

Investigating the Mainshock and Aftershock of the May 2006 Earthquake in Central Java for Aerothropolis Development at Yogyakarta International Airport

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March 9, 2023

Abstract

Research on the analysis of the source mechanism of the mainshock and aftershock events of the May 27, 2006, Yogyakarta earthquake, which is thought to have originated from the Opak fault and analysis of receiver function data to model the subsurface velocity P of the Central Java subsurface, to obtain a geological form model of the Opak fault. This research aims to support the development of the Yogyakarta Aerothropolis area in terms of disaster analysis. The data used in this study are remote Teleseismic receiver function data from the MERAMEX station installed in 2004, and data for the Bantul earthquake event and its aftershock event in 2006. The results obtained from the analysis are that the Yogyakarta area is shaped like a half-graben close to Yogyakarta International Airport. The fault that separates the western part of Yogyakarta is still not identified. Based on the results of the rupture process analysis of the source along the Opak fault plane, some zones have not yet released their energy. The distribution of aftershocks due to the mainshock on 27 May 2006 is spread around the Opak fault, which is heading North-South, and West-East, which is thought to have activated the minor fault to the east of the Opak fault. The opak fault rupture area can be analyzed to have a Low Anomaly velocity P value from the receiver function data and is the same as the aftershock event obtained.

1 **Investigating the Mainshock and Aftershock of the May 2006 Earthquake in**
2 **Central Java for Aerothropolis Development at Yogyakarta International**
3 **Airport**

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28

29**Key Points:**

- 30 • Analysis of source mechanisms of mainshock and aftershock events of May 27, 2006, earthquake in
31 Yogyakarta, Central Java.
- 32 • Use of receiver function data to model the subsurface velocity P and obtain a geological form model of the
33 Opak fault.

- 34 • Research aimed at supporting the development of the Yogyakarta Aerothropolis area in terms of disaster
35 analysis.

36

37**Abstract**

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39 May 27, 2006, Yogyakarta earthquake, which is thought to have originated from the Opak fault
40 and analysis of receiver function data to model the subsurface velocity P of the Central Java
41 subsurface, to obtain a geological form model of the Opak fault. This research aims to support
42 the development of the Yogyakarta Aerothropolis area in terms of disaster analysis. The data
43 used in this study are remote Teleseismic receiver function data from the MERAMEX station
44 installed in 2004, and data for the Bantul earthquake event and its aftershock event in 2006. The
45 results obtained from the analysis are that the Yogyakarta area is shaped like a half-graben close
46 to Yogyakarta International Airport. The fault that separates the western part of Yogyakarta is still
47 not identified. Based on the results of the rupture process analysis of the source along the Opak
48 fault plane, some zones have not yet released their energy. The distribution of aftershocks due to
49 the mainshock on 27 May 2006 is spread around the Opak fault, which is heading North-South,
50 and West-East, which is thought to have activated the minor fault to the east of the Opak fault.
51 The opak fault rupture area can be analyzed to have a Low Anomaly velocity P value from the
52 receiver function data and is the same as the aftershock event obtained.

53

54 Keywords: Rupture, aftershock, receiver function, opak fault, aerothropolis.

55

56**Plain Language Summary**

57 This study analyzes the source mechanism of the main earthquake and its aftershocks that
58 occurred on May 27, 2006, in Yogyakarta, Indonesia, believed to have originated from the Opak
59 fault. The researchers used receiver function data to model the Central Java subsurface velocity P
60 and obtained a geological form model of the Opak fault. The study aimed to support the
61 development of the Yogyakarta Aerothropolis area in terms of disaster analysis. The results
62 showed that the Yogyakarta area is shaped like a half-graben close to Yogyakarta International
63 Airport. The fault that separates the western part of Yogyakarta has not yet been identified. The
64 distribution of aftershocks due to the mainshock was spread around the Opak fault, which is
65 heading North-South and West-East, and it is thought to have activated the minor fault to the east
66 of the Opak fault. The study also found that the Opak fault rupture area can be analyzed to have a
67 low anomaly velocity P value from the receiver function data and is the same as the aftershock
68 event obtained. This study provides valuable information for disaster analysis in the Yogyakarta
69 Aerothropolis area.

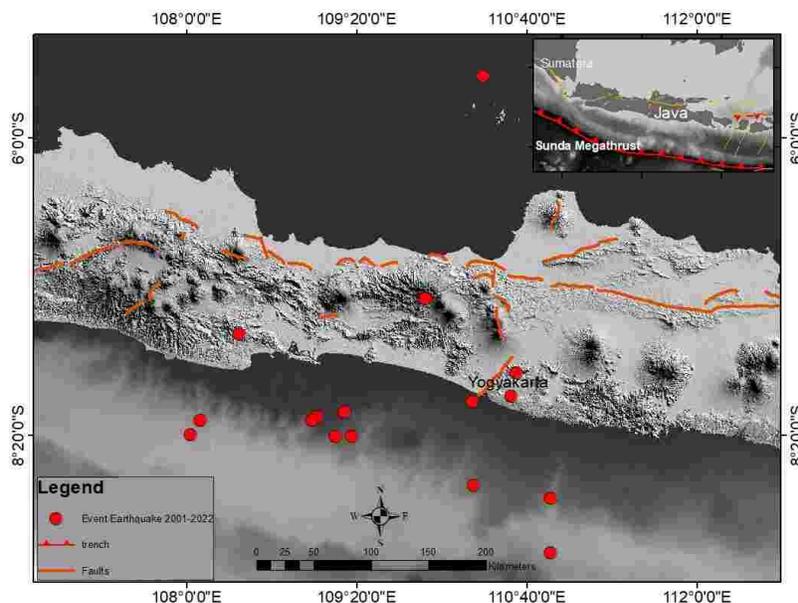
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711 Introduction

72 Yogyakarta is an area that has its uniqueness and charm, so many tourists come, both
 73 local and foreign. This has a positive impact on regional growth. The area around the airport,
 74 with many business activities or commercial services, is the basis for forming the aerotropolis
 75 area concept (Kasarda, 2016). This concept was developed in the Yogyakarta International
 76 Airport area to increase business competitiveness which will have an impact on the community's
 77 economy so that it will have a large effect (Sumarata, 2021). This condition is supported by
 78 proper planning, including security from the threat of disaster.

79 Yogyakarta is not only a popular tourist destination but also an area that is prone to
 80 earthquake disasters due to its location on the active tectonic plate. According to the National
 81 Disaster Management Agency (BNPB), Yogyakarta has a high seismic hazard level, with 21
 82 active faults identified in the region (BNPB, 2021). Therefore, proper disaster risk management
 83 is crucial in aerotropolis planning to minimize the potential impact of earthquakes on the
 84 community's economy and livelihoods. This can be achieved through a comprehensive analysis
 85 of the earthquake risk, including the identification of vulnerable areas and the implementation of
 86 appropriate measures to mitigate the risk. In addition, community participation and awareness-
 87 raising programs should be incorporated into the disaster risk management plan to ensure the
 88 sustainability of the aerotropolis area development. By considering the earthquake risk and
 89 implementing appropriate measures, the aerotropolis area in Yogyakarta can develop sustainably
 90 and contribute to the economic growth of the region while ensuring the safety and well-being of
 91 the local community.

92 The principles of aerotropolis planning consist of regional spatial structure, distance,
 93 zoning, land use, designation of the main functions of the area, integration, and connectivity
 94 (Kurniawan, 2016; Yujin, 2013; Wang, 2013; Greis, 2011). To find out good spatial planning in
 95 aerotropolis planning, an in-depth analysis of disasters in the area is also needed (Surya et al.,
 96 2020). Both seasonal disasters, such as floods and tornadoes, and large, unpredictable disasters,
 97 such as volcanoes, earthquakes, and tsunamis (Ammon et al., 2006). Therefore, this research was
 98 conducted to support one side of the earthquake disaster, which can be analyzed from many
 99 events around the area (Kuncoro et al., 2020).



100 Figure 1. Earthquake distribution map for 2001-2022 around Central Java and Yogyakarta. The
101 red circle is the mainshock distribution that has been relocated. The brown lines are the
102 distribution of faults spread across Java island.

103 The feasibility study has succeeded in reconstructing the source model in detail and
104 analyzing the source mechanism and the rupture process of the May 27, 2006, Yogyakarta
105 earthquake by knowing the coseismic process and the seismotectonic characteristics of the study
106 area. So by knowing the physical characteristics of the May 27, 2006, Yogyakarta earthquake, we
107 can understand the destructive nature of the earthquake, so it can be used as a guide for
108 earthquake mitigation, especially the Yogyakarta earthquake (Saputra et al., 2021). However, to
109 obtain information on the source mechanism and phenomenon of May 27, 2006, Yogyakarta
110 earthquake, it is recommended to use other earthquake data. So that a more detailed analysis of
111 the mechanism of the source and the process of earthquake rupture in the research area will be
112 obtained so that the characteristics of the Opak fault will be known in more detail, especially
113 after 2020, Yogyakarta International Airport, which is located in Temon District, Kulon Progo
114 Regency, has been operating. The airport area will be built with the Aerotropolis concept in its
115 development. In the Yogyakarta Regional Regulation number 5 of 2019 concerning the 2019-
116 2039 Regional Spatial Plan for the Special Region of Yogyakarta, the development of the
117 Temon-Prambanan Area is projected to become one of the province's strategic areas from an
118 economic point of view. One form is the development of the aerotropolis area. Aerotropolis can
119 increase economic growth in a country or region (Peoples, 2014). The concept of the Yogyakarta
120 aerotropolis with a radius of about 15 km from the airport (Syaifuddin et al., 2021).

