A multi-chemistry modelling framework to enable flexible and reproducible water quality simulations in existing hydro-models: 2. The OpenWQ-SUMMA and OpenWQ-CRHM model implementations and testing.

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Abstract

This work advances the cross-model deployment of ecological and biogeochemical simulation capabilities in existing processbased hydro-modeling tools, which we term "Open Water Quality" (OpenWQ). The companion paper details aspects of the OpenWQ architecture that enables its plug-in type incorporation into existing models, along with its innovative aspects that enable biogeochemistry lab-like capabilities. OpenWQ's innovative aspects allow modelers to define the pollution problem(s) of interest, the appropriate complexity of the biogeochemistry routines, test different modeling hypotheses, and deploy them across different hydro-models. In this second paper, we implemented the coupling recipe described in the first paper to integrate OpenWQ into two hydro-models, SUMMA and CRHM. Here we explain how the implemented coupling interface between the two models provides water quality simulation capacities in the host hydro-models but, more importantly, establishes a direct and permanent link for the transfer of innovation between the modeling communities. Example applications of different pollution studies enabled by our coupling recipe are also provided to address some of these fundamental water quality modeling challenges.

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11 Abstract

This work advances the cross-model deployment of ecological and biogeochemical sim-12 ulation capabilities in existing process-based hydro-modeling tools, which we term "Open 13 Water Quality" (OpenWQ). The companion paper details aspects of the OpenWQ architec-14 ture that enables its plug-in type incorporation into existing models, along with its innova-15 tive aspects that enable biogeochemistry lab-like capabilities. OpenWQ's innovative aspects 16 allow modelers to define the pollution problem(s) of interest, the appropriate complexity of 17 the biogeochemistry routines, test different modeling hypotheses, and deploy them across 18 19 different hydro-models. In this second paper, we implemented the coupling recipe described in the first paper to integrate OpenWQ into two hydro-models, SUMMA and CRHM. Here 20 we explain how the implemented coupling interface between the two models provides water 21 quality simulation capacities in the host hydro-models but, more importantly, establishes a 22 direct and permanent link for the transfer of innovation between the modeling communities. 23 Example applications of different pollution studies enabled by our coupling recipe are also 24 provided to address some of these fundamental water quality modeling challenges. 25

²⁶ 1 Introduction

Climate change is expected to affect many hydro-biogeochemical processes along the 27 land-river-lake continuum (Costa et al., 2022). However, the specific local impacts of climate 28 change on hydrology and water quality will vary regionally and seasonally. On the one hand, 29 extreme winter river flows arising from increased extreme precipitation events are expected 30 to occur with greater frequency worldwide (Alexander et al., 2020; Field & Barros, 2014; 31 Luo et al., 2020; Onishi et al., 2020; "Effects of lake water level fluctuation due to drought 32 and extreme winter precipitation on mixing and water quality of an alpine lake, Case Study: 33 Lake Arrowhead, California", 2020; Shi et al., 2011; Suddick et al., 2013; H. Wilson et al., 34 2019; Yang et al., 2019; Zheng et al., 2020). On the other hand, summer river flows are 35 projected to decrease across many regions around the world, including North America (Rood 36 et al., 2008), China (Shi et al., 2011; Ervinia et al., 2020; Zheng et al., 2020), and western 37 Europe (Charlton et al., 2018; Raimonet et al., 2018). In arid climates, the flow reduction 38 is predicted to be particularly strong (Jarsjö et al., 2017; Shi et al., 2011; Y. Wu et al., 39 2012; Ye & Grimm, 2013; Bussi et al., 2021), reducing dilution of pollution point sources, 40 including wastewater outlets (Bussi et al., 2017; Charlton et al., 2018; Raimonet et al., 2018; 41 Ockenden et al., 2016; Whitehead et al., 2009). 42

The increase in climate variability is projected to lead to an overall increase in precipi-43 tation and snowmelt runoff, which, combined with changes in intensified cropping patterns, 44 is anticipated to lead to greater freshwater nutrient pollution in temperate regions of North 45 America, Europe and Japan (Alexander et al., 2020; Jeppesen et al., 2011; Onishi et al., 46 2020; H. F. Wilson et al., 2019). Increased nutrient uptake rates from terrestrial plants 47 due to climate change are not expected to facilitate the increase in nutrient runoff losses 48 (Luo et al., 2020; Sperotto et al., 2019; Suddick et al., 2013; Zheng et al., 2020). Erosivity 49 from agricultural and urban nonpoint sources is also projected to increase with more intense 50 runoff events because of their greater capacity to mobilize particulate and soluble N and P 51 (El-Khoury et al., 2015; Ockenden et al., 2016, 2017; J. Wu & Malmström, 2015; Crossman 52 et al., 2014). 53

Understanding the key factors controlling the fate of pollution in aquatic ecosystems is 54 challenging because it requires capturing multiple climate, hydrological, and biogeochemical 55 patterns across different spatial scales in models (Costa et al., 2022). Whilst there have 56 been numerous modelling advances over the last decade, process representation, model inter-57 comparability, model configuration (input data), and model validation remain key challenges 58 in water quality. Another considerable difficulty with the application of water quality models 59 is the often-limited monitoring data available for validation, which can compromise the 60 confidence of scientists and decision-makers in their simulation results. 61

This two-part paper presents a methodology termed "OpenWQ" that tackles some of 62 these fundamental modeling challenges by streamlining the transfer of knowledge, technol-63 ogy, and innovation between the different scientific and modeling communities related to 64 water quality. Some of these scientific and modelling communities include biogeochemistry, 65 soil science, hydrology, hydrodynamics, and hydrogeology. This work aims to provide a con-66 crete, transformative framework for enhancing collaborative efforts and co-creation that can 67 benefit the entire environmental and ecological research and management community. Our 68 intent for the first paper was to advance OpenWQ's concept as a coupler and introduce its 69 customizable biogeochemistry modeling framework. This customizable framework enables 70 the cross-model deployment of eco-lab water quality modeling capabilities. This second 71 paper builds on the motivation for OpenWQ introduced in the first paper and outlines the 72 initial implementation of OpenWQ as a modeling system. The goals of this second paper 73 are to 74

Demonstrate how OpenWQ has been integrated into two hydro-models using the coupling procedures laid out in the first paper. The hydro-models have been selected for their likewise flexible modeling frameworks, namely the Structure for Unifying Multiple Modeling Alternatives (SUMMA, Clark et al., 2015a,b) and the Cold Regions Hydrological Model and its water quality modules (CRHM/CRHM-WQ, J. W. Pomeroy et al., 2007; Costa et al., 2021).

and SUMMA to examine a variety of common pollution problems.

2. Apply the multi-chemistry, eco-lab framework of OpenWQ incorporated with CRHM

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⁸³ 2 Coupling to hydro-models

OpenWQ has been coupled to two hydro-models, SUMMA (Clark et al., 2015a,b) and CRHM/CRHM-WQ (J. Pomeroy et al., 2022; Costa et al., 2021; J. W. Pomeroy et al., 2007). In companion paper 1 (*CITE_paper_1*), we explain how OpenWQ has been designed as a flexible coupler to streamline its integration in existing hydro-models. A coupling recipe has been developed to lay out the steps involved in the coupling process, which includes linking four coupling functions. Here, in paper 2, we explain how this coupling recipe was deployed in two separate coupling efforts and used to examine different pollution problems.

SUMMA SUMMA is written in Fortran, while OpenWQ is written in C++. Thus, it 91 was necessary to create wrapper functions that allow for the interaction between a Fortran 92 driver program and a C++ coupled program at runtime. Because it is OpenWQ that is 93 integrated into SUMMA, the coupled model was constructed in a manner to allow SUMMA 94 to manage the simulation timesteps. For this, we used the standard intrinsic Fortran module 95 *iso_c_binding* that defines constants, types, and procedures for C interoperability. The 96 *iso_c_binding* module is part of the language standard and makes the interface between 97 Fortran and C standard and portable. Essentially, Fortran-to-C wrapper functions were 98 created to allow SUMMA to interface with OpenWQ and, from the point of view of SUMMA, aq OpenWQ is an object to which it held a pointer upon initialization of OpenWQ. The wrapper 100 functions allowed SUMMA to interface with OpenWQ's coupler functions as if they were 101 method calls. Fortran-to-C wrapper functions were constructed for creating a reference to 102 a OpenWQ along with each of the four OpenWQ C++ coupler functions (openwq::decl, 103 openwq::run_time_start, openwq::run_space, and openwq::run_time_end; see Section 3.3.2 of 104 the first paper (CITE_paper_1). 105

CRHM Similar to OpenWQ, CRHM is written in C++. This avoids the need for
 wrapper interface functions as required in the SUMMA-OpenWQ coupling. CRHM already
 had a water quality module named CRHM-WQ that enables the simulation of the nitrogen
 (N) and phosphorous (P) cycles (Costa et al., 2021). Thus, CHRM already contains a so lute transport simulation routine that moves contaminants through the hydrological cycle
 represented by the model. The native implementation of solute transport in CRHM-WQ is

limited to 6 chemical species related to the N and P cycles, namely nitrate (NO_3-N) , ammonium (NH_4-N) , dissolved organic nitrogen (DON), soluble reactive phosphorous (SRP), and particulate phosphorous (PP). Therefore, instead of coupling all four OpenWQ coupler functions to CRHM-WQ, we considered it more appropriate to use CRHM's native transport routine, which has already been tested, and adjust it for a more dynamic allocation of the number of chemical species when OpenWQ is activated, leaving OpenWQ's coupling routine for transport (*openwq::run_space*) out of the coupling process.

¹¹⁹ 3 Case studies

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3.1 Overal science questions

¹²¹ OpenWQ enables exploring multiple working hypotheses related to multi-chemistry ¹²² biogeochemical cycling in different seasons, climate zones, and landscapes. The work aims ¹²³ to respond to critical challenges with water quality modeling as an interdisciplinary science, ¹²⁴ which have been highlighted in the companion paper (*CITE_paper_1*). The specific modeling ¹²⁵ challenges addressed are to:

1. Identify appropriate biogeochemical frameworks: what are the key biogeochemical processes and chemical species that should be represented in models to capture the fate and controls of pollution in a way that is (1) adequate for the quantity and quality of the available data for model evaluation, (2) scientifically sound from a biogeochemical perspective, and (3) useful for the objectives of the study and helpful for decision making?

- 2. Differentiate controlling biogeochemical transformations across hydrological compart-132 ments (e.g., snow, groundwater, etc.): should biogeochemical processes be differen-133 tiated across hydrological compartments and regions, e.g., although the processes 134 involved in the nitrogen cycle are the same regardless of the environment they occur 135 in (e.g., nitrification, denitrification, assimilation, ammonification, nitrogen fixation), 136 certain environmental conditions tend to favor particular processes (e.g., oxygen lev-137 els in groundwater tend to be low, leading to reduced nitrification rates because this 138 is a process involving dissolved oxygen)? 139
- Characterize model uncertainty: can we better understand the sources of uncertainty related to key modeling processes and methods used so that we can incrementally improve our models, increase predictability and model fidelity, and reduce parameter equifinality?

Some of these questions have also been raised by the hydrological community (Clark et al., 2015b; Reggiani et al., 1998; Beven, 2006; Renard et al., 2010; Wood et al., 2011), but the ones listed above extend to aspects specific to water quality modeling. The general concept of OpenWQ is to facilitate a test-improve workflow so that models can be flexibly and incrementally improved based on model results and inputs from different disciplines.

¹⁴⁹ **3.2 Model tests**

The model was evaluated using a series of tests designed to expose the accuracy and 150 robustness of the mathematical formulation and numerical solutions deployed for examining 151 a variety of pollution problems. In recent years, it has been recognized that such basic 152 tests are an essential step for proper model scrutiny, falsification, and acceptance (Clark et 153 al., 2021a). In this study, we pioneer the extension of such important principles for water 154 quality simulations where the spatiotemporal complexity and interplay of process generally 155 often give rise to highly non-linear response patterns that are difficult to disentangle in 156 real-world applications for proper model examination (Costa et al., 2022, 2020). The case 157 study tests devised are representative of a variety of environmental pollution problems for 158

Transport Scheme	Designation	Description			
No transport	1_noTrans_1species_1storder 2_noTrans_1species_2ndorder 3_noTrans_2species 4_noTrans_3species 5_noTrans_nitrogencycle 6_noTrans_oxygenBODcycle	Batch reactor involving 1 chemical species subject to first-order decay Batch reactor involving 1 chemical species subject to second-order decay Batch reactor involving 2 chemical species subject to first-order decay Batch reactor involving 3 chemical species subject to first-order decay Batch reactor representing the nitrogen cycle Batch reactor representing the dissolved oxygen cycle			
Advection- Dispersion	7_trans_contS_PorMedia 8_trans_contS_PorMedia_1storder	Non-reactive transport of 1 chemical species in porous media subject to a continuous load Reactive (linear decay) transport of 1 chemical species in porous media subject to a continuous load			
	9_trans_instS_PorMedia 10_trans_instS_PorMedia_1storder	Non-reactive transport of 1 chemical species in porous media subject to an instantaneous load Reactive (linear decay) transport of 1 chemical species in porous media subject to an instantaneous load			

Table 1: Case study tests used to test OpenWQ and its coupling to SUMMA and CRHM

which analytical solutions have been either specially derived for this paper or taken from the literature.

We divide the tests into reactive and non-reactive. This is an important distinction 161 because basic mass balance principles cannot be directly applied to reactive chemical con-162 stituents without proper consideration because they are subject to biogeochemical transfor-163 mation in addition to physical transport. We prepared a total of ten tests, of which two are 164 non-reactive and eight are reactive. Each of these tests has well-defined initial conditions, 165 boundary conditions, sink and source terms, and is subject to specific biogeochemistry as 166 described below. Based on these conditions, analytical solutions have been derived for com-167 parison with the model results. In order to benefit the comparison of model results, the 168 simulations are divided into two groups as shown in Table 1. 169

The case studies include six batch reactor-type tests, specifically two single-species 170 tests and four multi-species tests. The single-species tests are the simplest and represent 171 cases involving first- (Test 1) and second-order (Test 2) linear decay processes. The multi-172 species tests include (a) a two-species reaction network subject to linear decay (Test 3), (b) 173 a three-species reaction network subject to linear decay (Test 4), and reaction networks to 174 represent the (c) nitrogen (Test 5) and (d) dissolved oxygen (Test 6) cycles. The remaining 175 four tests involve the transport of chemical constituents, specifically two non-reactive tests, 176 one considering a continuous point source (Test 7), one considering an instantaneous points 177 source (Test 8), and two reactive tests that are similar to the homologous non-reactive 178 tests, but where chemical species undergo biogeochemical transformations (Tests 9 and 10, 179 respectively). 180

3.3 Multi-chemistry lab set up

The case studies involve different pollution problems. In some cases, they are generic 182 in the sense that they are representative of types of biogeochemical processes and could be 183 applied to examine a variety of water quality problems. In other cases, the tests are more 184 specific, focusing on environmental issues involving specific biogeochemical cycles, such as 185 the nitrogen and phosphorous cycles. As mentioned in the first paper, the biogeochemistry 186 calculation engine is flexible and easily extendible, supporting numerous forms of functional, 187 logical, and vector processing semantics. It is enabled by the comprehensive C++ Mathe-188 matical Expression Toolkit Library (*ExprTk*) developed by Arash Partow (Partow, 1999). 189 The implementation of ExprTk in OpenWQ is simple to use. It provides an efficient run-190

time mathematical expression parser and evaluation engine that is ideal for testing different modeling hypotheses to inform incremental improvements to the model setup.

In order to perform the various pollution case study tests described in Table 1, a series 193 of OpenWQ input files have been prepared. The biogeochemistry JSON files (see Section 194 ?? of the first paper) are where the multi-chemistry lab simulations are characterized. We 195 defined one biogeochemistry file per case study test, which we deployed in both SUMMA-196 OpenWQ and CRHM-OpenWQ coupled models. From paper 1, one of the transformative 197 innovations of OpenWQ is the ability for cross-model deployment of eco-lab co-creation 198 modeling capabilities using the same OpenWQ biogeochemistry JSON files. This allows the 199 implementation of multiple hydro-biogeochemical modeling hypotheses by simply loading 200 the desired JSON biogeochemistry file. The input files prepared for the different case study 201 tests are publicly available on GitHub. They can be used as the basis for validation of any 202 new coupling of OpenWQ to a hydro-model. 203

3.4 Deriving analytical solutions

In this section we provide details about the derivation of the analytical solutions used to verifying the performence of OpenWQ when coupled to SUMMA and CRHM. Each test is presented in a separate section, which starts with a diagram of relevent reaction network. The diagrams are then used to define the governing equations, which can be ordinary or partial differential equations. Finally, the derived analytical solution is presented. Details about the derivations is provided in the Appendix A-G.

3.4.1 Batch reactor: Single-species (Tests 1 and 2)

The single-species batch reactor tests involve (a) first-order and (b) second-order decay kinetics. Fig. 1 depicts the biogeochemical process.



Figure 1: Single-species reaction network where transf_A is a first-order biogeochemical decay reaction.

The analytical solution of a single species subject to first-order decay can be obtained by integrating Eq. 24 with initial condition at time t = 0 to obtain Eq. 25. In turn, the analytical solution of a single species subject to second-order decay can be obtained by integrating Eq. 26 with initial condition at time t = 0 to obtain Eq. 27. See Appendix A for more details.

$$\frac{dc_A}{dt} = -kc_A \tag{1}$$

$$c_A(t) = c_A(0)e^{-kt} \tag{2}$$

$$\frac{dc_A}{dt} = -kc_A^2 \tag{3}$$

$$c_A(t) = \frac{1}{\frac{1}{c_A(0)} + kt}$$
(4)

3.4.2 Batch reactor: Two-species reaction network (Test 3)

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The two-species reaction network with linear decay is represented schematically in Fig. 2.



Figure 2: Two-species reaction network where transf_A and transf_B are first-order biogeochemical transformations.

Similar to the single-species case with first-order decay (Eq. 25), the fate of Species A
in a two-species reaction network only depends on its own concentration and is given by Eq.
28.

$$c_A(t) = c_A(0)e^{-k_A t} \tag{5}$$

However, the fate of Species B depends on the time-evolution of concentrations of bothSpecies A and B. This evolution can be expressed as:

$$\frac{dc_B}{dt} = k_A c_A - k_B c_B \tag{6}$$

The analytical solution of Eq. 29 with initial condition at time t = 0 can be obtained by using the integrating factor, $e^{k_B t}$. Details on this derivation are provided in the Appendix B. The solution is provided in Eq. 38.

$$c_B(t) = c_B(0)e^{-k_B t} + \frac{k_A c_A(0)}{k_B - k_A} \left\{ e^{-k_A t} - e^{-k_B t} \right\}$$
(7)

230 3.4.3 Batch reactor: Three-species reaction network (Test 4)

A three-species reaction network with linear decay is represented schematically in Fig. 2.

Concentrations of Species A and B are given by respectively Eqs. 28 and 38. The
 differential equation for the concentration of Species C can be written as:

$$\frac{dc_C}{dt} = k_B c_B - k_C c_C \tag{8}$$



Figure 3: Three-species reaction network where transf_A, transf_B, and transf_C are first-order biogeochemical transformations.

234 235 The analytical solution Eq 40 can again be obtained by using the integrating factor $e^{k_C t}$. Details on this derivation are provided in the Appendix. The final solution is

$$c_{C}(t) = c_{C}(0)e^{-k_{C}t} - \frac{k_{B}}{k_{B} - k_{C}}c_{B}(0)\left(e^{-k_{B}t} - e^{-k_{C}t}\right) - \frac{k_{B}k_{A}}{(k_{B} - k_{A})(k_{C} - k_{A})}c_{A}(0)\left\{\left(e^{-k_{B}t} - e^{-k_{A}t}\right) - \frac{e^{-k_{C}t}}{k_{C} - k_{B}}\right\}.$$
(9)

²³⁶ 3.4.4 Batch reactor: Nitrogen cycle (Test 5)

The simulation of the nitrogen cycle is common in freshwater water quality studies because it allows for addressing issues with nutrient pollution and eutrophication. The nitrogen cycle simulated involves different nitrogen species (Fig. 4). The conceptual cycling framework used is similar to that used in catchment-scale models like HYPE Lindström et al. (2010), CRHM-WQ (Costa et al., 2021) models, SWAT (Arnold et al., 2012), and INCA ("A semi-distributed ntegrated itrogen model for multiple source assessment in tchments (INCA): Part I — model structure and process equations", 1998).



Figure 4: Nitrogen cycle as represented in several popular catchment-scale models. The nitrogen species Nref refers to refractory nitrogen, Nlab refers to labile nitrogen, DON refers to dissolved organic nitrogen, and DIN refers to dissolved innorganic nitrogen. The transformation degr refers to degration; diss1 and diss2 refers to dissolution from Nref and Nlab, respectively; min refers to mineralization; denitr refers to denitrification; and plantup refers to plant uptake.

The analytical solutions for the different nitrogen species can be obtained from Eq. 28, Eq. 38, or Eq. 47, depending on their positioning in the reaction chain:

Refractory N (Nref) Similar to Eq. 28, which for this specific processes and N species reads as

$$c_{Nref}(t) = c_{Nref}(0)e^{-(k_{degr} + k_{diss1})t}$$
 (10)

Labile N (Nref) Similar to Eq. 38, which for this specific processes and N species reads as

$$c_{Nlab}(t) = c_{Nlab}(0)e^{-(k_{min} + k_{diss2})t} + \frac{k_{degr}c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \left\{ e^{-k_{degr}t} - e^{-(k_{min} + k_{diss2})t} \right\}.$$
(11)

$$\frac{dc_{DON}}{dt} = k_{diss1}c_{Nref} + k_{diss2}c_{Nlab} \tag{12}$$

Due to similarities in the relative positioning of the chemical species in the reaction network, analytical solutions suitable for c_{Nref} and c_{Nlab} are given by Eqs. 28 and 38, respectively. The derivation of Eq. 50 is provided in Appendix D, and its solution reads

$$c_{DON}(t) = c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}} - \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}}e^{-(k_{degr} + k_{diss1})t} - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}}e^{-(k_{min} + k_{diss2})t} + \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2}} \left\{ \frac{e^{(k_{min} + k_{diss2})t} - 1}{k_{min} + k_{diss2}} - \frac{e^{-k_{degr}t}}{k_{degr} - 1} \right\}.$$
(13)

Dissolved Inorganic N (DIN) The ordinary differential equation for DIN can be
 written as

$$\frac{dc_{DIN}}{dt} = k_{min}c_{Nlab} - (k_{denit} + k_{plantup})c_{DIN}.$$
(14)

The analytical form for the concentration of $c_{Nlab}(t)$ is given in Eq. 49 and can be substituted in Eq. 56 to initiate the derivation that is provided in the Appendix. The analytical solution to Eq. 56 reads

$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]} = c_{DIN}(0)$$

$$+k_{min}c_{Nlab}(0)\left\{\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}-1}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right\}$$

$$+\frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\left\{\frac{e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]}-1}{k_{denitr}+k_{plantup}-k_{degr}}\cdot-\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}-1}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right\}$$

$$(15)$$

3.4.5 Batch reactor: Dissolved oxygen cycle (Test 6)

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Dissolved oxygen (DO) depletion in rivers and lakes is a common consequence of pollution. It typically occurs near discharge outlets of untreated wastewater and can lead to high levels of organic matter that, during decomposition by aerobic bacteria, causes a DOsag in the water column. For the purposes of DO dynamics, such organic matter is typically translated into units of biological oxygen demand (BOD). Fig. 5 depicts the processes involved, and the Streeter–Phelps equation (Streeter & Phelps, 1925) has long been used for such problems. It reads

$$\frac{dD}{dt} = k_1 L_t - k_2 D,\tag{16}$$

where *D* is the saturation deficit in g/m^3 , which can be derived from the dissolved oxygen concentration at saturation minus the actual dissolved oxygen concentration ($D = DO_{sat} - DO$); k_1 is the deoxygenation rate, usually in d^{-1} ; $L_t = L_a e^{-k_1 t}$; L_a is the initial oxygen demand of organic matter in the water, also called the ultimate *BOD* (*BOD* at time $t = \infty$); and k_2 is the reaeration rate, usually in d^{-1} . The unit of L_a is g/m^3 ; k_1 ; L_t is the oxygen demand remaining at time t, and t is the elapsed time, usually in days.



