Assessing Storm Surge Multi-Scenarios based on Ensemble Tropical Cyclone Forecasting

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Abstract

Ensemble forecasting is a promising tool to aid in making informed decisions against risks of coastal storm surges. Although tropical cyclone (TC) ensemble forecasts are commonly used in operational numerical weather prediction systems, their potential for disaster prediction has not been maximized. Here we present a novel, efficient, and practical method to utilize a large ensemble forecast of 1000 members to analyze storm surge scenarios toward effective decision making such as evacuation planning and issuing surge warnings. We perform the simulation of TC Hagibis (2019) using the Japan Meteorological Agency's (JMA) non-hydrostatic model. The simulated atmospheric predictions were utilized as inputs for a statistical surge model named the Storm Surge Hazard Potential Index (SSHPI) to estimate peak surge heights along the central coast of Japan. We show that Pareto optimized solutions from an ensemble storm surge forecast can describe potential worst (maximum) and optimum (minimum) storm surge scenarios while exemplifying a diversity of trade-off surge outcomes among different coastal places. For example, some of the Pareto optimized solutions that illustrate worst surge scenarios for inner bay locations are not necessarily accountable for bringing severe surge cases in open coasts. We further emphasize that an in-depth evaluation of Pareto optimal solutions can shed light on how meteorological variables such as track, intensity, and size of TCs influence the worst and optimum surge scenarios, which is not clearly quantified in current multi-scenario assessment methods such as those used by JMA/National Hurricane Center in the United States.

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2 3	Forecasting
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9	Key Points:
10 11	• The potential of ensemble tropical cyclone forecasting for assessing storm surge multi- scenarios is shown.
12 13	• Pareto optimized solutions from an ensemble storm surge forecast can efficiently illustrate potential worst and minimum storm surge scenarios.
14 15 16	• Analyses of meteorological variables of ensemble members in Pareto frontiers help understand the impact of a tropical cyclone on predicted storm surge multi-scenarios.

17 Abstract

Ensemble forecasting is a promising tool to aid in making informed decisions against risks of 18 coastal storm surges. Although tropical cyclone (TC) ensemble forecasts are commonly used in 19 operational numerical weather prediction systems, their potential for disaster prediction has not 20 been maximized. Here we present a novel, efficient, and practical method to utilize a large 21 22 ensemble forecast of 1000 members to analyze storm surge scenarios toward effective decision making such as evacuation planning and issuing surge warnings. We perform the simulation of TC 23 Hagibis (2019) using the Japan Meteorological Agency's (JMA) non-hydrostatic model. The 24 simulated atmospheric predictions were utilized as inputs for a statistical surge model named the 25 Storm Surge Hazard Potential Index (SSHPI) to estimate peak surge heights along the central coast 26 of Japan. We show that Pareto optimized solutions from an ensemble storm surge forecast can 27 describe potential worst (maximum) and optimum (minimum) storm surge scenarios while 28 29 exemplifying a diversity of trade-off surge outcomes among different coastal places. For example, some of the Pareto optimized solutions that illustrate worst surge scenarios for inner bay locations 30 are not necessarily accountable for bringing severe surge cases in open coasts. We further 31 emphasize that an in-depth evaluation of Pareto optimal solutions can shed light on how 32 meteorological variables such as track, intensity, and size of TCs influence the worst and optimum 33 surge scenarios, which is not clearly quantified in current multi-scenario assessment methods such 34 35 as those used by JMA/National Hurricane Center in the United States.

36 Plain Language Summary

Ensemble forecasting generates multiple predictions of a weather event with various possible 37 outcomes based on varying initial conditions, model parameters, and physics. The potential of 38 ensemble tropical cyclone (TC) forecasting for assessing storm surge multi-scenarios has largely 39 been overlooked previously. Enhanced analysis can unlock and maximize the benefit of ensemble 40 forecasting. This study simulated an extremely large ensemble (=1000 members) to reforecast past 41 TC Hagibis which hit the central coast of Japan in 2019 and utilized the results to predict storm 42 surges. We propose that Pareto optimality can identify good ensemble members that reasonably 43 represent potential worst/minimum storm surge scenarios, meaning no other ensemble members 44 45 can represent better than those. Comprehensive analyses of Pareto members can give forecasters and decision makers a better understanding of how the predicted track, wind intensity, and size of 46 a TC can impact the worst and best storm surge scenarios. This type of analysis is expected to 47 improve the planning of evacuations and the issuing of storm surge warnings. 48

49 **1 Introduction**

50 Since 1737, 29 coastal storm surge events have claimed at least 5,000 people globally. Two 51 of these events happened in the 21st century and ranked as two of the five worst coastal disasters 52 in the running millennium (Needham et al., 2015; Takagi et al., 2022). Rappaport (2014) has shown 53 that 49% of tropical cyclone (TC)-induced deaths are directly attributed to storm surges. Hence, it 54 is crucially important to improve the understanding of storm surge and their associated risk as it is 55 among the deadliest and most destructive natural disasters.

In recent years, forecast services have likely reduced TC-induced deaths relative to historical standards. For example, several countries have already adopted a dynamical TC ensemble prediction system (EPS) to capture forecast uncertainties and reduce sampling errors in the three-

59 dimensional meteorological simulation (Sharma et al., 2022). Numerical weather prediction

centers such as Japan Meteorological Agency (JMA), National Centers for Environmental 60 Prediction in the United States (US), European Centre for Medium-Range Weather Forecasts 61 generate TC track forecasts from their ensemble forecast models and utilize them in their 62 operational settings (Swinbank et al., 2016). Yamaguchi et al. (2015) have shown that EPS can 63 provide skillful guidance of TC genesis forecasts with a forecast lead time extending to two weeks 64 in seven TC basins. Nevertheless, there is a great potential to maximize the use of this EPS not 65 only in TC activity (e.g., track, intensity) forecast but also in forecasting hazards (e.g., storm 66 surge), aiding end users to be prepared better before the dangerous situation (Kobayashi et al., 67 2020; Duc et al., 2021). 68

Titley et al. (2019) have recently conducted a questionnaire survey at operational TC forecast 69 centers worldwide to understand the current and potential use of EPS in operational TC 70 forecasting. They reported that over 90% of respondents used an ensemble forecast for TC track 71 72 forecast, followed by genesis and intensity forecasts. In contrast, less than 10% of surveyed forecasters use ensemble products for hazard (e.g., storm surge) forecasting. Deterministic 73 forecasts are often used for hazard forecasting as it is produced using the best available TC data 74 and unperturbed models. In some cases, ensemble mean (e.g., track and intensity of TC) is used as 75 inputs for hazard forecast to compare the result with the deterministic forecasts, although the full 76 use of EPS in hazard forecasting remains challenging (Titley et al., 2019). A lack of detailed 77 78 analysis of ensemble members (beyond ensemble mean/median analysis) and less technical expertise on ensemble-based hazard forecasts hinder its' application among hazard forecasters. 79 Wilson et al. (2019) reported that a deterministic mindset resulted in tendencies to modify 80 81 understanding of probabilistic concepts when presented with different meteorological variables. Furthermore, local authorities responsible for hazard forecasting avoid EPS information as citizens 82 and emergency managers habitually trust a single forecast only, and they are not sufficiently 83 educated to deal with the probabilistic prediction (Lombardi et al., 2018). These findings highlight 84 that ensemble-based hazard (e.g., storm surge) forecast is unfamiliar in disaster risk management 85 86 communities.

Notwithstanding the challenges mentioned above, ensemble surge prediction system (ESPS) has 87 recently received considerable attention from both the research and operational communities. For 88 89 instances, Flowerdew et al. (2013), Greenslade et al. (2017), and Kristensen et al. (2022) have 90 successfully developed and evaluated the performance of an operational ESPS for United Kingdom, Australia, and Norway, respectively. Along the coastline of Canada, it was found that 91 20-member ESPS could reasonably estimate both the uncertainty in peak surge height and timing 92 of surge events resulting from imperfectly forecast atmospheric conditions six days before (Bernier 93 & Thompson, 2015). A 50-member ensemble simulation of 10 surge events during 2010 in Venice 94 by Mel & Lionello (2014) has shown that the distribution of maximum sea level is acceptably 95 realistic with respect to the deterministic forecast. They also found that the uncertainty became its 96 maximum during storm surge peaks and increased linearly with the forecasting lead time. 97 Although these ensemble simulation studies paved the way for a robust surge hazard assessment 98 over a single forecast-based assessment, they considered ensemble TC forecast information only 99 for developing and evaluating the performance (skill and accuracy) of an ESPS. In addition to 100 quantifying the uncertainty of surge height, ensemble-based storm surge multi-scenario (e.g., 101 worst/optimum case) analysis is equally important, aiding disaster risk managers in evacuation 102 planning (Kohno et al., 2018). 103

To the best of our knowledge, the potential of ensemble TC forecasting for assessing storm surge 104 multi-scenarios has largely been overlooked previously. However, recent developments have seen 105 the introduction of multi-scenario storm surge predictions, such as the worst-case scenario from 106 six typical TC tracks by the JMA (H. Hasegawa et al., 2017) and the maximum storm tide height 107 by the National Hurricane Center in the US (NHC, n.d.). These worst-case scenarios are composite 108 products, representing the maxima among all scenarios. Therefore, it is possible that the worst-109 case values for two adjacent locations may have come from two different ensemble TC track run. 110 Therefore, the users (e.g., emergency managers) cannot understand which forecasted TC track or 111 which combination of forecasted TC meteorological variables (track, intensity, size, translation 112 speed) may trigger the worst surge scenario for a particular location based on a composite product. 113 114 This can make it difficult for decision-makers to determine the appropriate level of storm surge warning and evacuation orders. In addition, storm surge is spatially heterogeneous because of its' 115 dependency on a TC characteristic and coastal geometry. It is entirely plausible that the worst case 116 scenario may not occur everywhere within a forecasted TC threat zone (Islam & Takagi, 2020a, 117 2020b). If the decision makers in cities/tourist districts with highly valuable economies issue a 118 higher warning level without any concrete understating over a worst event, they will inevitably 119 suffer significant economic losses because of false alarming (in case the area has not affected by 120 worst storm surge) and eventually can lower citizens trust over official warning (Sawada et al., 121 2022; Takagi et al., 2018). 122

Here we present Pareto optimality - a novel way of assessing storm surge multi-scenarios based 123 on ensemble TC forecasts. Our approach is more advanced than existing assessments. We 124 employed a multi-objective function to determine possible worst/optimum cases to quantify the 125 hazards in a large region. Our approach involved a comprehensive analysis of Pareto optimal 126 solutions in understanding the combination of forecasted TC meteorological variables - such as 127 track, intensity, size, and translation speed of TC - that could result in the worst/optimum surge 128 scenario. We utilized an extremely large ensemble (=1000 member) forecasts of TC Hagibis that 129 made landfall in central Japan in 2019. Our Pareto-based optimal solutions provide an 130 instantaneous overall assessment of storm surge multi-scenarios without any computational 131 burdens. The proposed method will allow forecasters to predict storm surge multi-scenarios 132 harnessing ensemble TC forecasts efficiently and help emergency responders as means of 133 quantifying surge hazards effectively. 134

135 **2 Data and Methods**

136 2.1 TC Hagibis and ensemble forecast

TC Hagibis in 2019, one of the most destructive and deadliest TC that hit Japan in decades (Shimozono et al., 2020; Ma et al., 2021), has been chosen to demonstrate our multi-scenario storm surge assessment. Hagibis was formed in the western North Pacific Ocean on 2 October 2019 and made landfall in central Japan on 12 October 2019 (around 0900 UTC), as depicted in Figure 1. At the landfall time, its maximum wind speeds sustained at 80 kt. This combined with heavy rainfall, resulted in high storm surges and severe flooding in the area (Shimozono et al., 2020; Ma et al., 2021; JMA, 2021).

The atmospheric ensemble forecasts of TC Hagibis were obtained by running JMA's former operational limited-area model called NHM (non-hydrostatic model; Saito et al., 2006). The integration domain (see Figure S1) had a grid spacing of 5 km consisting of 817×661 horizontal grid points and 50 vertical levels. Boundary conditions were interpolated to the NHM domain from
 JMA's global model forecasts and the forecast perturbations of JMA's operational one-week EPS.

