# Solar Wind Energy Input: The Primary Control Factor of Magnetotail Reconnection Site

Tsugunobu Nagai<sup>1</sup> and Iku Shinohara<sup>2</sup>

<sup>1</sup>No <sup>2</sup>Japan Aerospace Exploration Agency

November 22, 2022

#### Abstract

In this paper, we examine the solar wind energy input (expressed by  $-Vx \times Bs$ , where Vx is the x component of the solar wind velocity and Bs is the southward component of the interplanetary magnetic field IMF Bz) for an onset of magnetic reconnection in the near-Earth magnetotail. There are 41 events in which in situ observations of magnetic reconnection were made by Geotail. Magnetic reconnection in the postmidnight (premidnight) sector of the plasma sheet occurred under strong (weak) solar wind energy input conditions. Furthermore, we study temporal variations in the solar wind energy input with two different approaches using ground magnetic field observations and proton injections at geosynchronous altitude. These two analyses confirmed the preference of the postmidnight sector for the onset of magnetic reconnection under the strong solar wind energy input conditions. It is also found that the medium and weak solar wind energy input moves the onset location to earlier magnetic local times. The onset location of magnetic reconnection in the near-Earth magnetotail is controlled by the solar wind energy input through the global magnetospheric dynamics during the loading period.











1	Solar Wind Energy Input: The Primary Control Factor of Magnetotail
2	Reconnection Site
3	
4	Tsugunobu Nagai <sup>1</sup> and Iku Shinohara <sup>2</sup>
5	
6	<sup>1</sup> No.
7	<sup>2</sup> The Institute of Space and Astronautical Science, JAXA.
8	
9	Corresponding author: Tsugunobu Nagai (nagai@stp.isas.jaxa.jp)
10	
11	Key Points:
12	• The strong solar wind energy input produces favorable conditions for an
13	onset of magnetic reconnection in the postmidnight sector of the plasma
14	sheet
15	• The medium and weak solar wind energy input forms favorable conditions
16	for the onset of magnetic reconnection in earlier magnetic local times
17	• Duskside preference of the tail dynamics is caused by the solar wind
18	conditions at the Earth distance.

#### 20 Abstract

- In this paper, we examine the solar wind energy input (expressed by  $-Vx \times Bs$ ,
- 22 where Vx is the x component of the solar wind velocity and Bs is the southward
- component of the interplanetary magnetic field IMF Bz) for an onset of magnetic
- reconnection in the near-Earth magnetotail. There are 41 events in which in situ
   observations of magnetic reconnection were made by Geotail. Magnetic
- observations of magnetic reconnection were made by Geotail. Magnetic
   reconnection in the postmidnight (premidnight) sector of the plasma sheet occurred
- under strong (weak) solar wind energy input conditions. Furthermore, we study
- temporal variations in the solar wind energy input conditions. Furthermore, we study temporal variations in the solar wind energy input with two different approaches
- using ground magnetic field observations and proton injections at geosynchronous
- <sup>30</sup> altitude. These two analyses confirmed the preference of the postmidnight sector
- for the onset of magnetic reconnection under the strong solar wind energy input
- 32 conditions. It is also found that the medium and weak solar wind energy input
- moves the onset location to earlier magnetic local times. The onset location of
- magnetic reconnection in the near-Earth magnetotail is controlled by the solar wind
- <sup>35</sup> energy input through the global magnetospheric dynamics during the loading
- 36 period.

## 37 **1 Introduction**

Magnetic reconnection occurs in the near-Earth magnetotail in association 38 with the onset of magnetospheric substorms. It converts magnetic energy that is 39 previously stored in the magnetotail to plasma kinetic and thermal energy and 40 produces various dynamic phenomena in the magnetosphere-ionosphere system. It 41 produces fast plasma flows, which are observed more frequently in the 42 premidnight (dusk) sector of the plasma sheet relative to those in the postmidnight 43 (dawn) sector (e.g., Nagai et al., 1998). Satellite observations reveal that the 44 occurrence of auroral brightening is high in the 21–24 magnetic local time (MLT) 45 range (e.g., Liou et al., 2001; Frey et al., 2004). Hence, the tail fast plasma flow 46 results are consistent with the occurrence of auroral brightening corresponding to 47 an onset of substorms. 48

49

Nagai & Shinohara (2021) reported distribution of in situ magnetic
 reconnection observations in the near-Earth magnetotail (at radial distances of 20–
 30 R<sub>E</sub>) in association with substorm onsets made by Geotail over the period of
 1994–2019. Magnetic reconnection can be observed in the 21–02 MLT

- 54 (corresponding to the region of  $Y_{GSM} = +15$  to  $-10 R_E$ ). The important finding of
- 55 their study is that magnetic reconnection has a short dawn-dusk X-line
- <sup>56</sup> corresponding to 1-h MLT. There are several examples of magnetic reconnection
- 57 in the postmidnight sector of the plasma sheet, and they are mostly observed in
- highly active geomagnetic conditions and storms by Geotail (Nagai et al., 2013,

2015). The Cluster mission observed the ion diffusion region in magnetic 59 reconnection near  $Y_{GSM} = -5 R_E$  (Eastwood et al., 2010), and the postmidnight 60 event on August 21, 2002 was observed during a storm. During an intense storm 61 on May 28, 2017, the Magnetospheric Multiscale mission made in situ 62 observations of the ion diffusion region at ( $X_{GSM} = -19.3 R_E$ ,  $Y_{GSM} = -11.8 R_E$ , 63  $Z_{GSM} = 0.78 R_E$ ) (Rogers et al., 2019). Hence, magnetic reconnection can be 64 formed in the postmidnight sector of the plasma sheet. It may be conceivable that 65 magnetic reconnection observed in the postmidnight sector is a dawnward 66 extension of the X-line initially formed in the premidnight sector. However, 67 magnetic reconnection in the postmidnight sector produces a substorm current 68 system in the same sector. Any dawnward or duskward extension of the X-line is 69 not confirmed (Nagai & Shinohara, 2021). 70

71

It is reasonable to attribute an onset of magnetic reconnection to pre-72 conditions in the plasma sheet during the loading period (the growth phase of 73 substorms). In the growth phase of a substorm, magnetic field lines are transported 74 to the magnetotail as the loading process, and the magnetic field intensity increases 75 in the tail lobes (e.g., Caan et al., 1975, Shukhtina et al., 2014). The plasma sheet 76 thinning is a well-known characteristic of the growth phase (e.g., Baumjohann et 77 al., 1991, 1992). It is accompanied by an increase in the total pressure of the 78 plasma sheet, which is mainly caused by an increase in the plasma density (e.g., 79 Nagai et al., 1997). However, it is not simply caused by any pressure balance effect 80 (e.g., Sergeev et al., 2011, Saito et al., 2011, Yushkov et al., 2021), and plasma 81 transport processes should operate (e.g., Hsieh & Otto, 2015). Since it is difficult 82 to sample many plasma flows during the pure growth phase (not affected by 83 previous substorm activities) by spacecraft in the plasma sheet, the plasma flow 84 pattern producing the dynamics during the growth phase is not well explored. 85 86

