A spatial assessment of current and future foliar Hg uptake fluxes across European forests

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Abstract

Atmospheric mercury (Hg) is deposited to land surfaces mainly through vegetation uptake. Foliage stomatal gas exchange plays an important role for net vegetation Hg uptake, because foliage assimilates Hg via the stomata. Here, we use empirical relationships of foliar Hg uptake by forest tree species to produce a spatially highly resolved (1 km2) map of foliar Hg fluxes to European forests over one growing season. The modelled forest foliar Hg uptake flux is 23 ± 12 Mg Hg season-1, which agrees with previous estimates from literature. We spatially compare forest Hg fluxes with modelled fluxes of the chemistry-transport model GEOS-Chem and find a good overall agreement. For European pine forests, stomatal Hg uptake was shown to be sensitive to prevailing conditions of relatively high ambient water vapor pressure deficit (VPD). We tested a stomatal uptake model for the total pine needle Hg uptake flux during four previous growing seasons (1994, 2003, 2015/2017, 2018) and two climate change scenarios (RCP 4.5 and RCP 8.5). The resulting modelled total European pine needle Hg uptake fluxes are in a range of 8.0 - 9.3 Mg Hg season-1 (min - max). The lowest pine forest needle Hg uptake flux to Europe (8 Mg Hg season-1) among all investigated growing seasons is associated with unusually hot and dry ambient conditions in the European summer 2018, highlighting the sensitivity of the investigated flux to prolonged high VPD. We conclude, that stomatal modelling is particularly useful to investigate changes in Hg deposition in the context of extreme climate events.

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Key Points:

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11	•	Extreme hot and dry atmospheric conditions have the potential to reduce stom-
12		atal uptake of ambient mercury by pine trees in Europe
13	•	Atmospheric drought controls on stomatal mercury uptake should be accounted
14		for in mercury transport models like GEOS-Chem
15	•	Forest foliar mercury uptake fluxes to Europe from a bottom-up model generally
16		agree well with results derived from literature and GEOS-Chem

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17 Abstract

Atmospheric mercury (Hg) is deposited to land surfaces mainly through vegetation uptake. Foliage stomatal gas exchange plays an important role for net vegetation Hg uptake, because foliage assimilates Hg via the stomata. Here, we use empirical relationships of foliar Hg uptake by forest tree species to produce a spatially highly resolved (1 km²) map of foliar Hg fluxes to European forests over one growing season. The modelled forest foliar Hg uptake flux is 23 ± 12 Mg Hg season⁻¹, which agrees with previous estimates from literature.

We spatially compare forest Hg fluxes with modelled fluxes of the chemistry-transport 25 model GEOS-Chem and find a good overall agreement. For European pine forests, stom-26 atal Hg uptake was shown to be sensitive to prevailing conditions of relatively high am-27 bient water vapor pressure deficit (VPD). We tested a stomatal uptake model for the to-28 tal pine needle Hg uptake flux during four previous growing seasons (1994, 2003, 2015/2017, 29 2018) and two climate change scenarios (RCP 4.5 and RCP 8.5). The resulting modelled 30 total European pine needle Hg uptake fluxes are in a range of 8.0 - $9.3 \text{ Mg Hg season}^{-1}$ 31 (min - max). The lowest pine forest needle Hg uptake flux to Europe (8 Mg Hg season⁻¹) 32 among all investigated growing seasons is associated with unusually hot and dry ambi-33 ent conditions in the European summer 2018, highlighting the sensitivity of the inves-34 tigated flux to prolonged high VPD. We conclude, that stomatal modelling is particu-35 larly useful to investigate changes in Hg deposition in the context of extreme climate events. 36

37 1 Introduction

Mercury (Hg) is a toxic pollutant that is transported globally through the atmo-38 sphere and deposited from air to land surfaces mainly through vegetation uptake of am-39 bient gaseous elemental Hg(0) (Demers et al., 2013; Jiskra et al., 2015; Enrico et al., 2016; 40 Obrist et al., 2017; Feinberg et al., 2022). Consequently, vegetation uptake has the po-41 tential to lower atmospheric Hg(0) transport and Hg deposition to oceans, where Hg can 42 be methylated and bioaccumulated in marine seafood for human consumption (Zhou et 43 al., 2021). In order to assess and improve the effectiveness of mitigation policies for hu-44 man exposure, it is thus necessary to constrain environmental drivers of vegetation Hg(0)45 uptake. Furthermore, process understanding of vegetation Hg(0) uptake is essential for 46 assessing future human Hg exposure in the context of global change (Sonke et al., 2023). 47

Global vegetation and soil uptake of Hg(0) has been estimated to amount to 2850 48 \pm 500 Mg year⁻¹ (Obrist et al., 2021; Zhou et al., 2021; Feinberg et al., 2022), exceed-49 ing approximate direct anthropogenic emissions to the air of 2200 Mg Hg year⁻¹ (Sonke 50 et al., 2023). Forests contain 80 % of the global plant biomass (Pan et al., 2013), there-51 for representing a major vector for Hg(0) drawdown from the atmosphere. In forests, 52 half of the total Hg(0) net deposition is estimated to be stored in tree foliage, while the 53 other half is estimated to be transferred to vascular tissues (e.g. stem, branches, roots), 54 or taken up by understory vegetation (e.g. shrubs, grasses) or nonvascular plants (lichen 55 and mosses) (Zhou et al., 2021; Obrist et al., 2021; Zhou & Obrist, 2021). In tree foliage, 56 Hg concentrations increase linearly between foliage emergence and senescence (Rea et 57 al., 2002; Laacouri et al., 2013; Blackwell et al., 2014; Wohlgemuth et al., 2020; Pleijel 58 et al., 2021) implying a net foliar Hg deposition flux, albeit Hg re-emission from foliar 59 surfaces of up to 30% of gross foliage Hg(0) deposition had been observed in a subtrop-60 ical forest in China (W. Yuan, Sommar, et al., 2019). The bulk (90-96%) of Hg is stored 61 in foliage tissues as opposed to leaf surfaces and correlates with leaf stomatal density (Laacouri 62 et al., 2013). Studies on Hg stable isotopes in foliage (Demers et al., 2013; Zhou et al., 63 2021), enriched isotope tracer experiments (Rutter et al., 2011) and the vertical varia-64 tion of net foliar Hg uptake in forest canopies (Wohlgemuth et al., 2020) strongly sug-65 gest a diffusive uptake pathway of atmospheric Hg(0) to foliage interiors via the stom-66 ata (Liu et al., 2021). In this way, foliar Hg(0) uptake is linked to foliage stomatal aper-67 ture for atmospheric gas exchange (Wohlgemuth et al., 2022). 68

Trees regulate foliage stomatal aperture to balance the inward diffusion of CO_2 for 69 photosynthesis with the risk of desiccation caused by excessive outward diffusion of wa-70 ter vapor (Körner, 2013). The degree of stomatal aperture depends on atmospheric CO_2 71 levels and hydrological conditions (soil water availability and atmospheric evaporative 72 demand) and varies among foliage-specific traits (age, tree species-specific evolutionary 73 metabolic strategy and water use efficiency) (Körner, 2013). Pine, for instance, is an iso-74 hydric tree species capable of closing foliage stomata under warm and dry atmospheric 75 conditions relatively early compared to tree species like oak and spruce (Lagergren & Lin-76 droth, 2002; Zweifel et al., 2007, 2009), resulting in a reduced stomatal conductance for 77 pine needle diffusive gas exchange (Panek & Goldstein, 2001). Consistently, Hg(0) up-78 take rates by pine needles in Europe were found to be lower at forest sites across Europe, 79 where prolonged warm and dry atmospheric conditions prevailed over a given growing 80 season during daytime (Wohlgemuth et al., 2022). 81

Species-specific stomatal response strategies to meteorological conditions are par-82 ticularly relevant for projections of future foliar Hg uptake under climate change. Increas-83 ing global atmospheric temperatures driven by rising levels of greenhouse gases will re-84 sult in an increased frequency of droughts (Grossiord et al., 2020) and higher soil mois-85 ture deficits (Berg & Sheffield, 2018; Stocker et al., 2019) in various regions of the world. 86 These climatic conditions may decrease foliar Hg(0) uptake fluxes due to lower stomatal 87 conductance (Wohlgemuth et al., 2022). A reduced plant Hg sink could further be am-88 plified by deforestation and forest diebacks, particularly in the tropics (Allen et al., 2015; 89 Brando et al., 2019; Feinberg et al., 2023). Other regions of the world are projected to 90 become wetter through an increase in precipitation rates under climate change (IPCC, 91 2021a), which might lead to higher foliage stomatal conductance relative to the present 92 and thus higher foliar Hg uptake. With continuing anthropogenic carbon emissions, an 93 elevated atmospheric CO_2 level might have an antagonizing effect on the foliar stomatal 94 Hg(0) uptake flux: foliar Hg(0) uptake could decline with decreasing stomatal conductance under CO₂ fertilization (Norby & Zak, 2011), or, the opposite, the vegetation sink 96 for Hg(0) could increase with intensified biomass growth and higher soil C contents (Hararuk 97 et al., 2013; Jiskra et al., 2018; H. Zhang et al., 2016). In order to make projections of 98 the foliar Hg uptake flux in the next decades, these climate change impacts need to be 99 further investigated and potentially implemented into global and regional Hg cycle mod-100 els. 101

Current and future Hg fluxes are modelled in Global Chemical Transport Models 102 (CTMs). CTMs like GEOS-Chem (Selin et al., 2008) apply resistance-based algorithms 103 (Wesely, 2007) for modelling Hg(0) deposition fluxes from the atmosphere to vegetated 104 ecosystems and are often based on parameters like leaf area indices (LAIs), temperature 105 and wind speed. The resistance components for leaf stomata within CTMs commonly 106 represent consensus values optimized to fit observations of Hg deposition velocities over 107 vegetated surfaces (Selin et al., 2008; L. Zhang et al., 2009; Smith-Downey et al., 2010; 108 H. Zhang et al., 2016), without taking stomatal feedback to environmental conditions 109 into account (Wu et al., 2011; Khan et al., 2019). Consequently, forest tree species-specific 110 stomatal responses to climate change at foliage level are not parameterized in CTMs. 111 An additional problem related to CTMs is the uncertainty of modelled Hg(0) deposition 112 fluxes due to insufficient model evaluation against dry deposition measurements (Feinberg 113 et al., 2022). This issue of model validation was highlighted in a recent revision of GEOS-114 Chem parameterization after matching the GEOS-Chem model design to various exper-115 imental Hg(0) deposition measurements, which resulted in a doubling of the modelled 116 global flux of Hg(0) dry deposition to land compared to previous model outcomes (Feinberg 117 et al., 2022). 118

In this study, we assess the spatial variation of forest foliar Hg uptake fluxes across Europe by producing a spatially highly resolved map of foliar Hg uptake fluxes to European forests using a bottom-up model that incorporates pine tree stomatal responses to climate conditions. We compare these spatially resolved fluxes to forest dry deposition fluxes modelled in GEOS-Chem in order to identify spatial discrepancies between GEOS-Chem and the bottom-up model used here. We investigate the sensitivity of an empirical stomatal response model of pine to different climatic conditions during past growing seasons and for two climate change projections of the years 2068 - 2082 in order to outline the potential of incorporating a stomatal response function into CTMs.

¹²⁸ 2 Materials and Methods

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2.1 Description of datasets

For creating maps of foliar and pine needle Hg uptake fluxes in Europe applying a bottom-up model (Sect. 2.2 and 2.3), we drew on multiple data sources:

• Foliar Hg data. A dataset of foliar Hg uptake rates was derived from Hg mea-132 surements in foliage of tree canopies at 272 forest sites of the UNECE International 133 Co-operative Programme on Assessment and Monitoring of Air Pollution Effects 134 on Forests (ICP Forests). Forest sites are mostly located in Central and North-135 ern Europe (+737 sites in Austria from the Austrian Bio-Indicator Grid) and har-136 monized foliage sampling methods were employed. All foliage samples within this 137 dataset were harvested at the end of the growing seasons 2015 or 2017. Therefore, 138 average foliage values of 2015/2017 constitute reference values of forest foliar Hg 139 uptake fluxes relative to respective fluxes during investigated years of this study. 140 The dataset is publicly available and contains 3569 foliar Hg concentrations of 23 141 tree species and is described in detail in (Wohlgemuth et al., 2022). 142

- Meteorological data. Values on ambient temperature and relative humidity at surface air pressure (1000 hPa) in Europe (spatial resolution: 0.25° x 0.25°) originate from ERA5 hourly reanalysis data and were downloaded from the Copernicus Climate Data Store (Hersbach et al., 2018). The applied time frame includes hourly daytime (07:00 18:00 LT) values during the respective growing seasons (April October) of 1994, 2003, 2015, 2017, and 2018.
- Climate change data. Regional climate simulation data of air temperature and 149 relative humidity at 2 m above surface level for the years 2068 - 2082 and two dif-150 ferent climate change scenarios (Representative Concentration Pathway (RCP) 151 4.5 and RCP 8.5 (IPCC, 2021b)) were obtained from the Coordinated Regional 152 Climate Downscaling Experiment (CORDEX) (Jacob et al., 2020) framework for 153 the European domain with a spatial resolution of $0.11^{\circ} \ge 0.11^{\circ}$ and a temporal res-154 olution of 3hourly daytime (09:00 - 18:00 LT) values. For representing a range of 155 different climate model outputs, we calculated average values from multiple re-156 gional climate models (RCMs) downscaled from global climate models (GCMs) 157 depending on availability for download from the Copernicus Climate Data Store 158 (C3S, 2022). In total, we incorporated data of 15 combinations of 4 RCMs and 159 6 GCMs for RCP 4.5 and of 13 combinations of 6 RCMs and 8 GCMs for RCP 160 8.5 (see Table SI 3) for an overview of models and ensemble members). 161
- **European tree species distribution**. We used a map of spatial proportions of 162 tree species groups per km² land area from (Brus et al., 2012). For use in calcu-163 lating pine foliar Hg uptake fluxes (see Sect. 2.3), we summed up spatial relative 164 abundance values of Pinus sylvestris, Pinus pinaster, Pinus nigra and Pinus halepen-165 sis from European forest inventories (Mauri et al., 2017; Buras & Menzel, 2019) 166 and multiplied these pine relative abundances with the respective total forest area 167 per km^2 derived from (Brus et al., 2012) to obtain pine areal proportions. We per-168 formed the same calculation (sum of values of Pinus sylvestris, Pinus pinaster, Pi-169 nus nigra and Pinus halepensis and subsequent multiplication with respective to-170 tal forest area) to estimate the distribution of pine in Europe under climate change 171 using relative abundance probabilities projected from climate analogues for the 172 time period 2061 - 2090 and RCP 4.5 and RCP 8.5 by (Buras & Menzel, 2019). 173

Leaf Area Indices (LAIs) and Leaf Mass per Area (LMA) values. We
used the LAI satellite product (spatial resolution: 330 m) of PROBA-V (Dierckx
et al., 2014; Fuster et al., 2020) to upscale foliar Hg uptake rates at each ICP Forests
site to foliar Hg uptake fluxes (see Sect. 2.2), along with average LMA values per
tree species from (Forrester et al., 2017).

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2.2 Calculation of forest foliage Hg uptake fluxes

We determined forest foliar Hg uptake fluxes to European forests on a 1 km² spatial resolution applying three basic computational steps: 1) calculation of tree speciesspecific daily Hg uptake fluxes per m² ground area using a bottom-up model; 2) upscaling of respective foliar Hg fluxes per tree species to the European forested area using the areal distribution of corresponding tree species; 3) multiplication of daily forest foliar Hg uptake fluxes per latitude with latitude-dependent growing season length in order to obtain the forest foliar Hg uptake fluxes over one growing season.

Computational step 1) is based on the premise, that foliar Hg uptake rates are tree 187 species-specific (Laacouri et al., 2013; Wohlgemuth et al., 2022; Pleijel et al., 2021). For 188 this reason, we calculated median daily foliar Hg uptake fluxes per tree species group (see 189 Table SI 2 for details) of all forest sites from the foliar Hg dataset (Sect. 2.1). The bottom-190 up modeling approach for calculating daily foliar Hg uptake fluxes from daily foliar Hg 191 uptake rates is described in detail in Wohlgemuth et al., (2020) (Wohlgemuth et al., 2020). 192 Briefly, daily foliar Hg uptake rates per gram foliage dry weight (units of ng Hg $g_{d.w.}^{-1}$ d⁻¹) 193 were multiplied with tree species-specific LMA values (Sect. 2.1) to obtain daily foliar 194 Hg uptake rates per foliage surface area (ng Hg m_{leaf}^{-2} d⁻¹). Subsequently, values of daily 195 foliar Hg uptake rates per foliage surface area are multiplied with values of LAI (m_{leaf}^2) 196 m_{ground}^{-2} ; Sect. 2.1), resulting in daily foliar Hg fluxes per unit ground area (ng Hg m_{ground}^{-2} 197 d^{-1}). LAI values of coniferous forests are relatively constant during the active growing 198 season after the initial growth phase of current-season needles (R. Wang et al., 2017), 199 while LAI values of temperate deciduous forests increase rapidly at the beginning of the 200 growing season (leaf flushing) and climax at peak season (June – August, northern hemi-201 sphere) (Q. Wang et al., 2005). For coniferous tree species, we used the maximum LAI 202 value during the constant period at each forest site of the ICP Forests dataset to cal-203 culate needle foliar Hg uptake fluxes. For deciduous tree species, we calculated foliar Hg 204 uptake fluxes as a temporal sequence at every LAI value available over the growing sea-205 son and subsequently used median foliar Hg uptake flux values of the growing season. 206 For LAI values larger than 3, we applied a species-specific tree height correction factor, 207 to account for lower foliar Hg uptake fluxes of shaded leaves in the lower canopy (Wohlgemuth 208 et al., 2020) (refer to Table SI 1 for utilized tree height correction factors). For conifer-209 ous species, we multiplied Hg uptake fluxes of current-season needles with a species-specific 210 needle age correction factor to account for lower Hg uptake rates of older needle age classes 211 (Wohlgemuth et al., 2020) (refer to Table SI 1 for utilized needle age correction factors). 212

Computational step 2) involves the multiplication of the proportion of each tree 213 species per km² land area with the respective species-specific median daily foliar Hg up-214 take fluxes. We matched tree species-specific Hg data with the areal forest distribution 215 of the respective tree species (Brus et al., 2012). In the few cases of rare European tree 216 species, where specific Hg data was lacking, we pooled Hg or forest distribution data by 217 tree species group (see Table SI 2 for an overview of matched tree species groups between 218 the two datasets). Subsequently, we added up all tree species-specific daily foliar Hg up-219 take fluxes within each km^2 and obtained one forest foliar daily Hg uptake flux per km^2 . 220

In computational step 3) we calculated forest foliar Hg uptake fluxes per km² and one growing season by multiplying each daily foliar Hg uptake flux per km² with the growing season length in days following a simple latitudinal model (CLRTAP, 2017). The latitudinal model of growing season determines a growing season length of 192 days at latitude 50° and decreases by 3.5 days per 1° of latitude moving north and increases by 3.5 days per 1° of latitude moving south.

