Drag Coefficient Constraints for Space Weather Observations in the Upper Thermosphere

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Abstract

The space weather research community relies heavily on thermospheric density data to understand long-term thermospheric variability, construct assimilative, empirical, and semi-empirical global atmospheric models, and validate model performance. One of the challenges in resolving accurate thermospheric density datasets from satellite orbital drag measurements is modeling appropriate physical aerodynamic drag force coefficients. The drag coefficient may change throughout the thermospheric environment due to model dependencies on composition and altitude. As such, existing drag coefficient model errors may be altitude and solar cycle dependent, with greater errors at higher altitudes around 500 km near the oxygen-to-helium transition region. This can lead to errors in orbit-derived density datasets and models. In this paper, inter-satellite density comparisons at 500 km are evaluated to constrain drag coefficient modeling assumptions. Density consistency results indicate that drag coefficient models with incomplete energy and momentum accommodation produce the most consistent densities, while the standard diffuse modeling approach may not be appropriate at these altitudes. Models with momentum accommodation between 0.5 - 0.9 and energy accommodation between 0.83 - 0.96 may be the most appropriate at upper thermospheric altitudes. Modeling drag coefficients with diffuse gas-surface interactions could lead to errors in derived density of $^{25\%}$ and in-track satellite orbit prediction uncertainty during solar maximum conditions on the order of hundreds of meters.

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Key Points:

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10	•	Drag coefficient errors lead to altitude-dependent biases in thermospheric densi-
11		ties
12	•	Drag coefficient model assumptions are constrained by evaluating multi-satellite
13		density consistency
14	•	Aerodynamic drag modeling is improved near 500 km by assuming incomplete mo-
15		mentum accommodation

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16 Abstract

The space weather research community relies heavily on thermospheric density data to 17 understand long-term thermospheric variability, construct assimilative, empirical, and 18 semi-empirical global atmospheric models, and validate model performance. One of the 19 challenges in resolving accurate thermospheric density datasets from satellite orbital drag 20 measurements is modeling appropriate physical aerodynamic drag force coefficients. The 21 drag coefficient may change throughout the thermospheric environment due to model 22 dependencies on composition and altitude. As such, existing drag coefficient model er-23 rors may be altitude and solar cycle dependent, with greater errors at higher altitudes 24 around 500 km near the oxygen-to-helium transition region. This can lead to errors in 25 orbit-derived density datasets and models. In this paper, inter-satellite density compar-26 isons at ~ 500 km are evaluated to constrain drag coefficient modeling assumptions. Den-27 sity consistency results indicate that drag coefficient models with incomplete energy and 28 momentum accommodation produce the most consistent densities, while the standard 29 diffuse modeling approach may not be appropriate at these altitudes. Models with mo-30 mentum accommodation between 0.5 - 0.9 and energy accommodation between 0.83 -31 32 0.96 may be the most appropriate at upper thermospheric altitudes. Modeling drag coefficients with diffuse gas-surface interactions could lead to errors in derived density of 33 $\sim 25\%$ and in-track satellite orbit prediction uncertainty during solar maximum condi-34 tions on the order of hundreds of meters. 35

³⁶ Plain Language Summary

The Earth's upper atmosphere exerts forces on satellites that can change their paths. 37 It is critical to understand how these atmospheric drag forces work in order to measure 38 atmospheric variability and predict satellite orbital paths in an increasingly crowded near-39 Earth space. The atmospheric drag force depends on interactions between atmospheric 40 particles and the surface of a satellite. Gas particles can impact a satellite surface and 41 scatter in a variety of ways. Depending on the speed and direction of the scattered par-42 ticles, the atmospheric drag force on a satellite can change. The appropriate type of scat-43 tering for atmospheric particles interacting with a satellite surface in orbit remains un-44 certain. In this work, the authors try to infer the nature of gas-surface scattering by com-45 paring atmospheric densities derived from orbital changes measured for satellites of dif-46 ferent shapes. The density comparisons suggest that standard assumptions about sur-47 face scattering in orbit may not be appropriate, but rather assuming scattering that is 48 less random may be better at upper altitudes near and above 500 km. Assuming the wrong 49 scattering physics could lead to errors in densities derived from satellite orbital data of 50 $\sim 25\%$ and satellite orbit prediction variations of 10s-100s of meters along the satellite 51 track. 52

53 1 Introduction

Modeling accurate physical aerodynamic drag coefficients for satellites in orbit is 54 critical for understanding atmospheric drag variability and maintaining space domain 55 awareness in the Low Earth Orbit (LEO) environment. Drag coefficients (C_D) are used 56 to calibrate density models and extract atmospheric density and wind information from 57 on-orbit observations (Doornbos, 2012; Sutton et al., 2007; P. M. Mehta et al., 2017; March, 58 Visser, et al., 2019). Furthermore, inferred wind speeds and directions are sensitive to 59 even small changes in energy accommodation in C_D models (Ching et al., 1977). Accu-60 rate drag coefficients in orbit are also needed to improve satellite orbital lifetime esti-61 mates. The drag coefficient may change throughout the geospace environment due to 62 model dependencies on composition and altitude. Drag coefficient models tend to diverge 63 at altitudes above 400 km (Pilinski et al., 2013; Pardini et al., 2009), which is where there 64 are greater uncertainties in gas-surface interaction assumptions. 65