It is anticipated that the plasma flow pattern in the magnetotail changes with 87 the state of the solar wind. Indeed, Nagai et al. (2005) showed that the solar wind 88 energy input is the most influential factor that determines the radial distance of 89 magnetic reconnection in the magnetotail. Here, the solar wind energy input is 90 expressed by  $-Vx \times Bs$ , where Vx is the x component of the solar wind velocity 91 and Bs is the southward component of the interplanetary magnetic field IMF Bz. 92 This value, which is expressed as VBs in this study, is similar to that of the solar 93 wind electric field; however, it can more effectively express the energy input from 94 the solar wind to the magnetosphere through dayside magnetic reconnection. We 95 examined various solar wind parameters during the magnetic reconnection events 96 studied by Nagai & Shinohara (2021). The solar wind energy input manifests the 97

most prominent characteristics for magnetic reconnection in the postmidnight
 sector.

100

In this paper, we study temporal variations of the solar wind energy input in 101 three different ways to examine its impact on the onset location of magnetic 102 reconnection in the magnetotail. First, we investigate temporal variations of the 103 solar wind energy input for 41 magnetic reconnection events observed in the 104 magnetotail by Geotail. Unfortunately, this sample number is considered small. 105 Nagai & Shinohara (2021) discovered that each magnetic reconnection forms a 106 substorm current system and that its center is located just west of the magnetic 107 reconnection site. The center of the substorm current system can be used as a proxy 108 for the meridian of the magnetic reconnection site. Second, we examine 414+211 109 positive bays observed at mid-latitude ground magnetic field stations. Nagai & 110 Shinohara (2021) also discovered that a sharp dipolarization in the magnetic field 111 with proton injections occurs at geosynchronous altitude  $(6.6 R_E)$  in the meridian 112 of the magnetic reconnection site. Finally, we examine 371 proton injections at 113 geosynchronous altitude. The results of these three different investigations show 114 that the strong solar wind energy input produces favorable conditions for the onset 115 of magnetic reconnection in the postmidnight sector. It is also found that the 116 continuous medium and weak solar wind energy input moves the onset location to 117 earlier MLTs. 118

119

The remainder of this paper is categorized as follows: Section 2 describes the data sets used in this study. Section 3 describes the main analyses. Section 4 discusses the significance of the present results for the magnetotail dynamics. Section 5 gives the conclusions.

124

#### 125 **2 Data**

The Geotail spacecraft was used to conduct magnetic field and plasma 126 observations in the plasma sheet for the period of 1998–2020. All magnetic 127 reconnection events used in this study are the same as those selected by Nagai & 128 Shinohara (2021). Any magnetic reconnection event that met the selection criteria 129 by Nagai & Shinohara (2021) was not obtained in 2020. Magnetic field data were 130 obtained through the magnetic field experiment MGF (Kokubun et al., 1994), and 131 ion and electron data were obtained through the low-energy plasma experiment 132 LEP (Mukai et al., 1994). 133

134

Ground magnetic field data consist of 1-s (from Japanese and US stations)
 and 1-min (from other stations) digital data. This study presents magnetic field data

- using H (northward) and D (eastward) components. Even when digital data are 137 presented as X- and Y-component data, they are used as H- and D-component data. 138 The data used in this study were obtained from mid- and low-latitude stations; 139 therefore, there is no significant discrepancy between these two coordinate 140 systems. The station name and ABB (abbreviation) code were used according to 141 the World Data Center for Geomagnetism, Kyoto, Data Catalog, No. 32. 142 Information about the geographic and geomagnetic locations of the ground stations 143 is shown in the Data Catalog. The geomagnetic indices AU and AL (from Kyoto 144 University before 2014 and from SuperMAG after 2015) were used. 145 146 We also used data obtained by GOES-13 at 75° W, GOES-14 at 105° W, 147 and GOES-15 at 135° W (through November 2018) and 128° W (after December 148 2018). Magnetic field data in the VDH coordinate system were used. In this 149 system, H (northward) is antiparallel to the Earth's dipole axis, D (azimuthal east) 150 is orthogonal to H and a radius vector to the satellite, and V (nearly radial outward) 151 completes the Cartesian coordinate system. Therefore, the directions of the H and 152 D components are the same as those used for ground magnetic field data. The 153 Energetic Particle Sensor MAGnetospheric Proton Detector and Energetic Particle 154 Sensor MAGnetospheric Electron Detector provided proton (>80 keV) and 155 electron (>30 keV) fluxes, respectively, in five channels. More detailed 156 information is available in Nagai et al. (2019). Energetic electron (>200 keV) 157
- 158 fluxes observed by the geosynchronous meteorological spacecraft Himawari-8
- $(140^{\circ} \text{ E})$  were also used to monitor electron injections and particle trapping
- boundary motion (Walker et al., 1976) to identify substorm onsets.
- 161

The solar wind data were obtained by the spacecraft ACE. The OMNI 1-min data were used to examine any changes in the solar wind dynamic pressure and to exclude any possible ambiguity in the solar wind traveling time from the L1 point to the Earth. We also examined the Wind and Geotail data when necessary..

## 167 3 Analyses of the solar wind energy input for magnetic reconnection

- 168 3.1 Geotail in situ magnetic reconnection observations
- 169

Nagai & Shinohara (2021) examined 56 magnetic reconnection events in the near-Earth magnetotail (at radial distances of  $20-30 R_E$ ) observed by Geotail in 1994–2019. Here, 41 of those events were used in which ACE solar wind observations were available. Figure 1 shows the distribution of the Geotail footpoint locations in MLT. This distribution is not different from that using the original 56 events (Nagai & Shinohara, 2021) and that using 71 events (Nagai et



Figure 1. (a) MLT distribution of 41 magnetic reconnection events observed by

Geotail. (b) MLT distributions of 414 positive bays in 2015–2019 and 211 positive
bays in 2020–2021 (dashed line). (c) MLT distribution of 371 proton injection

bays in 2020–2021 (dashed line). (c) MLT distribution (
events observed by GOES-15.



