An assessment of extra-tropical cyclone precipitation extremes over the Southern Hemisphere using ERA5

Cameron McErlich¹, Adrian J. McDonald¹, James Arthur Renwick², and Alex J Schuddeboom¹

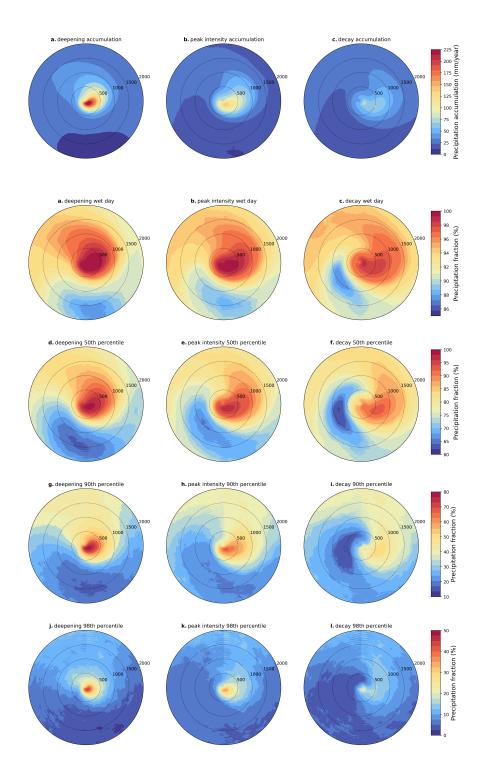
¹University of Canterbury ²Victoria University of Wellington

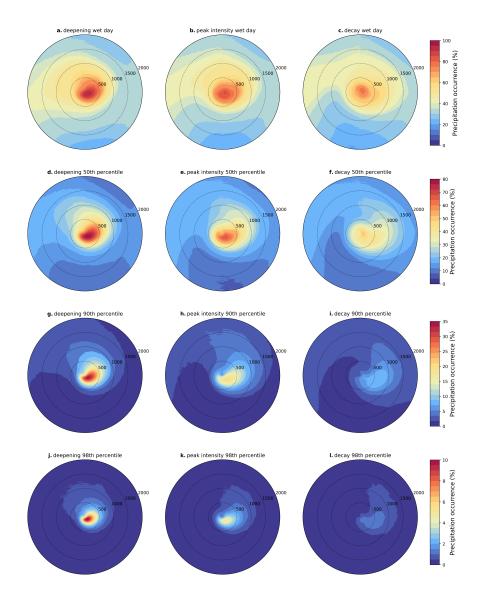
April 29, 2023

Abstract

ERA5 reanalysis is used to examine extreme precipitation using a spatially dependent precipitation threshold applied within a cyclone compositing framework. This is used to account for regional variation in precipitation generating processes within Southern Hemisphere mid-latitude cyclones across the cyclone lifecycle. The spatial extent of extreme precipitation is limited to a smaller region around the cyclone centre compared to non-extreme precipitation, though extreme precipitation displays a good spatial correlation with non-extreme precipitation. Extreme precipitation occurs more often during the deepening phase of the cyclone before it reaches a maximum depth. Precipitation occurrence at the 90th and 98th percentiles reduces to 46% and 30% of the deepening value across the cyclone lifecycle, averaged over the composite. Precipitation fraction at the 90th and 98th percentile reduces to 80% and 60% of the deepening value. Our methodology provides a quantitative assessment of precipitation extremes both spatially and temporally, within a cyclone compositing framework.

a. precipitation occurrence correlation				b. precipitation fraction correlation			
wet day : 50th	0.97	0.97	0.96	wet day : 50th	0.97	0.96	0.92
wet day : 90th	0.9	0.85	0.65	wet day : 90th	0.93	0.9	0.68
wet day : 98th	0.85	0.83	0.68	wet day : 98th	0.83	0.82	0.76
50th : 90th	0.97	0.94	0.79	50th : 90th	0.97	0.96	0.86
50th : 98th	0.92	0.91	0.81	50th : 98th	0.85	0.85	0.82
90th : 98th	0.98	0.97	0.97	90th : 98th	0.94	0.91	0.85
	Deepening	Peak intensity	Decay		Deepening	Peak intensity	Decay
c. precipitation occurrence average							
	c. precipita	tion occurren	ce average		d. precipi	tation fractio	n average
wet day		tion occurren 0.47 (97%)	ce average 0.45 (92%)	wet day	d. precipi 0.95 (100%)	tation fraction	n average 0.94 (100%)
wet day 50th	0.49 (100%)			,			
,	0.49 (100%) 0.28 (100%)	0.47 (97%)	0.45 (92%)	50th	0.95 (100%)	0.95 (100%)	0.94 (100%)
50th	0.49 (100%) 0.28 (100%) 0.06 (100%)	0.47 (97%) 0.26 (93%)	0.45 (92%) 0.23 (81%)	50th 90th	0.95 (100%) 0.82 (100%)	0.95 (100%) 0.82 (100%)	0.94 (100%)





An assessment of extra-tropical cyclone precipitation extremes over the Southern Hemisphere using ERA5

Cameron McErlich¹, Adrian McDonald^{1,2}, James Renwick³, and Alex Schuddeboom¹

¹School of Physical and Chemical Sciences, University of Canterbury, Christchurch, New Zealand ²Gateway Antarctica, University of Canterbury, Christchurch, New Zealand ³Victoria University of Wellington, School of Geography, Environment and Earth Science, Wellington, New Zealand

Key Points:

3

5 6 7

8

9

10	•	We detail a new methodology to assess precipitation extremes within cyclone com-
11		posites using a spatially dependent precipitation threshold
12	•	Extreme precipitation occurs preferentially and makes up a larger fraction of to-
13		tal accumulation before cyclones reach maximum depth
14	•	As extremes increase in intensity precipitation is constrained closer to the cyclone
15		centre and weakens more rapidly as the cyclone evolves

Corresponding author: Cameron McErlich, cameron.mcerlich@pg.canterbury.ac.nz

16 Abstract

ERA5 reanalysis is used to examine extreme precipitation using a spatially dependent 17 precipitation threshold applied within a cyclone compositing framework. This is used 18 to account for regional variation in precipitation generating processes within Southern 19 Hemisphere mid-latitude cyclones across the cyclone lifecycle. The spatial extent of ex-20 treme precipitation is limited to a smaller region around the cyclone centre compared 21 to non-extreme precipitation, though extreme precipitation displays a good spatial cor-22 relation with non-extreme precipitation. Extreme precipitation occurs more often dur-23 ing the deepening phase of the cyclone before it reaches a maximum depth. Precipita-24 tion occurrence at the 90th and 98th percentiles reduces to 46% and 30% of the deep-25 ening value across the cyclone lifecycle, averaged over the composite. Precipitation frac-26 tion at the 90th and 98th percentile reduces to 80% and 60% of the deepening value. Our 27 methodology provides a quantitative assessment of precipitation extremes both spatially 28 and temporally, within a cyclone compositing framework. 29

³⁰ Plain Language Summary

Extra-tropical cyclones play a major role in the circulation within the atmosphere, 31 acting to transfer heat towards the poles. Here we assess the representation of extreme 32 precipitation within extra-tropical cyclones. By applying a threshold for precipitation 33 that changes with geographic location, we are able to determine how extreme precipi-34 tation varies within a cyclone-centred coordinate system. When breaking cyclones into 35 lifecycle stages representing deepening, peak intensity and decay, we find that extreme 36 precipitation occurs most often as the cyclone is developing. The area of the cyclone rel-37 evant for extremes reduces towards the cyclone centre as the threshold for determining 38 extreme precipitation increases. Extreme precipitation weakens at a higher rate as cy-39 clones becomes more intense, highlighting the importance of extremes in the growth phase 40 of the cyclone. 41

