

# Non-monotonic trend analysis using Mann-Kendall with self-quantiles

Sadık Alashan<sup>1</sup>

<sup>1</sup>Bingol University

February 27, 2023

## Abstract

Recently, climate change makes itself felt at increasing levels due to rising temperatures, irregular precipitation patterns and changing weather events. Although the frequently used Mann-Kendall (MK) method has disadvantages such as needing serial independence, it helps to detect monotonic trends to investigate climate change effects on a given time series. Climate change may have different features on different levels such as the lows and highs of a given time series, leading to non-monotonic trends. Innovative trend analysis (ITA) as an innovative trend analysis method detects non-monotonic trends, which MK cannot. In this study, MK method is improved to detect non-monotonic trends (non-monotonic MK) and applied for Murat River basin, a branch of Euphrates River, precipitation series at Bingöl, Muş, and Ağrı meteorological stations. Although classical MK method cannot detect any trend on the river basin, non-monotonic MK (NMK) method detects two important decreasing (increasing) trends on the low (high) values of Bingöl and Muş (Bingöl) stations. Also, stationarity analysis is applied through the statistical significance level concept for the river basin precipitation series using the NMK method. Bingöl station has a non-stationary precipitation series with a value of 3.07 and 95% confidence level, while Muş station has a remarkable value of 1.58, Ağrı station conserves its stationarity characteristic on the precipitation series. It is hoped that the newly developed NMK method will help to understand the effects of climate change on hydro-meteorological historical records and predict future events for more efficient hydraulic structure designs.

## Hosted file

956693\_0\_art\_file\_10723216\_rqb320.docx available at <https://authorea.com/users/588698/articles/625840-non-monotonic-trend-analysis-using-mann-kendall-with-self-quantiles>

# Non-monotonic trend analysis using Mann-Kendall with self-quantiles

Sadık Alashan

<sup>1</sup> Civil Engineering Department, Bingöl University, Bingöl, Turkey, Email: sadikalashan@bingol.edu.tr

**Abstract:** Recently, climate change makes itself felt at increasing levels due to rising temperatures, irregular precipitation patterns and changing weather events. Although the frequently used Mann-Kendall (MK) method has disadvantages such as needing serial independence, it helps to detect monotonic trends to investigate climate change effects on a given time series. Climate change may have different features on different levels such as the lows and highs of a given time series, leading to non-monotonic trends. Innovative trend analysis (ITA) as an innovative trend analysis method detects non-monotonic trends, which MK cannot. In this study, MK method is improved to detect non-monotonic trends (non-monotonic MK) and applied for Murat River basin, a branch of Euphrates River, precipitation series at Bingöl, Muş, and Ağrı meteorological stations. Although classical MK method cannot detect any trend on the river basin, non-monotonic MK (NMK) method detects two important decreasing (increasing) trends on the low (high) values of Bingöl and Muş (Bingöl) stations. Also, stationarity analysis is applied through the statistical significance level concept for the river basin precipitation series using the NMK method. Bingöl station has a non-stationary precipitation series with a  $z_{NMK}$  value of 3.07 and 95% confidence level, while Muş station has a remarkable  $z_{NMK}$  value of 1.58, Ağrı station conserves its stationarity characteristic on the precipitation series. It is hoped that the newly developed NMK method will help to understand the effects of climate change on hydro-meteorological historical records and predict future events for more efficient hydraulic structure designs.

**Keywords:** Climate change, Mann-Kendall, hydro-meteorology, trend, non-monotonic.

## 25    **Introduction**

26            Climate change affects hydro-meteorological events with increasing temperatures  
27    dependent on greenhouse gases level increments in the atmosphere. Hydro-meteorological  
28    series have trend components as well as random features. To calculate trends on a time series,  
29    linear regression analysis (Haan, 1977), Mann-Kendall (Kendall, 1975; Mann, 1945), Sen  
30    slope estimator (Sen, 1968), Spearman's Rho (Spearman, 1987), and lastly Şen innovative  
31    trend analyses (ITA) methods are used in the literature (Şen, 2012). Although Mann-Kendall  
32    (MK) method is the most widely used method in the literature, it has some restrictive  
33    assumptions such as normality and independence of time series variable. Dependent series  
34    may tend to give significant trend components in the MK method even if there is no trend  
35    (Bayazit and Önöz, 2007; Cox and Stuart, 1955; Hamed and Ramachandra Rao, 1998).

36            Non-monotonic trends (low and high values have different positive or negative trend  
37    directions in a time series) are lastly come into question with Şen ITA method (Şen-ITA).  
38    Some researchers (Alifujiang et al., 2021; Berhail et al., 2022; Dabanli et al., 2021; Şan et al.,  
39    2021; Saplıoğlu and Güçlü, 2022; Şişman and Kizilöz, 2021) claim that Şen-ITA method is  
40    superior to MK because of that MK method cannot detect non-monotonic trends on time  
41    series while Şen-ITA can. Research results support these claims using visual evaluations on a  
42    1:1 graph. Also, Faulkner et al., (2020) examined flood flows in Great Britain using a non-  
43    stationary distribution, and MK methods, the non-stationary distribution is preferred at 68 out  
44    of 166 stations where MK cannot detect a trend. Nigussie and Altunkaynak, (2019) apply MK  
45    method to detect trends of extreme rainfall indices. Although the rainfall indices have  
46    increasing trends, MK method does not find significant trends even at 0.10 significance  
47    levels.

