High-Pressure Melting Curve of FeH: Implications for Eutectic Melting between Fe and Non-Magnetic FeH

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Abstract

While hydrogen could be an important light alloying element in planetary iron cores, phase relations in the Fe-FeH system remain largely unknown at high pressures and temperatures (*P*-*T*). A speculative Fe-H₂ phase diagram has been proposed assuming continuous solid solution between Fe and FeH and eutectic melting between FeH and H₂. Recent studies revealed that stoichiometric FeH becomes non-magnetic above ~40 GPa, which might affect its melting behavior. Here we examined the melting curve of non-magnetic FeH between 43 and 152 GPa by a combination of laser-heated diamond-anvil cell (DAC) techniques and synchrotron X-ray diffraction (XRD) analyses. The melting temperature was determined by employing the appearance of additional hazy XRD signals upon quenching temperature as a melting curve upon the loss of magnetism and extrapolates the experimental constraints to inner core pressures. The XRD data showed that non-magnetic FeH melts congruently at temperatures higher than the known eutectic melting curve for FeH_x (x > 1). Combined with the fact that the endmembers exhibit different crystal structures, these results indicate that Fe and non-magnetic FeH form a eutectic system. The dT/dP slope of the FeH melting curve is comparable to that for Fe, suggesting that the eutectic liquid composition of FeH_{0.42} (Fe + 0.75 wt% H) previously estimated at ~40 GPa changes little with increasing pressure.

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12	Key Points:				
13	• We determined the melting curve of non-magnetic FeH to 152 GPa based on				
14	synchrotron XRD measurements.				
15	• XRD data indicated congruent melting of non-magnetic FeH, suggesting eutectic				
16	melting between Fe and FeH above ~40 GPa.				
17	• Similar dT/dP slopes between the melting curves of Fe and FeH imply little				
18	pressure dependence of the Fe-FeH eutectic liquid composition.				
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23	continuous solid solution between Fe and FeH and eutectic melting between FeH and H ₂ .				
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28	temperature was determined by employing the appearance of additional hazy XRD				
29	signals upon quenching temperature as a melting criterion. We also performed				
30	thermodynamic modeling, which well reproduces the change in the curvature of FeH				
31	melting curve upon the loss of magnetism and extrapolates the experimental constraints				
32	to inner core pressures. The XRD data showed that non-magnetic FeH melts congruently				
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34	with the fact that the endmembers exhibit different crystal structures, these results				
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40 Plain Language Summary It is likely that a large amount of water was transported to 41 the Earth during its accretion and hydrogen was incorporated into core-forming metals. 42 Indeed, recent calculations found that hydrogen could be an important light element in 43 both the outer and inner core to explain the observed density and velocities. Nevertheless, 44 experimental study of Fe-H alloys has been challenging, in part because hydrogen is 45 almost insoluble in iron at 1 bar. Here we determined the melting curve of stoichiometric 46 FeH in the non-magnetic state from 43 to 152 GPa based on XRD measurements. The 47 melting temperature of FeH increases rapidly with compression upon the loss of local spin moment. The XRD spectra show that stoichiometric FeH melts congruently, 48 49 suggesting eutectic melting between Fe and FeH that is supported by the different crystal 50 structures between these two endmembers. Since the dT/dP slopes of the Fe and FeH 51 melting curves are similar, it is likely that the Fe-FeH eutectic liquid composition is little 52 dependent on pressure and could be around FeH_{0.42} at inner core conditions, which gives 53 the upper bound for the hydrogen concentration in the outer core.

54 1. Introduction

55 Hydrogen (H) is likely to be an important light element in the cores of terrestrial planets 56 (Hirose et al., 2021; Yoshizaki & McDonough, 2020). According to planet formation theories (Raymond et al., 2007; Walsh et al., 2011; Sato et al., 2016), water may have 57 58 been delivered to the growing Earth and Mars in large quantities. The high metal/silicate 59 partition coefficient of H found in recent experimental and computational works suggest 60 that most of the water transported to our planet was sequestered mainly as up to ~1 wt% 61 H in the core (Tagawa et al., 2021; Li et al., 2020; Yuan & Steinle-Neumann, 2020). 62 Indeed, H-bearing iron (Fe) alloys can explain the observed density and seismic-wave 63 velocities in both the liquid outer core (Umemoto & Hirose, 2015, 2020) and the solid 64 inner core (Wang et al., 2021; He et al., 2022). In addition, the recent marsquake 65 observations by the InSight probe revealed that the Martian core is less dense than 66 previously thought, possibly including 1–2 wt% H in addition to sulfur (S) (Stähler et al., 67 2021).

68 The Fe-FeH phase diagram is of great importance to understand the present state and 69 composition of planetary metallic cores. However, it is still vaguely known. Fe-H alloys 70 must be examined under high pressure, because the solubility of H in Fe is negligibly

small at 1 bar and rapidly increases with increasing pressure (e.g., Fukai & Suzuki, 1986).

72 Fukai (1992) speculated on the melting and subsolidus phase relations in the Fe-H₂

73 system at >100 GPa, assuming that Fe and FeH form a continuous solid solution (Figure

1a). While earlier experiments on the Fe-H(\pm Ni) system supported Fukai's phase diagram

75 below ~20 GPa (Sakamaki et al., 2009; Shibazaki et al., 2014), it has not been verified at

76 higher pressures. Experiments on FeH should be performed under H-undersaturated

conditions otherwise FeH₂ and FeH₃ form above 60 GPa (Pépin et al., 2014).

78 The crystal structure and melting temperature of stoichiometric FeH is a key to the Fe-79 FeH phase diagram. Previous studies have shown that it adopts the face-centered cubic 80 (fcc) structure above ~50 GPa and ~1000 K (Kato et al., 2020; Tagawa et al., 2022), which 81 is different from the hexagonal close-packed (hcp) structure for Fe at core conditions 82 (Komabayashi et al., 2009). Tagawa et al. (2022) also demonstrated that FeH loses its 83 local spin moment above ~40 GPa. The melting curve of non-magnetic fcc FeH is not 84 known. If Fe and FeH form a continuous solid solution as supposed by Fukai (1992), the 85 solidus temperature of FeH corresponds to the eutectic temperature in the FeH-H₂ system 86 (Sakamaki et al., 2009; Shibazaki et al., 2014; Hirose et al., 2019) (Figure 1a). However, 87 the different crystal structure between Fe and FeH above ~50 GPa will probably change 88 the Fe-FeH phase diagram significantly.