Figure 5: Representation of the transformations involved in the BOD - DO cycle

The equation states that the total rate of change in oxygen deficit (D) is equal to the difference between the two rates of deoxygenation and reaeration at any time. See more details in Appendix E. The solution reads

$$D = \frac{k_1 L_a}{k_2 - k_1} \left\{ e^{-k_1 t} - e^{-k_2 t} \right\} + D_a e^{-k_2 t}.$$
(17)

where D_a is the initial oxygen deficit $[g/m^3]$.

268 3.4.6 Reactive and non-reactive transport: Instantaneous point source 269 (Tests 7 and 8)

These tests were only performed with SUMMA-OpenWQ because, as mentioned before, while OpenWQ's biogeochemical simulation modules have been integrated in both SUMMA-OpenWQ and CRHM-OpenWQ, OpenWQ's solute transport simulation modules were only

integrated in SUMMA-OpenWQ, but not in CRHM-OpenWQ, since CRHM has it's own solute transport routines (Costa et al., 2021). The general 1D diffusion-advection-reaction partial differential equation is given by

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c, \qquad (18)$$

where c is the concentration of the tracer, t is time, D_x is the diffusion coefficient in the flow direction, x is the distance in the flow direction, ν is the flow velocity in the x-direction, and λ is a linear reaction term that can be set to zero for non-reactive tracers. Various analytical solutions for Eq. 18 have been derived and compiled by Wexler (1992) for different idealized scenarios. The modified equation for an instantaneous point source reads

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c + \frac{Q}{n} dt c_0 \delta \left\{ x - x_c \right\} \delta \left\{ t - t'_c \right\},\tag{19}$$

where δ is the Dirac delta (impulse) function, x_c is the location of the instantaneous point source, t'_c is the time at which the instantaneous point source is activated, Q is the fluid injection rate, and n is the aquifer porosity. The analytical solution to Eq. 67 has been derived by Bear (1979) considering the following boundary conditions and initial conditions.

- 2791. Boundary conditions (Neumann or second type):280 $c, \nabla c = 0, \quad x = \pm \infty$ 2812. Initial conditions:282 $c(x,t=0) = 0, \quad -\infty < x < \infty.$
- The analytical solution of Eq. 67 is given by

$$c(x) = \frac{c_0 Q}{4n\pi\gamma D_x} \exp\left[\frac{\nu(x-x_0)}{2D_x}\right]$$

$$\int_0^t \frac{1}{t} \exp\left[-\left(\frac{\nu^2}{4D_x}\right)\tau - \frac{(x-x_c)^2}{4D_x\tau}\right].$$
(20)

and can be used for solutes that are non-reactive (Test 7) or reactive (Test 8) by setting $\lambda = 0$ in Test 7.

The equation assumes that (1) the fluid has constant density and viscosity, (2) the 286 solute may be subject to first-order chemical transformation, (3) the flow is uniform in the 287 x-direction only, (4) the flow velocity is constant, and (5) the longitudinal diffusion coefficient 288 (D_x) is constant. The solution has a mathematical singularity in the point source region 289 (Abdelaziz et al., 2013) because it takes the form of the exponential integral when both 290 $x - x_c$ approaches zero, hence becoming infinite at $\tau = 0$. However, according to Wexler 291 (1989) the solution is still valid as long as $(x - x_c)^2$ is larger than ν^2 , as it is the case in our 292 293 tests.

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3.4.7 Reactive and non-reactive transport: Continuous point source (Tests 9 and 10)

Similarly to Tests 7 and 8, the tests here have only been performed for SUMMA-OpenWQ due to the same reasons highlighted above in Section 3.4.6. The analytical solution for the transport of solutes with a continuous point source has been derived in Wexler (1992) after modification from Bear (1979) and Genuchten & Alves (1982).

			Biogeochemistry			
Transport Scheme	Designation	Reference	Initial Conditions	Reaction kinetics	Transport	
	1_noTrans_1species_1storder	Fig. 1	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.01 \ ^{1}/day$	NA	
Patah Dasatan	2_noTrans_1species_2ndorder	Fig. 1	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.01^{1/day}$	NA	
	3_noTrans_2species	Fig. 2	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.03 \ ^{1}/day$	NA	
Datch Reactor			Species B ($c_0 = 0 \text{ mg/l}$)	$k_{transf_B} = 0.01 \ ^{1}/day$		
	4_noTrans_3species	Fig. 3	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.03 \ 1/day$	NA	
			Species B $(c_0 = 0 \text{ mg/l})$	$k_{transf_B} = 0.01 \ \frac{1}{day}$		
			Species C ($c_0 = 0 \text{ mg/l}$)	$k_{transf_{-C}} = 0.005 \ \frac{1}{day}$		
	5_noTrans_nitrogencycle	Fig. 4	Nref ($c_0 = 10 \text{ mg/l}$)	$k_{degrd} = 0.006 \ \frac{1}{day}$	NA	
			Nlab $(c_0 = 10 \text{ mg/l})$	$k_{dissol-1} = 0.0002 \ \frac{1}{day}$		
			$DIN (c_0 = 5 \text{ mg/l})$	$k_{dissol-2} = 0.0003 \frac{1}{day}$		
			DON $(c_0 = 2 \text{ mg/l})$	$k_{miner} = 0.003 \frac{1}{day}$		
				$k_{denitr} = 0.001 \frac{1}{day}$		
	a T DOD I	D: -		$k_{plantup} = 0.001 \frac{1}{day}$	27.4	
	6_noTrans_oxygenBODcycle	Fig. 5	BOD $(c_0 = 10 \text{ mg/l})$	$k_{BODdegrad} = 0.1 \frac{1}{day}$	NA	
		<u> </u>	$DO(c_0 = 12 \text{ mg/l})$	$\kappa_{reaertion} = 0.3 \frac{1}{day}$		
Advection-Dispersio	7_trans_contS_PorMedia	NA	Species A $(c_0 = 0 \text{ mg/l})$	NA	$D_z = 0.0001 \ m^2/s$	
	8_trans_contS_PorMedia_1storder	Fig. 1	Species A $(c_0 = 0 \text{ mg/l})$	$k_{transf_A} = 0.01 \ ^{1/day}$	$D_z = 0.0001 \ m^2/s$	
	9_trans_instS_PorMedia	NA	Species A $(c_0 = 0 \text{ mg/l})$	NA	$D_z = 0.0001 \ m^2/s$	
	10_trans_instS_PorMedia_1storder	Fig. 1	Species A $(c_0 = 0 \text{ mg/l})$	$k_{transf_A} = 0.01 \ ^{1}/day$	$D_z = 0.0001 \ m^2/s$	

Table 2: Model configuration for the different test cases

$$c(x) = \frac{c_0}{2} \left\{ \exp\left[\frac{x}{2D_x}(\nu - U)\right] \operatorname{erfc}\left[\frac{x - Ut}{2\sqrt{D_x t}}\right] + \exp\left[\frac{x}{2D_x}(\nu + U)\right] \operatorname{erfc}\left[\frac{x + Ut}{2\sqrt{D_x t}}\right] \right\},$$
(21)

where c_0 is the concentration of the continuous point source, $U = \sqrt{\nu^2 + 4\lambda D}$ and erfc is the complementary error function. Parameter λ is the first-order reaction rate used for Test 10, which can be set to zero for Test 9. See more details in Appendix G.

303

3.5 Model configuration and input files

The configuration of OpenWQ for the different test cases is summarized in Table 2. 304 The analytical solutions described in Section 3.4 have been set up using the same initial and 305 boundary conditions. The configuration of OpenWQ involves the characterization of biogeo-306 chemistry (initial conditions and reaction kinetics) and transport processes. The OpenWQ 307 input files set up to translate the model configurations described in Table 2 are publicly 308 available at https://github.com/ue-hydro/synthetic_tests. These input files can be 309 used with any hydro-model coupled to OpenWQ and constitute an essential verification 310 step in any new coupling project. 311

This public repository also contains the input files of the host hydro-models CRHM 312 and SUMMA prepared for each model test, which have, in essence, been set up to provide 313 no-flow conditions for Tests 1 to 6 (batch-reactor simulations) and steady-state conditions 314 for Tests 7 to 10 (reactive and non-reactive transport simulations). Setting up proper flow 315 conditions was a critical step to ensure the validity of the analytical solutions used to verify 316 the numerical results. In the case of Tests 7 to 10 (transport simulations), which were only 317 performed for SUMMA-OpenWQ as explained before, were based on the soil unsaturated 318 flow test cases formulated by Celia et al. (1990), which have been set up in SUMMA as 319 described in Clark et al. (2021b). The soil properties of these tests are described in Table 2 320 of that paper. 321

322 4 Results

The model results have been compared to the respective analytical solutions to verify the accuracy of OpenWQ in examining a variety of pollution problems. These tests contemplate several biogeochemical processes that are often the object of study in environmental studies.

The Nash–Sutcliffe efficiency (NSE) coefficient and the Root-Mean-Square Error (RMSE)

³²⁷ were used to quantify the accuracy of the numerical results.

$$NSE = 1 - \frac{\sum (X_{obs} - X_{mod})^2}{\sum (X_{obs} - \mu_{obs})^2}$$
(22)

$$RMSE = \sqrt{\frac{\sum (X_{obs} - X_{mod})^2}{n}}$$
(23)

where X_{obs} and X_{mod} are the observed and simulated concentrations, μ_{obs} is the concentration average, and n is the number of observations. NSE values equal to one indicate a perfect match between observations and model results, and NSE values equal to zero indicate model predictions as accurate as the mean of all observations.

4.1 Batch reactor: Single-species (Tests 1 and 2)

The results of Tests 1 and 2 contemplate the general cases of single chemical species subject to first- or second-order decay as presented in Fig. 6. Such reaction types are commonly used in sequence or parallel to represent complex reaction networks.



(a) Test 1 (1_noTrans_1species_1storder)



Figure 6: Comparison between model results and analytical solution for Tests 1 and 2

The results show good model predictability and reduced numerical dispersion. For 336 Test 1, the Nash-Sutcliffe and RMSE values obtained for summa-openwog were respectively 337 0.999999998259312 and 8.804e-05 mg/l and for crhm-openwq were 0.99999998259309 and 338 8.804e-05 mg/l. For Test 2, the Nash-Sutcliffe and RMSE values obtained for summa-330 openwq were 0.999999900146240 and 3.0319e-04 mg/l, and for crhm-openwq were respec-340 tively 0.999999900145974 and 3.0319e-04 mg/l. The results obtained with summa-openwq 341 and crhm-openwq are similar, which was expected since these tests deal only with biogeo-342 chemical transformations and nutrient transport, which is driven by the host-models, is not 343 activated. The RMSE is one order of magnitude higher for Test 2, which was also antic-344 ipated given that the biogeochemical formulation in this test contains an exponent (i.e., 345 power of two), which magnifies the propagation of round-off errors. 346

4.2 Batch reactor: Two-species reaction network (Test 3)

The results of Test 3 are presented in Fig. 7 and compared against analytical solutions. Two-species reaction networks are also common in water quality models and often integrated within comprehensive biogeochemical cycling representations.



Figure 7: Comparison between model results and analytical solution for Test 3

The model shows good agreement with the analytical solution for Test 3, including 351 the results for both Species A and B. For Species A, the Nash-Sutcliffe and RMSE values 352 obtained for summa-openwq were 0.999999987172389 and 1.525e-04 mg/l, and for crhm-353 openwq were respectively 0.999999987172368 and 1.525e-04 mg/l. For Species B, the Nash-354 Sutcliffe and RMSE values obtained for summa-openwq were 0.999999989757722 and 1.747e-355 04 mg/l, and for crhm-openwq were respectively 0.999999989757706 and 1.747e-04 mg/l. 356 The RMSE calculated for Species B is slightly higher than that calculated for Species A. 357 This was anticipated since the reaction network sequence starts with Species A; thus, the 358 effect of changes in the concentration of Species A will take one time step to travel to Species 359 B. Reducing the length of the time step reduces the problem. 360

361

4.3 Batch reactor: Three-species reaction network (Test 4)

Fig. 8 shows the model results for Test 4, comparing the numerical solution with the corresponding analytical solution. This constitutes a more complex reaction chain that is often integrated into more complex biogeochemical cycling frameworks.

Similarly to the previous tests, the model demonstrates good accuracy and reduced 365 numerical dispersion. For Species A, the Nash-Sutcliffe and RMSE values obtained for 366 summa-openwq were respectively 0.999999987172389 and 1.524e-04 mg/l and for crhm-367 openwq were 0.999999987172368 and 1.524e-04 mg/l. For Species B, the Nash-Sutcliffe 368 and RMSE values obtained for summa-openwq were respectively 0.999999989757722 and 369 1.747e-04 mg/l and for crhm-openwq were 0.999999989757706 and 1.747e-04 mg/l. For 370 Species C, the Nash-Sutcliffe and RMSE values obtained for summa-openwq were respec-371 tively 0.999899420454185 and 1.560e-02 mg/l and for crhm-openwq were 0.999899420454600372 and 1.560e-02 mg/l. 373



Figure 8: Comparison between model results and analytical solution for Test 4

A small phase shift can be noticed for Species C, which is also caused by the reaction 374 network sequence issue mentioned in the previous test (Test 3). In OpenWQ, the hydro-375 models (CRHM and SUMMA) control the model time step. In these tests, the time step 376 used was 15 minutes and was dictated by the data time intervals. The larger the time step is, 377 the larger the shifts are expected to become since the computation of the reactions depends 378 on the concentrations in the antecedent time step. So, in this case, a change in Species A 379 has taken 30 min (2 times 15 minutes) to numerically travel to Species C, producing a slight 380 phase shift. Reducing the length of the time step reduces the problem. 381

4.4 Batch reactor: Nitrogen cycle (Test 5)

382

The results of the nitrogen cycle simulation show good agreement with the numerical solution (Fig. 9). This is a reaction network involving four nitrogen species and it is commonly used in popular models such as HYPE, INCA, and SWAT.

The model shows good agreement with the analytical solution, reinforcing the confi-386 dence in OpenWQ's numerical solution and the implementation of the expression evaluator 387 based on the comprehensive C++ Mathematical Expression Toolkit Library (*ExpTk*) de-388 veloped by Arash Partow (Partow, 1999). For Nref, the Nash-Sutcliffe and RMSE values 389 obtained for summa-openwq were respectively 0.99999999184531 and 6.927e-05 mg/l and 390 for crhm-openwq were 0.99999999184529 and 6.927e-05 mg/l. For Nlab, the Nash-Sutcliffe 391 and RMSE values obtained for summa-openwq were respectively 0.999136304357914 and 392 9.426e-02 mg/l and for crhm-openwq were 0.999136304357383 and 9.426e-02 mg/l. For 393 DON, the Nash-Sutcliffe and RMSE values obtained for summa-openwog were respectively 394 0.999463220618209 and 1.3089e-02 mg/l and for crhm-openwq were 0.999463220618758 and 395 1.309e-02 mg/l. For DIN, the Nash-Sutcliffe and RMSE values obtained for summa-396 openwq were respectively 0.996669866632908 and 7.181e-02 mg/l and for crhm-openwq were 397 0.996669866633266 and 7.181e-02 mg/l. As in the previous reaction networks tested, the 398 RMSE increases as the information travels through the network sequence: DIN ; Nref. 399 This is caused by the same issue raised above related to the time of travel of information 400 through the network sequence computation. 401



Figure 9: Comparison between model results and analytical solution for Tests 5

4.5 Batch reactor: Dissolved oxygen cycle (Test 6)

402

The simulation of the dissolved oxygen cycle is presented in Fig. 10 with two water quality variables (1) dissolved oxygen (DO) and (2) biological oxygen demand (BOD). This is also a common pollution problem simulated by water quality models to evaluate mitigation strategies for contamination caused by urban wastewater discharges in rivers and lakes.



Figure 10: Comparison between model results and analytical solution for Tests 6

The model shows good agreement with observations, capturing the typical DO 407 sag often observed near discharge outlets as predicted by the analytical solution. For 408 DO, the Nash-Sutcliffe and RMSE values obtained for summa-openwq were respectively 409 0.999998916342720 and 1.416e-04 mg/l and for crhm-openwq were 0.999998916342720 and 410 1.416e-04 mg/l. For BOD, the Nash-Sutcliffe and RMSE values obtained for summa-411 openwq were respectively 0.999999874118595 and 2.783e-04 mg/l and for crhm-openwq 412 were 0.999999874118595 and 2.783e-04 mg/l. As in the previous cases, the RMSE increases 413 through the reaction network sequence, which is higher for BOD, despite that the Nash-414 Sutcliffe value remains very high for both species. 415

416 417

4.6 Reactive and non-reactive transport: Continuous point source (Tests 7 and 8)

Fig. 11 compares model results against analytical solutions for Tests 7 and 8, which 418 represent non-reactive and reactive scenarios of transport of a continuous pollution source 419 through porous media. The results show the traveling of solute through the soil profile driven 420 by advection and dispersion processes. In the case of Test 8 in Fig 11b, there is an additional 421 first-order biogeochemical reaction causing a linear decay in the concentrations. Each figure 422 shows the computed concentrations, as well as the model internal derivatives that have 423 been passed into the numerical solver to compute the concentration update. This is useful 424 information to understand the key processes controlling the concentration of the solute as 425 time progresses, as well as the relative contributions of transport and biogeochemistry. 426

Comparing Figs 11a and 11b, it becomes clear that the superposition of the different 427 processes (e.g., biogeochemistry, physical transport) gives rise to a highly non-linear re-428 sponse that the model can capture accurately. For Test 7, the Nash-Sutcliffe and RMSE 429 values obtained were respectively 0.9890 and 0.0747e-05 mg/l, and for Test 8 were respec-430 tively 0.9953 and 0.0392 mg/l. Results show that the Nash-Sutcliffe values remain high, but 431 they are slightly lower when compared to the purely biogeochemical kinetic Tests 1-6. The 432 opposite is observed for RMSE, where values have slightly increased. This minor deterio-433 ration of the numerical results with the activation of the transport routines was expected 434 and it remains within acceptable values. This is caused by the initial conditions required to 435 translate in the host-model to represent those set up in the analytical solutions, as well as 436 the velocity fields computed by the host model and associated numerical dispersion in both 437 OpenWQ and the host-model. 438

439 440

4.7 Reactive and non-reactive transport: Instantaneous point source (Tests 9 and 10)

The non-reactive and reactive tests for an instantaneous point source problem are shown in respectively Figs. 12a and 12b. Like in the previous case there is an additional first-order biogeochemical reaction causing a linear decay in the concentrations of Test 10 in Fig 12b.

For Test 9, the Nash-Sutcliffe and RMSE values obtained were respectively 0.9535 and 4.614e-05 mg/l, and for Test 10 were respectively 0.9521 and 1.409e-05 mg/l. The effect of biogeochemistry in the resulting concentrations can be also clearly observed in the uppermiddle panel of Fig. 12a as compared to the upper-middle panel of Fig. 12b. Similarly to the previous cases of Tests 7 and 8, results show that the Nash-Sutcliffe values remain high, but they are slightly lower when compared to the purely biogeochemical kinetic Tests 1-6.

450 5 Conclusions

This two-part paper describes and applies a methodology for incorporating and crossmodel deploying multi-biogeochemistry and ecological simulations in existing process-based hydro-modeling tools, the OpenWQ framework. OpenWQ was developed based on the motivation that (1) hydrology has a strong impact on the fate of pollution in the environ-



(b) Test 8

Figure 11: Comparison between model results and analytical solution for Tests 7 and 8

ment and (2) more should be done to streamline collaboration between the hydrological
and biogeochemical communities. The first paper describes OpenWQ's concept as a coupler
and customizable biogeochemistry modeling framework to allow cross-model deployment
of eco-lab and co-creation modeling capabilities, providing a transformative direction for
innovation in water quality modeling. This approach aims to address the inherent challenge of water quality modeling of being at the intersection between several fields, including
biogeochemistry, soil science, hydrology, hydrodynamics, and hydrogeology.

This second paper describes the integration of OpenWQ into two hydro-models, SUMMA and CRHM, illustrating how coupling interfaces between the two models have not only enabled water quality simulation capacities in the host hydro-models but, even more importantly, enabled establishing a direct and permanent link for transfer of knowledge, innovation, and technology between modeling communities. Example applications of



(b) Test 10

Figure 12: Comparison between model results and analytical solution for Tests 9 and 10

⁴⁶⁷ pollution studies enabled by the coupling of the tools are provided to begin to address some
⁴⁶⁸ of these fundamental modeling challenges.

The key points of this paper are as follows:

1. The coupling of OpenWQ to CRHM and SUMMA illustrates how the coupling recipe 470 presented in Paper 1 can be deployed to link the model to existing hydro-models 471 even if displaying drastically different modeling structures. The coupling steps, 472 coupling functions, and wrapper interface functions have been designed to optimize 473 the coupling process. It was also demonstrated that new updates to OpenWQ can 474 be integrated into the coupled models without requiring changes in the coupler 475 functions. This provides a permanent, concrete link between the hydrological and 476 biogeochemical research communities, where innovations on either side can be readily 477

integrated through the coupler into a unified modeling framework. 478 479 2. The model development and applications presented illustrate how OpenWQ's flexible 480 multi-chemistry lab capabilities enable addressing a variety of environmental and 481 pollution problems. In Paper 1, we described how such model flexibility has been ma-482 terialized. In this paper, we deployed the model to illustrate how it can be used to look at the nitrogen cycle and eutrophication problems, BOD-DO and DO sag issues, and 484 other environmental problems involving reaction networks run in parallel or sequence. 485 486 3. The model applications presented aim to illustrate how biogeochemical modeling 487 frameworks can be readily deployed across different hydro-models using the same 488 input files. It has been shown how this can be extremely helpful in removing barriers 489 in the selection of the most appropriate modeling framework for specific landscapes 490 and climate regions. 491 492 4. Separating the numerical solver from the physicochemical calculations enables more 493 controlled simulations where the specific impact of different processes can not only 494 be properly quantified but also added or removed as needed for uncertainty analysis. 495 Such model architecture is critical for model scalability by enabling the integration 496 of new modules in a controlled manner.

