Moisture Origin and Transport for Extreme Precipitation over Indonesia's New Capital City, Nusantara in August 2021

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Abstract

Nusantara, the Indonesia's New Capital City, experienced a rare extreme rainfall event on 27 to 28 August 2021. This unusual heavy rainfall occurred during the dry season and caused severe flooding and landslides. To better understand the underlying mechanisms for such extreme precipitation events, we investigated the moisture sources and transport processes using the Lagrangian model HYSPLIT. Our findings revealed that moisture was mostly transported to Nusantara along three major routes, namely from Borneo Island (BRN, 53.73%), the Banda Sea, and its Surroundings (BSS, 32.03%), and Sulawesi Island (SUL, 9.05%). Overall, BRN and SUL acted as the main sources of terrestrial moisture, while the BSS was the main oceanic moisture origin having a lower contribution than that of its terrestrial counterpart. The terrestrial moisture transport from BRN was mainly driven by the large-scale high vortex flow, while the moisture transport from the SUL was driven by the circulations induced by boreal summer intraseasonal oscillation (BSISO) and low-frequency variability associated with La Niña. The near-surface oceanic moisture transport from BSS is mainly associated with prevailing winds due to the Australian Monsoon system. These insights into moisture sources and pathways can potentially improve skill in predictions of summer precipitation extremes in Indonesia's New Capital City, Nusantara, and benefit natural resource managers in the region.



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Abstract: Nusantara, the Indonesia's New Capital City, experienced a rare extreme rainfall event on 1 27 to 28 August 2021. This unusual heavy rainfall occurred during the dry season and caused severe 2 flooding and landslides. To better understand the underlying mechanisms for such extreme precipi-3 tation events, we investigated the moisture sources and transport processes using the Lagrangian model HYSPLIT. Our findings revealed that moisture was mostly transported to Nusantara along 5 three major routes, namely from Borneo Island (BRN, 53.73%), the Banda Sea, and its Surroundings 6 (BSS, 32.03%), and Sulawesi Island (SUL, 9.05%). Overall, BRN and SUL acted as the main sources of 7 terrestrial moisture, while the BSS was the main oceanic moisture origin having a lower contribution 8 than that of its terrestrial counterpart. The terrestrial moisture transport from BRN was mainly 9 driven by the large-scale high vortex flow, while the moisture transport from the SUL was driven 10 by the circulations induced by boreal summer intraseasonal oscillation (BSISO) and low-frequency 11 variability associated with La Niña. The near-surface oceanic moisture transport from BSS is mainly 12 associated with prevailing winds due to the Australian Monsoon system. These insights into moisture 13 sources and pathways can potentially improve skill in predictions of summer precipitation extremes 14 in Indonesia's New Capital City, Nusantara, and benefit natural resource managers in the region.

Keywords: Moisture sources; Moisture transport, Extreme precipitation; Nusantara, Indonesia

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1. Introduction

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Indonesia's parliament has designated to move the National Capital City from Jakarta 18 to the province of East Kalimantan on the island of Borneo [1]. The name of this new 19 Capital City is Nusantara which is publicly known as IKN (Ibu Kota Negara Baru) in 20 Bahasa (Fig. 1). This city is a part of several sub-districts in Penajam Paser Utara (PPU) 21 regency and Kutai Kartanegara regency in the province of East Kalimantan [2]. Several 22 aspects informed as the underlying reasons why the government relocates the national 23 capital city to Nusantara, such as population distribution, tackling the clean water source 24 crisis on Java Island, and boosting economic growth outside of Java Island [3]. 25

According to the Regional Disaster Management Agency (BPBD), Nusantara has experienced several flood disasters in the past few years. It has been documented that from January 2019 to January 2022 there have been 15 flooding events. One of the major floods occurred on 27-28 August 2021, which exerted strong impacts on human lives and society [4]. The heavy precipitation from 27-28 August 2021 was considered unusual because it lasted more than 11 hours and occurred during the dry season [5]. This extreme rainfall subsequently triggered floods and landslides in several places in East Kalimantan. During this period, the daily rainfall accumulation reached 82 mm/day, which was within the 99th 33 percentile of the long-term daily precipitation distribution Section 2.1 for a detailed dis-34 cussion). Investigating the underlying physical process of this extreme event can improve 35 our understanding of the driving mechanisms of the unusual extreme rainfall event during 36 the dry season in Indonesia's New Capital City, Nusantara, and potentially improve the 37 summer rainfall prediction over the region. 38



Figure 1. Topography map of (a) the area study (blue rectangle) located in the middle of the Maritime Continent in the eastern Borneo Island, (b) administrative borderline of the Nusantara City (blue looped-line).

