# Linkages between mineral element composition of soils and sediments with hyporheic zone dissolved organic matter chemistry across the contiguous United States

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### Abstract

The hyporheic zone is a hotspot for biogeochemical cycling where interactions with mineral metals preserve the release and biodegradation of organic matter (OM). A small fraction of OM can still be exchanged between localized sediments and the overlying water column, and recent evidence suggest there exists a longitudinal structuring in sediment dissolved OM (DOM) chemistry across the continental United States (CONUS). In this study, we tested a hypothesis that water extractable sediment DOM chemistry could be explained by sediment metal contents and integrative watershed scale features at the CONUS scale. Crowdsourced samples were characterized for high resolution mass spectrometry and coupled with sediment metals determined via x-ray fluorescence as well as with land cover and soil elemental information obtained from national databases. Our results highlight weak relationships between DOM chemistry and elemental composition at the CONUS scale indicating limited transferability of organo-metal linkages into multi-scale hydrobiogeochemical models.

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 dissolved organic matter chemistry across the contiguous United States

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## 30 Plain language summary

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32 The hyporheic zone is a vital area for chemical cycling that influences the breakdown of organic 33 matter due to interactions with minerals and metals. Some organic matter moves between 34 sediment and water. Recent findings show a pattern in dissolved organic matter chemistry in 35 sediment across the US. The study tested whether sediment's organic matter chemistry could be 36 explained by its metals and surrounding features. Crowdsourced samples underwent analysis 37 along with sediment metals, land cover, and soil elements. Results revealed weak connections 38 between dissolved organic matter chemistry and elements, indicating limits in applying these 39 connections to continental scale models. 40 41 **Key Points** 42 43 1. Sediment dissolved organic matter chemistry is longitudinally structured across the continental 44 United States. 45 46 2. Sediment and soil mineral elemental compositions were weakly correlated with sediment 47 dissolved organic matter composition 48 49 3. Localized metal-organic matter linkages likely have limited capacity to inform continental-50 scale hydrobiogeochemical models. 51 52 53 Abstract 54 55 The hyporheic zone is a hotspot for biogeochemical cycling where interactions with mineral 56 metals preserve the release and biodegradation of organic matter (OM). A small fraction of OM 57 can still be exchanged between localized sediments and the overlying water column, and recent 58 evidence suggest there exists a longitudinal structuring in sediment dissolved OM (DOM) 59 chemistry across the continental United States (CONUS). In this study, we tested a hypothesis

60	that water extractable sediment DOM chemistry could be explained by sediment metal contents
61	and integrative watershed scale features at the CONUS scale. Crowdsourced samples were
62	characterized for high resolution mass spectrometry and coupled with sediment metals
63	determined via x-ray fluorescence as well as with land cover and soil elemental information
64	obtained from national databases. Our results highlight weak relationships between DOM
65	chemistry and elemental composition at the CONUS scale indicating limited transferability of
66	organo-metal linkages into multi-scale hydrobiogeochemical models.
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68	Key Words
69	hyporheic zone; continental; crowdsourced data; organo-mineral interactions; river corridor;
70	hydro-biogeochemistry
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77	1. Introduction
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79	Hyporheic zones within river corridors are considered major hotspots for biogeochemical cycling
80	and are important for nutrient availability, water quality, and ecosystem function (Boano et al.,
81	2014; Findlay, 1995). Within hyporheic zones, organic matter (OM) and nutrients are exchanged
82	between surface sediments and the overlying water column. A fraction of this OM is transported
83	in the dissolved phase (DOM) and is easily metabolized, whereas other fractions are selectively
84	preserved in sediments through interactions with sediment minerals (Lalonde et al., 2012). While
85	a majority of the OM stored in sediments is considered insoluble, a small fraction sorbed onto
86	sediments can be detached and is easily solubilized (Kang & Xing, 2008; Tao & Lin, 2000). This
87	water-soluble OM represents a significant source of sediment contributed DOM to rivers that
88	may be further consumed through microbially mediated processes (Tao, 1998; Zhou et al., 2019),
89	thus, implicating organo-mineral interactions as key contributors to biogeochemical processes in

90 hyporheic zones.

