

Return from Dormancy: Rapid inflation and seismic unrest at Mt. Edgecumbe (L'ux Shaa) Volcano, Alaska

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Abstract

In April 2022 a seismic swarm near Mt. Edgecumbe in southeast Alaska suggests renewed activity at this dormant volcano located in a transform fault setting. Oral Tlingit history describes low-level basaltic eruptions ≈ 800 years ago. Thin rhyolitic tephra deposited 5-4 ka. We analyze synthetic aperture radar data from 2014-2022 and resolve rapid inflation up to 8.7 cm/yr beginning in August 2018. Bayesian modeling suggests a gently westward dipping sill opened 0.65 m between 7.6 km to 5.3 km depth, centered about 2-3 km east of Mt. Edgecumbe. Reanalyzed seismicity, recorded 25 km away, shows increased activity since July 2019. We hypothesize mafic magma ascent through ductile material, accumulating below a silicic seal or in a silicic reservoir, and triggering seismicity in the overburden. Cloud-native open data and workflows enabled discovery and analysis of this rapid inflation within days after going unnoticed for > 3 years.

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2 **unrest at Mt. Edgecumbe (L'úx Shaa) Volcano, Alaska**

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8 **Key Points:**

- 9 • InSAR time series reveal up to 8.7 cm/yr of inflation at Mt. Edgecumbe in satel-
10 lite line-of-sight
11 • Constant inflation begins in August 2018, microseismicity in July 2019, larger earth-
12 quakes in 2020
13 • Bayesian modeling suggests magma recharge into a 14 deg westward dipping sill
14 ranging from 5.3-7.6 km depth

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Abstract

In April 2022 a seismic swarm near Mt. Edgecumbe in southeast Alaska suggests renewed activity at this dormant volcano located in a transform fault setting. Oral Tlingit history describes low-level basaltic eruptions ≈ 800 years ago. Thin rhyolitic tephra deposited 5-4 ka. We analyze synthetic aperture radar data from 2014-2022 and resolve rapid inflation up to 8.7 cm/yr beginning in August 2018. Bayesian modeling suggests a gently westward dipping sill opened 0.65 m between 7.6 km to 5.3 km depth, centered about 2-3 km east of Mt. Edgecumbe. Reanalyzed seismicity, recorded 25 km away, shows increased activity since July 2019. We hypothesize mafic magma ascent through ductile material, accumulating below a silicic seal or in a silicic reservoir, and triggering seismicity in the overburden. Cloud-native open data and workflows enabled discovery and analysis of this rapid inflation within days after going unnoticed for >3 years.

Plain Language Summary

In April 2022 a cluster of earthquakes detected near Mt. Edgecumbe in Southeast Alaska suggested magmatic activity. As the volcano is near the major Queen-Charlotte Fairweather fault that separates the Pacific and North American plates, similar earthquakes were previously assumed to be tectonic in nature. Since magmatic activity at volcanoes is often accompanied by ground deformation, we analyzed available satellite radar data going back to fall of 2014. This reveals crustal motion toward the satellite of up to 8.7 cm/yr starting in August 2018, which totalled up to 27 cm by November 2021. Modeling of the deformation suggests that magma is intruding between 5-8 km depth into a gently dipping tabular body. Since the nearest seismograph is 25 km away, we record only large seismicity, which increases in July 2019, but we perhaps missed some lower energy seismicity due to this distance. We believe that we are observing magma rising through somewhat malleable crust into an existing magmatic system and that the observed earthquakes are created as the overlying rock adjusts to the increased magmatic pressure. The observed activity is rare, especially in similar tectonic settings, and presents an opportunity to better understand the reactivation of dormant volcanoes.

1 Introduction

Located in Southeast Alaska, the home to about 73,000 people, or 10% of the state's population, Mount Edgecumbe (L'ux Shaa in Tlingit) is part of the Mt. Edgecumbe Volcanic Field on Kruzof Island on the west side of Sitka Sound (Figure 1). The eastern shore of Sitka Sound, about 25 km away from Mount Edgecumbe, is home to the almost 8,500 residents of Sitka, and destination of projected 200,000-400,000 cruise passengers in 2022 (Woolsey, 2021). Like most communities in Southeast Alaska, including the state capital Juneau, Sitka is only accessible via air and sea, as no connections to the highway system exist, and thus quite vulnerable to hazards due to earthquakes, tsunamis and volcanic eruptions, especially volcanic ash.

