The value of large-scale climatic indices for monthly forecasting severity of widespread flooding using dilated convolutional neural networks

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¹³ Key Points:

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14	• An adapted Wavenet model conditioned by climatic indices is applied to forecast
15	the monthly severity of widespread flooding in Germany
16	• Predictability of rain-on-snow floods is higher than of rainfall-generated floods show
17	ing the potential for forecasting severe winter floods
18	• Feature attribution reveals variable importance of large-scale climatic indices and
19	antecedent wetness conditions across German regions

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20 Abstract

Spatially co-occurring floods pose a threat to the resilience and recovery of the commu-21 nities. Their timely forecasting plays a crucial role for increasing flood preparedness and 22 limiting associated losses. In this study we investigated the potential of a dilated Con-23 volutional Neural Network (dCNN) model conditioned on large-scale climatic indices and 24 antecedent precipitation to forecast monthly severity of widespread flooding (i.e., spa-25 tially co-occurring floods) in Germany with one month lead time. The severity was es-26 timated from 63 years of daily streamflow series as the sum of concurrent exceedances 27 of at-site two-year return periods within a given month across 172 mesoscale catchments 28 (median area 516 $\rm km^2$). The model was trained individually for the whole country and 29 three diverse hydroclimatic regions to provide insights on heterogeneity of model per-30 formance and flood drivers. Our results showed a considerable potential for forecasting 31 widespread flood severity using dCNN especially as the length of training series increases. 32 However, event-based evaluation of model skill indicates large underestimation for rainfall-33 generated floods during dry conditions despite overall lower severity of these events com-34 pared to the rain-on-snow floods. Feature attribution and wavelet coherence analyses both 35 indicated considerable difference in the major flood drivers in three regions. While the 36 flooding in North-Eastern region is strongly affected by the Baltic Sea, the North-Western 37 region is affected more by global patterns associated with the El-Niño activity. In the 38 Southern region in addition to global patterns we detected the effect of the Mediterranean 39 Sea, while antecedent precipitation plays a less important role in this region. 40

⁴¹ Plain Language Summary

Floods that occur simultaneously at different locations in space are associated with 42 higher damages and are more difficult for communities to cope. Forecasting of such events 43 still remains a challenge. Here we investigate the potential of recently developed deep 44 learning networks for forecasting widespread flooding one month ahead. We test this model 45 on a large dataset of long streamflow time series in Germany and its three geographi-46 cal regions using climatic indices and past precipitation as supporting variables. We show 47 that the model has a considerable potential for flood forecasting, especially for rain-on-48 snow floods, but its performance is limited for rain-on-dry events. Finally, using explain-49 able machine learning we provide insights on contrasting differences in climatic flood drivers 50 in different German regions. 51

52 1 Introduction

River floods are one of the most devastating natural hazards in the world, annu-53 ally causing considerable fatalities and socio-economical losses that are expected to in-54 crease with advancing climate change (B. Merz et al., 2021). Their spatial co-occurrence, 55 i.e., when river flooding occurs simultaniously at multiple locations, is an increasingly 56 recognized threat for efficient flood adaptation (Brunner et al., 2020; Zscheischler et al., 57 2020). It limits preparedness and the resilience of the communities (Kreibich et al., 2017). 58 and poses a major challenge for the insurance and re-insurance industries (Zanardo et 59 al., 2019). Spatial co-occurrence of floods in Europe and globally is often associated with 60 the activity of large scale climatic patterns and teleconnections, such as North Atlantic 61 Oscillation (Zanardo et al., 2019) and El Niño Southern Oscillation (Ward et al., 2014). 62 At the same time changing spatial extent of floods (i.e., number of simultaneously af-63 fected river gauges) in Europe (Berghuijs et al., 2019; Kemter et al., 2020) is strongly 64 related to changes in dominant flood generation processes. 65

In Germany, as in the USA (Brunner et al., 2020), spatially co-occurring floods (i.e.,
 widespread flooding) occur in winter season (Uhlemann et al., 2010) and most frequently
 caused by the combination of intensive snowmelt and rainfall that are modulated by the
 dynamics of North Atlantic Oscillation and Scandinavian Pattern (Krug et al., 2020).

However, around 36% of all trans-basin floods in Germany in the second half of the 20th 70 century occurred in summer (Uhlemann et al., 2010). The two recent very severe sum-71 mer widespread floods in 2002 and 2013 were associated with the north-east propagat-72 ing cyclone originating from the Mediterranean Sea (Blöschl et al., 2013). Moreover, there 73 is evidence that spatial co-occurrence of flood-rich periods in Germany is mostly caused 74 by elevated antecedent wetness conditions (B. Merz et al., 2018) indicating that addi-75 tional information on land surface conditions or on prior rainfall that could be related 76 to the same atmospheric circulation mechanism might be useful for predicting widespread 77 flooding (Nakamura et al., 2013; Nied et al., 2017). 78

Timely forecasting of catastrophic widespread flooding is very important for suc-79 cessful flood adaptation in changing climate. Currently global climate models are able 80 to issue medium-range and seasonal forecasts (i.e., from several weeks to several months) 81 of the large-scale climatic patterns and teleconnections, such as North Atlantic Oscil-82 lation and El Niño Southern Oscillation, with relatively high fiedelity (Ludescher et al., 83 2014; Feng et al., 2021), compared to the limited forecasting skill of precipitation and 84 floods (Slater et al., 2019). Therefore, large-scale climatic indices might become useful 85 for improving accuracy and increasing lead time for forecasting of widespread flooding 86 events using data-driven models. 87

Climatic indices are based on the large scale (i.e., regional to global) geophysical 88 states and atmospheric circulations that are often used to characterize and investigate 89 the teleconnections between large scale circulation patterns and local scale phenomena 90 (e.g., precipitation, floods). In this study we investigate the value of ten standard cli-91 matic indices for forecasting the severity of widespread flooding in Germany. Since there 92 93 are clear indications that the emergence of floods (R. Merz et al., 2020) and particularly the occurrence of spatially coherent floods (Nied et al., 2017; B. Merz et al., 2018) are 94 related to antecedent wetness conditions we also investigate the role of antecedent pre-95 cipitation as a proxy of wetness in the forecasting of severity of widespread flooding. 96

In recent years deep learning neural networks have emerged as a promising tool for 97 time series forecasting, especially for medium range climate and weather forecast (Rasp 98 & Thuerey, 2020; Schultz et al., 2021). Currently, Recurrent Neural Networks (RNN), 99 particularly the Long Short-Term Memory (LSTM) models are recognized as effective 100 neural networks for time series forecasting (Hsu, 2017). Their recurrent connections that 101 allow the network to use the entire history of time series and to capture recurrent pat-102 terns or dynamic structures at different time scales (Hochreiter & Schmidhuber, 1997). 103 Recent work shows that a more parsimonious and interpretable representations of such 104 dynamics can be obtained by using multiple dilated convolutional layers (Yu & Lin, 2015) 105 Comparison of the dilated Convolutional Networks (dCNN) to the stat-of-the-art RNNs 106 (particularly to LSTM) using synthetic and real world examples showed comparable and 107 even superior performance of dCNN models for conditional time series forecasting in cases 108 when availability of long series is limited (Borovykh et al., 2017; Y. Chen et al., 2019; 109 Benhaddi & Ouarzazi, 2021). Moreover, dCNN models are much easier to train, faster 110 to converge and require less memory compared to the state-of-the-art RNN models (Benhaddi 111 & Ouarzazi, 2021) since the convolutional structure of the network reduces the number 112 of trainable parameters (Borovykh et al., 2017). 113

