Data assimilation experiments inform monitoring needs for near-term ecological forecasts in a eutrophic reservoir

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Abstract

Ecosystems around the globe are experiencing increased variability due to land use and climate change. In response, ecologists are increasingly using near-term, iterative ecological forecasts to predict how ecosystems will change in the future. To date, many near-term, iterative forecasting systems have been developed using high temporal frequency (minute to hourly resolution) data streams for assimilation. However, this approach may be cost-prohibitive or impossible for forecasting ecological variables that lack high-frequency sensors or have high data latency (i.e., a delay before data are available for modeling after collection). To explore the effects of data assimilation frequency on forecast skill, we developed water temperature forecasts for a eutrophic drinking water reservoir and conducted data assimilation experiments by selectively withholding observations to examine the effect of data availability on forecast accuracy. We used in-situ sensors, manually collected data, and a calibrated water quality ecosystem model driven by forecasted weather data to generate future water temperature forecasts using FLARE (Forecasting Lake And Reservoir Ecosystems), an open-source water quality forecasting system. We tested the effect of daily, weekly, fortnightly, and monthly data assimilation on the skill of 1 to 35-day-ahead water temperature forecasts. We found that forecast skill varied depending on the season, forecast horizon, depth, and data assimilation frequency, but overall forecast performance was high, with a mean 1-day-ahead forecast root mean square error (RMSE) of 0.94°C, mean 7-day RMSE of 1.33°C, and mean 35-day RMSE of 2.15°C. Aggregated across the year, daily data assimilation yielded the most skillful forecasts at 1-7-day-ahead horizons, weekly data assimilation resulted in the most skillful forecasts at 8-35-day-ahead horizons. Within a year, daily to fortnightly data assimilation substantially outperformed monthly data assimilation in the stratified summer period, whereas all data assimilation frequencies resulted in skillful forecasts across depths in the mixed spring/autumn periods for shorter forecast horizons. Our results suggest that lower-frequency data (i.e., weekly) may be adequate for developing accurate forecasts in some applications, further enabling the development of forecasts broadly across ecosystems and ecological variables without high-frequency sensor data.

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41	Key Words: data collection frequency; FLARE; high-frequency sensors; initial conditions;
42	observations; uncertainty; water temperature
43	
44	Introduction

45 In the face of increasing ecological variability due to climate and land use change (e.g.,
46 Gilarranz et al., 2022, Malhi et al., 2020), ecological forecasting is increasingly being used for

47 understanding and predicting future ecological change (Carey et al., 2022d, Lewis et al., 2022). Here, we define ecological forecasts as predictions of future environmental conditions with 48 quantified uncertainty (see Carey et al., 2022d, Lewis et al., 2022). Applications of ecological 49 50 forecasts can improve understanding of ecosystem processes (e.g., carbon cycling, Bett et al., 51 2020), quantify predictability of environmental variables (e.g., rodent abundances, White et al., 52 2019), and inform management of ecosystem services (e.g., fisheries management, Lindegren et 53 al., 2010). Because of their broad utility, forecasts are increasingly being developed by the 54 research community to predict population, community, and ecosystem dynamics (Lewis et al., 55 2022). For example, an ongoing, community-based forecasting challenge organized by the Ecological Forecasting Initiative's Research Coordination Network has received thousands of 56 57 ecological forecast submissions of National Ecological Observatory Network (NEON) data (e.g., 58 lake water temperature, tick abundances, forest net ecosystem production, beetle communities) 59 before the data have been collected (Thomas et al., 2023a). 60 Many near-term (daily to decadal) ecological forecasts are produced using the iterative, 61 near-term forecasting cycle, in which models are updated as new observational data become 62 available to generate forecasts into the future with quantified uncertainty (Dietze et al., 2018). 63 The process of updating forecast models with newly available data, termed data assimilation 64 (DA), is a critical component of the iterative, near-term forecast cycle (Dietze et al., 2018, Luo et 65 al., 2011). DA allows for iterative updating of ecological hypotheses and models as forecasts are 66 continuously assessed and updated with the most recent ecosystem observations (Dietze et al., 2018; White et al., 2019). DA can also improve forecast accuracy by updating forecast model 67 68 initial conditions (i.e., starting values given to the model), states, and/or parameters at the

timestep that the new observations become available (Cho et al., 2020, Gottwald and Reich,

2021, Luo et al., 2011, Niu et al., 2014). For example, McClure et al. (2021) developed 1 to 2week-ahead forecasts of reservoir methane emissions with and without weekly DA and found
that the accuracy of forecasts with DA was 44 - 128% higher than forecasts without DA over a
five-month forecasting period. Despite the usefulness of DA for improving forecasts, however,
the optimal frequency of observations for updating ecological models to produce skillful
forecasts is not well characterized.

76 While there are a number of best practices proposed for applying the near-term, iterative 77 forecast cycle in ecology (e.g., Clark et al., 2001, Harris et al., 2018, Lewis et al., 2022, White et 78 al., 2019), few recommendations exist for choosing the optimal frequency of DA to produce 79 accurate forecasts. Specifically, determining the appropriate frequency of observations for DA 80 across a range of ecological variables is needed to improve the scalability of ecological 81 forecasting, particularly if accurate forecasts can be developed using lower frequency 82 observations. For example, if weekly or fortnightly DA yielded similarly accurate lake dissolved 83 oxygen forecasts as daily DA, then water quality forecasting systems could be developed for 84 lakes that have weekly or fortnightly routine monitoring program data without needing expensive 85 high-frequency sensors, thereby enabling forecasts to be generated for many waterbodies 86 globally.

Currently, many automated ecological forecasting systems rely on high-frequency
sensors to assimilate data at each time step and generate accurate forecasts (e.g., Baracchini et
al., 2020b, Corbari et al., 2019, Marj and Meijerink, 2011, Page et al., 2018, Tanut et al., 2021),
but it is possible that high-frequency sensor data collection may not be needed for DA.
Moreover, deployment of high-frequency sensors is not always feasible for all ecological
variables (e.g., zooplankton abundance, biogeochemical concentrations, Marcé et al., 2016) and

some remote locations have additional logistical constraints for maintaining autonomous sensor
operation (Steere et al., 2000). Furthermore, some remotely sensed variables may only be
available as satellite orbits and weather conditions (e.g., cloud cover) allow (e.g., Herrick et al.,
2023). Thus, identifying how best to integrate observational data collected at different temporal
frequencies into forecast models has emerged as a critical need for ecological forecasters
(LaDeau et al., 2017).

99 Studies on the frequency of DA for environmental forecasts have generally shown that 100 more temporally frequent DA improves forecast accuracy, but not always, which may be related 101 to the sensitivity of forecasts to model initial conditions. For example, DA occurring every 24 102 hours using in-situ snow data (e.g., snow depth, density, snow water equivalent) resulted in 103 better predictions of these snow variables in an alpine snowpack model compared to DA 104 occurring every 3 hours (Piazzi et al., 2018). Conversely, DA 'experiments' performed for 105 NOAA (National Oceanic and Atmospheric Administration)'s Global Forecasting System using 106 meteorological observations collected at different frequencies showed that DA occurring every 2 107 hours resulted in more accurate air temperature and wind speed forecasts compared to DA 108 occurring every 6 hours (He et al., 2020). These differences are likely because uncertainty in 109 meteorological forecasts is primarily driven by the forecast model's initial conditions. Thus, 110 more frequent DA, which constrains the model's initial conditions, will almost always improve 111 the skill of meteorological forecasts (e.g., Clark et al., 2016, He et al., 2020, Simonin et al., 112 2017). In contrast, for forecasts of environmental systems in which model process uncertainty 113 and model driver data uncertainty are more important sources of uncertainty (e.g., Dietze, 2017a, 114 Heilman et al., 2022, Lofton et al., 2022, Thomas et al., 2020), it is unknown whether more

frequent DA can improve forecast skill by generating initial conditions more consistent withobservations.

117 To the best of our knowledge, there have been only a few ecological DA experiments that 118 have tested the effects of different observation frequencies on forecast skill (e.g., Massoud et al., 119 2018, Piazzi et al., 2018, Weng and Luo, 2011, Ziliani et al., 2019), and none that have 120 considered how the frequency of data used for assimilation affects forecast skill across both 121 spatial and temporal scales. Weng & Luo (2011) assimilated eight different carbon datasets (e.g., 122 root biomass, litter fall, soil respiration), each with different collection frequencies, to identify 123 the relative importance of these data sources in constraining long-term carbon dynamics, but did 124 not consider how different frequencies of the same dataset could affect forecast skill. Piazzi et al. 125 (2018) assimilated multiple snow observations at two different frequencies (3 and 24 hours) for 126 predicting different snow-related variables (e.g., depth, density, and snow water equivalent), and 127 Ziliani et al. (2019) performed DA tests using 1-20 second assimilation of water depth data to 128 assess water level forecast skill, but neither considered the effect of less frequent assimilation 129 (e.g., >24 hours). Massoud et al. (2018) performed DA tests using a wider range of temporal 130 frequencies (e.g., ~3-34-day abundance data) to predict plankton community dynamics, but did 131 not consider the effects of DA across spatial scales (i.e., how DA affects forecast skill across 132 multiple sites or depths in an aquatic ecosystem). As a result, further work is needed to quantify 133 the utility of increased observation and DA frequency over both time and space to forecast 134 performance in ecological systems with varying sensitivities to initial conditions. 135 Among ecosystems, freshwater lakes and reservoirs are particularly important systems

for developing near-term forecasts because they provide essential ecosystem services, including
drinking water, food, irrigation, and recreation (Carpenter et al., 2011, Meyer et al., 1999,

138	Williamson et al., 2016). Because freshwaters are experiencing greater variability and adverse
139	water quality issues in response to land use and climate change (e.g., O'Reilly et al., 2015, Paerl
140	and Paul, 2012, Woolway et al., 2021), some water managers have used forecasts to
141	preemptively address poor water quality events (reviewed by Lofton et al., 2023). To date,
142	iterative, near-term freshwater forecasts have been developed for a number of water quality
143	variables, including water temperature (e.g., Carey et al., 2022d, Thomas et al., 2023b),
144	dissolved oxygen (e.g., Wang et al., 2016), and phytoplankton (e.g., Page et al., 2017, Woelmer
145	et al., 2022). These forecasts have been developed using DA with observations collected by
146	high-frequency sensors at intervals ranging from 4 minutes to 24 hours. However, most manual-
147	sampling water quality monitoring programs collect observations on weekly to fortnightly scales
148	(e.g., Francy et al., 2015, Kirchner and Neal, 2013, Romero et al., 2002), currently precluding the
149	scaling of existing forecasting systems broadly and underscoring the need to determine whether
150	less frequent observations can be used to produce accurate forecasts.
151	To quantify how DA at different frequencies affects forecast skill up to 35 days into the
152	future, we performed DA experiments in which we separately assimilated daily, weekly,
153	fortnightly, and monthly data into reservoir water temperature forecasts. Water temperature
154	forecasts are used to inform management decisions on water extraction depth and preemptive
155	water quality interventions (Georgakakos et al., 2005; Kehoe et al., 2015; Mi et al., 2020), and
156	thus our study has much utility for both informing how best to forecast complex ecosystem
157	dynamics, as well as manage drinking water supplies. Our research questions were: 1) Which
158	frequency of DA generates the most skillful water temperature forecasts? 2) How does forecast
159	skill vary across time (specifically focusing on the mixed vs. stratified seasons within a year) and
160	space (i.e., reservoir depth)? and 3) How does DA frequency influence total forecast uncertainty

and what is the relative contribution of initial condition uncertainty to total forecast uncertainty?
As previous work has suggested that reservoir water temperature forecasts can sometimes exhibit
sensitivity to initial conditions (Thomas et al. 2020), we expected that less frequent DA would
result in decreased forecast skill and increased total uncertainty. In addition, we expected that
forecast skill would be better at deeper depths, especially during thermally-stratified periods
(e.g., Mercado-Bettín et al. 2021; Thomas et al. 2020).

- 167
- 168 Methods

169 *Forecasting system overview*

We applied the Forecasting Lake And Reservoir Ecosystems (FLARE) forecasting
system (Thomas et al., 2020) to Beaverdam Reservoir, Virginia, USA to produce daily water
temperature forecasts for 1-35 days into the future (hereafter referred to as forecast horizon)
during 1 January 2021 - 31 December 2021. FLARE is an open-source forecasting system that
incorporates real-time water quality sensor data, DA, ensemble-based forecasts, and uncertainty
quantification to predict near-term water quality conditions (Thomas et al., 2020).