On April 11, 2022, a Sitka resident noted that the openly available Alaska Earthquake Center's (AEC) location for an M2.1 earthquake was under Mt. Edgecumbe and inquired with Alaska Volcano Observatory (AVO) contacts whether this earthquake is related to the volcano or, as previously suggested for seismicity in the region, purely tectonic. As this dormant volcano is not monitored with a dedicated geophysical instrument network, the review of seismicity relies on the broadly distributed regional seismic network, and the related location uncertainties. Activity recorded by the closest seismograph SIT, 25 km away in Sitka, indicated that a seismic swarm started about 02:00 am AKDT on April 11, 2022. Earthquakes located by the AEC range in magnitude from 1.0-2.7 and appear broadly distributed to the NE of Mt. Edgecumbe near the eastern rim of Crater Ridge (Figure 1, red circles in inset). This swarm, with estimated depths ranging from about 1-12 km, was preceded by a burst of located seismic activity in 2020 that began

65 with an M3.0 earthquake on 2020-01-02 and lasted until about mid-2021. With the caveat
66 of kilometer-scale uncertainties in the published locations, the depth ranges of both swarms
67 appear similar, but epicenter locations seem more tightly clustered and slightly shifted
68 southward in 2022 than before (compare blue and red circles in Figure 1, inset). No other
69 seismicity in this region has been large enough to allow for location estimation through
70 the regional network before 2020 while seismicity of similar magnitudes has been con-
71 sistently recorded offshore (Figure 1).

72 To investigate the origin of these earthquakes and their relation to the volcano, we
73 use Sentinel-1 (Torres et al., 2012) synthetic aperture radar (SAR) observations to per-
74 form cloud-based interferometric SAR (InSAR) time series analysis on Mt. Edgecumbe.
75 Using seasonal data that omits winter acquisitions from 2014-10-17 until 2021-11-27, we
76 find rapid inflation that began abruptly in August 2018. Ground motion in the line of
77 sight of the satellite reaches velocities of 8.7 cm/yr. We perform a Bayesian inversion of
78 the InSAR velocity field, which suggests magma accumulation in a dipping body at about
79 5-7 km depth beneath the volcano. Our reanalysis of microseismicity on the closest seis-
80 mograph in Sitka indicates an uptick in recorded seismicity beginning in July 2019.

81 2 Background

82 Perhaps popularly best known for a very elaborate 1974 April Fools joke, when the
83 local Oliver “Porky” Bickar set 70 old tires on fire in Mt. Edgecumbe’s crater, making
84 Sitka’s residents believe that the volcano was erupting, the Mt. Edgecumbe volcanic field
85 has a rich history of volcanism, including felsic lavas and pyroclastic flows. Vents are lin-
86 early organized, trending northeast from the southwestern tip of Kruzof Island to Mud
87 Bay in the northeast, perhaps aligned with a crustal fissure (Riehle et al., 1992). The
88 most recent report of activity is part of Tlingit oral history and reports “a mountain blink-
89 ing, spouting fire and smoke” 800-900 years ago (Kitka, n.d.). During the Holocene, for
90 5.8 and 4.2 ka, Riehle et al. (1992) report two eruptions that created thin rhyolite lay-
91 ers on Kruzof Island, which they attribute to Crater Ridge. Prior to that, from about
92 14.6-13.1 ka Praetorius et al. (2016) document increased volcanism with a recurrence of
93 about 1.5 events per century that they attribute to last glacial maximum ice loss as re-
94 ported for, e.g., Iceland (e.g., Jull & McKenzie, 1996; Maclennan et al., 2002; Pagli &
95 Sigmundsson, 2008) and California (Jelinek et al., 2004). This activity ranged from basaltic
96 to late erupting rhyolite, the latter also found in adjacent marine sediment cores in tephra-
97 fall and pyroclastic flow deposits (Addison et al., 2010). Based on the eruptive sequence
98 of basalt and andesite throughout the existence of the field and late erupting rhyolite,
99 Riehle et al. (1992) suggest basaltic underplating created the system and produced a strat-
100 ified magma chamber with a low density high viscosity rhyolite cap.

101 In addition to volcanic activity and deformation induced by Last Glacial Maximum
102 deglaciation, this region also experiences more recent Glacial Isostatic Adjustment (GIA).
103 Larsen et al. (2005) attribute the rapid uplift since the late 1700s to viscoelastic relax-
104 ation of the mantle following the retreat of Little Ice Age glaciers. Ongoing glacial re-
105 treat continues to induce an elastic and viscoelastic response of the Earth measured by
106 GPS (Larsen et al., 2004; J. L. Elliott et al., 2010). Hu and Freymueller (2019) link on-
107 going variations in this uplift to accelerations in ice mass loss between the 1990s and 2012,
108 as documented elsewhere (Compton et al., 2015).

109 Tectonically, the Edgecumbe volcanic field is located near the Fairweather-Queen
110 Charlotte transform fault (45-46 mm/yr slip rate) to the west, accommodating most of
111 the Pacific-North America relative plate motions (J. Elliott & Freymueller, 2020a). As
112 such, this fault system experiences frequent large strike slip earthquakes, such as the 1972
113 M7.6 Sitka earthquake just south of Kruzof Island (Schell & Ruff, 1989). This region is
114 located on the slightly northward translating (about 3 mm/yr) Baranof block (J. Elliott
115 & Freymueller, 2020a). While this region has generally been understood to also exhibit