The work presented here represents the first application of OpenWQ to improve water 498 quality modeling, but there are many additional opportunities to build on this work. The 499 case studies simulated here cover a range of pollution problems and hydrologic processes, 500 but they are naturally limited in terms of the environments and contamination problems 501 covered. The model simulations presented are also based on simple perturbation experi-502 ments, which were necessary for validation purposes; however, more comprehensive model 503 applications should be carried out in the future for comparison with competing modeling ap-504 proaches. The work presented also does not examine different numerical integration options. Research is needed to continue exploring the potential opportunities that OpenWQ creates 506 for testing different biogeochemical, biophysical, and hydrologic processes and their model 507 representations. Access to the OpenWQ source code and the case study tests is provided 508 through the SUMMA website at openwg.readthedocs.io. 509

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511

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516 Appendices

A Deriving analytical solutions for Tests 1 and 2: Single-species batch reaction

The single-species batch reactor tests contemplated involve (a) first-order decay and (b) second-order decay kinetics. Fig. 1 depicts the biogeochemical process. The analytical solution of a single species subject to first-order decay can be obtained by integrating Eq. 24 obtaining Eq. 25. In turn, the analytical solution of a single species subject to second-order decay can be obtained by integrating Eq. 26 obtaining Eq. 27.

$$\frac{dc_A}{dt} = -kc_A \tag{24}$$

$$\frac{dc_A}{dt} = -kc_A \Leftrightarrow \frac{1}{c_A} dc_A = -kdt \Leftrightarrow \int_0^t \frac{1}{c_A} dc_A = \int_0^t -kdt \Leftrightarrow \ln(c_A) \Big|_0^t = -kt \Big|_0^t \Leftrightarrow \ln(c_A(t)) - \ln(c_A(0)) = -k\Delta t \Leftrightarrow \ln\left(\frac{c_A(t)}{c_A(0)}\right) = -k\Delta t \Leftrightarrow \frac{c_A(t)}{c_A(0)} = -k\Delta t \Leftrightarrow c_A(t) = c_A(0)e^{-k\Delta t}$$
(25)

$$\frac{dc_A}{dt} = -kc_A^2 \tag{26}$$

$$\frac{dc_A}{dt} = -kc_A^2 \Leftrightarrow \frac{1}{c_A^2} dc_A = -kdt \Leftrightarrow$$

$$\int_0^t c_A^{-2} dc_A = \int_0^t -kdt \Leftrightarrow -c_A^{-1} \Big|_0^t = -kt \Big|_0^t \Leftrightarrow$$

$$c_A^{-1} \Big|_0^t = kt \Big|_0^t \Leftrightarrow c_A(t)^{-1} - c_A(0)^{-1} = k\Delta t \Leftrightarrow$$

$$\frac{1}{c_A(t)} = \frac{1}{c_A(0)} + k\Delta t \Leftrightarrow c_A(t) = \frac{1}{\frac{1}{c_A(0)} + k\Delta t} \Leftrightarrow$$
(27)

⁵²⁴ B Deriving analytical solution for Test 3: Two-species reaction network

The two-species reaction network with linear decay is represented schematically in Fig. 2. Similarly to the single-species case above with first-order decay (Eq. 25), the fate of Species A in a two-species reaction network only depends on its own concentration and is given by Eq. 28.

$$c_A(t) = c_A(0)e^{(-k_A\Delta t)} \tag{28}$$

However, the fate of Species B depends on the time-evolution of concentrations of both Species A and B. This can be expressed as:

$$\frac{dc_B}{dt} = k_A c_A - k_B c_B \tag{29}$$

To derive the analytical solution of Eq. 29, we can start by replacing $c_A(t)$ by its analytical solution derived previously in Eq. 25. We obtain:

$$\frac{dc_B}{dt} = k_A c_A(0) e^{(-k_A t)} - k_B c_B(t) \Leftrightarrow$$

$$\frac{dc_B}{dt} + k_B c_B = k_A c_A(0) e^{(-k_A t)}$$
(30)

The integrating factor, $e^{(k_B t)}$, can be used to integrate Eq. 30 by parts (i.e., d(uv) = udv + vdu). Multiplying Eq. 30 by the integrating factor, we obtain:

$$\frac{dc_B}{dt}e^{(k_Bt)} + k_Bc_B(t)e^{(k_Bt)} = k_Ac_A(0)e^{(-k_At)}e^{(k_Bt)} \Leftrightarrow \frac{dc_B}{dt}e^{(k_Bt)} + k_Bc_B(t)e^{(k_Bt)} = k_Ac_A(0)e^{(k_B-k_A)t}$$
(31)

Using the general expression for integration by parts and considering that functions uand v are given by $u = c_B$ and $v = e^{(k_B t)}$, we obtain:

$$d\left(c_B e^{(k_B t)}\right) = c_B d\left(e^{(k_B t)}\right) + e^{(k_B t)} dc_B \tag{32}$$

Replacing the derivative of $e^{(k_B t)}$ in Eq. 32, where $d(e^{(k_B t)}) = k_B e^{(k_B t)} dt$, leads to:

$$d(uv) = udv + vdu \Leftrightarrow$$

$$d(c_B e^{(k_B t)}) = c_B [k_B e^{(k_B t)} dt] + e^{(k_B t)} dc_B$$
(33)

or

537

$$\frac{d}{dt}(c_B e^{(k_B t)}) = c_B k_B e^{(k_B t)} + e^{(k_B t)} \frac{dc_B}{dt}$$
(34)

The right-hand side of Eq. 34 is the same as the left-hand side of Eq. 31, so the same applies to the other side of the equations. Thus, we can re-write the equations as

$$\frac{d}{dt} \left(c_B e^{(k_B t)} \right) = k_A c_A(0) e^{(-k_A t)} e^{(k_B t)} \Leftrightarrow
\frac{d}{dt} \left(c_B e^{(k_B t)} \right) = k_A c_A(0) e^{\left[(k_B - k_A) t \right]} \Leftrightarrow
d \left(c_B e^{(k_B t)} \right) = k_A c_A(0) e^{\left[(k_B - k_A) t \right]} dt$$
(35)

540 Integrating both sides of the equation, we obtain

$$\int_{0}^{t} d(c_{B}e^{(k_{B}t)}) = \int_{0}^{t} k_{A}c_{A}(0)e^{\left[(k_{B}-k_{A})t\right]}dt \Leftrightarrow c_{B}(t)e^{(k_{B}t)} = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}e^{\left[(k_{B}-k_{A})t\right]} + C',$$
(36)

where C' is the integration constant. At t = 0, c_B becomes $c_B(0)$, so Eq. 36 can be written as

$$c_B(0)e^0 = \frac{k_A c_A(0)}{k_B - k_A} e^0 \Leftrightarrow$$

$$\Leftrightarrow C' = c_B(0) - \frac{k_A c_A(0)}{k_B - k_A},$$
(37)

⁵⁴³ which can be replaced in Eq. 36 to give:

$$c_{B}(t)e^{(k_{B}t)} = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}e^{\left[(k_{B}-k_{A})t\right]} + c_{B}(0) - \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \Leftrightarrow$$

$$c_{B}(t)e^{(k_{B}t)} = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{\left[(k_{B}-k_{A})t\right]} - 1\right] + c_{B}(0) \Leftrightarrow$$

$$c_{B}(t) = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{e^{\left[(k_{B}-k_{A})t\right]}}{e^{(k_{B}t)}} - \frac{1}{e^{(k_{B}t)}}\right] + \frac{c_{B}(0)}{e^{(k_{B}t)}} \Leftrightarrow$$

$$c_{B}(t) = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{(-k_{A}t)} - e^{(-k_{B}t)}\right] + c_{B}(0)e^{(-k_{B}t)}$$
(38)

C Deriving analytical solution for Test 4: Three-species reaction network

A three-species reaction network with linear decay is represented schematically in Fig.
Concentrations of Species A and Species B are given by respectively Eqs. 28 and 36.
The partial differential equation for the concentration of Species C can be written as:

$$\frac{dc_C}{dt} = k_B c_B - k_C c_C \tag{39}$$

The expression for $c_B(t)$ can be obtained from Eq. 36 and replaced in Eq. 40. We obtain:

$$\frac{dc_C}{dt} = k_B \left(\frac{k_A c_A(0)}{k_B - k_A} \left[e^{(-k_A t)} - e^{(-k_B t)} \right] + c_B(0) e^{(-k_B t)} \right) -k_C c_C \Leftrightarrow \frac{dc_C}{dt} + k_C c_C = k_B \left(\frac{k_A c_A(0)}{k_B - k_A} \left[e^{(-k_A t)} - e^{-k_B t} \right] + c_B(0) e^{(-k_B t)} \right)$$
(40)

The integrating factor used in Section B to enable integration of parts, $e^{(k_C t)}$, can be used here as well (i.e., d(uv) = udv + vdu). Multiplying Eq. 40 by this integrating factor, we obtain:

$$\frac{dc_C}{dt}e^{(k_Ct)} + k_Cc_C(t)e^{(k_Ct)} = k_B\left(\frac{k_Ac_A(0)}{k_B - k_A}\left[e^{(-k_At)} - e^{(-k_Bt)}\right] + c_B(0)e^{(-k_Bt)}\right)e^{(k_Ct)}$$
(41)

We can use the same expression derived to enable integration by parts in the previous case (Eq. 34) but replacing Species B by Species C: $\frac{d}{dt}\left(c_{C}e^{(k_{C}t)}\right) = c_{C}k_{C}e^{(k_{C}t)} + \frac{dc_{C}}{dt}e^{k_{C}t}$. Also, similarly to the previous case, the right-hand side of this equation is the same as the

$_{\tt 557}$ $\,$ left-hand side of Eq. 41, so the same applies to the other side of the equations. Thus, we

⁵⁵⁸ can re-write the equations as follows:

$$\frac{d}{dt}\left(c_{C}e^{k_{C}t}\right) = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{-k_{A}t} - e^{-k_{B}t}\right] + c_{B}(0)e^{-k_{B}t}\right)e^{(k_{C}t)} \Leftrightarrow d\left(c_{C}e^{k_{C}t}\right) = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{-k_{A}t} - e^{-k_{B}t}\right] + c_{B}(0)e^{-k_{B}t}\right)e^{(k_{C}t)}dt \Leftrightarrow d\left(c_{C}e^{k_{C}t}\right) = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{(k_{C}-k_{A})t} - e^{(k_{C}-k_{B})t}\right] + c_{B}(0)e^{(k_{C}-k_{B})t}\right] + c_{B}(0)e^{(k_{C}-k_{B})t}\right)dt$$
(42)

⁵⁵⁹ Integrating both sides of the equation we obtain:

$$\int_{0}^{t} d(c_{C}e^{(k_{C}t)}) =$$

$$\int_{0}^{t} \left[k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[e^{\left[(k_{C}-k_{A})t \right]} - e^{\left[(k_{C}-k_{B})t \right]} \right] \right] +$$

$$c_{B}(0)e^{\left[(k_{C}-k_{B})t \right]} \right] \Rightarrow \qquad (43)$$

$$c_{C}(t)e^{(k_{C}t)} = k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{1}{k_{C}-k_{A}} e^{\left[(k_{C}-k_{A})t \right]} -$$

$$\frac{1}{k_{C}-k_{B}} e^{\left[(k_{C}-k_{B})t \right]} \right] + \frac{c_{B}(0)}{k_{C}-k_{B}} e^{\left[(k_{C}-k_{B})t \right]} \right) + C'',$$

where C" is the integration constant. At t = 0, c_C becomes $c_C(t)$, so Eq. 43 becomes

$$c_{C}(0)e^{(k_{C}0)} = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}e^{0}-\frac{1}{k_{C}-k_{B}}e^{0}\right]\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}e^{0}\right) + C'' \Leftrightarrow$$

$$c_{C}(0)1 = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}1-\frac{1}{k_{C}-k_{B}}1\right]\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}1\right) + C'' \Leftrightarrow C'' =$$

$$c_{C}(0) - k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}-\frac{1}{k_{C}-k_{B}}\right]\right) + \frac{c_{B}(0)}{k_{C}-k_{B}}\right),$$
(44)

$_{561}$ which can be replaced in Eq. 43 to give:

$$c_{C}(t)e^{(k_{C}t)} = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}e^{\left[(k_{C}-k_{A})t\right]}-\frac{1}{k_{C}-k_{B}}e^{\left[(k_{C}-k_{B})t\right]}\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}e^{\left[(k_{C}-k_{B})t\right]}\right) +$$
(45)
$$c_{C}(0) - k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}-\frac{1}{k_{C}-k_{B}}\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}\right)$$

This can be rearranged to give:

$$c_{C}(t) = k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{1}{k_{C}-k_{A}} e^{(-k_{A}t)} - \frac{1}{k_{C}-k_{B}} e^{(-k_{B}t)} \right] + \frac{c_{B}(0)}{k_{C}-k_{B}} e^{(-k_{B}t)} \right] + c_{C}(0)e^{(-k_{C}t)} - k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{1}{k_{C}-k_{A}} - \frac{1}{k_{C}-k_{B}} \right] + \frac{c_{B}(0)}{k_{C}-k_{B}} \right) e^{(-k_{C}t)},$$

$$(46)$$

₅₆₂ and then simplified to:

$$c_{C}(t) = k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{e^{(-k_{A}t)}}{k_{C}-k_{A}} - \frac{e^{(-k_{B}t)}}{k_{C}-k_{B}} \right] + \frac{c_{B}(0)e^{(-k_{B}t)}}{k_{C}-k_{B}} \right) + c_{C}(0)e^{(-k_{C}t)} - (47)$$

$$k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{e^{(-k_{C}t)}}{k_{C}-k_{A}} - \frac{e^{(-k_{C}t)}}{k_{C}-k_{B}} \right] + \frac{c_{B}(0)e^{(-k_{C}t)}}{k_{C}-k_{B}} \right)$$

⁵⁶³ D Deriving analytical solution for Test 5: Nitrogen cycle

The simulation of the nitrogen cycle is very common in freshwater water quality studies 564 because it allows for addressing issues with nutrient pollution and eutrophication. The 565 nitrogen cycle simulated involves different nitrogen species (Fig. 4). The conceptual cycling 566 framework used is similar to that used in catchment-scale models like HYPE Lindström et 567 al. (2010), CRHM-WQ (Costa et al., 2021) models, SWAT (Arnold et al., 2012), and INCA 568 ("A semi-distributed ntegrated itrogen model for multiple source assessment in tchments 569 (INCA): Part I — model structure and process equations", 1998). The analytical solutions 570 for the different nitrogen species can be obtained directly (or derived) from Eq. 28, Eq. 38, 571 or Eq. 47 depending on their positioning in the reaction chain: 572

1. **refractoryN** (*Nref*): Similar to Eq. 28, which for this specific process and N species reads as:

$$c_{Nref}(t) = c_{Nref}(0)e^{\left[-(k_{degr} + k_{diss1})t\right]}$$

$$\tag{48}$$

2. labile N (Nref): Similar to Eq. 36, which for this specific processes and N species reads as:

$$c_{Nlab}(t) = \frac{k_{degr}c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \begin{bmatrix} e^{(-k_{degr}t)} - e^{\left[-(k_{min} + k_{diss2})t\right]} \\ + c_{Nlab}(0)e^{\left[-(k_{min} + k_{diss2})t\right]} \end{bmatrix}$$
(49)

3. **DON**: Needs specific derivation using Eqs. 28 and 36 (see below)

4. **DIN**: Needs specific derivation using Eqs. 49 and 28 (see below)

 $_{575}$ **DON** The concentration of DON does not depend on itself, but rather on the evolution of the concentration of Nref and Nlab. The differential equation can be written $_{577}$ as:

$$\frac{dc_{DON}}{dt} = k_{diss1}c_{Nref} + k_{diss2}c_{Nlab} \tag{50}$$

Analytical solutions suitable for c_{Nref} and c_{Nlab} , due to their positioning in the rection chain, have been derived before (Eqs. 28 and 38, respectively). These solutions can be replaced in Eq. 50.

$$\frac{dc_{DON}}{dt} = k_{diss1} \left[c_{Nref} e^{\left[-(k_{degr} + k_{diss1})t \right]} \right] + k_{diss2} \left[\frac{k_{degr} c_{Nref}}{k_{min} + k_{diss2} - k_{degr}} \left(e^{\left(-k_{degr}t \right)} - e^{\left[-(k_{min} + k_{diss2})t \right]} \right) + c_{Nlab}(0) e^{\left[-(k_{min} + k_{diss2})t \right]} \right]$$
(51)

Since this solution is independent of the concentration of DON, it can be integrated in its current state.

$$dc_{DON} = k_{diss1} \left[c_{Nref} e^{\left[-(k_{degr} + k_{diss1})t \right]} \right] dt$$

$$+k_{diss2} \left[\frac{k_{degr} c_{Nref}}{k_{min} + k_{diss2} - k_{degrd}} \left(e^{\left(-k_{degr}t \right)} - e^{\left[-(k_{min} + k_{diss2})t \right]} \right] \right] + c_{Nlab}(0) e^{\left[-(k_{min} + k_{diss2})t \right]} dt \Leftrightarrow$$

$$c_{DON}(t) = \frac{k_{diss1} c_{Nref}(t)}{-(k_{degr} + k_{diss1})} e^{\left[-(k_{degr} + k_{diss1})t \right]} + \frac{k_{diss2} k_{degr} c_{Nref}(t)}{-k_{degr}} \left(\frac{e^{\left(-k_{degr}t \right)}}{-(k_{min} + k_{diss2})t \right]} - \frac{e^{\left[-(k_{min} + k_{diss2})t \right]}}{-(k_{min} + k_{diss2})} \right) + \frac{k_{diss2} c_{Nlab}(0)}{-(k_{min} + k_{diss2})} e^{\left[-(k_{min} + k_{diss2})t \right]} + C'''$$
(52)

Considering that at t = 0, $c_{DON}(t)$ becomes $c_{DON}(0)$, we can determine C''' as follows:

$$C''' = c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{1}{-k_{degr}} - \frac{1}{-(k_{min} + k_{diss2})}\right) + \frac{k_{diss2}c_{Nlab}(0)}{-(k_{min} + k_{diss2})}$$
(53)

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Replacing Eq. 53 into Eq. 52 we obtain:

$$c_{DON}(t) = \frac{k_{diss1}c_{Nref}(t)}{-(k_{degr} + k_{diss1})} e^{\left[-(k_{degr} + k_{diss1})t\right]} + \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{e^{(-k_{degr}t)}}{-k_{degr}} - \frac{e^{\left[-(k_{min} + k_{diss2})t\right]}}{-(k_{min} + k_{diss2})}\right) + \frac{k_{diss2}c_{Nlab}(0)}{-(k_{min} + k_{diss2})}$$
(54)
$$e^{\left[-(k_{min} + k_{diss2})t\right]} + c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{1}{-k_{degr}} - \frac{1}{-(k_{min} + k_{diss2})}\right) + \frac{k_{diss2}c_{Nlab}(0)}{-(k_{min} + k_{diss2})}$$

Eq. 54 can be simplified to obtain:

$$c_{DON}(t) = -\frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} e^{\left[-(k_{degr} + k_{diss1})t\right]} + \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{e^{\left[k_{min} + k_{diss2}t\right]}}{k_{min} + k_{diss2}} - \frac{e^{\left(-k_{degr}t\right)}}{k_{degr}}\right) - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}}$$

$$e^{\left[-(k_{min} + k_{diss2})t\right]} + c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2}}$$

$$\left(\frac{1}{k_{min} + k_{diss2}} - \frac{1}{k_{degr}}\right) - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}}$$

586

$$\frac{dc_{DIN}}{dt} = k_{min}c_{Nlab} - (k_{denit} + k_{plantup})c_{DIN}$$
(56)

The concentration of $c_{Nlab}(t)$ has been derived above in Eq. 49 and can be replaced in Eq. 56 to give:

$$\frac{dc_{DIN}}{dt} = k_{min} \left(\frac{k_{degr} c_{(Nref,0)}}{k_{min} + k_{diss2} - k_{degr}} \left[e^{(-k_{degr}t)} - e^{\left[- (k_{min} + k_{diss2})t \right]} \right] + c_{Nlab}(0) e^{\left[- (k_{min} + k_{diss2})t \right]} \right) - (k_{denit} + k_{plantup}) c_{DIN}(t) \Leftrightarrow \left(\frac{dc_{DIN}}{dt} = \frac{k_{min} k_{degr} c_{(Nref,0)}}{k_{min} + k_{diss2} - k_{degr}} \left[e^{(-k_{degr}t)} - e^{\left[- (k_{min} + k_{diss2})t \right]} \right] + k_{min} c_{Nlab}(0) e^{\left[- (k_{min} + k_{diss2})t \right]} - (k_{denit} + k_{plantup}) c_{DIN}$$

$$(57)$$

The same integrating factor used before, which in this context reads $e^{\left[(k_{denit}+k_{plantup})t\right]}$, can be used here to allow similar integration by parts (i.e., d(uv) = udv + vdu). Multiplying Eq. 57 by the integration factor, we obtain:

$$\frac{dc_{DIN}}{dt}e^{\left[(k_{denit}+k_{plantup})t\right]} + (k_{denit}+k_{plantup})$$

$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]}$$

$$= \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\left[e^{(-k_{degr}t)} - e^{\left[-(k_{min}+k_{diss2})t\right]}\right]$$

$$e^{\left[(k_{denit}+k_{plantup})t\right]} + k_{min}c_{Nlab}(0)e^{\left[-(k_{min}+k_{diss2})t\right]}$$

$$e^{\left[(k_{denit}+k_{plantup})t\right]}$$
(58)

Now we can use the same expression derived for integration by parts of previous cases (i.e., Eq. 34) and adapt it to *DIN* and the respective transformations:

$$\frac{d}{dt} \Big[c_{DIN} e^{\left((k_{denit} + k_{plantup})t \right)} \Big] = \frac{dc_{DIN}}{dt} e^{\left[(k_{denit} + k_{plantup})t \right]} + (k_{denit} + k_{plantup})c_{DIN} e^{\left[(k_{denit} + k_{plantup})t \right]}$$
(59)

Similarly to previous cases, the right-hand side of this equation is the same as the lefthand side of Eq. 57, so the same applies to the other side of the equations so that we can re-write the equations as:

$$\frac{d}{dt} \left(c_{DIN} e^{\left[(k_{denit} + k_{plantup})t \right]} \right) = \frac{k_{min} k_{degr} c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \\
\left[e^{(-k_{degr}t)} - e^{\left[- (k_{min} + k_{diss2})t \right]} \right] e^{\left[(k_{denit} + k_{plantup})t \right]} + k_{min} c_{Nlab}(0) e^{\left[- (k_{min} + k_{diss2})t \right]} e^{\left[(k_{denit} + k_{plantup})t \right]} \Leftrightarrow d\left(c_{DIN} e^{\left[(k_{denit} + k_{plantup})t \right]} \right) = \left(\frac{k_{min} k_{degr} c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \right) \\
\left[e^{\left[((k_{denit} + k_{plantup}) - k_{degr})t \right]} - e^{\left[(k_{denit} + k_{plantup} - k_{min} - k_{diss2})t \right]} \right] \\
+ k_{min} c_{Nlab}(0) e^{\left[(k_{denit} + k_{plantup} - k_{min} - k_{diss2})t \right]} \right) dt$$
(60)