92 The relationships between DOM and sediment metals have been well studied (Kleber et al., 93 2021), and the transferability across scales is crucial to ensure the meaningful adaptation into 94 multi-scale river corridor hydrobiogeochemical models. Broadly, DOM-metal relationships are 95 driven by strong adsorptive processes at the sediment particle surface and are selective for DOM 96 fractions based on the overall sediment metal composition (J.-G. Lin & Chen, 1998). For 97 example, adsorption of DOM onto sediments containing iron oxides and hydroxides involves a 98 combination of both physical and chemical interactions and exhibit considerable adsorption 99 capacity for highly aromatic DOM (Chen et al., 2020; Li et al., 2023; Shields et al., 2016; 100 Wordofa et al., 2019). Similar observations have also been observed involving other metals such 101 as aluminum, manganese, and copper under varying environmental conditions (Coutinho et al., 102 2021; Lee et al., 2022; Scheel et al., 2007). Our current understanding of such organo-mineral 103 interactions is typically based on laboratory simulations and site level studies, which provide a 104 comprehensive understanding of these interactions at localized scales. However, limitations 105 remain in describing the transferability of such interactions across scales.

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107 There is some evidence to indicate that sediment DOM chemistry could be linked with sediment 108 metal concentrations at continental scales. Across the continental United States (CONUS), 109 geological formations have largely driven a longitudinal structuring in the elemental composition 110 of soil minerals and sediments (Grossman et al., 2004). Soils in the Western CONUS, for 111 example, are reported to have elevated aluminum and iron concentrations (Shacklette and 112 Boerngen 1984). Similarly, recent evidence suggest that the composition of DOM extracted from 113 hyporheic zone sediments also exhibits longitudinal structuring with clear shifts in 114 thermodynamic favorability along the east-west gradient (Vanessa A. Garayburu-Caruso et al., 115 2020); however, explanations for this observation have not yet been well constrained.

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In this study, we tested a hypothesis that longitudinal structuring of water extractable DOM from hyporheic zone sediments across the CONUS is linked with sediment metal distributions. More specifically, because iron is known to play a key role in the stabilization of aromatic, carboxylic rich DOM in sediments, we hypothesized that sediment iron would explain more variation in DOM chemistry than other sediment metals. To test these hypotheses, we used crowdsourced

122 data generated as part of the Worldwide Hydrobiogeochemical Observation Network for 123 Dynamic River Systems (WHONDRS) consortium in summer 2019 (Borton et al., 2022). Data 124 available included sediment elemental composition and ultra-high resolution mass spectrometry 125 for water extracted sediment DOM for samples across the CONUS. In addition, we extended our 126 analysis beyond site level elemental composition to include additional watershed scale features, 127 which would be representative of a more integrative regional scale contributions to the 128 downstream transport and deposition of particulate material to local sediments. As such, we used 129 land cover as an indicator of organic matter source that we hypothesized would be strongly 130 linked with the sediment DOM composition (Liu et al., 2019). Additionally, we included soil 131 elemental data from the US Geological Survey that may serve as an indicator of upstream 132 hyporheic zone sediment sources that interact with DOM during transport within the watershed. 133 For this, we hypothesized that soil elemental composition would not only correlate to hypothesized 134 zone sediment elemental composition (Salomão et al., 2021) but may also account for additional 135 variability in the DOM composition. Our results describe significant though relatively weak 136 linkages between both sediment and soil metals and DOM at the CONUS scale and have 137 significant implications for our conceptualization of organo-mineral interactions and their 138 scalability from localized processes into multi-basin predictive frameworks of hyporheic zone 139 biogeochemical function.

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141 2. Methods

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- 143 2.1 Sample Collection and Data Availability
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145 The data for this work was obtained from a publicly available database accessed on the 146 Environmental System Science Data Infrastructure for a Virtual Ecosystem (ESS-DIVE) 147 repository (Goldman et al., 2020). The samples were collected as a part of a CONUS scale 148 crowdsourced sampling campaign by the WHONDRS consortium in the Summer 2019 (Stegen 149 & Goldman, 2018). Complete methodological details for sample collections are available at 150 Garayburu-Caruso et al. (2020). Briefly, participants were sent a kit that included basic sampling 151 equipment for sediment collection along with detailed written instructions and a video to ensure 152 standardized sample collections. Within a 6-week period between August and September 2019,

153 sediments were collected according to NEON protocols (NEON.DOC.001193) at 97 sites, with 3 154 replicates taken from separate depositional locations within each site. Samples were shipped to 155 the Pacific Northwest National Laboratory within 24 hours of collection. Upon receiving, 156 sediments were sieved to <2mm, flash frozen, and stored at -80 °C until ultra-high resolution 157 mass spectrometry analysis. Sediment samples collected for mineralogy were stored at -20 °C.

