# Projected Changes and Time of Emergence of Temperature Extremes over Australia in CMIP5 and CMIP6

Xu Deng<sup>1,1</sup> and Sarah Perkins-Kirkpatrick<sup>1,1</sup>

<sup>1</sup>University of New South Wales

November 30, 2022

#### Abstract

This study focuses on the projections and time of emergence (TOE) for temperature extremes over Australian regions in the phase 6 of Coupled Model Intercomparison Project (CMIP6) models. The model outputs are based on the Shared Socioeconomic Pathways (SSPs) from the Tier 1 experiments (i.e., SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5) in the Scenario Model Intercomparison Project (ScenarioMIP), which is compared with the Representative Concentration Pathways (RCPs) in CMIP5 (i.e., RCP2.6, RCP4.5 and RCP8.5). Furthermore, two large ensembles (LEs) in CMIP6 are used to investigate the effects of internal variability on the projected changes and TOE. As shown in the temporal evolution and spatial distribution, the strongest warming levels are projected under the highest future scenario and the changes for some extremes follow a "warm-get-warmer" pattern over Australia. Over subregions, tropical Australia usually shows the highest warming. Compared to the RCPs in CMIP5, the multi-model medians in SSPs are higher for some indices and commonly exhibit wider spreads, likely related to the different forcings and higher climate sensitivity in a subset of the CMIP6 models. Based on a signal-to-noise framework, we confirm that the emergence patterns differ greatly for different extreme indices and the large uncertainty in TOE can result from the inter-model ranges of both signal and noise, for which internal variability contributes to the determination of the signal. We further demonstrate that the internally-generated variations influence the noise. Our findings can provide useful information for mitigation strategies and adaptation planning over Australia.

1	Projected Changes and Time of Emergence of Temperature Extremes over Australia
2	in CMIP5 and CMIP6
3	
4	Xu Deng <sup>1,2*</sup> and Sarah E. Perkins-Kirkpatrick <sup>1,2</sup>
5	<sup>1</sup> School of Science, University of New South Wales, Canberra, ACT, Australia.
6 7	<sup>2</sup> ARC Centre of Excellence for Climate Extremes, University of New South Wales, Canberra, ACT, Australia.
8	
9	Corresponding author: Xu Deng (xu.deng@student.adfa.edu.au)
10	
11	
12	Key Points:
13 14	• There indicates a "warm-get-warmer" pattern for some extremes over Australia and tropical regions usually show the highest warming
15 16	• Compared to CMIP5, the higher warming for some extremes in CMIP6 can lead to earlier time of emergence under the highest scenario
17 18	• Internal variability influences the determination of the noise

### 19 Abstract

This study focuses on the projections and time of emergence (TOE) for temperature extremes 20 over Australian regions in the phase 6 of Coupled Model Intercomparison Project (CMIP6) 21 models. The model outputs are based on the Shared Socioeconomic Pathways (SSPs) from the 22 Tier 1 experiments (i.e., SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5) in the Scenario Model 23 Intercomparison Project (ScenarioMIP), which is compared with the Representative 24 25 Concentration Pathways (RCPs) in CMIP5 (i.e., RCP2.6, RCP4.5 and RCP8.5). Furthermore, two large ensembles (LEs) in CMIP6 are used to investigate the effects of internal variability on 26 the projected changes and TOE. As shown in the temporal evolution and spatial distribution, the 27 28 strongest warming levels are projected under the highest future scenario and the changes for some extremes follow a "warm-get-warmer" pattern over Australia. Over subregions, tropical 29 Australia usually shows the highest warming. Compared to the RCPs in CMIP5, the multi-model 30 medians in SSPs are higher for some indices and commonly exhibit wider spreads, likely related 31 to the different forcings and higher climate sensitivity in a subset of the CMIP6 models. Based 32 on a signal-to-noise framework, we confirm that the emergence patterns differ greatly for 33 different extreme indices and the large uncertainty in TOE can result from the inter-model ranges 34 of both signal and noise, for which internal variability contributes to the determination of the 35 36 signal. We further demonstrate that the internally-generated variations influence the noise. Our findings can provide useful information for mitigation strategies and adaptation planning over 37 Australia. 38

# 40 **1 Introduction**

41	Anthropogenic climate change will lead to more severe temperature extremes, which
42	have significant impacts on society and natural systems (Intergovernmental Panel on Climate
43	Change, 2021). To assess possible climate futures, projections by global climate models from the
44	Scenario Model Intercomparison Project (ScenarioMIP; O'Neill et al., 2016) as part of the
45	Coupled Model Intercomparison Project phase 6 (CMIP6; Eyring et al., 2016) are useful
46	resources, and may provide new insights into how temperature extremes are projected to change
47	under climate change (e.g., Alexander & Arblaster, 2017; Grose et al., 2020; Sillmann, Kharin,
48	Zwiers, et al., 2013; Thibeault & Seth, 2014).
49	Over Australia, Alexander and Arblaster (2017) indicated that significant increases
50	(decreases) are projected for the occurrence of warm (cold) extremes by the end of this century
51	under the intermediate- and highest-emission scenarios in CMIP5, and that these changes are
52	most distinct in the tropics. Compared to 29 CMIP5 models, Grose et al. (2020) documented that
53	projected changes in temperature extremes over Australia are more distinct and span narrower
54	ranges in seven CMIP6 models. However, the smaller number of models used in this study may
55	lead to misleading conclusions. Recently, Tebaldi et al. (2021) demonstrated that the CMIP6
56	ensemble projects higher warming and larger spread for global mean temperature compared with
57	CMIP5, which could result from both a wider range of radiative forcing and higher climate
58	sensitivity in a subset of CMIP6 models. In the present study, to obtain a more reasonable
59	comparison with CMIP5, more models are included in the CMIP6 ensemble to analyze the
60	projected changes of temperature extremes over Australia.

In addition, detecting the time of emergence (TOE) for extremes over Australia needs
investigation. TOE is defined as the time when the externally forced climate signal (i.e., forced

response) emerges from the noise (i.e., natural variability), suggesting that a significant change is 63 detected and a novel climate regime become evident (e.g., Hawkins et al., 2020; Hawkins & 64 Sutton, 2012; King, Donat, et al., 2015). Estimating TOE can provide insights for mitigation 65 strategies, adaptation planning and scientific community, as the forced response relative to the 66 background noise may be more relevant for the assessment of climate impacts, compared to the 67 68 absolute change (Beaumont et al., 2011; Deutsch et al., 2008; Hawkins et al., 2020; Hawkins & Sutton, 2012; Ossó et al., 2021). For example, similar absolute changes in extreme temperature 69 can result in different ecological impacts since extratropical ecosystems are usually more 70 71 resilient than tropical ecosystems, as they are adapted to a more variable climate (Beaumont et al., 2011; Deutsch et al., 2008). 72

Previous studies have concluded that for mean temperature there is earlier TOE over 73 tropical regions than that in the extratropics where the noise is generally larger (e.g., Giorgi & 74 Bi, 2009; Hawkins et al., 2020; Hawkins & Sutton, 2012; Mahlstein et al., 2012; Mahlstein et al., 75 76 2011). Furthermore, for warm and cold extremes that display larger variability, the signals for these indices tend to emerge later over both the tropics and extratropics (e.g., King, Donat, et al., 77 2015; Tan et al., 2018) relative to mean temperature. Currently, most studies on TOE have been 78 79 conducted at global levels, with less detailed analyses over smaller-scale regions (e.g., Batibeniz 80 et al., 2020; Gaetani et al., 2020; Ossó et al., 2021), especially for Australia (King, Donat, et al., 81 2015). Under different future scenarios, we aim to investigate the TOE of extreme temperatures over Australia at the subregional scale. 82

A variety of methods have been used in TOE assessment, which can lead to a source of uncertainty (Abatzoglou et al., 2019; Gaetani et al., 2020). A recent study (Gaetani et al., 2020) found that compared to Kolmogorov-Smirnov (KS) non-parametric test (King, Donat, et al.,

86	2015), the signal-to-noise ratio (SNR) frameworks exhibit increased uncertainty and later times
87	for TOE over West Africa (Gaetani et al., 2020). However, the SNR methods facilitate the
88	separation between signal and noise, and identifying both components and their interaction
89	physically (e.g., slow-varying ocean conditions and the modes of internal variability) can deepen
90	our understanding in climate change (e.g., Barnes et al., 2019; Barsugli & Battisti, 1998). In this
91	study, we adopt the method by Hawkins and Sutton (2012) and Hawkins et al. (2020) to address
92	the TOE assessment, which is widely used and allows more cross-study comparisons (e.g.,
93	Abatzoglou et al., 2019; Gaetani et al., 2020; Hawkins et al., 2020; Hawkins & Sutton, 2012;
94	Ossó et al., 2021). For the uncertainty in the detection of TOE in this method, it can arise from
95	inter-model spread not only in the signal, but also from noise (Hawkins & Sutton, 2012).
96	Furthermore, as internal variability can also be an important source of uncertainty for
97	regional climate (Dai & Bloecker, 2019; Deser, Knutti, et al., 2012; Deser, Phillips, et al., 2012;
98	Hawkins & Sutton, 2009; Lehner et al., 2020), single-model initial-condition large ensembles
99	(SMILEs; hereafter LEs) are an important tool to investigate the consequences of the intrinsic
100	variability on the uncertainty in projected changes and TOE of extreme temperatures over
101	Australia, of which external forcing and model structure are identical among the members (e.g.,
102	Dai & Bloecker, 2019; Deser, 2020; Deser et al., 2020; Lehner et al., 2020; Mankin et al., 2020;
103	Perkins-Kirkpatrick et al., 2017; Xie et al., 2015).
104	Previous research evaluated the ability of CMIP6 models to simulate extreme

Previous research evaluated the ability of CMIP6 models to simulate extreme temperatures over Australian regions in the historical period (1950-2014), compared these results to the CMIP5 ensemble, and investigated the effects of internal variability on the corresponding trends based on the LEs in CMIP6 (Deng et al., 2021). Following from this research, the purposes of this study are: to assess future climate changes of the extremes and the TOE over Australian regions in both the CMIP6 and CMIP5 models, and to explore the effects of internal
variability on the projected changes and TOE based on LEs in CMIP6.