89 In this study, we examined the melting curve of fcc stoichiometric FeH in a pressure range 90 between 43 and 152 GPa in a laser-heated DAC, based on synchrotron XRD 91 measurements *in-situ* at high *P-T*. The results demonstrate that the dT/dP slope of the 92 melting curve of non-magnetic FeH is substantially larger than that observed for the 93 magnetic phase below 20 GPa (Sakamaki et al., 2009). In addition, XRD spectra indicated 94 that melting of FeH is congruent. It indicates that non-magnetic FeH melts at a 95 temperature maximum and forms a eutectic system with Fe, which is supported by the 96 different crystal structures between these two endmembers. The high dT/dP slope of the 97 FeH melting curve is similar to that of Fe, suggesting that the Fe-FeH eutectic liquid 98 composition that places the upper bound for H concentration in the Earth's liquid outer 99 core is little dependent on pressure.

100 2. Experimental Methods

101 **2.1. High-pressure Experiments**

High *P-T* conditions were generated by using laser-heated DAC techniques up to 152
GPa and 4630 K (Figure 2). We employed diamond anvils with beveled 300, 200, and

104 120 µm culet size. Sample configuration was similar to that reported in Tagawa et al. 105 (2016). Inside the preindented 20–25 µm thick rhenium (Re) gasket, the NaCl inner gasket 106 prepared with a Focused Ion Beam (FIB) was employed to prevent H loss to the Re gasket. 107 The surface of the diamond anvils was coated with titanium such that we could avoid the 108 failure of diamond anvils (Ohta et al., 2015). We loaded a ~10 µm thick pure Fe foil 109 (>99.999% purity, Toho Zinc) between the NaCl plates, which served not only as a 110 pressure medium but also as a pressure marker and a hydrogen insulator. After drying a 111 whole DAC in an oven, we cryogenically loaded liquid H₂ at temperatures below 20 K 112 (Chi et al., 2011; Tagawa et al., 2016). The sample was then weakly compressed and 113 brought back to room temperature. After it was further compressed to 15-30 GPa, double 114 hcp (dhcp) (\pm fcc) stoichiometric FeH was synthesized by laser heating Fe + H₂ to ~1000 115 K under hydrogen-saturated conditions (Hirao et al., 2004; Tagawa et al., 2016; Kato et 116 al., 2020) (dhcp FeH was sometimes obtained without heating). Subsequently pressure 117 was fully released under liquid nitrogen temperature (~85 K) in an N₂ atmosphere, and 118 we opened the sample chamber to remove excess H₂ without decomposition of FeH (note 119 that FeH is metastably quenchable to 1 bar below ~200 K as demonstrated by Antonov et 120 al., 2019). The sample was repressurized to >5 GPa under cryogenic temperature and 121 further to a pressure of interest at 300 K.

122 Heating was made along with in-situ high P-T XRD measurements at the beamline 123 BL10XU, SPring-8 synchrotron facility (Hirao et al., 2020). The sample was heated from 124 both sides using a couple of 100 W single-mode Yb fiber lasers. A laser beam was 125 converted to one with a flat energy distribution by beam-shaping optics, and the laser-126 heated spot was 30-40 µm across. Temperature was measured by a spectro-radiometric 127 method. The mean sample temperature, T_{mean} , is the average over 6 μ m area at a laser-128 heated hot spot, which corresponds to the X-ray beam size, for both sides. In addition, 129 Table 1 also gives the highest temperature at the center of the hot spot, T_{max} , that is more 130 relevant to melting. The uncertainties in T_{mean} and T_{max} may be $\pm 5\%$ according to Mori et 131 al. (2017). A monochromatized incident X-ray beam with an energy of ~30 keV was 132 focused to 6 µm in diameter on a sample. Diffraction patterns were collected on a flat 133 panel detector (PerkinElmer) with exposure time of 1 sec at 300 K before and after heating 134 and at high temperatures. The XRD patterns also provided the unit-cell volume of B2-135 type NaCl, from which pressure was determined at both 300 K and high temperatures 136 (Dorogokupets & Dewaele, 2007) considering its effective temperature (Campbell et al., 2009); $T_{NaCl} = \frac{3 \times T_{mean} + 300}{4} \pm \frac{T_{mean} - 300}{4}$. This method for determining pressure at high 137

138 temperature based on the volume of the pressure medium has been validated in Tagawa

et al. (2022) by simultaneously using NaCl and KCl; since the thermal expansivity ofNaCl is much larger than that of KCl, the effect of temperature variation is larger, but

141 NaCl gave pressures at high temperatures very similar to those by KCl.

142 After melting experiments, the sample was recovered from the DAC in run #1. We

143 prepared its cross section at the center of a laser-heated spot parallel to the compression

144 axis by using a focused Ga ion beam, FIB (FEI, Versa 3D DualBeam). The melting texture

145 was examined by a field-emission-type scanning electron microscope (SEM).

146 **2.2. Thermodynamic Modeling**

147 Our thermodynamic model is based on low-pressure metallurgical models for Fe-Sm 148 hydride and Fe-Ti hydride alloys (Zinkevich et al., 2002; Kivilahti & Miettinen, 1987) 149 that we extended to high pressures. Solid iron hydrides FeH_x are not stoichiometric, 150 exhibiting H contents x in the range $0 \le x \le 1$ and possibly x > 1 (Fukai, 1992; Sugimoto

151 & Fukai, 1992; Hirose et al., 2019; Ikuta et al., 2019).

We model the melting experiments by calculating the *P* and *T* trajectory of the liquidsolid boundary at constant composition. Here, the free energy G(P,T) of the solid fcc FeH (*s*) and liquid FeH (*l*) are equal, or

$$\Delta G = G^l - G^s = 0. \tag{1}$$

Following Helffrich & Connolly (2009), G's P and T dependence is separated into a room pressure (1 bar) T dependence given by polynomial expressions in T (Zinkevich et al., 2002) and a Birch-Murnaghan dependence on P which is third-order in finite strain $f = (1/2)[(V_0/V)^{2/3} - 1]$. Fcc FeH also requires the magnetic contribution to free energy, Gmag. For a phase ϕ ,

161
$$G^{\phi}(P,T) = G^{\phi}(T) + \int_{0}^{P} V_{x}^{\phi}(T,p) \, dp + G_{\text{mag}}^{\phi}$$
(2)

162 V^{ϕ_x} is a composition-dependent volume that depends on *x*, the stoichiometric coefficient 163 in FeH_x, defined as