48            Under the light of increasing carbon content of the atmosphere, it is often questioned  
49    whether hydro-meteorological variables, whose trends are not classified as monotonic and

50 non-monotonic, have stationary properties. This discussion cannot be conducted effectively  
51 without the aforementioned trend classifications. Some researchers argue that stationarity is  
52 dead and not valid for future event predictions (Khaliq et al., 2006; Milly et al., 2008; Vogel  
53 et al., 2011). The other researchers reject this idea and defend stationarity because it is more  
54 understandable and applicable (Bayazit, 2015; Lins and Cohn, 2011; Matalas, 2012;  
55 Montanari and Koutsoyiannis, 2014; Serinaldi and Kilsby, 2015). In the non-monotonic trend  
56 conditions, the explanations of the first research group are reasonable because non-monotonic  
57 trends distort distribution parameters of variables and estimation of these parameters is hard  
58 for future events. The comments of the second group are accepted for the monotonic trend  
59 conditions because monotonic trends change distribution parameters regularly and these  
60 changes can be guessed for future events and relatively easier than non-monotonic trend  
61 calculations. Non-monotonic trends cannot be detected by classical trend methods because  
62 there are different effects climatic impacts on low and high hydro-meteorological variables.  
63 Alashan, (2018) proposes non-stationary cumulative functions to forecast future hydro-  
64 meteorological events effectively. Generally, non-monotonic trends do not change total  
65 amount of hydro-meteorological variables. Burić and Doderović, (2021) stated that the  
66 climate of Podgorica (Montenegro) is getting more arid but total precipitation has not  
67 significant trends. Trends may not be a sign of non-stationarity (Oruc, 2021; Serinaldi et al.,  
68 2018).

69 In this study, classical MK method is revised to detect non-monotonic trends (NMK),  
70 which enables more effective use of the classical MK method. NMK method separately  
71 detects trends on low and high values of a time series according to statistical significance  
72 levels. The same method can group series as stationary and non-stationary using basic MK  
73 processes. It is a novel and helpful approach to detect distorted hydrologic series by the  
74 climate change.

## 75 Methodology

76 Mann-Kendall (MK) is a non-parametric method that cannot detect non-monotonic  
 77 trends on hydro-meteorological data (Kendall, 1948; Mann, 1945). A new non-monotonic  
 78 Mann-Kendall (NMK) method is launched in the literature to use the traditional MK  
 79 effectively in this study. The revised method can be used simply to detect non-monotonic  
 80 hidden trends in low and high values of a time series. NMK method has the following steps.

- 81 • A time series ( $X$ ) is separated into two times sub-series such as low sub-series ( $X_{low}$ )  
 82 and high sub-series ( $X_{high}$ ) with a rank number. For example, a time series ( $X$ ) consists  
 83 of  $x_1, x_2, \dots, x_n$ . If  $x_1$  is greater than the mean of the time series, it is classified  
 84 as a high value otherwise a low value. This process is repeated throughout the time  
 85 series (Figure 1).
- 86 • These low and high values are recorded into a low ( $X_{low}$ ) and high ( $X_{high}$ ) sub-series  
 87 with their rank numbers.
- 88 • If the low and high sub-series can be considered independent and thus the classical  
 89 MK can be applied to these series. If they are dependent, pre-whitening, over-  
 90 whitening, and variance correction can be applied despite some of their disadvantages.
- 91 • The data numbers of the low and high sub-series are  $n_l$  and  $n_h$ . MK test statistics,  
 92  $S_{low}$ , can be calculated with Eq. 1, where,  $j > i$  and  $i$  from 1 to  $n_l - 1$  and  $j$  from  
 93  $i + 1$  to  $n_l$ . The  $sgn$  function in Eq. 2 represents the signal function. If  $X_{low,j} >$   
 94  $X_{low,i}$  then  $sgn(X_{low,j} - X_{low,i}) = 1$ , if  $X_{low,j} < X_{low,i}$  then  $sgn(X_{low,j} - X_{low,i}) =$   
 95  $-1$ , otherwise 0. The same processes are applied to the high sub-series and all series.

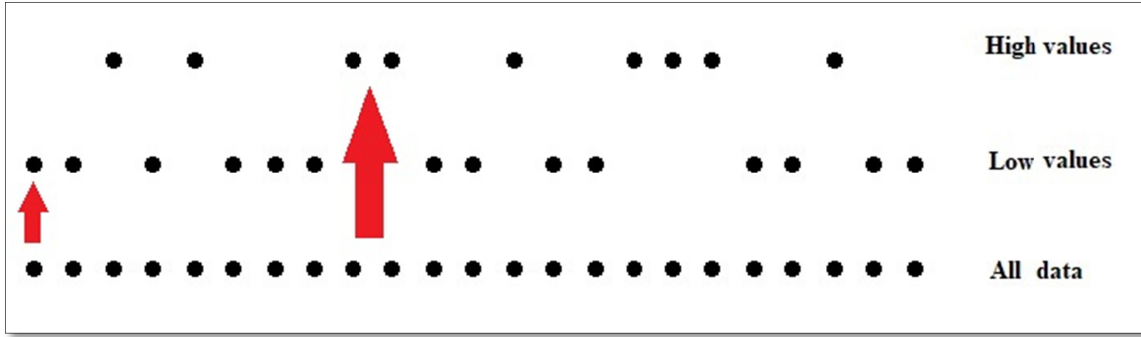
$$96 \quad S_{low} = \sum_{i=1}^{n_l-1} \sum_{j=i+1}^{n_l} sgn(X_{low,j} - X_{low,i}) \quad (1)$$

$$97 \quad sgn(X_{low,j} - X_{low,i}) = \begin{cases} 1 & \text{for } X_{low,j} > X_{low,i} \\ -1 & \text{for } X_{low,j} < X_{low,i} \\ 0 & \text{for } X_{low,j} = X_{low,i} \end{cases} \quad (2)$$

- According to  $H_0$  null hypothesis, the expectation value of the  $S_{low}$  parameter is zero (Eq. 3). Variance of the  $S_{low}$  test statistics is given in Eq. 4. NMK standardized test statistics,  $z_{low}$ , is calculated according to Eq. 5. If an absolute standardized  $z_{low}$  value is bigger than tabulated  $z$  values (1.65, 1.96 or 2.58) according to certain confidence levels (90%, 95% or 99%) then there is an increasing or a decreasing trend and  $H_0$ , null hypothesis, is rejected.  $E(S_{low}) = 0$

$$Var(S_{low}) = \frac{n_l(n_l-1)(2n_l+5)}{18} \quad (4)$$

$$z_{low} = \begin{cases} \frac{S_{low}-1}{Var(S_{low})} & \text{if } S_{low} > 0 \\ 0 & \text{if } S_{low} = 0 \\ \frac{S_{low}+1}{Var(S_{low})} & \text{if } S_{low} < 0 \end{cases} \quad (5)$$



**Figure 1** Separation of the main series into low and high sub-series to apply the non-monotonic MK method.

Although MK is a distribution-free test that is not affected by the true distribution of the data, the  $S$  parameter approaches the normal (Gaussian) probability distribution function (PDF) as the data length gets larger (Alashan, 2020; Hamed, 2009; Kendall and Gibbons, 1990). In the non-monotonic Mann-Kendall (NMK) method, there are two  $S$  parameters such as  $S_{low}$  and  $S_{high}$ . According to MK method, the null hypothesis,  $H_0$ , is accepted when  $S_{low}$  and  $S_{high}$  are close to zero. Thus, here is no trend and the mean of examined data is almost constant.