Physical satellite drag coefficients are characterized by gas-surface interactions (GSIs) describing the dynamics of the energy and momentum exchange between an impinging atmospheric atom or molecule and a satellite surface. GSI models characterize the angular distribution and velocity of the reemitted gas particles. These reemitted particle distributions can be described as diffuse, specular, or quasi-specular, with varying energy and momentum accommodation at the satellite surface. GSI physical dynamics are expected to be dependent on surface roughness, cleanliness, composition, and temperature as well as incident gas composition, temperature, velocity, and angle.

74 GSI models are typically characterized by parameters that constrain the velocity and angular distribution of reemitted gas particles, the most fundamental of which are 75 accommodation coefficient parameters. The accommodation coefficients that quantify 76 the energy and momentum exchange between the molecules and the surface include the 77 energy accommodation coefficient, α , and the normal and tangential momentum accom-78 modation coefficients, σ_n and σ_t . The energy accommodation coefficient controls the amount 79 of energy that the incident molecules lose upon reemission, as indicated by Eq. (1) (Wachman, 80 1962).81

 $\alpha = \frac{E_i - E_r}{E_i - E_s} = \frac{T_i - T_r}{T_i - T_s}$

In Eq. (1), E_i is the kinetic energy from incident molecules hitting the surface, E_r is the kinetic energy leaving the surface from reemitted molecules, and E_s is the energy that would be transported away from the surface from reemitted molecules in thermal equilibrium with the surface. T_i , T_r , and T_s are the kinetic temperatures corresponding to the kinetic energies described above, respectively.

The amount of energy transferred determines the velocity of the reflected molecules, so energy accommodation is related to the reemission velocity. If $\alpha = 0$, then $E_i = E_r$, and the incident molecules are reemitted with speeds equal to the incident relative speed between the atmosphere and the satellite. If $\alpha = 1$, then $E_r = E_s$, and the incident molecules are reemitted with slower speeds after having reached thermal equilibrium with the satellite surface. The energy accommodation coefficient is split into normal and tangential components in some models.

The momentum accommodation coefficient represents the change in momentum between the incident molecules and the reemitted molecules. The momentum accommodation coefficient is often separated into normal and tangential exchange components, as this choice is associated with better agreement between theoretical calculations and observations (Schaaf & Chambre, 1958).

100
$$\sigma_n = \frac{p_{i,n} - p_{r,n}}{p_{i,n} - p_{s,n}}$$

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$$\sigma_t = \frac{p_{i,t} - p_{r,t}}{p_{i,t} - p_{s,t}} \tag{3}$$

(1)

(2)

In Eqs. (2) and (3), the subscripts n and t signify normal and tangential components. The incident momentum flux is given by p_i , the reemitted momentum flux is given by p_r , and p_s is the momentum flux from diffusely reemitted molecules in thermal equilibrium with the surface (Storch, 2002).

The normal and tangential momentum accommodation coefficients are related to the reemitted velocity distribution (Walker et al., 2014). Diffuse scattering occurs when both σ components are equal to unity. Specular scattering occurs when both σ components are equal to 0. When σ_n and σ_t are valued between 0 and 1, reemission is quasispecular in nature. In general, larger momentum and energy accommodation leads to more diffusive scattering and smaller reflected velocities. Smaller momentum and energy accommodation coefficients yield more quasi-specular scattering.

Table 1. C_D Models and Parameters

C_D Model	Parameters	Reemitted Distribution Characteristics
Sentman Schamberg CLL	$egin{array}{lll} lpha \ lpha, u, \phi_0 \ \sigma_n, \sigma_t \end{array}$	Diffuse Diffuse, specular, or quasi-specular Diffuse, specular, or quasi-specular with scattering kernels