JIVIA's global model forecasts and the forecast perturbations of JMA's operational one-week EPS

Since we used NHM for all forecast members, the only source of uncertainty stemmed from initial 149 conditions. This uncertainty is encapsulated in error covariances of current atmospheric states 150 (analysis error covariances), estimated using a data assimilation system. An ensemble Kalman 151 filter (EnKF) was employed to sample from these error covariances and generate an analysis 152 ensemble. While operational forecast centers generally use around 100 ensemble members, a state-153 of-the-art data assimilation system with 1000 ensemble members, called the four-dimensional 154 variational-ensemble assimilation technique (4DEnVAR), was utilized in this study (Kobayashi et 155 al., 2020). Our 4DEnVAR system only applied horizontal localization, with the horizontal 156 localization length scales derived from the JMA's operational four-dimensional variational 157 assimilation system's climatological horizontal correlation length scales. This helped to remove 158 159 sampling noise in estimating forecast error covariances and maintain the coherent vertical structure between atmospheric fields, which is critical in predicting tropical cyclones. As the ensemble 160 member count was large (=1000), localization was relaxed by retaining vertical correlations and 161 removing horizontal correlations at distant locations (Duc et al., 2021). 162

163 Unlike EnKF, EnVAR solely estimates the means of analysis ensembles and not the analysis 164 ensembles themselves, even though this method heavily relies on forecast ensembles to estimate these means. To solve this issue, a common approach is to run a separate EnKF in parallel to 165 generate analysis ensembles. However, our 4DEnVAR system is unique in that an EnKF is not 166 167 necessary. Instead, the same EnVAR program was used to generate analysis perturbations, as suggested in the context of inflation functions (Duc et al., 2020), where we demonstrated that using 168 quadratic inflation functions implies using the Kalman gain to generate analysis perturbations. 169 Using the same program for analysis means and analysis perturbations is essential because it 170 ensures consistency between the two when the same background error covariance, localization, 171 and observations are utilized in both cases. The assimilation system commenced at 00UTC on 7 172 October 2019, with a 3-hour assimilation cycle and continued until 18:00UTC 10 October 2019. 173 The final analysis ensemble was then used as initial conditions for 39h forecasts with NHM. The 174 assimilation domain was chosen the same as the forecast domain in Figure S1 and we assimilated 175 all routine observations obtained from JMA's database. Here, we opted for a 39h forecast horizon 176 177 because JMA's operational Meso-scale Ensemble Prediction System (MEPS) also generates 39h forecasts at 6-hour intervals (JMA, 2023). 178

179 2.2 Ensemble storm surge forecast

We used storm surge hazard potential index (SSHPI; eq. 1), a statistical model to compute peak 180 storm surge height. While the coastal engineers and ocean modelers are interested in the forecast 181 182 of storm surge hydrograph, most of the decision makers responsible for issuing surge warning and relief measures have a primary interest in the value of predicted peak surge height. The SSHPI 183 uses meteorological variables sensitive to storm surge, including TC intensity (V_{max}), size (radius 184 of 50-kt wind; R_{50}), and translation speed (S). In addition, the SSHPI considers coastal geometry 185 (a = 1 = open coasts and a = -1 = bays), landfall location sensitivity (D_L) , and regional scale 186 bathymetry (L_{30}) . The SSHPI does not incorporate factors associated with wave set-up and 187 astronomic tide to keep the configuration simple. TC Hagibis ensemble forecasts (=1000 member; 188 see Section 2.1), particularly during landfall, was used as meteorological forcing of the SSHPI. 189 We produced corresponding 1000 perturbed surge forecasts with a lead time of 39h. The 190

bathymetry of the target region was obtained from the Japan Oceanographic Data Center (Japan 191 Oceanographic Data Center, 2020). The effectiveness of the SSHPI for predicting peak surge 192 hazard potential was discussed in Islam et al. (2021, 2022). The formulation of the SSHPI is the 193 194 following:

195

$$SSHPI = \left(\frac{V_{max}}{V_{ref}}\right)^2 \left(\frac{R_{50}}{R_{ref}}\right) \left(\frac{S}{S_{ref}}\right)^a \left(\frac{L_{30}}{L_*}\right) \left(D_L\right) \tag{1}$$

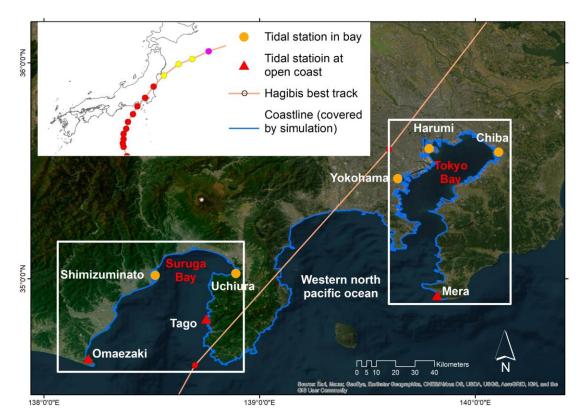
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$$197 \qquad \frac{R_{50}}{R_{ref}} = \begin{cases} 1.5 \quad if \ \frac{R_{50}}{R_{ref}} \ge 1.5\\ \frac{R_{50}}{R_{ref}} \ if \ 0.5 < \frac{R_{50}}{R_{ref}} < 1.5 \ ; \ (\frac{S}{S_{ref}})^a = \begin{cases} 1.5 \quad if \ (\frac{S}{S_{ref}})^a \ge 1.5\\ (\frac{S}{S_{ref}})^a \ if \ 0.5 < (\frac{S}{S_{ref}})^a < 1.5 \ ; \ \frac{L_{30}}{L_*} = \begin{cases} \frac{L_{30}}{L_*}, \ if \ \frac{L_{30}}{L_*} \ge 1\\ 1, \ if \ \frac{L_{30}}{L_*} \le 1\end{cases} \\ 1, \ if \ \frac{L_{30}}{L_*} \le 1\end{cases} \\ 0.5 \quad if \ (\frac{S}{S_{ref}})^a \le 0.5\end{cases}$$

$$D_L = \begin{cases} 1 \quad if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x \le 20\\ 0R \end{cases} \\ if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x \le 20\\ 1 - \frac{0.03(x - 20)}{20} \ if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x > 20\\ 1 - \frac{0.05(x - 10)}{10} \ if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x > 10 \end{cases}$$

 V_{ref} , R_{ref} , and S_{ref} , are reference constants as follows: 50-kt equivalents of the tropical storm 199 category, 95 NM (historical mean R₅₀ at the time of landfall in Japan mainland), and 35 km/h 200 (historical mean S at the time of landfall in Japan), respectively (Islam et al., 2021). L_{30} is the 201 horizontal distance (km) between the shoreline and the 30-m depth contour. L* was chosen to be 202 10 km. D_L is defined by different expressions depending on the surge estimated points' (e.g., tidal 203 station) position (right/left) respective to the TC track and horizontal distance (x in km) between 204 the TC landfall location and a surge estimated point. Compared to V_{max} , the upper and lower bounds 205 of R_{50} , S, and L_{30} in eq. 1 restrict their contribution in generating surge hazards and, thus, prevents 206 207 discrete jumps in the SSHPI.

Figure 1 shows a storm surge modeling domain and the position of tide gauges used for validating 208 surge model and predicting surge hazards in this study. There are two domains, covering Tokyo 209 Bay and Suruga Bay individually. Each domain has tide gauges located both in inner bays and 210 open coasts. It should be noted that the tide gauges chosen for this study are the only stations that 211 possess recorded (historical) storm surge data, which is kept by JMA (JMA, 2022) and Japan Coast 212 Guard (Japan Oceanography Data Center, 2021). The empirical relationship for expected storm 213 surge in each tide gauge was determined in Islam et al. (2021, 2022) by drawing a line of best fit 214 215 through the historical surge data and the SSHPI and thus, used for the surge forecasts in this study. 216



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Figure 1. Domain of the storm surge forecasts model and the locations of the tide gauges used for

219 model validation and surge forecasts.

220 2.3 Pareto optimality and assessing storm surge multi-scenarios

It is unrealistic to anticipate a "nice" forecast scenario that accurately predicts the exact intensity 221 of a hazard at all locations within a given domain for a particular condition (e.g., worst/optimum). 222 223 An improved forecast at one location is usually accompanied by a deterioration of forecast at another location and vice versa. The best we can do is to quantify the trade-off between different 224 objectives. Here, we conducted multi-objective optimization to select ensemble forecast members 225 (among 1000 ensemble forecasts; see Section 2.1 and 2.2) that reasonably characterize the 226 potential worst/optimum storm surge case for a particular location (e.g., Tokyo Bay) by computing 227 the Pareto frontier. The Pareto frontier captures the trade-offs between objectives. It is the set of 228 all Pareto-optimum solutions where a single Pareto optimal solution denotes a solution that is not 229 dominated by any other solution (Kochenderfer & Wheeler, 2019). 230

In this study, we analyzed a subset of 1000 forecasted surge scenarios for each tide gauge in a 231 specific domain, referred to as solution z. Each scenario is evaluated based on d objectives, 232 represented by the values $y^{1}(z), y^{2}(z), \dots, y^{d}(z)$. As an example, we considered a scenario in Tokyo 233 Bay where the objectives are to maximize the forecasted peak surge at four tide gauges: Harumi, 234 Chiba, Yokohama, and Mera (as shown in Figure 1). This objective function considers the potential 235 worst-case scenario in Tokyo Bay for a set of 1000 ensemble surge forecasts. We compared the 236 given two solutions z and z', if for every objective i, $y^i(z) \ge y^i(z')$ and the strict inequality holds 237 for at least for one objective, we considered that solution z dominates z'. In other words, if one 238 solution (e.g., ensemble member no. 10) provides 160 cm, 155 cm, 140 cm, 110 cm of forecasted 239 peak surge in Harumi, Chiba, Yokohama, and Mera, respectively, but another solution (e.g., 240

ensemble member no. 20) yields 160 cm, 155 cm, 140 cm, 90 cm of forecasted peak surge for the 241 same tide gauges, then the solution given by ensemble member no. 10 dominates the solution 242 provided by ensemble member no. 20. This is because a smaller storm surge (90 cm) in Mera is 243 undesirable as stated in the objective function. A similar objective function but minimizing the 244 forecasted peak surge at four tide gauges: Harumi, Chiba, Yokohama, and Mera (as shown in 245 Figure 1) was considered to determine ensemble forecasts member that characterizes a potential 246 optimum surge case. In order to select the most appropriate ensemble TC member that may cause 247 the worst surge case in the inner bay, but also results in the optimum surge at the open coast, we 248 further assume that objectives are to be maximized for the inner bay tide gauges (Harumi, Chiba, 249 and Yokohama), but minimized for open coast tide gauges (Mera) at the same time. 250

and Tokonama), but minimized for open coast tide gauges (Mera) at the same time.

- The details of the implementation of the algorithm used here are described in Tommy (2021). This
- algorithm can compute the Pareto frontier for four objectives within a minute, meaning the runtime
- should be acceptable to any operational hazard forecast settings.

254 **3 Results**

- 255 3.1 Model evaluation
- 256 3.1.1 TC ensemble forecasts validation

From the forecasts of atmospheric fields given by NHM, TC tracks and intensities were detected. Here, TC centers are defined as the average of the mean sea level pressure minima, geopotentials at 850 hPa and 700 hPa. Figure 2(b) shows 1000 track forecasts generated from 1000 initial

- conditions obtained from the 4DEnVAR data assimilation system, along with the ensemble mean
- 261 forecast, control forecast, and best track. The ellipses in the figure illustrate uncertainties of the
- 262 TC centers, which are determined by the forecast error covariances of the TC centers. The arrows
- show the distance errors between the observations (i.e., the best track) and the ensemble mean. For
- comparison, the operational 20-member JMA ensemble forecast (MEPS) is included in Figure 2a.
- As shown in Figure 2, 4DEnVAR outperforms MEPS in terms of track forecasts and its ensemble
- 266 mean is almost identical to the best track with only minor distance errors.

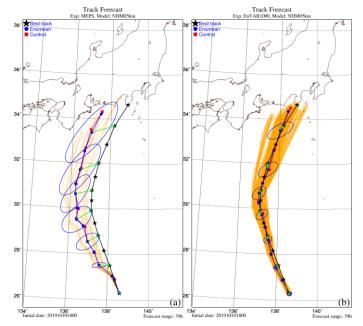


Figure 2. A comparison of 39h (at 1800 UTC on 10 October 2019) ensemble track forecasts for TC Hagibis issued by (a) the operational ensemble prediction system MEPS of JMA and (b) the 4DEnVAR data assimilation system. The ellipses represent forecast error covariances of the TC centers. The arrows denote the distance errors between the ensemble mean and best track.

Figure 3 presents the forecasted intensity of TC Hagibis as indicated by its central pressure. It is 272 evident from the figure that the 4DEnVAR (Figure 3b) surpasses JMA's operational MEPS (Figure 273 3a) in predicting the intensity. Even though both ensemble forecasts show overestimation of 274 intensity, the tendency of overestimation becomes more apparent with increasing forecast ranges 275 in MEPS (Figure 3a). Despite having a smaller number of ensemble members, MEPS exhibits 276 greater uncertainty in intensity forecast as compared to 4DEnVAR. This can be attributed to the 277 fact that JMA's operational MEPS employs singular vectors to generate initial conditions for 278 ensemble members, which maximizes their spread (JMA, 2023). 279

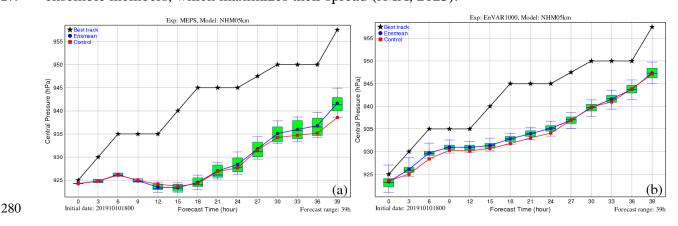
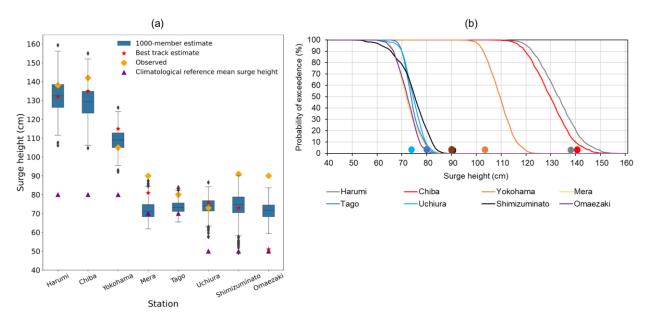


Figure 3. A comparison of 39h (at 1800 UTC on 10 October 2019) ensemble intensity forecast for TC Hagibis issued by (a) the operational ensemble prediction system MEPS of JMA and (b) the 4DEnVAR data assimilation system. The distributions of ensemble intensities are represented as box-and-whisker plots.

