

# Substorm Activity as a Driver of Energetic Pulsating Aurora

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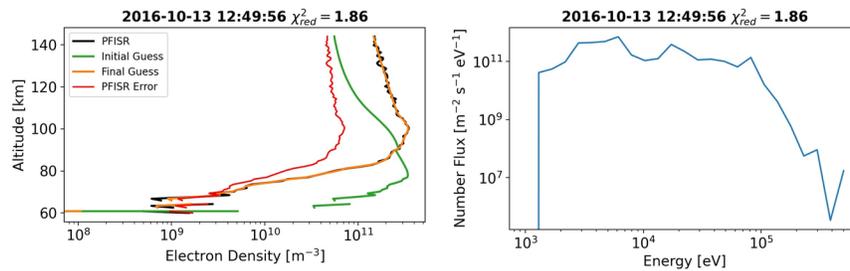
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## Abstract

Pulsating aurora are common diffuse-like aurora that appear as widely varying patches that blink on and off with periods up to 20 seconds. Rocket and incoherent scatter radar studies have suggested that energies of the responsible electrons are higher than other auroral types. However, there has yet to be a statistical study concerning the quantitative energy content of pulsating aurora. In this work, we analyzed the energy spectrum from 55 events. We obtained this by inverting the electron density profile as measured by the Poker Flat Incoherent Scatter Radar. We compared this to magnetic local time (MLT), AE index, and temporal proximity to substorm onset. There was a small propensity for higher energy fluxes between 2 and 4 MLT, but a stronger trend in relation to both temporal substorm proximity and AE index. We found that with rising AE, the average energy flux increased from 0.56 mW/m<sup>2</sup> for AE ≤ 200 to 2.24 mW/m<sup>2</sup> for an AE index > 600. Associated, is a spectral hardening where ≥ 30 keV electrons contribute 13% and 55% respectively. There was also an increase in total energy flux associated with closer temporal proximity to a substorm, although the higher energies remained present for approximately an hour. Our results confirm the high energy nature of pulsating aurora, demonstrate the connection to substorms, and imply their importance to coupling between the magnetosphere and atmosphere.



# Substorm Activity as a Driver of Energetic Pulsating Aurora

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## Key Points:

- We analyzed the inverted energies for 55 pulsating aurora events and found a close relationship to substorms and AE index.
- The average total energy flux and spectral hardness increase closer to substorm onset and for higher AE indices.
- The spectral hardness remains enhanced for approximately 1 hour after substorm onset.

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**Abstract**

Pulsating aurora are common diffuse-like aurora that appear as widely varying patches that blink on and off with periods up to 20 seconds. Rocket and incoherent scatter radar studies have suggested that energies of the responsible electrons are higher than other auroral types. However, there has yet to be a statistical study concerning the quantitative energy content of pulsating aurora. In this work, we analyzed the energy spectrum from 55 events. We obtained this by inverting the electron density profile as measured by the Poker Flat Incoherent Scatter Radar. We compared this to magnetic local time (MLT), AE index, and temporal proximity to substorm onset. There was a small propensity for higher energy fluxes between 2 and 4 MLT, but a stronger trend in relation to both temporal substorm proximity and AE index. We found that with rising AE, the average energy flux increased from  $0.56 \text{ mW} \cdot \text{m}^{-2}$  for  $\text{AE} \leq 200$  to  $2.24 \text{ mW} \cdot \text{m}^{-2}$  for an AE index  $> 600$ . Associated, is a spectral hardening where  $\geq 30 \text{ keV}$  electrons contribute 13% and 55% respectively. There was also an increase in total energy flux associated with closer temporal proximity to a substorm, although the higher energies remained present for approximately an hour. Our results confirm the high energy nature of pulsating aurora, demonstrate the connection to substorms, and imply their importance to coupling between the magnetosphere and atmosphere.

**Plain Language Summary**

Not all aurora (northern lights) are bright and defined curtains of light. Diffuse aurora are more modest. Barely visible to the naked eye, they spread across large portions of the night sky and can be easily overlooked. Pulsating aurora are a common and more playful type of diffuse aurora. In one of these displays, widely varying patches of aurora blink on and off with with periods ranging up to 20 seconds. While they aren't as bright, it has been suspected that the electrons which cause pulsating aurora are much more energetic than other types of aurora. Since energetic electrons move faster and thus can reach further into the atmosphere, it is possible that pulsating aurora may affect terrestrial climate. To study this, we first need a better understanding of pulsating aurora energies and how they can vary. In this study, we looked at the energy of 55 pulsating aurora events. In doing so, we confirmed that the energy of pulsating aurora is much higher than other types of aurora. We also found that the most energetic aurora happen close

47 in time to a magnetic disturbance known as a substorm and that a stronger disturbance  
48 leads to higher energies.

## 49 **1 Introduction**

50 Pulsating aurora are a stark contrast to the bright curtains of discrete aurora that  
51 often precede them. Diffuse and barely visible to the naked eye, this type of aurora is  
52 most often observed a few hours after magnetic midnight (e.g., Oguti et al., 1981; Jones  
53 et al., 2011). Often staying out for hours, pulsating aurora can cover large portions of  
54 the sky and in some cases expand over entire sections of the auroral region (Jones et al.,  
55 2013). Using SuperDarn and imager data, E. Bland et al. (2021) found that around half  
56 of pulsating aurora events extend between 4-5 hours of magnetic local time and between  
57  $62^\circ$  to  $70^\circ$  in magnetic latitude. Over this area, auroral patches blink on and off with  
58 periods ranging up to around 20 seconds (e.g., Davis, 1978; Lessard, 2012). Adding to  
59 the show, individual patches can be remarkably varied with differing periods, shapes, and  
60 sizes typically between 10s to 100s of kilometers (Johnstone, 1978; Lessard, 2012). See  
61 Figure 1 panels A1-A3 as an example.