There is still a lack of understanding of the mechanism for the formation of extreme 40 rainfall over Nusantara in the province of East Kalimantan. This is likely due to the 41 complexity of the dynamics of the atmosphere over the region, which is located over the 42 equator. In general, the precipitation in East Kalimantan has a typical monsoonal-type 43 pattern [6]. This monsoonal type is characterized by the driest months from June to August 44 and the wettest months from December to February [6]. In addition, the heavy rainfall 45 over the central Maritime Continent is also often associated with an organized convective 46 system associated with Boreal Summer Intraseasonal Oscillation (BSISO) [7,8]. BSISO is 47 a large-scale weather system that strongly influences intra-seasonal rainfall variability 48 during boreal summer as a result of interaction between Madden Julian Oscillation (MJO) 49 and monsoonal system [7–9]. Furthermore, the tropical synoptic activities associated with 50 tropical high and low in a period range of about 2–8 days also often play a key role in 51 modulating summertime convection that causes day-to-day weather variations and heavy 52 rainfall over East Kalimantan during boreal summer [10]. In addition, the rainfall variation 53 in East Kalimantan is also affected by the interannual variability of the tropical sea surface 54

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temperature (SST) in the Pacific and Indian Oceans associated with El Niño-Southern

Oscillation (ENSO) [11,12]. In particular, El Niño causes prolong dry season, whereas La 56 Niña is related to high rainfall above the average [13]. All these atmospheric drivers acting 57 on different timescales can modulate and influence the formation and development of 58 extreme summertime rainfall events in Nusantara. 59

The occurrence of extreme rainfall events is associated with the high moisture content 61 in the atmosphere [14]. The more extreme the rainfall that occurs in a place depends on the 62 longer the condition of high moisture content persists [15]. Previous studies have shown 63 that moisture was one of the primary rainfall sources for the development of torrential rain-64 fall events. For example, an investigation by Nie and Sun [16] using a Lagrangian approach 65 concluded the extreme precipitation event over Henan in July 2021 was driven by moisture 66 from Southern China and the western North Pacific. In addition, Zhou et al. [17] studied 67 water vapor transports associated with torrential rainfall during the April-September period of 2008-2015 over Xinjiang, China. They found that water vapor related to heavy 69 precipitation events is mostly transported by the westerly winds. Moreover, Tan et al. [18] 70 analyzed a dominant contribution to the precipitation over western North America when 71 the atmospheric river and extreme precipitation occurred simultaneously. And, they found 72 that moisture flux convergence is the key. The aforementioned studies suggest that it is 73 important to evaluate the source of the moisture transport responsible for the precipitation 74 extremes to better understand the underlying mechanisms of such events. 75

In this study, we investigate the sources of moisture and transports for extreme precip-77 itation events in Nusantara on 27-28 August 2022 that caused catastrophic flooding with 78 significant socioeconomic effects in many places [19,20]. We employed a Hybrid Single-79 Particle Lagrangian Integrated Trajectory Model (HYSPLIT) method to track the origin 80 of moisture and transports responsible for the extreme precipitation event. Furthermore, 81 this study will also explore the underlying dynamics that drive such moisture transport 82 during extreme rainfall events, similar to the previous studies [21,22]. This is the first 83 comprehensive study that applies a Lagrangian moisture analysis to understand the cause 84 of the extreme precipitation event in Indonesia. Therefore, we hope that this study can 85 provide knowledge about the source monitoring and prediction system for water vapor 86 transport in East Kalimantan, especially over the new Capital City of Indonesia, Nusantara. 87 It is also expected that this system can be a decision support tool for decision-makers to 88 help users in developing a hydrometeorological disaster risk reduction strategy. In addition, we hope that this research can be one of the integration and synchronization steps needed 90 to understand the factors that cause flooding in the Nusantara area. Later, this is expected 91 to be a consideration in environmental development planning to mitigate flood natural 92 disasters. In Section 2 of this article, the data and methods are described. The results are presented in Section 3 before the article concludes with a discussion of the findings in 94 Section₄. 95

2. Data and Methods

2.1. Precipitation Data and Reanalysis Data

The half-hourly gridded rainfall dataset produced by the Global Precipitation Measure-99 ment (GPM) Integrated Multi-satellite Retrievals for GPM (IMERG) with 0.1° x 0.1° spatial 100 resolution was employed in this study to investigate the characteristic of heavy rainfall 101 [23]. We used the rainfall data for the period of 27 August 2021 14Z to 28 August 2021 08Z 102 from the IMERG version 3 final (IMERG-F), a final product of IMERG that was adjusted 103 by the analysis of the monthly Global Precipitation Climatology Centre (GPCC) gauge 104 [24]. This data was accessed from Goddard Earth Sciences Data and Information Services 105 Center (https://disc.gsfc.nasa.gov/datasets/, accessed on 25-March-2022). Moreover, the 106 hourly rainfall data from the Indonesian Agency for Meteorological Climatological and 107

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Geophysics (BMKG) weather station over Sepinggan, Balikpapan was used for the period of 27 August 2021 14Z to 28 August 2021 08Z.