285 3.1.2 Storm surge ensemble forecasts validation

A comparison of the forecasted and measured peak storm surge (Japan Oceanography Data Center, 286 2021; JMA, 2022) at eight different tide gauges is shown in Figure 4a. It is noted that a significant 287 deviation from the climatological mean surge height is observed at all stations during TC Hagibis. 288 We evaluate JMA best track (JMA, 2021) as an ideal meteorological forcing input as well as our 289 39h ensemble TC forecasts. Both ensemble median forecasts and best track estimates 290 systematically underestimate the observed peak levels. Nevertheless, the observed peak surge 291 values are enveloped by the full ensemble of the forecasts in most tide gauges, implying that the 292 ensemble spread is large enough to represent the uncertainty in the prediction. The average mean 293 absolute error of the eight stations is 11.3 cm. In the case of Mera and Omaezaki, wave set-up is 294 often the dominant driver for generating storm surges (Islam et al., 2018, 2022), which is not 295 considered in the SSHPI. Therefore, the ensemble surge forecast underestimated observed surges. 296



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Figure 4. (a) A comparison of the 39h (at 1800 UTC on 10 October 2019) ensemble peak surge forecasts and the observed peak surge height for TC Hagibis at eight tide gauges. (b) Probability of exceeding (Y-axis) for a given storm surge threshold (X-axis) during TC Hagibis landfall time (at 0900 UTC, 12 October 2019). It was estimated by fitting ensemble forecasts empirically. A circle that shares the same color as a line represents the peak surge height recorded at a specific tide gauge.

The probability of surpassing a specific surge threshold during the landfall time of TC Hagibis is 304 illustrated in Figure 4b, as determined by the 39h ESPS. The observed peak surge levels are within 305 the predicted range, except for Mera and Omaezaki. For example, the probability of surpassing the 306 observed peak surge for Harumi (138 cm) is 26.4%. In general, Figure 4 indicates that the SSHPI 307 and its corresponding 39h peak surge forecasts are comparable in quality to those produced by 308 numerical surge models such as Liu et al. (2021). The latter study reported an RMSE of ~10 cm 309 when predicting the maximum total water level in Tokyo Bay during TC Hagibis with a 72h 310 forecast horizon, using atmospheric forcing fields from the Global Forecast System and a 311 hydrodynamic model known as the Semi-implicit Cross-scale Hydroscience Integrated System, 312 which has a nearshore resolution of ~150 m. 313

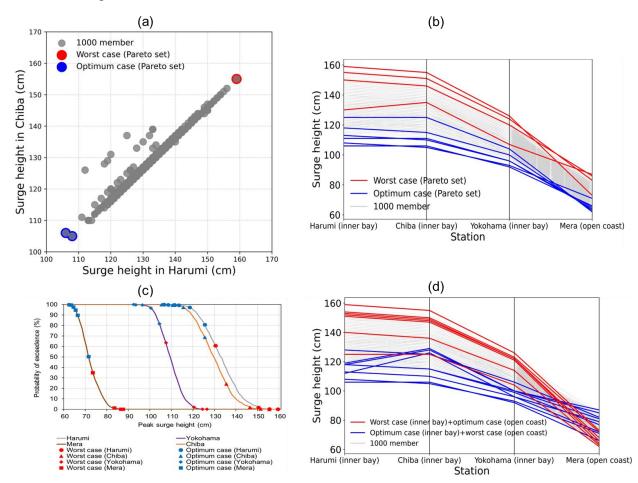
314 3.2 Multi-scenario analysis

315 3.2.1 Pareto optimal multi-scenarios

316 The Pareto-optimal frontier, as shown in Figure 5, illustrates a group of solutions that depict the forecasted potential worst and optimum storm surge scenarios for TC Hagibis in Tokyo Bay. The 317 two-dimensional Pareto frontiers (Figure 5a), allow for a straightforward evaluation of trade-offs 318 among the forecasted peak storm surge levels. The results identify the best one or two members 319 from the 1000 TC forecasts to represent the potential worst (Harumi, Chiba: ~155 cm) or optimum 320 (Harumi, Chiba: ~106 cm) surge scenario in the inner Tokyo Bay. Here, the tide gauges (Harumi, 321 322 Chiba) possess similar coastal geometry features, including bathymetry, and are situated in close proximity to each other (Figure 1). Therefore, the predicted surge response (from 1000 TC 323

forecasts; Figure 5a) between them is almost linear.

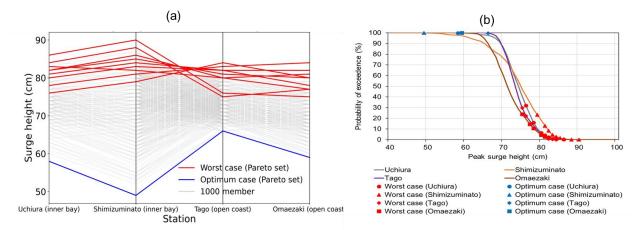
Figure 5b reveals a diversity of trade-off surge outcomes in the Pareto frontier for Tokyo Bay. It 325 includes four Pareto frontiers for worst-surge scenarios and six Pareto frontiers for optimum surge 326 scenarios. This diversity is due to the emergence of trade-offs among tide gauges with distinct 327 coastal geometry characteristics, including bathymetry. For instance, some of the identified most 328 severe surge scenarios for open coastlines in Tokyo Bay do not result in high surge levels in the 329 inner bay. Specifically, a Pareto optimal solution in Figure 5b predicts that Mera would experience 330 the worst surge levels of 90 cm (<1% exceedance probability; Figure 5c), while Harumi and Chiba 331 would experience approximately 135 cm (>25% exceedance probability; Figure 5c) of highest 332 surge levels under the same scenario. This storm surge level (~135 cm) in Harumi and Chiba is 333 substantially less than the worst surge levels (~155 cm) predicted by other optimal solutions. 334 Additionally, the surge intensity may vary across Tokyo Bay, depending on the characteristics of 335 the approaching TC and the impact can be much more severe in some places compared to others. 336 For example, several Pareto optimal solutions shown in Figure 5d predict that the inner Tokyo Bay 337 such as Harumi would witness surge levels higher than 150 cm, while it would keep as minimum 338 as 70 cm along the open coastline (e.g., Mera). Owing to such surge incongruence among the 339 coastal locations, creating multiple scenarios for different coastal places can lead to multiple 340 optimal solutions, as seen in Figures 5b and 5d. This emphasizes the importance of considering 341 multiple scenarios when issuing warnings and assessing the risks posed by extreme weather events 342 like storm surges. 343



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Figure 5. Forecasted Pareto optimal multi-scenarios due to TC Hagibis, apply for (a) Harumi and 345 Chiba in Tokyo Bay [objective function (red dot): max surge height in Harumi and Chiba; 346 objective function (blue dots): min surge height in Harumi and Chiba]; (b) Tokyo Bay (Harumi, 347 Chiba, Yokohama and Mera) using the parallel coordinate plot [objective function (red lines): max 348 surge height in Harumi, Chiba, Yokohama, and Mera; objective function (blue lines): min surge 349 height in Harumi, Chiba, Yokohama, and Mera]; (c) Probability of exceeding (Y-axis) a storm 350 surge threshold (X-axis) determined from each Pareto optimal solution in Figure 5b, at TC Hagibis 351 landfall time (at 0900 UTC, 12 October 2019). It was estimated by fitting ensemble forecasts 352 empirically; (d) Parallel coordinate plot including forecasted Pareto optimal solutions where 353 forecasted peak surge intensity varies across Tokyo Bay [objective function (red lines): max surge 354 355 height in Harumi, Chiba, and Yokohama + min surge height in Mera; objective function (blue lines): min surge height in Harumi, Chiba, and Yokohama + max surge height in Mera]. 356

357 The results displayed in Figure 6a are comparable to those in Figure 5b but for Suruga Bay. It includes nine Pareto frontiers for worst-surge scenarios and one Pareto frontier for optimum surge 358 scenario. Figure 6a predicts that all selected tide gauges would experience the worst surge levels 359 of ~85 cm (<1% exceedance probability; Figure 6b), while the minimum surge height would be 360 ~55 cm (>99% exceedance probability; Figure 6b). It is noteworthy that only one worst-surge 361 scenario is found to be shared by both the Pareto frontiers of Tokyo Bay (Figure 5b) and Suruga 362 Bay (Figure 6a), which represents a potential worst-case across the Japanese coastline during TC 363 Hagibis. This indicates that among 1000 TC ensemble forecasts, a particular TC ensemble member 364 has the potential to bring severe surge levels to both Bays. The reason behind this commonality 365 will be discussed in Section 3.2.2. 366



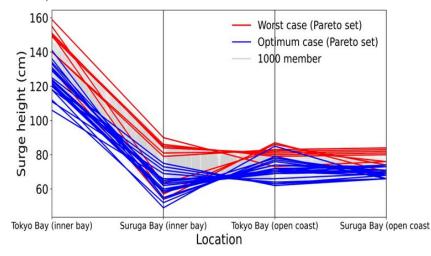
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Figure 6. (a) Parallel coordinate plot with Pareto optimal multi-scenarios for Suruga Bay, at TC Hagibis landfall time (at 0900 UTC, 12 October 2019) [objective function (red lines): max surge height in Uchiura, Shimizuminato, Tago, and Omaezaki; objective function (blue lines): min surge height in Uchiura, Shimizuminato, Tago, and Omaezaki]; (b) Probability of exceeding (Y-axis) a storm surge threshold (X-axis) determined from each Pareto optimal solution in Figure 6a. It was estimated by fitting ensemble forecasts empirically.

In addition to Pareto Frontiers illustrated in Figures 5b and 6a, we further noticed many distinct trade-offs when both bays were considered together (Figure 7). For instance, we incorporated the representative tide gauges for each category of coastal geometry, such as inner bay (Harumi in

Tokyo Bay and Shimizuminato in Suruga Bay) and open coast (Mera in Tokyo Bay and Tago in

Suruga Bay) in the objective function. This resulted in a large number of Pareto frontiers (worst 378 case: nine; optimum case: twenty-two), with multiple overlapping solutions between worst and 379 optimum cases. Thus, 33% of Pareto optimal solutions (red lines) predict that inner Tokyo Bay 380 will experience the worst surge levels of ~150 cm, while inner Suruga Bay will witness ~60 cm, 381 equivalent to minimum surge cases predicted by 41% of Pareto optimal solutions (blue lines; 382 Figure 7). Although we maximize the potential of the large (i.e., 1000) ensemble forecasts, our 383 proposed method is also found to be useful with the small ensemble size. We repeated the same 384 analysis with 36 ensemble size. Although some worst and minimum storm surge scenarios were 385 missed, a meaningful set of Pareto optimal solutions was still obtained (see Figure S2 in the 386 supplementary section). 387



388

Figure 7. Parallel coordinate plot with Pareto optimal multi-scenarios determined for both Tokyo
 Bay and Suruga Bay [objective function (red lines): max surge height in Harumi, Shimizuminato,
 Mera, and Tago; objective function (blue lines): min surge height in Harumi, Shimizuminato,

392 Mera, and Tago]

393 3.2.2 TC track and meteorological variables analysis of Pareto optimal solutions

It would be interesting to analyze the tracks and associated meteorological variables of the 394 identified Pareto optimal solutions in Figures 5b and 6a. For example, Figure 8 reveals the strong 395 sensitivity of the storm surge scenarios to the landfall location of TC Hagibis, leading to 396 contrasting Pareto optimal solutions for Tokyo Bay (Figure 5b) and Suruga Bay (Figure 6a). It 397 also demonstrates that the forecasted TC tracks for worst and optimum surge scenarios are 398 significantly different from one other in both bays. For example, TC tracks (red lines; Figure 8a) 399 that run parallel to the longitudinal axis of Tokyo Bay and pass over it would result in severe storm 400 surges than TCs (blue lines; Figure 8a) that would travel 100 km or more to the west of the axis. 401 Prior to landfall, under the worst case, easterly wind (Figure 9a) is forecasted to cause a buildup 402 of water on the west coast (e.g., Yokohama) and initial draw-down in the north-eastern end of the 403 Tokyo Bay (e.g., Chiba). Later, a surge level difference (0.6–0.8 m; Figure 5b) between the inner 404 and lower ends of the bay is projected to occur (due to strong southerly winds) during the peak 405 storm surge at the inner bays. During TC makes landfall under the optimum case (Figure 8a), the 406 destructive right-side semicircle of the TC (Figure 9f) will interact with the vast land area rather 407 than the ocean water, leading to less water being pushed towards Tokyo Bay (Figure 5b). 408