In this study, we apply the convolutional model Wavenet (van den Oord et al., 2016) 114 originally developed for audio forecasting and adapted by Borovykh et al. (2017) for con-115 ditional forecasting of the severity of widespread flooding. The model is based on the 116 dilated convolutions applied to both the input time series (i.e., index of widespread flood 117 severity) and to the covariates (e.g., large-scale climatic indices, antecedent wetness) al-118 lowing the model not only learn the inherent patterns from the input time series, but 119 also to learn its dependence on the covariates at different time scales. Previously the Wavenet 120 model was successfully applied for conditional forecasting of financial time series (Borovykh 121 et al., 2017), online retail time series (Y. Chen et al., 2019) and air quality time series 122

(Benhaddi & Ouarzazi, 2021). However, to the best of our knowledge adapted Wavenet
was not previously used for the conditional forecasting of any hydrological time series.
Moreover, the utility of large-scale climatic patterns and antecedent wetness for forecasting of widespread flooding in Germany, as well as the relative predictability of floods in
different hydroclimatic regions and of floods generated by different processes (e.g., rainon-snow vs rainfall) is still unclear. Given this background, this study has the following objectives:

- 130 1. test adapted Wavenet for forecasting of widespread floods in Germany
 - 2. examine variability of its performance across hydroclimatologically distinct regions in Germany
 - 3. examine variability of its performance for forecasting floods of different generation processes
 - 4. compare the importance of climatic indices and antecedent wetness conditions obtained using a model-based feature attribution method (explainable machine learning techniques) with the results of the model-independent wavelet coherence analysis
- 139 **2 Data**

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2.1 Severity index of widespread flooding and flood generation processes

To quantify the severity of widespread flooding across multiple gauges in Germany, 141 we construct a severity index (Figure 1a) in line with the index used by Nied et al. (2014) 142 First, we identify all streamflow events regardless of their peak magnitude from contin-143 uous daily time series of streamflow recorded at each gauge station using the method of 144 Tarasova et al. (2018). From these, we select events whose peak magnitude exceeds the 145 threshold of a local (i.e., at-site) 2-year return period (often associated with the full river 146 bank) defined using annual maxima observed for the period from 1951 to 2013 (the records 147 have minimum length 37 years, median length 61 years, maximum length 63 years). Fi-148 nally, we sum up the identified exceedences (i.e., ratio of peak magnitude and the mag-149 nitude of 2-year return period) for all catchments for a given month and normalize it by 150 the number of available gauges to account for a variable length of the observation time 151 series (Figure 1b). Therefore, the derived index comprises the information about two fea-152 tures, (a) the number of gauges that in a given month report an exceedence of the 2-year 153 return period (i.e., flooding), and (b) the magnitude of the exceedence of the at-site 2-154 year return period flow. In case none of the gauges report an exceedence above at-site 155 2-year return period the resulting severity index corresponds to 0 (Figure S1). 156

We compute the index of widespread flood severity for the whole country by com-157 bining the observations in all 172 streamflow stations across the country and we derive 158 regional indices for each of the three hydroclimatic regions in Germany: North Western 159 (NW, 76 catchments), North Eastern (NE, 28 catchments) and Southern (South, 68 catch-160 ments) regions (Figure 1a). The North-Western region (the Lower Rhine, Weser Rivers) 161 is associated with winter flooding and is strongly affected by the Atlantic storms (Conticello 162 et al., 2020). The North-Eastern region (the Elbe and Oder Rivers) is often associated 163 with moderate winter floods with rare but severe summer floods (Petrow et al., 2007; 164 Tarasova, Basso, & Merz, 2020). Finally, the Southern region (the Upper Rhine, Main 165 and Danube Rivers) is associated with a mixed seasonality of floods, while in the right 166 tributaries of the Danube summer floods are even more frequent compared to all other 167 seasons (Beurton & Thieken, 2009). For the analysis we only consider mesoscale catch-168 ments with area range from 31 to $23,700 \text{ km}^2$ (median 516 km²). 169

Additionally, to every value of the monthly severity of widespread flooding we assign a corresponding type of flood generation process that is identified as the most frequent event type among all affected catchments (Figure 1b, color-coding). Flood gen-



Figure 1. a) Study area and three distinct hydroclimatic regions (NW - North-Western, NE - North-Eastern and South - Southern regions); b) Derivation of the schematic index for widespread flooding (HQ2 stands for a local flood magnitude associated with 2-year return period); c) a hierarchical decision tree for classification of event types based on the event classification framework of Tarasova, Basso, Wendi, et al. (2020) (modified from Tarasova et al. (2023)), $M_{x,y,t}$ and $P_{x,y,t}$ stand for catchment- and event-averaged snowmelt and total precipitation respectively, $P_{x,y}$ represents a set of daily catchment-averaged total precipitation during the event, $SM_{x,y}(t_0)$ stands for antecedent soil moisture on the day prior to the event begin, $max(\kappa)$ indicates the value of soil moisture that corresponds to the point of maximum curvature of the function that describes the relation between event runoff coefficients and antecedent soil moisture (see Tarasova et al. (2018) for more details). Final event types are indicated as colored boxes. For a detailed description of indicators, classification thresholds and event types please refer to Tarasova, Basso, Wendi, et al. (2020)

eration processes are identified (see Figure 1c for the assignment process) using the in-173 formation on the proportion of catchment- and event-averaged snowmelt $(M_{x,y,t})$ in the 174 total volume of precipitation $(P_{x,y,t})$, catchment-averaged soil moisture on the day be-175 fore the event has started $(SM_{x,y}(t_0))$, proportion of maximum precipitation intensity 176 during the event $(\max(P_{x,y}(t)))$ and event-averaged total precipitation $(P_{x,y,t})$ and vari-177 ability of daily catchment-averaged precipitation sums during the event (var(Px,y)) (Tarasova, 178 Basso, Wendi, et al., 2020). Daily observed gridded (1 km) precipitation data are ob-179 tained from the REGNIE dataset (Rauthe et al., 2013), daily observed temperature grid-180 ded (4 km) is obtained from Zink et al. (2017). Daily gridded (4 km) soil moisture and 181 snowmelt are simulated by the mesoscale Hydrological Model (mHM) (Samaniego et al., 182 2010; Kumar et al., 2013; Zink et al., 2017). All these datasets are available for the whole 183 study period from 1951 to 2013. Using this information we distinguish five different event 184 types (Figure 1c): Rain.Snow (combination of rainfall and snowmelt when the propor-185 tion of the latter in the total precipitation sum is at least 30%), Rain.Wet.Intensity (intensity-186 dominated (i.e., most of precipitation has occurred in a single time step during event) 187 rainfall-induced event with wet antecedent conditions), Rain.Wet.Volume (volume-dominated 188 rainfall-induced event with wet antecedent conditions), Rain.Dry.Intensity (intensity-dominated 189 rainfall-induced event with dry antecedent conditions), Rain.Dry.Volume (volume-dominated 190 rainfall-induced event with dry antecedent conditions (Tarasova et al., 2023). The thresh-191 old between wet and dry conditions $(\max(\kappa))$ is defined as the value of catchment-averaged 192 soil moisture that corresponds to the point of maximum curvature of the non-linear func-193 tion that capture the increase in event runoff coefficients with the increase of soil mois-194 ture (Tarasova et al., 2018). The sensitivity analysis performed in Tarasova, Basso, Wendi, 195 et al. (2020) indicates that the parametric uncertainty and the choice of the hydrolog-196 ical model have only minor effect on classification results. For more details on the in-197 dicators and thresholds used for streamflow event classification refer to Tarasova, Basso, 198 Wendi, et al. (2020) and for the details on its application for classification of flood events 199 to Tarasova et al. (2023). 200

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2.2 Covariates: large-scale climatic indices and antecedent precipitation