121 This research aims to analyze the physical characteristics and the source mechanism of
122 the May 27, 2006, Yogyakarta earthquake. The urgency of this study is to provide a better
123 understanding of the destructive nature of the earthquake and to provide guidance for earthquake
124 mitigation, especially in the Yogyakarta region. The results of this study can also provide
125 important information for the development of the aerotropolis area around Yogyakarta
126 International Airport. Understanding the seismotectonic characteristics and potential risks of
127 earthquakes in the area is crucial for ensuring the safety and sustainability of the airport and the
128 surrounding region. Additionally, this research can contribute to the advancement of earthquake
129 science and the development of effective strategies for disaster risk reduction.

1302 Materials and Methods

1312.1 Receiver function data

132 The method used in this study is the aftershock relocation method to determine the
133 aftershock event in the 2006 earthquake that occurred to obtain a rupture zone (Saputra, 2021).
134 The relocation method was conducted using the hypoDD code (Waldhauser and Ellsworth, 2000)
135 and the double-difference algorithm (Zhang et al., 1889). This event data is combined with the
136 subsurface velocity model data resulting from the Receiver Function method (Figure 2) using the
137 MERAMEX network installed in 2004 (Amukti, 2019). The MERAMEX network is a dense
138 seismographic network that consists of 24 broadband and short-period seismometers, covering an
139 area of approximately 200 km radius around the study area. The subsurface structure is then
140 analyzed using the tomographic inversion method, which allows for a more detailed study of the
141 subsurface velocity structure and can provide insight into the characteristics of the active fault in
142 the study area.

143

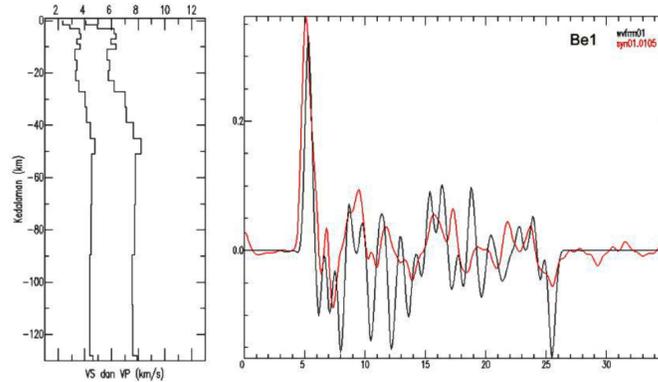


Figure 2. Example of receiver function data

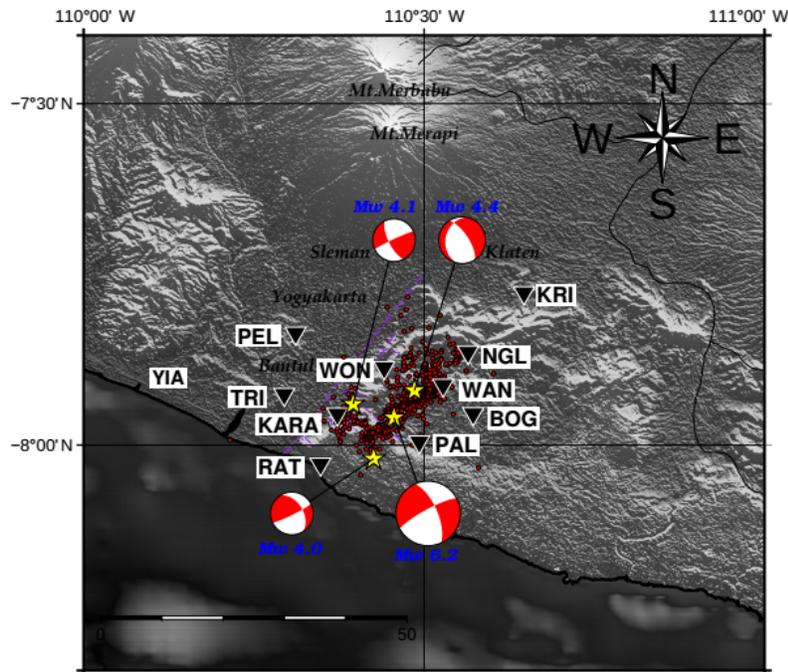
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146

1472.2 Aftershocks data

148 This study also uses relocated aftershock data obtained from temporary seismometers
 149 installed after the 2006 Yogyakarta mainshock. In this analysis, the study employs the aftershock
 150 relocation method to determine the precise location of each aftershock event and reduce
 151 uncertainty in earthquake location determination.

152



154 Figure 3. Distribution of relocated aftershocks taken from temporary seismometer recordings.

155 The red circle is the relocation aftershock distribution. The grey triangles are
 156 seismometers. The red and white focal ball is the mainshock and aftershock issued by
 157 the USGS.

158

159 Figure 3 shows the enlargement between the mainshock and the three aftershocks, namely on
 160 June 8, 9 and 16, which had different types of displacement. This shows the complexity of the
 161 faults around the Opak fault. The results of the mainshock rupture process analysis in the study

162of Saputra et al. (2021) show two weak zones and can be seen from the distribution of
163aftershocks scattered around the asperity zone (weak zone). Between the weak zones, there is a
164seismic gap, indicated as a zone that has not yet released its energy.

165

1663 Data

167 This study utilized the earthquake aftershock location method to determine the aftershock
168events of the 2006 earthquake in order to obtain the fault zone. The location method was
169performed using the hypoDD code (Waldhauser & Ellsworth, 2000) and the double-difference
170algorithm (Zhang et al., 1889). The earthquake data was combined with subsurface velocity
171model data generated from the Receiver Function method (Figure 2) using the MERAMEX
172network installed in 2004 (Amukti, 2019). The MERAMEX network is a dense seismograph
173network consisting of 24 broad-band and short-period seismometers covering an area of about
174200 km radius around the study area. The subsurface structure was then analyzed using the
175tomographic inversion method, which allows for a more detailed study of the subsurface velocity
176structure and provides insights into the characteristics of active faults in the study area.

177 The aftershock data was also used in this study, obtained from temporary seismometers
178installed after the main earthquake in Yogyakarta in 2006. In this analysis, the study used the
179earthquake location method to determine the precise location of each aftershock event and
180reduce uncertainty in determining the earthquake location. Figure 3 shows the enlargement
181between the main earthquake and three aftershocks on June 8, 9, and 16, which have different
182displacement types. This indicates the complexity of faults around the Opak fault. The results of
183the faulting process analysis in the main earthquake by Saputra et al. (2021) show the presence
184of two weak zones, which can be seen from the distribution of aftershocks scattered around the
185asperity zone (weak zone). Between the weak zones, there is a seismic gap, which is indicated as
186a zone that has not yet released its energy.

187 In this study, the Receiver Function method was used to generate the subsurface velocity
188model, which was then combined with the aftershock location data to obtain insights into the
189characteristics of active faults in the study area. The tomographic inversion method was used to
190obtain more detailed information about the subsurface structure. In addition, the earthquake
191location method was also used to determine the precise location of each aftershock event, which
192helps in understanding the complexity of faults around the Opak fault. The results of this study
193show the presence of two weak zones and a seismic gap between the weak zones. These results
194can provide important information for understanding the potential earthquake hazards in the
195region.

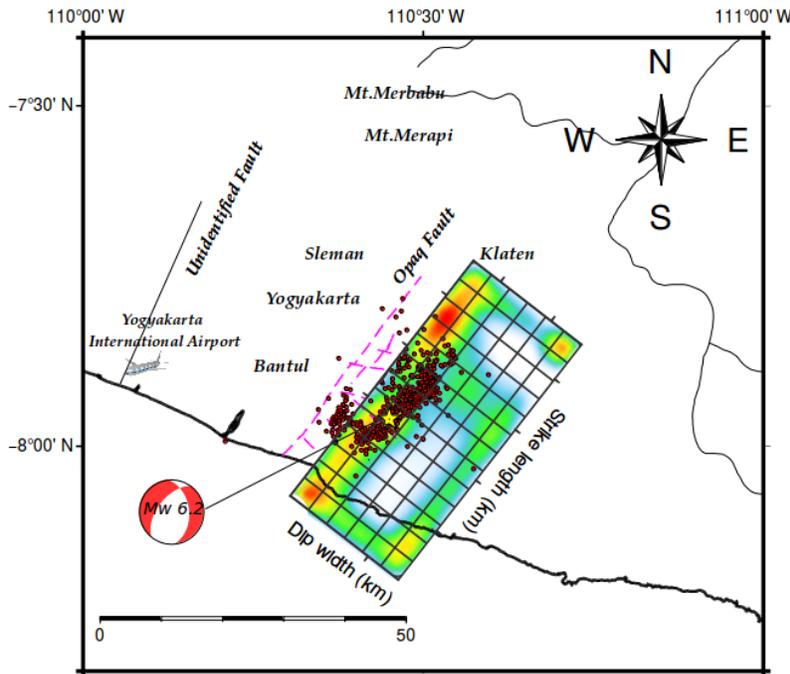
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1974 Results, or a descriptive heading about the results

1984.1 Analysis of the earthquake source rupture process

199 Analyzing the earthquake source rupture process is critical to seismological research, as it
200helps understand the mechanisms that trigger earthquakes. In this study, the researchers used the
201finite fault inversion method to analyze the earthquake source rupture process of the Yogyakarta
202earthquake. The results showed that the rupture process was a unilateral propagation from the
203southwest to the northeast. Moreover, the fault length was approximately 28 km, and the
204maximum slip was estimated to be around 1.4 meters. The results are crucial in improving our

205 understanding of the characteristics of the Yogyakarta earthquake and provide valuable
206 information for developing effective earthquake mitigation strategies.