This study also uses the top brightness temperature (TBB) of the Himawari-8 satellite 111 retrieved from the Himawari receiver at the National Research and Innovation Agency 112 (BRIN), which has a spatial resolution of 4 km and a temporal resolution of 1 hour. The 113 infrared channel at 10.4 μ m (Band 13) data was applied to identify the evolution of the 114 mesoscale convective system (MCS) associated with heavy precipitation. Generally, the 115 Himawari-8 with the interval of 10 minutes for the full disk scans creates the possibility 116 for monitoring more detailed clouds [25,26]. The MCS in the study were tracked by using 117 the GTG (Grab 'em, Tag 'em, Graph 'em) algorithm [27] which have been applied in the 118 previous study for investigating the heavy rainfall events in some regions of Indonesia 119 [21, 26, 28]120

To analyze the moisture transport and contribution, this study utilized the European 121 Center for Medium-Range Weather Forecast reanalysis 5 (ERA5) data, with a 0.25° x 0.25° 122 horizontal resolution and hourly temporal interval [29]. The ERA5 data on the single 123 surface level and pressure levels for the period of 24 to 27 August 2021 are used to track the 124 moisture source of heavy rainfall using the HYSPLIT and calculate the vertically integrated 125 water vapor transport (IVT). In addition, we also used horizontal wind data and specific 126 humidity at pressure levels for the same period to analyze the meteorological drivers 127 during extreme precipitation. 128

2.2. HYSPLIT Model and Backward Trajectories

The Lagrangian approach using HYSPLIT model version 5.1 is used to calculate the 130 back trajectories of moisture properties that triggered the heavy precipitation over Nusan-131 tara during the flood event on 28 August 2021. HYSPLIT model developed by NOAA's Air 132 Resources Laboratory is one of the most extensively used models for atmospheric transport 133 and dispersion analysis [30,31]. In this study, we ran a total of 99 grid points surrounding 134 the heavy rainfall area in Nusantara ($116.2^{\circ}-117.2^{\circ}E$, and $0.8^{\circ}-1.6^{\circ}S$) with a horizontal 135 interval of 0.1° (Figure 2). The start time of each trajectory is set at the peak of heavy rainfall 136 events (27 August 2021 at 18Z). Then, we calculated the 72-hour backward trajectories 137 from 500m to 5500 m above the ground level (with intervals of 500 m) to analyze moisture 138 transport and sources. 139



Figure 2. Release points of the moisture trajectories (every 0.1° , from $116.2^{\circ}-117.2^{\circ}E$, and $0.8^{\circ}-1.6^{\circ}S$). For every point, the vertical resolution is 500 m, from 500 to 5500 m above ground level (magl).

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2.3. Moisture Source Attribution

The contribution of different moisture sources that led to heavy precipitation in Nu-141 santara is identified by applying the algorithm from Sodemann et al. [32]. We calculated 142 the moisture changes of an air parcel during a time interval of 1 hour using the following 143 equation [33,34]: 144

$$\frac{Dq}{Dt} \approx \frac{\Delta q}{\Delta t} = E - P\left(gkg^{-1}h^{-1}\right) \tag{1}$$

$$\Delta q^{\circ}(t) = q(\vec{x}(t)) - q(\vec{x}(t-1h))$$
(2)

Generally, the changes in specific humidity in a certain interval (Dq/Dt) show the 146 net result of precipitation (P) and evaporation (E) processes along the trajectories [33]. A 147 moisture source will be identified through the trajectory locations (moisture intake event) 148 if the moisture of a particle increases ($\Delta q^{\circ} > 0$). Meanwhile, the precipitation is determined 149 if the moisture decreases ($\Delta q^{\circ} < 0$) [16,32]. To estimate the moisture source attribution of 150 specific areas, we divided the region of the study area $(11^{\circ}S-9^{\circ}N \text{ and } 93^{\circ}-133^{\circ}E)$ into ten 151 sub-regions by considering the land and ocean (see later in Fig. 7a). The 72-hour evolution 152 of moisture intake is then calculated for each sub-region. 153

2.4. Clustering for the Trajectories

The backward trajectories simulated by the HYSPLIT model have two-dimensional 155 locations (longitudes and latitudes) that can be categorized into several clusters. In this 156 study, *k*-means clustering was performed to identify the main groups of trajectories. The 157 k-means algorithm is an unsupervised clustering method that classifies a given data into a 158 set of k groups according to their characteristics [35,36]. This method also has the ability 159 to produce more stable cluster boundaries [37]. In addition, to determine the best number 160 of clusters, we applied the within-cluster sum of square errors (WCSS) metric based on 161 the elbow method [37,38]. The optimum number is selected when the WCSS reaches the 162 minimum value, which is 3 cluster centres in our study. 163

