# The effect of faceting on olivine wetting properties

Yongsheng Huang<sup>1</sup>, Takayuki Nakatani<sup>1</sup>, Sando Sawa<sup>1</sup>, Michihiko Nakamura<sup>1</sup>, and Catherine Mccammon<sup>2</sup>

<sup>1</sup>Tohoku University <sup>2</sup>University of Bayreuth

November 22, 2022

#### Abstract

Grain-scale pore geometry primarily controls the fluid distribution in rocks, affecting material transport and geophysical response. The dihedral angle ( $\Theta$ ) in the olivine-fluid system is a key parameter determining the pore fluid geometry in mantle wedges. Both curved and faceted olivine-fluid interfaces define  $\Theta$  in the system, generating the faceted-faceted (FF), facetedcurved (FC), and curved-curved (CC) angles. However, the effect of faceting on  $\Theta$  under various pressure and temperature (P-T) conditions and fluid compositions have not been constrained, and its mineralogical understanding is unresolved. This study evaluates the facet-bearing  $\Theta$  and their proportions in olivine-multicomponent aqueous fluid systems. Our results show that 1/3 of olivine-fluid  $\Theta$  are facet-bearing angles irrelative to the P-T conditions and fluid compositions. Faceting produces larger dihedral angles than the CC angles. The grain boundary plane (GBP) distribution reveals that the GBPs of faceted interfaces at triple junctions were subjected to low Miller Index faces ((100), (010), and (101)). Moreover, calculating the FF angles from two adjacent low Miller index planes highly reproduces measured angle values based on the olivine crystal habit. Therefore, our study suggests that the FF angle is strongly affected by olivine crystallography. The presence of faceting increases  $\Theta$  and critical fluid fraction ( $\Phi$ c) for percolation, thus decreases the permeability. In the mantle wedge, where olivine crystallographic preferred orientation (CPO) is expected, increasing the FF angle proportion with associated changes in fluid pore morphology will lead to the permeability anisotropy and consequent geophysical anomalies.

1	The effect of faceting on olivine wetting properties
2	Yongsheng Huang <sup>1*, a</sup> , Takayuki Nakatani <sup>1, b</sup> , Sando Sawa <sup>1, c</sup> , Michihiko Nakamura <sup>1</sup> , Catherine
3	McCammon <sup>2</sup>
4	1. Department of Earth Science, Graduate School of Science, Tohoku University, Aramaki–Aza–
5	Aoba, Aoba–Ku, Sendai, Miyagi 980–8578, Japan
6	2. Bayerisches Geoinstitut, University of Bayreuth, 95440 Bayreuth, Germany
7	a. Current address: Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, 511
8	Kehua Street, Wushan, Tianhe, Guangzhou 510640, China.
9	b. Current address: Geological Survey of Japan, AIST Central 7, Higashi $1-1-1$ , Tsukuba,
10	Ibaraki 305–8567, Japan.
11	c. Current address: Division of Advanced Mechanical Systems Engineering, Institute of
12	Engineering, Tokyo University of Agriculture and Technology, Nakamachi 2-24-16, Koganei,
13	Tokyo, 184-0012, Japan.
14	* Corresponding author. Email: huangyongsheng@gig.ac.cn
15	

# 16 Key points

17	1.	Faceting produces larger dihedral angles than those with curved interfaces in the olivine-
18		fluid system.

- Grain boundary planes are frequently composed of low-index planes, and the facet-facet
   angles are explained by olivine crystallography.
- 3. The facet proportion is consistent with electrical conductivity data; faceting effects could
  be magnified under a sheared-wedge mantle.
- 23
- 24
- 25

# 26 Abstract

27 Grain-scale pore geometry primarily controls the fluid distribution in rocks, affecting material transport and geophysical response. The dihedral angle ( $\theta$ ) in the olivine-fluid system is 28 29 a key parameter determining the pore fluid geometry in mantle wedges. Both curved and faceted 30 olivine–fluid interfaces define  $\theta$  in the system, generating the faceted–faceted (FF), faceted–curved 31 (FC), and curved–curved (CC) angles. However, the effect of faceting on  $\theta$  under various pressure 32 and temperature (P-T) conditions and fluid compositions have not been constrained, and its mineralogical understanding is unresolved. This study evaluates the facet-bearing  $\theta$  and their 33 34 proportions in olivine-multicomponent aqueous fluid systems. Our results show that 1/3 of 35 olivine–fluid  $\theta$  are facet-bearing angles irrelative to the P–T conditions and fluid compositions. 36 Faceting produces larger dihedral angles than the CC angles. The grain boundary plane (GBP) 37 distribution reveals that the GBPs of faceted interfaces at triple junctions were subjected to low 38 Miller Index faces ((100), (010), and (101)). Moreover, calculating the FF angles from two 39 adjacent low Miller index planes highly reproduces measured angle values based on the olivine 40 crystal habit. Therefore, our study suggests that the FF angle is strongly affected by olivine 41 crystallography. The presence of faceting increases  $\theta$  and critical fluid fraction ( $\phi_c$ ) for percolation, 42 thus decreases the permeability. In the mantle wedge, where olivine crystallographic preferred 43 orientation (CPO) is expected, increasing the FF angle proportion with associated changes in fluid 44 pore morphology will lead to the permeability anisotropy and consequent geophysical anomalies.

45

# 47 Plain Language Summary

48 Pore geometry controls fluid-bearing rocks' material transport and geophysical properties. 49 affecting elemental cycling and the interpretation of magnetotelluric and seismological 50 observations in subduction zones. The dihedral angle ( $\theta$ ) in the olivine-fluid system is a key 51 parameter determining fluid geometry in a mantle wedge. Both curved and faceted olivine-fluid 52 interfaces appear, but the effect of faceting on  $\theta$  under various P-T conditions and fluid 53 compositions is poorly constrained. Here, we measured the facet-bearing  $\theta$  and their proportions 54 in olivine–multicomponent aqueous fluid systems. The results show that 1/3 of olivine–fluid  $\theta$  are facet-bearing irrelative of the P–T conditions and fluid compositions. The θ values of the facet-55 56 bearing angle are comparable to, or larger than, the curved-curved angle. Both grain boundary 57 planes of faceted olivine-fluid interfaces are preferentially subjected to low Miller Index faces; thus, the FF angle is strongly affected by olivine crystallography. The presence of faceting 58 59 increases  $\theta$ , the critical fluid fraction ( $\phi_c$ ), and decreases the permeability. In the mantle wedge, 60 where olivine crystallographic preferred orientation (CPO) is expected, the presence of an FF angle 61 with associated changes in fluid pore morphology will lead to permeability anisotropy and 62 geophysical anomalies.

Key words: dihedral angle, faceted plane, Miller Index, crystallographic orientation, fluid
 connectivity, permeability anisotropy.

65

66

# 68 1. Introduction

69 Pore geometry significantly controls the distribution of geological fluid (i.e., aqueous fluid 70 and silicate melt) in a deep mantle wedge, thereby affecting element cycling and geophysical 71 response in subduction zones (Watson & Brenan, 1987; Hermann et al., 2006; Iwamori, 1998; van 72 Keken et al., 2011; Pommier & Evans, 2017; Worzewski et al., 2011; Zheng et al., 2016). Although 73 channelized fluid flow has often been inferred from the field studies (Angiboust et al., 2014), the 74 pervasive grain-scale fluid flow may be the most plausible fluid migration regime at high pressure 75 (P) and high temperature (T) conditions where dissolution-precipitation intensively operates and 76 interfacial energy minimization ("textural equilibrium") is quickly attained. Moreover, the 77 pervasive nature may be suitable for explaining the resistivity anomalies observed at a 78 magnetotelluric (MT) grid-scale (commonly >10 km), because it would be required for the 79 channelized flows to be distributed continuously and nearly isotropically over this length scale. In 80 an olivine-dominant mantle rock, the olivine-fluid dihedral angle ( $\theta$ ) is the primary parameter 81 controlling the grain-scale fluid connectivity (Huang et al., 2019, 2020; Mibe et al., 1999; 82 Toramaru & Fujii, 1986). Therefore, a precise constraint on  $\theta$  in the olivine-fluid system is 83 important for a complete understanding of the fluid distribution and migration in subduction zones.

The dihedral angle is a consequence of the fluid–mineral interaction, which changes the fluid pore geometry through dissolution and precipitation processes to minimize the interfacial energy in the system. It is defined as the ratio of the grain boundary energy ( $\gamma_{ss}$ ) to the solid–fluid interfacial energy ( $\gamma_{sf}$ ) (von Bargen & Waff, 1986; Watson & Brenan, 1987) as follows:

88 
$$2\cos(\theta/2) = \gamma_{ss}/\gamma_{sf}$$
(1).

89 In an isotropic system, where solid-fluid interfaces are smoothly curved with a constant mean 90 curvature, the geometry of the intergranular fluid is solely determined by the dihedral angle. Fluids 91 can wet the grain edges well and migrate along with interconnected tubular networks at  $\theta < 60^{\circ}$ 92 irrespective of the fluid fraction ( $\phi$ ). In contrast, the fluid is distributed as isolated pockets along 93 the grain edges, corners, and boundaries at  $\theta > 60^\circ$  when  $\phi$  is lower than the critical value ( $\phi_c$ ) 94 (Watson & Brenan, 1987; Holness, 1992, 1993), which is dependent on  $\theta$  (von Bargen & Waff, 95 1986). However, the fluid distribution in the realistic rocks can deviate from the ideal distribution 96 (Laporte & Watson, 1995; Waff & Faul, 1992). In fact, Huang et al. (2021) measured the electrical 97 conductivity of the texturally equilibrated forsterite-saline fluid aggregate at 800°C and 1 GPa 98 showing that the conductivity is significantly smaller than that expected from the interconnected 99 tube model especially at low  $\phi$ . The synchrotron X-ray microtomography (CT) of the post-run 100 products revealed that the fluid pores gradually get interconnected with each other, increasing  $\phi$ 101 above ~1.0 % even though  $\theta$  defined by the curved–curved interface is smaller than 60° under the 102 experimental conditions (Huang et al., 2019).

103 In a realistic mineral-fluid system with interfacial energy anisotropy, facet planes, 104 crystallographically controlled planar solid-liquid interfaces, are often present along with the 105 curved interface (Watson & Lupulescu, 1993; Price et al., 2006; Wark & Watson, 2000; Watson, 106 1999). The coexistence of both curved and faceted interfaces generates three types of  $\theta$ : curved– 107 curved (CC), faceted-curved (FC), and faceted-faceted (FF)  $\theta$  (Yoshino et al., 2006). Pores 108 surrounded by facet planes are difficult to connect unless  $\phi$  increase above the critical value which 109 depends on  $\theta$  defined by the faceted interfaces (Price et al., 2006). Therefore, the facet plane may 110 affect the fluid connectivity depending on its  $\theta$  values, which helps to explain the results of

111 electrical conductivity measurements in Huang et al. (2021). Price et al. (2006) reported that facet 112 bearing angles are larger than those of the curved-curved interfaces in the quartz-fluid and 113 tremolite-fluid systems. If this is the case, facet-bearing pores may require the relatively high  $\phi$ 114 for the fluid interconnection. Although many studies have been conducted to investigate the P–T, 115 fluid composition, and mineral assemblage dependence of  $\theta$  in the olivine–fluid system (Watson 116 & Brenan, 1987; Huang et al., 2019, 2020; Mibe et al., 1998, 1999; Yoshino et al., 2006), the effect 117 of the angle-type variation on  $\theta$  is poorly understood. Yoshino et al. (2006) systematically 118 investigated the effect of faceting on pore geometry in texturally equilibrated rocks and discussed 119 its implications for permeability in several representative systems, including the San Carlos 120 olivine-MORB melt. However, olivine (forsterite)-aqueous fluid systems have not yet been 121 studied. Huang et al. (2019, 2020) studied the dihedral angle between olivine and multicomponent 122 aqueous fluids over a wide range of pressures and temperatures, but they primarily focused on the 123 CC  $\theta$ . Laporte and Provost (2000) theoretically investigated the  $\theta$  in the system with simplified 124 surface energy anisotropy and reported that the relationship between the mean equilibrium  $\theta$  and 125 the ratio of the grain boundary to the surface energy is close to the isotropic case. Nevertheless, 126 empirical investigation of the facet-bearing dihedral angles concerning crystallographic 127 orientations remains unresolved for actual mineral-fluid systems.

Investigating the grain boundary plane distribution (GBPD) can provide a mineralogical
understanding of angle-types in terms of the FF, FC, and CC angles. Some researchers previously
investigated GBPD in the fluid-free olivine polycrystals and found that low Miller Index planes
preferentially appears in the grain boundaries (Faul & Fitz Gerald, 1999; Marquardt et al., 2015).
If such specific grain boundary planes (GBPs) preferentially appear at the facet-bearing triple
junctions in the mineral-fluid system, they may significantly control θ of the facet-bearing angles

and its resultant fluid pore geometry. However, previous studies did not distinguish the grain
boundary type in terms of the angle–types (i.e., FF, FC, and CC), and the GBPD has not been
examined in the fluid-bearing system.

137 In this study, we investigated the facet-bearing (i.e., FF and FC)  $\theta$  in the run products of 138 the olivine–fluid systems previously obtained by piston–cylinder experiments at various pressures 139 (P: 1–3 GPa), temperatures (T: 800–1100 °C), and fluid compositions (pure H<sub>2</sub>O, H<sub>2</sub>O–NaCl, and 140 H<sub>2</sub>O–CO<sub>2</sub> systems) (Huang et al., 2019, 2020). These results were compared with the CC  $\theta$  to 141 clarify the effect of faceting. Moreover, we examined the crystallographic orientation of the olivine 142 aggregate and identified the GBPD at triple junctions with different angle-types to evaluate the 143 influence of the crystallographic orientations on  $\theta$ . Based on the results, we discussed the origin 144 of faceting and its effect on  $\theta$  and inferred the consequences of faceting on the fluid connectivity, 145 fluid distribution, and permeability anisotropy in the olivine–fluid system under static and sheared 146 mantle conditions.

#### 147 **2. Methods**

#### 148 **2.1. Samples**

We analyzed the run products of the olivine–fluid systems previously obtained by Huang et al. (2020, 2019). To constrain the P–T and fluid composition dependency of the facet-bearing  $\theta$ , we selected 19 samples that cover a wide range of experimental P–T conditions (1–3 GPa and 800–1100 °C) and fluid compositions (H<sub>2</sub>O, H<sub>2</sub>O–CO<sub>2</sub> with X<sub>(CO2)</sub> = CO<sub>2</sub>/(H<sub>2</sub>O+CO<sub>2</sub>) = 0.5 in molar ratio), and H<sub>2</sub>O–NaCl with 5.0 and 27.5 wt.%NaCl). The run products containing magnesite and orthopyroxene due to olivine carbonation (Huang et al., 2020) were not included in the present study to avoid the effect of mineral species other than olivine.

156 Herein, we summarize the experimental procedures employed in a previous study (Huang 157 et al., 2020, 2019) to synthesize the samples. High P–T experiments were conducted at Bayerisches Geoinstitut, University of Bayreuth, using end-loaded piston-cylinder apparatus. Piston diameters 158 159 of 3/4 inch and 1/2 inches were used for the experiments at pressures of 1-2 GPa and 3 GPa, 160 respectively, along with a standard Talc–Pyrex assembly (Keppler et al., 2003). The starting 161 material was San Carlos olivine powder ( $Fo_{91}Fa_{9}$ ) with grain sizes of 38–53 µm. Deionized and 162 distilled water were used as the pure H<sub>2</sub>O sources. Pure oxalic acid dihydrate (C<sub>2</sub>H<sub>2</sub>O<sub>4</sub>·2H<sub>2</sub>O) was 163 used as the CO<sub>2</sub> source. To obtain the NaCl solution, we dissolved reagent-grade NaCl (99.99% 164 NaCl) in deionized and distilled water at room temperature (~25°C) and atmospheric pressure. 165 Water and hydrous oxalic acid were mixed to obtain the target  $X_{(CO2)}$ . The olivine powder, 166 combined with around 3.0 wt% fluid, was loaded into the end-welded noble metal capsule that was 167 sealed by arc welding. The Au capsules (2.2 mm outer diameter) and Au<sub>80</sub>Pd<sub>20</sub> alloy capsules (2.0 168 mm outer diameter) were used for experiments at 800-1000 °C and 1100°C, respectively. The 169 detailed procedures employed for the high P–T experiments were provided by Huang et al. (2019, 170 2020). The run duration ranged from 72 h to 211 h, depending on the target temperature. The 171 experimental conditions and results are summarized in Table 1. In our experiments,  $CO_2$  was 172 assumed to be the predominant C species based on previous studies (Allen & C., 1972; Huang et 173 al., 2020; Médard et al., 2008).