207 Figure 4. The results of the analysis of the rupture process of the Yogyakarta earthquake on May
208 27, 2006, show the slip distribution from the opaque fault plane. The rectangular
209 contour is the distribution of slip on the opaque fault plane. The red colour circle is the
210 aftershock distribution. The yellow star is the Yogyakarta earthquake mainshock,
211 resulting from a joint inversion research by Saputra et al. (2021). The purple dotted line
212 is an opaque fault, and on the east side of the main fault, there are several minor faults.
213 The black line is the boundary of the graben zone, where the fault has not been
214 identified. The red and white focal ball results from a joint inversion calculation from
215 Saputra et al. 2021 study (modification from Saputra et al. Research, 2021).

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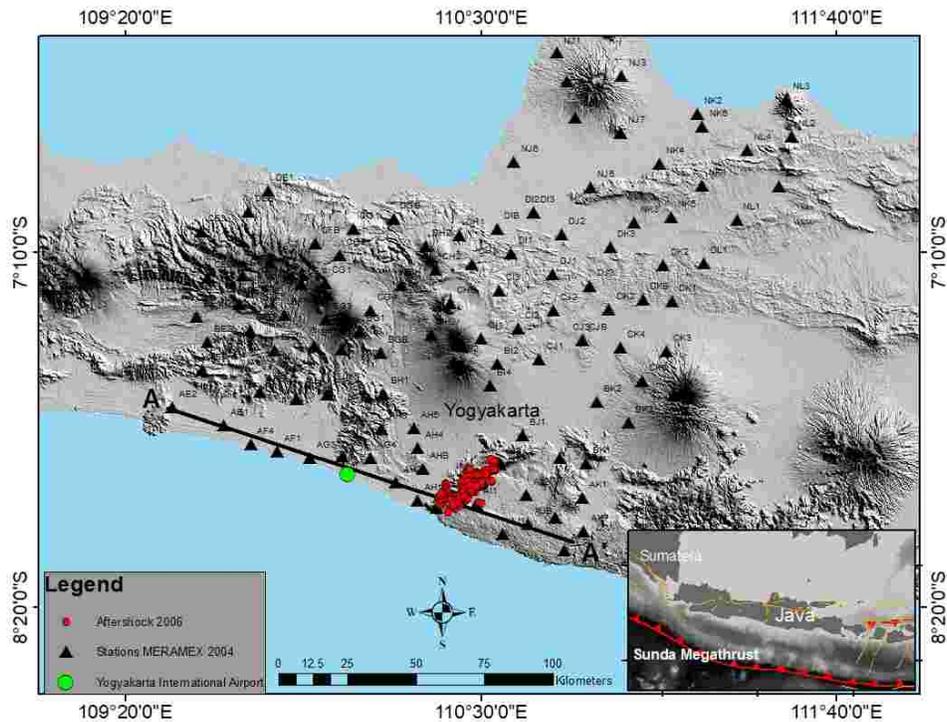
217 Figure 4 shows the coseismic slip distribution of the May 27, 2006, Yogyakarta
218 earthquake along the opaque fault plane overlaid with the aftershock distribution, which is the
219 result of research by Saputra et al. (2021). The aftershock distribution pattern converges in the
220 area of maximum slip. The concentration of slip that collects in the aftershock zone area
221 indicates that there is a continuous release of energy (Sykes, 2021). In the middle of the slip
222 zone, there is a seismic gap that shows the behaviour of the fault. The asperity zone is very
223 important to know as a basis for the initial analysis of earthquake hazards (Corbi et al., 2017;
224 Lay et al., 1981; Abercrombie et al., 2001). If you look at the displacement pattern between the
225 mainshock and aftershock, as shown in Figure 2, it is quite clear that the displacement pattern is
226 different. So it can be concluded that the source of the earthquake came from a different fault
227 area. This shows the complexity of the Opak fault. The slip distribution pattern shown in Figure
228 3 also shows that there is a white colour in the middle of the Opak fault plane. This indicates that
229 this zone is a zone that has not yet released its energy, one day, it can trigger other fault fields
230 around the main opaque fault. Several earthquake events that occurred after 2006 are still widely
231 distributed around the Opak fault. On the western fault boundary, it can be seen that there is a

232 weak field that has not yet released energy. This boundary is very close to the location of
 233 Yogyakarta International Airport.

234

235 4.2 Receiver function data

236 The receiver function method is used to model the subsurface with teleseismic earthquake
 237 events (Amukti, 2019). Figure 5 shows the distribution of MERAMEX stations which were
 238 installed in 2004 with a black triangle symbol, the Bantul earthquake aftershock event in 2006
 239 with a red circle symbol, and also the Jogjakarta International Airport, which is coloured in a
 240 green circle.



242 Figure 5. Aftershock distribution of the 2006 Yogyakarta earthquake recorded from the Meramex
 243 station installed in 2004. The red circles are the distribution of aftershocks. The black
 244 triangle is the distribution of Meramex stations. The red and yellow lines on the map
 245 index are subduction zones and local faults on the mainland, respectively. The black
 246 line is the velocity (V_p) model slice from A-A.'

247

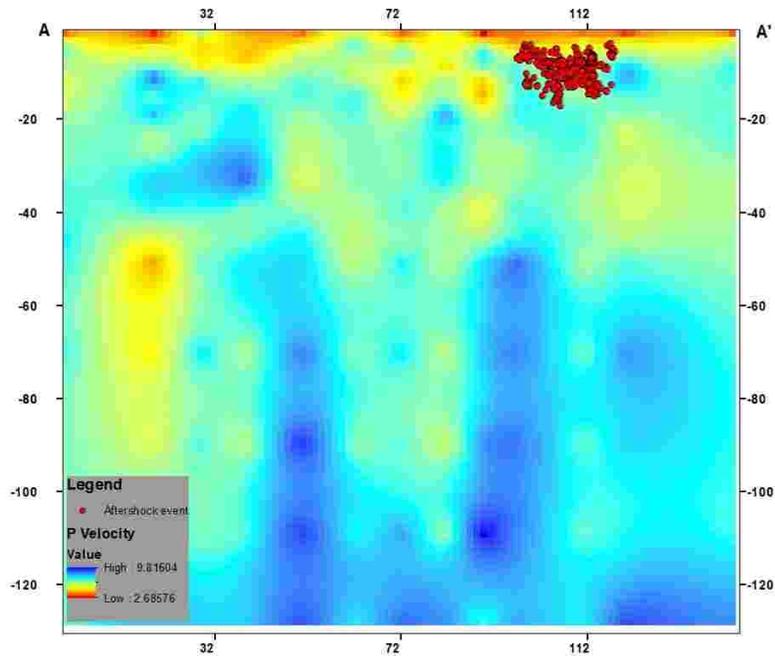
248 The velocity (V_p) results from the receiver function method are modelled to obtain an
 249 overview of the subsurface slices between points A to A' as shown in Figure (6). The aftershock
 250 event is seen at a depth of 5-20 Km and is located in the Low-Velocity Anomaly model. To view
 251 laterally, a velocity model is created at a depth of 11 km (Figure 6).

252

253

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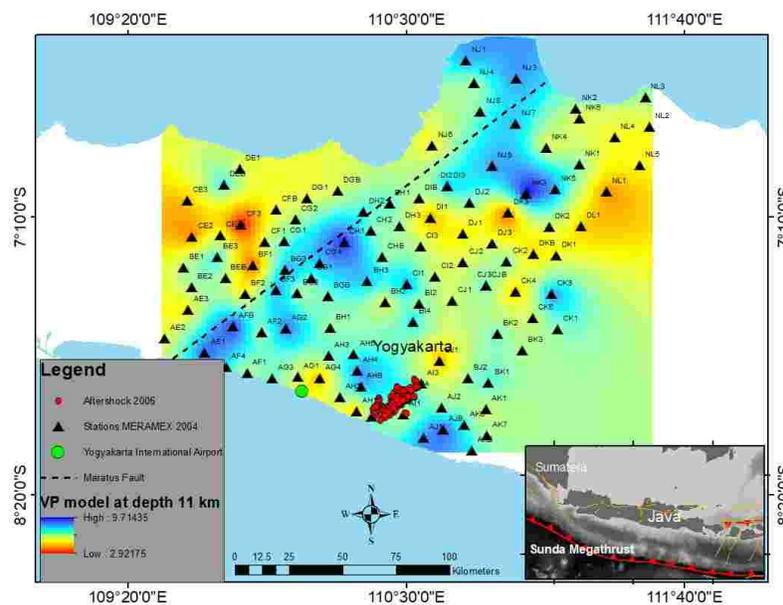
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256 Figure 6. Model Velocity (V_p) of the Receiver Function and combined with the aftershock event
 257 data on the A-A slice. The red circle is the aftershock distribution of the depth function.
 258

259 Figure 7 shows the distribution of P waves in Central Java with a depth of 11 Km. If
 260 further analyzed, it will be seen the distribution of P waves which have high speeds of up to 8
 261 Km/s, will form fragments (Amukti, 2019). These results explain that there is a fault in central
 262 Java known as the Maratus zone, which continues from Borneo Island to Java Island. This
 263 analysis is based on previous research conducted by Wakita (2000), Smith et al. 1 (2005), Hall
 264 and Sevastjanova (2012), Haberland et al. (2014), and Wölbern and Rumpker (2015). Figure 7
 265 also shows that the Velocity P model and aftershock data have a relationship, where the
 266 aftershock event is right in the low anomaly area of velocity in the area around the Opak fault.