176 Forecast generation via FLARE can be summarized by four steps (Figure 1). First, 10-177 min resolution water temperature data were collected by sensors deployed in the reservoir 178 (Figure 1 step 1). Second, these data were transferred to the cloud and stored in a GitHub 179 repository, where they were downloaded daily and made available for DA (Figure 1 step 2). 180 Simultaneously, 1 to 35-day-ahead NOAA meteorological forecasts were downloaded daily as 181 driver data for the reservoir hydrodynamic model to generate the water temperature forecasts. 182 Third, during the forecast generation step, DA was used to update initial conditions and 183 parameters with the most recent observations using an ensemble Kalman filter, a numerical

approach that allows for the updating of model states and parameters using data (Evensen, 2003)
(Figure 1 step 3a). Following DA, the reservoir hydrodynamic model was initialized with the
updated model states and parameters to produce 1-35-day-ahead forecasts for each 0.5 m depth
interval across the water column (Figure 1 step 3b). Finally, forecast skill was assessed by
comparing observed vs. predicted water temperatures for each daily forecast at each depth
(Figure 1 step 4). We repeated steps 3a-4 for daily, weekly, fortnightly, and monthly intervals of
DA throughout the year as part of the DA experiments to compare forecast skill over time.

191

192 *Study site and monitoring*

193 Beaverdam Reservoir (BVR) is a small (0.28 km²), shallow ($Z_{max} = 11$ m), dimictic, eutrophic reservoir in southwestern Virginia, USA (37.31° N, 79.82° W; Figure 2). BVR is 194 195 managed by the Western Virginia Water Authority as a secondary drinking water supply and is 196 located in a deciduous forest catchment (Doubek et al., 2019). During a typical year, BVR is 197 stratified from mid-March to late October and mixed from November to early March (Hounshell 198 et al., 2021). BVR experiences summer hypolimnetic anoxia and cyanobacterial blooms, both of 199 which are controlled by water temperature and thermal stratification (Doubek et al., 2019; Hamre 200 et al., 2018), making forecasts of water temperature important for water quality management.

Water quality monitoring of BVR includes both manual sampling and high-frequency
sensors. From 2014-present, manual water quality sampling occurred weekly to fortnightly
during the summer stratified period and fortnightly to monthly during the remainder of the year
(Carey et al., 2022c). Starting in June 2020, high-frequency sensors were deployed in the
reservoir, enabling a range of DA frequencies to be compared in this study. We deployed
NexSens T-Node FR Temperature Sensors (NexSens Technology, Fairborn, OH, USA) at 1 m

207 intervals from the surface to sediments and a YSI EXO2 sonde (YSI Incorporated, Yellow 208 Springs, OH, USA) that monitored temperature at 1.5 m at the deepest site in BVR (Figure 1; see 209 Carey et al., 2023 for sensor information). These sensors collected data every 10 minutes, which 210 was transmitted every 3 to 9 hours via secure sensor gateways to a Git repository in the cloud 211 (Carey et al., 2023, Daneshmand et al., 2021). We removed observations collected during 212 periods of sensor maintenance, as well as depth-adjusted the data using an offset calculated from 213 a CS451 Stainless-Steel Pressure Transducer (Campbell Scientific, Logan, UT, USA) to account 214 for water level changes (Wander et al., 2023b). Because of this range in latency, or the time that 215 it takes for data to become available for modeling after they are initially collected, we used the 216 daily mean in our forecasting application. Following quality checks, these data were integrated 217 into the FLARE forecasting system to produce depth-specific daily water temperature forecasts.

218

219 Hydrodynamic model configuration

For modeling reservoir hydrodynamics, we used the General Lake Model (GLM) v.3.3.0 (Hipsey et al., 2022) to forecast water temperature in BVR. GLM is an open source, 1-D processbased hydrodynamic model commonly used within the freshwater research community to simulate water quality in lakes and reservoirs (Hipsey et al., 2019). GLM uses a Lagrangian approach for simulating different water layers and has been applied to a variety of lakes worldwide for modeling (e.g., Bruce et al., 2018, Read et al., 2014) and forecasting hydrodynamics (e.g., Thomas et al., 2020, 2023b).

We configured GLM for BVR using historical bathymetric data (Carey et al., 2022b) and water temperature observations for initial conditions (Carey et al., 2023). We configured GLM with two sediment zones to simulate epilimnetic (surface) and hypolimnetic (bottom) sediment 230 temperature dynamics following Carey et al. (2022a). GLM requires meteorological and 231 reservoir inflow observations as driver data to run the model. Because we were applying GLM 232 for forecasting, meteorological forecasts, not observed meteorology, were used as driver data in 233 the model, as described below. Additionally, we set the inflow to equal outflow in this study 234 given limited inflow data for validation and the relatively short forecast horizons (\leq 35 days). We 235 initiated the model using its default parameter set (Hipsey et al., 2019) and performed calibration 236 via a 35-day spin-up period with DA to tune parameters before the start of our focal forecasting 237 period (described below).

238

239 FLARE configuration for DA and uncertainty

240 We configured FLARE for BVR following its application to other lakes and reservoirs 241 (Thomas et al., 2020, 2023b). We set the number of forecast ensemble members to 256 to ensure 242 an adequate representation of uncertainty and prevent the ensemble Kalman filter from 243 developing erroneous correlations among ensemble members that can occur with low ensemble 244 sizes (Duc et al., 2021, Machete and Smith, 2016). While we used default values for most GLM 245 parameters, we used the ensemble Kalman filter in FLARE to tune three model parameters that 246 we identified as important for water temperature simulations using GLM in a similar, nearby 247 reservoir (Carey et al., 2022a, Thomas et al., 2020): 1) the longwave radiation scaling factor 248 (hereafter, *longwave*); 2) epilimnetic sediment temperature parameter (hereafter, *epi sed temp*); 249 and 3) hypolimnetic sediment temperature parameter (hereafter, *hypo sed temp*). 250 We used state augmentation to tune the three parameters in the ensemble Kalman filter

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251 (Thomas et al. 2020). Specifically, correlations between the parameter values and the model
252 states with observations (i.e., water temperatures at the depths with sensor observations) were

253 used to adjust parameters to be consistent with the most recent data used in DA. The three tuned 254 parameters were initially calibrated during a spin-up period from 27 November to 31 December 255 2020 and were subsequently updated via DA throughout the forecasting period. To avoid the 256 common issue of artificially low parameter uncertainty in sequential DA (Dietze, 2017a), we 257 specified the standard deviation of a normal distribution for each parameter $(1.0^{\circ}C \text{ for the})$ 258 sediment temperature parameters and 0.02 for the longwave radiation scaling factor). Initial 259 exploration of parameter fitting in this study indicated that the application of FLARE over the 260 full year resulted in low parameter uncertainty, necessitating us to specify the standard deviation 261 a priori rather than estimating it using DA. The distributions we chose were adapted from a prior 262 application of FLARE that estimated the standard deviation of parameter distributions across six 263 lakes (Thomas et al., 2023b).

264 FLARE uses a numerical ensemble-based approach to simulate and propagate forecast 265 uncertainty (Thomas et al., 2020). We represented the contribution of uncertainty from 266 meteorological driver data, initial conditions, model process, and model parameters using the 267 256-member ensemble, following Thomas et al. (2020). First, to represent the contribution of 268 meteorological driver data uncertainty, we assigned each of the 256 FLARE ensemble members 269 one of the 30 ensemble members from the 1-35-day-ahead meteorological forecasts (National 270 Oceanic and Atmospheric Administration's Global Ensemble Forecasting System) to drive GLM 271 for forecasting. Second, we represented uncertainty in the initial conditions of the forecasts using 272 the spread in model states among the 256 ensemble members on the first day of each forecast. 273 This spread was determined by either using the prior day's forecast as a starting point for the 274 next day's forecast (when no data were available for DA) or the updated states following DA 275 (when data were available for DA). We set the observation uncertainty standard deviation to

276 0.1°C, determined from the standard deviation of temperature observations and following prior 277 applications of FLARE (Thomas et al. 2020). Third, we represented model process uncertainty 278 by adding random noise to the water temperature predictions from each of the 256 FLARE 279 ensemble members at each daily time-step in a 1-35-day-ahead forecast horizon. The random 280 noise for each modeled depth within an ensemble member was drawn from a normal distribution 281 with a standard deviation of 0.75°C, as used in a previous application of FLARE that reported 282 well-calibrated forecast uncertainty (Thomas et al. 2020). The random noise was spatially 283 correlated so that it was most similar for nearby depths and most different for further-apart 284 depths. The strength of the spatial correlation was determined by the exponential decay of the 285 correlation strength with distance (Thomas et al., 2023b). Fourth, we represented parameter 286 uncertainty using the standard deviations of the distributions for the three tuned GLM parameters 287 described above. A unique parameter value drawn from each of the three distributions was 288 assigned to each of the 256 FLARE ensemble members. The parameter value assigned to an 289 ensemble member was only updated when DA occurred. Parameters not tuned by the ensemble 290 Kalman filter were assumed to have fixed values and uncertainty in these parameters was not 291 calculated.

To determine whether there was a relationship between the magnitude of initial conditions uncertainty and the sensitivity of forecast skill to more frequent DA (following Clark et al., 2016, He et al., 2020, Simonin et al., 2017), we quantified the contribution of initial conditions uncertainty to total forecast uncertainty in our DA forecasts for all DA frequencies. For this analysis, we isolated the magnitude of initial conditions uncertainty by generating the water temperature forecasts for all 365 days with and without initial conditions uncertainty and compared the variance among all 256 ensemble members. We also calculated the proportion of

initial conditions uncertainty within total forecast uncertainty for all depths, horizons, andstratified vs. mixed periods.

301

302 Data assimilation experiments

303 To quantify the effect of DA at different frequencies on forecast skill, we conducted DA 304 experiments in BVR from 1 January to 31 December 2021 (n = 365 days). As noted above, we 305 used a spin-up period from 27 November - 31 December 2020 (n = 35 days) during which DA 306 occurred, but no forecasts were generated. During the one-year forecast period in 2021, we 307 forecasted daily water temperature at 23 depths in the reservoir (spanning 0.1 to 11 m depth at 308 0.5 m intervals) and assessed forecast performance relative to observations across each forecast's 309 daily predictions for 1 to 35-day-ahead horizons and depth intervals. We focused on three focal 310 depths (1, 5, and 9 m) when reporting results, as these depths are representative of the surface, 311 middle, and bottom layers of the water column, respectively. We chose 9 m to represent the 312 bottom of the reservoir because deeper depths were not always observed due to variability in 313 water levels throughout the year (within ~ 1 m due to seasonality in flows). 314 We performed DA experiments using four different DA frequencies (daily, weekly, 315 fortnightly, and monthly) to represent different data collection latencies that are commonly used 316 by water quality monitoring programs (e.g., Engelhardt and Kirillin, 2014, Francy et al., 2015, 317 Kirchner and Neal, 2013, Liu et al., 2019, Romero et al., 2002). We assimilated water 318 temperature data across different temporal frequencies by downsampling from the high-319 frequency observations collected by our sensors. This resulted in four different temporal 320 frequencies for which DA occurred, corresponding to either daily (representing standard FLARE 321 DA), weekly, fortnightly, or monthly DA (see Data Assimilation Experiments box in Figure 1

and Appendix S1: Figure S1 for visualization of DA frequencies). For example, for the weekly
DA frequency, observations were selected every seven days starting on 4 January 2021 and
ending on 31 December 2021. In this example, DA only occurred once per week; the forecasts
that were generated on the six other days in the same week did not include DA (i.e., no DA
occurred during 5 January - 10 December 2021 even though forecasts were still generated daily
during this interval; Figure 1). Fortnightly and monthly DA occurred every 14 days and 30 days,
respectively, throughout the year.

We generated 365 daily forecasts starting on 1 January 2021 for each of the four DA frequencies. While we recognize that we are producing hindcasts for a historical period, because the model was forced with only forecasted drivers and out-of-sample forecast evaluation occurred, we refer to these retrospective forecasts or hindcasts as forecasts throughout for consistency (following Jolliffe and Stephenson, 2012).

334

335 Analysis

336 *Question 1:* For all n = 1460 forecasts produced (365 forecasts generated daily over a 337 year for four different DA frequencies), we used root mean square error (RMSE) and continuous 338 ranked probability score (CRPS; Gneiting et al., 2005) to quantify forecast skill. We defined 339 skillful water temperature forecasts as those with an $RMSE < 2^{\circ}C$, a commonly-used threshold 340 for lake and reservoir hydrodynamic modeling following Bruce et al. (2018), Read et al. (2014), 341 and many others. Mean full water column RMSE was calculated for each of the 35 days across 342 all forecast horizons for each DA frequency regardless of whether data were assimilated the day 343 the forecast was generated. We aggregated RMSE across depths and dates to determine the 344 lowest temporal frequency of DA required to generate the most skillful water temperature

forecasts across the full water column and throughout the entire forecast period. We focus onRMSE in the results and all CRPS values are reported in the SI.