62 Some studies have attempted to classify different types of pulsating aurora. For in-  
63 stance, Royrvik and Davis (1977) classified events into patches, arcs, and arc segments.  
64 More recently, Grono and Donovan (2018) made a distinction between the quickly vary-  
65 ing amorphous pulsating aurora, more regular patchy pulsating aurora, and non-pulsating  
66 patchy aurora. We included all of these when making a general identification of pulsat-  
67 ing aurora.

68 Numerous studies have shown that the electrons responsible for pulsating aurora  
69 originate in the equatorial region of the outer Van Allen radiation belt. These electrons  
70 are pitch-angle scattered into the upper-atmosphere through wave-particle interactions,  
71 most likely with lower-band chorus waves (Nishimura et al., 2010, 2011; Jaynes et al.,  
72 2013; Kasahara et al., 2018; Hosokawa et al., 2020). Previous studies have found that  
73 the energy of these particles is substantially higher than other auroral types, ranging be-  
74 tween 10s to 100s of keV (e.g., Whalen et al., 1971; Sandahl et al., 1980). This energy  
75 can vary substantially, even within individual events. Jones et al. (2009) notes often see-  
76 ing a decrease in energy throughout an event. Hosokawa and Ogawa (2015) found, us-  
77 ing the European Incoherent Scatter Radar, that the energy spectrum of pulsating au-

78 rora is harder when a patch is “on” versus when it is “off” with only background aurora  
79 present.

80 Several papers concerning the height of pulsating aurora indicate that there may  
81 also be some relation between energy and substorm onset. In the two events that they  
82 analyzed, Oyama et al. (2017) found that the atmospheric electron densities associated  
83 with pulsating aurora dropped to a lower altitude following a substorm. This would in-  
84 dicate an influx of higher energy electrons capable of penetrating further into the atmo-  
85 sphere. These results match up with the statistical study of Hosokawa and Ogawa (2015)  
86 who showed that the electron density profile of pulsating aurora extends lower in alti-  
87 tude during periods with a large AE index ( $> 500$ ). This previous work is a strong in-  
88 dicator of the increase in higher energy electrons, or spectral hardening, during geomag-  
89 netically perturbed conditions. However, the results are qualitative as altitude is only  
90 a proxy for energy. Wing et al. (2013) did conduct a statistical study of auroral ener-  
91 gies associated with substorm onset. They made distinctions between broadband (Alfvén  
92 accelerated) electrons, monoenergetic (parallel electric field accelerated) electrons, and  
93 diffuse (whistler mode wave scattered) electrons. They found that energies increase in  
94 association to substorms for all types, with the largest for diffuse electrons. However, they  
95 made no distinction between general diffuse and pulsating aurora. Thus, there is a need  
96 for a statistical study concerning the energy of pulsating aurora and how it can predictably  
97 vary with magnetospheric driving.

## 98 **2 Data**

99 This paper presents a data set of 57 pulsating aurora events (due to missing model  
100 indices we couldn’t invert two of these) captured over 51 days with the Poker Flat Re-  
101 search Range All Sky Imager (PFRR ASI). A table of these days can be found in the  
102 supplemental material. This instrument takes an image approximately every 12 seconds  
103 at 428 nm, 557 nm, and 630 nm. We used the 428 nm images because they corresponds  
104 to a lower altitude that the higher energy electrons of pulsating aurora more often reach  
105 to. It is worth noting that despite the 12 second period of the camera, we can still ac-  
106 curately identify pulsating aurora, see Figure 1 panels A1-A3.

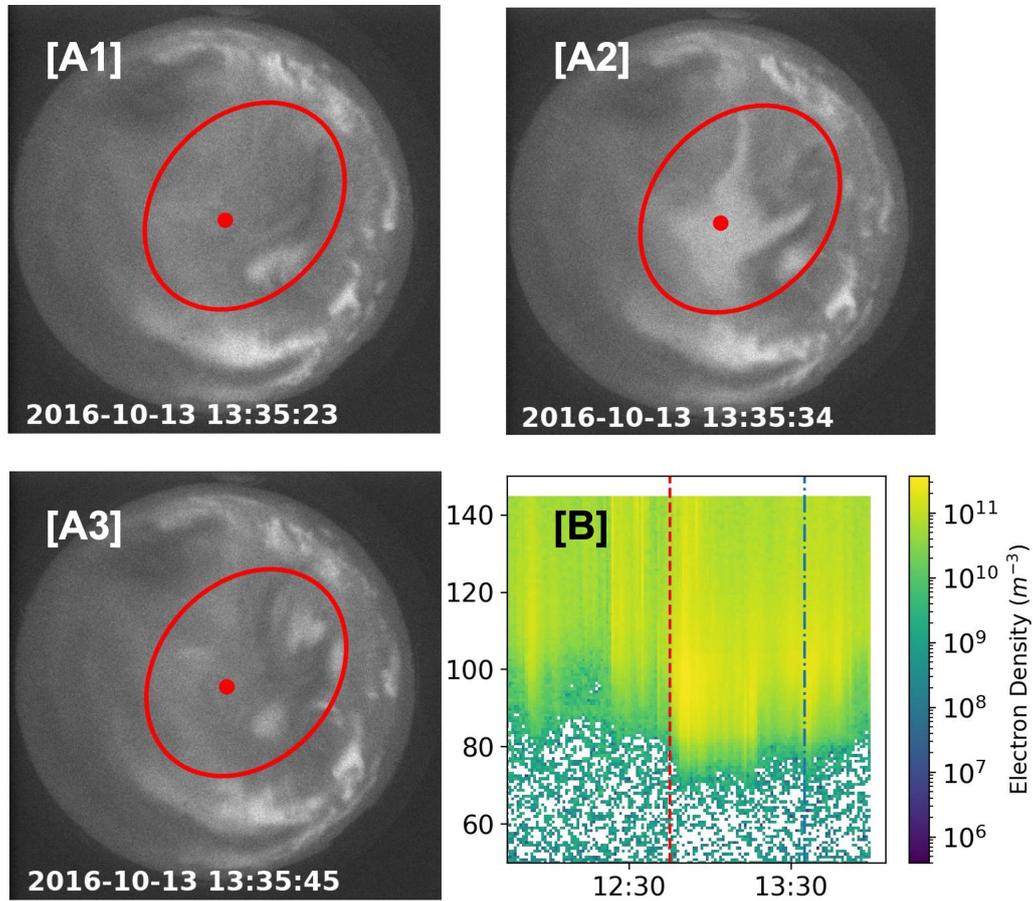
107 For each of these pulsating aurora events, the Poker Flat Incoherent Scatter Radar  
108 (PFISR) ran a MSWinds experiment (Kaepler et al., 2020). This mode is tuned for the