2.3 Calculation of pine foliar Hg uptake fluxes

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Daily foliar Hg uptake rates of pine tree species were calculated taking into account 228 the empirical dependence of needle Hg uptake fluxes to atmospheric VPD. Pine needle 229 daily Hg uptake rates (upR_{pine}; ng Hg $g_{d.w.}^{-1}$ d⁻¹) were found to be lower at forest sites, 230 where the daytime fraction of water VPD > 1.2 kPa during the respective sample life 231 period (proportion_{dayVPD} > 1.2kPa) was relatively high (Wohlgemuth et al., 2022). The 232 negative correlation of pine needle Hg uptake with timespan of elevated atmospheric VPD 233 was explained by a stomatal closure upon VPD threshold exceedance and thus a high 234 235 stomatal resistance suppressing the diffusive uptake of Hg(0) from the atmosphere. The linear regression of daily foliar Hg uptake rates with proportion_{dayVPD} > 1.2 kPa is: upR_{pine} 236 $= 0.116 - 0.13 \text{ x} (\text{proportion}_{dayVPD} > 1.2 \text{ kPa})$ (Wohlgemuth et al., 2022). We applied 237 this linear relationship to calculate the pine foliar Hg uptake rates of the forest area of 238 Europe during four different growing seasons in 1994, 2003, an average of 2015 and 2017, 239 2018, and projected for the time period 2068 - 2082 under RCP 4.5 and RCP 8.5 (IPCC, 240 2021b). We calculated hourly or 3hourly daytime VPD values from ERA5 (Hersbach et 241 al., 2018) or CORDEX data (Sect. 2.1) on surface temperature and relative humidity 242 using the Auguste-Roche-Magnus formula (W. Yuan, Zheng, et al., 2019) and subsequently 243 determining the fraction of daytime hours when the VPD was above the threshold of 1.2 244 kPa over the respective latitudinal growing season length. Calculations with climate data 245 were performed at sciCORE scientific computing center at University of Basel. We de-246 fined growing season length per latitude using a latitudinal model ((CLRTAP, 2017), see 247 Sect. 2.2). In 2068 - 2082 we assumed the beginning of the growing season to be 3 days 248 earlier and the end of the growing season to be 3 days later to take increases in grow-249 ing season length under climate change into account (Jeong et al., 2011; Garonna et al., 250 2014). The underlying areal distribution of pine is based on European forest inventories 251 and projections of pine abundances based on climate analogues under RCP 4.5 and RCP 252 8.5 by (Buras & Menzel, 2019) (see Sect. 2.1). 253

2.4 GEOS-Chem forest deposition flux calculation

GEOS-Chem is a global 3-D chemistry transport model, which includes a comprehensive Hg cycle (Selin et al., 2008). Table 1 gives an overview of the methodological approach and input parameters for calculating the respective Hg fluxes of GEOS-Chem and the bottom-up model (Sect. 2.2), which we compared spatially in this study.

	bottom-up model	GEOS-Chem
model input pa- rameters	 spatial forest distribution (Brus et al., 2012); leaf area indices (LAIs) (Dierckx et al., 2014; Fuster et al., 2020); leaf mass per area (LMA) (Forrester et al., 2017); meteorological parameter: day-time VPD (Hersbach et al., 2018); foliar Hg uptake rates (Wohlgemuth et al., 2022) 	spatial forest distribution (Gibbs, 2006); leaf area indices (LAIs) (H. Yuan et al., 2011); atmospheric Hg(0) levels (GEOS-Chem v12.8.1 simu- lation 2015); meteorological parameters: air temperature, pressure, solar ra- diation, cloud cover, wind speed (GEOS-FP) (Lucchesi, 2018)
spatial resolution	1 km x 1 km	$0.25 \ge 0.3125^{\circ}$
basic methodolog- ical approach for Hg flux calculation	spatial upscaling of measured foliar Hg uptake rates (Wohlgemuth et al., 2020)	in-series calculation of Hg dry deposition velocity from pa- rameterized resistance values (Wesely, 2007)
foliage stomatal Hg uptake flux component	calculated for pine based on daytime vapor pressure deficit (VPD) values (Sect. 2.3)	calculated within the canopy resistance component as a func- tion of land type, leaf area indices (LAIs), and solar radia- tion
model output com- pared in this study	tree-species specific forest foliar Hg(0) uptake fluxes	Hg(0) dry deposition fluxes to coniferous and deciduous forest land cover

Table 1. Caption

We used an offline version of the GEOS-Chem dry deposition code (Feinberg, 2022) 259 to be able to calculate dry deposition velocities at higher resolution and only for certain 260 land use types (i.e., forest areas). The offline dry deposition code computes deposition 261 velocities using a resistance-based approach (Y. Wang et al., 1998; Wesely, 2007). In-262 put variables (Table 1) are gridded hourly GEOS-FP meteorological data for (e.g., air 263 temperature, wind speed, solar radiation, and cloud cover) and weekly LAI values based 264 on MODIS (H. Yuan et al., 2011) for the year 2015. The model calculates the Hg(0) dry 265 deposition velocity based on species-specific parameters including its biological reactiv-266 ity (f₀ = 10⁻⁵) and Henry's Law Constant (H^{*} = 0.11 M atm⁻¹). To isolate the uptake 267 of Hg(0) to forests, we calculated the dry deposition velocity only over coniferous and 268 deciduous land cover types from the Olson land map (Gibbs, 2006). The offline calcu-269 lations output hourly dry deposition velocities over the European domain at $0.25 \times 0.3125^{\circ}$ 270 resolution. We converted the calculated Hg(0) deposition velocities to fluxes by multi-271 plying with hourly surface Hg(0) concentrations from a GEOS-Chem v12.8.1 simulation 272 for 2015. For this study, we compared the GEOS-Chem Hg(0) dry deposition fluxes to 273 forests with foliar Hg(0) uptake fluxes calculated using the bottom-up model. For both 274 models, Hg fluxes were averaged over the latitude-dependent growing season length in 275 days and cropped to the same spatial extent. As GEOS-Chem and the bottom-up model 276 differ in their geographic resolution (GEOS-Chem: $0.25^{\circ} \ge 0.31^{\circ} \sim 955 \text{ km}^2$ vs. bottom-277 up: 1 km^2), we downsampled daily forest foliar Hg uptake fluxes from the bottom-up 278 model through bilinear interpolation. 279

2.5 Uncertainty analysis of foliar Hg uptake fluxes

The relative uncertainty value per tree species group depended on propagated un-281 certainties of calculation parameters used to derive the respective foliar Hg uptake flux 282 per tree species group (see Table SI 4 for details and values). Subsequently, we calcu-283 lated one relative uncertainty value per geographic tile of our European flux map (Fig. 284 1) by summarizing the relative uncertainty of each foliar Hg uptake flux per tree species 285 group within each tile according to error propagation principles (Ku, 1966; Papula, 2003). 286 We obtained the relative uncertainty for the total foliar Hg uptake flux to European forests 287 (Fig. 1) by propagating all relative uncertainty values per tile. The final relative uncer-288 tainty value of total foliar Hg uptake flux to European forests and the reference grow-289 ing seasons 2015/2017 is 0.52. 290

²⁹¹ **3** Results and Discussion

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3.1 Spatial distribution of forest foliar Hg uptake fluxes across Europe

Figure 1 visualizes forest foliar Hg uptake fluxes per growing season at a spatial 293 resolution of 1 km^2 (g Hg km⁻² season⁻¹) in Europe. Forest foliar Hg uptake fluxes gen-294 erally follow a spatial distribution of European forests, because this map (Fig. 1) is based 295 on the proportion of forest tree species per land area (Brus et al., 2012). Consequently, 296 the largest forest foliage Hg uptake fluxes in terms of area are on the Scandinavian Penin-297 sula with dense forest land cover. Outside of Scandinavia, forest foliage Hg uptake fluxes 298 fall along large contiguous forested areas, e.g. in the Carpathian Mountains, the South-299 Eastern Alps, the Balkans, or forested low mountain areas like the Black Forest. 300



Figure 1. Spatial distribution of forest foliar Hg uptake fluxes (g Hg km² growing season⁻¹) to Europe based on a bottom-up extrapolation of foliar Hg concentrations, that were measured and averaged over the 2015 and 2017 growing seasons. Dark grey areas represent excluded non-forested areas (e.g. surface waters or non-vegetated mountain areas).

The sum of forest foliar Hg uptake fluxes over the land area of Europe as displayed in Figure 1 equals 23 ± 12 Mg Hg season⁻¹. This total flux agrees within uncertainty

with a previous estimate for the total foliar Hg uptake flux to Europe of 20 ± 3 Mg Hg 303 over the 2018 growing season based on foliar Hg uptake fluxes at four forested sites (Wohlgemuth 304 et al., 2020). (Zhou & Obrist, 2021) evaluated a median global foliar Hg assimilation of 305 28 Mg yr⁻¹ for deciduous broadleaf forests and 61 Mg yr⁻¹ for evergreen needleleaf forests 306 by combining foliar Hg concentrations with annual net foliar biomass production data 307 of the respective forest types. From these global assimilation estimates by (Zhou & Obrist, 308 2021), we calculated a total foliar Hg assimilation of 29 Mg yr^{-1} to the deciduous and 309 coniferous forest land area of Europe (for details see SI, Text S1), which is slightly higher 310 but still within the uncertainty of the 23 ± 12 Mg Hg season⁻¹ from this study. How-311 ever, foliar Hg uptake fluxes based on net primary foliar biomass production by Zhou 312 and Obrist, (2021) (Zhou & Obrist, 2021) does not correct for lower foliar Hg uptake rates 313 by shade leaves and multiyear old needles (see Sect. 2.2) relative to sun leaves and younger 314 needles (Wohlgemuth et al., 2020), likely resulting in a systematic over-estimation. We 315 assume, that the different time reference (seasonal vs. annual) of the flux from this study 316 $(23 \pm 12 \text{ Mg Hg season}^{-1})$ and the flux derived from Zhou and Obrist, (2021) (Zhou & 317 Obrist, 2021) (29 Mg Hg yr⁻¹) only plays a minor role for explaining the difference be-318 tween the two fluxes, since we expect a small net foliar biomass production in Europe 319 in winter outside of the growing season. 320

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3.2 Pine foliar Hg uptake fluxes under different VPD scenarios

Figure 2 shows total pine forest foliar Hg uptake fluxes to Europe calculated un-322 der different conditions of atmospheric surface-level water VPD during four past grow-323 ing seasons (1994, 2003, 2015/2017, 2018) and simulated for the years 2068 - 2082 as an 324 average of multiple climate model outputs (see Sect. 2.2) under two different climate change 325 scenarios (RCP 4.5 and RCP 8.5). The leftmost bar (Fig. 2) represents a theoretical base-326 line pine needle Hg uptake flux in absence of VPD induced stomatal control (potential 327 maximum transpiration rates) on the pine needle Hg uptake flux. The total pine nee-328 dle Hg uptake flux to Europe during the reference growing season 2015/2017 (Sect. 2.2) 329 is 9.3 \pm 3.7 Mg Hg representing 70% of the baseline flux of 13.3 \pm 5.3 Mg Hg season⁻¹. 330 Thus, based on the pine needle Hg uptake model used in this study (Sect. 2.3), the VPD 331 effect reduces the total pine needle Hg uptake flux to Europe by approximately 30%. 332



Figure 2. Pine needle Hg uptake flux to European pine forests (Mg Hg season⁻¹) calculated from atmospheric surface water vapor pressure deficit (VPD) conditions during the growing seasons 1994, 2003, 2015/2017, 2018 and projected for the years 2068 - 2082 under RCP 4.5 and RCP 8.5. Bar on the left represents a baseline pine forest needle Hg uptake flux with no VPD exceedance of 1.2 kPa throughout the growing season.

The relative standard deviation of modelled total pine needle Hg uptake fluxes for 333 the investigated growing seasons (1994, 2003, 2015/2017, 2018, 2068 - 2082) was 0.07. 334 Consequently, modelled total European pine needle Hg uptake fluxes hardly differed from 335 each other among growing seasons. The total pine needle Hg uptake flux in Europe de-336 pend on VPD conditions in areas where pine forests prevail. Pine forests are primarily 337 located in Northern Europe (SI Fig. 1), where hourly ambient VPD was > 1.2 kPa dur-338 ing 30% or less of daytime in the growing seasons 1994, 2003 and 2015/2017 due to rel-339 atively cool and moist ambient conditions as compared to Central and Southern Europe 340 (see e.g. VPD conditions during reference time period 2015/2017 Fig. 3a). In contrast 341 to previous years, the European summer hydrological condition of 2018 has been described 342 as an intense hot drought, during which pronounced stomatal closure of coniferous forests 343 in response to high VPD were recorded in Switzerland (Gharun et al., 2020). In South-344 ern Fennoscandia, conditions of ambient hourly VPD > 1.2 kPa prevailed over excep-345 tionally long time proportions (around 40%) during the summer of 2018 (see Fig. 3b, 346 (Buras et al., 2020)). As a result, the modelled total pine needle Hg uptake flux in Eu-347 rope in 2015/2017 (9.3 Mg Hg season⁻¹) was by a factor of 1.16 higher than the respec-348 tive flux in 2018 (8.0 Mg Hg season⁻¹). We conclude that hot and dry summer condi-349 tions (Fig. 2) in Fennoscandia crucially impact modelled past total pine needle Hg up-350 take fluxes in Europe. According to the model results, an average amount of 1.3 Mg Hg 351 was not deposited via pine needle uptake in 2018 compared to 2015/17, potentially re-352 maining in the atmosphere, where it can be long-range transported to the ocean (Zhou 353 et al., 2021). These 1.3 Mg Hg are more than three times larger than the reported an-354 thropogenic Hg emissions of Sweden in 2021 ??, highlighting the quantitative impact, 355 that hot droughts can have on the pine needle Hg uptake flux. 356



Figure 3. Average daytime proportion of surface level atmospheric water VPD > 1.2 kPa during a) the reference growing season 2015/2017, and b) the growing season 2018. All VPD values were calculated from hourly reanalysis data of ERA5 ambient air temperature and relative humidity (Sect. 2.1).

3.3 Projected pine forest needle Hg uptake fluxes under climate change scenarios

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The projected total pine forest needle Hg uptake flux for 2068 - 2082 (RCP 4.5: 9.3 359 \pm 5.5 Mg Hg season⁻¹; RCP 8.5: 8.1 \pm 4.9 Mg Hg season⁻¹) was in the same range as 360 the corresponding average flux for the years 1994, 2003, 2015 and 2017 of 9.1 \pm 0.2 Mg 361 Hg season⁻¹ (mean \pm sd), but slightly higher than the corresponding flux in the year 362 of 2018 (8.0 \pm 3.2 Mg Hg season⁻¹), during which Fennoscandia experienced a summer 363 of relatively long hot and dry ambient conditions. Figure 4 maps the absolute deviation 364 of the pine forest needle Hg uptake flux projected for 2068 - 2082 (simulated future flux) 365 from the corresponding 2018 flux in Europe. Under RCP 4.5, the simulated future flux 366 is higher (blue area in Fig. 4 a) than the 2018 flux in 65% of total area. Under RCP 8.5, 367 the simulated future flux is higher (blue area in Fig. 4 b) than the 2018 flux in 43% of 368 total area. In most area of Fennoscandia, where a majority of pine forests in Europe are 369 located (SI Fig. 1a), the future flux is projected to be larger than in 2018. For both cli-370 mate change scenarios, the projection predicts lower pine needle Hg fluxes to the Balkans 371 and to the Southern Iberian Peninsula than in 2018 (Fig. 4). 372



Figure 4. Absolute deviation of projected pine forest foliar Hg uptake fluxes for 2068 - 2082 (under RCP 4.5 (a) and RCP 8.5 (b)) from the corresponding flux modelled for 2018. In blue areas, the projected future flux under the two climate change scenarios is higher than the respective 2018 flux, in orange areas, this deviation is reversed.

The pine forest needle Hg uptake flux for 2068 - 2082 simulated here is a function 373 of both modelled ambient VPD conditions during the growing season and the projected 374 distribution of pine forests in Europe depending on climate analogs (Buras & Menzel, 375 2019). While the pine forest cover in Southern Sweden is projected to decrease under 376 the climate change scenarios RCP 4.5 and RCP 8.5 from around 50% km⁻² to around 377 25% km⁻², forest cover in Central and Northern Fennoscandia is projected to be rela-378 tively steady for climate analogs of both climate change scenarios (compare SI Fig. 1 a 379 - c). Average long-term precipitation rates are projected to increase in Scandinavia, along 380 with a decrease of meteorological drought in the coming decades under different climate 381 change scenarios (Forzieri et al., 2014; Samaniego et al., 2018; Kellomäki et al., 2018), 382 which could result in an increase of atmospheric humidity and a decrease of VPD in north-383 ern Europe (Oksanen et al., 2019). Under this scenario of wetter forest environments, the Hg sink of Scandinavian pine forest needles would not be significantly diminished. 385 However, drought trends in Fennoscandia are still inconsistent and extreme drought events 386 like in 2018 might occur more frequently under the current rate of climate change (IPCC, 387 2021a). The summer of 2018 was a record hot drought in Europe (Buras et al., 2020), 388 while climate simulations for 2068 - 2082 are averaged over multiple climate models (SI 389 Table 3), possibly averaging out extreme events. In a scenario, where the maximum pro-390 portion of daytime VPD > 1.2 kPA per growing season averaged over 2068 - 2082 pre-391 vails at each spatial unit, the total pine forest needle Hg uptake flux to Europe reduces 392 to 6.9 Mg Hg season⁻¹ for RCP 4.5 and 5.0 Mg Hg season⁻¹ for RCP 8.5, which cor-393 responds to 74% and 62% of the respective flux derived from an average VPD daytime 394 proportion. We therefore suggest that extreme climate events of extended time periods 395 of ambient daytime VPD > 1.2 kPa like during the growing season 2018 (Fig. 3b) could 396 reduce the pine forest needle Hg uptake flux in Fennoscandia in future even compared 397 to average long-term VPD projections (Fig. 4). 398

A source of model uncertainty of the future forest foliar Hg uptake flux under climate change arises from atmospheric Hg(0) concentrations that depend on anthropogenic emissions, re-emissions of mobilized legacy Hg and future global deposition fluxes under climate and land use change (Sonke et al., 2023; Feinberg et al., 2023), which we could not account for in this study. However, our model outputs call attention to the sensi-

tivity of the pine needle Hg uptake flux to extreme hot and dry ambient conditions, which 404 should be accounted for in chemistry-transport models under varying atmospheric Hg(0)405 levels. The impact of the hot and dry conditions on the pine Hg uptake fluxes might have 406 implications for Hg inputs into aquatic ecosystems. In a recent review on Hg cycling in the context of global change, (Sonke et al., 2023) highlighted the potential of legacy Hg 408 (i.e. actively cycling Hg that was mobilized in the past) to cause contamination by mo-409 bilization of Hg from soils to wetlands and coastal ecosystems via riverine systems. While 410 most soil Hg enters riverine systems by soil erosion from agricultural lands, contaminated 411 sites, and deforested woodland (Panagos et al., 2021; Sonke et al., 2023), a reduced for-412 est foliar Hg uptake and subsequent deposition to forest soils may decrease the amount 413 of runoff Hg from forest soils in the long-term, while long-range Hg transport to the open 414 ocean via the atmosphere might be enhanced (Zhou et al., 2021). 415

3.4 Comparison of bottom-up model with GEOS-Chem

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Figure 5 depicts spatial ratios of daily forest Hg uptake fluxes of the bottom-up model to GEOS-Chem. Absolute difference values of the two model outputs are shown in SI Fig. 2.