347 Question 2: Using RMSE and CRPS, we compared forecast skill across depths and 348 seasons to identify how the frequency of DA affected forecast accuracy over space and time. To 349 quantify spatial forecast performance, we calculated RMSE and CRPS for each depth (1-11 m) at 350 each forecast horizon (1-35 days ahead) and DA frequency in BVR. To quantify temporal 351 forecast performance, we compared forecast skill at each horizon aggregated within thermally-352 stratified vs. mixed periods in BVR. The stratified period began on the first day that the water 353 density difference between the reservoir surface (0.1 m) and the maximum depth observed for 354 the reservoir on each day (e.g., between 9-11 m) was $\ge 0.1 \text{ kg/m}^3$ for at least three consecutive 355 days (following Ladwig et al., 2021). Conversely, the mixed period began on the first day that 356 surface and bottom water density differences were $< 0.1 \text{ kg/m}^3$ for at least three consecutive 357 days. Altogether, we compared forecast skill between stratified vs. mixed periods; among depths 358 (1, 5, and 9 m), and among forecast horizons (focusing in on 1, 7, and 35-day-ahead forecasts) 359 for each of the four DA frequencies.

Question 3: We quantified total forecast uncertainty for each day in the 1-35-day forecast horizon using the variance of the 256-member FLARE ensemble. The relative contribution of initial condition uncertainty to total forecast uncertainty was calculated for each forecasted day by comparing the variance in the 256-member FLARE ensemble between the set of forecasts with initial condition uncertainty included and the set without initial condition uncertainty.

All statistical analyses were conducted in R v.4.2.0 (R Core Team, 2022). All R code and data files used to run these analyses are archived and available in the Zenodo repository (Wander et al., 2023a, 2023b).

368 **Results:**

369 *BVR water temperature dynamics*

BVR exhibited typical annual water temperature dynamics during the forecasting period 370 371 in 2021. Water temperature throughout the water column ranged from 1.4 to 29.9°C during the 372 year. The summer stratified period began on 12 March and ended on 7 November 2021, and the 373 reservoir was mixed from 1 January - 11 March and 8 November - 31 December (Figure 3). 374 Thermocline deepening occurred throughout the summer stratified period, starting at 1.5 m in 375 March with stratification onset and deepening to 9.5 m in November before fall turnover (Figure 376 3). During the winter, there were three brief periods of ice cover of one to three days in duration 377 in January and February when inverse stratification occurred (Figure 3; Carey and Breef-Pilz, 378 2022). We removed these few ice-cover days from the analysis and grouped mixed (n = 118) 379 days) vs. summer stratified data (n = 241 days) for analysis. 380

381 *Data assimilation frequency altered forecast output and parameters over time*

We were able to successfully forecast water temperature throughout the water column over the year using DA to update model states and parameters (Figures 4-5). Across all depths, DA constrained uncertainty by updating initial conditions with the most recent water temperature observations. Forecast uncertainty for the lower DA frequencies was strongly dependent on the time since last assimilation (Figure 4). On average, forecast variance at the one-day horizon across 2021 for forecasts with daily DA was 1.56°C while mean forecast variance at the one-day horizon for forecasts with monthly DA was 3.25°C.

We observed that DA frequency altered parameter evolution of the forecasts (Figure 5).
The daily DA frequency resulted in more variable parameter estimates through time for all three

391 tuned parameters, reflecting the more frequent adjustment that occurred each time data were 392 assimilated. Importantly, parameter evolution for forecasts with daily DA yielded very different 393 estimates than the weekly, fortnightly, and monthly DA forecast frequencies (Figure 5). For 394 example, the evolution of the longwave radiation scaling parameter (longwave) over the 365-day 395 forecast period showed that forecasts with weekly, fortnightly, and monthly DA frequencies 396 converged at ~ 0.91 by December 2021, whereas the longwave parameter for forecasts with daily 397 DA was at ~ 0.85 by the end of the year (Figure 5a). Similarly, the parameter controlling the 398 surface layer sediment temperature (epi sed temp) in daily DA forecasts began to diverge from 399 the other DA frequencies in early April (Figure 5c). The non-daily DA frequencies (i.e., weekly, 400 fortnightly, monthly DA) surface sediment layer temperature parameter (epi sed temp) values 401 ranged from 13.51 to 15.62°C, whereas the daily DA frequency epi sed temp values ranged 402 from 13.29°C to 17.0°C during April-December. For the parameter controlling the bottom layer 403 sediment (hypo sed temp), daily DA forecasts exhibited much more variable values (ranging 404 from 10.24°C to 11.21°C) than forecasts for any other DA frequency (range 10.65°C to 10.72°C; 405 Fig 5b) from April to December.

406

407 *Question 1: Which frequency of data assimilation generates the most skillful water temperature*408 *forecasts?*

Aggregated among depths and time periods, weekly DA resulted in the most skillful water temperature forecasts of the four DA frequencies for the greatest number of 1-35-dayahead horizons (Figure 6). Among horizons, we observed that the frequency of DA needed to produce skillful forecasts varied (Figure 6). At shorter horizons (1-7 days ahead), daily DA resulted in the most skilled forecasts, but at longer horizons (8-35 days ahead), weekly DA
resulted in the most skilled forecasts (Figure 6).

415 The skill of all forecasts degraded as the forecast horizon increased, but the decrease in 416 performance was greatest for daily DA forecasts, such that forecasts generated using monthly, 417 fortnightly, and weekly DA all outperformed daily DA forecasts by the 19-day forecast horizon 418 (Figure 6), when aggregating across all depths and time periods. The daily DA forecasts 419 exceeded the 2°C RMSE metric of skill on the 28-day-ahead horizon, whereas the weekly, 420 fortnightly, and monthly forecasts never exceeded that metric for any of the 1-35-day-ahead 421 horizons. These results were consistent across forecast evaluation metrics, including the CRPS 422 metric that evaluates the full ensemble forecast (Appendix S1: Figure S2). 423 424 *Question 2: How does forecast skill vary across time and space?* 425 Aggregated across depths, horizons, and DA frequencies over the year, forecast skill 426 overall was high, with a mean water temperature forecast RMSE of 1.53 ± 1.86 °C (1 S.D.).

427 Forecast skill was generally best at 9 m regardless of horizon or DA frequency. Aggregated 9 m

forecast skill was $1.29 \pm 1.80^{\circ}$ C, followed by aggregated 5 m forecast skill ($1.63 \pm 1.85^{\circ}$ C), and

429 aggregated 1 m forecast skill ($1.69 \pm 1.82^{\circ}$ C). As expected, forecast skill generally decreased

430 with horizon, with a mean 1-day-ahead forecast RMSE of $0.80 \pm 1.20^{\circ}$ C, mean 7-day RMSE of

431 $1.15 \pm 1.60^{\circ}$ C, and mean 35-day RMSE of $1.99 \pm 2.17^{\circ}$ C. However, we observed an exception to

this pattern for 1 m mixed forecasts, which is further described below.

428

433 On average, forecast skill was slightly better (as indicated by smaller RMSE) during the 434 stratified period than during the mixed period, aggregated among all depths and horizon 435 regardless of DA frequency (aggregated mixed RMSE = 1.59 ± 1.57 °C, stratified RMSE = 1.46

436	\pm 2.13°C; Figure 7). Forecast skill was more variable among forecast horizons than depths in the
437	mixed period, whereas forecast skill was variable across both depths and horizons in the
438	stratified period (Figure 7). In the stratified period, forecast skill was best at 9 m, with relatively
439	similar skill over the forecast horizon (Figure 7f). In the mixed period, forecast skill varied very
440	little among depths aggregated across horizons (Figure 7a, c, e), with consistently greater
441	decreases in skill with increasing horizon than in the stratified period, except for at 1 m. Forecast
442	skill at 1 m decreased rapidly until ~the 19-day horizon, after which forecast skill remained
443	constant for the daily DA and increased for the weekly, fortnightly, and monthly DA frequencies
444	until the end of the forecast horizon (Figure 7a).
445	While daily DA always resulted in the best forecast skill for 1-day-ahead horizons, lower
446	frequency DA typically outperformed daily DA as the forecast horizon increased. An exception
447	was for 9 m stratified forecasts, when daily DA resulted in the lowest RMSE for all forecast
448	horizons and never exceeded 1.51°C for the duration of the 35-day forecast horizon (Figure 7f).
449	Additionally, 9 m stratified forecasts were the only forecasts where skillful (RMSE $< 2^{\circ}$ C)
450	forecasts were produced for all DA frequencies and horizons (Figure 7f).
451	
452	Question 3. How does DA frequency influence total forecast uncertainty and what is the relative
453	contribution of initial condition uncertainty to total forecast uncertainty?
454	Lower frequency DA forecasts consistently had more total uncertainty (Figure 8). We
455	found that the differences between uncertainty for daily and monthly DA were largest at 1-day-
456	ahead horizons and largely converged by the end of the 35-day horizon (Figure 8). At 1 m depth,
457	total uncertainty was similar between the mixed and stratified periods across the 35-day horizon,
458	but at 5 and 9 m, total uncertainty was on average higher in the stratified than mixed period. Both

RMSE and total variance were similar for forecasts run with and without initial conditionsuncertainty included (Appendix S1: Figures S3-S4).

Forecasts with less frequent DA had a greater contribution of initial condition uncertainty 461 462 to total forecast uncertainty during the first few days of the forecast horizon. However, overall, 463 initial conditions uncertainty contributed a minimal proportion of the total uncertainty for 464 forecasts generated with daily DA (Figure 9). At the 1-day-ahead forecast horizon, daily DA 465 initial conditions uncertainty was 0% of total uncertainty, whereas initial conditions uncertainty 466 contributed 55 - 71% of total forecast uncertainty in forecasts for all other DA frequencies 467 (Figure 9). The role of initial conditions uncertainty for all depths in the mixed period and 468 surface forecasts in the stratified period was minimal (< 1%) across all DA frequencies after the 469 10-day horizon (Figure 9a-c, e). Conversely, initial conditions uncertainty made up a larger 470 proportion of total forecast uncertainty for stratified 5 m and 9 m forecasts for forecast horizons 471 between 10 and 20 days (ranging from 5-10%; Figure 9d, f).

472

473 Discussion:

474 Across a year of water temperature forecasts in our focal reservoir, we found that weekly 475 DA generally resulted in the most skillful water temperature forecasts. However, skill varied 476 among depths, forecast horizons, and time of year, suggesting that DA frequency should be 477 chosen based on the specific forecast application. For example, if water temperature forecasts are 478 specifically needed to guide decision-making that involves the deeper reservoir layers (e.g., 5 m 479 or 9 m) at short time horizons (e.g., <5 days ahead), daily DA might be most advantageous 480 (Figure 7). Conversely, if water temperature forecasts are needed for the surface water at 20-35 481 day-ahead horizons, then weekly to monthly DA may be sufficient (Figure 7). Despite the

usefulness of DA for improving forecast skill, more frequent DA did not always lead to more
skillful water temperature forecasts, in part because initial conditions uncertainty only comprised
a significant proportion of total forecast uncertainty within the first few days of the forecast
horizon (Figure 9). Below, we interpret our results for each research question and make
recommendations for considering which DA frequency might be appropriate for different
ecological forecast applications.

488

489 *Q1: Which frequency of data assimilation generates the most skillful water temperature*490 *forecasts?*

491 In this study we found that less frequent DA (e.g., weekly, fortnightly, and monthly DA) 492 sometimes led to more skillful water temperature forecasts than daily DA for all depths during 493 the mixed period. This pattern of weekly DA outperforming daily DA forecast skill during the 494 mixed period is likely because daily DA led to parameter overfitting, as indicated by the greater 495 short-term variability in parameter estimates over time (Figure 5). Because water temperatures 496 are fairly stable at deeper depths, and thus daily observations can consistently predict tomorrow's 497 water temperature accurately, parameter overfitting was less problematic for daily DA at 498 hypolimnetic depths (Figure 7f). As a result, hypolimnetic forecast skill was best with daily DA 499 during stratified conditions, but this pattern did not extend to other depths or the mixed period 500 (Figure 7).

501 Our work is consistent with studies that have found that the optimal DA frequency often 502 matches that of the forecast model timestep (e.g., Derot et al., 2020, Woelmer et al., 2022). For 503 example, during both the mixed and stratified periods, daily DA was always better for 1-day-504 ahead forecasts, but was often outperformed by weekly DA at 8-day-ahead forecast horizons

505 (Figure 6). Because water temperatures were homogenous among all depths during the mixed 506 period, water temperature variability among all depths was likely driven by air temperature 507 variability, ultimately making it more challenging to predict water temperature across depths as 508 the forecast horizon increased. During the stratified period, however, less frequent DA could still 509 generate accurate surface and mid-depth water temperature forecasts. The increased importance 510 of daily DA at bottom depths during the stratified period is likely because of the increased 511 thermal stability at bottom depths associated with thermal stratification (Figure 7). This pattern is 512 in contrast with other water temperature forecasting studies that have found daily DA necessary 513 for improving the skill of forecasts in the middle of the water column around the thermocline 514 (Baracchini et al., 2020a), but is likely explained by the overfitting of both the daily longwave 515 radiation and the epilimnetic sediment temperature parameters (Figure 5).

