

# Hybrid Modeling of Evapotranspiration: Inferring Stomatal and Aerodynamic Resistances Using Combined Physics-Based and Machine Learning

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November 30, 2022

## Abstract

The process of evapotranspiration transfers water vapour from vegetation and soil surfaces to the atmosphere, the so-called latent heat flux ( $\lambda E$ ), and thus crucially modulates Earth’s energy, water, and carbon cycles. Vegetation controls  $\lambda E$  through regulating the leaf stomata (i.e., surface resistance  $r_s$ ) and through altering surface roughness (aerodynamic resistance  $r_a$ ). Estimating  $r_s$  and  $r_a$  across different vegetation types proves to be a key challenge in predicting  $\lambda E$ . Here, we propose a hybrid modeling approach (i.e., combining mechanistic modeling and machine learning) for  $\lambda E$  where neural networks independently learn the resistances from observations as intermediate variables. In our hybrid modeling setup, we make use of the Penman-Monteith equation based on the Big Leaf theory in conjunction with multi-year flux measurements across different forest and grassland sites from the FLUXNET database. We follow two conceptually different strategies to constrain the hybrid model to control for equifinality arising when estimating the two resistances simultaneously. One strategy is to impose an a priori constraint on  $r_a$  based on our mechanistic understanding (theory-driven strategy), while the other strategy makes use of more observational data and adds a constraint in predicting  $r_a$  through multi-task learning of the latent as well as the sensible heat flux ( $H$ ; data-driven strategy). Our results show that all hybrid models exhibit a fairly high predictive skill for the target variables with  $R^2 = 0.82-0.89$  for grasslands and  $R^2 = 0.70-0.80$  for forests sites at the mean diurnal scale. The predictions of  $r_s$  and  $r_a$  show physical consistency across the two regularized hybrid models, but are physically implausible in the under-constrained hybrid model. The hybrid models are robust in reproducing consistent results for energy fluxes and resistances across different scales (diurnal, seasonal, interannual), reflecting their ability to learn the physical dependence of the target variables on the meteorological inputs. As a next step, we propose to test these heavily observation-informed parameterizations derived through hybrid modeling as a substitute for overly simple ad hoc formulations in Earth system models.

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Received xxxxxx; Accepted for publication xxxxxx; Published xxxxxx

## Abstract

The process of evapotranspiration evaporates liquid water from vegetation and soil surfaces to the atmosphere, the so-called latent heat flux ( $Q_{LE}$ ), and modulates Earth's energy, water, and carbon cycle. Vegetation controls  $Q_{LE}$  by regulating leaf stomata opening (surface resistance  $r_s$  in the Big Leaf approach) and by altering surface roughness (aerodynamic resistance  $r_a$ ). Estimating  $r_s$  and  $r_a$  across different vegetation types is a key challenge in predicting  $Q_{LE}$ . We propose a hybrid approach that combines mechanistic modeling and machine learning for modeling  $Q_{LE}$ . The hybrid model combines a feed-forward neural network which estimates the resistances from observations as intermediate variables and a mechanistic model in an end-to-end setting. In the hybrid modeling setup, we make use of the Penman-Monteith equation based on the Big Leaf approximation in conjunction with multi-year flux measurements across different forest and grassland sites from the FLUXNET database. This hybrid model setup is successful in predicting  $Q_{LE}$ , however, this approach yields equifinality. We follow two different strategies to constrain the hybrid model to control for equifinality arising when estimating the two resistances simultaneously. One strategy is to impose an *a priori* constraint on  $r_a$  based on mechanistic understanding (theory-driven strategy), while the other strategy makes use of more

observational data and adds a constraint in predicting  $r_a$  through multi-task learning of both latent and sensible heat flux ( $Q_H$ ; data-driven strategy) together. Our results show that all hybrid models exhibit a high predictive skill for the target variables with  $R^2 = 0.82-0.89$  for grasslands and  $R^2 = 0.70-0.80$  for forest sites at the mean diurnal scale. The predicted  $r_s$  and  $r_a$  show strong physical consistency across the two regularized hybrid models, but are physically implausible in the under-constrained hybrid model. The hybrid models are robust in reproducing consistent results for energy fluxes and resistances across different scales (diurnal, seasonal, interannual), reflecting their ability to learn the physical dependence of the target variables on the meteorological inputs. As a next step, we propose to test these heavily observation-informed parameterizations derived through hybrid modeling as a substitute for *ad hoc* formulations in Earth system models.

Keywords: Hybrid modeling, physics-constrained, machine learning, deep learning, multi-task learning, evapotranspiration, surface resistance, aerodynamic resistance

## 1. Introduction

Evapotranspiration, i.e. the surface latent heat flux ( $Q_{LE}$ ), plays a key role in driving Earth's energy, water, and carbon cycles, and is primarily controlled by dynamic meteorological conditions and soil water conditions as well as more static properties such as soil characteristics and plant traits (Jung et al., 2010; Dou & Yang, 2018; Ajami, 2021). The characterization of  $Q_{LE}$ , however, remains challenging as our understanding of the underlying processes that regulate the exchange flux of water between land and atmosphere is still limited (Friedl, 1996; Sellers et al., 1997; Wang & Dickinson, 2012; Chen et al., 2014; Massmann et al., 2019). While the physical drivers that cause water to evaporate are well described and understood, the influence of the biological control on  $Q_{LE}$ , mainly the transpirative water flux, is more difficult to assess. The key problem is that we cannot easily formulate universally valid mechanistic laws to describe ecosystem land-atmosphere interactions in the presence of changing atmospheric and soil conditions. As a consequence, empirical formulations, especially for surface and aerodynamic resistance, remain used in process-based models, which can lead to large uncertainties in predicting  $Q_{LE}$  (Polhamus et al., 2013). In this study, we propose a hybrid modeling (physics + machine learning) approach that allows inference of these biophysical controls based on observational data of  $Q_{LE}$  across ecosystems, while adhering to known physical laws (Reichstein et al., 2022).

Plants critically influence  $Q_{LE}$  mainly through their direct control of transpiration, but also through shaping aerodynamic surface properties (i.e. roughness). Plants use their leaf stomata to dynamically regulate the water loss to the atmosphere, which not only depends on the atmospheric water demand, but also on soil water availability (Damour et al., 2010; Kennedy et al., 2019; Carminati & Javaux, 2020). Simultaneously, plants use stomata to take up atmospheric  $CO_2$  for photosynthesis (Schulze, 1986; Chaves et al., 2016). To this end, most formulations of stomatal conductance (or the inverse, stomatal resistance  $r_s$ ) are empirical or rely on optimality concepts, such as minimizing the water loss while maximizing carbon assimilation (e.g. Tan et al., 2021). As such, these concepts do not take into account the active transpiration mechanism that some plants use to down-regulate leaf temperature through evaporative cooling to prevent leaf overheating at high irradiance and air temperature (Lin et al., 2017; Drake et al., 2018). Other empirical approaches, e.g., the Jarvis–Stewart formulation, Ball–Berry model, and Leuning model aim to derive parametrizations based on statistical correlations between  $r_s$  (or canopy resistance) and the key environmental variables (Jarvis, 1976; Stewart, 1988; Leuning et al., 1991; Leuning, 1995). These *ad hoc* formulations have several drawbacks, e.g., they are considered too rigid, especially when evaluated in a coupled system of atmosphere-biosphere feedbacks where some of the environmental variables are actually also a function of  $r_s$  (Ronda et al., 2001).

Formulations of how plants affect  $Q_{LE}$  via surface roughness and associated aerodynamic properties are considered less uncertain, but vary considerably among vegetation types (Shaw & Pereira, 1982; Nakai et al., 2008; Maurer et al., 2015). Generally, near-surface wind enhances turbulent mixing and thus the exchange of momentum, mass and heat between the surface and the atmosphere. The surface roughness lengths influence the mechanical turbulence as well as the near-surface atmospheric thermal structure (Vila-Guerau de Arellano et al., 2015). These relationships are formulated in the aerodynamic resistance  $r_a$ , which is conventionally assumed to scale inversely (hyperbola-type function) with wind speed, frictional velocity, and atmospheric instability based on the diagnostic empirical Monin–Obukhov similarity theory (Knauer et al., 2018). Several studies (Chehbouni et al., 1996; Liu et al., 2006; Su et al., 2021; Trebs et al., 2021) demonstrated that these parameterizations might work under controlled settings in the laboratory, yet they show large discrepancies when applied to other real landscapes and vegetation types. Overall, these empirical representations for  $r_s$  and  $r_a$  in deterministic models for  $Q_{LE}$  generally obey physical laws and phenomenological behaviour (Krasnopolsky, 2013; de Bezenac et al., 2017). Yet, they exhibit limited

89 capability to adapt to other or changing vegetation composition or long-term climatic conditions,  
90 especially with respect to soil moisture (Damour et al., 2010; Medlyn et al., 2011; Kennedy et al., 2019).

91 Statistical models have been proposed as alternative approaches to reliably estimate  $Q_{LE}$  due to  
92 their data-adaptiveness (Tramontana et al., 2016; Dou & Yang, 2018; Carter & Liang, 2019; ). In  
93 particular, approaches that use machine learning (ML) techniques are gaining traction because they can  
94 implicitly learn unknown latent processes and constitute a more complete statistical representation of  
95 the processes that influence  $Q_{LE}$  at different scales in space and time (Dou & Yang, 2018; Jung et al.,  
96 2009, 2020). However, these data-driven models are subject to several drawbacks, such as the need for  
97 large amounts of high-quality data, their limited physical consistency, and their lack of mechanistic  
98 interpretability (Karpatne, et al., 2017a,b).

99 The combination of ML and mechanistic modeling, here denoted hybrid modeling, allows to  
100 combine the strengths of both techniques: ensure physical consistency while efficiently harvesting the  
101 growing resource of observational data (Reichstein et al., 2019, 2022). Several studies have successfully  
102 applied hybrid modeling in hydrological applications, such as the characterization of the different  
103 known and unknown variables governing the global water cycle (Kraft et al., 2020, 2022), simulation  
104 of lake temperature dynamics (Jia et al., 2020), and the modeling of global extreme flooding events  
105 (Yang et al., 2019). Other studies focusing on land-atmosphere interactions of ecosystem fluxes, such  
106 as  $Q_{LE}$  (Zhao et al., 2019), showed that these hybrid approaches allow for better extrapolation and  
107 generalization capabilities during extreme conditions.

108 In the methods section 2.1-2.2, we use a hybrid modeling approach and develop different models  
109 of  $Q_{LE}$  using the Penman-Monteith equation (Penman, 1948; Monteith, 1965) and eddy covariance flux  
110 measurements from several grassland as well as forest sites (Baldocchi et al., 2001; Li et al., 2018). Our  
111 hybrid models not only seek to yield accurate predictions of  $Q_{LE}$ , but more importantly should enable  
112 us to learn (interpretability) the functioning and influence of biophysical processes on  $Q_{LE}$ , expressed  
113 through the surface and aerodynamic resistances. We present and explore the problem of equifinality  
114 in our setting (Sec. 2.3.2) (i.e., different combinations of  $r_a$  and  $r_s$  may result in the same  $Q_{LE}$ ) and  
115 propose two conceptually different solutions (theory- versus data-driven) to this issue (Sec. 2.3.3). We  
116 evaluate the predictions of our hybrid models for  $Q_{LE}$ ,  $r_a$  and  $r_s$  against purely statistical models as well  
117 as against established mechanistic models in Sec. 3.

## 2. Methodology

In this section we describe the data pre-processing methods and different model setups taken. Sec. 2.1 describes the data used and processing. Sec. 2.2 defines the physics-based component of the hybrid model, and Sec. 2.3 provides an overview of all the models.

### 2.1 FLUXNET 2015 Data

The global flux network (FLUXNET; <https://fluxnet.org>), a global network of eddy covariance (EC) towers, provides estimates of energy, water and carbon fluxes at the land surfaces across climate regimes and plant functional types (Baldocchi et al., 2001; Li et al., 2018). The measurements in the FLUXNET 2015 Tier 1 dataset are resolved at a half-hourly frequency. Following Reichstein et al., (2005), we select only measured data and omit gap-filled data. Further, we restrict our analysis to energy-balance-corrected measurements, because the EC data do not satisfy the energy balance budget closure which potentially introduces high uncertainty/systematic bias in our results (Wilson et al., 2002). Daytime values are selected based on a threshold of sensible heat flux  $Q_H > 5 \text{ Wm}^{-2}$  and incoming short-wave radiation  $SW_{in} > 50 \text{ Wm}^{-2}$  to avoid stable boundary layer conditions following Lin et al., (2018) and Li et al., (2019). Only positive values are selected for the latent heat flux ( $Q_{LE}$ ), net radiation ( $R_n$ ), soil heat flux ( $Q_G$ ), and vapor pressure deficit (VPD) for daylight data according to Zhou et al. (2016). Winter months between October and March are excluded to focus on surface heat fluxes when the vegetation is active following Zhao et al. (2019). The FLUXNET sites chosen include three forest and three grassland sites with varying climates, site properties and long-term data (Table 1).