111

#### 112 **2 Data and Methods**

113 2.1 Model Data

Although the scenarios in the ScenarioMIP consist of two tiers, we only use the Tier 1 114 experiments based on the Shared Socioeconomic Pathway (SSP) scenarios: SSP1-2.6, SSP2-4.5, 115 SSP3-7.0 and SSP5-8.5, as these sample a varying range of possible emission futures and contain 116 117 relatively large number of model outputs. Among them, SSP1-2.6, SSP2-4.5 and SSP5-8.5 indicate the same nominal stratospheric-adjusted radiative forcing  $(2.6, 4.5 \text{ and } 8.5 \text{ W m}^{-2})$ 118 reached in 2100, compared to the scenarios based on Representative Concentration Pathways 119 (RCPs) used in CMIP5 (i.e., RCP2.6, RCP4.5 and RCP8.5); and SSP3-7.0 fills a gap between 120 medium and high end in the range of future forcing pathways, not included in previous CMIP 121 generations (O'Neill et al., 2016; Tebaldi et al., 2021). Despite the similarity among the future 122 scenarios in CMIP6 and CMIP5, it is noted that there are some differences, such as the 123 composition of some radiatively active gases or species (e.g.,  $CO_2$  and  $CH_4$ ) and aerosol 124 emissions, making the resulting effective radiative forcing (ERF) different (Lurton et al., 2020; 125 Riahi et al., 2017; Tebaldi et al., 2021). 126

As one aim of this study is to compare the two CMIP ensembles in projected changes and TOE in extremes, we do not consider the interdependence among the models and use emergent constraints or any other ways of model weighting to reduce the differences between CMIP6 and CMIP5 (e.g., Tokarska et al., 2020), which is similar to the practice by Seneviratne and Hauser (2020). Similar to Deng et al. (2021), only one ensemble member (typically the first member) in

131

132	each model is considered for the main part of analysis. There are 25 models in CMIP6 and 26
133	models in CMIP5 for at least one of the future scenarios. In addition, two LEs under SSP5-8.5
134	and SSP1-2.6 in CMIP6 are used to investigate the impacts of internal variability on the
135	projected changes and TOE of the extremes: CanESM5-LE and MIROC6-LE, which contain 25
136	members and 50 members, respectively. Detailed information on the simulations from CMIP6
137	and CMIP5 models are listed in the Tables S1 and S2, respectively.
138	2.2 Temperature indices
139	As in Deng et al. (2021), based on daily maximum and minimum temperatures (TX and
140	TN), the annualized temperature extremes defined by the Expert Team on Climate Change
141	Detection and Indices (ETCCDI; Zhang et al., 2011) are used, which forms a continuous and
142	comprehensive investigation of changes in extremes, similar to other studies for CMIP5 (e.g.,
143	Alexander & Arblaster, 2017; Sillmann, Kharin, Zhang, et al., 2013; Sillmann, Kharin, Zwiers, et
144	al., 2013; Thibeault & Seth, 2014). Besides diurnal temperature range (DTR), other extreme
145	indices for temperatures are classified into four categories: absolute indices (hottest day [TXx],
146	coldest day [TXn], warmest night [TNx] and coldest night [TNn]), threshold indices (summer
147	days [SU], tropical nights [TR] and frost days [FD]), percentile-based indices (warm days
148	[TX90p], cold days [TX10p], warm nights [TN90p] and cold nights [TN10p]), and duration
149	indices (warm spell duration index [WSDI] and cold spell duration index [CSDI]). The bootstrap
150	resampling procedure by Zhang et al. (2005) is applied to the percentile-based and duration

151 indices, among which the spells crossing year boundaries are taken into consideration for WSDI

and CSDI. Since the definitions of growing season length (GSL) and ice days (ID) are not

suitable over most of Australia (Alexander & Arblaster, 2017), we do not use them in this study.
Detailed information on the indices can be found in Table S3.

155 2.3 Time of Emergence

The TOE is determined using the signal-to-noise framework as detailed by Hawkins and Sutton (2012) and Hawkins et al. (2020), which is considered as the first year when the signal-tonoise ratio (SNR) is larger than nominated thresholds (e.g., 1 and 2). As suggested by Frame et al. (2017), we consider SNR=1 as the threshold for an "unusual" climate and SNR=2 as "unfamiliar". This approach linearly regresses annual local variations in temperature extremes onto global mean surface temperature change ( $\Delta GMST$ ), relative to the base period:

162 
$$\hat{L}(t) = \alpha G(t) + \beta$$

where  $\hat{L}(t)$  represents the regressed L(t), denoting annual local changes in extremes over time; 163 G(t) is a smoothed version of  $\Delta GMST$  over the same period;  $\alpha$  defines the linear scaling between 164 165  $\hat{L}(t)$  and G(t); and  $\beta$  is a constant.  $\Delta GMST$  is smoothed with a "Locally Weighted Scatterplot" Smoothing" filter (LOWESS; Cleveland, 1979) of 21 years, which filters out interannual 166 variability (though retaining multi-decadal variability). The signal of local climate change 167 described by  $\Delta GMST$  is  $\alpha G(t)$ , and the noise is defined as the standard deviation of the residuals 168  $(L(t) - \alpha G(t))$ . The method implies that local variations for some variables scale well with 169  $\Delta GMST$  (Fischer et al., 2014; Seneviratne & Hauser, 2020; Sutton et al., 2015). It is also noted 170 that internal variability can contribute to the determination of signal, which may introduce 171 further uncertainty in the estimate of TOE (Gaetani et al., 2020; Kumar & Ganguly, 2018; 172 Lehner et al., 2020). 173

174	To compare observed SNR with the simulations, Berkeley Earth Surface Temperatures
175	(BEST; Rohde, Muller, Jacobsen, Muller, et al., 2013; Rohde, Muller, Jacobsen, Perlmutter, et
176	al., 2013) is used in this study. Although TN in BEST is biased over Australia (Deng et al.,
177	2021), the TX and TN in BEST show higher correlation compared to Australian gridded climate
178	data (AGCD, previously termed Australian Water Availability Project [AWAP]; Jones et al.,
179	2009), which is better than other global datasets, including National Centers for Environmental
180	Prediction/National Center for Atmospheric Research (NCEP/NCAR) Reanalysis 1 (NCEP1;
181	Kalnay et al., 1996), NCEP/Department of Energy (DOE) Reanalysis 2 (NCEP2; Kanamitsu et
182	al., 2002), Twentieth Century Reanalysis (20CR; Compo et al., 2011), and European Centre for
183	Medium-Range Weather Forecasts (ECMWF) Reanalysis version 5 (ERA5) with preliminary
184	extension to 1950 (Bell et al., 2021; Hersbach et al., 2020) (not shown).
185	2.4 Regional Assessment
186	According to climatological and geographical conditions (Perkins et al., 2014;
187	http://www.bom.gov.au/climate/change/about/temp_timeseries.shtml), Australia is divided into
188	nine sub-regions: AUS (Australia), NA (Northern Australia), SA (Southern Australia), SEA
189	(South East Australia), MEA (Middle Eastern Australia), TA (Tropical Australia), SWA (South
190	West Australia), SSA (Southern South Australia), CAU (Central Australia), and MWA (Mid-
191	Western Australia), shown in Table S4 and Fig. S1, which allows a detailed assessment over
192	smaller subregions. And the base period is from 1961 to 1990, which is commonly used and
193	allows us to analyze TOE with respect to a recent period. Still, we regrid TX and TN to $1^{\circ} \times 1^{\circ}$
194	resolution using bilinear interpolation, and then calculate extreme indices. In addition, grid boxes
195	containing less than 75% land are masked out (King, van Oldenborgh, et al., 2015).

196	In the next section, temporal variations from 1950 to 2100 for the ETCCDI indices in
197	different future scenarios are first analyzed, followed by the spatial patterns of the changes in the
198	indices over 2071-2011 and 2031-2060. Then, the SNR and TOE for TXx and TNn is calculated
199	to address when a novel climate for temperature extremes emerges. For consistency among
200	CMIP6, CMIP5 and BEST, we calculate the noise in SNR for the period 1950-2005, as the
201	estimation of noise can stabilize over longer timescale (Dai & Bloecker, 2019; Santer et al.,
202	2011). Finally, we use two LEs to check the effects of internal variability on the projected
203	responses of extremes and TOE.
204	
205	3 Results
206	3.1 Projected changes
207	Relative to the base period 1961-1990, Figs. 1 and 2 indicate time series of the anomalies
208	for the 14 ETCCDI indices averaged over Australia (10-45°S, 110-155°E) during the period
209	1950-2100 under different future scenarios in CMIP6 (SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-
210	8.5) and CMIP5 (RCP2.6, RCP4.5 and RCP8.5). For the multi-model medians (Fig. 1),
211	consistent with RCPs in CMIP5 (Fig. 2), the Tier 1 experiments in ScenarioMIP show projected
212	increases in the absolute indices (TXx, TXn, TNx and TNn) and in the warm extremes for
213	percentile-based, duration and threshold indices (TX90p, TN90p, WSDI, SU and TR); in
214	contrast, there are decreases in other cold extremes (TX10p, TN10p, CSDI and FD).
215	Among the scenarios, the indices under SSP5-8.5 and RCP8.5 generally show larger
216	warming evolution, especially by the end of the century. Moreover, except for DTR, CSDI and
217	FD (Fig. 1e, k and n), extremes under the SSP3-7.0 fill the gap between SSP2-4.5&RCP4.5 and

- 218 SSP5-8.5&RCP8.5. For example, in the year 2100, the median of TXx under SSP3-7.0 is
- 4.58°C, lower than 5.78°C&5.82°C in SSP5-8.5&RCP8.5 and higher than 3.24°C&2.67°C in
- 220 SSP2-4.5&RCP4.5. In the lower emission scenarios (SSP1-2.6&RCP2.6) there is a stabilization
- for the extremes in the second half of 21<sup>st</sup> century, achieving lowest warming (e.g.,
- 222 2.23°C&1.92°C for TXx in 2100). This result implies the benefits of mitigation strategies
- associated with these scenarios (O'Neill et al., 2016). However, the separation for the adjacent
- 224 pathways (e.g., SSP5-8.5&SSP3-7.0, SSP3-7.0&SSP2-4.5 and SSP2-4.5&SSP1-2.6) usually
- 225 occurs after 2060s for most indices over Australia. In particular, compared to SSP5-
- 8.5&RCP8.5, if a more aggressive mitigation policy is undertaken (e.g., SSP1-2.6&RCP2.6), it
- 227 may still take one or two decades to notice its effects on projected changes in temperature
- 228 extremes over Australia.