109 D-region of the atmosphere and provides electron density as a function of altitude and  
 110 time. The spatial scale of the measurements range between 40 and 144 km with a 0.75  
 111 km resolution. Temporally, the data is integrated over one minute (Janches et al., 2009).  
 112 The MSWinds experiment measures in four beam directions. We used the vertical beam  
 113 as it has the best statistics and is a negligible angle away from the parallel magnetic field  
 114 direction ( $< 20^\circ$ ). Figure 1 panel B is an example of electron densities measured by PFISR  
 115 MSWinds23 during a period of typical pulsating aurora on October 13, 2016. This event  
 116 began just after a substorm and continued until the end of the PFISR experiment. This  
 117 data is also an excellent example of the electron density profile pushing to lower altitudes  
 118 during pulsating aurora. For additional PFISR information see Appendix B.

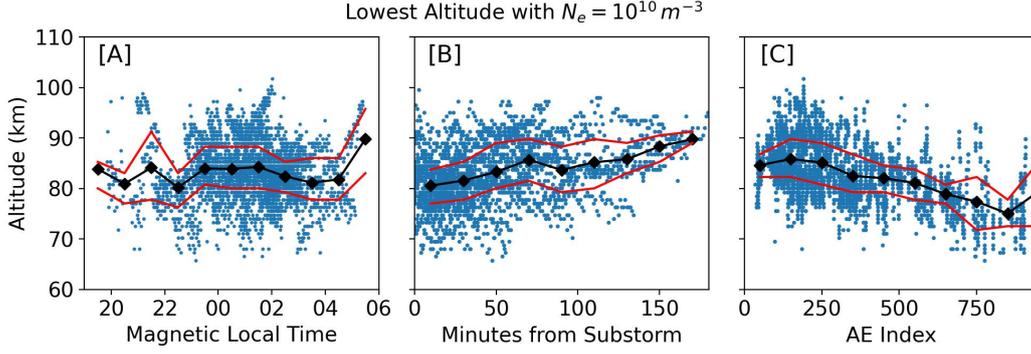
### 119 **3 Analysis**

120 In this study, we attempt to quantify the energy spectra of pulsating aurora, in par-  
 121 ticular, the higher energy portion of the spectrum. Much of the previous work has in-  
 122 dicated that the energy of pulsating aurora varies significantly both within and between  
 123 events (Jones et al., 2009; Wing et al., 2013; Hosokawa & Ogawa, 2015). Based on these  
 124 results, we chose to examine variations related to magnetic local time (MLT), AE index,  
 125 and an epoch associated with temporal substorm proximity. We set an epoch time of 0  
 126 to substorm onsets taken from lists created by Newell and Gjerloev (2011), Forsyth et  
 127 al. (2015), and Ohtani and Gjerloev (2020). We chose these three lists because they cover  
 128 a time period that encompasses our data. Each method identifies substorms in a slightly  
 129 different way, so by including all three we can identify more events over a broader range  
 130 of criteria. We limited these substorms to those that occurred within  $\pm 15^\circ$  longitude and  
 131  $\pm 8^\circ$  latitude of the Poker Flat Research Range. For the AE indices, we used archived  
 132 10-minute predicted values (Luo et al., 2013).

133 As a proxy for energy, we chose the lower altitude boundary that PFISR measured  
 134 a number density of  $N_e = 10^{10} \text{ m}^{-3}$  for each 1-minute integrated altitude profile. Ad-  
 135 ditionally, to meet this criteria, the associated error had to be less than  $5 \times 10^9 \text{ m}^{-3}$ .  
 136 We chose these values somewhat arbitrarily given that it is a round number near the de-  
 137 tection limit of PFISR. However, we did test the sensitivity and found them to be ac-  
 138 ceptably insensitive. We then plotted the lower boundaries against MLT, substorm prox-  
 139 imity, and AE index.