42 **1** Introduction

Climate change is experienced by society through a variety of ways, including ex-43 treme weather events. Unlike long term climate trends which seem distant and occur more 44 gradually, extremes are direct and occur in everyday life. Howe et al. (2014) suggests that 45 people tend to accurately recall and report experiences with extreme weather, with in-46 creasing likelihood based on proximity and magnitude of an event. The impacts extremes 47 have are numerous. Economic impacts include closure of roads, outages of power grids, 48 water shut-offs, and physical damage to buildings, bridges, crops and livestock (Jahn, 49 2015). Environmental impacts include coastal erosion, changes in water supply and land 50 coverage (Seneviratne et al., 2012; Seddon et al., 2016; Seneviratne et al., 2021). Soci-51 etal impacts include food and water availability, loss of life, increasing insurance costs 52 and changes in property values (Morss et al., 2011; Bell et al., 2018; Zscheischler et al., 53 2018; Konapala et al., 2020). Extreme weather has intensified in recent decades, and will 54 continue to have a disproportionately large impact on the environment, society and the 55 economy (Seneviratne et al., 2021). 56

On a global scale, precipitation extremes are predicted to increase in intensity and 57 frequency as the climate warms (e.g. Zhang et al., 2007; Min et al., 2011; Hirsch & Arch-58 field, 2015). Work detailed in Kotz et al. (2022) has recently identified that increases in 59 extreme rainfall reduce economic growth rates. A study by Pendergrass and Knutti (2018) 60 investigated the uneven nature of precipitation, finding that half the annual precipita-61 tion occurs during the wettest 12 days of the year. When assessing output from CMIP5 62 climate models, they also found a shortening of the average number of days needed to 63 reach half the annual precipitation highlighting the increased importance of extreme events. 64

Extra-tropical cyclones (hereafter referred to as cyclones) are key components of 65 the atmospheric general circulation due to their ability to transport large quantities of 66 heat, moisture, and momentum. Cyclones are an important contributor to extreme weather 67 events as their passage is associated with strong winds, precipitation, and temperature 68 changes (Papritz et al., 2014). Studies quantifying cyclone-associated precipitation find 69 that up to 90% of precipitation in the mid-latitude storm tracks is associated with frontal 70 systems and their associated cyclones (Catto et al., 2012; Hawcroft et al., 2012; Utsumi 71 et al., 2017). Pfahl and Wernli (2012) has also identified that a high percentage of pre-72 cipitation extremes are directly related to cyclones, with some locations having up to 80%73 of their precipitation extremes associated with cyclones. Utsumi et al. (2017) also iden-74 tify that large amounts of extreme precipitation in mid-latitude regions are associated 75 with cyclones. McErlich, McDonald, Schuddeboom, et al. (2023) has also recently demon-76 strated that larger precipitation extremes are associated with regions where large scale 77 precipitation processes, such as cyclones, dominate. 78

In this study we use a regionally dependent precipitation threshold to classify precipitation extremes relative to the cyclone centre. Cyclone composites for both average and extreme precipitation are calculated from the ERA5 reanalysis. Composites are then partitioned into different stages of the lifecycle to assess the spatial and temporal evolution of extremes. The rate at which extremes precipitation changes throughout the cyclone lifecycle is then quantified.

⁸⁵ 2 Datasets and Methods

2.1 ERA5

We use output from the ERA5 reanalysis to identify cyclones over the Southern 87 Hemisphere for the years 1980 - 2019 inclusive. ERA5 is available on a 0.25° latitude/longitude 88 grid and at an hourly temporal resolution, however three hourly data was used in this 89 study. Previous work (McDonald & Cairns, 2020; McErlich, McDonald, Schuddeboom, 90 et al., 2023; McErlich, McDonald, Renwick, & Schuddeboom, 2023) shows that ERA5 91 is consistent with a number of satellite and reanalysis datasets for determining precip-92 itation. This past research also shows that ERA5 can be used to successfully examine 93 cyclonic structure within a compositing framework for 10m wind speeds, total column 94 water vapour, cloud liquid water, and precipitation. 95

96

86

2.2 Cyclone tracking and compositing methodology

Work by Crawford and Serreze (2016) introduces and details the mean sea level
pressure (MSLP) based cyclone tracking algorithm used in this study. Previous work in
McErlich, McDonald, Renwick, and Schuddeboom (2023) also describes application of
the cyclone tracking algorithm over the Southern Hemisphere using ERA5, but a short
summary is included in the following paragraphs.

Local minima in the ERA5 mean sea level pressure (MSLP) field are used to iden-102 tity cyclone centres. This covers the time period between 1980 - 2019 with a temporal 103 resolution of three hour and cyclones are identified using local minima in the MSLP field. 104 A radius-based threshold is also used to identify whether it is a closed low pressure sys-105 tem, and thus can be characterised as a cyclone. A maximum propagation speed of 150 106 kmhr⁻¹ is used to join related low pressure centres into continuous cyclone tracks. Cri-107 teria are also applied to reject systems that have a lifespan shorter than 24 hours, a track 108 length less than 100km, or do not spend some part of their lifetime at latitudes south 109 of 30°S. This causes cyclone tracks to predominately be concentrated over the South-110 ern Ocean (not shown; see Figure 1 McErlich, McDonald, Renwick, & Schuddeboom, 2023); 111 This is in agreement with previous research that analyzed Southern hemisphere cyclone 112 tracks (e.g. Hoskins & Hodges, 2005; Bengtsson et al., 2006; Hodges et al., 2011). 113

The identified cyclone centers are then used to transform ERA5 data into a cyclone 114 centered-coordinate system in the form of cyclone composites as detailed in McErlich, 115 McDonald, Renwick, and Schuddeboom (2023). Cyclone composites are calculated us-116 ing a radius of 2000 km, which is commonly used in previous work (e.g. Field & Wood, 117 2007; Field et al., 2008; Naud et al., 2012; Booth et al., 2018; McErlich, McDonald, Ren-118 wick, & Schuddeboom, 2023). Individual composites are rotated so that the direction 119 of propagation of the cyclone is chosen to be travelling eastward. Given the zonal west-120 erly winds over the Southern Ocean many cyclones require little rotation. This step ap-121 proximately aligns the position of the warm/cold fronts and the area of warm, moist air 122 associated with them. While not all fronts will be at the same position relative to the 123 direction of the cyclone, this rotation acts to focus the structure of the composite (Govekar 124 et al., 2011). 125

2.3 Analysis of cyclone lifecycle

To better understand changes in precipitation as the cyclone evolves, we partition 127 the cyclones into three distinct developmental phases. We classify cyclones relative to 128 the time of their maximum depth, which is defined as the time of maximum difference 129 between the edge pressure and central pressure of the cyclone. We define three phases 130 to represent periods of deepening, peak intensity and decay within the cyclone. The phase 131 of peak intensity is defined as 6 hours either side of the time of maximum depth. The 132 deepening phase is defined as measurements between 6 hours and 18 hours previous to 133 the time of maximum depth. The decay phase is defined as measurements between 6 and 134 18 hours after the time of maximum depth. 135

In order to partition the cyclones into phases of deepening, peak intensity, and de-136 cay, a further criterion based on the deepening rate $(\frac{\partial p}{\partial t}, \text{ scaled by latitude})$ was also as-137 sessed. Cyclone tracks were kept if the deepening rate changed from positive (indicat-138 ing strengthening) during the deepening phase to negative (indicating weakening) dur-139 ing the decay phase. Cyclone track that pass this criterion are masked to only include 140 data within the previously defined cyclone phases. Tracks without measurements 18 hours 141 before and after the point of peak intensity are rejected, causing a minimum cyclone lifes-142 pan of 36 hours within the subset of cyclone tracks used in this analysis. 143

144

2.4 Identification of cyclone extremes associated with cyclones

In order to identify regions of the cyclone that correspond to precipitation extremes, 145 we use a methodology introduced in McErlich, McDonald, Schuddeboom, et al. (2023). 146 This produces regionally dependent thresholds for extreme precipitation which are then 147 applied to a cyclone compositing framework. Firstly, a spatial map of precipitation wet 148 day frequency is produced using a 1 mm/day threshold to define a wet day over the South-149 ern hemisphere. This threshold is commonly used within the community (Polade et al., 150 2014; Schär et al., 2016) and is also used in a number of extreme precipitation indices 151 as defined by the Expert Team on Climate Change Detection and Indices (ETCCDI; Zhang 152 et al., 2011). 153

Secondly, rainfall data from regions of the same precipitation frequency are grouped
 together; This data is then aggregated to produce cumulative precipitation intensity dis tributions. McErlich, McDonald, Schuddeboom, et al. (2023) establishes this method ology forms coherent precipitation groupings, even through it connects spatially disparate
 regions together. The variability within wet day frequency regions has also been shown
 to be comparable to that within geographic regions.