A great deal of theoretical groundwork on GSI physics has been developed, with 113 fundamental contributions in the 1960s laying the foundation for updated modeling frame-114 works today. Early efforts from Sentman (1961) proposed diffuse scattering with com-115 plete or near-complete energy accommodation. Schamberg's model (Schamberg, 1959) 116 could vary the reemission angle via the parameter ν and the lobe width via the param-117 eter ϕ_0 to achieve a wide range of diffuse, specular, and quasi-specular reemitted par-118 ticle distributions. Advancements by Goodman (Goodman & Wachman, 1967), and more 119 recently the body of work from Moe and Moe (K. Moe et al., 1972; M. M. Moe et al., 120 1993; K. Moe & Moe, 2005), consider the effect of surface contamination or cleanliness. 121 This led to the implementation of variable energy accommodation coefficients based on 122 surface coverage by adsorbed gas particles, particularly atomic oxygen. Diffuse reflec-123 tion with variable or incomplete accommodation (DRIA) is the current predominant method 124 for computing drag coefficients for satellites in LEO. Many recent approaches (Sutton 125 et al., 2007; Doornbos, 2012; March, Visser, et al., 2019) select the most appropriate value 126 for energy accommodation in a DRIA model based on data consistency metrics or other 127 insights from orbital measurements. Following the work of Moe and Moe, Pilinski et al. 128 (2013) implemented DRIA with Langmuir isotherm-based variable energy accommoda-129 tion based on atomic oxygen number density in the Semi-Empirical Satellite Accommo-130 dation Model (SESAM). Recent interest in the quasi-specular scattering model which 131 implements scattering kernels and allows for variable accommodation known as the Cercignani-132 Lampis-Lord (CLL) model (Cercignani & Lampis, 1971, 1997; Lord, 1991) marks a sig-133 nificant shift in the computation of aerodynamic drag coefficients. The CLL model moves 134 away from fully diffuse scattering and the DRIA framework that has characterized much 135 of the orbital drag coefficient modeling advancements since the 1960s. Walker et al. (2014) 136 and P. M. Mehta et al. (2014) implemented the CLL model with variable Langmuir-based 137 energy accommodation to achieve scattering with narrower angular distribution lobes. 138 Table 1 provides a summary of the relevant foundational C_D models examined in this 139 analysis, their corresponding GSI parameters, and achievable reemitted distribution char-140 acteristics. 141

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1.1 Insights from Laboratory Molecular Beam Experiments

Historical, current, and near-future laboratory molecular beam experiments pro-143 vide insight into and estimates of physical gas-surface interactions. Studies by Cook (1965) 144 led to the widely-used fixed drag coefficient of 2.2. Historically, many studies have ei-145 ther relied on laboratory measurements from clean, pristine surfaces that likely do not 146 fully represent in-orbit conditions or have not controlled for surface contamination leav-147 ing it difficult to apply the results to orbital environments. Quasi-specular scattering has 148 been demonstrated for clean materials under ultra-high vacuum conditions at orbital ve-149 locity conditions (Murray et al., 2017). 150

Some laboratory experiments have provided insight into the ideas that 1) gas-surface
 interactions may depend on the incident angle between the flow and the satellite surface
 normal, and 2) different gas constituents may scatter in unique ways. There has been
 some work that supports the idea that energy and momentum accommodation coefficients are expected to be a function of incident angle and constituent mass. Goodman

and Wachman (1967) and Wachman (1962) derived a formula for energy accommodation as a function of the ratio of the mass of the incident molecule to the mass of the surface material based on lattice theory and experimental data.

In general, laboratory experiments investigating helium scattering have concluded 159 that helium reemission is characterized by low energy and momentum accommodation. 160 Seidl and Steinheil (1974) derived a tangential momentum accommodation coefficient 161 of 0.012 for helium scattered off of sapphire (Al_2O_3) , which was the most similar surface 162 coating to an oxygen-covered surface. Gaposchkin (1994) infers the reflected scattering 163 angle for helium from the data presented in the Seidl and Steinheil (1974) study and shows 164 that θ_r increases as the incident angle θ_i increases. As the incident angle approaches 90°, 165 θ_r is nearly equal to θ_i , which is consistent with specular reflection. Gaposchkin's work 166 also suggests that the scattering distribution direction and lobe width depends on the 167 incident velocity. For intermediate velocities (between 3 and 6 km/s), the scattering dis-168 tribution is more diffuse. For lower velocities, scattering is more specular. Liu et al. (1979) 169 found tangential momentum accommodation to be 0.046 for helium off of aluminum and 170 anodized aluminum plates. 171

Laboratory molecular beam experiments with heavier gas constituents typically re-172 port higher energy and momentum accommodation. Minton et al. (2004) derived energy 173 accommodation coefficients of ~ 0.91 for CO₂ and Ar scattering at low incident angles. 174 N_2 scattering experiments by N. A. Mehta et al. (2018) indicate that normal momen-175 tum accommodation is high (between 0.4 and 0.6), while tangential momentum accom-176 modation is low (between 0 and 0.2). Murray et al. (2017) reports O and O_2 scattering 177 with near-complete energy accommodation associated with thermal desorption off of an 178 Au surface, while O and O_2 scattering off of SiO₂ and flat graphite surfaces appeared 179 more specular with lower energy and momentum accommodation. 180

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1.2 Gas-Surface Interactions in Orbit