Figure 2. Average IMF Bz, solar wind energy input VBs, and auroral electrojet
index AL variations for Geotail magnetic reconnection events for the period from
-240 min to +120 min for 27 events in the 19–24 MLT range (a), and for 14 events
in the 00–03 MLT range (b).

al., 2015). The plasma sheet in the near-Earth magnetotail of  $Y_{GSM} = +10$  to  $-10 R_E$ 

are almost equally sampled so that the distribution can represent the occurrence

192 frequency of magnetic reconnection. Although magnetic reconnection is observed

with high occurrence in the premidnight sector, there are events in the

194 postmidnight sector.

195

Here, solar wind conditions were examined for two MLT groups: the events 196 in the premidnight sector (27 events) and those in the postmidnight sector (14 197 events). Figure 2 shows the average variations of IMF Bz, solar wind energy input 198 VBs (expressed by  $-Vx \times Bs$ , where Vx is the x component of the solar wind 199 velocity and Bs is the southward component of the interplanetary magnetic field 200 Bz), and auroral electrojet index AL. The zero epoch is the time when Geotail 201 detected tailward plasma flows leading to electron heating (the most essential 202 indication of magnetic reconnection). The AL index results show that the zero 203 epoch well corresponds to an onset of substorms. Although there are no significant 204 differences in the magnitudes of substorms in AL, clear differences exist in the 205 solar wind conditions before the substorm onset. The solar wind energy input VBs 206 is strong prior to the onset of magnetic reconnection occurring in the postmidnight 207 sector. IMF Bz is continuously southward in the time from -240 min for magnetic 208 reconnection occurring in the premidnight sector. Since the number of magnetic 209 reconnection events is small, these two findings will be further examined using the 210 two different approaches. 211

212

213 3.2 Location of the substorm current wedge

214

McPherron et al. (1973) proposed a substorm current wedge, which can be 215 used to model a substorm current system. A substorm current wedge is composed 216 of downward (into the ionosphere) field-aligned currents in the eastern part and 217 upward (from the ionosphere) field-aligned currents in the western part. These 218 field-aligned currents produce a positive bay signature in the northward 219 component, H, of the magnetic field at mid- and low-latitudes on the ground. 220 Furthermore, they cause changes in the east-west component, D, of the magnetic 221 field at mid-latitudes on the ground and in the vicinity of the geosynchronous 222 altitude in space. The eastern downward field-aligned currents produce negative D 223 variations (the western deflection), whereas the western upward field-aligned 224 currents produce positive D variations (the eastward deflection) in the Northern 225 Hemisphere. The D sign is the opposite in the Southern Hemisphere. Based on the 226 analyses by Nagai & Shinohara (2021), the zero D deflection meridian is used as 227 the center of the substorm current system in this study. 228 229



Figure 3. The solar wind energy input VBs, IMF Bz, and Super MAG auroral 231 electrojet indices AU and AL, 80-110, 110-170, and 170-250 keV proton fluxes 232 and 30-50, 50-100, 100-200, 200-350, and 350-600 keV electron fluxes at 233 GOES-14 for the period of 04:00–09:00 UT on December 28, 2018. The unit of the 234 fluxes is  $cm^{-2} s^{-1} sr^{-1} keV^{-1}$ . The magnetic field (H and D) is at GOES-14 and 235 GOES-15, and ground magnetic field data (H and D components) are from Fresno 236 (FRN), Boulder (BOU) Stennis (BSL), and Fredericksburg (FRD). The vertical 237 line on the right side corresponds to 40 nT for ground magnetic field data. Vertical 238 dashed lines show 07:17 UT and 07:42 UT onset times. 239

Figure 3 presents a clear-cut event that was observed on December 28, 2018. 242 This event demonstrates that the center of the substorm current system can form in 243 the postmidnight sector, even for a well-isolated substorm, after the absolutely 244 quiet period, not during storms. IMF Bz maintained its northward direction within 245 the period from 04:40 UT to 06:09 UT (time-shifted) and the auroral electrojet 246 activity subsided. At 06:10 UT, IMF Bz became southward, and the solar wind 247 energy input VBs exceeded more than 4.0 mV/m for one hour. There were two 248 successive onsets at 07:17 UT and 07:42 UT (vertical dashed lines in Figure 3). 249 For the first onset, GOES-14 near 00:30 MLT detected a sharp dipolarization with 250 a positive D spike in the magnetic field (Figure 3f) and proton injections (Figure 251 3d). For the second onset, GOES-15 near 23:18 MLT detected a sharp 252 dipolarization (Figure 3g) (no proton and electron data were available). Figure 3h 253 shows the ground magnetic field observations from four U.S. stations Fresno 254 (FRN), Boulder (BOU), Stennis (BSL), and Fredericksburg (FRD). A positive H 255 bay started at 07:17 UT and another started at 07:42 UT. For the first onset, the D 256 deflection was almost zero at Stennis (BSL) near 01:20 MLT, whereas for the 257 second onset the zero D meridian was located west of Fresno near 23:30 MLT. For 258 the second onset, a positive D deflection is seen at Honolulu near 21:07 MLT. At 259 the first onset, a sharp negative bay caused by the westward electrojet started at 260 Fort Churchill (00:33 MLT), and at the second onset, another sharp negative bay 261 started at Fort Simpson (22:40 MLT) in the auroral zone (not shown here). These 262 westward electrojet activities produced a two-step decrease in the AL index. The 263 ground magnetitic observations are fairly consistent with the duskward shift of the 264 sharp dipolarization region at geosynchronous altitude. In this substorm activity, 265 the first onset occurred in the postmidnight sector, while the second onset position 266 shifted duskward. The first onset in the postmidnight sector evolved into a 267 medium-sized substorm, and it was not any pseudo-onset. This event suggests that 268 it is important to consider the temporal development of substorm activity. 269

270

It is not easy to obtain many substorm onsets without any selection biases. 271 Furthermore, good coverage of ground magnetic field observations is needed to 272 determine the center of the substorm current system. We collected well-separated 273 dipolarization events in the magnetic field at geosynchronous altitude using the 274 GOES-13, GOES-14, and GOES-15 data in the period from January 2015 to 275 December 2019, in which at least two spacecraft made observations at 2-h or 4-h 276 separated longitudes. This procedure mainly covers events in the 03:00-09:00 UT 277 range, corresponding to the US nighttime. The ground stations FRN, BOU, BSL, 278 and FRD were used for most cases. Then, we identify a positive bay and its zero D 279 deflection meridian according to the method adopted by Nagai & Shinohara 280