Climatic indices are a low dimensional proxy of the dynamics of the atmosphere 202 and the ocean from seasonal to multi-decadal scales (Domeisen et al., 2018). We selected 203 ten different climatic indices that capture the effects of global climatic variability (e.g., 204 Pacific North American Pattern, Southern Oscillation Index, North Pacific, Antarctic 205 Oscillation), patterns that capture more local effect of the Baltic Sea (Scandinavian Pat-206 tern, East Atlantic) and the effect of blocking conditions (Greenland Blocking Index, North Atlantic Oscillation, Arctic Oscillation) that were previously related to flood occurrence 208 in Europe (e.g., Zanardo et al. (2019)). We also include the Mediterranean Oscillation 209 Index that captures the effect of the Mediterranean Sea (Trigo et al., 2002). We provide 210 a short description of each index that was used in this study. 211

Pacific North American Pattern (PNA) is one of the strongest extratropical teleconnections (Wallace & Gutzler, 1981) and is recognized as a major mode of planetaryscale atmospheric variability evident at all time scales over North America during boreal winter (Leathers et al., 1991). PNA modes are defined by the anomalies in the geopotential height fields over the North Pacific Ocean and North American continent and are strongly influenced by the El Niño-Southern Oscillation (Li et al., 2019).

Southern Oscillation index (SOI) is based on the sea level pressure differences
between Tahiti and Darwin, Australia, and represents air pressure differences between
western and eastern tropical Pacific. Periods of negative index correspond to El Niño
episodes and abnormally warm ocean in the eastern tropical Pacific (Harrison & Larkin,
1998). The opposite is representative of La Niña episodes. These irregular occurring episodes
have a major effect of the Earth's climate system and result in devastating hydrometeorological extremes in different regions of the world (Ward et al., 2016). Although El Niño-

Southern Oscillation teleconnections show only weak effect on European region compared
 to tropical regions and Pacific-North American sector, at seasonal time scales it remains
 one of the strongest predictors of European climate (Brönnimann, 2007).

North Pacific (NP) Oscillation describes the seesaw pattern of sea level pressure over subtropical and mid-latitude North Pacific Ocean (Wallace & Gutzler, 1981).
It impacts temperature and precipitation over Eurasia and North America (S. Chen et al., 2021) and is believed to be a trigger of El Niño-Southern Oscillation events (Park et al., 2013).

Antarctic Oscillation (AAO) also known as Southern Annual Mode strongly
affects the climate in high-latitudes of the the Southern Hemisphere (Thompson And &
Wallace, 2000). However, several studies have indicates that seasonally variable effect
of the El Niño-Southern Oscillation on extratropical circulation is strongly modulated
by AAO (Silvestri & Vera, 2003; L'heureux & Thompson, 2006).

Arctic oscillation (AO) variability depends strongly on the strength of the Icelandic and Aleutian Lows that are in turn correspond to the variability of the North American Oscillation and Pacific North American patterns respectively (Thompson & Wallace, 1998). Although Arctic Oscillation has a strong correlation with North Atlantic Oscillation, it does not show the same changes in summer period as the latter in the last few decades (Hanna et al., 2015).

North Atlantic Oscillation (NAO) is a hemispheric meridional oscillation represented as the surface pressure fluctuation between the Azores and Iceland, which governs major weather patterns in Europe (Visbeck et al., 2001). NAO is strongly related to the occurrence and intensity of the blocking high pressure over Greenland (Hanna et al., 2015).

East Atlantic (EA) is structurally similar to NAO and is characterized by north south dipole anomaly centers over North Atlantic. Compared to NAO the centers of the
 anomalies are shifted southeastward. Differently from NAO, EA has a strong subtrop ical link (Barnston & Livezey, 1987). EA affects the temperature and precipitation in
 Europe and North America (Comas-Bru & Mcdermott, 2014).

Greenland Blocking Index (GBI) is the mean 500 hPa geopotential height for the 60-80°N, 20-80°W region. GBI represent the blocking over Greenland which has an impact on the climate and weather in the Northern Hemisphere (Hanna et al., 2016). Variability of the GBI is related to the North Atlantic Oscillation variations, as well as to changes in East Atlantic and Scandinavian patterns (Scherrer et al., 2006; Hanna et al., 2015).

Scandinavian Pattern (SCA) is a low-frequency teleconnection that represents geopotential height anomalies over the extratropical Northern Hemisphere, particularly centered around the Scandinavian Peninsula, northeastern Atlantic and central Siberia (Bueh & Nakamura, 2007). The positive phase is related to the major blocking anticyclones over Scandinavia and Russia. This teleconnection is often related to temperature and precipitation in Europe with the contrasting effect on the Southern and Northern parts (Liu et al., 2014).

Mediterranean Oscillation Index (MOI) is a patterns that captures pressure differences between the West and East Mediterranean (i.e., Algiers and Cairo). The MOI is related to the activity of cyclogenesis over the Mediterranean Sea (Trigo et al., 2002). Since this process is triggered by the cold fronts from the Atlantic, there is a relation between MOI variability and the North Atlantic Oscillation, as well as Arctic Oscillation (Dünkeloh & Jacobeit, 2003). The monthly time series of all climatic indices, their wavelet and power spectrums (see Section 3.3.2) used in this study are displayed in Figures S2-S3. The links to the time series of climatic indices sources is provided in the Open Research Section.

The time series of antecedent wetness conditions for the whole study period are also considered as a possible covariate and approximated as the monthly mean of catchmentaveraged precipitation of all study catchments for the forecasting of the country-wide severity index and of catchments belonging to the respectful region (i.e., North-Western, North-Eastern and Southern) for forecasting of the regional severity indices (Figure S4). Catchment-averaged precipitation was obtained from daily observed gridded precipitation dataset REGNIE (Rauthe et al., 2013).

283 3 Methods

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3.1 Architecture of the dilated Convolutional Neural Networks

Convolutional Neural Networks (CNN) are deep learning models that consist of the 285 stacked convolutional layers where the output is connected to local regions of the input 286 (Figure 2b) by means of a convolution (i.e., dot product) between the input and sliding 287 filters (i.e., matrix of weights). This architecture results in much smaller number of train-288 able parameters compared to the fully-connected networks (Figure 2a). Due to their ability to effectively recognize patterns in the input series, CNN networks are particularly 290 successful in image recognition and time series classification (Krizhevsky et al., 2012; Le-291 cun et al., 2015; Wang et al., 2016). To allow the CNN network to access larger portion 292 of time series beyond local nodes and to learn recurring patterns at both shorter and longer 293 time scales, a dilation can be added to each convolutional layer (Figure 2c). The dila-294 tion essentially means applying the same sliding filter, but skipping input values with 295 a certain step (van den Oord et al., 2016). At each subsequent convolutional layer the 296 dilation progressively increases by a factor of two (Figure 2c), hence allowing model to 297 efficiently learn connections even between far-apart points while keeping number of train-298 able parameters low (Borovykh et al., 2017). In case of time series forecasting the un-299 derlying idea of the dilated convolutional architecture is to learn repeating patterns in 300 the time series and use them to predict future values. Multi-layer structure of these net-301 works acts similar to wavelet transformation and is effective for discerning low- and high-302 frequency components from time series while reducing noise propagation with each layer 303 (Aussem & Murtagh, 1997). 304