The post-run capsules were cut using a diamond wire saw to expose the run products. The run products were then impregnated with an epoxy resin under a vacuum. We first polished these products using the sandpaper and diamond past with particle sizes down to 1.0 µm and then a 0.06 µm colloidal silica suspension. For electron backscattered diffraction (EBSD) analysis, the 178 samples were polished with colloidal silica suspension for more than 10 h using an automatic179 vibratory polishing machine (VibroMet, Buehler Ltd.).

#### 180 **2.2. Scanning Electron Microscope (SEM)**

We observed polished cross-sections of the run charges using a field-emission type scanning electron microscope (FE–SEM; JSM–7100F, JEOL Ltd.) with an accelerating voltage of 15 keV. The mineral phases and micro-texture of samples were observed. We took FF and FC type triple junctions for quantitative analysis. More than 300 backscattered electron (BSE) or secondary electron (SE) images of  $1280 \times 960$  pixels were obtained for each run product with high magnifications and resolution, depending on the pore size.

#### 187 **2.3.**

#### 2.3. Dihedral angle measurement

188 The apparent FF and FC  $\theta$  were measured at the triple junctions on the SEM images using 189 the Image–J software. We followed the method used for the CC angle measurements in Huang et 190 al. (2019, 2020). The median of measured angles was regarded as true  $\theta$  (Jurewicz & Jurewicz, 191 1986). Generally, the median of measured angles is close to the true  $\theta$  value to the greatest extent 192 when a sufficient number of angles are measured by using high magnification and high-resolution 193 SEM images. However, a deviation between the median and true angle could have occurred when 194 the sample suffered from severe plucking during cutting and polishing, which locally reduced the 195 randomness of the apparent angle distribution. Although Cmíral et al. (1998) demonstrated that  $\theta$ 196 values obtained with transmission electron microscopy (TEM) was smaller than those measured 197 from low magnification SEM images, our FE-SEM images were taken at a high magnification (up 198 to 150,000) and provided clear images comparable with those by TEM. In this study, more than 199 100 angles were measured on the acquired SE images for each angle-type for each sample. The

statistical error of the median value was estimated to be less than 1.5°, as discussed by Huang et al. (2020). The details of  $\theta$  selection and measurement are given in Huang et al. (2019).

202 **2.4. Electron Backscattered Diffraction (EBSD)** 

203 To identify the crystallographic orientation of the olivine aggregate and its consequences 204 on  $\theta$ , we mapped 10 representative samples by using FE–SEM equipped with an electron 205 backscatter diffraction (EBSD; HKL Channel5, Oxford Instruments plc.) at Tohoku University. 206 Analyses were conducted at an accelerating voltage of 15.0 kV. Two types of EBSD data were 207 collected for each sample. One was obtained with lower magnification ( $\times 200-500$ ) to investigate 208 the crystallographic preferred orientation (CPO) of the entire sample, and the other with higher 209 magnification ( $\times 5,000-15,000$ ) to identify the crystallographic orientation near the fluid pool. The 210 EBSD data were then processed using the MTEX MATLAB toolbox. Grain boundaries were 211 detected using a threshold misorientation angle of 10°. Grains smaller than two or three pixels 212 were removed, depending on the step size (250 nm). The grain-size distribution and CPO were 213 constructed from one point per data point.

214

215

216

217

218

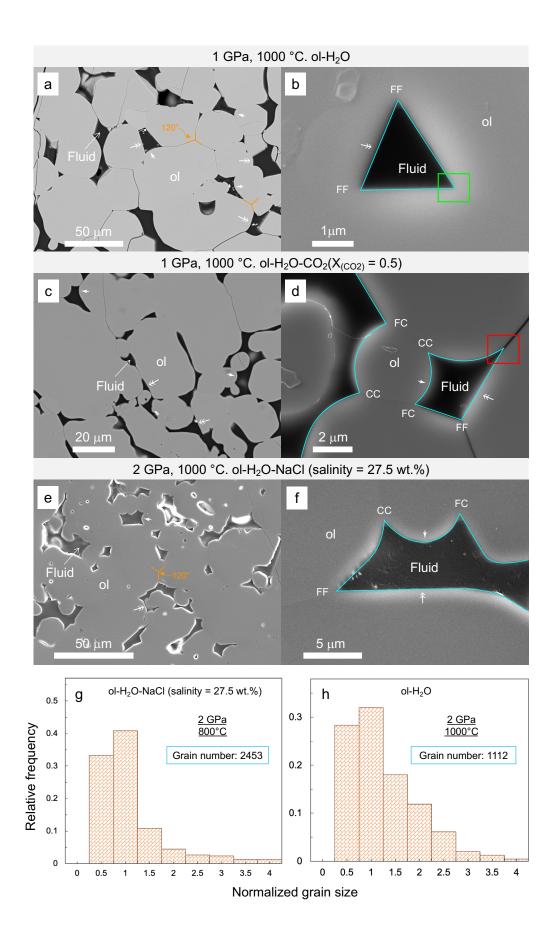
20	$\mathbf{n}$
L	<u>′</u> U

Table 1. Experimental conditions and results

Name         Gen         (b)         Union         Parame         Margin of the promote								Statistic angle					Number of									
1       1       800       192       H <sub>2</sub> O       0       -       187       32.6       -       -       CC       20       10       70       <	Run no.	P (GPa)	т (°С)	Duration (h)	Fluid phase	Salinity (wt%)	X(CO <sub>2</sub> )	number for	Faceting proportion ( % )	EBSD	GBPD	Angle Type	angle	Median (Degree)	I&D group							
1         1         800         192         H <sub>2</sub> O         0         -         193         32.0         -         -         OCC         OCC         712         1           2         1         800         192         H <sub>2</sub> O-NaCl         5         -         218         31.9         -         -         FC         100         63.5         0           3         1         800         192         H <sub>2</sub> O-NaCl         27.5         -         309         32.9         -         -         -         CC <sup>2</sup> 222         63.5         0           4         1         1000         120         H <sub>2</sub> O-NaCl         27.5         -         302         32.0         ·         ·         FC         1131         64.5         1           5         1         1000         120         H <sub>2</sub> O-NaCl         27.5         -         302         32.0         ·         ·         FC         1131         66.5         1         1000         120         H <sub>2</sub> O-NaCl         27.5         -         201         33.3         ·         ·         FC         1132         68.5         1         168         68.1         0         168         68.1												FF		70.8								
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	1	1	800	102	H-O	٥	_	187	32.6	_	_				1							
2         1         800         192         H <sub>0</sub> O-NaCl         5         -         218         31.9         -         -         -         0         0         0         0           3         1         800         192         H <sub>0</sub> O-NaCl         27.5         -         309         32.9         - <td></td> <td></td> <td>000</td> <td>152</td> <td>1120</td> <td>0</td> <td>-</td> <td>107</td> <td>52.0</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td>			000	152	1120	0	-	107	52.0													
2       1       800       192       H <sub>0</sub> O-NaCl       5       -       218       31.9       -       -       FC       122       63.5       0         3       1       800       192       H <sub>0</sub> O-NaCl       27.5       -       309       32.9       -       -       FC       123       63.3       0         4       1       1000       120       H <sub>0</sub> O-NaCl       27.5       -       509       30.4       ·!       ·!       ·!       FC       133       64.1       ·!       ·!         5       1       1000       120       H <sub>0</sub> O-NaCl       27.5       -       302       32.0       ·!																						
2         1         000         192         H <sub>2</sub> O-NaCl         2         1         3         -         -         OCC <sup>1</sup> 22         66.5         0           3         1         800         192         H <sub>2</sub> O-NaCl         27.5         -         309         32.9         -         -         -         CC <sup>2</sup> 207         63.3         0           4         1         1000         120         H <sub>2</sub> O-NaCl         27.5         -         309         32.9         -         -         -         CC <sup>2</sup> 207         63.5         0           5         1         1000         120         H <sub>2</sub> O-CO <sub>2</sub> 0         0.5         383         32.1         ·         ·         FF         145         72.2         60.6         77.8         60.0         77.8         7.8         60.0         77.8         7.8																						
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	2	1	800	192	H <sub>2</sub> O-NaCl	5	-	218	31.9	-	-				D							
3         1         800         192         H <sub>0</sub> O-NaCl         27.5         -         209         32.9         -																						
3         1         800         192         H <sub>2</sub> O-NaCl         27.5         -         309         32.9         -         -																						
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$															_							
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	3	1	800	192	H <sub>2</sub> O-NaCl	27.5	-	309	32.9	-	-				D							
4       1       1000       120       H <sub>0</sub> O       0       -       509       30.4       ·i       ·i       -       C       100       433       64.2       i         5       1       1000       120       H <sub>0</sub> O-CO2       0       0.5       383       32.1       ·i       ·i       -       60       713       77.2       1       60       77       2       800       120       H <sub>0</sub> O-NaCl       27.5       -       302       32.0       ·i       ·i       -       66       76       77       66       77       72       800       122       H <sub>0</sub> O-NaCl       75       -       205       31.8       ·i       -       -       66       76       776       68.6       1         8       2       800       192       H <sub>0</sub> O-NaCl       75       -       201       33.3       ·i       ·i       -       66       77       66.5       77       66.5       77       76       776       76       777       76       76       777       76       76       777       76       76       777       76       76       777       76       76       777       76       776       776 <td></td> <td>453</td> <td>60.6</td> <td></td>													453	60.6								
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	4	1	1000	120	H-O	0	_	509	30.4	1	~				1							
5         1         1000         120         H <sub>Q</sub> O-CO <sub>2</sub> 0         0.5         383         32.1         i         i         i         180         72         171         74.4         1           6         1         1000         120         H <sub>Q</sub> O-NaCl         27.5         -         302         32.0         i         i         16         174         66         7         0         67         174         66         174         66         1         66         1         66         1         66         1         66         1         66         7         7         66         7         7         7         67         176         667         10         66         1         66	-		1000	120	1120	0	-	000	50.4	,	,											
5       1       1000       120       H <sub>2</sub> O-CO <sub>2</sub> 0       0.5       383       32.1       4'       4' $\frac{FG}{CC}$ 111       78.4       1         6       1       1000       120       H <sub>2</sub> O-NaCl       27.5       -       302       32.0       4'       4' $\frac{FG}{CC}$ 111       78.4       1         7       2       800       120       H <sub>2</sub> O-NaCl       57.5       -       302       32.0       4'       4' $\frac{FG}{CC}$ 117.8       68.7       1       1       10       <																						
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $																						
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	5	1	1000	120	H <sub>2</sub> O-CO <sub>2</sub>	0	0.5	383	32.1	$\checkmark$	$\checkmark$				I.							
6         1         1000         120         H <sub>2</sub> O-NaCl         27.5         -         302         32.0         V         V         FC         1174         907         D           7         2         800         212         H <sub>2</sub> O         0         -         205         31.8         V         -         FC         174         68.8         FC         174         68.8         FC         174         68.8         FC         174         68.8         FC         176         57.8         FC         176         57.8         FC         176         57.8         FC         178         55.0         FC																						
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $																						
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	6	1	1000	120	H <sub>2</sub> O-NaCl	27.5	-	302	32.0		$\checkmark$				D							
7         2         800         212         H <sub>2</sub> O         0         -         205         31.8         4         -         FFC CC <sup>2</sup> 136 CC <sup>2</sup> 68.7 CC <sup>2</sup> 136 CC <sup>2</sup> 68.7 CC <sup>2</sup> 137         68.7 CC <sup>2</sup> 138         66.8 CC <sup>2</sup> 138         66.8 CC <sup>2</sup> 138         139         139         139         139         139         139         139         149         139         139         149         139         149 </td <td></td>																						
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$																						
7       2       800       212       H <sub>2</sub> O       0       -       205       31.8       ×       -       CC <sup>2</sup> 212       69.9       1         8       2       800       192       H <sub>2</sub> O-NaCl       5       -       201       33.3       √       √       √       FF       118       64.3         9       2       800       210       H <sub>2</sub> O-NaCl       27.5       -       269       31.3       √       √       FF       117       68.3       D         10       2       1000       120       H <sub>2</sub> O-NaCl       27.5       -       269       31.3       √       √       FF       107       68.3       D         11       2       1000       120       H <sub>2</sub> O-NaCl       27.5       -       269       33.5       √       √       FF       101       50.7       FF       101       50.7       FF       101       50.7       FF       102       57.0       FF       101       50.7       FF       102       57.0       FF       120	_	_				_				,												
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	7	2	800	212	H <sub>2</sub> O	0	-	205	31.8		-				1							
8       2       800       192       H <sub>2</sub> O-NaCl       5       -       201       33.3       V       V																						
8       2       800       192       H <sub>2</sub> O-NaCl       5       -       201       33.3       V       V       FC       172       57.9       p         9       2       800       210       H <sub>2</sub> O-NaCl       27.5       -       269       31.3       V       V       FC       172       57.9       p         9       2       800       210       H <sub>2</sub> O-NaCl       27.5       -       269       31.3       V       V       FC       172       55.4       p         10       2       1000       120       H <sub>2</sub> O-NaCl       5       -       314       32.8       V       V       FC       120       55.6       1         11       2       1000       133       H <sub>2</sub> O-NaCl       5       -       314       32.8       V       V       FC       160       45.5         11       2       1000       72       H <sub>2</sub> O-NaCl       5       -       316       33.8       V       V       FC       160       56.6       16.5       16.5       16.5       16.5       16.5       17.5       16.5       17.5       16.5       17.5       16.5       17.5       16.5       16.5												FF										
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	~		~			400		-		001		,	1				-					
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	8	2	800	192	H <sub>2</sub> O-NaCI	5	-	201	33.3	V	N		247		D							
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$												All	537	56.4								
$\begin{array}{c c c c c c c c c c c c c c c c c c c $												FF	107									
$ \begin{array}{ccccc} & 210 & 96.4 \\ \hline All & 495 & 96.0 \\ \hline FF & 102 & 86.4 \\ \hline FF & 102 & 86.0 \\ \hline FF & 102 & 86.4 \\ \hline FF & 102 & 86.4 \\ \hline FF & 102 & 86.0 \\ \hline CC^{*} & 213 & 86.0 \\ \hline CC^{*} & 213 & 86.0 \\ \hline CC^{*} & 211 & 86.0 \\ \hline H & 50.7 \\ \hline CC^{*} & 211 & 43.4 \\ \hline FF & 101 & 50.7 \\ \hline CC^{*} & 211 & 43.4 \\ \hline FF & 101 & 50.7 \\ \hline CC^{*} & 211 & 43.4 \\ \hline FF & 101 & 50.7 \\ \hline CC^{*} & 211 & 43.4 \\ \hline FF & 102 & 57.0 \\ \hline CC^{*} & 211 & 43.4 \\ \hline FF & 102 & 57.0 \\ \hline CC^{*} & 211 & 43.4 \\ \hline FF & 120 & 57.0 \\ \hline CC^{*} & 304 & 51.7 \\ \hline CC^{*} & 304 & 51.7 \\ \hline FF & 134 & 99.1 \\ \hline CC^{*} & 304 & 51.7 \\ \hline FF & 134 & 59.1 \\ \hline FF & 134 & 59.1 \\ \hline CC^{*} & 304 & 51.7 \\ \hline FF & 134 & 59.1 \\ \hline FF & 135 & 56.0 \\ \hline FF & 107 & 55.0 \\ \hline FF & 107 & 55.0 \\ \hline FF & 103 & 64.0 \\ \hline FF & 104 & 60.0 \\ \hline FF & 107 & 55.0 \\ \hline FF & 102 & 64.0 \\ \hline FF & 107 & $	0	2	800	210		27 5		260	21.2	1	2		178	59.0	D							
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	9	2				21.5	-	269	31.3	N	v	CC <sup>a</sup>										
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $																						
$ \begin{array}{ccccc} 100 & 120 & 120 & 120 & 120 & 120 & 120 & 0 & - & 582 & 33.5 & 4 & 4 & 7 & 7 & 7 & 7 & 7 & 7 & 7 & 7$		2	1000	120	H₂O																	
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	10					0	_	582	33.5	V					I.							
$ \begin{array}{ccccccc} 11 & 2 & 1000 & 133 & $						1120				1120			20	. 120	20		Ũ		002			
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$												All										
$\begin{array}{ccccccc} 11 & 2 & 1000 & 133 & H_2O-NaCl & 5 & - & 314 & 32.8 & v & v & Cc^a & 201 & 43.4 & D \\ & & & & & & & & & & & & & & & & \\ & & & & & & & & & & & & & & & \\ & & & & & & & & & & & & & & \\ & & & & & & & & & & & & & & \\ & & & & & & & & & & & & & \\ & & & & & & & & & & & & & \\ & & & & & & & & & & & & & \\ & & & & & & & & & & & \\ & & & & & & & & & & & \\ & & & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & & \\ & & & & & & \\ & & & & & & \\ & & & & & & \\ & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & \\ & & & & & & & & & \\ & & & $			1000	133					32.8						D							
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	11	2			H <sub>2</sub> O-NaCl	5	-	314		$\checkmark$	$\checkmark$											
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$																						
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$												FE										
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		_					-	136														
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	12	2	1000	120	H <sub>2</sub> O-NaCl	27.5			33.8		V				D							
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$																						
$ \begin{array}{cccccc} 13 & 2 & 1100 & 72 & H_2O & 0 & - & 284 & 34.2 & - & - & \begin{array}{ccccccccccccccccccccccccccccccccccc$												-										
$ \begin{array}{cccccc} & 210 & 53.8 \\ & All & 497 & 55.9 \\ & FF & 104 & 58.3 \\ & FC & 224 & 52.6 \\ & FC & 224 & 52.6 \\ & FC & 224 & 52.6 \\ & FF & 102 & 50.2 \\ & & & & & & \\ & & & & & & \\ & & & & $	10	2	2 1100	) 72	H <sub>2</sub> O	ц <u>о</u>	0		004				FC			5						
$ \begin{array}{ccccccc} & 1100 & 72 & H_2 O-NaCl & 27.5 & - & 201 & 30.0 & - & - & - & - & - & - & - & - & - & $	15	2				U	-	284	34.2	-	-	CC <sup>a</sup>	210	53.8	U							
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$																	55.9					
$ \begin{array}{ccccccc} 14 & 2 & 1100 & 12 & 12 & 12 & 12 & 12 & 1$		2					H₀O-NaCl															
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	14		2	1100	72	H <sub>2</sub> O-NaCl		H <sub>2</sub> O-NaCl	H₀O-NaCl	H <sub>2</sub> O-NaCl	27 5	-	201	30.0	-	-				р		
$ \begin{array}{cccccccc} 15 & 3 & 800 & 211 & H_2O & 0 & - & 100 & 35.8 & - & - & - & - & - & - & - & - & - & $	14				12	00 72	12	12	12	12	1120 11001	21.0		201	50.0						D	
$ \begin{array}{ccccccc} 15 & 3 & 800 & 211 & H_2O & 0 & - & 100 & 35.8 & - & - & \begin{array}{ccccccccccccccccccccccccccccccccccc$																						
$ \begin{array}{c ccccc} 15 & 3 & 800 & 211 & H_2O & 0 & - & 100 & 33.8 & - & - & \\ & & & & & & & \\ 16 & 3 & 800 & 192 & H_2O-NaCl & 5 & - & 166 & 33.6 & - & - & & \\ & & & & & & & \\ 17 & 3 & 800 & 211 & H_2O-NaCl & 27.5 & - & 166 & 33.6 & - & - & & \\ & & & & & & & \\ & & & & &$		3																				
$ \begin{array}{c cccc} & 200 & 66.2 \\ \hline All & 488 & 66.0 \\ \hline All & 488 & 66.0 \\ \hline FF & 107 & 55.0 \\ \hline FC & 132 & 53.2 \\ \hline CC^{2} & 242 & 50.2 \\ \hline All & 481 & 52.4 \\ \hline FF & 103 & 64.0 \\ \hline CC^{2} & 200 & 56.4 \\ \hline FF & 103 & 64.0 \\ \hline CC^{2} & 200 & 56.4 \\ \hline FF & 103 & 64.0 \\ \hline CC^{2} & 200 & 56.4 \\ \hline FF & 103 & 64.0 \\ \hline CC^{2} & 200 & 56.4 \\ \hline FF & 100 & 60.6 \\ \hline CC^{2} & 200 & 56.4 \\ \hline FF & 100 & 60.6 \\ \hline CC^{2} & 200 & 56.4 \\ \hline FF & 100 & 60.6 \\ \hline CC^{2} & 200 & 56.4 \\ \hline FF & 100 & 60.6 \\ \hline F$	15		3 800	3 800	3 800	3 800	211	211	211	211	0 211	H <sub>2</sub> O	0	-	100	35.8	-	-				1
$ \begin{array}{cccccc} & & & & & \\ 16 & 3 & 800 & 192 & H_2 O-NaCl & 5 & - & 166 & 33.6 & - & - & \begin{array}{ccccccccccccccccccccccccccccccccccc$														2 -								
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$																						
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$		3																				
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	16		800	192	H <sub>2</sub> O-NaCl	5	-	166	33.6	-	-				D							
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$																						
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		3	800	211	H <sub>2</sub> O-NaCl																	
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$								157														
$\begin{array}{c c c c c c c c c c c c c c c c c c c $	17					27.5	-		36.9	-	-				D							
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$																						
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$					120 H <sub>2</sub> O																	
$\begin{array}{c ccccc} 10 & 120 & H_2O & 0 & - & 344 & 20.3 & - & - & & \\ & & & & & & & & \\ & & & &$	40					•		344	28.3	-	-				-							
$\begin{array}{c cccc} & All & 514 & 53.7 \\ \hline FF & 102 & 61.4 \\ \hline 19 & 3 & 1000 & 120 & H_2 O-NaCl & 27.5 & - & 158 & 30.0 &  &  & \frac{FC}{CC^3} & 220 & 47.4 & D \\ \hline \end{array}$	18					0	-								D							
$\begin{array}{ccccccc} & & & & & & & & & & & & & & & &$																						
19 3 1000 120 H_2O-NaCl 27.5 - 158 30.0 $\sqrt{\begin{array}{c} FC \\ C^a \end{array}}$ 147 55.5 D CC <sup>a</sup> 220 47.4 D						07.5																
$19  5  1000  120  n_20 - n_3cc  27.5  -  150  50.0  v  v  cc^a  220  47.4$	10							450	00.0	-1					5							
	19		1000	120	H <sub>2</sub> O-NaCl	27.5	-	158	30.0	N	N				U							
												All	469	51.6								