229 230 Figure 1. Time series of the anomalies (base period: 1961-1990) for the 14 ETCCDI indices averaged over

231 Australia (10°S-45°S, 110°E-155°E) from 1950 to 2100, under the historical simulations and Tier 1 experiments of

232 ScenarioMIP in CMIP6: Hist (grey), SSP1-2.6 (green), SSP2-4.5 (blue), SSP3-7.0 (yellow) and SSP5-8.5 (red) (the

233 number of models indicated in parentheses in the legend). Solid lines represent the multi-model medians and

234 shading indicates the full range across the models for each experiment.





Figure 2. Same as Fig.1, but for CMIP5: Hist (grey), RCP2.6 (green), RCP4.5 (blue), and RCP8.5 (red).

To illustrate the spreads and medians of the projected climatological changes in extremes over Australian regions in detail, boxplots for SSP5-8.5&RCP8.5 and SSP1-2.6&RCP2.6 are shown in Figs. 3 and 4, and Figs. S2 and S3 for SSP3-7.0 and SSP2-4.5&RCP4.5. Over the regions, the spreads of the indices in SSPs and RCPs tend to be larger with higher emission pathways and over time, among which some regions such as NA and TA commonly span relatively wider ranges. Compared to RCP8.5, the spreads in SSP5-8.5 are usually larger,

- especially over the period 2071-2100. As for the multi-model medians, most indices display
- 245 larger warming trends over TA and lower warming over southern Australian regions (e.g., SSA
- and SWA); while for other indices (e.g., TXx, TNx and TN10p), there are relatively similar
- 247 warming levels across the 10 regions. Relative to RCPs, the warming levels for some indices
- 248 (e.g., TXx, TNn, and WSDI) tends to be higher under the SSPs; in contrast, the relative
- 249 magnitudes of some indices between RCPs and SSPs, such as TXn and TNx (e.g., Fig. 3b, c and
- Fig. S3b, c), differ among the regions and the levels of radiative forcing.







Generally, the spatial patterns for the extremes in both CMIP6 and CMIP5 (Figs. 5 and 6;
Figs. S4-S15 in the supplementary material) are similar to previous studies (Alexander &

Arblaster, 2017). In the highest scenarios for CMIP6 and CMIP5, the extreme indices show a 263 warmer Australia than other pathways, especially in the end of this century. For most indices 264 (expect DTR in Fig. S6 and FD in Fig. S15), most models (at least 75%) in both CMIP 265 ensembles project significant changes in extreme temperature indices over most regions of 266 Australia, both in the middle and the end of the century. However, there are different warming 267 268 patterns for some indices. For example, as shown in Fig. 5, the warming pattern in TXx is relatively consistent among the regions, with the highest warming over central Australia; while 269 for TNn, Northern Australia displays the most marked warming (Fig. 6). 270



Figure 5. Multi-model median changes in TXx for 2071–2100 (a-d; i-k) and 2031–2060 (e-h; l-n) relative to the 273 base period 1961–2010, under different future scenarios in CMIP6 (SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5) 274 and CMIP5 (RCP2.6, RCP4.5 and RCP8.5). Hatching indicates that at least 75% of the models for each future 275 scenario project significant changes at 95% level, based on the two-tailed Student's t-test.

276



277 278 279

Figure 6. Same as Fig. 5, but for TNn.

Compared to RCP scenarios in CMIP5, the higher projected warming for some extremes 280 (e.g., TXx and TNn) and the larger spreads in CMIP6 (especially under SSP5-8.5) by the end of 281 the 21st century is likely related to the different forcings in the SSPs and higher ECS in some 282 CMIP6 models (e.g., Fyfe et al., 2021; Palmer et al., 2021; Tebaldi et al., 2021). Although there 283 are similar levels of stratospheric-adjusted radiative forcing in 2100 in RCPs and SSPs, aerosol 284 emissions, the composition of gases and some radiatively active species (e.g., CO<sub>2</sub> and CH<sub>4</sub>) and 285

286	the resulting ERF in the pathways can be very different (Fyfe et al., 2021; Lurton et al., 2020;
287	Smith et al., 2020; Tebaldi et al., 2021). In addition, the wider inter-model spread of the
288	projected changes under stronger external forcing can result from higher climate sensitivity
289	(Lehner et al., 2020; Tebaldi et al., 2021). As documented in Meehl et al. (2020), 12 of the 39
290	CMIP6 models show higher ECS than the CMIP5 models, some of which can contribute to the
291	wider ranges of projected changes in this study.
292	3.2 Signal-to-Noise Ratio and Time of Emergence
293	The maps of SNR for TXx and TNn in the year 2005 are plotted for BEST, CMIP6 and
294	CMIP5 (Fig. 7), the corresponding signal and noise of which are shown in Figs. S16 and S17,
295	respectively. Although the spatial patterns of noise are relatively similar (Fig. S17), the signals of
296	TXx and TNn show noticeable differences between the observation and the two CMIP
297	ensembles (Fig. S16), which means the resulting SNR in BEST and the two CMIP ensembles
298	differ greatly (Fig. 7). The largest observed SNR for TXx (> 1.2) occurs over central and
299	southwestern regions (Fig. 7a), and for TNn there exhibit negative SNR values ( $< -0.2$ ) over
300	southwest, northern and southeast parts in Australia (Fig. 7d). In contrast, the SNR of TXx and
301	TNn for both CMIP6 and CMIP5 in 2005 tend to be between 0.2 and 0.8. Although there are
302	differences in the observations and the simulations, the low SNR values in 2005 suggest that the
303	signal for the two temperature extremes over most Australia regions has not emerged from the
304	noise.



Figure 7. Signal-to-noise ratio (SNR) in the year 2005 for temperature extremes in BEST, CMIP6 and CMIP5. (a)
SNR in TXx for BEST; (b) SNR in TXx for the multi-model medians in CMIP6; and (c) SNR in TXx for the multi-model medians in CMIP5. (d-f) Same as (a-c), but for TNn.

As spatial aggregation or averaging may reduce the impact of internal variability (Deser, 312 Knutti, et al., 2012; Hawkins & Sutton, 2009; Lehner et al., 2020), Figs. 8 and 9 show the times 313 series (1950-2100) of SNR for TXx and TNn, which are averaged over each region before the 314 calculation of SNR (the corresponding signal and noise are in the supplementary Figs. S18-S20). 315 For the temporal variations of median SNR over the period 1950-2014, the signal and SNR for 316 TXx in BEST can be within the spread of the two CMIP ensembles over some regions (Fig. 8 317 and Fig. S18). However, for TNn the signal and SNR are usually outside the ranges of CMIP6 318 and CMIP5 at the beginning of this century (Fig. 9 and Fig. S19). Despite the influence of 319 observational uncertainty in BEST over Australia (Deng et al., 2021), the above results suggest 320 that the differences between the observed and simulated signal and SNR are mostly related to 321 internal variability (Dai & Bloecker, 2019). In the study by Dai and Bloecker (2019), they 322

concluded that comparing the trends of the observed and modelled precipitation (a variable also 323 exhibiting relatively large variability), which can represent the signal in some studies (e.g., 324 Gaetani et al., 2020), is not appropriate over short timescales and at local and regional scales, as 325

the observed precipitation changes are still dominated by internal variability. 326

327





Figure 8. Time series of signal-to-noise ratio (SNR) in TXx from 1950-2100 over 10 Australian regions for BEST 330 (black), SSP5-8.5 (red) and RCP8.5 (blue) (the number of models indicated in parentheses in the legend). Solid lines

331 represent the multi-model medians and shading indicates the full range across the models for each experiment.



Fig. 10 exhibits the spatial distributions of the multi-model median SNR for TXx and TNn under SSP5-8.5 and RCP8.5 in the year 2050, for which the signal is in Figs. S21. Under both SSP5-8.5 and RCP8.5, despite exhibiting different spatial patterns, the magnitudes of SNR for TXx and TNn are already above 1 over most Australian regions in 2050. For TXx (Fig. 10a, c), there are larger SNR values (>2) over northwest Australia and lower SNR values (>1) over

southwest regions. In contrast, the SNR for TNn (Fig. 10b, d) is more than 2 over western and 341 central Australia and indicates lower values (>1) over tropical and southeast regions. As 342 described in Frame et al. (2017), around mid-century, the regions exhibiting SNR > 1 suggest 343 that there would be "unusual" climate compared to the recent climate over 1950-2005; and for 344 TXx over northwest Australia and TNn over western and central regions, the new climate for the 345 extremes would be "unfamiliar" (SNR > 2). Compared to RCP8.5, SSP5-8.5 in CMIP6 generally 346 displays stronger SNR and the corresponding signal for the two indices, which is valid for other 347 SSPs and RCPs (Figs. S22 and S23). 348

349



Figure 10. Median signal-to-noise ratio (SNR) for TXx and TNn under SSP5-8.5 and RCP8.5 in the year 2050. (a)
SNR for TXx under SSP5-8.5 in the year 2050; (b) SNR for TNn under SSP5-8.5 in the year 2050; (c, d) same as (a)
and (b), but for RCP8.5.