164
$$V_x^{\phi} = V_{\text{FeH}_x}^{\phi}(x) = V_{\text{Fe}}^{\phi} + xV_{\text{H}}^{\phi}.$$
 (3)

165 A polynomial representation for the thermal expansivity, $\alpha(T)$, provides the temperature 166 part of the volume dependence (see Table 2), which is itself composition dependent. The 167 pressure contribution to the free energy in equation (2) is evaluated through a 168 compositionally-dependent bulk modulus K_x and its pressure derivative K_x ' (Table 2). 169 The assumption that $K = \delta_T$ provides *K*'s temperature dependence (Helffrich & Connolly, 170 2009),

171

180

$$K_{x}(T) = K_{x,0} exp\left[-\delta_{T} \int_{T_{0}}^{T} \alpha_{x}(t) dt\right]$$
(4)

172 The Birch-Murnaghan relation's *P* dependence is implicit in

173
$$P = 3K_x(T)f(1+2f)^{5/2}(1+2\psi f),$$
 (5)

174 with $\psi = 3(K'_x - 4)/4$. The pressure contribution to *G* is obtained by finding *f* for a 175 given *P* and then integrating (5) by parts,

176
$$\int V dP = [PV]_{0,T}^{P,T} - \int P dV = [PV]_{0,T}^{P,T} - \int P \frac{dV}{df} df.$$
 (6)

177 The magnetic contribution to the free energy depends on the compositionally dependent

178 Curie/Néel temperature $T_{C,x}$ and magnetic moment β_x , and a magnetic enthalpy parameter

179 *p* that only depends on structure

$$G_{\max(x)} = RT \log(\beta_x + 1) \times g(T/T_{C,x}, p)$$
(7)

181 whose rationale and expression for g may be found in Hillert & Jarl (1978).

182 The polynomials for the 1 bar G temperature dependence are from Zinkevich et al. (2002). To refine the physical properties of pure iron, we use those data and modify the volume 183 184 integral contributions to fit a suite of experimental constraints. For iron, we use 185 Komabayashi & Fei's (2010) experimental data points for the bcc-fcc boundary, the body-186 centered cubic (bcc)-hcp boundary, and the fcc-hcp boundary, and the fcc-hcp boundary 187 from Komabayashi (2014), and the experimental melting points below the liquid-fcc-hcp 188 triple point from Anzellini et al. (2013). We use a nonlinear minimization scheme (R Core 189 Team, 2018) for the liquid parameters V_0 , α , K, and K that yield $\Delta G = 0$ at the 190 experimentally determined P and T conditions, plus auxiliary constraints that prevent bcc 191 and fcc iron from becoming stable at inner core conditions.

192 G(T) for compounds like FeH_x are built from element contributions and additional terms 193 (Zinkevich et al. 2002). Since we are only interested in the FeH composition, we can 194 simplify these terms to the endmember stoichiometric component FeH and dispense 195 with the complexity of nonideal mixing.

- 196 Hydrogen solubility in liquid Fe is more complex. Previous eutectic melting experiments
- 197 in Fe-H (Hirose et al., 2019) suggested that melts yielded FeH_x liquid compositions with
- 198 $0 \le x \le 2$. Our experiments (Table 1) provide iso-composition melting points of fcc FeH,
- 199 so we specialize the liquid properties to that of x = 1. Along with Sakamaki et al.'s (2009)
- 200 data for fcc FeH_x (with $x \approx 1$), these provide a way to estimate the required thermophysical

201	properties of FeH _x ($x \le 1$) liquids — V_0 , α , K , and K — and for solid fcc FeH — β_x and	ıd
202	T_{Cx} . The values are listed in Table 2.	

203 **3. Results**

204 Melting was detected at six P-T conditions from 43 and 152 GPa in five separate runs 205 (Table 1). In each experiment, the dhcp FeH sample underwent a full transformation to 206 the fcc structure when it was first heated to >1500 K, consistent with previous theoretical 207 and experimental studies (Isaev et al., 2007; Thompson et al., 2018; Kato et al., 2020). 208 Neither FeH₂ nor FeH₃ was formed upon heating above 60 GPa (Pépin et al., 2014), which 209 proved no excess hydrogen remained in the sample chamber. As demonstrated by earlier 210 total energy calculations (Tagawa et al., 2022), fcc stoichiometric FeH is non-magnetic 211 in this pressure range while it is magnetic at lower pressures. The previous experiments 212 by Tagawa et al. (2022) showed that fcc FeH becomes less compressible above 41 GPa 213 at room temperature, supporting the magnetic transition to non-magnetic. We confirmed 214 each time at 300 K before and after melting the sample that the unit-cell volumes of the 215 sample, including those of crystals that formed upon quenching temperature, were 216 consistent with the compression curve of stoichiometric FeH (Tagawa et al., 2022).

In the first three runs, we determined the melting temperature of stoichiometric FeH while collecting the *P-V-T* data for the fcc phase. Such *P-V-T* data obtained in these three experiments and the high-temperature equation of state of fcc non-magnetic FeH have been reported in Tagawa et al. (2022) (the same run numbers were used). The *P-T* conditions at which the volume data were collected for the solid sample are plotted in Figure 2 (see small blue circles).

223 In run #1, we started heating at 62 GPa and 1800 K and observed a complete 224 transformation from dhcp to fcc FeH (Figure 3a). We increased temperature stepwise to 225 $T_{\text{max}} = 2480$ K and then rapidly quenched to 300 K by shutting down the laser beam. The 226 unit-cell volume of the sample found at room temperature was on the compression curve 227 of non-magnetic fcc FeH (Tagawa et al., 2022). Subsequently this sample was heated to 228 higher temperature, $T_{\text{max}} = 2740$ K. Upon quenching temperature, additional "hazy" 229 diffraction signals appeared close to the fcc peaks in the 2D and 1D XRD spectra (Figures 230 3c, d), which was not observed when quenching the sample from $T_{\text{max}} = 2480$ K (Figure 231 3b). This additional signal may have originated from incomplete fcc-like crystals with 232 irregular planar stacking that were formed from liquid upon rapid temperature quenching. 233 We recovered this sample and examined its texture on its cross section at the center of a 234 laser-heated spot. The NaCl grains were found within FeH metal, which clearly indicates