Note: The fluid fraction in each experiment was ~3.0 wt.%. The true  $\theta$  value for each system was a median value  $\pm$ 223 1.5°. The analytical error for the faceting proportion was 0.5%. a the CC is cited from Huang et al. (2019); b the CC 224 is cited from Huang et al. (2020). The run products that were employed for the electron backscattered diffraction 225 (EBSD) and grain boundary plane distribution (GBPD) are marked in the table.  $X_{(CO2)} = CO_2/(H_2O+CO_2)$  in mole.

- 226 Abbreviations: ol = olivine, FF = faceted–faceted angle, FC = faceted–curved angle, CC = curved–curved angle, All
- 227 = all measured angles, I = angle-type independent group, D = angle-type dependent group.



229 Figure 1. Representative SEM images and grain size distribution of the run products. a Backscattered electron 230 image of the run product in the olivine-H<sub>2</sub>O system at 1000°C and 1 GPa for 120 h. b High-magnification secondary 231 electron image of a typical triple junction in the olivine-H<sub>2</sub>O system at 1000°C and 1 GPa, which shows apparent FF 232 angles. c Backscattered electron image of the run product in the H<sub>2</sub>O–CO<sub>2</sub> system ( $X_{(CO2)} = 0.5$ ) at 1000°C and 1 GPa 233 for 211 h. **d** High-magnification secondary electron image of the typical apparent  $\theta$  in the H<sub>2</sub>O–CO<sub>2</sub> system (X<sub>(CO2)</sub>= 234 0.5) at 1000°C and 1 GPa for 211 h, which shows the coexistence of three types of apparent  $\theta$ . e Backscattered electron 235 image of the run product in the H<sub>2</sub>O–NaCl system (salinity = 27.5 wt.%) at 2 GPa and 1000°C for 120 h. f High-236 magnification secondary electron image of typical apparent  $\theta$  in the H<sub>2</sub>O–NaCl system (salinity = 27.5 wt.%) at 2 GPa 237 and 1000°C for 120 h. The run products are composed of mineral grains and epoxy resin-filled pores previously filled 238 with aqueous fluid during the experiment. The orange marks denote the olivine-olivine-olivine triple junctions with 239 intersection angles of  $\sim 120^{\circ}$ , indicating the attainment of textural equilibrium. The white single and double arrows 240 represent the curved and faceted interfaces, respectively. For the CC, FC, FF angles, the interfaces in **b**, **d**, and **f** are 241 highlighted by cyan curves. The green and red rectangles illustrate suitable and unsuitable angles (i.e., with a crack) 242 for measurement, respectively. g, Histogram of the grain size distribution in the H<sub>2</sub>O–NaCl system (salinity = 27.5 243 wt.%) at 2 GPa and 800°C. h, Histogram of the grain size distribution in the H<sub>2</sub>O system at 2 GPa and 1000°C. The 244 grain size distribution was normalized by the mean grain size of the recovered sample. The grain size is concentrated 245 with a peak around the mean grain size. Abbreviations: ol = olivine, FF = faceted-faceted angle, FC = faceted-curved 246 angle, CC= curved–curved angle.

# 247 **3. Results**

248 **3.1. Product phases and microstructures** 

In all systems, the recovered samples are composed of olivine aggregates and intergranular fluid pores which are mostly filled with epoxy resin (Figure 1). The grain size of olivine in the run products increased with an enhancement of temperature, reaching ~110  $\mu$ m at 1100°C through grain growth by Ostwald ripening and coalescence of two adjacent grains. The olivine grains were homogeneous in chemical compositions without obvious compositional zoning. Fluid-filled pores

254 were generally encompassed by three or more grains. Curved interfaces often coexisted with flat 255 interfaces, even within a single pore (Figure 1). The curved interface results from the attainment 256 of a constant mean curvature to minimize the surface energy by minimizing the surface area 257 (Bargen & Waff, 1986; Waff & Faul, 1992), whereas the flat interface is attributed to the 258 crystallographically controlled minimum interfacial energy (Yoshino et al., 2006). These two 259 kinds of interfaces produce the three types of apparent  $\theta$ : CC, FC, and FF (Figure 1b, d, and f). 260 The attainment of local interfacial energy minimization via balancing of interfacial tensions at 261 triple junctions (i.e., textural equilibration) was demonstrated by 1) the occurrence of many olivine triple junctions with angles of ~120° (Figure 1a, e) (e.g., Liu et al., 2018), 2) the cumulative 262 263 frequency curves for apparent  $\theta$  that shows good agreement with the predicted curve for an 264 equilibrated texture (Elliott et al., 1997), and 3) the concentrated normalized grain size distribution 265 on the mean grain size (Figure 1g, h; Figure S3 of the Supporting Information) (Faul, 1997; Huang 266 et al., 2021). This is supported by the fact that the experimental durations (72–211 h) were 267 sufficiently long for attaining textural equilibrium compared with those in previous studies (e.g., 268 12 h for a grain size of 10 µm at 727°C, Holness & Siklos, 2000). We measured the apparent FF 269 and FC angles with clear interfaces (e.g., angles denoted by green rectangles in Figure 1b) and 270 excluded those affected by cracks (e.g., the angle indicated by the red rectangle in Figure 1d).

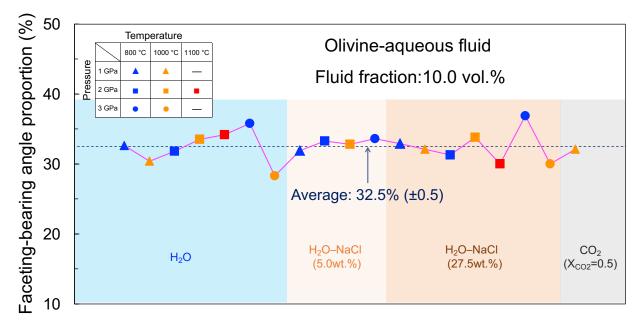


Figure 2. The faceting-bearing angle proportion measured from the olivine–fluid system. The results were calculated from 19 run products with a 10.0 vol.% fluid fraction. The average of 19 values is shown in the panel. The analytical error of each value is 0.5%. The P–T condition and fluid composition for each sample are shown in the panel.

276

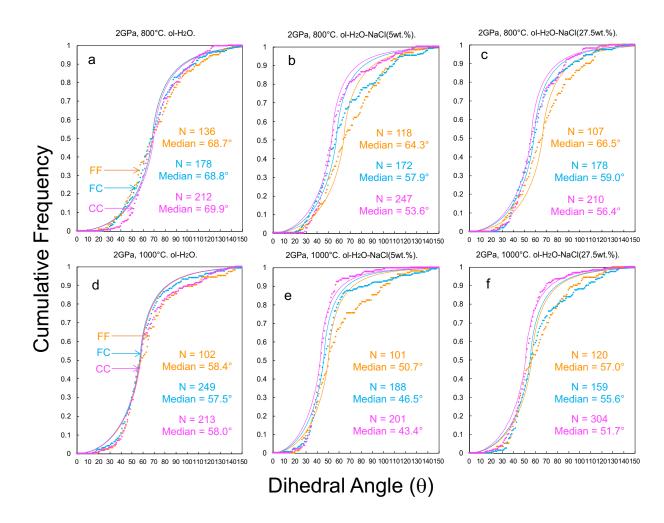
### **3.2. Proportion of faceting-bearing angles**

277 The proportion of facet-bearing angles (i.e., FF and FC angles) was evaluated from the 278 SEM images of each recovered sample using Image-J (Figure 2). A total of 5025 angles were 279 counted from the 19 samples (Table 1). To avoid the potential effect of heterogeneity in the angle-280 type distribution, we processed several images and obtained an average value for each sample. 281 Although minor fluctuation occurs, the facet-bearing angle proportion is almost constant at 282 approximately 32.5% (±0.5) without systematic P–T and fluid composition dependences. Namely, 283 the FF and FC angles constituted approximately 1/3 of the dihedral angles. It is well demonstrated 284 that the proportion of faceting is primarily controlled by  $\phi$  and that the faceting proportion 285 increases with an increase in the liquid fraction in the solid-liquid system (Watson & Lupulescu, 286 1993; Wark & Watson, 2000; Watson, 1999; Yoshino et al., 2005, 2006). Yoshino et al. (2005,

2006) reported that the faceting interface fraction in the olivine–basaltic melt system mostly fell in the range of 30.0–35.0 % with the melt fraction of 10.1–16.5 vol.%. Given an initial  $\phi$  of approximately 10.0 vol.% in our study, our calculated faceting proportion is roughly consistent with the previous research (Yoshino et al., 2006).