### 2.2 The physically-based component: Penman-Monteith equation

Various process-based models exist for the estimation of  $Q_{LE}$ . They can be subdivided into energy, mass transfer-based methods, water balance methods, and aerodynamic methods (Brutsaert, 2005; Zhao et al., 2013). One prominent example is the Penman-Monteith (PM) equation (Penman, 1948; Monteith, 1965) that provides the theoretical basis for determining  $Q_{LE}$  and its response to changing climate and vegetation conditions (Monteith & Unsworth, 2013). The estimation of  $Q_{LE}$  can be traced back to the model proposed by Penman (1948), which combines the energy balance and mass

transfer approaches to estimate evaporation from open water surfaces. The model was later extended to vegetative surfaces (Monteith, 1985; Monteith & Unsworth, 2013; Vialet-Chabrand & Lawson, 2019). The PM equation

$$Q_{LE} = \frac{s_c(R_n - Q_G) + \frac{\rho_a c_p (e_s - e_a)}{r_a}}{s_c + \gamma(1 + \frac{r_s}{r_a})}, \quad (1)$$

describes the latent heat flux  $Q_{LE}$  ( $\text{Wm}^{-2}$ ), where  $R_n$  and  $Q_G$  are measured in ( $\text{Wm}^{-2}$ ),  $r_s$  and  $r_a$  are measured in ( $\text{sm}^{-1}$ ),  $s_c$  is the slope of the saturation vapor pressure-temperature relationship ( $\text{kPa C}^{-1}$ ),  $e_s - e_a$  is the VPD of air ( $\text{kPa}$ ),  $\rho_a$  is the mean air density at constant pressure ( $\text{kg m}^{-3}$ ),  $c_p$  is the specific heat of dry air at constant pressure ( $1004.834 \text{ J kg}^{-1} \text{ C}^{-1}$ ), and  $\gamma$  is the psychrometric constant ( $\text{kPa C}^{-1}$ ).

## 2.3 Overview of models

The following subsections present the different models used that differ in their approach towards being more data- or theory-driven. Each subsection describes in detail the structure and difference between each model. All models were randomly initialized and drawn from a uniform distribution.

### 2.3.1 Inverted Penman-Monteith and pure machine learning model

The PM equation is considered to be physics-based, since core physiological and aerodynamic factors describe the evaporative process (Jain et al., 2008). The equation highlights the relationship between evapotranspiration and surface conductance, which is regulated by the leaf stomata to minimize the water loss to the atmosphere (Hetherington & Woodward, 2003; Damour et al., 2010; Gerosa et al., 2012). Different approaches exist to model surface conductance at the leaf level with various success. The determination of surface conductance at the canopy scale, however, is even more challenging due to canopy heterogeneity and variability in microclimate within the canopy (Bonan et al., 2011; Lin et al., 2018). A common approach is to invert the Penman-Monteith equation for  $r_s$  to obtain the bulk surface resistance and understand its variations

$$r_s = \frac{r_a s_c (R_n - Q_G) + \rho_a c_p (e_s - e_a) - r_a Q_{LE} (s_c + \gamma)}{\gamma Q_{LE}}, \quad (2)$$

assuming that the aerodynamic resistance  $r_a$  is known; a strong assumption as we will revisit later. The inverted PM equation (PM Inv) is used to quantify canopy parameters and expresses the relative significance of advective and radiative energy for  $Q_{LE}$  as a function of the ratio of surface to aerodynamic resistance (Kelliher et al., 1992; Köstner et al., 1992; Zeppel & Eamus, 2008; Zhang et al., 2016).

As a result of the inversion of the PM equation, this leads to highly unstable estimates of the resistances. Therefore, we restrict surface and aerodynamic resistance values derived using Penman-Monteith inversion and empirical formulations (Knauer et al., 2018) based on intervals that are physically realistic (0-2000  $\text{sm}^{-1}$  and 0-500  $\text{sm}^{-1}$ , respectively).

The estimates for  $r_s$  from Eq. 2 derived through inverting the PM equation are referred to here as the PM Inv model. Values for  $r_a$  are estimated using the Big Leaf formulation from Knauer et al. (2018), which calculates  $r_a$  as the sum of aerodynamic resistance for momentum ( $r_{am}$ ) and canopy boundary layer resistance for heat ( $r_{bh}$ )

$$r_{am} = WS/U^{*2}, \quad (3)$$

$$r_{bh} = 6.2 U^{*-0.667}, \quad (4)$$

and

$$r_a = r_{am} + r_{bh}, \quad (5)$$

where WS is wind speed ( $\text{ms}^{-1}$ ) and  $U^*$  is friction velocity ( $\text{ms}^{-1}$ ). The PM Inv model represents a baseline physical model for comparison against pure data-driven models for  $Q_{LE}$ . The pure ML model for  $Q_{LE}$  is set up to evaluate predictions against hybrid models. The pure ML model consists of a feed-forward neural network (FNN) and details about the hyperparameters of the model are found in Table 2 of the Supp. Info. The  $r_s$  is calculated from  $Q_{LE}$  predictions from the pure ML model by using PM



Inv, and  $r_a$  is estimated using the *ad hoc* formulation (Eq. 5) approach. This model is purely data-driven and does not contain any physical constraint regarding  $Q_{LE}$ .

### 2.3.2 Under-constrained hybrid model

The hybrid model estimates  $Q_{LE}$  using the PM equation (Eq. 1), where the two intermediate variables  $r_s$  and  $r_a$  are estimated by two FNNs (Fig. 1). The variables used for predicting  $r_s$  are air temperature (TA), water availability index (WAI), incoming shortwave radiation ( $SW_{in}$ ), mean incoming shortwave potential ( $SW_{pot\ sm}$ ), VPD, and  $R_n$ . The WAI is calculated as the annual cumulative difference between  $Q_{LE}$  and precipitation ( $P$ ). The WAI at time  $t$  ( $WAI_t$ ) is calculated from the difference between  $Q_{LE_t}$  and  $P_t$  added to WAI at the previous time step ( $WAI_{t-1}$ )

$$WAI_t = P_t - Q_{LE_t} + WAI_{t-1} . \quad (6)$$

The variables for predicting  $r_a$  are WS and  $U^*$ . The predictors are normalized using the mean and standard deviation of the training dataset. Thus, the hybrid model predicts first the intermediate (or *latent*) variables  $r_s$  and  $r_a$  and uses them to estimate  $Q_{LE}$  based on the PM equation. The hybrid model predicts  $Q_{LE}$  in end-to-end manner, whereby the loss function minimizes the difference between predicted and observed  $Q_{LE}$ . The loss function is hence defined as the mean absolute difference between the model predictions and observations with  $n$  sample size, and parameters  $\theta$  for  $r_s$  and  $r_a$

$$\min_{\theta_{r_a}, \theta_{r_s}} \sum_{i=1}^n |\hat{Q}_{LE_i} - Q_{LE_i}| . \quad (7)$$

We use the mean absolute error as opposed to mean squared error as it is less sensitive to outliers. Although the two FNNs for  $r_a$  and  $r_s$  take different predictor variables, the hybrid model is under-constrained when simultaneously estimating the two intermediate variables using only one target  $Q_{LE}$ . The proposed hybrid model thus suffers from an equifinality problem. The issue of equifinality, or non-uniqueness, occurs when different model parametrization or structures result in equivalent representations of the system (Beven, 2006; Schmidt et al., 2020).

Thus, many different combinations of  $r_s$  and  $r_a$  can result in the same  $Q_{LE}$  value (Fig. 2).

### 2.3.3 Constrained hybrid models: *a priori* and multi-task learning models

The identification and elimination of equifinality, non-uniqueness, in the physics-based component is one of the key challenges in hybrid modeling (Kraft et al., 2022). One way to reduce equifinality is to restrict the parameter space through model regularization (Fig. 3). This can be achieved through two approaches; either by including additional theory or data in the loss function. The integration of *a priori* knowledge in the loss function (i.e., a regularization) induces an *a priori* constraint on  $r_a$  in the hybrid model based on the empirical formulation presented in Eq. 5 as the formulation for  $r_a$  is considered to be more robust than for  $r_s$ . To do so we regularize the loss function by adding a constraint on the loss minimizing aerodynamic resistance  $\text{Loss}(r_a, \hat{r}_a) / \phi$ . The relative importance of  $r_a$  in the new loss is regulated by  $\phi$ , which is varied between high influence to low influence of theory. Based on multiple model runs, the  $\phi$  value is selected  $\phi$  with minor influence based on prior knowledge in the loss function.

Another way of restricting the parameter space is by extending the framework to model auxiliary target variables whereby auxiliary tasks help regularize the problem objective (Liebel & Körner, 2018). Since the sensible heat flux ( $Q_H$ ) is also dependent on the aerodynamic resistance  $r_a$ , we explore multi-task learning approach by restricting the parameter space through modeling auxiliary variables in a multi-task setting. The multi-task learning approach here uses an intermediate variable regularization by adding  $Q_H$  as an auxiliary target variable in addition to  $Q_{LE}$  (Fig. 3). The estimation of  $Q_H$  is based on the resistance formulation

$$Q_H = \frac{\rho_a c_p (TS - TA)}{r_a}, \quad (8)$$

where TS and TA are surface and air temperature respectively. The TS is estimated using the Stefan-Boltzmann equation

$$TS = \sqrt[4]{\frac{Q_{LW_{out}}}{\sigma \epsilon}}, \quad (9)$$

Where  $Q_{LW_{out}}$  is the outgoing longwave radiation ( $\text{Wm}^{-2}$ ),  $\sigma$  is the Stefan-Boltzmann constant ( $5.789 \times 10^{-8} \text{ Wm}^{-2}\text{K}^{-4}$ ) and  $\epsilon$  is emissivity (dimensionless). The emissivity ranges between 0-1, and the values chosen were based on selecting models with the highest predictive accuracy.

## 2.4 Evaluation

We consider one pure machine learning model, one under-constrained hybrid model (i.e. with no strategy to decouple  $r_a$  and  $r_s$ ), and two constrained hybrid models which make four models in total. The constrained hybrid models consist of either an *a priori* constraint on  $r_a$  or using a multi-task learning approach. For a baseline comparison, we use a pure ML model predicting latent heat flux directly without intermediate resistances and the estimation of the inverted PM equation to evaluate the predictions of the hybrid models. The network architectures and hyperparameters used are similar for the different models (Table 2 in the supplementary information) for a fair comparison. Evaluation metrics such as the root mean square error (RMSE) and mean absolute error (MAE), and coefficient of determination ( $R^2$ ) are used to evaluate the model predictions. To highlight the impact of noise on model performance, we evaluate the model predictions at the half-hourly and 7-day mean aggregated scale. The intermediate variables are assessed against the key meteorological predictor variables to scrutinize physical consistency and plausibility. The target variables are assessed against observations as well as the key meteorological predictor variables to estimate model performance and interpretability. We conduct five model runs with random initializations for each of the hybrid models and for one forest site (DE-Tha) as well as, one grassland site (DE-Gri) to evaluate model robustness at the mean diurnal scale. More information can be found in Table 3 of the supplementary information.

### 3. Results and discussion

#### 3.1 Statistical performance and mechanistic plausibility of the models

We evaluate predicted  $Q_{LE}$  ( $\hat{Q}_{LE}$ ) from all the hybrid models and the pure ML model against observed  $Q_{LE}$  ( $Q_{LE_{obs}}$ ) at half-hourly scale and at 7-day mean aggregates (mean diurnal) for forest (Fig. 4) and grassland (Fig. 5) sites. All models reproduce similar  $Q_{LE}$  patterns compared to observations with minor differences in performance. For forests (Fig. 4), the more flexible models, the under-constrained hybrid model and pure ML model, exhibit a slightly higher performance ( $R^2 = 0.49$ ) in comparison to the multi-task learning model ( $R^2 = 0.48$ ) and the *a priori* constraint model ( $R^2 = 0.46$ ). For grasslands, the performance of all models is generally higher than for forests. We find that the performance of the multi-task learning model exceeds the performance of the *a priori* constraint model and is similar to the pure ML model ( $R^2 = 0.74-0.75$ ) (Fig. 5). This finding could indicate that our theory-based constraint for  $r_a$  might be too rigid and is not supported by the flux observations. Overall, the RMSE ranges from 70-73  $Wm^{-2}$  for forests and 60-71  $Wm^{-2}$  for grasslands at a half-hourly scale

for all models. The MAE at half-hourly measurements range between 50-53  $\text{Wm}^{-2}$  for forests and 43-48  $\text{Wm}^{-2}$  for grasslands for all models. The multi-task learning model provides predictions for  $Q_H$  ( $\hat{Q}_H$ ) (Fig. 6) of similar accuracy compared to the  $Q_{LE}$  predictions for all sites (Fig. 4-5), reaching  $R^2 = 0.53$  for forests and  $R^2 = 0.68$  for grasslands sites at half-hourly scale.

Our results at half-hourly scale are impacted by random measurement noise in the EC data. So that there is plateauing effect in terms of fit of the models due to the irreducible instrument and observation noise. To reduce the effect of this instrumental noise source, we aggregate half-hourly predictions in a 7-day window and calculate the mean diurnal cycle. The results presented in this noise-corrected manner demonstrate an even higher fit between  $Q_{LE_{obs}}$  versus  $\hat{Q}_{LE}$  (Fig. 4-5) and  $Q_{H_{obs}}$  versus  $\hat{Q}_H$  (Fig. 6) for forests and grasslands. The  $R^2$  coefficient increases across all models by 53-70% for forests and 15-25% for grasslands sites based on the aggregated mean diurnal predictions. Further, the RMSE drops by 47-52% for forests, and by 43-48% for grasslands, while MAE also decreases by 47-52% for forests and 42-46% for grasslands. Adjusting noise in  $\hat{Q}_H$  in the same manner also increases  $R^2$  from 0.68 to 0.87 for grasslands, and  $R^2$  from 0.53 to 0.69 for forests (Fig. 6).

To assess the physical plausibility of the presented models, we evaluate their predictions of  $\hat{Q}_{LE}$  against the key predictor for atmospheric dryness, VPD. In all models,  $\hat{Q}_{LE}$  increases sharply at relatively low values of VPD (0-1 kPa), but starts to stabilize and eventually decreases for higher values of VPD ( $> 1$  kPa; Fig 7). This behavior of the models aligns well with other studies that have shown that the transpiration rate increases with increasing VPD at the low and medium range, but starts to decrease again when VPD reaches high values (Buckley, 2005; Massmann et al., 2019; Monteith, 1995; Mott & Peak, 2013). This plant response could reflect their ability to downregulate stomatal conductance as a preemptive measure to decrease water losses and to circumvent damages arising from intense dehydration of the canopy when the lower atmosphere becomes too dry (Farquhar, 1978; Massmann et al., 2019; Vico et al., 2013). Generally, grasslands sites reach higher  $\hat{Q}_{LE}$  values than forest sites for the same VPD range. Again, this result is related to the different plant responses to VPD, since grasses are assumed to exhibit higher surface conductance (lower surface resistance  $r_s$ , respectively) compared to forests, resulting in higher transpiration rates (Garratt, 1992; Jarvis & Stewart, 1979). This aspect is discussed further in Sec. 3.2 when evaluating the learned resistances,  $r_s$  and  $r_a$ .