**Figure 1.** Panels A1-A3 show a series of 428 nm images from the Poker Flat Research Range All Sky Imager with several pulsating aurora patches of differing sizes. Even though the imaging rate is 12 seconds, we can still identify pulsating aurora. The red dot indicates the center of each image and thus the approximate location of the vertical PFISR beam. Panel B is the PFISR electron number density data for a pulsating aurora event on October 13, 2016. The data is plotted vs. altitude in km and universal time. The dashed red line indicates the start of pulsating aurora. The dashed and dotted blue line indicates when the images were taken. The radar stopped taking data before the pulsating aurora ended.



**Figure 2.** Lowest altitude PFISR measurements during pulsating aurora with  $N_e = 10^{10} \text{ m}^{-3}$  plotted versus magnetic local time [A], time from the nearest substorm onset [B], and AE index [C]. The black diamonds indicate the average altitude for the surrounding hour, 20 minutes, 200 AE units respectively. The red lines indicate the 25% and 75% quartiles.

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### 3.1 Magnetic Local Time

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Figure 2 panel A shows the altitude boundary values compared to MLT as calculated from the IGRF model for 2020. As we would expect, a majority of the measurements occurred several hours after magnetic midnight. Previous studies have shown that this is the most common time for pulsating aurora (Oguti et al., 1981; Jones et al., 2011). However, our data is biased towards common pulsating aurora times, as this is when we requested runs. The hourly averages shown by the black diamonds centered on each hour indicate that there may be a dip between 2 to 4 MLT, similar to previous results (Hosokawa & Ogawa, 2015; Partamies et al., 2017; E. C. Bland et al., 2019; Nanjo et al., 2021). However, due to the wide scatter of data and limited statistics for several time bins, it's difficult to say how significant this behavior is in our data.

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### 3.2 Substorm Proximity and AE index

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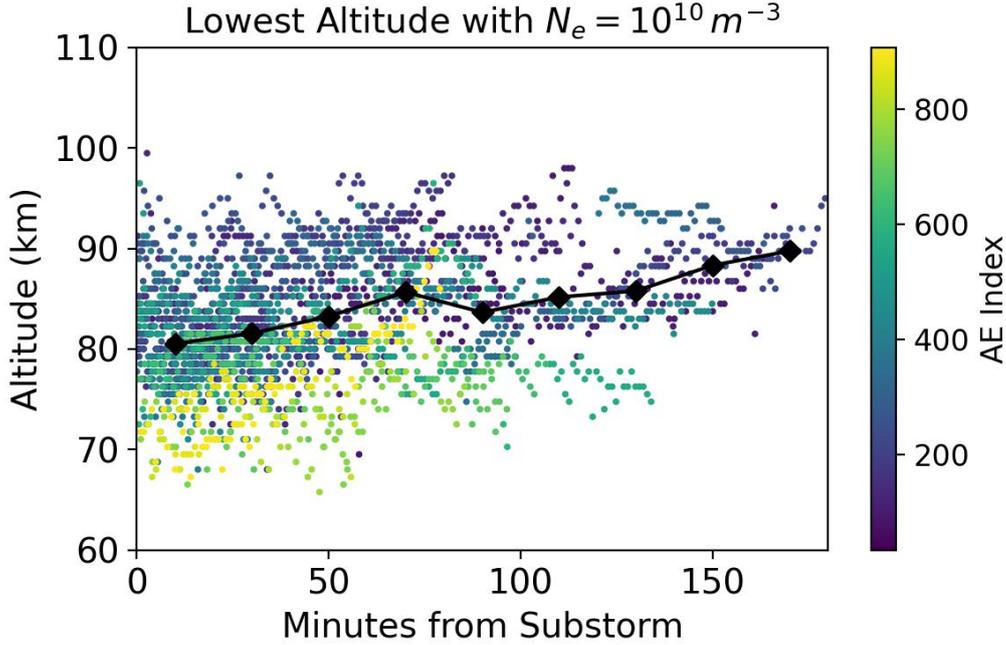
Figure 2 panel B shows the altitude boundary values compared to temporal substorm proximity. Here we see that lower altitudes are more common closer to the start of a substorm, indicating a hardening of the spectrum.

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Figures 2B and 2C shows the altitude boundary values compared to AE index. Similar to substorm proximity, there is a clear relation between a higher AE value and lower altitudes. This matches well with Hosokawa and Ogawa (2015) who found that the peak



**Figure 3.** Lowest altitude PFISR measurements during pulsating aurora with  $N_e = 10^{10} \text{ m}^{-3}$  plotted versus time from the nearest substorm onset. The markers are colored based on AE index. The black diamonds indicate the average altitude for the surrounding 20 minutes.

158 height of pulsating aurora lowers during higher AE indices. However, our data may be  
 159 more representative of the higher energy side of the spectrum since we used a lower bound-  
 160 ary value.

161 We combined Figures 2B and 2C to produce Figure 3. Here we have colored the  
 162 markers of Figure 2 panel B based on AE index. This result shows that both temporal  
 163 substorm proximity and AE index play a role in varying the lower altitude boundary.  
 164 The lowest altitudes tend to occur with both a high AE index and close temporal prox-  
 165 imity to a substorm.