**3.3. Cumulative frequency of apparent dihedral angles** 

292 The representative cumulative frequencies of the measured FF and FC angles in the 293 olivine-fluid systems at 2 GPa are shown in Figure 3, in which the curves of the apparent CC 294 angles from Huang et al. (2019, 2020) are plotted for comparison. The rest of the cumulative 295 frequencies for the other conditions and histograms for all systems in this study are shown in 296 Figures S4 and S5 of the Supporting Information. For all systems investigated, the cumulative 297 frequencies mostly showed a sharp enhancement around the median  $\theta$ , and the frequency 298 distribution histogram also showed a concentrated distribution of measured angles around the 299 median value. In most cases, the CC angles sharply increase around the median angle, which is in 300 accord with the theoretical prediction for the system with one true  $\theta$ . In contrast, the FF angles 301 gradually increased around the median angle. This could indicate the expanded range of true  $\theta$ 302 owing to surface energy anisotropy on the facet-bearing angles (Laporte & Provost, 2000). In some 303 cases (Figure 3), the data largely deviated from the theoretical curve in regions of high apparent  $\theta$ . 304 This deviation was, first, attributed to the relatively small number of measured apparent FF angles. 305 Second, this can be attributed to the presence of a very large  $\theta$  associated possibly with sub-grain 306 boundaries, in which the misorientation between two adjacent grains is very small (Laporte et al., 307 1997). These two possible reasons cause the FF and FC angles to deviate more from the theoretical 308 curve than CC angle. Nevertheless, the median angles in such cases are assumed to represent the309 true value because the angles smaller than the median fit the theoretical line well.

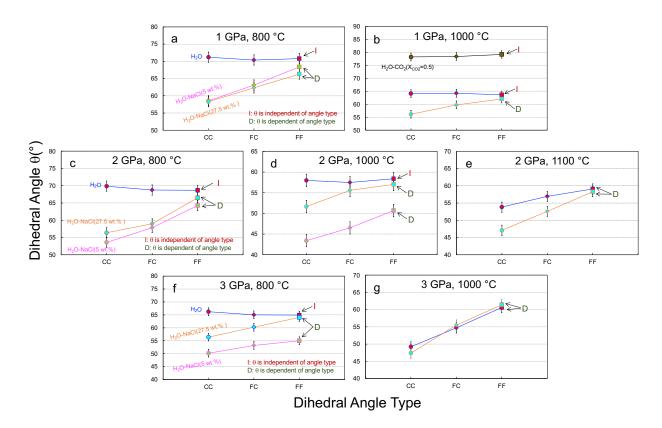


310

Figure 3. Representative cumulative frequency curves of measured apparent dihedral angles ( $\theta$ ) in olivine– fluid systems at 2 GPa and 800–1000 °C. The P–T condition and fluid composition are shown in the top of each panel. The facet-bearing angles (FC and FF) were measured in this study, while the data of the CC angles are from Huang et al. (2019, 2020). The median angle value and number (N) of the measured angles are shown for each case. The thin, solid curves represent the theoretical cumulative frequency curves obtained from the isotropic system with one true  $\theta$ . Abbreviations: FF = faceted–faceted angle, FC = faceted–curved angle, CC = curved–curved angle.

#### **318 3.4. Faceting effect on the median dihedral angle**

319 The angle-type dependence of the median  $\theta$  in the olivine-fluid systems at 1–3 GPa and 320 800-1100 °C is shown in Figure 4. Two angle-type dependencies were recognized according to 321 the fluid compositions and P–T conditions. The angle values were comparable among the CC, FC, 322 and FF in the H<sub>2</sub>O systems at relatively low P-T conditions and the H<sub>2</sub>O-CO<sub>2</sub> system. In contrast, 323 the median angle of the facet-bearing angles was higher than that of the CC angle in the H<sub>2</sub>O 324 systems under higher P–T conditions and the H<sub>2</sub>O–NaCl systems. Hereafter, the former and latter 325 systems are called as I group (angle-type independent group) and D group (angle-type dependent 326 group), respectively. The corresponding groups for each run product are summarized in Table 1. 327 Under each P–T condition, FC and CC angles showed a relatively wide variation in  $\theta$  between the 328 H<sub>2</sub>O and H<sub>2</sub>O–NaCl systems, except for the results at 3 GPa and 1100°C, whereas FF angle was 329 generally characterized by limited  $\theta$  variations. The  $\theta$  in the H<sub>2</sub>O–CO<sub>2</sub> system at 1GPa and 1000°C 330 was larger than those of the other fluid compositions irrespective of the angle-type. In the H<sub>2</sub>O-331 NaCl (5.0 wt.% NaCl) systems, the  $\theta$  values of the FF angle are similar (Figure 4a and c) or smaller by 5-10° than those in the H<sub>2</sub>O and H<sub>2</sub>O-NaCl (27.5 wt.% NaCl) systems at the same P-T 332 333 conditions (Figure 4d and f).



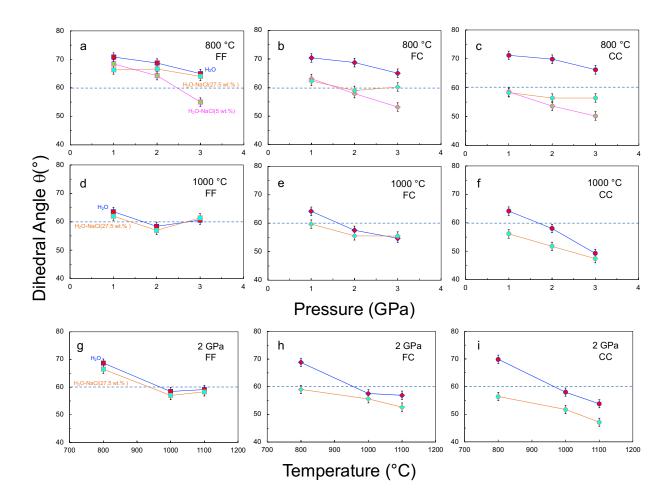
334

Figure 4. Faceting dependence of the median dihedral angles ( $\theta$ ) in the olivine–fluid system at 1–3 GPa and 800–1100 °C. Different colors in each panel denote the fluid compositions. The CC data are cited from Huang et al. (2019, 2020). The error bar of ±1.5° is shown in the panels. Abbreviations: FF = faceted–faceted angle, FC = faceted– curved angle, CC = curved–curved angle, I = angle–type independent group, D = angle–type dependent group.

## 339 **3.5.** P–T dependence of the median dihedral angles in the different fluid systems

The P–T dependence of the median FF and FC angle in the olivine–fluid systems measured in this study is shown in Figure 5, along with the CC angle reported by Huang et al. (2019). The CC angle decreases with increasing pressure and temperature (Figure 5c, f, and i), corresponding to the increase in olivine solubility under high P–T conditions (Huang et al., 2019). On the other hand, the P–T dependence of the FC and FF angles was not obvious compared to that of the CC angle. In particular, the FF angles showed a stepwise change in  $\theta$  but had similar  $\theta$  values (Figure 5a, d, and g). These findings suggest that factors other than solubility variation dominantly

347 controlled  $\theta$  of the FF angles. More discussion is given in Section 4.1 based on the EBSD results. 348 It is worth noting that the median values of the FF angle were larger than, or close to, 60° in most 349 cases.



350

Figure 5. Pressure and temperature dependence of the median dihedral angles ( $\theta$ ) in the olivine–fluid system. a–f Pressure dependence at 800–1000 °C. g–i. Temperature dependence at 2 GPa. The P–T condition and fluid composition are shown in each panel. The CC data are from Huang et al. (2019). The error bar of ±1.5° is shown along with the median angle. The blue dash line represents  $\theta$  of 60°. Abbreviations: FF = faceted–faceted angle, FC = faceted–curved angle, CC = curved–curved angle.

- 356
- 357

#### 358 **3.6.** Crystallographic orientation of olivine

359 Figure 6 shows the representative EBSD maps and pole figures obtained from the five 360 recovered samples. Figure S6 of the Supporting Information shows the rest of maps and pole 361 figures. Our samples exhibited a merely weak (010) CPO (i.e., the b axis was slightly parallel to 362 the compaction direction), regardless of the P-T conditions and fluid compositions. It was not as 363 intense as the strong CPO that developed in the deformed olivine aggregate (Pommier et al., 2018). 364 The weak CPO in this study is reasonable because the cell assembly of the piston-cylinder 365 experiments with materials softened at high P-T conditions have been developed to avoid intense 366 differential stress. Figure 7 shows the olivine grains' high-magnification orientation maps and the 367 corresponding SE images. Additional figures related to the EBSD data processing are provided in 368 Figures S7–S10 of the Supporting Information. If two adjacent grains have the same orientation, 369 the apparent  $\theta$  between the two grains will be large because of the small grain boundary energy 370 (white circles in Figure 7a1–d1). Further analyses on the crystallographic orientations of the GBPs 371 are described in the discussion.

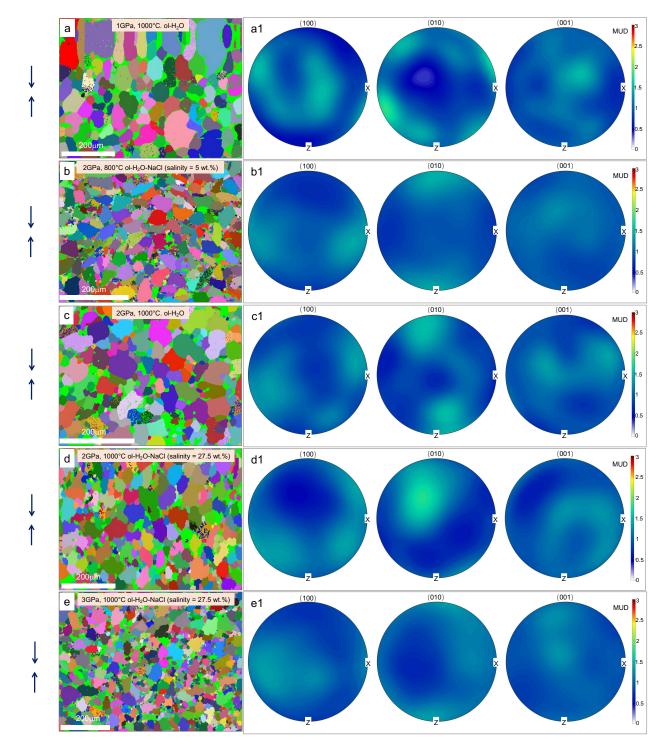
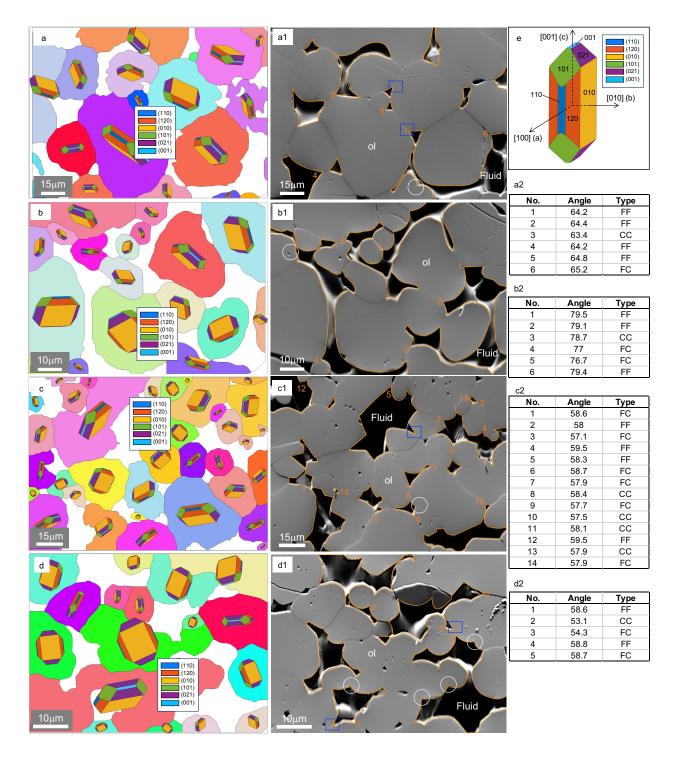




Figure 6. Representative EBSD maps and corresponding pole figures. a–e, EBSD maps of recovered olivine aggregate in olivine–fluid systems. The small points within grains are attributed to the noise, crystal defects, and fluid inclusions, which are mostly removed via denoise processes. In these systems, all grains are olivine with different orientations. a1–e1, Pole figures to show the crystallographic orientations of (100), (010), and (001) corresponding to

a-e. The intensities in the color bar are multiples of the uniform distribution (MUD). The P-T condition and fluid
 composition are shown at the top of a-e. The arrow on the left side represents the compaction direction (parallel to
 the direction of piston movement).



382 Figure 7. The olivine grains' representative 3-D crystal orientation and corresponding SE image in the olivine-383 fluid system. EBSD mappings with 3-D crystal orientation of olivine in the H<sub>2</sub>O system at 1 GPa and 1000°C (a), in 384 the H<sub>2</sub>O–CO<sub>2</sub>( $X_{(CO2)}$ =0.5) system at 1 GPa and 1000°C (b), in the H<sub>2</sub>O system at 2 GPa and 1000°C (c), and in the 385 H<sub>2</sub>O–NaCl (27.5wt%) system at 2 GPa and 1000°C (d). Colored and white areas denote olivine grains and fluid pools, 386 respectively. The 3-D crystal orientation was visualized for each grain in the high magnification images by showing 387 the oriented olivine crystal with an idealized morphology. Even though the grain configurations in **a**-**d** are slightly 388 altered during EBSD data processes such as denoising and binarization, those essentially have no effects on the 389 orientation identification. e, Crystal habit of single olivine crystal. a1-d1, Secondary electron images of recovered 390 olivine aggregates corresponding to **a**-**d**. Olivine grains have a grey color, and fluid pools are black areas that are 391 sometimes filled by resin. An orange outline visually emphasizes the interface between the olivine and fluids, and the 392 apparent angles with a sequenced number are examples of analyzed angles and corresponding grain boundaries. a2-393 d2, Dihedral angle-types, and their values in terms of FF, FC, CC corresponding to measured angles in a1-d1. These 394 angles have values close to the corresponding median  $\theta$ . The white circle in **a1–d1** represents  $\theta$  of two grains that have 395 the same orientation, leading to a larger  $\theta$  due to low misorientation of two adjacent grains. The blue square in a1, c1, 396 and d1 represents  $\theta$  that has one shared flat plane for both grain boundary and interfacial boundary. Abbreviations: ol 397 = olivine, FF = faceted–faceted angle, FC = faceted–curved angle, CC = curved–curved angle.

398 **3.7. Grain boundary plane distribution** 

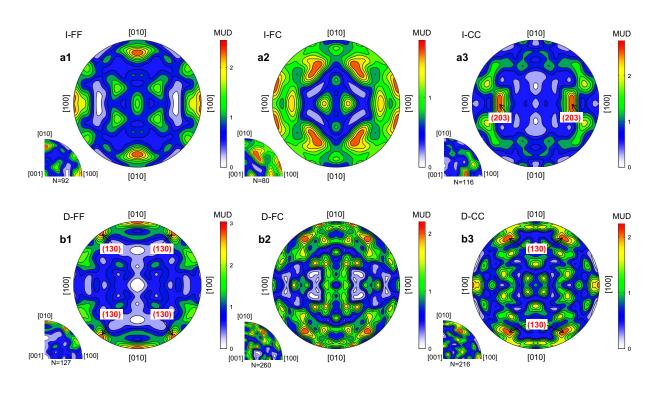
399 To clarify the crystallographic orientation of olivine on GBPs, we evaluated the GBPD for 400 the grain boundaries at the FF, FC, and CC angles on a high magnification orientation map (Figure 401 7). In general, GBPs are not vertical; they tend to incline at various degrees with respect to the 402 polished cross-section of the sample. In the two-dimensional (2-D) SEM images, we cannot 403 recognize how much these planes inclined. To mitigate this problem, we focused on the olivine-404 fluid-olivine triple junction with apparent dihedral angles lower than the median value  $+5^{\circ}$  and 405 assumed that their grain boundaries were subvertical to the polished section. For example, in the 406 system with one true  $\theta$  of 60°, 71% of the apparent dihedral angles fell within the range from 0 to

65°, in which 68% of GBPs formed an angle greater than or equal to 67° with respect to the 407 408 sectioning plane. Thus, we inferred the representative errors of ~23° in our GBPD analyses. More 409 details about the errors of the GBPD analyses are provided in the Supporting Information Note 1 410 and Figures S1-S2. At such triple junctions, we determined the crystallographic orientations of 411 olivine sharing the assumed vertical GBPs in the nine samples, including both I and D groups. In 412 the calculation, we used the Euler angles derived from the EBSD analysis and the trend of the 413 GBPs with respect to the horizontal side of the corresponding SEM image. The equivalent olivine 414 orientations obtained for each angle-type were stereologically projected in the crystal reference 415 frame, as shown in Figure 8.