117 If the values of  $S_{low}$  and  $S_{high}$  are sufficiently different from the zero then there is a  
 118 trend and the null hypothesis is rejected. For  $S_{low} \approx S_{high}$ , the mean of the dataset changes  
 119 but the variances are almost constant. This situation is called as the stationary process and the  
 120 distribution of data is similar to the previous one.

121 If the values of  $S_{low}$  and  $S_{high}$  are statistically different from zero and each other,  
 122 thus the mean and variance of data change and the dataset is non-stationary. If the values of  
 123  $S_{low}$  and  $S_{high}$  have opposite signs, close to each other and sufficiently different from zero  
 124 ( $S_{low} \approx -S_{high}$ ), then there is sometimes no monotonic trend on the dataset and MK cannot  
 125 detect non-monotonic trends. Non-stationary conditions are valid for the dataset, and the  
 126 distribution of the data changes over time. MK should be modified to detect non-stationary  
 127 situations and calculate non-monotonic trends on a given time series.

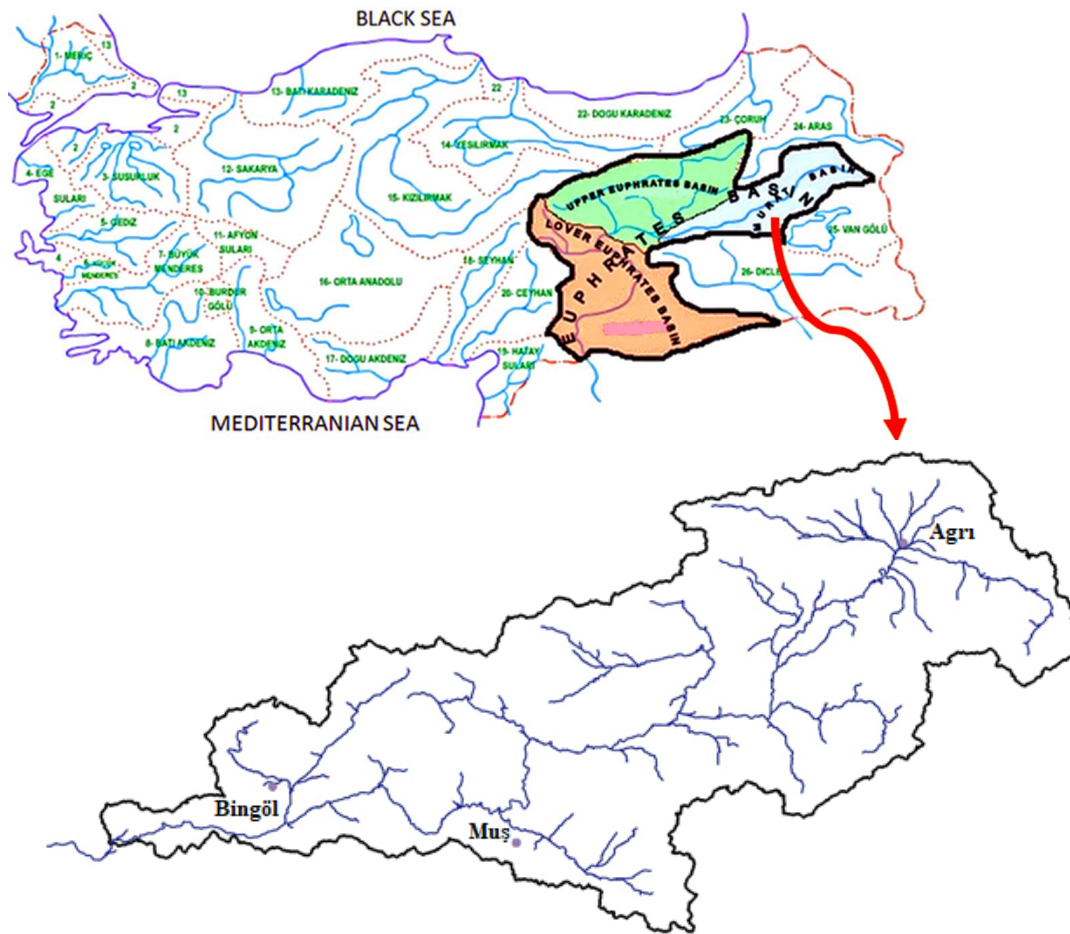
128 In the stationary conditions, the values of  $S_{low}$  and  $S_{high}$  must be close to each  
 129 other ( $S_{high} - S_{low} \approx 0$ ). Non-monotonic trend conditions are valid when the values of  
 130  $S_{high} - S_{low}$  increase. To overcome this problem, a new test statistic ( $z_{NMK}$ ) must be  
 131 launched for MK method. As a main rule, if the  $S_{low}$  and  $S_{high}$  parameter values have a  
 132 normal PDF separately then the sum and difference of these parameters must also have a  
 133 normal PDF (Cramér, 1936). According to this rule, the  $z_{NMK}$  parameter has a normal PDF  
 134 and can be calculated according to Eq. 6. If the calculated  $z_{NMK}$  values are bigger than  
 135 tabulated z values (1.65, 1.96, and 2.58) according to statistical confidence levels (90%, 95%,  
 136 or 99%), then there is a non-monotonic trend in the dataset. For the first time, this approach  
 137 (NMK) allows us to detect non-monotonic trends in time series data using the MK method.

$$138 \quad z_{NMK} = \frac{S_{high} - S_{low}}{\sqrt{Var(S_{high}) + Var(S_{low})}} \quad (6)$$

139

## Study Area and Application

In this study, the Murat River is selected as the study area (Figure 2). It is a branch of the Euphrates River and located in eastern Turkey. The river basin ranges from northeastern to southwestern and has approximately 28,850 km<sup>2</sup> area (Güneş, 2011). The elevation of the river basin changes from 819 m to 4,038 m above mean sea level with a 1,834 m average altitude (Alashan et al., 2016).