We next evaluate the hybrid models' consistency with respect to the interannual variability of  $Q_{LE}$  for the different sites. The interannual anomalies are calculated as the difference between the average annual estimates of  $Q_{LE_{obs}}$  in the training dataset and the annual estimates of  $Q_{LE_{obs}}$  and  $\hat{Q}_{LE}$  in the validation and test dataset for the EC data and models, respectively, to evaluate the predictive capacity of the different models (Jung et al., 2009; Besnard et al., 2019). Figures 4 and 5 show the overall fit and performance of the models in predicting interannual anomalies of  $\hat{Q}_{LE}$  compared to observed anomalies of  $Q_{LE_{obs}}$ . The values of  $R^2$  range between 0.47-0.49 for the interannual  $\hat{Q}_{LE}$  anomalies for forests and thus exhibit a comparable performance as at the half-hourly frequency ( $R^2$  ranges between 0.46-0.49) (Fig. 4). We observe a similar behavior at grassland site:  $R^2$  ranges between 0.65-0.75 at the half-hourly scale and between 0.62-0.74 for the interannual  $Q_{LE}$  anomalies (Fig. 5). Overall, the evaluation of the models at multiple temporal scales shows that the models are capable of learning not only the predominant structure of the diurnal and seasonal cycle, but also the subtler year-to-year anomalies. The presented consistency reflects that the models learn the physically correct dependence of the meteorological predictor variables controlling  $Q_{LE}$ .

### 3.2 Evaluation of the learned latent variables $\hat{r}_s$ and $\hat{r}_a$

Next, we evaluate the impact of the  $Q_{LE}$ -controlling resistances  $\hat{r}_s$  and  $\hat{r}_a$  which are treated as intermediate variables in our hybrid approach. First, we plot the inferred estimates of  $\hat{r}_s$  and  $\hat{r}_a$  against the key meteorological drivers, namely VPD and the frictional velocity  $U^*$ , respectively (Fig. 8-9). The behavior of  $\hat{r}_s$  against VPD is consistent across all the models and reflects a similar behavior as presented for  $\hat{Q}_{LE}$ . The predicted  $\hat{r}_s$  shows a gentle increase at lower ranges of VPD, so the stomata are still open for gas exchange with the atmosphere. However, as VPD increases to higher values, the stomata start to close and thus the surface resistance increases sharply (Massmann et al., 2019). Further, we find that  $\hat{r}_s$  is generally lower for grasslands, which explains the generally higher estimates of  $Q_{LE}$  compared to forests, as discussed above (Fig. 7). Another striking finding is that the models seem to be able to identify differences in the physiological functioning across different plant types in controlling  $\hat{r}_s$ . For instance, the inferred relationship of  $\hat{r}_s$  and VPD is very similar for the two forest sites DE-Tha and FR-LBr, which are dominated by evergreen needle-leaf trees, however, is quite different for the more arid site FR-Pue, which is dominated by evergreen broad-leaf trees (Fig. 8 a-c). There, the hybrid models show that on average  $r_s$  rises more steeply with increasing VPD but flattens out at very high VPD (compare fit lines in Fig. 8 a-c). Future research is needed to determine whether this behavior

actually reflects the plants' mechanism for preventing leaf overheating by maintaining some evaporative cooling through the stomata (Lin et al. 2017), or whether it is just an artifact of too sparse data at high VPD. Overall, the inferred  $\hat{r}_s$  through hybrid modeling (Fig. 8 a-c) is much more precise than its conventional derivation by inverting the Penman-Monteith equation while making assumptions for  $r_a$  (Fig. 8d). This aspect constitutes a key advantage of our hybrid approach as opposed to the inversion method, where artificial noise in the flux measurements directly propagates into the inverted estimates of  $\hat{r}_s$  resulting in high artificial variability and a bias in  $\hat{r}_s$  ranging between 0-30% (Wehr & Saleska, 2021).

The inferred relationship for  $\hat{r}_a$  against its key driver  $U^*$  is not consistent across the hybrid models. The two constrained hybrid models, i.e., multi-task learning (Fig. 8f) and *a priori* constraint (Fig. 8g), consistently reflect the expected negative logarithmic relationship of  $\hat{r}_a$  against  $U^*$  (Fig. 8-9). In particular, in the case of the hybrid multi-tasking model, this result is promising because the relationship emerges from the observational data alone, without inducing any prespecified knowledge. Furthermore, the two constrained hybrid models show variations of the  $\hat{r}_a$  relationship across the sites (Fig. 8f, g and Fig. 9f, g). Thus, they are capable of capturing the canopy heterogeneity across sites and are more flexible than the conventional rigid parameterizations shown in Fig 8h (forests) and Fig. 9h (grasslands), where  $r_a$  is a homogenous function of  $U^*$  across the different sites.

The under-constrained hybrid model (Fig. 8e), however, illustrates the risk of equifinality and physics-violating results in this approach. In other words,  $\hat{r}_a$  exhibits physically inconsistent relationships in the under-constrained model across the sites (Fig. 8e), while the predicted  $\hat{r}_s$  and  $\hat{Q}_{LE}$  retain physically plausible estimates (Fig. 8a and Fig. 7 g-i, respectively). The issue of equifinality is more prominent in forests than in grasslands, likely because aerodynamic resistance is less dominant in controlling  $Q_{LE}$  in forests (Fig 8e and 9e; Chen & Liu, 2020).

The aerodynamic resistance  $r_a$  constitutes a critical link in the surface energy balance especially under different environmental and stability conditions, as it has a bearing on both,  $Q_{LE}$  and  $Q_H$ . There uncertainties in  $Q_{LE}$  and  $Q_H$  mainly arise from the uncertainty in estimating in  $r_a$  for both dense and sparse canopy, and particularly for arid and semi-arid conditions (Trebs et al., 2021). Our multi-task learning hybrid model, however, is able to provide a fairly high accuracy for  $Q_{LE}$  and  $Q_H$  predictions for grasslands under unstable and semi-arid conditions without overestimating  $r_a$ , which has been

proven difficult in other modeling efforts (Trebs et al., 2021). For example, the predictions for  $Q_{LE}$  (Fig. 5) and  $Q_H$  (Fig. 6c, d) at the US-Var grassland site, characterized by a dry Mediterranean-type climate (Xu & Baldocchi, 2004; De Kauwe et al., 2017), are fairly accurate and relate to physically consistent  $r_a$  predictions.

To get an estimate of the structural (epistemic) uncertainty for the inferred relationships for  $r_s$  and  $r_a$ , we train each model five times with random initializations (refer to Sec. 2.3). The hybrid models show consistent predictions for the relationships for  $r_s$  and  $r_a$  at mean diurnal scale across the model runs with different initializations. The under-constrained hybrid model is consistent in producing physically uninterpretable  $r_a$  for all initializations, especially for forests while the constrained hybrid models are able to reproduce consistently the physically plausible relationships for  $r_s$  and  $r_a$ . Hence, our hybrid modeling approach yields robust predictions, yet, we stress the caveats related to equifinality in these under-constrained model setups.

Lastly, we compare the behavior of surface conductance ( $g_s$ ) against  $Q_{LE_{obs}}$  with varying VPD at the mean diurnal scale for the multi-task learning model, the most promising approach, and the conventionally analyzed inverted PM equation for selected sites (Fig. 10). Both agree on a quasi-linear relationship between  $g_s$  and  $Q_{LE_{obs}}$  with a gradient in  $g_s$  (y direction) with changing VPD. So, as VPD increases, the  $g_s$  decreases for the same level of evapotranspiration. This is consistent with the findings of Monteith (1995) whereby model estimates reflect the surface feedback response where a decrease in  $g_s$  as VPD increases is a result of a direct increase in transpiration lowering leaf water potential (Streck, 2003; Mallick et al., 2013, 2016). The general behavior of  $g_s$  is similar between the multi-task learning (Fig. 10b, d) model and the PM Inv model (Fig. 10a, c), however, the estimation of  $g_s$  alongside changing  $Q_{LE_{obs}}$  in the multi-task learning model is less sensitive to noise at low  $Q_{LE_{obs}}$  compared to the PM Inv. Overall,  $g_s$  based on the inverted PM equation is considerably higher than based on the hybrid modeling approach. The higher estimation could constitute a systematic bias in  $g_s$  rooted in the inversion of PM. In particular, for dense canopies, the overestimation could be related to the non-linear relationship of the stomata to light, as is the case for the DE-Tha forest (Fig. 10a) (Campbell, G. S., & Norman, 1998; Irmak, S. et al., 2008). In grasslands, like DE-Gri (Fig. 10c), the overestimation could be attributed to the propagation of measurement error in deriving the energy balance (Wohlfahrt et al., 2009; Knauer et al., 2018). In summary, the multi-task learning model not only provides more confined but also lower estimates for  $g_s$  in contrast to widely used inversion method.

## 4 Conclusions

We present a new approach for an end-to-end hybrid modeling of latent heat fluxes that can simultaneously retrieve the two controlling intermediate variables — the surface ( $r_s$ ) and aerodynamic resistance ( $r_a$ ) — while maintaining physical consistency across different vegetation types. The hybrid models provide reliable predictions against measurements of latent heat fluxes at different time scales, ranging from daily to seasonal to interannual variations. This cross-scale consistency shows that our model framework is able to learn the physically consistent dependencies between the meteorological input variables and the target fluxes, rather than just the dominant structure of diurnal and seasonal cycles.

The main novelty and outcome of this study are data-driven parameterizations for  $r_s$  and  $r_a$  jointly estimated by two separate neural networks, which can lead new insights on biophysical regulation of surface evaporation. We show that the neural networks together can provide many solutions (non-uniqueness) and lead to physically plausible predictions for  $Q_{LE}$  fluxes, while presenting physically implausible relationships to the predictors. This non-uniqueness can be mitigated by introducing either more data or theory into the loss function of the hybrid model. Specifically, we make use of two different approaches (*a priori* constraint and multi-task learning) to regularize the parameter space for the neural networks. The resulting relationships for  $r_s$  and  $r_a$  not only show physically consistent behavior across scales, but also reveal new insights into how the varying resistances control surface energy fluxes.

In the determination of  $r_a$ , we find considerable variation between sites compared to the very uniform empirical formulations conventionally used. This inter-site spread in the observation-based parameterizations suggests that the conventional empirical formulations are too rigid and do not account for the variability caused by the vegetation canopy structure. Also, in the determination of  $r_s$ , the parameterizations derived from hybrid modeling show differences between sites, highlighting in particular the different physiological functions of the different plant types. In addition, we detect that these learned parameterizations in the hybrid models exhibit lower stomatal conductance, suggesting that the  $r_s$  values usually obtained by inversion of the Penman-Monteith equation may be systematically overestimated.

Several approaches have already been proposed to use the growing number of observations to constrain uncertainty in mechanistic model simulations, especially for key unknown plant behavior in the coupled Earth system (Lian et al., 2018; Winkler et al., 2019a,b; Varney et al., 2020). As a next step, we propose to derive parameterizations directly from observations using hybrid modeling, as presented in this study, to replace these *ad hoc* formulations in Earth system models. This approach will not only help reduce



uncertainty, but also advance significantly the understanding of biogeochemical processes in land-atmosphere coupling.

#### **Code and data availability**

All data used in this study are available from public databases or the literature, which can be found with the references provided in the respective “Data and methods” subsection. Processed data and analysis scripts are available from the corresponding author upon request, and the repository will be published together with this article. Correspondence and requests for materials should be addressed to Reda ElGhawi (relghawi@bgc-jena.mpg.de).

#### **Acknowledgements**

This research was funded by the European Research Council (ERC) Synergy Grant “Understanding and modeling the Earth System with Machine Learning (USMILE)” under the Horizon 2020 research and innovation programme (Grant agreement No. 855187). Gentine acknowledges funding from the National Science Foundation grant, Learning the Earth with Artificial intelligence and Physics (LEAP).

#### **Author contributions**

R.E.G., A.J.W. and M.R. designed the study. R.E.G. conducted the analysis. B.K. provided technical support in setting up the hybrid modelling framework. C.R. and M.K. contributed to the conceptual and technical machine learning aspect of the study. All authors contributed ideas and to the interpretation of the results. R.E.G. and A.J.W. drafted the manuscript with inputs from all authors.