516 We note that there are many ways to quantify skill beyond the 2°C RMSE threshold used 517 here. We chose to use RMSE because it is a commonly used metric by lake modelers to 518 determine the deviation between observed vs. modeled values (Bruce et al., 2018, Read et al., 519 2014). However, forecast skill could also be quantified by other metrics, such as the correlation 520 coefficient, Nash-Sutcliffe model efficiency coefficient, percent relative error, normalized mean 521 absolute error, or others (Bennett et al., 2013). While 2°C RMSE is a subjective criterion of 522 forecast skill, we note that CRPS results followed similar patterns as our RMSE metric 523 (Appendix S1: Figure S2), further supporting our results when evaluating the full distribution of 524 the forecasts.

525

526 *Q2: How does forecast skill vary across time and space?*

527 Our use of a 35-day forecast horizon allowed us to compare water temperature

528 predictability across multiple horizons at different depths and times of year, thereby elucidating 529 patterns in ecosystem predictability across both space and time. We generally observed expected 530 declines in forecast skill with increasing horizon, as noted in many other studies. However, 1-m 531 forecast skill during the mixed period increased with lower frequency DA (weekly, fortnightly, 532 and monthly), while forecast skill with daily DA leveled off at 19-35-day horizons (Figure 7a). 533 Improved forecast skill for lower frequency DA suggests that our forecasts are capturing surface 534 water temperature dynamics in the mixed period at longer horizons better than other depths, 535 particularly those during the stratified period. This may be due to the smaller range in water 536 temperature variation that occurs in the mixed period relative to the stratified period over a 35-537 day interval, allowing variance to level off at longer forecast horizons as water temperature 538 observations better matched predicted values (Figures 7-8). While improved forecast skill at 539 longer horizons has been observed in the literature (e.g., Wheeler et al., 2023), this pattern is 540 often associated with variables that have predictable, cyclical patterns at long horizons (e.g., 541 annual tree leaf-out).

542 Overall, we observed generally high forecast skill across all depths and times of year for 543 most forecast horizons. Across DA frequencies, depths, and times of year, RMSE was only 544 consistently above the 2°C threshold for daily DA at 28-35-day horizons (Figure 6). By the end 545 of the 35-day forecast horizon, daily DA forecast skill for most depths and times of the year was 546 >2°C, except 9 m stratified forecasts, which had a mean RMSE of 1.29 ± 1.8 °C across DA 547 frequencies. The higher forecast skill at 9 m is likely because fluctuations in bottom water 548 temperatures were minimal during stratification (Figure 3). 549 Our findings are similar to other water temperature lake and reservoir forecasting studies.

549 Our findings are similar to other water temperature lake and reservoir forecasting studies.
550 First, the pattern of increased forecast skill in the bottom waters is consistent with Mercado-

551 Bettin et al. (2021) and Thomas et al. (2020), who both found that the bottom water forecasts 552 were more skillful than surface water forecasts. This is likely because bottom waters are not 553 changing as much as surface waters throughout the year due to less atmospheric exchange. 554 However, Clayer et al. (2023) found that surface water temperatures were more accurately 555 simulated than bottom water temperatures, suggesting that the complex lake characteristics that 556 control bottom water temperatures were not captured as well as the air temperature dynamics 557 controlling surface water temperatures. Second, our finding that forecast skill was greater in the 558 stratified period rather than mixed period is similar to the results of Thomas et al. (2020), likely 559 due to the fact that water temperature dynamics were changing less among depths in stratified 560 than mixed periods (Figure 3). Because of the variability in water temperature dynamics among 561 seasons and depths, determining the conditions in which we can most accurately forecast water 562 temperature can improve our understanding of ecosystem processes and functioning. Moreover, 563 accurately forecasting water temperature is critical for forecasting additional lake and reservoir 564 variables that are strongly driven by water temperature, such as phytoplankton biomass, 565 dissolved oxygen concentrations, and greenhouse gas emissions (e.g., McClure et al. 2021).

566

567 *Q3:* How does DA frequency influence forecast uncertainty?

We found that initial conditions uncertainty contributed a substantial proportion of total uncertainty for weekly, fortnightly, and monthly DA, but only during the first few days of the forecast horizon. From 6-23 day-ahead horizons, the contribution of initial conditions decreased to <1% across all DA frequencies, depths, and times of year (Figure 9). We observed that highfrequency DA was required for skillful 9 m stratified forecasts, while weekly DA was sufficient for other depths and times. This finding may be because the contribution of initial conditions

574 uncertainty decreases more rapidly within the first few days of the forecast horizon for the daily 575 DA forecasts at 9 m in the stratified period. For all other depths and times of the year, the rate at 576 which initial conditions uncertainty decreases is greater for weekly, fortnightly, and monthly 577 DA, resulting in more similar performance of daily and weekly DA early in the forecast horizon (Figure 9). However, more frequent DA may not always improve forecast performance, 578 579 especially when initial conditions uncertainty is not the dominant source of uncertainty, as seen 580 at longer horizons. Given that initial conditions uncertainty predominated at the beginning of the 581 forecast horizon, it is likely that total forecast uncertainty at longer horizons was primarily 582 influenced by uncertainty in model process, model parameters, and/or meteorological driver data 583 (Figure 9). Conversely, the dominant source of uncertainty for weather forecasting is typically 584 initial conditions uncertainty given the inherent instability of atmospheric processes (Dietze, 585 2017b), which is why more frequent DA often substantially improves meteorological forecast 586 skill.

587 Other lake and reservoir water quality forecasting studies have found that model driver 588 data and process uncertainty were the dominant sources of total forecast uncertainty (Lofton et 589 al., 2022, McClure et al., 2021, Thomas et al., 2020). Therefore, constraining other sources of 590 uncertainty by using an ensemble approach or different forecasting models would likely further 591 improve water temperature forecast skill. Additionally, using a different DA technique that uses 592 a Bayesian approach to estimate a posterior distribution, rather than assuming that the parameters 593 and model states are normally distributed, may also reduce uncertainty (e.g., particle filter; Wang 594 et al., 2023). Because the dominant source of uncertainty in ecological forecasts will likely differ 595 depending on the variable being forecasted, different DA techniques may not improve forecast 596 skill equally among all ecological variables.

597

598 *Recommendations for setting up DA for other forecasting systems*

599 Determining whether an ecological forecasting application requires high-frequency 600 sensors is necessary for increasing the scalability of ecological forecasting across ecosystems and 601 variables. While high-frequency sensor data may improve forecast skill in some cases, sensor 602 deployment is often costly, which limits the application of high-frequency data in some 603 forecasting systems. Moreover, even if high-frequency sensors are deployed, identifying the 604 minimum frequency of data required to make skillful ecological forecasts can be a useful 605 exercise because high-frequency sensors malfunction and require maintenance, which can result 606 in data gaps (e.g., Herrick et al., 2023). Many water quality forecasting applications to date have 607 relied on high-frequency sensor data for assimilation to produce skillful forecasts of different 608 variables (Cho and Park, 2019, Derot et al., 2020, Page et al., 2018). In this study, we found that 609 daily DA only produced the most skillful 9 m stratified period water temperature forecasts, 610 whereas weekly DA generally produced the most skillful surface and middle layer water 611 temperature forecasts (Figure 7). Our findings indicate that high-frequency sensors may not be 612 needed for accurate mixed period water temperature forecasts or surface layer forecasts in the 613 stratified period.

The minimum frequency of DA needed to set up fully operational forecasting systems is likely to vary based on the ecosystem or forecast variable of interest. Depending on the water quality forecast application, different frequencies of data collection may be necessary to fully understand and predict water quality dynamics over time. For example, George and Hurley (2004) found that fortnightly observations were required to discern gradual trends in phytoplankton productivity, but monthly data were adequate for capturing declines in

620 phytoplankton biomass over a 30-year period. Despite many successful applications of high-621 frequency DA in the literature for forecasting (e.g., Cho et al., 2020, Gottwald and Reich, 2021, 622 Luo et al., 2011, Niu et al., 2014), not all ecological variables benefit from frequent DA, as not 623 all variables are similarly forecastable. 624 In addition to the frequency of data collection, data latency can also affect the frequency 625 of DA. Even for forecasting systems with high-frequency sensor data, data latency may reduce 626 forecast skill if data are not immediately transmitted to forecasting workflows (e.g., they require 627 a manual download) (Dietze et al., 2018). In cases with high data latency of the forecast variable 628 (e.g., microscope counts of phytoplankton requiring laboratory analysis), data fusion approaches 629 that assimilate multiple data sources may improve forecast skill (e.g., Baracchini et al., 2020b, 630 Chen et al., 2021). For example, some studies have assimilated both in-situ measurements and 631 remote sensing data to forecast reservoir water quality variables, including chlorophyll a and 632 conductivity (Abdul Wahid and Arunbabu, 2022, Chen et al., 2021). 633 Finally, understanding the contributions of different sources of uncertainty can be useful 634 for determining the DA frequency that generates the most skillful forecasts. Specifically, 635 knowing the relative contribution of initial conditions uncertainty can inform sampling frequency 636 needed to improve ecological forecast skill. For forecasts with total uncertainty dominated by 637 process, parameter, or driver uncertainty, improving forecast skill may require modifying 638 processes used for forecasting the ecological variable of interest, further constraining parameters 639 by collecting more data, or improving weather forecast driver data (e.g., Grönquist et al., 2021). 640 641

643 Study Limitations

644 Our results suggest that weekly DA may suffice for some lake and reservoir water 645 temperature forecasting applications, with the caveat that more frequent DA often improved 646 water temperature forecast performance at short forecast horizons. However, we only assessed 647 forecast skill for a single reservoir and ecological variable for only one year, and therefore note 648 the limitations of extending these results to other systems and variables. Additionally, updating 649 model parameters and initial conditions too regularly can lead to overprediction biases when 650 forecasting, which may explain why weekly rather than daily DA resulted in more skillful water 651 temperature forecasts in the mixed period (see Lin et al., 2021). Finally, because we did not 652 quantify the contribution of all sources of uncertainty, we can only identify the relative role that 653 DA has on reducing initial conditions uncertainty. Future studies that consider the role of other 654 sources of uncertainty will improve our understanding of DA on total forecast uncertainty.

655

656 *Conclusions*

657 This study emphasizes the importance of DA for improving ecological forecast skill and 658 has implications for forecasting efforts among a wide range of ecosystems and ecological 659 variables. We argue that weekly observations of water temperature are likely "good enough" to 660 set up a skillful forecasting system for many reservoir management applications, while daily DA 661 would be most useful for applications requiring high forecast accuracy in the bottom waters or at 662 short (< 5 - 7 day) forecast horizons. Because water temperature dynamics control many 663 biological, chemical, and physical lake processes (Magnuson et al., 1979, Read et al., 2019, 664 Yvon-Durocher et al., 2012), water temperature must be accurately forecasted before we can 665 forecast other water quality variables. Therefore, determining ways to improve water

temperature forecasts will have broad utility for advancing the development of many additionalwater quality forecasting systems.

668	Because near-term, iterative forecasts are particularly well suited to address ecological
669	questions (Carey et al., 2022d, Dietze et al., 2018, White et al., 2019), determining how best to
670	design and deploy ecological near-term, iterative forecasting systems is a pressing need (Diez et
671	al., 2012, Ibáñez et al., 2013, Moustahfid et al., 2021). With the increasing deployment of high-
672	frequency sensor networks (e.g., National Ecological Observatory Network (NEON) and Global
673	Lake Ecological Observatory Network (GLEON); Mantovani et al., 2020, Marcé et al., 2016,
674	Park et al., 2020) comes a growing need to understand how best to use these sensor data for
675	forecasting. In response, we advocate for using DA experiments across ecosystems and
676	ecological variables to determine how best to integrate observational data into iterative
677	forecasting systems.
678	
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680	
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682	are publicly available on Zenodo (https://doi.org/10.5281/zenodo.7951402;
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694	oversaw sensor data collection. HLW wrote the initial draft of the manuscript with CCC and
695	MEL; all coauthors reviewed the manuscript and approved its final version.
696	
697	References
698	Abdul Wahid, A. and E. Arunbabu (2022) Forecasting water quality using seasonal ARIMA
699	model by integrating in-situ measurements and remote sensing techniques in Krishnagiri
700	reservoir, India. Water Pract. Technol., 17, 1230-52.
701	Baracchini, T., P. Y. Chu, J. Šukys, G. Lieberherr, S. Wunderle, A. Wüest, and D. Bouffard
702	(2020a) Data assimilation of in situ and satellite remote sensing data to 3D hydrodynamic
703	lake models: a case study using Delft3D-FLOW v4.03 and OpenDA v2.4. Geosci. Model
704	<i>Dev.</i> , 13 , 1267–84.
705	Baracchini, T., A. Wüest, and D. Bouffard (2020b) Meteolakes: An operational online three-
706	dimensional forecasting platform for lake hydrodynamics. Water Res., 172, 115529.
707	Bennett, N. D., B. F. Croke, G. Guariso, J. H. Guillaume, S. H. Hamilton, A. J. Jakeman, S.
708	Marsili-Libelli, L. T. Newham, et al. (2013) Characterising performance of
709	environmental models. Environ. Model. Softw., 40, 1–20.
710	Bett, P. E., K. E. Williams, C. Burton, A. A. Scaife, A. J. Wiltshire, and R. Gilham (2020)

