Airborne observations of surface winds, waves and currents from meso to submesoscales

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Abstract

In this work we present a unique set of coincident and collocated high- resolution observations of surface currents and directional properties of surface waves collected from an airborne instrument, the Modular Aerial Sensing System (MASS), collected off the coast of Southern California. High-resolution observations of near surface current profiles and shear are obtained using a new instrument, DoppVis, capable of capturing horizontal spatial current variability down to 128m resolution. This data set provides a unique opportunity to examine how currents at scales ranging from 1-100 km modulate bulk (e.g. significant wave height), directional and spectral properties of surface gravity waves. Such observations are a step toward developing better understanding of the underlying physics of submesoscale processes (e.g. frontogenesis and frontal arrest) and the nature of transitions between mesoscale and submesoscale dynamics.

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Key Points:

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13	٠	Unique coincident and collocated airborne observations of SST, surface currents
14		and properties of surface waves across submesoscales features
15	•	A new airborne instrument enables observations of surface currents, vertical and
16		horizontal shear to capture quickly evolving ocean features
17	•	Such observations are crucial to develop better understanding of the physics of sub-
18		mesoscale processes and wave-current interaction

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19 Abstract

In this work we present a unique set of coincident and collocated high- resolution obser-20 vations of surface currents and directional properties of surface waves collected from an 21 airborne instrument, the Modular Aerial Sensing System (MASS), collected off the coast 22 of Southern California. High-resolution observations of near surface current profiles and 23 shear are obtained using a new instrument, DoppVis, capable of capturing horizontal spa-24 tial current variability down to 128m resolution. This data set provides a unique oppor-25 tunity to examine how currents at scales ranging from 1-100 km modulate bulk (e.g. sig-26 nificant wave height), directional and spectral properties of surface gravity waves. Such 27 observations are a step toward developing better understanding of the underlying physics 28 of submesoscale processes (e.g. frontogenesis and frontal arrest) and the nature of tran-29 sitions between mesoscale and submesoscale dynamics. 30

³¹ Plain Language Summary

In recent years, through improvement of computational resolution of global ocean 32 models, scientists have begun to suspect that kilometer-scale eddies, whirlpools and fronts, 33 called "submesoscale" variability, make important contributions to horizontal and ver-34 tical exchange of climate and biological variables in the upper ocean. Such features are 35 challenging to analyze, because of their size (and how quickly they evolve; within hours), 36 they are too large to study from a research vessel but smaller than regions typically stud-37 ied with satellite measurements. In this work, we use a research aircraft instrumented 38 to characterize ocean currents, temperature, color (in turn chlorophyll concentration) 30 and the properties of surface waves over an area large enough to capture submesoscale 40 processes. This approach is a step forward in understanding and quantifying the under-41 lying physics of submesoscale processes, and in turn develops parameterization that can 42 help improve the fidelity of weather and climate models. 43

44 1 Introduction

The transfer of mass, momentum, and energy between the atmosphere and ocean 45 are complex due to their interactions across a broad range of space and time scales (Melville, 46 1996). A better understanding of the physics of these processes is fundamental for im-47 proved parameterizations used in coupled air-sea models of weather and climate, par-48 ticularly as Earth's climate changes (Cavaleri et al., 2012). For example, although the 49 importance of surface waves in these models has long been acknowledged, only relatively 50 recently have global models included physics-based models of their effects (see, for ex-51 ample McWilliams & Restrepo, 1999; Sullivan & McWilliams, 2010; Li et al., 2016). Specif-52 ically, the effects of the non-breaking surface wave induced transport that catalyses Lang-53 muir circulations have been shown to reduce errors in sea surface temperature (Belcher 54 et al., 2012), crucial to climate modelling. 55

Submession currents have horizontal scales on the order of 0.1-10 km and 56 have recently been hypothesized to make important contributions to vertical exchanges 57 of climate and biological variables in the upper ocean as well as provide a pathway from 58 energetically rich large scale flows to small scale dissipation. Model studies and limited 59 observations (e.g. D'Asaro et al., 2018) show that submesoscale vertical exchange is con-60 centrated near kilometer-scale fronts, jets, and eddies (McWilliams, 2016). Submesoscale 61 physics are at the smallest scales that have been resolved in global ocean models, where 62 their net effect on heat exchange between the ocean and atmosphere has shown to be 63 much larger than mesoscale eddies (Su et al., 2018). However, these simulations are sen-64 sitive to the parameterized physics of the smaller scale motion, which remains poorly un-65 derstood. To address the fundamental questions of the observed nature of submesoscale 66 dynamics and the interactions between submesoscale dynamics and smaller scale surface 67 wave processes, a comprehensive set of novel, coincident and collocated measurements 68

69 of the dynamical variables is needed to improve state-of-the-art high-resolution simula-70 tions of weather and climate, and better understand vertical exchanges of heat and bio-

71 geochemical tracers.

Due to non-linear coupling between oceanic and atmospheric processes, including 72 currents, winds, and waves, coincident observations are necessary to understand these 73 dynamics. A number of model and observational studies have demonstrated the variety 74 of ways by which air-sea interaction can induce horizontal divergence of surface currents 75 and thus force vertical velocities. Coincident observations of surface vector winds and 76 77 currents are needed to better understand the coupling of winds and currents and in turn improve surface flux parameterizations (Bourassa et al., 2019). The wind stress can be 78 modified by SST and velocity gradients (Dewar & Flierl, 1987; Fairall et al., 1996; Chel-79 ton et al., 2004; O'Neill et al., 2005) while the resulting convergence of the ocean Ek-80 man layer can be modified by surface vorticity (Stern, 1965; McGillicuddy Jr et al., 2007). 81 At submesoscales, these effects are expected to increase in intensity. Frontal structures 82 are strongly affected by the relative direction of the wind with *downfront* Ekman trans-83 port sharpening the fronts and inducing vertical exchange while *upfront* transport subdues the front and stratifies the upper ocean (Thomas et al., 2005). Surface waves and 85 wave breaking can be strongly modulated at fronts (Romero et al., 2017; Vrećica et al., 86 2022), suggesting that even the basic formulation of air-sea exchange in terms of sim-87 ple bulk coefficients will likely break down at sufficiently small scales. As such, an un-88 derstanding of submesoscale structure and vertical velocity requires that air-sea inter-89 action parameters be observed simultaneously with submesoscale measurements. Of par-90 ticular interest are observations of surface and near surface currents, wave breaking (Vrećica 91 et al., 2022), and the directional properties of ocean surface waves. While much progress 92 have been made on characterizing the latter (e.g. Herbers et al., 2012; Lenain & Melville, 93 2014; Melville et al., 2016; Lenain & Melville, 2017; Lenain & Pizzo, 2020), collecting 94 observations of near surface currents, i.e. from the surface down to several meters depth, 95 remains challenging and spatially limited in large part due to the presence of waves, which 96 induce platform motions and additional sources of background noise. An alternative ap-97 proach to traditional in-situ techniques is to infer current profiles remotely based on ob-98 servations of the spatio-temporal evolution of surface waves that follow a dispersion re-99 lationship, pursued in the present study. 100