354

As for the temporal evolution of SNR (Figs. 8 and 9), in general, the multi-model medians of SNR in TXx and TNn are slightly larger in SSP5-8.5 than RCP8.5 (e.g., 3.22 under SSP5-8.5 and 2.77 under RCP8.5 in 2050 over AUS); while over some southern regions for TNn (e.g., SA, SSA, SWA), the two CMIP ensemble show higher similarity. In addition, the medians

of signal and noise for the two indices are also comparable in the two scenarios (Figs. S18-S20). 359 It is noted that the differences in signal between CMIP6 and CMIP5 in the end of the century 360 resemble that shown in Figs. 3a and 3d, which may further imply that the regional climate 361 sensitivity in CMIP6 and CMIP5 is comparable indicated in previous studies (Palmer et al., 362 2021; Seneviratne & Hauser, 2020). In terms of the inter-model spread, although the spreads of 363 364 the signals for TXx and TNn in SSP5-8.5 are commonly larger than RCP8.5, in which there are more models showing stronger signal in SSP5-8.5 (Figs. S18 and S19), the ranges in noise (Fig. 365 S20) also contribute to the uncertainty of SNR. Consequently, the relative magnitudes of SNR in 366 SSP5-8.5 and RCP8.5 may change (e.g., Fig. 8a), compared to the signal (e.g., Fig. S18a). For 367 example, the spread of the signal in SSP5-8.5 is slightly larger in the end of the century than 368 RCP8.5; however, influenced by the noise, the resulting range of SNR in SSP5-8.5 becomes 369 narrower. Over the regions, the ranges of SNR for TXx are usually narrower over southern 370 regions (e.g., SSA and SEA); in contrast, for TNn, northern regions such as TA exhibit less 371 uncertainty for SNR and TOE. In other scenarios (Figs. S24-S27), the medians in SNR for TXx 372 and TNn are lower, compared to SSP5-8.5&RCP8.5; and the medians in SSPs are generally still 373 higher than that in RCPs. Also, the spreads of SNR and signal in the lower forcing pathways is 374 375 generally narrower, consistent with the time series of projected changes.

To estimate the TOE for TXx and TNn, we use SNR > 1 and SNR > 2 as the thresholds (Hawkins and Sutton 2012; Frame et al. 2017; Hawkins et al. 2020) and present the spatial patterns for multi-model median TOE under SSP5-8.5 and RCP8.5 (Fig. 11). As TOE occurring at the end of the century may be a temporary change, which is considered as "pseudoemergence", we exclude the TOE occurring after the year 2050 (Abatzoglou et al., 2019; Diffenbaugh & Scherer, 2011; Hawkins et al., 2014; King, Donat, et al., 2015).





383 384

Figure 11. Median time of emergence (TOE) for TXx and TNn based on SNR thresholds under SSP5-8.5 and
RCP8.5. (a) TOE for TXx under SSP5-8.5 when SNR > 1; (b) TOE for TXx under SSP5-8.5 when SNR > 2; (c, d)
same as (a) and (b), but for TNn; (e-h) same as (a-d), but for RCP8.5.

Over some central and tropical parts of Australia, the multi-model median TOE in TXx 388 for SNR > 1 can occur as early as the second decade of this century (2010-2020). Generally, the 389 signal emerges earlier over northwestern region than the southeast for both thresholds (Fig. 11a, 390 b, e and f), in which the signal emerges in 2020s for SNR > 1 and 2040s for SNR > 2, as there 391 indicate relative smaller noise and larger signal (Figs. S17 and S21). Over the southeast regions, 392 the TOE occurs within 2030-2050 for SNR > 1. In contrast, for TNn, the signal emerges from the 393 noise in 2020s over Australia (SNR > 1; Fig. 11c and g); while for SNR > 2, the TOE is within 394 the fifth decade (2040-2050) over western and central regions (Fig. 11d and h). Compared to 395 396 RCP8.5, the multi-model medians of TOE for TXx and TNn in CMIP6 show earlier TOE over more regions based on the threshold SNR > 2, implying the larger median SNR in the middle of 397 this century as shown in Figs. 8 and 9. However, the uncertainty surrounding these TOE 398 estimates remains large (Figs. 8 and 9). For example, for SNR = 2, the range (inter-model 399 spread) of TOE for TXx over AUS can be from 2010s to 2060s (Fig. 8a). For lower scenarios, 400

#### manuscript submitted to Earth's Future

the multi-model medians of TOE commonly occur later and over smaller regions than the 401 stronger pathways. For example, TOE (SNR > 1) for TNn under SSP2-4.5&RCP4.5 usually can 402 403 be 10 years later over some southeast regions (Fig. S28g and k) than that shown in SSP5-8.5&RCP4.5 (Fig. 11c and g), as the signal is lower compared to that in higher pathways. 404 The analysis on SNR and TOE has useful implications for Australia. Under the highest-405 emission scenarios SSP5-8.5&RCP8.5, in which the medians of the signal for TXx are 406 407 comparable, the early emergence over northwest Australia suggests that there is less time for stakeholders and policy makers to implement effective measures, compared to southeast 408 Australia. In contrast, if under lower scenarios, the TOE for TXx can be postponed, especially 409 410 for southeast regions which exhibit larger variability for TXx in the extratropics. However, the adaptation policy may change for different extremes, even under same future pathways. For 411 TNn, the TOE (SNR > 1) can occur over most regions even under lower-emission scenarios; 412 while the "unfamiliar" climate (SNR > 2) can be largely postponed if taking a more sustainable 413 414 pathway (lower emission). It is also noted that the large uncertainty in the estimate of SNR and TOE highlights further challenges for stakeholders and policy makers. 415 416 3.3 Large Ensembles in CMIP6

Previous research has demonstrated the model uncertainty in estimating the effects of internal variability on the TXx and TNn trends, shown in LEs during 1950-2014 over Australian regions (Deng et al., 2021). Therefore, how internal variability influences the projected changes and TOE/SNR (including signal and the noise) needs further investigation. In Fig. 12, which represents the boxplots of projected changes in TXx and TNn for CanESM5-LE and MIROC6-LE over Australian regions under SSP5-8.5 and SSP1-2.6, model uncertainty for representing internal variability still exists, and the relative magnitudes of the spreads for projected changes resemble the results in Figs. 12 and 13 in Deng et al. (2021). The projected changes in TXx for MIROC6-LE span larger ranges than CanESM5-LE by a factor of ~3 or more over the regions, which can be larger than that in Fig. 3a. Moreover, there exhibit larger ranges of the projected changes for TXx over SEA, MEA, and SSA, and relatively narrower spreads over TA for CanESM5-LE and SWA for MIROC6-LE. For TNn, the relative magnitude for the two LEs are comparable over the regions. The different effects of internal variability for different LEs and regions complicate the assessment of the uncertainty on projected changes.





432 433

Figure 12. Boxplots of projected changes in TXx and TNn over 2071–2100 (bold color) and 2031-2060 (light color) relative to the base period 1961–1990 across 10 Australian regions, for CanESM5-LE (cyan) and MIROC6-LE (green). (a) TXx under SSP5-8.5; (b) TNn under SSP5-8.5; (c, d) same as (a, b) but for SSP1-2.6. The boxes indicate the interquartile spreads (ranges between the 25th and 75th percentiles), the black lines within the boxes are the multi-member medians, the whiskers extend to the edges of 1.5 × interquartile ranges and "outliers" outside of the whiskers are denoted by diamonds.

440	The temporal evolution of signal and the boxplots of noise for TXx and TNn over
441	Australian regions under SSP5-8.5 are shown in Figs. 13-15, and the resulting SNR in Figs. S29
442	and S30. The relative magnitudes of the ranges in signal and noise over the regions between the
443	two LEs also resemble that for the spread of the TXx and TNn trends shown in Deng et al.
444	(2021). This suggests that internal variability has impacts not only on the uncertainty of signal,
445	but also on the ranges of noise, making the resulting spread of SNR (Figs. S29 and S30) wider or
446	narrower than that for the corresponding signal (Figs. 13 and 14), which introduce further
447	uncertainty in the ranges of TOE. Although the effects of internal variability on TXx and TNn
448	are similar under SSP1-2.6, the temporal evolution of the SNR and the signal for TXx and TNn
449	stabilizes and there are narrower spreads for SNR compared to SSP5-8.5, which is due to the
450	lower magnitude in signal under the lower scenario (Figs. S31-34).



452 Time (year)
453 Figure 13. Time series of signal (unit: K) in TXx from 1950-2100 over 10 Australian regions under SSP5-8.5 for
454 CanESM5-LE (cyan) and MIROC6-LE (green) (the number of members indicated in parentheses in the legend).
455 Solid lines represent the multi-member medians and shading indicates the full range across the members for each

456 LE.





τÅ SÉA MÉA SWA SSA CAU MWA AUS т'n SWA AÚS NA SA NA SA SĖA MĖA CÁU 461 Figure 15. Boxplots of noise (unit: K) in TXx (a) and TNn (b) calculated over the period 1950-2005 across 10 462 Australian regions, for CanESM5-LE (cyan) and MIROC6-LE (green). The boxes indicate the interquartile spreads 463 (ranges between the 25th and 75th percentiles), the black lines within the boxes are the multi-member medians, the 464 whiskers extend to the edges of  $1.5 \times$  interquartile ranges and "outliers" outside of the whiskers are denoted by 465 466 diamonds. 467

#### 468 **4 Conclusions**

In this study, we analyzed the projected changes for the temperature extremes under 469 470 future scenarios (SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5) from the Tier 1 experiment in ScenarioMIP, which is compared with RCP2.6, RCP4.5 and RCP8.5 in CMIP5. We then use an 471 SNR framework to estimate the time when the signal of climate change for TXx and TNn 472 emerges from the internal variability in the two CMIP ensemble. In addition, two LEs in CMIP6 473 are employed to estimate the effect of internal variability on the projected changes and 474 TOE/SNR. 475 The projected changes for the multi-model medians of the extremes under the highest 476