Early efforts to constrain gas-surface interactions in space relied on experimental 182 laboratory data. As space-based missions increased and on-orbit measurement techniques 183 advanced, more studies were developed which used observational methods to determine 184 the aerodynamic coefficients of different spacecraft from the attitude or orbital trajec-185 tory changes. In these studies, GSI and surface accommodation are typically investigated 186 by considering the spacecraft attitude and geometry and comparing to different mod-187 els. The on-orbit drag force is typically measured by either accelerometer data or ob-188 served orbital decay, where the product of density and C_D is determined from the ob-189 servations. Paddlewheel satellite studies (K. Moe, 1966; Beletskii, 1970; Imbro et al., 1975) 190 fall into this category, where paddlewheel satellites measure both spin and orbital de-191 cay. Spherical satellites (Bowman & Moe, 2005; K. Moe & Bowman, 2005; Pilinski et 192 al., 2011) and satellites with complex geometries (Ching et al., 1977; Pardini et al., 2009; 193 March et al., 2021) have also been used in these types of observational studies. Many 194 of these studies have primarily assumed diffuse reflection with incomplete energy accom-195 modation (DRIA) and have attempted to constrain the energy accommodation coeffi-196 cient. In general, these studies have concluded that energy accommodation should be 197 near-complete at altitudes near 200 km. At higher altitudes near 400-500 km, energy ac-198 commodation coefficients have been estimated to be closer to 0.85 (March, Visser, et al., 199 2019; March et al., 2021). A similar approach has been employed in Bernstein et al. (2020) 200 through multi-satellite density comparisons, which is expanded upon in this paper. It 201 is also important to recognize that results obtained from these types of observational meth-202 ods can be dependent on modeled atmospheric densities and are thus subject to the model 203 biases and uncertainties. 204

Much of the work towards understanding on-orbit GSI dynamics has focused on the role of atomic oxygen. Orbital measurements in the 1960s and 1970s from pressure

gauges and mass spectrometers led to the idea that satellite surfaces in orbit are con-207 taminated with adsorbed molecules (Carter et al., 1969; Hedin et al., 1973). These early 208 insights formed the foundation for the body of work of Moe and Moe (K. Moe et al., 1972; 209 K. Moe & Moe, 1992; M. M. Moe et al., 1993, 1995; K. Moe & Moe, 2005), who mod-210 eled the adsorption of a monolayer of atomic oxygen using the Langmuir isotherm (K. Moe 211 et al., 1972). Bowman and Moe (2005) and K. Moe and Bowman (2005) theorize that 212 GSIs and the associated energy and momentum accommodation coefficients are depen-213 dent on atomic oxygen adsorption, which is expected to control the level of surface clean-214 liness. The result of these studies is that on-orbit reemission is considered to be primar-215 ily diffuse with high energy accommodation (>0.8) at altitudes studied (<300 km) due 216 to surface contamination by adsorbed atomic oxygen. Additional support for this line 217 of reasoning comes from a direct measurement of the reemission angle of scattered atomic 218 oxygen from a vitreous carbon surface on the STS-8 Space Shuttle flight. Diffuse ree-219 mission with a nearly cosine distribution and near full accommodation was observed at 220 an altitude near 225 km (Gregory & Peters, June 1986). Near-fully diffuse angular dis-221 tributions of reemitted molecules were also measured near 180 km (Beletskii, 1970). At 222 higher altitudes (800-1000 km), evidence of more quasi-specular reemission behavior has 223 been observed for spacecraft materials (Harrison & Swinerd, 1996). To support these re-224 sults, it has been proposed that surface contamination is lower at higher altitudes. 225

The separation of on-orbit force coefficients into drag and lift components signi-226 fies the dependence of the drag force on incident angle of the flow. Drag coefficients cor-227 respond to the force parallel to the flow, while lift coefficients correspond to the force 228 perpendicular to the flow. Schamberg (1959) provides drag and lift coefficient expres-229 sions. Drag and lift coefficient expressions for a flat surface element as a function of in-230 cident angle are also derived by Schaaf and Chambre (1958). Ching et al. (1977) stud-231 ied the ratio of lift to drag on the S3-1 satellite, which orbited at an altitude of 175 km. 232 They found that the ratio of lift to drag force increases as energy accommodation de-233 creases for all cases from specular to diffuse. 234

Recent missions highlight the ongoing interest in investigating GSI models and drag 235 coefficients in orbit. The $\Delta Dsat$ CubeSat mission (Virgili & Roberts, 2013) was designed 236 to study how the drag coefficient changes for different satellite surface materials using 237 steerable fins. This study only examines the drag forces for the surface area exposed to 238 the flow, with no net lift forces, which allows the drag coefficient to be determined from 239 the variation in the trajectory over time based on a free-parameter fitting scheme in the 240 orbit determination process. The SOAR satellite (Crisp et al., 2021) was designed specif-241 ically to investigate how incident angle and surface material impact gas-surface inter-242 actions in orbit. Different panels will have different surface coatings to represent a range 243 of expected GSI dynamics. For the SOAR experiment, data will only provide informa-244 tion for altitudes below 400 km. Thus, this is an area of emerging research and these in-245 quiries currently do not have satisfactory answers. 246