Thanks to the convolutional structure of the dCNN model (Figure 2b), it is much 305 easier to train, faster to converge and requires less memory compared to the state-of-the-306 art RNN models (Benhaddi & Ouarzazi, 2021). Moreover, the adapted Wavenet model 307 performs dilated convolutions on both the response time series (i.e., index of widespread 308 flood severity) and on the covariates (e.g., large-scale climatic indices, antecedent pre-309 cipitation) (Figure 2d) allowing the model to learn not only the inherent patterns from 310 the input time series, but also to learn its dependencies with the covariates at different 311 time scales. The adapted version uses rectified linear unit (ReLU) instead of the gated 312 activation function (Figure 2d,e). The parametrized skip connection at the first layer (Fig-313 ure 2d) ensures that the Wavenet network can learn to discard a covariate if it does not 314 improve the forecast. The result from the first layer is the input for the next convolu-315 tional layer (Figure 2e), this step is repeated till the last convolutional layer that cor-316 responds to the selected depth of the network. The output of the last convolutional layer 317 is passed through a 1x1 convolution to flatten the output and produce one-dimensional 318 forecast (Figure 2f). For a detailed description of the adapted Wavenet model refer to 319 Borovykh et al. (2017). 320



Architecture of dilated Convolutional Neural Networks and the structure of the Figure 2. adapted Wavenet model: a) example of a fully-connected network where all nodes of the output layer are connected to all nodes of the input resulting in many trainable parameters; b) example of a convolutional network where each node is connected only to a local region of the input and the weights are shared, i.e., reducing number of trainable parameters compared to the fullyconnected network and allowing more efficient training, but only learning local (i.e., short-term) dependencies); c) example of a dilated convolutional network where regular local convolutions as in panel b are substituted by dilations to allow output to be influenced by more nodes from the input and efficiently learn the connections even between far-apart data points, i.e., in case of time series this allows to account for short-term (i.e., local) and long-term structures in the data; d) network structure of the adapted Wavenet model: dilation of the input (i.e., autoregressive part) and covariate time series with ReLU activation function (non-linearity term) and parametrized skip connection at the first convolutional layer that allows the model learn to discard covariates if they do not improve the forecast; e) dilation at each subsequent convolutional layer i of the network; f) 1x1 convolution of the last convolutional layer L to reduce dimensionality and obtain one-dimentional forecast. Differently from fully-connected networks (panel a) where the data is passed through all layers sequentially, residual connections at each convolutional layer allow data to bypass some layers resulting in very efficient training (modified from Borovykh et al. (2017))

Table 1.	Model set-ups and corresponding covariates to forecast severity of widespread floods
at time step	p t

Model	Structure	$\parallel \#$ covariates
auto	$[severity_1, severity_{t-1}]^{a}$	0
season	$[severity_1, severity_{t-1}] + [sin_1, sin_{t-1}] + [cos_1, cos_{t-1}]^{b}$	2
precip	$[severity_1, severity_{t-1}] + [sin_1, sin_{t-1}] + [cos_1, cos_{t-1}] + [P_1, P_{t-1}]^{c}$	3
climate	$[severity_1, severity_{t-1}] + [sin_1, sin_{t-1}] + [cos_1, cos_{t-1}] + [Clim_1, Clim_{t-1}]^d$	12
full	$[severity_1, severity_{t-1}] + [sin_1, sin_{t-1}] + [cos_1, cos_{t-1}] + [P_1, P_{t-1}] + [Clim_1, Clim_{t-1}]$	13

^a Forecast of flood severity for time step t using only past observations of flood severity index from time

step 1 to t-1 (i.e., autoregressive part)

 b Sine and cosine of the month of the year as a proxy of seasonality

 c Mean monthly country-averaged or region-averaged observed precipitation

^d Ten different climatic indexes are considered simultaneously: PNA, SOI, NP, AAO, AO, NAO, EA, GBI, SCA, MOI

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3.2 Setup of the adapted Wavenet model

We use a dCNN network with four dilation layers, which means that 32 input time 322 steps (i.e., 2^5 months) are used to produce a single output time step (i.e., forecast) (Fig-323 ure 2c). This choice leverages the length of the available time series and the number of 324 trainable weights. The total length of monthly time series (756 months, i.e., 63 years from 325 January 1951 to December 2013) was split into training (32+640 months), i.e., ca. 2.7+53.3326 years) and test samples (32+52 months, i.e., 2.7+4.3 years). The test sample includes 327 one of the most severe floods (June 2013) recorded in instrumental observation period 328 in Germany that affected the Elbe River and the Danube River catchments (Blöschl et 329 al., 2013) (the North-Eastern and the Southern regions respectively). We use time-series-330 split cross-validation that resembles K-fold cross-validation for ordered time series and 331 is designed specifically for time series forecasting (i.e., validation split never preceeds train-332 ing splits). In our case we use 7 time series splits (i.e., folds) for the training set. We use 333 adaptive model estimation (Adam) optimizer (Kingma & Ba, 2014) with the initial learn-334 ing rate of 0.005 to train the model and mean absolute error as a measure of model per-335 formance. The maximum number of epochs was set to 150. Additionally, we used an early 336 stopping strategy to avoid overfitting. 337

To efficiently analyze the value of different covariates used in this study we exam-338 ine models of different complexity that correspond to a different number of covariates 339 used for conditioning the forecast of widespread flood severity. The simplest model (auto 340 model) does not use any covariates and makes the forecast purely based on the struc-341 ture of the patterns identified from the past time series of flood severity index (Table 1). 342 The seasonal model uses two covariates, namely the sine and cosine of the month of the 343 year to allow the model learning of the seasonality dynamics (Jiang et al., 2022). The 344 precipitation model (precipitation time series, i.e., 345 it can use observed precipitation from previous months for the forecast (Table 1). Es-346 sentially, it only uses antecedent precipitation since the model structure is causal (i.e., 347 only past data is used for generating the forecast (Y. Chen et al., 2019)). The climate 348 model does not use precipitation as a covariate (Table 1), but instead includes ten dif-349 ferent climatic indices listed in the Section 2.2. Again, similar to precipitation only cli-350 matic indexes from previous months (Table 1) are used for the forecast given the causal 351 structure of the model. The full model uses all above mentioned information (i.e., to-352 tal of 13 covariates, Table 1). 353

To account for the stochastic nature of the training of the deep learning models we 354 initialize training of each of the above mentioned models 10 times independently of each 355 other (Jiang et al., 2022). Therefore, we obtain an ensemble of 10 realizations for train-356 ing and also for test period in forecasting mode for each examined model complexity. To 357 avoid overfitting of the model during training we use L1 type of regularization that re-358 duces the weights of the less influential covariates to zero (i.e., it only selects covariates 359 that are essential for the forecasting of flood severity). We identify an optimal regular-360 ization parameter value for each model complexity by systematically varying L1 param-361 eter value and examining corresponding model performance (Figure S13). Too low a value 362 of L1 parameter results in model overfitting (i.e., higher errors for test samples compared 363 to train samples in cross-validation), while too high values of L1 parameter results in a 364 very strict penalty and can prevent model from learning the variability beyond the long-365 term mean and seasonality (Figure S13). 366

Although from the theoretical point of view dCNN models are specifically designed 367 to effectively learn patterns from short time series (Borovykh et al., 2017), the length 368 of the input data available for training can still affect model performance. To evaluate 369 the effect of the length of time series available for training on model performance we per-370 form an experiment by systematically reducing the length of the training series, while 371 keeping the length of the test series unchanged. In all examined cases we compare the 372 performance with the baseline (i.e., mean monthly flood severity index derived from the 373 whole observed series) to identify minimum requirements for the duration of training data 374 for flood forecasting using dCNN models. 375