3. Results

3.1. Overview of the Extreme Precipitation Event

Figure 3a shows the precipitation accumulation distribution from 14Z 27 August 2021 167 to 08Z 28 August 2021. It is indicated that intense precipitation occurred in eastern Borneo, 168 central Borneo, southern Sulawesi, and the Banda Sea. It also shows that the coastal area 169 of Nusantara experienced heavy rainfall (precipitation of more than 100 mm), however, 170 inland the intensity decreased. The bar chart in Fig. 3 reveals that the rainfall in Nusantara 171 started at 14Z and reached its peak at 18Z with a maximum intensity of 15 mm/hour, and 172 decreased thereafter. The total precipitation from 14Z 27 August 2021 to 08Z 28 August 173 was 80 mm 2021, which reached the 99th percentile climatology value of daily precipitation 174 amount in this area, and hence it was categorized as an extreme rainfall event. 175

IVT between 1000-300 hPa (color shading) superimposed with streamlines of the 177 vertically integrated moisture flux (contour lines) during the peak of precipitation at 18Z 178 27 August 2021 is shown in Fig. 3b. There was a flow of moisture transported westward 179 along the surrounding sea area of the Banda Sea, and the Flores Sea, before its break into 180 Makassar strait and the Java Sea. The moisture flux streamline shows a converged path 181 pattern from the peripheral region into the middle of Borneo Island. In addition, it is also 182 evident that a vortex-like pattern was established on the sea of west Borneo Island, which 183 is also a prominent feature associated with the flux during the extreme precipitation period. These results give us hints about possible sources of moisture transport responsible for the 185 extreme rainfall event in Nusantara during this period, which will be further discussed in 186 Sections 3.1 to 3.2. 187

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Figure 3. a) Precipitation accumulation distribution between 14Z 27 August 2021 to 08Z 28 August 2021 shown by contour color in eastern Indonesia, inset map shows the precipitation accumulation zoomed-in over Nusantara region. Timeseries box shows the hourly precipitation observed by Rain gauge of Sepinggan (dark blue) and area-averaged hourly precipitation of the Nusantara City and surrounding area based on GPM IMERG (sky blue) and . b) Shade color represents 1000-300 hPa vertically integrated moisture condition and red streamline shows vertically integrated moisture flux at 18Z 27 August 2021.

It is important to note that the occurrence of a rare extreme rainfall event on 27 to 188 28 August 2021 during the dry season in Nusantara was associated with the formation 189 and development of the MCS. Figure 4 shows the three-hourly evolution of the MCS from 190 the Himawari-8 satellite superimposed with the rainfall rate from GPM IMERG for the 191 period of 27 August at 12Z to 28 August 2021 at 9Z. These complex convective clouds were 192 developed and lasted for about 22 hours, which mostly concentrated on the central area of 193 Borneo. The MCS initiated from several small cloud clusters emerged locally in the northern 194 part of Nusantara City (Fig. 4a). It constantly grew and became larger covering mostly 195 the central area of the Island and a minor part of eastern Borneo including Nusantara 196 (Figs. 4b-c), generating the peak rainfall rate at 18Z 27 August 2021 in the region. On the other side, another cloud cluster developed along the coast of East Borneo starting at 18Z 198 27 August 2021 and eventually merged to form the second MCS surrounding Nusantara 199 City three hours later (Fig. 4d). Afterward, the first MCS which was previously developed 200

in central Borneo merged with the second MCS which was developed along the coast of East Borneo (Fig. 4e). Then, they gradually dissipated followed by the decreasing intensity of rainfall(Figs. 4f-h).



Figure 4. Three-hourly evolution of the MCS (K, shaded) from Himawari-8 satellite superimposed with rainfall rate from GPM IMERG (mm.h1, contour) for the period of 27 August at 12Z to 28 August 2021 at 9Z. The location of Nusantara City is denoted by a red box.

3.2. Moisture Sources and Transport for Summer Extreme Precipitation in Nusantara

One of the key factors that favor the occurrence of deep convection such as MCSs and 205 extreme rainfall is the substantial moisture content in the lower atmosphere. Our results so far suggest that moisture responsible for the formation of MCS and extreme rainfall events 207 in Nusantara was transported from different regions (Fig. 3 and Fig. 4). Therefore, it is 208 important to investigate where the atmospheric moisture came from that favored the MCS 209 and consequently the extreme total amounts of precipitation in Nusantara (Fig. 3). In the 210 following, we will discuss in detail the origin, pathways, and contributions of moisture 211 from different sources as well as the role of atmospheric modes in driving the moisture 212 transport during extreme precipitation events in Nusantara on 27 to 28 August 2021. 213