**Figure 2** The map of the study area (Murat River Basin).

The dataset used in this study on a 0.5°x0.5° grid produced by the University of East Anglia Climate Research Unit (CRU) is taken from the World Bank Group, Climate Change Knowledge Portal (<https://climateknowledgeportal.worldbank.org/download-data>). The

observed dataset covers the period 1901-2020 and quality controlled (Harris et al., 2020).

Statistical features of the selected stations are given in Table 1.

**Table 1.** Statistical parameters of the selected stations annual total precipitation (mm)

Stations	Latitude	Longitude	Observed Years	Minimum (mm)	Mean (mm)	Maximum (mm)	Standard deviation (mm)	Skewness
Bingöl	38°53'N	40°29'E	1901-2020	505.24	765.38	1146.13	112.08	0.54
Muş	38°44'N	41°30'E	1901-2020	469.10	703.34	1030.56	105.53	0.40
Ağrı	39°43'N	43°02'E	1901-2020	318.93	514.68	743.41	82.12	0.32

MK and NMK methods are conducted to determine trends on precipitation series of the Murat River basin at Bingöl, Muş, and Ağrı stations. According to MK results, there is no monotonic trend at the 90% confidence level at all stations (Table 2). Ağrı has a positive  $z$  value (0.60), Bingöl and Muş have negative  $z$  values (-0.44 and -0.41). All  $z$  values are smaller than  $\pm 1.65$  and there is no monotonic trend even at the 90% confidence level.

NMK method gives important decreasing trends ( $z_{low} > 1.96$ ) on low sub-series of Bingöl and Muş stations and increasing trend ( $z_{high} > 1.65$ ) on high sub-series of Bingöl station. There is no non-monotonic trend on the high series of Muş station ( $z_{high} = 0.00$ ). Ağrı station has no monotonic or non-monotonic trend on main, low, and high series ( $z = 0.60$ ,  $z_{low} = -1.40$ , and  $z_{high} = -0.04$ ).

NMK method can also detect whether a time series is stationary (Table 2). Among the three stations, only Bingöl station has lost stationarity at the 95% confidence level ( $z_{NMK} = 3.07$ ). Positive  $z_{NMK}$  values represent an increase in the variance of the precipitation series.

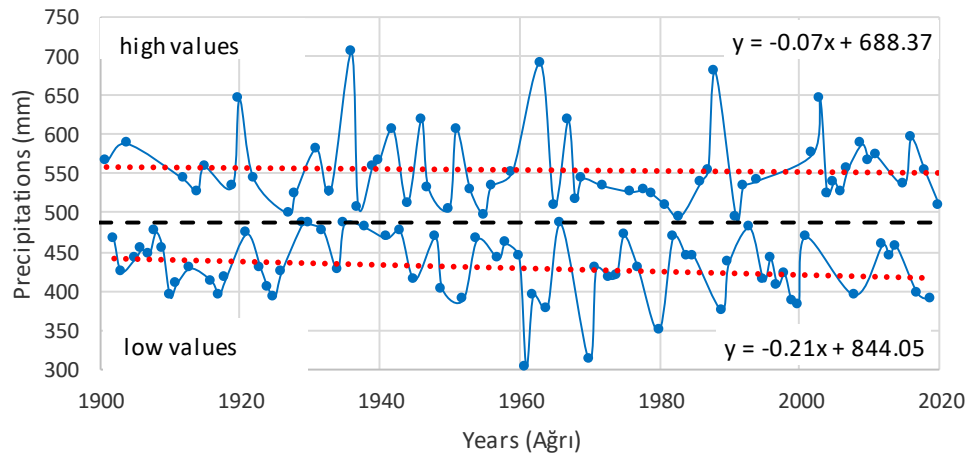
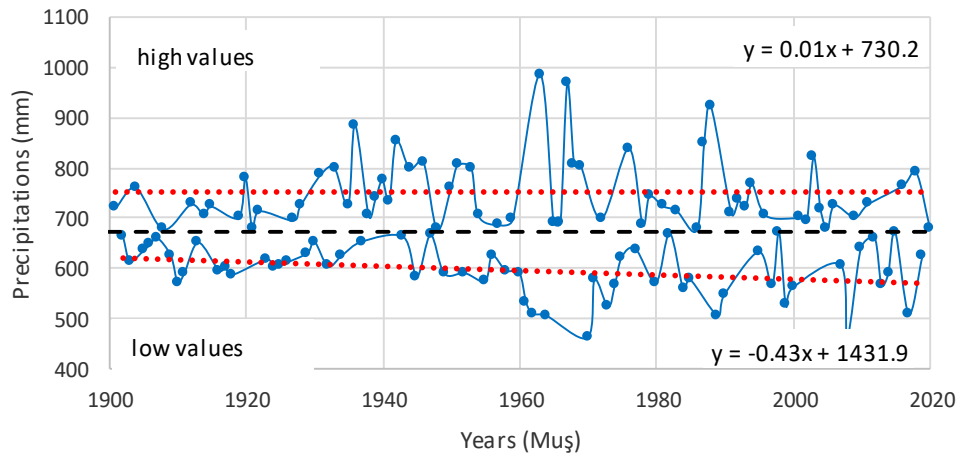
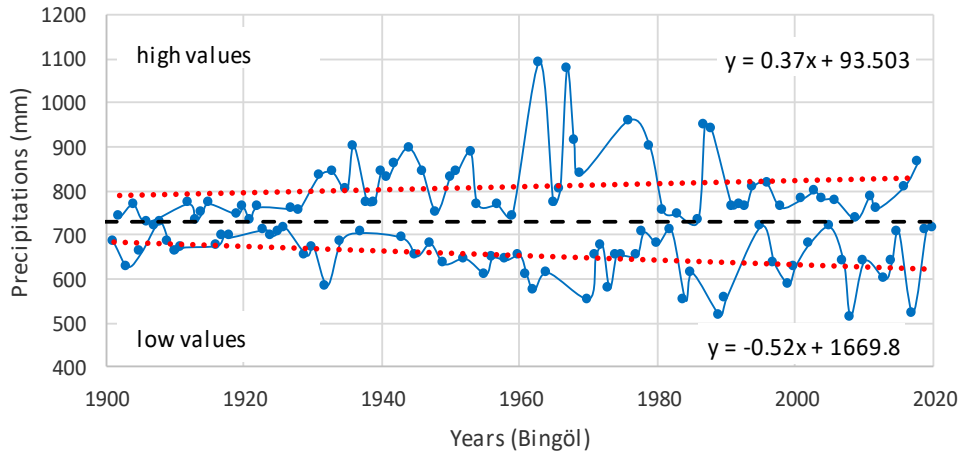
**Table 2** Mann-Kendall test results for Precipitation Series of Murat River Basin