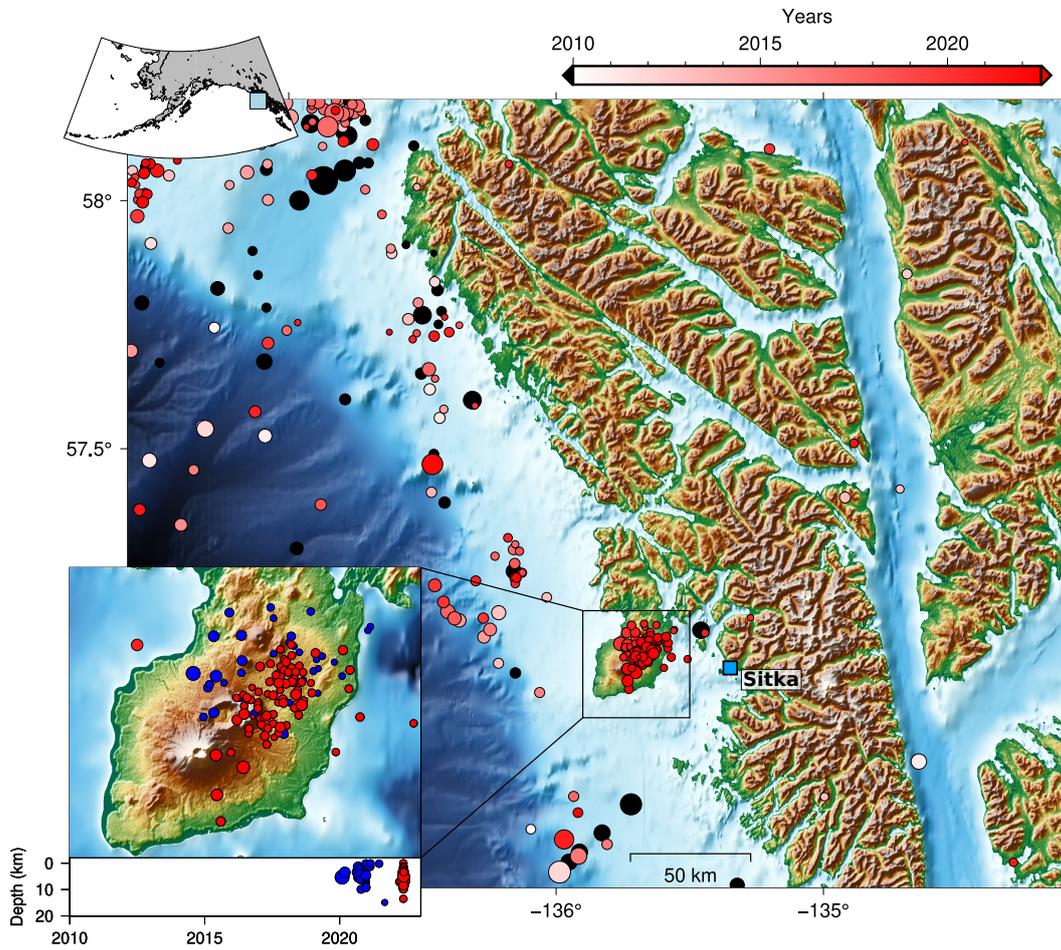


Figure 1. Mt. Edgecumbe overview map. Main map shows topography and regional seismicity from 1990-01-01 until 2022-04-18 (earthquake locations from Alaska Earthquake Center retrieved via USGS ANSS catalog). Earthquakes before 2010 are colored black, from 2010 onward according to colorbar. Blue square shows location of town of Sitka. Bottom inset shows close up of Kruzof Island seismicity mainly NE of Mt. Edgecumbe with a time-depth plot illustrating the absence of earthquakes in that region before 2020, blue earthquakes occurred before 2022, red earthquakes in 2022. Top inset shows location of main map in SE Alaska.

116 small rates of fault normal convergence (e.g., J. Elliott & Freymueller, 2020a), more re-
 117 cent data and modeling allow for small amounts of extension in this area (J. Elliott &
 118 Freymueller, 2020b), perhaps supporting the volcanism here.

119 3 Data & Methods

120 We use synthetic aperture radar data acquired between 2014-10-17 and 2021-11-
 121 27 by the European Space Agency’s Sentinel-1 mission (Torres et al., 2012) on ascend-
 122 ing path 174 and frame 402. Recently, Sentinel-1B was tasked with the observation of
 123 this scene, but it remains unobserved since December 2021 due to an anomaly in the satel-
 124 lite’s power system (ESA, 2022). Other scenes do not capture the full island, but more
 125 importantly, they do not reach back in time until 2014. Hence, we utilize this observa-
 126 tion geometry in this project.