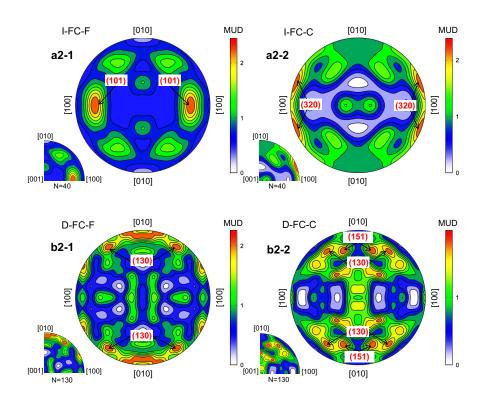
416 In the I group, the low Miller Index planes such as (100) and (010) were dominant at the 417 grain boundaries of the FF triple junctions, whereas the GBPs of the CC triple junctions were often 418 characterized by a higher Miller Index such as (203). Note that the multiples of uniform density 419 (MUD) spots indexed as (101) and (100) in Figure 8a1 are indistinguishable because of the tilt of 420 the GBP. In the D group, the GBPs of the FF triple junctions were focused on (010), but equally 421 around (130), which was interpreted as a broad concentration around (010). Although the differences in these planes were out of the possible errors of  $\pm \sim 20^{\circ}$  in our analyses, the large 422 423 variations in the true  $\theta$  at the FF angle especially in the D groups (Figure 3b, c, e, f, Section 3.3) 424 could cause the apparent broadening of the concentration around (010) beyond the assumed errors. 425 At the CC triple junctions, the high Miller Index planes around (130) is dominant, and the (100) is 426 relatively weak. The GBPD at the FC triple junctions tended to exhibit mixed characteristics of 427 the FF and CC results in both groups. Although the amount of the data was reduced, the GBPs at 428 the faceted and curved sides were separately analyzed at the FC triple junctions, as shown in Figure 429 9. As in the FF and CC junctions, the GBPs at the faceted side were dominated by low Miller Index

planes such as (101) and (010), whereas the GBPs at the curved side rather preferred the high
Miller Index planes such as (320) and (151). This indicates that the GBPDs observed in the FF and
CC triple junctions are likely true. The weak CPO developed in the run products did not
significantly affect the results.

434 Marquardt et al. (2015) found that the fluid-free olivine aggregates have a preferred (100) 435 plane of the grain boundary, which is different from the preferential appearance of the (100), 436 (010), and (101) planes on the grain boundaries at the FF angle in present study. This discrepancy 437 is most likely due to the high  $\phi$  in our study. The crystal habit of olivine grown freely in the fluid-438 rich system is characterized by the dominant (010) plane (Waff & Faul, 1992). Previous studies 439 have shown that the (010) plane of olivine has the lowest energy, followed by the (100) and (001)440 planes in fluid/melt systems (Deer et al., 2013; Gurmani et al., 2011; de Leeuw et al., 2000; Watson 441 et al., 1997), indicating a low grain boundary energy in the FF triple junctions.



443 Figure 8. Pole figure of grain boundary plane distribution. a1-a3, Grain boundary plane distribution in the I group 444 (i.e., the systems where the  $\theta$  values are independent of the angle-type). **b1-b3**, Grain boundary plane distribution in 445 the **D** group (i.e., the systems where the faceting increases  $\theta$ ). The equivalent olivine orientations obtained in the 446 analysis were rearranged into the first quadrant to make the results more evident. Subsequently, by assuming that the 447 grain boundary planes were equally distributed in each quadrant, the orientations in the first quadrant were copied in 448 the other quadrants to obtain the pole figure. The analyzed number (N) in the first quadrant is shown at the lower left 449 in each pole figure. MUD is the multiples of uniform density, shown by the color bar's intensities. The Miller Indexes 450 were marked around the high MUD. Abbreviations: FF = faceted-faceted angle, FC = faceted-curved angle, CC = 451 curved–curved angle, I = angle-type independent group, D = angle-type dependent group.



452

Figure 9. Pole figure of FC angles. Pole figure of grain boundary plane distributions (GBPDs) of faceted (a2–1 and b2–1) and curved (a2–2 and b2–2) sides at the FC triple junction for I group (i.e., the systems where the  $\theta$  values are independent of the angle–type; a2–1 and a2–2) and D group (i.e., the systems where the faceting increases  $\theta$ ; b2–1 and b2–2). The data were plotted in the same way as in Figure 8. The analyzed number (N) in the first quadrant is shown at the lower left in each pole figure. MUD represents the multiples of uniform density, which shows the

intensities in the color bar. The Miller Indexes were marked around the high MUD. Abbreviations: FC = facetedcurved angle, F = grain boundary at the faceting side, C = grain boundary at the curved side, I = angle-type independent group, D = angle-type dependent group.

# 461 **4. Discussion**

In section 4.1, we calculated the theoretical FF angle based on the ideal crystal habit and discussed the primary factor controlling  $\theta$  of FF angle. Based on our results, we further discuss the effect of the faceted interface on the fluid distribution under static conditions in Section 4.2 and the permeability anisotropy in the sheared mantle in Section 4.3.

466 **4.1 Primary factor controlling θ of FF angle** 

467 Under high P-T conditions, enhanced olivine solubility significantly decreased the 468 interfacial energy with the fluid, resulting in an obvious P–T dependence of  $\theta$  of the CC angle 469 (Huang et al., 2019, 2020). However,  $\theta$  of the FF angle was less sensitive to P–T conditions than 470 that of the CC angle, with discrete values. The variation cannot explain such characteristics in 471 interfacial energy under P-T conditions. Laporte and Provost (2000) theoretically investigated an 472 anisotropic system and showed that  $\theta$  of the FF angle was controlled by the crystallographic 473 orientation of two adjacent minerals. In this study, our GBPD analyses revealed that low Miller 474 Index planes, such as (100), (010), and (101), preferentially appeared at the GBPs of the FF angle 475 (Figures 8 and 9). Because the faceted mineral-fluid interfaces appeared to have low Miller indices, 476 the  $\theta$  value of the FF angle can be estimated from the angles between the GBPs of (100), (010), 477 (101), and the other low Miller Index olivine surfaces. To test this inference, we calculated the 478 angles between the GBPs of (100), (010), and (101) and the low-energy faceted planes of (001), 479 (011), (110), (101), (120), and (021) based on the crystal planes appearing in the ideal habit of

480	olivine crystals (Figure. 7e). The calculations were made for asymmetrical configurations and
481	symmetrical ones, because the olivine aggregate sintered in the present study showed a weak CPO
482	(i.e., the random alignment of grains). Triple junctions that one extended grain boundary plane
483	acts as one of interfaces, were considered as an extreme case (Flat face in Table 2). All the
484	calculated configurations are presented in Table 1 of the Supporting Information and the angles
485	consistent with the experimentally obtained FF $\theta$ (i.e., 50–55°, 55–65°, 65–70°, and 75–80°) are
486	summarized in Table 2. The calculated candidates cover these experimental values in Table 2. For
487	instance, the measured I–Type FF $\theta$ in the H <sub>2</sub> O–CO <sub>2</sub> system at 1 GPa and 1000°C, is 79.2°. This
488	may correspond to the calculated dihedral angles of 77.0° and 81.4° from a symmetrical triple
489	junction composed of (100)–(101) and (010)–(021) planes, respectively, or $81.5^{\circ}$ , $79.2^{\circ}$ , and $80.1^{\circ}$
490	from an asymmetrical triple junction composed of the (100)–(101) and (100)–(120), (010)–(021)
491	and (100)–(101), and (101)–(120) and (100)–(110), respectively. In fact, the GBPs can be slightly
492	inclined from the assumed planes, which may have resulted in a relatively broad distribution of
493	the apparent FF $\theta$ as shown in Figure 4. For a symmetrical triple angle configuration, the two
494	grains come in contact with the same low Miller index planes and share the lattice, whereas the
495	lattice may not be highly shared for an asymmetrical triple angle configuration. The discrepancy
496	in the grain boundary energy for these two cases is most likely negligible because the low Miller
497	index GBPs have intrinsically low-surface energy (e.g., Yoshino et al., 2006). Thus, the $\theta$ value of
498	the FC angle may be controlled by both interfacial energy and crystallographic orientation,
499	resulting in the FC angle showing intermediate characteristics between the FF and CC angles.

501 502

Table 2. Theoretical FF angle between two crystal planes

Crystal	GBCP	IBCP	ACP(°)	Calculated FF Angle(°)	Configuration	Measured FF angle(°)
C1	100	110	25.0	50.0	Symmetrical	
C2	100	110	25.0	00.0	Gymmethoar	_
C1	101	001	51.5	51.5	Flat face	50–55
C1	101	120	55.1	55.1	Flat face	-
C1	100	110	25.0	63.5	Asymmetrical	55 65
C2	100	101	38.5	03.5	Asymmetrical	55–65
C1	100	110	25.0	65.7	Asymmetrical	
C2	100	021	40.7	00.7	Asymmetrical	_
C1	101	021	66.2	66.2	Flat face	- 65–70
C1	100	110	25.0	68.0	Asymmetrical	
C2	010	120	43.0	08.0		
C1	100	110	25.0	69.8	Asymmetrical	_
C2	101	110	44.8	03.0	Asymmetrical	
C1	100	110	25.0	76.5	Asymmetrical	
C2	101	110	51.5	70.0	Asymmetrical	_
C1	100	101	38.5	77.0	Symmetrical	
C2	100	101	38.5	11.0	Symmetrical	_
C1	100	101	38.5	79.2	Asymmetrical	
C2	010	021	40.7	10.2	Asymmetrical	- 75–80
C1	100	110	25.0	80.1	Asymmetrical	75-00
C2	101	120	55.1	00.1	, loyninication	_
C1	101	021	40.7	81.4	Symmetrical	
C2	101	021	40.7	<del>-</del>	Cymmetrical	_
C1	100	101	38.5	81.5	Asymmetrical	
C2	100	120	43.0	01.0	Asymmetrical	

#### 506 507

Note: The calculated FF angle was obtained by summing the two angles between two crystal planes (ACP). This configuration shows the geometry of the calculated dihedral angle. We assume that the two crystals (crystal 1, C1; crystal 2, C2) come in contact with same (symmetrical configuration) or different (asymmetrical configuration) low– Miller index planes to form the FF angle. Additionally, we show extreme cases where one flat plane is shared for both grain boundary and interfacial boundary. GBCP, grain boundary crystal plane; IBCP, interfacial boundary crystal plane. The crystal cell parameters (a=4.7540 Å, b=10.1971 Å, and c=5.9806 Å) employed for the angle calculation were obtained from Deer et al. (2013).

# 516 4.2. Consequences of faceting on the fluid connectivity in the undeformed olivine–fluid 517 system

518 Our study demonstrated that  $\sim 1/3$  of the dihedral angles in the olivine-fluid system are 519 facet-bearing, irrespective of P-T conditions and fluid compositions (Table 1). The fluid pores 520 surrounded by faceted interfaces are difficult to connect each other even at  $\theta < 60^\circ$ , which requires 521 a threshold  $\phi$  for the establishment of a fluid network, as in the case of a CC angle > 60° (Price et 522 al., 2006). Thus, a system that includes both curved and faceted interfaces with low  $\phi$ , the bulk 523 permeability may be reduced. Huang et al. (2021) measured the electrical conductivity of the fluid-524 bearing forsterite aggregate with various  $\phi$  under the textural equilibrium states at 1 GPa and 800°C 525 in the H<sub>2</sub>O-NaCl system with 5.0 wt.% NaCl. The electrical conductivity measurements and the 526 synchrotron X-ray computed CT imaging of the post run products showed that fluid pores were 527 not interconnected at  $\phi$  of 0.51 vol.%. In contrast, they started to form the fluid network at  $\phi$  above 528 2.14 vol.%. Although the CC angle can be lower than 60° under this experimental P–T condition (Huang et al., 2019), fluid interconnection was not established at  $\phi$  below ~1.0–2.0 vol.%. This 529 530 could most likely be attributed to a presence of substantial number of faceted interfaces that 531 increases  $\phi_c$  and decreases the permeability, as pointed out by Price et al. (2006).

Toramaru and Fujii (1986) examined the melt connectivity in peridotites composed of olivine, clinopyroxene, and orthopyroxene based on a bond percolation model with the melt stability at the grain edges and corners, namely, dihedral angles. They found that the partial melt was not stable at pyroxene-dominated grain edges and that the interconnection was established when pyroxene modal composition was lower than  $\sim$ 25–20 vol.% when the grain size of olivine and pyroxenes are similar. If we assume that the faceting triple junctions hampers the fluid interconnection, an analogous discussion will be possible for the fluid connectivity in facet-bearing olivine aggregates. Given the slightly higher proportion of faceting triple junctions (28–36%; Table 1) than the above pyroxene modal composition of Toramaru and Fujii (1986), the electrical conductivity results of Huang et al. (2021), in which the fluid percolation was prohibited at  $\phi$ =0.51 vol.% but established at a small critical fraction (2.14 vol.%) seems consistent with the model prediction.

# 544 **4.3.** Preferential appearance of faceted fluid pores in the sheared mantle

Our study demonstrated that faceted olivine-fluid interfaces are preferentially 545 546 accompanied by low Miller index GBPs such as (100), (010), and (101). The extensive occurrence 547 of faceted crystallographic faces in deep-seated rocks can change the bulk permeability, elastic, 548 anelastic, and electrical properties (Waff & Faul, 1992). Waff and Faul (1992) investigated the 549 melt distribution in the texturally equilibrated olivine-melt system and found that the presence of 550 the melt film along the pervasive faceted crystal interface was making  $\theta$  much smaller, and 551 increasing the permeability of the bulk rock. However, the effect of the faceted interface on pore 552 morphology and permeability obtained from olivine-fluid systems in the present study is different 553 from that of olivine-melt system. Our results showed that the grain boundaries associated to the 554 flat interface are dry and the FF and FC  $\theta$  were comparable to or larger than the CC angles, working 555 against the establishment of fluid connectivity. This effect could be magnified under the shear 556 stress such as the corner flow in subduction zones. In the deformed olivine aggregate with strong 557 CPO, the crystal axes (i.e., a, b, and c axes) of olivine grains are aligned in specific directions 558 depending on the temperature, olivine water content, and stress state (Jung et al., 2006; Jung & 559 Karato, 2001; Karato et al., 2008; Kneller et al., 2005, 2008). This alignment may lead to low

560 Miller index planes dominating grain boundaries, thereby increasing the proportion of facet-561 bearing angles in the sheared mantle.

562 Jung and Karato (2001) examined the water-saturated olivine fabric under shear strain and 563 found that the c axis was subparallel to the shear direction, and the b axis was perpendicular to the 564 shear direction (B-type fabric). Therefore, in water-rich subduction zone, a B-type fabric is 565 expected, in which the c axis is subparallel to the subduction direction and the b axis is 566 perpendicular to the plate interface. Liu and Zhao (2017) detected the Vp anisotropy in the mantle 567 wedge beneath Japan, supporting the hypothesis that the B-type fabric is dominant in the fore-arc 568 mantle wedge. This type of grain alignment could produce abundant grain boundaries composed 569 of the (100) and (010) planes parallel to the subduction direction to form many FF angles 570 comparable to or larger than the CC angle, decreasing the permeability and electrical conductivity 571 along the subduction direction. That is to say, the presence of faceting may contribute to the 572 anisotropy of permeability and conductivity, which is important for understanding fluid flow, flux 573 melting, anomalies of the electrical conductivity, and seismic wave velocities in subduction 574 systems.