- that the sample was molten (Figure 4). These observations indicate that the sample did
 not melt to 2480 K but melted at 2740 K at 67–69 GPa (Figure 2). We also noted that FeH
 melted congruently (liquid FeH coexisted with solid FeH) since no Fe-H alloys other than
 stoichiometric FeH formed from liquid upon quenching temperature; liquid Fe-H
 crystallized Fe-H with almost equivalent H concentration at quenching (Hirose et al.,
 2019; Tagawa et al., 2021).
- 241 Similarly in runs #2-#5, we explored the melting of stoichiometric FeH, by employing 242 the appearance of the additional hazy diffraction signal observed upon quenching 243 temperature as a melting criterion. In run #2, the sample was compressed to 138 GPa and 244 then heated. The hazy XRD signal was not observed when the sample temperature was 245 quenched from $T_{\text{max}} = 3900$ K at 149 GPa. It was detected in the next heating cycle when 246 the sample was quenched from $T_{\text{max}} = 4630$ K at 152 GPa, the highest *P*-*T* condition in 247 the present study (Figure 2). In run #3, we collected the volume data for solid fcc FeH in 248 a wide P-T range from 40 GPa/1600 K to 136 GPa/3600 K, repeating a number of heating 249 cycles (Tagawa et al., 2022). Melting of the sample was not observed when quenched 250 from 3560 K/113 GPa and 3610 K/135 GPa. On the other hand, the hazy signal from 251 quenched liquid was obtained when the sample was heated to 3860 K/121 GPa and 4210 252 K/142 GPa. Note that since FeH melted congruently, melting did not cause heterogeneity 253 in hydrogen concentration in the sample. Indeed, we observed only peaks from 254 stoichiometric FeH along with the hazy signal and those from the NaCl pressure medium; 255 the unit-cell volumes of FeH collected at room temperature were always plotted in a 256 single *P-V* curve (Tagawa et al., 2022).
- Runs #4 and #5 were carried out at 40–50 GPa, in order to determine the melting temperature near conditions where stoichiometric FeH loses its local spin moment (Tagawa et al., 2022). In run #4, we observed no signs of melting when quenching from $T_{max} = 1830$ K at 43 GPa but clearly obtained the additional hazy XRD signal after heating to $T_{max} = 2110$ K at 45 GPa (Figure 2). We rapidly quenched the sample only when heating
- 262 to $T_{\text{max}} = 2170$ K at 47 GPa in run #5.
- The melting curve of FeH was obtained by fitting the thermodynamic model to these notmelting and melting data along with those previously reported below 20 GPa (Sakamaki et al., 2009). Its dT/dP slope becomes substantially larger due to the loss of magnetism above ~40 GPa (Tagawa et al., 2022). The melting point of non-magnetic FeH is higher than the eutectic melting temperature of FeH_x (1 < x < 2) observed between 43 and 127 GPa (Hirose et al., 2019) (Figure 2). Note that the melting temperature of magnetic FeH found below 20 GPa corresponds to the eutectic temperature between FeH and H₂ (Fukai,

270 1992) (Figure 1a).

271 We do not model a change in volume at the magnetic to non-magnetic transition because

- it was not clearly observed in the compression curve of FeH reported by Tagawa et al.
- 273 (2022). Omitting this detail simplifies the model at the cost of poorer constraints on fcc
- 274 FeH's magnetic properties and the liquid bulk modulus: μ_B and K trade off since the
- 275 melting curve slope is $dP/dT = \Delta S/\Delta V$ and μ_B affects ΔS .

276 **4. Discussion**

277 The present experiments suggest that the loss of local spin moment stabilizes solid fcc 278 FeH to higher temperatures. FeH is not stable at ambient pressure because of the 279 negligible solubility of hydrogen in iron (e.g., Fukai & Suzuki, 1986). Above ~5 GPa 280 where FeH is stabilized (Sakamaki et al., 2009), Fe and FeH probably form a continuous 281 solid solution in the fcc structure (Fukai, 1992; Shibazaki et al., 2014) (Figure 1a). At 282 pressures greater than ~15 GPa, hcp Fe appears instead of the bcc phase at relatively low 283 temperatures, which complicates the Fe-FeH phase diagram (Figure 1b). Stoichiometric 284 FeH becomes non-magnetic at 41 GPa at 300 K (Tagawa et al., 2022). The melting 285 temperature of the non-magnetic phase increases rapidly with increasing pressure. The 286 present XRD observations of the congruent melting of FeH, along with different crystal 287 structures between Fe and FeH, indicate that the Fe-FeH becomes a eutectic system above 288 ~40 GPa (Figure 1c). The recent experiments by Tagawa et al. (2022) demonstrated that 289 fcc FeH is stable at least to 142 GPa and 3660 K. The stability of fcc FeH over the hcp 290 and dhcp structures has been shown by theory (Isaev et al., 2007). It is therefore likely 291 that hcp Fe and fcc FeH form a binary eutectic system also at inner core conditions.

292 The melting curve of FeH is extrapolated to inner core pressures by our thermodynamic 293 model (Figure 5a). It shows that the melting point of stoichiometric FeH is 5500 K at 330 294 GPa of the inner core boundary (ICB) pressure. Fe forms a binary eutectic system also 295 with FeSi, FeO, Fe₂S (Tateno et al., 2019), and Fe₇C₃ (Lord et al., 2009; Mashino et al., 296 2019) at the ICB pressure. The melting curve of FeH is compared to those of these Fe 297 alloys/compounds; FeSi (Lord et al., 2010), FeO (Fischer et al., 2010), FeS (Boehler, 298 1992) (instead of Fe₂S), and Fe₃C (liquidus curve, Liu et al., 2016) (instead of Fe₇C₃). 299 Note that the dT/dP slope of the melting curve of non-magnetic FeH is larger than those 300 of other endmembers except Fe₃C. On the other hand, it is similar to that of pure Fe, 301 which was obtained also by the present thermodynamic modeling (Figure 5b). FeH melts 302 at only ~350 K below Fe at both the core-mantle boundary (CMB) and the ICB.

303 The possible eutectic liquid composition in the Fe-FeH system has been estimated to be

304 $FeH_{0.42}$ (Fe + 0.75 wt% H) at ~40 GPa, based on the liquidus phase relations in Fe-O-H 305 along with the subsolidus phase equilibria of an Fe-H alloy (Oka et al., 2022). The binary 306 eutectic composition is controlled in a large part by the melting points of Fe and FeH 307 endmembers. The similar dT/dP slopes between their melting curves (Figure 5b) suggest 308 that the depression of Fe-FeH eutectic temperature compared to the melting points of Fe 309 and FeH endmembers is little dependent on pressure. The eutectic liquid composition of 310 $FeH_{0.42}$ estimated at ~40 GPa may also change little with increasing pressure. The 311 observed outer core density and velocity can be explained with liquid Fe containing 1.0 312 wt% H when hydrogen is a single light element in the core (Umemoto & Hirose, 2015, 313 2020). FeH_{0.42} (Fe + 0.75 wt% H) could be the maximum hydrogen concentration in the 314 liquid core because otherwise FeH would crystallize at the ICB and not form a solid inner 315 core that is denser than the liquid outer core (Tagawa et al., 2022).