- 711 Skillful seasonal prediction of key carbon cycle components: NPP and fire risk. *Environ*.
 712 *Res. Commun.*, 2, 055002.
- 713 Bruce, L. C., M. A. Frassl, G. B. Arhonditsis, G. Gal, D. P. Hamilton, P. C. Hanson, A. L.
- Hetherington, J. M. Melack, *et al.* (2018) A multi-lake comparative analysis of the
- General Lake Model (GLM): Stress-testing across a global observatory network. *Environ*. *Model. Softw.*, **102**, 274–91.
- 717 Carey, C. C. and A. Breef-Pilz (2022) Ice cover data for Falling Creek Reservoir and Beaverdam
- 718 Reservoir, Vinton, Virginia, USA for 2013-2022. *Environmental Data Initiative*
- 719 *repository*. DOI: 10.6073/pasta/917b3947d91470eecf979e9297ed4d2e
- 720 Carey, C. C., A. Breef-Pilz, B. J. Bookout, R. P. McClure, and J. H. Wynne (2023) Time series
- of high-frequency sensor data measuring water temperature, dissolved oxygen,
- conductivity, specific conductance, total dissolved solids, chlorophyll a, phycocyanin,
- fluorescent dissolved organic matter, and turbidity at discrete depths in Beaverdam
- Reservoir, Virginia, USA in 2016-2022. *Environmental Data Initiative repository*. DOI:
- 725 10.6073/pasta/4182de376fde52e15d493fdd9f26d0c7
- 726 Carey, C. C., P. C. Hanson, R. Q. Thomas, A. B. Gerling, A. G. Hounshell, A. S. L. Lewis, M. E.
- T27 Lofton, R. P. McClure, *et al.* (2022a) Anoxia decreases the magnitude of the carbon,
- nitrogen, and phosphorus sink in freshwaters. *Glob. Change Biol.*, **28**, 4861–81.
- 729 Carey, C. C., A. S. L. Lewis, D. W. Howard, W. M. Woelmer, P. A. Gantzer, K. A. Bierlein, J.
- 730 C. Little, and WVWA (2022b) Bathymetry and watershed area for Falling Creek
- 731 Reservoir, Beaverdam Reservoir, and Carvins Cove Reservoir. *Environmental Data*
- 732 *Initiative repository*. DOI: 10.6073/pasta/352735344150f7e77d2bc18b69a22412
- 733 Carey, C. C., A. S. Lewis, R. P. McClure, A. B. Gerling, A. Breef-Pilz, and A. Das (2022c) Time

734	series of high-frequency profiles of depth, temperature, dissolved oxygen, conductivity,
735	specific conductance, chlorophyll a, turbidity, pH, oxidation-reduction potential,
736	photosynthetic active radiation, and descent rate for Beaverdam Reservoir, Carvins Cove
737	Reservoir, Falling Creek Reservoir, Gatewood Reservoir, and Spring Hollow Reservoir
738	in Southwestern Virginia, USA 2013-2021. Environmental Data Initiative repository.
739	DOI: 10.6073/pasta/c4c45b5b10b4cb4cd4b5e613c3effbd0
740	Carey, C. C., W. M. Woelmer, M. E. Lofton, R. J. Figueiredo, B. J. Bookout, R. S. Corrigan, V.
741	Daneshmand, A. G. Hounshell, et al. (2022d) Advancing lake and reservoir water quality
742	management with near-term, iterative ecological forecasting. Inland Waters, 12, 107-20.
743	Carpenter, S. R., E. H. Stanley, and M. J. Vander Zanden (2011) State of the world's freshwater
744	ecosystems: physical, chemical, and biological changes. Annu. Rev. Environ. Resour., 36,
745	75–99.
746	Chen, C., Q. Chen, G. Li, M. He, J. Dong, H. Yan, Z. Wang, and Z. Duan (2021) A novel multi-
747	source data fusion method based on Bayesian inference for accurate estimation of
748	chlorophyll-a concentration over eutrophic lakes. Environ. Model. Softw., 141, 105057.
749	Cho, H. and H. Park (2019) Merged-LSTM and multistep prediction of daily chlorophyll-a
750	concentration for algal bloom forecast. IOP Conf. Ser. Earth Environ. Sci., 351, 012020.
751	Cho, K. H., Y. Pachepsky, M. Ligaray, Y. Kwon, and K. H. Kim (2020) Data assimilation in
752	surface water quality modeling: A review. Water Res., 186, 116307.
753	Clark, J. S., S. R. Carpenter, M. Barber, S. Collins, A. Dobson, J. A. Foley, D. M. Lodge, M.
754	Pascual, et al. (2001) Ecological forecasts: an emerging imperative. science, 293, 657-
755	60.
756	

757	permitting models: a step-change in rainfall forecasting. Meteorol. Appl., 23, 165-81.
758	Clayer, F., L. Jackson-Blake, D. Mercado-Bettín, M. Shikhani, A. French, T. Moore, J. Sample,
759	M. Norling, et al. (2023) Sources of skill in lake temperature, discharge and ice-off
760	seasonal forecasting tools. Hydrol. Earth Syst. Sci., 27, 1361-81.
761	Corbari, C., R. Salerno, A. Ceppi, V. Telesca, and M. Mancini (2019) Smart irrigation forecast
762	using satellite LANDSAT data and meteo-hydrological modeling. Agric. Water Manag.,
763	212 , 283–94.
764	Daneshmand, V., A. Breef-Pilz, C. C. Carey, Y. Jin, YJ. Ku, K. C. Subratie, R. Q. Thomas, and
765	R. J. Figueiredo (2021) Edge-to-cloud Virtualized Cyberinfrastructure for Near Real-time
766	Water Quality Forecasting in Lakes and Reservoirs. 2021 IEEE 17th International
767	Conference on eScience (eScience). IEEE, pp. 138–48.
768	DeChant, C. M. and H. Moradkhani (2011) Improving the characterization of initial condition
769	for ensemble streamflow prediction using data assimilation. Hydrol. Earth Syst. Sci., 15,
770	3399–3410.
771	Derot, J., H. Yajima, and F. G. Schmitt (2020) Benefits of machine learning and sampling
772	frequency on phytoplankton bloom forecasts in coastal areas. Ecol. Inform., 60, 101174.
773	Dietze, M. C. (2017a) Ecological forecasting. Princeton University Press.
774	Dietze, M. C. (2017b) Prediction in ecology: a first-principles framework. Ecol. Appl., 27, 2048–
775	60.
776	Dietze, M. C., A. Fox, L. M. Beck-Johnson, J. L. Betancourt, M. B. Hooten, C. S. Jarnevich, T.
777	H. Keitt, M. A. Kenney, et al. (2018) Iterative near-term ecological forecasting: Needs,
778	opportunities, and challenges. Proc. Natl. Acad. Sci., 115, 1424-32.
779	Diez, J. M., I. Ibáñez, A. J. Miller-Rushing, S. J. Mazer, T. M. Crimmins, M. A. Crimmins, C. D.

780	Bertelsen, and D. W. Inouye (2012) Forecasting phenology: from species variability to
781	community patterns. Ecol. Lett., 15, 545–53.
782	Doubek, J. P., K. L. Campbell, M. E. Lofton, R. P. McClure, and C. C. Carey (2019)
783	Hypolimnetic hypoxia increases the biomass variability and compositional variability of
784	crustacean zooplankton communities. Water, 11, 2179.
785	Duc, L., K. Saito, and D. Hotta (2021) Analysis and design of covariance inflation methods
786	using inflation functions. Part 2: adaptive inflation. Q. J. R. Meteorol. Soc., 147, 2375-
787	94.
788	Engelhardt, C. and G. Kirillin (2014) Criteria for the onset and breakup of summer lake
789	stratification based on routine temperature measurements. Fundam. Appl. Limnol., 184,
790	183–94.
791	Evensen, G. (2003) The Ensemble Kalman Filter: theoretical formulation and practical
792	implementation. Ocean Dyn., 53, 343–67.
793	Francy, D. S., J. L. Graham, E. A. Stelzer, C. D. Ecker, A. M. Brady, P. Struffolino, and K. A.
794	Loftin (2015) Water quality, cyanobacteria, and environmental factors and their relations
795	to microcystin concentrations for use in predictive models at ohio lake erie and inland
796	lake recreational sites, 2013-14. US Geological Survey Scientific Investigations Report
797	2015–5120, 58 p.
798	Georgakakos, K. P., N. E. Graham, T. M. Carpenter, and H. Yao (2005) Integrating climate-
799	hydrology forecasts and multi-objective reservoir management for northern California.
800	Eos Trans. Am. Geophys. Union, 86, 122–27.
801	George, D. G. and M. A. Hurley (2004) The influence of sampling frequency on the detection of
802	long-term change in three lakes in the English Lake District. Aquat. Ecosyst. Health

803 *Manag.*, 7, 1–14.

- Gilarranz, L. J., A. Narwani, D. Odermatt, R. Siber, and V. Dakos (2022) Regime shifts, trends,
 and variability of lake productivity at a global scale. *Proc. Natl. Acad. Sci.*, 119,
 e2116413119.
- Gneiting, T., A. E. Raftery, A. H. Westveld, and T. Goldman (2005) Calibrated probabilistic
 forecasting using ensemble model output statistics and minimum CRPS estimation. *Mon. Weather Rev.*, 133, 1098–1118.
- Gottwald, G. A. and S. Reich (2021) Supervised learning from noisy observations: Combining
 machine-learning techniques with data assimilation. *Phys. Nonlinear Phenom.*, 423,
- 812 132911.
- B13 Grönquist, P., C. Yao, T. Ben-Nun, N. Dryden, P. Dueben, S. Li, and T. Hoefler (2021) Deep
 B14 learning for post-processing ensemble weather forecasts. *Philos. Trans. R. Soc. Math.*B15 *Phys. Eng. Sci.*, **379**, 20200092.
- 816 Hamre, K. D., M. E. Lofton, R. P. McClure, Z. W. Munger, J. P. Doubek, A. B. Gerling, M. E.
- 817 Schreiber, and C. C. Carey (2018) In situ fluorometry reveals a persistent, perennial
- hypolimnetic cyanobacterial bloom in a seasonally anoxic reservoir. *Freshw. Sci.*, 37,
 483–95.
- Harris, D. J., S. D. Taylor, and E. P. White (2018) Forecasting biodiversity in breeding birds
 using best practices. *PeerJ*, 6, e4278.
- He, H., L. Lei, J. S. Whitaker, and Z.-M. Tan (2020) Impacts of Assimilation Frequency on
- 823 Ensemble Kalman Filter Data Assimilation and Imbalances. J. Adv. Model. Earth Syst.,
 824 12, e2020MS002187.
- Heilman, K. A., M. C. Dietze, A. A. Arizpe, J. Aragon, A. Gray, J. D. Shaw, A. O. Finley, S.

826	Klesse, et al. (2022) Ecological forecasting of tree growth: Regional fusion of tree-ring
827	and forest inventory data to quantify drivers and characterize uncertainty. Glob. Change
828	<i>Biol.</i> , 28 , 2442–60.