#### **References**

- Ajami, H. (2021). Geohydrology: Global Hydrological Cycle. In *Encyclopedia of Geology* (pp. 393–398). Elsevier. <https://doi.org/10.1016/b978-0-12-409548-9.12387-5>
- Baldocchi, D., Falge, E., Gu, L., Olson, R., Hollinger, D., Running, S., Anthoni, P., Bernhofer, C., Davis, K., Evans, R., Fuentes, J., Goldstein, A., Katul, G., Law, B., Lee, X., Malhi, Y., Meyers, T., Munger, W.,

435 Oechel, W., ... Wofsy, S. (2001). FLUXNET: A New Tool to Study the Temporal and Spatial Variability  
436 of Ecosystem-Scale Carbon Dioxide, Water Vapor, and Energy Flux Densities. *Bulletin of the American*  
437 *Meteorological Society*, 82(11), 2415–2434. [https://doi.org/10.1175/1520-](https://doi.org/10.1175/1520-0477(2001)082<2415:FANTTS>2.3.CO;2)  
438 0477(2001)082<2415:FANTTS>2.3.CO;2

439 Besnard, S., Carvalhais, N., Arain, M. A., Black, A., Brede, B., Buchmann, N., Chen, J., Clevers, J. G. P. W.,  
440 Dutrieux, L. P., Gans, F., Herold, M., Jung, M., Kosugi, Y., Knohl, A., Law, B. E., Paul-Limoges, E.,  
441 Lohila, A., Merbold, L., Rouspard, O., ... Reichstein, M. (2019). Memory effects of climate and  
442 vegetation affecting net ecosystem CO<sub>2</sub> fluxes in global forests. *PLOS ONE*, 14(2), e0211510.  
443 <https://doi.org/10.1371/journal.pone.0211510>

444 Beven, K. (2006). A manifesto for the equifinality thesis. *Journal of Hydrology*, 320(1–2), 18–36.  
445 <https://doi.org/10.1016/J.JHYDROL.2005.07.007>

446 Bonan, G. B., Lawrence, P. J., Oleson, K. W., Levis, S., Jung, M., Reichstein, M., Lawrence, D. M., &  
447 Swenson, S. C. (2011). Improving canopy processes in the Community Land Model version 4 (CLM4)  
448 using global flux fields empirically inferred from FLUXNET data *Journal of Geophysical Research:*  
449 *Biogeosciences* 116, no. G2 (2011). *Journal of Geophysical Research: Biogeosciences*, 116(G2).  
450 <https://doi.org/10.1029/2010JG001593>

451 Brutsaert, W. (2005). *Hydrology: an introduction*. Cambridge University Press.

452 Buckley, T. N. (2005). The control of stomata by water balance. *New Phytologist*, 168(2), 275–292.  
453 <https://doi.org/10.1111/J.1469-8137.2005.01543.X>

454 Campbell, G. S., & Norman, J. M. (1998). *An Introduction to Environmental Biophysics*. Springer-Verlag.

455 Carminati, A., & Javaux, M. (2020). Soil Rather Than Xylem Vulnerability Controls Stomatal Response to  
456 Drought. *Trends in Plant Science*, 25(9), 868–880. <https://doi.org/10.1016/J.TPLANTS.2020.04.003>

457 Carter, C., & Liang, S. (2019). Evaluation of ten machine learning methods for estimating terrestrial  
458 evapotranspiration from remote sensing. *International Journal of Applied Earth Observation and*  
459 *Geoinformation*, 78, 86–92. <https://doi.org/10.1016/J.JAG.2019.01.020>

460 Chaves, M. M., Costa, J. M., Zarrouk, O., Pinheiro, C., Lopes, C. M., & Pereira, J. S. (2016). Controlling  
461 stomatal aperture in semi-arid regions—The dilemma of saving water or being cool? *Plant Science*, 251,  
462 54–64. <https://doi.org/10.1016/J.PLANTSCI.2016.06.015>

463 Chehbouni, A., Lo Seen, D., Njoku, E. G., & Monteny, B. M. (1996). Examination of the difference between  
464 radiative and aerodynamic surface temperatures over sparsely vegetated surfaces. *Remote Sensing of*  
465 *Environment*, 58(2), 177–186. [https://doi.org/10.1016/S0034-4257\(96\)00037-5](https://doi.org/10.1016/S0034-4257(96)00037-5)

466 Chen, J. M., & Liu, J. (2020). Evolution of evapotranspiration models using thermal and shortwave remote  
467 sensing data. *Remote Sensing of Environment*, 237, 111594. <https://doi.org/10.1016/J.RSE.2019.111594>

468 Chen, Y., Xia, J., Liang, S., Feng, J., Fisher, J. B., Li, X., Li, X., Liu, S., Ma, Z., Miyata, A., Mu, Q., Sun, L.,  
469 Tang, J., Wang, K., Wen, J., Xue, Y., Yu, G., Zha, T., Zhang, L., ... Yuan, W. (2014). Comparison of  
470 satellite-based evapotranspiration models over terrestrial ecosystems in China. *Remote Sensing of*  
471 *Environment*, 140, 279–293. <https://doi.org/10.1016/j.rse.2013.08.045>

472 Damour, G., Simonneau, T., Cochard, H., & Urban, L. (2010). An overview of models of stomatal conductance  
473 at the leaf level. In *Plant, Cell and Environment* (Vol. 33, Issue 9, pp. 1419–1438). Plant Cell Environ.  
474 <https://doi.org/10.1111/j.1365-3040.2010.02181.x>

475 de Bezenac, E., Pajot, A., & Gallinari, P. (2017). Deep Learning for Physical Processes: Incorporating Prior  
476 Scientific Knowledge. *ArXiv*. <http://arxiv.org/abs/1711.07970>

477 De Kauwe, M. G., Medlyn, B. E., Knauer, J., & Williams, C. A. (2017). Ideas and perspectives: How coupled is  
478 the vegetation to the boundary layer? *Biogeosciences*, 14(19), 4435–4453. [https://doi.org/10.5194/BG-14-](https://doi.org/10.5194/BG-14-4435-2017)  
479 4435-2017

480 Dou, X., & Yang, Y. (2018). Evapotranspiration estimation using four different machine learning approaches in  
481 different terrestrial ecosystems. *Computers and Electronics in Agriculture*, 148, 95–106.  
482 <https://doi.org/10.1016/j.compag.2018.03.010>

483 Drake, J. E., Tjoelker, M. G., Vårhammar, A., Medlyn, B. E., Reich, P. B., Leigh, A., Pfautsch, S., Blackman,  
484 C. J., López, R., Aspinwall, M. J., Crous, K. Y., Duursma, R. A., Kumarathunge, D., De Kauwe, M. G.,  
485 Jiang, M., Nicotra, A. B., Tissue, D. T., Choat, B., Atkin, O. K., & Barton, C. V. M. (2018). Trees tolerate  
486 an extreme heatwave via sustained transpirational cooling and increased leaf thermal tolerance. *Global*  
487 *Change Biology*, 24(6), 2390–2402. <https://doi.org/10.1111/GCB.14037>

488 Farquhar, G. D. (1978). Feedforward Responses of Stomata to Humidity. *Functional Plant Biology*, 5(6), 787–  
489 800. <https://doi.org/10.1071/PP9780787>

490 Friedl, M. A. (1996). Relationships among remotely sensed data, surface energy balance, and area-averaged  
491 fluxes over partially vegetated land surfaces. *Journal of Applied Meteorology and Climatology*, 35(11),  
492 2091–2103. [https://doi.org/10.1175/1520-0450\(1996\)035<2091:RARSDDS>2.0.CO;2](https://doi.org/10.1175/1520-0450(1996)035<2091:RARSDDS>2.0.CO;2)

493 Garratt, J. R. (1992). *The Atmospheric Boundary Layer*. Cambridge University Press.

494 Gerosa, G., Mereu, S., Finco, A., & Marzuoli, R. (2012). Stomatal Conductance Modeling to Estimate the  
495 Evapotranspiration of Natural and Agricultural Ecosystems. In *Evapotranspiration - Remote Sensing and*  
496 *Modeling*. InTech. <https://doi.org/10.5772/21825>

497 Halladay, K., & Good, P. (2017). Non-linear interactions between CO<sub>2</sub> radiative and physiological effects on  
498 Amazonian evapotranspiration in an Earth system model. *Climate Dynamics*, 49(7–8), 2471–2490.  
499 <https://doi.org/10.1007/s00382-016-3449-0>

500 Hetherington, A. M., & Woodward, F. I. (2003). The role of stomata in sensing and driving environmental  
501 change. *Nature*, 424(6951), 901–908. <https://doi.org/10.1038/nature01843>

502 Irmak, S., M., Irmak, A., Arkebauer, T. J., Weiss, A., Martin, D. L., & Eisenhauer, D. E. (2008). On the scaling  
503 up leaf stomatal resistance to canopy resistance using photosynthetic photon flux density. *Agricultural and*  
504 *Forest Meteorology*, 148(6–7), 1034–1044.

505 Jain, S. K., Nayak, P. C., & Sudheer, K. P. (2008). Models for estimating evapotranspiration using artificial  
506 neural networks, and their physical interpretation. *Hydrological Processes*, 22(13), 2225–2234.  
507 <https://doi.org/10.1002/hyp.6819>

- 508 Jarvis, P. G. (1976). The interpretation of the variations in leaf water potential and stomatal conductance found  
509 in canopies in the field. *Philosophical Transactions of the Royal Society of London. B, Biological*  
510 *Sciences*, 273(927), 593–610. <https://doi.org/10.1098/RSTB.1976.0035>
- 511 Jarvis, P. G., & Stewart, J. B. (1979). Evaporation of water from plantation forest. *D. Ford, D.C. Malcolm, J.*  
512 *Attersson (Eds.), The Ecology of Even-Aged Forest Plantations, Institute of Terrestrial Ecology,*  
513 *Cambridge*, 327–350.
- 514 Jia, X., Zwart, J., Sadler, J., Appling, A., Oliver, S., Markstrom, S., Willard, J., Xu, S., Steinbach, M., Read, J.,  
515 & Kumar, V. (2020). Physics-Guided Recurrent Graph Model for Predicting Flow and Temperature in  
516 River Networks. *ArXiv Preprint*.
- 517 Jung, M., Reichstein, M., & Bondeau, A. (2009). Towards global empirical upscaling of FLUXNET eddy  
518 covariance observations: Validation of a model tree ensemble approach using a biosphere model.  
519 *Biogeosciences*, 6(10), 2001–2013. <https://doi.org/10.5194/BG-6-2001-2009>
- 520 Jung, M., Reichstein, M., Ciais, P., Seneviratne, S. I., Sheffield, J., Goulden, M. L., Bonan, G., Cescatti, A.,  
521 Chen, J., De Jeu, R., Dolman, A. J., Eugster, W., Gerten, D., Gianelle, D., Gobron, N., Heinke, J.,  
522 Kimball, J., Law, B. E., Montagnani, L., ... Zhang, K. (2010). Recent decline in the global land  
523 evapotranspiration trend due to limited moisture supply. *Nature*, 467(7318), 951–954.  
524 <https://doi.org/10.1038/nature09396>
- 525 Jung, M., Schwalm, C., Migliavacca, M., Walther, S., Camps-Valls, G., Koirala, S., Anthoni, P., Besnard, S.,  
526 Bodesheim, P., Carvalhais, N., Chevallier, F., Gans, F., S Goll, D., Haverd, V., Köhler, P., Ichii, K., K  
527 Jain, A., Liu, J., Lombardozzi, D., ... Reichstein, M. (2020). Scaling carbon fluxes from eddy covariance  
528 sites to globe: Synthesis and evaluation of the FLUXCOM approach. *Biogeosciences*, 17(5), 1343–1365.  
529 <https://doi.org/10.5194/bg-17-1343-2020>
- 530 Karpatne, A., Atluri, G., Faghmous, J. H., Steinbach, M., Banerjee, A., Ganguly, A., Shekhar, S., Samatova, N.,  
531 & Kumar, V. (2017). Theory-guided data science: A new paradigm for scientific discovery from data.  
532 *IEEE Transactions on Knowledge and Data Engineering*, 29(10), 2318–2.
- 533 Karpatne, A., Watkins, W., Read, J., & Kumar, V. (2017). Physics-guided Neural Networks (PGNN): An  
534 Application in Lake Temperature Modeling. *ArXiv*. <http://arxiv.org/abs/1710.11431>
- 535 Kelliher, F. M., Köstner, B. M. M., Hollinger, D. Y., Byers, J. N., Hunt, J. E., McSeveny, T. M., Meserth, R.,  
536 Weir, P. L., & Schulze, E. D. (1992). Evaporation, xylem sap flow, and tree transpiration in a New  
537 Zealand broad-leaved forest. *Agricultural and Forest Meteorology*, 62(1–2), 53–73.  
538 [https://doi.org/10.1016/0168-1923\(92\)90005-O](https://doi.org/10.1016/0168-1923(92)90005-O)
- 539 Kennedy, D., Swenson, S., Oleson, K. W., Lawrence, D. M., Fisher, R., Lola da Costa, A. C., & Gentine, P.  
540 (2019). Implementing Plant Hydraulics in the Community Land Model, Version 5. *Journal of Advances in*  
541 *Modeling Earth Systems*, 11(2), 485–513. <https://doi.org/10.1029/2018MS001500>
- 542 Knauer, J., El-Madany, T. S., Zaehle, S., & Migliavacca, M. (2018). Bigleaf—An R package for the calculation  
543 of physical and physiological ecosystem properties from eddy covariance data. *PLOS ONE*, 13(8),  
544 e0201114. <https://doi.org/10.1371/JOURNAL.PONE.0201114>
- 545 Köstner, B. M. M., Schulze, E. D., Kelliher, F. M., Hollinger, D. Y., Byers, J. N., Hunt, J. E., McSeveny, T. M.,

546 Meserth, R., & Weir, P. L. (1992). Transpiration and canopy conductance in a pristine broad-leaved forest  
547 of Nothofagus: an analysis of xylem sap flow and eddy correlation measurements. *Oecologia*, 91(3), 350–  
548 359. <https://doi.org/10.1007/BF00317623>

549 Kraft, B., Jung, M., Körner, M., Koirala, S., & Reichstein, M. (2022). Towards hybrid modeling of the global  
550 hydrological cycle. *Hydrology and Earth System Sciences*, 26(6), 1579–1614.  
551 <https://doi.org/10.5194/HESS-26-1579-2022>

552 Kraft, B., Jung, M., Körner, M., & Reichstein, M. (2020). Hybrid modeling: Fusion of a deep approach and  
553 physics-based model for global hydrological modeling. *The International Archives of Photogrammetry,*  
554 *Remote Sensing and Spatial Information Sciences*, 43, 1537–1544. [https://doi.org/10.5194/isprs-archives-](https://doi.org/10.5194/isprs-archives-XLIII-B2-2020-1537-2020)  
555 [XLIII-B2-2020-1537-2020](https://doi.org/10.5194/isprs-archives-XLIII-B2-2020-1537-2020)