127 We use the on-demand InSAR capabilities embedded in NASA Alaska Satellite Fa-
 128 cility’s (ASF) Hybrid Pluggable Processing Pipeline (HyP3) (Hogenson et al., 2016; John-
 129 ston et al., 2022) that employs GAMMA (Wegmüller et al., 2016) to process 126 SAR
 130 scenes into 405 interferograms with a temporal baseline of 48 days within one season.
 131 We assume decorrelation due to winter weather to be large and hence omit data between
 132 December 1 and March 1. We bridge each winter season with at least 3 InSAR pairs.
 133 The first image in this stack, acquired on 2014-10-17, serves as the reference image that
 134 all remaining repeat acquisitions are aligned to. We use a MintPy (Yunjun et al., 2019)
 135 workflow run within ASF’s cloud-based OpenSARLab platform (Hogenson et al., 2021;
 136 Meyer et al., 2021) to perform SBAS-based (Berardino et al., 2002) time series analy-
 137 sis of the interferogram stack. Our workflow rejects interferograms below 0.7 temporal
 138 and 0.7 spatial average coherence, and performs atmospheric correction using the ERA-
 139 5 atmospheric models (Hersbach et al., 2020). Residual topographic errors were estimated
 140 and corrected using the approach by Fattahi and Amelung (2013), and phase unwrap-
 141 ping error correction was performed using a phase bridging method in which coherent
 142 components with the smallest distance from each other are assumed connected and a smooth
 143 phase variation across them is enforced (Chen & Zebker, 2002).

144 To isolate and model volcanic signals, we reference the resulting InSAR phase time
 145 series to a location on Kruzof Island, but away from the volcanic deformation (Figure 2).
 146 We then downsample the cumulative LOS displacements using a variance-based algo-
 147 rithm (Agram & Jolivet, 2012) to appropriately weight near and far field pixels accord-
 148 ing to the information they carry. The downsampled displacement field with 3613 pix-
 149 els (Figure 4) is then used as input to a Markov-Chain Monte-Carlo (MCMC) based Bayesian
 150 inversion approach that utilizes Metropolis-Hastings sampling (e.g., Haario et al., 2001;
 151 Aster et al., 2019) implemented in the pymc (Patil et al., 2010) library. We assume a uni-
 152 form prior distribution of the input parameter space and generate 1 million sample so-
 153 lutions after a 10^5 sample burn-in period. To reduce autocorrelation, we only use every
 154 10^4 sample to estimate the posterior distribution. We test several analytical solutions
 155 that relate subsurface pressure changes to surface deformation including a pressure point
 156 source (Mogi, 1958), a more general spheroid (Yang et al., 1988), and circular (Fialko
 157 et al., 2001) and rectangular (Okada, 1992) sill-like sources. For the circular sill model
 158 (Fialko et al., 2001) we generated only 10^5 sample solutions following a 10^4 sample burn-
 159 in period using only every 10^2 sample to estimate the solution due to the computational
 160 cost of calculating each solution. The overall preferred geometry is determined based on
 161 misfit minimization between observations and the most-likely solution, as well as statisti-
 162 cal methods (F-test, e.g., Press et al. (2007)) to ensure the fit to the data for any given
 163 model improves appropriately with the additional degrees of freedom introduced with
 164 each model.

165 In order to investigate the possibility of contemporaneous seismic unrest not cap-
 166 tured in existing seismic catalogs, we implement the REDPy detection algorithm (Hotovec-

167 Ellis & Jeffries, 2016) to analyze data recorded at the closest seismic station SIT. REDPy
 168 utilizes a standard short-time average/long-time average (STA/LTA) algorithm (Allen,
 169 1978) to trigger on events, before cross-correlating events with successive triggers in an
 170 attempt to group them into families. REDPy has been used to quantify the temporal
 171 evolution of repeating seismicity at several other volcanic settings (e.g., Hotovec-Ellis et
 172 al., 2022; Wellik et al., 2021). To focus on swarm-like clusters of events, we opt for a re-
 173 relaxed STA/LTA trigger ratio of 4 on data filtered between 1 and 10 Hz, and a high cross-
 174 correlation threshold of 0.7 to group up events. We also manually inspect each cluster
 175 and remove all clusters that are dominated by false detections (details in Text S1).

176 4 Results

177 The InSAR time series analysis (sample wrapped phase observations in Figures S1-
 178 S4; all re-wrapped phase observations in Figure S5) resolves a circular feature of motion
 179 toward the satellite (uplift and horizontal motion toward the satellite), perhaps slightly
 180 elongated along the NS axis, with a diameter of approximately 17 km (Figure 2). In the
 181 satellite line-of-sight, the deformation appears centered about 2-3 km to the east of Mt.
 182 Edgumbe (Figure 2a). The LOS shortening reaches a maximum cumulative displace-
 183 ment of about 27 cm, tapering off roughly symmetrically toward the edges of the deforma-
 184 tion signal. Figures 2(1-4) show time series of 4 representative pixels marked in Fig-
 185 ure 2a. Except for possible seasonal deformation, likely driven by seasonal hydrologic or
 186 cryospheric effects (e.g., Heki, 2001; Grapenthin et al., 2006; Borsa et al., 2014), the time
 187 series show no substantial deformation from October 2014 until August 2018. At that
 188 time, effectively linear deformation began at rates of up to 8.7 cm/yr in the satellite line-
 189 of-sight. Supplementary Figure S6 shows the time series of displacement accumulation
 190 in map view for all 126 time steps.