Mann-Kendall for all series							
Stations	Data length ( $n$ )	Test statistics, $S$	$Var(S)$	Standard test statistics, $z$	Significance level	Tabulated test statistics, $z_{tab}$	Decision
Bingöl	120	-194	194367	-0.44	90%	$\pm 1.65$	No trend
Muş	120	-178	194367	-0.41	90%	$\pm 1.65$	No trend
Ağrı	120	264	194367	0.60	90%	$\pm 1.65$	No trend
Non-monotonic Mann-Kendall for low series							
Stations	Data length ( $n$ )	Test statistics, $S_{low}$ (Eq. 1)	$Var(S_{low})$ (Eq. 4)	Standard test statistics, $z_{low}$ (Eq. 5)	Significance level	Tabulated test statistics, $z_{tab}$	Decision
Bingöl	61	-390	25823	-2.43	95%	$\pm 1.96$	Important decreasing trend
Muş	59	-351	23384	-2.30	95%	$\pm 1.96$	Important decreasing trend
Ağrı	64	-240	29792	-1.40	90%	$\pm 1.65$	No trend
Non-monotonic Mann-Kendall for high series							
Stations	Data length ( $n$ )	Test statistics, $S_{high}$ (Eq. 1)	$Var(S_{high})$ (Eq. 4)	Standard test statistics, $z_{high}$ (Eq. 5)	Significance level	Tabulated test statistics, $z_{tab}$	Decision
Bingöl	59	291	23384	1.90	90%	$\pm 1.65$	Increasing trend
Muş	61	1	25823	0.00	90%	$\pm 1.65$	No trend
Ağrı	56	-4	20020	-0.04	90%	$\pm 1.65$	No trend
Non-monotonic Mann-Kendall for stationarity test							
Stations	Test statistics, $S_{low}$ (Eq. 1)	Test statistics, $S_{high}$ (Eq. 1)	$Var(S_{low})$ (Eq. 4)	$Var(S_{high})$ (Eq. 4)	Standard test statistics, $z_{NMK}$ (Eq. 6)	Tabulated test statistics, $z_{tab}$	Decision
Bingöl	-390	291	25823	23384	3.07	$\pm 1.96$	Non-Stationary
Muş	-351	1	23384	25823	1.58	$\pm 1.65$	Stationary
Ağrı	-240	-4	29792	20020	1.05	$\pm 1.65$	Stationary

Although the Muş station has a noteworthy  $z_{NMK}$  value of 1.58, it does not exceed the critical  $z_{tab}$  value of 1.65 at the 90% confidence level and therefore has stationarity. The minimum  $z_{NMK}$  value of 1.05 is calculated at the Ağrı station. To visualize non-monotonic trends on Murat River Basin (Euphrates) low and high precipitation series and calculate trend slopes, quantiles linear regression method (LR) is used as in Figure 3 to calculate trend slopes at different intervals (Solaimani, 2022).

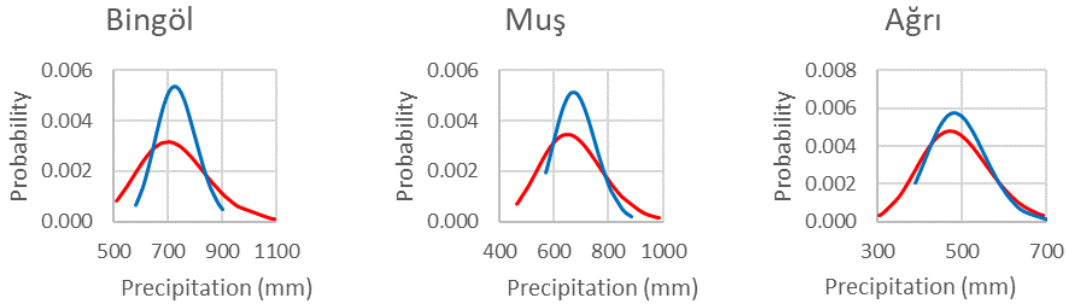
As in Figure 3, Bingöl station has an important increasing trend (0.37 mm/year) on high precipitation series and decreasing trend (-0.52 mm/year) on low series. These values fit NMK method but MK gives no trend on all series. Muş station has an insignificantly increasing trend (0.01 mm/year) on the high series and an important decreasing trend (-0.43 mm/year) on the low series, which is detectable by NMK method ( $z_{low} = -2.30$  and  $z_{high} = 0.00$ ). There is an insignificantly increasing trend (-0.07 mm/year) on the high series although a noteworthy decreasing trend (-0.21 mm/year) exists on the low series of Ağrı station. NMK method results for Ağrı station are again consistent with quantiles linear regression ( $z_{low} = -1.40$  and  $z_{high} = -0.04$ ).