DRIA is the standard model used to compute drag coefficients for satellites in or-247 bit for research purposes. Studies which derive energy accommodation coefficients tend 248 to suggest that, given diffuse reemission, energy accommodation decreases with increas-249 ing altitude and decreasing atomic oxygen (M. M. Moe et al., 1995; Pardini et al., 2009). 250 However, this conclusion may still be limited by the fact that diffuse reemission is as-251 sumed. It is also possible that, as altitude increases and atomic oxygen decreases, ree-252 mission gets more quasi-specular with less tangential momentum exchange. This is some-253 thing that the DRIA model framework is not able to test. Additionally, it is possible that 254 different gas constituent species, like atomic oxygen and helium, scatter in different ways. 255 DRIA with fixed α is not able to account for these scattering differences. Some DRIA 256 implementations, like SESAM (Pilinski et al., 2013) and RSM (P. M. Mehta et al., 2014), 257 try to partially incorporate this effect by implementing variable energy accommodation 258 according to atomic oxygen number density, but these are still constrained by the assump-259

tion of complete tangential momentum accommodation. Understanding the potential errors associated with the DRIA model is particularly significant when trying to calibrate atmospheric density models and datasets. The construction of thermospheric density and wind models from on-orbit observations depends on drag coefficients. Errors from C_D models, like DRIA, can lead to different biases at different altitudes/pressure levels.

Additionally, there is a disconnect in the theoretical foundations between the cur-265 rent state-of-the-art DRIA C_D models with their dependence on atomic oxygen adsorp-266 tion and the laboratory experiment research on the effects of incident mass and angle. 267 This motivates the present study into appropriate GSI modeling assumptions across a 268 wide range of orbital conditions. Some work has been done to try to directly extrapo-269 late laboratory experimental data to orbital drag coefficient models (Poovathingal et al., 270 2019; N. A. Mehta et al., 2018). The study presented in this paper does not directly fit 271 laboratory data to an orbital drag coefficient model, but rather the authors consider in-272 sights from laboratory experimental results to constrain C_D model parameters which can 273 then be applied to orbital density data processing and checked for consistency. 274

Following the body of theoretical and observational studies into GSI physics sum-275 marized in this section, the drag coefficient remains uncertain at upper thermospheric 276 altitudes. This is a result of a lack of on-orbit measurements as well as variability in C_D 277 model assumptions related to GSI dynamics and atmospheric constituents, which are fur-278 ther explored in Section 2. It is important to constrain this drag coefficient uncertainty 279 because it can cause altitude-dependent density biases. In order to determine which C_D 280 model assumptions may be most appropriate at upper thermospheric altitudes, the au-281 thors evaluate inter-satellite density consistency. The authors derive and compare den-282 sities for different shaped satellites, including GRACE (an elongated satellite) as well 283 as a set of compact satellites, for a variety of different C_D models. This method is fur-284 ther described in Section 3. In Section 4, the authors present results highlighting which 285 C_D models produce the most consistent densities across the atmospheric conditions sam-286 pled as well as further interpretation of the relevant C_D model assumptions. Conclud-287 ing discussion is included in Section 5. 288

²⁸⁹ 2 Current Drag Coefficient Model Formulations and Sensitivities

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2.1 Incident Angle and Tangential Momentum Accommodation

When considering C_D sensitivity to GSI dynamics, incident angle of incoming molecules plays an important role that in many current models is not fully incorporated. Incidence angle describes the angle between the flow and the satellite surface normal. In this section, the influence of the incident angle is described. This helps to show why the drag coefficient for elongated satellites, compared to compact satellites, may be more sensitive to the nature of gas-surface interactions at high incidence angles.

Accommodation assumptions in drag coefficient models can explicitly impact the 297 increase or decrease in drag. Making assertions about the effect on the drag requires care-298 ful specification of the orientation of the satellite surface relative to the incident flow, 299 or the incident angle. The force on the satellite is given by the change in momentum along 300 the velocity direction of the incident gas molecules. To help conceptualize the drag ef-301 fects, the case of a flat plate moving perpendicular to the flow can be examined. Fig. 1 302 illustrates this scenario, showing both specular scattering (Fig. 1a) and diffuse scatter-303 ing (Fig. 1b) cases for an incident molecule with velocity V_i impacting the plate and reemit-304 ting with velocity V_r . The incident molecule impacts the flat surface with an incidence 305 angle of 0° from the surface normal. In the specular scattering case, energy E is con-306 served such that $E_i = E_r$ for the incident and reflected molecules. Then it follows that 307 $|V_i| = |V_r|$. This means that in the flow direction normal to the satellite surface, the 308 change in momentum which corresponds to the force on the satellite is maximized be-309



Figure 1. Specular scattering (a) and diffuse scattering (b) for a flat plate moving perpendicular to the flow direction V_i . Specular scattering (c) and diffuse scattering (d) for a flat plate moving parallel to the flow direction V_i .