267



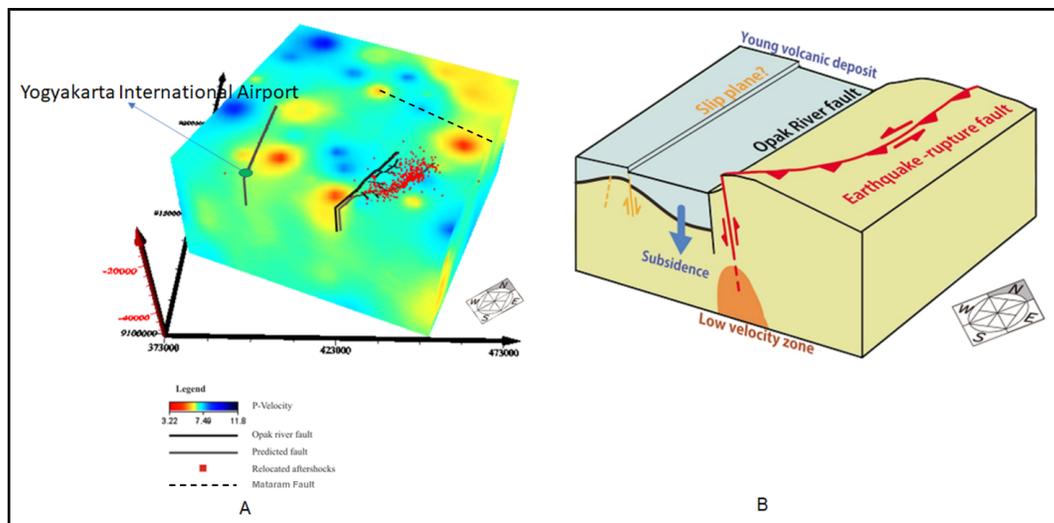
268 Figure 7. Model velocity (V_p) combined with the aftershock event at a depth of 11 Km. The red
 269 circle is the relocated aftershock distribution. The black triangle is the distribution of
 270 Meramex stations. The green circle is the international position of Yogyakarta airport.

271

272 Another interesting thing is that right below Yogyakarta International Airport has a low
 273 anomaly velocity as well, so a more detailed analysis and interpretation is needed in this area, so
 274 a 3-dimensional model was made (Figure 7).

275 Figure 8 shows a reconstructed 3-dimensional model of the Receiver Function and
 276 Aftershock Event, and a geological model is created that can explain this situation. The opaque
 277 fault is seen in the low-velocity anomaly, but from the aftershock event data, it is explained that
 278 the ruptured fault does not persist in the Opak fault but is more to the east and has a curved
 279 shape. To the north of the Opak fault, there is the Mataram fault which was confirmed by the
 280 National Earthquake Center in 2022.

281



282 Figure 8. 3D Model Reconstruction. A. It is a 3-dimensional model of the data velocity (V_p)
 283 receiver function. B. It is a geological model interpretation image of the Opak Fault.

284

285 What is interesting is that to the west, there is a low-speed anomaly that can be
 286 interpreted as a prediction of aircraft errors/slips, and this anomaly is right under the construction
 287 of the Yogyakarta International Airport (Figure 8. a). However, this still needs to be confirmed
 288 with other geophysical research methods such as the Gravity, Electromagnetic and geological
 289 observation accuracy surveys in the field.

290 On May 26, 2006, at 22:54 UTC, Yogyakarta was rocked by an earthquake. Many
 291 scientists debate the sources and mechanisms of how earthquakes occur. The opinion most often
 292 expressed by scientists is the source of the earthquake originated from the Opak fault activity.
 293 The German Task Force (GTF), together with the Seismological Division of the Meteorological
 294 and Geophysics Agency (BMG), undertook the installation of a seismic station around the Opak
 295 Fault to record aftershock events (Walter et al, 2008). The recording results show that the
 296 epicentres of the aftershocks were not aligned along the Opak River Fault but 10 km further to
 297 the east (Walter et al., 2008).

298 Tsuji et al. (2009) observed the Yogyakarta earthquake with SAR interferometry. His
 299 research results show that surface deformation occurred 10 km east of the Opak River Fault

300 which is suspected as the source of the May 2006 event. He modelled a schematic diagram of the
301 relationship between the earthquake fault, the Opak River Fault, and subsidence in young
302 volcanic deposits. The conclusion from the model is that the earthquake displacement has a
303 reverse slip component in addition to a strike slip along the eastern oblique fault plane.

304 In order to prove this, this research conducted modelling of the P wave velocity structure in the
305 area around the Opak River Fault using MERAMEX data which was installed in 2004, and the
306 Yogyakarta earthquake occurred in 2006. To analyze the Opak Fault, we added aftershock data
307 from Anggraini (2014) and Saputra (2021).

308 Geologically, the Opak Fault is an active fault with a long continuity (almost North-
309 South), and its movement is controlled by subduction in the southern part of Java Island. Its
310 position, which is near the surface and intersects urban areas, makes this Fault very dangerous
311 (during an earthquake) because it has a direct impact on humans.

312 The Opak Fault is a fault that has an oblique (horizontal-vertical) movement (Saputra et al.,
313 2021). This oblique movement is indeed common in faults with wide dimensions to
314 accommodate the large energy. It moves in a sinistral direction and is followed by the rising east
315 block. However, when viewed from the surface, this Fault is not clearly visible because it has
316 been covered by young volcanic deposits.

317 Apart from the Opak Fault, right below the location of YIA Airport, there is also an
318 “unidentified” fault. This Fault is thought to still be part of the Opak Fault, with the shape of the
319 Fault almost resembling a half-graben. To the north of the Opak Fault, there is the Mataram fault
320 which was confirmed by the National Earthquake Center in 2022. The Opak Fault and the faults
321 around it are very likely to be found because it is estimated that the Opak Fault is on the
322 continental boundary line at the age of the Oligocene-Miocene (Susilohadi, 2020).

3235 Conclusions

324 Therefore this research was conducted to support one side of the earthquake disaster
325 aspect, which can be analyzed from many events that have occurred around the area. This data
326 shows that the area around Yogyakarta experienced a major earthquake in 2006, which caused
327 many fatalities. The feasibility study has succeeded in reconstructing the source model in detail
328 and analyzing the source mechanism and the rupture process of the May 27, 2006, Yogyakarta
329 earthquake by knowing the coseismic process and the seismotectonic characteristics of the study
330 area. So that a more detailed analysis of the mechanism of the source and the process of
331 earthquake rupture in the research area will be obtained so that the characteristics of the Opak
332 fault will be known in more detail, especially after 2020, Yogyakarta International Airport, which
333 is located in Temon District, Kulon Progo Regency, has been operating.

334 The method used in this study is the aftershock relocation method to determine the
335 aftershock event in the 2006 earthquake that occurred to obtain a rupture zone (Saputra, 2021),
336 then this event data is combined with the subsurface velocity model data resulting from the
337 Receiver Function method (Figure 2) with using the MERAMEX network which was installed in
338 2004 (Amukti, 2019). Figure 4 shows the coseismic slip distribution of the May 27, 2006,
339 Yogyakarta earthquake along the opaque fault plane overlaid with the aftershock distribution,
340 which is the result of research by Saputra et al. (2021). This indicates that this zone is a zone that
341 has not yet released its energy, one day, it can trigger other fault fields around the main opaque
342 fault. Figure 5 shows the distribution of MERAMEX stations which were installed in 2004 with
343 a black triangle symbol, the Bantul earthquake aftershock event in 2006 with a red circle symbol,
344 and also the Jogjakarta International Airport, which is coloured in a green circle. Figure 8 shows

345a reconstructed 3-dimensional model of the Receiver Function and Aftershock Event, and a
346geological model is created that can explain this situation. Opaque faults are seen in the low-
347velocity anomaly, but from the aftershock event data, it is clear that the ruptured fault is not
348exactly at the opaque fault but rather to the east and with a curved shape. What is interesting is
349that to the west, there is a low-velocity anomaly which can be interpreted as a predicted fault/slip
350plane, and this anomaly is right under the construction of the Yogyakarta International Airport.
351To prove this, in this study, a P-wave velocity structure modelling was carried out in the area
352around the Opak River Fault using MERAMEX data installed in 2004 and the Yogyakarta
353earthquake occurred in 2006.

354

355**Acknowledgments**

356The authors want to express our sincere gratitude to the funding agency “Riset Kolaborasi
357Indonesia” for the financial support provided for this research. We also appreciate IRIS, BMKG,
358and GFZ for providing the waveform data used in this study. Our special thanks go to all authors
359namely: RK, SPS, MAA, APS, S, BEN, RL, and SK for their contribution in critical comments,
360performing data curation, and acquisition, examining the results and editing the draft article.
361Lastly, we thank the Postgraduate School of Airlangga University and BRIN for providing
362facilities and support throughout this research project.

