (2001). The dependence of acoustic properties of
666	a crack on the resonance mode and geometry. Geophysical research letters,
667	28(17), 3325-3328. doi: 10.1029/2001GL013025
668	Kumagai, H., Miyakawa, K., Negishi, H., Inoue, H., Obara, K., & Suetsugu, D.
669	(2003). Magmatic dike resonances inferred from very-long-period seismic

-31-

670	signals. Science, 299(5615), 2058–2061. Retrieved from http://science
671	.sciencemag.org/content/299/5615/2058 doi: 10.1126/science.1081195
672	Li, L., & Lee, S. H. (2008). Efficient field-scale simulation of black oil in a nat-
673	urally fractured reservoir through discrete fracture networks and homoge-
674	nized media. SPE Reservoir evaluation & engineering, $11(04)$, 750–758. doi:
675	10.2118/103901-PA
676	Liang, C., Karlstrom, L., & Dunham, E. M. (2020). Magma oscillations in a
677	conduit-reservoir system, application to very long period (vlp) seismicity at
678	basaltic volcanoes: 1. theory. Journal of Geophysical Research: Solid Earth,
679	125(1), e2019JB017437. doi: 10.1029/2019JB017437
680	Liang, C., O'Reilly, O., Dunham, E. M., & Moos, D. (2017). Hydraulic fracture di-
681	agnostics from krauklis-wave resonance and tube-wave reflections. <i>Geophysics</i> ,
682	82(3), D171-D186.doi: 10.1190/geo2016-0480.1
683	Liang, C., Peng, J., Ampuero, JP., Shauer, N., & Dai, K. (2023, August). Dataset
684	for "Resonances in fluid-filled cracks of complex geometry and application
685	to very long period (VLP) seismic signals at Mayotte submarine volcano".
686	Zenodo. Retrieved from https://doi.org/10.5281/zenodo.8275079 doi:
687	10.5281/zenodo.8275079
688	Lipovsky, B. P., & Dunham, E. M. (2015). Vibrational modes of hydraulic frac-
689	tures: Inference of fracture geometry from resonant frequencies and attenua-
690	tion. Journal of Geophysical Research: Solid Earth, 120(2), 1080–1107. doi:
691	10.1002/2014JB011286
692	Lokmer, I., Saccorotti, G., Di Lieto, B., & Bean, C. J. (2008). Temporal evolution
693	of long-period seismicity at etna volcano, italy, and its relationships with the
694	2004–2005 eruption. Earth and Planetary Science Letters, 266(1-2), 205–220.
695	doi: 10.1016/j.epsl.2007.11.017
696	Maeda, Y., & Kumagai, H. (2013). An analytical formula for the longitudinal res-
697	onance frequencies of a fluid-filled crack. Geophysical Research Letters, $40(19)$,
698	5108–5112. doi: $10.1002/grl.51002$
699	Maeda, Y., & Kumagai, H. (2017). A generalized equation for the resonance fre-
700	quencies of a fluid-filled crack. Geophysical Journal International, $209(1)$, 192–
701	201. doi: 10.1093/gji/ggx019
702	McQuillan, M., & Karlstrom, L. (2021). Fluid resonance in elastic-walled englacial