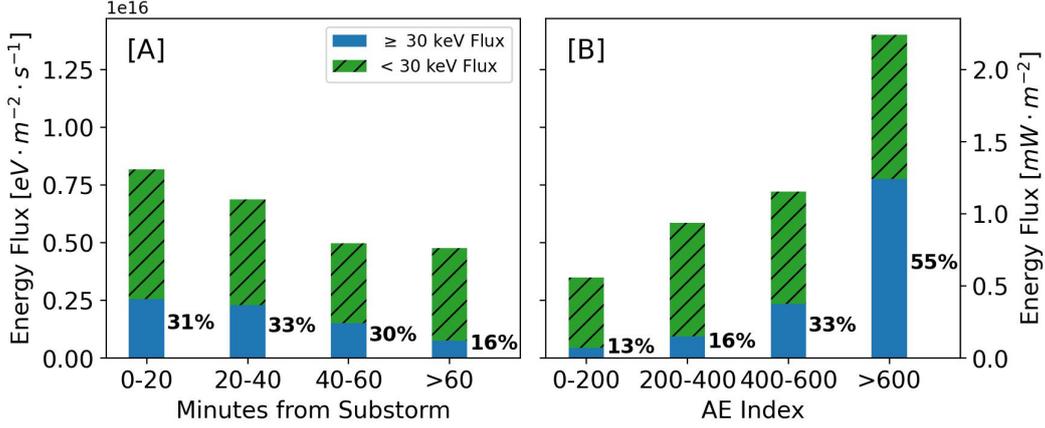
### 166 3.3 Energy Spectra from Electron Density Inversion

167 Our analysis of the lower altitude boundary with  $N_e = 10^{10} \text{ m}^{-3}$  indicates that  
 168 both AE index and temporal substorm proximity have significant impacts on how hard  
 169 the pulsating aurora energy spectrum can be. However, this metric is only a proxy for

170 energy. To investigate further, we solved the inverted problem required to convert the  
171 PFISR electron densities into energy spectra. To do this, we used the process outlined  
172 in Semeter and Kamalabadi (2005). In doing so, we assumed that the pitch angle dis-  
173 tribution was isotropic (Whalen et al., 1971; Sandahl et al., 1980), and that the electron  
174 density varies slowly compared to our time scales. We describe our exact implementa-  
175 tion of the inversion process in Appendix A. We also included an example fitted elec-  
176 tron density and the associated energy spectrum as a supplemental figure. In an anal-  
177 ysis like this, there are multiple spectra that could result in a reasonably good fit of the  
178 density profile, making the problem ill-defined. To help mitigate this, we chose the so-  
179 lution that maximized the Berg Entropy. As Semeter and Kamalabadi (2005) states, this  
180 solution “may be viewed as the most noncommittal approach with respect to the unavail-  
181 able information.” To further reduce uncertainty, we chose an energy threshold of 30 keV  
182 to separate the low and high portions of the energy spectrum and integrated the two re-  
183 gions. This gives us an average low and high energy flux and limits the dependency of  
184 our results on the spectral shape.

185 The largest source of error in the inversion process is likely the assumed atmospheric  
186 chemistry that connects PFISR observations to an ionization rate. This is not well known,  
187 especially for the D-region. As our primary chemistry model we used the Glukhov-Pasko-  
188 Ina (GPI) model (Glukhov et al., 1992; Lehtinen & Inan, 2007). This has been shown  
189 to perform well for the D-region (Marshall et al., 2019). However, it is not well defined  
190 in the E-region. To account for this, we set the values above 90 km to those calculated  
191 by Gledhill (1986) for nighttime aurora. The Gledhill model is suitably close that of Vickrey  
192 et al. (1982) above 90 km and the Vickrey model has been shown to perform well in this  
193 region (Sivadas et al., 2017). While we could have used the Vickrey model, we believe  
194 the Gledhill model is slightly better for this data. However, both models are only rough  
195 estimates. We refer to this adjusted model as GPI+. To provide context to our results  
196 calculated using GPI+, we inverted each density profile using three additional chemistry  
197 models. These results can be found in Appendix A.

198 After performing the inversions, we found the geometric mean for  $\geq 30$  keV and  
199  $< 30$  keV electrons in bins relative to substorm onset and AE index. Figure 4 shows the  
200 results. While not shown here, we found similar relative behavior in analyses for energy  
201 thresholds of 50 keV and 100 keV. For  $< 20$  min the high energy contributions were 13%  
202 and 1.3% respectively. For  $> 600$  AE these were 35% and 2.6% respectively.



**Figure 4.** The high ( $\geq 30$  keV) and low ( $< 30$  keV) energy flux contributions to pulsating aurora events occurring in four temporal bins relative to substorm onset [A] and AE index [B].

## 4 Discussion

Before discussing our results, it is worth mentioning possible sources of systematic error. One, several previous studies found that the electron density and energy spectra shifts towards higher energies during the on phase of pulsating aurora (Hosokawa & Ogawa, 2015; Whalen et al., 1971). Our data is integrated over one minute, so these variations will likely be smoothed out, thus reducing the spectral hardness. Two, we are not capturing the full spectrum of electron precipitation. Ionization associated with electron energies less than about 1 keV, peaks above the altitudes that PFISR measures in the D-region mode (Fang et al., 2010). If the energy flux for this portion of the spectrum is significant, we could be overestimating the spectral hardness and underestimating the total energy. Three, the instrument sensitivity limits our ability to detect higher energy particles with lower fluxes. If populations such as these are present, we could be underestimating the spectral hardness. Four, we only selected pulsating aurora that were in the center of the imager, but we didn't account for times that the PFISR beam wasn't directly on a pulsating patch. If the precipitating flux is highly local, we could be underestimating the energy flux during such periods.

Figure 4 shows how the energy composition of pulsating aurora varies with respect to both substorm proximity [A] and AE index [B]. Within an hour of a substorm around a third of the total energy flux is carried by  $\geq 30$  keV electrons. At  $> 60$  minutes this drops to around a sixth. Interestingly, while the total energy flux climbs closer to the

223 substorm, the energy composition remains similar all the way out to an hour after on-  
 224 set. This indicates that the initial substorm “kick” hardens the spectrum and it remains  
 225 hard up to an hour afterwards, even as the total energy flux decreases.