In this work, we present a unique set of coincident and collocated observations of 101 high-resolution surface currents and directional properties of surface waves collected from 102 an airborne instrument, the Modular Aerial Sensing System (MASS), off the coast of South-103 ern California in May 2021 as part of the ONR-funded Task Force Ocean (TFO) research 104 initiative and in October 2021 as part of the NASA-funded Submesoscale Ocean Dynam-105 ics Experiment (S-MODE) research initiative. Two of the TFO research flights were ded-106 icated to collecting observations across two small counter-rotating eddies separated by 107 approximately 100 km. One of the S-MODE flights collected observations across a sharp 108 SST front. This data set provides a unique opportunity to examine how currents at scales 109 ranging from 1-100 km modulate surface gravity waves, i.e. bulk, directional and spec-110 tral properties. 111

This paper is structured as follows: the overview of the experiment and processing techniques is given in section 2 and 3. In sections 4 through 6, analysis of meso to submesoscales surface kinematics and currents collected during two experiments in Southern California is discussed. Section 7 discusses potential implications for submesoscale and air-sea interaction studies and presents some summary points.

117 **2** Experiments

In this study we consider observations collected during two distinct experiments.
 The first was conducted as part of the "Platform Centric ASW Processing with Through-

the-Sensor Data Assimilation and Fusion" project, funded through the ONR Task Force
Ocean (TFO) initiative with the aim of collecting a simultaneous combination of acoustic, air-sea interaction and oceanographic measurements. Observations from a research
vessel, a drifting instrument array, autonomous surface vehicles and a research aircraft
were collected in May 2021, approximately 45km offshore of San Diego, CA, in the vicinity of CalCoFi Line #90.

The second experiment was conducted as part of the NASA S-MODE program, a project that aims to characterize the contribution of submesoscale ocean dynamics to vertical and horizontal transport in the upper ocean by employing a combination of aircraftbased remote sensing measurements of the ocean surface, in-situ measurements from research vessels and a variety of autonomous oceanographic platforms, and numerical modeling (Farrar et al., 2020). The "pilot" experiment considered here was conducted in the fall of 2021 off the coast of San Francisco, CA.

Data collected from instrumented Wave Gliders (Grare et al., 2021) and an airborne instrument, the SIO Modular Aerial Sensing System (MASS, Melville et al., 2016) during these two field programs are considered in the analysis. During both experiments the MASS instrument was installed on a Twin Otter DHC-6 aircraft (Twin Otter International, Grand Junction, CO).

3 SIO-MASS DoppVis instrument: Enabling novel airborne observations of near-surface currents

The Modular Aerial Sensing System (MASS) is an airborne instrument developed 140 at the Air-Sea Interaction Laboratory (SIO) to simultaneously collect observations of sea 141 surface temperature and ocean color (Melville et al., 2016; Lenain & Pizzo, 2021a), winds 142 and mean-square slope (Lenain et al., 2019), surface waves (Lenain & Melville, 2017; Lenain 143 & Pizzo, 2020), and ocean topography (Villas Bôas et al., 2022), at horizontal scales rang-144 ing from sub-meter to mesoscales. Over the past 11 years, the instrument was flown for 145 more than 30 missions, covering a broad range of environmental conditions, locations, 146 and applications. Details on the system performance and various applications of the MASS 147 can be found in Melville et al. (2016); Lenain and Melville (2017); Lenain et al. (2019); 148 Lenain and Pizzo (2020); Vrećica et al. (2022). 149

In 2020, we started the development and integration of a new sensor into the MASS 150 instrument, called "DoppVis", to obtain coincident observations of surface currents along-151 side the MASS observations listed above. The approach used is to infer currents from 152 optical observations of the spatio-temporal evolution of surface waves, whose dispersion 153 is altered by the presence of an underlying current. This technique has been primarily 154 used with radar technology (e.g. Stewart & Joy, 1974; Campana et al., 2016; Lund et 155 al., 2015) then later applied to airborne video imagery (Dugan et al., 2001; Dugan & Pi-156 otrowski, 2003; Anderson et al., 2013). 157

Starting from the dispersion relation for small-amplitude linear waves propagating on top of a depth-varying current,

$$\omega(\mathbf{k}) = \omega_0(k) + \mathbf{c}(k) \cdot \mathbf{k},\tag{1}$$

where ω is the wave frequency, ω_0 is the frequency in the absence of currents, i.e equal to \sqrt{gk} in deep water, $\mathbf{k} = (k_x, k_y)$ is the wavenumber, $k = |\mathbf{k}|$, and \mathbf{c} is the Doppler shift velocity due to the underlying current. Following Stewart and Joy (1974), assuming the waves are in deep-water, \mathbf{c} can be approximated as a weighted average of the current profile as a function of depth such that

$$\mathbf{c}(k) = 2k \int_{-\infty}^{0} \mathbf{U}(z) e^{2kz} \mathrm{d}z, \qquad (2)$$

where $\mathbf{U}(z) = (U, V)$ is the Lagrangian mean current profile as a function of depth z(Pizzo et al., 2022). Based on this relationship, one can assign an effective depth z_e to the measured Doppler velocities $\mathbf{c}(k)$ by finding the depth at which the Doppler velocity is equal to the current (Stewart & Joy, 1974; Smeltzer et al., 2019), such that $z_e(k) = -1/2k$. This is referred to as the Effective Depth Method (EDM) in Smeltzer et al. (2019).

The DoppVis instrument collects visible imagery of the ocean surface using a Nikon 163 D850 camera with 14mm lens mounted with a 90 degree rotation (long edge of image 164 parallel with flight track) and a 30 degree positive pitch angle from nadir (pointing slightly 165 ahead of aircraft). The camera is synchronized to a coupled GPS/IMU system collect-166 ing images at a 2Hz frame rate. Raw images are carefully calibrated for lens distortion 167 and boresight misalignment with the GPS/IMU over a hard terrestrial target, then geo-168 referenced and exported with reference to WGS84 datum with a UTM zone 10 projec-169 tion (EPSG 32610) at 50cm horizontal resolution. Each image is then interpolated on 170 a regular grid, to enable the generation of 3D cubes of imagery (time, UTM X, UTM 171 Y) of set duration and dimension (N_x, N_y) , where $N_x = N_y$, typically in the range of 172 128 to 512m. The number of collected data cubes in the cross and along track direction 173 of the aircraft varies as a function of aircraft altitude. All data presented here were col-174 lected at 1500 AMSL, corresponding to approximately two $256 \times 256 \text{ m}^2$ cubes in the cross-175 track direction. Following the same approach described in (Smeltzer et al., 2019), all cubes 176 of space-time data are converted to wavenumber-frequency space using a 3D FFT. Each 177 of these 3D spectra are then averaged in the cross-track and along-track direction (1 km 178 bin) to improve SNR. Doppler shift velocities are extracted from the spectrum as a func-179 tion of wavenumber by masking the spectrum into wavenumber magnitude bins (bin half-180 width of $4\pi/N_x$, where for each bin the current $\mathbf{c}(k)$ is estimated using a normalized 181 scalar product method (Huang et al., 2016; Streßer et al., 2017) with a Gaussian char-182 acteristic function (Smeltzer et al., 2019) peaked along the linear dispersion relation. 183