# 575 **5. Conclusion**

In this study, we determined quantitatively the effect of faceting on the olivine–fluid  $\theta$  in different fluid systems (H<sub>2</sub>O, H<sub>2</sub>O–CO<sub>2</sub> (X<sub>(CO2)</sub>= 0.5), H<sub>2</sub>O–NaCl (5.0 and 27.5 wt.% NaCl)) at 1– 3 GPa and 800–1100 °C. The results show that 1/3 of olivine–fluid  $\theta$  are faceted plane-bearing angles irrelative to the P–T conditions and fluid compositions. The  $\theta$  measurement shows that in the H<sub>2</sub>O systems at relatively low P–T conditions and in the H<sub>2</sub>O–CO<sub>2</sub> system, the facet-bearing angle values (i.e., FF, FC) are comparable to those of the CC angle. However, in the H<sub>2</sub>O systems 582 at high P–T conditions and in the H<sub>2</sub>O–NaCl systems, the facet-bearing angle values (i.e., FF, FC) 583 are larger than those of the CC angle. EBSD analyses show that the run products do not have an 584 intense crystallographic preferred orientation (CPO) corresponding to static compression 585 conditions. Strikingly, the grain boundary plane distribution (GBPD) revealed that the faceted and 586 curved interfaces at facet-bearing triple junctions have grain boundary planes (GBPs) subjected to 587 the low (e.g., (100), (010), (101)) and high (e.g., (130), (203), (320)) Miller Index faces, 588 respectively. The calculation of  $\theta$  values between two adjacent crystal planes highly reproduces 589 the measured values of the FF angles, which further supports the results of our GBPD analyses. 590 Therefore, our results suggest the importance of crystallographic orientation in determine the 591 origin of the FF angle. The presence of the FF angle and associated changes in fluid pore 592 morphology will require a high fluid fraction for establishing the fluid networks. It will further 593 lead to the permeability anisotropy and changes in geophysical characteristics, particularly for the 594 mantle wedge settings where olivine CPO is expected.

#### 595 Acknowledgments

596 This work was supported by JSPS KAKENHI Grant Nos. JP16H06348 and JP16K13903 597 awarded to M. Nakamura, JSPS Japanese-German Graduate Externship, International Joint 598 Graduate Program in Earth and Environmental Sciences, Tohoku University (GP-EES), and by 599 the Ministry of Education, Culture, Sports, Science and Technology (MEXT) of Japan under its 600 Earthquake and Volcano Hazards Observation and Research Program, the Core Research Cluster 601 of Disaster Science in Tohoku University (Designated National University), and by the 602 Tuguangchi Award for Excellent Young Scholar (E1510316) in Guangzhou Institute of 603 Geochemistry, Chinese Academy of Sciences. We provide the datasets on figshare 604 (https://doi.org/10.6084/m9.figshare.19786252.v1), including dataset of pole figure derived from 605 EBSD data and dataset of FF angle calculation, to support our research.

#### 607 **Reference**

- Allen, C., J. (1972). The role of water in the mantle of the earth: the stability of amphiboles and
   micas. 24th International Geological Congress, Montreal, 2, 231–240.
- Angiboust, S., Pettke, T., de Hoog, J. C. M., Caron, B., & Oncken, O. (2014). Channelized Fluid
  Flow and Eclogite-facies Metasomatism along the Subduction Shear Zone. *Journal of Petrology*, 55(5), 883–916. https://doi.org/10.1093/PETROLOGY/EGU010
- 613 Cmíral, M., Fitz Gerald, J. D., Faul, U. H., & Green, D. H. (1998). A close look at dihedral angles
  614 and melt geometry in olivine–basalt aggregates: A TEM study. *Contributions to Mineralogy*615 *and Petrology*, 130(3–4), 336–345. https://doi.org/10.1007/s004100050369
- de Leeuw, N. H., Parker, S. C., Catlow, C. R. A., & Price, G. D. (2000). Modelling the effect of
  water on the surface structure and stability of forsterite. *Physics and Chemistry of Minerals*,
  27, 332–341. https://doi.org/10.1007/S002690050262
- Deer et al. (2013). An Introduction to the Rock-Forming Minerals (third edition). *The Canadian Mineralogist*, *51*. https://doi.org/10.3749/CANMIN.51.4.663
- Elliott, M. T., Cheadle, M. J., & Jerram, D. A. (1997). On the identification of textural equilibrium
  in rocks using dihedral angle measurements. *Geology*, 25(4), 355–358.
  https://doi.org/10.1130/0091–7613
- Faul, U. H. (1997). Permeability of partially molten upper mantle rocks from experiments and
  percolation theory. *Journal of Geophysical Research: Solid Earth*, *102*(B5), 10299–10311.
  https://doi.org/10.1029/96jb03460
- Faul, U. H., & Fitz Gerald, J. D. (1999). Grain misorientations in partially molten olivine
  aggregates: an electron backscatter diffraction study. *Physics and Chemistry of Minerals*,
  26(3), 187–197. https://doi.org/10.1007/S002690050176
- Gurmani, S. F., Jahn, S., Brasse, H., & Schilling, F. R. (2011). Atomic scale view on partially
  molten rocks: Molecular dynamics simulations of melt-wetted olivine grain boundaries. *Journal of Geophysical Research: Solid Earth*, *116*(B12), 12209.
  https://doi.org/10.1029/2011JB008519
- Hermann, J., Spandler, C., Hack, A., & Korsakov, A. v. (2006). Aqueous fluids and hydrous melts
  in high-pressure and ultra-high pressure rocks: Implications for element transfer in
  subduction zones. *Lithos*, 92(3-4), 399-417. https://doi.org/10.1016/j.lithos.2006.03.055
- Holness, M. B., & Siklos, S. T. C. (2000). The rates and extent of textural equilibration in high–
  temperature fluid-bearing systems. *Chemical Geology*, *162*, 137–153.
  https://doi.org/10.1016/S0009–2541(99)00124–2
- Holness, M. B. (1992). Equilibrium dihedral angles in the system quartz-CO<sub>2</sub>-H<sub>2</sub>O-NaCl at
  800°C and 1–15 kbar: the effects of pressure and fluid composition on the permeability of
  quartzites. *Earth and Planetary Science Letters*, *114*(1), 171–184.
  https://doi.org/10.1016/0012-821X(92)90159-S

- Holness, M. B. (1993). Temperature and pressure dependence of quartz–aqueous fluid dihedral
  angles: the control of adsorbed H<sub>2</sub>O on the permeability of quartzites. *Earth and Planetary Science Letters*, *117*(3–4), 363–377. https://doi.org/10.1016/0012–821X(93)90090–V
- Huang, Y., Nakatani, T., Nakamura, M., & McCammon, C. (2019). Saline aqueous fluid
  circulation in mantle wedge inferred from olivine wetting properties. *Nature Communications*, *10*(1), 5557. https://doi.org/10.1038/s41467–019–13513–7
- Huang, Y., Nakatani, T., Nakamura, M., & McCammon, C. (2020). Experimental constraint on
  grain–scale fluid connectivity in subduction zones. *Earth and Planetary Science Letters*, 552,
  116610. https://doi.org/10.1016/j.epsl.2020.116610
- 653 Huang, Y., Guo, H., Nakatani, T., Uesugi, K., Nakamura, M., & Keppler, H. (2021). Electrical 654 Conductivity in Texturally Equilibrated Fluid-Bearing Forsterite Aggregates at 800°C and 655 1 GPa: Implications for the High Electrical Conductivity Anomalies in Mantle Wedges. 656 Geophysical Research: Solid Earth, 126(4), e2020JB021343. Journal of 657 https://doi.org/10.1029/2020JB021343
- Iwamori, H. (1998). Transportation of H<sub>2</sub>O and melting in subduction zones. *Earth and Planetary Science Letters*, *160*(1–2), 65–80. https://doi.org/10.1016/S0012–821X(98)00080–6
- Jung, H., & Karato, S. I. (2001). Water-Induced Fabric Transitions in Olivine. *Science*, 293(5534),
   1460–1463. https://doi.org/10.1126/SCIENCE.1062235
- Jung, H., Katayama, I., Jiang, Z., Hiraga, T., & Karato, S. (2006). Effect of water and stress on the
  lattice-preferred orientation of olivine. *Tectonophysics*, 421, 1–22.
  https://doi.org/10.1016/j.tecto.2006.02.011
- Jurewicz, S. R., & Jurewicz, A. J. G. (1986). Distribution of apparent angles on random sections
  with emphasis on dihedral angle measurements. *Journal of Geophysical Research*, *91*(B9),
  9277. https://doi.org/10.1029/jb091ib09p09277
- Karato, S. I., Jung, H., Katayama, I., & Skemer, P. (2008). Geodynamic Significance of Seismic
  Anisotropy of the Upper Mantle: New Insights from Laboratory Studies. *Annual Review of Earth* and *Planetary Sciences*, *36*, 59–95.
  https://doi.org/10.1146/ANNUREV.EARTH.36.031207.124120
- van Keken, P. E., Hacker, B. R., Syracuse, E. M., & Abers, G. A. (2011). Subduction factory: 4.
  Depth-dependent flux of H<sub>2</sub>O from subducting slabs worldwide. *Journal of Geophysical*
- 674 *Research*, *116*(B1), B01401. https://doi.org/10.1029/2010JB007922
- von Bargen, N., & Waff, H. S. (1986). Permeabilities, interfacial areas and curvatures of partially
  molten systems: Results of numerical computations of equilibrium microstructures. *Journal of Geophysical Research*, *91*(B9), 9261. https://doi.org/10.1029/JB091IB09P09261
- 678Kneller, E. A., van Keken, P. E., Karato, S. I., & Park, J. (2005). B-type olivine fabric in the679mantle wedge: Insights from high-resolution non-Newtonian subduction zone models. Earth680andPlanetaryScienceLetters,237(3-4),781-797.
- 681 https://doi.org/10.1016/J.EPSL.2005.06.049

- Kneller, E. A., Long, M. D., & van Keken, P. E. (2008). Olivine fabric transitions and shear wave
  anisotropy in the Ryukyu subduction system. *Earth and Planetary Science Letters*, 268, 268–
  282. https://doi.org/10.1016/j.epsl.2008.01.004
- Laporte, D., & Provost, A. (2000). Equilibrium geometry of a fluid phase in a polycrystalline
  aggregate with anisotropic surface energies: Dry grain boundaries. *Journal of Geophysical Research: Solid Earth*, 105(B11), 25937–25953. https://doi.org/10.1029/2000jb900256
- Laporte, D., & Watson, E. B. (1995). Experimental and theoretical constraints on melt distribution
   in crustal sources: the effect of crystalline anisotropy on melt interconnectivity. *Chemical Geology*, 124(3–4). https://doi.org/10.1016/0009–2541(95)00052–N
- Laporte, D., Rapaille, C. and Provost, A., (1997). Wetting angles, equilibrium melt geometry, and
  the permeability threshold of partially molten crustal protoliths. *In Granite: from segregation of melt to emplacement fabrics*, 31–54. Springer, Dordrecht. https://doi.org/10.1007/978–94–
  017–1717–5–3
- Liu, X., & Zhao, D. (2017). P-wave anisotropy, mantle wedge flow and olivine fabrics beneath
  Japan. *Geophysical Journal International*, 210, 1410–1431.
  https://doi.org/10.1093/GJI/GGX247
- Liu, X., Matsukage, K.N., Li, Y., Takahashi, E., Suzuki, T., & Xiong, X., (2018). Aqueous Fluid
  Connectivity in Subducting Oceanic Crust at the Mantle Transition Zone Conditions. *Journal of Geophysical Research: Solid Earth*, *123*, 6562–6573
  https://doi.org/10.1029/2018JB015973
- Marquardt, K., Rohrer, G. S., Morales, L., Rybacki, E., Marquardt, H., & Lin, B. (2015). The most
  frequent interfaces in olivine aggregates: the GBCD and its importance for grain boundary
  related processes. *Contributions to Mineralogy and Petrology*, *170*(4), 1–17.
  https://doi.org/10.1007/S00410–015–1193–9
- Médard, E., McCammon, C. A., Barr, J. A., & Grove, T. L. (2008). Oxygen fugacity, temperature
  reproducibility, and H<sub>2</sub>O contents of nominally anhydrous piston–cylinder experiments using
  graphite capsules. *American Mineralogist*, 93(11–12), 1838–1844.
  https://doi.org/10.2138/am.2008.2842
- Mibe, K., Fujii, T., & Yasuda, A. (1998). Connectivity of aqueous fluid in the Earth's upper mantle. *Geophysical Research Letters*, 25(8), 1233–1236. https://doi.org/10.1029/98GL00872
- Mibe, K., Fujii, T., & Yasuda, A. (1999). Control of the location of the volcanic front in island
  arcs by aqueous fluid connectivity in the mantle wedge. *Nature*, 401(6750), 259–262.
  https://doi.org/10.1038/45762
- Pommier, A., & Evans, R. L. (2017). Constraints on fluids in subduction zones from
  electromagnetic data. *Geosphere*, 13(4), 1026–1049. https://doi.org/10.1130/GES01473.1
- Pommier, A., Kohlstedt, D. L., Hansen, L. N., Mackwell, S., Tasaka, M., Heidelbach, F., &
  Leinenweber, K. (2018). Transport properties of olivine grain boundaries from electrical
  conductivity experiments. *Contributions to Mineralogy and Petrology*, *173*(5), 41.
  https://doi.org/10.1007/s00410-018-1468-z

Price, J. D., Wark, D. A., Watson, E. B., & Smith, A. M. (2006). Grain-scale permeabilities of
faceted polycrystalline aggregates. *Geofluids*, 6(4), 302–318. https://doi.org/10.1111/j.1468–
8123.2006.00149.x

Toramaru, A., & Fujii, N. (1986). Connectivity of melt phase in a partially molten peridotite. *Journal of Geophysical Research: Solid Earth*, 91(B9), 9239–9252.
https://doi.org/10.1029/JB091IB09P09239

- Waff, H. S., & Faul, U. H. (1992). Effects of crystalline anisotropy on fluid distribution in
  ultramafic partial melts. *Journal of Geophysical Research*, 97(B6), 9003.
  https://doi.org/10.1029/92JB00066
- Wark, D. A., & Watson, E. B. (2000). Effect of grain size on the distribution and transport of deep–
  seated fluids and melts. *Geophysical Research Letters*, 27(14), 2029–2032.
  https://doi.org/10.1029/2000GL011503
- Watson, E., & Brenan, J. M. (1987). Fluids in the lithosphere, 1. Experimentally-determined
  wetting characteristics of CO<sub>2</sub>–H<sub>2</sub>O fluids and their implications for fluid transport, host-rock
  physical properties, and fluid inclusion formation. *Earth and Planetary Science Letters*, 85(4),
  497–515. https://doi.org/10.1016/0012–821X(87)90144–0
- Watson, E., & Lupulescu, A. (1993). Aqueous fluid connectivity and chemical transport in
  clinopyroxene-rich rocks. *Earth and Planetary Science Letters*, 117(1–2), 279–294.
  https://doi.org/10.1016/0012–821X(93)90133–T
- Watson, E., (1999). Lithogic partitioning of fluids and melts. *American Mineralogist*, 84(11–12),
  1693–1710. https://doi.org/10.2138/AM–1999–11–1201
- Watson, G. W., Oliver, P. M., & Parker, S. C. (1997). Computer simulation of the structure and
  stability of forsterite surfaces. *Physics and Chemistry of Minerals*, 25(1), 70–78.
  https://doi.org/10.1007/S002690050088
- Worzewski, T., Jegen, M., Kopp, H., Brasse, H., & Taylor Castillo, W. (2011). Magnetotelluric
  image of the fluid cycle in the Costa Rican subduction zone. *Nature Geoscience*, 4(2), 108–
  111. https://doi.org/10.1038/ngeo1041
- Yoshino, T., Takei, Y., Wark, D. A., & Watson, E. B. (2005). Grain boundary wetness of texturally
  equilibrated rocks, with implications for seismic properties of the upper mantle. *Journal of Geophysical Research: Solid Earth*, *110*(B8), 1–16. https://doi.org/10.1029/2004JB003544
- Yoshino, T., Price, J. D., Wark, D. A., & Watson, E. B. (2006). Effect of faceting on pore geometry
  in texturally equilibrated rocks: implications for low permeability at low porosity. *Contributions to Mineralogy and Petrology*, 152(2), 169–186.
  https://doi.org/10.1007/S00410-006-0099-Y
- Zheng, Y. F., Chen, R. X., Xu, Z., & Zhang, S. B. (2016). The transport of water in subduction
  zones. *Science China Earth Sciences*, *59*, 651–682. https://doi.org/10.1007/s11430–015–
  5258–4
- 758