703	transport networks. Journal of Glaciology, $67(266),999{-}1012.$ doi: 10.1017/jog
704	.2021.48
705	Mercury, N., Lemoine, A., Doubre, C., Bertil, D., van Der Woerd, J., Hoste-
706	Colomer, R., & Battaglia, J. (2022). Onset of a submarine eruption east of
707	mayotte, comoros archipelago: the first ten months seismicity of the seismo-
708	volcanic sequence (2018–2019). Comptes Rendus. Géoscience, 354 (S2), 105–
709	136. doi: 10.5802/crgeos.191
710	Métaxian, JP., Araujo, S., Mora, M., & Lesage, P. (2003). Seismicity related to the
711	glacier of cotopaxi volcano, ecuador. Geophysical Research Letters, $30(9)$. doi:
712	10.1029/2002GL016773
713	Mittal, T., Jordan, J. S., Retailleau, L., Beauducel, F., & Peltier, A. (2022). May-
714	otte 2018 eruption likely sourced from a magmatic mush. Earth and Planetary
715	Science Letters, 590, 117566. doi: 10.1016/j.epsl.2022.117566
716	Moinfar, A., Varavei, A., Sepehrnoori, K., & Johns, R. T. (2013, 07). Develop-
717	ment of an Efficient Embedded Discrete Fracture Model for 3D Compositional
718	Reservoir Simulation in Fractured Reservoirs. SPE Journal, $19(02)$, 289-303.
719	doi: 10.2118/154246-PA
720	Nakagawa, S., Nakashima, S., & Korneev, V. A. (2016). Laboratory measurements
721	of guided-wave propagation within a fluid-saturated fracture. Geophysical
722	Prospecting, $64(1)$, 143–156. doi: 10.1111/1365-2478.12223
723	Nakano, M., & Kumagai, H. (2005). Response of a hydrothermal system to
724	magmatic heat inferred from temporal variations in the complex frequen-
725	cies of long-period events at kusatsu-shirane volcano, japan. Journal of
726	volcanology and geothermal research, 147(3-4), 233–244. doi: 10.1016/
727	j.jvolgeores.2005.04.003
728	Nikkhoo, M., & Walter, T. R. (2015). Triangular dislocation: an analytical, artefact-
729	free solution. Geophysical Journal International, 201(2), 1119–1141. doi: 10
730	.1093/gji/ggv035
731	Niu, J., & Song, TR. A. (2020). Real-time and in-situ assessment of conduit
732	permeability through diverse long-period tremors beneath as volcano,
733	japan. Journal of Volcanology and Geothermal Research, 401, 106964. doi:
734	10.1016/j.jvolgeores.2020.106964
735	Okada, Y. (1985). Surface deformation due to shear and tensile faults in a half-
736	space. Bulletin of the seismological society of America, $75(4)$, 1135–1154.
-----	--
737	Okada, Y. (1992). Internal deformation due to shear and tensile faults in a half-
738	space. Bulletin of the Seismological Society of America, $82(2)$, 1018–1040.
739	O'Reilly, O., Dunham, E. M., & Nordström, J. (2017). Simulation of wave propa-
740	gation along fluid-filled cracks using high-order summation-by-parts operators
741	and implicit-explicit time stepping. SIAM Journal on Scientific Computing,
742	39(4), B675–B702. Retrieved from https://doi.org/10.1137/16M1097511
743	doi: 10.1137/16M1097511
744	Paillet, F. L., & White, J. E. (1982, 08). Acoustic modes of propagation in
745	the borehole and their relationship to rock properties. $Geophysics, 47(8),$
746	1215-1228. Retrieved from https://doi.org/10.1190/1.1441384 doi:
747	10.1190/1.1441384
748	Pointer, T., Liu, E., & Hudson, J. A. (1998). Numerical modelling of seismic
749	waves scattered by hydrofractures: application of the indirect boundary el-
750	ement method. Geophysical Journal International, 135(1), 289–303. doi:
751	10.1046/j.1365-246X.1998.00644.x
752	Retailleau, L., Saurel, JM., Laporte, M., Lavayssière, A., Ferrazzini, V., Zhu,
753	W., \ldots others (2022). Automatic detection for a comprehensive view of
754	mayotte seismicity. Comptes Rendus. Géoscience, $354(S2)$, $153-170$. doi:
755	10.5802/crgeos.133
756	Rice, J. R. (1993). Spatio-temporal complexity of slip on a fault. Journal of Geo-
757	physical Research: Solid Earth, $98(B6)$, $9885-9907$. doi: $10.1029/93$ JB00191
758	Rona, A. (2007). The acoustic resonance of rectangular and cylindrical cavities.
759	Journal of Algorithms & Computational Technology, $1(3)$, $329-356$. doi:
760	10.1260/174830107782424110
761	Saurel, JM., Jacques, E., Aiken, C., Lemoine, A., Retailleau, L., Lavayssière, A.,
762	\ldots others (2021). May otte seismic crisis: building knowledge in near real-time
763	by combining land and ocean-bottom seismometers, first results. $Geophysical$
764	Journal International, $228(2)$, 1281–1293. doi: 10.1093/gji/ggab392
765	Segall, P. (2010). Earthquake and volcano deformation. Princeton University Press.
766	doi: $10.1515/9781400833856$
767	Shauer, N., Desmond, K. W., Gordon, P. A., Liu, F., & Duarte, C. A. (2021). A
768	three-dimensional generalized finite element method for the simulation of wave