Precipitation around the cyclone composite is then assessed to determine if it is
 an extreme for a given latitude/longitude based on the corresponding wet day frequency.
 Wet day frequency regions are identified within the cyclone composites using the geo-

graphic position of the cyclone centres, as determined by the tracking algorithm. A spa-163 tially dependent threshold for precipitation extremes defined by the locations wet day 164 frequency is then applied to the cyclone composites, removing data below the nth per-165 centile value of the aggregated precipitation. This process is repeated across each wet 166 day frequency region to account for changes in the underlying processes that generate 167 precipitation. This methodology produces masked cyclone composites consisting of only 168 precipitation at locations where it is above the spatially dependent nth percentile thresh-169 old. 170

171 Setting this threshold to the upper tail of the precipitation distribution determines cyclone composites for the extremes. Here we assess both the 90th and the 98th percentiles 172 allowing us to examine precipitation extremes. We also mask precipitation by a 1 mm 173 per day wet day threshold (Zhang et al., 2011) and 50th percentile value. While not in-174 dicative of extremes, these sets of composites provide data which the extreme compos-175 ites can be compared with, allowing us to understand unique features associated with 176 extreme precipitation. We note that the wet day frequency has a narrower range (see 177 Figure 1 McErlich, McDonald, Schuddeboom, et al., 2023) over the Southern Ocean where 178 the concentration of cyclones is the highest. This means there will be less variability in 179 the extreme precipitation thresholds over the Southern Ocean than for a global analy-180 181 sis.

182 **3 Results**

Figure 1 shows the occurrence of precipitation for the wet day precipitation, 50th, 90th and 98th percentile masked composites in the cyclone centered coordinate system. Here precipitation occurrence is defined within each masked subset of composites as the percentage of time precipitation is identified. To highlight the changes in the structure, each set of composites are shown on different colour scales. Because of the rotation applied to the cyclone composites, the top of the composites may not align with north, so cardinal directions are not used to describe cyclone features.

Figure 1a-c shows the precipitation occurrence for the wet day precipitation dur-190 ing the deepening phase, with rates up to 100% relative to the cyclone. Occurrence then 191 decreases slightly throughout the cyclone lifecycle. Given the potential for significant di-192 abatic heating/latent heat release during the deepening phase (Wernli et al., 2002; Lud-193 wig et al., 2014; Binder et al., 2016; Messmer & Simmonds, 2021) this pattern is expected. 194 The spatial structure shows high precipitation occurrence about the cyclone centre, that 195 extends in a tail towards the left side of the composite. This tail rotates clockwise through-196 out the lifecycle, likely highlighting the warm seclusion identified in idealised models. 197

Looking at the 50th percentile masked composites, Figure 1d-f similarly shows high-198 est precipitation occurrence during the deepening phase, of up to 80%, before decreas-199 ing in later lifecycle stages. The spatial region associated with high occurrences shows 200 a reduced extent from that seen in 1a-c, with a more pronounced comma structure in 201 the upper left quadrant of the cyclone. This comma structure in precipitation has long 202 been identified in conceptual models of cyclones (see Semple, 2003) and is likely related 203 to the warm conveyor belt. This comma rotates clockwise as the cyclone evolves and the 204 drier poleward area of the cyclone moves equatorward towards the cyclone centre. 205

Comparable patterns are observed when looking at the 90th percentile masked cyclone composites for extreme precipitation occurrence (Figure 1g-i) and the 98th percentile masked composites (Figure 1j-l). Precipitation occurrence is once again greatest during deepening, before decreasing in the peak intensity and decay phases. The structure of the high occurrence comma also rotates clockwise through the cyclone lifecycle. A further reduction in the spatial extent of high occurrence regions from Figure 1d-f is observed, such that the occurrence of precipitating extremes outside the comma is very

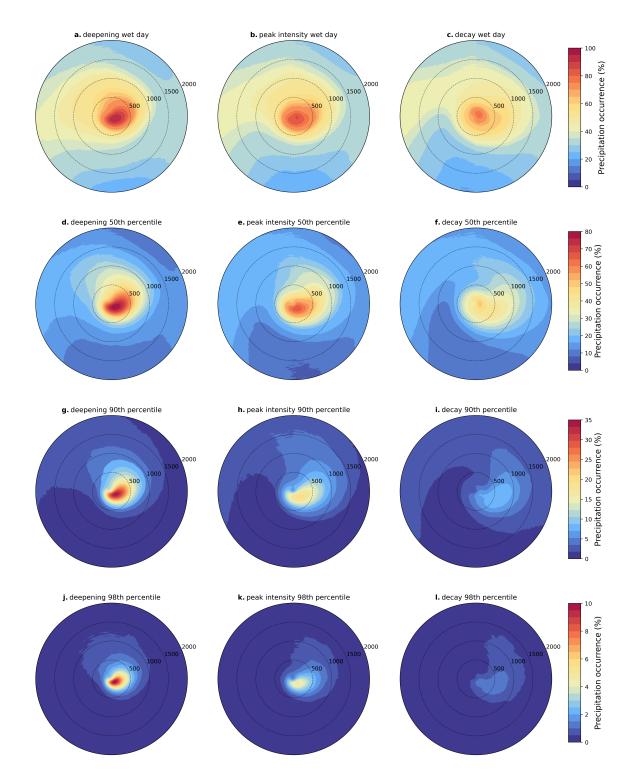


Figure 1. ERA5 cyclone composites of precipitation occurrence between 1980 - 2019 broken into the deepening, peak intensity and decay phases for (a - c) wet day precipitation (d - f) precipitation masked by the 50th percentile value (g - i) precipitation masked by the 90th percentile value (k - l) precipitation masked by the 98th percentile value.

low. For the 90th percentile masked composites precipitation extremes rarely occur within
 the drier poleward region of the cyclone; For the 98th percentile masked composites pre cipitation extremes rarely occur outside of the comma structure.

To examine extreme precipitation from a different perspective, the fraction of the 216 total precipitation fraction associated with each threshold is derived in Figure 2. This 217 fraction is defined as the ratio of precipitation accumulated above a given threshold and 218 the total accumulated precipitation (Supplementary Figure 1). For the wet day percentiles, 219 this accumulation is determined using a 1 mm per day threshold. The nth percentile thresh-220 221 old is used for the 50th, 90th and 98th percentile masked cyclone composites. Figure 2 shows precipitation fraction calculated for each point across the cyclone composite, for 222 the wet day precipitation and 50th, 90th and 98th percentile masked composites. Note 223 the different colour scales on each row of subplots used to distinguish structure. 224

Figure 2a-c shows areas within the cyclone composites where almost 100% of pre-225 cipitation is above the 1 mm wet day threshold. Higher precipitation fraction is concen-226 trated in the comma region of the cyclone composite and in the warm equatorward re-227 gion of the cyclone. Precipitation fraction is greatest during the deepening phase, de-228 creases and rotates clockwise in the peak intensity phase, then weakens during the de-229 cay phase. The area of lowest precipitation fraction occurs in the cold poleward region 230 of the composite, where more of the rainfall is below the 1 mm wet day threshold sug-231 gesting a region dominated by drizzle. This region moves equatorward up the left flank 232 of the cyclone into the upper left quadrant as the cyclone evolves. An almost identical 233 pattern is seen for the 50th percentiles masked composites on Figure 2d-f. Differences 234 include a lesser precipitation fraction, and a reduced spatial extent similar to that seen 235 between Figure 1a-c and d-f. 236