Figure 5. Ratios per spatial unit of daily forest Hg uptake fluxes averaged over the latitude-specific growing season length of the bottom-up model to GEOS-Chem.

Results of average daily foliar Hg uptake fluxes from GEOS-Chem and the bottom-420 up model were geographically comparable: In 59% of the spatial domain with values > 421 0, average daily foliar Hg uptake fluxes from the two models differed by factor of 1 - 2422 from each other, in 37% of the domain, model values differed by a factor of 2 - 10 from 423 each other, and in 4% of the domain respective values differed by a factor of > 10 from 424 each other (Fig 5). We examined if differences in modelled average daily foliar Hg up-425 take fluxes at the same geographic location originate from differences in the underlying forest distribution maps of the two compared models. In 78% of spatial tiles with val-427 ues > 0, the ratio of average daily foliar Hg uptake fluxes of the bottom-up model to GEOS-428 Chem agreed in range (Fig. 5) with the ratio of the forest fraction of the bottom-up model 429

to GEOS-Chem per respective spatial tile. We thus hypothesize that the bottom-up model 430 and GEOS-Chem generally produce similar foliar Hg flux values per spatial unit given 431 the same forest distribution. Reasons for minor differences in model outputs are chal-432 lenging to identify, since the two models are based on different approaches, parameters 433 and underlying maps (Sect. 2.4). For future assessment of model accuracy, we therefore 434 suggest to compare model results to actual measurements of the forest foliar Hg uptake 435 flux (Obrist et al., 2021; Feinberg et al., 2022). The total foliar Hg uptake flux to the 436 European forested area (Fig. 1 and 5) was 22 Mg Hg season⁻¹ for GEOS-Chem which 437 almost equals the total flux of 23 ± 12 Mg Hg season⁻¹ for the bottom-up model (Sect. 438 3.1).439

4 Conclusion

440

We created a highly resolved (1 km^2) map (Fig. 1), which visualizes the spatial vari-441 ation of foliar Hg uptake fluxes to European forests. The highest foliar Hg uptake fluxes 442 receive Fennoscandia, densely forested areas in Central and Southern Europe, e.g. the 443 Carpathian Mountains, the Balkans, or multiple low mountain areas. We suggest, that 444 this map (Fig. 1) can guide decisions on European background Hg monitoring of the terrestrial environment. The total forest foliar Hg uptake flux over the course of one grow-446 ing season agrees well with Hg flux estimates derived from literature and from the chem-447 ical transport model GEOS-Chem for the same land area of Europe (Fig. 5). This pre-448 cision among modelling results on a European scale using different approaches gives us 449 confidence that the bottom-up model is overall able to represent the seasonal forest fo-450 liar Hg uptake flux. We suggest that the accuracy of modelling results have to be fur-451 ther determined using direct forest foliar Hg flux measurements. 452

Using an empirical relationship between Hg needle uptake rates of pine trees and 453 VPD threshold exceedance, we found a reduction in modelled pine forest needle Hg up-454 take flux during the relatively hot and dry growing season in Fennoscandia in 2018 com-455 pared to the growing seasons in 1994, 2003 and 2015/2017 (Fig. 2). The modelled av-456 erage amount of Hg, that was not deposited via pine needle uptake in 2018 compared 457 to the reference time period of 2015/17 exceeded the reported anthropogenic Hg emis-458 sions of Sweden in 2021, highlighting the quantitative significance of stomatal Hg up-459 take. If these hot summer droughts occurred more frequently in Fennoscandia under cli-460 mate change, the pine forest needle Hg uptake flux would be diminished while these ex-461 treme conditions prevail, potentially increasing the Hg burden of the ocean via long-range 462 atmospheric transport. In order to better represent the impact of extreme climate events 463 on the pine forest needle Hg uptake flux, we therefore advise to incorporate a stomatal 464 component of the pine needle Hg uptake flux into chemical transport models like GEOS-465 Chem. 466

5 Open Research 467

Calculated forest foliar Hg uptake fluxes to Europe (Fig. 1) and GEOS-Chem sim-468 ulation data aggregated to seasonal values are publicly available for download from Zen-469 odo at https://zenodo.org/record/7851718#.ZFUeLM5Bw2w and https://zenodo.org/ 470 record/7900753#.ZFUgqM5Bw2w respectively. All input datasets to the bottom-up model 471 are described in detail in Section 2.1, along with their respective publications and databases, 472 from which the datasets can be accessed. The offline dry deposition code from GEOS-473 Chem is accessible for download (Feinberg, 2022) and model output data from GEOS-474 Chem can be obtained from the corresponding author upon request. All calculations and 475 visualizations were done in R, Version 4.0.3. 476

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- Part of the ICP Forests data was co-financed by the European Commission. ERA5 and
- 486 CORDEX data were obtained from the Copernicus Climate Change Service (C3S) Cli-
- mate Data Store (https://cds.climate.copernicus.eu/cdsapp#!/home). PROBA-V leaf
- area index values were downloaded from the VITO Product Distribution Portal. Climate
- calculations under two climate change scenarios (Fig. 2, Fig. 3) were performed at sci-
- ⁴⁹⁰ CORE (http://scicore.unibas.ch/) scientific computing center at University of Basel.

491 References

- Allen, C. D., Breshears, D. D., & McDowell, N. G. (2015). On underestimation of global vulnerability to tree mortality and forest die-off from hotter drought in the Anthropocene. *Ecosphere*, 6(8), art129. Retrieved 2021-10-22, from https://onlinelibrary.wiley.com/doi/abs/10.1890/ES15-00203.1 doi: 10.1890/ES15-00203.1
 Berg, A., & Sheffield, J. (2018, June). Climate Change and Drought: the Soil Mois-
- 498
 ture Perspective.
 Curr Clim Change Rep. 4(2), 180–191.
 Retrieved 2021

 499
 10-22, from https://doi.org/10.1007/s40641-018-0095-0
 doi: 10.1007/

 500
 s40641-018-0095-0
- Blackwell, B. D., Driscoll, C. T., Maxwell, J. A., & Holsen, T. M. (2014, June).
 Changing climate alters inputs and pathways of mercury deposition to forested
 ecosystems. *Biogeochemistry*, 119(1-3), 215–228. Retrieved 2020-03-10,
 from http://link.springer.com/10.1007/s10533-014-9961-6
 doi: 10.1007/s10533-014-9961-6
- Brando, P. M., Paolucci, L., Ummenhofer, C. C., Ordway, E. M., Hartmann,
- H., Cattau, M. E., ... Balch, J. (2019, May). Droughts, Wildfires, and Forest Carbon Cycling: A Pantropical Synthesis. Annu. Rev. Earth Planet. Sci., 47(1), 555–581. Retrieved 2021-10-22, from https:// www.annualreviews.org/doi/10.1146/annurev-earth-082517-010235
 doi: 10.1146/annurev-earth-082517-010235
- Brus, D. J., Hengeveld, G. M., Walvoort, D. J. J., Goedhart, P. W., Heidema, A. H.,
 Nabuurs, G. J., & Gunia, K. (2012, January). Statistical mapping of tree
 species over Europe. European Journal of Forest Research, 131(1), 145–
 157. Retrieved 2019-08-26, from http://link.springer.com/10.1007/
 \$10342-011-0513-5 doi: 10.1007/s10342-011-0513-5
- Buras, A., & Menzel, A. (2019). Projecting Tree Species Composition Changes
 of European Forests for 2061–2090 Under RCP 4.5 and RCP 8.5 Scenarios. Frontiers in Plant Science, 9, 1986. Retrieved 2021-09-22, from
 https://www.frontiersin.org/article/10.3389/fpls.2018.01986 doi:
 10.3389/fpls.2018.01986
- Buras, A., Rammig, A., & Zang, C. S. (2020, March). Quantifying impacts of the
 2018 drought on European ecosystems in comparison to 2003. *Biogeosciences*,
 17(6), 1655–1672. Retrieved 2021-12-07, from https://bg.copernicus.org/
 articles/17/1655/2020/ doi: 10.5194/bg-17-1655-2020
- C3S. (2022). Copernicus Climate Change Service, Climate Data Store (CDS):
 CORDEX regional climate model data on single levels. Retrieved from
 https://cds.climate.copernicus.eu/cdsapp#!/dataset/10.24381/
 cds.bc91edc3?tab=overview doi: 10.24381/cds.bc91edc3

530	CLRTAP. (2017). Revised Chapter 3 of the Manual on Methodologies and Cri-
531	teria for Modelling and Mapping Critical Loads and Levels and Air Pol-
532	lution Effects, Risks and Trends: Mapping Critical Levels for Vegetation.
533	Umweltbundesamt. Retrieved from https://www.umweltbundesamt.de/en/
534	manual-for-modelling-mapping-critical-loads-levels doi: $https://www$
535	. unweltbundes amt. de/en/manual-for-modelling-mapping-critical-loads-levels
536	Demers, J. D., Blum, J. D., & Zak, D. R. (2013, March). Mercury isotopes in a
537	forested ecosystem: Implications for air-surface exchange dynamics and the
538	global mercury cycle. Global Biochem. Cy., 27(1), 222–238. Retrieved 2018-
539	11-14, from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/
540	gbc.20021 doi: 10.1002/gbc.20021
541	Dierckx, W., Sterckx, S., Benhadj, I., Livens, S., Duhoux, G., Van Achteren, T.,
542	Saint, G. (2014, April). PROBA-V mission for global vegetation
543	monitoring: standard products and image quality. Int. J. Remote Sens.,
544	35(7), $2589-2014$. Retrieved 2021-08-30, from https://doi.org/10.1080/
545	01431161.2014.883097 doi: 10.1080/01431161.2014.883097
546	Enrico, M., Roux, G. L., Marusczak, N., Heimburger, LE., Claustres, A., Fu, X.,
547	dominated by gasaous elemental marcury dry deposition <u>Fraviran</u> Sai Tach
548	a dominated by gaseous elemental mercury dry deposition. Entrino 3.52 and 5.62 and 5.62 and 2.61 and 2.62 from http://puba.aca.org/
549	doi/(10, 1021/acg, est, 5b06058, doi: 10.1021/acg, est, 5b06058
550	Four $A = (2022)$ Code reference: prifein /offline dryden: Offline dry denosition
551	model from GEOS-Chem v1.0 (v1.0)
553	doi: https://doi.org/10.5281/zenodo.6498126
555	Feinberg A Dlamini T Jiskra M Shah V & E Selin N (2022) Eval-
555	uating atmospheric mercury (Hg) uptake by vegetation in a chemistry-
556	transport model. Environmental Science: Processes & Impacts. Retrieved
557	2022-07-25. from https://pubs.rsc.org/en/content/articlelanding/
558	2022/em/d2em00032f (Publisher: Royal Society of Chemistry) doi:
559	10.1039/D2EM00032F
560	Feinberg, A., Jiskra, M., Borrelli, P., Biswakarma, J., & Selin, N. E. (2023, Jan-
561	uary). Land use change as an anthropogenic driver of mercury pollution. Re-
562	trieved 2023-04-15, from https://eartharxiv.org/repository/view/4963/
563	(Publisher: EarthArXiv)
564	Forrester, D. I., Tachauer, I. H. H., Annighoefer, P., Barbeito, I., Pretzsch, H., Ruiz-
565	Peinado, R., Sileshi, G. W. (2017, July). Generalized biomass and leaf area
566	allometric equations for European tree species incorporating stand structure,
567	tree age and climate. Forest Ecology and Management, 396, 160–175. Re-
568	trieved 2020-06-10, from http://www.sciencedirect.com/science/article/
569	pii/S0378112717301238 doi: 10.1016/j.foreco.2017.04.011
570	Forzieri, G., Feyen, L., Rojas, R., Flörke, M., Wimmer, F., & Bianchi, A. (2014,
571	January). Ensemble projections of future streamflow droughts in Europe.
572	Hydrology and Earth System Sciences, 18(1), 85–108. Retrieved 2021-12-17,
573	from https://hess.copernicus.org/articles/18/85/2014/ (Publisher:
574	Copermicus GmbH) doi: $10.5194/ness-18-85-2014$
575	Fuster, B., Sanchez-Zapero, J., Camacho, F., Garcia-Santos, V., Verger, A., La-
576	LAL fADAD and fCOVED collection 200 m products of Conomicus Clobal
577	LAI, IAPAR and ICOVER conection 500 in products of Coperincus Global
578	Land Service. Include Sens. $12(0)$, 1017. Retrieved 2021-03-17, Irollin https://www.mdpi.com/2072-4292/12/6/1017. doi: 10.3200/rs12061017
5/9	Caronna I Jong R d Wit Δ I W d Milehor C Λ Schmid R & Schoon
58U	man M E (2014) Strong contribution of autumn phenology to changes in
582	man, m. E. (2011). Shong controlution of autumn phonology to changes in
	satellite-derived growing season length estimates across Europe (1982–2011)
583	satellite-derived growing season length estimates across Europe (1982–2011). <i>Glob. Change Biol.</i> , 20(11), 3457–3470. Retrieved 2019-10-17 from
583 584	satellite-derived growing season length estimates across Europe (1982–2011). <i>Glob. Change Biol.</i> , 20(11), 3457–3470. Retrieved 2019-10-17, from https://onlinelibrary.wiley.com/doi/abs/10.1111/gcb.12625 doi:

585	$10.1111/{ m gcb}.12625$
586	Gharun, M., Hörtnagl, L., Paul-Limoges, E., Ghiasi, S., Feigenwinter, I., Burri,
587	S., Buchmann, N. (2020, September). Physiological response of Swiss
588	ecosystems to 2018 drought across plant types and elevation. <i>Philosophical</i>
589	Transactions of the Royal Society B: Biological Sciences, 375(1810), 20190521.
590	Retrieved 2023-04-16, from https://royalsocietypublishing.org/
591	doi/full/10.1098/rstb.2019.0521 (Publisher: Royal Society) doi:
592	10.1098/rstb.2019.0521
593	Gibbs, H. K. (2006). Olson's Major World Ecosystem Complexes Ranked by Carbon
594	in Live Vegetation: an Updated Database Using the GLC2000 Land Cover
595	Product (NDP-017b).
596	doi: https://www.osti.gov/biblio/1389498
597	Grossiord, C., Buckley, T. N., Cernusak, L. A., Novick, K. A., Poulter, B., Sieg-
598	wolf, R. T. W., McDowell, N. G. (2020). Plant responses to rising vapor
599	pressure deficit. New Phytol., 226(6), 1550–1566. Retrieved 2020-09-07, from
600	https://nph.onlinelibrary.wiley.com/doi/abs/10.1111/nph.16485 doi:
601	10.1111/nph.16485
602	Hararuk, O., Obrist, D., & Luo, Y. (2013, April). Modelling the sensitivity of
603	soil mercury storage to climate-induced changes in soil carbon pools. Bio-
604	<i>geosciences</i> , 10(4), 2393–2407. Retrieved 2021-10-27, from https://bg
605	.copernicus.org/articles/10/2393/2013/ doi: 10.5194/bg-10-2393-2013
606	Hersbach, H., Bell, B., Berrisford, P., Biavati, G., Horányi, A., Muñoz Sabater, J.,
607	Thépaut, JN. (2018). ERA5 hourly data on pressure levels from 1979
608	to present. Copernicus Climate Change Service (C3S) Climate Data Store
609	(CDS).
610	doi: 10.24381/cds.bd0915c6
611	IPCC. (2021a). Climate change 2021. The physical science basis. Working group I
612	contribution to the Sixth Assessment Report of the Intergovernmental Panel
613	on Climate Change.
614	IPCC. (2021b). Summary for Policymakers. Climate Change 2021: The Physical
615	Science Basis. Contribution of Working Group I to the Sixth Assessment Re-
616	port of the Intergovernmental Panel on Climate Change. Cambride University
617	Press.
618	Jacob, D., Teichmann, C., Sobolowski, S., Katragkou, E., Anders, L. Belda, M.,
619	Wulfmever, V. (2020, April). Regional climate downscaling over Eu-
620	rope: perspectives from the EURO-CORDEX community. Reg Environ
621	<i>Change</i> , 20(2), 51. Retrieved 2021-11-03, from https://doi.org/10.1007/
622	s10113-020-01606-9 doi: 10.1007/s10113-020-01606-9
623	Jeong, SJ., Ho, CH., Gim, HJ., & Brown, M. E. (2011). Phenology shifts at
624	start vs. end of growing season in temperate vegetation over the Northern
625	Hemisphere for the period 1982–2008. Global Change Biology, 17(7), 2385–
626	2399. Retrieved 2021-12-20, from https://onlinelibrary.wiley.com/doi/
627	abs/10.1111/j.1365-2486.2011.02397.x doi: 10.1111/j.1365-2486.2011
628	.02397.x
629	Jiskra, M., Sonke, J. E., Obrist, D., Bieser, J., Ebinghaus, R., Myhre, C. L.,
630	Dommergue, A. (2018). A vegetation control on seasonal variations
631	in global atmospheric mercury concentrations. Nat. Geosci., 1–7. Re-
632	trieved from http:https://doi.org/10.1038/s41561-018-0078-8 doi:
633	10.1038/s41561-018-0078-8
634	Jiskra, M., Wiederhold, J. G., Skyllberg, U., Kronberg, RM., Hajdas, I., & Kret-
635	zschmar, R. (2015). Mercury deposition and re-emission pathways in boreal
636	forest soils investigated with Hg isotope signatures. Environ. Sci. Technol.,
637	<i>49</i> (12), 7188–7196.
638	Kellomäki, S., Strandman, H., Heinonen, T., Asikainen, A., Venäläinen, A., & Pel-
639	tola, H. (2018, March). Temporal and Spatial Change in Diameter Growth