cause V_r is equal in magnitude and opposite in direction to V_i . The change in momen-310 tum in this case becomes $2m_iV_i$. In the diffuse scattering case, the angular probability 311 distribution of the reemitted molecule is a cosine distribution, meaning that the molecule 312 has some probability of scattering in any direction. As such, integration over the prob-313 ability distribution is required. The change in momentum between the reflected molecule 314 and the incident molecule will be smaller than $2m_iV_i$ due to this cosine angular prob-315 ability distribution of reflection. For the case of a flat plate moving perpendicular to the 316 flow, diffuse scattering leads to smaller drag coefficients than those produced from spec-317 ular scattering. 318

An additional scenario to consider is a molecule impacting a flat plate surface with 319 an incident angle approaching 90° . This scenario can be visualized as a flat plate mov-320 ing parallel to the flow, also shown in Fig. 1. In the completely specular scattering case 321 shown in Fig. 1c, momentum and energy accommodation is 0. This means that the change 322 in momentum is 0, and there is no drag from the incident molecule. In the diffuse scat-323 tering case shown in Fig. 1d, the cosine angular probability distribution of reflection sug-324 gests that the reemitted molecule will scatter in a random direction which will impart 325 some amount of non-zero momentum transfer to the flat plate surface. For the case of 326 a flat plate moving parallel to the flow, diffuse scattering leads to larger drag coefficients 327 than those produced from specular scattering. 328

This example is helpful to show how an elongated satellite with extended surface area parallel to the flow is likely to be more sensitive to tangential momentum accommodation and the scattering direction than compact shaped satellites, especially as helium increases in the orbital environment. This is an important clarification to make as it highlights the significance of side scattering in C_D computations. As helium increases in the atmosphere, the thermal velocity increases due to helium's low molecular mass. This causes the ratio of the relative velocity between the satellite and flow to the ther-



Figure 2. Force contributions of the side plates relative to the front plate for GRACE using the Sentman diffuse model (a), the CLL model (b), and the Schamberg model (c) in an atomic oxygen atmosphere (blue) and a helium atmosphere (orange).

mal velocity, known as the speed ratio, to decrease. Side scattering becomes increasingly 336 important as the speed ratio decreases in helium-rich atmospheric conditions and ran-337 dom thermal motion increases the interactions of atmospheric gas with the surfaces par-338 allel to the flow. As a result of increased impacts to the satellite, the drag coefficient in-339 creases by an amount dependent on the tangential momentum accommodation. Addi-340 tionally, helium may have different gas-surface interaction dynamics with the satellite 341 surface than other, heavier atmospheric constituents like atomic oxygen, which can im-342 pact the nature of the scattering. These concepts provide motivation for the study pre-343 sented in this paper. 344

The effect of side scattering for an elongated satellite can be examined by compar-345 ing the relative force contributions of different panels on the Gravity Recovery and Cli-346 mate Experiment (GRACE) satellite. The authors compute C_D for GRACE using the 347 panel method, which is an analytical approach to geometry modeling that uses several 348 flat plates to estimate a simplified satellite geometry and does not account for multiple 349 scattering instances. An alternative approach to the panel method would be to use Di-350 rect Simulation Monte Carlo (DSMC) methods, which simulate rarefied gas flows to ad-351 dress shadowing and handle multiple scattering events. As the geometry of GRACE is 352 relatively simple, the panel method should be appropriate for these conditions. 353

Fig. 2 shows the force contribution of the side panels relative to the front panel for 354 GRACE for one day of the year in 2003 using three different C_D models. The force con-355 tribution for each panel is given by the product of the drag coefficient and the surface 356 area A of the satellite panel. The ratio of the combined force contributions from the side 357 panels to the front panel force contribution is shown for an atmosphere composed en-358 tirely of oxygen in blue and for an atmosphere composed entirely of helium in orange. 359 Panel force contributions are computed using the Sentman diffuse C_D model in Fig. 2a, 360 the CLL C_D model with 75% normal and tangential momentum accommodation in Fig. 361 2b, and the Schamberg quasi-specular C_D model with 50% energy accommodation in 362 Fig. 2c. 363

For all three C_D models in Fig. 2, comparison of the blue and orange lines show 364 that the side plates contribute approximately twice as much to the total force in a he-365 lium atmosphere than in an atomic oxygen atmosphere. This supports the notion that 366 side scattering has a greater effect in helium-rich atmospheres. Additionally, compar-367 ison of the C_D models in Fig. 2 shows that as tangential momentum accommodation in-368 creases, the relative force contribution of the side plates increases. Tangential momen-369 tum accommodation is lowest in the quasi-specular Schamberg C_D model used in Fig. 370 2c, increases to 0.75 in the CLL C_D model in Fig. 2b, and maximizes to full accommo-371 dation in the Sentman diffuse C_D model in Fig. 2a. Tangential momentum accommo-372 dation significantly impacts the scattering angular distribution of a C_D model. As σ_t de-373

creases, the scattering lobe becomes more specular. Side drag is reduced by decreasing σ_t . Drag coefficient sensitivity to GSI parameters are further explored in the following subsection.