When looking at the precipitation fraction for the 90th percentile masked compos-237 ites, Figure 2g-i shows a decrease compared to the 50th percentile threshold as would 238 be expected. The fraction of the total precipitation linked to events above this thresh-239 old are still highest during the deepening phase, but the greatest precipitation fraction 240 (almost 80%) is lower compared to Figure 2a/d. That up to 80% of precipitation is as-241 sociated with the top 10% of the precipitation distribution is meaningful, and highlights 242 the importance of cyclones for extreme precipitation in general. Regions of high precip-243 itation fraction are more concentrated in the centre of the cyclone compared to Figure 244 2a-f, but with a broad region of 40% to 50% precipitation fraction within the equator-245 ward region of the cyclone highlighting the importance of extreme precipitation in the 246 overall accumulation. A clockwise rotation in regions of highest precipitation fraction 247 is also observed throughout the cyclone evolution, as the precipitation fraction weakens 248 throughout the peak intensity and decay phases. 249

When applying the strictest threshold and masking by the 98th percentile value, Figure 2j-l shows a further decrease in the fraction of the total precipitation associated with events above this threshold. Though, the greatest precipitation fraction is close to 50% and occurs during the deepening phase, before a reduction and clockwise rotation is seen similar to Figure 2g-i. Structurally the large precipitation fraction shows the smallest extent at the 98th percentile, being most concentrated within the comma region of near the cyclone centre.

Results observed for the precipitation occurrence and fraction show similarities in 257 structure between the masked cyclone composites. Figure 3a-b shows the Pearson cor-258 relation coefficients and spatial averages for the precipitation fraction across the cyclone 259 composite as shown on Figures 1 and 2. Correlation is calculated pairwise between each 260 set of cyclone composites, spatially across the composite. Results also show that precip-261 itation is greatest during the deepening phase, and then weakens as the cyclone evolves. 262 However, Figures 1 - 2 show this weakening happens at different rates. To investigate 263 the similarities between the cyclone composites and to quantify this weakening, the ra-264

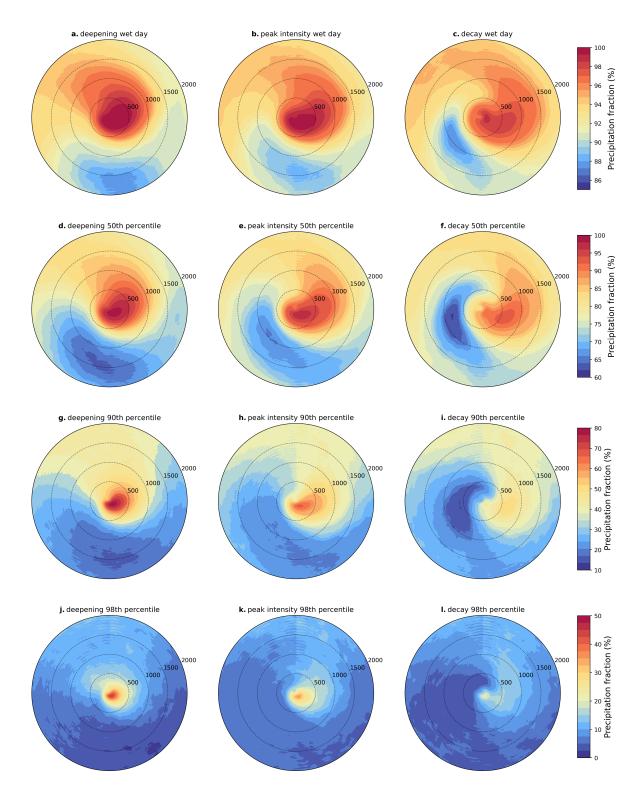


Figure 2. ERA5 cyclone composites of precipitation fraction between 1980 - 2019 broken into the deepening, peak intensity and decay phases for (a - c) wet day precipitation (d - f) precipitation masked by the 50th percentile value (g - i) precipitation masked by the 90th percentile value (k - l) precipitation masked by the 98th percentile value.

- tio of the averages in the peak intensity and decay phases relative to the deepening phase
- ²⁶⁶ are also shown in Figure 3c-d.

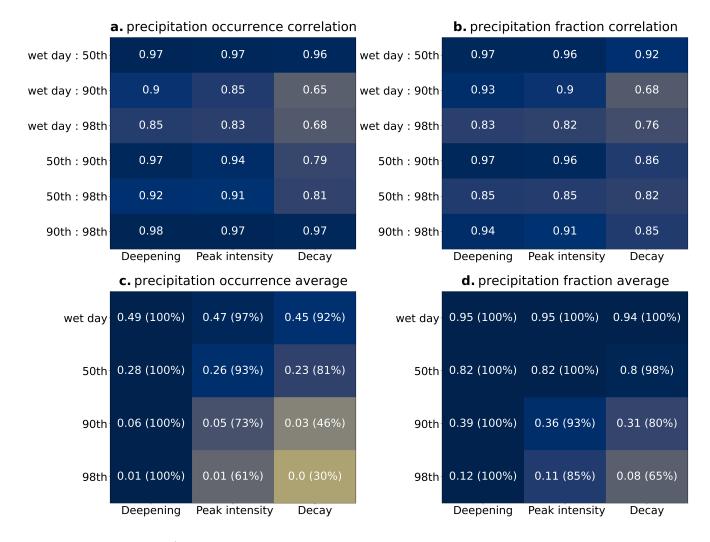


Figure 3. a) Pairwise Pearson correlation coefficients for cyclone composites within the deepening, peak intensity and decay phases for the precipitation occurrence b) same as a) but for precipitation fraction c) Precipitation averages across the cyclone composites within the deepening, peak intensity and decay phases for the precipitation occurrence d) same as c) but for precipitation fraction. The bracketed percentages for (c - d) indicate precipitation averages as a proportion of the corresponding deepening phase value.

Pairwise spatial correlations for the precipitation occurrence (Figure 3a) show strong 267 agreement between each masked composites at the same period in the cyclone lifecycle. 268 Correlation is strongest during the deepening phase in all cases. Correlation is lowest 269 between the wet day precipitation and 90/98th percentile masked composites, but still 270 strong within the deepening phase. When comparing the wet day precipitation/50th per-271 centiles and 90th/98th masked composites, the spatial correlation remains consistently 272 high across the cyclone lifecycle. When looking at the precipitation fraction, the pair-273 wise correlation shows a similar trend to that observed for the precipitation occurrence. 274 There is a strong correlation during the deepening phase and weaker correlation during 275 the decay phase. 276

Looking at the average precipitation occurrence across the cyclone (Figure 3c), not 277 only do the extremes have lower occurrence that the non-extreme precipitation, but drop 278 off significantly faster. The wet day precipitation and 50th percentile masked compos-279 ites decrease to 90% and 81% of the deepening value during the decay phase, respectively. 280 The 90th percentile composites decrease to 46% of the deepening value and the 98th per-281 centile composites to 30% of the deepening value during the decay phase. This indicates 282 as you increase the threshold for determining extremes, they will be disproportionately 283 be experienced most often during the deepening phase, as the occurrence weakens greatly 284 by the decay phase. Averages for the precipitation fraction (Figure 3d) display a sim-285 ilar trend. The 90th and 98th percentile masked composites decrease to 80% and 65%286 of the deepening value, respectively, while the wet day precipitation and 50th percentile 287 masked composites decrease negligibly. This highlights that as you increase the thresh-288 old for the extremes the deepening phase becomes more important, as a lower propor-289 tion of the total precipitation experienced in future cyclone phases can be defined as ex-290 treme. 291

²⁹² 4 Discussion and Conclusion

Using a spatially dependent threshold for precipitation, we have developed a sim-293 ple methodology to assess the contribution to accumulation and the occurrence of pre-294 cipitation extremes by masking data within a cyclone compositing framework. This method-295 ology is based on grouping regions of similar precipitation frequency together, which have 296 been shown to be influenced by the same underlying dynamic and thermodynamic pre-297 cipitation processes (McErlich, McDonald, Schuddeboom, et al., 2023). While some past 298 studies investigate precipitation extremes and how they connect to cyclones (e.g. Pfahl 299 & Wernli, 2012; Catto & Dowdy, 2021; Messmer & Simmonds, 2021), this is the first in-300 vestigation from a cyclone centered perspective over the Southern Hemisphere, instead 301 of focusing on assessing the spatial distributions of cyclones. 302