556 Krasnopolsky, V. M. (2013). *The application of neural networks in the Earth system sciences*.

557 Leuning, R. (1995). A critical appraisal of a combined stomatal-photosynthesis model for C3 plants. *Plant, Cell*  
558 *& Environment*, 18(4), 339–355. <https://doi.org/10.1111/J.1365-3040.1995.TB00370.X>

559 Leuning, R., Kriedemann, P. E., & McMurtrie, R. E. (1991). Simulation of evapotranspiration by trees.  
560 *Agricultural Water Management*, 19(3), 205–221. [https://doi.org/10.1016/0378-3774\(91\)90042-H](https://doi.org/10.1016/0378-3774(91)90042-H)

561 Li, L., Wang, Y., Arora, V. K., Eamus, D., Shi, H., Li, J., Cheng, L., Cleverly, J., Hajima, T., Ji, D., Jones, C.,  
562 Kawamiya, T., Li, W., Tjiputra, J., Wiltshire, A., Zhang, L., & Yu, Q. (2018). Evaluating global land  
563 surface models in CMIP5: Analysis of ecosystem water- and light-use efficiencies and rainfall  
564 partitioning. *Journal of Climate*, 31(8), 2995–3008. <https://doi.org/10.1175/JCLI-D-16-0177.1>

565 Li, X., Gentine, P., Lin, C., Zhou, S., Sun, Z., Zheng, Y., Liu, J., & Zheng, C. (2019). A simple and objective  
566 method to partition evapotranspiration into transpiration and evaporation at eddy-covariance sites.  
567 *Agricultural and Forest Meteorology*, 265, 171–182.  
568 <https://doi.org/10.1016/J.AGRFORMET.2018.11.017>

569 Lian, X., Piao, S., Huntingford, C., Li, Y., Zeng, Z., Wang, X., Ciais, P., McVicar, T. R., Peng, S., Ottlé, C.,  
570 Yang, H., Yang, Y., Zhang, Y., & Wang, T. (2018). Partitioning global land evapotranspiration using  
571 CMIP5 models constrained by observations. *Nature Climate Change* 2018 8:7, 8(7), 640–646.  
572 <https://doi.org/10.1038/s41558-018-0207-9>

573 Liebel, L., & Körner, M. (2018). *Auxiliary Tasks in Multi-task Learning*.  
574 <https://doi.org/10.48550/arxiv.1805.06334>

575 Lin, C., Gentine, P., Huang, Y., Guan, K., Kimm, H., & Zhou, S. (2018). Diel ecosystem conductance response  
576 to vapor pressure deficit is suboptimal and independent of soil moisture. *Agricultural and Forest*  
577 *Meteorology*, 250–251, 24–34. <https://doi.org/10.1016/j.agrformet.2017.12.078>

578 Lin, H., Chen, Y., Zhang, H., Fu, P., & Fan, Z. (2017). Stronger cooling effects of transpiration and leaf  
579 physical traits of plants from a hot dry habitat than from a hot wet habitat. *Functional Ecology*, 31(12),  
580 2202–2211. <https://doi.org/10.1111/1365-2435.12923/SUPPINFO>

581 Liu, S., Mao, D., & Lu, L. (2006). Measurement and estimation of the aerodynamic resistance. *European*  
582 *Geosciences Union*, 3(3), 681–705. [www.hydrol-earth-syst-sci-discuss.net/3/681/2006/](http://www.hydrol-earth-syst-sci-discuss.net/3/681/2006/)

583 Mallick, K., Jarvis, A., Fisher, J. B., Tu, K. P., Boegh, E., & Niyogi, D. (2013). Latent Heat Flux and Canopy  
584 Conductance Based on Penman–Monteith, Priestley–Taylor Equation, and Bouchet’s Complementary  
585 Hypothesis. *Journal of Hydrometeorology*, 14(2), 419–442. <https://doi.org/10.1175/JHM-D-12-0117.1>

586 Mallick, K., Trebs, I., Boegh, E., Giustarini, L., Schlerf, M., Drewry, D. T., Hoffmann, L., Von Randow, C.,  
587 Kruijt, B., Araújo, A., Saleska, S., Ehleringer, J. R., Domingues, T. F., Ometto, J. P. H. B., Nobre, A. D.,  
588 Luiz Leal De Moraes, O., Hayek, M., William Munger, J., & Wofsy, S. C. (2016). Canopy-scale  
589 biophysical controls of transpiration and evaporation in the Amazon Basin. *Hydrology and Earth System  
590 Sciences*, 20(10), 4237–4264. <https://doi.org/10.5194/HESS-20-4237-2016>

591 Massmann, A., Gentine, P., & Lin, C. (2019). When Does Vapor Pressure Deficit Drive or Reduce  
592 Evapotranspiration? *Journal of Advances in Modeling Earth Systems*, 11(10), 3305–3320.  
593 <https://doi.org/10.1029/2019MS001790>

594 Maurer, K. D., Bohrer, G., Kenny, W. T., & Ivanov, V. Y. (2015). Large-eddy simulations of surface roughness  
595 parameter sensitivity to canopy-structure characteristics. *Biogeosciences*, 12(8), 2533–2548.  
596 <https://doi.org/10.5194/BG-12-2533-2015>

597 Medlyn, B. E., Duursma, R. A., Eamus, D., Ellsworth, D. S., Prentice, I. C., Barton, C. V. M., Crous, K. Y., De  
598 Angelis, P., Freeman, M., & Wingate, L. (2011). Reconciling the optimal and empirical approaches to  
599 modelling stomatal conductance. *Global Change Biology*, 17(6), 2134–2144.  
600 <https://doi.org/10.1111/J.1365-2486.2010.02375.X>

601 Monteith, J. L. (1995). A reinterpretation of stomatal responses to humidity. *Plant, Cell & Environment*, 18(4),  
602 357–364. <https://doi.org/10.1111/J.1365-3040.1995.TB00371.X>

603 Monteith, J. L. (1965). Evaporation and environment. *Fogg (Ed.), Symposium of the Society for Experimental  
604 Biology, The State and Movement of Water in Living Organisms, Vol. 19*, 205–234.

605 Monteith, J. L. (1985). Evaporation from land surfaces: progress in analysis and prediction since 1948. In:  
606 *Advances in Evapotranspiration, Proc. National Conference on Advances in Evapotranspiration. Am.:*  
607 *Soc. Agric. Eng., St. Joseph, MI, Chicago, IL*, 4–12. [https://agris.fao.org/agris-](https://agris.fao.org/agris-search/search.do?recordID=US8644525)  
608 [search/search.do?recordID=US8644525](https://agris.fao.org/agris-search/search.do?recordID=US8644525)

609 Monteith, J. L., & Unsworth, M. (2013). *Principles of environmental physics: plants, animals, and the  
610 atmosphere* (Fourth). Elsevier Ltd.

611 Mott, K. A., & Peak, D. (2013). Testing a vapour-phase model of stomatal responses to humidity. *Plant, Cell &  
612 Environment*, 36(5), 936–944. <https://doi.org/10.1111/PCE.12026>

613 Mystakidis, S., Davin, E. L., Gruber, N., & Seneviratne, S. I. (2016). Constraining future terrestrial carbon cycle  
614 projections using observation-based water and carbon flux estimates. *Global Change Biology*, 22(6),  
615 2198–2215. <https://doi.org/10.1111/gcb.13217>

616 Nakai, T., Sumida, A., Daikoku, K., Matsumoto, K., van der Molen, M. K., Kodama, Y., Kononov, A. V.,  
617 Maximov, T. C., Dolman, A. J., Yabuki, H., Hara, T., & Ohta, T. (2008). Parameterisation of aerodynamic  
618 roughness over boreal, cool- and warm-temperate forests. *Agricultural and Forest Meteorology*, 148(12),  
619 1916–1925. <https://doi.org/10.1016/J.AGRFORMET.2008.03.009>

620 Penman, H. L. (1948). Natural evaporation from open water, bare soil and grass. *Proceedings of the Royal*  
621 *Society of London. Series A, Mathematical and Physical Sciences*, 193(1032), 120–145.  
622 <https://doi.org/10.1098/rspa.1948.0037>

623 Polhamus, A., Fisher, J. B., & Tu, K. P. (2013). What controls the error structure in evapotranspiration models?  
624 *Agricultural and Forest Meteorology*, 169, 12–24. <https://doi.org/10.1016/j.agrformet.2012.10.002>

625 Reichstein, M., Ahrens, B., Kraft, B., Camps-Valls, G., Carvalhais, N., Gans, F., Gentine, P., & Winkler, A. J.  
626 (2022). Combining System Modeling and Machine Learning into Hybrid Ecosystem Modeling.  
627 *Knowledge-Guided Machine Learning*, 327–352. <https://doi.org/10.1201/9781003143376-14>

628 Reichstein, M., Camps-Valls, G., Stevens, B., Jung, M., Denzler, J., Carvalhais, N., & Prabhat. (2019). Deep  
629 learning and process understanding for data-driven Earth system science. *Nature*, 566(7743), 195–204.  
630 <https://doi.org/10.1038/s41586-019-0912-1>

631 Ronda, R. J., De Bruin, H. A. R., & Holtslag, A. A. M. (2001). Representation of the canopy conductance in  
632 modeling the surface energy budget for low vegetation. *Journal of Applied Meteorology*, 40(8), 1431–  
633 1444.

634 Schmidt, L., Heße, F., Attinger, S., & Kumar, R. (2020). Challenges in Applying Machine Learning Models for  
635 Hydrological Inference: A Case Study for Flooding Events Across Germany. *Water Resources Research*,  
636 56(5), e2019WR025924. <https://doi.org/10.1029/2019WR025924>

637 Schulze, E. D. (1986). Carbon dioxide and water vapor exchange in response to drought in the atmosphere and  
638 in the soil. *Annu. Rev. Plant Physiol.; (United States)*, 37(1), 247–274.  
639 <https://doi.org/10.1146/ANNUREV.PP.37.060186.001335>

640 Sellers, P. J., Dickinson, R. E., Randall, D. A., Betts, A. K., Hall, F. G., Berry, J. A., Collatz, G. J., Denning, A.  
641 S., Mooney, H. A., Nobre, C. A., Sato, N., Field, C. B., & Henderson-Sellers, A. (1997). Modeling the  
642 exchanges of energy, water, and carbon between continents and the atmosphere. In *Science* (Vol. 275,  
643 Issue 5299, pp. 502–509). American Association for the Advancement of Science.  
644 <https://doi.org/10.1126/science.275.5299.502>

645 Shaw, R. H., & Pereira, A. R. (1982). Aerodynamic roughness of a plant canopy: A numerical experiment.  
646 *Agricultural Meteorology*, 26(1), 51–65. [https://doi.org/10.1016/0002-1571\(82\)90057-7](https://doi.org/10.1016/0002-1571(82)90057-7)

647 Stewart, J. B. (1988). Modelling surface conductance of pine forest. *Agricultural and Forest Meteorology*,  
648 43(1), 19–35. [https://doi.org/10.1016/0168-1923\(88\)90003-2](https://doi.org/10.1016/0168-1923(88)90003-2)

649 Streck, N. A. (2003). Stomatal response to water vapor pressure deficit: an unsolved issue. *Current Agricultural*  
650 *Science and Technology*, 9(4), 317–322.

651 Su, Y., Zhang, C., Chen, X., Liu, L., Ciais, P., Peng, J., Wu, S., Wu, J., Shang, J., Wang, Y., Yuan, W., Yang,  
652 Y., Wu, Z., & Laforetza, R. (2021). Aerodynamic resistance and Bowen ratio explain the biophysical  
653 effects of forest cover on understory air and soil temperatures at the global scale. *Agricultural and Forest*  
654 *Meteorology*, 308–309, 108615. <https://doi.org/10.1016/J.AGRFORMET.2021.108615>

655 Tan, S., Wang, H., Prentice, I. C., & Yang, K. (2021). Land-surface evapotranspiration derived from a first-  
656 principles primary production model. *Environmental Research Letters*, 16(10), 104047.