477 scenario show the strongest warming, and the warming for the indices under SSP3-7.0 fills the 478 gap between SSP2-4.5and SSP5-8.5, with SSP1-2.6 showing the least warming, especially in the 479 end of this century. For some extreme indices (TXx, TXn, TNx, TNn, WSDI and CSDI), 480 although the spatial patterns of warming can be different, there usually projects "warm-get-481 warmer" pattern over Australia. As for the spread in the projections of temperature extremes,

they broadly span narrower envelopes for most indices under lower scenarios in the end of this 482 century. If we take a more sustainable pathway (SSP1-2.6), although it may take two or three 483 decades to take effects, the narrower spreads and weaker projected changes pose relatively less 484 challenge for adaptation decisions compared to other scenarios. Compared to other regions, TA 485 usually shows highest warming. However, as the performance of the models over TA usually 486 487 shows lower scores (Deng et al., 2021), the projected changes for the medians and the spread for the extremes may not be robust (Pierce et al., 2009), which is also applied to other regions such 488 as SSA and SEA. 489

Compared to the counterpart future pathways in CMIP5, the spread in the CMIP6 SSPs 490 491 are commonly wider than RCPs; and for some extremes (e.g., TXx and TNn), the multi-model medians in SPPs are usually higher as well. This is likely caused by different forcings and higher 492 ECS in some CMIP6 models (e.g., Fyfe et al., 2021; Palmer et al., 2021; Tebaldi et al., 2021). 493 For example, Fyfe et al. (2021) concluded that despite the partly countervailing effect by the 494 495 background stratospheric aerosols, the higher amount of  $CO_2$  can lead to stronger warming in SSPs. In this study, we also find that for some indices (e.g., TXx), it is the models with higher 496 ECS that usually show warmer evolution than the multi-model medians in SSP5-8.5 (not shown). 497 498 To further figure out relative importance of each factor, more experiments based on CMIP6 499 models forced by CMIP5 RCP scenarios and/or CMIP5 models forced by CMIP6 SSP scenarios 500 needed be conducted and added to the collection in ScenarioMIP (Fyfe et al., 2021; Tebaldi et al., 2021). 501

We also demonstrate that the medians of SNR for both TXx and TNn in SSPs are commonly higher than in RCPs; and the uncertainty for the SNR of TNn is wider. It is noted that the spreads of SNR for both indices decrease under lower scenarios, which confirms the benefits

505	of lower emission future pathways. Furthermore, the large uncertainty in time of emergence
506	(TOE) result from the inter-model spread of both signal and noise, which is consistent with
507	Hawkins and Sutton (2012). As previous studies concluded that the statistical fit used in the SNR
508	framework can attribute internal variability to the signals (e.g., Hawkins & Sutton, 2012; Kumar
509	& Ganguly, 2018; Lehner et al., 2020), we further illustrate that internal variability can also
510	influence the ranges of noise. To better isolate forced response, dynamical adjustment or LEs can
511	be used (e.g., Lehner et al., 2020; Merrifield et al., 2020). In contrast, using the mean across the
512	range of noise in a LE may be a more appropriate way to represent the expected noise for the
513	model, which needs further investigation.
514	This study suggests that for different extreme temperature indices, the patterns for
515	projected changes and TOE over Australia can be different, which poses large challenge for
516	stakeholders and policymakers. A further effort is to improve the climate models in simulating
517	the physical processes and the internal variability. Unless they are better understood and
518	constrained, the uncertainty of projected changes and TOE will likely continue over future model
519	generations.

## 521 **Conflict of Interest**

522 The authors declare no financial or other conflicts of interests that could have appeared to 523 influence the work reported in this paper.

# 525 Acknowledgments

526	We acknowledge two anonymous reviewers for their constructive comments. We thank
527	Edward Hawkins for feedback and comments. This research/project was undertaken with the
528	assistance of resources and services from the National Computational Infrastructure (NCI),
529	which is supported by the Australian Government. We thank the World Climate Research
530	Programme's Working Group on Coupled Modelling, which is responsible for CMIP and
531	coordinated CMIP5 and CMIP6. We further acknowledge the climate modeling groups for
532	producing and making available their model output, the Earth System Grid Federation (ESGF)
533	for archiving the data and providing access, and the multiple funding agencies who support
534	CMIP and ESGF. S.E.P-K. is supported by ARC grant number FT170100106 and CLEX grant
535	number CE170100023.
536	
537	Data Availability Statement
538	The BEST dataset is obtained from http://berkeleyearth.org/data, and the methodological
539	details are provided in the references: Rohde, Muller, Jacobsen, Muller, et al. (2013) and Rohde,
540	Muller, Jacobsen, Perlmutter, et al. (2013). The CMIP6 and CMIP5 outputs can be downloaded
541	from the Earth System Grid Federation ( <u>https://esgf-node.llnl.gov/search/cmip6/</u> and <u>https://esgf-</u>
542	node.llnl.gov/search/cmip5/). Code for the temperature extremes in the ETCCDI indices is
543	archived at https://doi.org/10.5281/zenodo.4903200 (Deng, 2021).

544

### 545 **References**

Abatzoglou, J. T., Williams, A. P., & Barbero, R. (2019). Global emergence of anthropogenic climate change in fire
 weather indices. *Geophysical Research Letters*, 46(1), 326-336. <u>https://doi.org/10.1029/2018gl080959</u>