640	of Boreal Scots Pine, Norway Spruce, and Birch under Recent-Generation
641	(CMIP5) Global Climate Model Projections for the 21st Century. Forests,
642	9(3), 118. Retrieved 2021-12-16, from https://www.mdpi.com/1999-4907/9/
643	3/118_doi: 10.3390/f9030118
644	Khan, T., Obrist, D., Agnan, Y., E. Selin, N., & A. Perlinger, J. (2019).
645	Atmosphere-terrestrial exchange of gaseous elemental mercury: parameteriza-
646	tion improvement through direct comparison with measured ecosystem fluxes
040	Environmental Science: Processes & Imnacte 21(10) 1600-1712 Botrioved
647	2021 10 20 from https://www.arg.arg/out.org/out.org/arg/arg/arg/
648	2021-10-20, from https://pubs.rsc.org/en/content/articlelanding/
649	2019/em/c9em00341j doi: 10.1039/C9EM00341J
650	Ku, H. (1966). Notes on the use of propagation of error formulas. Journal of Re-
651	search of the National Bureau of Standards, $70C(4)$.
652	Körner, C. (2013). Plant–Environment Interactions. In A. Bresinsky, C. Körner,
653	J. W. Kadereit, G. Neuhaus, & U. Sonnewald (Eds.), Strasburger's Plant
654	Sciences: Including Prokaryotes and Fungi (pp. 1065–1166). Berlin, Hei-
655	delberg: Springer. Retrieved 2019-11-25, from https://doi.org/10.1007/
656	978-3-642-15518-5 12
657	Laacouri A Nater E A & Kolka B K (2013 September) Distribution and
650	untaka dynamics of moreury in loaves of common deciduous tree species in
058	Minnesota USA Environ Sai Technol $\frac{7}{12}$ 10462 10470 Botrioved
659	$\begin{array}{c} \text{Nillinesota, 0.5.A.} \text{Entrion. Sci. rectifiel., 47(10), 10402-10470.} \text{Retrieved} \\ 2010 06 20 from http://www.sci.exe.exe.exe.exe.exe.exe.exe.exe.exe.ex$
660	2019-00-20, from http://pubs.acs.org/doi/abs/10.1021/es4013572 doi:
661	10.1021/es401357z
662	Lagergren, F., & Lindroth, A. (2002, August). Transpiration response to soil mois-
663	ture in pine and spruce trees in Sweden. Agr. Forest Meteorol., $112(2)$, 67–85.
664	Retrieved 2021-08-31, from https://www.sciencedirect.com/science/
665	article/pii/S0168192302000606 doi: $10.1016/S0168-1923(02)00060-6$
666	Liu, Y., Liu, G., Wang, Z., Guo, Y., Yin, Y., Zhang, X., Jiang, G. (2021, Octo-
667	ber). Understanding foliar accumulation of atmospheric Hg in terrestrial veg-
668	etation: Progress and challenges. Critical Reviews in Environmental Science
669	and Technology, $\theta(0)$, 1-22. Retrieved 2021-10-18, from https://doi.org/10
670	.1080/10643389.2021.1989235 doi: 10.1080/10643389.2021.1989235
671	Lucchesi B (2018) File Specification for GEOS FP GMAO Office Note No. 4 (Ver-
671	sion 1.2) Batriaved from http://gmaa.gsfc.paga.gov/pubs/office.patas
072	Mauri A Strong C & San Miguel Awang L (2017 January) EU Fanget a high
673	Mauri, A., Strona, G., & San-Wiguel-Ayanz, J. (2017, January). EU-Forest, a high-
674	resolution tree occurrence dataset for Europe. Sci Data, 4. Retrieved 2020-01-
675	29, from https://www.ncbi.nlm.nih.gov/pmc/articles/PMC5215058/ doi:
676	10.1038/sdata.2016.123
677	Norby, R. J., & Zak, D. R. (2011). Ecological Lessons from Free-Air CO2
678	Enrichment (FACE) Experiments. Annual Review of Ecology, Evo-
679	lution, and Systematics, $42(1)$, 181–203. Retrieved 2021-10-22, from
680	https://doi.org/10.1146/annurev-ecolsys-102209-144647 doi:
681	10.1146/annurev-ecolsys-102209-144647
682	Obrist, D., Agnan, Y., Jiskra, M., Olson, C. L., Colegrove, D. P., Hueber, J.,
683	Helmig, D. (2017, July). Tundra uptake of atmospheric elemental mercury
684	drives Arctic mercury pollution. Nature, 547(7662), 201–204. Retrieved
685	2019-06-24 from http://www.nature.com/articles/nature22997 doi:
696	10 1038/nature22997
080	Obriet D. Dey, F. M. Harrison, I. I. Kryong, C. F. Mungor, I. W. Moogmüller
687	UDISt, D., Roy, E. M., Harrison, J. L., Rwong, C. F., Munger, J. W., Mooshuner,
688	π ., Commane, κ . (2021, July). Previously unaccounted atmospheric mer-
689	cury deposition in a midiatitude deciduous forest. <i>PNAS</i> , 118(29). Retrieved
690	2021-07-26, from https://www.pnas.org/content/118/29/e2105477118 doi:
691	10.1073/pnas.2105477118
692	Oksanen, E., Lihavainen, J., Keinänen, M., Keski-Saari, S., Kontunen-Soppela, S.,
693	Sellin, A., & Söber, A. (2019). Northern Forest Trees Under Increasing Atmo-
694	spheric Humidity. In F. M. Cánovas, U. Lüttge, R. Matyssek, & H. Pretzsch

695 696 697	 (Eds.), Progress in Botany Vol. 80 (pp. 317-336). Cham: Springer International Publishing. Retrieved 2021-12-17, from https://doi.org/10.1007/124_2017_15
698 699	Pan, Y., Birdsey, R. A., Phillips, O. L., & Jackson, R. B. (2013, November). The Structure, Distribution, and Biomass of the World's Forests. <i>Annu. Rev.</i>
700	Ecol. Evol. Syst., 44(1), 593-622. Retrieved 2022-01-10, from https://
701	www.annualreviews.org/doi/10.1146/annurev-ecolsys-110512-135914
702	doi: 10.1146/annurev-ecolsys-110512-135914
703	Panagos, P., Jiskra, M., Borrelli, P., Liakos, L., & Ballabio, C. (2021, October).
704	Mercury in European topsoils: Anthropogenic sources, stocks and fluxes. En-
705	vironmental Research, 201, 111556. Retrieved 2021-07-29, from https://
706	www.sciencedirect.com/science/article/pii/S0013935121008501 doi:
707	10.1016/j.envres.2021.111556
708	Panek, J. A., & Goldstein, A. H. (2001, March). Response of stomatal conduc-
709	tance to drought in ponderosa pine: implications for carbon and ozone uptake.
710	<i>Tree Physiol.</i> , 21(5), 337–344. Retrieved 2021-07-22, from https://doi.org/
711	10.1093/treephys/21.5.337 doi: 10.1093/treephys/21.5.337
712	Papula, L. (2003). Mathematische Formelsammlung für Ingenieure und Naturwis-
713	senschaftler (8th ed.). Vieweg.
714	Pleijel, H., Klingberg, J., Nerentorp, M., Broberg, M. C., Nyirambangutse, B.,
715	Munthe, J., & Wallin, G. (2021, December). Mercury accumulation
716	in leaves of different plant types – the significance of tissue age and spe- aifie leaf area $Riagonalizing and 18(23) 6313 6328 Betriousd 2021 12$
717	Diogeosciences, 10(25), 0515-0526. Refineved 2021-12- 23 from https://bg.copernicus.org/articles/18/6313/2021/
718	10 5194/bg-18-6313-2021
719	Rea A W Lindberg S E Scherbatskov T & Keeler G J (2002) Mercury ac-
720	cumulation in foliage over time in two northern mixed-hardwood forests Wa-
722	ter Air Soil Poll., 133, 49–67.
723	Rutter, A. P., Schauer, J. J., Shafer, M. M., Creswell, J. E., Olson, M. R., Robinson,
724	M., Mallek, J. L. (2011, February). Dry deposition of gaseous elemen-
725	tal mercury to plants and soils using mercury stable isotopes in a controlled
726	environment. <i>Atmos. Environ.</i> , 45(4), 848–855. Retrieved 2019-09-24, from
727	http://www.sciencedirect.com/science/article/pii/S1352231010009805
728	doi: 10.1016/j.atmosenv.2010.11.025
729	Samaniego, L., Thober, S., Kumar, R., Wanders, N., Rakovec, O., Pan, M.,
730	Marx, A. (2018, May). Anthropogenic warming exacerbates European soil
731	moisture droughts. Nature Clim Change, 8(5), 421–426. Retrieved 2021-
732	12-17, from https://www.nature.com/articles/s41558-018-0138-5 doi:
733	10.1038/s41558-018-0138-5
734	Selin, N. E., Jacob, D. J., Yantosca, R. M., Strode, S., Jaeglé, L., & Sunderland,
735	E. M. (2008). Global 3-D land-ocean-atmosphere model for mercury:
736	Present-day versus preindustrial cycles and anthropogenic enrichment fac-
737	22 from https://ownersha.com/aci/obs/2020-09-
738	22, non https://agupubs.oniineiibrary.witey.com/doi/abs/10.1029/ 2007CB003040 doi: 10.1020/2007CB003040
739	Smith Downow N V Sunderland F M f_r Lacob D I (2010) Anthro
740	pogenic impacts on global storage and emissions of mercury from terres-
741	trial soils: Insights from a new global model
743	search: Biogeosciences, 115(G3). Retrieved 2019-06-20. from https://
744	agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2009JG001124 doi:
745	10.1029/2009JG001124
746	Sonke, J. E., Angot, H., Zhang, Y., Poulain, A., Björn, E., & Schartup, A. (2023,
747	May). Global change effects on biogeochemical mercury cycling. Ambio,
748	52(5), 853-876. Retrieved 2023-04-15, from https://doi.org/10.1007/
749	s13280-023-01855-y doi: 10.1007/s13280-023-01855-y

750	Stocker, B. D., Zscheischler, J., Keenan, T. F., Prentice, I. C., Seneviratne, S. I.,
751	& Peñuelas, J. (2019, April). Drought impacts on terrestrial primary pro-
752	duction underestimated by satellite monitoring. Nat. Geosci., 12(4), 264–
753	270. Retrieved 2021-10-22, from https://www.nature.com/articles/
754	s41561-019-0318-6 doi: 10.1038/s41561-019-0318-6
755	Wang, Q., Tenhunen, J., Dinh, N., Reichstein, M., Otieno, D., Granier, A., & Pile-
756	garrd, K. (2005, June). Evaluation of seasonal variation of MODIS derived
757	leaf area index at two European deciduous broadleaf forest sites. Remote
758	Sensing of Environment, $96(3-4)$, $475-484$. Retrieved 2019-05-15, from
759	https://linkinghub.elsevier.com/retrieve/pii/S0034425705001252
760	doi: 10.1016/j.rse.2005.04.003
761	Wang, R., Chen, J. M., Liu, Z., & Arain, A. (2017, August). Evaluation of seasonal
762	variations of remotely sensed leaf area index over five evergreen confierous
763	201 Detrieved 2020 10 12 from http://www.aciencediment.com/acience/
764	201. Refleved 2020-10-13, from http://www.sciencedifect.com/science/
765	Wang V Jacob D I is Logan I A (1008) Clobal simulation of transcription
766	O3 NO x hydrocarbon chemistry: 1 Model formulation – <i>Journal of Ceenhus</i>
769	ical Research: Atmospheres 103(D9) 10713-10725 Betrieved 2022-05-25
760	from https://onlinelibrary_wiley_com/doi/abs/10_1029/98 ID00158
770	(eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1029/98.ID00158) doi:
771	10.1029/98JD00158
772	Weselv, M. L. (2007, January). Parameterization of surface resistances
773	to gaseous dry deposition in regional-scale numerical models. Atmo-
774	spheric Environment, 41, 52–63. Retrieved 2021-10-21, from https://
775	www.sciencedirect.com/science/article/pii/S1352231007009740 doi:
776	10.1016/j.atmosenv.2007.10.058
777	Wohlgemuth, L., Osterwalder, S., Joseph, C., Kahmen, A., Hoch, G., Alewell, C., &
778	Jiskra, M. (2020, December). A bottom-up quantification of foliar mercury
779	uptake fluxes across Europe. $Biogeosciences, 17(24), 6441-6456$. Retrieved
780	2021-01-14, from https://bg.copernicus.org/articles/17/6441/2020/
781	doi: https://doi.org/10.5194/bg-17-6441-2020
782	Wohlgemuth, L., Rautio, P., Ahrends, B., Russ, A., Vesterdal, L., Waldner, P.,
783	Jiskra, M. (2022, March). Physiological and climate controls on foliar mercury $10(5)$ 1225 1252 D 4
784	2022 02 14 from https://bn.com/outi-alag/(10/1225/0000/
785	(Publisher: Concernique CmbH) doi: 10.5104/bg.10.1335.2022
786	Wu Z Wang X Chan E Turningood A A Cuenthar A B Niveri D
787	Alapaty K (2011 May) Evaluating the calculated dry denosition velocities of
790	reactive nitrogen oxides and ozone from two community models over a temper-
790	ate deciduous forest. Atmospheric Environment, 45(16), 2663–2674. Retrieved
791	2021-10-21, from https://www.sciencedirect.com/science/article/pii/
792	S1352231011002275 doi: 10.1016/j.atmosenv.2011.02.063
793	Yuan, H., Dai, Y., Xiao, Z., Ji, D., & Shangguan, W. (2011, May). Reprocessing
794	the MODIS Leaf Area Index products for land surface and climate mod-
795	elling. Remote Sensing of Environment, 115(5), 1171–1187. Retrieved
796	2023-04-16, from https://www.sciencedirect.com/science/article/pii/
797	S0034425711000149 doi: 10.1016/j.rse.2011.01.001
798	Yuan, W., Sommar, J., Lin, CJ., Wang, X., Li, K., Liu, Y., Feng, X. (2019,
799	January). Stable isotope evidence shows re-emission of elemental mercury
800	vapor occurring after reductive loss from foliage. Environ. Sci. Technol.,
801	53(2), 651–660. Retrieved 2019-11-29, from https://doi.org/10.1021/
802	acs.est.8b04865 doi: 10.1021/acs.est.8b04865
803	Yuan, W., Zheng, Y., Piao, S., Ciais, P., Lombardozzi, D., Wang, Y., Yang,
804	S. (2019, August). Increased atmospheric vapor pressure deficit re-

805	duces global vegetation growth. Sci. Adv., 5(8). Retrieved 2019-09-06,
806	from https://advances.sciencemag.org/content/5/8/eaax1396 doi:
807	10.1126/sciadv.aax1396
808	Zhang, H., Holmes, C., & Wu, S. (2016, September). Impacts of changes
809	in climate, land use and land cover on atmospheric mercury. Atmo-
810	spheric Environment, 141, 230–244. Retrieved 2021-10-21, from https://
811	linkinghub.elsevier.com/retrieve/pii/S1352231016304897 doi:
812	10.1016/j.atmosenv. $2016.06.056$
813	Zhang, L., Wright, L. P., & Blanchard, P. (2009, December). A review of cur-
814	rent knowledge concerning dry deposition of atmospheric mercury. At-
815	mos. Environ., 43(37), 5853–5864. Retrieved 2019-10-03, from http://
816	www.sciencedirect.com/science/article/pii/S1352231009007407 doi:
817	10.1016/j.atmosenv.2009.08.019
818	Zhou, J., & Obrist, D. (2021). Global mercury assimilation by vegetation. <i>Environ</i> .
819	Sci. Technol Retrieved 2021-10-08, from https://doi.org/10.1021/acs.est
820	.1c03530 doi: $10.1021/acs.est.1c03530$
821	Zhou, J., Obrist, D., Dastoor, A., Jiskra, M., & Ryjkov, A. (2021, March). Vege-
822	tation uptake of mercury and impacts on global cycling. Nat. Rev. Earth Env-
823	iron., 1-16. Retrieved 2021-03-16, from https://www.nature.com/articles/
824	s43017-021-00146-y doi: 10.1038/s43017-021-00146-y
825	Zweifel, R., Rigling, A., & Dobbertin, M. (2009). Species-specific stomatal response
826	of trees to drought – a link to vegetation dynamics? J. Veg. Sci., $20(3)$, 442–
827	454. Retrieved 2021-09-08, from https://onlinelibrary.wiley.com/doi/
828	abs/10.1111/j.1654-1103.2009.05701.x doi: 10.1111/j.1654-1103.2009
829	.05701.x
830	Zweifel, R., Steppe, K., & Sterck, F. J. (2007, June). Stomatal regulation by micro-
831	climate and tree water relations: interpreting ecophysiological field data with a
832	hydraulic plant model. J. Exp. Bot., $58(8)$, $2113-2131$. Retrieved 2021-04-19,
833	from https://doi.org/10.1093/jxb/erm050 doi: 10.1093/jxb/erm050

1	https://www.overleaf.com/project/623c4a3b9809e1d1d1117124
2	
3	A spatial assessment of current and future foliar Hg
4	uptake fluxes across European forests
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Key Points:

10

11	•	Extreme hot and dry atmospheric conditions have the potential to reduce stom-
12		atal uptake of ambient mercury by pine trees in Europe
13	•	Atmospheric drought controls on stomatal mercury uptake should be accounted
14		for in mercury transport models like GEOS-Chem
15	•	Forest foliar mercury uptake fluxes to Europe from a bottom-up model generally
16		agree well with results derived from literature and GEOS-Chem

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17 Abstract

Atmospheric mercury (Hg) is deposited to land surfaces mainly through vegetation uptake. Foliage stomatal gas exchange plays an important role for net vegetation Hg uptake, because foliage assimilates Hg via the stomata. Here, we use empirical relationships of foliar Hg uptake by forest tree species to produce a spatially highly resolved (1 km²) map of foliar Hg fluxes to European forests over one growing season. The modelled forest foliar Hg uptake flux is 23 ± 12 Mg Hg season⁻¹, which agrees with previous estimates from literature.