657 <https://doi.org/10.1088/1748-9326/AC29EB>

658 Tramontana, G., Jung, M., Schwalm, C. R., Ichii, K., Camps-Valls, G., Ráduly, B., Reichstein, M., Arain, M.  
659 A., Cescatti, A., Kiely, G., Merbold, L., Serrano-Ortiz, P., Sickert, S., Wolf, S., & Papale, D. (2016).  
660 Predicting carbon dioxide and energy fluxes across global FLUXNET sites with regression algorithms.  
661 *Biogeosciences*, 13(14), 4291–4313. <https://doi.org/10.5194/BG-13-4291-2016>

662 Trebs, I., Mallick, K., Bhattarai, N., Sulis, M., Cleverly, J., Woodgate, W., Silberstein, R., Hinko-Najera, N.,  
663 Beringer, J., Meyer, W. S., Su, Z., & Boulet, G. (2021). The role of aerodynamic resistance in thermal  
664 remote sensing-based evapotranspiration models. *Remote Sensing of Environment*, 264, 112602.  
665 <https://doi.org/10.1016/J.RSE.2021.112602>

666 Varney, R. M., Chadburn, S. E., Friedlingstein, P., Burke, E. J., Koven, C. D., Hugelius, G., & Cox, P. M.  
667 (2020). A spatial emergent constraint on the sensitivity of soil carbon turnover to global warming. *Nature*  
668 *Communications* 2020 11:1, 11(1), 1–8. <https://doi.org/10.1038/s41467-020-19208-8>

669 Vialet-Chabrand, S., & Lawson, T. (2019). Dynamic leaf energy balance: deriving stomatal conductance from  
670 thermal imaging in a dynamic environment. *Journal of Experimental Botany*, 70(10), 2839.  
671 <https://doi.org/10.1093/JXB/ERZ068>

672 Vico, G., Manzoni, S., Palmroth, S., Weih, M., & Katul, G. (2013). A perspective on optimal leaf stomatal  
673 conductance under CO<sub>2</sub> and light co-limitations. *Agricultural and Forest Meteorology*, 182–183, 191–  
674 199. <https://doi.org/10.1016/J.AGRFORMET.2013.07.005>

675 Vila-Guerau de Arellano, J., C. van Heerwaarden, C., J. H. van Stratum, B., & van den Dries, K. (2015).  
676 *Atmospheric Boundary Layer*. Cambridge University Press. <https://doi.org/10.1017/CBO9781316117422>

677 Wang, K., & Dickinson, R. E. (2012). A review of global terrestrial evapotranspiration: Observation, modeling,  
678 climatology, and climatic variability. *Reviews of Geophysics*, 50(2).  
679 <https://doi.org/10.1029/2011RG000373>

680 Wehr, R., & Saleska, S. R. (2021). Calculating canopy stomatal conductance from eddy covariance  
681 measurements, in light of the energy budget closure problem. *Biogeosciences*, 18(1), 13–24.  
682 <https://doi.org/10.5194/BG-18-13-2021>

683 Winkler, A. J., Myneni, R. B., Alexandrov, G. A., & Brovkin, V. (2019). Earth system models underestimate  
684 carbon fixation by plants in the high latitudes. *Nature Communications* 2019 10:1, 10(1), 1–8.  
685 <https://doi.org/10.1038/s41467-019-08633-z>

686 Winkler, A. J., Myneni, R. B., & Brovkin, V. (2019). Investigating the applicability of emergent constraints.  
687 *Earth System Dynamics*, 10(3), 501–523. <https://doi.org/10.5194/ESD-10-501-2019>

688 Wohlfahrt, G., Haslwanter, A., Hörtnagl, L., Jasoni, R. L., Fenstermaker, L. F., Arnone, J. A., & Hammerle, A.  
689 (2009). On the consequences of the energy imbalance for calculating surface conductance to water vapour.  
690 *Agricultural and Forest Meteorology*, 149(9), 1556–1559.  
691 <https://doi.org/10.1016/J.AGRFORMET.2009.03.015>

692 Xu, L., & Baldocchi, D. D. (2004). Seasonal variation in carbon dioxide exchange over a Mediterranean annual  
693 grassland in California. *Agricultural and Forest Meteorology*, 123(1–2), 79–96.



694 <https://doi.org/10.1016/J.AGRFORMET.2003.10.004>

695 Yang, T., Sun, F., Gentine, P., Liu, W., Wang, H., Yin, J., Du, M., & Liu, C. (2019). Evaluation and machine  
696 learning improvement of global hydrological model-based flood simulations. *Environmental Research*  
697 *Letters*, 14(11), 114027. <https://doi.org/10.1088/1748-9326/ab4d5e>

698 Zeppel, M., & Eamus, D. (2008). Coordination of leaf area, sapwood area and canopy conductance leads to  
699 species convergence of tree water use in a remnant evergreen woodland. *Australian Journal of Botany*,  
700 56(2), 97–108. <https://doi.org/10.1071/BT07091>

701 Zhang, Z. Z., Zhao, P., McCarthy, H. R., Zhao, X. H., Niu, J. F., Zhu, L. W., Ni, G. Y., Ouyang, L., & Huang,  
702 Y. Q. (2016). Influence of the decoupling degree on the estimation of canopy stomatal conductance for  
703 two broadleaf tree species. *Agricultural and Forest Meteorology*, 221, 230–241.  
704 <https://doi.org/10.1016/J.AGRFORMET.2016.02.018>

705 Zhao, L., Xia, J., Xu, C. yu, Wang, Z., Sobkowiak, L., & Long, C. (2013). Evapotranspiration estimation  
706 methods in hydrological models. *Journal of Geographical Sciences* 2013 23:2, 23(2), 359–369.  
707 <https://doi.org/10.1007/S11442-013-1015-9>

708 Zhao, W. L., Gentine, P., Reichstein, M., Zhang, Y., Zhou, S., Wen, Y., Lin, C., Li, X., & Qiu, G. Y. (2019).  
709 Physics-Constrained Machine Learning of Evapotranspiration. *Geophysical Research Letters*, 46(24),  
710 14496–14507. <https://doi.org/10.1029/2019GL085291>

711 Zhou, S., Yu, B., Zhang, Y., Huang, Y., & Wang, G. (2016). Partitioning evapotranspiration based on the  
712 concept of underlying water use efficiency. *Water Resources Research*, 52(2), 1160–1175.  
713 <https://doi.org/10.1002/2015WR017766>

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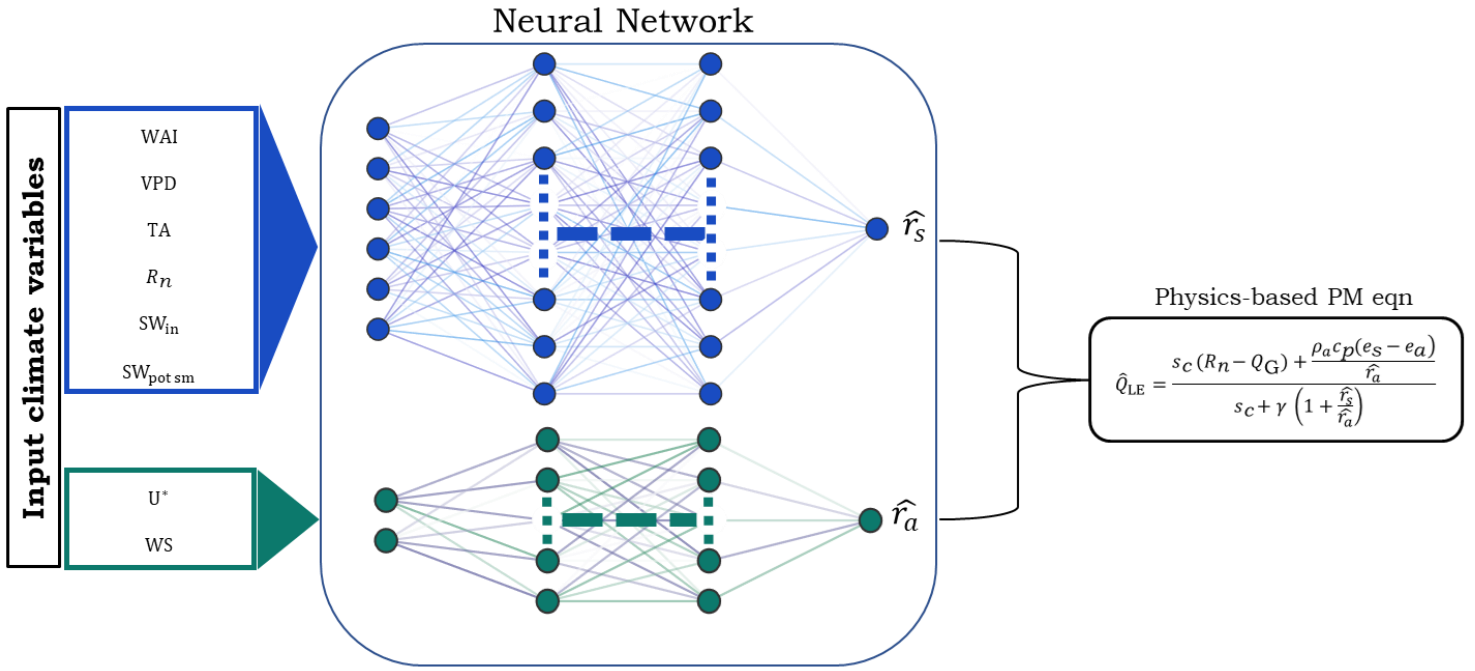
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726 Table 1: Detailed description of each site used derived from the FLUXNET 2015 Tier 1 data.

Site ID	IGBP	Elevation (m)	Mean Annual Temperature (°C)	Mean Annual Precipitation (mm)	Data Availability	DOI
DE-Tha	ENF <sup>1</sup>	385	8.2	843	19 years (1996 - 2014)	Christian Bernhofer, Thomas Grünwald, Uta Moderow, Markus Hehn, Uwe Eichelmann, Heiko Prasse, Udo Postel (1996-2014) FLUXNET2015 DE-Tha Tharandt, Dataset. <a href="https://doi.org/10.18140/FLX/1440152">https://doi.org/10.18140/FLX/1440152</a>
FR-Pue	EBF <sup>2</sup>	270	13.5	883	15 years (2000 - 2014)	Jean-Marc Ourcival, Karim Piquemal, Richard Joffre, Limousin Jean-Marc (2000-2014) FLUXNET2015 FR-Pue Puechabon, Dataset. <a href="https://doi.org/10.18140/FLX/1440164">https://doi.org/10.18140/FLX/1440164</a>
FR-LBr	ENF <sup>1</sup>	61	13.6	900	12 years (1996 - 2008)	Paul Berbigier, Jean Bonnefond, Alexandre Bosc, Pierre Trichet, Denis Loustau (1996-2008) FLUXNET2015 FR-LBr Le Bray, Dataset. <a href="https://doi.org/10.18140/FLX/1440163">https://doi.org/10.18140/FLX/1440163</a>
CH-Cha	GRA <sup>3</sup>	393	9.5	1136	10 years (2005 - 2014)	Lutz Merbold, Kathrin Fuchs, Nina Buchmann, Lukas Hörtnagl (2012-2016) FLUXNET-CH4 CH-Cha Chamau, Dataset. <a href="https://doi.org/10.18140/FLX/1669629">https://doi.org/10.18140/FLX/1669629</a>
DE-Gri	GRA <sup>3</sup>	385	7.8	901	11 years (2004 - 2014)	Christian Bernhofer, Thomas Grünwald, Uta Moderow, Markus Hehn, Uwe Eichelmann, Heiko Prasse, Udo Postel () FLUXNET2015 DE-Gri , Dataset. <a href="https://doi.org/10.18140/FLX/1440147">https://doi.org/10.18140/FLX/1440147</a>
US-Var	GRA <sup>3</sup>	129	15.8	559	15 years (2000 - 2014)	(2000-2014) FLUXNET2015 US-Var Vaira Ranch- Ione, Dataset. <a href="https://doi.org/10.18140/FLX/1440094">https://doi.org/10.18140/FLX/1440094</a>

- 727 1. ENF (Evergreen Needleleaf Forests: Lands dominated by woody vegetation with a percent cover >60% and height  
 728 exceeding 2 meters. Almost all trees remain green all year. Canopy is never without green foliage).  
 729 2. EBF (Evergreen Broadleaf Forests: Lands dominated by woody vegetation with a percent cover >60% and height  
 730 exceeding 2 meters. Almost all trees and shrubs remain green year-round. Canopy is never without green foliage).  
 731 3. GRA (Grasslands: Lands with herbaceous types of cover. Tree and shrub cover is less than 10%. Permanent  
 732 wetlands lands with a permanent mixture of water and herbaceous or woody vegetation. The vegetation can be  
 733 present in either salt, brackish, or fresh water.)

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Figure 1: Architecture of the basic hybrid model consists of two neural networks, which estimate  $\hat{r}_s$  and  $\hat{r}_a$  individually with independent input climate variables. The latent variables are used in the Penman-Monteith equation to estimate the latent heat flux ( $Q_{LE}$ ), and the objective function minimizes losses for  $Q_{LE}$ . WS is wind speed ( $\text{ms}^{-1}$ ), and  $U^*$  is friction velocity ( $\text{ms}^{-1}$ ).  $R_n$  is the net radiation ( $\text{Wm}^{-2}$ ), VPD, is the vapor pressure deficit of air (kPa), WAI is the water availability index calculated in Eq. 6, TA is air temperature ( $^{\circ}\text{C}$ ),  $SW_{in}$  is incoming shortwave radiation ( $\text{Wm}^{-2}$ ), and  $SW_{pot\ sm}$  is mean incoming shortwave potential ( $\text{Wm}^{-2}$ ).