(2021). Therefore, we use well-isolated events, and we take the first onset when 281 there are successive onsets. Considering the December 28, 2018 event (Figure 3), 282 the first onset at 07:17 UT is taken, but the second onset at 07:42 UT is not 283 selected. It is difficult to use events in highly active conditions (storm activities) 284 and the period when the solar wind dynamic pressure changes significantly, since 285 H and D variations are fairly irregular. These limitations are inevitable when the 286 ground magnetic field data are used. We can identify the onset meridian for 414 287 events. Out of these 414 events, 320 events (77%) were obtained in the time period 288 of 05-09 UT. The MLT distribution of these events is presented in Figure 1. The 289 center of the substorm current wedge forms mostly in the 22–24 MLT range, and 290 there are 91 events (22.0 %) in the postmidnight sector. The characteristics seen in 291 the MLT distribution are very similar to those in the occurrence of the auroral 292 brightening (e.g., Frey et al., 2004). For example, 82.9 % of the events are 293 distributed in the 21.0–24.5 MLT range and the median is 23.0 MLT. Hence, the 294 procedure adopted here samples reasonably representative events. 295 296 Figure 4 presents the average variations of IMF Bz, solar wind energy input 297 VBs, and auroral electrojet index AL for four MLT groups: 88 events in 19-22 298 MLT, 119 events in 22-23 MLT, 116 events in 23-24 MLT, and 91 events in 00-299 04 MLT. The AL index results show that the zero epoch well corresponds to an 300 onset of substorms. The next three characteristics emerge. 301 1. The strong solar energy input before the onset is seen for the postmidnight sector. 302 2. The solar wind energy input before the onset becomes weaker in earlier MLTs. 303 3. The IMF Bz is continuously southward in earlier MLTs, resulting in continuous 304 substorm activities. 305 306 We also collected energetic flux recovery events using the Himawari-8 data 307 to obtain a positive bay. The high-energy electron flux data can be used to identify 308 an onset of the well-isolated substorm even in other UT ranges (Nagai, 1982). We 309 identified the onset meridian for 211 events in 2020–2021. The MLT distribution 310 of these events is presented in Figure 1. Out of 211 events, 139 events (66%) were 311

obtained in the time period of 08-13 UT, and the ground magnetic stations

Honolulu (HON), Eyrewell (EYR), Canberra (CNB), Memambetsu (MMB), and

- Kakioka (KAK) were used for most cases. Figure 5 shows the average variations
   of IMF Bz, solar wind energy input VBs, and auroral electrojet index AL presented
- in the same format used in Figure 4. There are 49 events in 20–22 MLT, 60 events
- in the same format used in Figure 4. There are 49 events in 20 22 MET, 60 events in 22–23 MLT, 62 events in 23–24 MLT, and 40 events in 00–03 MLT. The three
- characteristics obtained in the previous analysis can be found in the results using
- the totally different data set.
- 320



Figure 4. Average IMF Bz, solar wind energy input VBs, and auroral electrojet index AL variations for four MLT groups of mid-latitude positive bay events for the period from -240 min to +120 min for the data set in 2015–2019.



Figure 5. Average IMF Bz, solar wind energy input VBs, and auroral electrojet index AL variations for four MLT groups of mid-latitude positive bay events for

the period from -240 min to +120 min for the data set in 2020–2021.

333

### 332 3.3 Location of proton injections

Proton (>80 keV) injections are observed with a sharp dipolarization in the 334 magnetic field with a positive D deflection indicating the upward field-aligned 335 currents (Nagai et al., 2019). GOES-15 was located at 5° geomagnetic latitude at 336 geosynchronous altitude and can minimize flux changes due to the particle 337 trapping boundary motion (Walker et al., 1976). Furthermore, the proton 338 observations by GOES-15 were nearly continuous over the period from January 339 2015 to February 2020. We collect proton injection events by GOES-15 according 340 to the method used by Nagai et al. (2019). However, when proton injections were 341 observed earlier by GOES-14 (2-h later MLT) or GOES-13 (4-h later MLT), the 342 proton flux-increase events at GOES-15 were discarded because they were events 343 far from the injection meridian. Several proton flux-increase events are associated 344 with a depression in the total magnetic field. Such events were discarded since they 345 are caused by drifting protons (e.g., Nagai, 1982). We identified 371 proton 346 injection events. Proton injection can be identified even under highly active 347 conditions. Figure 6 shows the average variations in proton flux, electron flux, and 348 magnetic fields. The average variations are fairly consistent with the results 349 presented by Nagai et al. (2019). Figure 1 shows the MLT distribution of the 350 proton injection events. Half of the proton injections occur in the 22-24 MLT 351 range (185 events). There are 90 events in the postmidnight sector and 96 events in 352 the earlier MLT region (19–22 MLT). 353

354

Figure 7 presents the average variations of IMF Bz, solar wind energy input 355 VBs, and auroral electrojet index AL for four MLT ranges. The AL index results 356 show that the zero epoch well corresponds to an onset of substorms. IMF Bz is 357 continuously southward in all four MLT ranges. This implies that the proton 358 injections can be selected even during highly active conditions, and the analysis 359 using the proton injection can complement the analysis using the ground magnetic 360 field data. The most important finding is that the results obtained using the proton 361 injection events have the three characteristics highlighted in Section 3.2. A strong 362 solar wind energy input is seen in the 00–04 MLT group. The solar wind energy 363 input becomes weaker in earlier MLTs, and the prolonged southward IMF Bz 364 period becomes prominent in earlier MLTs. 365

366

367 4. The role of the solar wind energy input

368

The three analyses presented in Section 3 demonstrate that the strong solar wind energy input provides favorable conditions for the onset of magnetic



Figure 6. Average proton flux, electron flux, and magnetic field variations derived from 371 proton injection events observed by GOES-15 for the period from -240

- min to +240 min. The unit of the fluxes is  $cm^{-2} s^{-1} sr^{-1} keV^{-1}$ .
- 375



Figure 7. Average IMF Bz, solar wind energy input VBs, and auroral electrojet

index AL variations for four MLT groups of proton injection events observed by GOES-15 for the period from -240 min to +120 min.