226 The energy spectrum associated with AE index varies even more dramatically. In  
 227 highly perturbed times of  $AE > 600$  over a half of the average energy flux is carried by  
 228 the  $\geq 30$  keV electrons. This again drops to just over a tenth for quiet periods of  $AE$   
 229  $\leq 200$ . Assumptions about the atmospheric chemistry can vary these values, but for ev-  
 230 ery model we found the same relative behavior. The relative behavior was also the same  
 231 when we used threshold values of 50 keV and 100 keV. Thus, we speculate with a high  
 232 level of confidence that pulsating auroral energies are varied by both the strength of a  
 233 substorm as well as temporal proximity to it.

234 Combining our results with those of E. Bland et al. (2021), we can perform a back-  
 235 of-the-envelope calculation to estimate the incoming power of a typical pulsating aurora  
 236 event. We will assume an event extending between  $62^\circ$  and  $70^\circ$  magnetic latitude and  
 237 4 hours of magnetic local time. Using this, approximately 5.4 gigawatts (GW) of power  
 238 would be entering the atmosphere during periods with  $AE > 600$  with 2.7 GW coming  
 239 from  $\geq 30$  keV electrons. For periods  $< 20$  minutes after substorm onset and all AE  
 240 indices these values are 2.9 GW and 0.9 GW respectively.

241 Our results are significant as they indicate a process connecting substorms and pul-  
 242 sating aurora. There is a well documented relation between substorm activity post-midnight  
 243 and whistler-mode wave generation near the equator (Tsurutani & Smith, 1974; Thorne  
 244 et al., 1974). The proposed mechanism connecting them is Doppler-shifted cyclotron res-  
 245 onance with 10-100 keV substorm injected electrons (Dungey, 1963; Kennel & Petschek,  
 246 1966). In addition, the amplitude of already present whistler-mode waves can vary with  
 247 substorm injection. Meredith et al. (2000) showed that between  $3.8 < L < 6$  whistler-  
 248 mode amplitudes increased after a substorm and then decayed with a timescale of  $\tau \approx$   
 249 1.1 hours. That value is similar to the timescale over which we see a decrease in the spec-  
 250 tral hardness. Given that whistler-mode waves are known to drive pulsating aurora, this  
 251 is one likely explanation. Additionally, the MLT dependence of substorm driven whistler-  
 252 mode waves could explain the slight increase in energetic pulsating aurora events we see  
 253 post-midnight (Tsurutani & Smith, 1977).

254 Our results also confirm, as previous studies have hinted at, that the energetic na-  
 255 ture of pulsating aurora is inherent to the phenomenon and not just a result of a few ex-  
 256 treme events. This is important because pulsating aurora are very common (Oguti et  
 257 al., 1981) and can be long-lasting (Jones et al., 2013). Thus they represent a relatively  
 258 large transfer of energy between the magnetosphere and lower ionosphere. When con-  
 259 sidering the effects of this transfer, the total energy flux is clearly important, but so too  
 260 is the hardness of the energy spectrum. Higher energy electrons reach further into the  
 261 atmosphere and thus have a higher probability of influencing terrestrial climate through  
 262 processes like  $\text{NO}_x$  based ozone depletion (Turunen et al., 2016; Verronen et al., 2021,  
 263 & and references therein). We found that the hardest spectra occur close in time to a  
 264 substorm and for high AE indices. In short, our results can be used to more accurately  
 265 parameterize the atmospheric consequences of pulsating aurora.

## 266 5 Summary

- 267 • The energy flux of pulsating aurora correlates strong with the temporal substorm  
 268 proximity and AE index.
- 269 • In relation to temporal substorm proximity the total energy flux varies between  
 270 1.31 and  $0.76 \text{ mW} \cdot \text{m}^{-2}$  for  $\leq 20$  and  $> 60$  minutes. The associated contribu-  
 271 tion to the total energy flux from  $\geq 30$  keV electrons are 31% and 16%.
- 272 • In relation to substorms, the energy spectrum remains hard out to 1 hour after  
 273 onset before softening.
- 274 • In relation to AE index the total energy flux varies between  $2.24$  and  $0.56 \text{ mW} \cdot$   
 275  $\text{m}^{-2}$  for  $> 600$  and  $\leq 200$  AE indices. The associated contributions to the to-  
 276 tal energy flux from  $\geq 30$  keV electrons are 55% and 13%.
- 277 • We estimate that for a typically pulsating auroral event occurring  $< 20$  min af-  
 278 ter substorm onset (AE  $> 600$ ), approximately 2.9 (5.4) GW of power enters the  
 279 atmosphere. The contributions from  $\geq 30$  keV electrons are 0.9 (2.7) GW.

## 280 Appendix A Inversion Technique

To solve the inverted problem of extracting an energy spectrum from electron den-  
 sities, we used the process outlined in Semeter and Kamalabadi (2005). We assumed the  
 pitch angle distribution of the incoming electrons was isotropic and used the universal

energy dissipation function ( $\Lambda$ ) given in the paper. We took our range-energy function from Barrett and Hays (1976) as

$$R = 4.7 \times 10^{-6} + 5.36 \times 10^{-5} K^{1.67} - 0.38 \times 10^{-7} K^{-0.7} \quad [\text{kg} \cdot \text{m}^{-2}]$$

where  $K$  is the electron energy in keV. Using these, we can construct a matrix  $A$ , where

$$A_{ij} = \frac{\Lambda \left( \frac{s(z_i)}{R(K_j)} \right) \rho(z_i) K_j \Delta K_j}{35.5 R(K_j)}$$

281 where  $s(z_i) = \sec(\theta) \int_{z_0}^{\infty} \rho(z) dz$  is the mass distance traveled by an electron as a func-  
 282 tion of altitude. We assumed the dip angle of the magnetic field,  $\theta \approx 0$ . We calculated  
 283 the neutral atmospheric density  $\rho(z)$  using the NRLMSISE00 model and approximated  
 284  $z \rightarrow \infty$  as  $z = 1000$  km (Hedin, 1991).