Figure 1(a) shows an example of current profiles (U, V) collected from DoppVis dur-184 ing an overflight of an instrumented Wave Glider during the S-MODE experiment on Novem-185 ber 4, 2021 at 17:20 UTC. The Wave Glider was instrumented with an upward-looking 186 Nortek Signature 1000 ADCP (orange squares) and a downward-looking ADCP (Tele-187 dyne RDI Workhorse 300kHz), carefully motion compensated using an onboard GPS/IMU 188 system (Grare et al., 2021). Observations from DoppVis and the wave glider were col-189 lected within 5min and no further than 500m from each other to minimize any error as-190 sociated with natural spatial and temporal variability. We find good agreement between 191 in-situ and remotely sensed observations of near-surface current (U, V), with a bias = 192 -0.014 m/s and rms deviation = 0.052 m/s, and a coefficient of determination $R^2 = 0.96$. 193

Finally, airborne observations of SST collected from MASS on November 4, 2021, 194 over the entire domain along with current estimates (1km along-track resolution) from 195 DoppVis at two depths, $z = -1.5 \pm 0.5$ m (black arrows) and $z = -0.4 \pm 0.1$ m (red ar-196 rows) are shown in figure 1(c). Note the correlation between features present in the SST 197 fields and the surface currents from DoppVis. Throughout the domain, we consistently 198 find larger magnitudes of the eastern component of the current closer to the surface, likely 199 caused by Stokes drift included in the Lagrangian current observed by DoppVis (Pizzo 200 et al., 2022). Wind and waves were coming from the west at the time of the flight. 201

4 Wave-current interactions from meso- to submesoscales

The collocated observations from the multiple instruments on the MASS allow for investigation of interactions between currents and other oceanographic properties such as waves, heat, and biological communities. We now examine observations collected during the "Platform Centric ASW Processing with Through-the-Sensor Data Assimilation and Fusion" TFO project in May 2021 to illustrate the importance of high spatial resolution, collocated observations for studying wave-current observations.



Figure 1. (a) Current profiles (U, V) collected from DoppVis (gray triangles) during an overflight of an instrumented Wave Glider during the S-MODE experiment on November 4, 2021 at 17:20 UTC. The Wave Glider is equipped with an upward-looking Nortek Signature 1000 ADCP (orange squares) and a downward-looking ADCP (Teledyne RDI Workhorse 300kHz; blue circles). The location of the Wave Glider with respect to the DoppVis observations is shown in (c). (b) Comparison between in-situ observations of currents collected at a 2 to 3m water depth obtained from DoppVis and two Wave Gliders, named WHOI43 and STOKES, for the times and locations of overflights (within 500m) during the entire S-MODE experiment. (c) Airborne observations of SST collected from MASS on November 4, 2021 over the entire domain along with current estimates (1km along-track resolution) from DoppVis at two depths, deep, $z = -1.5 \pm 0.5m$ (black arrows) and shallow, $z = -0.4 \pm 0.1m$ (red arrows).

Several research flights were dedicated to collecting observations across two counter-209 rotating eddies separated by approximately 100 km. Figure 2(a) shows depth-averaged 210 (0.3-2m) surface currents (250 m along-track resolution) along with significant wave height 211 measurements collected coincidentally with the MASS instrument. The black contours 212 represent the sea surface height (ssh, AVISO) at the time of the observations, while cor-213 responding surface currents from HYCOM (GOF 3.1, GLBy0.08-expt93.0) are shown as 214 gray quivers. We find the HYCOM estimated currents to be in good agreement with DoppVis 215 observations at the mesoscale, but strikingly miss many of the submesoscale features through-216 out the domain. We also find significant modulation of the surface wave properties across 217 the domain that is not caused by temporal (e.g. inertial) variability. 218

Currents modulate the properties of surface gravity waves both through wave refraction (e.g. Ardhuin et al., 2017; Romero et al., 2017, 2020; Bôas et al., 2020; Pizzo & Salmon, 2021) and local effects (Rascle et al., 2016, 2017; Lenain & Pizzo, 2021b), therefore high-resolution measurements of currents and spectral wave properties is needed to better understand wave-current interaction from meso- to submesoscales.

Figure 2 shows a subset of these observations, focusing on an area of the domain with significant submesoscale variability in ocean current magnitude and direction. We find large modulation of surface currents, at scales of 1-8 km associated with submesoscale variations within the mesoscale eddy. Panel (b) shows a sharp SST front at 32.885°N with a change in temperature of approximately $0.5^{\circ}C$, associated with a velocity gradient $\partial U/\partial x \approx 10f$, where f is the Coriolis frequency.

Despite approximately constant surface wind (figure 2c) and significant wave height 230 (figure 2d) as a function of latitude on this transect, the mean spectral saturation $\langle B \rangle$ 231 (Romero et al., 2017) had significant variation (figure 2e). The saturation spectrum is 232 defined as $B = \phi(k)k^3$, where ϕ is the omnidirectional wave spectrum (Phillips, 1984, 233 1985). The mean saturation is obtained by averaging B over the saturation range of the 234 wave spectrum, with a lower bound k_n as defined in Lenain and Pizzo (2020). At that 235 time, wind waves were approximately oriented in the same direction as the current V236 (figure 2d). 237

We find that the mean saturation $\langle B \rangle$ generally follow the same evolution as V. While this comparison remains qualitative, it hints at the importance of incorporating wave-current interaction in wave spectrum parameterization, along the line of Lenain and Pizzo (2021a) that investigated the use of WKB to predict the modulation of surface gravity waves by internal wave currents.

Though beyond the scope of this work, but a focus of further studies in particular as part of the S-MODE program, we expect that these collocated, coincident observations will lead to new insights about coherent features that arise from wind-wave-current coupling, or some combination of those processes, which may be particularly important for the flux of climatically-important variables such as heat and biogeochemical tracers between atmosphere and ocean (Li et al., 2016; Smith et al., 2016; Verma & Sarkar, 2021; Freilich & Mahadevan, 2021; Gula et al., 2022).