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159 2.2 Soil Elemental and Land Use Data

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161 Watershed land cover was extracted for each site from the 2016 National Land Cover Database 162 (Dewitz, 2020) and represents the average relative land cover for the upstream drainage area for 163 each site (Garayburu-Caruso et al., 2022). Primary land cover classes considered include 164 croplands, grasslands, shrubs, urban, and forests. Soil chemical data was obtained from the 2004 165 National Geochemical Survey provided by the United States Geological Survey (Grossman et al., 166 2004). The spatial spread of the sample point locations that make up the National Geochemical 167 Survey allowed for a 500m resolution interpolation of the chemical composition across the 168 continental U.S using ordinary Kriging technique (Wackernagel, 1995). Kriging uses the known 169 variance of measured points in area to estimate the continuous composition of a larger region 170 (Pang et al., 2009). Interpolation was completed using point locations of calcium (Ca), copper 171 (Cu), iron (Fe), silicon (Si), and phosphorus (P). Average elemental concentrations within the 172 upstream catchment area from the sampling location were the extracted using the Geospatial 173 Data Puller for Waters in the Contiguous Unites States tool for R (Willi & Ross, 2023).

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# 176 2.3 X-ray Fluorescence

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Sediment elemental analysis was performed using a Bruker M4 plus x-ray florescence (XRF). Sediment samples were prepared by drying at 60 °C for 48 hours and were then ground and homogenized using a ball mill. Ground samples were randomized and packed into the center of custom acrylic disks and sealed with Kapton tape. The XRF samples were run with a XFlash VH50C-LE. Primary analytes measured included silica (Si), phosphorus (P), calcium (Ca), manganese (Mn), iron (Fe), and copper (Cu).

## 185 2.4 Sediment Extractions

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Extractions were performed on thawed sediments at a 1:2 ratio of sediment to MilliQ water while shaking in the dark at 375 rpm at 21°C for 2 hours. The samples were then centrifuged at 21 °C at 6000 rcf for 5 minutes. The supernatant was filtered through a 0.22µm polyethersulfone (PES) membrane filter. The sediment extracts were then normalized to DOC concentrations of 1.5 mg/L and acidified to pH 2 with 85% phosphoric acid. The DOM was then extracted using PPL (styrene divenyl benzene polymer, Agilent Bond Elut) cartridges using protocols described by Dittmar 2008.

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195 2.5 Fourier Transform ion Cyclotron Resonance Mass Spectrometry

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197 Samples were characterized by Fourier-transform ion cyclotron resonance mass spectrometry 198 (FTICR-MS) analysis at the Environmental Molecular Sciences Laboratory in Richland, WA. 199 Detailed methodological information is provided in Garayburu-Caruso et al., 2020. Briefly, 200 samples were injected into a 12 T Burker SolariX FTICR mass spectrometer equipped with an 201 electrospray ionization source. Samples were analyzed in negative (-) mode with an ion 202 accumulation of 0.1-0.2s. One-hundred forty-four scans were co-added and internally calibrated 203 with an OM homologous series separated by 14 Da (-CH<sub>2</sub> groups). The mass accuracy was 204 typically within 1 ppm for the collected mass range (100-900 m/z). BrukerDaltonik Data 205 Analysis (version 4.2) was then used convert spectra into a list of m/z values with a signal to 206 noise threshold set a 7. Peaks were then aligned, and chemical formula were assigned using 207 Formularity (Tolić et al., 2017) and considering only the presence of C, H, O, N, S, and P. 208 Common molecular indices, such as the aromaticity index (Koch & Dittmar, 2006, 2016) were 209 further calculated with the R package ftmsRanalysis (Bramer et al., 2020). We further binned 210 identified peaks based on their presence-absence distributions within a sample into common 211 biomolecular classes (e.g. lignin-like, protein-like) (Kim et al., 2003). We note that placement of 212 a peak within these classes does not constitute a positive identification but is rather a broad 213 acknowledgement of stoichiometric similarity between the molecular formula of the identified 214 peak and the biomolecular class.