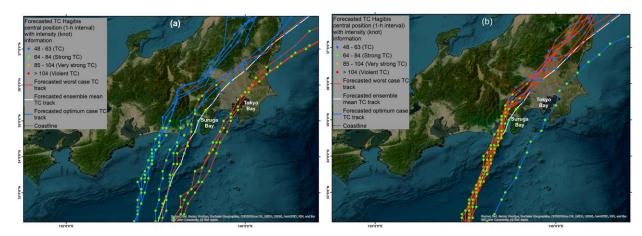
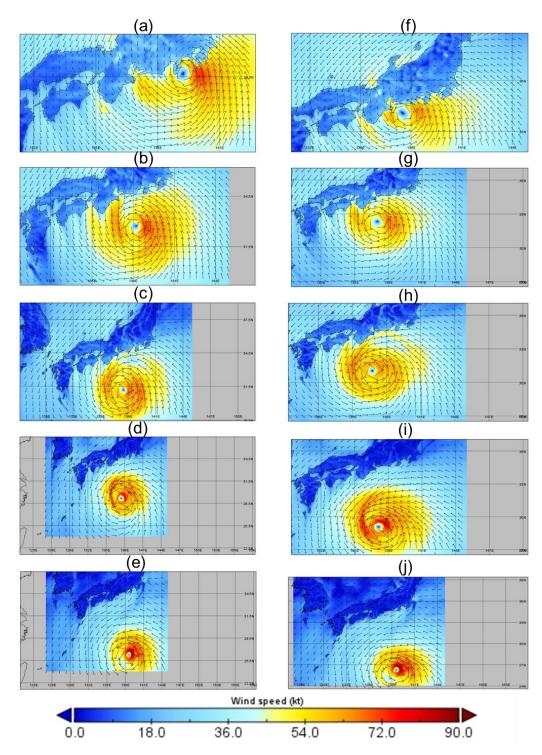




Figure 8. Forecasted TC Hagibis track in 39h lead time, respective to each Pareto frontier
determined for (a) Tokyo Bay (as shown in Figure 5b); (b) Suruga Bay (as shown in Figure 6a).
Red, blue, and white lines correspond to the forecasted worst, optimum, and ensemble mean TC
track.

Notably, the optimized ensemble TC members for minimum surge scenarios in Tokyo Bay are forecasted to be stronger and larger than the members belonging to the worst surge cases until 24 hours prior to landfall, despite being centered in the same location. This is evident in Figure 9d, 9e, as opposed to Figure 9i, 9j. Despite both sets (worst and minimum) of optimized ensemble TC members weaken as they approach the mainland of Japan (Figure 9c, 9h and Figure 10), the worst TC members intensify by 7-kt (V_{max}) and remain large (R_{50} : ~120 NM) in the last 12 hours before landfall (Figure 9a, 9b and Figure 10). This large swath of strong winds is forecasted to affect a

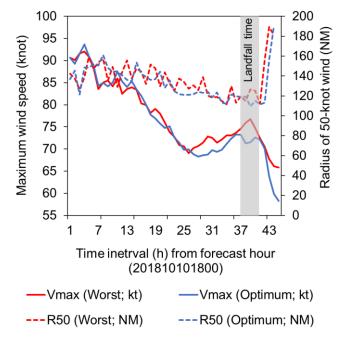
- 420 greater sea area and induce a motion in a greater quantity of water in Tokyo Bay.
- 422



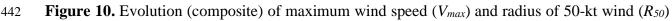
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Figure 9. Forecasted composite 10 m wind field (kt-vectors), generated from optimized ensemble
TC members for worst surge scenarios in Tokyo Bay (as shown in Figure 5b) during (a) landfall
time (at 0900 UTC, 12 October 2019); (b) 6-h before landfall (at 0300 UTC, 12 October 2019);
(c) 12-h before landfall (at 2100 UTC, 11 October 2019); (d) 24-h before landfall (at 0900 UTC,
11 October 2019); (e) 39-h before landfall (at 1800 UTC, 10 October 2019); (f-j) same as (a-e) but

- 429 generated from optimized ensemble TC members for minimum surge scenarios in Tokyo Bay (as
- 430 shown in 5b).
- 431 Figure 8a highlights that in addition to TC tracks that pass directly over Tokyo Bay, one particular
- 432 worst TC forecast (red line) makes landfall around 70 km west of the longitudinal axis of Tokyo
- Bay. This specific track is forecasted to cause severe storm surges in both Tokyo Bay (Figure 5b)
- and Suruga Bay (Figure 6a). This particular ensemble forecast has a wider range of intense winds
- 435 (R_{50} : ~140 NM) across a larger area and a slower movement speed (S: ~32 km/h), despite having 436 a similar landfall wind intensity (V_{max} : ~75-kt) compared to other worst-case TC forecasts. This
- 437 unique phenomenon corroborates earlier numerical analyses that propose the likelihood of a severe
- 438 storm surge scenario in the upper-bay region when a large and intense TC moves slowly, parallel
- to the longitudinal axis of Tokyo Bay, after making landfall 25 km southwest. (Islam & Takagi,
- 440 2020b).



441



for worst and optimum TC forecasts in Tokyo Bay (as shown in Figure 5b) in 39-h lead time.

444 4 Conclusions and discussion

The application of ensemble TC forecasting in hazard prediction, such as storm surge, has been 445 greatly overlooked despite its use in forecasting TC track, intensity, and genesis. Enhanced 446 analysis can unlock and maximize the benefit of ensemble forecasting. Here, we proposed Pareto 447 optimality – a novel and practical way to identify potential ensemble TC (Hagibis) forecast from 448 449 an extremely large ensemble (=1000 member) that can effectively assess storm surge multiscenarios, including possible worst and optimum cases for a coastal location. The variability in 450 storm surge intensity across the coastline makes it challenging for decision-makers to plan 451 effective evacuation measures. To address this, we have demonstrated that a diversity of trade-off 452 surge outcomes among coastal places can be identified by choosing the Pareto optimized forecasts. 453 The in-depth evaluation of Pareto optimal solutions can shed light on how meteorological variables 454 such as track, intensity, and size of TCs influence the worst and optimum surge scenarios, which 455

456 are not well understood by emergency managers using current multi-scenario assessment methods457 (such as those used by JMA and NHC).

The significance of evaluating trade-offs based on Pareto optimization has long been 458 acknowledged in the context of sustainable development goals and management of ecosystem 459 services (Flecker et al., 2022). However, its application in disaster-risk communities is noticeably 460 lacking. During a coastal storm surge event, effective evacuation planning and warning issuance 461 involve multi-criteria problems such as storm surge intensity, coastal population vulnerability, and 462 available evacuation resources. Traditionally, this decision-making process has relied on a hazard 463 map, which typically depicts the severity of the predicted storm surge (e.g., exceeding a critical 464 surge height; J. Hasegawa et al., 2017). However, this approach does not fully capture the diversity 465 of potential storm surge scenarios across the coastlines, which can lead to ineffective evacuation 466 planning. A recent TC Fani in 2019 serves as evidence to support this statement. TC Fani struck 467 the southeastern part of India, approximately 450 km from the southwest coast of Bangladesh, as 468 a Category 4 TC (in Saffir-Simpson Hurricane Wind Scale). Prior to TC Fani reaching the 469 Bangladesh coast as a tropical storm, the Bangladesh Meteorological Department (BMD) issued 470 'danger' signal number seven (out of ten), which led to the evacuation of one million people 471 (Bangladesh Meteorological Department, 2021). Later, the catastrophe, such as storm surge level 472 (~1 m), did not hit the danger level as anticipated. BMD's evacuation order for the entire southwest 473 474 coast was not based on a specific surge scenario (e.g., worst case) and forecasted meteorological conditions associated with it, leading to an excessive number of evacuees. Such a false alarm 475 demotivated people to seek shelter when a Category 2 TC Amphan caused ~2.75 m storm surges 476 and claimed at least 26 lives in 2020, despite the issuance of 'great danger' signal number ten 477 (Raju, 2019; ReliefWeb, 2021; Alam et al., 2023). It seems that BMD took a safer and conservative 478 decision during TC Fani by issuing signal no. 7, nevertheless, this cannot be considered effective 479 decision-making. While such a complex decision-making process can certainly be improved by 480 quantifying the uncertainty through an ensemble multi-scenario forecast, incorporating Pareto 481 optimality can further maximize the benefits of it. 482

Pareto optimal solution provides an effective first filter to identify ensemble multi-scenario surge 483 forecasts. This information can be presented visually to enhance the understanding of the 484 uncertainty in the forecast. The median of Pareto optimal solutions could be utilized given a series 485 of worst/minimum surge estimations for a specific location by several ensemble members. For 486 example, Figure 5b identifies four worst scenarios for Harumi in Tokyo Bay where the median 487 surge level is 150 cm. Although we stress the importance of diversity in trade-offs surge outcomes, 488 a certain scenario (e.g., median of Pareto optimal solutions) can be given more weight depending 489 on the values of society and decision-makers. In the decision-making process, a user-defined 490 acceptable level of uncertainty or reference surge height (e.g., 25-year return period of surge) can 491 be set for a specific location (e.g., Harumi). The forecaster can then determine if the height of the 492 Pareto optimal solution exceeds this acceptable level. Subsequently, a relevant warning signal can 493 be issued in a forecast horizon (e.g., 39-h lead time). The warning signal can be tailored to a 494 specific location, if a diversity in the trade-off between surge outcomes exists among Pareto 495 frontiers. For example, Pareto optimal solutions in Figure 5b predicted that TC Hagibis will bring 496 worst surge level as maximum as 160 cm in the inner Tokyo Bay, requiring the issuance of an 497 emergency warning, closing flood gates, and large-scale evacuation of the coastal population 498 living below the storm surge height. On the other hand, those living along open coastlines are 499 advised to stay indoors as the predicted worst surge level (90 cm) does not meet the criteria for 500 issuing an emergency warning. Once this forecast becomes available, decision-makers (e.g., 501

emergency managers) can start evacuation planning based on the forecasted worst TC track, wind 502 intensity, and peak surge height. For instance, TC track that forecasted to bring severe storm surges 503 of 160 cm in the inner Tokyo Bay, would make landfall 70 km west of the central bay axis with a 504 landfall V_{max} of 75 kt, which is stronger by 11-kt from the historical mean (64-kt; (Islam et al., 505 2022)). Furthermore, it is projected to be twice as large as the historical average (65 NM) and 506 move at a slower speed by 9 km/h compared to the average translation speed (41 km/h) in Tokyo 507 Bay. The forecasted landfall location and meteorological conditions of the worst TC indicate that 508 Tokyo Bay would be situated in the destructive right-side semicircle of the TC track, resulting in 509 prolonged exposure to severe storm surges and strong winds. Emergency managers can utilize this 510 information to disseminate surge warnings to residents and commence evacuation procedures with 511 a 39-hour lead time. This evacuation can be done by dividing coastal regions into different zones 512 depending on their vulnerability. Although disaster planning is not so straightforward as explained 513 here, our proposed ensemble-based storm surge multi-scenario analysis is expected to motivate 514 forecasters and risk management practitioners to explore new ways to assess storm surge hazards 515 and reduce the associated risk. 516

517 Finally, we acknowledge that this study focuses exclusively on peak surge height while 518 determining total sea water level that includes the influence of astronomic tide, wave set-up, and 519 river discharge are also critical and can be done utilizing a full physical numerical model. 520 Furthermore, several algorithms are currently available to determine a Pareto frontier. We 521 encourage researchers from multiple disciplines to build on our approach to help us reach an 522 improved understanding of Pareto optimality based multi-scenario analysis.

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530 **Open Research**

Observed surge data be downloaded from the JMA 531 storm can (https://www.data.jma.go.jp/kaiyou/db/tide/genbo/index.php) and JODC 532 (https://jdoss1.jodc.go.jp/vpage/tide.html) websites. Predicted tide data can be obtained from the 533 JMA (https://www.data.jma.go.jp/kaiyou/db/tide/suisan/index.php) website. TC best track data 534 can be derived from the JMA (https://www.jma.go.jp/jma/jma-eng/jma-center/rsmc-hp-pub-535 eg/trackarchives.html) website. Ensemble forecast data may be available upon request. 536

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1	Assessing Storm Surge Multi-Scenarios based on Ensemble Tropical Cyclone
2 3	Forecasting
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9	Key Points:
10 11	• The potential of ensemble tropical cyclone forecasting for assessing storm surge multi- scenarios is shown.
12 13	• Pareto optimized solutions from an ensemble storm surge forecast can efficiently illustrate potential worst and minimum storm surge scenarios.
14 15 16	• Analyses of meteorological variables of ensemble members in Pareto frontiers help understand the impact of a tropical cyclone on predicted storm surge multi-scenarios.