377

2.2 Drag Coefficient Sensitivities for Different Shaped Satellites

Satellites of different shapes are sensitive in different ways to GSI parameters. This 378 idea has already been introduced in the discussion of incident angle and lift effects in Sec-379 tion 2.1. To improve C_D modeling across LEO conditions, the GSI assumptions must 380 be appropriate across satellite geometries. It is important to recognize how C_D for dif-381 ferent geometries is expected to change according to various GSI assumptions and in dif-382 ferent environmental conditions. Understanding these sensitivities will help with infer-383 ring C_D model limitations and inconsistencies from multi-satellite density comparisons, 384 which is what the authors are undertaking with this study. For this reason, in this sec-385 tion C_D sensitivities to different modeling parameters are examined for both GRACE 386 (an elongated satellite) and a compact sphere. 387

First, energy accommodation in the commonly-used DRIA C_D modeling framework 388 can be examined. DRIA C_D modeling assumes diffuse reemission with $\alpha < 1$, where 389 α is the single GSI input parameter. Many C_D modeling approaches which rely on DRIA 390 assume a fixed value for α (Sutton et al., 2007; March, Visser, et al., 2019), while oth-391 ers allow α to vary according to certain atmospheric inputs (Pilinski et al., 2013). Be-392 cause DRIA depends on only one GSI parameter, it is a simple and advantageous ap-393 proach to physical C_D modeling. Fig. 3 shows Sentman DRIA C_D sensitivity to α for 394 GRACE, represented by the blue lines, and a sphere, represented by the orange lines. 395 In Fig. 3a C_D has been computed for an atmosphere composed entirely of atomic oxy-396 gen, while C_D in Fig. 3b has been computed for helium-dominated atmospheric condi-397 tions. Percentages included in Fig. 3 indicate C_D sensitivity to the range of α values from 0 to 1, computed as $\frac{\max(C_D) - \min(C_D)}{0.5*(\max(C_D) + \min(C_D))}*100$. The drag coefficient decreases for both GRACE and the sphere as α increases, since increasing energy accommodation results 398 399 400 in smaller reflected velocities. Drag coefficient sensitivity percentages are similar for both 401 GRACE and the sphere. Comparison of C_D in atomic oxygen-dominated conditions in 402 Fig. 3a and helium-dominated conditions in Fig. 3b reveals that C_D for the sphere re-403 mains nearly unchanged as the atmospheric composition decreases in mass. In contrast, 404 C_D for GRACE increases significantly in the helium-dominated atmosphere compared 405 to the oxygen-dominated atmosphere due to side scattering effects discussed in Section 406 2.1. Thus, GRACE is more sensitive to incident mass effects than the sphere for the Sent-407 man DRIA C_D model. 408

In contrast to DRIA, the CLL C_D model has two input GSI parameters: the nor-409 mal and tangential momentum accommodation coefficients. CLL C_D modeling allows 410 reemission to be diffuse ($\sigma_n = \sigma_t = 1$), specular ($\sigma_n = \sigma_t = 0$), or quasi-specular 411 $(0 < \sigma_n < 1 \text{ and } 0 < \sigma_t < 1)$. CLL can also be parameterized for back-scattering by 412 setting $1 < \sigma_t < 2$ (Lord, 1991), however this approach is not considered to be rele-413 vant in the context of physical on-orbit GSIs and thus does not apply to the work in this 414 paper. Fig. 4 shows CLL C_D sensitivity to σ_t , where σ_n has been fixed to 0.5, for GRACE 415 and a sphere represented by blue and orange lines, respectively. Just like in Fig. 3, Fig. 416 4a represents atomic oxygen atmospheric conditions and Fig. 4b represents helium at-417 mospheric conditions. Percentages in Fig. 4 are computed the same way as in Fig. 3. 418 As σ_t increases, C_D increases for both GRACE and the sphere. As σ_t approaches 1, the 419 reemission angle gets closer to the surface normal, causing drag to increase for high-incident 420 angle impacts. Drag coefficient sensitivity to σ_t for the sphere remains nearly unchanged 421 between the atomic oxygen and helium conditions. GRACE C_D sensitivity to σ_t , how-422 ever, significantly increases in the helium-dominated atmospheric conditions. In other 423 words, GRACE C_D changes more in response to changing σ_t in helium-rich conditions 424



Figure 3. Sentman DRIA C_D sensitivity to the energy accommodation coefficient parameter for GRACE (blue curve) and a sphere (orange curve) in 100% atomic oxygen conditions (a) and 100% helium conditions (b).



Figure 4. CLL C_D sensitivity to the tangential momentum accommodation parameter σ_t for GRACE (blue curve) and a sphere (orange curve) in 100% atomic oxygen conditions (a) and 100% helium conditions (b). Fixed $\sigma_n = 0.5$.

than in atomic oxygen-rich conditions. Just like for the DRIA model, the CLL C_D model for GRACE is more sensitive to incident mass effects than for the sphere.