# 216 2.6 Statistical Analysis

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218 All statistical analyses were performed with the R statistical platform version 4.2.2 (R Core 219 Team, 2015). Prior to analysis, XRF and the relative proportion of binned biomolecular classes 220 for the 3 replicates collected at each site were averaged. General linear relationships among 221 variables were then described with Pearson correlations. Redundancy analysis (RDA) were 222 performed with the vegan package (Okansen et al., 2020) to describe variability among FTICR-223 MS biomolecular classes (response variables) with sediment elemental composition and other 224 environmental parameters (predictor variables). The final model was built after downselecting 225 important variables using the 'ordistep' function, a stepwise model building function for 226 constrained ordination analyses (Okansen et al, 2020).

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228 3. Results and Discussion

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230 3.1 Sediment DOM is significantly but weakly related with sediment elemental composition

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232 At the CONUS scale, we observed that the composition of water extractable DOM released from 233 hyporheic zone sediments was highly variable and exhibited a longitudinal structuring that was 234 weakly correlated with local sediment elemental composition. In the eastern United States, 235 sediment DOM had a proportionally higher diversity of aromatic compounds compared to the 236 west coast, where more saturated and potentially more bioavailable protein and carbohydrate-like 237 molecular formula were detected (Figure 1). This distribution was not explicitly mimicked for 238 the sediment elemental composition as most elements had a poor relationship with longitude 239 (Figure 1). Given such, these elements (Si, P, Ca, Mn & Fe) also exhibited poor or insignificant 240 relationships with the diversity of measured DOM biochemical classes (Figure 2). Alternatively, 241 Cu was the primary element that provided consistent relationships with the observed CONUS 242 scale spatial distribution of DOM. More specifically, there was a significant relationship between 243 Cu and the relative fraction of molecular formula identified as highly aromatic, such as ligninlike ( $r^2 = 0.20$ , p < 0.001) and tannin-like ( $r^2 = 0.09$ , p < 0.05), and an inverse relationship with 244 the relative fraction of molecular formula identified as protein-like ( $r^2 = 0.16$ , p < 0.001) and 245

carbohydrate-like ( $r^2 = 0.27$ , p < 0.001). This is further highlighted in a redundancy analysis where Cu was the only element to provide significant explanatory power for the diversity of sediment DOM molecules present (Figure 3, Table 1).

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250 While the primary relationships with Cu alone within the RDA were somewhat surprising, it is 251 notable the impact of increased Cu can have on river corridor biogeochemistry that may explain 252 some of these trends. For instance, elevated Cu can have a significant impact on microbial 253 community structure and negatively impact heterotrophic respiration in river sediments (Ahmed 254 et al., 2018), and the availability of DOM for hyporheic exchange may be driven by broad scale 255 heterogeneity in microbial turnover rates and/or diversity of microbial communities (Radke & 256 Maier, 2014; Schaper et al., 2018). Such processes have been directly tied to the release and 257 preferential consumption of protein- and lipid-like DOM within the hyporheic zone (Zhou et al., 258 2019). Furthermore, organo-metal complexation involving Cu preferentially target aromatic rich 259 DOM fractions (Coutinho et al., 2021; Huang et al., 2021; Skrabal et al., 1997; Xu et al., 2019), 260 and Cu binding ligands involving aromatic humic substances can be easily mobilized out of 261 sediments into overlying surface waters (Lehman & Mills, 1994; Nissenbaum & Swaine, 1976). 262 While we did not directly measure organo-Cu complexation, this potential explanation would be 263 consistent with the linkages between Cu and a higher diversity of aromatic-like molecular 264 formula observed in this study. These relationships highlight Cu as a potential important linkage 265 to our conceptual understanding of DOM exchange within the hyporheic zone.

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267 Organo-metal interactions have been recognized as an important component of local sediment 268 biogeochemical processes. While our results indicate a potential importance of Cu over the 269 chemistry of DOM released from hyporheic zone sediments, we more directly hypothesized that 270 Fe would have stronger explanatory power given its strong affinity for DOM interaction. For 271 example, the adsorption to reactive Fe-oxyhydroxide surfaces is considered as an important 272 mechanism for OM preservation in sediments (Lalonde et al., 2012), and co-precipitation of 273 select aromatic DOM fractions with Fe during aquatic transport can provide a fresh source of 274 terrestrial OM to sediments (Chen et al., 2014; Du et al., 2018). Thus, we found it surprising that 275 there were no clear linkages between the observed DOM chemistry and Fe in this study (Figure 276 2). As such, this observation suggests a systematic disconnect in the ability to detect predictable

