# A novel proxy for tracking the provenance of dust based on paired E1'-peroxy paramagnetic defect centres in fine-grained quartz

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#### Abstract

Crystal lattice defects in quartz have long been exploited for age determination, yet also show potential for sediment provenance studies. Here we introduce a novel method for tracking aeolian dust provenance by utilising the natural accumulation of E1' and peroxy defect centres in quartz. Our approach is based on the previously observed premise that E1' and peroxy centres arise from Frenkel defect pairs, and that their concentration increases with age of the quartz-bearing source rock. We propose that these defect centres can be utilised as a characteristic feature of the source rock and consequently, for fingerprinting sediments derived from it. We successfully apply our new protocol to distinguish fine-grained quartz extracted from loess deposits from two regions in Central Asia which are known to derive from different source material of differing age. Our method offers strong potential for identifying variability in source, both spatially and through time down sedimentary sequences.

## A novel proxy for tracking the provenance of dust based on paired E<sub>1</sub>'-peroxy paramagnetic defect centres in fine-grained quartz

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### 20 Key Points

- New sediment provenance tool exploits E<sub>1</sub>' and peroxy paramagnetic defects in quartz
- New proxy successfully differentiates quartz in loess from two different basins in
   Central Asia
- Potential applications for identifying climate-driven source change through time in
   loess and other sedimentary sequences
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### 27 Abstract

Crystal lattice defects in quartz have long been exploited for age determination, yet also show 28 potential for sediment provenance studies. Here we introduce a novel method for tracking 29 aeolian dust provenance by utilising the natural accumulation of E<sub>1</sub>' and peroxy defect centres 30 in quartz. Our approach is based on the previously observed premise that E<sub>1</sub>' and peroxy centres 31 arise from Frenkel defect pairs, and that their concentration increases with age of the quartz-32 bearing source rock. We propose that these defect centres can be utilised as a characteristic 33 feature of the source rock and consequently, for fingerprinting sediments derived from it. We 34 successfully apply our new protocol to distinguish fine-grained quartz extracted from loess 35 deposits from two regions in Central Asia which are known to derive from different source 36

material of differing age. Our method offers strong potential for identifying variability in
source, both spatially and through time down sedimentary sequences.

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#### 40 Plain Language Summary

41 Identifying the origins of dust deposits allows us to reconstruct sediment transport pathways which are essential for understanding past atmospheric circulation patterns. Here we propose 42 to exploit the characteristics of two naturally occurring defect centres in crystalline quartz, the 43 E<sub>1</sub>' and peroxy centres, as a means to distinguish sediment deriving from different origins. 44 These centres occur as pairs and are hypothesised to increase with the age of the quartz-bearing 45 rock. By this logic, the E<sub>1</sub>' and peroxy centres can be used to determine the lithic origins of 46 47 sedimentary quartz in a similar way to detrital zircon-based provenance techniques, while analysing a more ubiquitous mineral (quartz). We apply our approach, which uses a simplified 48 49 protocol for measurement in contrast to earlier studies, to successfully distinguish between loess (wind-blown dust deposits) from two different basins in Central Asia. Our new method 50 51 holds great potential in its application to loess sequences as well as other sedimentary archives.

52 Keywords: Provenance, Quartz, Defect centres, Loess, Electron Spin Resonance

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#### 54 **1. Introduction**

Identifying the original source rocks of sediments is important for understanding sediment 55 56 cycling. For aeolian sediments, pinpointing provenance has additional advantages of facilitating the reconstruction of transport pathways that relate to atmospheric circulation, and 57 58 thereby changes in climate dynamics through time. Aeolian loess deposits have long been recognised as valuable archives of past climates in terrestrial environments (Kukla, 1988; 59 Schaetzl et al., 2018). Loess sequences represent long-term accumulation of aeolian dust and 60 thus identifying their provenance provides an important proxy for reconstructing changes in 61 atmospheric circulation through time. A number of established tools are used to identify the 62 source of dust in loess deposits. These include grain-size analysis, in particular grain sorting 63 and end-member modelling, to elucidate transport modes and likely source area types 64 (Vandenberghe, 2013; Li et al., 2018); major and trace elemental composition of bulk dust (Sun 65 et al., 2002; Újvári et al., 2008); radiogenic isotope signatures (Sr and Nd) characteristic of 66