⁵⁹⁷ Now we can integrate the equation with respect to $c_{DIN}(t)$ and t:

$$\int d\left(c_{DIN}e^{\left[(k_{denit}+k_{plantup})t\right]}\right) = \int \left(\frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\right)$$

$$\left[e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]} - e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}\right]$$

$$+k_{min}c_{Nlab}(0)e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}\right)dt \Leftrightarrow$$

$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]} = \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}$$

$$\left[e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]} - \frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}}{(k_{denit}+k_{plantup})-k_{min}-k_{diss2}}\right]$$

$$+k_{min}c_{Nlab}(0)\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}}{(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t}$$

⁵⁹⁸ Considering that at t = 0, $c_{DIN}(t)$ becomes $c_{DIN}(0)$, we can determine the constant ⁵⁹⁹ C'''' as follows:

$$C'''' = c_{DIN}(0) - \frac{k_{miner}k_{degr}c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \begin{bmatrix} \frac{1}{(k_{denit} + k_{plantup}) - k_{degr}} - \frac{1}{(k_{denit} + k_{plantup}) - k_{min} - k_{diss2}} \end{bmatrix}$$
(62)
$$-k_{miner}c_{Nlab}(0) \frac{1}{(k_{denit} + k_{plantup}) - k_{min} - k_{diss2}}$$

Replacing the integrating constant in Eq. 61 we obtain:

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$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]} = \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}$$

$$\left[\frac{e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]}}{k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]} + k_{min}c_{Nlab}(0)\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right]$$

$$+c_{DIN}(0) - \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\left[\frac{1}{k_{denit}+k_{plantup}-k_{degr}} - \frac{1}{k_{denit}+k_{plantup}-k_{diss2}}\right] - k_{min}c_{Nlab}(0)$$

$$\left(\frac{1}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right)$$

$$(63)$$

⁶⁰¹ E Deriving analytical solution for Test 6: DO-BOD cycle

Dissolved oxygen (DO) depletion in rivers and lakes is a common consequence of pollution. It typically occurs near discharge outlets of untreated wastewater, which can lead to high levels of organic matter that, during decomposition by aerobic bacteria, causes a DOsag in the water column. For purposes of DO dynamics, such organic matter is typically translated into units of biological oxygen demand (BOD). Fig. 5 depicts the processes involved, and the Streeter–Phelps equation (Streeter & Phelps, 1925) has long been derived for such problems. It reads:

$$\frac{dD}{dt} = k_1 L_t - k_2 D \tag{64}$$

The equation states that the total change in oxygen deficit (D) is equal to the difference between the two rates of deoxygenation and reaeration at any time. The derivation of the Streeter-Phelps equation is well known, so it won't be repeated here.

$$D = \frac{k_1 L_a}{k_2 - k_1} (e^{(-k_1 t)} - e^{(-k_2 t)}) + D_a e^{(-k_2 t)}$$
(65)

where D is the saturation deficit in g/m^3 , which can be derived from the dissolved oxygen concentration at saturation minus the actual dissolved oxygen concentration ($D = DO_{sat} - DO$); k_1 is the deoxygenation rate, usually in d^{-1} ; k_2 is the reaeration rate, usually in d^{-1} ; L_a is the initial oxygen demand of organic matter in the water, also called the ultimate BOD (BOD at time $t = \infty$). The unit of L_a is g/m^3 ; k_1 ; L_t is the oxygen demand remaining at time t, $L_t = L_a e^{(-k_1 t)}$; D_a is the initial oxygen deficit $[g/m^3]$; and t is the elapsed time in days.

⁶¹⁹ F Reactive and non-reactive transport: Instantaneous point source (Tests 7 and 8)

The general 1D advection-dispersion-reaction partial differential equation is given by

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c \tag{66}$$

where c is the concentration of the tracer, t is time, D_x is the dispersion coefficient in the flow direction, x is the distance in the flow direction, ν is the flow velocity in the x direction, and λ is a linear reaction term that can be set to zero for non-reactive tracers. Various analytical solutions for Eq. 66 have been derived and compiled by Wexler (1992) for different idealized scenarios.

Eq. 66 has been modified by Wexler (1992) to include an instantaneous point source and it reads:

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c + \frac{Q}{n} \partial t c_0 \delta(x - x_c) \delta(t - t'_c)$$
(67)

where t'_c is the time at which the instantaneous point source is activated, Q is the fluid injection rate, n is the aquifer porosity, δ is the Dirac delta (impulse) function, and \mathbf{x}_c is the location of the instantaneous point source. The analytical solution to Eq. 67 has been derived by Bear (1979) considering the following boundary conditions and initial conditions.

6321. Boundary conditions (Neumann or second-type):633 $c, \nabla c = 0, \quad x = \pm \infty; \ y = \pm \infty$ 6342. Initial conditions:635 $c^{t=0} = 0, \quad -\infty < x < \infty; -\infty < y < \infty,$

Eq. 68 shows the resulting analytical solution that can be used for solutes that are non-reactive (Test 7) or reactive (Test 8) by simply setting $\lambda = 0$ in the former (Test 7).

$$c(x) = \frac{c_0 Q}{4n\pi\gamma D_x} \exp\left[\frac{\nu(x-x_0)}{2D_x}\right]$$

$$\frac{1}{t} \exp\left[-\left(\frac{\nu^2}{4D_x}\right)\tau - \frac{(x-Xc)^2}{4D_x\tau}\right],$$
(68)

The equation assumes that (1) the fluid has constant density and viscosity, (2) the 638 solute may be subject to first-order chemical transformation, (3) the flow is uniform in 639 the x-direction only, (4) the flow velocity is constant, and (5) the longitudinal dispersion 640 coefficient (D) is constant. The solution has a mathematical singularity in the point source 641 region (Abdelaziz et al., 2013) because it takes the form of the exponential integral when 642 both $x - x_c$ approaches zero, hence becoming infinite at $\tau = 0$. However, according to 643 Wexler (1989) the solution is still valid as long as $(x - x_c)^2$ is larger than ν^2 , as it is the 644 case in our simulation tests. 645

G Reactive and non-reactive transport: Continuous point source (Tests 9 and 10)

The analytical solution for the transport of solutes with a continuous point source has been derived in Wexler (1992) after modification from Bear (1979) and Genuchten & Alves (1982).

$$c(x) = \frac{c_0}{2} \left\{ \exp\left[\frac{x}{2D_x}(\nu - U)\right] \operatorname{erfc}\left[\frac{x - Ut}{2\sqrt{D_x t}}\right] + \exp\left[\frac{x}{2D_x}(\nu + U)\right] \operatorname{erfc}\left[\frac{x + Ut}{2\sqrt{D_x t}}\right] \right\}$$
(69)

where c_0 is the concentration of the continuous point source, $U = \sqrt{\nu^2 + 4\lambda D}$ and erfc is the complementary error function. Parameter λ is the first-order reaction rate used for Test 10, which can be set to zero for Test 9.

654 References

- Abdelaziz, R., Pearson, A., & Merkel, B. (2013). Lattice Boltzmann modeling for tracer
 test analysis in a fractured Gneiss aquifer. *Natural Science*, 5, 368-374. doi: 10.4236/
 ns.2013.53050
- Alexander, A. C., Levenstein, B., Sanderson, L. A., Blukacz-Richards, E. A., & Chambers,
 P. A. (2020). How does climate variability affect water quality dynamics in canada's oil
 sands region? Science of The Total Environment, 732, 139062. doi: https://doi.org/
 10.1016/j.scitotenv.2020.139062
- Arnold, J. G., Moriasi, D. N., Gassman, P. W., Abbaspour, K. C., White, M. J., Srinivasan,
 R., ... Jha, M. K. (2012). SWAT: Model Use, Calibration, and Validation. Trans.
 ASABE, 55(4), 1491–1508.
- Bear, J. (1979). *Hydraulics of groundwater*. McGraw-Hill series in water resources and environmental engineering. McGraw-Hill, New York.
- Beven, K. (2006). A manifesto for the equifinality thesis. *Journal of Hydrology*, 320(1), 18-36. (The model parameter estimation experiment) doi: https://doi.org/10.1016/j.jhydrol .2005.07.007
- Bussi, G., Janes, V., Whitehead, P. G., Dadson, S. J., & Holman, I. P. (2017). Dynamic response of land use and river nutrient concentration to long-term climatic changes. *The Science of the total environment*, 590-591, 818-831. doi: 10.1016/j.scitotenv.2017.03.069
- ⁶⁷³ Bussi, G., Whitehead, P. G., Jin, L., Taye, M. T., Dyer, E., Hirpa, F. A., ... Charles,
 ⁶⁷⁴ K. J. (2021). Impacts of climate change and population growth on river nutrient loads
 ⁶⁷⁵ in a data scarce region: The upper awash river (ethiopia). Sustainability, 13. doi: 10.3390/su13031254
- Celia, M. A., Bouloutas, E. T., & Zarba, R. L. (1990). A general mass-conservative numerical solution for the unsaturated flow equation. Water Resources Research, 26(7), 1483-1496.
 doi: https://doi.org/10.1029/WR026i007p01483
- Charlton, M. B., Bowes, M. J., Hutchins, M. G., Orr, H. G., Soley, R., & Davison, P. (2018).
 Mapping eutrophication risk from climate change: Future phosphorus concentrations in
 english rivers. Science of The Total Environment, 613-614, 1510-1526. doi: https://
 doi.org/10.1016/j.scitotenv.2017.07.218
- Clark, M. P., Nijssen, B., Lundquist, J. D., Kavetski, D., Rupp, D. E., Woods, R. A., ...
 Rasmussen, R. M. (2015a). A unified approach for process-based hydrologic modeling: 1.
 modeling concept. *Water Resources Research*, 51(4), 2498-2514. Retrieved from https://
 agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/2015WR017198
 doi.org/10.1002/2015WR017198
- ⁶⁸⁹ Clark, M. P., Nijssen, B., Lundquist, J. D., Kavetski, D., Rupp, D. E., Woods, R. A.,
 ⁶⁹⁰ ... Marks, D. G. (2015b). A unified approach for process-based hydrologic model ⁶⁹¹ ing: 2. model implementation and case studies. Water Resources Research, 51(4), 2515 ⁶⁹² 2542. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/
 ⁶⁹³ 2015WR017200 doi: https://doi.org/10.1002/2015WR017200
- ⁶⁹⁴ Clark, M. P., Zolfaghari, R., Green, K. R., Trim, S., Knoben, W. J. M., Bennett, A., ⁶⁹⁵ ... Spiteri, R. J. (2021a). The numerical implementation of land models: Problem

- ⁶⁹⁶ formulation and laugh tests. Journal of Hydrometeorology, 22(6), 1627 1648. doi: ⁶⁹⁷ 10.1175/JHM-D-20-0175.1
- ⁶⁹⁸ Clark, M. P., Zolfaghari, R., Green, K. R., Trim, S., Knoben, W. J. M., Bennett, A.,
 ⁶⁹⁹ ... Spiteri, R. J. (2021b). The numerical implementation of land models: Problem
 ⁷⁰⁰ formulation and laugh tests. *Journal of Hydrometeorology*, 22(6), 1627 1648. doi:
 ⁷⁰¹ https://doi.org/10.1175/JHM-D-20-0175.1
- Costa, D., Baulch, H., Elliott, J., Pomeroy, J., & Wheater, H. (2020). Modelling nutrient dynamics in cold agricultural catchments: A review. *Environmental Modelling & Software*, 124, 104586. doi: https://doi.org/10.1016/j.envsoft.2019.104586
- Costa, D., Pomeroy, J. W., Brown, T., Baulch, H., Elliott, J., & Macrae, M. (2021).
 Advances in the simulation of nutrient dynamics in cold climate agricultural basins:
 Developing new nitrogen and phosphorus modules for the Cold Regions Hydrological
 Modelling Platform. Journal of Hydrology, 603, 126901. doi: https://doi.org/10.1016/
 j.jhydrol.2021.126901
- ⁷¹⁰ Costa, D., Sutter, C., Shepherd, A., Jarvie, H., Wilson, H., Elliott, J., ... Macrae, M. ⁷¹¹ (2022). Impact of climate change on catchment nutrient dynamics: insights from around ⁷¹² the world. *Environmental Reviews*, $\theta(0)$, null. doi: 10.1139/er-2021-0109
- Crossman, J., Futter, M. N., Whitehead, P. G., Stainsby, E., Baulch, H. M., Jin, L., ... Dillon, P. J. (2014). Flow pathways and nutrient transport mechanisms drive hydrochemical sensitivity to climate change across catchments with different geology and topography. *Hydrol. Earth Syst. Sci.*, 18, 5125-5148. doi: 10.5194/hess-18-5125-2014
- Effects of lake water level fluctuation due to drought and extreme winter precipitation
 on mixing and water quality of an alpine lake, case study: Lake arrowhead, california.
 (2020). Science of The Total Environment, 714, 136762. doi: https://doi.org/10.1016/
 j.scitotenv.2020.136762
- El-Khoury, A., Seidou, O., Lapen, D. R., Que, Z., Mohammadian, M., Sunohara, M., & Bahram, D. (2015). Combined impacts of future climate and land use changes on discharge, nitrogen and phosphorus loads for a canadian river basin. *Journal of Environmental Management*, 151, 76-86. doi: https://doi.org/10.1016/j.jenvman.2014.12.012
- Ervinia, A., Huang, J., & Zhang, Z. (2020). Nitrogen sources, processes, and associated impacts of climate and land-use changes in a coastal china watershed: Insights from the inca-n model. *Marine Pollution Bulletin*, 159, 111502. doi: https://doi.org/10.1016/j.marpolbul.2020.111502
- Field, C. B., & Barros, V. R. (2014). Climate change 2014: Impacts, adaptation, and vulnerability: Working group ii contribution to the fifth assessment report of the intergovernmental panel on climate change (Tech. Rep.).
- Genuchten, M. T., & Alves, W. (1982). Analytical solutions of the one-dimensional convective-dispersive solute transport equation. US Department of Agriculture.
- Jarsjö, J., Törnqvist, R., & Su, Y. (2017). Climate-driven change of nitrogen reten tion-attenuation near irrigated fields: multi-model projections for central asia. Envi ronmental Earth Sciences, 76, 117. doi: 10.1007/s12665-017-6418-y
- Jeppesen, E., Kronvang, B., Olesen, J. E., Audet, J., Søndergaard, M., Hoffmann, C. C., ...
 Özkan, K. (2011). Climate change effects on nitrogen loading from cultivated catchments
 in europe: implications for nitrogen retention, ecological state of lakes and adaptation.
 Hydrobiologia, 663, 1-21. doi: 10.1007/s10750-010-0547-6
- Lindström, G., Pers, C., Rosberg, J., Strömqvist, J., & Arheimer, B. (2010). Development and testing of the HYPE (Hydrological Predictions for the Environment) water quality model for different spatial scales. *Hydrology Research*, 41 (3-4), 295. doi: 10.2166/nh.2010
 .007
- Luo, C., Li, Z., Liu, H., Li, H., Wan, R., Pan, J., & Chen, X. (2020). Differences in the responses of flow and nutrient load to isolated and coupled future climate and land use changes. *Journal of Environmental Management*, 256, 109918. doi: https://doi.org/10.1016/j.jenvman.2019.109918

Ockenden, M. C., Deasy, C. E., Benskin, C., Beven, K. J., Burke, S., Collins, A. L., ... 749 Haygarth, P. M. (2016). Changing climate and nutrient transfers: Evidence from high 750 temporal resolution concentration-flow dynamics in headwater catchments. Science of 751 The Total Environment, 548-549, 325-339. doi: https://doi.org/10.1016/j.scitotenv.2015 752 .12.086 753

- Ockenden, M. C., Hollaway, M. J., Beven, K. J., Collins, A. L., Evans, R., Falloon, P. D., ... 754 Haygarth, P. M. (2017). Major agricultural changes required to mitigate phosphorus losses 755 under climate change. Nature Communications, 8, 161. doi: 10.1038/s41467-017-00232-0 756
- Onishi, T., Yoshino, J., Hiramatsu, K., & Somura, H. (2020). Developing a hydro-chemical 757 model of ise bay watersheds and the evaluation of climate change impacts on discharge 758 and nitrate-nitrogen loads. Limnology, 21, 465-486. doi: 10.1007/s10201-020-00622-2 759
- Partow, A. (1999). ExprTk C++ Mathematical Expression Library. 760
- Pomeroy, J., Brown, T., Fang, X., Shook, K., Pradhananga, D., Armstrong, R., ... Lopez 761 Moreno, J. (2022). The cold regions hydrological modelling platform for hydrological 762 diagnosis and prediction based on process understanding. Journal of Hydrology, 615, 763 128711. doi: https://doi.org/10.1016/j.jhydrol.2022.128711 764
- Pomerov, J. W., Grav, D. M., Brown, T., Hedstrom, N. R., Quinton, W. L., Granger, R. J., 765 & Carey, S. K. (2007). The cold regions hydrological model: A platform for basing process 766 representation and model structure on physical evidence. Hydrological Processes, 21(19), 767 2650-2667. doi: {10.1002/hyp.6787} 768
- Raimonet, M., Thieu, V., Silvestre, M., Oudin, L., Rabouille, C., Vautard, R., & Garnier, 769 J. (2018). Landward perspective of coastal eutrophication potential under future climate 770 change: The seine river case (france). Frontiers in Marine Science, 5, 136. 771
- Reggiani, P., Sivapalan, M., & Majid Hassanizadeh, S. (1998). A unifying framework for 772 watershed thermodynamics: balance equations for mass, momentum, energy and entropy, and the second law of thermodynamics. Advances in Water Resources, 22(4), 367-398. 774 doi: https://doi.org/10.1016/S0309-1708(98)00012-8 775
- Renard, B., Kavetski, D., Kuczera, G., Thyer, M., & Franks, S. W. (2010). Understand-776 ing predictive uncertainty in hydrologic modeling: The challenge of identifying input 777 and structural errors. Water Resources Research, 46(5). doi: https://doi.org/10.1029/ 778 2009WR008328 779
- Rood, S. B., Pan, J., Gill, K. M., Franks, C. G., Samuelson, G. M., & Shepherd, A. 780 (2008). Declining summer flows of rocky mountain rivers: Changing seasonal hydrol-781 ogy and probable impacts on floodplain forests. Journal of Hydrology, 349, 397-410. doi: 782 https://doi.org/10.1016/j.jhydrol.2007.11.012 783

784

795

796

797

- A semi-distributed ntegrated itrogen model for multiple source assessment in tchments (INCA): Part I — model structure and process equations. (1998). Science of The Total 785 Environment, 210-211, 547-558. doi: 10.1016/S0048-9697(98)00037-0 786
- Shi, X., Mao, J., Thornton, P. E., Hoffman, F. M., & Post, W. M. (2011). The impact of 787 climate, co2, nitrogen deposition and land use change on simulated contemporary global 788 river flow. Geophysical Research Letters, 38. (https://doi.org/10.1029/2011GL046773) 789 doi: https://doi.org/10.1029/2011GL046773 790
- Sperotto, A., Molina, J. L., Torresan, S., Critto, A., Pulido-Velazquez, M., & Marcomini, 791 A. (2019). A bayesian networks approach for the assessment of climate change impacts 792 on nutrients loading. Environmental Science & Policy, 100, 21-36. doi: https://doi.org/ 793 10.1016/j.envsci.2019.06.004 794
 - Streeter, H., & Phelps, E. (1925). A study of the pollution and natural purification of the ohio river. iii. factors concerned in the phenomena of oxidation and reaeration. U.S. Government Printing Office.
- Suddick, E. C., Whitney, P., Townsend, A. R., & Davidson, E. A. (2013). The role of nitrogen 798 in climate change and the impacts of nitrogen-climate interactions in the united states: 799 foreword to thematic issue. Biogeochemistry, 114, 1-10. doi: 10.1007/s10533-012-9795-z 800
- Wexler, E. J. (1989). Analytical solutions for one-, two-, and three-dimensional solute 801
- transport in ground-water systems with uniform flow (Tech. Rep.). doi: 10.3133/ofr8956 802

- Wexler, E. J. (1992). Analytical solutions for one-, two-, and three-dimensional solute transport in ground-water systems with uniform flow. US Government Printing Office.
- Whitehead, P. G., Wilby, R. L., Battarbee, R. W., Kernan, M., & Wade, A. J. (2009). A review of the potential impacts of climate change on surface water quality. *Hydrological Sciences Journal*, 54, 101-123. (doi: 10.1623/hysj.54.1.101) doi: 10.1623/hysj.54.1.101
- Wilson, H., Elliott, J., Macrae, M., & Glenn, A. (2019). Near-surface soils as a source of phosphorus in snowmelt runoff from cropland. *Journal of Environmental Quality*, 48, 921-930. doi: 10.2134/jeq2019.04.0155
- Wilson, H. F., Casson, N. J., Glenn, A. J., Badiou, P., & Boychuk, L. (2019). Landscape controls on nutrient export during snowmelt and an extreme rainfall runoff event
 in northern agricultural watersheds. *Journal of Environmental Quality*, 48, 841-849. doi:
 10.2134/jeq2018.07.0278
- Wood, E. F., Roundy, J. K., Troy, T. J., van Beek, L. P. H., Bierkens, M. F. P., Blyth, E.,
 Whitehead, P. (2011). Hyperresolution global land surface modeling: Meeting a grand challenge for monitoring earth's terrestrial water. *Water Resources Research*, 47(5). doi: https://doi.org/10.1029/2010WR010090
- Wu, J., & Malmström, M. E. (2015). Nutrient loadings from urban catchments under climate
 change scenarios: Case studies in stockholm, sweden. Science of The Total Environment,
 518-519, 393-406. doi: https://doi.org/10.1016/j.scitotenv.2015.02.041
- Wu, Y., Liu, S., & Gallant, A. L. (2012). Predicting impacts of increased co2 and climate change on the water cycle and water quality in the semiarid james river basin of the midwestern usa. *Science of The Total Environment*, 430, 150-160. doi: https://doi.org/ 10.1016/j.scitotenv.2012.04.058
- Yang, Q., Zhang, X., Almendinger, J. E., Huang, M., Chen, X., Leng, G., ... Li, X. (2019). Climate change will pose challenges to water quality management in the st. croix river basin. *Environmental Pollution*, 251, 302-311. doi: https://doi.org/10.1016/ j.envpol.2019.04.129
- Ye, L., & Grimm, N. B. (2013). Modelling potential impacts of climate change on water
 and nitrate export from a mid-sized, semiarid watershed in the us southwest. *Climatic Change*, 120, 419-431. doi: 10.1007/s10584-013-0827-z
- Zheng, W., Wang, S., Tan, K., & Lei, Y. (2020). Nitrate accumulation and leaching
 potential is controlled by land-use and extreme precipitation in a headwater catchment
 in the north china plain. Science of The Total Environment, 707, 136168. doi: https://
 doi.org/10.1016/j.scitotenv.2019.136168

A multi-chemistry modelling framework to enable flexible and reproducible water quality simulations in existing hydro-models: 2. The OpenWQ-SUMMA and OpenWQ-CRHM model implementations and testing