5 Transitions from mesoscale to submesoscale motion

High-resolution measurements of surface currents enable the characterization of the 251 transition between geostrophically balanced motion at larger scales and unbalanced (ageostrophic 252 and wave) motion at smaller scales (Chereskin et al., 2019). In particular, determining 253 the dynamics that predominate at the submesoscale has important implications for the 254 vertical structure of the upper-ocean (Cronin et al., 2019) and the spatio-temporal dis-255 tribution of energy dissipation (Buckingham et al., 2019; Dong et al., 2020; Schubert et 256 al., 2020; Ajayi et al., 2021), and can lead to improvement of ocean-atmosphere coupled 257 models (Li et al., 2016). While DoppVis observations are accurate enough to resolve sub-258



(a) MASS significant wave height H_s and DoppVis surface current observations Figure 2. (250m along-track resolution) collected during the TFO experiment on May 19, 2021. The black contours represent the sea surface height (ssh, AVISO) at the time of the observations, while corresponding surface currents from HYCOM are shown as gray quivers. The larger scale trends of the surface currents are in general agreement with the HYCOM product, though the latter completely misses significant submesoscale features. We find significant modulation of the wave conditions across the domain, by up to 20%. (b) Subset of the data presented in (a) (nortwestern part of the flight) with the track colorcoded for measured SST. Note the strong modulation of surface currents over very short distances (1-8 km), coinciding in places with sharp SST fronts (for example at latitude around 32.885° in the northern part of the operation area). (c) and (d) show the wind speed U_{10} and significant wave height H_s collected along the NNW-SSE track shown in panel (b). (e) North component of the current V and mean spectral saturation $\langle B \rangle$, all plotted over the same range of latitude as in panel (b). Note the modulation of $\langle B \rangle$ associated with changes in V, the component of the current aligned with the waves, hinting at local wavecurrent interaction.



Figure 3. Comparison of Kinetic energy (KE) spectra collected from the DoppVis Lagrangian surface currents on May 19 2021 during one of the long East - West transects shown in figure 2(a) compared with in-situ ADCP observations (20-m depth) and LLC4320 hourly model products from Chereskin et al. (2019).

mesoscale currents (horizontal resolution of less than 500 m), the airborne platform also 259 travels quickly enough to collect nearly synoptic observations at the mesoscale (100 km). 260 This is particularly valuable because it allows for examination of spatial variability across 261 a range of scales and cross-scale interactions by not aliasing temporal variability. Fig-262 ure 3 shows kinetic energy (KE) spectra computed from the DoppVis surface current ob-263 servations collected during one of the long East-West transects shown in figure 2(a) com-264 pared with the average of 11 years of in-situ ADCP observations (20-m depth) and 1 year 265 of LLC4320 hourly model products from Chereskin et al. (2019). The agreement at low 266 wavenumbers (>10 km) is remarkable, as well as the fact that the kinetic energy spec-267 trum from DoppVis maintains a nearly continuous slope for higher wavenumbers, down 268 to 1 km scale. Feedbacks on ocean currents from wind and waves is likely important for 269 determining dynamics at scales of less than 10 km where submesoscale processes pre-270 dominate (Haney et al., 2015; Suzuki & Fox-Kemper, 2016; Yuan & Liang, 2021). Un-271 derstanding these couplings needs to be advanced by direct observations that span scales 272 from submesoscale to mesoscale. Characterization of the transition between geostroph-273 ically balanced motion and unbalanced dynamics is further explored in forthcoming work. 274

²⁷⁵ 6 Velocity gradient statistics

Submesoscale dynamics are defined by the large velocity gradients at these scales, with O(1) Rossby number (McWilliams, 2016). Consistent with the small spatial scale of submesoscale dynamics, they also have fast temporal scales (Callies et al., 2020). Observations of not just velocity, but 2D velocity gradients of quickly evolving submesoscale features are therefore important for understanding submesoscale dynamics, but require



Figure 4. (a) Surface vorticity–strain and (b) divergence-strain JPDFs computed from 128m resolution DoppVis products collected on October 5 2021 during the S-MODE pilot program. (c) and (d) Surface vorticity–strain for $\Delta/f > 0.8$ and $\Delta/f < -0.8$ respectively. The dashed lines represents the $\sigma = \zeta$ lines.

synoptic observations over scales of tens of kilometers, which is not feasible with relatively slow-moving ship-based observations. Previous observations of submesoscale velocity gradients have used two ships (Shcherbina et al., 2013). Aircraft-based observations offer a well-suited platform for studying submesoscale dynamics due to their fast
speed relative to ships and much higher resolution measurements than what is afforded by satellites.

In recent years, through numerical (Balwada et al., 2021) and observational (Shcherbina 287 et al., 2013) studies, the surface vorticity-strain Joint Probability Density Function (JPDF) 288 has been demonstrated as a useful tool to distinguish between different flow regimes, in 289 particular in the context of submesoscale dynamics. In figure 4, we show vorticity-strain 290 and divergence-strain JPDFs computed from high-resolution DoppVis surface current 291 observations collected during the S-MODE pilot experiment on October 2021 (60km long 292 section). During that portion of the flight, the aircraft flew at 3000ft AMSL, increasing 293 the swath width of surface current observation up to 2 km, with a horizontal resolution 294 (cross and along-track) of 128 m. Vorticity ζ , strain rate σ and divergence Δ are then 295 computed from the gradients of the measured surface currents. These velocities are first 296 low-pass filtered (3-point tophat) prior to computing gradients. Overall, the distribu-297 tion is skewed toward positive ζ/f values, with $\sigma/f > \zeta/f$ consistent with frontal struc-298 tures, and classical submesoscale frontogenesis. This is particularly evident in panel (d), 299 where the vorticity-strain JPDF for $\Delta/f < -0.8$ (i.e. convergence) is shown to be skewed 300 to positive values. The same JPDF for positive values of Δ/f , larger than 0.8 (panel c) 301 is not skewed toward negative ζ/f values, perhaps implying the contribution of other 302 processes, e.g. surface wave modulation properties and associated Stokes drift (recall that 303 DoppVis observes Lagrangian current). 304

7 Discussions and summary

Direct spatial and temporal observations of the lower atmosphere, sea surface and 306 upper ocean are crucial for improved knowledge of air-sea interaction. However, the broad 307 range of scales, or equivalently the strong spatial and temporal variability of theses in-308 teractions (see figure 1), make this a formidable theoretical, numerical, and observational 309 challenge. Traditional in-situ assets such as moorings and buoys are limited by their spa-310 tial coverage or their potential spatial biasing, in particular near an ocean front where 311 buoys or drifters can cluster, while satellite imagery estimates important quantities like 312 313 wind, significant wave height and currents through indirect methods (often based on σ_0 relations) and can also only sample at sparse time intervals. Observations and models 314 have revealed two-way coupling processes between ocean currents and wind (Chelton & 315 Xie, 2010; Wenegrat & Arthur, 2018) and ocean currents and waves (Marechal & de Marez, 316 2022; Wang et al., 2020) at both the mesoscale and submesoscale. 317

In this work, we present unique coincident and collocated suborbital high-resolution observations of SST, surface currents and directional properties of surface waves collected from an airborne instrument (MASS), off the coast of Southern California, across submesoscales features.

Central to this work is the development of a novel airborne instrument, DoppVis, that enables high-resolution observations of surface currents, vertical and horizontal shear alongside the other MASS instruments to capture quickly evolving features (both in time and space) such as submesoscale fronts, but also tidal estuarine flows, and more generally ocean phenomena that require frequent revisits (e.g oil spills).

The combination of surface waves, currents, and SST remotely observed from this 327 instrument provides a unique opportunity to examine how currents at scales ranging from 328 1-100 km modulate surface gravity waves, i.e. bulk (e.g. significant wave height), direc-329 tional and spectral properties. Observations of near-surface currents at such high spa-330 tial resolution also enable the investigation of the transition between balanced and un-331 balanced motion, and the flow structure (vorticity-strain space), crucial to developing 332 a better understanding of the underlying physics of submesoscale processes such as fron-333 togenesis and frontal arrest. 334

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³⁴⁰ Open research

Data availability: All presented data will be found at the UCSD Library Digital Collection, https://doi.org/10.6075/J0F76CRK (data archiving is underway).