clay minerals (Chen et al., 2007; Rao et al., 2015; Ben-Israel et al., 2015); and detrital zircon 67 age profiles (Pullen et al., 2011). Whilst these methods are widely used to identify relative 68 changes in dust sources through time, they rely either on materials found in very low 69 concentrations (e.g. detrital zircons) or aggregated bulk measurements from multiple size 70 fractions which are likely to have been transported from both distal and proximal sources (e.g. 71 major and trace elements, radiogenic isotopes). Furthermore, post-depositional weathering can 72 alter in situ chemical signatures within loess, rendering certain analytical techniques, if 73 uncorrected, inaccurate for provenance (Yang et al., 2001). 74

Given the limitations of the above-mentioned provenance techniques, there has been increasing 75 76 interest in the characteristics of mineral quartz as a tool for linking dust sources and sinks. There are obvious advantages in using quartz as a provenance tool: not only is it ubiquitous, 77 78 but also highly resistant to weathering and diagenesis (Goldich, 1938). A number of quartzspecific petrographic, isotopic and geochemical provenance methods have been proposed 79 80 (Bernet & Basset, 2005; Nagashima et al., 2007, 2017; Shimada et al., 2013, Ackerson et al., 2015). Of these, electron spin resonance (ESR) signal of various defect centres in quartz has 81 82 been increasingly applied to fingerprint sources in sedimentary settings (Nagashima et al, 2007, 2011; Tissoux et al, 2015; Wei et al., 2020). 83

ESR measures the intensity of paramagnetic species (containing an unpaired electron) in a 84 material. Lattice defects and impurities in quartz give rise to various paramagnetic defect 85 centres. E<sub>1</sub>' is one such centre (Weil, 1984), comprising an unpaired electron in an oxygen 86 87 vacancy ( $\equiv$ Si, Fiegl et al., 1974) that are known to arise from diamagnetic oxygen vacancies (Si=Si, Jani et al., 1983). The most commonly used ESR based provenance protocol utilises 88 89 the heat treated- $E_1$ ' (hereafter referred to as HT- $E_1$ ') intensity of quartz by measuring the intensity of the  $E_1$ ' centre following gamma ( $\gamma$ ) irradiation and thermal treatment (Toyoda & 90 91 Hattori, 2000; Toyoda et al., 2016). It is based on the premise that  $\gamma$ -irradiation and subsequent 92 heating facilitate conversion of quartz diamagnetic oxygen vacancies into paramagnetic  $E_1$ ' 93 centres resolvable by ESR and expressed as HT-E<sub>1</sub>'. The HT-E<sub>1</sub>' intensity has been observed to increase with rock age (Toyoda et al., 1992) and is assumed to reflect the total number of 94 95 oxygen vacancies, which is characteristic of the rock type and consequently of the source rock from which the quartz is derived (Toyoda & Hattori, 2000; Toyoda et al., 2016). The HT-E<sub>1</sub>' 96 intensity is utilised in combination with the crystallinity index of quartz as a common 97

provenance tool (Nagashima et al, 2007, 2011; Toyoda et al., 2016), since the combination of
these is interpreted to reflect the age, formation and crystallisation conditions of quartz in the
source rock.

Here we investigate a simplified new provenance method based solely on the analysis of two 101 naturally occurring paramagnetic defect centres in quartz, the E<sub>1</sub>' and peroxy centres, using 102 ESR. The formation and characteristics of  $E_1$ ' and peroxy centres in natural and artificial quartz 103 have long been the subject of empirical study (Weeks, 1956; McMorris, 1970; Stapelbroek et 104 al., 1979; Friebele et al., 1979). Odom and Rink (1989), however, were the first to observe that 105 E<sub>1</sub>' and peroxy signals in quartz increase with the age of granitic host rocks, and are positively 106 107 correlated. Following which, Rink and Odom (1991) proposed that these defect centres arise from Frenkel defect pairs formed by alpha-recoil nuclei, emitted by alpha-emitting elements 108 109 (U and Th) and accumulated in rocks through time. If this principle holds true for all rock types and the mechanisms for defect formation are understood, then the E<sub>1</sub>' and peroxy signals have 110 obvious applications not only as geochronometers (Odom & Rink, 1989) but also for sediment 111 provenance. In this study, we investigate and apply this new method to distinguish quartz in 112 113 loess sediments from inland Asia. Additionally, we compare the natural  $E_1$ ' with HT- $E_1$ ' characteristics to elucidate differences between our new method and the previous approach 114 (Toyoda and Hattori, 2000). 115