376

3.3 Explainable machine learning

Recently, explainable machine learning methods (XAI) (Buhrmester et al., 2021) 377 have been shown to be useful for understanding the structure of the otherwise black-box 378 deep learning models. Interpretability and explainability of deep learning models rests 379 on the ability of the post hoc explanation of model prediction strategy by understand-380 ing the importance of the features used for predictions (Mamalakis et al., 2022b) and 381 by evaluating the physical plausibility of the identified importances (Ebert-Uphoff & Hilburn, 382 2020). In our case, this means evaluating the role of different climatic indices and an-383 tecedent precipitation for the forecasting of widespread flood severity using a model de-384 pendent feature attribution method (i.e., Shapley values, see Section 3.3.1) and compar-385 ing the results with the model-independent wavelet coherence analysis between these variables (see Section 3.3.2), as well as analyzing the results on the importance of different 387 features obtained from this study with the reports from previous studies (see Section 5). 388

389

3.3.1 Feature attribution: Shapley value

To evaluate the usefulness of different covariates used in this study (i.e., large-scale 390 climatic indices, antecedent precipitation) we used the SHAP method (Lundberg et al., 391 2017) implemented in DeepExplainer library that approximates Shapley values (Shapley, 392 1952). Shapley values quantify the average marginal contribution of a feature value (e.g., 303 antecedent precipitation) across all possible coalitions (i.e., all possible combinations of 394 the features). The Shapley value of each feature shows how the forecast will change when 395 this feature is added to the set of all other considered features. We implemented SHAP 396 method in a global setting, i.e., by computing Shapley values not for a single forecast 397 (i.e., month or event), but for the full sample to understand the overall effect of each co-398 variate on model decisions (Buhrmester et al., 2021). 399

400

3.3.2 Wavelet transformation and wavelet coherence

Given the inherent similarity of the dCNN and wavelet transform (Aussem & Murtagh, 1997) we compare the results of the model-dependent feature attribution method described

in the previous Section with the model-independent continuous wavelet transform and 403 biwavelet analysis of the flood severity index and the covariates used in this study. Wavelet 404 analysis is an effective tool to transform 1D-time series into 2D-time-frequency spectrum. 405 We performed wavelet analysis using a continuous wavelet transform with the Morlet wavelet 406 function (Torrence & Compo, 1998) of the monthly time series of climatic indices (Fig-407 ure S2-S3), as well as the flood severity index (Figure S1) and precipitation (Figure S4) 408 for the whole country and for the three hydroclimatic regions individually. The significance level of the wavelet spectrum is computed in relation to red noise (Torrence & Compo, 410 1998). 411

Biwavelet analysis is performed using Wavelet Coherence Transform (R package 'biwavelet') to identify the relation between two time series at different scales. In this case the coherence is quantified by a cross-wavelet power (the measure of the coherence in a time frequency domain) that resembles standard correlation coefficient (Torrence & Compo, 1998).

417 **4 Results**

418 419

4.1 Forecasting performance in Germany and its three hydroclimatic regions

All investigated models reduce the mean absolute error (MAE) of the forecast when 420 compared to the baseline (i.e., mean monthly observed flood severity index) (Figure 3). 421 However, more complex models that include climatic indices (i.e., climate and full model) 422 are able to predict the widespread flood severity index more accurately when compared 423 to the models with fewer covariates (R^2 increases from 0.04 to 0.32, Figure 4). Partic-424 ularly, more complex models reduce the phase shift that is apparent for simpler mod-425 els (Figure 4), although there is still a considerable underestimation of flood severity, es-426 pecially for the most extreme events (Figure S14). The forecasting skill of the best model 427 (i.e., climate or full model depending on the hydroclimatic region) varies across differ-428 ent regions when compared to the corresponding baseline: the models show higher skill 429 in the North-Western and the North-Eastern regions (Figure 3), indicating that the pre-430 dictability of widespread flooding in these regions from climate indices and antecedent 431 precipitation might be higher. 432

433

4.2 The role of antecedent precipitation for forecast performance

There is a clear reduction of mean absolute error with increasing model complex-434 ity when forecasting flood severity index in Germany, although it is worth to mention 435 that the full model is associated with the highest level of performance variability as may 436 be expected given the increase in the number of effective parameters (Figure 3). When 437 comparing the performance of the climate and full models we observe considerable re-438 gional differences. As for the country-wide forecast, in the North-Eastern region the full 439 model outperforms the climate model that does not account for antecedent precipita-440 tion. In the Southern and North-Western region the full model shows much higher un-441 certainty (i.e., high variability in the performance among different realizations) compared 442 to the climate model, with considerably better median performance of the climate model 443 in the Southern region (Figure 3). Variable regional performance of the model that ac-444 counts for antecedent wetness conditions (i.e., full model, Table 1) confirms a variable 445 role of antecedent wetness for generation of widespread flooding in different German re-446 gions suggested in R. Merz et al. (2020) who showed that especially in the drier North-447 Eastern parts of Germany antecedent soil moisture is an even more important feature 448 for generation of annual maximum floods than the volume of precipitation events that 449 induces the corresponding flood. 450



Figure 3. Mean absolute error (MAE) for the forecast of monthly widespread flood severity in Germany (DE) and its three hydroclimatic regions (NW: North-Western region; NE: North-Eastern region; South: Southern region) for the test period (August 2009-December 2013) using models of different complexity (ensemble of 10 realizations for each case) (see Table 1). Grey dashed line indicates MAE for the baseline case (i.e., mean monthly index of flood severity for the whole study period from 1951 to 2013). The tendency in performance improvement of the auto, season and precip models in Germany and individual regions are very similar, therefore, only the performance of climate and full models are shown for the three hydroclimatic regions



Figure 4. One month ahead forecast of monthly severity of widespread flooding in Germany for test period (August 2009-December 2013) using models of increasing complexity (see Table 1). The black solid line indicates observed flood severity index. Blue solid line represent the mean of the ensemble of 10 realizations for each model. Light shaded area shows the spread of the ensemble simulations with darker color indicating the spread between 25th and 75th percentiles.

451 452

4.3 Temporal variability of performance: forecasting floods generated by different processes

We also detect considerable differences in the forecast skill in time (Figure 4). For example, the forecast for summer 2013 when one of the most extreme flood events in Germany occurred was very poor for all models (Figure 4). On the other hand, the forecast of monthly severity corresponding to the winter-spring flood of 2011 was fairly accurate (Figure 4), especially in individual regions (Figure S15). This indicates that there might be considerable differences in the predictability of floods in the examined sample.

As floods can be generated by different atmospheric and hydrological processes (Hirschboeck, 459 1987; Tarasova et al., 2019) and might be characterized by considerably different con-460 tribution of different ingredients (e.g., volume and intensity of inducing precipitation events, 461 wetness states) (R. Merz et al., 2020), their intrinsic predictability from climatic indices 462 and antecedent precipitation might indeed be different. Therefore, we compare the abil-463 ity of the full model to forecast events that correspond to different event types according to Tarasova, Basso, Wendi, et al. (2020). The results show that during the test phase 465 there is on average higher underestimation (in terms of percent bias) of rainfall-generated 466 events compared to the events that are generated by mixture of rainfall and snowmelt 467 (Rain.Snow) (Figure 5d), although the variability in the performance for individual events 468 is rather high. Despite their rarity only rainfall intensity-dominated events (Rain.Dry.Intensity, 469 Figure 5b) are associated with smaller bias than mixtures of rainfall and snowmelt in 470 the cross-validation phase (Figure 5c). However, these events also correspond to the low-471 est severity (in terms of the number of affected gauges and the magnitude of the events) (Figure 5a), while Rain.Snow events together with the Rain.Wet.Volume are associated 473 474 with the highest severity. Interestingly, although Rain.Wet.Intensity and Rain.Dry.Volume that usually occur in autumn and summer (Tarasova, Basso, Wendi, et al., 2020) are as-475 sociated with smaller severity, they are strongly underestimated (Figure 5). Therefore, 476 it seems that winter/spring processes can be better predicted using monthly climatic in-477 dices than summer processes when convection might contribute considerably to the gen-478 eration of rainfall (Conticello et al., 2020). 479