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11 Abstract

This work advances the cross-model deployment of ecological and biogeochemical sim-12 ulation capabilities in existing process-based hydro-modeling tools, which we term "Open 13 Water Quality" (OpenWQ). The companion paper details aspects of the OpenWQ architec-14 ture that enables its plug-in type incorporation into existing models, along with its innova-15 tive aspects that enable biogeochemistry lab-like capabilities. OpenWQ's innovative aspects 16 allow modelers to define the pollution problem(s) of interest, the appropriate complexity of 17 the biogeochemistry routines, test different modeling hypotheses, and deploy them across 18 19 different hydro-models. In this second paper, we implemented the coupling recipe described in the first paper to integrate OpenWQ into two hydro-models, SUMMA and CRHM. Here 20 we explain how the implemented coupling interface between the two models provides water 21 quality simulation capacities in the host hydro-models but, more importantly, establishes a 22 direct and permanent link for the transfer of innovation between the modeling communities. 23 Example applications of different pollution studies enabled by our coupling recipe are also 24 provided to address some of these fundamental water quality modeling challenges. 25

²⁶ 1 Introduction

Climate change is expected to affect many hydro-biogeochemical processes along the 27 land-river-lake continuum (Costa et al., 2022). However, the specific local impacts of climate 28 change on hydrology and water quality will vary regionally and seasonally. On the one hand, 29 extreme winter river flows arising from increased extreme precipitation events are expected 30 to occur with greater frequency worldwide (Alexander et al., 2020; Field & Barros, 2014; 31 Luo et al., 2020; Onishi et al., 2020; "Effects of lake water level fluctuation due to drought 32 and extreme winter precipitation on mixing and water quality of an alpine lake, Case Study: 33 Lake Arrowhead, California", 2020; Shi et al., 2011; Suddick et al., 2013; H. Wilson et al., 34 2019; Yang et al., 2019; Zheng et al., 2020). On the other hand, summer river flows are 35 projected to decrease across many regions around the world, including North America (Rood 36 et al., 2008), China (Shi et al., 2011; Ervinia et al., 2020; Zheng et al., 2020), and western 37 Europe (Charlton et al., 2018; Raimonet et al., 2018). In arid climates, the flow reduction 38 is predicted to be particularly strong (Jarsjö et al., 2017; Shi et al., 2011; Y. Wu et al., 39 2012; Ye & Grimm, 2013; Bussi et al., 2021), reducing dilution of pollution point sources, 40 including wastewater outlets (Bussi et al., 2017; Charlton et al., 2018; Raimonet et al., 2018; 41 Ockenden et al., 2016; Whitehead et al., 2009). 42

The increase in climate variability is projected to lead to an overall increase in precipi-43 tation and snowmelt runoff, which, combined with changes in intensified cropping patterns, 44 is anticipated to lead to greater freshwater nutrient pollution in temperate regions of North 45 America, Europe and Japan (Alexander et al., 2020; Jeppesen et al., 2011; Onishi et al., 46 2020; H. F. Wilson et al., 2019). Increased nutrient uptake rates from terrestrial plants 47 due to climate change are not expected to facilitate the increase in nutrient runoff losses 48 (Luo et al., 2020; Sperotto et al., 2019; Suddick et al., 2013; Zheng et al., 2020). Erosivity 49 from agricultural and urban nonpoint sources is also projected to increase with more intense 50 runoff events because of their greater capacity to mobilize particulate and soluble N and P 51 (El-Khoury et al., 2015; Ockenden et al., 2016, 2017; J. Wu & Malmström, 2015; Crossman 52 et al., 2014). 53

Understanding the key factors controlling the fate of pollution in aquatic ecosystems is 54 challenging because it requires capturing multiple climate, hydrological, and biogeochemical 55 patterns across different spatial scales in models (Costa et al., 2022). Whilst there have 56 been numerous modelling advances over the last decade, process representation, model inter-57 comparability, model configuration (input data), and model validation remain key challenges 58 in water quality. Another considerable difficulty with the application of water quality models 59 is the often-limited monitoring data available for validation, which can compromise the 60 confidence of scientists and decision-makers in their simulation results. 61

This two-part paper presents a methodology termed "OpenWQ" that tackles some of 62 these fundamental modeling challenges by streamlining the transfer of knowledge, technol-63 ogy, and innovation between the different scientific and modeling communities related to 64 water quality. Some of these scientific and modelling communities include biogeochemistry, 65 soil science, hydrology, hydrodynamics, and hydrogeology. This work aims to provide a con-66 crete, transformative framework for enhancing collaborative efforts and co-creation that can 67 benefit the entire environmental and ecological research and management community. Our 68 intent for the first paper was to advance OpenWQ's concept as a coupler and introduce its 69 customizable biogeochemistry modeling framework. This customizable framework enables 70 the cross-model deployment of eco-lab water quality modeling capabilities. This second 71 paper builds on the motivation for OpenWQ introduced in the first paper and outlines the 72 initial implementation of OpenWQ as a modeling system. The goals of this second paper 73 are to 74

Demonstrate how OpenWQ has been integrated into two hydro-models using the coupling procedures laid out in the first paper. The hydro-models have been selected for their likewise flexible modeling frameworks, namely the Structure for Unifying Multiple Modeling Alternatives (SUMMA, Clark et al., 2015a,b) and the Cold Regions Hydrological Model and its water quality modules (CRHM/CRHM-WQ, J. W. Pomeroy et al., 2007; Costa et al., 2021).

and SUMMA to examine a variety of common pollution problems.

2. Apply the multi-chemistry, eco-lab framework of OpenWQ incorporated with CRHM

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⁸³ 2 Coupling to hydro-models

OpenWQ has been coupled to two hydro-models, SUMMA (Clark et al., 2015a,b) and CRHM/CRHM-WQ (J. Pomeroy et al., 2022; Costa et al., 2021; J. W. Pomeroy et al., 2007). In companion paper 1 (*CITE_paper_1*), we explain how OpenWQ has been designed as a flexible coupler to streamline its integration in existing hydro-models. A coupling recipe has been developed to lay out the steps involved in the coupling process, which includes linking four coupling functions. Here, in paper 2, we explain how this coupling recipe was deployed in two separate coupling efforts and used to examine different pollution problems.

SUMMA SUMMA is written in Fortran, while OpenWQ is written in C++. Thus, it 91 was necessary to create wrapper functions that allow for the interaction between a Fortran 92 driver program and a C++ coupled program at runtime. Because it is OpenWQ that is 93 integrated into SUMMA, the coupled model was constructed in a manner to allow SUMMA 94 to manage the simulation timesteps. For this, we used the standard intrinsic Fortran module 95 *iso_c_binding* that defines constants, types, and procedures for C interoperability. The 96 *iso_c_binding* module is part of the language standard and makes the interface between 97 Fortran and C standard and portable. Essentially, Fortran-to-C wrapper functions were 98 created to allow SUMMA to interface with OpenWQ and, from the point of view of SUMMA, aq OpenWQ is an object to which it held a pointer upon initialization of OpenWQ. The wrapper 100 functions allowed SUMMA to interface with OpenWQ's coupler functions as if they were 101 method calls. Fortran-to-C wrapper functions were constructed for creating a reference to 102 a OpenWQ along with each of the four OpenWQ C++ coupler functions (openwq::decl, 103 openwq::run_time_start, openwq::run_space, and openwq::run_time_end; see Section 3.3.2 of 104 the first paper (CITE_paper_1). 105

CRHM Similar to OpenWQ, CRHM is written in C++. This avoids the need for
 wrapper interface functions as required in the SUMMA-OpenWQ coupling. CRHM already
 had a water quality module named CRHM-WQ that enables the simulation of the nitrogen
 (N) and phosphorous (P) cycles (Costa et al., 2021). Thus, CHRM already contains a so lute transport simulation routine that moves contaminants through the hydrological cycle
 represented by the model. The native implementation of solute transport in CRHM-WQ is

limited to 6 chemical species related to the N and P cycles, namely nitrate (NO_3-N) , ammonium (NH_4-N) , dissolved organic nitrogen (DON), soluble reactive phosphorous (SRP), and particulate phosphorous (PP). Therefore, instead of coupling all four OpenWQ coupler functions to CRHM-WQ, we considered it more appropriate to use CRHM's native transport routine, which has already been tested, and adjust it for a more dynamic allocation of the number of chemical species when OpenWQ is activated, leaving OpenWQ's coupling routine for transport (*openwq::run_space*) out of the coupling process.

¹¹⁹ 3 Case studies

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3.1 Overal science questions

¹²¹ OpenWQ enables exploring multiple working hypotheses related to multi-chemistry ¹²² biogeochemical cycling in different seasons, climate zones, and landscapes. The work aims ¹²³ to respond to critical challenges with water quality modeling as an interdisciplinary science, ¹²⁴ which have been highlighted in the companion paper (*CITE_paper_1*). The specific modeling ¹²⁵ challenges addressed are to:

1. Identify appropriate biogeochemical frameworks: what are the key biogeochemical processes and chemical species that should be represented in models to capture the fate and controls of pollution in a way that is (1) adequate for the quantity and quality of the available data for model evaluation, (2) scientifically sound from a biogeochemical perspective, and (3) useful for the objectives of the study and helpful for decision making?

- 2. Differentiate controlling biogeochemical transformations across hydrological compart-132 ments (e.g., snow, groundwater, etc.): should biogeochemical processes be differen-133 tiated across hydrological compartments and regions, e.g., although the processes 134 involved in the nitrogen cycle are the same regardless of the environment they occur 135 in (e.g., nitrification, denitrification, assimilation, ammonification, nitrogen fixation), 136 certain environmental conditions tend to favor particular processes (e.g., oxygen lev-137 els in groundwater tend to be low, leading to reduced nitrification rates because this 138 is a process involving dissolved oxygen)? 139
- Characterize model uncertainty: can we better understand the sources of uncertainty related to key modeling processes and methods used so that we can incrementally improve our models, increase predictability and model fidelity, and reduce parameter equifinality?

Some of these questions have also been raised by the hydrological community (Clark et al., 2015b; Reggiani et al., 1998; Beven, 2006; Renard et al., 2010; Wood et al., 2011), but the ones listed above extend to aspects specific to water quality modeling. The general concept of OpenWQ is to facilitate a test-improve workflow so that models can be flexibly and incrementally improved based on model results and inputs from different disciplines.

¹⁴⁹ **3.2 Model tests**

The model was evaluated using a series of tests designed to expose the accuracy and 150 robustness of the mathematical formulation and numerical solutions deployed for examining 151 a variety of pollution problems. In recent years, it has been recognized that such basic 152 tests are an essential step for proper model scrutiny, falsification, and acceptance (Clark et 153 al., 2021a). In this study, we pioneer the extension of such important principles for water 154 quality simulations where the spatiotemporal complexity and interplay of process generally 155 often give rise to highly non-linear response patterns that are difficult to disentangle in 156 real-world applications for proper model examination (Costa et al., 2022, 2020). The case 157 study tests devised are representative of a variety of environmental pollution problems for 158

Transport Scheme	Designation	Description			
No transport	1_noTrans_1species_1storder 2_noTrans_1species_2ndorder 3_noTrans_2species 4_noTrans_3species 5_noTrans_nitrogencycle 6_noTrans_oxygenBODcycle	Batch reactor involving 1 chemical species subject to first-order decay Batch reactor involving 1 chemical species subject to second-order decay Batch reactor involving 2 chemical species subject to first-order decay Batch reactor involving 3 chemical species subject to first-order decay Batch reactor representing the nitrogen cycle Batch reactor representing the dissolved oxygen cycle			
Advection- Dispersion	7_trans_contS_PorMedia 8_trans_contS_PorMedia_1storder	Non-reactive transport of 1 chemical species in porous media subject to a continuous load Reactive (linear decay) transport of 1 chemical species in porous media subject to a continuous load			
	9_trans_instS_PorMedia 10_trans_instS_PorMedia_1storder	Non-reactive transport of 1 chemical species in porous media subject to an instantaneous load Reactive (linear decay) transport of 1 chemical species in porous media subject to an instantaneous load			

Table 1: Case study tests used to test OpenWQ and its coupling to SUMMA and CRHM

which analytical solutions have been either specially derived for this paper or taken from the literature.

We divide the tests into reactive and non-reactive. This is an important distinction 161 because basic mass balance principles cannot be directly applied to reactive chemical con-162 stituents without proper consideration because they are subject to biogeochemical transfor-163 mation in addition to physical transport. We prepared a total of ten tests, of which two are 164 non-reactive and eight are reactive. Each of these tests has well-defined initial conditions, 165 boundary conditions, sink and source terms, and is subject to specific biogeochemistry as 166 described below. Based on these conditions, analytical solutions have been derived for com-167 parison with the model results. In order to benefit the comparison of model results, the 168 simulations are divided into two groups as shown in Table 1. 169

The case studies include six batch reactor-type tests, specifically two single-species 170 tests and four multi-species tests. The single-species tests are the simplest and represent 171 cases involving first- (Test 1) and second-order (Test 2) linear decay processes. The multi-172 species tests include (a) a two-species reaction network subject to linear decay (Test 3), (b) 173 a three-species reaction network subject to linear decay (Test 4), and reaction networks to 174 represent the (c) nitrogen (Test 5) and (d) dissolved oxygen (Test 6) cycles. The remaining 175 four tests involve the transport of chemical constituents, specifically two non-reactive tests, 176 one considering a continuous point source (Test 7), one considering an instantaneous points 177 source (Test 8), and two reactive tests that are similar to the homologous non-reactive 178 tests, but where chemical species undergo biogeochemical transformations (Tests 9 and 10, 179 respectively). 180

3.3 Multi-chemistry lab set up

The case studies involve different pollution problems. In some cases, they are generic 182 in the sense that they are representative of types of biogeochemical processes and could be 183 applied to examine a variety of water quality problems. In other cases, the tests are more 184 specific, focusing on environmental issues involving specific biogeochemical cycles, such as 185 the nitrogen and phosphorous cycles. As mentioned in the first paper, the biogeochemistry 186 calculation engine is flexible and easily extendible, supporting numerous forms of functional, 187 logical, and vector processing semantics. It is enabled by the comprehensive C++ Mathe-188 matical Expression Toolkit Library (*ExprTk*) developed by Arash Partow (Partow, 1999). 189 The implementation of ExprTk in OpenWQ is simple to use. It provides an efficient run-190

time mathematical expression parser and evaluation engine that is ideal for testing different modeling hypotheses to inform incremental improvements to the model setup.

In order to perform the various pollution case study tests described in Table 1, a series 193 of OpenWQ input files have been prepared. The biogeochemistry JSON files (see Section 194 ?? of the first paper) are where the multi-chemistry lab simulations are characterized. We 195 defined one biogeochemistry file per case study test, which we deployed in both SUMMA-196 OpenWQ and CRHM-OpenWQ coupled models. From paper 1, one of the transformative 197 innovations of OpenWQ is the ability for cross-model deployment of eco-lab co-creation 198 modeling capabilities using the same OpenWQ biogeochemistry JSON files. This allows the 199 implementation of multiple hydro-biogeochemical modeling hypotheses by simply loading 200 the desired JSON biogeochemistry file. The input files prepared for the different case study 201 tests are publicly available on GitHub. They can be used as the basis for validation of any 202 new coupling of OpenWQ to a hydro-model. 203

3.4 Deriving analytical solutions

In this section we provide details about the derivation of the analytical solutions used to verifying the performence of OpenWQ when coupled to SUMMA and CRHM. Each test is presented in a separate section, which starts with a diagram of relevent reaction network. The diagrams are then used to define the governing equations, which can be ordinary or partial differential equations. Finally, the derived analytical solution is presented. Details about the derivations is provided in the Appendix A-G.

3.4.1 Batch reactor: Single-species (Tests 1 and 2)

The single-species batch reactor tests involve (a) first-order and (b) second-order decay kinetics. Fig. 1 depicts the biogeochemical process.



Figure 1: Single-species reaction network where transf_A is a first-order biogeochemical decay reaction.

The analytical solution of a single species subject to first-order decay can be obtained by integrating Eq. 24 with initial condition at time t = 0 to obtain Eq. 25. In turn, the analytical solution of a single species subject to second-order decay can be obtained by integrating Eq. 26 with initial condition at time t = 0 to obtain Eq. 27. See Appendix A for more details.

$$\frac{dc_A}{dt} = -kc_A \tag{1}$$

$$c_A(t) = c_A(0)e^{-kt} \tag{2}$$

$$\frac{dc_A}{dt} = -kc_A^2 \tag{3}$$

$$c_A(t) = \frac{1}{\frac{1}{c_A(0)} + kt}$$
(4)

3.4.2 Batch reactor: Two-species reaction network (Test 3)

219

The two-species reaction network with linear decay is represented schematically in Fig. 2.



Figure 2: Two-species reaction network where transf_A and transf_B are first-order biogeochemical transformations.

Similar to the single-species case with first-order decay (Eq. 25), the fate of Species A
in a two-species reaction network only depends on its own concentration and is given by Eq.
28.

$$c_A(t) = c_A(0)e^{-k_A t} \tag{5}$$

However, the fate of Species B depends on the time-evolution of concentrations of bothSpecies A and B. This evolution can be expressed as:

$$\frac{dc_B}{dt} = k_A c_A - k_B c_B \tag{6}$$

The analytical solution of Eq. 29 with initial condition at time t = 0 can be obtained by using the integrating factor, $e^{k_B t}$. Details on this derivation are provided in the Appendix B. The solution is provided in Eq. 38.

$$c_B(t) = c_B(0)e^{-k_B t} + \frac{k_A c_A(0)}{k_B - k_A} \left\{ e^{-k_A t} - e^{-k_B t} \right\}$$
(7)

230 3.4.3 Batch reactor: Three-species reaction network (Test 4)

A three-species reaction network with linear decay is represented schematically in Fig. 2.

Concentrations of Species A and B are given by respectively Eqs. 28 and 38. The differential equation for the concentration of Species C can be written as:

$$\frac{dc_C}{dt} = k_B c_B - k_C c_C \tag{8}$$



Figure 3: Three-species reaction network where transf_A, transf_B, and transf_C are first-order biogeochemical transformations.

234 235 The analytical solution Eq 40 can again be obtained by using the integrating factor $e^{k_C t}$. Details on this derivation are provided in the Appendix. The final solution is

$$c_{C}(t) = c_{C}(0)e^{-k_{C}t} - \frac{k_{B}}{k_{B} - k_{C}}c_{B}(0)\left(e^{-k_{B}t} - e^{-k_{C}t}\right) - \frac{k_{B}k_{A}}{(k_{B} - k_{A})(k_{C} - k_{A})}c_{A}(0)\left\{\left(e^{-k_{B}t} - e^{-k_{A}t}\right) - \frac{e^{-k_{C}t}}{k_{C} - k_{B}}\right\}.$$
(9)

²³⁶ 3.4.4 Batch reactor: Nitrogen cycle (Test 5)

The simulation of the nitrogen cycle is common in freshwater water quality studies because it allows for addressing issues with nutrient pollution and eutrophication. The nitrogen cycle simulated involves different nitrogen species (Fig. 4). The conceptual cycling framework used is similar to that used in catchment-scale models like HYPE Lindström et al. (2010), CRHM-WQ (Costa et al., 2021) models, SWAT (Arnold et al., 2012), and INCA ("A semi-distributed ntegrated itrogen model for multiple source assessment in tchments (INCA): Part I — model structure and process equations", 1998).



Figure 4: Nitrogen cycle as represented in several popular catchment-scale models. The nitrogen species Nref refers to refractory nitrogen, Nlab refers to labile nitrogen, DON refers to dissolved organic nitrogen, and DIN refers to dissolved innorganic nitrogen. The transformation degr refers to degration; diss1 and diss2 refers to dissolution from Nref and Nlab, respectively; min refers to mineralization; denitr refers to denitrification; and plantup refers to plant uptake.

The analytical solutions for the different nitrogen species can be obtained from Eq. 28, Eq. 38, or Eq. 47, depending on their positioning in the reaction chain:

Refractory N (Nref) Similar to Eq. 28, which for this specific processes and N species reads as

$$c_{Nref}(t) = c_{Nref}(0)e^{-(k_{degr} + k_{diss1})t}$$
 (10)

Labile N (Nref) Similar to Eq. 38, which for this specific processes and N species reads as

$$c_{Nlab}(t) = c_{Nlab}(0)e^{-(k_{min} + k_{diss2})t} + \frac{k_{degr}c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \left\{ e^{-k_{degr}t} - e^{-(k_{min} + k_{diss2})t} \right\}.$$
(11)

$$\frac{dc_{DON}}{dt} = k_{diss1}c_{Nref} + k_{diss2}c_{Nlab} \tag{12}$$

Due to similarities in the relative positioning of the chemical species in the reaction network, analytical solutions suitable for c_{Nref} and c_{Nlab} are given by Eqs. 28 and 38, respectively. The derivation of Eq. 50 is provided in Appendix D, and its solution reads

$$c_{DON}(t) = c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}} - \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}}e^{-(k_{degr} + k_{diss1})t} - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}}e^{-(k_{min} + k_{diss2})t} + \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2}} \left\{ \frac{e^{(k_{min} + k_{diss2})t} - 1}{k_{min} + k_{diss2}} - \frac{e^{-k_{degr}t}}{k_{degr} - 1} \right\}.$$
(13)

Dissolved Inorganic N (DIN) The ordinary differential equation for DIN can be
 written as

$$\frac{dc_{DIN}}{dt} = k_{min}c_{Nlab} - (k_{denit} + k_{plantup})c_{DIN}.$$
(14)

The analytical form for the concentration of $c_{Nlab}(t)$ is given in Eq. 49 and can be substituted in Eq. 56 to initiate the derivation that is provided in the Appendix. The analytical solution to Eq. 56 reads

$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]} = c_{DIN}(0)$$

$$+k_{min}c_{Nlab}(0)\left\{\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}-1}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right\}$$

$$+\frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\left\{\frac{e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]}-1}{k_{denitr}+k_{plantup}-k_{degr}}\cdot-\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}-1}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right\}$$

$$(15)$$

3.4.5 Batch reactor: Dissolved oxygen cycle (Test 6)

257

Dissolved oxygen (DO) depletion in rivers and lakes is a common consequence of pollution. It typically occurs near discharge outlets of untreated wastewater and can lead to high levels of organic matter that, during decomposition by aerobic bacteria, causes a DOsag in the water column. For the purposes of DO dynamics, such organic matter is typically translated into units of biological oxygen demand (BOD). Fig. 5 depicts the processes involved, and the Streeter–Phelps equation (Streeter & Phelps, 1925) has long been used for such problems. It reads

$$\frac{dD}{dt} = k_1 L_t - k_2 D,\tag{16}$$

where D is the saturation deficit in g/m^3 , which can be derived from the dissolved oxygen concentration at saturation minus the actual dissolved oxygen concentration $(D = DO_{sat} - DO)$; k_1 is the deoxygenation rate, usually in d^{-1} ; $L_t = L_a e^{-k_1 t}$; L_a is the initial oxygen demand of organic matter in the water, also called the ultimate BOD (BOD at time $t = \infty$); and k_2 is the reaeration rate, usually in d^{-1} . The unit of L_a is g/m^3 ; k_1 ; L_t is the oxygen demand remaining at time t, and t is the elapsed time, usually in days.