Two parameter Gamma PDF is fitted to precipitation series at Bingöl, Muş, and Ağrı stations to compare stationarity between the first period (1901-1960, blue line) and the second period (1961-2020, red line) (Figure 4). Even though the probabilities of low and high values increase in Bingöl station, the probabilities of middle values decrease. The shape of the Bingöl precipitation distribution changes and accordingly the series loses its stationarity ( $z_{NMK} = 3.07$ ).



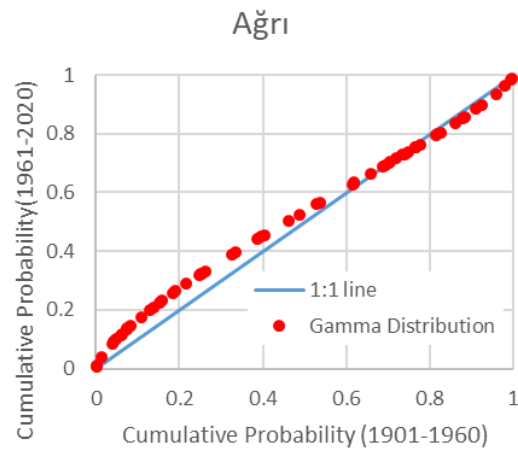
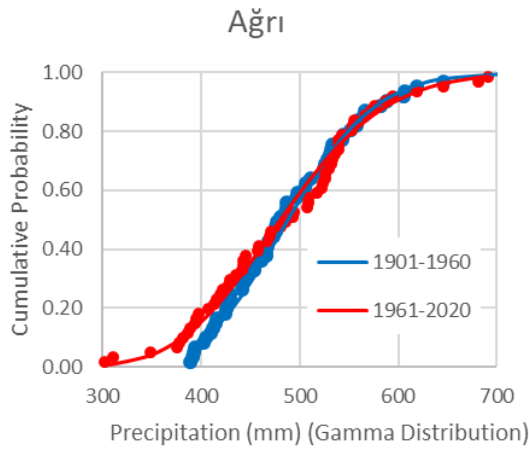
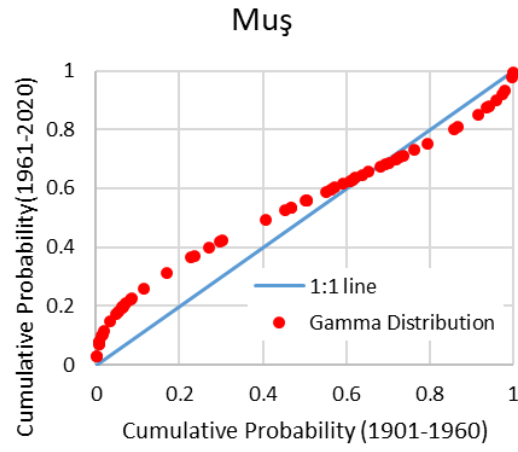
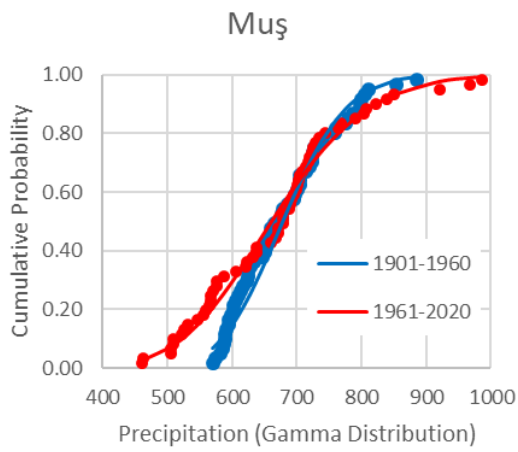
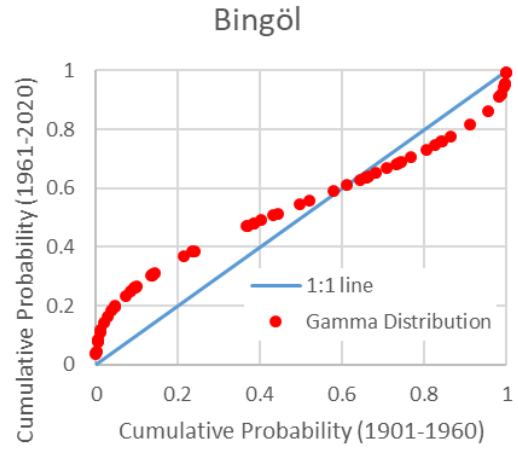
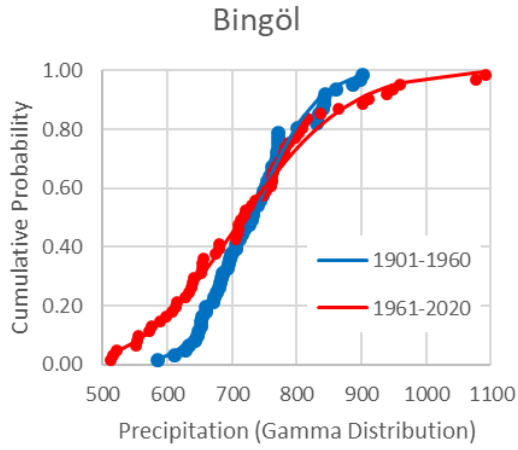
**Figure 3** The application of quantiles linear regression method for selected precipitation stations on the Murat River basin.

There is a minor change in the probability of Muş station precipitation high values although there is an important increase in low values probabilities ( $z_{low} = -2.30$ ). The stationarity of the precipitation series at Muş station is considerable with a  $z_{NMK}$  value of 1.58. Ağrı station maintains stationarity ( $z_{NMK} = 1.05$ ) but there is a remarkable increase in low precipitation probabilities ( $z_{low} = -1.40$ ).



**Figure 4** Gamma PDFs of precipitation series at selected stations (Blue lines represent the period 1901-1960, and red lines represent the period 1961-2020).