-34-

769	propagation in fluid-filled fractures. Computer Methods in Applied Mechanics
770	and Engineering, 386, 114136. doi: 10.1016/j.cma.2021.114136
771	Stuart, G., Murray, T., Brisbourne, A., Styles, P., & Toon, S. (2005). Seismic emis-
772	sions from a surging glacier: Bakaninbreen, svalbard. Annals of Glaciology, 42,
773	151–157. doi: 10.3189/172756405781812538
774	Sun, F., Gong, Y., & Dong, C. (2020). A novel fast direct solver for 3d elas-
775	tic inclusion problems with the isogeometric boundary element method.
776	Journal of Computational and Applied Mathematics, 377, 112904. doi:
777	10.1016/j.cam.2020.112904
778	Tang, X., & Cheng, C. (1988). Wave propagation in a fluid-filled fracture—an exper-
779	imental study. Geophysical Research Letters, $15(13)$, 1463–1466. doi: 10.1029/
780	GL015i013p01463
781	Tang, X., & Cheng, C. (1989). A dynamic model for fluid flow in open borehole frac-
782	tures. Journal of Geophysical Research: Solid Earth, $94(B6)$, 7567–7576. doi:
	10,1020/IB004;B06p07567
783	10.1029/3D0341D00p07307
783 784	Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of
783 784 785	Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments.
783 784 785 786	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi:
783 784 785 786 787	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904
783 784 785 786 787 788	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-
783 784 785 786 787 788 788	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir sim-
783 784 785 786 787 788 788 789 790	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi:
783 784 785 786 787 788 789 790 791	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA
783 784 785 786 787 788 789 790 791 792	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the
783 784 785 786 787 788 789 790 791 792 793	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>,
783 784 785 786 787 788 789 790 791 792 793 794	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>, 174(3), 1174–1186. doi: 10.1111/j.1365-246X.2008.03871.x
 783 784 785 786 787 788 789 790 791 792 793 794 795 	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>, 174(3), 1174–1186. doi: 10.1111/j.1365-246X.2008.03871.x Zheng, Y., Malallah, A. H., Fehler, M. C., & Hu, H. (2016). 2d full-waveform model-
 783 784 785 786 787 788 789 790 791 792 793 794 795 796 	 Tary, JB., Van der Baan, M., & Eaton, D. W. (2014). Interpretation of resonance frequencies recorded during hydraulic fracturing treatments. <i>Journal of Geophysical Research: Solid Earth</i>, 119(2), 1295–1315. doi: 110.1002/2013JB010904 Xu, Y., Cavalcante Filho, J., Yu, W., & Sepehrnoori, K. (2017). Discrete-fracture modeling of complex hydraulic-fracture geometries in reservoir simulators. <i>SPE Reservoir Evaluation & Engineering</i>, 20(02), 403–422. doi: 10.2118/183647-PA Yamamoto, M., & Kawakatsu, H. (2008). An efficient method to compute the dynamic response of a fluid-filled crack. <i>Geophysical Journal International</i>, 174 (3), 1174–1186. doi: 10.1111/j.1365-246X.2008.03871.x Zheng, Y., Malallah, A. H., Fehler, M. C., & Hu, H. (2016). 2d full-waveform modeling of seismic waves in layered karstic media. <i>Geophysics</i>, 81(2), T25–T34. doi:

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Appendix A Matrices D and T for a simple crack intersection

In this section, we show step by step how to construct matrices **D** and **T** for a simple crack intersection shown in Figure A1. The element number and positive flux direc-

- tion of each active connection as labeled. The boundary edges have zero flux and they
- $_{802}$ do not contribute to **D** and **T**. Thus, we have five elements and five active connections
- numbered as $\{2 \rightarrow 1, 3 \rightarrow 2, 4 \rightarrow 2, 5 \rightarrow 2, 5 \rightarrow 4\}$, where $i \rightarrow j$ defines the positive flux direction. The size of both **D** and **T** are 5 by 5.
 - T_{54} T_{52} T_{32} (3) $(4) \leftarrow 2$ T_{42} T_{21} (1)

Figure A1. Geometry of a simple crack intersection. The element number and the positive flow direction of each active connection (non-zero flux) are indicated by the circled number and arrow, respectively. The scalar transmisibilities are labled near each connection.