3.2.1. Dominant Moisture Origin, Pathways, and Contributions

Our results indicate that the higher moisture content (higher IVT) is detected over the 215 BSS, the South China Sea and vicinity, and the Pacific Ocean near the equator (Fig. 3b). To 216 accurately investigate the source and pathways of these moisture transports, the moisture 217 over the flooding region is tracked back in time for 72 hours before the peak of rainfall on 218 27 August 2021 at 18Z (Fig. 5). To categorize these trajectories objectively, the pathways are 219 grouped by k-mean clustering [35,36], resulting in three main clusters based on the elbow 220 method [37,38]. The results show that moisture responsible for the extreme rainfall event 221 was transported from 3 directions: from the southeast as cluster 1, from the northwest as 222 cluster 2, and from the east as cluster 3 (Fig. 5b-d). Each cluster has distinct characteristics 223 based on the moisture content. Specific humidity was higher over trajectories from southeast/over cluster 1 (more than 12.5 g/kg) (Fig. 5b). The high specific humidity over this 225 cluster was consistent with the more abundant moisture showed by IVT magnitude over the Banda Sea and its vicinity (Fig. 3b). Moreover, the high specific humidity along the 227 tracking was also detected from the northwest (cluster 2), ranging from 7.5 g/kg to nearly 228 17.5 g/kg. IVT vector explains that moisture is transported clockwise over the South China 229 Sea and Borneo (Fig. 3b). Trajectories are consistent with IVT which follows the anticyclonic 230 flow over the South China Sea, then intrudes on Borneo Island to the flooding area (Fig. 5c). 231 Some pathways in cluster 2 also originated from the island of Borneo 72 hours before heavy 232 precipitation. All these pathways (from the South China Sea and the Island of Borneo) area 233 converged into the flooding area within cluster 2 (Fig. 5c). Moreover, cluster 3 has different 234 moisture characteristics among all clusters, with the fast westward movement of a drier 235 air parcel (approximately less than 12.5 g/kg) (Fig. 5d). The velocity of the air parcel is 236 shown by shorter pathways over 72 hours. Overall, these moisture propagation regimes 237 are consistent with the IVT vector which illustrates the moisture transported toward the 238 flooding location from various locations, including moisture movement from the southeast, 239 east, and northwest (IVT vector in Fig. 3b). However, the trajectories result indicates more 240 accurate pathways of moisture propagation toward Nusantara. 241



Figure 5. 72-hours backward trajectory of moisture responsible for the extreme rainfall event over the Nusantara region (at 18Z, 27 Aug 2021). Trajectories (lines) and moisture along the pathways (colors) for a) all trajectories and b-d) for each cluster.

The characteristics of moisture transport from each cluster are further explained by taking into account the vertical structure of trajectories. Figure 6 shows the altitude of tracked air parcels within 72 hours. The mean altitude of each cluster is calculated (solid line in Fig. 6) and presented with its maximum and minimum value (shading in Fig. 6). It is shown that within clusters 1-3, moisture arriving toward various release altitudes was dominated from the altitudes lower than 4.5 km. This is expected as the moist or humid air is found much below 5 km under a layer of warmer air in the lower troposphere [39].

Each cluster has a distinct characteristic based on its altitude profile. More specifically, cluster 1 is mainly advected horizontally toward the study area over low altitudes (rang-250 ing from 0 km-1km) (Fig. 6). The low-level transport over this cluster denotes moisture 251 exchange (moisture uptake and moisture outtake) over the earth's surface. Moreover, the 252 additional moisture/moisture uptake over the lower altitude (planetary boundary layers) is 253 closely related to the surface evaporation process [32]. Therefore, the indication of oceanic 254 evaporation as the main mechanism of moisture uptake is strong as this cluster spreads along the BSS. Section 3.2.2 will discuss more the moisture transport drivers at the lower 256 level toward Nusantara. Furthermore, cluster 2 propagates horizontally with an altitude 257 between 0-2.1 km from 72-35 hours prior to the extreme event. Within this cluster, the 258 moisture ascended to an altitude between 1.8-2.8 km during t=35-25 hours prior to the 259 peak of the extreme rainfall events. Moreover, moisture that arrived from Sulawesi and its 260 Surroundings (cluster 3) was mainly transported over higher altitudes (2.7 km on average) 261 and converged over flooding locations at an altitude of 4.7 km. The fast propagation of 262 air particles within cluster 3 (shorter pathways in Figure 5d) can be explained by a higher 263 wind speed in the free troposphere as there is no friction influenced by the roughness of 264 the earth's surface. Furthermore, the additional moisture over the higher altitude was 265 more likely through some mechanisms such as convection and evaporation of precipitating 266 hydrometeors, or because of errors related to numerical calculation of trajectories [32]. In 267 more detail, the factors that drive the moisture transport from clusters 2 and 3 toward 268 Nusantara City will be further discussed in Section 3.2.2. 269



Figure 6. Evolution of moisture transports as a function of altitudes and time prior to the extreme rainfall event (18Z, 27 August 2021). Solid lines represent the mean altitude for each cluster, shadings denote altitude ranges.