363

364**Open Research**

365**Data Availability Statement**

366The Agency of Meteorology, Climatology, and Geophysics of Indonesia maintained the data,
367which was available upon request. Maps were created using Generic Mapping Tools (GMT)
368version 6 (Wessel et al., 2019a, 2019b), licensed under LGPL version 3 or later, available at
369<https://www.genericmapping-tools.org/>.

370

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1 **Investigating the Mainshock and Aftershock of the May 2006 Earthquake in**
2 **Central Java for Aerothropolis Development at Yogyakarta International**
3 **Airport**

4

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29**Key Points:**

- 30 • Analysis of source mechanisms of mainshock and aftershock events of May 27, 2006, earthquake in
31 Yogyakarta, Central Java.
- 32 • Use of receiver function data to model the subsurface velocity P and obtain a geological form model of the
33 Opak fault.

- 34 • Research aimed at supporting the development of the Yogyakarta Aerothropolis area in terms of disaster
35 analysis.

36

37**Abstract**

38Research on the analysis of the source mechanism of the mainshock and aftershock events of the
39May 27, 2006, Yogyakarta earthquake, which is thought to have originated from the Opak fault
40and analysis of receiver function data to model the subsurface velocity P of the Central Java
41subsurface, to obtain a geological form model of the Opak fault. This research aims to support
42the development of the Yogyakarta Aerothropolis area in terms of disaster analysis. The data
43used in this study are remote Teleseismic receiver function data from the MERAMEX station
44installed in 2004, and data for the Bantul earthquake event and its aftershock event in 2006. The
45results obtained from the analysis are that the Yogyakarta area is shaped like a half-graben close
46to Yogyakarta International Airport. The fault that separates the western part of Yogyakarta is still
47not identified. Based on the results of the rupture process analysis of the source along the Opak
48fault plane, some zones have not yet released their energy. The distribution of aftershocks due to
49the mainshock on 27 May 2006 is spread around the Opak fault, which is heading North-South,
50and West-East, which is thought to have activated the minor fault to the east of the Opak fault.
51The opak fault rupture area can be analyzed to have a Low Anomaly velocity P value from the
52receiver function data and is the same as the aftershock event obtained.

53

54Keywords: Rupture, aftershock, receiver function, opak fault, aerothropolis.

55

56**Plain Language Summary**

57This study analyzes the source mechanism of the main earthquake and its aftershocks that
58occurred on May 27, 2006, in Yogyakarta, Indonesia, believed to have originated from the Opak
59fault. The researchers used receiver function data to model the Central Java subsurface velocity P
60and obtained a geological form model of the Opak fault. The study aimed to support the
61development of the Yogyakarta Aerothropolis area in terms of disaster analysis. The results
62showed that the Yogyakarta area is shaped like a half-graben close to Yogyakarta International
63Airport. The fault that separates the western part of Yogyakarta has not yet been identified. The
64distribution of aftershocks due to the mainshock was spread around the Opak fault, which is
65heading North-South and West-East, and it is thought to have activated the minor fault to the east
66of the Opak fault. The study also found that the Opak fault rupture area can be analyzed to have a
67low anomaly velocity P value from the receiver function data and is the same as the aftershock
68event obtained. This study provides valuable information for disaster analysis in the Yogyakarta
69Aerothropolis area.

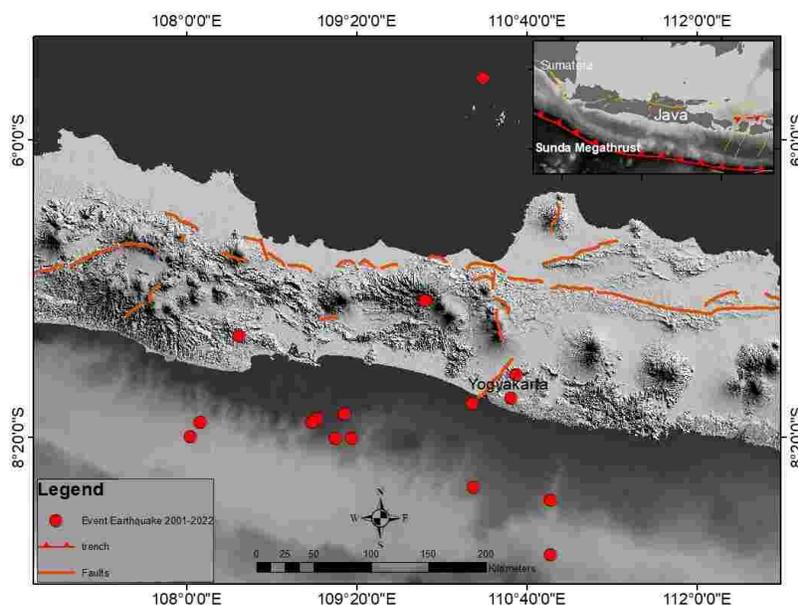
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711 Introduction

72 Yogyakarta is an area that has its uniqueness and charm, so many tourists come, both
 73 local and foreign. This has a positive impact on regional growth. The area around the airport,
 74 with many business activities or commercial services, is the basis for forming the aerotropolis
 75 area concept (Kasarda, 2016). This concept was developed in the Yogyakarta International
 76 Airport area to increase business competitiveness which will have an impact on the community's
 77 economy so that it will have a large effect (Sumarata, 2021). This condition is supported by
 78 proper planning, including security from the threat of disaster.

79 Yogyakarta is not only a popular tourist destination but also an area that is prone to
 80 earthquake disasters due to its location on the active tectonic plate. According to the National
 81 Disaster Management Agency (BNPB), Yogyakarta has a high seismic hazard level, with 21
 82 active faults identified in the region (BNPB, 2021). Therefore, proper disaster risk management
 83 is crucial in aerotropolis planning to minimize the potential impact of earthquakes on the
 84 community's economy and livelihoods. This can be achieved through a comprehensive analysis
 85 of the earthquake risk, including the identification of vulnerable areas and the implementation of
 86 appropriate measures to mitigate the risk. In addition, community participation and awareness-
 87 raising programs should be incorporated into the disaster risk management plan to ensure the
 88 sustainability of the aerotropolis area development. By considering the earthquake risk and
 89 implementing appropriate measures, the aerotropolis area in Yogyakarta can develop sustainably
 90 and contribute to the economic growth of the region while ensuring the safety and well-being of
 91 the local community.

92 The principles of aerotropolis planning consist of regional spatial structure, distance,
 93 zoning, land use, designation of the main functions of the area, integration, and connectivity
 94 (Kurniawan, 2016; Yujin, 2013; Wang, 2013; Greis, 2011). To find out good spatial planning in
 95 aerotropolis planning, an in-depth analysis of disasters in the area is also needed (Surya et al.,
 96 2020). Both seasonal disasters, such as floods and tornadoes, and large, unpredictable disasters,
 97 such as volcanoes, earthquakes, and tsunamis (Ammon et al., 2006). Therefore, this research was
 98 conducted to support one side of the earthquake disaster, which can be analyzed from many
 99 events around the area (Kuncoro et al., 2020).



100 Figure 1. Earthquake distribution map for 2001-2022 around Central Java and Yogyakarta. The
101 red circle is the mainshock distribution that has been relocated. The brown lines are the
102 distribution of faults spread across Java island.

103 The feasibility study has succeeded in reconstructing the source model in detail and
104 analyzing the source mechanism and the rupture process of the May 27, 2006, Yogyakarta
105 earthquake by knowing the coseismic process and the seismotectonic characteristics of the study
106 area. So by knowing the physical characteristics of the May 27, 2006, Yogyakarta earthquake, we
107 can understand the destructive nature of the earthquake, so it can be used as a guide for
108 earthquake mitigation, especially the Yogyakarta earthquake (Saputra et al., 2021). However, to
109 obtain information on the source mechanism and phenomenon of May 27, 2006, Yogyakarta
110 earthquake, it is recommended to use other earthquake data. So that a more detailed analysis of
111 the mechanism of the source and the process of earthquake rupture in the research area will be
112 obtained so that the characteristics of the Opak fault will be known in more detail, especially
113 after 2020, Yogyakarta International Airport, which is located in Temon District, Kulon Progo
114 Regency, has been operating. The airport area will be built with the Aerotropolis concept in its
115 development. In the Yogyakarta Regional Regulation number 5 of 2019 concerning the 2019-
116 2039 Regional Spatial Plan for the Special Region of Yogyakarta, the development of the
117 Temon-Prambanan Area is projected to become one of the province's strategic areas from an
118 economic point of view. One form is the development of the aerotropolis area. Aerotropolis can
119 increase economic growth in a country or region (Peoples, 2014). The concept of the Yogyakarta
120 aerotropolis with a radius of about 15 km from the airport (Syaifuddin et al., 2021).