The matrix  $A$  relates the ion production rate ( $q$ ) and the differential number flux ( $\phi$ ) via

$$q_i = A_{ij} \frac{\phi_j}{\Delta K_j}$$

285 As Fang et al. (2010) showed, using a range-energy function gives poor estimates of the  
 286 ion production rate from electrons below 1 keV. However, the altitude range of the PFISR  
 287 data means that there is very little, if any, contribution from these energies. Therefore,  
 288 we assume that the range-energy function is a good enough estimate in this case.

Important atmospheric chemistry is encapsulated in the conversion of electron density measured by PFISR to an ion production rate. This is especially relevant below 85 km, where the chemistry of ion production becomes increasingly complex (Mitra, 1981). There are several ways of handling the chemistry. The simplest models describe this relation with the continuity equation

$$\frac{dn}{dt} = q - \alpha n^2$$

289 Assuming the temporal change of the electron density, as measured by PFISR, is small  
 290 compared to the timescales we are studying, we can say that  $q = \alpha n^2$ , where  $\alpha$  is the  
 291 effective recombination coefficient. From our experience, this steady state assumption  
 292 is good for pulsating aurora, at least when integrated over 1 minute like the PFISR data  
 293 is.

294 In a model that uses this description, the chemistry is encapsulated in the coeffi-  
 295 cient. However, obtaining an accurate estimate of  $\alpha$  is difficult in practice. For our pri-

296      mary results we used the Glukhov-Pasko-Ina (GPI) model (Glukhov et al., 1992; Lehti-  
 297      nen & Inan, 2007). This uses the specific conditions as measured by PFISR, and mod-  
 298      eled by the International Reference Ionosphere (IRI-2016) and NRLMSISE-00. Previ-  
 299      ous work has shown that GPI performs well for the D-region (Marshall et al., 2019). How-  
 300      ever, it is not well defined in the E-region. To account for this, we set the values above  
 301      90 km to those of Gledhill (1986) for nighttime aurora. The Gledhill model is suitably  
 302      close that of Vickrey et al. (1982) above 90 km and the Vickrey model performs well in  
 303      this region (Sivadas et al., 2017). While we could have used the Vickrey model, we be-  
 304      lieve the Gledhill model is slightly more accurate to this data. We refer to this adjusted  
 305      model as GPI+. Given that the the chemistry in this region of the atmosphere is not  
 306      well known, we also performed our analysis with three additional models to provide con-  
 307      text.

1. The best fit from Vickrey et al. (1982) of multiple observations from several au-  
 thors of  $\alpha$  in the E-region.

$$\alpha(h) = 2.5 \times 10^{-12} e^{-h_{\text{km}}/51.2} \quad [\text{m}^3 \cdot \text{s}^{-1}]$$

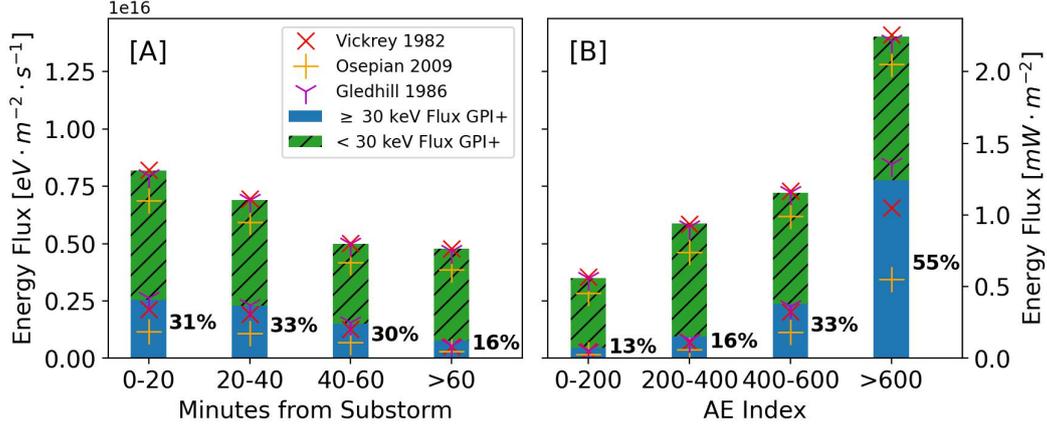
308      To use this model we needed to extend it into the D-region, where it is not well  
 309      defined.

2. The observations of Osepian et al. (2009) during a solar proton event on January  
 17, 2005 at 9:50 UT. While these observations cover the D-region, they must be  
 extended into the E-region. They also only cover a single event and that event is  
 not pulsating aurora.
3. The best fit of Gledhill (1986) for nighttime aurora covering the E-region and D-  
 region.

$$\alpha(h) = 4.3 \times 10^{-6} e^{-2.42 \times 10^{-2} h_{\text{km}}} + 8.16 \times 10^{12} e^{-0.524 h_{\text{km}}} \quad [\text{cm}^3 \cdot \text{s}^{-1}]$$

314      Figure A1 shows how these three additional chemistry model compare with our anal-  
 315      ysis. They are represented by scatter points around each bar. These points can be con-  
 316      sidered as rough bounds on our results.