277 patterns of DOM exchange at the CONUS scale based on known physiochemical associations 278 relevant at local scales. There are several plausible explanations for this observation. First, 279 organic-Fe linkages are strongly driven by the presence of iron(III) (Lalonde et al., 2012). While 280 we assume there is high probability the bulk of iron in these sediments is in the form of iron(III), 281 this was not directly measured. Additionally, the analytical approach taken reflects bulk sediment 282 elemental composition rather than reflecting elements only on the reactive sediment surface 283 where DOM availability would be controlled by surface complexation with iron(III). Organo-Fe 284 linkages are also considered fairly stable, and the dissolution of such complexes generally 285 requires high pH and extreme redox conditions (Grybos et al., 2009). Thus, sediment DOM 286 extracted in MQ may be best described as that which is readily available and easily exchanged at 287 the sediment water interface (Tao & Lin, 2000), which is unlikely to impact stable organo-metal 288 complexes such as those involving iron. On the other hand, Cu-binding organic matter is more 289 easily mobilized from sediments (Skrabal et al., 1997) and is consistent with the observed 290 linkages between Cu and the sediment extracted DOM composition in this study.

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Section 3.2: Sediment DOM is significantly but weakly related to land use and soil elementalcomposition

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295 While the CONUS scale relationships between Cu and DOM composition were significant, we recognize such relationships are statistically weak (e.g.,  $r^2 < 0.3$  in most cases) indicating 296 297 significant variability across the dataset that remains unexplained. Like the sediment elemental 298 composition, watershed soil elemental composition provided weak relationships with the 299 individual DOM biochemical classes (Figure 2). Within multivariate space, however, soil Cu 300 again was the only significant addition in terms of explanatory power in the spatial variation in 301 sediment DOM composition and displayed similar linkages to aromatic type molecular formula. 302 However, this added explanatory power was <4% (Table 1) and may not be environmentally 303 relevant on its own.

304

While soil Cu added a small amount of explanatory power, its relationship with DOM chemistry was significant in both univariate (Figure 2) and multivariate (Figure 3) analyses. This is consistent with the sediment elemental data, and thus indicates Cu as potentially important for 308 understanding hyporheic zone DOM chemistry across the CONUS. Elements like Cu may play a 309 pivotal role in bioavailability and the residence time of aromatic DOM through a watershed from 310 the upland soils to the hyporheic zone sediments (Liu et al., 2019; Salomão et al., 2021). Such 311 relationships are subject to a myriad of complex interactions, including the impact of human 312 activities (urbanization, mining, agriculture) which can lead to increases in Cu concentrations in 313 streams compared to the surrounding soils (Lin et al., 2022; Wang et al., 2021).

314

315 In addition to the soil elemental composition, we included land cover as a source indicator for 316 organic matter traversing the aquatic network that may undergo deposition and exchange at the 317 sediment-water interface. Of the five major land cover classes, shrublands and grasslands 318 provided the only significant relationships with any of the DOM biochemical classes (Figure 2). 319 For instance, each were both inversely related to the relative richness of highly aromatic-like 320 formulae and positively related with higher relative richness of carbohydrate- and protein-like 321 formulae. This result is generally consistent with biomarker research that highlights an 322 enrichment of hemicellulose and protein rich materials in grasslands and a depletion of lignin in 323 comparison to woody forest soils (Bianchi & Canuel, 2011; Dümig et al., 2013). This 324 relationship is further emphasized in multivariate space giving a bidirectional pull along both 325 RDA1 and RDA2, indicating its linkage to more aliphatic type biochemical classes. It is notable 326 however, that land cover provided around 10% of the explanatory power (Table 1) in how 327 sediment DOM is longitudinally structured at the CONUS scale. This observation was surprising 328 given the strong historical support of linkages between landcover and the composition of OM 329 traversing through aquatic realms in both dissolved (S. Chen et al., 2021; Roebuck et al., 2020; 330 Williams et al., 2010; Wilson & Xenopoulos, 2009) and particulate phases (Jung et al., 2015; 331 Longworth et al., 2007; Lu et al., 2014).