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Airborne observations of surface winds, waves and currents from meso to submesoscales

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Key Points:

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13	٠	Unique coincident and collocated airborne observations of SST, surface currents
14		and properties of surface waves across submesoscales features
15	•	A new airborne instrument enables observations of surface currents, vertical and
16		horizontal shear to capture quickly evolving ocean features
17	•	Such observations are crucial to develop better understanding of the physics of sub-
18		mesoscale processes and wave-current interaction

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19 Abstract

In this work we present a unique set of coincident and collocated high- resolution obser-20 vations of surface currents and directional properties of surface waves collected from an 21 airborne instrument, the Modular Aerial Sensing System (MASS), collected off the coast 22 of Southern California. High-resolution observations of near surface current profiles and 23 shear are obtained using a new instrument, DoppVis, capable of capturing horizontal spa-24 tial current variability down to 128m resolution. This data set provides a unique oppor-25 tunity to examine how currents at scales ranging from 1-100 km modulate bulk (e.g. sig-26 nificant wave height), directional and spectral properties of surface gravity waves. Such 27 observations are a step toward developing better understanding of the underlying physics 28 of submesoscale processes (e.g. frontogenesis and frontal arrest) and the nature of tran-29 sitions between mesoscale and submesoscale dynamics. 30

³¹ Plain Language Summary

In recent years, through improvement of computational resolution of global ocean 32 models, scientists have begun to suspect that kilometer-scale eddies, whirlpools and fronts, 33 called "submesoscale" variability, make important contributions to horizontal and ver-34 tical exchange of climate and biological variables in the upper ocean. Such features are 35 challenging to analyze, because of their size (and how quickly they evolve; within hours), 36 they are too large to study from a research vessel but smaller than regions typically stud-37 ied with satellite measurements. In this work, we use a research aircraft instrumented 38 to characterize ocean currents, temperature, color (in turn chlorophyll concentration) 30 and the properties of surface waves over an area large enough to capture submesoscale 40 processes. This approach is a step forward in understanding and quantifying the under-41 lying physics of submesoscale processes, and in turn develops parameterization that can 42 help improve the fidelity of weather and climate models. 43

44 1 Introduction

The transfer of mass, momentum, and energy between the atmosphere and ocean 45 are complex due to their interactions across a broad range of space and time scales (Melville, 46 1996). A better understanding of the physics of these processes is fundamental for im-47 proved parameterizations used in coupled air-sea models of weather and climate, par-48 ticularly as Earth's climate changes (Cavaleri et al., 2012). For example, although the 49 importance of surface waves in these models has long been acknowledged, only relatively 50 recently have global models included physics-based models of their effects (see, for ex-51 ample McWilliams & Restrepo, 1999; Sullivan & McWilliams, 2010; Li et al., 2016). Specif-52 ically, the effects of the non-breaking surface wave induced transport that catalyses Lang-53 muir circulations have been shown to reduce errors in sea surface temperature (Belcher 54 et al., 2012), crucial to climate modelling. 55

Submession currents have horizontal scales on the order of 0.1-10 km and 56 have recently been hypothesized to make important contributions to vertical exchanges 57 of climate and biological variables in the upper ocean as well as provide a pathway from 58 energetically rich large scale flows to small scale dissipation. Model studies and limited 59 observations (e.g. D'Asaro et al., 2018) show that submesoscale vertical exchange is con-60 centrated near kilometer-scale fronts, jets, and eddies (McWilliams, 2016). Submesoscale 61 physics are at the smallest scales that have been resolved in global ocean models, where 62 their net effect on heat exchange between the ocean and atmosphere has shown to be 63 much larger than mesoscale eddies (Su et al., 2018). However, these simulations are sen-64 sitive to the parameterized physics of the smaller scale motion, which remains poorly un-65 derstood. To address the fundamental questions of the observed nature of submesoscale 66 dynamics and the interactions between submesoscale dynamics and smaller scale surface 67 wave processes, a comprehensive set of novel, coincident and collocated measurements 68

69 of the dynamical variables is needed to improve state-of-the-art high-resolution simula-70 tions of weather and climate, and better understand vertical exchanges of heat and bio-

71 geochemical tracers.

Due to non-linear coupling between oceanic and atmospheric processes, including 72 currents, winds, and waves, coincident observations are necessary to understand these 73 dynamics. A number of model and observational studies have demonstrated the variety 74 of ways by which air-sea interaction can induce horizontal divergence of surface currents 75 and thus force vertical velocities. Coincident observations of surface vector winds and 76 77 currents are needed to better understand the coupling of winds and currents and in turn improve surface flux parameterizations (Bourassa et al., 2019). The wind stress can be 78 modified by SST and velocity gradients (Dewar & Flierl, 1987; Fairall et al., 1996; Chel-79 ton et al., 2004; O'Neill et al., 2005) while the resulting convergence of the ocean Ek-80 man layer can be modified by surface vorticity (Stern, 1965; McGillicuddy Jr et al., 2007). 81 At submesoscales, these effects are expected to increase in intensity. Frontal structures 82 are strongly affected by the relative direction of the wind with *downfront* Ekman trans-83 port sharpening the fronts and inducing vertical exchange while *upfront* transport subdues the front and stratifies the upper ocean (Thomas et al., 2005). Surface waves and 85 wave breaking can be strongly modulated at fronts (Romero et al., 2017; Vrećica et al., 86 2022), suggesting that even the basic formulation of air-sea exchange in terms of sim-87 ple bulk coefficients will likely break down at sufficiently small scales. As such, an un-88 derstanding of submesoscale structure and vertical velocity requires that air-sea inter-89 action parameters be observed simultaneously with submesoscale measurements. Of par-90 ticular interest are observations of surface and near surface currents, wave breaking (Vrećica 91 et al., 2022), and the directional properties of ocean surface waves. While much progress 92 have been made on characterizing the latter (e.g. Herbers et al., 2012; Lenain & Melville, 93 2014; Melville et al., 2016; Lenain & Melville, 2017; Lenain & Pizzo, 2020), collecting 94 observations of near surface currents, i.e. from the surface down to several meters depth, 95 remains challenging and spatially limited in large part due to the presence of waves, which 96 induce platform motions and additional sources of background noise. An alternative ap-97 proach to traditional in-situ techniques is to infer current profiles remotely based on ob-98 servations of the spatio-temporal evolution of surface waves that follow a dispersion re-99 lationship, pursued in the present study. 100

In this work, we present a unique set of coincident and collocated observations of 101 high-resolution surface currents and directional properties of surface waves collected from 102 an airborne instrument, the Modular Aerial Sensing System (MASS), off the coast of South-103 ern California in May 2021 as part of the ONR-funded Task Force Ocean (TFO) research 104 initiative and in October 2021 as part of the NASA-funded Submesoscale Ocean Dynam-105 ics Experiment (S-MODE) research initiative. Two of the TFO research flights were ded-106 icated to collecting observations across two small counter-rotating eddies separated by 107 approximately 100 km. One of the S-MODE flights collected observations across a sharp 108 SST front. This data set provides a unique opportunity to examine how currents at scales 109 ranging from 1-100 km modulate surface gravity waves, i.e. bulk, directional and spec-110 tral properties. 111

This paper is structured as follows: the overview of the experiment and processing techniques is given in section 2 and 3. In sections 4 through 6, analysis of meso to submesoscales surface kinematics and currents collected during two experiments in Southern California is discussed. Section 7 discusses potential implications for submesoscale and air-sea interaction studies and presents some summary points.