reconnection in the postmidnight sector. However, it is still unknown whether or

not there exists any threshold of the total amount of the solar wind energy input

leading to an onset of magnetic reconnection. Nagai et al. (2005) examined

magnetic reconnection events near the midnight meridian, and they showed that 384 magnetic reconnection forms close to (far from) the Earth in the magnetotail for 385 strong (weak) solar wind energy input conditions. Magnetic reconnection can start 386 even with a small amount of solar wind energy input. Indeed, when the substorm 387 magnitude is determined by the AL index magnitude, both small and large 388 substorms can be recorded. It is expected that an onset of magnetic reconnection 389 forms in the region where the current sheet in the plasma sheet becomes extremely 390 thin. The current sheet thinning is likely controlled by the plasma convection in the 391 near-Earth magnetotail. Under strong solar wind energy input conditions, the 392 plasma convection probably transports plasmas away from the midnight meridian 393 immediately. Under weak solar wind energy input conditions, the plasma 394 convection tends to operate more effectively in the premidnight sector and makes 395 favorable conditions there. The IMF Bz is continuously southward for the events in 396 the 19–22 MT range (Figures 4 and 5). 397

398

To test these findings, we conducted analyses for the well-isolated substorm 399 events. We selected 131 events (out of the 414 positive bay events) in which any 400 previous substorm activities are not discernible in the AL index for the 3-h period 401 before the onset. This selection rule can sample the events similar to the December 402 28, 2018 event (Figure 3). Figure 8 presents the results of the analyses for isolated 403 events and all events. With this procedure, the AL index becomes small before the 404 onset (the magnitude of AL is less than 100 nT). The IMF Bz becomes northward, 405 and the solar wind energy input becomes small for the events in the 00-04 MLT 406 range for the period of -240 min to -60 min. We envisage that magnetic 407 reconnection occurs in the postmidnight sector during storms and during highly 408 active conditions. This phenomenon is true. However, storm activity is not a 409 necessary condition for an onset of magnetic reconnection in the postmidnight 410 sector. The strong solar wind energy input leads to an onset of magnetic 411 reconnection in the postmidnight sector even after the quiet period. The results of 412 the analyses for the isolated events also show that in earlier MLTs, the duration of 413 the southward IMF Bz becomes longer, while the magnitude of the northward IMF 414 Bz becomes smaller, indicating that the weak convection produces the thinning in 415 the premidnight sector of the plasma sheet. 416

417

It is interesting to note the numbers of the events. In the 23–04 MLT range, 419 42.5 % of the events are adopted for the isolated events (43 out of 116 in the 23–24 420 MLT range and 45 out of 91 in the 00-04 MLT range). However, only 12.5 % and



- Figure 8. Average IMF Bz, solar wind energy input VBs, and auroral electrojet
- index AL variations for four MLT groups of well-isolated mid-latitude positive bay
- events (thick curves) for the period from -240 min to +120 min. The results from
- all events (shown in Figure 4) are also presented with thin curves for comparison.
- <sup>426</sup> The event number is given in the upper right corner of each panel (the number in
- 427 parenthesis is the event number for all events).

26.9 % are selected for the isolated events in the 19–23 MLT range (11 out 429 of 88 in the 19-22 MLT and 32 out of 119 in the 23-23 MLT). This result 430 indicates that when there is enough solar wind energy input, the favorable 431 condition for the onset of magnetic reconnection can be formed near the midnight 432 meridian. However, when the solar wind energy input is relatively weak, the onset 433 of magnetic reconnection is delayed and the magnetic convection produces the 434 favorable conditions for the onset of magnetic reconnection in earlier MLTs. 435 Hence, the observed high occurrence frequency of magnetic reconnection in the 436 premidnight sector is most likely caused by the fact that the medium solar wind 437 energy input is more common in the solar wind near the Earth than close to the 438 sun. 439

440

It is also interesting to compare the IMF Bz temporal variations for the 441 isolated event (thick curves of Figure 8) and those for all events (thin curves of 442 Figure 8). The IMF Bz variations for the period of -90 min to 0 min are almost 443 similar for each MLT sector, although the event numbers are quite different. This 444 indicates that the time history of the solar wind energy input contributes to form 445 the favorable conditions for the onset of magnetic reconnection in the plasma 446 sheet. There are ambiguities in the solar wind and substorm parameters used in the 447 analyses. The solar wind observations at L1 may not correctly represent the solar 448 wind conditions for the solar wind-magnetosphere interaction in some cases. The 449 separation between EYR and CNB is approximately 2-h in the geomagnetic 450 longitude and the separation between CNB and MMB is approximately 1-h. It is 451 not simple to determine the center of the current system precisely. Hence, we 452 should take into account the limitations of the present analyses. In the present 453 stage, we cannot deduce any single parameter (for example, the sum of the solar 454 wind energy input for the 1-h period before the onset) for predicting the onset 455 location. 456

457

There might be other factors for determining the MLT location of an onset 458 of magnetic reconnection. The auroral break-up position moves to earlier MLTs in 459 the Northern Hemisphere and to later MLTs in the Southern Hemisphere for IMF 460 By > 0 (e.g., Liou et al., 2001, Wang et al., 2007, Liou & Newell, 2010, Milan et 461 al., 2010). However, the magnitude of the shift is less than one hour. The influence 462 of the IMF By is asymmetric in the northern and southern tail lobes (e.g., Ohma et 463 al., 2019). Thus, it is unlikely that the IMF By controls any onset conditions 464 occurring in the equatorial plane (see also Elhawary et al., 2022). The observed 465 shift in the auroral break-up position due to IMF By is probably caused by the field 466 line mapping. 467

Substorm activities intrude into earlier MLTs (the evening sector) during 469 continuous activities (e.g., Wiens & Rostoker, 1975, Iijima & Potemra, 1978). 470 Hence, the ionospheric convection pattern has skewness, which is coupled to 471 convection flows in the magnetotail. Vasyliunas (1970), in his pioneering work, 472 proposed that the skewness is attributed to the high ionospheric conductivity of the 473 auroral oval. Yasuhara et al. (1983) and Barbosa (1984) presented conceptual 474 models, and Lotko et al. (2014) implemented a more sophisticated ionosphere-475 magnetosphere global simulation. These studies demonstrated that the rotation of 476 the convection system increases as the Hall conductivity increases. In this 477 mechanism, the ionosphere controls the magnetotail convection through the 478 ionosphere-magnetosphere coupling. This mechanism might work in medium and 479 weak continuous activities. However, this mechanism cannot produce the onset of 480 magnetic reconnection in the postmidnight sector, especially during storms. 481 Furthermore, the onset location of magnetic reconnection can shift dawnward in 482 the premidnight sector during the continuously high activity. In the double-onset 483 substorm event on September 16, 2017, studied by Nagai & Shinohara (2021), the 484 first onset occurred near the 22:00 MLT meridian at 04:32 UT, while the second 485 onset occurred near 24:00 MLT at 05:05 UT. Geotail made in situ observations of 486 magnetic reconnection only for the second onset. Hence, there was the dawnward 487 jump of the magnetic reconnection site in the two successive onsets. There are 488 other examples for the dawnward jump of the onset meridian in Geotail 489 observations, which will be investigated in a separate paper. Hence, the 490 magnetosphere-ionosphere coupling might change the magnetotail convection 491 pattern; however, it cannot become the most effective. 492