- Alexander, L. V., & Arblaster, J. M. (2017). Historical and projected trends in temperature and precipitation
   extremes in Australia in observations and CMIP5. *Weather and Climate Extremes*, 15, 34-56.
   <u>https://doi.org/10.1016/j.wace.2017.02.001</u>
- Barnes, E. A., Hurrell, J. W., Ebert-Uphoff, I., Anderson, C., & Anderson, D. (2019). Viewing forced climate
  patterns through an AI Lens. *Geophysical Research Letters*, 46(22), 13389-13398.
  https://doi.org/10.1029/2019gl084944
- Barsugli, J. J., & Battisti, D. S. (1998). The basic effects of atmosphere-ocean thermal coupling on midlatitude
   variability. *Journal of the Atmospheric Sciences*, 55(4), 477-493. <u>https://doi.org/10.1175/1520-</u>
   0469(1998)055<0477:Tbeoao>2.0.Co;2
- Batibeniz, F., Ashfaq, M., Diffenbaugh, N. S., Key, K., Evans, K. J., Turuncoglu, U. U., et al. (2020). Doubling of
  U.S. population exposure to climate extremes by 2050. *Earth's Future*, 8(4), e2019EF001421.
  https://doi.org/10.1029/2019ef001421
- Beaumont, L. J., Pitman, A., Perkins, S., Zimmermann, N. E., Yoccoz, N. G., & Thuiller, W. (2011). Impacts of
   climate change on the world's most exceptional ecoregions. *Proceedings of the National Academy of Sciences of the United States of America*, 108(6), 2306-2311. https://doi.org/10.1073/pnas.1007217108
- Bell, B., Hersbach, H., Simmons, A., Berrisford, P., Dahlgren, P., Horányi, A., et al. (2021). The ERA5 global
   reanalysis: Preliminary extension to 1950. *Quarterly Journal of the Royal Meteorological Society, 147*(741), 4186-4227. <u>https://doi.org/10.1002/qj.4174</u>
- 566 Cleveland, W. S. (1979). Robust locally weighted regression and smoothing scatterplots. *Journal of the American* 567 *Statistical Association*, 74(368), 829-836. <u>https://doi.org/10.2307/2286407</u>
- Compo, G. P., Whitaker, J. S., Sardeshmukh, P. D., Matsui, N., Allan, R. J., Yin, X., et al. (2011). The twentieth
   century reanalysis project. *Quarterly Journal of the Royal Meteorological Society*, *137*(654), 1-28.
   <u>https://doi.org/10.1002/qj.776</u>
- Dai, A., & Bloecker, C. E. (2019). Impacts of internal variability on temperature and precipitation trends in large
   ensemble simulations by two climate models. *Climate Dynamics*, 52(1-2), 289-306.
   https://doi.org/10.1007/s00382-018-4132-4
- 574 Deng, X. (2021). Code for temperature extremes in ETCCDI indices based on the NCAR Command Language.
   575 Zenodo. <u>https://doi.org/10.5281/zenodo.4903200</u>
- Deng, X., Perkins-Kirkpatrick, S. E., Lewis, S. C., & Ritchie, E. A. (2021). Evaluation of extreme temperatures over
   Australia in the historical simulations of CMIP5 and CMIP6 models. *Earth's Future*, 9(7), e2020EF001902.
   <u>https://doi.org/10.1029/2020ef001902</u>
- Deser, C. (2020). "Certain uncertainty: The role of internal climate variability in projections of regional climate
   change and risk management". *Earth's Future*, 8(12), e2020EF001854.
   https://doi.org/10.1029/2020ef001854
- Deser, C., Knutti, R., Solomon, S., & Phillips, A. S. (2012). Communication of the role of natural variability in
   future North American climate. *Nature Climate Change*, 2(11), 775-779.
   <u>https://doi.org/10.1038/nclimate1562</u>
- Deser, C., Lehner, F., Rodgers, K. B., Ault, T., Delworth, T. L., DiNezio, P. N., et al. (2020). Insights from Earth
   system model initial-condition large ensembles and future prospects. *Nature Climate Change*, 10(4), 277 <u>https://doi.org/10.1038/s41558-020-0731-2</u>
- Deser, C., Phillips, A., Bourdette, V., & Teng, H. Y. (2012). Uncertainty in climate change projections: the role of internal variability. *Climate Dynamics*, *38*(3-4), 527-546. <u>https://doi.org/10.1007/s00382-010-0977-x</u>
- Deutsch, C. A., Tewksbury, J. J., Huey, R. B., Sheldon, K. S., Ghalambor, C. K., Haak, D. C., et al. (2008). Impacts
   of climate warming on terrestrial ectotherms across latitude. *Proceedings of the National Academy of Sciences of the United States of America*, 105(18), 6668-6672. https://doi.org/10.1073/pnas.0709472105
- Diffenbaugh, N. S., & Scherer, M. (2011). Observational and model evidence of global emergence of permanent,
   unprecedented heat in the 20th and 21st centuries. *Climatic Change*, 107(3-4), 615-624.
   https://doi.org/10.1007/s10584-011-0112-y
- Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., et al. (2016). Overview of the Coupled
   Model Intercomparison Project Phase 6 (CMIP6) experimental design and organization. *Geoscientific Model Development*, 9(5), 1937-1958. <u>https://doi.org/10.5194/gmd-9-1937-2016</u>
- Fischer, E. M., Sedláček, J., Hawkins, E., & Knutti, R. (2014). Models agree on forced response pattern of
   precipitation and temperature extremes. *Geophysical Research Letters*, 41(23), 8554-8562.
   <u>https://doi.org/10.1002/2014gl062018</u>
- Frame, D., Joshi, M., Hawkins, E., Harrington, L. J., & de Roiste, M. (2017). Population-based emergence of
   unfamiliar climates. *Nature Climate Change*, 7(6), 407-411. <u>https://doi.org/10.1038/nclimate3297</u>
- Fyfe, J. C., Kharin, V. V., Santer, B. D., Cole, J. N. S., & Gillett, N. P. (2021). Significant impact of forcing
   uncertainty in a large ensemble of climate model simulations. *Proceedings of the National Academy of Sciences of the United States of America*, 118(23), e2016549118. https://doi.org/10.1073/pnas.2016549118
- Gaetani, M., Janicot, S., Vrac, M., Famien, A. M., & Sultan, B. (2020). Robust assessment of the time of emergence
   of precipitation change in West Africa. *Scientific Reports*, 10(1), 7670. <u>https://doi.org/10.1038/s41598-020-63782-2</u>
- Giorgi, F., & Bi, X. (2009). Time of emergence (TOE) of GHG-forced precipitation change hot-spots. *Geophysical Research Letters*, 36(6), L06709. <u>https://doi.org/10.1029/2009gl037593</u>
- Grose, M. R., Narsey, S., Delage, F. P., Dowdy, A. J., Bador, M., Boschat, G., et al. (2020). Insights From CMIP6
  for Australia's Future Climate. *Earth's Future*, 8(5), e2019EF001469.
  https://doi.org/10.1029/2019ef001469
- Hawkins, E., Anderson, B., Diffenbaugh, N., Mahlstein, I., Betts, R., Hegerl, G., et al. (2014). Uncertainties in the
   timing of unprecedented climates. *Nature*, 511(7507), E3-E5. <u>https://doi.org/10.1038/nature13523</u>
- Hawkins, E., Frame, D., Harrington, L., Joshi, M., King, A., Rojas, M., et al. (2020). Observed emergence of the
  climate change signal: From the familiar to the unknown. *Geophysical Research Letters*, 47(6),
  e2019GL086259. <u>https://doi.org/10.1029/2019gl086259</u>
- Hawkins, E., & Sutton, R. (2009). The potential to narrow uncertainty in regional climate predictions. *Bulletin of the American Meteorological Society*, *90*(8), 1095-1108. <u>https://doi.org/10.1175/2009bams2607.1</u>
- Hawkins, E., & Sutton, R. (2012). Time of emergence of climate signals. *Geophysical Research Letters*, 39(1),
   L01702. <u>https://doi.org/10.1029/2011gl050087</u>
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., et al. (2020). The ERA5 global
   reanalysis. *Quarterly Journal of the Royal Meteorological Society*, *146*(730), 1999-2049.
   https://doi.org/10.1002/qj.3803
- Intergovernmental Panel on Climate Change. (2021). Climate Change 2021: The physical science basis.
  Contribution of Working Group I to the sixth assessment report of the Intergovernmental Panel on Climate
  Change (V. Masson-Delmotte, P. Zhai, A. Pirani, S. L. Connors, C. Péan, S. Berger, N. Caud, Y. Chen, L.
  Goldfarb, M. I. Gomis, M. Huang, K. Leitzell, E. Lonnoy, J. B. R. Matthews, T. K. Maycock, T.
  Waterfield, O. Yelekçi, R. Yu, & B. Zhou Eds.). Cambridge, UK and New York, NY, USA: Cambridge
  University Press.
- Jones, D. A., Wang, W., & Fawcett, R. (2009). High-quality spatial climate data-sets for Australia. *Australian Meteorological and Oceanographic Journal*, 58(4), 233-248. <u>https://doi.org/10.22499/2.5804.003</u>
- Kalnay, E., Kanamitsu, M., Kistler, R., Collins, W., Deaven, D., Gandin, L., et al. (1996). The NCEP/NCAR 40 year reanalysis project. *Bulletin of the American Meteorological Society*, *77*(3), 437-472.
   https://doi.org/10.1175/1520-0477(1996)077<0437:Tnyrp>2.0.Co;2
- Kanamitsu, M., Ebisuzaki, W., Woollen, J., Yang, S. K., Hnilo, J. J., Fiorino, M., et al. (2002). NCEP–DOE AMIPII Reanalysis (R-2). *Bulletin of the American Meteorological Society*, *83*(11), 1631-1643.
  https://doi.org/10.1175/bams-83-11-1631(2002)083<1631:Nar>2.3.Co;2
- King, A. D., Donat, M. G., Fischer, E. M., Hawkins, E., Alexander, L. V., Karoly, D. J., et al. (2015). The timing of anthropogenic emergence in simulated climate extremes. *Environmental Research Letters*, 10(9), 094015.
   <u>https://doi.org/10.1088/1748-9326/10/9/094015</u>
- King, A. D., van Oldenborgh, G. J., Karoly, D. J., Lewis, S. C., & Cullen, H. (2015). Attribution of the record high
   Central England temperature of 2014 to anthropogenic influences. *Environmental Research Letters, 10*(5).
   <u>https://doi.org/10.1088/1748-9326/10/5/054002</u>
- Kumar, D., & Ganguly, A. R. (2018). Intercomparison of model response and internal variability across climate model ensembles. *Climate Dynamics*, *51*(1-2), 207-219. <u>https://doi.org/10.1007/s00382-017-3914-4</u>
- Lehner, F., Deser, C., Maher, N., Marotzke, J., Fischer, E. M., Brunner, L., et al. (2020). Partitioning climate
   projection uncertainty with multiple large ensembles and CMIP5/6. *Earth System Dynamics*, 11(2), 491 508. <u>https://doi.org/10.5194/esd-11-491-2020</u>
- Lurton, T., Balkanski, Y., Bastrikov, V., Bekki, S., Bopp, L., Braconnot, P., et al. (2020). Implementation of the
   CMIP6 forcing data in the IPSL-CM6A-LR model. *Journal of Advances in Modeling Earth Systems, 12*(4),
   e2019MS001940. <u>https://doi.org/10.1029/2019ms001940</u>
- Mahlstein, I., Hegerl, G., & Solomon, S. (2012). Emerging local warming signals in observational data. *Geophysical Research Letters*, 39(21), L21711. <u>https://doi.org/10.1029/2012gl053952</u>
- Mahlstein, I., Knutti, R., Solomon, S., & Portmann, R. W. (2011). Early onset of significant local warming in low latitude countries. *Environmental Research Letters*, 6(3), 034009. <u>https://doi.org/10.1088/1748-9326/6/3/034009</u>