Here we see that the greatest precipitation occurrence and fraction of total precip-303 itation occur before the cyclone reaches its peak intensity. The precipitation accumu-304 lation (Supplementary Figure 1) is also greatest during the deepening phase, meaning 305 precipitation extremes will be experienced most acutely during the deepening phase. This 306 is not a new result, as Booth et al. (2018) and McErlich, McDonald, Renwick, and Schud-307 deboom (2023) also show that greater precipitation occurs before the cyclone reaches peak 308 intensity. These studies also show a weakening in the comma structure of precipitation 309 through the cyclone lifecycle. However, we show that extreme precipitation also displays 310 a similar trend. For both the 90th and 98th masked percentile composites, the precip-311 itation occurrence and fraction is greatest during the deepening phase and then decays 312 slightly as the cyclone evolves. Many studies have shown support for the concept that 313 the release of latent heating associated with precipitation leads to the intensification of 314 a cyclone (Wernli et al., 2002; Ludwig et al., 2014; Binder et al., 2016; Messmer & Sim-315 monds, 2021), which is consistent with our findings. 316

Figure 3c-d shows that precipitation occurrence and precipitation fraction weakens at a faster rate as you increase from the 90th to 98th percentile of precipitation. This suggests a larger diabatic heating and subsequent intensification of the cyclone from extreme precipitation events. Figure 3c-d also shows that the precipitation occurrence shows a larger change from the deepening to the decay phase than the precipitation fraction at every extreme threshold. This suggests that during the decay phase a larger amount of rainfall is associated with a smaller number of extreme events.

Looking at the spatial pattern of precipitation within cyclone composites, Figures 1 and 2 show that as you apply a stricter threshold to mask precipitation there is a reduction in the spatial extent of both high precipitation occurrence and fraction. As you move towards the extremes, precipitation is more concentrated towards the centre of the cyclone within the comma region linked to the warm conveyor belt.

Figure 3a-b shows that for the precipitation occurrence and fraction, the 90th and 329 98th percentile masked composites show strong spatial correlation across all states of the 330 cyclone lifecycle. The wet day precipitation and 50th and 90th/98th percentile masked 331 cyclone composites shows strong agreement for precipitation occurrence and fraction dur-332 ing the deepening and peak intensity phases, but weaker agreement during the decay phase. 333 These correlations show that the spatial regions of the cyclone where precipitation ex-334 335 tremes are important remain similar across the cyclone lifecycle, suggesting that knowing the median precipitation pattern could provide insight into the upper tail of the dis-336 tributions, at least for the precipitation occurrence and fraction. 337

This work has determined the spatial structure of extreme precipitation relative 338 to cyclone centres, and provided a quantification of how these extremes change as the 339 cyclone evolves. However, the underlying processes that determine precipitation have not 340 been assessed to provide a physical justification for results seen in this work. McErlich, 341 McDonald, Schuddeboom, et al. (2023) has shown that vertical velocity and convective 342 available potential energy are important drivers of global precipitation and that precip-343 itation is determined, in part, by the occurrence of these precipitation-generating pro-344 cesses. Future work will examine these processes to determine how they influence the 345 behaviour of precipitation and precipitation extremes in cyclones through their lifecy-346 cle. 347

348 Acknowledgments

The ERA5 reanalysis products were obtained from the Copernicus Climate Data Store (https://cds.climate.copernicus.eu/). Data used to visualise the figures is available at https://zenodo.org/record/7787119. We would also like to acknowledge funding from the Ministry of Business, Innovation and Employment (MBIE), New Zealand, through the Whakahura project (grant number E7141).

354 **References**

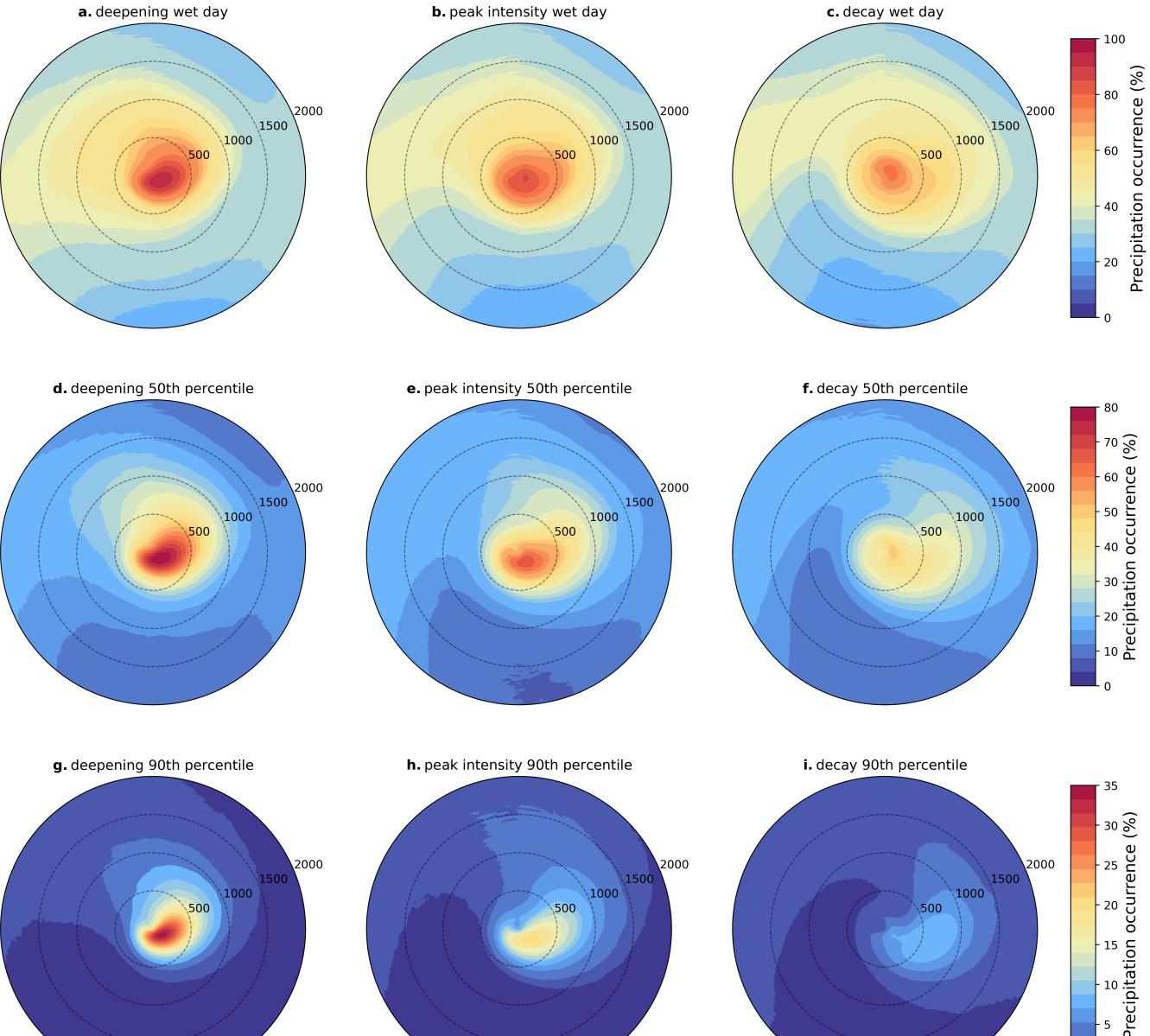
- Bell, J. E., Brown, C. L., Conlon, K., Herring, S., Kunkel, K. E., Lawrimore, J., ...
 Uejio, C. (2018). Changes in extreme events and the potential impacts on human health. Journal of the Air & Waste Management Association, 68(4), 265-287. doi: 10.1080/10962247.2017.1401017
- Bengtsson, L., Hodges, K. I., & Roeckner, E. (2006). Storm tracks and climate change. Journal of Climate, 19(15), 3518 3543. Retrieved from https://journals.ametsoc.org/view/journals/clim/19/15/jcli3815.1.xml doi: 10.1175/JCLI3815.1
- Binder, H., Boettcher, M., Joos, H., & Wernli, H. (2016). The role of warm conveyor belts for the intensification of extratropical cyclones in northern hemi sphere winter. Journal of the Atmospheric Sciences, 73(10), 3997 4020. doi: 10.1175/JAS-D-15-0302.1
- Booth, J. F., Naud, C. M., & Jeyaratnam, J. (2018). Extratropical cyclone precipitation life cycles: A satellite-based analysis. Geophysical Research Letters, 45(16), 8647-8654. Retrieved from https://agupubs.onlinelibrary
 .wiley.com/doi/abs/10.1029/2018GL078977 doi: https://doi.org/10.1029/
 2018GL078977
- 372Catto, J. L., & Dowdy, A.(2021).Understanding compound hazards from a373weather system perspective.Weather and Climate Extremes, 32, 100313.374Retrieved from https://www.sciencedirect.com/science/article/pii/375S2212094721000116doi: https://doi.org/10.1016/j.wace.2021.100313
- ³⁷⁶ Catto, J. L., Jakob, C., Berry, G., & Nicholls, N. (2012). Relating global precipita-