316 5. Conclusions

317 We carried out high P-T experiments using a laser-heated DAC and determined the 318 melting curve of non-magnetic stoichiometric FeH in a pressure range from 43 and 152 319 GPa. Melting was recognized on the basis of the appearance of additional hazy signals in 320 2D XRD images upon quenching temperature, which we confirmed is consistent with the 321 observation of a melting texture on the cross section of a recovered sample. The results 322 demonstrate that the dT/dP slope of the melting curve of non-magnetic FeH is 323 substantially larger than that of the magnetic phase which was determined previously 324 below 20 GPa. Our experiments also show that FeH melts congruently at a temperature 325 maximum in the Fe-H₂ system in the pressure range explored.

326 These results indicate that the loss of local spin moment in FeH expands its stability with 327 respect to liquid. They also suggest that Fe and non-magnetic FeH form a eutectic system 328 above ~40 GPa, which is supported by the fact that these two adopt different crystal 329 structures, hcp and fcc, respectively. The dT/dP slope of the melting curve of non-330 magnetic FeH is similar to that of Fe, suggesting that the Fe-FeH eutectic liquid 331 composition, which was estimated to be $FeH_{0.42}$ (Fe + 0.75 wt% H) at ~40 GPa (Oka et 332 al., 2022), will only change to a minor extent with increasing pressure. It gives the upper 333 bound for hydrogen concentration in the outer core, while the observed density and 334 velocity allow the presence of up to 1.0 wt% H in the liquid core (Umemoto & Hirose, 335 2015, 2020).

336 Data Availability Statement

337 Data for this research are found in Tables 1 and 2 available online (from
338 https://doi.org/10.5281/zenodo.6342458).

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343 **References**

- 344 Antonov, V. E., Gurev, V. M., Kulakov, V. I., Kuzovnikov, M. A., Sholin, I. A., & Zuykova,
- 345 V. Y. (2019). Solubility of deuterium and hydrogen in fcc iron at high pressures and
- 346 temperatures. *Physical Review Materials*, *3*, 113604.
- 347 https://doi.org/10.1103/PhysRevMaterials.3.113604
- 348 Anzellini, S., Dewaele, A., Mezouar, M., Loubeyre, P., & Morard, G. (2013). Melting of
- iron at Earth's inner core boundary based on fast X-ray diffraction. *Science*, *340*, 464–
 466. https://doi.org/10.1126/science.1233514
- 351 Boehler, R. (1992). Melting of the Fe-FeO and the Fe-FeS systems at high pressure:
- Constraints on core temperatures. *Earth and Planetary Science Letters*, *111*, 217–227.
 https://doi.org/10.1016/0012-821X(92)90180-4
- 354 Campbell, A. J., Danielson, L., Righter, K., Seagle, C. T., Wang, Y., & Prakapenka, V. B.
- 355 (2009). High pressure effects on the iron-iron oxide and nickel-nickel oxide oxygen
- 356 fugacity buffers. *Earth and Planetary Science Letters*, 286, 556–564.
- 357 https://doi.org/10.1016/j.epsl.2009.07.022
- Chi, Z., Nguyen, H., Matsuoka, T., Kagayama, T., Hirao, N., Ohishi, Y., & Shimizu, K.
- 359 (2011). Cryogenic implementation of charging diamond anvil cells with H_2 and D_2 .
- 360 The Review of Scientific Instruments, 82, 105109.
- 361 https://doi.org/10.1063/1.3652981
- 362 Dinsdale, A. T. (1991). SGTE data for pure elements. *Calphad*, 15, 317–425.
- 363 https://doi.org/10.1016/0364-5916(91)90030-N
- 364 Dorogokupets, P. I., & Dewaele, A. (2007). Equations of state of MgO, Au, Pt, NaCl-B1,
- and NaCl-B2: Internally consistent high-temperature pressure scales. *High Pressure Research*, 27, 431–446. https://doi.org/10.1080/08957950701659700
- Fischer, R., & Campbell, A. J. (2010). High-pressure melting of wüstite. *American Mineralogist*, 95, 1473–1477. https://doi.org/10.2138/am.2010.3463
- 369 Fukai, Y. (1992). Some properties of the Fe-H system at high pressures and temperatures,
- and their implications for the Earth's core. In Y. Syono, M.H. Manghnani (Eds.), *High*-

- 371 pressure research: Applications to Earth and planetary sciences (Vol. 67, pp. 373–
 372 385). https://doi.org/10.1029/GM067p0373
- Fukai, Y., & Suzuki, T. (1986). Iron-water reaction under high pressure and its
 implication in the evolution of the Earth. *Journal of Geophysical Research*, *91*, 9222–
 9230. https://doi.org/10.1029/JB091iB09p09222
- He, Y., Sun, S., Kim, D.-Y., Jang, B.-G., Li, H., & Mao, H.-K. (2022). Superionic iron
 alloys and their seismic velocities in Earth's inner core. *Nature*, 602, 258–262.
- 378 https://doi.org/10.1038/s41586-021-04361-x
- Helffrich, G., & Connolly, J. A. D. (2009). Physical contradictions and remedies using
 simple polythermal equations of state. *American Mineralogist*, *94*, 1616–1619.
 https://doi.org/10.2138/am.2009.3262
- 382 Hillert, M., & Jarl, M. (1978). A model for alloying effects in ferromagnetic metals.
 383 *Calphad*, 2, 227–238. https://doi.org/10.1016/0364-5916(78)90011-1
- Hirao, N., Kondo, T., Ohtani, E., Takemura, K., & Kikegawa, T. (2004). Compression of
 iron hydride to 80 GPa and hydrogen in the Earth's inner core. *Geophysical Research Letters*, *31*, L06616. https://doi.org/10.1029/2003GL019380
- 387 Hirao, N., Kawaguchi, S. I., Hirose, K., Shimizu, K., Ohtani, E., & Ohishi, Y. (2020).
 388 New developments in high-pressure X-ray diffraction beamline for diamond anvil cell
- 388 New developments in high-pressure X-ray diffraction beamline for diamond
 389 at SPring-8. *Matter and Radiation at Extremes*, 5, 1–10.
- 390 https://doi.org/10.1063/1.5126038
- 391 Hirose, K., Tagawa, S., Kuwayama, Y., Sinmyo, R., Morard, G., Ohishi, Y., & Genda, H.
- 392 (2019). Hydrogen limits carbon in liquid iron. *Geophysical Research Letters*, 46,
 393 5190–5197. https://doi.org/10.1029/2019GL082591
- Hirose, K., Wood, B., & Vočadlo, L. (2021). Light elements in the Earth's core. *Nature Reviews Earth & Environment*, 2, 645–658. https://doi.org/10.1038/s43017-021 00203-6
- 397 Ikuta, D., Ohtani, E., Sano-Furukawa, A., Shibazaki, Y., Terasaki, H., Yuan, L., & Hattori,
 398 T. (2019). Interstitial hydrogen atoms in face-centered cubic iron in the Earth's core.