Figure 5: Representation of the transformations involved in the BOD - DO cycle

The equation states that the total rate of change in oxygen deficit (D) is equal to the difference between the two rates of deoxygenation and reaeration at any time. See more details in Appendix E. The solution reads

$$D = \frac{k_1 L_a}{k_2 - k_1} \left\{ e^{-k_1 t} - e^{-k_2 t} \right\} + D_a e^{-k_2 t}.$$
(17)

where D_a is the initial oxygen deficit $[g/m^3]$.

268 3.4.6 Reactive and non-reactive transport: Instantaneous point source 269 (Tests 7 and 8)

These tests were only performed with SUMMA-OpenWQ because, as mentioned before, while OpenWQ's biogeochemical simulation modules have been integrated in both SUMMA-OpenWQ and CRHM-OpenWQ, OpenWQ's solute transport simulation modules were only

integrated in SUMMA-OpenWQ, but not in CRHM-OpenWQ, since CRHM has it's own solute transport routines (Costa et al., 2021). The general 1D diffusion-advection-reaction partial differential equation is given by

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c, \qquad (18)$$

where c is the concentration of the tracer, t is time, D_x is the diffusion coefficient in the flow direction, x is the distance in the flow direction, ν is the flow velocity in the x-direction, and λ is a linear reaction term that can be set to zero for non-reactive tracers. Various analytical solutions for Eq. 18 have been derived and compiled by Wexler (1992) for different idealized scenarios. The modified equation for an instantaneous point source reads

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c + \frac{Q}{n} dt c_0 \delta \left\{ x - x_c \right\} \delta \left\{ t - t'_c \right\},\tag{19}$$

where δ is the Dirac delta (impulse) function, x_c is the location of the instantaneous point source, t'_c is the time at which the instantaneous point source is activated, Q is the fluid injection rate, and n is the aquifer porosity. The analytical solution to Eq. 67 has been derived by Bear (1979) considering the following boundary conditions and initial conditions.

- 2791. Boundary conditions (Neumann or second type):280 $c, \nabla c = 0, \quad x = \pm \infty$ 2812. Initial conditions:282 $c(x,t=0) = 0, \quad -\infty < x < \infty.$
- The analytical solution of Eq. 67 is given by

$$c(x) = \frac{c_0 Q}{4n\pi\gamma D_x} \exp\left[\frac{\nu(x-x_0)}{2D_x}\right]$$

$$\int_0^t \frac{1}{t} \exp\left[-\left(\frac{\nu^2}{4D_x}\right)\tau - \frac{(x-x_c)^2}{4D_x\tau}\right].$$
(20)

and can be used for solutes that are non-reactive (Test 7) or reactive (Test 8) by setting $\lambda = 0$ in Test 7.

The equation assumes that (1) the fluid has constant density and viscosity, (2) the 286 solute may be subject to first-order chemical transformation, (3) the flow is uniform in the 287 x-direction only, (4) the flow velocity is constant, and (5) the longitudinal diffusion coefficient 288 (D_x) is constant. The solution has a mathematical singularity in the point source region 289 (Abdelaziz et al., 2013) because it takes the form of the exponential integral when both 290 $x - x_c$ approaches zero, hence becoming infinite at $\tau = 0$. However, according to Wexler 291 (1989) the solution is still valid as long as $(x - x_c)^2$ is larger than ν^2 , as it is the case in our 292 293 tests.

294 295

3.4.7 Reactive and non-reactive transport: Continuous point source (Tests 9 and 10)

Similarly to Tests 7 and 8, the tests here have only been performed for SUMMA-OpenWQ due to the same reasons highlighted above in Section 3.4.6. The analytical solution for the transport of solutes with a continuous point source has been derived in Wexler (1992) after modification from Bear (1979) and Genuchten & Alves (1982).

			Biogeochemistry			
Transport Scheme	Designation	Reference	Initial Conditions	Reaction kinetics	Transport	
	1_noTrans_1species_1storder	Fig. 1	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.01 \ ^{1}/day$	NA	
Patah Dasatan	2_noTrans_1species_2ndorder	Fig. 1	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.01^{1/day}$	NA	
	3_noTrans_2species	Fig. 2	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.03 \ ^{1}/day$	NA	
Datch Reactor			Species B ($c_0 = 0 \text{ mg/l}$)	$k_{transf_B} = 0.01 \ ^{1}/day$		
	4_noTrans_3species	Fig. 3	Species A $(c_0 = 10 \text{ mg/l})$	$k_{transf_A} = 0.03 \ 1/day$	NA	
			Species B $(c_0 = 0 \text{ mg/l})$	$k_{transf_B} = 0.01 \ \frac{1}{day}$		
			Species C ($c_0 = 0 \text{ mg/l}$)	$k_{transf_{-C}} = 0.005 \ \frac{1}{day}$		
	5_noTrans_nitrogencycle	Fig. 4	Nref ($c_0 = 10 \text{ mg/l}$)	$k_{degrd} = 0.006 \ \frac{1}{day}$	NA	
			Nlab $(c_0 = 10 \text{ mg/l})$	$k_{dissol-1} = 0.0002 \ \frac{1}{day}$		
			$DIN (c_0 = 5 \text{ mg/l})$	$k_{dissol-2} = 0.0003 \frac{1}{day}$		
			DON $(c_0 = 2 \text{ mg/l})$	$k_{miner} = 0.003 \frac{1}{day}$		
				$k_{denitr} = 0.001 \frac{1}{day}$		
	a T DOD I	D: -		$k_{plantup} = 0.001 \frac{1}{day}$	27.4	
	6_noTrans_oxygenBODcycle	Fig. 5	BOD $(c_0 = 10 \text{ mg/l})$	$k_{BODdegrad} = 0.1 \frac{1}{day}$	NA	
		<u> </u>	$DO(c_0 = 12 \text{ mg/l})$	$\kappa_{reaertion} = 0.3 \frac{1}{day}$		
Advection-Dispersio	7_trans_contS_PorMedia	NA	Species A $(c_0 = 0 \text{ mg/l})$	NA	$D_z = 0.0001 \ m^2/s$	
	8_trans_contS_PorMedia_1storder	Fig. 1	Species A $(c_0 = 0 \text{ mg/l})$	$k_{transf_A} = 0.01 \ ^{1/day}$	$D_z = 0.0001 \ m^2/s$	
	9_trans_instS_PorMedia	NA	Species A $(c_0 = 0 \text{ mg/l})$	NA	$D_z = 0.0001 \ m^2/s$	
	10_trans_instS_PorMedia_1storder	Fig. 1	Species A $(c_0 = 0 \text{ mg/l})$	$k_{transf_A} = 0.01 \ ^{1}/day$	$D_z = 0.0001 \ m^2/s$	

Table 2: Model configuration for the different test cases

$$c(x) = \frac{c_0}{2} \left\{ \exp\left[\frac{x}{2D_x}(\nu - U)\right] \operatorname{erfc}\left[\frac{x - Ut}{2\sqrt{D_x t}}\right] + \exp\left[\frac{x}{2D_x}(\nu + U)\right] \operatorname{erfc}\left[\frac{x + Ut}{2\sqrt{D_x t}}\right] \right\},$$
(21)

where c_0 is the concentration of the continuous point source, $U = \sqrt{\nu^2 + 4\lambda D}$ and erfc is the complementary error function. Parameter λ is the first-order reaction rate used for Test 10, which can be set to zero for Test 9. See more details in Appendix G.

303

3.5 Model configuration and input files

The configuration of OpenWQ for the different test cases is summarized in Table 2. 304 The analytical solutions described in Section 3.4 have been set up using the same initial and 305 boundary conditions. The configuration of OpenWQ involves the characterization of biogeo-306 chemistry (initial conditions and reaction kinetics) and transport processes. The OpenWQ 307 input files set up to translate the model configurations described in Table 2 are publicly 308 available at https://github.com/ue-hydro/synthetic_tests. These input files can be 309 used with any hydro-model coupled to OpenWQ and constitute an essential verification 310 step in any new coupling project. 311

This public repository also contains the input files of the host hydro-models CRHM 312 and SUMMA prepared for each model test, which have, in essence, been set up to provide 313 no-flow conditions for Tests 1 to 6 (batch-reactor simulations) and steady-state conditions 314 for Tests 7 to 10 (reactive and non-reactive transport simulations). Setting up proper flow 315 conditions was a critical step to ensure the validity of the analytical solutions used to verify 316 the numerical results. In the case of Tests 7 to 10 (transport simulations), which were only 317 performed for SUMMA-OpenWQ as explained before, were based on the soil unsaturated 318 flow test cases formulated by Celia et al. (1990), which have been set up in SUMMA as 319 described in Clark et al. (2021b). The soil properties of these tests are described in Table 2 320 of that paper. 321

322 4 Results

The model results have been compared to the respective analytical solutions to verify the accuracy of OpenWQ in examining a variety of pollution problems. These tests contemplate several biogeochemical processes that are often the object of study in environmental studies.

The Nash–Sutcliffe efficiency (NSE) coefficient and the Root-Mean-Square Error (RMSE)

³²⁷ were used to quantify the accuracy of the numerical results.

$$NSE = 1 - \frac{\sum (X_{obs} - X_{mod})^2}{\sum (X_{obs} - \mu_{obs})^2}$$
(22)

$$RMSE = \sqrt{\frac{\sum (X_{obs} - X_{mod})^2}{n}}$$
(23)

where X_{obs} and X_{mod} are the observed and simulated concentrations, μ_{obs} is the concentration average, and n is the number of observations. NSE values equal to one indicate a perfect match between observations and model results, and NSE values equal to zero indicate model predictions as accurate as the mean of all observations.

4.1 Batch reactor: Single-species (Tests 1 and 2)

The results of Tests 1 and 2 contemplate the general cases of single chemical species subject to first- or second-order decay as presented in Fig. 6. Such reaction types are commonly used in sequence or parallel to represent complex reaction networks.



(a) Test 1 (1_noTrans_1species_1storder)



Figure 6: Comparison between model results and analytical solution for Tests 1 and 2

The results show good model predictability and reduced numerical dispersion. For 336 Test 1, the Nash-Sutcliffe and RMSE values obtained for summa-openwog were respectively 337 0.999999998259312 and 8.804e-05 mg/l and for crhm-openwq were 0.99999998259309 and 338 8.804e-05 mg/l. For Test 2, the Nash-Sutcliffe and RMSE values obtained for summa-330 openwq were 0.999999900146240 and 3.0319e-04 mg/l, and for crhm-openwq were respec-340 tively 0.999999900145974 and 3.0319e-04 mg/l. The results obtained with summa-openwq 341 and crhm-openwq are similar, which was expected since these tests deal only with biogeo-342 chemical transformations and nutrient transport, which is driven by the host-models, is not 343 activated. The RMSE is one order of magnitude higher for Test 2, which was also antic-344 ipated given that the biogeochemical formulation in this test contains an exponent (i.e., 345 power of two), which magnifies the propagation of round-off errors. 346

4.2 Batch reactor: Two-species reaction network (Test 3)

The results of Test 3 are presented in Fig. 7 and compared against analytical solutions. Two-species reaction networks are also common in water quality models and often integrated within comprehensive biogeochemical cycling representations.



Figure 7: Comparison between model results and analytical solution for Test 3

The model shows good agreement with the analytical solution for Test 3, including 351 the results for both Species A and B. For Species A, the Nash-Sutcliffe and RMSE values 352 obtained for summa-openwq were 0.999999987172389 and 1.525e-04 mg/l, and for crhm-353 openwq were respectively 0.999999987172368 and 1.525e-04 mg/l. For Species B, the Nash-354 Sutcliffe and RMSE values obtained for summa-openwq were 0.999999989757722 and 1.747e-355 04 mg/l, and for crhm-openwq were respectively 0.999999989757706 and 1.747e-04 mg/l. 356 The RMSE calculated for Species B is slightly higher than that calculated for Species A. 357 This was anticipated since the reaction network sequence starts with Species A; thus, the 358 effect of changes in the concentration of Species A will take one time step to travel to Species 359 B. Reducing the length of the time step reduces the problem. 360

361

4.3 Batch reactor: Three-species reaction network (Test 4)

Fig. 8 shows the model results for Test 4, comparing the numerical solution with the corresponding analytical solution. This constitutes a more complex reaction chain that is often integrated into more complex biogeochemical cycling frameworks.

Similarly to the previous tests, the model demonstrates good accuracy and reduced 365 numerical dispersion. For Species A, the Nash-Sutcliffe and RMSE values obtained for 366 summa-openwq were respectively 0.999999987172389 and 1.524e-04 mg/l and for crhm-367 openwq were 0.999999987172368 and 1.524e-04 mg/l. For Species B, the Nash-Sutcliffe 368 and RMSE values obtained for summa-openwq were respectively 0.999999989757722 and 369 1.747e-04 mg/l and for crhm-openwq were 0.999999989757706 and 1.747e-04 mg/l. For 370 Species C, the Nash-Sutcliffe and RMSE values obtained for summa-openwq were respec-371 tively 0.999899420454185 and 1.560e-02 mg/l and for crhm-openwq were 0.999899420454600372 and 1.560e-02 mg/l. 373



Figure 8: Comparison between model results and analytical solution for Test 4

A small phase shift can be noticed for Species C, which is also caused by the reaction 374 network sequence issue mentioned in the previous test (Test 3). In OpenWQ, the hydro-375 models (CRHM and SUMMA) control the model time step. In these tests, the time step 376 used was 15 minutes and was dictated by the data time intervals. The larger the time step is, 377 the larger the shifts are expected to become since the computation of the reactions depends 378 on the concentrations in the antecedent time step. So, in this case, a change in Species A 379 has taken 30 min (2 times 15 minutes) to numerically travel to Species C, producing a slight 380 phase shift. Reducing the length of the time step reduces the problem. 381

4.4 Batch reactor: Nitrogen cycle (Test 5)

382

The results of the nitrogen cycle simulation show good agreement with the numerical solution (Fig. 9). This is a reaction network involving four nitrogen species and it is commonly used in popular models such as HYPE, INCA, and SWAT.

The model shows good agreement with the analytical solution, reinforcing the confi-386 dence in OpenWQ's numerical solution and the implementation of the expression evaluator 387 based on the comprehensive C++ Mathematical Expression Toolkit Library (*ExpTk*) de-388 veloped by Arash Partow (Partow, 1999). For Nref, the Nash-Sutcliffe and RMSE values 389 obtained for summa-openwq were respectively 0.99999999184531 and 6.927e-05 mg/l and 390 for crhm-openwq were 0.99999999184529 and 6.927e-05 mg/l. For Nlab, the Nash-Sutcliffe 391 and RMSE values obtained for summa-openwo were respectively 0.999136304357914 and 392 9.426e-02 mg/l and for crhm-openwq were 0.999136304357383 and 9.426e-02 mg/l. For 393 DON, the Nash-Sutcliffe and RMSE values obtained for summa-openwog were respectively 394 0.999463220618209 and 1.3089e-02 mg/l and for crhm-openwq were 0.999463220618758 and 395 1.309e-02 mg/l. For DIN, the Nash-Sutcliffe and RMSE values obtained for summa-396 openwq were respectively 0.996669866632908 and 7.181e-02 mg/l and for crhm-openwq were 397 0.996669866633266 and 7.181e-02 mg/l. As in the previous reaction networks tested, the 398 RMSE increases as the information travels through the network sequence: DIN ; Nref. 399 This is caused by the same issue raised above related to the time of travel of information 400 through the network sequence computation. 401



Figure 9: Comparison between model results and analytical solution for Tests 5

4.5 Batch reactor: Dissolved oxygen cycle (Test 6)

402

The simulation of the dissolved oxygen cycle is presented in Fig. 10 with two water quality variables (1) dissolved oxygen (DO) and (2) biological oxygen demand (BOD). This is also a common pollution problem simulated by water quality models to evaluate mitigation strategies for contamination caused by urban wastewater discharges in rivers and lakes.



Figure 10: Comparison between model results and analytical solution for Tests 6

The model shows good agreement with observations, capturing the typical DO 407 sag often observed near discharge outlets as predicted by the analytical solution. For 408 DO, the Nash-Sutcliffe and RMSE values obtained for summa-openwq were respectively 409 0.999998916342720 and 1.416e-04 mg/l and for crhm-openwq were 0.999998916342720 and 410 1.416e-04 mg/l. For BOD, the Nash-Sutcliffe and RMSE values obtained for summa-411 openwq were respectively 0.999999874118595 and 2.783e-04 mg/l and for crhm-openwq 412 were 0.999999874118595 and 2.783e-04 mg/l. As in the previous cases, the RMSE increases 413 through the reaction network sequence, which is higher for BOD, despite that the Nash-414 Sutcliffe value remains very high for both species. 415

416 417

4.6 Reactive and non-reactive transport: Continuous point source (Tests 7 and 8)

Fig. 11 compares model results against analytical solutions for Tests 7 and 8, which 418 represent non-reactive and reactive scenarios of transport of a continuous pollution source 419 through porous media. The results show the traveling of solute through the soil profile driven 420 by advection and dispersion processes. In the case of Test 8 in Fig 11b, there is an additional 421 first-order biogeochemical reaction causing a linear decay in the concentrations. Each figure 422 shows the computed concentrations, as well as the model internal derivatives that have 423 been passed into the numerical solver to compute the concentration update. This is useful 424 information to understand the key processes controlling the concentration of the solute as 425 time progresses, as well as the relative contributions of transport and biogeochemistry. 426

Comparing Figs 11a and 11b, it becomes clear that the superposition of the different 427 processes (e.g., biogeochemistry, physical transport) gives rise to a highly non-linear re-428 sponse that the model can capture accurately. For Test 7, the Nash-Sutcliffe and RMSE 429 values obtained were respectively 0.9890 and 0.0747e-05 mg/l, and for Test 8 were respec-430 tively 0.9953 and 0.0392 mg/l. Results show that the Nash-Sutcliffe values remain high, but 431 they are slightly lower when compared to the purely biogeochemical kinetic Tests 1-6. The 432 opposite is observed for RMSE, where values have slightly increased. This minor deterio-433 ration of the numerical results with the activation of the transport routines was expected 434 and it remains within acceptable values. This is caused by the initial conditions required to 435 translate in the host-model to represent those set up in the analytical solutions, as well as 436 the velocity fields computed by the host model and associated numerical dispersion in both 437 OpenWQ and the host-model. 438

439 440

4.7 Reactive and non-reactive transport: Instantaneous point source (Tests 9 and 10)

The non-reactive and reactive tests for an instantaneous point source problem are shown in respectively Figs. 12a and 12b. Like in the previous case there is an additional first-order biogeochemical reaction causing a linear decay in the concentrations of Test 10 in Fig 12b.

For Test 9, the Nash-Sutcliffe and RMSE values obtained were respectively 0.9535 and 4.614e-05 mg/l, and for Test 10 were respectively 0.9521 and 1.409e-05 mg/l. The effect of biogeochemistry in the resulting concentrations can be also clearly observed in the uppermiddle panel of Fig. 12a as compared to the upper-middle panel of Fig. 12b. Similarly to the previous cases of Tests 7 and 8, results show that the Nash-Sutcliffe values remain high, but they are slightly lower when compared to the purely biogeochemical kinetic Tests 1-6.

450 5 Conclusions

This two-part paper describes and applies a methodology for incorporating and crossmodel deploying multi-biogeochemistry and ecological simulations in existing process-based hydro-modeling tools, the OpenWQ framework. OpenWQ was developed based on the motivation that (1) hydrology has a strong impact on the fate of pollution in the environ-



(b) Test 8

Figure 11: Comparison between model results and analytical solution for Tests 7 and 8

ment and (2) more should be done to streamline collaboration between the hydrological
and biogeochemical communities. The first paper describes OpenWQ's concept as a coupler
and customizable biogeochemistry modeling framework to allow cross-model deployment
of eco-lab and co-creation modeling capabilities, providing a transformative direction for
innovation in water quality modeling. This approach aims to address the inherent challenge of water quality modeling of being at the intersection between several fields, including
biogeochemistry, soil science, hydrology, hydrodynamics, and hydrogeology.

This second paper describes the integration of OpenWQ into two hydro-models, SUMMA and CRHM, illustrating how coupling interfaces between the two models have not only enabled water quality simulation capacities in the host hydro-models but, even more importantly, enabled establishing a direct and permanent link for transfer of knowledge, innovation, and technology between modeling communities. Example applications of



(b) Test 10

Figure 12: Comparison between model results and analytical solution for Tests 9 and 10

⁴⁶⁷ pollution studies enabled by the coupling of the tools are provided to begin to address some
⁴⁶⁸ of these fundamental modeling challenges.

The key points of this paper are as follows:

1. The coupling of OpenWQ to CRHM and SUMMA illustrates how the coupling recipe 470 presented in Paper 1 can be deployed to link the model to existing hydro-models 471 even if displaying drastically different modeling structures. The coupling steps, 472 coupling functions, and wrapper interface functions have been designed to optimize 473 the coupling process. It was also demonstrated that new updates to OpenWQ can 474 be integrated into the coupled models without requiring changes in the coupler 475 functions. This provides a permanent, concrete link between the hydrological and 476 biogeochemical research communities, where innovations on either side can be readily 477

integrated through the coupler into a unified modeling framework. 478 479 2. The model development and applications presented illustrate how OpenWQ's flexible 480 multi-chemistry lab capabilities enable addressing a variety of environmental and 481 pollution problems. In Paper 1, we described how such model flexibility has been ma-482 terialized. In this paper, we deployed the model to illustrate how it can be used to look at the nitrogen cycle and eutrophication problems, BOD-DO and DO sag issues, and 484 other environmental problems involving reaction networks run in parallel or sequence. 485 486 3. The model applications presented aim to illustrate how biogeochemical modeling 487 frameworks can be readily deployed across different hydro-models using the same 488 input files. It has been shown how this can be extremely helpful in removing barriers 489 in the selection of the most appropriate modeling framework for specific landscapes 490 and climate regions. 491 492 4. Separating the numerical solver from the physicochemical calculations enables more 493 controlled simulations where the specific impact of different processes can not only 494 be properly quantified but also added or removed as needed for uncertainty analysis. 495 Such model architecture is critical for model scalability by enabling the integration 496 of new modules in a controlled manner.