We spatially compare forest Hg fluxes with modelled fluxes of the chemistry-transport 25 model GEOS-Chem and find a good overall agreement. For European pine forests, stom-26 atal Hg uptake was shown to be sensitive to prevailing conditions of relatively high am-27 bient water vapor pressure deficit (VPD). We tested a stomatal uptake model for the to-28 tal pine needle Hg uptake flux during four previous growing seasons (1994, 2003, 2015/2017, 29 2018) and two climate change scenarios (RCP 4.5 and RCP 8.5). The resulting modelled 30 total European pine needle Hg uptake fluxes are in a range of 8.0 - 9.3 Mg Hg season⁻¹ 31 (min - max). The lowest pine forest needle Hg uptake flux to Europe (8 Mg Hg season⁻¹) 32 among all investigated growing seasons is associated with unusually hot and dry ambi-33 ent conditions in the European summer 2018, highlighting the sensitivity of the inves-34 tigated flux to prolonged high VPD. We conclude, that stomatal modelling is particu-35 larly useful to investigate changes in Hg deposition in the context of extreme climate events. 36

37 1 Introduction

Mercury (Hg) is a toxic pollutant that is transported globally through the atmo-38 sphere and deposited from air to land surfaces mainly through vegetation uptake of am-39 bient gaseous elemental Hg(0) (Demers et al., 2013; Jiskra et al., 2015; Enrico et al., 2016; 40 Obrist et al., 2017; Feinberg et al., 2022). Consequently, vegetation uptake has the po-41 tential to lower atmospheric Hg(0) transport and Hg deposition to oceans, where Hg can 42 be methylated and bioaccumulated in marine seafood for human consumption (Zhou et 43 al., 2021). In order to assess and improve the effectiveness of mitigation policies for hu-44 man exposure, it is thus necessary to constrain environmental drivers of vegetation Hg(0)45 uptake. Furthermore, process understanding of vegetation Hg(0) uptake is essential for 46 assessing future human Hg exposure in the context of global change (Sonke et al., 2023). 47

Global vegetation and soil uptake of Hg(0) has been estimated to amount to 2850 48 \pm 500 Mg year⁻¹ (Obrist et al., 2021; Zhou et al., 2021; Feinberg et al., 2022), exceed-49 ing approximate direct anthropogenic emissions to the air of 2200 Mg Hg year⁻¹ (Sonke 50 et al., 2023). Forests contain 80 % of the global plant biomass (Pan et al., 2013), there-51 for representing a major vector for Hg(0) drawdown from the atmosphere. In forests, 52 half of the total Hg(0) net deposition is estimated to be stored in tree foliage, while the 53 other half is estimated to be transferred to vascular tissues (e.g. stem, branches, roots), 54 or taken up by understory vegetation (e.g. shrubs, grasses) or nonvascular plants (lichen 55 and mosses) (Zhou et al., 2021; Obrist et al., 2021; Zhou & Obrist, 2021). In tree foliage, 56 Hg concentrations increase linearly between foliage emergence and senescence (Rea et 57 al., 2002; Laacouri et al., 2013; Blackwell et al., 2014; Wohlgemuth et al., 2020; Pleijel 58 et al., 2021) implying a net foliar Hg deposition flux, albeit Hg re-emission from foliar 59 surfaces of up to 30% of gross foliage Hg(0) deposition had been observed in a subtrop-60 ical forest in China (W. Yuan, Sommar, et al., 2019). The bulk (90-96%) of Hg is stored 61 in foliage tissues as opposed to leaf surfaces and correlates with leaf stomatal density (Laacouri 62 et al., 2013). Studies on Hg stable isotopes in foliage (Demers et al., 2013; Zhou et al., 63 2021), enriched isotope tracer experiments (Rutter et al., 2011) and the vertical varia-64 tion of net foliar Hg uptake in forest canopies (Wohlgemuth et al., 2020) strongly sug-65 gest a diffusive uptake pathway of atmospheric Hg(0) to foliage interiors via the stom-66 ata (Liu et al., 2021). In this way, foliar Hg(0) uptake is linked to foliage stomatal aper-67 ture for atmospheric gas exchange (Wohlgemuth et al., 2022). 68

Trees regulate foliage stomatal aperture to balance the inward diffusion of CO_2 for 69 photosynthesis with the risk of desiccation caused by excessive outward diffusion of wa-70 ter vapor (Körner, 2013). The degree of stomatal aperture depends on atmospheric CO_2 71 levels and hydrological conditions (soil water availability and atmospheric evaporative 72 demand) and varies among foliage-specific traits (age, tree species-specific evolutionary 73 metabolic strategy and water use efficiency) (Körner, 2013). Pine, for instance, is an iso-74 hydric tree species capable of closing foliage stomata under warm and dry atmospheric 75 conditions relatively early compared to tree species like oak and spruce (Lagergren & Lin-76 droth, 2002; Zweifel et al., 2007, 2009), resulting in a reduced stomatal conductance for 77 pine needle diffusive gas exchange (Panek & Goldstein, 2001). Consistently, Hg(0) up-78 take rates by pine needles in Europe were found to be lower at forest sites across Europe, 79 where prolonged warm and dry atmospheric conditions prevailed over a given growing 80 season during daytime (Wohlgemuth et al., 2022). 81

Species-specific stomatal response strategies to meteorological conditions are par-82 ticularly relevant for projections of future foliar Hg uptake under climate change. Increas-83 ing global atmospheric temperatures driven by rising levels of greenhouse gases will re-84 sult in an increased frequency of droughts (Grossiord et al., 2020) and higher soil mois-85 ture deficits (Berg & Sheffield, 2018; Stocker et al., 2019) in various regions of the world. 86 These climatic conditions may decrease foliar Hg(0) uptake fluxes due to lower stomatal 87 conductance (Wohlgemuth et al., 2022). A reduced plant Hg sink could further be am-88 plified by deforestation and forest diebacks, particularly in the tropics (Allen et al., 2015; 89 Brando et al., 2019; Feinberg et al., 2023). Other regions of the world are projected to 90 become wetter through an increase in precipitation rates under climate change (IPCC, 91 2021a), which might lead to higher foliage stomatal conductance relative to the present 92 and thus higher foliar Hg uptake. With continuing anthropogenic carbon emissions, an 93 elevated atmospheric CO_2 level might have an antagonizing effect on the foliar stomatal 94 Hg(0) uptake flux: foliar Hg(0) uptake could decline with decreasing stomatal conductance under CO₂ fertilization (Norby & Zak, 2011), or, the opposite, the vegetation sink 96 for Hg(0) could increase with intensified biomass growth and higher soil C contents (Hararuk 97 et al., 2013; Jiskra et al., 2018; H. Zhang et al., 2016). In order to make projections of 98 the foliar Hg uptake flux in the next decades, these climate change impacts need to be 99 further investigated and potentially implemented into global and regional Hg cycle mod-100 els. 101

Current and future Hg fluxes are modelled in Global Chemical Transport Models 102 (CTMs). CTMs like GEOS-Chem (Selin et al., 2008) apply resistance-based algorithms 103 (Wesely, 2007) for modelling Hg(0) deposition fluxes from the atmosphere to vegetated 104 ecosystems and are often based on parameters like leaf area indices (LAIs), temperature 105 and wind speed. The resistance components for leaf stomata within CTMs commonly 106 represent consensus values optimized to fit observations of Hg deposition velocities over 107 vegetated surfaces (Selin et al., 2008; L. Zhang et al., 2009; Smith-Downey et al., 2010; 108 H. Zhang et al., 2016), without taking stomatal feedback to environmental conditions 109 into account (Wu et al., 2011; Khan et al., 2019). Consequently, forest tree species-specific 110 stomatal responses to climate change at foliage level are not parameterized in CTMs. 111 An additional problem related to CTMs is the uncertainty of modelled Hg(0) deposition 112 fluxes due to insufficient model evaluation against dry deposition measurements (Feinberg 113 et al., 2022). This issue of model validation was highlighted in a recent revision of GEOS-114 Chem parameterization after matching the GEOS-Chem model design to various exper-115 imental Hg(0) deposition measurements, which resulted in a doubling of the modelled 116 global flux of Hg(0) dry deposition to land compared to previous model outcomes (Feinberg 117 et al., 2022). 118

In this study, we assess the spatial variation of forest foliar Hg uptake fluxes across Europe by producing a spatially highly resolved map of foliar Hg uptake fluxes to European forests using a bottom-up model that incorporates pine tree stomatal responses to climate conditions. We compare these spatially resolved fluxes to forest dry deposition fluxes modelled in GEOS-Chem in order to identify spatial discrepancies between GEOS-Chem and the bottom-up model used here. We investigate the sensitivity of an empirical stomatal response model of pine to different climatic conditions during past growing seasons and for two climate change projections of the years 2068 - 2082 in order to outline the potential of incorporating a stomatal response function into CTMs.

¹²⁸ 2 Materials and Methods

129

2.1 Description of datasets

For creating maps of foliar and pine needle Hg uptake fluxes in Europe applying a bottom-up model (Sect. 2.2 and 2.3), we drew on multiple data sources:

• Foliar Hg data. A dataset of foliar Hg uptake rates was derived from Hg mea-132 surements in foliage of tree canopies at 272 forest sites of the UNECE International 133 Co-operative Programme on Assessment and Monitoring of Air Pollution Effects 134 on Forests (ICP Forests). Forest sites are mostly located in Central and North-135 ern Europe (+737 sites in Austria from the Austrian Bio-Indicator Grid) and har-136 monized foliage sampling methods were employed. All foliage samples within this 137 dataset were harvested at the end of the growing seasons 2015 or 2017. Therefore, 138 average foliage values of 2015/2017 constitute reference values of forest foliar Hg 139 uptake fluxes relative to respective fluxes during investigated years of this study. 140 The dataset is publicly available and contains 3569 foliar Hg concentrations of 23 141 tree species and is described in detail in (Wohlgemuth et al., 2022). 142

- Meteorological data. Values on ambient temperature and relative humidity at surface air pressure (1000 hPa) in Europe (spatial resolution: 0.25° x 0.25°) originate from ERA5 hourly reanalysis data and were downloaded from the Copernicus Climate Data Store (Hersbach et al., 2018). The applied time frame includes hourly daytime (07:00 18:00 LT) values during the respective growing seasons (April October) of 1994, 2003, 2015, 2017, and 2018.
- Climate change data. Regional climate simulation data of air temperature and 149 relative humidity at 2 m above surface level for the years 2068 - 2082 and two dif-150 ferent climate change scenarios (Representative Concentration Pathway (RCP) 151 4.5 and RCP 8.5 (IPCC, 2021b)) were obtained from the Coordinated Regional 152 Climate Downscaling Experiment (CORDEX) (Jacob et al., 2020) framework for 153 the European domain with a spatial resolution of $0.11^{\circ} \ge 0.11^{\circ}$ and a temporal res-154 olution of 3hourly daytime (09:00 - 18:00 LT) values. For representing a range of 155 different climate model outputs, we calculated average values from multiple re-156 gional climate models (RCMs) downscaled from global climate models (GCMs) 157 depending on availability for download from the Copernicus Climate Data Store 158 (C3S, 2022). In total, we incorporated data of 15 combinations of 4 RCMs and 159 6 GCMs for RCP 4.5 and of 13 combinations of 6 RCMs and 8 GCMs for RCP 160 8.5 (see Table SI 3) for an overview of models and ensemble members). 161
- **European tree species distribution**. We used a map of spatial proportions of 162 tree species groups per km² land area from (Brus et al., 2012). For use in calcu-163 lating pine foliar Hg uptake fluxes (see Sect. 2.3), we summed up spatial relative 164 abundance values of Pinus sylvestris, Pinus pinaster, Pinus nigra and Pinus halepen-165 sis from European forest inventories (Mauri et al., 2017; Buras & Menzel, 2019) 166 and multiplied these pine relative abundances with the respective total forest area 167 per km^2 derived from (Brus et al., 2012) to obtain pine areal proportions. We per-168 formed the same calculation (sum of values of Pinus sylvestris, Pinus pinaster, Pi-169 nus nigra and Pinus halepensis and subsequent multiplication with respective to-170 tal forest area) to estimate the distribution of pine in Europe under climate change 171 using relative abundance probabilities projected from climate analogues for the 172 time period 2061 - 2090 and RCP 4.5 and RCP 8.5 by (Buras & Menzel, 2019). 173

Leaf Area Indices (LAIs) and Leaf Mass per Area (LMA) values. We
used the LAI satellite product (spatial resolution: 330 m) of PROBA-V (Dierckx
et al., 2014; Fuster et al., 2020) to upscale foliar Hg uptake rates at each ICP Forests
site to foliar Hg uptake fluxes (see Sect. 2.2), along with average LMA values per
tree species from (Forrester et al., 2017).

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2.2 Calculation of forest foliage Hg uptake fluxes

We determined forest foliar Hg uptake fluxes to European forests on a 1 km² spatial resolution applying three basic computational steps: 1) calculation of tree speciesspecific daily Hg uptake fluxes per m² ground area using a bottom-up model; 2) upscaling of respective foliar Hg fluxes per tree species to the European forested area using the areal distribution of corresponding tree species; 3) multiplication of daily forest foliar Hg uptake fluxes per latitude with latitude-dependent growing season length in order to obtain the forest foliar Hg uptake fluxes over one growing season.

Computational step 1) is based on the premise, that foliar Hg uptake rates are tree 187 species-specific (Laacouri et al., 2013; Wohlgemuth et al., 2022; Pleijel et al., 2021). For 188 this reason, we calculated median daily foliar Hg uptake fluxes per tree species group (see 189 Table SI 2 for details) of all forest sites from the foliar Hg dataset (Sect. 2.1). The bottom-190 up modeling approach for calculating daily foliar Hg uptake fluxes from daily foliar Hg 191 uptake rates is described in detail in Wohlgemuth et al., (2020) (Wohlgemuth et al., 2020). 192 Briefly, daily foliar Hg uptake rates per gram foliage dry weight (units of ng Hg $g_{d.w.}^{-1}$ d⁻¹) 193 were multiplied with tree species-specific LMA values (Sect. 2.1) to obtain daily foliar 194 Hg uptake rates per foliage surface area (ng Hg m_{leaf}^{-2} d⁻¹). Subsequently, values of daily 195 foliar Hg uptake rates per foliage surface area are multiplied with values of LAI (m_{leaf}^2) 196 m_{ground}^{-2} ; Sect. 2.1), resulting in daily foliar Hg fluxes per unit ground area (ng Hg m_{ground}^{-2} 197 d^{-1}). LAI values of coniferous forests are relatively constant during the active growing 198 season after the initial growth phase of current-season needles (R. Wang et al., 2017), 199 while LAI values of temperate deciduous forests increase rapidly at the beginning of the 200 growing season (leaf flushing) and climax at peak season (June – August, northern hemi-201 sphere) (Q. Wang et al., 2005). For coniferous tree species, we used the maximum LAI 202 value during the constant period at each forest site of the ICP Forests dataset to cal-203 culate needle foliar Hg uptake fluxes. For deciduous tree species, we calculated foliar Hg 204 uptake fluxes as a temporal sequence at every LAI value available over the growing sea-205 son and subsequently used median foliar Hg uptake flux values of the growing season. 206 For LAI values larger than 3, we applied a species-specific tree height correction factor, 207 to account for lower foliar Hg uptake fluxes of shaded leaves in the lower canopy (Wohlgemuth 208 et al., 2020) (refer to Table SI 1 for utilized tree height correction factors). For conifer-209 ous species, we multiplied Hg uptake fluxes of current-season needles with a species-specific 210 needle age correction factor to account for lower Hg uptake rates of older needle age classes 211 (Wohlgemuth et al., 2020) (refer to Table SI 1 for utilized needle age correction factors). 212

Computational step 2) involves the multiplication of the proportion of each tree 213 species per km² land area with the respective species-specific median daily foliar Hg up-214 take fluxes. We matched tree species-specific Hg data with the areal forest distribution 215 of the respective tree species (Brus et al., 2012). In the few cases of rare European tree 216 species, where specific Hg data was lacking, we pooled Hg or forest distribution data by 217 tree species group (see Table SI 2 for an overview of matched tree species groups between 218 the two datasets). Subsequently, we added up all tree species-specific daily foliar Hg up-219 take fluxes within each km^2 and obtained one forest foliar daily Hg uptake flux per km^2 . 220

In computational step 3) we calculated forest foliar Hg uptake fluxes per km² and one growing season by multiplying each daily foliar Hg uptake flux per km² with the growing season length in days following a simple latitudinal model (CLRTAP, 2017). The latitudinal model of growing season determines a growing season length of 192 days at latitude 50° and decreases by 3.5 days per 1° of latitude moving north and increases by 3.5 days per 1° of latitude moving south.

2.3 Calculation of pine foliar Hg uptake fluxes

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Daily foliar Hg uptake rates of pine tree species were calculated taking into account 228 the empirical dependence of needle Hg uptake fluxes to atmospheric VPD. Pine needle 229 daily Hg uptake rates (upR_{pine}; ng Hg $g_{d.w.}^{-1}$ d⁻¹) were found to be lower at forest sites, 230 where the daytime fraction of water VPD > 1.2 kPa during the respective sample life 231 period (proportion_{dayVPD} > 1.2kPa) was relatively high (Wohlgemuth et al., 2022). The 232 negative correlation of pine needle Hg uptake with timespan of elevated atmospheric VPD 233 was explained by a stomatal closure upon VPD threshold exceedance and thus a high 234 235 stomatal resistance suppressing the diffusive uptake of Hg(0) from the atmosphere. The linear regression of daily foliar Hg uptake rates with proportion_{dayVPD} > 1.2 kPa is: upR_{pine} 236 $= 0.116 - 0.13 \text{ x} (\text{proportion}_{dayVPD} > 1.2 \text{ kPa})$ (Wohlgemuth et al., 2022). We applied 237 this linear relationship to calculate the pine foliar Hg uptake rates of the forest area of 238 Europe during four different growing seasons in 1994, 2003, an average of 2015 and 2017, 239 2018, and projected for the time period 2068 - 2082 under RCP 4.5 and RCP 8.5 (IPCC, 240 2021b). We calculated hourly or 3hourly daytime VPD values from ERA5 (Hersbach et 241 al., 2018) or CORDEX data (Sect. 2.1) on surface temperature and relative humidity 242 using the Auguste-Roche-Magnus formula (W. Yuan, Zheng, et al., 2019) and subsequently 243 determining the fraction of daytime hours when the VPD was above the threshold of 1.2 244 kPa over the respective latitudinal growing season length. Calculations with climate data 245 were performed at sciCORE scientific computing center at University of Basel. We de-246 fined growing season length per latitude using a latitudinal model ((CLRTAP, 2017), see 247 Sect. 2.2). In 2068 - 2082 we assumed the beginning of the growing season to be 3 days 248 earlier and the end of the growing season to be 3 days later to take increases in grow-249 ing season length under climate change into account (Jeong et al., 2011; Garonna et al., 250 2014). The underlying areal distribution of pine is based on European forest inventories 251 and projections of pine abundances based on climate analogues under RCP 4.5 and RCP 252 8.5 by (Buras & Menzel, 2019) (see Sect. 2.1). 253

2.4 GEOS-Chem forest deposition flux calculation

GEOS-Chem is a global 3-D chemistry transport model, which includes a comprehensive Hg cycle (Selin et al., 2008). Table 1 gives an overview of the methodological approach and input parameters for calculating the respective Hg fluxes of GEOS-Chem and the bottom-up model (Sect. 2.2), which we compared spatially in this study.

	bottom-up model	GEOS-Chem
model input pa- rameters	 spatial forest distribution (Brus et al., 2012); leaf area indices (LAIs) (Dierckx et al., 2014; Fuster et al., 2020); leaf mass per area (LMA) (Forrester et al., 2017); meteorological parameter: day-time VPD (Hersbach et al., 2018); foliar Hg uptake rates (Wohlgemuth et al., 2022) 	spatial forest distribution (Gibbs, 2006); leaf area indices (LAIs) (H. Yuan et al., 2011); atmospheric Hg(0) levels (GEOS-Chem v12.8.1 simu- lation 2015); meteorological parameters: air temperature, pressure, solar ra- diation, cloud cover, wind speed (GEOS-FP) (Lucchesi, 2018)
spatial resolution	1 km x 1 km	$0.25 \ge 0.3125^{\circ}$
basic methodolog- ical approach for Hg flux calculation	spatial upscaling of measured foliar Hg uptake rates (Wohlgemuth et al., 2020)	in-series calculation of Hg dry deposition velocity from pa- rameterized resistance values (Wesely, 2007)
foliage stomatal Hg uptake flux component	calculated for pine based on daytime vapor pressure deficit (VPD) values (Sect. 2.3)	calculated within the canopy resistance component as a func- tion of land type, leaf area indices (LAIs), and solar radia- tion
model output com- pared in this study	tree-species specific forest foliar Hg(0) uptake fluxes	Hg(0) dry deposition fluxes to coniferous and deciduous forest land cover

Table 1. Caption

We used an offline version of the GEOS-Chem dry deposition code (Feinberg, 2022) 259 to be able to calculate dry deposition velocities at higher resolution and only for certain 260 land use types (i.e., forest areas). The offline dry deposition code computes deposition 261 velocities using a resistance-based approach (Y. Wang et al., 1998; Wesely, 2007). In-262 put variables (Table 1) are gridded hourly GEOS-FP meteorological data for (e.g., air 263 temperature, wind speed, solar radiation, and cloud cover) and weekly LAI values based 264 on MODIS (H. Yuan et al., 2011) for the year 2015. The model calculates the Hg(0) dry 265 deposition velocity based on species-specific parameters including its biological reactiv-266 ity (f₀ = 10⁻⁵) and Henry's Law Constant (H^{*} = 0.11 M atm⁻¹). To isolate the uptake 267 of Hg(0) to forests, we calculated the dry deposition velocity only over coniferous and 268 deciduous land cover types from the Olson land map (Gibbs, 2006). The offline calcu-269 lations output hourly dry deposition velocities over the European domain at $0.25 \times 0.3125^{\circ}$ 270 resolution. We converted the calculated Hg(0) deposition velocities to fluxes by multi-271 plying with hourly surface Hg(0) concentrations from a GEOS-Chem v12.8.1 simulation 272 for 2015. For this study, we compared the GEOS-Chem Hg(0) dry deposition fluxes to 273 forests with foliar Hg(0) uptake fluxes calculated using the bottom-up model. For both 274 models, Hg fluxes were averaged over the latitude-dependent growing season length in 275 days and cropped to the same spatial extent. As GEOS-Chem and the bottom-up model 276 differ in their geographic resolution (GEOS-Chem: $0.25^{\circ} \ge 0.31^{\circ} \sim 955 \text{ km}^2$ vs. bottom-277 up: 1 km^2), we downsampled daily forest foliar Hg uptake fluxes from the bottom-up 278 model through bilinear interpolation. 279

2.5 Uncertainty analysis of foliar Hg uptake fluxes

The relative uncertainty value per tree species group depended on propagated un-281 certainties of calculation parameters used to derive the respective foliar Hg uptake flux 282 per tree species group (see Table SI 4 for details and values). Subsequently, we calcu-283 lated one relative uncertainty value per geographic tile of our European flux map (Fig. 284 1) by summarizing the relative uncertainty of each foliar Hg uptake flux per tree species 285 group within each tile according to error propagation principles (Ku, 1966; Papula, 2003). 286 We obtained the relative uncertainty for the total foliar Hg uptake flux to European forests 287 (Fig. 1) by propagating all relative uncertainty values per tile. The final relative uncer-288 tainty value of total foliar Hg uptake flux to European forests and the reference grow-289 ing seasons 2015/2017 is 0.52. 290

²⁹¹ **3** Results and Discussion

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3.1 Spatial distribution of forest foliar Hg uptake fluxes across Europe

Figure 1 visualizes forest foliar Hg uptake fluxes per growing season at a spatial 293 resolution of 1 km^2 (g Hg km⁻² season⁻¹) in Europe. Forest foliar Hg uptake fluxes gen-294 erally follow a spatial distribution of European forests, because this map (Fig. 1) is based 295 on the proportion of forest tree species per land area (Brus et al., 2012). Consequently, 296 the largest forest foliage Hg uptake fluxes in terms of area are on the Scandinavian Penin-297 sula with dense forest land cover. Outside of Scandinavia, forest foliage Hg uptake fluxes 298 fall along large contiguous forested areas, e.g. in the Carpathian Mountains, the South-299 Eastern Alps, the Balkans, or forested low mountain areas like the Black Forest. 300



Figure 1. Spatial distribution of forest foliar Hg uptake fluxes (g Hg km² growing season⁻¹) to Europe based on a bottom-up extrapolation of foliar Hg concentrations, that were measured and averaged over the 2015 and 2017 growing seasons. Dark grey areas represent excluded non-forested areas (e.g. surface waters or non-vegetated mountain areas).