660	Mankin, J. S., Lehner, F., Coats, S., & McKinnon, K. A. (2020). The value of initial condition large ensembles to
661	robust adaptation decision-making. Earth's Future, 8(10), e2012EF001610.
662	https://doi.org/10.1029/2020ef001610
663	Meehl, G. A., Senior, C. A., Eyring, V., Flato, G., Lamarque, J. F., Stouffer, R. J., et al. (2020). Context for
664	interpreting equilibrium climate sensitivity and transient climate response from the CMIP6 Earth system
665	models. Science Advances, 6(26), eaba1981. https://doi.org/10.1126/sciadv.aba1981
666	Merrifield, A. L., Brunner, L., Lorenz, R., Medhaug, I., & Knutti, R. (2020). An investigation of weighting schemes
667	suitable for incorporating large ensembles into multi-model ensembles. Earth System Dynamics, 11(3),
668	807-834. https://doi.org/10.5194/esd-11-807-2020
669	O'Neill, B. C., Tebaldi, C., van Vuuren, D. P., Eyring, V., Friedlingstein, P., Hurtt, G., et al. (2016). The Scenario
670	Model Intercomparison Project (ScenarioMIP) for CMIP6. Geoscientific Model Development, 9(9), 3461-
671	3482. https://doi.org/10.5194/gmd-9-3461-2016
672	Ossó, A., Allan, R. P., Hawkins, E., Shaffrey, L., & Maraun, D. (2021). Emerging new climate extremes over
673	Europe, Climate Dynamics, https://doi.org/10.1007/s00382-021-05917-3
674	Palmer, T. E., Booth, B. B. B., & McSweeney, C. F. (2021). How does the CMIP6 ensemble change the picture for
675	European climate projections? <i>Environmental Research Letters</i> , 16(9), 094042.
676	https://doi.org/10.1088/1748-932.6/ac1ed9
677	Perkins, S. E., Moise, A., Whetton, P., & Katzfey, J. (2014), Regional changes of climate extremes over Australia –
678	A comparison of regional dynamical downscaling and global climate model simulations. <i>International</i>
679	iournal of climatology 34(12) 3456-3478 https://doi.org/10.1002/joc.3927
680	Perkins-Kirknatrick S E Fischer E M Angélil O & Gibson P B (2017) The influence of internal climate
681	variability on heatwave frequency trends <i>Environmental Research Letters</i> 12(4)
682	https://doi.org/10.1088/1748-9326/aa63fe
683	Pierce D W Barnett T P Santer B D & Gleckler P I (2009) Selecting global climate models for regional
684	climate change studies. Proceedings of the National Academy of Sciences of the United States of America
685	106(21) 8441-8446 https://doi.org/10.1073/pnas.0900094106
686	Riahi K van Vuuren D P Kriegler F Edmonds I O'Neill B C Eujimori S et al (2017) The Shared
687	Socioeconomic Pathways and their energy land use and greenhouse gas emissions implications: An
688	overview Global Environmental Change 42 153-168 https://doi.org/10.1016/j.gloenycha.2016.05.009
689	Rohde R Muller R Jacobsen R Muller F Perlmutter S Rosenfeld A et al (2013) A new estimate of the
690	average Earth surface land temperature spanning 1753 to 2011 <i>Geoinformatics and Geostatistics: An</i>
691	Overview 1(1) https://doi.org/10.4172/2327-4581.1000101
692	Robde R. Muller R. Jacobsen R. Perlmutter S. Rosenfeld A. Wurtele I. et al. (2013). Berkeley Earth
693	temperature averaging process. Geoinformatics and Geostatistics: An Overview 1(2)
60/	https://doi.org/10.4172/2327_4581_1000103
695	Santer B. D. Mears C. Doutriaux C. Caldwell P. Gleckler P. I. Wigley, T. M. L. et al. (2011). Separating
606	signal and noise in atmospheric temperature changes: The importance of timescale. <i>Journal of Geophysical</i>
607	Basagrah: Atmospheres, 116(D22), D22105, https://doi.org/10.1020/2011id016263
608	Sensuirotne, S. J. & Houser, M. (2020). Degional climate sensitivity of climate extremes in CMID6 versus CMID5
600	multi model ensembles. Easth's Eutura 8(0), e2010EE001474, https://doi.org/10.1020/2010EE001474
700	Sillmann I. Kharin V. V. Zhang Y. Zwiers F. W. & Bronaugh D. (2012). Climate extremes indices in the
700	CMID5 multimodel encemble: Dart 1. Model evaluation in the present climate. <i>Journal of Geophysical</i>
701	Paragraphi Atmospheres, 118(A), 1716, 1722, https://doi.org/10.1002/jard.50202
702	Sillmann I. Kharin V. V. Zwiere F. W. Zhang V. & Pronough D. (2012). Climate extremes indices in the
703	CMID5 multimodal anomhla: Dart 2. Eutura alimata projections. <i>Journal of Coophysical Pasagraph</i> :
704	Atmospheres, 119(6), 2472-2402, https://doi.org/10.1002/j.ord.50189
705	Almospheres, 110(0), 24/5-2495. <u>https://doi.org/10.1002/jgrd.50186</u>
700	faming and adjustments in CMID6 models. Atwasphavis Chamistry, and Dhusiag. 20(16), 0501, 0619
707	https://doi.org/10.5104/com 20.0501.2020
700	$\frac{\text{Intps://doi.org/10.5194/acp-20-9591-2020}}{System D. System E. & Hendrice E. (2015). What does also be required to the set of the set o$
709	Dhilosophical Transactions of the Poyal Society A: Mathematical Division and Engineering Sciences
710	1 nuosopnicai 1 ransactions of the Royal Society A. Mathematical, Physical and Engineering Sciences, 273(2054), 20140426, https://doi.org/10.1008/rets.2014.0426
712	J/J(20J+J), 20140420. IIIIPS.//001.01g/10.1070/1813.2014.0420 Ton V. Gon T. V. & Horton D. F. (2018). Deviated timing of personinghis changes in climate sufference for
712 713	tarrestrial and marine ecosystems. <i>Clobal Change Pialory</i> 24(10), 4606 4709
717	$\frac{1}{1000}$ $\frac{1}{10000}$ $\frac{1}{10000}$ $\frac{1}{100000}$ $\frac{1}{10000000000000000000000000000000000$
/14	<u>nups//doi.01g/10.1111/gc0.14527</u>

- Tebaldi, C., Debeire, K., Eyring, V., Fischer, E., Fyfe, J., Friedlingstein, P., et al. (2021). Climate model projections
   from the Scenario Model Intercomparison Project (ScenarioMIP) of CMIP6. *Earth System Dynamics*,
   *12*(1), 253-293. <u>https://doi.org/10.5194/esd-12-253-2021</u>
- Thibeault, J. M., & Seth, A. (2014). Changing climate extremes in the Northeast United States: observations and projections from CMIP5. *Climatic Change*, *127*(2), 273-287. <u>https://doi.org/10.1007/s10584-014-1257-2</u>
- Tokarska, K. B., Stolpe, M. B., Sippel, S., Fischer, E. M., Smith, C. J., Lehner, F., et al. (2020). Past warming trend
   constrains future warming in CMIP6 models. *Science Advances*, 6(12), eaaz9549.
   <u>https://doi.org/10.1126/sciadv.aaz9549</u>
- Xie, S. P., Deser, C., Vecchi, G. A., Collins, M., Delworth, T. L., Hall, A., et al. (2015). Towards predictive understanding of regional climate change. *Nature Climate Change*, 5(10), 921-930.
   <u>https://doi.org/10.1038/nclimate2689</u>
- Zhang, X., Alexander, L., Hegerl, G. C., Jones, P., Tank, A. K., Peterson, T. C., et al. (2011). Indices for monitoring
   changes in extremes based on daily temperature and precipitation data. *Wiley Interdisciplinary Reviews: Climate Change*, 2(6), 851-870. https://doi.org/10.1002/wcc.147
- Zhang, X., Hegerl, G., Zwiers, F. W., & Kenyon, J. (2005). Avoiding inhomogeneity in percentile-based indices of
   temperature extremes. *Journal of Climate, 18*(11), 1641-1651. <u>https://doi.org/10.1175/jcli3366.1</u>

	<b>AGU</b> PUBLICATIONS
1	
2	Earth's Future
3	Supporting Information for
4 5	Projected Changes and Time of Emergence of Temperature Extremes over Australia in CMIP5 and CMIP6
6	Xu Deng <sup>1,2</sup> and Sarah E. Perkins-Kirkpatrick <sup>1,2</sup>
7	<sup>1</sup> School of Science, University of New South Wales, Canberra, ACT, Australia
8	<sup>2</sup> ARC Centre of Excellence for Climate Extremes, University of New South Wales, Canberra, ACT, Australia
9	
10	
11 12 13 14 15 16 17 18 19 20 22 22 22 22 22 22 22 22 22 22 22 22	Contents of this file Tables S1 to S4 Figures S1 to S34

				,	
Model	Historical	SSP1-2.6	SSP2-4.5	SSP3-7.0	SSP5-8.5
1. ACCESS-CM2	rlilplfl	rlilplfl	rlilplfl	rlilplfl	rlilplfl
2. ACCESS-ESM1-5	rlilplfl	-	rlilplfl	rlilplfl	rlilplfl
3. AWI-CM-1-1-MR	rlilplfl	rlilplfl	rlilplfl	rlilp1f1	rlilplfl
4. AWI-ESM-1-1-LR	rlilplfl	-	-	-	-
5. BCC-CSM2-MR	rlilplfl	rlilplfl	rlilplfl	rli1p1f1	rlilplfl
6. BCC-ESM1	rlilplfl	-	-	-	-
7. CanESM5	rlilplfl	rlilplfl	rlilplfl	rli1p1f1	rlilplfl
	~	~			~
	r25i1p1f1	r25i1p1f1			r25i1p1f1
8. CMCC-ESM2	rlilplfl	rlilplfl	rlilplfl	rlilp1f1	rlilplfl
9. CNRM-CM6-1	rlilplf2	rlilplf2	rlilplf2	rlilp1f2	rlilplf2
10. CNRM-CM6-1-	rlilplf2	rlilplf2	-	-	rlilplf2
HR					
11. CNRM-ESM2-1	rli1p1f2	rlilp1f2	rlilplf2	r1i1p1f2	rlilplf2
12. FGOALS-f3-L	rlilplfl	-	-	-	-
13. FGOALS-g3	rlilplfl	rlilplfl	rlilplfl	rlilplfl	rlilplfl
14. GFDL-CM4	rlilplfl	-	rlilplfl	-	rlilplfl
15. GFDL-ESM4	rlilplfl	rlilplfl	rlilplfl	rlilp1f1	rlilplfl
16. GISS-E2-1-G	rlilplfl	-	-	-	-
17. HadGEM3-	rli1p1f3	rlilp1f3	rlilp1f3	-	rlilp1f3
GC31-LL					
18. HadGEM3-	rli1p1f3	rlilp1f3	-	-	rlilp1f3
GC31-MM					
19. INM-CM4-8	rlilplfl	rlilplfl	rlilplfl	rlilplfl	rlilplfl
20. INM-CM5-0	rlilplfl	rlilplfl	rlilplfl	r1i1p1f1	rlilplfl
21. IPSL-CM6A-LR	rlilplfl	rlilplfl	rlilplfl	r1i1p1f1	rlilplfl
22. MIROC-ES2L	rli1p1f2	rlilp1f2	rlilp1f2	r1i1p1f2	rlilp1f2
23. MIROC6	rlilplfl	rlilplfl	rlilplfl	r1i1p1f1	rlilplfl
	~	~			~
	r50i1p1f1	r50i1p1f1			r50i1p1f1
24. MPI-ESM-1-2-	rlilplfl	-	-	-	-
HAM					
25. MPI-ESM1-2-HR	rlilplfl	rli1p1f1	rlilplfl	rli1p1f1	rlilplfl
26. MPI-ESM1-2-LR	rlilplfl	rlilplfl	rlilplfl	rlilp1f1	rlilplfl
27. MRI-ESM2-0	rlilplfl	rli1p1f1	rlilplfl	rli1p1f1	rlilplfl
28. NorCPM1	rlilplfl	-	-	-	-
29. NorESM2-LM	rlilplfl	rlilplfl	rlilplfl	r1i1p1f1	rlilplfl
30. NorESM2-MM	rlilplfl	rlilplfl	rlilplfl	r1i1p1f1	rlilplfl
31. SAM0-UNICON	rlilplfl	-	-	-	-
32. UKESM1-0-LL	rli1p1f2	rlilp1f2	rlilp1f2	r1i1p1f2	rlilp1f2
Number of models	32	23	23	21	25