117 **2** Experiments

In this study we consider observations collected during two distinct experiments.
 The first was conducted as part of the "Platform Centric ASW Processing with Through-

the-Sensor Data Assimilation and Fusion" project, funded through the ONR Task Force
Ocean (TFO) initiative with the aim of collecting a simultaneous combination of acoustic, air-sea interaction and oceanographic measurements. Observations from a research
vessel, a drifting instrument array, autonomous surface vehicles and a research aircraft
were collected in May 2021, approximately 45km offshore of San Diego, CA, in the vicinity of CalCoFi Line #90.

The second experiment was conducted as part of the NASA S-MODE program, a project that aims to characterize the contribution of submesoscale ocean dynamics to vertical and horizontal transport in the upper ocean by employing a combination of aircraftbased remote sensing measurements of the ocean surface, in-situ measurements from research vessels and a variety of autonomous oceanographic platforms, and numerical modeling (Farrar et al., 2020). The "pilot" experiment considered here was conducted in the fall of 2021 off the coast of San Francisco, CA.

Data collected from instrumented Wave Gliders (Grare et al., 2021) and an airborne instrument, the SIO Modular Aerial Sensing System (MASS, Melville et al., 2016) during these two field programs are considered in the analysis. During both experiments the MASS instrument was installed on a Twin Otter DHC-6 aircraft (Twin Otter International, Grand Junction, CO).

3 SIO-MASS DoppVis instrument: Enabling novel airborne observations of near-surface currents

The Modular Aerial Sensing System (MASS) is an airborne instrument developed 140 at the Air-Sea Interaction Laboratory (SIO) to simultaneously collect observations of sea 141 surface temperature and ocean color (Melville et al., 2016; Lenain & Pizzo, 2021a), winds 142 and mean-square slope (Lenain et al., 2019), surface waves (Lenain & Melville, 2017; Lenain 143 & Pizzo, 2020), and ocean topography (Villas Bôas et al., 2022), at horizontal scales rang-144 ing from sub-meter to mesoscales. Over the past 11 years, the instrument was flown for 145 more than 30 missions, covering a broad range of environmental conditions, locations, 146 and applications. Details on the system performance and various applications of the MASS 147 can be found in Melville et al. (2016); Lenain and Melville (2017); Lenain et al. (2019); 148 Lenain and Pizzo (2020); Vrećica et al. (2022). 149

In 2020, we started the development and integration of a new sensor into the MASS 150 instrument, called "DoppVis", to obtain coincident observations of surface currents along-151 side the MASS observations listed above. The approach used is to infer currents from 152 optical observations of the spatio-temporal evolution of surface waves, whose dispersion 153 is altered by the presence of an underlying current. This technique has been primarily 154 used with radar technology (e.g. Stewart & Joy, 1974; Campana et al., 2016; Lund et 155 al., 2015) then later applied to airborne video imagery (Dugan et al., 2001; Dugan & Pi-156 otrowski, 2003; Anderson et al., 2013). 157

Starting from the dispersion relation for small-amplitude linear waves propagating on top of a depth-varying current,

$$\omega(\mathbf{k}) = \omega_0(k) + \mathbf{c}(k) \cdot \mathbf{k},\tag{1}$$

where ω is the wave frequency, ω_0 is the frequency in the absence of currents, i.e equal to \sqrt{gk} in deep water, $\mathbf{k} = (k_x, k_y)$ is the wavenumber, $k = |\mathbf{k}|$, and \mathbf{c} is the Doppler shift velocity due to the underlying current. Following Stewart and Joy (1974), assuming the waves are in deep-water, \mathbf{c} can be approximated as a weighted average of the current profile as a function of depth such that

$$\mathbf{c}(k) = 2k \int_{-\infty}^{0} \mathbf{U}(z) e^{2kz} \mathrm{d}z, \qquad (2)$$

where $\mathbf{U}(z) = (U, V)$ is the Lagrangian mean current profile as a function of depth z(Pizzo et al., 2022). Based on this relationship, one can assign an effective depth z_e to the measured Doppler velocities $\mathbf{c}(k)$ by finding the depth at which the Doppler velocity is equal to the current (Stewart & Joy, 1974; Smeltzer et al., 2019), such that $z_e(k) = -1/2k$. This is referred to as the Effective Depth Method (EDM) in Smeltzer et al. (2019).

The DoppVis instrument collects visible imagery of the ocean surface using a Nikon 163 D850 camera with 14mm lens mounted with a 90 degree rotation (long edge of image 164 parallel with flight track) and a 30 degree positive pitch angle from nadir (pointing slightly 165 ahead of aircraft). The camera is synchronized to a coupled GPS/IMU system collect-166 ing images at a 2Hz frame rate. Raw images are carefully calibrated for lens distortion 167 and boresight misalignment with the GPS/IMU over a hard terrestrial target, then geo-168 referenced and exported with reference to WGS84 datum with a UTM zone 10 projec-169 tion (EPSG 32610) at 50cm horizontal resolution. Each image is then interpolated on 170 a regular grid, to enable the generation of 3D cubes of imagery (time, UTM X, UTM 171 Y) of set duration and dimension (N_x, N_y) , where $N_x = N_y$, typically in the range of 172 128 to 512m. The number of collected data cubes in the cross and along track direction 173 of the aircraft varies as a function of aircraft altitude. All data presented here were col-174 lected at 1500 AMSL, corresponding to approximately two $256 \times 256 \text{ m}^2$ cubes in the cross-175 track direction. Following the same approach described in (Smeltzer et al., 2019), all cubes 176 of space-time data are converted to wavenumber-frequency space using a 3D FFT. Each 177 of these 3D spectra are then averaged in the cross-track and along-track direction (1 km 178 bin) to improve SNR. Doppler shift velocities are extracted from the spectrum as a func-179 tion of wavenumber by masking the spectrum into wavenumber magnitude bins (bin half-180 width of $4\pi/N_x$, where for each bin the current $\mathbf{c}(k)$ is estimated using a normalized 181 scalar product method (Huang et al., 2016; Streßer et al., 2017) with a Gaussian char-182 acteristic function (Smeltzer et al., 2019) peaked along the linear dispersion relation. 183