The work presented here represents the first application of OpenWQ to improve water 498 quality modeling, but there are many additional opportunities to build on this work. The 499 case studies simulated here cover a range of pollution problems and hydrologic processes, 500 but they are naturally limited in terms of the environments and contamination problems 501 covered. The model simulations presented are also based on simple perturbation experi-502 ments, which were necessary for validation purposes; however, more comprehensive model 503 applications should be carried out in the future for comparison with competing modeling ap-504 proaches. The work presented also does not examine different numerical integration options. Research is needed to continue exploring the potential opportunities that OpenWQ creates 506 for testing different biogeochemical, biophysical, and hydrologic processes and their model 507 representations. Access to the OpenWQ source code and the case study tests is provided 508 through the SUMMA website at openwg.readthedocs.io. 509

510 Acknowledgments

511

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516 Appendices

A Deriving analytical solutions for Tests 1 and 2: Single-species batch reaction

The single-species batch reactor tests contemplated involve (a) first-order decay and (b) second-order decay kinetics. Fig. 1 depicts the biogeochemical process. The analytical solution of a single species subject to first-order decay can be obtained by integrating Eq. 24 obtaining Eq. 25. In turn, the analytical solution of a single species subject to second-order decay can be obtained by integrating Eq. 26 obtaining Eq. 27.

$$\frac{dc_A}{dt} = -kc_A \tag{24}$$

$$\frac{dc_A}{dt} = -kc_A \Leftrightarrow \frac{1}{c_A} dc_A = -kdt \Leftrightarrow \int_0^t \frac{1}{c_A} dc_A = \int_0^t -kdt \Leftrightarrow \ln(c_A) \Big|_0^t = -kt \Big|_0^t \Leftrightarrow \ln(c_A(t)) - \ln(c_A(0)) = -k\Delta t \Leftrightarrow \ln\left(\frac{c_A(t)}{c_A(0)}\right) = -k\Delta t \Leftrightarrow \frac{c_A(t)}{c_A(0)} = -k\Delta t \Leftrightarrow c_A(t) = c_A(0)e^{-k\Delta t}$$
(25)

$$\frac{dc_A}{dt} = -kc_A^2 \tag{26}$$

$$\frac{dc_A}{dt} = -kc_A^2 \Leftrightarrow \frac{1}{c_A^2} dc_A = -kdt \Leftrightarrow$$

$$\int_0^t c_A^{-2} dc_A = \int_0^t -kdt \Leftrightarrow -c_A^{-1} \Big|_0^t = -kt \Big|_0^t \Leftrightarrow$$

$$c_A^{-1} \Big|_0^t = kt \Big|_0^t \Leftrightarrow c_A(t)^{-1} - c_A(0)^{-1} = k\Delta t \Leftrightarrow$$

$$\frac{1}{c_A(t)} = \frac{1}{c_A(0)} + k\Delta t \Leftrightarrow c_A(t) = \frac{1}{\frac{1}{c_A(0)} + k\Delta t} \Leftrightarrow$$
(27)

⁵²⁴ B Deriving analytical solution for Test 3: Two-species reaction network

The two-species reaction network with linear decay is represented schematically in Fig. 2. Similarly to the single-species case above with first-order decay (Eq. 25), the fate of Species A in a two-species reaction network only depends on its own concentration and is given by Eq. 28.

$$c_A(t) = c_A(0)e^{(-k_A\Delta t)} \tag{28}$$

However, the fate of Species B depends on the time-evolution of concentrations of both Species A and B. This can be expressed as:

$$\frac{dc_B}{dt} = k_A c_A - k_B c_B \tag{29}$$

To derive the analytical solution of Eq. 29, we can start by replacing $c_A(t)$ by its analytical solution derived previously in Eq. 25. We obtain:

$$\frac{dc_B}{dt} = k_A c_A(0) e^{(-k_A t)} - k_B c_B(t) \Leftrightarrow$$

$$\frac{dc_B}{dt} + k_B c_B = k_A c_A(0) e^{(-k_A t)}$$
(30)

The integrating factor, $e^{(k_B t)}$, can be used to integrate Eq. 30 by parts (i.e., d(uv) = udv + vdu). Multiplying Eq. 30 by the integrating factor, we obtain:

$$\frac{dc_B}{dt}e^{(k_Bt)} + k_Bc_B(t)e^{(k_Bt)} = k_Ac_A(0)e^{(-k_At)}e^{(k_Bt)} \Leftrightarrow \frac{dc_B}{dt}e^{(k_Bt)} + k_Bc_B(t)e^{(k_Bt)} = k_Ac_A(0)e^{(k_B-k_A)t}$$
(31)

Using the general expression for integration by parts and considering that functions uand v are given by $u = c_B$ and $v = e^{(k_B t)}$, we obtain:

$$d\left(c_B e^{(k_B t)}\right) = c_B d\left(e^{(k_B t)}\right) + e^{(k_B t)} dc_B \tag{32}$$

Replacing the derivative of $e^{(k_B t)}$ in Eq. 32, where $d(e^{(k_B t)}) = k_B e^{(k_B t)} dt$, leads to:

$$d(uv) = udv + vdu \Leftrightarrow$$

$$d(c_B e^{(k_B t)}) = c_B [k_B e^{(k_B t)} dt] + e^{(k_B t)} dc_B$$
(33)

or

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$$\frac{d}{dt}(c_B e^{(k_B t)}) = c_B k_B e^{(k_B t)} + e^{(k_B t)} \frac{dc_B}{dt}$$
(34)

The right-hand side of Eq. 34 is the same as the left-hand side of Eq. 31, so the same applies to the other side of the equations. Thus, we can re-write the equations as

$$\frac{d}{dt} \left(c_B e^{(k_B t)} \right) = k_A c_A(0) e^{(-k_A t)} e^{(k_B t)} \Leftrightarrow
\frac{d}{dt} \left(c_B e^{(k_B t)} \right) = k_A c_A(0) e^{\left[(k_B - k_A) t \right]} \Leftrightarrow
d \left(c_B e^{(k_B t)} \right) = k_A c_A(0) e^{\left[(k_B - k_A) t \right]} dt$$
(35)

540 Integrating both sides of the equation, we obtain

$$\int_{0}^{t} d(c_{B}e^{(k_{B}t)}) = \int_{0}^{t} k_{A}c_{A}(0)e^{\left[(k_{B}-k_{A})t\right]}dt \Leftrightarrow c_{B}(t)e^{(k_{B}t)} = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}e^{\left[(k_{B}-k_{A})t\right]} + C',$$
(36)

where C' is the integration constant. At t = 0, c_B becomes $c_B(0)$, so Eq. 36 can be written as

$$c_B(0)e^0 = \frac{k_A c_A(0)}{k_B - k_A} e^0 \Leftrightarrow$$

$$\Leftrightarrow C' = c_B(0) - \frac{k_A c_A(0)}{k_B - k_A},$$
(37)

⁵⁴³ which can be replaced in Eq. 36 to give:

$$c_{B}(t)e^{(k_{B}t)} = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}e^{\left[(k_{B}-k_{A})t\right]} + c_{B}(0) - \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \Leftrightarrow$$

$$c_{B}(t)e^{(k_{B}t)} = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{\left[(k_{B}-k_{A})t\right]} - 1\right] + c_{B}(0) \Leftrightarrow$$

$$c_{B}(t) = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{e^{\left[(k_{B}-k_{A})t\right]}}{e^{(k_{B}t)}} - \frac{1}{e^{(k_{B}t)}}\right] + \frac{c_{B}(0)}{e^{(k_{B}t)}} \Leftrightarrow$$

$$c_{B}(t) = \frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{(-k_{A}t)} - e^{(-k_{B}t)}\right] + c_{B}(0)e^{(-k_{B}t)}$$
(38)

C Deriving analytical solution for Test 4: Three-species reaction network

A three-species reaction network with linear decay is represented schematically in Fig.
Concentrations of Species A and Species B are given by respectively Eqs. 28 and 36.
The partial differential equation for the concentration of Species C can be written as:

$$\frac{dc_C}{dt} = k_B c_B - k_C c_C \tag{39}$$

The expression for $c_B(t)$ can be obtained from Eq. 36 and replaced in Eq. 40. We obtain:

$$\frac{dc_C}{dt} = k_B \left(\frac{k_A c_A(0)}{k_B - k_A} \left[e^{(-k_A t)} - e^{(-k_B t)} \right] + c_B(0) e^{(-k_B t)} \right) -k_C c_C \Leftrightarrow \frac{dc_C}{dt} + k_C c_C = k_B \left(\frac{k_A c_A(0)}{k_B - k_A} \left[e^{(-k_A t)} - e^{-k_B t} \right] + c_B(0) e^{(-k_B t)} \right)$$
(40)

The integrating factor used in Section B to enable integration of parts, $e^{(k_C t)}$, can be used here as well (i.e., d(uv) = udv + vdu). Multiplying Eq. 40 by this integrating factor, we obtain:

$$\frac{dc_C}{dt}e^{(k_Ct)} + k_Cc_C(t)e^{(k_Ct)} = k_B\left(\frac{k_Ac_A(0)}{k_B - k_A}\left[e^{(-k_At)} - e^{(-k_Bt)}\right] + c_B(0)e^{(-k_Bt)}\right)e^{(k_Ct)}$$
(41)

We can use the same expression derived to enable integration by parts in the previous case (Eq. 34) but replacing Species B by Species C: $\frac{d}{dt}\left(c_{C}e^{(k_{C}t)}\right) = c_{C}k_{C}e^{(k_{C}t)} + \frac{dc_{C}}{dt}e^{k_{C}t}$. Also, similarly to the previous case, the right-hand side of this equation is the same as the

$_{\tt 557}$ $\,$ left-hand side of Eq. 41, so the same applies to the other side of the equations. Thus, we

⁵⁵⁸ can re-write the equations as follows:

$$\frac{d}{dt}\left(c_{C}e^{k_{C}t}\right) = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{-k_{A}t} - e^{-k_{B}t}\right] + c_{B}(0)e^{-k_{B}t}\right)e^{(k_{C}t)} \Leftrightarrow d\left(c_{C}e^{k_{C}t}\right) = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{-k_{A}t} - e^{-k_{B}t}\right] + c_{B}(0)e^{-k_{B}t}\right)e^{(k_{C}t)}dt \Leftrightarrow d\left(c_{C}e^{k_{C}t}\right) = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[e^{(k_{C}-k_{A})t} - e^{(k_{C}-k_{B})t}\right] + c_{B}(0)e^{(k_{C}-k_{B})t}\right] + c_{B}(0)e^{(k_{C}-k_{B})t}\right)dt$$
(42)

⁵⁵⁹ Integrating both sides of the equation we obtain:

$$\int_{0}^{t} d(c_{C}e^{(k_{C}t)}) =$$

$$\int_{0}^{t} \left[k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[e^{\left[(k_{C}-k_{A})t \right]} - e^{\left[(k_{C}-k_{B})t \right]} \right] \right] +$$

$$c_{B}(0)e^{\left[(k_{C}-k_{B})t \right]} \right] \Rightarrow \qquad (43)$$

$$c_{C}(t)e^{(k_{C}t)} = k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{1}{k_{C}-k_{A}} e^{\left[(k_{C}-k_{A})t \right]} -$$

$$\frac{1}{k_{C}-k_{B}} e^{\left[(k_{C}-k_{B})t \right]} \right] + \frac{c_{B}(0)}{k_{C}-k_{B}} e^{\left[(k_{C}-k_{B})t \right]} \right) + C'',$$

where C" is the integration constant. At t = 0, c_C becomes $c_C(t)$, so Eq. 43 becomes

$$c_{C}(0)e^{(k_{C}0)} = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}e^{0}-\frac{1}{k_{C}-k_{B}}e^{0}\right]\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}e^{0}\right) + C'' \Leftrightarrow$$

$$c_{C}(0)1 = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}1-\frac{1}{k_{C}-k_{B}}1\right]\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}1\right) + C'' \Leftrightarrow C'' =$$

$$c_{C}(0) - k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}-\frac{1}{k_{C}-k_{B}}\right]\right) + \frac{c_{B}(0)}{k_{C}-k_{B}}\right),$$
(44)

$_{561}$ which can be replaced in Eq. 43 to give:

$$c_{C}(t)e^{(k_{C}t)} = k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}e^{\left[(k_{C}-k_{A})t\right]}-\frac{1}{k_{C}-k_{B}}e^{\left[(k_{C}-k_{B})t\right]}\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}e^{\left[(k_{C}-k_{B})t\right]}\right) +$$
(45)
$$c_{C}(0) - k_{B}\left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}}\left[\frac{1}{k_{C}-k_{A}}-\frac{1}{k_{C}-k_{B}}\right] + \frac{c_{B}(0)}{k_{C}-k_{B}}\right)$$

This can be rearranged to give:

$$c_{C}(t) = k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{1}{k_{C}-k_{A}} e^{(-k_{A}t)} - \frac{1}{k_{C}-k_{B}} e^{(-k_{B}t)} \right] + \frac{c_{B}(0)}{k_{C}-k_{B}} e^{(-k_{B}t)} \right] + c_{C}(0)e^{(-k_{C}t)} - k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{1}{k_{C}-k_{A}} - \frac{1}{k_{C}-k_{B}} \right] + \frac{c_{B}(0)}{k_{C}-k_{B}} \right) e^{(-k_{C}t)},$$

$$(46)$$

₅₆₂ and then simplified to:

$$c_{C}(t) = k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{e^{(-k_{A}t)}}{k_{C}-k_{A}} - \frac{e^{(-k_{B}t)}}{k_{C}-k_{B}} \right] + \frac{c_{B}(0)e^{(-k_{B}t)}}{k_{C}-k_{B}} \right) + c_{C}(0)e^{(-k_{C}t)} - (47)$$

$$k_{B} \left(\frac{k_{A}c_{A}(0)}{k_{B}-k_{A}} \left[\frac{e^{(-k_{C}t)}}{k_{C}-k_{A}} - \frac{e^{(-k_{C}t)}}{k_{C}-k_{B}} \right] + \frac{c_{B}(0)e^{(-k_{C}t)}}{k_{C}-k_{B}} \right)$$

⁵⁶³ D Deriving analytical solution for Test 5: Nitrogen cycle

The simulation of the nitrogen cycle is very common in freshwater water quality studies 564 because it allows for addressing issues with nutrient pollution and eutrophication. The 565 nitrogen cycle simulated involves different nitrogen species (Fig. 4). The conceptual cycling 566 framework used is similar to that used in catchment-scale models like HYPE Lindström et 567 al. (2010), CRHM-WQ (Costa et al., 2021) models, SWAT (Arnold et al., 2012), and INCA 568 ("A semi-distributed ntegrated itrogen model for multiple source assessment in tchments 569 (INCA): Part I — model structure and process equations", 1998). The analytical solutions 570 for the different nitrogen species can be obtained directly (or derived) from Eq. 28, Eq. 38, 571 or Eq. 47 depending on their positioning in the reaction chain: 572

1. **refractoryN** (*Nref*): Similar to Eq. 28, which for this specific process and N species reads as:

$$c_{Nref}(t) = c_{Nref}(0)e^{\left[-(k_{degr} + k_{diss1})t\right]}$$

$$\tag{48}$$

2. labileN (*Nref*): Similar to Eq. 36, which for this specific processes and N species reads as:

$$c_{Nlab}(t) = \frac{k_{degr}c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \begin{bmatrix} e^{(-k_{degr}t)} - e^{\left[-(k_{min} + k_{diss2})t\right]} \\ + c_{Nlab}(0)e^{\left[-(k_{min} + k_{diss2})t\right]} \end{bmatrix}$$
(49)

3. **DON**: Needs specific derivation using Eqs. 28 and 36 (see below)

4. **DIN**: Needs specific derivation using Eqs. 49 and 28 (see below)

 $_{575}$ **DON** The concentration of DON does not depend on itself, but rather on the evolution of the concentration of Nref and Nlab. The differential equation can be written $_{577}$ as:

$$\frac{dc_{DON}}{dt} = k_{diss1}c_{Nref} + k_{diss2}c_{Nlab} \tag{50}$$

Analytical solutions suitable for c_{Nref} and c_{Nlab} , due to their positioning in the rection chain, have been derived before (Eqs. 28 and 38, respectively). These solutions can be replaced in Eq. 50.

$$\frac{dc_{DON}}{dt} = k_{diss1} \left[c_{Nref} e^{\left[-(k_{degr} + k_{diss1})t \right]} \right] + k_{diss2} \left[\frac{k_{degr} c_{Nref}}{k_{min} + k_{diss2} - k_{degr}} \left(e^{\left(-k_{degr}t \right)} - e^{\left[-(k_{min} + k_{diss2})t \right]} \right) + c_{Nlab}(0) e^{\left[-(k_{min} + k_{diss2})t \right]} \right]$$
(51)

Since this solution is independent of the concentration of DON, it can be integrated in its current state.

$$dc_{DON} = k_{diss1} \left[c_{Nref} e^{\left[-(k_{degr} + k_{diss1})t \right]} \right] dt$$

$$+k_{diss2} \left[\frac{k_{degr} c_{Nref}}{k_{min} + k_{diss2} - k_{degrd}} \left(e^{\left(-k_{degr}t \right)} - e^{\left[-(k_{min} + k_{diss2})t \right]} \right] \right] + c_{Nlab}(0) e^{\left[-(k_{min} + k_{diss2})t \right]} dt \Leftrightarrow$$

$$c_{DON}(t) = \frac{k_{diss1} c_{Nref}(t)}{-(k_{degr} + k_{diss1})} e^{\left[-(k_{degr} + k_{diss1})t \right]} + \frac{k_{diss2} k_{degr} c_{Nref}(t)}{-k_{degr}} \left(\frac{e^{\left(-k_{degr}t \right)}}{-(k_{min} + k_{diss2})t \right]} - \frac{e^{\left[-(k_{min} + k_{diss2})t \right]}}{-(k_{min} + k_{diss2})} \right) + \frac{k_{diss2} c_{Nlab}(0)}{-(k_{min} + k_{diss2})} e^{\left[-(k_{min} + k_{diss2})t \right]} + C'''$$
(52)

Considering that at t = 0, $c_{DON}(t)$ becomes $c_{DON}(0)$, we can determine C''' as follows:

$$C''' = c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{1}{-k_{degr}} - \frac{1}{-(k_{min} + k_{diss2})}\right) + \frac{k_{diss2}c_{Nlab}(0)}{-(k_{min} + k_{diss2})}$$
(53)

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Replacing Eq. 53 into Eq. 52 we obtain:

$$c_{DON}(t) = \frac{k_{diss1}c_{Nref}(t)}{-(k_{degr} + k_{diss1})} e^{\left[-(k_{degr} + k_{diss1})t\right]} + \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{e^{(-k_{degr}t)}}{-k_{degr}} - \frac{e^{\left[-(k_{min} + k_{diss2})t\right]}}{-(k_{min} + k_{diss2})}\right) + \frac{k_{diss2}c_{Nlab}(0)}{-(k_{min} + k_{diss2})}$$
(54)
$$e^{\left[-(k_{min} + k_{diss2})t\right]} + c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{1}{-k_{degr}} - \frac{1}{-(k_{min} + k_{diss2})}\right) + \frac{k_{diss2}c_{Nlab}(0)}{-(k_{min} + k_{diss2})}$$

Eq. 54 can be simplified to obtain:

$$c_{DON}(t) = -\frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} e^{\left[-(k_{degr} + k_{diss1})t\right]} + \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2} - k_{degr}} \left(\frac{e^{\left[k_{min} + k_{diss2}t\right]}}{k_{min} + k_{diss2}} - \frac{e^{\left(-k_{degr}t\right)}}{k_{degr}}\right) - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}}$$

$$e^{\left[-(k_{min} + k_{diss2})t\right]} + c_{DON}(0) + \frac{k_{diss1}c_{Nref}(t)}{k_{degr} + k_{diss1}} - \frac{k_{diss2}k_{degr}c_{Nref}(t)}{k_{min} + k_{diss2}}$$

$$\left(\frac{1}{k_{min} + k_{diss2}} - \frac{1}{k_{degr}}\right) - \frac{k_{diss2}c_{Nlab}(0)}{k_{min} + k_{diss2}}$$

586

$$\frac{dc_{DIN}}{dt} = k_{min}c_{Nlab} - (k_{denit} + k_{plantup})c_{DIN}$$
(56)

The concentration of $c_{Nlab}(t)$ has been derived above in Eq. 49 and can be replaced in Eq. 56 to give:

$$\frac{dc_{DIN}}{dt} = k_{min} \left(\frac{k_{degr} c_{(Nref,0)}}{k_{min} + k_{diss2} - k_{degr}} \left[e^{(-k_{degr}t)} - e^{\left[- (k_{min} + k_{diss2})t \right]} \right] + c_{Nlab}(0) e^{\left[- (k_{min} + k_{diss2})t \right]} \right) - (k_{denit} + k_{plantup}) c_{DIN}(t) \Leftrightarrow \left(\frac{dc_{DIN}}{dt} = \frac{k_{min} k_{degr} c_{(Nref,0)}}{k_{min} + k_{diss2} - k_{degr}} \left[e^{(-k_{degr}t)} - e^{\left[- (k_{min} + k_{diss2})t \right]} \right] + k_{min} c_{Nlab}(0) e^{\left[- (k_{min} + k_{diss2})t \right]} - (k_{denit} + k_{plantup}) c_{DIN}$$
(57)

The same integrating factor used before, which in this context reads $e^{\left[(k_{denit}+k_{plantup})t\right]}$, can be used here to allow similar integration by parts (i.e., d(uv) = udv + vdu). Multiplying Eq. 57 by the integration factor, we obtain:

$$\frac{dc_{DIN}}{dt}e^{\left[(k_{denit}+k_{plantup})t\right]} + (k_{denit}+k_{plantup})$$

$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]}$$

$$= \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\left[e^{(-k_{degr}t)} - e^{\left[-(k_{min}+k_{diss2})t\right]}\right]$$

$$e^{\left[(k_{denit}+k_{plantup})t\right]} + k_{min}c_{Nlab}(0)e^{\left[-(k_{min}+k_{diss2})t\right]}$$

$$e^{\left[(k_{denit}+k_{plantup})t\right]}$$
(58)

Now we can use the same expression derived for integration by parts of previous cases (i.e., Eq. 34) and adapt it to *DIN* and the respective transformations:

$$\frac{d}{dt} \Big[c_{DIN} e^{\left((k_{denit} + k_{plantup})t \right)} \Big] = \frac{dc_{DIN}}{dt} e^{\left[(k_{denit} + k_{plantup})t \right]} + (k_{denit} + k_{plantup})c_{DIN} e^{\left[(k_{denit} + k_{plantup})t \right]}$$
(59)

Similarly to previous cases, the right-hand side of this equation is the same as the lefthand side of Eq. 57, so the same applies to the other side of the equations so that we can re-write the equations as:

$$\frac{d}{dt} \left(c_{DIN} e^{\left[(k_{denit} + k_{plantup})t \right]} \right) = \frac{k_{min} k_{degr} c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \\
\left[e^{(-k_{degr}t)} - e^{\left[- (k_{min} + k_{diss2})t \right]} \right] e^{\left[(k_{denit} + k_{plantup})t \right]} + k_{min} c_{Nlab}(0) e^{\left[- (k_{min} + k_{diss2})t \right]} e^{\left[(k_{denit} + k_{plantup})t \right]} \Leftrightarrow d\left(c_{DIN} e^{\left[(k_{denit} + k_{plantup})t \right]} \right) = \left(\frac{k_{min} k_{degr} c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \right) \\
\left[e^{\left[((k_{denit} + k_{plantup}) - k_{degr})t \right]} - e^{\left[(k_{denit} + k_{plantup} - k_{min} - k_{diss2})t \right]} \right] \\
+ k_{min} c_{Nlab}(0) e^{\left[(k_{denit} + k_{plantup} - k_{min} - k_{diss2})t \right]} \right) dt$$
(60)

⁵⁹⁷ Now we can integrate the equation with respect to $c_{DIN}(t)$ and t:

$$\int d\left(c_{DIN}e^{\left[(k_{denit}+k_{plantup})t\right]}\right) = \int \left(\frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\right)$$

$$\left[e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]} - e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}\right]$$

$$+k_{min}c_{Nlab}(0)e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}\right)dt \Leftrightarrow$$