Inland Asia represents one of the world's major atmospheric dust source, both past and present 116 117 (Kok et al., 2021), and has been the regional focus for investigations using ESR signals of quartz as a provenance technique (Sun et al., 2007, 2008). The basins to the north, east and 118 west of the Asian high mountains (the Pamirs, Alai-Altai and Tien Shan) lie in topographic 119 120 rain shadows and represent substantial sinks for glacially- and fluvially-derived sediment from the uplands (Schaetzl et al., 2018; Figure 1). In particular, the loess deposits of arid Central 121 122 Asia (ACA) and the Chinese Loess Plateau (CLP) reach hundreds of metres in thickness (Liu, 123 1985; Li et al., 2015) representing long-term substantial accumulation, and by extension, likely distal sources of global aeolian dust. Distinguishing between various potential dust sources in 124 inland Asia is important at multiple scales. At regional levels, identifying changes in source 125 126 down loess sequences enables reconstruction of variability in dust transport pathways and atmospheric circulation through time. At global scales, it can help identify the relative 127 contributions of different basins to the generic "Asian" mineral dust identified in Greenland 128

ice cores (Svensson et al., 2000; Bory et al., 2003) and Pacific Ocean marine sediments (Nakai

130 et al., 1993; Letelier et al., 2019).





Figure 1. Regional setting and location of the loess sites under study. The elevation map was
created using SRTM (Shuttle Radar Topography Mission) data provided by AW3D of the
Japan Aerospace Exploration Agency.

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In this study, we characterise the natural E<sub>1</sub>' and peroxy centre intensities of loess from two basins in ACA, the Ili basin and Tajik depression of southeast (SE) Kazakhstan and Tajikistan respectively (Figure 1). Recent studies based on geochemical fingerprinting and back trajectory analysis indicate that these two loess regions derive sediment from different source areas (Li et al., 2018; Li et al., 2019; Fitzsimmons et al., 2020). The potential significance of the region for global atmospheric dust loads past and present, coupled with the likely discrete sources for the two basins, make them well suited for targeted spatial and temporal comparison of quartzcharacteristics for provenance using our new method.

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#### 146 **2. Material and Methods**

We undertook measurements on fine-grained (4-11 µm) quartz extracted from loess samples 147 collected from five loess sections in Central Asia (Figure 1 and Figure S1). Four of these 148 sections, Panfilov (PAN), Ashubulak (ASH), Taukaraturyuk (TAU) and Malubai (MAL), are 149 located along a c. 200 km east-west transect of the Zalisky-Alatau range in the Ili basin of SE 150 Kazakhstan. The fifth site, Karamaidan (KAR), is a c. 60 m-thick partial section of a c. 130 m-151 thick loess-paleosol sequence, located in the foothills of the Gissar mountain range on the 152 northern margins of the Tajik Depression, in southern Tajikistan (Figure 1). We analysed 114 153 samples; 59 from SE Kazakhstan and 55 from Tajikistan. A detailed account of sampling, site 154 description and age-range, sample preparation, instrumentation and measurement protocols are 155 provided in the Supporting Information (SI). 156

157 We performed two sets of experiments on fine-grained quartz:

(i) To test our new approach, we measured the intensity of E<sub>1</sub>' and peroxy centres for all
samples. These measurements were conducted on natural quartz samples as is, without any
prior treatment and we hereafter refer to these measurements as the natural E<sub>1</sub>' and peroxy
intensity.