399 Scientific Reports, 9, 7108. https://doi.org/10.1038/s41598-019-43601-z

- 400 Isaev, E. I., Skorodumova, N. V, Ahuja, R., Vekilov, Y. K., & Johansson, B. (2007).
- 401 Dynamical stability of Fe-H in the Earth's mantle and core regions. *Proceedings of the*
- 402 National Academy of Sciences of the United States of America, 104, 9168–9171.
- 403 https://doi.org/10.1073/pnas.0609701104
- Kato, C., Umemoto, K., Ohta, K., Tagawa, S., Hirose, K., & Ohishi, Y. (2020). Stability
 of fcc phase FeH to 137 GPa. *American Mineralogist*, *105*, 917–921.
- 406 https://doi.org/10.2138/am-2020-7153
- 407 Kivilahti, J. K., & Miettinen, J. M. (1987). A thermodynamic analysis of the Ti-H system.

408 *Calphad, 11*, 187–188.

- Komabayashi, T. (2014). Thermodynamics of melting relations in the system Fe-FeO at
 high pressure: Implications for oxygen in the Earth's core. *Journal of Geophysical Research, 119*, 4164–4177. https://doi.org/10.1002/2014JB010980.
- 412 Komabayashi, T., & Fei, Y. (2010). Internally consistent thermodynamic database for
- 413 iron to the Earth's core conditions. *Journal of Geophysical Research*, 115,
- 414 doi:10.1029/2009JB006442.
- Komabayashi, T., Fei, Y., Meng, Y., & Prakapenka, V. (2009). In-situ X-ray diffraction
 measurements of the γ-ε transition boundary of iron in an internally-heated diamond
 anvil cell. *Earth and Planetary Science Letters*, 282, 252–257.
- 418 https://doi.org/10.1016/j.epsl.2009.03.025
- Li, Y., Vočadlo, L., Sun, T., & Brodholt, J. P. (2020). The Earth's core as a reservoir of
 water. *Nature Geoscience*, *13*, 453–458. https://doi.org/10.1038/s41561-020-0578-1
- Liu, J., Lin, J., Prakapenka, V.B., Prescher, C., & Yoshino, T. (2016). Phase relations of
 Fe₃C and Fe₇C₃ up to 185 GPa and 5200 K: Implication for the stability of iron carbide
 in the Earth's core. *Geophysical Research Letters*, 43, 12415–12422. https://doi.org/
 10.1002/2016GL071353.
- 425 Lord, O. T., Walter, M. J., Dasgupta, R., Walker, D., & Clark, S. M. (2009). Melting in
- the Fe–C system to 70 GPa. *Earth and Planetary Science Letters*, 284, 157–167.
 https://doi.org/10.1016/j.epsl.2009.04.017.
- 427 https://doi.org/10.1010/j.epsi.2009.04.017.
- 428 Lord, O. T., Walter, M. J., Dobson, D. P., Armstrong, L., Clark, S. M., & Kleppe, A.
- 429 (2010). The FeSi phase diagram to 150 GPa. *Journal of Geophysical Research*, *115*,
 430 B06208. doi:10.1029/2009JB006528.
- 431 Mashino, I., Miozzi, F., Hirose, K., Morard, G., & Sinmyo, R. (2019). Melting
 432 experiments on the Fe-C binary system up to 255 GPa: Constraints on the carbon
 433 E. H. E.
- 433 content in the Earth's core. *Earth and Planetary Science Letters*, *515*, 135–144.
- 434 https://doi.org/10.1016/j.epsl.2019.03.020
- 435 Mori, Y., Ozawa, H., Hirose, K., Sinmyo, R., Tateno, S., Morard, G., & Ohishi, Y. (2017).
- 436 Melting experiments on Fe–Fe₃S system to 254 GPa. *Earth and Planetary Science*

437 *Letters*, 464, 135–141. https://doi.org/10.1016/j.epsl.2017.02.021

- 438 Ohta, K., Ichimaru, K., Einaga, M., Kawaguchi, S., Shimizu, K., Matsuoka, T., et al.
 439 (2015). Phase boundary of hot dense fluid hydrogen. *Scientific Reports*, *5*, 16560.
 440 https://doi.org/10.1038/srep16560
- 441 Oka, K., Tagawa, S., Hirose, K., & Ohishi, Y. (2022). Melting experiments on Fe-O-H:
- 442 Evidence for eutectic melting in Fe-FeH and implications for hydrogen in the core.
- 443 *Earth and Space Science Open Archive*. https://doi.org/10.1002/essoar.10510359.1
- 444 Pépin, C. M., Dewaele, A., Geneste, G., Loubeyre, P., & Mezouar, M. (2014). New iron