Figure 1(a) shows an example of current profiles (U, V) collected from DoppVis dur-184 ing an overflight of an instrumented Wave Glider during the S-MODE experiment on Novem-185 ber 4, 2021 at 17:20 UTC. The Wave Glider was instrumented with an upward-looking 186 Nortek Signature 1000 ADCP (orange squares) and a downward-looking ADCP (Tele-187 dyne RDI Workhorse 300kHz), carefully motion compensated using an onboard GPS/IMU 188 system (Grare et al., 2021). Observations from DoppVis and the wave glider were col-189 lected within 5min and no further than 500m from each other to minimize any error as-190 sociated with natural spatial and temporal variability. We find good agreement between 191 in-situ and remotely sensed observations of near-surface current (U, V), with a bias = 192 -0.014 m/s and rms deviation = 0.052 m/s, and a coefficient of determination $R^2 = 0.96$. 193

Finally, airborne observations of SST collected from MASS on November 4, 2021, 194 over the entire domain along with current estimates (1km along-track resolution) from 195 DoppVis at two depths, $z = -1.5 \pm 0.5$ m (black arrows) and $z = -0.4 \pm 0.1$ m (red ar-196 rows) are shown in figure 1(c). Note the correlation between features present in the SST 197 fields and the surface currents from DoppVis. Throughout the domain, we consistently 198 find larger magnitudes of the eastern component of the current closer to the surface, likely 199 caused by Stokes drift included in the Lagrangian current observed by DoppVis (Pizzo 200 et al., 2022). Wind and waves were coming from the west at the time of the flight. 201

4 Wave-current interactions from meso- to submesoscales

The collocated observations from the multiple instruments on the MASS allow for investigation of interactions between currents and other oceanographic properties such as waves, heat, and biological communities. We now examine observations collected during the "Platform Centric ASW Processing with Through-the-Sensor Data Assimilation and Fusion" TFO project in May 2021 to illustrate the importance of high spatial resolution, collocated observations for studying wave-current observations.



Figure 1. (a) Current profiles (U, V) collected from DoppVis (gray triangles) during an overflight of an instrumented Wave Glider during the S-MODE experiment on November 4, 2021 at 17:20 UTC. The Wave Glider is equipped with an upward-looking Nortek Signature 1000 ADCP (orange squares) and a downward-looking ADCP (Teledyne RDI Workhorse 300kHz; blue circles). The location of the Wave Glider with respect to the DoppVis observations is shown in (c). (b) Comparison between in-situ observations of currents collected at a 2 to 3m water depth obtained from DoppVis and two Wave Gliders, named WHOI43 and STOKES, for the times and locations of overflights (within 500m) during the entire S-MODE experiment. (c) Airborne observations of SST collected from MASS on November 4, 2021 over the entire domain along with current estimates (1km along-track resolution) from DoppVis at two depths, deep, $z = -1.5 \pm 0.5m$ (black arrows) and shallow, $z = -0.4 \pm 0.1m$ (red arrows).

Several research flights were dedicated to collecting observations across two counter-209 rotating eddies separated by approximately 100 km. Figure 2(a) shows depth-averaged 210 (0.3-2m) surface currents (250 m along-track resolution) along with significant wave height 211 measurements collected coincidentally with the MASS instrument. The black contours 212 represent the sea surface height (ssh, AVISO) at the time of the observations, while cor-213 responding surface currents from HYCOM (GOF 3.1, GLBy0.08-expt93.0) are shown as 214 gray quivers. We find the HYCOM estimated currents to be in good agreement with DoppVis 215 observations at the mesoscale, but strikingly miss many of the submesoscale features through-216 out the domain. We also find significant modulation of the surface wave properties across 217 the domain that is not caused by temporal (e.g. inertial) variability. 218

Currents modulate the properties of surface gravity waves both through wave refraction (e.g. Ardhuin et al., 2017; Romero et al., 2017, 2020; Bôas et al., 2020; Pizzo & Salmon, 2021) and local effects (Rascle et al., 2016, 2017; Lenain & Pizzo, 2021b), therefore high-resolution measurements of currents and spectral wave properties is needed to better understand wave-current interaction from meso- to submesoscales.

Figure 2 shows a subset of these observations, focusing on an area of the domain with significant submesoscale variability in ocean current magnitude and direction. We find large modulation of surface currents, at scales of 1-8 km associated with submesoscale variations within the mesoscale eddy. Panel (b) shows a sharp SST front at 32.885°N with a change in temperature of approximately $0.5^{\circ}C$, associated with a velocity gradient $\partial U/\partial x \approx 10f$, where f is the Coriolis frequency.

Despite approximately constant surface wind (figure 2c) and significant wave height 230 (figure 2d) as a function of latitude on this transect, the mean spectral saturation $\langle B \rangle$ 231 (Romero et al., 2017) had significant variation (figure 2e). The saturation spectrum is 232 defined as $B = \phi(k)k^3$, where ϕ is the omnidirectional wave spectrum (Phillips, 1984, 233 1985). The mean saturation is obtained by averaging B over the saturation range of the 234 wave spectrum, with a lower bound k_n as defined in Lenain and Pizzo (2020). At that 235 time, wind waves were approximately oriented in the same direction as the current V236 (figure 2d). 237

We find that the mean saturation $\langle B \rangle$ generally follow the same evolution as V. While this comparison remains qualitative, it hints at the importance of incorporating wave-current interaction in wave spectrum parameterization, along the line of Lenain and Pizzo (2021a) that investigated the use of WKB to predict the modulation of surface gravity waves by internal wave currents.

Though beyond the scope of this work, but a focus of further studies in particular as part of the S-MODE program, we expect that these collocated, coincident observations will lead to new insights about coherent features that arise from wind-wave-current coupling, or some combination of those processes, which may be particularly important for the flux of climatically-important variables such as heat and biogeochemical tracers between atmosphere and ocean (Li et al., 2016; Smith et al., 2016; Verma & Sarkar, 2021; Freilich & Mahadevan, 2021; Gula et al., 2022).

5 Transitions from mesoscale to submesoscale motion

High-resolution measurements of surface currents enable the characterization of the 251 transition between geostrophically balanced motion at larger scales and unbalanced (ageostrophic 252 and wave) motion at smaller scales (Chereskin et al., 2019). In particular, determining 253 the dynamics that predominate at the submesoscale has important implications for the 254 vertical structure of the upper-ocean (Cronin et al., 2019) and the spatio-temporal dis-255 tribution of energy dissipation (Buckingham et al., 2019; Dong et al., 2020; Schubert et 256 al., 2020; Ajayi et al., 2021), and can lead to improvement of ocean-atmosphere coupled 257 models (Li et al., 2016). While DoppVis observations are accurate enough to resolve sub-258



(a) MASS significant wave height H_s and DoppVis surface current observations Figure 2. (250m along-track resolution) collected during the TFO experiment on May 19, 2021. The black contours represent the sea surface height (ssh, AVISO) at the time of the observations, while corresponding surface currents from HYCOM are shown as gray quivers. The larger scale trends of the surface currents are in general agreement with the HYCOM product, though the latter completely misses significant submesoscale features. We find significant modulation of the wave conditions across the domain, by up to 20%. (b) Subset of the data presented in (a) (nortwestern part of the flight) with the track colorcoded for measured SST. Note the strong modulation of surface currents over very short distances (1-8 km), coinciding in places with sharp SST fronts (for example at latitude around 32.885° in the northern part of the operation area). (c) and (d) show the wind speed U_{10} and significant wave height H_s collected along the NNW-SSE track shown in panel (b). (e) North component of the current V and mean spectral saturation $\langle B \rangle$, all plotted over the same range of latitude as in panel (b). Note the modulation of $\langle B \rangle$ associated with changes in V, the component of the current aligned with the waves, hinting at local wavecurrent interaction.