(ii) To understand differences between our new approach and the previous ESR-based 162 provenance method, we measured the HT-E<sub>1</sub>' intensity following published protocols 163 (Toyoda & Hattori, 2000; Nagashima et al., 2007) and compared them to the natural  $E_1$ ' 164 intensity for all our samples. This involved irradiating all the samples with a  $\gamma$ -dose of 165 2000 Gy, followed by heating them to 350 °C for 15 minutes to obtain the HT- $E_1$ ' intensity. 166 Further, we test our approach and evaluate the need for  $\gamma$ -irradiation and thermal treatment 167 prior to  $E_1$ ' measurement, by investigating the variation of ESR centres ( $E_1$ ', peroxy and 168 Al-hole) with  $\gamma$ -irradiation (varying from 0 to 40000 Gy) and temperature (300, 350 and 169 400 °C) for two representative samples, one from Kazakhstan (A0016, TAU) and the other 170

from Tajikistan (A0329, KAR). Detailed descriptions of experimental parameters aregiven in the SI.

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#### 174 **3.** Results and Discussion

# 3.1. Paired E<sub>1</sub>'-peroxy centres in quartz: Methodological considerations linking quartz crystal defect dynamics to provenance

Our measurements of the natural  $E_1$ ' and peroxy intensities of 114 Kazakh and Tajik samples indicate that the natural intensity of  $E_1$ ' and peroxy centres yield a positive correlation, with a Pearson coefficient of 0.72 (Figure 2a). This supports the hypothesis that these ESR centres indeed arise from Frenkel defect pairs in quartz as per Odom and Rink (1989), and justifies our exploration of both the natural  $E_1$ ' and peroxy signals for their application as indicators of provenance.

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Figure 2. (a) Cross-plot of natural  $E_1$ ' and peroxy centre intensities of fine-grained quartz from Kazakhstan and Tajikistan based on our new approach; (b) Comparison between natural and heat-treated (HT)  $E_1$ ' and peroxy intensities.

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189 Although intuitive, the measurement of 'natural' signals for provenance remains largely 190 unexplored. Our provenance approach, based on natural  $E_1$ ' and peroxy signals, represents a simpler protocol than the hitherto established method based on HT-E<sub>1</sub>' intensity (Toyoda et al., 2016, and references therein). Figure 2b compares the natural E<sub>1</sub>' and HT-E<sub>1</sub>' results for all samples. We observe that  $\gamma$ -irradiation and heating simply increases the signal intensity beyond the natural E<sub>1</sub>' in all samples. This indicates that the natural E<sub>1</sub>' intensity produces the same inherited provenance characteristics as HT-E<sub>1</sub>' intensity, and implies that  $\gamma$ -irradiation and heating may not be necessary for our samples.

The proposed mechanism for formation of the  $HT-E_1$ ' in the established provenance protocols 197 (Toyoda & Ikeya, 1991; Toyoda & Hattori, 2000) suggests that γ-irradiation creates hole-198 supplying Al-hole centres (which arise from Al impurities in quartz), and post-irradiation 199 200 heating causes the migration of holes from the Al-hole centres. The released holes recombine with one of the two electrons of the diamagnetic oxygen vacancies, giving rise to an oxygen 201 202 vacancy with an unpaired electron, i.e. an  $E_1$ ' centre. Therefore,  $\gamma$ -irradiation and thermal treatment essentially converts all diamagnetic oxygen vacancies - which are characteristic of a 203 given rock - into E<sub>1</sub>' centres, known as HT-E<sub>1</sub>' centres (Toyoda et al., 2016, and references 204 therein). Meanwhile our results on natural E<sub>1</sub>' measurements suggest that sample pre-treatment 205 206 by  $\gamma$ -irradiation and heating is unnecessary for ESR-based quartz provenance methods. Therefore, we undertook an additional series of irradiation and heating experiments on two 207 208 representative samples, A0016 (Kazakhstan) and A0329 (Tajikistan), to systematically assess 209 the effects of  $\gamma$ -irradiation and thermal treatment on ESR centres (E<sub>1</sub>', peroxy, Al-hole centre), and whether such pre-treatment is necessary for ESR-based provenance studies. 210