The sum of forest foliar Hg uptake fluxes over the land area of Europe as displayed in Figure 1 equals 23 ± 12 Mg Hg season⁻¹. This total flux agrees within uncertainty

with a previous estimate for the total foliar Hg uptake flux to Europe of 20 ± 3 Mg Hg 303 over the 2018 growing season based on foliar Hg uptake fluxes at four forested sites (Wohlgemuth 304 et al., 2020). (Zhou & Obrist, 2021) evaluated a median global foliar Hg assimilation of 305 28 Mg yr⁻¹ for deciduous broadleaf forests and 61 Mg yr⁻¹ for evergreen needleleaf forests 306 by combining foliar Hg concentrations with annual net foliar biomass production data 307 of the respective forest types. From these global assimilation estimates by (Zhou & Obrist, 308 2021), we calculated a total foliar Hg assimilation of 29 Mg yr^{-1} to the deciduous and 309 coniferous forest land area of Europe (for details see SI, Text S1), which is slightly higher 310 but still within the uncertainty of the 23 ± 12 Mg Hg season⁻¹ from this study. How-311 ever, foliar Hg uptake fluxes based on net primary foliar biomass production by Zhou 312 and Obrist, (2021) (Zhou & Obrist, 2021) does not correct for lower foliar Hg uptake rates 313 by shade leaves and multiyear old needles (see Sect. 2.2) relative to sun leaves and younger 314 needles (Wohlgemuth et al., 2020), likely resulting in a systematic over-estimation. We 315 assume, that the different time reference (seasonal vs. annual) of the flux from this study 316 $(23 \pm 12 \text{ Mg Hg season}^{-1})$ and the flux derived from Zhou and Obrist, (2021) (Zhou & 317 Obrist, 2021) (29 Mg Hg yr⁻¹) only plays a minor role for explaining the difference be-318 tween the two fluxes, since we expect a small net foliar biomass production in Europe 319 in winter outside of the growing season. 320

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3.2 Pine foliar Hg uptake fluxes under different VPD scenarios

Figure 2 shows total pine forest foliar Hg uptake fluxes to Europe calculated un-322 der different conditions of atmospheric surface-level water VPD during four past grow-323 ing seasons (1994, 2003, 2015/2017, 2018) and simulated for the years 2068 - 2082 as an 324 average of multiple climate model outputs (see Sect. 2.2) under two different climate change 325 scenarios (RCP 4.5 and RCP 8.5). The leftmost bar (Fig. 2) represents a theoretical base-326 line pine needle Hg uptake flux in absence of VPD induced stomatal control (potential 327 maximum transpiration rates) on the pine needle Hg uptake flux. The total pine nee-328 dle Hg uptake flux to Europe during the reference growing season 2015/2017 (Sect. 2.2) 329 is 9.3 \pm 3.7 Mg Hg representing 70% of the baseline flux of 13.3 \pm 5.3 Mg Hg season⁻¹. 330 Thus, based on the pine needle Hg uptake model used in this study (Sect. 2.3), the VPD 331 effect reduces the total pine needle Hg uptake flux to Europe by approximately 30%. 332



Figure 2. Pine needle Hg uptake flux to European pine forests (Mg Hg season⁻¹) calculated from atmospheric surface water vapor pressure deficit (VPD) conditions during the growing seasons 1994, 2003, 2015/2017, 2018 and projected for the years 2068 - 2082 under RCP 4.5 and RCP 8.5. Bar on the left represents a baseline pine forest needle Hg uptake flux with no VPD exceedance of 1.2 kPa throughout the growing season.

The relative standard deviation of modelled total pine needle Hg uptake fluxes for 333 the investigated growing seasons (1994, 2003, 2015/2017, 2018, 2068 - 2082) was 0.07. 334 Consequently, modelled total European pine needle Hg uptake fluxes hardly differed from 335 each other among growing seasons. The total pine needle Hg uptake flux in Europe de-336 pend on VPD conditions in areas where pine forests prevail. Pine forests are primarily 337 located in Northern Europe (SI Fig. 1), where hourly ambient VPD was > 1.2 kPa dur-338 ing 30% or less of daytime in the growing seasons 1994, 2003 and 2015/2017 due to rel-339 atively cool and moist ambient conditions as compared to Central and Southern Europe 340 (see e.g. VPD conditions during reference time period 2015/2017 Fig. 3a). In contrast 341 to previous years, the European summer hydrological condition of 2018 has been described 342 as an intense hot drought, during which pronounced stomatal closure of coniferous forests 343 in response to high VPD were recorded in Switzerland (Gharun et al., 2020). In South-344 ern Fennoscandia, conditions of ambient hourly VPD > 1.2 kPa prevailed over excep-345 tionally long time proportions (around 40%) during the summer of 2018 (see Fig. 3b, 346 (Buras et al., 2020)). As a result, the modelled total pine needle Hg uptake flux in Eu-347 rope in 2015/2017 (9.3 Mg Hg season⁻¹) was by a factor of 1.16 higher than the respec-348 tive flux in 2018 (8.0 Mg Hg season⁻¹). We conclude that hot and dry summer condi-349 tions (Fig. 2) in Fennoscandia crucially impact modelled past total pine needle Hg up-350 take fluxes in Europe. According to the model results, an average amount of 1.3 Mg Hg 351 was not deposited via pine needle uptake in 2018 compared to 2015/17, potentially re-352 maining in the atmosphere, where it can be long-range transported to the ocean (Zhou 353 et al., 2021). These 1.3 Mg Hg are more than three times larger than the reported an-354 thropogenic Hg emissions of Sweden in 2021 ??, highlighting the quantitative impact, 355 that hot droughts can have on the pine needle Hg uptake flux. 356



Figure 3. Average daytime proportion of surface level atmospheric water VPD > 1.2 kPa during a) the reference growing season 2015/2017, and b) the growing season 2018. All VPD values were calculated from hourly reanalysis data of ERA5 ambient air temperature and relative humidity (Sect. 2.1).

3.3 Projected pine forest needle Hg uptake fluxes under climate change scenarios

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The projected total pine forest needle Hg uptake flux for 2068 - 2082 (RCP 4.5: 9.3 359 \pm 5.5 Mg Hg season⁻¹; RCP 8.5: 8.1 \pm 4.9 Mg Hg season⁻¹) was in the same range as 360 the corresponding average flux for the years 1994, 2003, 2015 and 2017 of 9.1 \pm 0.2 Mg 361 Hg season⁻¹ (mean \pm sd), but slightly higher than the corresponding flux in the year 362 of 2018 (8.0 \pm 3.2 Mg Hg season⁻¹), during which Fennoscandia experienced a summer 363 of relatively long hot and dry ambient conditions. Figure 4 maps the absolute deviation 364 of the pine forest needle Hg uptake flux projected for 2068 - 2082 (simulated future flux) 365 from the corresponding 2018 flux in Europe. Under RCP 4.5, the simulated future flux 366 is higher (blue area in Fig. 4 a) than the 2018 flux in 65% of total area. Under RCP 8.5, 367 the simulated future flux is higher (blue area in Fig. 4 b) than the 2018 flux in 43% of 368 total area. In most area of Fennoscandia, where a majority of pine forests in Europe are 369 located (SI Fig. 1a), the future flux is projected to be larger than in 2018. For both cli-370 mate change scenarios, the projection predicts lower pine needle Hg fluxes to the Balkans 371 and to the Southern Iberian Peninsula than in 2018 (Fig. 4). 372



Figure 4. Absolute deviation of projected pine forest foliar Hg uptake fluxes for 2068 - 2082 (under RCP 4.5 (a) and RCP 8.5 (b)) from the corresponding flux modelled for 2018. In blue areas, the projected future flux under the two climate change scenarios is higher than the respective 2018 flux, in orange areas, this deviation is reversed.

The pine forest needle Hg uptake flux for 2068 - 2082 simulated here is a function 373 of both modelled ambient VPD conditions during the growing season and the projected 374 distribution of pine forests in Europe depending on climate analogs (Buras & Menzel, 375 2019). While the pine forest cover in Southern Sweden is projected to decrease under 376 the climate change scenarios RCP 4.5 and RCP 8.5 from around 50% km⁻² to around 377 25% km⁻², forest cover in Central and Northern Fennoscandia is projected to be rela-378 tively steady for climate analogs of both climate change scenarios (compare SI Fig. 1 a 379 - c). Average long-term precipitation rates are projected to increase in Scandinavia, along 380 with a decrease of meteorological drought in the coming decades under different climate 381 change scenarios (Forzieri et al., 2014; Samaniego et al., 2018; Kellomäki et al., 2018), 382 which could result in an increase of atmospheric humidity and a decrease of VPD in north-383 ern Europe (Oksanen et al., 2019). Under this scenario of wetter forest environments, the Hg sink of Scandinavian pine forest needles would not be significantly diminished. 385 However, drought trends in Fennoscandia are still inconsistent and extreme drought events 386 like in 2018 might occur more frequently under the current rate of climate change (IPCC, 387 2021a). The summer of 2018 was a record hot drought in Europe (Buras et al., 2020), 388 while climate simulations for 2068 - 2082 are averaged over multiple climate models (SI 389 Table 3), possibly averaging out extreme events. In a scenario, where the maximum pro-390 portion of daytime VPD > 1.2 kPA per growing season averaged over 2068 - 2082 pre-391 vails at each spatial unit, the total pine forest needle Hg uptake flux to Europe reduces 392 to 6.9 Mg Hg season⁻¹ for RCP 4.5 and 5.0 Mg Hg season⁻¹ for RCP 8.5, which cor-393 responds to 74% and 62% of the respective flux derived from an average VPD daytime 394 proportion. We therefore suggest that extreme climate events of extended time periods 395 of ambient daytime VPD > 1.2 kPa like during the growing season 2018 (Fig. 3b) could 396 reduce the pine forest needle Hg uptake flux in Fennoscandia in future even compared 397 to average long-term VPD projections (Fig. 4). 398

A source of model uncertainty of the future forest foliar Hg uptake flux under climate change arises from atmospheric Hg(0) concentrations that depend on anthropogenic emissions, re-emissions of mobilized legacy Hg and future global deposition fluxes under climate and land use change (Sonke et al., 2023; Feinberg et al., 2023), which we could not account for in this study. However, our model outputs call attention to the sensi-

tivity of the pine needle Hg uptake flux to extreme hot and dry ambient conditions, which 404 should be accounted for in chemistry-transport models under varying atmospheric Hg(0)405 levels. The impact of the hot and dry conditions on the pine Hg uptake fluxes might have 406 implications for Hg inputs into aquatic ecosystems. In a recent review on Hg cycling in the context of global change, (Sonke et al., 2023) highlighted the potential of legacy Hg 408 (i.e. actively cycling Hg that was mobilized in the past) to cause contamination by mo-409 bilization of Hg from soils to wetlands and coastal ecosystems via riverine systems. While 410 most soil Hg enters riverine systems by soil erosion from agricultural lands, contaminated 411 sites, and deforested woodland (Panagos et al., 2021; Sonke et al., 2023), a reduced for-412 est foliar Hg uptake and subsequent deposition to forest soils may decrease the amount 413 of runoff Hg from forest soils in the long-term, while long-range Hg transport to the open 414 ocean via the atmosphere might be enhanced (Zhou et al., 2021). 415

3.4 Comparison of bottom-up model with GEOS-Chem

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Figure 5 depicts spatial ratios of daily forest Hg uptake fluxes of the bottom-up model to GEOS-Chem. Absolute difference values of the two model outputs are shown in SI Fig. 2.



Figure 5. Ratios per spatial unit of daily forest Hg uptake fluxes averaged over the latitude-specific growing season length of the bottom-up model to GEOS-Chem.

Results of average daily foliar Hg uptake fluxes from GEOS-Chem and the bottom-420 up model were geographically comparable: In 59% of the spatial domain with values > 421 0, average daily foliar Hg uptake fluxes from the two models differed by factor of 1 - 2422 from each other, in 37% of the domain, model values differed by a factor of 2 - 10 from 423 each other, and in 4% of the domain respective values differed by a factor of > 10 from 424 each other (Fig 5). We examined if differences in modelled average daily foliar Hg up-425 take fluxes at the same geographic location originate from differences in the underlying forest distribution maps of the two compared models. In 78% of spatial tiles with val-427 ues > 0, the ratio of average daily foliar Hg uptake fluxes of the bottom-up model to GEOS-428 Chem agreed in range (Fig. 5) with the ratio of the forest fraction of the bottom-up model 429

to GEOS-Chem per respective spatial tile. We thus hypothesize that the bottom-up model 430 and GEOS-Chem generally produce similar foliar Hg flux values per spatial unit given 431 the same forest distribution. Reasons for minor differences in model outputs are chal-432 lenging to identify, since the two models are based on different approaches, parameters 433 and underlying maps (Sect. 2.4). For future assessment of model accuracy, we therefore 434 suggest to compare model results to actual measurements of the forest foliar Hg uptake 435 flux (Obrist et al., 2021; Feinberg et al., 2022). The total foliar Hg uptake flux to the 436 European forested area (Fig. 1 and 5) was 22 Mg Hg season⁻¹ for GEOS-Chem which 437 almost equals the total flux of 23 ± 12 Mg Hg season⁻¹ for the bottom-up model (Sect. 438 3.1).439

4 Conclusion

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We created a highly resolved (1 km^2) map (Fig. 1), which visualizes the spatial vari-441 ation of foliar Hg uptake fluxes to European forests. The highest foliar Hg uptake fluxes 442 receive Fennoscandia, densely forested areas in Central and Southern Europe, e.g. the 443 Carpathian Mountains, the Balkans, or multiple low mountain areas. We suggest, that 444 this map (Fig. 1) can guide decisions on European background Hg monitoring of the terrestrial environment. The total forest foliar Hg uptake flux over the course of one grow-446 ing season agrees well with Hg flux estimates derived from literature and from the chem-447 ical transport model GEOS-Chem for the same land area of Europe (Fig. 5). This pre-448 cision among modelling results on a European scale using different approaches gives us 449 confidence that the bottom-up model is overall able to represent the seasonal forest fo-450 liar Hg uptake flux. We suggest that the accuracy of modelling results have to be fur-451 ther determined using direct forest foliar Hg flux measurements. 452

Using an empirical relationship between Hg needle uptake rates of pine trees and 453 VPD threshold exceedance, we found a reduction in modelled pine forest needle Hg up-454 take flux during the relatively hot and dry growing season in Fennoscandia in 2018 com-455 pared to the growing seasons in 1994, 2003 and 2015/2017 (Fig. 2). The modelled av-456 erage amount of Hg, that was not deposited via pine needle uptake in 2018 compared 457 to the reference time period of 2015/17 exceeded the reported anthropogenic Hg emis-458 sions of Sweden in 2021, highlighting the quantitative significance of stomatal Hg up-459 take. If these hot summer droughts occurred more frequently in Fennoscandia under cli-460 mate change, the pine forest needle Hg uptake flux would be diminished while these ex-461 treme conditions prevail, potentially increasing the Hg burden of the ocean via long-range 462 atmospheric transport. In order to better represent the impact of extreme climate events 463 on the pine forest needle Hg uptake flux, we therefore advise to incorporate a stomatal 464 component of the pine needle Hg uptake flux into chemical transport models like GEOS-465 Chem. 466

5 Open Research 467

Calculated forest foliar Hg uptake fluxes to Europe (Fig. 1) and GEOS-Chem sim-468 ulation data aggregated to seasonal values are publicly available for download from Zen-469 odo at https://zenodo.org/record/7851718#.ZFUeLM5Bw2w and https://zenodo.org/ 470 record/7900753#.ZFUgqM5Bw2w respectively. All input datasets to the bottom-up model 471 are described in detail in Section 2.1, along with their respective publications and databases, 472 from which the datasets can be accessed. The offline dry deposition code from GEOS-473 Chem is accessible for download (Feinberg, 2022) and model output data from GEOS-474 Chem can be obtained from the corresponding author upon request. All calculations and 475 visualizations were done in R, Version 4.0.3. 476

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- Part of the ICP Forests data was co-financed by the European Commission. ERA5 and
- 486 CORDEX data were obtained from the Copernicus Climate Change Service (C3S) Cli-
- mate Data Store (https://cds.climate.copernicus.eu/cdsapp#!/home). PROBA-V leaf
- area index values were downloaded from the VITO Product Distribution Portal. Climate
- calculations under two climate change scenarios (Fig. 2, Fig. 3) were performed at sci-
- ⁴⁹⁰ CORE (http://scicore.unibas.ch/) scientific computing center at University of Basel.