Figure 3. Comparison of Kinetic energy (KE) spectra collected from the DoppVis Lagrangian surface currents on May 19 2021 during one of the long East - West transects shown in figure 2(a) compared with in-situ ADCP observations (20-m depth) and LLC4320 hourly model products from Chereskin et al. (2019).

mesoscale currents (horizontal resolution of less than 500 m), the airborne platform also 259 travels quickly enough to collect nearly synoptic observations at the mesoscale (100 km). 260 This is particularly valuable because it allows for examination of spatial variability across 261 a range of scales and cross-scale interactions by not aliasing temporal variability. Fig-262 ure 3 shows kinetic energy (KE) spectra computed from the DoppVis surface current ob-263 servations collected during one of the long East-West transects shown in figure 2(a) com-264 pared with the average of 11 years of in-situ ADCP observations (20-m depth) and 1 year 265 of LLC4320 hourly model products from Chereskin et al. (2019). The agreement at low 266 wavenumbers (>10 km) is remarkable, as well as the fact that the kinetic energy spec-267 trum from DoppVis maintains a nearly continuous slope for higher wavenumbers, down 268 to 1 km scale. Feedbacks on ocean currents from wind and waves is likely important for 269 determining dynamics at scales of less than 10 km where submesoscale processes pre-270 dominate (Haney et al., 2015; Suzuki & Fox-Kemper, 2016; Yuan & Liang, 2021). Un-271 derstanding these couplings needs to be advanced by direct observations that span scales 272 from submesoscale to mesoscale. Characterization of the transition between geostroph-273 ically balanced motion and unbalanced dynamics is further explored in forthcoming work. 274

²⁷⁵ 6 Velocity gradient statistics

Submesoscale dynamics are defined by the large velocity gradients at these scales, with O(1) Rossby number (McWilliams, 2016). Consistent with the small spatial scale of submesoscale dynamics, they also have fast temporal scales (Callies et al., 2020). Observations of not just velocity, but 2D velocity gradients of quickly evolving submesoscale features are therefore important for understanding submesoscale dynamics, but require



Figure 4. (a) Surface vorticity–strain and (b) divergence-strain JPDFs computed from 128m resolution DoppVis products collected on October 5 2021 during the S-MODE pilot program. (c) and (d) Surface vorticity–strain for $\Delta/f > 0.8$ and $\Delta/f < -0.8$ respectively. The dashed lines represents the $\sigma = \zeta$ lines.

synoptic observations over scales of tens of kilometers, which is not feasible with relatively slow-moving ship-based observations. Previous observations of submesoscale velocity gradients have used two ships (Shcherbina et al., 2013). Aircraft-based observations offer a well-suited platform for studying submesoscale dynamics due to their fast
speed relative to ships and much higher resolution measurements than what is afforded by satellites.

In recent years, through numerical (Balwada et al., 2021) and observational (Shcherbina 287 et al., 2013) studies, the surface vorticity-strain Joint Probability Density Function (JPDF) 288 has been demonstrated as a useful tool to distinguish between different flow regimes, in 289 particular in the context of submesoscale dynamics. In figure 4, we show vorticity-strain 290 and divergence-strain JPDFs computed from high-resolution DoppVis surface current 291 observations collected during the S-MODE pilot experiment on October 2021 (60km long 292 section). During that portion of the flight, the aircraft flew at 3000ft AMSL, increasing 293 the swath width of surface current observation up to 2 km, with a horizontal resolution 294 (cross and along-track) of 128 m. Vorticity ζ , strain rate σ and divergence Δ are then 295 computed from the gradients of the measured surface currents. These velocities are first 296 low-pass filtered (3-point tophat) prior to computing gradients. Overall, the distribu-297 tion is skewed toward positive ζ/f values, with $\sigma/f > \zeta/f$ consistent with frontal struc-298 tures, and classical submesoscale frontogenesis. This is particularly evident in panel (d), 299 where the vorticity-strain JPDF for $\Delta/f < -0.8$ (i.e. convergence) is shown to be skewed 300 to positive values. The same JPDF for positive values of Δ/f , larger than 0.8 (panel c) 301 is not skewed toward negative ζ/f values, perhaps implying the contribution of other 302 processes, e.g. surface wave modulation properties and associated Stokes drift (recall that 303 DoppVis observes Lagrangian current). 304

7 Discussions and summary

Direct spatial and temporal observations of the lower atmosphere, sea surface and 306 upper ocean are crucial for improved knowledge of air-sea interaction. However, the broad 307 range of scales, or equivalently the strong spatial and temporal variability of theses in-308 teractions (see figure 1), make this a formidable theoretical, numerical, and observational 309 challenge. Traditional in-situ assets such as moorings and buoys are limited by their spa-310 tial coverage or their potential spatial biasing, in particular near an ocean front where 311 buoys or drifters can cluster, while satellite imagery estimates important quantities like 312 313 wind, significant wave height and currents through indirect methods (often based on σ_0 relations) and can also only sample at sparse time intervals. Observations and models 314 have revealed two-way coupling processes between ocean currents and wind (Chelton & 315 Xie, 2010; Wenegrat & Arthur, 2018) and ocean currents and waves (Marechal & de Marez, 316 2022; Wang et al., 2020) at both the mesoscale and submesoscale. 317

In this work, we present unique coincident and collocated suborbital high-resolution observations of SST, surface currents and directional properties of surface waves collected from an airborne instrument (MASS), off the coast of Southern California, across submesoscales features.

Central to this work is the development of a novel airborne instrument, DoppVis, that enables high-resolution observations of surface currents, vertical and horizontal shear alongside the other MASS instruments to capture quickly evolving features (both in time and space) such as submesoscale fronts, but also tidal estuarine flows, and more generally ocean phenomena that require frequent revisits (e.g oil spills).

The combination of surface waves, currents, and SST remotely observed from this 327 instrument provides a unique opportunity to examine how currents at scales ranging from 328 1-100 km modulate surface gravity waves, i.e. bulk (e.g. significant wave height), direc-329 tional and spectral properties. Observations of near-surface currents at such high spa-330 tial resolution also enable the investigation of the transition between balanced and un-331 balanced motion, and the flow structure (vorticity-strain space), crucial to developing 332 a better understanding of the underlying physics of submesoscale processes such as fron-333 togenesis and frontal arrest. 334

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³⁴⁰ Open research

Data availability: All presented data will be found at the UCSD Library Digital Collection, https://doi.org/10.6075/J0F76CRK (data archiving is underway).

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