	<b><i>CAGU</i></b> PUBLICATIONS
1	
2	Journal of Geophysical Research: Solid Earth
3	Supporting Information for
4	The effect of faceting on olivine wetting properties
5 6	Yongsheng Huang <sup>1, a*</sup> , Takayuki Nakatani <sup>1, b</sup> , Sando Sawa <sup>1. c</sup> , Michihiko Nakamura <sup>1</sup> , Catherine McCammon <sup>2</sup>
7 8	1. Department of Earth Science, Graduate School of Science, Tohoku University, Aramaki– Aza–Aoba, Aoba–Ku, Sendai, Miyagi 980–8578, Japan
9	2. Bayerisches Geoinstitut, University of Bayreuth, 95440 Bayreuth, Germany
10 11	a. Current address: Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, 511 Kehua Street, Wushan, Tianhe, Guangzhou 510640, China.
12	b. Current address: Geological Survey of Japan, AIST Central 7, Higashi 1–1–1, Tsukuba,
13	Ibaraki 305–8567, Japan.
14	c. Current address: Division of Advanced Mechanical Systems Engineering, Institute of
15	Engineering, Tokyo University of Agriculture and Technology, Nakamachi 2-24-16,
16	Koganei, Tokyo, 184-0012, Japan.
17	* Corresponding author. Email: huangyongsheng@gig.ac.cn
18	
19	Contents of this file
20 21	1. Note 1. Errors of the GBPD analyses
22	2. Figures S1 to S10. The Supporting Information Figures
23	3. Table 1. List of calculated FF type dihedral angles
24 25	
26	1. Note 1. Errors of the GBPD analyses
27	Our GBPD analyses focused on the olivine–olivine–fluid triple junction with apparent dihedral
28	angles lower than the median value + 5° and assumed a vertical grain boundary plane there. Based on
29	the simple theoretical calculation, we show that grain boundary planes at such triple junctions are

30 dominantly subvertical with respect to the polished section. Although we cannot exactly determine 31 the extent of grain boundary plane tilting in the cross-sectional image, the apparent dihedral angles 32 can be used to constrain the extent of tilting statistically.

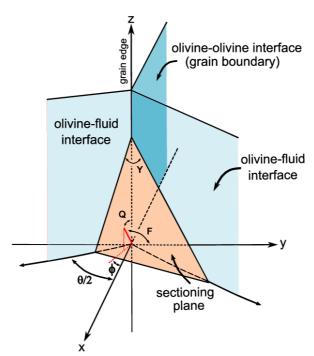
33 Following the method of Harker and Parker (1945), we can calculate e apparent dihedral 34 angle, Y on an arbitrary sectioning plane at the mineral-mineral-fluid triple junction in an isotropic 35 system with one true dihedral angle  $\theta$ . This method is identical to that used to compute the theoretical 36 cumulative frequency curve of the apparent dihedral angle, as shown in Figure 3. A schematic of the 37 triple junction with a sectioning plane is shown in Figure S1. The unit normal of the sectioning plane 38 is defined in angular coordinates Q and  $\phi$  (Q,  $\phi$  = 0–90°), and Y is a function of  $\theta$ , Q, and  $\phi$  (Harker & 39 Parker, 1945). In Figure S2a, the contours of Y for a representative  $\theta$  of 60° are shown in the sin2Q 40 versus  $\phi$  diagram. In this diagram, the area fraction of angles  $\leq$  Y corresponds to the probability that 41 the observed apparent dihedral angles become less than, or equal to, Y (Harker & Parker, 1945). The 42 apparent dihedral angles around  $\theta$  were more likely to be observed on the polished section than the 43 other angles. The median of the Y values closely corresponds to  $\theta$  (Jurewicz & Jurewicz, 1986). We 44 note that a Y smaller than  $\theta$  requires a smaller  $\phi$ , and vice versa. Increasing sin2Q (i.e., Q) tends to 45 cause Y values to deviate from the median (i.e.,  $\theta$ ).

46 The angle between the grain boundary plane and the arbitrary sectioning plane,  $F(F = 0-90^\circ)$ 47 can be calculated from their normals. F is dependent on Q and f, but independent of  $\theta$  (Figure S1). In 48 Figure S2b, the contours of F are shown in the sin2Q versus f diagram. As in the case of Y, the area 49 fraction of the angles  $\geq$  F should correspond to the probability that the angles become more than or 50 equal to F. At sin2Q (i.e., Q) = 0 or  $\phi$  = 0, the grain boundary plane is vertical (F = 90°). With increasing 51 Q and  $\phi$ , F tends to deviate from 90°: the grain boundary plane becomes more tilted. Therefore, 52 subvertical (i.e., F close to 90°) grain boundary planes can be expected at the triple junction with Y 53 smaller than the median, because such Y values can only be observed at low  $\phi$ .

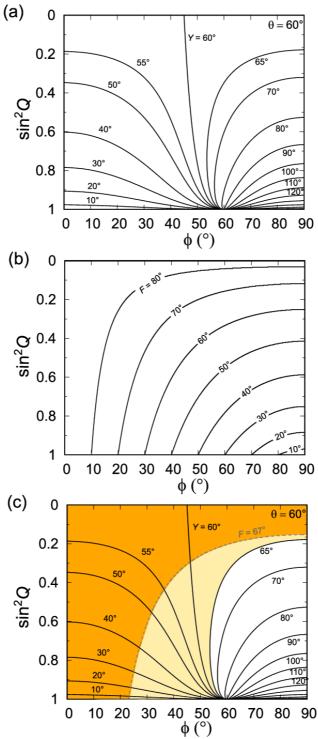
54 Combining the Y and F contours in the sin2Q versus  $\phi$  diagram allows us to compute the 55 probability of observing subvertical grain boundary planes at the triple junctions in an arbitrary Y 56 window on the polished section. We regarded the minimum deviation of F from 90°, which satisfies 57 probability of more than  $\sim$ 68%, as the representative error (1s) of our GBPD analyses. In figure S2c, 58 the area of  $F \ge 67^\circ$  in our preferred Y window from 0° to 65° (i.e., median + 5°) is shown in the sin2Q 59 versus f diagram for  $\theta$  = 60°. 71% of the apparent dihedral angles fell within the range of 0° ≤ Y ≤ 65° 60 in which 68% of grain boundary planes formed an angle greater than or equal to 67° with respect to 61 the sectioning plane. Thus, we inferred the representative errors of ~23° in our GBPD analyses. 62 Although this value slightly increased and decreased at lower and higher  $\theta$ , respectively, it was not 63 significantly dependent on  $\theta$  in the range of interest (23–24° at  $\theta$  = 50–80°). If we do not use dihedral 64 angle constraints (i.e., Y window of 0–180°), the probability of  $F \ge 67^{\circ}$  decreases to 49% and the 65 estimated error becomes 35°.

66

### 67 **2.** Figures S1 to S10. The Supporting Information Figures



- 69 Figure S1. Schematic olivine–fluid triple junction with a sectioning plane after Jurewicz and
- **Jurewicz (1986).**  $\theta$  is the true dihedral angle formed by two olivine–fluid interfaces (pale blue planes).
- 71 Y is the apparent dihedral angle observed on the sectioning plane (pale orange plane). The bold red
- 72 line represents the unit normal of the sectioning plane defined in the angular coordinates Q and  $\phi$ . F
- 73 is the angle formed by the sectioning plane and grain boundary plane (deep sky–blue plane).
- 74



**Figure S2. Results of sectioning calculation at the olivine–olivine–fluid triple junction. a,** Contours of the apparent dihedral angle, Y in the sin2Q versus  $\phi$  diagram calculated according to Harker and Parker (1945) assuming the true dihedral angle,  $\theta$  of 60°. **b**, Contours of the angle formed by grain boundary plane and sectioning plane, F in the sin2Q versus  $\phi$  diagram. **c**, The area of F ≥ 67° within the Y window of 0–65° in the sin2Q versus  $\phi$  diagram (orange). The ratio of this area to the area of Y = 0– 60° (orange + pale yellow) yields the probability of F ≥ 67° in the selected Y window.

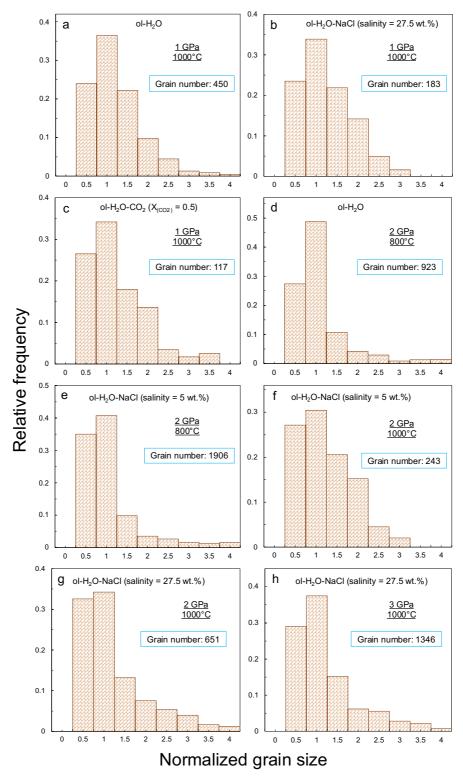
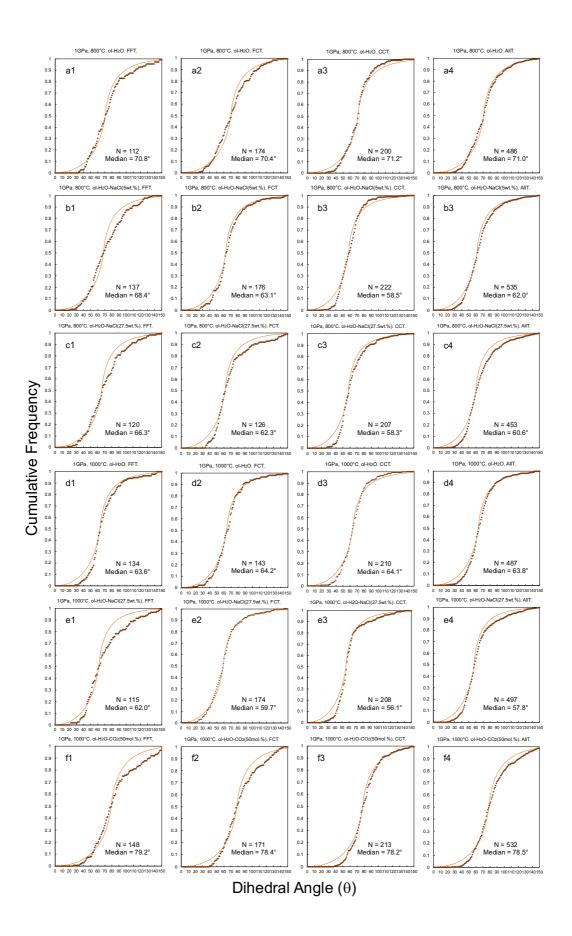
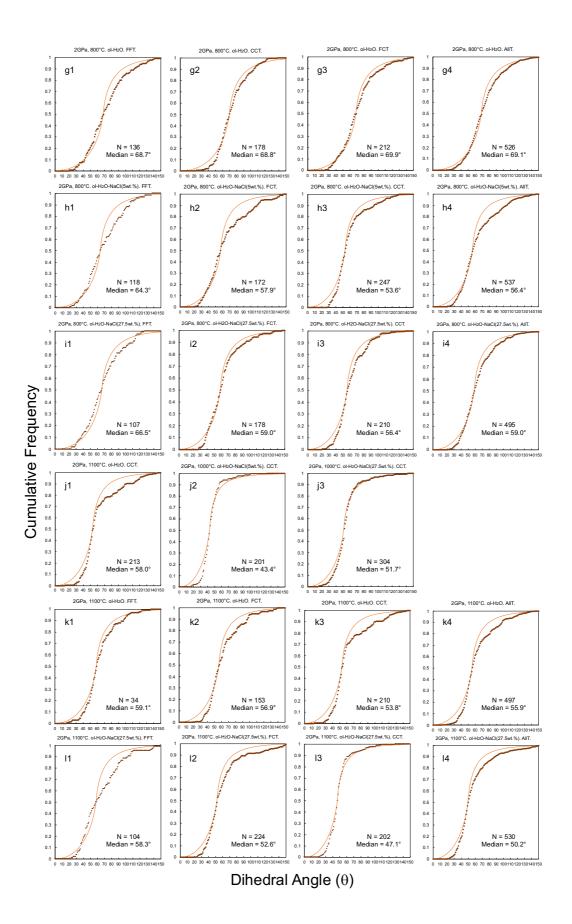




Figure S3. Histogram of the grain size distribution in the olivine-fluid system. The grain size 86 distribution was normalized by the mean grain size of the recovered sample. The grain size is 87 concentrated with a peak around the mean grain size. Abbreviation: ol = olivine.





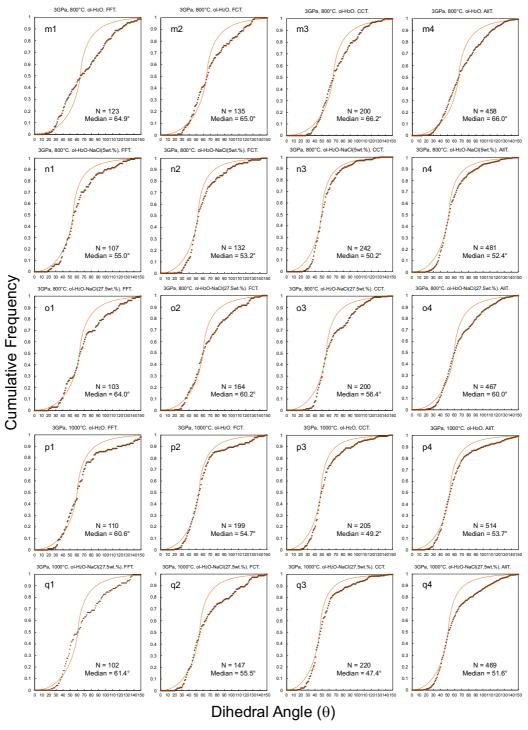


Figure S4. Cumulative frequency curves of measured apparent dihedral angles ( $\theta$ ) in olivine–fluid

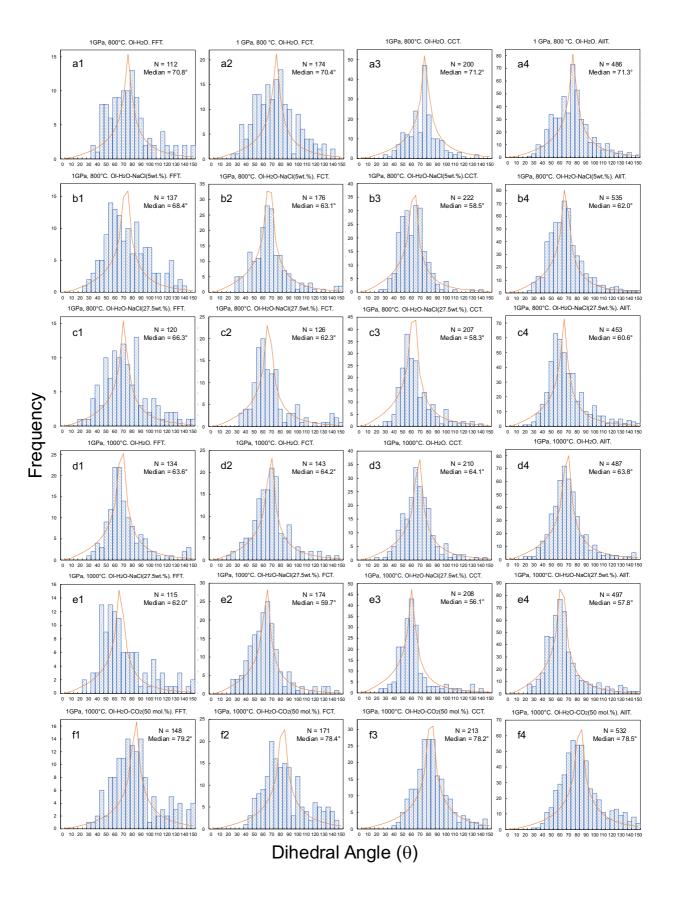
92 systems at 1–3 GPa and 800–1100 °C. The median value and number (N) of the measured angles are