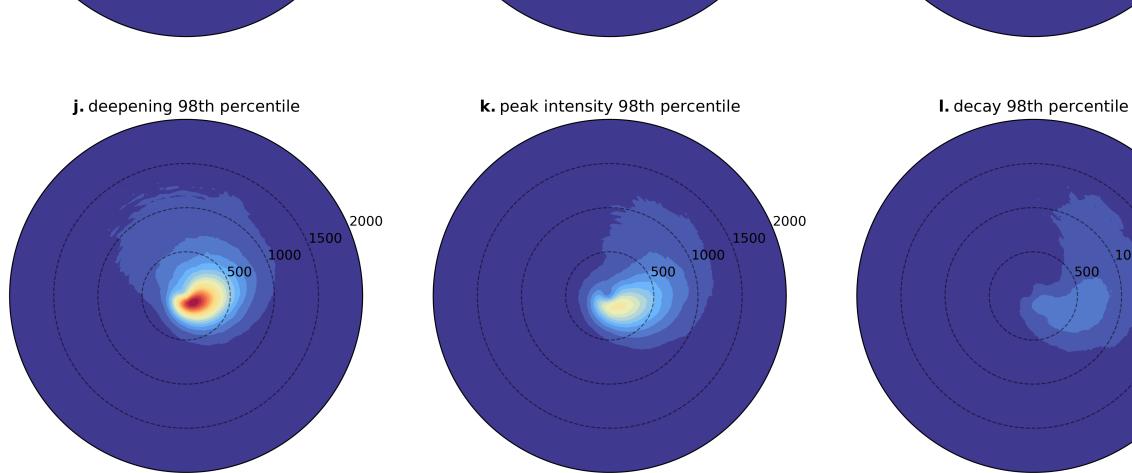
377	tion to atmospheric fronts. Geophysical Research Letters, $39(10)$. doi: https://
378	doi.org/10.1029/2012GL051736
379	Crawford, A. D., & Serreze, M. C. (2016). Does the summer arctic frontal zone in-
380	fluence arctic ocean cyclone activity? Journal of Climate, 29(13), 4977 - 4993.
381	doi: 10.1175/JCLI-D-15-0755.1
382	Field, P. R., Gettelman, A., Neale, R. B., Wood, R., Rasch, P. J., & Morrison, H.
383	(2008). Midlatitude cyclone compositing to constrain climate model behavior
384	using satellite observations. Journal of Climate, 21(22), 5887 - 5903. doi:
385	10.1175/2008JCLI2235.1
386	Field, P. R., & Wood, R. (2007). Precipitation and cloud structure in midlatitude
387	cyclones. Journal of Climate, 20(2), 233 - 254. doi: 10.1175/JCLI3998.1
388	Govekar, P. D., Jakob, C., Reeder, M. J., & Haynes, J. (2011). The three-
389	dimensional distribution of clouds around southern hemisphere extratropical
390	cyclones. Geophysical Research Letters, 38(21). doi: https://doi.org/10.1029/
	2011GL049091
391	Hawcroft, M., Shaffrey, L., Hodges, K., & Dacre, H. (2012, 12). How much northern
392	
393	hemisphere precipitation is associated with extratropical cyclones? <i>Geophysical</i>
394	Research Letters, 39 , 24809 doi: 10.1029/2012GL053866
395	Hirsch, R. M., & Archfield, S. A. (2015, Mar 01). Not higher but more often. <i>Nature</i>
396	Climate Change, 5(3), 198-199. doi: 10.1038/nclimate2551
397	Hodges, K. I., Lee, R. W., & Bengtsson, L. (2011). A comparison of extratrop-
398	ical cyclones in recent reanalyses era-interim, nasa merra, ncep cfsr, and
399	jra-25. Journal of Climate, 24(18), 4888 - 4906. Retrieved from https://
400	journals.ametsoc.org/view/journals/clim/24/18/2011jcli4097.1.xml
401	doi: 10.1175/2011JCLI4097.1
402	Hoskins, B. J., & Hodges, K. I. (2005). A new perspective on southern hemisphere
403	storm tracks. Journal of Climate, 18(20), 4108 - 4129. doi: 10.1175/JCLI3570
404	.1
404	Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Map-
404 405	Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Map- ping the shadow of experience of extreme weather events. <i>Climatic Change</i> ,
404 405 406	Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i> , 127(2), 381-389. doi: 10.1007/s10584-014-1253-6
404 405 406 407 408	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional
404 405 406 407 408 409	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi
404 405 406 407 408 409 410	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005
404 405 406 407 408 409 410 411	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will
404 405 406 407 408 409 410 411 412	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal pre-
404 405 406 407 408 409 410 411 412 413	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1),
404 405 406 407 408 409 410 411 412 413 414	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi .org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w
404 405 406 407 408 409 410 411 412 413 414	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes
404 405 406 407 408 409 410 411 412 413 414 415 416	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021
404 405 406 407 408 409 410 411 412 413 414	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8
404 405 406 407 408 409 410 411 412 413 414 415 416	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi .org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous
404 405 406 407 408 409 410 411 412 413 414 415 416 417	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi .org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive
404 405 406 407 408 409 410 411 412 413 414 415 416 417 418	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi .org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. <i>Quarterly Journal of the Royal Meteorolog-</i>
404 405 406 407 408 409 410 411 412 413 414 415 416 417 418	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi .org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. <i>Quarterly Journal of the Royal Meteorological Society</i>, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253
404 405 406 407 408 409 410 411 412 413 414 415 416 417 418 419 420	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. <i>Quarterly Journal of the Royal Meteorological Society</i>, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanaly-
404 405 406 407 408 409 410 411 412 413 414 415 416 417 418 419 420 421	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. Climatic Change, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. Weather and Climate Extremes, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. Nature Communications, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. Nature, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. Quarterly Journal of the Royal Meteorological Society, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice
404 405 406 407 408 409 410 411 412 413 414 415 416 417 418 419 420 421	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. Climatic Change, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. Weather and Climate Extremes, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. Nature Communications, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. Nature, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. Quarterly Journal of the Royal Meteorological Society, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice shelf region. Earth and Space Science, 7(1), e2019EA000794. Retrieved
404 405 406 407 408 410 411 412 413 414 415 416 417 418 419 420 421 422	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. <i>Quarterly Journal of the Royal Meteorological Society</i>, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice shelf region. <i>Earth and Space Science</i>, 7(1), e2019EA000794. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/
404 405 406 407 408 410 411 412 413 414 415 416 417 418 419 420 421 422 423 424	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. Climatic Change, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. Weather and Climate Extremes, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. Nature Communications, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. Nature, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. Quarterly Journal of the Royal Meteorological Society, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice shelf region. Earth and Space Science, 7(1), e2019EA000794. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2019EA000794
404 405 406 407 408 409 410 411 412 413 414 415 416 417 418 419 420 421 422 423 424	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. <i>Climatic Change</i>, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. <i>Weather and Climate Extremes</i>, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. <i>Nature Communications</i>, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. <i>Nature</i>, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. <i>Quarterly Journal of the Royal Meteorological Society</i>, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice shelf region. <i>Earth and Space Science</i>, 7(1), e2019EA000794. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/
404 405 406 407 408 410 411 412 413 414 415 416 417 418 419 420 421 422 423 424 425 426	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. Climatic Change, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. Weather and Climate Extremes, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. Nature Communications, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. Nature, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. Quarterly Journal of the Royal Meteorological Society, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice shelf region. Earth and Space Science, 7(1), e2019EA000794. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2019EA000794
404 405 406 407 408 410 411 412 413 414 415 416 417 418 419 420 421 422 423 424 425 426	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. Climatic Change, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. Weather and Climate Extremes, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. Nature Communications, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. Nature, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. Quarterly Journal of the Royal Meteorological Society, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice shelf region. Earth and Space Science, 7(1), e2019EA000794. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2019EA000794 McErlich, C., McDonald, A., Renwick, J., & Schuddeboom, A. (2023). An assess-
404 405 406 407 408 410 411 412 413 414 415 416 417 418 419 420 421 422 423 424 425 426 427 428	 Howe, P. D., Boudet, H., Leiserowitz, A., & Maibach, E. W. (2014, Nov 01). Mapping the shadow of experience of extreme weather events. Climatic Change, 127(2), 381-389. doi: 10.1007/s10584-014-1253-6 Jahn, M. (2015). Economics of extreme weather events: Terminology and regional impact models. Weather and Climate Extremes, 10, 29-39. doi: https://doi.org/10.1016/j.wace.2015.08.005 Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020). Climate change will affect global water availability through compounding changes in seasonal precipitation and evaporation [Journal Article]. Nature Communications, 11(1), 3044. doi: 10.1038/s41467-020-16757-w Kotz, M., Levermann, A., & Wenz, L. (2022, Jan 01). The effect of rainfall changes on economic production. Nature, 601(7892), 223-227. doi: 10.1038/s41586-021 -04283-8 Ludwig, P., Pinto, J. G., Reyers, M., & Gray, S. L. (2014). The role of anomalous sst and surface fluxes over the southeastern north atlantic in the explosive development of windstorm xynthia. Quarterly Journal of the Royal Meteorological Society, 140(682), 1729-1741. doi: https://doi.org/10.1002/qj.2253 McDonald, A. J., & Cairns, L. H. (2020). A new method to evaluate reanalyses using synoptic patterns: An example application in the ross sea/ross ice shelf region. Earth and Space Science, 7(1), e2019EA000794. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2019EA000794 McErlich, C., McDonald, A., Renwick, J., & Schuddeboom, A. (2023). An assessment of southern hemisphere extra-tropical cyclones in era5 using windsat. Au-