To double-check the stationarity of precipitation series at Bingöl, Muş, and Ağrı stations and compare those between the 1901-1960 and 1961-2020 periods, the Gamma cumulative distribution functions (CDF) are drawn in Figure 5. The CDF values for the periods 1901-1960 and 1961-2020 are shown on the left side of this figure and the comparison of these cumulative probability values with each other is on the right side. As in Figure 4, although the cumulative probabilities of low values are increasing, the cumulative probabilities of high values are decreasing between the 1901-1960 and 1961-2020 periods. The CDF probability values indicate that a certain probability value or a smaller value will occur. Therefore, while the probability of the resulting values being lower than the low values increases, the probability of the resulting values being lower than the high values decreases, or the risk of the resulting values being higher than the high values increases.



216

217 **Figure 5** Gamma CDFs of precipitation series and comparison of these values between 1921-

218 1960 and 1961-2020 periods at selected stations.

Accordingly, the recurrence times of droughts and floods are expected to decrease. The differences between the CDF probability values for the 1921-1960 and 1961-2020 periods of Bingöl station ( $z_{NMK} = 3.07$ ) are bigger than those of Muş station with  $z_{NMK}$  value of 1.58 (Figure 5). Ağrı station ( $z_{NMK} = 1.05$ ) has no important differences between cumulative probability values between 1921-1960 and 1961-2020 periods as in the figure. The probability of low and high values increases because all  $z_{NMK}$  values are positive.

## Results and Discussion

Classical MK is frequently used in the literature to detect monotonic trends on hydro-meteorological parameters, despite some restrictive assumptions such as independent time series and normality of test statistics. MK method cannot detect non-monotonic trends that have different effects on low and high values of a given time series. In this study, MK method is improved to detect non-monotonic trends and investigate whether a given time series is stationary. Non-monotonic Mann-Kendall (NMK) is applied in addition to MK and quantiles linear regression to detect trends on precipitation series at Bingöl, Muş, and Ağrı stations of the Murat River basin.

Classical MK cannot detect any trend on all the stations but NMK method gives two important decreasing trends on low values of the precipitation series at Bingöl and Muş stations and one increasing trend on high values of the precipitation series at Bingöl station. Ağrı station has a neither monotonic nor non-monotonic trend on the precipitation series. Quantiles linear regression yields consistent results with NMK method. There are trend slope values of -0.52 mm/year ( $z_{low} = -2.43$ ) and -0.43 mm/year ( $z_{low} = -2.30$ ) for low values of Bingöl and Muş precipitation series. There is also a 0.37 mm/year ( $z_{high} = 1.90$ ) trend slope value for high values of the Bingöl precipitation series. Ağrı station has small trend

slopes on the precipitation series such as -0.07 mm/year ( $z_{high} = -0.04$ ) on high values and -0.21 mm/year ( $z_{low} = -1.40$ ) on low values and again it is consistent with NMK method.

To investigate stationarity on the all precipitation series, NMK method gives non-stationarity for only the Bingöl precipitation series ( $z_{NMK} = 3.07$ ), which has an important trend on the low values and an increasing trend on the high values. Although Muş has a considerable  $z_{NMK}$  value of 1.58, Muş and Ağrı precipitation series are stationary. It is consistent with NMK values, because Muş has only important decreasing trend on the low values and the Ağrı precipitation series ( $z_{NMK} = 1.05$ ) has no trend. Positive  $z_{NMK}$  values represent increasing serial variance and uncertainty on the precipitation series of the Murat River.

Non-monotonic trends cannot be detected by classical methods, but NMK method is helpful to detect non-monotonic trends and enables better prediction of future events and support more efficient design procedure for hydraulic structures. The positive  $z_{NMK}$  values found for the Murat River basin indicate that lower values than the low (drought) values and higher values than the high (flood) values of the precipitation series are expected to emerge more frequently in the future.

## REFERENCE

- Alashan, S., 2020. Combination of modified Mann-Kendall method and Şen innovative trend analysis. Eng. Reports n/a, e12131. <https://doi.org/10.1002/eng2.12131>
- Alashan, S., 2018. Data Analysis in Nonstationary State. Water Resour. Manag. <https://doi.org/10.1007/s11269-018-1928-2>
- Alashan, S., Şen, Z., Toprak, Z.F., 2016. Hydroelectric Energy Potential of Turkey: A Refined Calculation Method. Arab. J. Sci. Eng. 41. <https://doi.org/10.1007/s13369-015-1982-5>
- Alifujiang, Y., Abuduwaili, J., Ge, Y., 2021. Trend analysis of annual and seasonal river runoff by using innovative trend analysis with significant test. Water (Switzerland). <https://doi.org/10.3390/w13010095>
- Bayazit, M., 2015. Nonstationarity of Hydrological Records and Recent Trends in Trend Analysis: A State-of-the-art Review. Environ. Process. <https://doi.org/10.1007/s40710->

272 015-0081-7

273 Bayazit, M., Önöz, B., 2007. To prewhiten or not to prewhiten in trend analysis? Hydrol. Sci.  
274 J. <https://doi.org/10.1623/hysj.52.4.611>

275 Berhail, S., Tourki, M., Merrouche, I., Bendekiche, H., 2022. Geo-statistical assessment of  
276 meteorological drought in the context of climate change: case of the Macta basin  
277 (Northwest of Algeria). Model. Earth Syst. Environ. [https://doi.org/10.1007/s40808-020-](https://doi.org/10.1007/s40808-020-01055-7)  
278 01055-7

279 Burić, D., Doderović, M., 2021. Changes in temperature and precipitation in the instrumental  
280 period (1951–2018) and projections up to 2100 in Podgorica (Montenegro). Int. J.  
281 Climatol. 41, E133–E149. <https://doi.org/10.1002/joc.6671>

282 Cox, D.R., Stuart, A., 1955. Some Quick Sign Tests for Trend in Location and Dispersion.  
283 Biometrika. <https://doi.org/10.2307/2333424>

284 Dabanli, I., Şişman, E., Güçlü, Y.S., Birpınar, M.E., Şen, Z., 2021. Climate change impacts  
285 on sea surface temperature (SST) trend around Turkey seashores. Acta Geophys.  
286 <https://doi.org/10.1007/s11600-021-00544-2>