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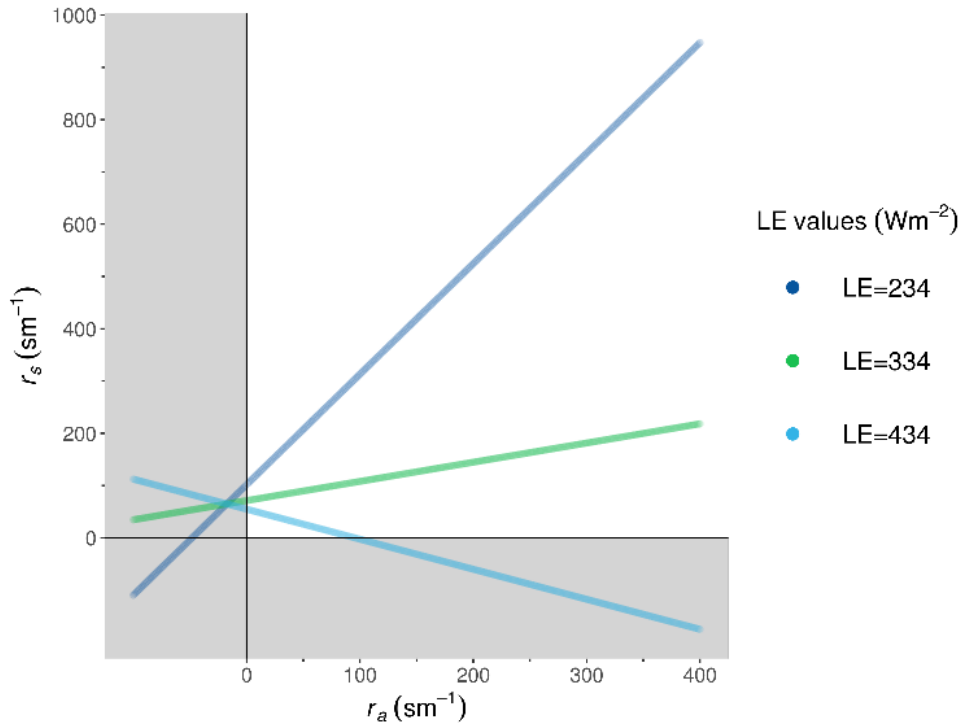
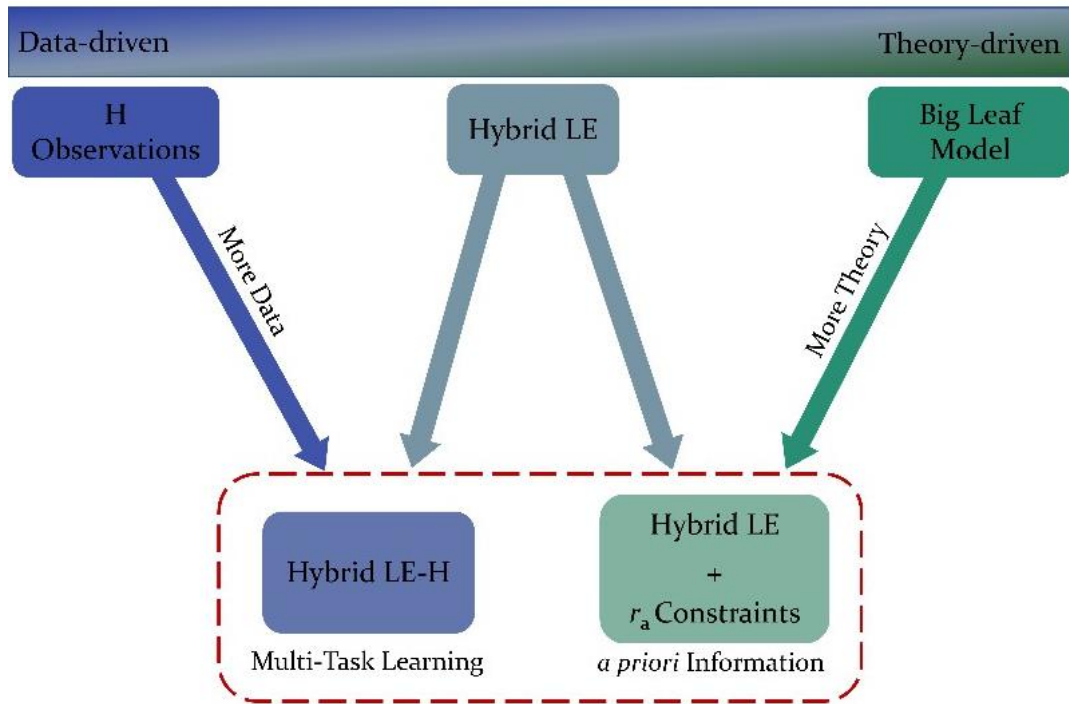


Figure 2: Equifinality in the physics-based component of hybrid model: The lines represent different  $Q_{LE}$  values that can exist for specific conditions (the actual  $Q_{LE}$  value is approximately  $334 \text{ Wm}^{-2}$ ). Fixing all parameters of the PM equation  $s_c = 0.175 \text{ kPaC}^{-1}$ ,  $R_n = 520.38 \text{ Wm}^{-2}$ ,  $Q_G = 18.51 \text{ Wm}^{-2}$ ,  $VPD = 1.333 \text{ kPa}$ ,  $\rho_a = 1.143 \text{ kg m}^{-3}$ ,  $c_p = 1004.834 \text{ J kg}^{-1} \text{ C}^{-1}$ ,  $\gamma = 0.0644 \text{ kPaC}^{-1}$ , the different combinations of  $r_s$  and  $r_a$  values lead to the same  $Q_{LE}$ . Shaded areas show the physically non-plausible and non-realistic values for  $r_s$  and  $r_a$  combinations, and non-shaded areas show physically plausible values.



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Figure 3: Proposed methods for constraining the hybrid model: Right-side shows the theory-driven hybrid model with *a priori* constraint for  $r_a$  from the Big Leaf model. Left-side shows data-driven hybrid model with more information from learning an additional target variable  $Q_H$  through multi-task learning.

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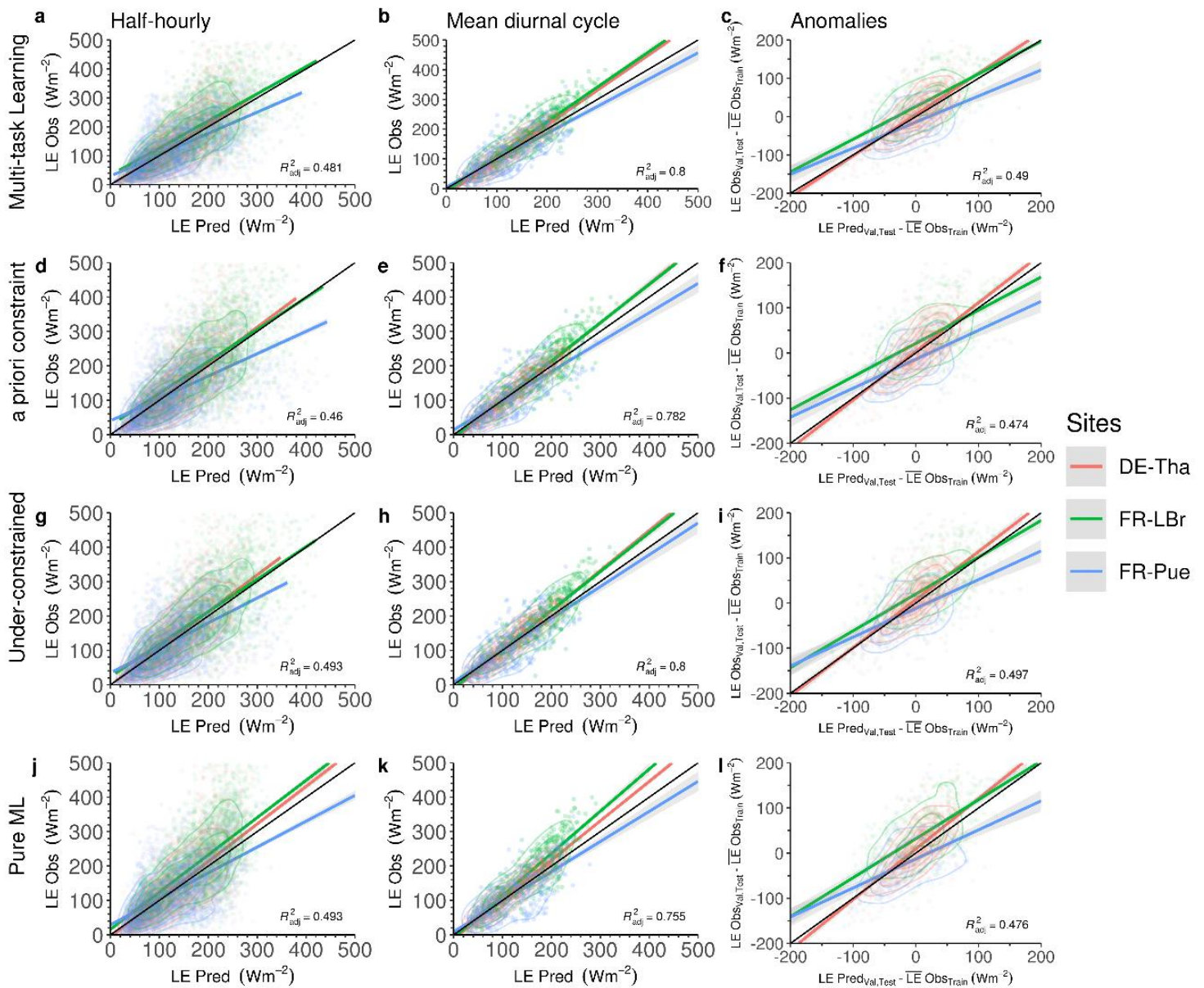


Figure 4: Evaluation of  $Q_{LE}$  observations and predictions at different temporal scales for forests. a,d,g,j show predictions against observations at a half-hourly scale for different models; b,e,h,k show predictions against observations at mean diurnal scale; c,f,i,l show  $Q_{LE}$  anomalies at interannual scale for the different models.



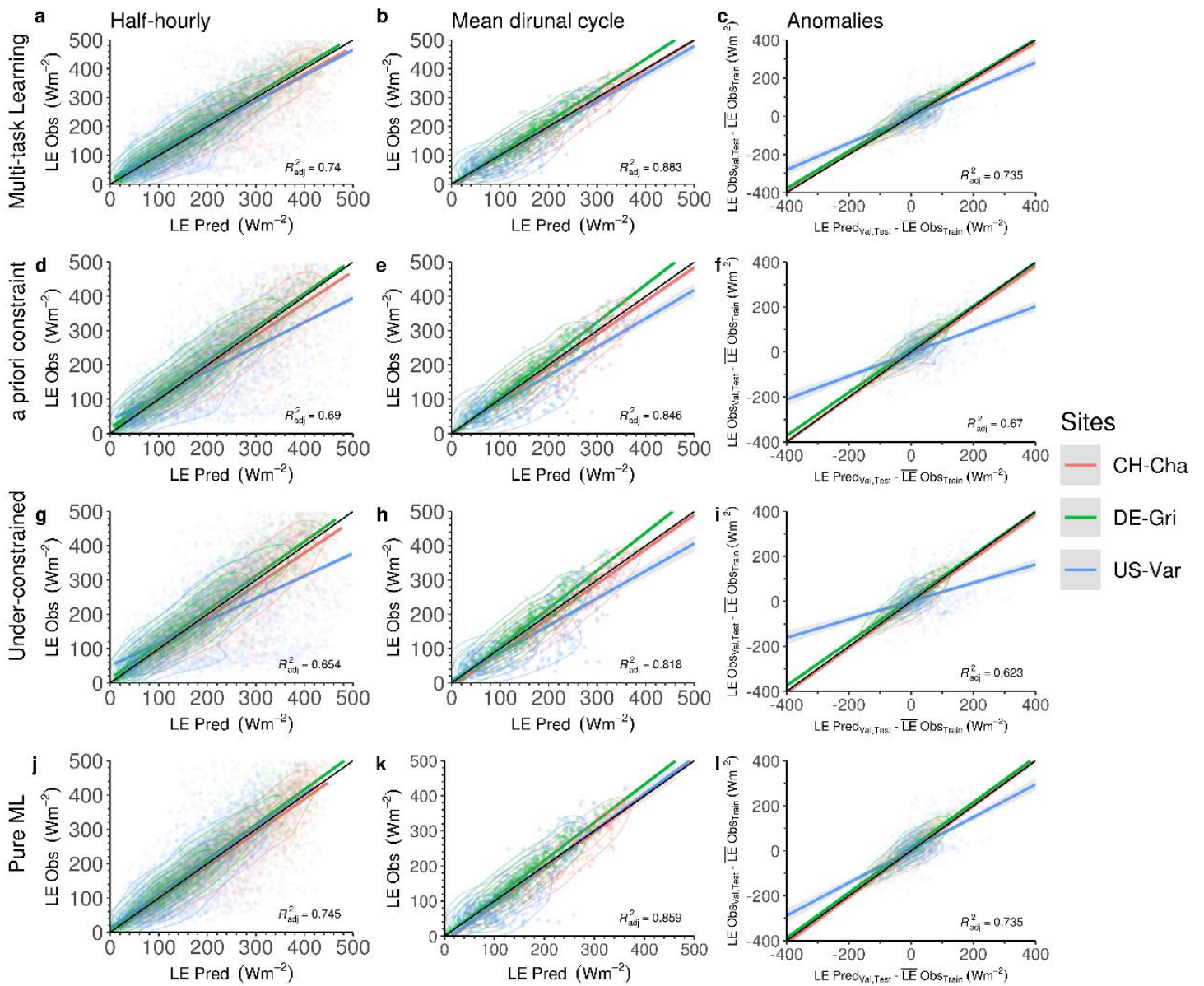
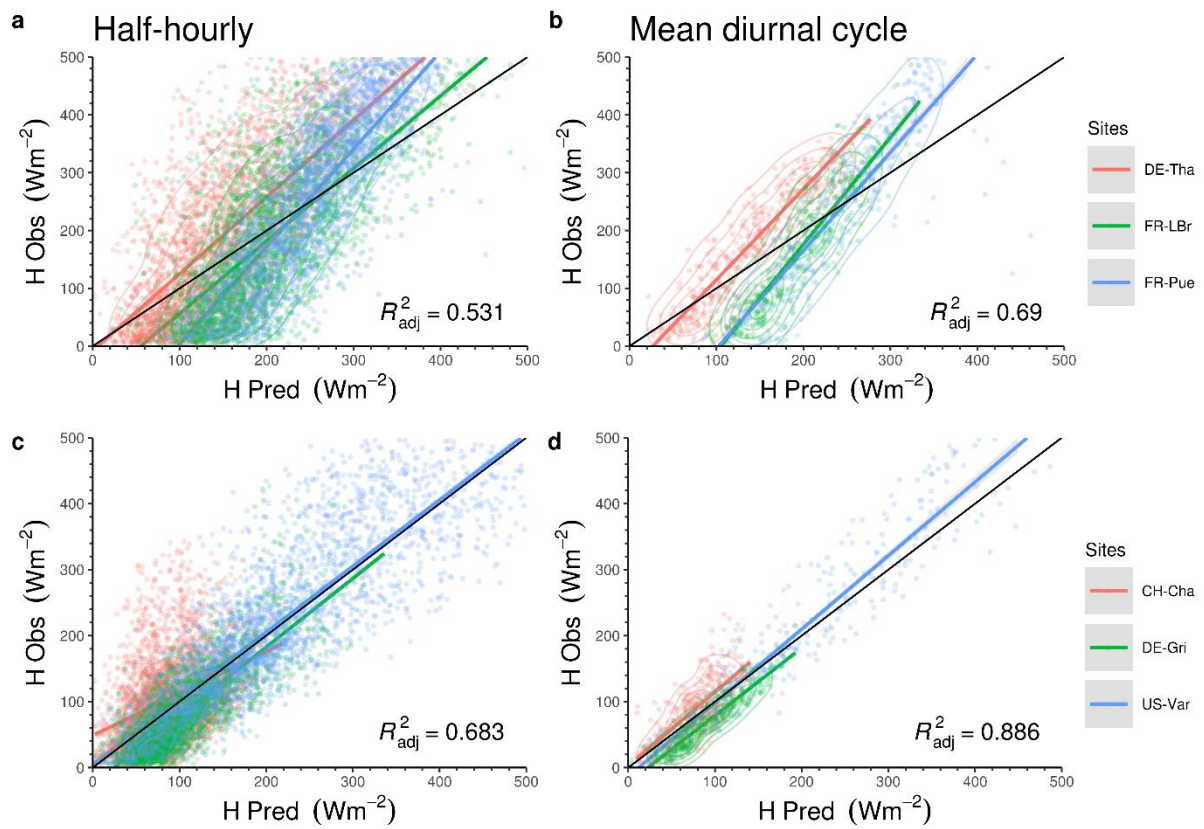


Figure 5: Evaluation of  $Q_{LE}$  observations and predictions at different temporal scales for grasslands. a,d,g,j show predictions against observations at a half-hourly scale for different models. b,e,h,k show predictions against observations at mean diurnal scale. c,f,i,l show  $Q_{LE}$  anomalies at interannual scale for the different models.