**Table S1.** Models and simulations in CMIP6 used in this study.

Model	Historical	RCP2.6	RCP4.5	RCP8.5
1. ACCESS1-0	rlilpl	-	r1i1p1	rlilpl
2. ACCESS1-3	rlilpl	-	rli1p1	rlilpl
3. bcc-csm1-1	rlilpl	rlilpl	rli1p1	rlilpl
4. BNU-ESM	rlilpl	rlilpl	rlilpl	rlilpl
5. CanESM2	rlilpl	rlilpl	rlilpl	rlilpl
6. CCSM4	rlilpl	rlilpl	rlilpl	rlilpl
7. CESM1-BGC	rlilpl	-	rlilpl	rlilpl
8. CMCC-CM	rlilpl	-	rlilpl	rlilpl
9. CNRM-CM5	rlilpl	rlilpl	rlilpl	rlilpl
10. CSIRO-Mk3-6-0	rlilpl	rlilpl	rlilpl	rlilpl
11. FGOALS-g2	rlilpl	-	rlilpl	rlilpl
12. GFDL-ESM2G	rlilpl	rlilpl	rlilpl	rlilpl
13. GFDL-ESM2M	rlilpl	rlilpl	rlilpl	rlilpl
14. GISS-E2-R	r6i1p1	-	r6i1p1	r2i1p1
15. HadGEM2-CC	rlilpl	-	rli1p1	rlilpl
16. HadGEM2-ES	rlilpl	rlilpl	rli1p1	rlilpl
17. IPSL-CM5A-LR	rlilpl	rlilpl	rli1p1	rlilpl
18. IPSL-CM5A-MR	rlilpl	rlilpl	rlilpl	rlilpl
19. IPSL-CM5B-LR	rlilpl	-	rlilpl	rlilpl
20. MIROC5	rlilpl	rlilpl	rlilpl	rlilpl
21. MIROC-ESM	rlilpl	rlilpl	rlilpl	rlilpl
22. MIROC-ESM-	rlilpl	rlilpl	rli1p1	rlilpl
CHEM				
23. MPI-ESM-LR	rlilpl	rlilpl	rlilpl	rlilpl
24. MPI-ESM-MR	rlilpl	rlilpl	rli1p1	rlilpl
25. MRI-CGCM3	rlilpl	rlilpl	rlilpl	rlilpl
26. NorESM1-M	rlilpl	rlilpl	rlilpl	rlilpl
Number of models	26	18	26	26

**Table S2.** Models and simulations in CMIP5 used in this study.
 44

 $\begin{array}{r} 45\\ 46\\ 47\\ 48\\ 49\\ 50\\ 51\\ 52\\ 53\\ 54\\ 55\\ 56\\ 57\\ 58\\ 960\\ 61\\ 62\\ \end{array}$ 

Labol	Index Neme	Description	Unit	
Laber	Index Ivanie	Description	Ullit	
TXx	Hottest day	Annual maximum value of daily maximum temperature	°С	
TXn	Coldest day	Annual minimum value of daily maximum temperature	°C	
TNx	Warmest night	Annual maximum value of daily minimum temperature	°C	
TNn	Coldest night	Annual minimum value of daily minimum temperature	°C	
DTD	Diurnal	Annual mean difference between daily maximum and	00	
DIK	temperature range	minimum temperature	÷C	
TYOO		Percentage of time when daily maximum temperature is	%	
1X90p	Warm days	greater than 90 <sup>th</sup> percentile (using running 5-day window)		
	K10p Cold days	Percentage of time when daily maximum temperature is less	%	
TX10p		than 10 <sup>th</sup> percentile (using running 5-day window)		
	p Warm nights	Percentage of time when daily minimum temperature is	%	
TN90p		greater than 90 <sup>th</sup> percentile (using running 5-day window)		
		Percentage of time when daily minimum temperature is less		
TN10p	Cold nights	than 10 <sup>th</sup> percentile (using running 5-day window)	%	
		Annual count when at least six consecutive days of		
WSDI	Warm spell	maximum temperature is greater than $00^{\text{th}}$ percentile (using	dave	
WSDI	duration index	maximum temperature is greater than 50° percentile (using	uays	
		running 5-day window)		
~~~	Cold spell	Annual count when at least six consecutive days of minimum		
CSDI	duration index	temperature is less than 10 <sup>th</sup> percentile (using running 5-day	days	
		window)		
SU	Summer days	Annual count when daily maximum temperature is greater	days	
50	Summer days	than 25°C	uujb	
тр	Tropical nights	Annual count when daily minimum temperature is greater	dave	
IK	riopical inglits	than 20°C	uays	
ED	Erect days	Annual count when daily minimum temperature is less than	dava	
FD	Frost days	0°C	days	

63 
 Table S3. Extreme temperature indices used in this study.

 $\begin{array}{c} 64\\ 65\\ 66\\ 67\\ 68\\ 69\\ 70\\ 71\\ 72\\ 73\\ 74\\ 75\\ 76\\ 77\\ 78\\ 80\\ 81\\ 82\\ 83\\ 84\\ 85\\ \end{array}$ 

Label	Region	Lat (°S)	Lon (°E)
1. AUS	Australia	10-45	110-155
2. NA	Northern Australia	10–26	110–155
3. SA	Southern Australia	26–45	110–155
4. SEA	South East Australia	32.5–45	140–155
5. MEA	Middle Eastern Australia	20-32.5	140–155
6. TA	Tropical Australia	10–20	110–155
7. SWA	South West Australia	27.5-40	110-127.5
8. SSA	Southern South Australia	30–40	127.5–140
9. CAU	Central Australia	20-30	127.5-140
10. MWA	Mid-Western Australia	20-27.5	110-127.5

**Table S4.** Latitude and longitude boundaries of Australian regions.



103 104 Fig. S1. Regions used in the study. Northern Australia (NA) and Southern Australia (SA) are 105 divided by the dashed line at 26°S, and solid lines denote the boundaries of other Australian 106 107 subregions.



**Fig. S2.** Boxplots of projected changes in the 14 ETCCDI indices over 2071–2100 (bold color) and 2031-2060 (light color) relative to the base period 1961–1990 across 10 Australian regions, under SSP3-7.0 (red). The boxes indicate the interquartile spreads (ranges between the 25th and 75th percentiles), the black lines within the boxes are the multi-model medians, the whiskers extend to the edges of 1.5 × interquartile ranges and "outliers" outside of the whiskers are denoted by diamonds.





**Fig. S4.** Multi-model median changes in TXn for 2071–2100 (a-d; i-k) and 2031–2060 (e-h; l-n)

120 relative to the base period 1961–2010, under different future scenarios in CMIP6 (SSP1-2.6, SSP2-

121 4.5, SSP3-7.0 and SSP5-8.5) and CMIP5 (RCP2.6, RCP4.5 and RCP8.5). Hatching indicates that at

least 75% of the models for each future scenario project significant changes at 95% level, basedon the two-tailed Student's t-test.













 $\begin{array}{c} 140\\ 141 \end{array}$ 

15













Fig. S16. Signal (unit: K) in the year 2005 for temperature extremes in BEST, CMIP6 and CMIP5. (a)
Signal in TXx for BEST; (b) Signal in TXx for the multi-model medians in CMIP6; and (c) Signal in
TXx for the multi-model medians in CMIP5. (d-f) Same as (a-c), but for TNn.

- 1/9







213 Fig. S18. Time series of signal (unit: K) in TXx from 1950-2100 over 10 Australian regions under SSP5-8.5 for BEST (black), CMIP6 (red) and CMIP5 (blue) (the number of models indicated in parentheses in the legend). Solid lines represent the multi-model medians and shading indicates the full range across the models for each experiment.





**Fig. S20.** Boxplots of noise (unit: K) in TXx (a) and TNn (b) calculated over the period 1950-2005 across 10 Australian regions, under SSP5-8.5 (red) and RCP8.5 (blue). The boxes indicate the interquartile spreads (ranges between the 25th and 75th percentiles), the black lines within the boxes are the multi-model medians, the whiskers extend to the edges of 1.5 × interquartile ranges and "outliers" outside of the whiskers are denoted by diamonds.



Fig. S21. Median signal (unit: K) for TXx and TNn under SSP5-8.5 and RCP8.5 in the years 2050. (a)
Signal for TXx under SSP5-8.5 in the year 2050; (b) Signal for TNn under SSP5-8.5 in the year
260 2050; (c, d) same as (a) and (b), but for RCP8.5.



289

290

291

**Fig. S22.** Median signal-to-noise ratio (SNR) for TXx and TNn under SSP3-7.0, SSP2-4.5&RCP4.5, and SSP1-2.6&RCP2.6 in the year 2050. (a) SNR for TXx under SSP3-7.0 in the year 2050; (b) SNR for TNn under SSP3-7.0 in the year 2050; (c, d) same as (a) and (b), but for SSP2-4.5; (e, f) same as (a) and (b), but for RCP4.5; (g, h) same as (a) and (b), but for SSP1-2.6; (i, j) same as (a) and (b), but for RCP2.6.



$\mathbf{r}$	n	0
7	9	0



302



**Fig. S24.** Time series of signal-to-noise ratio (SNR) in TXx from 1950-2100 over 10 Australian regions under SSP1-2.6 (green), SSP2-4.5 (blue), SSP3-7.0 (yellow) (the number of models indicated in parentheses in the legend). Solid lines represent the multi-model medians and shading indicates the full range across the models for each experiment.









Fig. S28. Median time of emergence (TOE) for TXx and TNn based on SNR thresholds under
SSP3-7.0, SSP2-4.5&RCP4.5, and SSP1-2.6&RCP2.6. (a) TOE for TXx under SSP3-7.0 when SNR > 1;
(b) TOE for TXx under SSP3-7.0 when SNR > 2; (c, d) same as (a) and (b), but for TNn; (e-h) same
as (a-d), but for SSP2-4.5; (i-l) same as (a-d), but for RCP4.5; (m-p) same as (a-d), but for SSP1-2.6;
(q-t) same as (a-d), but for SSP1-2.6.


Fig. S29. Time series of signal-to-noise ratio (SNR) in TXx from 1950-2100 over 10 Australian
regions under SSP5-8.5 for CanESM5-LE (cyan) and MIROC6-LE (green) (the number of members
indicated in parentheses in the legend). Solid lines represent the multi-member medians and
shading indicates the full range across the members for each LE.









- 359
- 361





 
 366
 1950
 2000 Time (year)
 2100

 367
 Fig. S34. Same as Fig. S29, but for signal in TNn under SSP1-2.6.