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Let's first consider the matrix \mathbf{D} , which sums the flux from active connections to 805 obtain the net out-flux from each element. We consider the first row of \mathbf{D} as an exam-806 ple, relevant for element 1. The only connection that contributes to the net out-flux of 807 element 1 is connection 1 with the positive direction of $2 \rightarrow 1$, the opposite to the out-808 flux direction. Thus, D(1,1) = -1 and other entries of the first row are zeros. How-809 ever, for element 2, the positive flux of connection 1 aligns with the outflux direction, 810 which leads to D(2,1) = 1. Similarly, other entries of matrix **D** can be determined and 811 the matrix \mathbf{D} is: 812

$$\mathbf{D} = \begin{bmatrix} -1 & 0 & 0 & 0 & 0 \\ 1 & -1 & -1 & -1 & 0 \\ 0 & 1 & 0 & 0 & 0 \\ 0 & 0 & 1 & 0 & -1 \\ 0 & 0 & 0 & 1 & 1 \end{bmatrix}.$$
 (A1)

We now proceed to construct the matrix \mathbf{T} , which computes the flux on each active connection from the pressure on each cell. Note that we only store the flux in the positive direction. For instance, the flux on the first connection is $Q_{2\rightarrow 1} = T_{21} (p_2 - p_1)$, which means $T(1,2) = -T(1,1) = T_{21}$. Similarly, other entries of the matrix \mathbf{T} can be com-

 $_{817}$ puted and the full expression of T is:

$$\mathbf{T} = \begin{bmatrix} -T_{21} & T_{21} & 0 & 0 & 0\\ 0 & -T_{32} & T_{32} & 0 & 0\\ 0 & -T_{42} & 0 & T_{42} & 0\\ 0 & -T_{52} & 0 & 0 & T_{52}\\ 0 & 0 & 0 & -T_{54} & T_{54} \end{bmatrix}.$$
 (A2)

Appendix B Resonant frequencies from time domain results by GFEM

In this section, we explain the procedure to obtain selective resonant frequencies 819 from the time domain simulation results using the GFEM code developed by Shauer et 820 al. (2021). As shown in Figure B1, we apply injection sources with a gaussian source time 821 function on the certain position on the crack (red stars), obtain the pressure time series 822 (duration of 50 s) on three receiving points (blue triangles), and then extract the res-823 onant frequencies at spectral peaks. For the rectangular crack, we place one source at 824 the upperleft corner, which manages to excite all the first eight modes, and three receivers 825 (R1, R2, and R3) at (-0.5, 0), (-0.20, 0.25), and (0, 0.25), respectively. Different receivers 826 sample different eigenmodes. For instance, receiver R1 samples modes 1, 2, 5, and 8 as 827 shown in Figure B1-c. The modes sampled by R2 and R3 are shown in Table 2. We make 828 this choice to selectively sample closely-spaced modes, for instance mode 2 and 3, at dif-829 ferent receivers to avoid ambiguity. 830

For the elliptical crack, we place two sources at the leftmost and uppermost ends, and three receivers at (-0.5, 0), (0, 0.25), (0, 0) respectively. Due to the excitation and monitoring geometry, we focus only sampling the longitudinal and transverse modes, which are clearly seperated peaks in the spectrum. The eigenmodes sampled by different receivers are shown in Figure B1-f and Table 3.



Figure B1. (a, d) The source and receiver positions. (b, e) Pressure time series at three receivers. (c, f) The normalized spectral amplitude of data at receiver R1. The vertial black dashed lines are the resonant frequencies (with mode number labelled) computed by BEM+FVM method.