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Figure 6: Evaluation of  $Q_H$  observations and predictions at half-hourly, and mean diurnal scale for forest (a,b) and grasslands (c,d) for multi-task learning hybrid model.  $Q_H$  predictions are similar in range compare to  $Q_{LE}$  predictions in figures 4-5 for forests and grasslands.

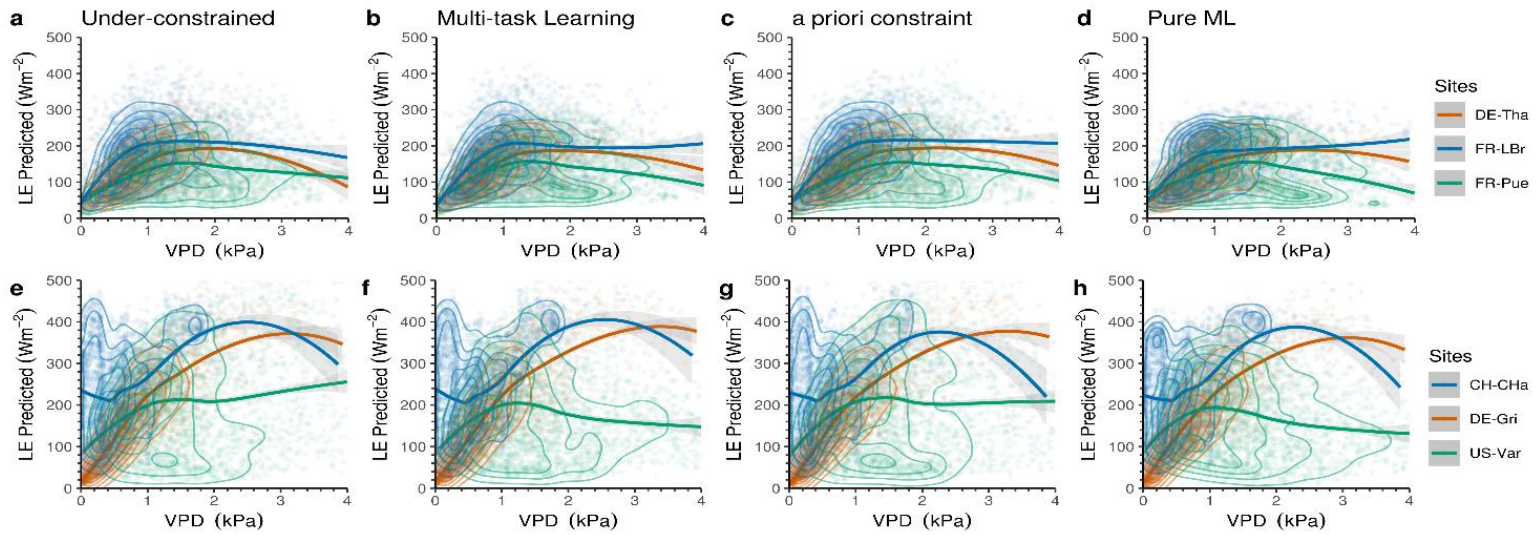
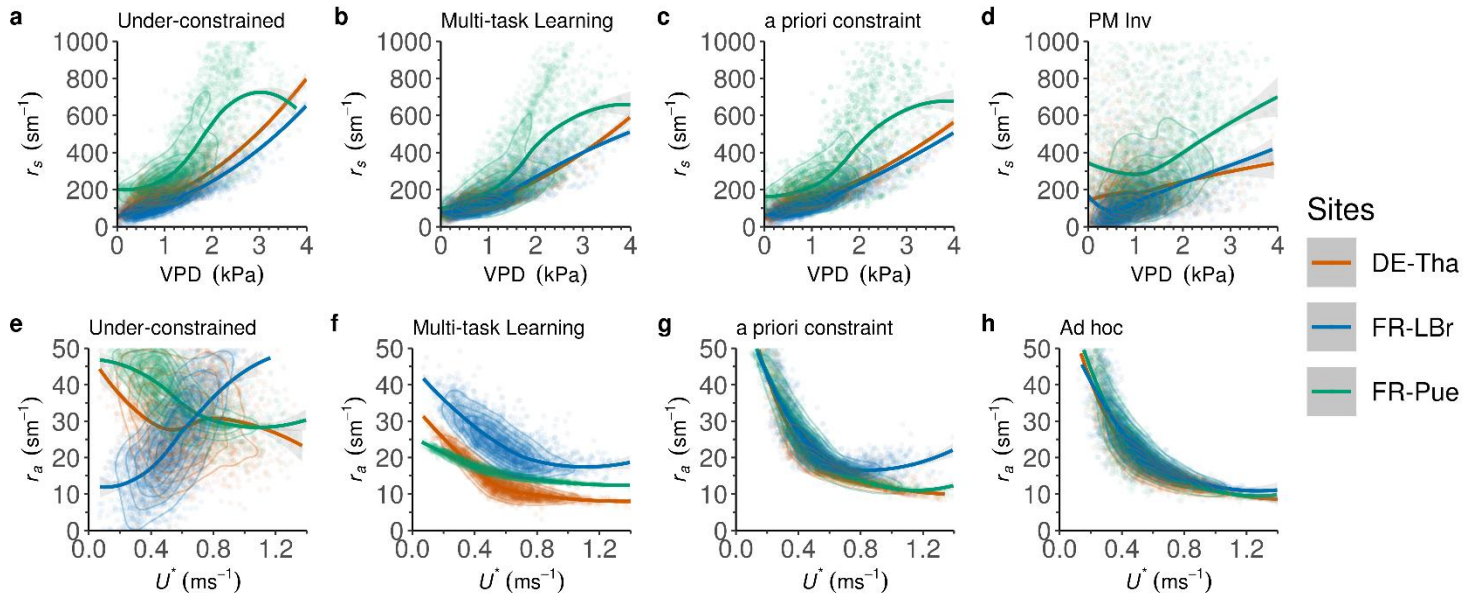


Figure 7: Evaluating  $Q_{LE}$  predictions against VPD for different models for forests (a-d) and grasslands (e-h). Higher evapotranspiration rates evident for grasslands compared to forests associated with higher stomatal conductance.

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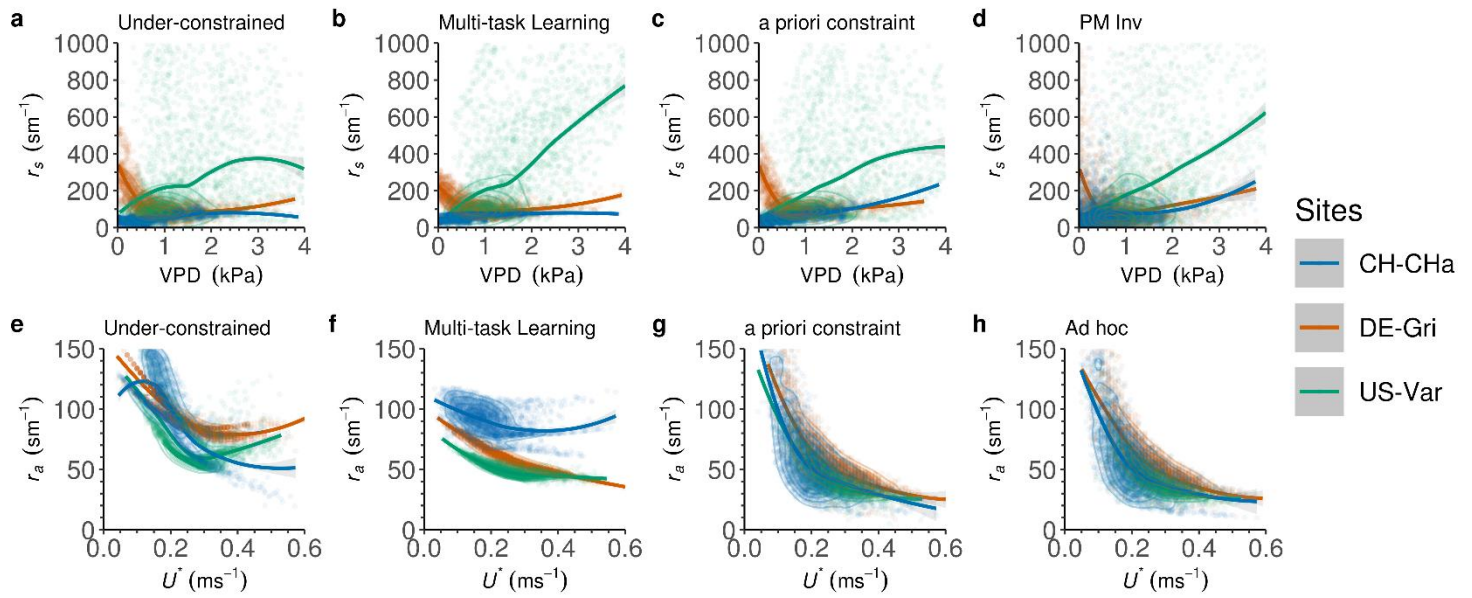


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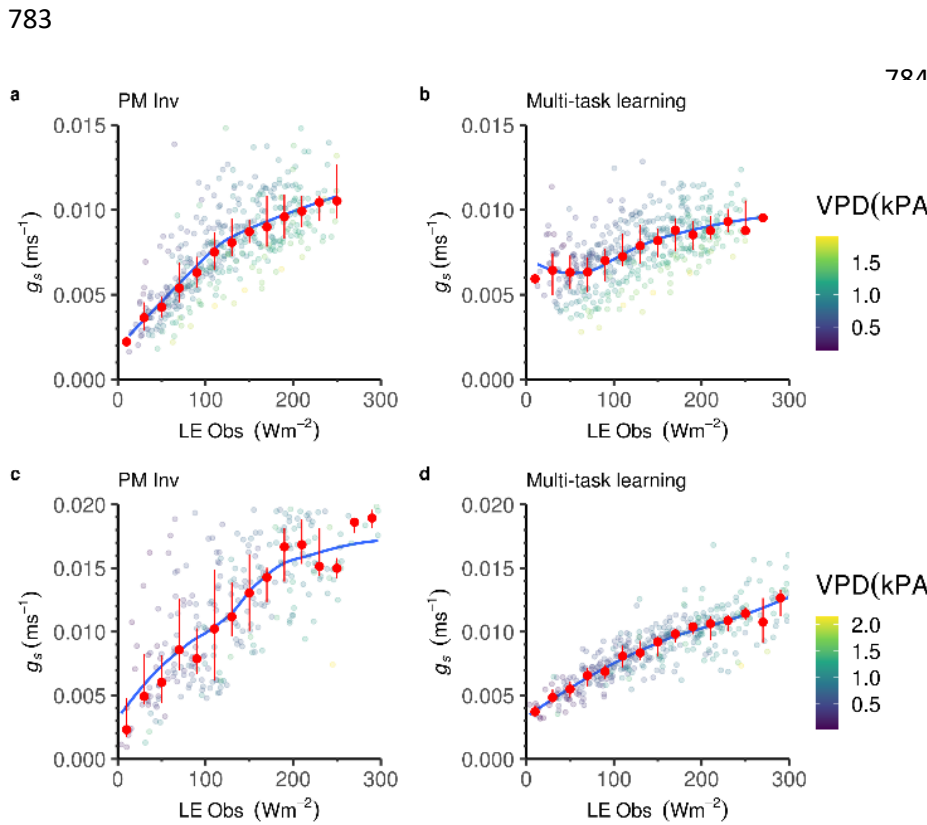
Figure 8: Assessing latent variables  $r_s$  and  $r_a$  against VPD and  $U^*$  respectively for different models in forests. Constrained hybrid models reveal physical consistency of latent variables compared to under-constrained model, especially under different environmental conditions.

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 Figure 9: Assessing latent variables  $r_s$  and  $r_a$  against VPD and  $U^*$  respectively for different models in grasslands. The constrained hybrid models yield more physically consistent results compared to under-constrained model, and able to capture the vegetation and climate heterogeneities.



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 Figure 10: Physical consistency of  $g_s$  and  $Q_{LE_{obs}}$  with VPD at mean diurnal scale of DE-Tha forest (a,b) and DE-Gri grassland (c,d). The multi-task learning model is able to capture the same patterns as shown by Penman-Monteith, while being more resistant to noise in the data which may cause overestimation of surface conductance due to the instability of the inversion.