493

Lu et al. (2018) suggested that the duskward preference of magnetic 494 reconnection is of the plasma sheet origin. The Hall electric field can be induced 495 when the plasma sheet is thinning. The Hall electric field (toward the neutral sheet) 496 may transport field lines with plasmas dawnward, resulting in the further thinning 497 of the duskside plasma sheet. However, in this mechanism the thinning process can 498 be hampered in the dawnside plasma sheet. Hence, this possibility contradicts the 499 dawnward jump of the magnetic reconnection site and the magnetic reconnection 500 occurring in the postmidnight sector during storms. 501

502

503 This paper demonstrates that the strong solar wind energy input makes the 504 favorable conditions for the onset of magnetic reconnection in the postmidnight 505 sector. Sun et al. (2016) reported the dawnward preference of fast plasma flows in 506 the Mercury magnetotail, which is opposite to the duskward preference in the near-507 Earth magnetotail. In Mercury, the solar wind energy input can always be strong due to its closeness to the sun, so the finding in this study would explain theMercury situation.

510

#### 511 5 Conclusions

This study examined the behavior of the solar wind energy input (expressed 512 by  $-Vx \times Bs$ ) for magnetic reconnection in the near-Earth magnetotail in 513 association with an onset of magnetospheric substorms. The in situ observations of 514 magnetic reconnection in the magnetotail were not sufficient yet in the present 515 stage. It is difficult to deduce the precise locations of magnetic reconnection in the 516 magnetotail for comprehensive statistical studies. The event number (41 events) 517 from the Geotail observations was not enough; however, there is the prominent 518 characteristic that magnetic reconnection can form in the postmidnight sector of 519 the plasma sheet in the strong solar wind energy input conditions. Two different 520 analyses were conducted using the ground magnetic field observations and the 521 proton injections at geosynchronous altitude. The preference of the postmidnight 522 sector for the onset of magnetic reconnection during the strong solar wind energy 523 input was confirmed. Furthermore, it was found that the medium and weaker solar 524 wind energy input moves the onset location to earlier MLTs. Onset conditions of 525 magnetic reconnection are likely regulated with the global dynamics of the 526 magnetosphere. This study provides a clue to further understanding of 527 preconditions for the onset of magnetic reconnection in the near-Earth 528 malgnetotail. 529

- 530
- 531 Acknowledgments
- 532 The work at ISAS/JAXA was supported by MEXT/JSPS KAKENHI Grant
- 533 17H06140.
- 534
- 535 Data Availability Statement
- 536
- 537 The data sets used in this study are publicly available. All Geotail data are from the
- 538 Data Archives and Transmission System (DARTS) of the Institute of Space and
- Astronautical Science (ISAS) (http://www.darts.isas.jaxa.jp) and they are easily
- obtained. The authors calculated the Geotail footpoint using
- 541 https://sscweb.gsfc.nasa.gov/cgi-bin/Locator.cgi. The GOES data are obtained
- 542 from NOAA National Centers for Environmental Information
- 543 (http://www.ngdc.noaa.gov/stp/satellite/goes/index.html) and the NASA/CDAWeb

- (http://cdaweb.gsfc.nasa.gov). The ACE and Wind magnetic field and NASA 544
- OMNI data are obtained from the NASA/CDAWeb (http://cdaweb.gsfc.nasa.gov). 545
- The Himawari-8 energetic electron data are obtained from https://aer-nc-546
- web.nict.go.jp/himawari-seda/. The digital ground magnetic field data and 547
- geomagnetic indices are provided by the World Data Center for Geomagnetism at 548
- Kyoto University (http://wdc.kugi.kyoto-u.ac.jp/index.html). Information on the 549
- ground magnetic stations can be found in the Data Catalog (pdf) of WDC at Kyoto. 550
- Some of the digital magnetic field data are provided by the THEMIS website 551
- (http://themis.ssl.berkeley.edu/index.shtml) and by Super MAG 552
- (http://supermag.jhuapl.edu/). Geomagnetic indices are also provided by Super 553 MAG.
- 554
- 555
- References 556
- Barbosa, D. D. (1984). An energy principle for high-latitude electrodynamics. 557
- Journal of Geophysical Research, 89(A5), 2881–2890. 558
- https://doi.org/10.1029/JA089iA05p02881 559
- 560

Baumjohann W., Paschmann, G., Nagai, T., & Lühr, H. (1991) Superposed epoch 561 analysis of the substorm plasma sheet, J. Geophys. Res., 96(A7), 11605-11608. 562 https://doi.org/10.1029/91JA00775 563

- 564
- Baumjohann W., Paschmann, G., & Nagai, T. (1992) Thinning and thickening of 565
- the plasma sheet, J. Geophys. Res., 97(A11), 17173-17175. 566
- https://doi.org/10.1029/92JA01519 567
- 568
- Caan, M. N., McPherron, R. L., & Russell, C. T. (1975). Substorm and 569
- interplanetary magnetic field effects on the geomagnetic tail lobes. Journal of 570
- Geophysical Research, 80, 191–194. https://doi.org/10.1029/JA080i001p00191 571
- 572
- Eastwood, J. P., Phan, T. D., Øieroset, M., & Shay, M. A. (2010). Average 573
- properties of the magnetic reconnection ion diffusion region in the Earth's 574
- magnetotail: The 2001–2005 Cluster observations and comparison with 575
- simulations. Journal of Geophysical Research, 115, A08215. 576
- https://doi.org/10.1029/2009JA014962 577
- 578
- Elhawary, R., Laundal, K. M., Reistad, J. P., & Hatch, S. M. (2022). Possible 579
- ionospheric influence on substorm onset location. Geophysical Research Letters, 580
- 49, e2021GL096691. https://doi.org/10.1029/2021GL096691 581