$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]} = \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}$$

$$\left[e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]} - \frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}}{(k_{denit}+k_{plantup})-k_{min}-k_{diss2}}\right]$$

$$+k_{min}c_{Nlab}(0)\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}}{(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t}$$

⁵⁹⁸ Considering that at t = 0, $c_{DIN}(t)$ becomes $c_{DIN}(0)$, we can determine the constant ⁵⁹⁹ C'''' as follows:

$$C'''' = c_{DIN}(0) - \frac{k_{miner}k_{degr}c_{Nref}(0)}{k_{min} + k_{diss2} - k_{degr}} \begin{bmatrix} \frac{1}{(k_{denit} + k_{plantup}) - k_{degr}} - \frac{1}{(k_{denit} + k_{plantup}) - k_{min} - k_{diss2}} \end{bmatrix}$$
(62)
$$-k_{miner}c_{Nlab}(0) \frac{1}{(k_{denit} + k_{plantup}) - k_{min} - k_{diss2}}$$

Replacing the integrating constant in Eq. 61 we obtain:

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$$c_{DIN}(t)e^{\left[(k_{denit}+k_{plantup})t\right]} = \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}$$

$$\left[\frac{e^{\left[(k_{denit}+k_{plantup}-k_{degr})t\right]}}{k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]} + k_{min}c_{Nlab}(0)\frac{e^{\left[(k_{denit}+k_{plantup}-k_{min}-k_{diss2})t\right]}}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right]$$

$$+c_{DIN}(0) - \frac{k_{min}k_{degr}c_{Nref}(0)}{k_{min}+k_{diss2}-k_{degr}}\left[\frac{1}{k_{denit}+k_{plantup}-k_{degr}} - \frac{1}{k_{denit}+k_{plantup}-k_{diss2}}\right] - k_{min}c_{Nlab}(0)$$

$$\left(\frac{1}{k_{denit}+k_{plantup}-k_{min}-k_{diss2}}\right)$$

$$(63)$$

⁶⁰¹ E Deriving analytical solution for Test 6: DO-BOD cycle

Dissolved oxygen (DO) depletion in rivers and lakes is a common consequence of pollution. It typically occurs near discharge outlets of untreated wastewater, which can lead to high levels of organic matter that, during decomposition by aerobic bacteria, causes a DOsag in the water column. For purposes of DO dynamics, such organic matter is typically translated into units of biological oxygen demand (BOD). Fig. 5 depicts the processes involved, and the Streeter–Phelps equation (Streeter & Phelps, 1925) has long been derived for such problems. It reads:

$$\frac{dD}{dt} = k_1 L_t - k_2 D \tag{64}$$

The equation states that the total change in oxygen deficit (D) is equal to the difference between the two rates of deoxygenation and reaeration at any time. The derivation of the Streeter-Phelps equation is well known, so it won't be repeated here.

$$D = \frac{k_1 L_a}{k_2 - k_1} (e^{(-k_1 t)} - e^{(-k_2 t)}) + D_a e^{(-k_2 t)}$$
(65)

where D is the saturation deficit in g/m^3 , which can be derived from the dissolved oxygen concentration at saturation minus the actual dissolved oxygen concentration ($D = DO_{sat} - DO$); k_1 is the deoxygenation rate, usually in d^{-1} ; k_2 is the reaeration rate, usually in d^{-1} ; L_a is the initial oxygen demand of organic matter in the water, also called the ultimate BOD (BOD at time $t = \infty$). The unit of L_a is g/m^3 ; k_1 ; L_t is the oxygen demand remaining at time t, $L_t = L_a e^{(-k_1 t)}$; D_a is the initial oxygen deficit $[g/m^3]$; and t is the elapsed time in days.

⁶¹⁹ F Reactive and non-reactive transport: Instantaneous point source (Tests 7 and 8)

The general 1D advection-dispersion-reaction partial differential equation is given by

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c \tag{66}$$

where c is the concentration of the tracer, t is time, D_x is the dispersion coefficient in the flow direction, x is the distance in the flow direction, ν is the flow velocity in the x direction, and λ is a linear reaction term that can be set to zero for non-reactive tracers. Various analytical solutions for Eq. 66 have been derived and compiled by Wexler (1992) for different idealized scenarios.

Eq. 66 has been modified by Wexler (1992) to include an instantaneous point source and it reads:

$$\frac{\partial c}{\partial t} = D_x \frac{\partial^2 c}{\partial x^2} - \nu \frac{\partial c}{\partial x} - \lambda c + \frac{Q}{n} \partial t c_0 \delta(x - x_c) \delta(t - t'_c)$$
(67)

where t'_c is the time at which the instantaneous point source is activated, Q is the fluid injection rate, n is the aquifer porosity, δ is the Dirac delta (impulse) function, and \mathbf{x}_c is the location of the instantaneous point source. The analytical solution to Eq. 67 has been derived by Bear (1979) considering the following boundary conditions and initial conditions.

6321. Boundary conditions (Neumann or second-type):633 $c, \nabla c = 0, \quad x = \pm \infty; \ y = \pm \infty$ 6342. Initial conditions:635 $c^{t=0} = 0, \quad -\infty < x < \infty; -\infty < y < \infty,$

Eq. 68 shows the resulting analytical solution that can be used for solutes that are non-reactive (Test 7) or reactive (Test 8) by simply setting $\lambda = 0$ in the former (Test 7).

$$c(x) = \frac{c_0 Q}{4n\pi\gamma D_x} \exp\left[\frac{\nu(x-x_0)}{2D_x}\right]$$

$$\frac{1}{t} \exp\left[-\left(\frac{\nu^2}{4D_x}\right)\tau - \frac{(x-Xc)^2}{4D_x\tau}\right],$$
(68)

The equation assumes that (1) the fluid has constant density and viscosity, (2) the 638 solute may be subject to first-order chemical transformation, (3) the flow is uniform in 639 the x-direction only, (4) the flow velocity is constant, and (5) the longitudinal dispersion 640 coefficient (D) is constant. The solution has a mathematical singularity in the point source 641 region (Abdelaziz et al., 2013) because it takes the form of the exponential integral when 642 both $x - x_c$ approaches zero, hence becoming infinite at $\tau = 0$. However, according to 643 Wexler (1989) the solution is still valid as long as $(x - x_c)^2$ is larger than ν^2 , as it is the 644 case in our simulation tests. 645

G Reactive and non-reactive transport: Continuous point source (Tests 9 and 10)

The analytical solution for the transport of solutes with a continuous point source has been derived in Wexler (1992) after modification from Bear (1979) and Genuchten & Alves (1982).

$$c(x) = \frac{c_0}{2} \left\{ \exp\left[\frac{x}{2D_x}(\nu - U)\right] \operatorname{erfc}\left[\frac{x - Ut}{2\sqrt{D_x t}}\right] + \exp\left[\frac{x}{2D_x}(\nu + U)\right] \operatorname{erfc}\left[\frac{x + Ut}{2\sqrt{D_x t}}\right] \right\}$$
(69)

where c_0 is the concentration of the continuous point source, $U = \sqrt{\nu^2 + 4\lambda D}$ and erfc is the complementary error function. Parameter λ is the first-order reaction rate used for Test 10, which can be set to zero for Test 9.

654 References

- Abdelaziz, R., Pearson, A., & Merkel, B. (2013). Lattice Boltzmann modeling for tracer
 test analysis in a fractured Gneiss aquifer. *Natural Science*, 5, 368-374. doi: 10.4236/
 ns.2013.53050
- Alexander, A. C., Levenstein, B., Sanderson, L. A., Blukacz-Richards, E. A., & Chambers,
 P. A. (2020). How does climate variability affect water quality dynamics in canada's oil
 sands region? Science of The Total Environment, 732, 139062. doi: https://doi.org/
 10.1016/j.scitotenv.2020.139062
- Arnold, J. G., Moriasi, D. N., Gassman, P. W., Abbaspour, K. C., White, M. J., Srinivasan,
 R., ... Jha, M. K. (2012). SWAT: Model Use, Calibration, and Validation. Trans.
 ASABE, 55(4), 1491–1508.
- Bear, J. (1979). *Hydraulics of groundwater*. McGraw-Hill series in water resources and environmental engineering. McGraw-Hill, New York.
- Beven, K. (2006). A manifesto for the equifinality thesis. *Journal of Hydrology*, 320(1), 18-36. (The model parameter estimation experiment) doi: https://doi.org/10.1016/j.jhydrol .2005.07.007
- Bussi, G., Janes, V., Whitehead, P. G., Dadson, S. J., & Holman, I. P. (2017). Dynamic response of land use and river nutrient concentration to long-term climatic changes. *The Science of the total environment*, 590-591, 818-831. doi: 10.1016/j.scitotenv.2017.03.069
- ⁶⁷³ Bussi, G., Whitehead, P. G., Jin, L., Taye, M. T., Dyer, E., Hirpa, F. A., ... Charles,
 ⁶⁷⁴ K. J. (2021). Impacts of climate change and population growth on river nutrient loads
 ⁶⁷⁵ in a data scarce region: The upper awash river (ethiopia). Sustainability, 13. doi: 10.3390/su13031254
- Celia, M. A., Bouloutas, E. T., & Zarba, R. L. (1990). A general mass-conservative numerical solution for the unsaturated flow equation. Water Resources Research, 26(7), 1483-1496.
 doi: https://doi.org/10.1029/WR026i007p01483
- Charlton, M. B., Bowes, M. J., Hutchins, M. G., Orr, H. G., Soley, R., & Davison, P. (2018).
 Mapping eutrophication risk from climate change: Future phosphorus concentrations in
 english rivers. Science of The Total Environment, 613-614, 1510-1526. doi: https://
 doi.org/10.1016/j.scitotenv.2017.07.218
- Clark, M. P., Nijssen, B., Lundquist, J. D., Kavetski, D., Rupp, D. E., Woods, R. A., ...
 Rasmussen, R. M. (2015a). A unified approach for process-based hydrologic modeling: 1.
 modeling concept. *Water Resources Research*, 51(4), 2498-2514. Retrieved from https://
 agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/2015WR017198
 doi.org/10.1002/2015WR017198
- ⁶⁸⁹ Clark, M. P., Nijssen, B., Lundquist, J. D., Kavetski, D., Rupp, D. E., Woods, R. A.,
 ⁶⁹⁰ ... Marks, D. G. (2015b). A unified approach for process-based hydrologic model ⁶⁹¹ ing: 2. model implementation and case studies. Water Resources Research, 51(4), 2515 ⁶⁹² 2542. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/
 ⁶⁹³ 2015WR017200 doi: https://doi.org/10.1002/2015WR017200
- ⁶⁹⁴ Clark, M. P., Zolfaghari, R., Green, K. R., Trim, S., Knoben, W. J. M., Bennett, A., ⁶⁹⁵ ... Spiteri, R. J. (2021a). The numerical implementation of land models: Problem

- ⁶⁹⁶ formulation and laugh tests. Journal of Hydrometeorology, 22(6), 1627 1648. doi: ⁶⁹⁷ 10.1175/JHM-D-20-0175.1
- ⁶⁹⁸ Clark, M. P., Zolfaghari, R., Green, K. R., Trim, S., Knoben, W. J. M., Bennett, A.,
 ⁶⁹⁹ ... Spiteri, R. J. (2021b). The numerical implementation of land models: Problem
 ⁷⁰⁰ formulation and laugh tests. *Journal of Hydrometeorology*, 22(6), 1627 1648. doi:
 ⁷⁰¹ https://doi.org/10.1175/JHM-D-20-0175.1
- Costa, D., Baulch, H., Elliott, J., Pomeroy, J., & Wheater, H. (2020). Modelling nutrient dynamics in cold agricultural catchments: A review. *Environmental Modelling & Software*, 124, 104586. doi: https://doi.org/10.1016/j.envsoft.2019.104586
- Costa, D., Pomeroy, J. W., Brown, T., Baulch, H., Elliott, J., & Macrae, M. (2021).
 Advances in the simulation of nutrient dynamics in cold climate agricultural basins: Developing new nitrogen and phosphorus modules for the Cold Regions Hydrological Modelling Platform. *Journal of Hydrology*, 603, 126901. doi: https://doi.org/10.1016/ j.jhydrol.2021.126901
- ⁷¹⁰ Costa, D., Sutter, C., Shepherd, A., Jarvie, H., Wilson, H., Elliott, J., ... Macrae, M. ⁷¹¹ (2022). Impact of climate change on catchment nutrient dynamics: insights from around ⁷¹² the world. *Environmental Reviews*, $\theta(0)$, null. doi: 10.1139/er-2021-0109
- Crossman, J., Futter, M. N., Whitehead, P. G., Stainsby, E., Baulch, H. M., Jin, L., ... Dillon, P. J. (2014). Flow pathways and nutrient transport mechanisms drive hydrochemical sensitivity to climate change across catchments with different geology and topography. *Hydrol. Earth Syst. Sci.*, 18, 5125-5148. doi: 10.5194/hess-18-5125-2014
- Effects of lake water level fluctuation due to drought and extreme winter precipitation
 on mixing and water quality of an alpine lake, case study: Lake arrowhead, california.
 (2020). Science of The Total Environment, 714, 136762. doi: https://doi.org/10.1016/
 j.scitotenv.2020.136762
- El-Khoury, A., Seidou, O., Lapen, D. R., Que, Z., Mohammadian, M., Sunohara, M., & Bahram, D. (2015). Combined impacts of future climate and land use changes on discharge, nitrogen and phosphorus loads for a canadian river basin. *Journal of Environmental Management*, 151, 76-86. doi: https://doi.org/10.1016/j.jenvman.2014.12.012
- Ervinia, A., Huang, J., & Zhang, Z. (2020). Nitrogen sources, processes, and associated impacts of climate and land-use changes in a coastal china watershed: Insights from the inca-n model. *Marine Pollution Bulletin*, 159, 111502. doi: https://doi.org/10.1016/j.marpolbul.2020.111502
- Field, C. B., & Barros, V. R. (2014). Climate change 2014: Impacts, adaptation, and vulnerability: Working group ii contribution to the fifth assessment report of the intergovernmental panel on climate change (Tech. Rep.).
- Genuchten, M. T., & Alves, W. (1982). Analytical solutions of the one-dimensional convective-dispersive solute transport equation. US Department of Agriculture.
- Jarsjö, J., Törnqvist, R., & Su, Y. (2017). Climate-driven change of nitrogen reten tion-attenuation near irrigated fields: multi-model projections for central asia. Envi ronmental Earth Sciences, 76, 117. doi: 10.1007/s12665-017-6418-y
- Jeppesen, E., Kronvang, B., Olesen, J. E., Audet, J., Søndergaard, M., Hoffmann, C. C., ...
 Özkan, K. (2011). Climate change effects on nitrogen loading from cultivated catchments
 in europe: implications for nitrogen retention, ecological state of lakes and adaptation.
 Hydrobiologia, 663, 1-21. doi: 10.1007/s10750-010-0547-6
- Lindström, G., Pers, C., Rosberg, J., Strömqvist, J., & Arheimer, B. (2010). Development and testing of the HYPE (Hydrological Predictions for the Environment) water quality model for different spatial scales. *Hydrology Research*, 41 (3-4), 295. doi: 10.2166/nh.2010
 .007
- Luo, C., Li, Z., Liu, H., Li, H., Wan, R., Pan, J., & Chen, X. (2020). Differences in the responses of flow and nutrient load to isolated and coupled future climate and land use changes. *Journal of Environmental Management*, 256, 109918. doi: https://doi.org/10.1016/j.jenvman.2019.109918

Ockenden, M. C., Deasy, C. E., Benskin, C., Beven, K. J., Burke, S., Collins, A. L., ... 749 Haygarth, P. M. (2016). Changing climate and nutrient transfers: Evidence from high 750 temporal resolution concentration-flow dynamics in headwater catchments. Science of 751 The Total Environment, 548-549, 325-339. doi: https://doi.org/10.1016/j.scitotenv.2015 752 .12.086 753

- Ockenden, M. C., Hollaway, M. J., Beven, K. J., Collins, A. L., Evans, R., Falloon, P. D., ... 754 Haygarth, P. M. (2017). Major agricultural changes required to mitigate phosphorus losses 755 under climate change. Nature Communications, 8, 161. doi: 10.1038/s41467-017-00232-0 756
- Onishi, T., Yoshino, J., Hiramatsu, K., & Somura, H. (2020). Developing a hydro-chemical 757 model of ise bay watersheds and the evaluation of climate change impacts on discharge 758 and nitrate-nitrogen loads. Limnology, 21, 465-486. doi: 10.1007/s10201-020-00622-2 759
- Partow, A. (1999). ExprTk C++ Mathematical Expression Library. 760
- Pomeroy, J., Brown, T., Fang, X., Shook, K., Pradhananga, D., Armstrong, R., ... Lopez 761 Moreno, J. (2022). The cold regions hydrological modelling platform for hydrological 762 diagnosis and prediction based on process understanding. Journal of Hydrology, 615, 763 128711. doi: https://doi.org/10.1016/j.jhydrol.2022.128711 764
- Pomerov, J. W., Grav, D. M., Brown, T., Hedstrom, N. R., Quinton, W. L., Granger, R. J., 765 & Carey, S. K. (2007). The cold regions hydrological model: A platform for basing process 766 representation and model structure on physical evidence. Hydrological Processes, 21(19), 767 2650-2667. doi: {10.1002/hyp.6787} 768
- Raimonet, M., Thieu, V., Silvestre, M., Oudin, L., Rabouille, C., Vautard, R., & Garnier, 769 J. (2018). Landward perspective of coastal eutrophication potential under future climate 770 change: The seine river case (france). Frontiers in Marine Science, 5, 136. 771
- Reggiani, P., Sivapalan, M., & Majid Hassanizadeh, S. (1998). A unifying framework for 772 watershed thermodynamics: balance equations for mass, momentum, energy and entropy, and the second law of thermodynamics. Advances in Water Resources, 22(4), 367-398. 774 doi: https://doi.org/10.1016/S0309-1708(98)00012-8 775
- Renard, B., Kavetski, D., Kuczera, G., Thyer, M., & Franks, S. W. (2010). Understand-776 ing predictive uncertainty in hydrologic modeling: The challenge of identifying input 777 and structural errors. Water Resources Research, 46(5). doi: https://doi.org/10.1029/ 778 2009WR008328 779
- Rood, S. B., Pan, J., Gill, K. M., Franks, C. G., Samuelson, G. M., & Shepherd, A. 780 (2008). Declining summer flows of rocky mountain rivers: Changing seasonal hydrol-781 ogy and probable impacts on floodplain forests. Journal of Hydrology, 349, 397-410. doi: 782 https://doi.org/10.1016/j.jhydrol.2007.11.012 783

784

795

796

797

- A semi-distributed ntegrated itrogen model for multiple source assessment in tchments (INCA): Part I — model structure and process equations. (1998). Science of The Total 785 Environment, 210-211, 547-558. doi: 10.1016/S0048-9697(98)00037-0 786
- Shi, X., Mao, J., Thornton, P. E., Hoffman, F. M., & Post, W. M. (2011). The impact of 787 climate, co2, nitrogen deposition and land use change on simulated contemporary global 788 river flow. Geophysical Research Letters, 38. (https://doi.org/10.1029/2011GL046773) 789 doi: https://doi.org/10.1029/2011GL046773 790
- Sperotto, A., Molina, J. L., Torresan, S., Critto, A., Pulido-Velazquez, M., & Marcomini, 791 A. (2019). A bayesian networks approach for the assessment of climate change impacts 792 on nutrients loading. Environmental Science & Policy, 100, 21-36. doi: https://doi.org/ 793 10.1016/j.envsci.2019.06.004 794
 - Streeter, H., & Phelps, E. (1925). A study of the pollution and natural purification of the ohio river. iii. factors concerned in the phenomena of oxidation and reaeration. U.S. Government Printing Office.
- Suddick, E. C., Whitney, P., Townsend, A. R., & Davidson, E. A. (2013). The role of nitrogen 798 in climate change and the impacts of nitrogen-climate interactions in the united states: 799 foreword to thematic issue. Biogeochemistry, 114, 1-10. doi: 10.1007/s10533-012-9795-z 800
- Wexler, E. J. (1989). Analytical solutions for one-, two-, and three-dimensional solute 801
- transport in ground-water systems with uniform flow (Tech. Rep.). doi: 10.3133/ofr8956 802

- Wexler, E. J. (1992). Analytical solutions for one-, two-, and three-dimensional solute transport in ground-water systems with uniform flow. US Government Printing Office.
- Whitehead, P. G., Wilby, R. L., Battarbee, R. W., Kernan, M., & Wade, A. J. (2009). A review of the potential impacts of climate change on surface water quality. *Hydrological Sciences Journal*, 54, 101-123. (doi: 10.1623/hysj.54.1.101) doi: 10.1623/hysj.54.1.101
- Wilson, H., Elliott, J., Macrae, M., & Glenn, A. (2019). Near-surface soils as a source of phosphorus in snowmelt runoff from cropland. *Journal of Environmental Quality*, 48, 921-930. doi: 10.2134/jeq2019.04.0155
- Wilson, H. F., Casson, N. J., Glenn, A. J., Badiou, P., & Boychuk, L. (2019). Landscape controls on nutrient export during snowmelt and an extreme rainfall runoff event
 in northern agricultural watersheds. *Journal of Environmental Quality*, 48, 841-849. doi:
 10.2134/jeq2018.07.0278
- Wood, E. F., Roundy, J. K., Troy, T. J., van Beek, L. P. H., Bierkens, M. F. P., Blyth, E.,
 Whitehead, P. (2011). Hyperresolution global land surface modeling: Meeting a grand challenge for monitoring earth's terrestrial water. *Water Resources Research*, 47(5). doi: https://doi.org/10.1029/2010WR010090
- Wu, J., & Malmström, M. E. (2015). Nutrient loadings from urban catchments under climate
 change scenarios: Case studies in stockholm, sweden. Science of The Total Environment,
 518-519, 393-406. doi: https://doi.org/10.1016/j.scitotenv.2015.02.041
- Wu, Y., Liu, S., & Gallant, A. L. (2012). Predicting impacts of increased co2 and climate change on the water cycle and water quality in the semiarid james river basin of the midwestern usa. *Science of The Total Environment*, 430, 150-160. doi: https://doi.org/ 10.1016/j.scitotenv.2012.04.058
- Yang, Q., Zhang, X., Almendinger, J. E., Huang, M., Chen, X., Leng, G., ... Li, X. (2019). Climate change will pose challenges to water quality management in the st. croix river basin. *Environmental Pollution*, 251, 302-311. doi: https://doi.org/10.1016/ j.envpol.2019.04.129
- Ye, L., & Grimm, N. B. (2013). Modelling potential impacts of climate change on water
 and nitrate export from a mid-sized, semiarid watershed in the us southwest. *Climatic Change*, 120, 419-431. doi: 10.1007/s10584-013-0827-z
- Zheng, W., Wang, S., Tan, K., & Lei, Y. (2020). Nitrate accumulation and leaching
 potential is controlled by land-use and extreme precipitation in a headwater catchment
 in the north china plain. Science of The Total Environment, 707, 136168. doi: https://
 doi.org/10.1016/j.scitotenv.2019.136168