287 Faulkner, D., Warren, S., Spencer, P., Sharkey, P., 2020. Can we still predict the future from  
288 the past? Implementing non-stationary flood frequency analysis in the UK. J. Flood Risk  
289 Manag. 13. <https://doi.org/10.1111/jfr3.12582>

290 Günek, H., 2011. Murat Nehri Havzasinin (Firat) Su Potansiyeli Ve Değerlendirilmesi. Doğu  
291 Coğrafya Derg.

292 Haan, C.T., 1977. Statistical methods in hydrology. Stat. methods Hydrol.  
293 <https://doi.org/10.1201/9780429423116-36>

294 Hamed, K.H., 2009. Exact distribution of the Mann-Kendall trend test statistic for persistent  
295 data. J. Hydrol. <https://doi.org/10.1016/j.jhydrol.2008.11.024>

296 Hamed, K.H., Ramachandra Rao, A., 1998. A modified Mann-Kendall trend test for  
297 autocorrelated data. J. Hydrol. 204, 182–196. [https://doi.org/10.1016/S0022-](https://doi.org/10.1016/S0022-1694(97)00125-X)  
298 1694(97)00125-X

299 Harris, I., Osborn, T.J., Jones, P., Lister, D., 2020. Version 4 of the CRU TS monthly high-  
300 resolution gridded multivariate climate dataset. Sci. Data 7.  
301 <https://doi.org/10.1038/s41597-020-0453-3>

302 Kendall, M.G., 1975. Rank Correlation Methods, Charles Griffin, London (1975). Google  
303 Sch.

304 Kendall, M.G., Gibbons, J.D., 1990. Rank Correlation Methods, fifth. ed. Griffin, London.

305 Khaliq, M.N., Ouarda, T.B.M.J., Ondo, J.C., Gachon, P., Bobée, B., 2006. Frequency analysis  
306 of a sequence of dependent and/or non-stationary hydro-meteorological observations: A  
307 review. J. Hydrol. <https://doi.org/10.1016/j.jhydrol.2006.03.004>

308 Lins, H.F., Cohn, T.A., 2011. Stationarity: Wanted dead or alive? J. Am. Water Resour.  
309 Assoc. <https://doi.org/10.1111/j.1752-1688.2011.00542.x>

310 Mann, H.B., 1945. Nonparametric Tests Against Trend. Econometrica 13, 245.  
311 <https://doi.org/10.2307/1907187>

312 Matalas, N.C., 2012. Comment on the Announced Death of Stationarity. J. Water Resour.  
313 Plan. Manag. [https://doi.org/10.1061/\(asce\)wr.1943-5452.0000215](https://doi.org/10.1061/(asce)wr.1943-5452.0000215)

- Milly, P.C.D., Bentacourt, J., Falkenmark, M., Robert, M., Hirsch, R.M., Kundzewicz, Z.W., Letternmaier, D.P., Stouffer, R.J., 2008. Stationarity is dead: Whither water management? *Science* (80-. ).
- Montanari, A., Koutsoyiannis, D., 2014. Modeling and mitigating natural hazards: Stationarity is immortal! *Water Resour. Res.* <https://doi.org/10.1002/2014WR016092>
- Nigussie, T.A., Altunkaynak, A., 2019. Impacts of climate change on the trends of extreme rainfall indices and values of maximum precipitation at Olimpiyat Station, Istanbul, Turkey. *Theor. Appl. Climatol.* 135, 1501–1515. <https://doi.org/10.1007/s00704-018-2449-x>
- Oruc, S., 2021. Trend and nonstationary relation of extreme rainfall: Central Anatolia, Turkey. *Acta Geophys.* <https://doi.org/10.1007/s11600-020-00518-w>
- Şan, M., Akçay, F., Linh, N.T.T., Kankal, M., Pham, Q.B., 2021. Innovative and polygonal trend analyses applications for rainfall data in Vietnam. *Theor. Appl. Climatol.* 144, 809–822. <https://doi.org/10.1007/s00704-021-03574-4>
- Saplıoğlu, K., Güçlü, Y.S., 2022. Combination of Wilcoxon test and scatter diagram for trend analysis of hydrological data. *J. Hydrol.* 612, 128132. <https://doi.org/https://doi.org/10.1016/j.jhydrol.2022.128132>
- Sen, P.K., 1968. Estimates of the Regression Coefficient Based on Kendall's Tau. *J. Am. Stat. Assoc.* 63, 1379–1389. <https://doi.org/10.1080/01621459.1968.10480934>
- Şen, Z., 2012. Innovative Trend Analysis Methodology. *J. Hydrol. Eng.* 17, 1042–1046. [https://doi.org/10.1061/\(ASCE\)HE.1943-5584.0000556](https://doi.org/10.1061/(ASCE)HE.1943-5584.0000556)
- Serinaldi, F., Kilsby, C.G., 2015. Stationarity is undead: Uncertainty dominates the distribution of extremes. *Adv. Water Resour.* <https://doi.org/10.1016/j.advwatres.2014.12.013>
- Serinaldi, F., Kilsby, C.G., Lombardo, F., 2018. Untenable nonstationarity: An assessment of the fitness for purpose of trend tests in hydrology. *Adv. Water Resour.* <https://doi.org/10.1016/j.advwatres.2017.10.015>
- Şişman, E., Kizilöz, B., 2021. The application of piecewise ITA method in Oxford, 1870–2019. *Theor. Appl. Climatol.* 145, 1451–1465. <https://doi.org/10.1007/s00704-021-03703-z>
- Solaimani, K., 2022. Seasonal relationship between climatic variables and evaporation based on Bayesian quantile regression method in southern Caspian region. *Arab. J. Geosci.* 15, 1000. <https://doi.org/10.1007/s12517-022-10263-7>
- Spearman, C., 1987. The proof and measurement of association between two things. By C. Spearman, 1904. *Am. J. Psychol.* 100, 441–471. <https://doi.org/10.2307/1422689>
- Vogel, R.M., Yaindl, C., Walter, M., 2011. Nonstationarity: Flood magnification and recurrence reduction factors in the united states. *J. Am. Water Resour. Assoc.* <https://doi.org/10.1111/j.1752-1688.2011.00541.x>