211 First, we tested the effect of  $\gamma$ -irradiation on the ESR intensity of natural E<sub>1</sub>', peroxy and Alhole centres. We irradiated 11 aliquots of each sample with a  $\gamma$ -dose varying from 0 to 40000 212 213 Gy and measured the intensity of each centre. We observe that, overall, the natural E<sub>1</sub>' intensity does not change with  $\gamma$ -irradiation (Figure 3a). Our observations corroborate with results 214 215 obtained from fine-grained quartz from other regions (eastern Europe and North America; Figure S2). Likewise, the natural intensity of the peroxy centre for both samples does not vary 216 with  $\gamma$ -dose (Figure S3). We note that the uncertainty for repeat measurements on the same 217 aliquot for peroxy signal (Figure S3) is 10-20% higher than for the natural E<sub>1</sub>', which has an 218 219 uncertainty of <1%. This is most likely due to the weak peroxy signal observed in our samples.

220 By contrast, the natural Al-hole centre intensity increases exponentially with increasing γ-dose

221 (Figure S4).



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**Figure 3.** Variation in natural  $E_1$ ' and HT- $E_1$ ' (350 °C for 15 min) intensity of fine-grained quartz (a) with  $\gamma$ -dose for sample (i) A0016 and (ii) A0329. The inset in each plot shows the variation over the lower dose range (0-2000 Gy) for the respective sample; (b) with depth (or increasing absorbed dose) at site KAR.

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Second, we tested the effect of temperature on irradiated quartz  $E_1$ ' and peroxy centres for the 228 same two representative samples (A0016 and A0329). All  $\gamma$ -irradiated aliquots (n=11) of each 229 sample were heated to 300, 350 and 400 °C for 15 min, followed by measurement of the 230 resulting  $E_1$ ' and peroxy centre intensities. In both samples,  $E_1$ ' intensity increases with 231 temperature, peaks at 350 °C, and thereafter decreases (Figure S5a, b). The E<sub>1</sub>' intensity at any 232 given temperature does not change with increasing  $\gamma$ -dose (Figure S5c, d), which can also be 233 seen in the constant ratio of  $HT-E_1$ ' to natural  $E_1$ ' intensity for both samples (Figure 3a). In 234 contrast, the peroxy signal shows minimal change in average intensity with increased 235 236 temperature (Figure S6). However, the generally weak peroxy signals produce high scatter in the data, rendering assessment of peroxy intensity response to increasing temperature difficult. 237

Our experiments on fine-grained quartz have two important implications for measuring natural 238  $E_1$ ' as a provenance signal. First,  $E_1$ ' and HT- $E_1$ ' intensity remains unchanged with increasing 239  $\gamma$ -dose, and the ratio between the two signals remains constant (Figure 3a). This suggests that 240 heating increases net E<sub>1</sub>' intensity, and irradiation has no effect. This is the case not only for 241 our samples from two regions in Central Asia, but also for detrital quartz from modern river 242 sediments (Wei et al., 2017). We conclude that natural  $E_1$ ' reflects the quartz characteristics 243 just as well as the HT- $E_1$ ' signal and argue that  $\gamma$ -irradiation and heating is not necessary. 244 Second, in both our samples, the intensity of the Al-hole centre increases with increasing  $\gamma$ -245 246 dose, in contrast to natural  $E_1$ ' and HT- $E_1$ '. If we assume the proposed formation mechanism of HT-E<sub>1</sub>' centre to be true (Toyoda & Ikeya, 1991; Toyoda & Hattori, 2000), our observations 247 imply that the number of diamagnetic oxygen vacancies in our samples is less than the number 248 of holes released from Al-centres upon heating, even for unirradiated samples. This further 249 sustains our view that the  $\gamma$ -irradiation step is redundant for our samples. 250

Our work is based on the premise that naturally occurring E<sub>1</sub>' and peroxy centres accumulate 251 252 in rocks over million-year time scales, primarily due to the effect of heavy particle irradiation and thereby show an increase with rock age (Odom and Rink, 1989). Therefore, the use of these 253 254 defect centres as provenance indicators relies on the fact that their intensity does not change significantly with ionising radiation received during transport and/ or burial as a result of the 255 256 short time spent by quartz in sedimentary settings as compared to that in the source rock. A recent study by Toyoda and Amimoto (2021) suggests that apart from ionising radiation 257 258 received in rocks, exposure to radiation during its sedimentary history is also likely to alter the