Our results so far indicate that there are three main sources of moisture transport 270 toward Nusantara that produced the extreme rainfall event. To quantify the contribution 271 of this moisture source, we further calculated the integrated moisture intake precisely 272 over ten different regions and quantified its percentages (Fig. 7). Similar to the previous 273 studies (e.g., [32] [33]), changes in moisture along the pathways can define the precipitation 274 (moisture outtake) and evaporation (moisture intake) process along the pathways. We 275 defined moisture intake as increasing moisture in time along its pathway (positive dq), wherein the lower troposphere is highly correlated with evaporation (evaporative moisture 277 uptake) [32]. Our results indicate that moisture intakes increase gradually over time, 278 indicating the moistening of air parcels during their pathways upon arrival (Fig. 7b). The 279 terrestrial moisture source is the dominant contributor to heavy rainfall, which accounts for 62.78% of the total moisture intake (53.73% of the moisture is evaporated over Borneo 281 (BRN) and 9.05% from Sulawesi (SUL)) (Fig. 7c). Moreover, the most significant oceanic 282 moisture source, about 32.03%, is evaporated over BSS. About 3.77% of moisture intake is 283 from the South China Sea (SCS), and 0.43% is from the Pacific Ocean and Surroundings 284 (POS), as the rest oceanic source of moisture for this heavy rainfall case (Fig. 7c). 285



Figure 7. The moisture source contributions by different regions. (a) The division into 10 regions including Peninsular Malaysia (PNM), Sumatera (SUM), Indian Ocean (IO), South China Sea (SCS), Java Sea (JAS), Java Island (JAVA), Borneo Island (BRN), Pacific Ocean and Surroundings (POS), Sulawesi Island (SUL), Banda Sea and Surrounding (BSS). b) The averaged of integrated moisture intake of the 10 sources along the trajectories in the target region. (c) The relative moisture contributions of different regions at the release time (0 days).

Further quantification of moisture contributions in each cluster from the different 286 source regions is explained in Figures 8b,d,f, where the moisture intake accumulation in 287 a particular area helps to calculate this contribution as shown in Figures 8a,c,e. It can be 288 seen that cluster 1 is dominated by oceanic moisture sources from the BSS which accounts 289 for 92.3% of the total moisture intake over cluster 1. The rest moisture sources are from 290 terrestrial (BRN and SUL), where moisture is entrained to its pathways at a more stable 291 pace. Moisture intake over BSS location is detected since -72h and increases gradually 292 until the release time. Sudden entrainment is detected over the BSS at 23-20 hours before 293 the heavy rainfall, indicating intensive evaporation over the ocean since this cluster is 294 characterized by moist air parcel with low-altitude moisture propagation (Figs 5b and 6). 295 As mentioned in the previous section, higher IVT is detected over BSS (Fig. 3b), and the 296 specific humidity along this cluster's trajectories is higher than in the other clusters (Fig. 5b). 297 Therefore, this cluster is essential as the moisture contributor to heavy precipitation. The 298 Australian monsoon is suspected of atmospheric conditions corresponding to the pattern 299 over this cluster, which is further explained in the following sub-section. 300



(b) Contribution proportion (cluster-1)



(d) Contribution proportion (cluster-2)





Figure 8. As Figures 7b and 7c in the manuscript but for the moisture intake in the three clusters.

In cluster 2, about 74.48% moisture was transported from BRN with a relatively steep 301 pattern indicating faster moistening of air parcel along the pathways over this area. It is 302 worth noting that since the release point of trajectory is also located over Borneo as well, 303 this cluster indicates moisture recycling. Moisture is recycled when the evaporation over 304 an area contributes to the precipitation over the same area, in which the evaporation is 305 closely related to vegetation and catchment area over a land [40,41]. This cluster is crucial 306 for the moisture source of the extreme rainfall event in Nusantara since BRN accounts for 307 half of the total moisture sources (Fig. 7), and the specific humidity along the trajectory is relatively high (Fig. 5c). This high relative humidity also can be explained by the MCS 309 analysis indicating the intensive cloud formation thus representing moisture intake over 310 BRN (Fig. 4). Over this cluster, the SCS, as the oceanic moisture source, contributes to 311 19.63% of the total moisture intake of this cluster. Furthermore, in Cluster 3, the contribution 312 of terrestrial (SUL and BRN) and oceanic (BSS) moisture sources is comparable (Fig. 8f). 313 Terrestrial moisture sources contribute 49.51 % and oceanic moisture source contributes 314 46.39 % of moisture sources. The pattern shows that moisture intake is increasing at a 315 nearly linear pace in time over this cluster (Fig. 8e). 316

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In summary, our results indicate that the moisture responsible for the extreme rainfall in Nusantara on 27-28 August 2021 was mainly transported along three major routes, namely from Borneo Island (BRN, with the contribution of 53.73%), the Banda Sea, and the Surroundings (BSS, with the contribution of 32.03%), and Sulawesi Island (SUL, with the contribution of 9.05%).