To determine the differential number flux ( $\phi$ ) we iterated using the maximum en-  
 tropy method outlined in Semeter and Kamalabadi (2005). We monitored convergence  
 through the  $\chi^2$  value between the modeled ion production rate and the rate calculated  
 from the PFISR measurements. We stopped iterating when the step difference in the  $\chi^2$



**Figure A1.** The high ( $\geq 30$  keV) and low ( $< 30$  keV) energy flux contributions to pulsating aurora events occurring in four temporal bins relative to substorm onset [A] and AE index [B]. We set the bar heights to the GPI+ model. The scatter points indicate the individual values from the three other chemistry models.

values was less than 0.01. This usually took between 100 and 1000 steps. From the spectra that converged, we took those with a  $1 \leq \chi_{\text{reduced}}^2 < 3$  to be suitably good models. To calculate  $\chi^2$  it is important to have an accurate description of the variances (errors) in the PFISR data. The data products contain absolute errors associated with the measured number density. To propagate this to the ion production rate we calculated an intermediary recombination coefficient,  $\alpha_{\text{chem}}(z) = q_{\text{chem}}(z)/n(z)^2$ . Our errors were then

$$\Delta q_{\text{chem}}(z) = 2\alpha_{\text{chem}}(z)n(z)\Delta n(z)$$

317 To determine  $\chi_{\text{reduced}}^2$  we need an estimate of the degrees of freedom in the model. We  
 318 set this as the number of altitude bins where the errors were less than the data (fitted  
 319 values) minus the number of energy bins (varied values).

320 When performing the inversion, we found that the differential number flux of the  
 321 highest energy bin was often over an order of magnitude larger than the next highest bin.  
 322 We believe this is not physical and instead an artifact due to the initial electron density  
 323 guess only needing to converge to the PFISR sensitivity ( $\sim 10^9 \text{ m}^{-3}$ ) and not zero for  
 324 lower altitudes. To mitigate this error, we only calculated our averages up to the sec-  
 325 ond highest energy bin.

## Appendix B PFISR Data

For this work, we used PFISR data that was taken during the MSWinds experiments. These experiments use a barker code pulse pattern that allows better sensitivity in the D-region atmosphere. In this mode, the instrument estimates electron density from the received power.

We assumed that any signal below 60 km was unphysical in regards to precipitating electrons and took the average signal between 55 and 60 km as our baseline background. We then subtracted this background from the entire electron density profile. In an ideal world, the background would average to 0, but due to hardware constraints after late 2016, the PFISR system has difficulty properly calibrating this and it can be as high as  $1 \times 10^9 \text{ m}^{-3}$ . These steps are required so we don't get errant estimates of the high energy portion of the spectrum.

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Date	Start Time (UTC)	End Time (UTC)
2012-03-28	10:25:28	12:46:49
2012-12-20	11:12:07	16:18:47
2013-03-24	10:52:26	12:38:36
2014-11-02	12:40:31	13:31:41
2015-01-13	11:26:42	11:59:50
2015-01-14	10:32:55	10:54:12
2015-01-26	7:49:39	8:18:41
2015-02-26	9:46:46	10:45:44
2015-02-26	12:29:40	14:03:29
2015-03-12	10:01:48	10:44:22
2016-10-10	11:11:54	11:59:55
2016-10-13	12:49:20	13:59:51
2016-10-16	11:47:46	12:59:54
2016-10-19	10:33:39	12:15:24
2016-11-02	12:10:36	12:59:52
2016-11-13	9:24:40	10:59:49
2016-11-25	10:00:02	10:59:52
2016-12-11	9:51:00	9:59:56
2016-12-20	8:59:47	10:59:56
2016-12-26	10:32:15	10:59:51
2017-01-06	8:32:02	8:59:51
2017-03-30	12:10:59	13:04:47
2017-04-14	12:00:04	12:58:46
2017-04-18	12:00:01	12:31:27
2017-08-17	8:08:43	8:43:34
2017-09-03	9:36:22	10:45:19
2017-09-03	11:10:50	13:00:10
2017-09-14	11:30:31	14:00:05
2017-09-18	8:34:19	8:59:50
2018-10-23	11:05:32	11:35:47
2018-12-30	11:10:22	11:53:21
2019-01-06	12:00:07	12:59:53
2019-01-07	12:00:00	12:59:54
2019-01-26	13:47:03	13:59:48
2019-01-31	13:00:10	13:59:54
2019-02-01	13:00:09	13:59:49
2019-02-28	13:50:41	16:01:17
2019-03-01	10:04:47	12:04:41
2019-03-02	6:44:08	7:20:41
2019-03-28	13:45:12	14:06:12
2019-03-31	13:00:02	13:39:53
2019-04-03	13:00:11	13:48:53
2019-09-01	7:36:54	8:09:04
2019-09-01	9:14:12	10:21:35

2019-09-01	11:39:50	13:15:49
2019-09-02	10:12:14	11:06:53
2019-09-03	12:28:43	13:19:55
2019-12-18	11:42:17	11:59:59
2019-12-19	8:23:32	9:46:10
2019-12-19	10:47:53	11:59:52
2020-01-04	11:36:36	12:47:11
2020-01-05	12:23:32	15:02:37
2020-01-31	11:51:52	13:22:28
2020-03-31	11:36:47	12:51:36
2020-03-31	13:00:22	13:54:27
2020-10-01	12:34:59	14:43:50
2021-01-13	11:08:54	14:26:32