Frey, H. U., Mende, S. B., Angelopoulos, V., & Donovan, E. F. (2004), Substorm 583 onset observations by IMAGE-FUV, J. Geophys. Res., 109, A10304. 584 https://doi.org/10.1029/2004JA010607 585 586 Hsieh, M.-S., & A. Otto (2015), Thin current sheet formation in response to the 587 loading and the depletion of magnetic flux during the substorm growth phase, J. 588 Geophys. Res. Space Physics, 120, 4264–4278. 589 https:/doi.org/10.1002/2014JA020925 590 591 Iijima, T., & Potemra, T. A., (1978) Large-scale characteristics of field-aligned 592 currents associated with substorms, J. Geophys. Res., 83(A2), 599-615. 593 https://doi.org/10.1029/JA083iA02p00599 594 595 Kokubun, S., Yamamoto, T., Acuña, M. H., Hayashi, K., Shiokawa, K., & 596 Kawano, H. (1994). The Geotail magnetic field experiment. Journal of 597 Geomagnetism and Geoelectricity, 46, 7–21. https://doi.org/10.5636/jgg.46.7 598 599 Liou, K., & Newell, P. T. (2010). On the azimuthal location of auroral breakup: 600 Hemispheric asymmetry. Geophysical Research Letters, 37, L23103. 601 https://doi.org/10.1029/2010GL045537 602 603 Liou, K., Newell, P. T., Sibeck, D. G., Meng, C. -I., Brittnacher, M., & Parks, G. 604 (2001). Observation of IMF and seasonal effects in the location of auroral 605 substorm onset. Journal of Geophysical Research: Space Physics, 106, 5799–5810. 606 https://doi.org/10.1029/2000JA003001 607 608 Lotko, W., Smith, R. H., Zhang, B., Ouellette, J. E., Brambles, O. J., & Lyon, J. G. 609 (2014). Ionospheric control of magnetotail reconnection. Science, 345, 184-187. 610 https://doi.org/10.1126/science.1252907 611 612 Lu, S., Pritchett, P. L., Angelopoulos, V., & Artemyev, A. V. (2018). Formation of 613 dawn-dusk asymmetry in Earth's Magnetotail thin current sheet: A three-614 dimensional particle-in-cell simulation. Journal of Geophysical Research: Space 615 Physics, 123, 2801–2814. https://doi.org/10.1002/2017JA025095 616 617 McPherron, R. L., Russell, C. T., & Aubry, M. P. (1973). Satellite studies of 618 magnetospheric substorms on August 15, 1968: 9. Phenomenological model for 619 substorms. Journal of Geophysical Research, 78, 3131–3149. 620 https://doi.org/10.1029/JA078i016p03131 621 622

Milan, S. E., A. Grocott, and B. Hubert (2010), A superposed epoch analysis of 623 auroral evolution during substorms: Local time of onset region, J. Geophys. Res., 624 115, A00I04. https://doi.org/10.1029/2010JA015663 625 626 Mukai, T., Machida, S., Saito, Y., Hirahara, M., Terasawa, T., Kaya, N., Obara, T., 627 Ejiri, M., & Nishida, A. (1994). The Low-Energy Particle (LEP) experiment 628 onboard the Geotail satellite. Journal of Geomagnetism and Geoelectricity, 46, 629 669-692. https://doi.org/10.5636/jgg.46.669 630 631 Nagai, T. (1982). Local time dependence of electron flux changes during 632 substorms derived from multi-satellite observations at synchronous orbit. Journal 633 of Geophysical Research, 87(A5), 3456-3468. 634 https://doi.org/10.1029/JA087iA05p03456 635 636 Nagai, T., & Shinohara, I. (2021). Dawn-dusk confinement of magnetic 637 reconnection site in the near-Earth magnetotail and its implication for 638 dipolarization and substorm current system. Journal of Geophysical Research: 639 Space Physics, 126, e2021. https://doi.org/10.1029/2021JA029691 640 641 Nagai, T., Mukai, T., Yamamoto, T., Nishida, A., Kokubun, S., & Lepping, R. P. 642 (1997). Plasma sheet pressure changes during the substorm growth phase. 643 Geophysical Research Letters, 24, 963–966. https://doi.org/10.1029/97GL00374 644 645 Nagai, T., Fujimoto, M., Saito, Y., Machida, S., Terasawa, T., Nakamura, R., 646 Yamamoto, T., Mukai, T., Nishida, A., & Kokubun, S. (1998). Structure and 647 dynamics of magnetic reconnection for substorm onsets with Geotail observations. 648 Journal of Geophysical Research: Space Physics, 103(A3), 4419–4440. 649 https://doi.org/10.1029/97JA02190 650 651 Nagai, T., Fujimoto, M., Nakamura, R., Baumjohann, W., Ieda, A., Shinohara, I., 652 Machida, S., Saito, Y., & Mukai, T. (2005), Solar wind control of the radial 653 distance of the magnetic reconnection site in the magnetotail, Journal of 654 Geophysical Research, 110, A09208. https://doi.org/10.1029/2005JA011207 655 656 Nagai, T., Shinohara, I., Zenitani, S., Nakamura, R., Nakamura, T. K. M., 657 Fujimoto, Saito, Y., & Mukai, T. (2013). Three-dimensional structure of magnetic 658 reconnection in the magnetotail from Geotail observations. Journal of Geophysical 659 Research: Space Physics, 118, 1667–1678. https://doi.org/10.1002/jgra.50247 660

- Nagai, T., Shinohara, I., & Zenitani, S. (2015). The dawn-dusk length of the X line
- in the near-Earth magnetotail: Geotail survey in 1994–2014. Journal of
- 664 Geophysical Research: Space Physics, 120, 8762–8773.
- 665 <u>https://doi.org/10.1002/2015JA021606</u>
- <sup>667</sup> Nagai, T., Shinohara, I., Singer, H. J., Rodriguez, J., & Onsager, T. G. (2019).
- <sup>668</sup> Proton and electron injection path at geosynchronous altitude. Journal of
- 669 Geophysical Research: Space Physics, 124, 4083–4103.
- 670 https://doi.org/10.1029/2018JA026281
- 671