191 The seismicity panel of Figure 2 presents the results of the seismic reanalysis from
 192 2018 until November 2021, the time span covered by the InSAR analysis; data until April
 193 2022 are included in Figure S7. The monthly event detections (gray bars) do not notably
 194 increase above background rates until July 2019. They taper off by the beginning of 2020
 195 only to increase again in the middle of 2020 from which point on seismic activity appears
 196 elevated until the end of 2021, as reflected in the gradient of the cumulative seismicity.
 197 The most recent swarm in April 2022 leads to the most detections compared to any prior
 198 time (Figure S7).

199 We model high coherence (> 0.3) components of the variance-based downsampled
 200 InSAR velocities from August 2018 until November 2021 with classic analytical solutions
 201 for subsurface volume changes. Figure 3 shows the results of the MCMC inversion in the
 202 form of posterior probability density functions (PDF) for each parameter of both sill mod-
 203 els along the respective diagonal (point source and spheroid results are in Figures S8-9).
 204 The off-diagonal panels show the combined PDFs of parameter pairs where circular pat-
 205 tern indicate largely independent parameters and deviations from circular to elliptical
 206 or linear highlight trade-offs between parameters. We find some trade-offs between rect-
 207 angular sill width, length, opening and depth where, for instance, greater depth would
 208 allow for slightly larger opening, and smaller area of the sill would require slightly larger
 209 opening. Overall, the ranges of these trade-offs are small (100s of meters for spatial pa-
 210 rameters, about 1 cm for opening). Distances of the horizontal location of the source (x ,
 211 y) are given with respect to the center of the displacement field at 57.0843 deg N and
 212 135.7196 deg W, a location on Crater Ridge (see Figure 4).

213 We use the most likely values for each parameter of each model to predict the sur-
 214 face deformation in the satellite LOS and show the results for the sill models in Figure 4
 215 (point source and spheroid results are in Figures S8-S9). The top two rows of this fig-
 216 ure show the variance-based downsampled deformation results, upon which the param-
 217 eter estimation and model selection is based, and the bottom two rows show full reso-

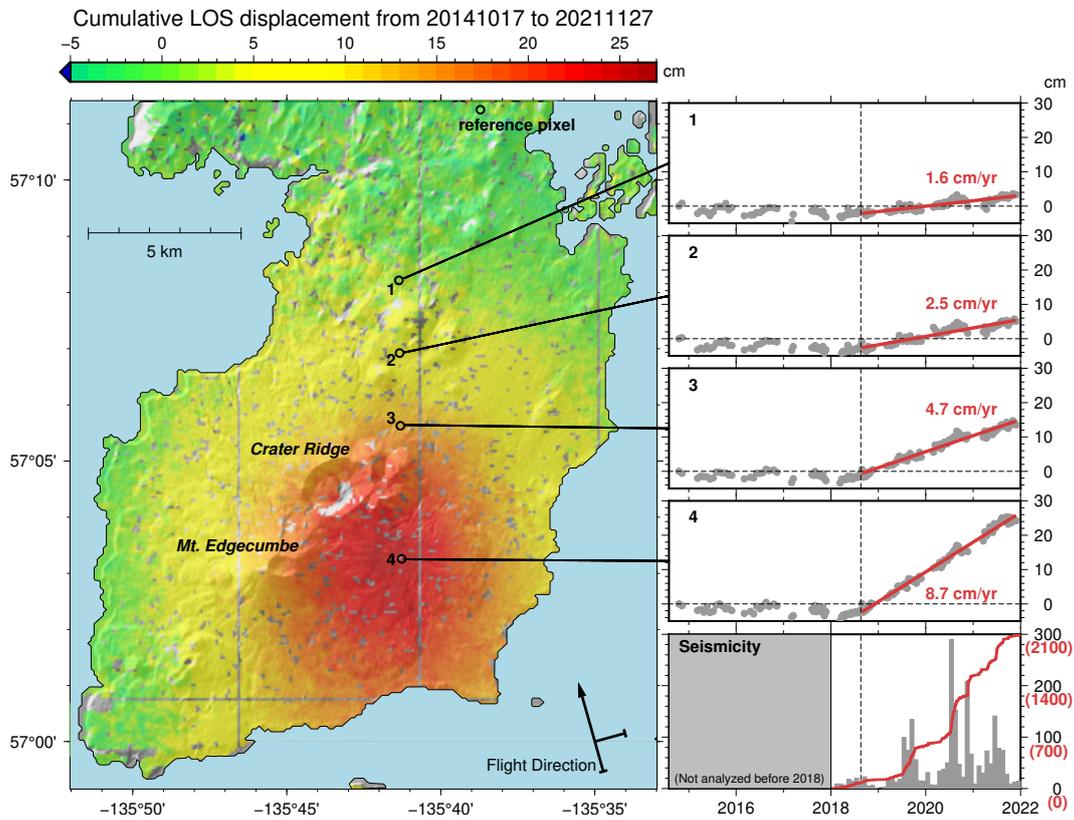


Figure 2. Cumulative line-of-sight displacement map inferred from path 174, frame 402 Sentinel-1A/B SAR data between 2014-10-17 and 2021-11-27. Regional deformation due to tectonics and GIA is largely removed by referencing the local deformation to the observations at the marked reference pixel. Points 1-4 indicate the locations for the time series to the right, showing no or only subtle seasonal deformation until August 2018 followed by a sharp inflection with annual deformation rates of up to 8.7 cm/yr. Right bottom panel shows monthly earthquake counts (gray) and cumulative seismicity (red) from Jan 2018-Dec 2021 on station AT.SIT inferred using a trigger and clustering technique (REDPy). Note that this time period does not include the April 2022 swarm as Sentinel 1B has not been acquiring data since December 2021. Figure S7 includes the most recent seismic data.