- 829 Herrick, C., B. G. Steele, J. A. Brentrup, K. L. Cottingham, M. J. Ducey, D. A. Lutz, M. W.
- Palace, M. C. Thompson, *et al.* (2023) lakeCoSTR: A tool to facilitate use of Landsat
 Collection 2 to estimate lake surface water temperatures. *Ecosphere*, 14, e4357.
- Hipsey, M. R., C. Boon, L. C. Bruce, Q. Thomas, M. Weber, L. Winslow, J. S. Read, and D. P.
 Hamilton (2022) AquaticEcoDynamics/glm-aed: v3.3.0.
- Hipsey, M. R., L. C. Bruce, C. Boon, B. Busch, C. C. Carey, D. P. Hamilton, P. C. Hanson, J. S.
- 835 Read, et al. (2019) A General Lake Model (GLM 3.0) for linking with high-frequency
- sensor data from the Global Lake Ecological Observatory Network (GLEON). *Geosci.*
- 837 *Model Dev.*, **12**, 473–523.
- Hounshell, A. G., R. P. McClure, M. E. Lofton, and C. C. Carey (2021) Whole-ecosystem
- 839 oxygenation experiments reveal substantially greater hypolimnetic methane
- concentrations in reservoirs during anoxia. *Limnol. Oceanogr. Lett.*, **6**, 33–42.
- 841 Ibáñez, I., E. S. Gornish, L. Buckley, D. M. Debinski, J. Hellmann, B. Helmuth, J.
- 842 HilleRisLambers, A. M. Latimer, *et al.* (2013) Moving forward in global-change ecology:
 843 capitalizing on natural variability. *Ecol. Evol.*, 3, 170–81.
- Jolliffe, I. T. and D. B. Stephenson (2012) *Forecast verification: a practitioner's guide in atmospheric science*. John Wiley & Sons.
- Kehoe, M. J., K. P. Chun, and H. M. Baulch (2015) Who Smells? Forecasting Taste and Odor in
 a Drinking Water Reservoir. *Environ. Sci. Technol.*, 49, 10984–92.
- 848 Kirchner, J. W. and C. Neal (2013) Universal fractal scaling in stream chemistry and its

- 849 implications for solute transport and water quality trend detection. *Proc. Natl. Acad. Sci.*,
 850 110, 12213–18.
- LaDeau, S. L., B. A. Han, E. J. Rosi-Marshall, and K. C. Weathers (2017) The Next Decade of
 Big Data in Ecosystem Science. *Ecosystems*, 20, 274–83.
- Ladwig, R., L. A. Rock, and H. A. Dugan (2021) Impact of salinization on lake stratification and
 spring mixing. *Limnol. Oceanogr. Lett,* 8, 93-102.
- 855 Lewis, A. S. L., W. M. Woelmer, H. L. Wander, D. W. Howard, J. W. Smith, R. P. McClure, M.
- E. Lofton, N. W. Hammond, *et al.* (2022) Increased adoption of best practices in
 ecological forecasting enables comparisons of forecastability. *Ecol. Appl.*, **32**, e2500.
- Lin, E., Y. Yang, X. Qiu, Q. Xie, R. Gan, B. Zhang, and X. Liu (2021) Impacts of the radar data
- assimilation frequency and large-scale constraint on the short-term precipitation forecast
 of a severe convection case. *Atmospheric Res.*, 257, 105590.
- 861 Lindegren, M., C. Möllmann, A. Nielsen, K. Brander, B. R. MacKenzie, and N. Chr. Stenseth
- 862 (2010) Ecological forecasting under climate change: the case of Baltic cod. *Proc. R. Soc.*863 *B Biol. Sci.*, 277, 2121–30.
- Liu, H., Y. D. Chen, T. Liu, and L. Lin (2019) The river chief system and river pollution control
 in China: A case study of Foshan. *Water*, 11, 1606.
- Lofton, M. E., J. A. Brentrup, W. S. Beck, J. A. Zwart, R. Bhattacharya, L. S. Brighenti, S. H.
- Burnet, I. M. McCullough, *et al.* (2022) Using near-term forecasts and uncertainty
- 868 partitioning to inform prediction of oligotrophic lake cyanobacterial density. *Ecol. Appl.*,
 869 32, e2590.
- 870 Lofton, M. E., D. W. Howard, R. Q. Thomas, and C. C. Carey (2023) Progress and opportunities
- 871 in advancing near-term forecasting of freshwater quality. *Glob. Change Biol.*, **29**,1691–

872 1714.

- 873 Luo, Y., K. Ogle, C. Tucker, S. Fei, C. Gao, S. LaDeau, J. S. Clark, and D. S. Schimel (2011) 874 Ecological forecasting and data assimilation in a data-rich era. Ecol. Appl., 21, 1429–42.
- 875 Machete, R. L. and L. A. Smith (2016) Demonstrating the value of larger ensembles in
- 876 forecasting physical systems. Tellus Dyn. Meteorol. Oceanogr., 68, 28393.
- 877 Magnuson, J. J., L. B. Crowder, and P. A. Medvick (1979) Temperature as an ecological 878 resource. Am. Zool., 19, 331-43.
- 879 Malhi, Y., J. Franklin, N. Seddon, M. Solan, M. G. Turner, C. B. Field, and N. Knowlton (2020)
- 880 Climate change and ecosystems: threats, opportunities and solutions. Philos. Trans. R. 881
- Soc. B Biol. Sci., 375, 20190104.
- 882 Mantovani, C., L. Corgnati, J. Horstmann, A. Rubio, E. Reyes, C. Quentin, S. Cosoli, J. L.
- 883 Asensio, et al. (2020) Best Practices on High Frequency Radar Deployment and 884 Operation for Ocean Current Measurement. Front. Mar. Sci., 7, 210.
- 885 Marcé, R., G. George, P. Buscarinu, M. Deidda, J. Dunalska, E. de Eyto, G. Flaim, H.-P.
- 886 Grossart, et al. (2016) Automatic High Frequency Monitoring for Improved Lake and 887 Reservoir Management. Environ. Sci. Technol., 50, 10780-94.
- 888 Marj, A. F. and A. M. J. Meijerink (2011) Agricultural drought forecasting using satellite 889 images, climate indices and artificial neural network. Int. J. Remote Sens., 32, 9707–19.
- 890 Massoud, E. C., J. Huisman, E. Benincà, M. C. Dietze, W. Bouten, and J. A. Vrugt (2018)
- 891 Probing the limits of predictability: data assimilation of chaotic dynamics in complex food webs. *Ecol. Lett.*, **21**, 93–103. 892
- 893 McClure, R. P., R. Q. Thomas, M. E. Lofton, W. M. Woelmer, and C. C. Carey (2021) Iterative
- 894 Forecasting Improves Near-Term Predictions of Methane Ebullition Rates. Front.

- *Environ. Sci.*, **9**, 756603.
- 896 Mercado-Bettín, D., F. Clayer, M. Shikhani, T. N. Moore, M. D. Frías, L. Jackson-Blake, J.
- 897 Sample, M. Iturbide, *et al.* (2021) Forecasting water temperature in lakes and reservoirs
 898 using seasonal climate prediction. *Water Res.*, 201, 117286.
- Meyer, J. L., M. J. Sale, P. J. Mulholland, and N. L. Poff (1999) Impacts of climate change on
 aquatic ecosystem functioning and health 1. *JAWRA J. Am. Water Resour. Assoc.*, 35,
 1373–86.
- 902 Mi, C., T. Shatwell, J. Ma, Y. Xu, F. Su, and K. Rinke (2020) Ensemble warming projections in
- 903 Germany's largest drinking water reservoir and potential adaptation strategies. *Sci. Total*904 *Environ.*, 748, 141366.
- 905 Moustahfid, H., L. C. Hendrickson, A. Arkhipkin, G. J. Pierce, A. Gangopadhyay, H. Kidokoro,
- 906 U. Markaida, C. Nigmatullin, *et al.* (2021) Ecological-fishery forecasting of squid stock

907 dynamics under climate variability and change: review, Challenges, and

- 908 Recommendations. *Rev. Fish. Sci. Aquac.*, **29**, 682–705.
- Niu, S., Y. Luo, M. C. Dietze, T. F. Keenan, Z. Shi, J. Li, and F. S. C. Iii (2014) The role of data
 assimilation in predictive ecology. *Ecosphere*, 5, 1-16.
- 911 O'Reilly, C. M., S. Sharma, D. K. Gray, S. E. Hampton, J. S. Read, R. J. Rowley, P. Schneider,
- J. D. Lenters, *et al.* (2015) Rapid and highly variable warming of lake surface waters
 around the globe. *Geophys. Res. Lett.*, 42, 10–773.
- Paerl, H. W. and V. J. Paul (2012) Climate change: links to global expansion of harmful
 cyanobacteria. *Water Res.*, 46, 1349–63.
- 916 Page, T., P. J. Smith, K. J. Beven, I. D. Jones, J. A. Elliott, S. C. Maberly, E. B. Mackay, M. De
- 917 Ville, *et al.* (2018) Adaptive forecasting of phytoplankton communities. *Water Res.*, **134**,

918 74–85.

919	Page, T., P. J. Smith, K. J. Beven, I. D. Jones, J. A. Elliott, S. C. Maberly, E. B. Mackay, M. De	
920	Ville, et al. (2017) Constraining uncertainty and process-representation in an algal	
921	community lake model using high frequency in-lake observations. Ecol. Model., 357, 1-	
922	13.	
923	Park, J., K. T. Kim, and W. H. Lee (2020) Recent Advances in Information and Communications	
924	Technology (ICT) and Sensor Technology for Monitoring Water Quality. Water, 12, 510.	
925	Parrish, M. A., H. Moradkhani, and C. M. DeChant (2012) Toward reduction of model	
926	uncertainty: Integration of Bayesian model averaging and data assimilation. Water	
927	<i>Resour. Res.</i> , 48 , W03519.	
928	Piazzi, G., G. Thirel, L. Campo, and S. Gabellani (2018) A particle filter scheme for multivariate	
929	data assimilation into a point-scale snowpack model in an Alpine environment. The	
930	<i>Cryosphere</i> , 12 , 2287–2306.	
931	R Core Team (2022) R: A language and environment for statistical computing.	
932	Read, J. S., X. Jia, J. Willard, A. P. Appling, J. A. Zwart, S. K. Oliver, A. Karpatne, G. J. A.	
933	Hansen, et al. (2019) Process-Guided Deep Learning Predictions of Lake Water	
934	Temperature. Water Resour. Res., 55, 9173–90.	
935	Read, J. S., L. A. Winslow, G. J. A. Hansen, J. Van Den Hoek, P. C. Hanson, L. C. Bruce, and C.	
936	D. Markfort (2014) Simulating 2368 temperate lakes reveals weak coherence in	
937	stratification phenology. Ecol. Model., 291, 142-50.	
938	Romero, J. R., I. Kagalou, J. Imberger, D. Hela, M. Kotti, A. Bartzokas, T. Albanis, N.	
939	Evmirides, et al. (2002) Seasonal water quality of shallow and eutrophic Lake Pamvotis,	
940	Greece: implications for restoration. Hydrobiologia, 474, 91-105	

- 941 Rousso, B. Z., E. Bertone, R. Stewart, and D. P. Hamilton (2020) A systematic literature review
 942 of forecasting and predictive models for cyanobacteria blooms in freshwater lakes. *Water*943 *Res.*, 182, 115959.
- 944 Simonin, D., C. Pierce, N. Roberts, S. P. Ballard, and Z. Li (2017) Performance of Met Office

945 hourly cycling NWP-based nowcasting for precipitation forecasts. *Q. J. R. Meteorol.*946 *Soc.*, 143, 2862–73.

Steere, D. C., A. Baptista, D. McNamee, C. Pu, and J. Walpole (2000) Research challenges in
environmental observation and forecasting systems. *Proceedings of the 6th annual*

- 949 *international conference on Mobile computing and networking MobiCom '00.* ACM
- 950 Press, Boston, Massachusetts, United States, pp. 292–99.
- Tanut, B., R. Waranusast, and P. Riyamongkol (2021) High accuracy pre-harvest sugarcane yield
 forecasting model utilizing drone image analysis, data mining, and reverse design
 method. *Agriculture*, 11, 682.
- 954 Thomas, R. Q., C. Boettiger, C. C. Carey, M. C. Dietze, L. R. Johnson, M. A. Kenney, J. S.
- 955 McLachlan, J. A. Peters, *et al.* (2023a) The NEON Ecological Forecasting Challenge.
 956 *Front. Ecol. Environ.*, 21, 112–13.
- 957 Thomas, R. Q., R. J. Figueiredo, V. Daneshmand, B. J. Bookout, L. K. Puckett, and C. C. Carey

958 (2020) A Near-Term Iterative Forecasting System Successfully Predicts Reservoir

- 959 Hydrodynamics and Partitions Uncertainty in Real Time. *Water Resour. Res.*, **56**,
- 960 e2019WR026138.
- 961 Thomas, R. Q., R. P. McClure, T. N. Moore, W. M. Woelmer, C. Boettiger, R. J. Figueiredo, R.
- 962 T. Hensley, and C. C. Carey (2023b) Near-term forecasts of NEON lakes reveal gradients
- 963 of environmental predictability across the US. *Front. Ecol. Environ.*, n/a.

- 964 Wander, H. L., R. Q. Thomas, T. N. Moore, M. E. Lofton, A. Breef-Pilz, and C. C. Carey
- 965 (2023a) Data assimilation experiments inform monitoring needs for near-term ecological
- 966 forecasts in a eutrophic reservoir: data, forecasts, and scores [Data set]. Zenodo.
- 967 https://doi.org/10.5281/zenodo.7951402
- 968 Wander, H. L., R. Q. Thomas, T. N. Moore, M. E. Lofton, A. Breef-Pilz, and C. C. Carey
- 969 (2023b) hlwander/BVRE-forecast-code: Data assimilation experiments inform
- 970 monitoring needs for near-term ecological forecasts in a eutrophic reservoir: Code (v1.1).
 971 Zenodo. https://doi.org/10.5281/zenodo.7958471
- 972 Wang, S., N. Flipo, and T. Romary (2023) Which filter for data assimilation in water quality
- 973 models? Focus on oxygen reaeration and heterotrophic bacteria activity. *J. Hydrol.*, 620,
 974 129423.
- 975 Wang, X., J. Zhang, and V. Babovic (2016) Improving real-time forecasting of water quality
- 976 indicators with combination of process-based models and data assimilation technique.
- 977 *Ecol. Indic.*, **66**, 428–39.
- Weng, E. and Y. Luo (2011) Relative information contributions of model vs. data to short- and
 long-term forecasts of forest carbon dynamics. *Ecol. Appl.*, 21, 1490–1505.
- 980 Wheeler, K., M. Dietze, D. LeBauer, J. Peters, A. D. Richardson, R. Q. Thomas, K. Zhu, U.
- 981 Bhat, *et al.* (2023) Predicting Spring Phenology in Deciduous Broadleaf Forests: An
 982 Open Community Forecast Challenge. Available at SSRN: 10.2139/ssrn.4357147
- 983 White, E. P., G. M. Yenni, S. D. Taylor, E. M. Christensen, E. K. Bledsoe, J. L. Simonis, and S.
- 984 K. M. Ernest (2019) Developing an automated iterative near-term forecasting system for
 985 an ecological study. *Methods Ecol. Evol.*, 10, 332–44.
- 986 Williamson, C. E., E. P. Overholt, J. A. Brentrup, R. M. Pilla, T. H. Leach, S. G. Schladow, J. D.