- 445 hydrides under high pressure. *Physical Review Letters*, *113*, 265504.
- 446 https://doi.org/10.1103/PhysRevLett.113.265504
- 447 Piet, H., Chizmeshya, A. V. G., Chen, B., Chariton, S., Greenberg, E., Prakapenka, V. B.,
- & Shim, S.-H. (2021). Effect of nickel on the high-pressure phases in FeH. *Physical Review B*, *104*, 224106. https://doi.org/10.1103/PhysRevB.104.224106
- 450 Raymond, S. N., Quinn, T., & Lunine, J. I. (2007). High-resolution simulations of the
 451 final assembly of Earth-like planets. 2. Water delivery and planetary habitability.
 452 Astrobiology, 7, 66–84. https://doi.org/10.1089/ast.2006.06-0126.
- R Core Team (2018). *R: A language and environment for statistical computing*. R
 Foundation for Statistical Computing. Vienna, Austria.
- 455 Sakamaki, K., Takahashi, E., Nakajima, Y., Nishihara, Y., Funakoshi, K., Suzuki, T., &
- 456 Fukai, Y. (2009). Melting phase relation of FeH_x up to 20 GPa: Implication for the
- 457 temperature of the Earth's core. *Physics of the Earth and Planetary Interiors*, 174,
- 458 192–201. https://doi.org/10.1016/j.pepi.2008.05.017
- 459 Sato, T., Okuzumi, S., & Ida, S. (2016). On the water delivery to terrestrial embryos by
 460 ice pebble accretion. *Astronomy & Astrophysics*, 589, A15.
- 461 https://doi.org/10.1051/0004-6361/201527069
- 462 Shibazaki, Y., Terasaki, H., Ohtani, E., Tateyama, R., Nishida, K., Funakoshi, K., & Higo,
- 463 Y. (2014). High-pressure and high-temperature phase diagram for Fe_{0.9}Ni_{0.1}-H alloy.
- 464 *Physics of the Earth and Planetary Interiors*, 228, 192–201.
- 465 https://doi.org/10.1016/j.pepi.2013.12.013
- 466 Stähler, S. C., Khan, A., Banerdt, W. B., Lognonné, P., Giardini, D., Ceylan, S., et al.
- 467 (2021). Seismic detection of the martian core. *Science*, *373*, 443–448.
- 468 https://doi.org/10.1126/science.abi7730
- 469 Sugimoto, H., & Fukai, Y. (1992). Solubility of hydrogen in metals under high hydrogen
- 470 pressures: thermodynamical calculations. *Acta Metallurgica et Materialia*, 40, 2327–
 471 2336. https://doi.org/10.1016/0956-7151(92)90151-4
- 472 Tagawa, S., Ohta, K., Hirose, K., Kato, C., & Ohishi, Y. (2016). Compression of Fe-Si-H
- 473 alloys to core pressures. *Geophysical Research Letters*, *43*, 3686–3692.
- 474 https://doi.org/10.1002/2016GL068848
- 475 Tagawa, S., Sakamoto, N., Hirose, K., Yokoo, S., Hernlund, J., Ohishi, Y., & Yurimoto,
- 476 H. (2021). Experimental evidence for hydrogen incorporation into Earth's core. *Nature*477 *Communications*, *12*, 2588. https://doi.org/10.1038/s41467-021-22035-0
- 478 Tagawa, S., Gomi, H., Hirose, K., & Ohishi, Y. (2022). High-temperature equation of
- 479 state of FeH: Implications for hydrogen in Earth's inner core. *Geophysical Research*
- 480 Letters, 49, e2021GL096260. https://doi. org/10.1029/2021GL096260
- 481 Tateno, S., Ozawa, H., Hirose, K., Suzuki, T., I-Kawaguchi, S., & Hirao, N. (2019). Fe₂S:

- The most Fe-rich iron sulfide at the Earth's inner core pressures. *Geophysical Research Letters*, 46, 11944–11949. https://doi.org/ 10.1029/2019GL085248
- 484 Thompson, E. C., Davis, A. H., Bi, W., Zhao, J., Alp, E. E., Zhang, D., et al. (2018). High-
- 485 pressure geophysical properties of fcc phase FeH_x. *Geochemistry, Geophysics,* 486 *Geosystems, 19,* 305-314. https://doi.org/10.1002/2017GC007168
- 487 Umemoto, K., & Hirose, K. (2015). Liquid iron-hydrogen alloys at outer core conditions
 488 by first-principles calculations. *Geophysical Research Letters*, 42, 7513–7520.
- 489 https://doi.org/10.1002/2015GL065899
- 490 Umemoto, K., & Hirose, K. (2020). Chemical compositions of the outer core examined
 491 by first principles calculations. *Earth and Planetary Science Letters*, *531*, 116009.

492 https://doi.org/10.1016/j.epsl.2019.116009

- Walsh, K. J., Morbidelli, A., Raymond, S. N., O'brien, D. P., & Mandell, A. M. (2011). A
 low mass for Mars from Jupiter's early gas-driven migration. *Nature*, 475, 206–209.
 https://doi.org/10.1038/nature10201
- Wang, W., Li, Y., Brodholt, J. P., Vočadlo, L., Walter, M. J., & Wu, Z. (2021). Strong shear
 softening induced by superionic hydrogen in Earth's inner core. *Earth and Planetary*
- 498 Science Letters, 568, 117014. https://doi.org/10.1016/j.epsl.2021.117014
- 499 Yoshizaki, T., & McDonough, W. F. (2020). The composition of Mars. *Geochimica et*500 *Cosmochimica Acta*, 273, 137–162. https://doi.org/10.1016/j.gca.2020.01.011
- 501 Yuan, L., & Steinle-Neumann, G. (2020). Strong sequestration of hydrogen into the
- 502 Earth's core during planetary differentiation. *Geophysical Research Letters*, 47,
 503 e2020GL088303. https://doi.org/10.1029/2020GL088303
- 504 Zinkevich, M., Mattern, M., Handstein, A., & Gutfleisch, O. (2002). Thermodynamics of
- 505 Fe-Sm, Fe-H, and H-Sm systems and its application to the hydrogen-
- 506 disproportionation-desorption-recombination (HDDR) process for the system
- 507 Fe₁₇Sm₂-H₂. *Journal of Alloys and Compounds*, *339*, 118–139.
- 508 https://doi.org/10.1016/S0925-8388(01)01990-9

Run #	Pressure (GPa)	Temperature (K)	
		$T_{\rm max}$	T_{mean}
Melting			
1	69 (7)	2740 (120)	2520 (110)
2	152 (15)	4630 (220)	4490 (210)
3	121 (12)	3860 (180)	3630 (170)
	142 (14)	4210 (200)	3970 (180)
4	45 (4)	2110 (90)	2060 (90)
5	47 (5)	2170 (90)	2100 (90)
Not-me	lting		
1	67 (7)	2480 (110)	2340 (100)
2	149 (15)	3900 (180)	3550 (160)
3	113 (11)	3560 (160)	3410 (160)
	135 (13)	3610 (170)	3500 (160)
4	43 (4)	1830 (80)	1750 (70)