121 This research aims to analyze the physical characteristics and the source mechanism of
122 the May 27, 2006, Yogyakarta earthquake. The urgency of this study is to provide a better
123 understanding of the destructive nature of the earthquake and to provide guidance for earthquake
124 mitigation, especially in the Yogyakarta region. The results of this study can also provide
125 important information for the development of the aerotropolis area around Yogyakarta
126 International Airport. Understanding the seismotectonic characteristics and potential risks of
127 earthquakes in the area is crucial for ensuring the safety and sustainability of the airport and the
128 surrounding region. Additionally, this research can contribute to the advancement of earthquake
129 science and the development of effective strategies for disaster risk reduction.

1302 **Materials and Methods**

1312.1 **Receiver function data**

132 The method used in this study is the aftershock relocation method to determine the
133 aftershock event in the 2006 earthquake that occurred to obtain a rupture zone (Saputra, 2021).
134 The relocation method was conducted using the hypoDD code (Waldhauser and Ellsworth, 2000)
135 and the double-difference algorithm (Zhang et al., 1889). This event data is combined with the
136 subsurface velocity model data resulting from the Receiver Function method (Figure 2) using the
137 MERAMEX network installed in 2004 (Amukti, 2019). The MERAMEX network is a dense
138 seismographic network that consists of 24 broadband and short-period seismometers, covering an
139 area of approximately 200 km radius around the study area. The subsurface structure is then
140 analyzed using the tomographic inversion method, which allows for a more detailed study of the
141 subsurface velocity structure and can provide insight into the characteristics of the active fault in
142 the study area.

143

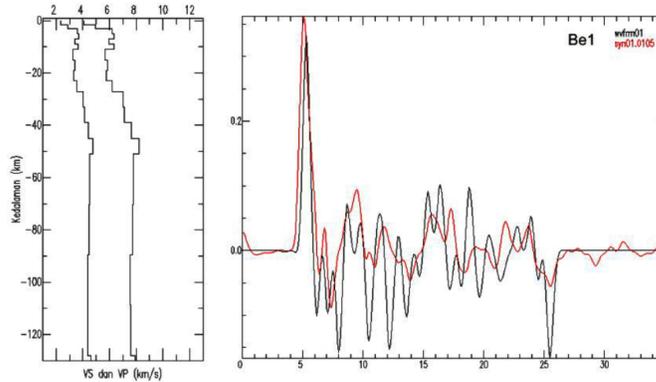


Figure 2. Example of receiver function data

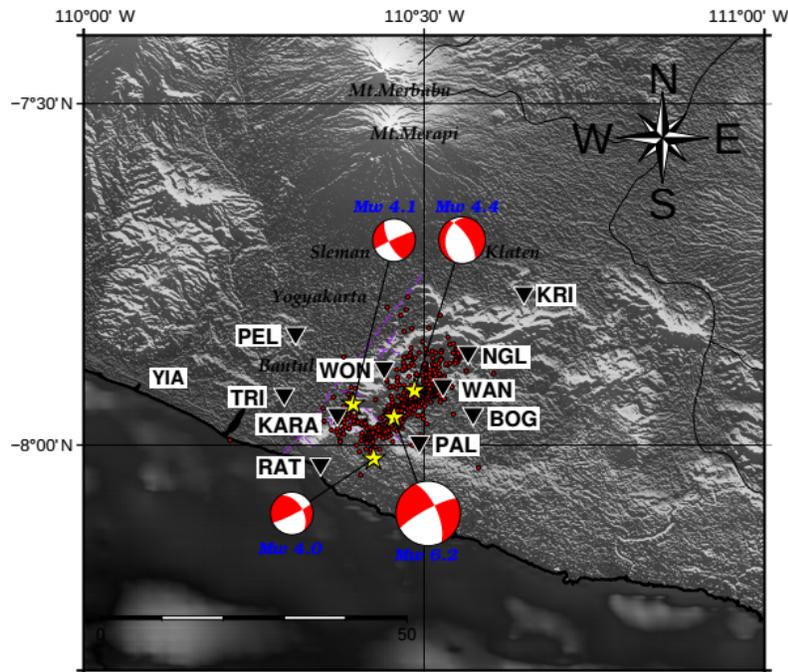
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1472.2 Aftershocks data

148 This study also uses relocated aftershock data obtained from temporary seismometers
 149 installed after the 2006 Yogyakarta mainshock. In this analysis, the study employs the aftershock
 150 relocation method to determine the precise location of each aftershock event and reduce
 151 uncertainty in earthquake location determination.

152



154 Figure 3. Distribution of relocated aftershocks taken from temporary seismometer recordings.

155 The red circle is the relocation aftershock distribution. The grey triangles are
 156 seismometers. The red and white focal ball is the mainshock and aftershock issued by
 157 the USGS.

158

159 Figure 3 shows the enlargement between the mainshock and the three aftershocks, namely on
 160 June 8, 9 and 16, which had different types of displacement. This shows the complexity of the
 161 faults around the Opak fault. The results of the mainshock rupture process analysis in the study

162of Saputra et al. (2021) show two weak zones and can be seen from the distribution of
163aftershocks scattered around the asperity zone (weak zone). Between the weak zones, there is a
164seismic gap, indicated as a zone that has not yet released its energy.

165

1663 Data

167 This study utilized the earthquake aftershock location method to determine the aftershock
168events of the 2006 earthquake in order to obtain the fault zone. The location method was
169performed using the hypoDD code (Waldhauser & Ellsworth, 2000) and the double-difference
170algorithm (Zhang et al., 1889). The earthquake data was combined with subsurface velocity
171model data generated from the Receiver Function method (Figure 2) using the MERAMEX
172network installed in 2004 (Amukti, 2019). The MERAMEX network is a dense seismograph
173network consisting of 24 broad-band and short-period seismometers covering an area of about
174200 km radius around the study area. The subsurface structure was then analyzed using the
175tomographic inversion method, which allows for a more detailed study of the subsurface velocity
176structure and provides insights into the characteristics of active faults in the study area.

177 The aftershock data was also used in this study, obtained from temporary seismometers
178installed after the main earthquake in Yogyakarta in 2006. In this analysis, the study used the
179earthquake location method to determine the precise location of each aftershock event and
180reduce uncertainty in determining the earthquake location. Figure 3 shows the enlargement
181between the main earthquake and three aftershocks on June 8, 9, and 16, which have different
182displacement types. This indicates the complexity of faults around the Opak fault. The results of
183the faulting process analysis in the main earthquake by Saputra et al. (2021) show the presence
184of two weak zones, which can be seen from the distribution of aftershocks scattered around the
185asperity zone (weak zone). Between the weak zones, there is a seismic gap, which is indicated as
186a zone that has not yet released its energy.

187 In this study, the Receiver Function method was used to generate the subsurface velocity
188model, which was then combined with the aftershock location data to obtain insights into the
189characteristics of active faults in the study area. The tomographic inversion method was used to
190obtain more detailed information about the subsurface structure. In addition, the earthquake
191location method was also used to determine the precise location of each aftershock event, which
192helps in understanding the complexity of faults around the Opak fault. The results of this study
193show the presence of two weak zones and a seismic gap between the weak zones. These results
194can provide important information for understanding the potential earthquake hazards in the
195region.

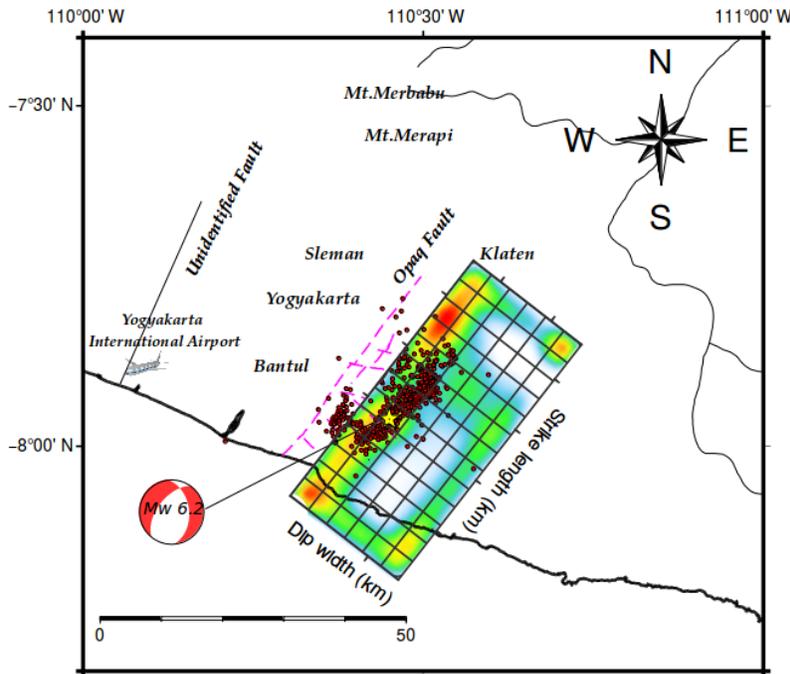
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1974 Results, or a descriptive heading about the results

1984.1 Analysis of the earthquake source rupture process

199 Analyzing the earthquake source rupture process is critical to seismological research, as it
200helps understand the mechanisms that trigger earthquakes. In this study, the researchers used the
201finite fault inversion method to analyze the earthquake source rupture process of the Yogyakarta
202earthquake. The results showed that the rupture process was a unilateral propagation from the
203southwest to the northeast. Moreover, the fault length was approximately 28 km, and the
204maximum slip was estimated to be around 1.4 meters. The results are crucial in improving our

205 understanding of the characteristics of the Yogyakarta earthquake and provide valuable
 206 information for developing effective earthquake mitigation strategies.