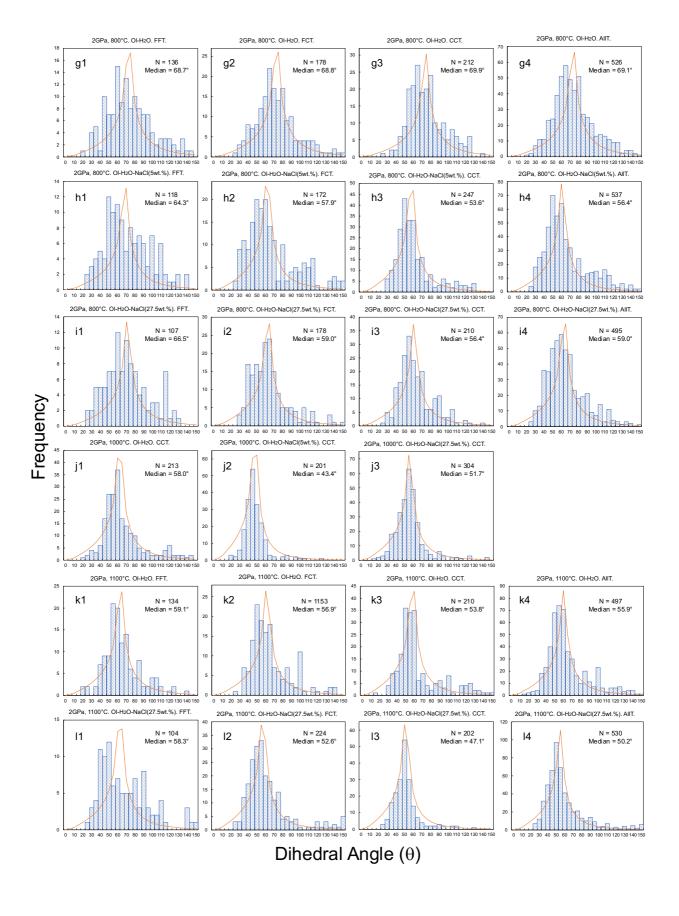
93 shown for each experimental condition. The solid lines represent the theoretical cumulative frequency 94 curves of the isotropic system with one true  $\theta$  (Jurewicz & Jurewicz, 1986). This angle is assumed to

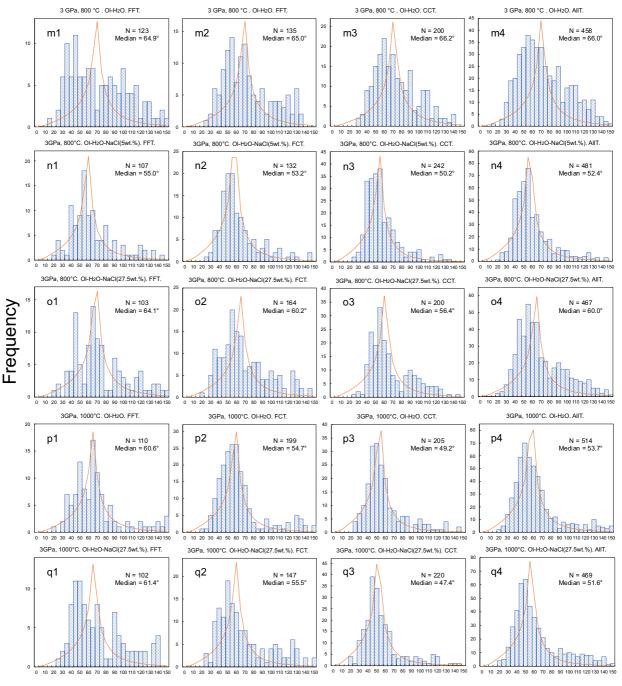
95 coincide with the obtained median value. The P–T and fluid composition are shown for each system.

96 Abbreviations: ol=olivine, FFT = faceted–faceted type angle, FCT = faceted–curved type angle, CCT=

97 curved–curved type angle, AlIT= all type of measured angle.





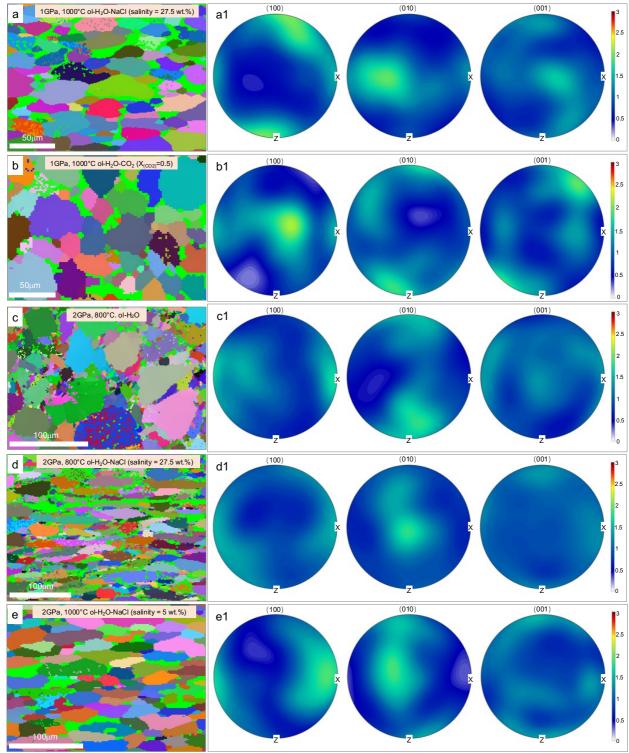


# Dihedral Angle ( $\theta$ )

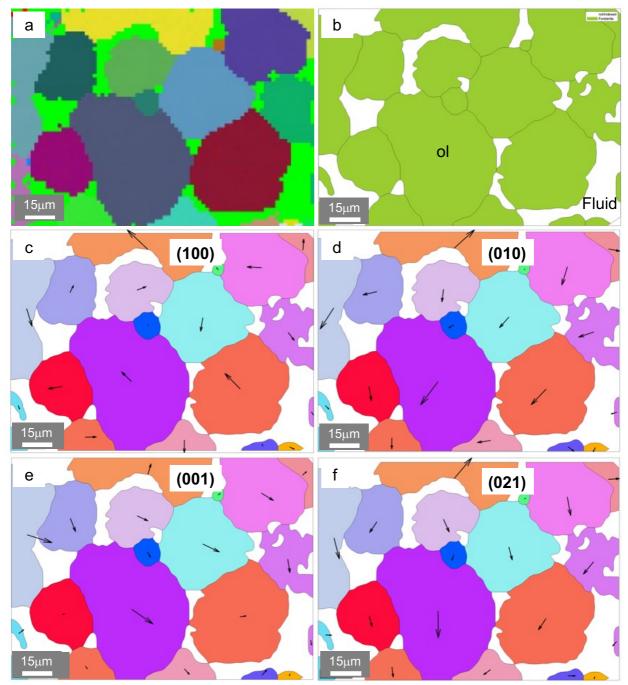
104 Figure S5. Frequency distribution histograms of measured apparent dihedral angles (θ) in olivine– 105 fluid systems at 1–3 GPa and 800–1100 °C. Theoretical distributions (orange curves) for the mono–

106 mineralic and isotropic systems are also shown in the histograms along with the median values

(Jurewicz & Jurewicz, 1986). The P–T and fluid composition are shown for each system. Abbreviations:
 ol=olivine, FFT = faceted–faceted type angle, FCT = faceted–curved type angle, CCT= curved–curved
 type angle, AllT= all type of measured angle.



112 Figure S6. EBSD maps and corresponding Pole figures under static compression conditions. a-e, 113 EBSD maps of recovered olivine aggregate in olivine–fluid systems. The small points within grain are 114 attributed to the noise, crystal defects, and fluid inclusions, which are mostly removed via denoise 115 processes while making pole figure. The crack in c was formed during the decompression, and the 116 mapping direction was calibrated during the EBSD data process. In these systems, all grains are olivine 117 with different orientations. a1-e1, Pole figures to show the crystallographic orientation of (100), 118 (010), and (001) corresponding to a-e. The intensities in the color bar are in multiples of the random 119 distribution. The P–T condition and fluid composition are shown along with the corresponding system. 120 Abbreviation: ol=olivine.



122 123

Figure S7. EBSD data and olivine orientation in the H2O system at 1 GPa and 1000°C. a, EBSD map 124 with denoise process. b, Grain image reconstruction produced by MTEX toolbox in MATLAB. c-f, 125 Special crystal orientation in 2–D grain image given by MTEX toolbox in MATLAB. The orientation is 126 shown for each system. The orientations sometime seem similar or overlapped in 2-D image, it is 127 necessary to check the 3–D spatial orientation while compilation. All grains in figure are olivine.

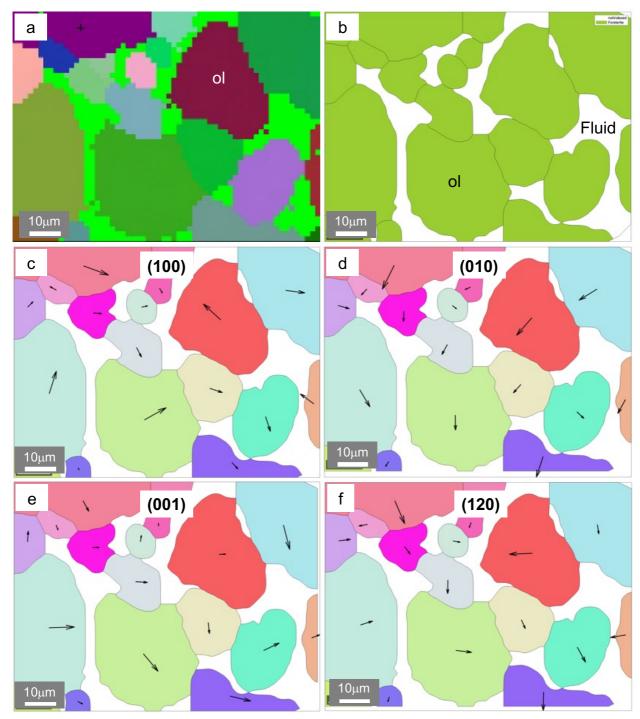
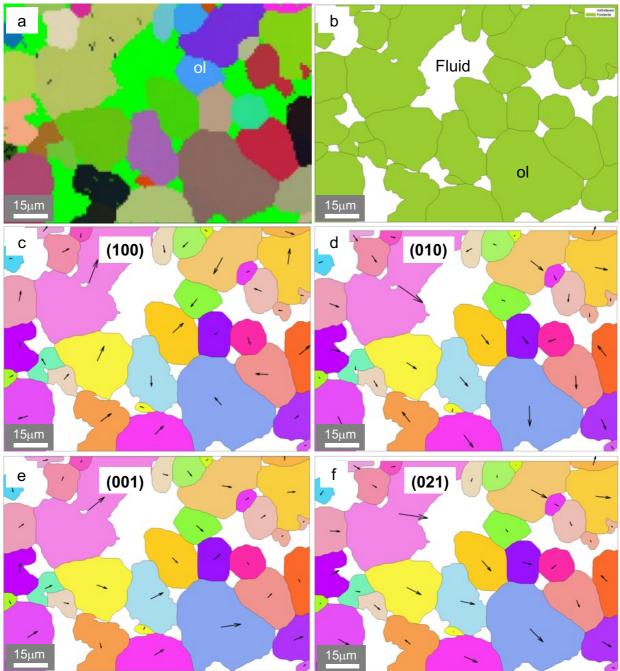


Figure S8. EBSD data and olivine orientation in the  $H_2O-CO_2(X_{(co2)=}0.5)$  system at 1 GPa and 1000°C.

a, EBSD map with denoise process. b, Grain image reconstruction produced by MTEX toolbox in
 MATLAB. c-f, Special crystal orientation in 2–D grain image given by MTEX toolbox in MATLAB. The
 orientation is shown for each system. The orientations sometime seem similar or overlapped in 2–D
 image, it is necessary to check the 3–D spatial orientation while compilation. All grains in figure are
 olivine.







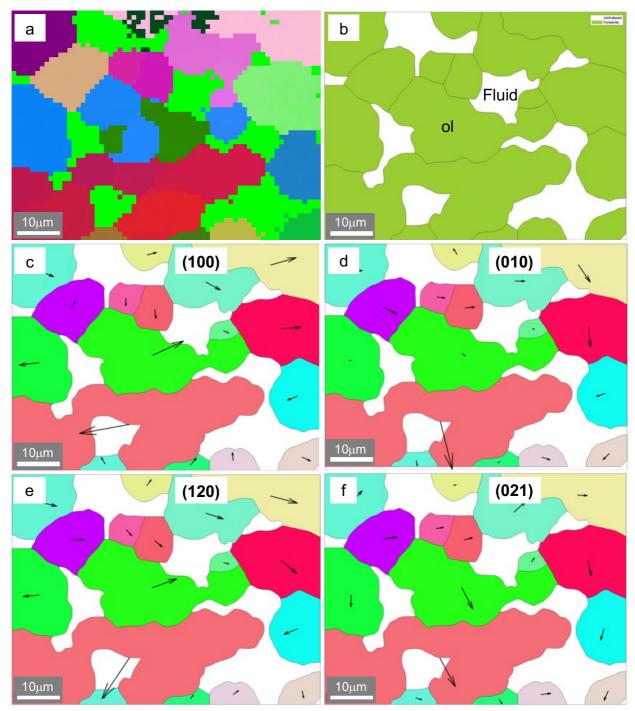


Figure S10. EBSD data and olivine orentation in the H<sub>2</sub>O–NaCl (27.5wt%) system at 2 GPa and 1000°C. a, EBSD map with denoise process. b, Grain image reconstruction produced by MTEX toolbox in MATLAB. c-f, Special crystal orientation in 2-D grain image given by MTEX toolbox in MATLAB. The orientation is shown for each system. The orientations sometime seem similar or overlapped in 2-D image, it is necessary to check the 3–D spatial orientation while compilation. All grains in figure are

- olivine.

## **3.** Table 1. List of calculated FF–type dihedral angles formed by low–Miller

Faceted Interface Grain Boundary C2 25.0 43.0 38.5 47.0 40.7 66.2 55.1 51.5 44.8 Flat face Plane 21 50.0\* 68.0 63.5 72.0 65.7 91.2 80.1 76.5 69.8 25.0 100 110 25.0 100 120 43.0 68.0\* 86.0 81.5 90.0 83.7 109.2 94.5 87.8 43.0 98.1 100 101 38.5 63.5\* 81.5\* 77.0\* 85.5 79.2 104.7 93.6 90.0 83.3 38.5 120 021 47.0 72.0 65.7\* **94.0** 87.7 47.0 40.7 010 90.0 85.5 87.7 113.2 102.1 98.5 91.8 010 40.7 83.7 79.2\* 81.4\* 106.9 95.8 92.2 85.5 40.7 66.2\* 55.1\* 91.2 80.1\* 76.5\* 101 021 66.2 109.2 104.7 113.2 106.9 132.4 121.3 117.7 111 95.8 92.2 101 101 120 001 55.1 51.5 98.1 94.5 102.1 98.5 121.3 117.7 **110.2** 106.6 106.6 **103** 93.6 99.9 90.0 96.3 51.5\* 101 110 44.8 69.8\* 87.8 83.3 91.8 85.5 111.0 99.9 96.3 89.6 44.8

### 157 Index grain boundary planes and interfaces with fluid at a triple junction

158 159

Bold represents symmetric configuration.

160 The star superscript represents calculated angles that are consistent with experimental data.