987	Warren, S. S. Urmy, et al. (2016) Sentinel responses to droughts, wildfires, and floods:
988	effects of UV radiation on lakes and their ecosystem services. Front. Ecol. Environ., 14,
989	102–9.
990	Woelmer, W. M., R. Q. Thomas, M. E. Lofton, R. P. McClure, H. L. Wander, and C. C. Carey
991	(2022) Near-term phytoplankton forecasts reveal the effects of model time step and
992	forecast horizon on predictability. Ecol. Appl., 32, e2642.
993	Woolway, R. I., E. Jennings, T. Shatwell, M. Golub, D. C. Pierson, and S. C. Maberly (2021)
994	Lake heatwaves under climate change. Nature, 589, 402–7.
995	Yvon-Durocher, G., J. M. Caffrey, A. Cescatti, M. Dossena, P. del Giorgio, J. M. Gasol, J. M.
996	Montoya, J. Pumpanen, et al. (2012) Reconciling the temperature dependence of
997	respiration across timescales and ecosystem types. Nature, 487, 472–76.
998	Ziliani, M. G., R. Ghostine, B. Ait-El-Fquih, M. F. McCabe, and I. Hoteit (2019) Enhanced flood
999	forecasting through ensemble data assimilation and joint state-parameter estimation. J.
1000	<i>Hydrol.</i> , 577 , 123924.
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1004	Figure	Captions
1004	rigure	Captions

Figure 1: Forecasting Lake And Reservoir Ecosystems (FLARE) workflow showing the step-bystep process for generating daily water temperature forecasts, starting with data collection from

- 1007 thermistors deployed in the reservoir (step 1), then data access for running the forecast model
- 1008 (step 2), then generation of forecasts with data assimilation (step 3), and ending with forecast
- skill assessment (step 4). During the data assimilation steps (3a-b), data assimilation experiments
- 1010 were performed with four different data assimilation frequencies (daily, weekly, fortnightly, and
- 1011 monthly; see dashed line box). Steps 1-4 occurred throughout the entire forecast period (1
- 1012 January 31 December 2021). Buoy figure via NexSens Technology Inc., CC by 2.0
- 1013 https://creativecommons.org/licenses/by/2.0/
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Figure 2: Map of Beaverdam Reservoir, Vinton, VA (37.31° N, 79.82° W). The map shows the
surrounding forested watershed; the point represents the reservoir monitoring site where highfrequency sensor data were collected.

1018

Figure 3: Observed water temperature for all depths with high-frequency sensors during the

- 1020 forecasting period of 1 January 31 December 2021 in Beaverdam Reservoir (BVR). The gray
- 1021 background indicates the mixed period (1 January 11 March, 8 November 31 December
- 1022 2021), while the white background indicates the thermally-stratified period (12 March 7
- 1023 November 2021), defined by a $<0.1 \text{ kg/m}^3$ density differential between surface and bottom
- 1024 layers.

1026 Figure 4: Example of water temperature forecasts at 1 m (a), 5 m (b), and 9 m (c) generated for 1027 1-35 days into the future in Beaverdam Reservoir. Data assimilation (DA) frequencies are 1028 depicted by colors; shading shows 95% confidence intervals around the mean predicted 1029 temperature for each day. Black points represent water temperature observations. Colored points 1030 represent the most recent day that data was assimilated for each DA frequency. In this example, 1031 data were most recently assimilated on the day that the forecasts were generated: 25 June for the 1032 monthly DA scenario, 9 July for the fortnightly DA scenario, 16 July for the weekly DA 1033 scenario, and 22 July for the daily DA scenario. 1034 1035 Figure 5: Parameter evolution during the forecast period (1 January - 31 December 2021) for 1036 daily, weekly, fortnightly, and monthly data assimilation (DA) frequencies at 1-day-ahead

forecast horizons. Longwave (a) is the longwave radiation scaling parameter, hypo_sed_temp (b)
is the hypolimnetic sediment temperature parameter, and epi_sed_temp (c) is the epilimnetic
sediment temperature parameter.

1040

Figure 6: Root mean square error (RMSE) of mean forecasted water temperature compared to observations for 1-35 day-ahead forecast horizons in Beaverdam Reservoir, aggregated for all depths in the water column and days within the 365-day forecast period. RMSE for each forecast horizon was averaged from forecasts generated during 1 January - 31 December 2021. Colored lines represent different data assimilation (DA) frequencies. The dotted line depicts the 2°C threshold for skillful water temperature forecasts.

1048 Figure 7: Root mean square error (RMSE) of mean forecasted water temperature compared to

1049 observations for 1-35-day-ahead forecast horizons in Beaverdam Reservoir during the mixed

1050 (panels a, c, e) vs. stratified (panels b, d, f) periods at 1 m (a, b), 5 m (c, d), and 9 m (e, f). RMSE

1051 for each forecast horizon was averaged across the 365-day forecast period (1 January - 31

1052 December 2021). Colored lines correspond to different data assimilation (DA) frequencies

1053 (daily, weekly, fortnightly, and monthly); dotted horizontal lines depict the 2°C threshold for

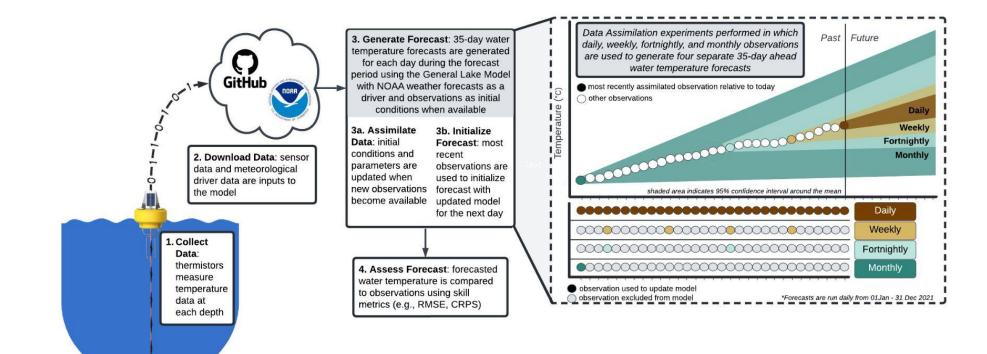
skillful forecasts.

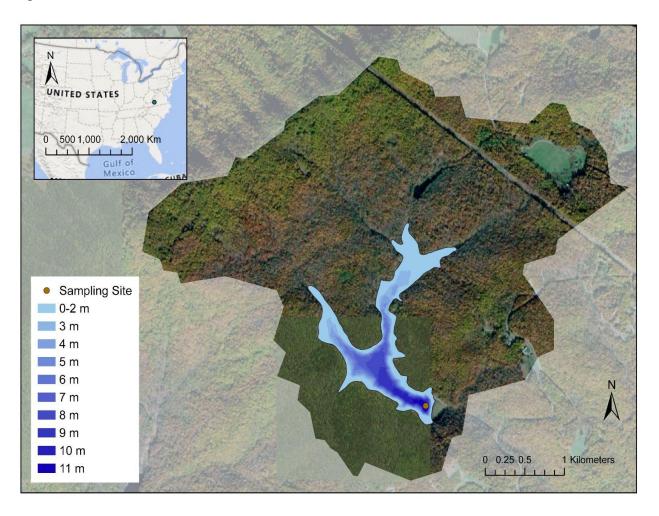
1055

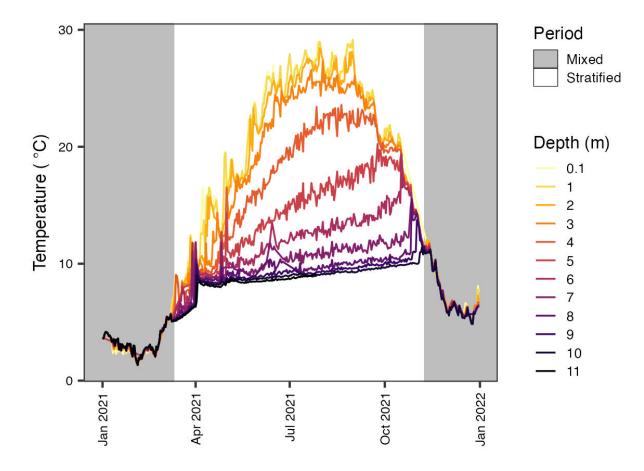
Figure 8: Mean water temperature forecast variance across horizons (1-35 days ahead) in
Beaverdam Reservoir during the mixed (panels a, c, e) vs. stratified (panels b, d, f) periods for 1
m (a, b), 5 m (c, d), and 9 m (e, f). Variance for each forecast horizon was averaged from all 365
forecasts generated during the forecast period (1 January - 31 December 2021). Colored lines
correspond to different data assimilation (DA) frequencies (daily, weekly, fortnightly, and
monthly).

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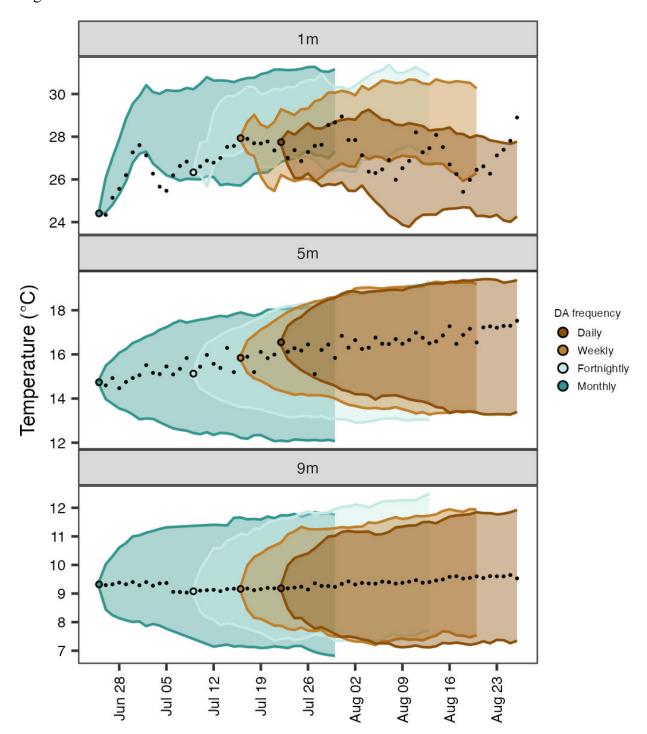
Figure 9: Proportion of initial conditions uncertainty relative to total forecast uncertainty
averaged across all forecasts generated with each DA frequency calculated using forecast
variance across 1 January - 31 December 2021. Colored lines depict data assimilation (DA)
frequencies (daily, weekly, fortnightly, and monthly). Panels a, c, and e represent mixed period
forecasts, panels b, d, and f represent stratified period forecasts. Depths (1, 5, and 9 m) are
indicated by gray facet labels.