17 Abstract

Ensemble forecasting is a promising tool to aid in making informed decisions against risks of 18 coastal storm surges. Although tropical cyclone (TC) ensemble forecasts are commonly used in 19 operational numerical weather prediction systems, their potential for disaster prediction has not 20 been maximized. Here we present a novel, efficient, and practical method to utilize a large 21 22 ensemble forecast of 1000 members to analyze storm surge scenarios toward effective decision making such as evacuation planning and issuing surge warnings. We perform the simulation of TC 23 Hagibis (2019) using the Japan Meteorological Agency's (JMA) non-hydrostatic model. The 24 simulated atmospheric predictions were utilized as inputs for a statistical surge model named the 25 Storm Surge Hazard Potential Index (SSHPI) to estimate peak surge heights along the central coast 26 of Japan. We show that Pareto optimized solutions from an ensemble storm surge forecast can 27 describe potential worst (maximum) and optimum (minimum) storm surge scenarios while 28 29 exemplifying a diversity of trade-off surge outcomes among different coastal places. For example, some of the Pareto optimized solutions that illustrate worst surge scenarios for inner bay locations 30 are not necessarily accountable for bringing severe surge cases in open coasts. We further 31 emphasize that an in-depth evaluation of Pareto optimal solutions can shed light on how 32 meteorological variables such as track, intensity, and size of TCs influence the worst and optimum 33 surge scenarios, which is not clearly quantified in current multi-scenario assessment methods such 34 35 as those used by JMA/National Hurricane Center in the United States.

36 Plain Language Summary

Ensemble forecasting generates multiple predictions of a weather event with various possible 37 outcomes based on varying initial conditions, model parameters, and physics. The potential of 38 ensemble tropical cyclone (TC) forecasting for assessing storm surge multi-scenarios has largely 39 been overlooked previously. Enhanced analysis can unlock and maximize the benefit of ensemble 40 forecasting. This study simulated an extremely large ensemble (=1000 members) to reforecast past 41 TC Hagibis which hit the central coast of Japan in 2019 and utilized the results to predict storm 42 surges. We propose that Pareto optimality can identify good ensemble members that reasonably 43 represent potential worst/minimum storm surge scenarios, meaning no other ensemble members 44 45 can represent better than those. Comprehensive analyses of Pareto members can give forecasters and decision makers a better understanding of how the predicted track, wind intensity, and size of 46 a TC can impact the worst and best storm surge scenarios. This type of analysis is expected to 47 improve the planning of evacuations and the issuing of storm surge warnings. 48

49 **1 Introduction**

50 Since 1737, 29 coastal storm surge events have claimed at least 5,000 people globally. Two 51 of these events happened in the 21st century and ranked as two of the five worst coastal disasters 52 in the running millennium (Needham et al., 2015; Takagi et al., 2022). Rappaport (2014) has shown 53 that 49% of tropical cyclone (TC)-induced deaths are directly attributed to storm surges. Hence, it 54 is crucially important to improve the understanding of storm surge and their associated risk as it is 55 among the deadliest and most destructive natural disasters.

In recent years, forecast services have likely reduced TC-induced deaths relative to historical standards. For example, several countries have already adopted a dynamical TC ensemble prediction system (EPS) to capture forecast uncertainties and reduce sampling errors in the three-

59 dimensional meteorological simulation (Sharma et al., 2022). Numerical weather prediction

centers such as Japan Meteorological Agency (JMA), National Centers for Environmental 60 Prediction in the United States (US), European Centre for Medium-Range Weather Forecasts 61 generate TC track forecasts from their ensemble forecast models and utilize them in their 62 operational settings (Swinbank et al., 2016). Yamaguchi et al. (2015) have shown that EPS can 63 provide skillful guidance of TC genesis forecasts with a forecast lead time extending to two weeks 64 in seven TC basins. Nevertheless, there is a great potential to maximize the use of this EPS not 65 only in TC activity (e.g., track, intensity) forecast but also in forecasting hazards (e.g., storm 66 surge), aiding end users to be prepared better before the dangerous situation (Kobayashi et al., 67 2020; Duc et al., 2021). 68

Titley et al. (2019) have recently conducted a questionnaire survey at operational TC forecast 69 centers worldwide to understand the current and potential use of EPS in operational TC 70 forecasting. They reported that over 90% of respondents used an ensemble forecast for TC track 71 72 forecast, followed by genesis and intensity forecasts. In contrast, less than 10% of surveyed forecasters use ensemble products for hazard (e.g., storm surge) forecasting. Deterministic 73 forecasts are often used for hazard forecasting as it is produced using the best available TC data 74 and unperturbed models. In some cases, ensemble mean (e.g., track and intensity of TC) is used as 75 inputs for hazard forecast to compare the result with the deterministic forecasts, although the full 76 use of EPS in hazard forecasting remains challenging (Titley et al., 2019). A lack of detailed 77 78 analysis of ensemble members (beyond ensemble mean/median analysis) and less technical expertise on ensemble-based hazard forecasts hinder its' application among hazard forecasters. 79 Wilson et al. (2019) reported that a deterministic mindset resulted in tendencies to modify 80 81 understanding of probabilistic concepts when presented with different meteorological variables. Furthermore, local authorities responsible for hazard forecasting avoid EPS information as citizens 82 and emergency managers habitually trust a single forecast only, and they are not sufficiently 83 educated to deal with the probabilistic prediction (Lombardi et al., 2018). These findings highlight 84 that ensemble-based hazard (e.g., storm surge) forecast is unfamiliar in disaster risk management 85 86 communities.

Notwithstanding the challenges mentioned above, ensemble surge prediction system (ESPS) has 87 recently received considerable attention from both the research and operational communities. For 88 89 instances, Flowerdew et al. (2013), Greenslade et al. (2017), and Kristensen et al. (2022) have 90 successfully developed and evaluated the performance of an operational ESPS for United Kingdom, Australia, and Norway, respectively. Along the coastline of Canada, it was found that 91 20-member ESPS could reasonably estimate both the uncertainty in peak surge height and timing 92 of surge events resulting from imperfectly forecast atmospheric conditions six days before (Bernier 93 & Thompson, 2015). A 50-member ensemble simulation of 10 surge events during 2010 in Venice 94 by Mel & Lionello (2014) has shown that the distribution of maximum sea level is acceptably 95 realistic with respect to the deterministic forecast. They also found that the uncertainty became its 96 maximum during storm surge peaks and increased linearly with the forecasting lead time. 97 Although these ensemble simulation studies paved the way for a robust surge hazard assessment 98 over a single forecast-based assessment, they considered ensemble TC forecast information only 99 for developing and evaluating the performance (skill and accuracy) of an ESPS. In addition to 100 quantifying the uncertainty of surge height, ensemble-based storm surge multi-scenario (e.g., 101 worst/optimum case) analysis is equally important, aiding disaster risk managers in evacuation 102 planning (Kohno et al., 2018). 103

To the best of our knowledge, the potential of ensemble TC forecasting for assessing storm surge 104 multi-scenarios has largely been overlooked previously. However, recent developments have seen 105 the introduction of multi-scenario storm surge predictions, such as the worst-case scenario from 106 six typical TC tracks by the JMA (H. Hasegawa et al., 2017) and the maximum storm tide height 107 by the National Hurricane Center in the US (NHC, n.d.). These worst-case scenarios are composite 108 products, representing the maxima among all scenarios. Therefore, it is possible that the worst-109 case values for two adjacent locations may have come from two different ensemble TC track run. 110 Therefore, the users (e.g., emergency managers) cannot understand which forecasted TC track or 111 which combination of forecasted TC meteorological variables (track, intensity, size, translation 112 speed) may trigger the worst surge scenario for a particular location based on a composite product. 113 114 This can make it difficult for decision-makers to determine the appropriate level of storm surge warning and evacuation orders. In addition, storm surge is spatially heterogeneous because of its' 115 dependency on a TC characteristic and coastal geometry. It is entirely plausible that the worst case 116 scenario may not occur everywhere within a forecasted TC threat zone (Islam & Takagi, 2020a, 117 2020b). If the decision makers in cities/tourist districts with highly valuable economies issue a 118 higher warning level without any concrete understating over a worst event, they will inevitably 119 suffer significant economic losses because of false alarming (in case the area has not affected by 120 worst storm surge) and eventually can lower citizens trust over official warning (Sawada et al., 121 2022; Takagi et al., 2018). 122

Here we present Pareto optimality - a novel way of assessing storm surge multi-scenarios based 123 on ensemble TC forecasts. Our approach is more advanced than existing assessments. We 124 employed a multi-objective function to determine possible worst/optimum cases to quantify the 125 hazards in a large region. Our approach involved a comprehensive analysis of Pareto optimal 126 solutions in understanding the combination of forecasted TC meteorological variables - such as 127 track, intensity, size, and translation speed of TC - that could result in the worst/optimum surge 128 scenario. We utilized an extremely large ensemble (=1000 member) forecasts of TC Hagibis that 129 made landfall in central Japan in 2019. Our Pareto-based optimal solutions provide an 130 instantaneous overall assessment of storm surge multi-scenarios without any computational 131 burdens. The proposed method will allow forecasters to predict storm surge multi-scenarios 132 harnessing ensemble TC forecasts efficiently and help emergency responders as means of 133 quantifying surge hazards effectively. 134

135 **2 Data and Methods**

136 2.1 TC Hagibis and ensemble forecast

TC Hagibis in 2019, one of the most destructive and deadliest TC that hit Japan in decades (Shimozono et al., 2020; Ma et al., 2021), has been chosen to demonstrate our multi-scenario storm surge assessment. Hagibis was formed in the western North Pacific Ocean on 2 October 2019 and made landfall in central Japan on 12 October 2019 (around 0900 UTC), as depicted in Figure 1. At the landfall time, its maximum wind speeds sustained at 80 kt. This combined with heavy rainfall, resulted in high storm surges and severe flooding in the area (Shimozono et al., 2020; Ma et al., 2021; JMA, 2021).

The atmospheric ensemble forecasts of TC Hagibis were obtained by running JMA's former operational limited-area model called NHM (non-hydrostatic model; Saito et al., 2006). The integration domain (see Figure S1) had a grid spacing of 5 km consisting of 817×661 horizontal grid points and 50 vertical levels. Boundary conditions were interpolated to the NHM domain from
 JMA's global model forecasts and the forecast perturbations of JMA's operational one-week EPS.

JIVIA's global model forecasts and the forecast perturbations of JMA's operational one-week EPS

Since we used NHM for all forecast members, the only source of uncertainty stemmed from initial 149 conditions. This uncertainty is encapsulated in error covariances of current atmospheric states 150 (analysis error covariances), estimated using a data assimilation system. An ensemble Kalman 151 filter (EnKF) was employed to sample from these error covariances and generate an analysis 152 ensemble. While operational forecast centers generally use around 100 ensemble members, a state-153 of-the-art data assimilation system with 1000 ensemble members, called the four-dimensional 154 variational-ensemble assimilation technique (4DEnVAR), was utilized in this study (Kobayashi et 155 al., 2020). Our 4DEnVAR system only applied horizontal localization, with the horizontal 156 localization length scales derived from the JMA's operational four-dimensional variational 157 assimilation system's climatological horizontal correlation length scales. This helped to remove 158 159 sampling noise in estimating forecast error covariances and maintain the coherent vertical structure between atmospheric fields, which is critical in predicting tropical cyclones. As the ensemble 160 member count was large (=1000), localization was relaxed by retaining vertical correlations and 161 removing horizontal correlations at distant locations (Duc et al., 2021). 162

163 Unlike EnKF, EnVAR solely estimates the means of analysis ensembles and not the analysis 164 ensembles themselves, even though this method heavily relies on forecast ensembles to estimate these means. To solve this issue, a common approach is to run a separate EnKF in parallel to 165 generate analysis ensembles. However, our 4DEnVAR system is unique in that an EnKF is not 166 167 necessary. Instead, the same EnVAR program was used to generate analysis perturbations, as suggested in the context of inflation functions (Duc et al., 2020), where we demonstrated that using 168 quadratic inflation functions implies using the Kalman gain to generate analysis perturbations. 169 Using the same program for analysis means and analysis perturbations is essential because it 170 ensures consistency between the two when the same background error covariance, localization, 171 and observations are utilized in both cases. The assimilation system commenced at 00UTC on 7 172 October 2019, with a 3-hour assimilation cycle and continued until 18:00UTC 10 October 2019. 173 The final analysis ensemble was then used as initial conditions for 39h forecasts with NHM. The 174 assimilation domain was chosen the same as the forecast domain in Figure S1 and we assimilated 175 all routine observations obtained from JMA's database. Here, we opted for a 39h forecast horizon 176 177 because JMA's operational Meso-scale Ensemble Prediction System (MEPS) also generates 39h forecasts at 6-hour intervals (JMA, 2023). 178