Table 1

Table 2 Physical properties								
Property	Value	Scale and Units	Source					
1 5			(if not this study)					
V_0 volume	V_0 volume							
(Fe bcc)	0.71000	J bar ^{-1}						
(Fe fcc)	0.68640	$J \text{ bar}^{-1}$						
(Fe hcp)	0.67357	J bar ^{-1}						
(Fe lia)	0.70205	J bar ⁻¹						
(H fcc)	0.12861	$I \text{ bar}^{-1} (\text{mol at H})^{-1}$	А					
(H lia)	0.06437	I bar ^{-1} (mol at H) ^{-1}						
α thermal expan	sivity							
h_{\perp} (Fe bcc)	3 05500	$10^{-6} K^{-1}$						
$b_{\rm T}$ (Fe bee)	0.80000	10^{-3}						
b_3 (Fe bee)	-9.80000	$10 10^{-6} V^{-1}$						
b_1 (Fe Icc)	9.4310	10 K 10^{-5} K^{-1}						
b_1 (Fe hep)	2 7720	10^{-5} K^{-1}						
b_1 (Fe liq) b_2 (H fcc)	7 989	10 K $\times 10^{-6} \text{ K}^{-1} \text{ (mol at H)}^{-1}$	Δ					
$b_{\rm T}$ (H fcc)	2 093	×10 K (mol. at. 11) ×10 ⁻⁸ K^{-2} (mol. at. H) ⁻¹	A					
b_2 (II Icc)	7 385	×10 K (1101. at. 11) ×10 ⁻⁵ K^{-1} (mol. at. H) ⁻¹	11					
$U_1(\Pi \Pi q)$ K bulk modulus	7.505	$\times 10$ K (1101. at. 11)						
(Fe bcc)	164	GPa	D					
(Fe fcc)	165 3	GPa	D					
(Fe hcn)	165	GPa	D					
(Fe lig)	148	GPa	D					
(I t liq) (H fcc)	-2 75	$GPa (mol at H)^{-1}$	А					
(H lia)	110.86	GPa (mol. at. H) $^{-1}$	11					
K' bulk modulu	s pressure derivative							
(Fe hcc)	5 29		D					
(Fe fcc)	5.5		D					
(Fe hcp)	4 97		D					
(Fe lig)	6 39		D					
(H fcc)	-0.99	$(mol at H)^{-1}$	А					
(H lia)	-1.90	(mol. at. H) $(mol. at. H)^{-1}$						
T_C Curie tempe	T_c Curie temperature							
(Fe bcc)	1043	К	В					
(Fe fcc)	67	K	B					
(H fcc)	2242	K						
β magnetic mon	nent							
(Fe bcc)	2.22	$\mu_{\rm B}$	В					
(Fe fcc)	0.7	$\mu_{\rm B}$	В					
(H fcc)	1.45	$\mu_{\rm B}$						
p magnetic enthalpy fraction								
(all fcc)	0.28		С					
(all bcc)	0.4		С					
Notes a (T	$(x) = h_1(x) + h_2(x)$	$T + h_{2}(\mathbf{r})/T \cdot h_{2}(\mathbf{r}) = h_{12}$	+ rh					

Notes. $\alpha_{\text{FeHx}}(T, x) = b_1(x) + b_2(x)T + b_3(x)/T$; $b_i(x) = b_{i,\text{Fe}} + xb_{i,\text{H}}$

 $V_{\text{FeHx}}(x) = V_{\text{Fe}} + xV_{\text{H}}; K_{\text{FeHx}}(x) = K_{\text{Fe}} + xK_{\text{H}}; K'_{\text{FeH}}x(x) = K'_{\text{Fe}} + xK'_{\text{H}}$

 $T_{C,x} = T_{C,\text{Fe}} + xT_{C,\text{H}}; \beta_x = \beta_{\text{Fe}} + x\beta_{\text{H}}.$

 $\mu_B\,$ - Bohr magneton unit.

Sources: A - Tagawa et al. (2022); B - Dinsdale (1991); C - Hillert & Jarl (1978); D - Komabayashi & Fei (2010).



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516 Figure 1. Changes in the Fe-H phase diagram with increasing pressure. (a) Fe and FeH 517 form continuous solid solution at ~10 GPa, where the solidus temperature of magnetic 518 FeH corresponds to the eutectic temperature between FeH and H₂ (Fukai, 1992). (b) The 519 appearance of the hcp phase on the Fe-rich portion leads to the gap in solid solution and 520 the peritectic points between Fe and FeH. (c) FeH loses the local spin moment and melts 521 congruently with a temperature maximum above ~40 GPa. Fe and FeH form eutectic 522 melting with the eutectic liquid composition of $FeH_{0.42}$ (Oka et al., 2022). (d) The 523 formation of FeH₂ (Pépin et al., 2014) results in eutectic melting between FeH and FeH₂. 524 The depression of Fe-FeH eutectic temperature and the eutectic liquid composition likely 525 change little with increasing pressure.



Figure 2. Experimental data and calculated melting curve of stoichiometric FeH. Present experimental conditions of melting (red, inverse triangles) and non-melting (blue, normal triangles) are plotted along with those where the unit-cell volume was determined for solid FeH in a previous study (Tagawa et al., 2022) (small blue circles). Those reported by recent experiments (Piet et al., 2021) are also shown (small squares). Dotted curve represents eutectic melting temperature between FeH and H₂/FeH₂ based on the solidus temperatures of magnetic FeH (open squares) (Sakamaki et al., 2009) and FeH_x ($1 \le x \le$ 2) (open circles) (Hirose et al., 2019). The dT/dP slope of the melting curve of non-magnetic FeH is substantially larger than that of the non-magnetic phase.



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Figure 3. XRD patterns collected in run #1. (a) Melting did not occur when heating FeH to $T_{\text{max}} = 2480$ K; additional XRD peaks were not observed upon quenching temperature to 300 K. (b) Partially molten sample was quenched from $T_{\text{max}} = 2740$ K to room temperature. The hazy diffraction signal appeared close to the fcc peaks because of the formation of incomplete crystals from liquid upon temperature quenching.





Figure 4. Scanning electron microscope image of the melting texture of FeH observed

- on the sample cross section recovered from run #1. The presence of NaCl grains (pressure
- 553 medium) in FeH metal showed that the sample was molten during heating.



Figure 5. (a) Comparison of the melting curve of FeH (gray) with those of FeSi (green) (Lord et al., 2010), FeO (brown) (Fischer et al., 2010), FeS (blue) (Boehler, 1992), and Fe₃C (yellow, liquidus curve) (Liu et al., 2016). Fe forms eutectic system with these alloys/compounds in each binary system at the ICB pressure of 330 GPa (Fe₂S and Fe₇C₃ instead of FeS and Fe₃C, respectively). (b) Calculated phase diagram for Fe (black thin lines) and FeH melting (gray bold line). Iron phase fields are labeled.