480 481

4.4 Model-independent regional wavelet coherence of the covariates and flood severity index

To increase interpretability of modeling results, first we analyze wavelet coherence 482 between model covariates and flood severity index that is independent from model per-483 formance. As expected, we observe a very strong coherence between precipitation and 484 flood severity index in Germany, especially at longer time scales (up to several years, Fig-485 ure 6) that corresponds well with the previously reported dependence between the oc-486 currences of multi-year flood-rich and flood-poor periods and antecedent wetness con-487 ditions in Germany (B. Merz et al., 2018). Interestingly, at the shorter time scales the 488 coherence is rather intermittent. We also detect considerable differences in coherence among different regions (Figure 6). Particularly in the Southern region the coherence with pre-490 cipitation is lower than in the rest of the country, which corresponds well with inferior 491 performance of the full model in this region (Figure 3). 492

493 Among climatic indices that reflect global-scale climatic patterns and are related with the El Niño Southern Oscillation, there is particularly persistent coherence between 494 PNA and NP and flood severity index in Germany (Figure 7, Figure S5). A very sim-495 ilar coherence pattern is observed in the North-Western region (Figure 7, Figure S6), while 496 in the North-Eastern the coherence is very intermittent indicating weaker effect of these 497 global patterns on the occurrence of widespread flooding in the region (Figure 7, Fig-498 ure S7). Generally, there is also a very weak coherence between flooding in the North-499 Eastern region and all considered global patterns compared to the North-Western region 500 (Figure S6 and S7). Interestingly, PNA seems to be very coherent with flood severity in 501



Figure 5. Severity of flood events generated by different processes in Germany (a) and underestimation (expressed as $bias = (severity_{obs} - severity_{mod})/severity_{obs})$ of their corresponding magnitudes during: c) cross-validation and d) test periods. Forecast is based on the ensemble of 10 full models that consider all covariates. Significance between event types is evaluated using pairwise Kruskal-Wallis test with false discovery rate correction. Significance levels: *** for p<0.01; ** for p<0.05; * for p<0.1. Panel b shows the proportions of events generated by different processes during cross-validation and test periods. Panel a is shown in log space. Panels b, c and d are shown in linear space.



Figure 6. Wavelet coherence (WC) between monthly flood severity index and antecedent precipitation in Germany (DE) and its three hydroclimatic regions (NW: North-Western region; NE: North-Eastern region; South: Southern region). x axis is a time axis in months from the beginning to the end of the time series. y axis indicates periodicity or scale at which two time series are coherent at a given moment in time of the x axis. Black outline highlights regions of significant coherence. Red colors correspond to higher coherence values, blue colors correspond to lower coherence values. The shaded part outlines the cone of influence. Wavelet spectrum of the corresponding monthly flood severity index and precipitation series can be found in Figure S1 and Figure S4.

the Southern region especially at multi-annual time scales and this relation is even more persistent than in the North-Western region (Figure 7, Figure S8).

Previous study indicated the importance of the climatic patterns associated with 504 the activity Atlantic Ocean, particularly of the Baltic Sea, and with the occurrence of 505 blocking situations for precipitation and flooding in the Northern Europe (Zanardo et 506 al., 2019; Comas-Bru & Mcdermott, 2014). Indeed, our results also indicate coherence 507 between widespread flooding and SCA in all regions (Figure 7). There is also consider-508 able coherence with NAO at annual and multi-annual scales in the North-Western and 509 North-Eastern regions and with EA at longer time scales (Figure 7). Instead, in the South-510 ern region the flooding is coherent with AO at longer time scales (Figure S8). 511

The coherence between widespread flooding and MOI seems to be rather similar between North-Western and North-Eastern regions where the dependence is strong at longer time scales. Differently, in the Southern region, the coherence is stronger at shorter time scales (Figure 7). Generally, wavelet coherence analysis between climatic indices and widespread flood severity shows clear differences in hydroclimatic controls across three German regions.

518

4.5 Model-dependent feature attribution using Shapley values

Since climate model showed a comparable or even better performance than the full model in the North-Western and Southern regions we performed feature attribution using Shapley values for these model complexity versions. We compare the differences between the feature rankings of two model versions to provide insights on the importance of different covariates and on potential of substituting antecedent precipitation solely by the information available from climatic indices in different hydroclimatic regions of Germany (Figure S9-S12).



Figure 7. Wavelet coherence (WC) between monthly flood severity index and large scale climatic indices in Germany (DE) and its three hydroclimatic regions (NW: North-Western region; NE: North-Eastern region; South: Southern region): PNA - Pacific North American Pattern; NAO - North Atlantic Oscillation; EA - East Atlantic; SCA - Scandinavian Pattern; MOI - Mediterranean Oscillation Index. x axis is a time axis in months from the beginning to the end of the time series. y axis indicates periodicity or scale at which two time series are coherent at a given moment in time of the x axis. Black outline highlights regions of significant coherence. Red colors correspond to higher coherence values, blue colors correspond to lower coherence values (see colorbar in Figure 6). The shaded part outlines the cone of influence. Wavelet spectrum of the corresponding monthly flood severity and climate indices can be found in Figures S1-S3.

Regardless of model complexity, sine and cosine of the month of the year corresponds to high Shapley values in all regions (Figure 8), as widespread flooding is associated with very pronounced seasonality, especially in the Northern regions (Figure S1). In the Southern region where floods occur throughout the year (Beurton & Thieken, 2009) and the seasonality of widespread flooding is not as pronounced (Figure S1), sine and cosine correspondingly are ranked lower (Figure 8).

532

4.5.1 Germany-wide model

For the forecasting of the Germany-wide flood severity index the full model that 533 includes antecedent precipitation performed better than climate model. Feature attri-534 bution shows that apart from antecedent precipitation and seasonality, global patterns, 535 such as PNA and AAO, patterns associated with the Baltic Sea (EA and SCA) and with 536 the Mediterranean Sea (MOI) are found important for the model performance (Figure 537 8). This corresponds well with the wavelet coherence analysis (Figure 7) that indicates 538 a coherence between these indices and the flood severity index at different time scales. 539 Interestingly, despite a clear coherence between Germany-wide flood severity indices and 540 NP pattern (Figure S5), it was not identified as an important predictor by the Wavenet 541 model (Figure 8). This might be explained by a strong relation between NP and other 542 global atmospheric patterns related to the El Niño Southern Oscillation (Park et al., 2013). 543 There is also a strong relation to the GBI index at annual scale for Germany-wide flood 544 severity (Figure S5), although GBI does not seem to be an important covariate for model 545 predictions (Figure 8). A very pronounced seasonality of GBI (Figure S3) might be al-546 ready accounted for by sine and cosine of month of the year that are used as support co-547 variates to improve seasonality learning (Jiang et al., 2022). Interestingly, when the pre-548 cipitation are not considered (i.e., in the case of climate model), the importance of the 549 NP and GBI indices that have strong coherence with precipitation at annual scale (Fig-550 ure S9) increases (Figure S16). 551

552

4.5.2 North-Western region

The climate model (i.e., the model without the antecedent precipitation) performs 553 slightly better than the full model in the North-Western region (Figure 3), indicating that 554 the information on antecedent precipitation can be efficiently substituted by monthly 555 variability of other covariates in this region. In fact, although antecedent precipitation 556 is ranked first among covariates of the full model, feature attribution of the climate model 557 shows that SCA index gains more importance in the absence of precipitation (Figure 8). 558 as they are both strongly related to each other in this region according to the model in-559 dependent wavelet coherence analysis (Figure S10). Also the importance of the dynam-560 ics of the historical flood severity index itself increases for climate model compared to 561 the full model of the North-Western region. 562