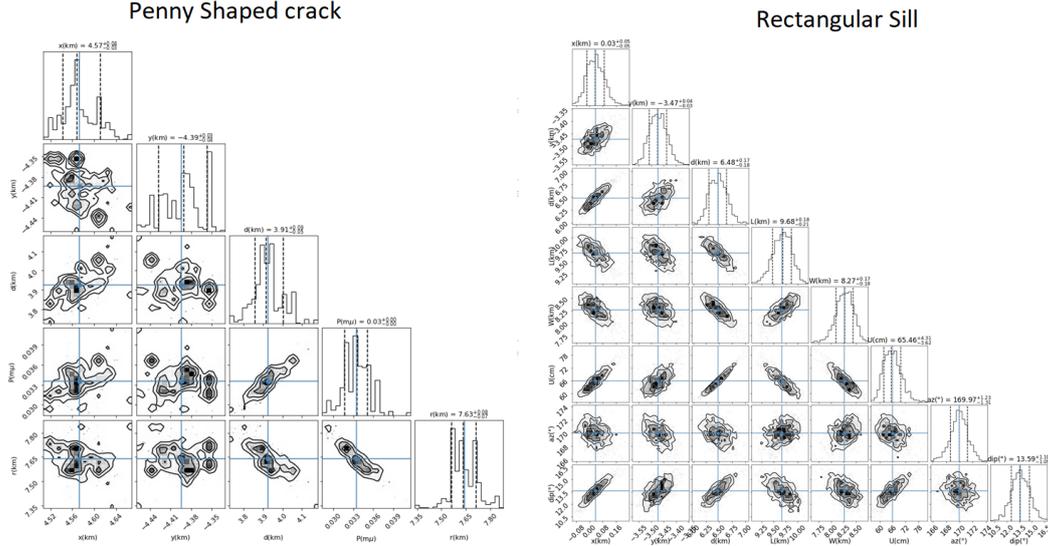


Figure 3.

218 lution results. Both the penny-shaped crack (Figure 4, middle column) and the rectan-
 219 gular sill (Figure 4, right column) visually reproduce the observations. Both models show
 220 slight residuals around Mt. Edgumbe, slightly more prominent in the circular sill model.

221 To formally evaluate the goodness-of-fit of the most likely parameters we calculate
 222 the reduced χ^2_ν statistic, which is slightly lower for the rectangular sill model (33.2 cm)
 223 than for the circular sill (37.2 cm). Point source and spheroid have much higher χ^2_ν val-
 224 ues (59.6 cm and 82.3 cm, respectively), thus we reject those models. An F-test (Tables
 225 S1-S2) confirms that the 3 additional parameters of a the rectangular sill model are jus-
 226 tified based on the improvement in fit to the data. Hence, our preferred model is a rectan-
 227 gular sill with most likely parameters values that its center about 3.5 km to the south
 228 of Crater Ridge and East of Mt. Edgumbe, dipping about 14 degrees to the West and
 229 striking about 11 degrees east of south. With its center at about 6.5 km depth, 9.7 km
 230 length and 8.3 km width the top and bottom edges of the sill at about 5.3 km and 7.6 km
 231 depth, respectively. The total opening from August 2018 until November 2018 is about
 232 65 cm. Assuming incompressible magma, this yields a total volume change of about $52.33 \times$
 233 10^6 m^3 .

234 **5 Discussion**

235 Given the large volume of highly coherent SAR data, about 4 years of effectively
 236 zero or “background” deformation, clear signals in the sample year-to-year wrapped in-
 237 terferograms, and the sharp change in deformation across the region at the same time,
 238 we have high confidence that the InSAR data reveal a physical signal related to local-
 239 ized ground deformation from volcanic activity. The lack of correlation between the sig-
 240 nal and the existing and relatively low topography over the island further suggests a phys-
 241 ical process. Topographic error correction using the approach by Fattahi and Amelung
 242 (2013) showed no significant indications of topographic error in the phase observations.
 243 We find that phase referencing to our reference pixel effectively removes tectonic and GIA
 244 related contributions. Given the large spatial wavelengths of these signals, we expect any
 245 residual motion across the observed area to be very small and non-consequential given
 246 the large magnitude of the observed signal. Further, our seismic analysis reveals a grad-