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333 4. Conclusions (or Implications)

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Our results show a continental scale gradient in the composition of water extractable sediment DOM that displayed significant but weak associations with local sediment elemental composition, as well as soil elemental and land cover composition of the upstream drainage area. Relative to soil elemental or land cover composition, sediment elemental composition had stronger relationships with relative diversity gradients within compound classes (Fig. 2) and overall DOM chemistry (Fig. 3). This indicates that highly localized processes, within sediment pores, have a stronger influence over hyporheic zone DOM than watershed-scale processes associated with hillslope soils and DOM sources. A conceptually similar inference was found via analyses of DOM biochemical transformations, which were fully decoupled between physically adjacent surface water and sediments (Stegen et al., 2022). Further insights on additional controls over hyporheic zone DOM chemistry remain elusive and further evaluation is needed.

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347 Our results further indicate that highly localized controls on DOM chemistry may be tied to 348 biogeochemical history of sediments in a manner that limit capacity to explain CONUS-scale 349 variation. For example, localized variation in hydrological flow (e.g., sediment transport), 350 sediment residence time (e.g. potential for microbial turnover), and depth of the overlying water 351 column (e.g. potential for photochemical impacts) may all contribute to the underlying 352 biogeochemical controls on DOM chemistry associated with hyporheic zone sediments. We infer 353 that our current understanding of DOM-metal linkages does not necessarily transfer across 354 CONUS streams/rivers, limiting the use of localized studies to inform CONUS-scale models. 355 The minimal explanatory power of both land cover and soil elemental composition further 356 highlights the difficulty in extrapolating site- and regional-scale biogeochemical mechanisms to 357 multi-basin and continental scales. Future studies should continue with efforts to identify broad 358 scale linkages between sediment DOM chemistry and localized biogeochemical process that 359 better enable predictive modeling of river corridors across diverse settings and large spatial 360 scales.

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- 363 5. Data Availability
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All high-resolution mass spectrometry and XRF data used for the curation of this manuscript can be obtained free from the Environmental System Science Data Infrastructure for a Virtual Ecosystem (ESS-DIVE) Data Repository (Goldman et al., 2020). Watershed land cover data is available at ESS-DIVE via V A Garayburu-Caruso et al. (2022). An additional data package containing all scripts relevant to this manuscript in addition to the interpolated National 370 Geochemical Survey data will be available at the time of revision. For the purposes of review,

- 371 we have provided a singular data file containing all relevant data and code required to reproduce
- 372 figures and statistical results.
- 373
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580 8. Tables

581

- Table 1: Redundancy analysis model output showing the proportion of variance in organic matter
- 583 chemistries explained by the different sediment and soil elements and land cover.
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Parameter	Degrees of Freedom	Variance	F	P (>F)
Sediment Copper (wt percent)	1	1.0647	11.3941	0.001
Soil Copper (ppm)	1	0.2909	3.1014	0.031
Percent Shrub	1	0.2360	2.5153	0.075
Percent Grass	1	0.1829	1.1498	0.115
Percent Forest	1	0.4093	4.3636	0.007
Residual	62	5.8162		

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589 Figure 1: Percent of FTICR-MS biomolecular classes as a function of longitude

- 592 Figure 2: Correlation Matrix to show  $r^2$  values for significant correlations (p < 0.05) between
- 593 FTICRMS biochemical classes with sediment and soil elemental composition as well as land
- 594 cover.
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	%Amino Sugar	%Carbohydrates	%Condensed Aromatics	%Lignin	%Lipid	%Protein	%Tannin	Unsaturated Hydrocarbon	Longtitude
Longitude	-0.11		0.22	0.13		-0.29	0.17	-0.06	1.00
%Si (sediment)		-0.10	0.11			-0.06	0.13		
%P (sediment)			0.11				0.07		
%Ca (sediment)		0.14					-0.06		
%Mn (sediment)									
%Fe (sediment)									
%Cu (sediment)	-0.23	-0.29	0.06	0.18		-0.16	0.08		
%Crop									
%Grass		0.12					-0.07		
%Shrub			-0.08			0.11	-0.08		-0.22
%Urban									0.11
%Forest									-0.07
Ca_wt_percent (soil)									
Cu_ppm (soil)	-0.07		0.08			-0.06			0.09
Fe_wt_percent (soil)	0.10		-0.08			0.10			-0.16
Mn_wt_percent (ppm)									
P_wt_percent (soil)	0.06		-0.06	-0.07		0.11			-0.34

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597 Figure 3: Redundancy Analysis generated with results from stepwise model building for 598 constrained ordination. Panel (a) is the RDA while panel (b) displays RDA 1 as a function of 599 longitude.



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