Apart from antecedent precipitation and seasonality, global patterns that are par-563 ticularly related to the El Niño Northern Oscillation (e.g., PNA, NP, AAO and SOI), 564 as well as the East Atlantic pattern that indicates the strong effect of the Baltic Sea, are 565 ranked high for model performance (Figure 8). Interestingly, according to the wavelet 566 coherence analysis there is a strong relation between PNA and SOI patterns and flood 567 severity index particularly at medium time scale (annual to bi-annual) (Figure 7, Fig-568 ure S6), at which the relation between the flood severity index and precipitation is the 569 lowest (Figure 6). AAO is related to flood severity index at much shorter scales, while 570 the coherence with the East Atlantic pattern is instead detected at longer (ca. 3 to 5 years) 571 time scales (Figure 7, Figure S6). 572



Figure 8. Shapley values of the covariates of the full model (for Germany and all regions) and for climate model (for the North-Western and Southern region where climate model performed comparable or better than the full model) computed for the whole time series from 1951 to 2013 (i.e., global settings) ordered by decreasing variable importance: auto - the historical time series of flood severity; sin - sine of the month of the year; cos - cosine of the month of the year; Precip - antecedent mean monthly precipitation; NP - North Pacific Patterns; PNA - Pacific North American Pattern; SOI - Southern Oscillation Index; AAO - Antarctic Oscillation; AO - Arctic Oscillation; NAO - North Atlantic Oscillation; GBI - Greenwich Blocking Index; SCA - Scandinavian Pattern; EA - East Atlantic; MOI_-Mediterranean Oscillation Index

573 4.5.3 North-Eastern region

In contrast to the North-Western region, feature attribution analysis does not in-574 dicate any major effect of global patterns on the occurrence of widespread flooding in 575 the North-Eastern region, where apart from very strong effect of seasonality, only an-576 tecedent precipitation and East Atlantic pattern contribute considerably to the perfor-577 mance of the full model (Figure 8). The strong effect of the latter on the flood severity 578 index especially at longer time scales is also apparent from the corresponding wavelet 579 coherence analysis (Figure 7). Interestingly, there is no difference in feature importance 580 of climatic indices between climate and full model of the North-Eastern region (Figure 581 S16), but the full model performs considerably better (Figure 3). 582

583 4.5.4 Southern region

In the Southern region, the PNA pattern has the highest contribution to model per-584 formance for both climate and full models (Figure 8), and is strongly related to the flood 585 severity index at longer time scales (Figure 7). At shorter time scales the effect of the 586 MOI pattern is apparent that is also associated with high Shapley values (Figure 8) and 587 might indicate the importance of the Mediterranean Sea for the occurrence of flooding 588 in this region (Figure 7). In this region the climate model performs considerably bet-589 ter than the full model and the feature attribution indicates that the importance of SCA, EA and NAO indices increases in the absence of precipitation covariate (Figure 8). Ac-591 cording to wavelet coherence analysis precipitation in the Southern region show strong 592 dependencies with SCA at longer time scales and with EA and NAO at shorter time scales 593 (Figure S12), indicating that the information on antecedent precipitation can be ef-594 fectively extracted from the climatic indices in this region. 595

596 5 Discussion

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598

5.1 Adapted Wavenet conditioned by climatic indices for monthly flood forecasting

Dilated convolutional neural networks provide a powerful tool for time series fore-599 casting when only limited length of time series is available for training (Borovykh et al., 600 2017), especially compared to the recurrent neural networks (Y. Chen et al., 2019). The 601 consistent improvement compared with the baseline (i.e., mean monthly values of widespread 602 flood severity index) for all model complexities investigated in our study shows the po-603 tential of dilated convolutional models for flood forecasting, although even the most com-604 plex model tends to underestimate the most extreme events (Figure S14). Moreover, un-605 fortunately, even for the adapted Wavenet there is a clear dependence between model 606 performance and the length of time series available for training. When less than 40 years of the data is available the model is not able to outperform the baseline based on mean 608 monthly flood severity for the whole study period (Figure S13). This indicates that ex-609 tending forecast lead times to seasonal and longer scales using the Wavenet model alone 610 does not seem feasible yet. However, using high-fidelity seasonal forecasts of climate in-611 dices produced by the global climate models (Feng et al., 2021) as the input for the Wavenet 612 model might be a promising avenue to extend lead times also for flood forecasting. 613

Apart from that, wavelet coherence analysis clearly shows that there are strong dependences between severity of widespread flooding and climatic indices even at long time scales (i.e., longer than 5 years) (Figure 7, Figure S5-S8). Rather shallow dilation depth (i.e., four layers) selected in our study due to the limited time series available for training might have impacted the ability of the adapted Wavenet model to account for such dependencies. Despite this, a good agreement between the model-independent wavelet coherence analysis and feature attribution analysis using Shapley values indicate that the model was able to identify the primary climatic drivers and might become a valuable forecasting tool especially as the length of flood observations increases.

Even with good agreement between wavelet coherence analysis and feature attri-624 bution method, it might be difficult to pre-select the features beforehand using wavelet 625 coherence analysis alone due to high inter-relation between the indices (Figure S9-S12). 626 Despite the comprehensiveness of monthly climatic indices they do not capture the full 627 picture of large scale climatic patterns that might be important for spatially co-occurring 628 floods. An alternative design of the Wavenet model including daily indices that resolve 629 shorter phenomena such as synoptic waves (Lu et al., 2013) might provide an opportu-630 nity for a more detailed near-time event-based widespread flood forecasting. Moreover, 631 a more comprehensive and flexible patterns (e.g., based on integrated vapor transport) 632 (Nakamura et al., 2013; Lima et al., 2017; Conticello et al., 2020), in combination with 633 the convolutional neural networks might also be a promising avenue for a more accurate 634 event-based forecasting of spatially co-occurring floods. 635

5.2 Predictability of floods generated by different processes

The predictability seems to vary depending on generation processes of individual 637 widespread flooding events (Figure 5). Rain-on-snow events (i.e., mixture of rainfall and 638 snowmelt) are associated with the highest predictability despite their severity. In Ger-639 many these events are associated with westerly and north-westerly circulation patterns 640 (Nied et al., 2014), and are often the cause of the widespread transboundary floods par-641 ticularly during negative NAO and SCA phases (Krug et al., 2020). Our results also in-642 dicate that SCA is one of the most important covariates for the forecasting of widespread 643 flooding in Germany (Figure 8). Moreover, Conticello et al. (2020) notes that high flow 644 events in the North-Western part of Germany, where floods almost exclusively occur in 645 winter and are usually generated by rain-on-snow processes (Tarasova, Basso, & Merz, 646 2020), are associated with higher predictability from large scale circulation patterns than 647 the events in the Eastern and Southern parts of the country. 648

In contrast, volume-dominated floods generated by rainfall during dry conditions 649 often occur in late summer and the beginning of the autumn (Tarasova, Basso, & Merz, 650 2020) and have limited predictability (Figure 5). On one hand it can be limited by the 651 complex patterns of soil moisture that modulate local occurrence of floods and hence di-652 rectly affect the spatial extent of flooding (Nied et al., 2017), while on the other hand 653 the meteorological phenomenon that are not well captured by the climatic indices, such 654 as local convective processes (R. Merz & Blöschl, 2003; Kemter et al., 2020) and the ac-655 tivity of the Vb cyclones that is associated with heavy precipitation and recent summer 656 floods in Central Europe and particularly in Germany (Hofstätter et al., 2018; Krug et 657 al., 2021), might play a more important role for the generation of these events. 658

⁶⁵⁹ Changing climatic conditions and changing flood generation processes indicate that
 ⁶⁶⁰ the predictability of widespread flooding might decrease as the number of dry rainfall
 ⁶⁶¹ generated events increases in Central, Eastern and Southern Germany (Tarasova et al.,
 ⁶⁶² 2023; Winter et al., 2022), which might deteriorate the efficiency of early warning sys ⁶⁶³ tems in these regions and urges the development of the methods that might integrate
 ⁶⁶⁴ the atmospheric phenomena and soil moisture patterns relevant for their forecasting.