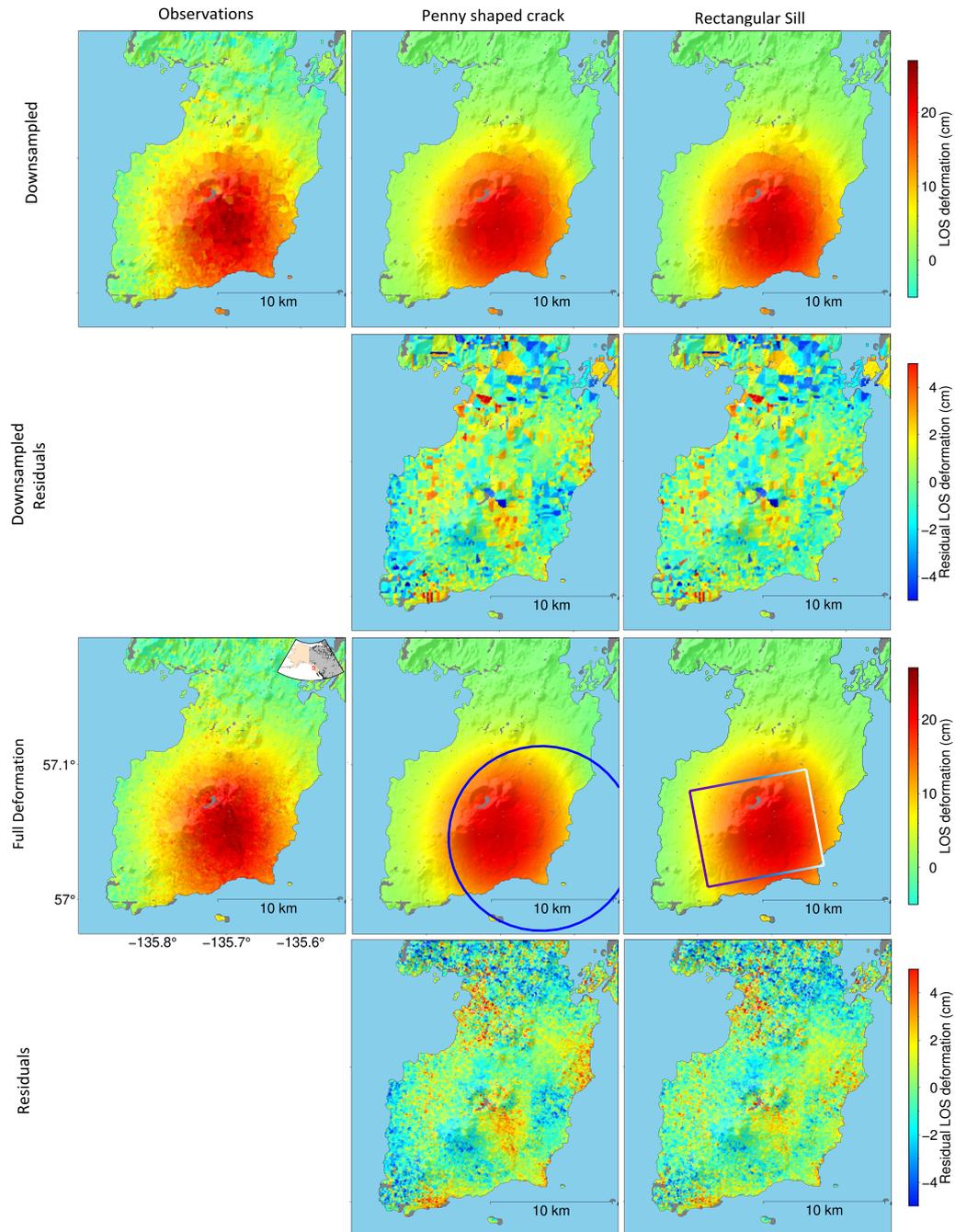


Figure 4.

Observations (left column), penny-shaped crack model predictions and data residuals (middle column), and rectangular sill model predictions and data residuals (right column). Top two rows are at variance-based down-sampled resolution, bottom two rows are at observation resolution. Model estimation and F-test are done for down-sampled residuals.

247 ual increase in seismicity above background starting from July 2019. Comparisons with
 248 earthquakes located by AEC and retrieved via the USGS ANSS catalog show coincident
 249 spikes in seismicity from 2020 onward as well, providing added credibility to our seismic
 250 analysis (Figure S7).

251 Modeling the deformation with a slightly dipping sill fits the observations well. Sta-
 252 tistical analysis of data residuals prefers this model over other geometries. Slight resid-
 253 uals in the center of the deformation signal may be due to planar geometry assumptions
 254 or, perhaps, smaller deformation sources projecting away from the sill. The comparison
 255 between the residuals of horizontal circular sill and the dipping rectangular sill shows
 256 that the misfit in this region is substantially reduced by the dip (as other tests with hor-
 257 izontal rectangular sills also showed).

258 Due to the long dormancy of the system we expect that the crust surrounding any
 259 remnant magma body is relatively cool and thus viscous effects seem unlikely over the
 260 observed time spans. This is also reflected in the InSAR time series, which show largely
 261 linear motion (with superimposed small amplitude seasonal effects) after August 2018,
 262 suggesting that the source geometry is stable over time and volume change rates are fairly
 263 constant.