179 2.2 Ensemble storm surge forecast

We used storm surge hazard potential index (SSHPI; eq. 1), a statistical model to compute peak 180 storm surge height. While the coastal engineers and ocean modelers are interested in the forecast 181 182 of storm surge hydrograph, most of the decision makers responsible for issuing surge warning and relief measures have a primary interest in the value of predicted peak surge height. The SSHPI 183 uses meteorological variables sensitive to storm surge, including TC intensity (V_{max}), size (radius 184 of 50-kt wind; R_{50}), and translation speed (S). In addition, the SSHPI considers coastal geometry 185 (a = 1 = open coasts and a = -1 = bays), landfall location sensitivity (D_L) , and regional scale 186 bathymetry (L_{30}) . The SSHPI does not incorporate factors associated with wave set-up and 187 astronomic tide to keep the configuration simple. TC Hagibis ensemble forecasts (=1000 member; 188 see Section 2.1), particularly during landfall, was used as meteorological forcing of the SSHPI. 189 We produced corresponding 1000 perturbed surge forecasts with a lead time of 39h. The 190

bathymetry of the target region was obtained from the Japan Oceanographic Data Center (Japan 191 Oceanographic Data Center, 2020). The effectiveness of the SSHPI for predicting peak surge 192 hazard potential was discussed in Islam et al. (2021, 2022). The formulation of the SSHPI is the 193 194 following:

195

$$SSHPI = \left(\frac{V_{max}}{V_{ref}}\right)^2 \left(\frac{R_{50}}{R_{ref}}\right) \left(\frac{S}{S_{ref}}\right)^a \left(\frac{L_{30}}{L_*}\right) \left(D_L\right) \tag{1}$$

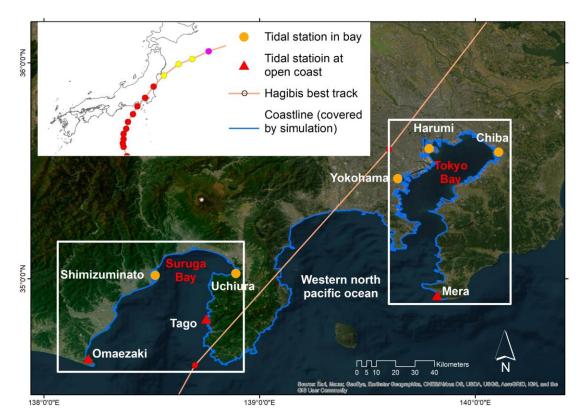
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$$197 \qquad \frac{R_{50}}{R_{ref}} = \begin{cases} 1.5 \quad if \ \frac{R_{50}}{R_{ref}} \ge 1.5\\ \frac{R_{50}}{R_{ref}} \ if \ 0.5 < \frac{R_{50}}{R_{ref}} < 1.5 \ ; \ (\frac{S}{S_{ref}})^a = \begin{cases} 1.5 \quad if \ (\frac{S}{S_{ref}})^a \ge 1.5\\ (\frac{S}{S_{ref}})^a \ if \ 0.5 < (\frac{S}{S_{ref}})^a < 1.5 \ ; \ \frac{L_{30}}{L_*} = \begin{cases} \frac{L_{30}}{L_*}, \ if \ \frac{L_{30}}{L_*} \ge 1\\ 1, \ if \ \frac{L_{30}}{L_*} \le 1\end{cases} \\ 1, \ if \ \frac{L_{30}}{L_*} \le 1\end{cases} \\ 0.5 \quad if \ (\frac{S}{S_{ref}})^a \le 0.5\end{cases}$$

$$D_L = \begin{cases} 1 \quad if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x \le 20\\ 0R \end{cases} \\ if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x \le 20\\ 1 - \frac{0.03(x - 20)}{20} \ if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x > 20\\ 1 - \frac{0.05(x - 10)}{10} \ if \ the \ surge \ estimated \ point \ falls \ right \ side \ of \ TC \ track \ and \ x > 10 \end{cases}$$

 V_{ref} , R_{ref} , and S_{ref} , are reference constants as follows: 50-kt equivalents of the tropical storm 199 category, 95 NM (historical mean R₅₀ at the time of landfall in Japan mainland), and 35 km/h 200 (historical mean S at the time of landfall in Japan), respectively (Islam et al., 2021). L_{30} is the 201 horizontal distance (km) between the shoreline and the 30-m depth contour. L* was chosen to be 202 10 km. D_L is defined by different expressions depending on the surge estimated points' (e.g., tidal 203 station) position (right/left) respective to the TC track and horizontal distance (x in km) between 204 the TC landfall location and a surge estimated point. Compared to V_{max} , the upper and lower bounds 205 of R_{50} , S, and L_{30} in eq. 1 restrict their contribution in generating surge hazards and, thus, prevents 206 207 discrete jumps in the SSHPI.

Figure 1 shows a storm surge modeling domain and the position of tide gauges used for validating 208 surge model and predicting surge hazards in this study. There are two domains, covering Tokyo 209 Bay and Suruga Bay individually. Each domain has tide gauges located both in inner bays and 210 open coasts. It should be noted that the tide gauges chosen for this study are the only stations that 211 possess recorded (historical) storm surge data, which is kept by JMA (JMA, 2022) and Japan Coast 212 Guard (Japan Oceanography Data Center, 2021). The empirical relationship for expected storm 213 surge in each tide gauge was determined in Islam et al. (2021, 2022) by drawing a line of best fit 214 215 through the historical surge data and the SSHPI and thus, used for the surge forecasts in this study. 216



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Figure 1. Domain of the storm surge forecasts model and the locations of the tide gauges used for

219 model validation and surge forecasts.

220 2.3 Pareto optimality and assessing storm surge multi-scenarios

It is unrealistic to anticipate a "nice" forecast scenario that accurately predicts the exact intensity 221 of a hazard at all locations within a given domain for a particular condition (e.g., worst/optimum). 222 223 An improved forecast at one location is usually accompanied by a deterioration of forecast at another location and vice versa. The best we can do is to quantify the trade-off between different 224 objectives. Here, we conducted multi-objective optimization to select ensemble forecast members 225 (among 1000 ensemble forecasts; see Section 2.1 and 2.2) that reasonably characterize the 226 potential worst/optimum storm surge case for a particular location (e.g., Tokyo Bay) by computing 227 the Pareto frontier. The Pareto frontier captures the trade-offs between objectives. It is the set of 228 all Pareto-optimum solutions where a single Pareto optimal solution denotes a solution that is not 229 dominated by any other solution (Kochenderfer & Wheeler, 2019). 230

In this study, we analyzed a subset of 1000 forecasted surge scenarios for each tide gauge in a 231 specific domain, referred to as solution z. Each scenario is evaluated based on d objectives, 232 represented by the values $y^{1}(z), y^{2}(z), \dots, y^{d}(z)$. As an example, we considered a scenario in Tokyo 233 Bay where the objectives are to maximize the forecasted peak surge at four tide gauges: Harumi, 234 Chiba, Yokohama, and Mera (as shown in Figure 1). This objective function considers the potential 235 worst-case scenario in Tokyo Bay for a set of 1000 ensemble surge forecasts. We compared the 236 given two solutions z and z', if for every objective i, $y^i(z) \ge y^i(z')$ and the strict inequality holds 237 for at least for one objective, we considered that solution z dominates z'. In other words, if one 238 solution (e.g., ensemble member no. 10) provides 160 cm, 155 cm, 140 cm, 110 cm of forecasted 239 peak surge in Harumi, Chiba, Yokohama, and Mera, respectively, but another solution (e.g., 240

ensemble member no. 20) yields 160 cm, 155 cm, 140 cm, 90 cm of forecasted peak surge for the 241 same tide gauges, then the solution given by ensemble member no. 10 dominates the solution 242 provided by ensemble member no. 20. This is because a smaller storm surge (90 cm) in Mera is 243 undesirable as stated in the objective function. A similar objective function but minimizing the 244 forecasted peak surge at four tide gauges: Harumi, Chiba, Yokohama, and Mera (as shown in 245 Figure 1) was considered to determine ensemble forecasts member that characterizes a potential 246 optimum surge case. In order to select the most appropriate ensemble TC member that may cause 247 the worst surge case in the inner bay, but also results in the optimum surge at the open coast, we 248 further assume that objectives are to be maximized for the inner bay tide gauges (Harumi, Chiba, 249 and Yokohama), but minimized for open coast tide gauges (Mera) at the same time. 250

and Tokonama), but minimized for open coast tide gauges (Mera) at the same time.

- The details of the implementation of the algorithm used here are described in Tommy (2021). This
- algorithm can compute the Pareto frontier for four objectives within a minute, meaning the runtime
- should be acceptable to any operational hazard forecast settings.

254 **3 Results**

- 255 3.1 Model evaluation
- 256 3.1.1 TC ensemble forecasts validation

From the forecasts of atmospheric fields given by NHM, TC tracks and intensities were detected. Here, TC centers are defined as the average of the mean sea level pressure minima, geopotentials at 850 hPa and 700 hPa. Figure 2(b) shows 1000 track forecasts generated from 1000 initial

- conditions obtained from the 4DEnVAR data assimilation system, along with the ensemble mean
- 261 forecast, control forecast, and best track. The ellipses in the figure illustrate uncertainties of the
- 262 TC centers, which are determined by the forecast error covariances of the TC centers. The arrows
- show the distance errors between the observations (i.e., the best track) and the ensemble mean. For
- comparison, the operational 20-member JMA ensemble forecast (MEPS) is included in Figure 2a.
- As shown in Figure 2, 4DEnVAR outperforms MEPS in terms of track forecasts and its ensemble
- 266 mean is almost identical to the best track with only minor distance errors.

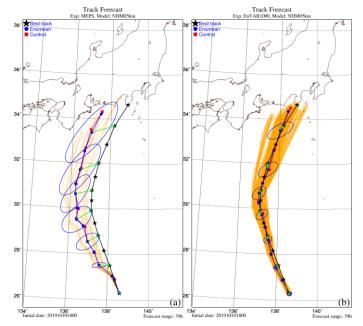


Figure 2. A comparison of 39h (at 1800 UTC on 10 October 2019) ensemble track forecasts for TC Hagibis issued by (a) the operational ensemble prediction system MEPS of JMA and (b) the 4DEnVAR data assimilation system. The ellipses represent forecast error covariances of the TC centers. The arrows denote the distance errors between the ensemble mean and best track.

Figure 3 presents the forecasted intensity of TC Hagibis as indicated by its central pressure. It is 272 evident from the figure that the 4DEnVAR (Figure 3b) surpasses JMA's operational MEPS (Figure 273 3a) in predicting the intensity. Even though both ensemble forecasts show overestimation of 274 intensity, the tendency of overestimation becomes more apparent with increasing forecast ranges 275 in MEPS (Figure 3a). Despite having a smaller number of ensemble members, MEPS exhibits 276 greater uncertainty in intensity forecast as compared to 4DEnVAR. This can be attributed to the 277 fact that JMA's operational MEPS employs singular vectors to generate initial conditions for 278 ensemble members, which maximizes their spread (JMA, 2023). 279

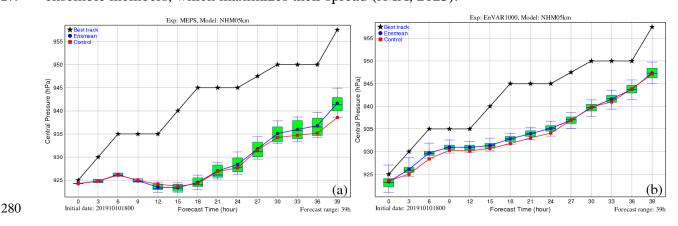
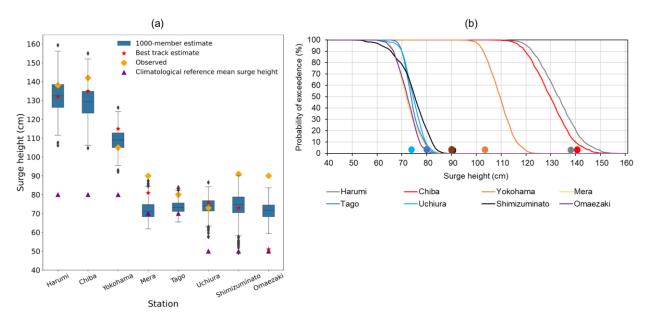


Figure 3. A comparison of 39h (at 1800 UTC on 10 October 2019) ensemble intensity forecast for TC Hagibis issued by (a) the operational ensemble prediction system MEPS of JMA and (b) the 4DEnVAR data assimilation system. The distributions of ensemble intensities are represented as box-and-whisker plots.

285 3.1.2 Storm surge ensemble forecasts validation

A comparison of the forecasted and measured peak storm surge (Japan Oceanography Data Center, 286 2021; JMA, 2022) at eight different tide gauges is shown in Figure 4a. It is noted that a significant 287 deviation from the climatological mean surge height is observed at all stations during TC Hagibis. 288 We evaluate JMA best track (JMA, 2021) as an ideal meteorological forcing input as well as our 289 39h ensemble TC forecasts. Both ensemble median forecasts and best track estimates 290 systematically underestimate the observed peak levels. Nevertheless, the observed peak surge 291 values are enveloped by the full ensemble of the forecasts in most tide gauges, implying that the 292 ensemble spread is large enough to represent the uncertainty in the prediction. The average mean 293 absolute error of the eight stations is 11.3 cm. In the case of Mera and Omaezaki, wave set-up is 294 often the dominant driver for generating storm surges (Islam et al., 2018, 2022), which is not 295 considered in the SSHPI. Therefore, the ensemble surge forecast underestimated observed surges. 296



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Figure 4. (a) A comparison of the 39h (at 1800 UTC on 10 October 2019) ensemble peak surge forecasts and the observed peak surge height for TC Hagibis at eight tide gauges. (b) Probability of exceeding (Y-axis) for a given storm surge threshold (X-axis) during TC Hagibis landfall time (at 0900 UTC, 12 October 2019). It was estimated by fitting ensemble forecasts empirically. A circle that shares the same color as a line represents the peak surge height recorded at a specific tide gauge.