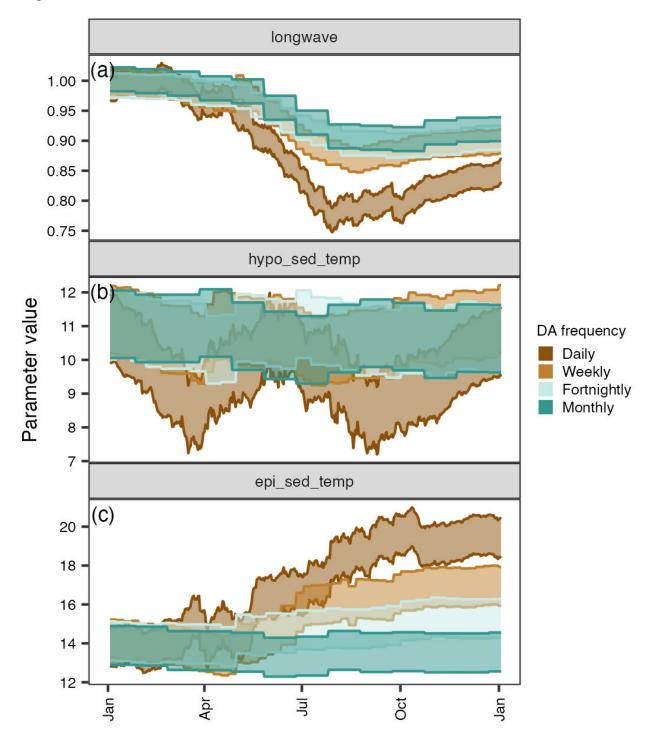


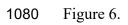


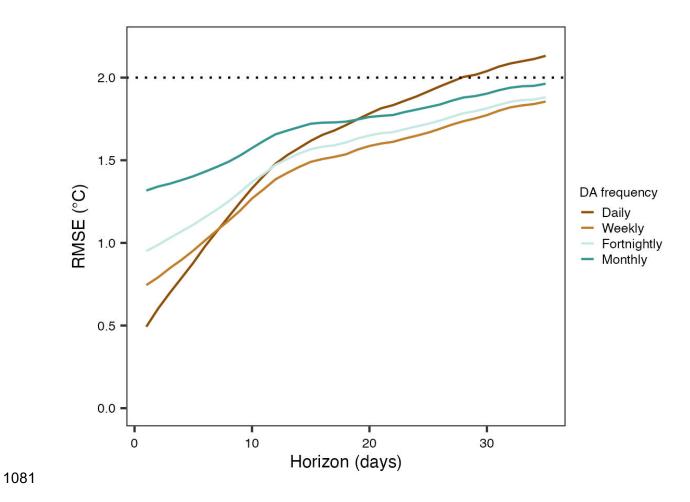
1076 Figure 4.



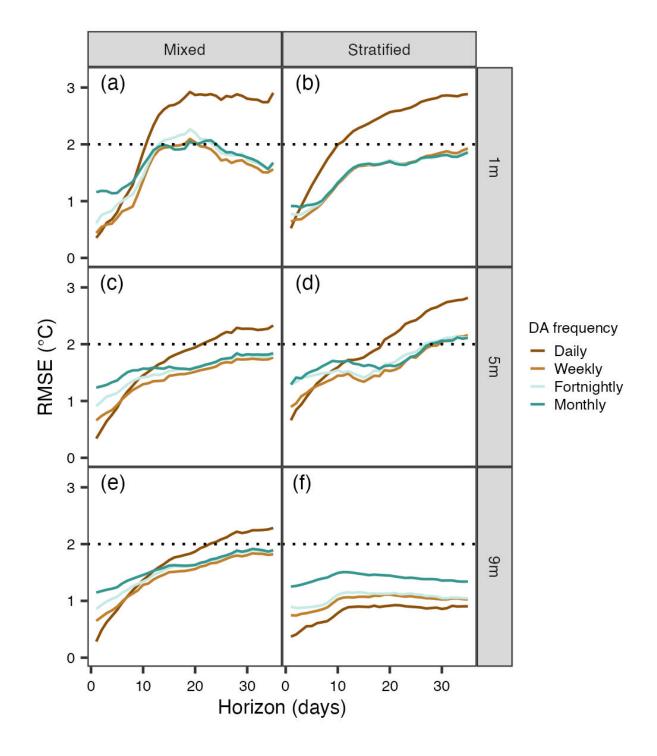
1078 Figure 5.

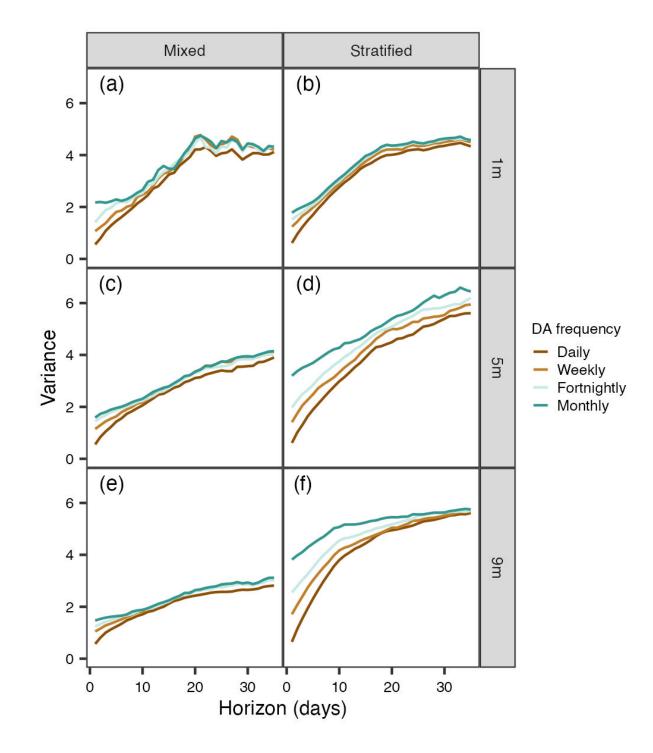


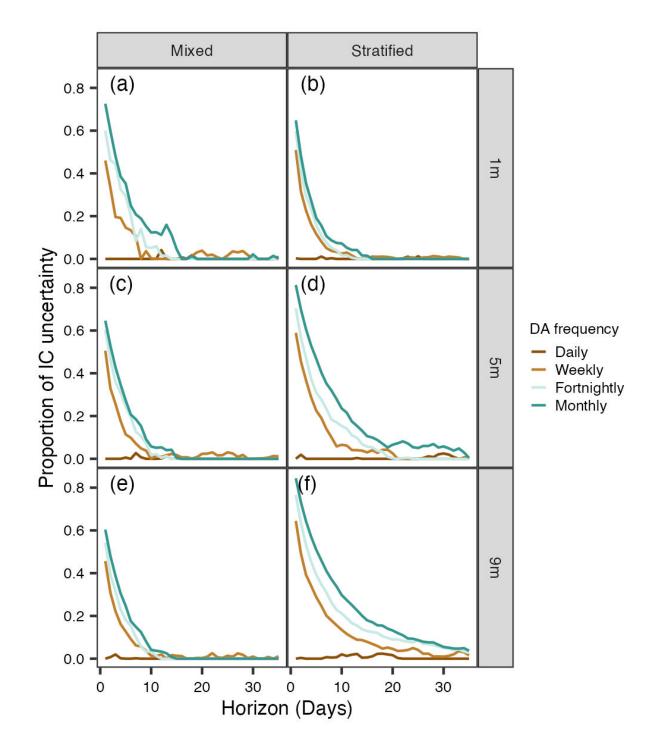




1082 Figure 7.







Supplement to: Data assimilation experiments inform monitoring needs for near-term ecological forecasts in a eutrophic reservoir

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Appendix S1

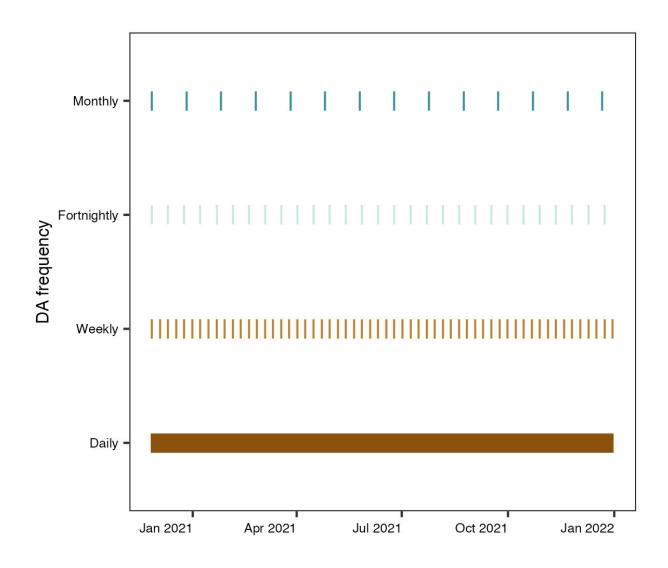


Figure S1: Frequencies for daily, weekly, fortnightly, and monthly data assimilation (DA); lines indicate the dates when DA occurred. For example, daily DA occurred every day from 27 November 2020 to 31 December 2021, whereas monthly DA occurred 14 times during the 14-month period of November 2020 to January 2022.

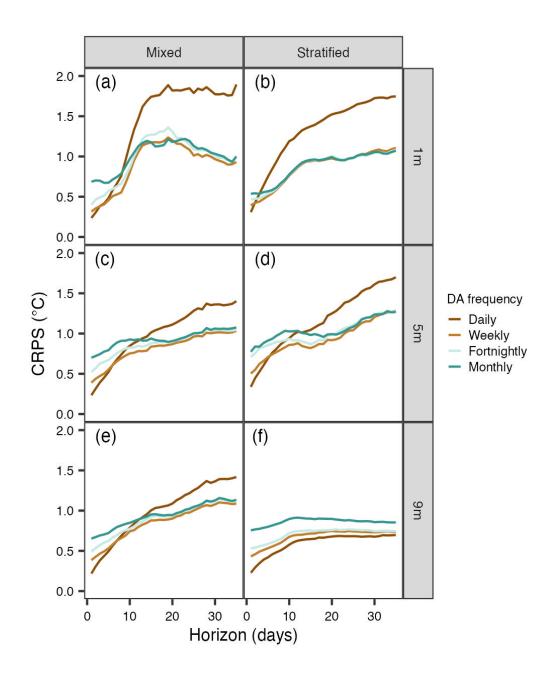


Figure S2: Water temperature forecast continuous ranked probability score (CRPS) across different depths (1 m: a, b; 5 m: c, d; and 9 m: e, f) and horizons in Beaverdam Reservoir during the mixed (a, c, e) vs. stratified (b, d, f) periods. Each grouping of bars represents a different data assimilation (DA) frequency (daily, weekly, fortnightly, monthly). Depths are indicated in the right y-axis labels; horizons are depicted by colored boxplots.

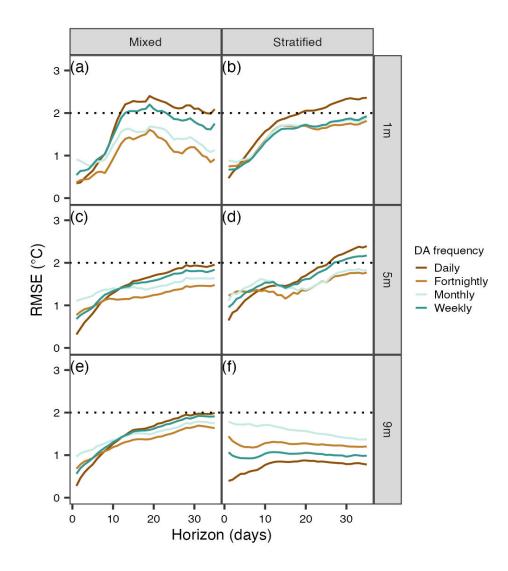


Figure S3: Root mean square error (RMSE) calculated from comparing water temperature observations with forecasts that did not include initial conditions uncertainty for 1-35-day-ahead forecasts in Beaverdam Reservoir during the mixed (a, c, e) vs. stratified (b, d, f) periods at 1 m (a, b), 5 m (c, d), and 9 m (e, f). RMSE for each forecast horizon was averaged across the 365-day forecast period (1 January - 31 December 2021). Colored lines correspond to different data assimilation (DA) frequencies (daily, weekly, fortnightly, and monthly); dotted horizontal lines depict the 2°C RMSE threshold for skillful forecasts.

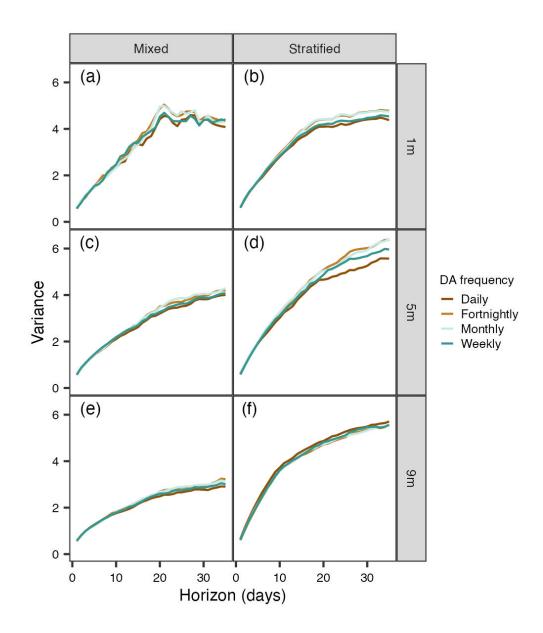


Figure S4: Variance for forecasts that did not include initial conditions uncertainty across horizons (1-35 days ahead) in Beaverdam Reservoir during the mixed (a, c, e) vs. stratified (b, d, f) periods for 1 m (a, b), 5 m (c, d), and 9 m (e, f). Variance for each forecast horizon was averaged from all 365 forecasts generated during the forecast period (1 January - 31 December 2021). Colored lines correspond to different data assimilation (DA) frequencies (daily, weekly, fortnightly, and monthly).