 $E_1$ ' signature of quartz. We, therefore provide further evidence that the  $E_1$ ' signal is 259 independent of ionising radiation received during transport and/ or burial by examining the 260 natural  $E_1$ ' and  $HT-E_1$ ' intensity variation with depth down a c. 60 m thick loss section at 261 KAR, Tajikistan. Based on previous (Forster & Heller, 1994) and our own magnetic 262 susceptibility measurements, we correlated the sampled part of the KAR profile to marine 263 oxygen isotope stages (MIS) 19-9 (c. 800-300 ka, Figure S7; Lisiecki & Raymo, 2005). This 264 correlation places the samples from this section beyond the limits of quartz optically stimulated 265 luminescence (OSL) dating. Hence, we estimated the minimum absorbed dose (natural burial 266 dose) received by the uppermost sample to c. 1000 Gy using the post-infrared infrared 267 stimulated luminescence protocol (Table S; Buylaert et al., 2012) on polymineral fine-grains 268 (Figure S8). Assuming a dose rate of 3-4 Gy/ka, which is typical for loess, all samples below 269 the uppermost sample would have received ionising radiation corresponding to an absorbed 270 dose of c. 1000 to 3000 Gy down the profile. The measurement of natural  $E_1$ ' and HT-  $E_1$ ' 271 intensity with depth at KAR shows no discerning patterns of incremental increase or decrease 272 (Figure 3b). Whilst, laboratory  $\gamma$ -irradiation experiments on the E<sub>1</sub>' signal shows minor 273 variations in the lower dose range (0-2000 Gy; wherein the E<sub>1</sub>' intensity first decreases and 274 then increases), unlike that at higher doses (>2000 Gy; Figure 3a). This raises an interesting 275 276 observation regarding the response of E<sub>1</sub>' signal to naturally absorbed doses down KAR, which are also likely to vary between c.1000-3000 Gy down the sequence. In Figure 3a, we 277 observe that  $E_1$ ' intensity in the lower dose range varies by 5-20% of the natural  $E_1$ ' values for 278 the representative samples, whereas the variation in E<sub>1</sub>' values with depth (and absorbed dose) 279 280 between loess and soil horizons at KAR ranges by c. 30-50% of the average E<sub>1</sub>' values (Figure 3b). Here we note that the average  $E_1$ ' value is biased by the number of samples taken from the 281 loess versus the soil horizons at KAR, hence the percent variation (if any) in E<sub>1</sub>' signature at 282 KAR is likely to be higher than suggested. This implies that the change in  $E_1$ ' value at KAR is 283 284 not an artefact of ionising radiation received during burial, but rather represents a change in source (see section 3.2). 285

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#### **3.2** Applications in aeolian environments: Examples from Central Asia

The ultimate aim of sedimentary provenance analysis in aeolian environments is to identify whether parent rocks or sedimentary sources can be linked to the loess or desert deposits in question, and additionally to determine the trajectory of dust transport pathways and by extension, the most likely climate circulation patterns prevailing at the time of transport and deposition. Our study provides a critical step towards achieving the aim of identifying provenance within sediments by establishing a simplified protocol exploiting characteristic signatures based on defect centres in quartz.

Our suite of 114 samples forms two distinct spectral clusters depending on geographic region 296 297 (Figure 2a). Of the two natural ESR signals measured, the natural E<sub>1</sub>' intensity yields the greatest difference between the two regions (Figure 2a). The more diffused peroxy signatures 298 are a result of inherently weak peroxy signal in our samples, and may also reflect the hypothesis 299 that 'peroxy' signals, as measured, represent the overlap of a peroxy radical ( $\equiv$ Si-O-O', POR) 300 and non-bridging oxygen hole centre ( $\equiv$ *Si*-O·, NBOHC) (based on observations by Salh, 2011; 301 Skuja et al., 2020; Figure S9). Nevertheless, the peroxy signals from the two regions are 302 statistically distinguishable, and along with the natural E<sub>1</sub>' intensity, can be used as provenance 303 indicators. 304