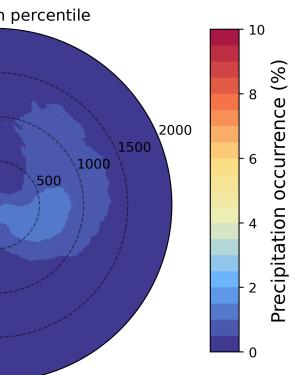
432	linked to universal precipitation drivers. Nature Geosience.
433	Messmer, M., & Simmonds, I. (2021). Global analysis of cyclone-induced com-
434	pound precipitation and wind extreme events. Weather and Climate Extremes,
435	32, 100324. doi: https://doi.org/10.1016/j.wace.2021.100324
436	Min, SK., Zhang, X., Zwiers, F. W., & Hegerl, G. C. (2011, Feb 01). Human con-
437	tribution to more-intense precipitation extremes. Nature, 470(7334), 378-381.
438	doi: 10.1038/nature09763
439	Morss, R. E., Wilhelmi, O. V., Meehl, G. A., & Dilling, L. (2011). Improving
440	societal outcomes of extreme weather in a changing climate: an integrated
441	perspective. Annual Review of Environment and Resources, 36, 1–25.
442	Naud, C. M., Posselt, D. J., & van den Heever, S. C. (2012). Observational analysis
443	of cloud and precipitation in midlatitude cyclones: Northern versus south-
444	ern hemisphere warm fronts. Journal of Climate, 25(14), 5135 - 5151. doi:
445	10.1175/JCLI-D-11-00569.1
446	Papritz, L., Pfahl, S., Rudeva, I., Simmonds, I., Sodemann, H., & Wernli, H.
447	(2014). The role of extratropical cyclones and fronts for southern ocean
448	freshwater fluxes. Journal of Climate, 27(16), 6205 - 6224. doi: 10.1175/
449	JCLI-D-13-00409.1
450	Pendergrass, A. G., & Knutti, R. (2018). The uneven nature of daily precipita-
451	tion and its change. Geophysical Research Letters, 45(21), 11,980-11,988.
452	Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/
453	10.1029/2018GL080298 doi: https://doi.org/10.1029/2018GL080298
454	Pfahl, S., & Wernli, H. (2012). Quantifying the relevance of cyclones for precipita-
455	tion extremes. Journal of Climate, 25(19), 6770 - 6780. doi: 10.1175/JCLI-D
456	-11-00705.1
457	Polade, S. D., Pierce, D. W., Cayan, D. R., Gershunov, A., & Dettinger, M. D.
458	(2014, Mar 13). The key role of dry days in changing regional climate and
459	precipitation regimes. Scientific Reports, 4(1), 4364. doi: 10.1038/srep04364
460	Schär, C., Ban, N., Fischer, E. M., Rajczak, J., Schmidli, J., Frei, C., Zwiers,
461	F. W. (2016, Jul 01). Percentile indices for assessing changes in heavy
462	precipitation events. Climatic Change, 137(1), 201-216. doi: 10.1007/
463	s10584-016-1669-2
464	Seddon, A. W. R., Macias-Fauria, M., Long, P. R., Benz, D., & Willis, K. J. (2016,
465	Mar 01). Sensitivity of global terrestrial ecosystems to climate variability. Na-
466	ture, 531(7593), 229-232. doi: 10.1038/nature16986
467	Semple, A. (2003). A review and unification of conceptual models of cyclogenesis.
468	Meteorological Applications, 10, 39-59.
469	Seneviratne, S., Nicholls, N., Easterling, D., Goodess, C., S. Kanae, J. K., Luo, Y.,
470	Zhang, X. (2012). Changes in climate extremes and their impacts on the
471	natural physical environment (C. Field et al., Eds.). Cambridge, UK, and New
472	York, NY, USA: Cambridge University Press.
473	Seneviratne, S., Zhang, X., Adnan, M., Badi, W., Dereczynsk, C., Luca, A. D.,
474	Zhou, B. (2021). Weather and Climate Extreme Events in a Changing Climate
475	(V. Masson-Delmotte et al., Eds.). Cambridge, UK, and New York, NY, USA:
476	Cambridge University Press.
477	Utsumi, N., Kim, H., Kanae, S., & Oki, T. (2017). Relative contributions of weather
478	systems to mean and extreme global precipitation. Journal of Geophysi-
479	cal Research: Atmospheres, 122(1), 152-167. doi: https://doi.org/10.1002/
480	2016JD025222
481	Wernli, H., Dirren, S., Liniger, M. A., & Zillig, M. (2002). Dynamical aspects
482	of the life cycle of the winter storm 'lothar' (24–26 december 1999). Quar-
483	terly Journal of the Royal Meteorological Society, 128(580), 405-429. doi:
484	https://doi.org/10.1256/003590002321042036
485	Zhang, X., Alexander, L., Hegerl, G. C., Jones, P., Tank, A. K., Peterson, T. C.,
486	Zwiers, F. W. (2011). Indices for monitoring changes in extremes based