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5.3 Regionally variable importance of climatic indices and antecedent wetness conditions for forecasting widespread floods

Feature attribution analysis implemented in our study shows clear differences in hydroclimatic drivers of widespread flooding across different regions in Germany. Par-

ticularly, there is apparent difference in the value of global patterns associated with El 669 Niño Southern Oscillation between Western and Eastern parts of Germany. While in the 670 North-West and in the Southern parts a global Pacific North American pattern is a very 671 important feature for model predictions (Figure 8), it has only secondary effect in the 672 North-Eastern part. Although there is few evidence about the effect of El Niño in Eu-673 rope (Brönnimann, 2007), there is a clear coherence between widespread floods in the 674 North-West of Germany and PNA patterns (Figure 7) that is also coherent with the find-675 ings of Conticello et al. (2020), who found that high discharge events in the Western part 676 of Germany are associated with intensive moisture transport from Tropics that is gov-677 erned by global patterns, such as PNA. 678

Feature attribution analysis indicates that AAO pattern might be important for forecasting widespread flood severity in the North-Western region as well. However, although AAO is able to modulate the effect of El Niño Southern Oscillation on the extratropical circulations (L'heureux & Thompson, 2006), the exact process chain of its effect on flooding in the Northern Hemisphere is not understood.

In the drier North-Eastern region we see a very pronounced role of seasonality and 684 antecedent wetness conditions (Figure 8). This is in line with R. Merz et al. (2020), who 685 noted that in this region the antecedent moisture conditions is even more important for 686 the emergence of local floods than the volume of inducing precipitation events. Also Nied 687 et al. (2017) found that in this region the spatial extent of moderate floods (i.e., 2-year 688 return periods) are modulated more by the specifics of the soil moisture patterns than 689 by weather patterns. This was later corroborated by B. Merz et al. (2018) highlighting 690 the ultimate role of antecedent wetness conditions in this region for the occurrence of 691 spatially coherent flood-rich periods. It is also important to note the effect of the East 692 Atlantic pattern for widespread flooding in this region (Figure 8) indicating the primary 693 role of the Baltic Sea for moisture transport and generation of flood-producing precip-694 itation in this region. 695

Interestingly, the results of feature attribution indicates relatively low importance 696 of the North Atlantic Oscillation on model predictions, that is also confirmed by the model-697 independent wavelet coherence analysis that shows rather low coherence between the sever-698 ity of widespread flooding and NAO (Figure 7-8). Previous studies have shown strong 699 dependence between NAO and winter precipitation in Europe (Haylock & Goodess, 2004; 700 Villarini et al., 2011). However, particularly for flood losses the effect of NAO is notably 701 weaker in Germany compared to the Atlantic coast of Great Britain and Norway and 702 the Mediterranean region (Zanardo et al., 2019). 703

In the Southern region in addition to the effect of global PNA pattern, the effect 704 of the Mediterranean Sea in terms of the Mediterranean Oscillation Index also becomes 705 apparent (Figure 8). Severe floods in this region are often associated with the activity 706 of the Vb cyclones propagating north-eastward from the Mediterranean Sea to Central 707 Europe (Blöschl et al., 2013; Hofstätter et al., 2018; B. Merz et al., 2018) that might be 708 additionally intensified by the Mediterranean Sea through pre-moistening of the conti-709 nental moisture (Krug et al., 2021). Therefore, adding more indices or patterns related 710 to the activity of the Mediterranean Sea might improve the forecasting of widespread 711 flooding, especially in the regions where summer floods occur. 712

It is worth to note that the relevance of the climatic indices discussed above is conditioned on the selected study period. The attribution results depend greatly on the chosen reference period (Mamalakis et al., 2022a) (i.e., in our case the period from 1951 to 2013) and the relevance of climatic indices might change for a different period of reference.

718 6 Conclusions

In this study we investigate the potential of the adapted Wavenet model to fore-719 cast the monthly severity of the widespread flooding in Germany and its three distinct 720 hydroclimatic regions individually using various large-scale climatic indices and antecedent 721 precipitation as covariates. We observe a consistent increase in model performance with 722 increasing model complexity (i.e., increasing number of covariates), but there is a clear 723 dependence between the length of the available time series for training and model skill. 724 This dependence indicates that even the dilated structure of the Wavenet model is not 725 able to reconcile the limited data availability and the number of trainable parameters, 726 which might be also a limiting factor for adapting the Wavenet model for flood forecast-727 ing at longer (e.g., annual) time scales. 728

The results of the model-dependent feature attribution based on Shapley values are 729 consistent with the model-independent wavelet coherence analysis indicating the relia-730 bility of trained Wavenet model for forecasting of widespread floods. Both methods high-731 light considerable differences in the large-scale hydroclimatic drivers of widespread flood-732 ing across different regions in Germany. While in the North-Western and the Southern 733 regions the effect of the global patterns associated with El Niño Southern Oscillation (i.e., 734 Pacific North American pattern) is apparent, the North Eastern region is affected by more 735 local processes related to the Baltic Sea activity (East Atlantic pattern). In the South-736 ern region, where summer floods are as frequent as winter events, we also detect the ef-737 fect of the Mediterranean Sea (Mediterranean Oscillation Index). 738

Moreover, the feature attribution based on the trained Wavenet model indicates variable importance of antecedent wetness conditions across German regions. Apart from the Southern region antecedent precipitation is ranked as one of the most important covariates. Especially in drier North-Eastern region model performance increases considerably when antecedent precipitation is added as a covariate.

We also see considerable differences in the predictability of flood events generated 744 by different processes. Despite their superior severity, events generated by the mixture 745 of rainfall and snowmelt are associated with low biases, while the predictability of volume-746 dominated rainfall events generated during dry conditions is low. Increasing frequency 747 of the latter in several regions might lead to the deterioration of the predictability of widespread 748 flooding in Germany when only using large scale climatic indices for the forecast. There-749 fore, more efforts should focus on developing sub-seasonal to seasonal forecasting approaches 750 based on convolutional neural networks conditioned on more flexible spatial climatic pat-751 terns that might be more relevant for flood generation in warmer and drier periods. 752

753 Open Research Section

Time series of SOI, NP, AAO, NAO, AO, PNA and GBI used in this study are avail-754 able from https://psl.noaa.gov/data/climateindices/list/ (no registration required, in-755 dividual download for each climatic index), MOI data (Algiers and Cairo) are available 756 from https://crudata.uea.ac.uk/cru/data//moi/ (no registration required, individual down-757 load for each computational variant of the index) and SCA and EA data are available 758 in Comas-Bru and Hernández (2018). Mean monthly precipitation per catchment were 759 obtained from the daily interpolated precipitation dataset available in Rauthe et al. (2013). 760 Severity index of spatial flooding is computed using streamflow events (available in Tarasova 761 et al. (2018)) identified from daily streamflow observations available for download from 762 Global Runoff Dataset (https://portal.grdc.bafg.de/applications/ no registration required, 763 individual download for each streamflow gauge or selected region) and Bavarian Min-764 istry for Environment (https://www.gkd.bayern.de/de/fluesse/abfluss no registration re-765 quired, individual download for each streamflow gauge). The post-processed data used 766 in this study: the severity index of widespread flooding, flood generation processes of each 767

flood event and monthly precipitation for Germany and its three hydroclimatic regions, 768 are available in Tarasova (2023). 769

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