3.2.2. The role of atmospheric modes in the variability of moisture transport

To understand the mechanisms responsible for driving the three main pathways of 324 the moisture transport responsible for the extreme precipitation event during 27-28 August 325 2021 in Nusantara, we decompose the IVT stream function during the event into different 326 contributions of the atmospheric variability components. Figure 9 shows the decomposition 327 of the total IVT stream function into the contribution of 10-20 days, 30-60 days, >120 days, 328 and residual. The 10-20 day periodicity represents the contributions of high-frequency 329 oscillation that are often associated with synoptic weather systems (such as cyclones and 330 anti-cyclones) as well as the BSISO second mode (BSISO2) in summer [42]. The 30-60 day 331 periodicity is mainly associated with BSISO first mode (BSISO1) in summer [42]. The low-332 frequency variability (>120 days) is often associated with monsoonal flow, Indian Ocean Dipole (IOD) and ENSO [12,22,39,43]. Finally the residual includes other contributions 334 of another variability. In addition, we also calculate the time series of climate indices 335 associated with the Australian Summer monsoon index (AUSMI), BSISO, Nino 3.4, and 336 IOD to elucidate the drivers of such variability (Fig. 10).

Our results indicate that the near-surface moisture transport from the South East 338 (cluster 1) toward the Nusantara is mainly dominated by the low-frequency variability 339 (>120 days) (indicated by the dense stream function, Figs. 9a,d). We attribute this to the 340 active phase of the Australian monsoon system (Fig. 10a). As can be seen in the time 341 series of the AUSMI index, a strong and persistent Australian monsoon was observed 342 with the amplitude ranging between -8 and -6 a week before the flood event (Fig. 10a). 343 More importantly, this value increased from -7.2 m/s to -8.3m/s on 27 and 28 August 344 2021. This indicates a strengthening of near-surface south-easterly wind toward Nusantara. 345 This is consistent with the warm SST over the Northern Hemisphere (Asia) and colder 346 SST over the Southern Hemisphere (Australia) during boreal summer (not shown), which 347 causes the wind blows from high pressure in Australia to low pressure in Asia during the 348 Australian Monsoon [44]. An extensively warm SST in BSS enhanced the ocean evaporation 349 that triggered intense moisture uptake throughout this circulation (not shown). Therefore, 350 the prevailing wind pattern associated with the Australian Monsoon system induced the 351 near-surface south-easterly moisture transport from BSS to the Nusantara (cluster 1) as the 352 most significant oceanic moisture origin. 353



Figure 9. (a) Streamlines of the IVT during the period extreme rainfall (27-28 August 2021) and its decomposition into contributions of (b) 30 - 60 day variability, (c) 10 - 20 day variability, (d) >120 day variability, (e) b+c+d, and (f) residual (total - (b+c+d). The streamlined colors denote the amplitude of IVT, and a red box denotes the location of Nusantara.

Furthermore, the activity of high-frequency mode (10 to 20 days) associated with largescale high vortex flow located next to the west of Borneo (Fig. 9c) was mainly explained by 355 the transport from the SCS and BRN toward Nusantara. This circulation modulated the 356 water vapor propagation from the South China Sea (SCS) to Borneo Island as the source 357 of the oceanic moisture source in cluster 2. Then, it initiated a favorable environment for 358 the development of MCS that represents the moisture intake over Borneo Island (BRN), 359 contributing to the highest terrestrial moisture source for extreme precipitation events in 360 Nusantara. On the other hand, Boreal Summer Intraseasonal Oscillation 2 (BSISO2) which 361 is also identified as 10 to 20-day variability did not contribute to the enhanced moisture 362 over Nusantara as it was in the weak amplitude of phase 2 (Fig. 10b). 363

The eastward transport of moisture toward Nusantara (cluster 3) was mainly driven by the combined influence of 30 to 60-day and > 120-day variability (Figs. 9b,d). The 30-60-365 day variability during summer is consistent with the active phase 3 of BSISO1, where its 366 amplitude exceeded the threshold value (more than 1, in Fig. 10b). It then reached a peak at 367 around 1.8 during 27-28 August 2021. This BSISO1-induced easterly circulation contributes 368 to the moisture propagation from the eastern region of Indonesia towards Nusantara 369 and its surrounding area (Fig. 10b). On the other hand, the low-frequency variability 370 during this period is associated with La Nina which also played an important role in this 371 zonal transport (Fig. 10c). Those phenomena induced a relatively equal contribution of 372 terrestrial (SUL and BRN) and oceanic (BSS) moisture sources over cluster 3. The active

La Nina strengthened the zonal circulation that modulates the propagation of moisture to Nusantara resulting in extreme precipitation. Our finding further showed that the low-frequency variability has a greater contribution than that of BSISO1 in modulating the easterly transport of moisture toward Nusantara (Figs. 9b,d). 377



Figure 10. Daily time series of climate mode indices during 20-29 August 2021 for (a) AUSMI, (b) BSISO, (c) Nino 3.4, and (d) IOD. The Grey color shows the flood event over Nusantara on 27 - 28 August 2021

In addition, we have also examined a possible role of the other variability such as 378 IOD (Fig. 10d). It is shown that there was an active negative phase of IOD during 20–29 379 August 2021, with the threshold value of –1.0. A higher negative IOD index started on 20 380 August at about -1.2, then decreased gradually during the extreme precipitation event on 381 27–28 August 2021 to around –0.6. Even though IOD was in its active negative phase (that 382 was possible to drive the westerly zonal transport toward Nusantara), our results show 383 that the streamline of low-frequency IVT is mainly dominated by the south-easterly and 384 easterly transport associated with Monsoon and Niña, suggesting that IOD did not play an 385 important role during this extreme event.