We observe higher natural  $E_1$ ' and peroxy signals in the Kazakh loess than the Tajik samples. 305 Since the  $E_1$ ' and peroxy centres are known to accumulate with time (Odom and Rink, 1989), 306 this suggests that the Kazakh quartz is sourced from older rocks than the Tajik quartz. The 307 308 suggestion that the Kazakh source rock is older is consistent with the rocks of the central Tien Shan, which are of Palaeozoic age or older (Tursungaziev & Petrov, 2008), and represent the 309 likely source material to the Kazakh loess piedmont (Li et al., 2018; Fitzsimmons et al., 2020). 310 These are older than the predominantly Mesozoic and younger rocks of the Gissar Mountains 311 and northwestern Pamirs (Vlasov et al., 1991), which provide the most likely source rocks for 312 the Tajik loess (Li et al., 2019). The formation of two separate clusters for two independently 313 sourced regions of different source rock age provides support for our proposed approach as a 314 provenance tool. 315

In addition to differentiating likely source regions for different loess sites, we ultimately aim to identify potential changes in source through time down long sedimentary (including loess) sequences such as those found in inland Asia. Figure 4a shows down-profile variations in

natural E<sub>1</sub>' intensity of loessic quartz at KAR in Tajikistan, a site which preserves multiple 319 primary loess and buried soil (paleosol) horizons (refer SI for details). We observe a marked 320 difference in the natural E<sub>1</sub>' and peroxy signature in loess versus paleosol horizons (Figure 4b). 321 Recent work using trace element concentrations of loess in Tajikistan, combined with 322 meteorological reanalysis, suggests that provenance is site-dependent and likely to be 323 dominated by proximal montane sites (Li et al., 2016), with some contribution from distal 324 sources such as the Karakum desert (Li et al., 2019). The relative contributions of distal and 325 proximal sources to loess deposits in Tajikistan may have changed over glacial-interglacial 326 327 timescales as a result of changes in the dynamics and intensity of atmospheric circulation. We suggest that our observed changes in paired E<sub>1</sub>'-peroxy characteristics represent variability in 328 the dominant source signature of quartz between the primary (glacial) loess and paleosol 329 (interglacial) horizons. At this stage, without more targeted investigations of potential source 330 rocks, we cannot identify whether the change in the dominant source signature through time at 331 KAR is a result of proximal and/ or distal transport of dust. Nevertheless, our observations hold 332 promise for more focused investigations of source signature based on our provenance method. 333

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Figure 4. (a) Down-profile variability in  $E_1$ ' intensity of fine-grained quartz at KAR, Tajikistan. The magnetic susceptibility data was measured in the field (refer SI) while LR04

benthic  $\delta^{18}$ O data was obtained from Lisiecki and Raymo (2005); (b) Natural E<sub>1</sub>' and peroxy variation in samples from various stratigraphic sections (identified here as loess, palaeosol, and weakly developed palaeosols) at KAR based on field stratigraphic description and magnetic susceptibility data.

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#### 343 4. Conclusion

344 We propose a new method for determining the provenance of sedimentary quartz based on the natural accumulation of E<sub>1</sub>' and peroxy centres. We confirm, based on empirical measurements 345 of 114 fine-grained loessic quartz samples, that firstly, the natural E<sub>1</sub>' and peroxy signals are 346 positively correlated and are likely to arise from Frenkel defect pairs. Secondly, the increase in 347 intensity of these centres with age of the source rock can be seen in the higher values obtained 348 from the Kazakh samples, which are derived from older rocks than the Tajik loess. Therefore, 349 quartz from the Ili basin of SE Kazakhstan yields signals distinct from quartz in the Tajik basin 350 in Tajikistan, indicating a difference in provenance consistent with previously published 351 studies. Furthermore, down-profile measurements at the KAR site in Tajikistan indicates a shift 352 in source between primary loess and paleosol horizons, most likely in response to changes in 353 atmospheric circulation associated with climatic oscillations. Thus, our observations suggest 354 this to be a robust new technique for sediment provenance, that can be applicable to a range of 355 settings, exploiting the characteristics of one of the most ubiquitous minerals found in nature: 356 357 quartz.

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