207 Figure 4. The results of the analysis of the rupture process of the Yogyakarta earthquake on May
 208 27, 2006, show the slip distribution from the opaque fault plane. The rectangular
 209 contour is the distribution of slip on the opaque fault plane. The red colour circle is the
 210 aftershock distribution. The yellow star is the Yogyakarta earthquake mainshock,
 211 resulting from a joint inversion research by Saputra et al. (2021). The purple dotted line
 212 is an opaque fault, and on the east side of the main fault, there are several minor faults.
 213 The black line is the boundary of the graben zone, where the fault has not been
 214 identified. The red and white focal ball results from a joint inversion calculation from
 215 Saputra et al. 2021 study (modification from Saputra et al. Research, 2021).
 216

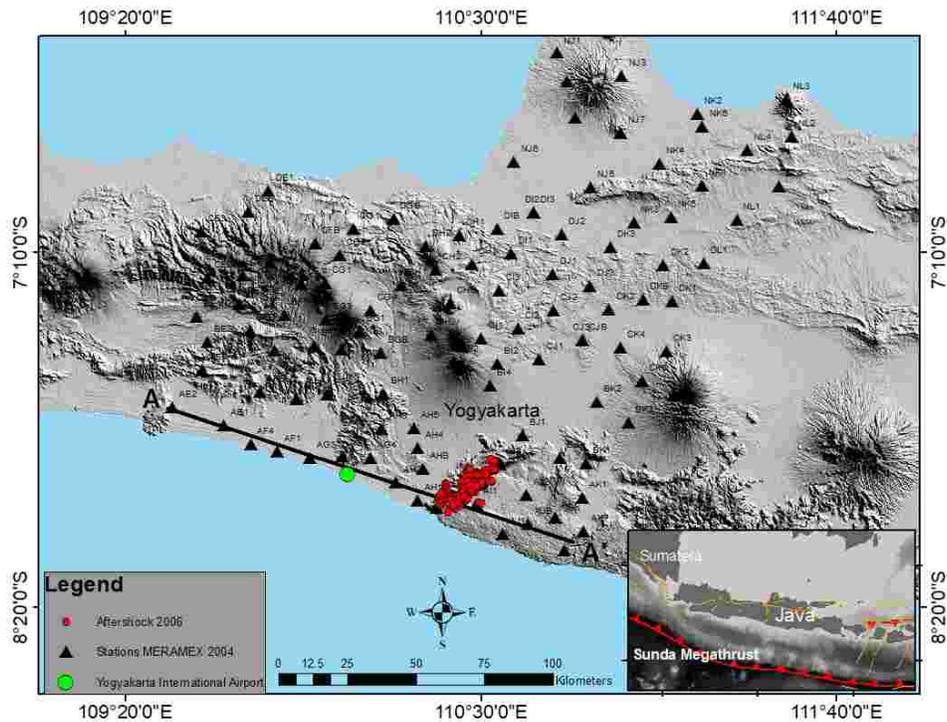
217 Figure 4 shows the coseismic slip distribution of the May 27, 2006, Yogyakarta
 218 earthquake along the opaque fault plane overlaid with the aftershock distribution, which is the
 219 result of research by Saputra et al. (2021). The aftershock distribution pattern converges in the
 220 area of maximum slip. The concentration of slip that collects in the aftershock zone area
 221 indicates that there is a continuous release of energy (Sykes, 2021). In the middle of the slip
 222 zone, there is a seismic gap that shows the behaviour of the fault. The asperity zone is very
 223 important to know as a basis for the initial analysis of earthquake hazards (Corbi et al., 2017;
 224 Lay et al., 1981; Abercrombie et al., 2001). If you look at the displacement pattern between the
 225 mainshock and aftershock, as shown in Figure 2, it is quite clear that the displacement pattern is
 226 different. So it can be concluded that the source of the earthquake came from a different fault
 227 area. This shows the complexity of the Opak fault. The slip distribution pattern shown in Figure
 228 3 also shows that there is a white colour in the middle of the Opak fault plane. This indicates that
 229 this zone is a zone that has not yet released its energy, one day, it can trigger other fault fields
 230 around the main opaque fault. Several earthquake events that occurred after 2006 are still widely
 231 distributed around the Opak fault. On the western fault boundary, it can be seen that there is a

232 weak field that has not yet released energy. This boundary is very close to the location of
 233 Yogyakarta International Airport.

234

235 4.2 Receiver function data

236 The receiver function method is used to model the subsurface with teleseismic earthquake
 237 events (Amukti, 2019). Figure 5 shows the distribution of MERAMEX stations which were
 238 installed in 2004 with a black triangle symbol, the Bantul earthquake aftershock event in 2006
 239 with a red circle symbol, and also the Jogjakarta International Airport, which is coloured in a
 240 green circle.



242 Figure 5. Aftershock distribution of the 2006 Yogyakarta earthquake recorded from the Meramex
 243 station installed in 2004. The red circles are the distribution of aftershocks. The black
 244 triangle is the distribution of Meramex stations. The red and yellow lines on the map
 245 index are subduction zones and local faults on the mainland, respectively. The black
 246 line is the velocity (V_p) model slice from A-A.'

247

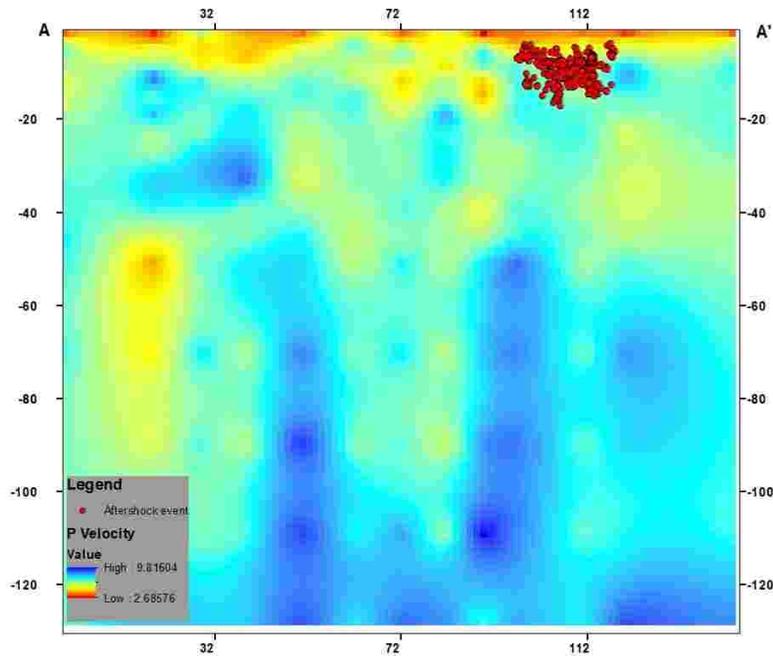
248 The velocity (V_p) results from the receiver function method are modelled to obtain an
 249 overview of the subsurface slices between points A to A' as shown in Figure (6). The aftershock
 250 event is seen at a depth of 5-20 Km and is located in the Low-Velocity Anomaly model. To view
 251 laterally, a velocity model is created at a depth of 11 km (Figure 6).

252

253

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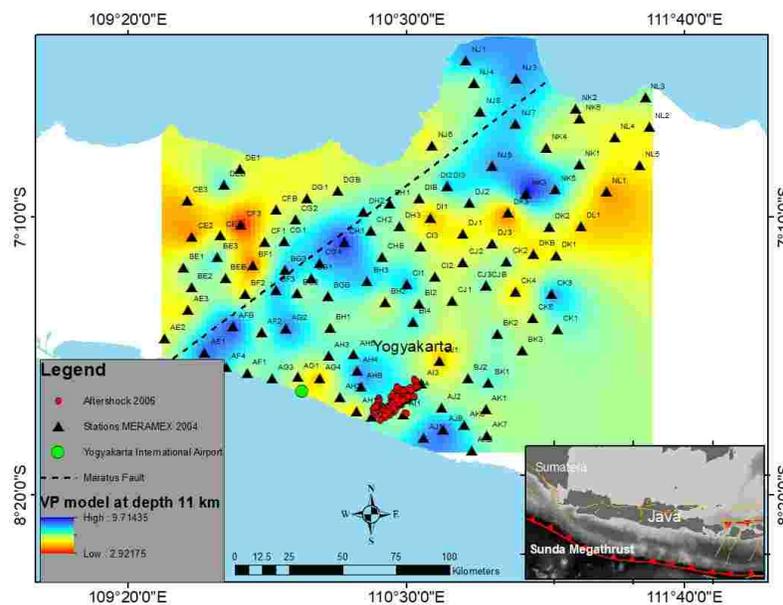
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256 Figure 6. Model Velocity (V_p) of the Receiver Function and combined with the aftershock event
 257 data on the A-A slice. The red circle is the aftershock distribution of the depth function.
 258

259 Figure 7 shows the distribution of P waves in Central Java with a depth of 11 Km. If
 260 further analyzed, it will be seen the distribution of P waves which have high speeds of up to 8
 261 261 Km/s, will form fragments (Amukti, 2019). These results explain that there is a fault in central
 262 Java known as the Maratus zone, which continues from Borneo Island to Java Island. This
 263 analysis is based on previous research conducted by Wakita (2000), Smith et al. 1 (2005), Hall
 264 and Sevastjanova (2012), Haberland et al. (2014), and Wölbern and Rumpker (2015). Figure 7
 265 also shows that the Velocity P model and aftershock data have a relationship, where the
 266 aftershock event is right in the low anomaly area of velocity in the area around the Opak fault.