- Ohma, A., Østgaard, N., Reistad, J. P., Tenfjord, P., Laundal, K. M., Moretto
- Jørgensen, T., Haaland, S. E., Krcelic, P., & Milan, S. (2019). Observations of
- asymmetric lobe convection for weak and strong tail activity. Journal of
- 675 Geophysical Research: Space Physics, 124, 9999–10017.
- 676 https://doi.org/10.1029/2019JA026773
- 677
- Rogers, A. J., Farrugia, C. J., & Torbert, R. B. (2019). Numerical algorithm for
- detecting ion diffusion regions in the geomagnetic tail with applications to MMS
- tail season 1 May to 30 September 2017. Journal of Geophysical Research: Space
- 681 Physics, 124, 6487–6503. <u>https://doi.org/10.1029/2018JA026429</u>
- 682
- Saito, M. H., Fairfield, D., Le, G., Hau, L.-N., Angelopoulos, V., McFadden, J. P.,
- Auster, U., Bonnell, J. W., & Larson, D. (2011). Structure, force balance, and
- evolution of incompressible cross-tail current sheet thinning. Journal of
- 686 Geophysical Research: Space Physics, 116, A10217.
- 687 <u>https://doi.org/10.1029/2011JA016654</u>
- 688
- 689 Sergeev, V., Angelopoulos, V., Kubyshkina, M., Donovan, E., Zhou, X.-Z.,
- Runov, A., Singer, H., McFadden, J., & Nakamura, R. (2011). Substorm growth
- and expansion onset as observed with ideal ground-spacecraft THEMIS coverage.
- Journal of Geophysical Research: Space Physics, 116, A00I26.
- 693 https://doi.org/10.1029/2010JA015689
- 694
- 695 Shukhtina, M. A., Dmitrieva, N. P., & Sergeev, V. A. (2014). On the conditions
- <sup>696</sup> preceding sudden magnetotail magnetic flux unloading. Geophysical Research
- 697 Letters, 41, 1093–1099. <u>https://doi.org/10.1002/2014GL059290</u>
- 698
- 699 Sun, W. J., S. Y. Fu, J. A. Slavin, J. M. Raines, Q. G. Zong, G. K. Poh, & T. H.
- 700 Zurbuchen (2016), Spatial distribution of Mercury's flux ropes and reconnection

701 702 703	fronts: MESSENGER observations, J. Geophys. Res. Space Physics, 121, 7590–7607. https://doi.org/10.1002/2016JA022787
704	Vasyliunas, V. M. (1970). Mathematical models of magnetospheric convection and
705	its coupling to the ionosphere, pp. 60–71, in Particles and Fields in the
706	Magnetosphere, Edited by McCormac, B. M. Springer, Dordrecht
707	
708	Walker, R. J., Erickson, K. N., Swanson, R. L., & Winckler, J. R. (1976).
709	Substorm-associated particle boundary motion at synchronous orbit. Journal of
710	Geophysical Research, 81, 5541–5550. <u>https://doi.org/10.1029/JA081i031p05541</u>
711	
712	Wang, H., Lühr, H., Ma, S. Y., & Frey, H. U. (2007). Interhemispheric comparison
713	of average substorm onset locations: Evidence for deviation from conjugacy.
714	Annales Geophysicae, 25, 989–999. http://www.ann-geophys.net/25/989/2007/
715	
716	Wiens, R. G., & Rostoker, G. (1975). Characteristics of the development of the
717	westward electrojet during the expansive phase of magnetospheric substorms,
718	Journal of Geophysical Research, 80, 2109-2128.
719	https://doi.org/10.1029/JA080i016p02109.
720	
721	Y asuhara, F., Greenwald, R., & Akasofu, SI. (1983). On the rotation of the polar
722	Cap potential pattern and associated polar phenomena. Journal of Geophysical
723	Research, 88(A7), 3775–3777. <u>https://doi.org/10.1029/JA0881A0/p03775</u>
724	Vushkov E V Patrukovich $\Lambda$ $\Lambda$ Artemvev $\Lambda$ V & Nakamura P (2021)
726	Thermodynamics of the magnetotail current sheet thinning Journal of Geophysical
720	Research: Space Physics 126 e2020 https://doi.org/10.1029/2020IA028969
728	Research. Space Thysics, 126, 62626. <u>https://doi.org/10.1025/2626511626565</u>
729	Figure Captions
730	
731	Figure 1. (a) MLT distribution of 41 magnetic reconnection events observed by
732	Geotail. (b) MLT distributions of 414 positive bays in 2015–2019 and 211 positive
733	bays in 2020–2021 (dashed line). (c) MLT distribution of 371 proton injection
734	events observed by GOES-15.
735	-
736	Figure 2. Average IMF Bz, solar wind energy input VBs, and auroral electrojet
737	index AL variations for Geotail magnetic reconnection events for the period from
738	-240 min to +120 min for 27 events in the 19–24 MLT range (a), and for 14 events
739	in the 00–03 MLT range (b).
740	

741	Figure 3. The solar wind energy input VBs, IMF Bz, and Super MAG auroral
742	electrojet indices AU and AL, $80-110$ , $110-170$ , and $170-250$ keV proton fluxes and $20, 50, 50, 100, 100, 200, 200, 350, and 350, 600 keV algebras fluxes at$
743	and $50-50$ , $50-100$ , $100-200$ , $200-550$ , and $550-000$ keV electron mixes at COES 14 for the period of 04:00, 00:00 UT on December 28, 2018. The unit of the
744	$OOES^{-14}$ for the period of 04.00–09.00 OT on December 28, 2018. The unit of the fluxes is $cm^{-2} s^{-1} sr^{-1} keV^{-1}$ . The magnetic field (H and D) is at COES 14 and
745	COES 15 and ground magnetic field data (H and D components) are from Erospo
740	(FPN) Boulder (BOL) Steppis (BSL) and Eredericksburg (FPD). The vertical
747	line on the right side corresponds to 40 nT for ground magnetic field data. Vertical
740	dashed lines show 07.17 UT and 07.42 UT onset times
749	dashed lines show 07.17 01 and 07.42 01 onset diffes.
751	Figure 4. Average IMF Bz, solar wind energy input VBs, and auroral electrojet
752	index AL variations for four MLT groups of mid-latitude positive bay events for
753	the period from $-240$ min to $+120$ min for the data set in 2015–2019.
754	
755	Figure 5. Average IMF Bz, solar wind energy input VBs, and auroral electrojet
756	index AL variations for four MLT groups of mid-latitude positive bay events for
757	the period from $-240$ min to $+120$ min for the data set in 2020–2021.
758	
759	Figure 6. Average proton flux, electron flux, and magnetic field variations derived
760	from 371 proton injection events observed by GOES-15 for the period from -240
761	min to +240 min. The unit of the fluxes is $cm^{-2} s^{-1} sr^{-1} keV^{-1}$ .
762	Figure 7 Average IME Bz solar wind energy input VBs, and surgral electroist
703	index AL variations for four MLT groups of proton injection events observed by
/04 765	GOES-15 for the period from $-240$ min to $\pm 120$ min
765	COLS-15 for the period from 240 min to +120 min.
767	Figure 8 Average IME Bz solar wind energy input VBs and auroral electroiet
768	index AL variations for four MLT groups of well-isolated mid-latitude positive bay
769	events (thick curves) for the period from $-240$ min to $\pm 120$ min. The results from
770	all events (shown in Figure 4) are also presented with thin curves for comparison.
771	The event number is given in the upper right corner of each panel (the number in
772	parenthesis is the event number for all events).