- on daily temperature and precipitation data. WIREs Climate Change, 2(6),
 851-870. doi: 10.1002/wcc.147
- 489 Zhang, X., Zwiers, F. W., Hegerl, G. C., Lambert, F. H., Gillett, N. P., Solomon,
- S., ... Nozawa, T. (2007, Jul 01). Detection of human influence on
 twentieth-century precipitation trends. Nature, 448(7152), 461-465. doi:
 10.1038/nature06025
- ⁴⁹³ Zscheischler, J., Westra, S., van den Hurk, B. J. J. M., Seneviratne, S. I., Ward,
- P. J., Pitman, A., ... Zhang, X. (2018, Jun 01). Future climate risk
 from compound events. Nature Climate Change, 8(6), 469-477. doi:
 10.1038/s41558-018-0156-3

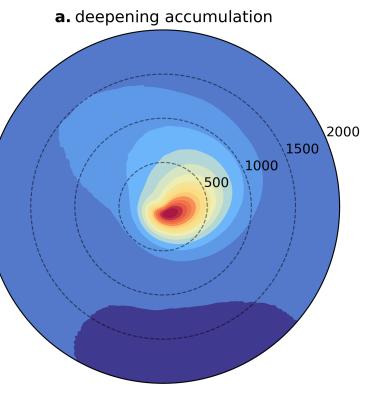
precipitation_occurrence.png.

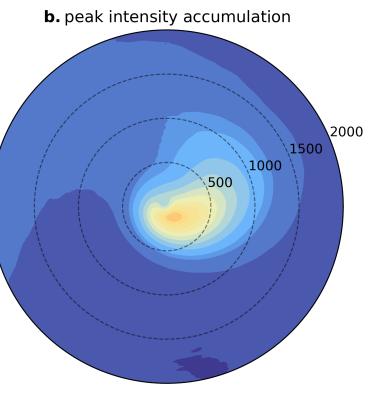


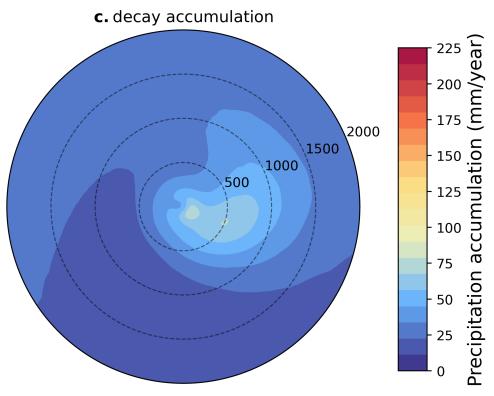




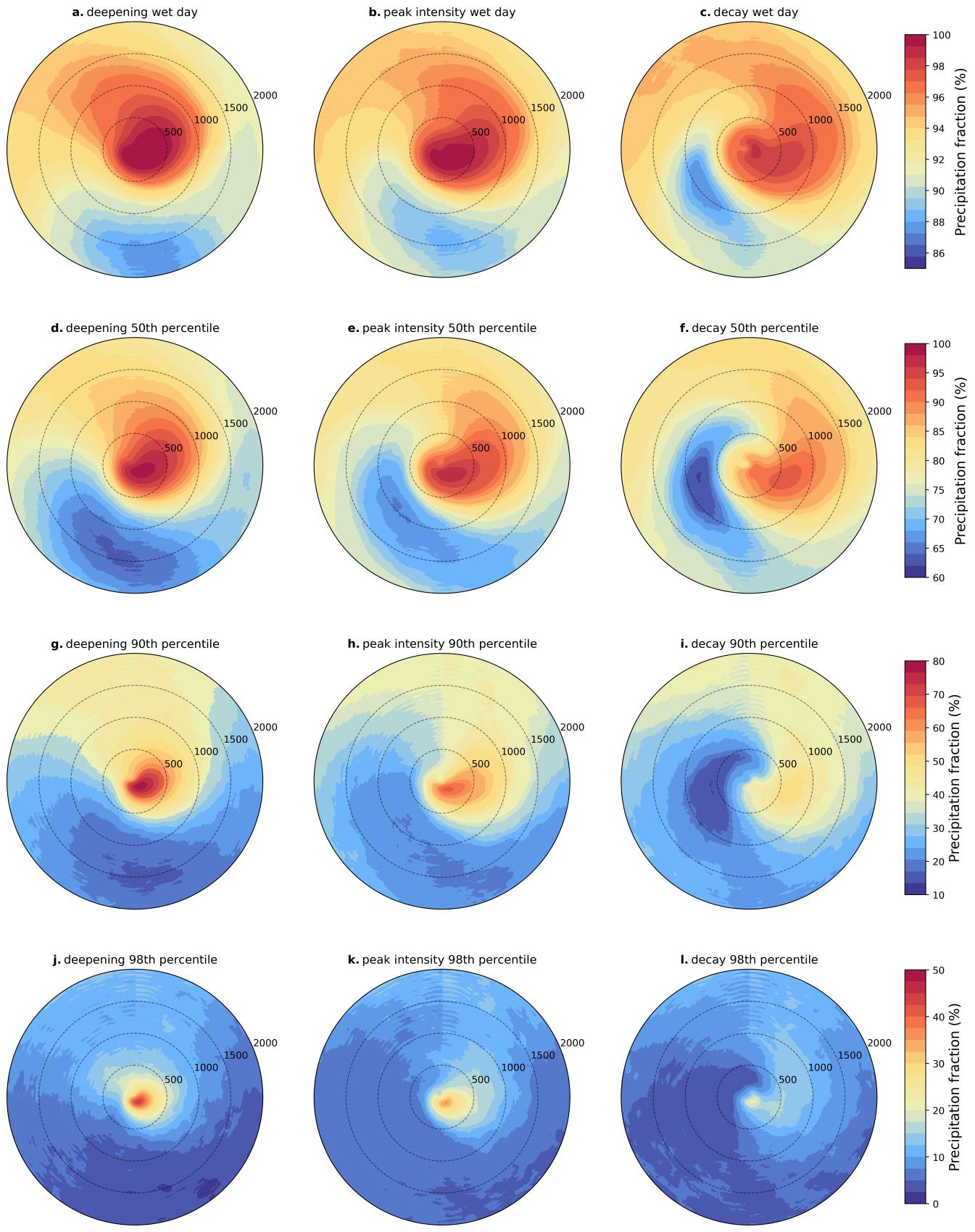
precipitation_accumulation.png.







precipitation_fraction.png.



combined_matrix.png.

a. precipitation occurrence correlation

	I I				I I		
wet day : 50th	0.97	0.97	0.96	wet day : 50th	0.97	0.96	0.92
wet day : 90th	0.9	0.85	0.65	wet day : 90th	0.93	0.9	0.68
wet day : 98th	0.85	0.83	0.68	wet day : 98th	0.83	0.82	0.76
50th : 90th	0.97	0.94	0.79	50th : 90th	0.97	0.96	0.86
50th : 98th	0.92	0.91	0.81	50th : 98th	0.85	0.85	0.82
90th : 98th	0.98	0.97	0.97	90th : 98th	0.94	0.91	0.85
	Deepening	Peak intensity	Decay		Deepening	Peak intensity	Decay
	c. precipita	tion occurren	ce average		d. precipi	tation fractio	n average
wet day	0.49 (100%)	0.47 (97%)	0.45 (92%)	wet day	0.95 (100%)	0.95 (100%)	0.94 (100%)
50th	0.28 (100%)	0.26 (93%)	0.23 (81%)	50th	0.82 (100%)	0.82 (100%)	0.8 (98%)
90th	0.06 (100%)	0.05 (73%)	0.03 (46%)	90th	0.39 (100%)	0.36 (93%)	0.31 (80%)
98th	0.01 (100%)	0.01 (61%)	0.0 (30%)	98th	0.12 (100%)	0.11 (85%)	0.08 (65%)
	Deepening	Peak intensity	Decay		Deepening	Peak intensity	Decay

b. precipitation fraction correlation