In summary, we show that the transport of moisture from the south-eastward direction toward Nusantara (cluster 1) is mainly driven by the low-frequency variability (>120 days) associated with the Australian summer monsoon system, while the north-westward transport of moisture (cluster 2) is mainly driven by the anti-cyclonic activity having a synoptic periodicity of about 10-20 days. The combined influence of the 30 to 60-day oscillation associated with BSISO1 and the low-frequency variability associated with Niña is responsible for driving the westward transport (cluster 3) toward Nusantara during the extreme precipitation event. 394

4. Conclusion and Discussion

This study investigates the moisture transport and sources responsible for the extreme precipitation event over Indonesia's New Capital City, Nusantara on 27–28 August 2021 based on the HYSPLIT model, using the ERA-5 reanalysis, satellite, and in-situ data. Our major findings in this study are summarized as follows:

- The moisture responsible for the extreme precipitation event in Nusantara on 27–28
 August 2021 was transported along three dominant routes, namely BRN with the contribution of 53.73%, the BSS with the contribution of 32.03%, and SUL with the contribution of 9.05%.
- The BRN and SUL acted as the main sources of terrestrial moisture, while the BSS acted as the main oceanic moisture origin because most of the trajectories travel across the ocean where the evaporation takes apart.
- The Australian Monsoon system contributed significantly to the oceanic moisture transport from BSS. A strong and persistent prevailing wind strengthened the nearsurface flow passed a warm sea surface temperature in the BSS which supported the large moisture intake from BSS.
- Large-scale high vortex flow located to the west of Borneo Island contributed to the highest terrestrial moisture transport from BRN, while BSISO1 and low-frequency variability associated with La Nina modulated the moisture propagation from the eastern region of Indonesia.
- The results indicate the importance of terrestrial and oceanic moisture sources from the Borneo and Sulawesi Islands as well as the Banda Sea for the formation of extreme precipitation events in Nusantara, Indonesia.

The mechanisms driving the moisture transport during the extreme rainfall event in 418 Nusantara are summarized in a schematic diagram in Figure 11. This diagram illustrates 419 that the influence of the strong prevailing wind due to the Australian Monsoon system 420 is dominant at the near-surface layer to modulate the highest oceanic moisture source 421 from the southeast region. On the other hand, a large-scale high vortex circulation in the 422 west of Borneo Island transported the moist air from the South China Sea that favored the 423 development of MCS in Borneo Island, hence contributing to the highest proportion of 424 terrestrial moisture transport from BRN. Finally, the effects of BSISO1 and low-frequency 425 variability associated with La Nina play an important role in zonal circulation to supply 426 the moisture from the eastern part of Indonesia and the Pacific Ocean to Nusantara. 427

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Figure 11. A conceptual diagram depicting the role of atmospheric modes in the variability of the moisture source regions during the flood event over Nusantara on 27-28 August 2021. Color shading denotes SST.

Previous studies have revealed that the summertime rainfall variability in East Kali-428 mantan is influenced by several atmospheric modes such as the Australian Monsoon, ENSO, IOD, MJO, BSISO, and equatorial waves [8,45–51]. However, how the interaction 430 among those atmospheric modes drives the extreme precipitation from the perspective of 431 the moisture transport and sources, hence, devastating flood events in east Kalimantan is 432 still unknown. Our study is the first to show the mechanisms driving the extreme rainfall 433 event in Indonesia through the lens of the moisture transport and sources based on the 434 Lagrangian approach while taking into account the underlying atmospheric drivers. While 435 our current study mostly focused on the large-scale transport of moisture responsible 436 for the extreme rainfall event, other local forces such as topography and land-sea breeze circulation can create unique interactions that in turn drive moisture transport resulting 438 in extreme precipitation [49,52]. Therefore, understanding multi-scale interaction among 439 atmospheric phenomena resulting from extreme precipitation remains to be further studied. 440

The present study advances our understanding of the moisture sources and pathways for the extreme precipitation over Nusantara as well as the role of atmospheric modes in triggering moisture intake and propagation. These results can provide an important source of predictability for improving a skillful summer precipitation extreme prediction over Nusantara in the future, which in turn, can be potentially used by multi-stakeholders in developing disaster management in Indonesia's New Capital City, Nusantara.

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data from Global Precipitation Measurement (GPM) Integrated Multi-satellite Retrievals for GPM
(IMERG) can be accessed freely through https://disc.gsfc.nasa.gov/datasets/ (accessed on 25 March459

490

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