491 References

- Allen, C. D., Breshears, D. D., & McDowell, N. G. (2015). On underestimation of global vulnerability to tree mortality and forest die-off from hotter drought in the Anthropocene. *Ecosphere*, 6(8), art129. Retrieved 2021-10-22, from https://onlinelibrary.wiley.com/doi/abs/10.1890/ES15-00203.1 doi: 10.1890/ES15-00203.1
 Berg, A., & Sheffield, J. (2018, June). Climate Change and Drought: the Soil Mois-
- 498
 ture Perspective.
 Curr Clim Change Rep. 4(2), 180–191.
 Retrieved 2021

 499
 10-22, from https://doi.org/10.1007/s40641-018-0095-0
 doi: 10.1007/

 500
 s40641-018-0095-0
- Blackwell, B. D., Driscoll, C. T., Maxwell, J. A., & Holsen, T. M. (2014, June).
 Changing climate alters inputs and pathways of mercury deposition to forested
 ecosystems. *Biogeochemistry*, 119(1-3), 215–228. Retrieved 2020-03-10,
 from http://link.springer.com/10.1007/s10533-014-9961-6
 doi: 10.1007/s10533-014-9961-6
- Brando, P. M., Paolucci, L., Ummenhofer, C. C., Ordway, E. M., Hartmann,
- H., Cattau, M. E., ... Balch, J. (2019, May). Droughts, Wildfires, and Forest Carbon Cycling: A Pantropical Synthesis. Annu. Rev. Earth Planet. Sci., 47(1), 555–581. Retrieved 2021-10-22, from https:// www.annualreviews.org/doi/10.1146/annurev-earth-082517-010235
 doi: 10.1146/annurev-earth-082517-010235
- 512
 Brus, D. J., Hengeveld, G. M., Walvoort, D. J. J., Goedhart, P. W., Heidema, A. H.,

 513
 Nabuurs, G. J., & Gunia, K. (2012, January). Statistical mapping of tree

 514
 species over Europe. European Journal of Forest Research, 131(1), 145–

 515
 157. Retrieved 2019-08-26, from http://link.springer.com/10.1007/

 516
 \$10342-011-0513-5
- Buras, A., & Menzel, A. (2019). Projecting Tree Species Composition Changes
 of European Forests for 2061–2090 Under RCP 4.5 and RCP 8.5 Scenarios. Frontiers in Plant Science, 9, 1986. Retrieved 2021-09-22, from
 https://www.frontiersin.org/article/10.3389/fpls.2018.01986 doi:
 10.3389/fpls.2018.01986
- Buras, A., Rammig, A., & Zang, C. S. (2020, March). Quantifying impacts of the
 2018 drought on European ecosystems in comparison to 2003. *Biogeosciences*,
 17(6), 1655–1672. Retrieved 2021-12-07, from https://bg.copernicus.org/
 articles/17/1655/2020/ doi: 10.5194/bg-17-1655-2020
- C3S. (2022). Copernicus Climate Change Service, Climate Data Store (CDS):
 CORDEX regional climate model data on single levels. Retrieved from
 https://cds.climate.copernicus.eu/cdsapp#!/dataset/10.24381/
 cds.bc91edc3?tab=overview doi: 10.24381/cds.bc91edc3

530	CLRTAP. (2017). Revised Chapter 3 of the Manual on Methodologies and Cri-
531	teria for Modelling and Mapping Critical Loads and Levels and Air Pol-
532	lution Effects, Risks and Trends: Mapping Critical Levels for Vegetation.
533	Umweltbundesamt. Retrieved from https://www.umweltbundesamt.de/en/
534	manual-for-modelling-mapping-critical-loads-levels doi: $https://www$
535	. unweltbundes amt. de/en/manual-for-modelling-mapping-critical-loads-levels
536	Demers, J. D., Blum, J. D., & Zak, D. R. (2013, March). Mercury isotopes in a
537	forested ecosystem: Implications for air-surface exchange dynamics and the
538	global mercury cycle. Global Biochem. Cy., 27(1), 222–238. Retrieved 2018-
539	11-14, from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/
540	gbc.20021 doi: 10.1002/gbc.20021
541	Dierckx, W., Sterckx, S., Benhadj, I., Livens, S., Duhoux, G., Van Achteren, T.,
542	Saint, G. (2014, April). PROBA-V mission for global vegetation
543	monitoring: standard products and image quality. Int. J. Remote Sens.,
544	35(7), $2589-2014$. Retrieved 2021-08-30, from https://doi.org/10.1080/
545	01431161.2014.883097 doi: 10.1080/01431161.2014.883097
546	Enrico, M., Roux, G. L., Marusczak, N., Heimburger, LE., Claustres, A., Fu, X.,
547	dominated by gasaous elemental marcury dry deposition <u>Fraviran</u> Sai Tach
548	a dominated by gaseous elemental mercury dry deposition. Entrino 3.52 and 5.62 and 5.62 and 2.61 and 2.62 from http://puba.aca.org/
549	doi/(10, 1021/acg, est, 5b06058, doi: 10.1021/acg, est, 5b06058
550	Four $A = (2022)$ Code reference: prifein /offline dryden: Offline dry denosition
551	model from GEOS-Chem v1.0 (v1.0)
553	doi: https://doi.org/10.5281/zenodo.6498126
555	Feinberg A Dlamini T Jiskra M Shah V & E Selin N (2022) Eval-
555	uating atmospheric mercury (Hg) uptake by vegetation in a chemistry-
556	transport model. Environmental Science: Processes & Impacts. Retrieved
557	2022-07-25. from https://pubs.rsc.org/en/content/articlelanding/
558	2022/em/d2em00032f (Publisher: Royal Society of Chemistry) doi:
559	10.1039/D2EM00032F
560	Feinberg, A., Jiskra, M., Borrelli, P., Biswakarma, J., & Selin, N. E. (2023, Jan-
561	uary). Land use change as an anthropogenic driver of mercury pollution. Re-
562	trieved 2023-04-15, from https://eartharxiv.org/repository/view/4963/
563	(Publisher: EarthArXiv)
564	Forrester, D. I., Tachauer, I. H. H., Annighoefer, P., Barbeito, I., Pretzsch, H., Ruiz-
565	Peinado, R., Sileshi, G. W. (2017, July). Generalized biomass and leaf area
566	allometric equations for European tree species incorporating stand structure,
567	tree age and climate. Forest Ecology and Management, 396, 160–175. Re-
568	trieved 2020-06-10, from http://www.sciencedirect.com/science/article/
569	pii/S0378112717301238 doi: 10.1016/j.foreco.2017.04.011
570	Forzieri, G., Feyen, L., Rojas, R., Flörke, M., Wimmer, F., & Bianchi, A. (2014,
571	January). Ensemble projections of future streamflow droughts in Europe.
572	Hydrology and Earth System Sciences, 18(1), 85–108. Retrieved 2021-12-17,
573	from https://hess.copernicus.org/articles/18/85/2014/ (Publisher:
574	Copernicus GmbH) doi: 10.5194/ness-18-85-2014
575	Fuster, B., Sanchez-Zapero, J., Camacho, F., Garcia-Santos, V., Verger, A., La-
576	LAL fAPAR and fCOVER collection 300 m products of Constraints Clobal
577	Land Service Remote Sense 19(6) 1017 Retrieved 2021 03 17 from
570	https://www.mdpi_com/2072-4292/12/6/1017_doi: 10.3390/rs12061017
213	Garonna I. Jong R. d. Wit A. I. W. d. Mücher, C. A. Schmid, R. & Schoop
581	man, M. E. (2014). Strong contribution of autumn phenology to changes in
582	satellite-derived growing season length estimates across Europe (1982–2011).
583	Glob. Change Biol., 20(11), 3457–3470. Retrieved 2019-10-17. from
584	https://onlinelibrary.wiley.com/doi/abs/10.1111/gcb.12625 doi:
504	

585	$10.1111/{ m gcb}.12625$
586	Gharun, M., Hörtnagl, L., Paul-Limoges, E., Ghiasi, S., Feigenwinter, I., Burri,
587	S., Buchmann, N. (2020, September). Physiological response of Swiss
588	ecosystems to 2018 drought across plant types and elevation. <i>Philosophical</i>
589	Transactions of the Royal Society B: Biological Sciences, 375(1810), 20190521.
590	Retrieved 2023-04-16, from https://royalsocietypublishing.org/
591	doi/full/10.1098/rstb.2019.0521 (Publisher: Royal Society) doi:
592	10.1098/rstb.2019.0521
593	Gibbs, H. K. (2006). Olson's Major World Ecosystem Complexes Ranked by Carbon
594	in Live Vegetation: an Updated Database Using the GLC2000 Land Cover
595	Product (NDP-017b).
596	doi: https://www.osti.gov/biblio/1389498
597	Grossiord, C., Buckley, T. N., Cernusak, L. A., Novick, K. A., Poulter, B., Sieg-
598	wolf, R. T. W., McDowell, N. G. (2020). Plant responses to rising vapor
599	pressure deficit. New Phytol., 226(6), 1550–1566. Retrieved 2020-09-07, from
600	https://nph.onlinelibrary.wiley.com/doi/abs/10.1111/nph.16485 doi:
601	10.1111/nph.16485
602	Hararuk, O., Obrist, D., & Luo, Y. (2013, April). Modelling the sensitivity of
603	soil mercury storage to climate-induced changes in soil carbon pools. Bio-
604	<i>geosciences</i> , 10(4), 2393–2407. Retrieved 2021-10-27, from https://bg
605	.copernicus.org/articles/10/2393/2013/ doi: 10.5194/bg-10-2393-2013
606	Hersbach, H., Bell, B., Berrisford, P., Biavati, G., Horányi, A., Muñoz Sabater, J.,
607	Thépaut, JN. (2018). ERA5 hourly data on pressure levels from 1979
608	to present. Copernicus Climate Change Service (C3S) Climate Data Store
609	(CDS).
610	doi: 10.24381/cds.bd0915c6
611	IPCC. (2021a). Climate change 2021. The physical science basis. Working group I
612	contribution to the Sixth Assessment Report of the Intergovernmental Panel
613	on Climate Change.
614	IPCC. (2021b). Summary for Policymakers. Climate Change 2021: The Physical
615	Science Basis, Contribution of Working Group I to the Sixth Assessment Re-
616	port of the Intergovernmental Panel on Climate Change. Cambride University
617	Press.
618	Jacob, D., Teichmann, C., Sobolowski, S., Katragkou, E., Anders, L. Belda, M.,
619	Wulfmever, V. (2020, April), Regional climate downscaling over Eu-
620	rope: perspectives from the EURO-CORDEX community. Rea Environ
621	Change, 20(2), 51. Retrieved 2021-11-03, from https://doi.org/10.1007/
622	s10113-020-01606-9 doi: 10.1007/s10113-020-01606-9
623	Jeong, SJ., Ho, CH., Gim, HJ., & Brown, M. E. (2011). Phenology shifts at
624	start vs. end of growing season in temperate vegetation over the Northern
625	Hemisphere for the period 1982–2008. Global Change Biology, 17(7), 2385–
626	2399. Retrieved 2021-12-20, from https://onlinelibrary.wilev.com/doi/
627	abs/10.1111/j.1365-2486.2011.02397.x doi: 10.1111/j.1365-2486.2011
628	.02397.x
629	Jiskra, M., Sonke, J. E., Obrist, D., Bieser, J., Ebinghaus, R., Myhre, C. L.
630	Dommergue, A. (2018). A vegetation control on seasonal variations
631	in global atmospheric mercury concentrations. Nat. Geosci., 1–7. Re-
632	trieved from http:https://doi.org/10.1038/s41561-018-0078-8 doi:
633	10.1038/s41561-018-0078-8
634	Jiskra, M., Wiederhold, J. G., Skyllberg, U., Kronberg, RM., Haidas, I., & Kret-
635	zschmar, R. (2015). Mercury deposition and re-emission pathways in boreal
636	forest soils investigated with Hg isotope signatures. Environ. Sci. Technol.,
637	49(12), 7188–7196.
638	Kellomäki, S., Strandman, H., Heinonen, T., Asikainen, A., Venäläinen, A., & Pel-
639	tola, H. (2018, March). Temporal and Spatial Change in Diameter Growth

640	of Boreal Scots Pine, Norway Spruce, and Birch under Recent-Generation
641	(CMIP5) Global Climate Model Projections for the 21st Century. Forests,
642	9(3), 118. Retrieved 2021-12-16, from https://www.mdpi.com/1999-4907/9/
643	3/118 doi: 10.3390/f9030118
644	Khan, T., Obrist, D., Agnan, Y., E. Selin, N., & A. Perlinger, J. (2019).
645	Atmosphere-terrestrial exchange of gaseous elemental mercury: parameteriza-
646	tion improvement through direct comparison with measured ecosystem fluxes
647	Environmental Science: Processes & Imnacts 21(10) 1609-1712 Retrieved
047	2021 10 20 from https://pubs.rsc.org/op/content/articlelanding/
648	2021-10-20, from fittps://pubs.isc.org/en/content/articleranding/
649	$V_{10} = U_{10}(C) = V_{10}(C) + V_{10}($
650	Ku, H. (1966). Notes on the use of propagation of error formulas. Journal of Re-
651	search of the National Bureau of Standards, 700 (4).
652	Korner, C. (2013). Plant-Environment Interactions. In A. Bresinsky, C. Korner,
653	J. W. Kadereit, G. Neunaus, & U. Sonnewald (Eds.), Strasourger's Plant
654	Sciences: Including Prokaryotes and Fungi (pp. 1065–1166). Berlin, Hei-
655	delberg: Springer. Retrieved 2019-11-25, from https://doi.org/10.100//
656	978-3-642-15518-5_12
657	Laacouri, A., Nater, E. A., & Kolka, R. K. (2013, September). Distribution and
658	uptake dynamics of mercury in leaves of common deciduous tree species in
659	Minnesota, U.S.A. Environ. Sci. Technol., $47(18)$, 10462–10470. Retrieved
660	2019-06-20, from http://pubs.acs.org/doi/abs/10.1021/es401357z doi:
661	10.1021/es401357z
662	Lagergren, F., & Lindroth, A. (2002, August). Transpiration response to soil mois-
663	ture in pine and spruce trees in Sweden. Agr. Forest Meteorol., 112(2), 67–85.
664	Retrieved 2021-08-31, from https://www.sciencedirect.com/science/
665	article/pii/S0168192302000606 doi: 10.1016/S0168-1923(02)00060-6
666	Liu, Y., Liu, G., Wang, Z., Guo, Y., Yin, Y., Zhang, X., Jiang, G. (2021, Octo-
667	ber). Understanding foliar accumulation of atmospheric Hg in terrestrial veg-
668	etation: Progress and challenges. Critical Reviews in Environmental Science
669	and Technology, $\theta(0)$, 1-22. Retrieved 2021-10-18, from https://doi.org/10
670	.1080/10643389.2021.1989235 doi: 10.1080/10643389.2021.1989235
671	Lucchesi, R. (2018). File Specification for GEOS FP. GMAO Office Note No. 4 (Ver-
672	sion 1.2). Retrieved from http://gmao.gsfc.nasa.gov/pubs/office_notes
673	Mauri, A., Strona, G., & San-Miguel-Ayanz, J. (2017, January). EU-Forest, a high-
674	resolution tree occurrence dataset for Europe. Sci Data, 4. Retrieved 2020-01-
675	29, from https://www.ncbi.nlm.nih.gov/pmc/articles/PMC5215058/ doi:
676	10.1038/sdata.2016.123
677	Norby, R. J., & Zak, D. R. (2011). Ecological Lessons from Free-Air CO2
678	Enrichment (FACE) Experiments. Annual Review of Ecology. Evo-
679	lution. and Sustematics, $42(1)$, $181-203$. Retrieved 2021-10-22, from
680	https://doi.org/10.1146/annurey-ecolsys-102209-144647 doi:
681	10.1146/annurev-ecolsvs-102209-144647
692	Obrist D. Agnan V. Jiskra M. Olson C. L. Colegrove D. P. Hueber, J.
692	Helmig D (2017 July) Tundra untake of atmospheric elemental mercury
684	drives Arctic mercury pollution Nature $5/7(7662)$ 201–204 Retrieved
695	2019-06-24 from http://www.nature.com/articles/nature22997 doi:
696	10 1038/nature22007
607	Obrist D Boy F M Harrison I L Kwong C F Munger I W Moosmüller
687	H Commane R (2021 July) Previously unaccounted atmospheric mer-
688	D_{11} , Communic, II. (2021, Jury). The robusty unaccounted atmospheric mer-
089	2021_07_{-26} from https://unu ppg.org/content/118/20/o2105/77119 doi:
090	10 1073/pnas 2105477118
091	Oksanon F. Libawainon I. Koinänon M. Koski Saeri S. Kontunon Sonnala S.
692	Sollin Δ is Sober Δ (2010) Northern Forest Trees Under Ingrossing Atmos
693	spheria Humidity In F. M. Conorrea, H. Lüttrea, D. Matyczalz, J. H. Destrach
694	spheric frumulty. In F. W. Canovas, C. Luttge, R. Matyssek, & H. Fletzsch

695 696 697	(Eds.), <i>Progress in Botany Vol. 80</i> (pp. 317–336). Cham: Springer Interna- tional Publishing. Retrieved 2021-12-17, from https://doi.org/10.1007/ 124_2017_15						
698 699	Pan, Y., Birdsey, R. A., Phillips, O. L., & Jackson, R. B. (2013, November). The Structure, Distribution, and Biomass of the World's Forests. <i>Annu. Rev.</i>						
700	Ecol. Evol. Syst., 44(1), 593-622. Retrieved 2022-01-10, from https://						
701	www.annualreviews.org/doi/10.1146/annurev-ecolsys-110512-135914						
702	doi: 10.1146/annurev-ecolsys-110512-135914						
703	Panagos, P., Jiskra, M., Borrelli, P., Liakos, L., & Ballabio, C. (2021, October).						
704	Mercury in European topsoils: Anthropogenic sources, stocks and fluxes. En-						
705	vironmental Research, 201, 111556. Retrieved 2021-07-29, from https://						
706	www.sciencedirect.com/science/article/pii/S0013935121008501 doi:						
707	10.1016/j.envres.2021.111556						
708	Panek, J. A., & Goldstein, A. H. (2001, March). Response of stomatal conduc-						
709	tance to drought in ponderosa pine: implications for carbon and ozone uptake.						
710	Tree Physiol., 21(5), 337-344. Retrieved 2021-07-22, from https://doi.org/						
711	10.1093/treephys/21.5.337 doi: 10.1093/treephys/21.5.337						
712	Papula, L. (2003). Mathematische Formelsammlung für Ingenieure und Naturwis-						
713	senschaftler (8th ed.). Vieweg.						
714	Pleijel, H., Klingberg, J., Nerentorp, M., Broberg, M. C., Nyirambangutse, B.,						
715	Munthe, J., & Wallin, G. (2021, December). Mercury accumulation						
716	in leaves of different plant types – the significance of tissue age and spe-						
717	cific leaf area. $Biogeosciences, 18(23), 6313-6328$. Retrieved 2021-12-						
718	23, from https://bg.copernicus.org/articles/18/6313/2021/ doi: 10.5104/b = 18.6212.2021						
719	10.3194/0g-18-0313-2021						
720	Rea, A. W., Lindberg, S. E., Scherbatskoy, I., & Keeler, G. J. (2002). Mercury ac-						
721	tor Air Soil Poll 122 40.67						
(22	Putter A P Schouer I I Shafer M M Crossvell I F Olson M P Behinson						
723	M Mallek I. L. (2011 February) Dry deposition of assesse elemen-						
724	tal mercury to plants and soils using mercury stable isotopes in a controlled						
726	environment. Atmos. Environ. 45(4), 848–855. Retrieved 2019-09-24. from						
727	http://www.sciencedirect.com/science/article/pii/S1352231010009805						
728	doi: 10.1016/j.atmosenv.2010.11.025						
729	Samaniego, L., Thober, S., Kumar, R., Wanders, N., Rakovec, O., Pan, M.,						
730	Marx, A. (2018, May). Anthropogenic warming exacerbates European soil						
731	moisture droughts. Nature Clim Change, 8(5), 421–426. Retrieved 2021-						
732	12-17, from https://www.nature.com/articles/s41558-018-0138-5 doi:						
733	10.1038/s41558-018-0138-5						
734	Selin, N. E., Jacob, D. J., Yantosca, R. M., Strode, S., Jaeglé, L., & Sunderland,						
735	E. M. (2008). Global 3-D land-ocean-atmosphere model for mercury:						
736	Present-day versus preindustrial cycles and anthropogenic enrichment fac-						
737	tors for deposition. Global Biogeochemical Cycles, $22(2)$. Retrieved 2020-09-						
738	22, from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/						
739	2007GB003040 doi: 10.1029/2007GB003040						
740	Smith-Downey, N. V., Sunderland, E. M., & Jacob, D. J. (2010). Anthro-						
741	pogenic impacts on global storage and emissions of mercury from terres-						
742	trial soils: Insights from a new global model. Journal of Geophysical Re- 145(GP) Detries 10000 and 0000 for the solution of the solution o						
743	search: Biogeosciences, 115(G3). Retrieved 2019-06-20, from https://						
744	agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2009JG001124 doi:						
745	10.1029/2009JG001124						
746	Sonke, J. E., Angot, H., Zhang, Y., Poulain, A., Bjorn, E., & Schartup, A. (2023,						
747	52(5) 853-876 Bothiored 2022 04 15 from https://doi.org/10.1007/						
748	$\sigma_{2}(0), \sigma_{2}=\sigma_{1}0.$ netheved 2020-04-10, non netps://doi.org/10.100//						
749	515200 025 01655 y doi: 10.1001/515260-025-01655-y						