The probability of surpassing a specific surge threshold during the landfall time of TC Hagibis is 304 illustrated in Figure 4b, as determined by the 39h ESPS. The observed peak surge levels are within 305 the predicted range, except for Mera and Omaezaki. For example, the probability of surpassing the 306 observed peak surge for Harumi (138 cm) is 26.4%. In general, Figure 4 indicates that the SSHPI 307 and its corresponding 39h peak surge forecasts are comparable in quality to those produced by 308 numerical surge models such as Liu et al. (2021). The latter study reported an RMSE of ~10 cm 309 when predicting the maximum total water level in Tokyo Bay during TC Hagibis with a 72h 310 forecast horizon, using atmospheric forcing fields from the Global Forecast System and a 311 hydrodynamic model known as the Semi-implicit Cross-scale Hydroscience Integrated System, 312 which has a nearshore resolution of ~150 m. 313

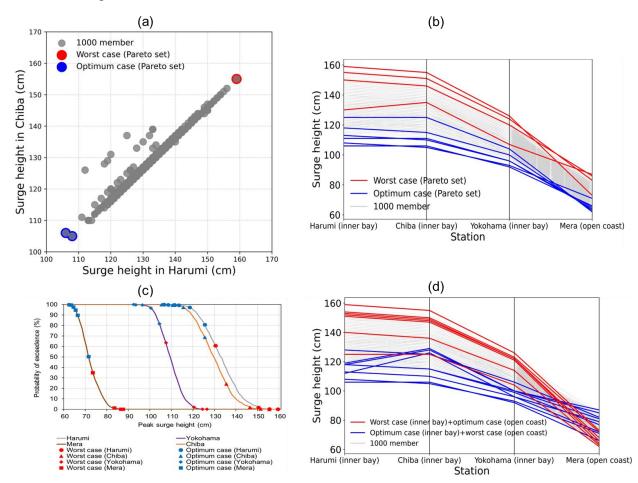
314 3.2 Multi-scenario analysis

315 3.2.1 Pareto optimal multi-scenarios

316 The Pareto-optimal frontier, as shown in Figure 5, illustrates a group of solutions that depict the forecasted potential worst and optimum storm surge scenarios for TC Hagibis in Tokyo Bay. The 317 two-dimensional Pareto frontiers (Figure 5a), allow for a straightforward evaluation of trade-offs 318 among the forecasted peak storm surge levels. The results identify the best one or two members 319 from the 1000 TC forecasts to represent the potential worst (Harumi, Chiba: ~155 cm) or optimum 320 (Harumi, Chiba: ~106 cm) surge scenario in the inner Tokyo Bay. Here, the tide gauges (Harumi, 321 322 Chiba) possess similar coastal geometry features, including bathymetry, and are situated in close proximity to each other (Figure 1). Therefore, the predicted surge response (from 1000 TC 323

forecasts; Figure 5a) between them is almost linear.

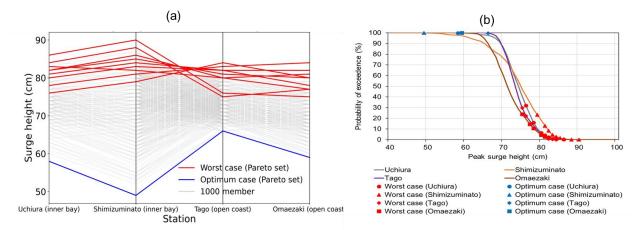
Figure 5b reveals a diversity of trade-off surge outcomes in the Pareto frontier for Tokyo Bay. It 325 includes four Pareto frontiers for worst-surge scenarios and six Pareto frontiers for optimum surge 326 scenarios. This diversity is due to the emergence of trade-offs among tide gauges with distinct 327 coastal geometry characteristics, including bathymetry. For instance, some of the identified most 328 severe surge scenarios for open coastlines in Tokyo Bay do not result in high surge levels in the 329 inner bay. Specifically, a Pareto optimal solution in Figure 5b predicts that Mera would experience 330 the worst surge levels of 90 cm (<1% exceedance probability; Figure 5c), while Harumi and Chiba 331 would experience approximately 135 cm (>25% exceedance probability; Figure 5c) of highest 332 surge levels under the same scenario. This storm surge level (~135 cm) in Harumi and Chiba is 333 substantially less than the worst surge levels (~155 cm) predicted by other optimal solutions. 334 Additionally, the surge intensity may vary across Tokyo Bay, depending on the characteristics of 335 the approaching TC and the impact can be much more severe in some places compared to others. 336 For example, several Pareto optimal solutions shown in Figure 5d predict that the inner Tokyo Bay 337 such as Harumi would witness surge levels higher than 150 cm, while it would keep as minimum 338 as 70 cm along the open coastline (e.g., Mera). Owing to such surge incongruence among the 339 coastal locations, creating multiple scenarios for different coastal places can lead to multiple 340 optimal solutions, as seen in Figures 5b and 5d. This emphasizes the importance of considering 341 multiple scenarios when issuing warnings and assessing the risks posed by extreme weather events 342 like storm surges. 343



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Figure 5. Forecasted Pareto optimal multi-scenarios due to TC Hagibis, apply for (a) Harumi and 345 Chiba in Tokyo Bay [objective function (red dot): max surge height in Harumi and Chiba; 346 objective function (blue dots): min surge height in Harumi and Chiba]; (b) Tokyo Bay (Harumi, 347 Chiba, Yokohama and Mera) using the parallel coordinate plot [objective function (red lines): max 348 surge height in Harumi, Chiba, Yokohama, and Mera; objective function (blue lines): min surge 349 height in Harumi, Chiba, Yokohama, and Mera]; (c) Probability of exceeding (Y-axis) a storm 350 surge threshold (X-axis) determined from each Pareto optimal solution in Figure 5b, at TC Hagibis 351 landfall time (at 0900 UTC, 12 October 2019). It was estimated by fitting ensemble forecasts 352 empirically; (d) Parallel coordinate plot including forecasted Pareto optimal solutions where 353 forecasted peak surge intensity varies across Tokyo Bay [objective function (red lines): max surge 354 355 height in Harumi, Chiba, and Yokohama + min surge height in Mera; objective function (blue lines): min surge height in Harumi, Chiba, and Yokohama + max surge height in Mera]. 356

357 The results displayed in Figure 6a are comparable to those in Figure 5b but for Suruga Bay. It includes nine Pareto frontiers for worst-surge scenarios and one Pareto frontier for optimum surge 358 scenario. Figure 6a predicts that all selected tide gauges would experience the worst surge levels 359 of ~85 cm (<1% exceedance probability; Figure 6b), while the minimum surge height would be 360 ~55 cm (>99% exceedance probability; Figure 6b). It is noteworthy that only one worst-surge 361 scenario is found to be shared by both the Pareto frontiers of Tokyo Bay (Figure 5b) and Suruga 362 Bay (Figure 6a), which represents a potential worst-case across the Japanese coastline during TC 363 Hagibis. This indicates that among 1000 TC ensemble forecasts, a particular TC ensemble member 364 has the potential to bring severe surge levels to both Bays. The reason behind this commonality 365 will be discussed in Section 3.2.2. 366



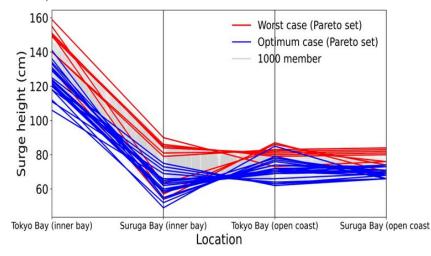
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Figure 6. (a) Parallel coordinate plot with Pareto optimal multi-scenarios for Suruga Bay, at TC Hagibis landfall time (at 0900 UTC, 12 October 2019) [objective function (red lines): max surge height in Uchiura, Shimizuminato, Tago, and Omaezaki; objective function (blue lines): min surge height in Uchiura, Shimizuminato, Tago, and Omaezaki]; (b) Probability of exceeding (Y-axis) a storm surge threshold (X-axis) determined from each Pareto optimal solution in Figure 6a. It was estimated by fitting ensemble forecasts empirically.

In addition to Pareto Frontiers illustrated in Figures 5b and 6a, we further noticed many distinct trade-offs when both bays were considered together (Figure 7). For instance, we incorporated the representative tide gauges for each category of coastal geometry, such as inner bay (Harumi in

Tokyo Bay and Shimizuminato in Suruga Bay) and open coast (Mera in Tokyo Bay and Tago in

Suruga Bay) in the objective function. This resulted in a large number of Pareto frontiers (worst 378 case: nine; optimum case: twenty-two), with multiple overlapping solutions between worst and 379 optimum cases. Thus, 33% of Pareto optimal solutions (red lines) predict that inner Tokyo Bay 380 will experience the worst surge levels of ~150 cm, while inner Suruga Bay will witness ~60 cm, 381 equivalent to minimum surge cases predicted by 41% of Pareto optimal solutions (blue lines; 382 Figure 7). Although we maximize the potential of the large (i.e., 1000) ensemble forecasts, our 383 proposed method is also found to be useful with the small ensemble size. We repeated the same 384 analysis with 36 ensemble size. Although some worst and minimum storm surge scenarios were 385 missed, a meaningful set of Pareto optimal solutions was still obtained (see Figure S2 in the 386 supplementary section). 387



388

Figure 7. Parallel coordinate plot with Pareto optimal multi-scenarios determined for both Tokyo
 Bay and Suruga Bay [objective function (red lines): max surge height in Harumi, Shimizuminato,
 Mera, and Tago; objective function (blue lines): min surge height in Harumi, Shimizuminato,

392 Mera, and Tago]

393 3.2.2 TC track and meteorological variables analysis of Pareto optimal solutions

It would be interesting to analyze the tracks and associated meteorological variables of the 394 identified Pareto optimal solutions in Figures 5b and 6a. For example, Figure 8 reveals the strong 395 sensitivity of the storm surge scenarios to the landfall location of TC Hagibis, leading to 396 contrasting Pareto optimal solutions for Tokyo Bay (Figure 5b) and Suruga Bay (Figure 6a). It 397 also demonstrates that the forecasted TC tracks for worst and optimum surge scenarios are 398 significantly different from one other in both bays. For example, TC tracks (red lines; Figure 8a) 399 that run parallel to the longitudinal axis of Tokyo Bay and pass over it would result in severe storm 400 surges than TCs (blue lines; Figure 8a) that would travel 100 km or more to the west of the axis. 401 Prior to landfall, under the worst case, easterly wind (Figure 9a) is forecasted to cause a buildup 402 of water on the west coast (e.g., Yokohama) and initial draw-down in the north-eastern end of the 403 Tokyo Bay (e.g., Chiba). Later, a surge level difference (0.6–0.8 m; Figure 5b) between the inner 404 and lower ends of the bay is projected to occur (due to strong southerly winds) during the peak 405 storm surge at the inner bays. During TC makes landfall under the optimum case (Figure 8a), the 406 destructive right-side semicircle of the TC (Figure 9f) will interact with the vast land area rather 407 than the ocean water, leading to less water being pushed towards Tokyo Bay (Figure 5b). 408

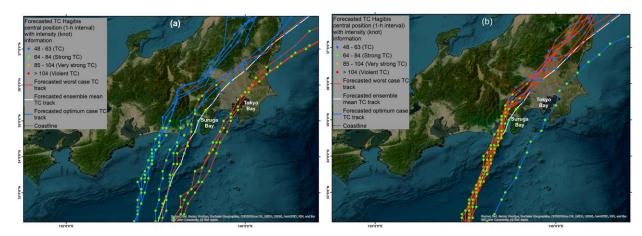
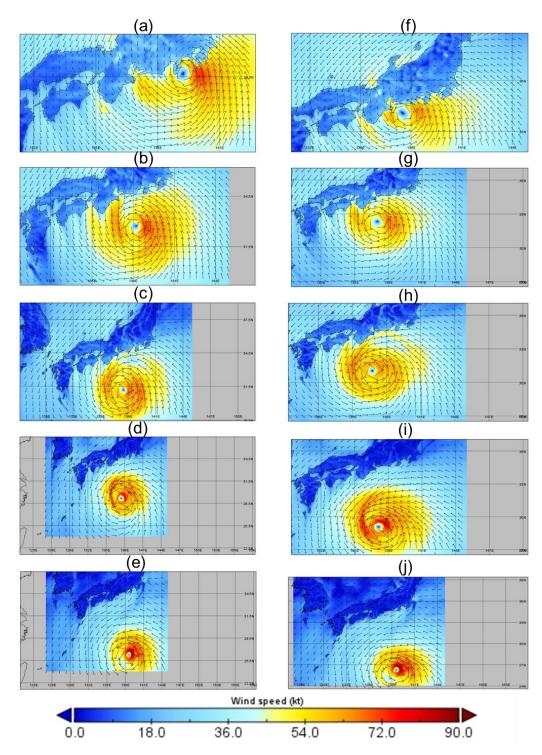




Figure 8. Forecasted TC Hagibis track in 39h lead time, respective to each Pareto frontier
determined for (a) Tokyo Bay (as shown in Figure 5b); (b) Suruga Bay (as shown in Figure 6a).
Red, blue, and white lines correspond to the forecasted worst, optimum, and ensemble mean TC
track.