264 Based on lacking basalt inclusions in Edgcumbe rhyolites and pyroclasts, Riehle
 265 et al. (1992) suggest a compositionally stratified magma chamber capped by a highly vis-
 266 cous, low density silicic seal. While this architecture may not fully reflect our current
 267 understanding of transcrustal magmatic systems (e.g., Cashman et al., 2017), our ob-
 268 servations of a relatively tabular intrusion, and the lack of any notable degassing sug-
 269 gest magma accumulation either below some barrier or into a pre-existing weakness. Com-
 270 bined with observations at, for instance, Laguna del Maule in Chile (e.g., Le Mével et
 271 al., 2021) which exhibits large and time varying deformation, we hypothesize that the
 272 observed deformation can similarly be explained by either an intrusion of mafic magma
 273 into a silicic reservoir or underplating of basalt below a silicic seal. The time delay be-
 274 tween onset of the intrusion and seismicity detectable at about 25 km distance is likely
 275 due to relatively ductile material at depth through which the mafic magma rises with-
 276 out creating noticeable seismicity at that distance and the inflation drives the fractur-
 277 ing of the overlying rock column, not the movement of magma. This is supported by the
 278 proximity to the Queen-Charlotte Fairweather Transform Fault, separating the oceanic
 279 Pacific Plate from the North American plate, supporting higher heat flux.

280 While it is not uncommon for volcanoes to deform, the activity at Mt. Edgcumbe
 281 is somewhat unique as reactivation of a dormant volcanic system, especially in large trans-
 282 form settings, is rarely observed. One recent similar case is perhaps the 2021 Icelandic
 283 Mt. Fagradalsfjall eruption on the Reykjanes Peninsula, an oblique rifting environment,
 284 which triggered strike slip earthquakes and released tectonic stresses accumulated over
 285 long time periods after 800 years with no eruption (Sigmundsson et al., 2021). In clas-
 286 sic subduction settings, Mt. Peulik in Alaska inflated from 1996-1998 at rates very sim-
 287 ilar to those we observed here, not resulting in an eruption (Lu et al., 2002). Mével et
 288 al. (2015) finds commonalities in the deformation between Laguna del Maule, Yellow-
 289 stone, Long Valley, and Three Sisters, suggesting that high rates of deformation over long
 290 timescales (more than 8 years) and varying displacement rates over time are not uncom-
 291 mon and do not necessarily result in an eruption.

292 6 Conclusions

293 A seismic swarm in 2022 near Mt. Edgcumbe (L'ux Shaa), Southeast Alaska, in-
 294 spired retrospective analysis of seismic and SAR data spanning the time period between
 295 2018-2022, and 2014-2022, respectively. Aided by open, cloud-native data and tools, we
 296 generated InSAR time series and stable deformation models just days after the onset of

297 seismic unrest, with most of the time spent on different kinds of re-analysis to ensure re-
 298 liability of our results. The InSAR time series analysis indicates no deformation until
 299 August of 2018 when rapid linear inflation at rates up to 8.7 cm/yr in the satellite LOS
 300 set on. The cumulative deformation over more than 3 years reaches as much as 27 cm
 301 in LOS centered about 2-3 km east of Mt. Edgecumbe and has a radius of about 17 km.
 302 We find the deformation to be consistent with a slightly westward dipping tabular body
 303 inflating between 5.3-7.6 km depth, striking north-northwest, and extending in length
 304 and width about 9 km and 8 km, respectively. With a cumulative opening of 0.65 m, we
 305 estimate about $52 \times 10^6 \text{ m}^3$ total magma intruded. The re-analysis of the seismicity shows
 306 that the regional network did not locate any earthquakes under Kruzof Island until Jan-
 307 uary 2020 and seismicity detected on the single closest station does not markedly increase
 308 until July 2019. Thus, we suggest a mafic intrusion into or below ductile rhyolite that
 309 acts as a seal, preventing further rise and degassing. Any seismicity related to the in-
 310 trusion may be too small to register at the distant seismograph, which records fractur-
 311 ing of the crust overlying the intrusion as strain built up through the inflation is released.
 312 This activity provides an excellent opportunity to understand the cyclicity of volcanic
 313 activity and the reawakening of dormant systems, especially in transform fault settings.

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 322 data from 2014-2022, retrieved from ASF DAAC, processed by ESA. Results generated
 323 using Copernicus Atmosphere Monitoring Service information.

324 Open Research

325 The SAR data analyzed are available through ESA and ASF DAAC. We use HyP3
 326 (Hogenson et al., 2016; Johnston et al., 2022) and OSL (Hogenson et al., 2021; Meyer
 327 et al., 2021) workflows. The seismic data analyzed are available through the IRIS Data
 328 Management Center under network code 'AT' and station 'SIT'. We use the ObsPy seis-
 329 mological toolbox (Beyreuther et al., 2010) and REDPy (Hotovec-Ellis & Jeffries, 2016)
 330 for the seismic analysis. Earthquake locations are available as part of the USGS Com-
 331 cat catalog (<https://earthquake.usgs.gov/data/comcat/>).

332 References

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