750	Stocker, B. D., Zscheischler, J., Keenan, T. F., Prentice, I. C., Seneviratne, S. I.,
751	& Peñuelas, J. (2019, April). Drought impacts on terrestrial primary pro-
752	duction underestimated by satellite monitoring. Nat. Geosci., 12(4), 264–
753	270. Retrieved 2021-10-22, from https://www.nature.com/articles/
754	s41561-019-0318-6 doi: 10.1038/s41561-019-0318-6
755	Wang, Q., Tenhunen, J., Dinh, N., Reichstein, M., Otieno, D., Granier, A., & Pile-
756	garrd, K. (2005, June). Evaluation of seasonal variation of MODIS derived
757	leaf area index at two European deciduous broadleaf forest sites. Remote
758	Sensing of Environment, $96(3-4)$, $475-484$. Retrieved 2019-05-15, from
759	https://linkinghub.elsevier.com/retrieve/pii/S0034425705001252
760	doi: 10.1016/j.rse.2005.04.003
761	Wang, R., Chen, J. M., Liu, Z., & Arain, A. (2017, August). Evaluation of seasonal
762	variations of remotely sensed leaf area index over five evergreen coniferous
763	torests. ISPRS Journal of Photogrammetry and Remote Sensing, 130, 187–
764	201. Retrieved 2020-10-13, from http://www.sciencedirect.com/science/
765	article/pii/S0924271616305998 doi: 10.1016/j.isprsjprs.2017.05.017
766	Wang, Y., Jacob, D. J., & Logan, J. A. (1998). Global simulation of tropospheric
767	03-NO x -hydrocarbon chemistry: 1. Model formulation. Journal of Geophys-
768	<i>ical Research: Atmospheres</i> , 103(D9), 10713–10725. Retrieved 2022-05-25,
769	(oprint: https://onlinelibrary.wiley.com/doi/abs/10.1029/98JD00158)
770	$(_eprint: nttps://oninenbrary.wney.com/doi/pdi/10.1029/965D00158) doi: 10.1020/08 ID00158$
771	Weedy M. L. (2007 January) Parameterization of surface resistances
772	to gaseous dry deposition in regional-scale numerical models
774	spheric Environment, 41, 52–63. Retrieved 2021-10-21, from https://
775	www.sciencedirect.com/science/article/pii/S1352231007009740 doi:
776	10.1016/j.atmosenv.2007.10.058
777	Wohlgemuth, L., Osterwalder, S., Joseph, C., Kahmen, A., Hoch, G., Alewell, C., &
778	Jiskra, M. (2020, December). A bottom-up quantification of foliar mercury
779	uptake fluxes across Europe. <i>Biogeosciences</i> , 17(24), 6441–6456. Retrieved
780	2021-01-14, from https://bg.copernicus.org/articles/17/6441/2020/
781	doi: https://doi.org/10.5194/bg-17-6441-2020
782	Wohlgemuth, L., Rautio, P., Ahrends, B., Russ, A., Vesterdal, L., Waldner, P.,
783	Jiskra, M. (2022, March). Physiological and climate controls on foliar mercury
784	uptake by European tree species. <i>Biogeosciences</i> , $19(5)$, 1335–1353. Retrieved
785	2022-03-14, from https://bg.copernicus.org/articles/19/1335/2022/
786	(Publisher: Copernicus GmbH) doi: 10.5194/bg-19-1335-2022
787	Wu, Z., Wang, X., Chen, F., Turnipseed, A. A., Guenther, A. B., Niyogi, D.,
788	Alapaty, K. (2011, May). Evaluating the calculated dry deposition velocities of
789	reactive nitrogen oxides and ozone from two community models over a temper- ate desiduous fenerat. Atmospheric Environment $15(16)$, 2662, 2674. Detriared
790	ate deciduous forest. Atmospheric Environment, 45(10), 2005–2014. Retrieved
791	2021-10-21, HOIH HUUPS://WWW.SCIENcedifect.com/Science/article/pii/
792	Vuon H. Dai V. Viao Z. Ji D. & Shangguan W. (2011 May) – Paprocessing
793	the MODIS Leef Area Index products for land surface and climate mod
794	elling Remote Sensing of Environment 115(5) 1171–1187 Retrieved
796	2023-04-16. from https://www.sciencedirect.com/science/article/pij/
797	S0034425711000149 doi: 10.1016/j.rse.2011.01.001
798	Yuan, W., Sommar, J., Lin, CJ., Wang, X., Li, K., Liu, Y., Feng, X. (2019)
799	January). Stable isotope evidence shows re-emission of elemental mercury
800	vapor occurring after reductive loss from foliage. Environ. Sci. Technol.,
801	53(2), 651-660. Retrieved 2019-11-29, from https://doi.org/10.1021/
802	acs.est.8b04865 doi: 10.1021/acs.est.8b04865
803	Yuan, W., Zheng, Y., Piao, S., Ciais, P., Lombardozzi, D., Wang, Y., Yang,
804	S. (2019, August). Increased atmospheric vapor pressure deficit re-

805	duces global vegetation growth. Sci. Adv., 5(8). Retrieved 2019-09-06,						
806	from https://advances.sciencemag.org/content/5/8/eaax1396 doi:						
807	10.1126/sciadv.aax1396						
808	Zhang, H., Holmes, C., & Wu, S. (2016, September). Impacts of changes						
809	in climate, land use and land cover on atmospheric mercury. Atmo-						
810	spheric Environment, 141, 230–244. Retrieved 2021-10-21, from https://						
811	linkinghub.elsevier.com/retrieve/pii/S1352231016304897 doi:						
812	10.1016/j.atmosenv. $2016.06.056$						
813	Zhang, L., Wright, L. P., & Blanchard, P. (2009, December). A review of cur-						
814	rent knowledge concerning dry deposition of atmospheric mercury. At -						
815	mos. Environ., 43(37), 5853–5864. Retrieved 2019-10-03, from http://						
816	www.sciencedirect.com/science/article/pii/S1352231009007407 doi:						
817	10.1016/j.atmosenv.2009.08.019						
818	Zhou, J., & Obrist, D. (2021). Global mercury assimilation by vegetation. <i>Environ</i> .						
819	Sci. Technol Retrieved 2021-10-08, from https://doi.org/10.1021/acs.est						
820	.1c03530 doi: 10.1021/acs.est.1c03530						
821	Zhou, J., Obrist, D., Dastoor, A., Jiskra, M., & Ryjkov, A. (2021, March). Vege-						
822	tation uptake of mercury and impacts on global cycling. Nat. Rev. Earth Env-						
823	iron., 1-16. Retrieved 2021-03-16, from https://www.nature.com/articles/						
824	s43017-021-00146-y doi: 10.1038/s43017-021-00146-y						
825	Zweifel, R., Rigling, A., & Dobbertin, M. (2009). Species-specific stomatal response						
826	of trees to drought – a link to vegetation dynamics? J. Veg. Sci., $20(3)$, 442–						
827	454. Retrieved 2021-09-08, from https://onlinelibrary.wiley.com/doi/						
828	abs/10.1111/j.1654-1103.2009.05701.x doi: 10.1111/j.1654-1103.2009						
829	.05701.x						
830	Zweifel, R., Steppe, K., & Sterck, F. J. (2007, June). Stomatal regulation by micro-						
831	climate and tree water relations: interpreting ecophysiological field data with a						
832	hydraulic plant model. J. Exp. Bot., $58(8)$, $2113-2131$. Retrieved 2021-04-19,						
833	from https://doi.org/10.1093/jxb/erm050 doi: 10.1093/jxb/erm050						

Supporting Information for "A spatial assessment of current and future foliar Hg uptake fluxes across European forests"

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- 2. Figures S1 and S2
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Text S1.

Zhou & Obrist, (2021) give an estimate of median global foliage Hg assimilation by evergreen needleleaf forests of 61 Mg Hg year⁻¹ and by deciduous broadleaf forests of 28 Mg Hg year⁻¹ (Zhou & Obrist, 2021, Table 1). This estimate is based on global data on foliar Hg concentrations and net foliar biomass production. The global land area of evergreen needleleaf forests is given as 6.17 Mio km² and of deciduous broadleaf forests as 1.12 Mio km² (Zhou & Obrist, 2021). Converted to the land area of coniferous (1.0 Mio

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 $\rm km^2$) and deciduous forests (0.73 Mio $\rm km^2$) in Europe, we obtain a total Hg assimilation of 10.2 Mg Hg year⁻¹ for European coniferous forests and of 18.4 Mg Hg year⁻¹ for European deciduous forests.

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Figure S1. Simulated areal cover (percent km⁻²) of pine forests in Europe (a) historically, (b) for the time period 2061 - 2090 under the climate change scenario RCP 4.5, and (c) RCP 8.5. Geographic distribution was derived from statistic mapping from Brus et al. 2012 and projected relative abundance probabilities under climate analogs from Buras et al. 2019 (see Sect. 2.1 for details).



Figure S2. Absolute difference map of model outputs of forest foliar Hg uptake fluxes (g Hg $\rm km^{-2} \ day^{-1}$) from the bottom-up model - GEOS-Chem.

Table S1. Overview of LAI correction factor for tree height (cf_{height}) and Hg uptake rate correction factor for needle age class (cf_{age}) obtained from Wohlgemuth et al., (2020).

:

Tree species group	\mathbf{cf}_{height}	\mathbf{cf}_{age}
Abies alba	0.68	0.79
mix of all broadleaf values	0.63	
Betula pendula	0.63	_
Carpinus betulus	0.63	_
mix of all conifer values	0.68	0.79
Fagus sylvatica	0.56	_
Fraxinus excelsior	0.63	_
Larix decidua	0.68	
Pinus cembra; Pinus mugo arborea; Pinus nigra	0.69	0.86
Pinus nigra subsp. laricio	0.08	0.80
Picea abies	0.68	0.79
Pinus pinaster	0.68	0.86
Pinus sylvestris	0.68	0.86
Pinus sylvestris	0.68	0.86
Pseudotsuga menziesii	0.68	0.86
Quercus cerris; Quercus frainetto;	0.7	
Quercus ilex; Quercus pubescens	0.7	-
Quercus robur; Quercus petraea	0.7	

Table S2. Dataset on proportion of tree species per land area matched to dataset on foliarHg uptake rates by tree species groups.

Matched and aggregated tree species groups					
Dataset on tree species Dataset on foliar Hg up-					
proportion of land area	take per tree species				
(Brus et al., 2012)	(Wohlgemuth et al., 2022)				
Abies spp	Abies alba				
Alnus spp; Broadleaved misc; Castanea spp;					
Eucalyptus spp; Populus spp; Robinia Spp	mix of an broadleaf values				
Betula spp	Betula pendula				
Carpinus spp	Carpinus betulus				
Conifers misc	mix of all conifer values				
Fagus spp	Fagus sylvatica				
Fraxinus spp	Fraxinus excelsior				
Larix spp	Larix decidua				
Picea spp	Picea abies				
Pipus mise	Pinus cembra; Pinus mugo arborea; Pinus nigra				
I mus misc	Pinus nigra subsp. laricio				
Pinus pinaster	Pinus pinaster				
Pinus sylvestris	Pinus sylvestris				
Pseudotsuga menziesii	Pseudotsuga menziesii				
Quoreus mise	Quercus cerris; Quercus frainetto;				
	Quercus ilex; Quercus pubescens				
Quercus robur & Quercus petraea	Quercus robur; Quercus petraea				

Table S3. Overview of all combinations of Global Climate Models (GCMs) – Regional Climate Models (RCMs) and ensemble members used for downloading simulated data of 2m air temperature and 2m relative humidity from the Copernicus Climate Data Store in the framework of the Coordinated Regional Climate Downscaling Experiment (CORDEX) for the two climate scenarios of Representative Concentration Pathways (RCPs) 4.5 and 8.5. RCM data on a high regional resolution (here: European domain; 0.11° x 0.11°) depend on output from GCMs for lateral and lower boundary conditions. Temporal resolution of downloaded data was 3 hours and time period was 2068-2082.

Climate scenario	GCM	RCM	ensemble member
	CNRM-CERFACS-CM5	KNMI-RACMO22E	r1i1p1
	ICHEC-EC-EARTH	DMI-HIRHAM5	r3i1p1
	ICHEC-EC-EARTH	GERICS-REMO2015	r12i1p1
	ICHEC-EC-EARTH	KNMI-RACMO22E	r1i1p1
	ICHEC-EC-EARTH	SMHI-RC4A	r12i1p1
	IPSL-CM5A-MR	SMHI-RC4A	r1i1p1
RCP 4.5	MOHC-HadGEM2-ES	GERICS-REMO2015	r1i1p1
	MOHC-HadGEM2-ES	KNMI-RACMO22E	r1i1p1
	MOHC-HadGEM2-ES	SMHI-RC4A	r1i1p1
	MPI-M-MPI-ESM-LR	SMHI-RC4A	r1i1p1
	NCC-NorESM1-M	DMI-HIRHAM5	r1i1p1
	NCC-NorESM1-M	GERICS-REMO2015	r1i1p1
	NCC-NorESM1-M	SMHI-RC4A	r1i1p1
	CCCma-CanESM2	CLMcom-CLM-CCLM4-8-17	r1i1p1
	CNRM-CERFACS-CM5	CLMcom-ETH-COSMO-crCLIM	r1i1p1
	CNRM-CERFACS-CM5	GERICS-REMO2015	r1i1p1
	ICHEC-EC-EARTH	CLMcom-ETH-COSMO-crCLIM	r12i1p1
	ICHEC-EC-EARTH	KNMI-RACMO22E	r12i1p1
	ICHEC-EC-EARTH	SMHI-RCA4	r12i1p1
RCP 8.5	IPSL-CM5A-MR	DMI-HIRHAM5	r1i1p1
	MIROC-MIROC5	CLMcom-CLM-CCLM4-8-17	r1i1p1
	MOHC-HadGEM2-ES	SMHI-RCA4	r1i1p1
	MPI-M-MPI-ESM-LR	DMI-HIRHAM5	r1i1p1
	NCC-NorESM1-M	CLMcom-ETH-COSMO-crCLIM	r1i1p1
	NCC-NorESM1-M	GERICS-REMO2015	r1i1p1
	NCC-NorESM1-M	KNMI-RACMO22E	r1i1p1

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Table S4. Overview of relative uncertainty values for different parameters per forest tree species group propagated by error propagation principle for every spatial tile of the forest foliar Hg uptake flux map (Fig. 1).

	relative uncertainty (ru)							
Tree species group"	LMA^{b}	DMA^c	$ $ LAI d	Forest a	$area^{e}$	cf_{age} f	pine^{g}	$ \operatorname{total}^h$
Abies spp	0.36					0.03	-	0.70
Betula spp	0.61					-	-	0.86
Broadleaved mixed	0.82					-	-	1.02
Carpinus spp	0.18					-	-	0.63
Fagus spp	0.47					-	-	0.77
Fraxinus spp	0.54	0.1	0.44		0.40	-	-	0.81
Larix spp	0.76	0.1	0.44		0.40	-	-	0.97
Picea spp	0.75					0.03	-	0.96
Pine	0.72					0.06	0.014	0.94
Pseudotsuga menziesii	0.69					0.06	-	0.92
Quercus misc	0.28					-	-	0.67
Quercus robur & Quercus petraea	0.52					-	-	0.80

a See Table S2 for an overview of tree species aggregated into tree species groups. For the group of mixed conifer species (Table S2), we calculated an average uncertainty value from total uncertainty values of coniferous needle tree species groups, which equals 0.90.

b Leaf mass per area (LMA) values per tree species group were obtained from Forrester et al., 2017. The relative uncertainty of LMA per tree species group was calculated including the range of all LMA values within each respective tree species group: (maximum LMA - minimum LMA)/(average LMA).

c Foliar Hg values were obtained from a dataset of Hg concentrations in foliage samples of the ICP Forests biomonitoring network and the Austrian Bio-Indicator Grid measured using a direct mercury analyzer (DMA) (see Sect. ?? and Wohlgemuth et al., 2022). A DMA measurement sequence of foliar Hg concentrations was accepted when primary liquid reference standards did not deviate by more than $\pm 10\%$ from target value.

d Relative root mean square deviation (RMSD) of the leaf area index (LAI) product from PROBA-V from LAI ground observations evaluated by Fuster et al. (2020).

e The uncertainty of the proportion of tree species per forest land area (spatial resolution: 1 km^2) was not evaluated by Brus et al., 2012 and depends on the heterogenous availability of national forest inventories in Europe. From the overall accuracy given in Brus et al., 2012, we estimated a relative uncertainty value of 0.4 per tree species and km².

f Uncertainty of the correction factor (cf_{age}) for upscaling Hg uptake rates of needles of different age classes to whole coniferous evergreen trees (see Wohlgemuth et al., 2020).

g Relative RMSD of the linear regression slope of the average daily pine needle Hg uptake rate vs. time proportion of VPD > 1.2 kPa (see Wohlgemuth et al., 2022).

h Total species uncertainty =
$$\sqrt{ru_{LMA}^2 + ru_{DMA}^2 + ru_{LAI}^2 + ru_{for.area}^2 + ru_{cfage}^2 + ru_{pine}^2}$$

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