Notably, the optimized ensemble TC members for minimum surge scenarios in Tokyo Bay are forecasted to be stronger and larger than the members belonging to the worst surge cases until 24 hours prior to landfall, despite being centered in the same location. This is evident in Figure 9d, 9e, as opposed to Figure 9i, 9j. Despite both sets (worst and minimum) of optimized ensemble TC members weaken as they approach the mainland of Japan (Figure 9c, 9h and Figure 10), the worst TC members intensify by 7-kt (V_{max}) and remain large (R_{50} : ~120 NM) in the last 12 hours before landfall (Figure 9a, 9b and Figure 10). This large swath of strong winds is forecasted to affect a

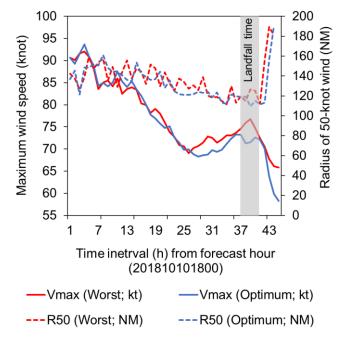
- 420 greater sea area and induce a motion in a greater quantity of water in Tokyo Bay.
- 422



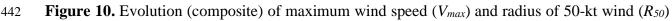
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Figure 9. Forecasted composite 10 m wind field (kt-vectors), generated from optimized ensemble
TC members for worst surge scenarios in Tokyo Bay (as shown in Figure 5b) during (a) landfall
time (at 0900 UTC, 12 October 2019); (b) 6-h before landfall (at 0300 UTC, 12 October 2019);
(c) 12-h before landfall (at 2100 UTC, 11 October 2019); (d) 24-h before landfall (at 0900 UTC,
11 October 2019); (e) 39-h before landfall (at 1800 UTC, 10 October 2019); (f-j) same as (a-e) but

- 429 generated from optimized ensemble TC members for minimum surge scenarios in Tokyo Bay (as
- 430 shown in 5b).
- 431 Figure 8a highlights that in addition to TC tracks that pass directly over Tokyo Bay, one particular
- 432 worst TC forecast (red line) makes landfall around 70 km west of the longitudinal axis of Tokyo
- Bay. This specific track is forecasted to cause severe storm surges in both Tokyo Bay (Figure 5b)
- and Suruga Bay (Figure 6a). This particular ensemble forecast has a wider range of intense winds
- 435 (R_{50} : ~140 NM) across a larger area and a slower movement speed (S: ~32 km/h), despite having 436 a similar landfall wind intensity (V_{max} : ~75-kt) compared to other worst-case TC forecasts. This
- 437 unique phenomenon corroborates earlier numerical analyses that propose the likelihood of a severe
- 438 storm surge scenario in the upper-bay region when a large and intense TC moves slowly, parallel
- to the longitudinal axis of Tokyo Bay, after making landfall 25 km southwest. (Islam & Takagi,
- 440 2020b).



441



for worst and optimum TC forecasts in Tokyo Bay (as shown in Figure 5b) in 39-h lead time.

444 4 Conclusions and discussion

The application of ensemble TC forecasting in hazard prediction, such as storm surge, has been 445 greatly overlooked despite its use in forecasting TC track, intensity, and genesis. Enhanced 446 analysis can unlock and maximize the benefit of ensemble forecasting. Here, we proposed Pareto 447 optimality – a novel and practical way to identify potential ensemble TC (Hagibis) forecast from 448 449 an extremely large ensemble (=1000 member) that can effectively assess storm surge multiscenarios, including possible worst and optimum cases for a coastal location. The variability in 450 storm surge intensity across the coastline makes it challenging for decision-makers to plan 451 effective evacuation measures. To address this, we have demonstrated that a diversity of trade-off 452 surge outcomes among coastal places can be identified by choosing the Pareto optimized forecasts. 453 The in-depth evaluation of Pareto optimal solutions can shed light on how meteorological variables 454 such as track, intensity, and size of TCs influence the worst and optimum surge scenarios, which 455

456 are not well understood by emergency managers using current multi-scenario assessment methods457 (such as those used by JMA and NHC).

The significance of evaluating trade-offs based on Pareto optimization has long been 458 acknowledged in the context of sustainable development goals and management of ecosystem 459 services (Flecker et al., 2022). However, its application in disaster-risk communities is noticeably 460 lacking. During a coastal storm surge event, effective evacuation planning and warning issuance 461 involve multi-criteria problems such as storm surge intensity, coastal population vulnerability, and 462 available evacuation resources. Traditionally, this decision-making process has relied on a hazard 463 map, which typically depicts the severity of the predicted storm surge (e.g., exceeding a critical 464 surge height; J. Hasegawa et al., 2017). However, this approach does not fully capture the diversity 465 of potential storm surge scenarios across the coastlines, which can lead to ineffective evacuation 466 planning. A recent TC Fani in 2019 serves as evidence to support this statement. TC Fani struck 467 the southeastern part of India, approximately 450 km from the southwest coast of Bangladesh, as 468 a Category 4 TC (in Saffir-Simpson Hurricane Wind Scale). Prior to TC Fani reaching the 469 Bangladesh coast as a tropical storm, the Bangladesh Meteorological Department (BMD) issued 470 'danger' signal number seven (out of ten), which led to the evacuation of one million people 471 (Bangladesh Meteorological Department, 2021). Later, the catastrophe, such as storm surge level 472 (~1 m), did not hit the danger level as anticipated. BMD's evacuation order for the entire southwest 473 474 coast was not based on a specific surge scenario (e.g., worst case) and forecasted meteorological conditions associated with it, leading to an excessive number of evacuees. Such a false alarm 475 demotivated people to seek shelter when a Category 2 TC Amphan caused ~2.75 m storm surges 476 and claimed at least 26 lives in 2020, despite the issuance of 'great danger' signal number ten 477 (Raju, 2019; ReliefWeb, 2021; Alam et al., 2023). It seems that BMD took a safer and conservative 478 decision during TC Fani by issuing signal no. 7, nevertheless, this cannot be considered effective 479 decision-making. While such a complex decision-making process can certainly be improved by 480 quantifying the uncertainty through an ensemble multi-scenario forecast, incorporating Pareto 481 optimality can further maximize the benefits of it. 482

Pareto optimal solution provides an effective first filter to identify ensemble multi-scenario surge 483 forecasts. This information can be presented visually to enhance the understanding of the 484 uncertainty in the forecast. The median of Pareto optimal solutions could be utilized given a series 485 of worst/minimum surge estimations for a specific location by several ensemble members. For 486 example, Figure 5b identifies four worst scenarios for Harumi in Tokyo Bay where the median 487 surge level is 150 cm. Although we stress the importance of diversity in trade-offs surge outcomes, 488 a certain scenario (e.g., median of Pareto optimal solutions) can be given more weight depending 489 on the values of society and decision-makers. In the decision-making process, a user-defined 490 acceptable level of uncertainty or reference surge height (e.g., 25-year return period of surge) can 491 be set for a specific location (e.g., Harumi). The forecaster can then determine if the height of the 492 Pareto optimal solution exceeds this acceptable level. Subsequently, a relevant warning signal can 493 be issued in a forecast horizon (e.g., 39-h lead time). The warning signal can be tailored to a 494 specific location, if a diversity in the trade-off between surge outcomes exists among Pareto 495 frontiers. For example, Pareto optimal solutions in Figure 5b predicted that TC Hagibis will bring 496 worst surge level as maximum as 160 cm in the inner Tokyo Bay, requiring the issuance of an 497 emergency warning, closing flood gates, and large-scale evacuation of the coastal population 498 living below the storm surge height. On the other hand, those living along open coastlines are 499 advised to stay indoors as the predicted worst surge level (90 cm) does not meet the criteria for 500 issuing an emergency warning. Once this forecast becomes available, decision-makers (e.g., 501

emergency managers) can start evacuation planning based on the forecasted worst TC track, wind 502 intensity, and peak surge height. For instance, TC track that forecasted to bring severe storm surges 503 of 160 cm in the inner Tokyo Bay, would make landfall 70 km west of the central bay axis with a 504 landfall V_{max} of 75 kt, which is stronger by 11-kt from the historical mean (64-kt; (Islam et al., 505 2022)). Furthermore, it is projected to be twice as large as the historical average (65 NM) and 506 move at a slower speed by 9 km/h compared to the average translation speed (41 km/h) in Tokyo 507 Bay. The forecasted landfall location and meteorological conditions of the worst TC indicate that 508 Tokyo Bay would be situated in the destructive right-side semicircle of the TC track, resulting in 509 prolonged exposure to severe storm surges and strong winds. Emergency managers can utilize this 510 information to disseminate surge warnings to residents and commence evacuation procedures with 511 a 39-hour lead time. This evacuation can be done by dividing coastal regions into different zones 512 depending on their vulnerability. Although disaster planning is not so straightforward as explained 513 here, our proposed ensemble-based storm surge multi-scenario analysis is expected to motivate 514 forecasters and risk management practitioners to explore new ways to assess storm surge hazards 515 and reduce the associated risk. 516

517 Finally, we acknowledge that this study focuses exclusively on peak surge height while 518 determining total sea water level that includes the influence of astronomic tide, wave set-up, and 519 river discharge are also critical and can be done utilizing a full physical numerical model. 520 Furthermore, several algorithms are currently available to determine a Pareto frontier. We 521 encourage researchers from multiple disciplines to build on our approach to help us reach an 522 improved understanding of Pareto optimality based multi-scenario analysis.

523 Acknowledgment

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530 **Open Research**

Observed surge data be downloaded from the JMA 531 storm can (https://www.data.jma.go.jp/kaiyou/db/tide/genbo/index.php) and JODC 532 (https://jdoss1.jodc.go.jp/vpage/tide.html) websites. Predicted tide data can be obtained from the 533 JMA (https://www.data.jma.go.jp/kaiyou/db/tide/suisan/index.php) website. TC best track data 534 can be derived from the JMA (https://www.jma.go.jp/jma/jma-eng/jma-center/rsmc-hp-pub-535 eg/trackarchives.html) website. Ensemble forecast data may be available upon request. 536

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Supporting Information for

Assessing Storm Surge Multi-Scenarios based on Ensemble Tropical Cyclone Forecasting

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Contents of this file

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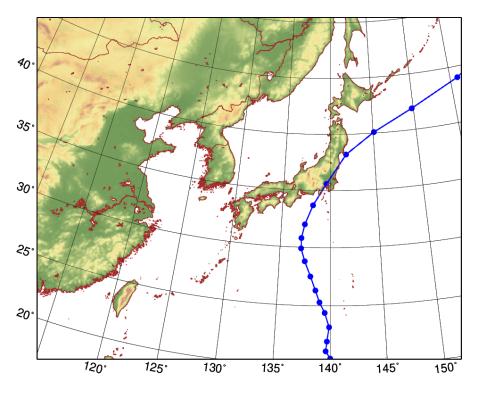


Figure S1. The analysis and forecast domain of the data assimilation system 4DEnVAR and the forecast model NHM. The best track of the TC Hagibis is also plotted.

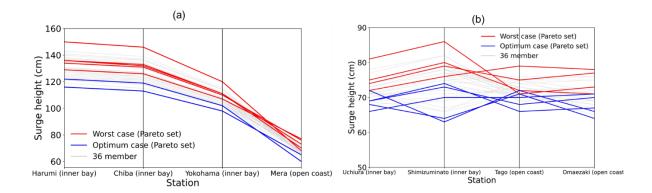


Figure S2. Parallel coordinate plot with Pareto optimal multi-scenarios based on 36 ensemble forecasts for (a) Tokyo Bay (Harumi, Chiba, Yokohama and Mera) [objective function (red lines): max surge height in Harumi, Chiba, Yokohama, and Mera; objective function (blue lines): min surge height in Harumi, Chiba, Yokohama, and Mera]; (b) Suruga Bay (Uchiura, Shimizuminato, Tago, and Omaezaki) [objective function (red lines): max surge height in Uchiura, Shimizuminato, Tago, and Omaezaki; objective function (blue lines): min surge height in Uchiura, Shimizuminato, Tago, and Omaezaki; objective function (blue lines): min surge height in Uchiura, Shimizuminato, Tago, and Omaezaki; objective function (blue lines): min surge height in Uchiura, Shimizuminato, Tago, and Omaezaki].