267



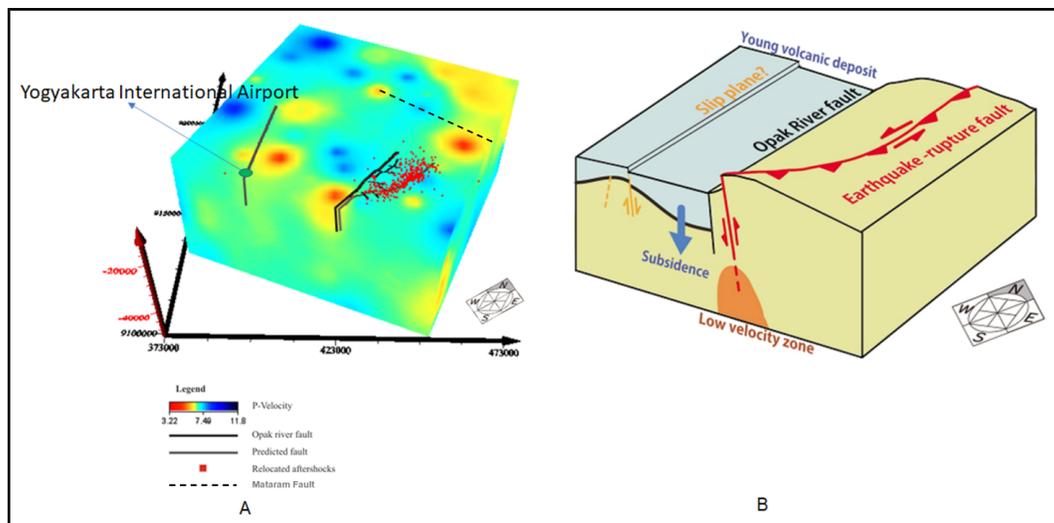
268 Figure 7. Model velocity (V_p) combined with the aftershock event at a depth of 11 Km. The red
 269 circle is the relocated aftershock distribution. The black triangle is the distribution of
 270 Meramex stations. The green circle is the international position of Yogyakarta airport.

271

272 Another interesting thing is that right below Yogyakarta International Airport has a low
 273 anomaly velocity as well, so a more detailed analysis and interpretation is needed in this area, so
 274 a 3-dimensional model was made (Figure 7).

275 Figure 8 shows a reconstructed 3-dimensional model of the Receiver Function and
 276 Aftershock Event, and a geological model is created that can explain this situation. The opaque
 277 fault is seen in the low-velocity anomaly, but from the aftershock event data, it is explained that
 278 the ruptured fault does not persist in the Opak fault but is more to the east and has a curved
 279 shape. To the north of the Opak fault, there is the Mataram fault which was confirmed by the
 280 National Earthquake Center in 2022.

281



282 Figure 8. 3D Model Reconstruction. A. It is a 3-dimensional model of the data velocity (V_p)
 283 receiver function. B. It is a geological model interpretation image of the Opak Fault.

284

285 What is interesting is that to the west, there is a low-speed anomaly that can be
 286 interpreted as a prediction of aircraft errors/slips, and this anomaly is right under the construction
 287 of the Yogyakarta International Airport (Figure 8. a). However, this still needs to be confirmed
 288 with other geophysical research methods such as the Gravity, Electromagnetic and geological
 289 observation accuracy surveys in the field.

290 On May 26, 2006, at 22:54 UTC, Yogyakarta was rocked by an earthquake. Many
 291 scientists debate the sources and mechanisms of how earthquakes occur. The opinion most often
 292 expressed by scientists is the source of the earthquake originated from the Opak fault activity.
 293 The German Task Force (GTF), together with the Seismological Division of the Meteorological
 294 and Geophysics Agency (BMG), undertook the installation of a seismic station around the Opak
 295 Fault to record aftershock events (Walter et al, 2008). The recording results show that the
 296 epicentres of the aftershocks were not aligned along the Opak River Fault but 10 km further to
 297 the east (Walter et al., 2008).

298 Tsuji et al. (2009) observed the Yogyakarta earthquake with SAR interferometry. His
 299 research results show that surface deformation occurred 10 km east of the Opak River Fault

300which is suspected as the source of the May 2006 event. He modelled a schematic diagram of the
301relationship between the earthquake fault, the Opak River Fault, and subsidence in young
302volcanic deposits. The conclusion from the model is that the earthquake displacement has a
303reverse slip component in addition to a strike slip along the eastern oblique fault plane.

304In order to prove this, this research conducted modelling of the P wave velocity structure in the
305area around the Opak River Fault using MERAMEX data which was installed in 2004, and the
306Yogyakarta earthquake occurred in 2006. To analyze the Opak Fault, we added aftershock data
307from Anggraini (2014) and Saputra (2021).

308 Geologically, the Opak Fault is an active fault with a long continuity (almost North-
309South), and its movement is controlled by subduction in the southern part of Java Island. Its
310position, which is near the surface and intersects urban areas, makes this Fault very dangerous
311(during an earthquake) because it has a direct impact on humans.

312The Opak Fault is a fault that has an oblique (horizontal-vertical) movement (Saputra et al.,
3132021). This oblique movement is indeed common in faults with wide dimensions to
314accommodate the large energy. It moves in a sinistral direction and is followed by the rising east
315block. However, when viewed from the surface, this Fault is not clearly visible because it has
316been covered by young volcanic deposits.

317 Apart from the Opak Fault, right below the location of YIA Airport, there is also an
318“unidentified” fault. This Fault is thought to still be part of the Opak Fault, with the shape of the
319Fault almost resembling a half-graben. To the north of the Opak Fault, there is the Mataram fault
320which was confirmed by the National Earthquake Center in 2022. The Opak Fault and the faults
321around it are very likely to be found because it is estimated that the Opak Fault is on the
322continental boundary line at the age of the Oligocene-Miocene (Susilohadi, 2020).

3235 Conclusions

324 Therefore this research was conducted to support one side of the earthquake disaster
325aspect, which can be analyzed from many events that have occurred around the area. This data
326shows that the area around Yogyakarta experienced a major earthquake in 2006, which caused
327many fatalities. The feasibility study has succeeded in reconstructing the source model in detail
328and analyzing the source mechanism and the rupture process of the May 27, 2006, Yogyakarta
329earthquake by knowing the coseismic process and the seismotectonic characteristics of the study
330area. So that a more detailed analysis of the mechanism of the source and the process of
331earthquake rupture in the research area will be obtained so that the characteristics of the Opak
332fault will be known in more detail, especially after 2020, Yogyakarta International Airport, which
333is located in Temon District, Kulon Progo Regency, has been operating.

334 The method used in this study is the aftershock relocation method to determine the
335aftershock event in the 2006 earthquake that occurred to obtain a rupture zone (Saputra, 2021),
336then this event data is combined with the subsurface velocity model data resulting from the
337Receiver Function method (Figure 2) with using the MERAMEX network which was installed in
3382004 (Amukti, 2019). Figure 4 shows the coseismic slip distribution of the May 27, 2006,
339Yogyakarta earthquake along the opaque fault plane overlaid with the aftershock distribution,
340which is the result of research by Saputra et al. (2021). This indicates that this zone is a zone that
341has not yet released its energy, one day, it can trigger other fault fields around the main opaque
342fault. Figure 5 shows the distribution of MERAMEX stations which were installed in 2004 with
343a black triangle symbol, the Bantul earthquake aftershock event in 2006 with a red circle symbol,
344and also the Jogjakarta International Airport, which is coloured in a green circle. Figure 8 shows

345a reconstructed 3-dimensional model of the Receiver Function and Aftershock Event, and a
346geological model is created that can explain this situation. Opaque faults are seen in the low-
347velocity anomaly, but from the aftershock event data, it is clear that the ruptured fault is not
348exactly at the opaque fault but rather to the east and with a curved shape. What is interesting is
349that to the west, there is a low-velocity anomaly which can be interpreted as a predicted fault/slip
350plane, and this anomaly is right under the construction of the Yogyakarta International Airport.
351To prove this, in this study, a P-wave velocity structure modelling was carried out in the area
352around the Opak River Fault using MERAMEX data installed in 2004 and the Yogyakarta
353earthquake occurred in 2006.

354

355**Acknowledgments**

356The authors want to express our sincere gratitude to the funding agency “Riset Kolaborasi
357Indonesia” for the financial support provided for this research. We also appreciate IRIS, BMKG,
358and GFZ for providing the waveform data used in this study. Our special thanks go to all authors
359namely: RK, SPS, MAA, APS, S, BEN, RL, and SK for their contribution in critical comments,
360performing data curation, and acquisition, examining the results and editing the draft article.
361Lastly, we thank the Postgraduate School of Airlangga University and BRIN for providing
362facilities and support throughout this research project.

363

364**Open Research**

365**Data Availability Statement**

366The Agency of Meteorology, Climatology, and Geophysics of Indonesia maintained the data,
367which was available upon request. Maps were created using Generic Mapping Tools (GMT)
368version 6 (Wessel et al., 2019a, 2019b), licensed under LGPL version 3 or later, available at
369<https://www.genericmapping-tools.org/>.

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