Fig. 5 shows CLL C_D sensitivity to both σ_n on the x-axis and σ_t on the y-axis for GRACE in atomic oxygen conditions (Fig. 5a), GRACE in helium conditions (Fig. 5b), a sphere in atomic oxygen conditions (Fig. 5c), and a sphere in helium conditions (Fig. 5d). The drag coefficient is highly sensitive to both σ_n and σ_t . GRACE C_D in heliumrich conditions, shown in Fig. 5b, exhibits the greatest variability of the four panels. Additionally, C_D for GRACE changes more significantly than for a sphere as conditions transition from an atomic oxygen atmosphere to a helium atmosphere.

Figures 3, 4, and 5 show that both the elongated satellite (GRACE) and the sphere 434 are more sensitive to tangential momentum accommodation than energy accommoda-435 tion, especially in helium-rich conditions. This discussion motivates using satellites with 436 different shapes for evaluating C_D model assumptions. This research approach follows 437 in the steps of previous investigations of lift-to-drag measurements. The behavior of low-438 incidence angle interactions compared to high-incidence angle interactions leads to dif-439 ferences in the C_D response to gas-surface interaction parameters. For this research, C_D 440 model assumptions are evaluated in order to constrain appropriate on-orbit gas-surface 441 interaction parameters. 442



Figure 5. CLL C_D sensitivity to normal and tangential momentum accommodation for GRACE in atomic oxygen conditions (a), GRACE in helium conditions (b), a sphere in atomic oxygen conditions (c), and a sphere in helium conditions (d).

443 **3 Method**

The analysis presented in this paper involves evaluating inter-satellite density con-444 sistency in order to constrain and improve physical gas-surface interaction assumptions 445 in C_D models, especially in helium-rich regimes. The density comparisons are made be-446 tween GRACE (an elongated satellite) and a set of compact, or near-spherical, satellites 447 in order to evaluate the incidence angle effect. This analysis method extends the approach 448 in Bernstein et al. (2020), in which density discrepancies between GRACE and a selec-449 tion of compact satellites were analyzed over four distinct week-long time periods. Den-450 451 sity discrepancies are defined by Eq. (4):

452

$$\Delta R = \left(\frac{\rho_{obs}}{\rho_{mod}}\right)_{GRACE} - \left(\frac{\rho_{obs}}{\rho_{mod}}\right)_{sphere} \tag{4}$$

The density values in Eq. (4) represent weighted averages along the orbit, referred 453 to as effective densities. The density discrepancy metric ΔR in Eq. (4) is also approx-454 imately equal to a function of C_D ratios (see Section III in Bernstein et al. (2020)). Ex-455 amining observed-to-modeled density ratios allows for the comparison of analogous ob-456 served quantities for GRACE and the compact satellites while removing differences in 457 densities due to the satellites not sampling identical altitudes. This is based on the as-458 sumption that atmospheric model bias should be the same for GRACE and the compact 459 satellites in the comparisons. To better account for altitude-dependent model biases, a 460 fitting scheme is implemented which is discussed more in Section 3.2. 461

Effective densities for the compact satellites are obtained from orbital Two-Line 462 Element (TLE) data following the procedure in Picone et al. (2005). The United States 463 Space Command (USSPACECOM) 18th Space Control Squadron (18SPCS) archives or-464 bital tracking data for most known Earth-orbiting artificial satellites in the form of TLE 465 entries, which provide an expansive record of historical orbital element information. Ef-466 fective densities are computed from TLEs based on how the satellite orbit changes over 467 time, represented by the change in mean mean motion. Effective densities for GRACE 468 come from accelerometer-derived density datasets from Sutton et al. (2007). These den-469 sities are scaled with force coefficients modeled using a variety of C_D modeling frame-470 works and integrated along time intervals consistent with the TLE-derived effective den-471 sities. 472

The time periods of analysis in Bernstein et al. (2020) were selected to sample a 473 variety of solar and geomagnetic conditions while benefiting from the fidelity of the Ther-474 mosphere Ionosphere Electrodynamics General Circulation Model (TIE-GCM) (Qian et 475 al., 2014). The results indicated that tangential momentum accommodation plays a sig-476 nificant role in deriving consistent densities from drag data. However, various limitations 477 were introduced into the analysis as a result of these time periods and data/model con-478 siderations. These include the signal-to-noise ratio associated with the TLE data, TLE 479 and accelerometer dataset differences (Bernstein et al., 2021), local time and latitude dif-480 ferences between the orbits of GRACE and the compact objects, sensitivity to atmospheric 481 model error, and challenges associated with making statistically significant conclusions 482 given the small amount of data. This work expands upon the study in Bernstein et al. 483 (2020) by addressing each of the above-mentioned limitations and proposing improve-484 ments to gas-surface interaction modeling based on density consistency results. 485

First, it was important to expand the datasets temporally. Among other advantages, this enables the authors to have enough data to find satellite orbit conjunctions for more reliable density comparisons. It also is helpful for increasing some of the TLE signal-to-noise through averaging, which enables the use of solar minimum TLE data for which atmospheric helium is in greater abundance.

Refining the density discrepancy analysis from Bernstein et al. (2020) to be robust against uncertainty can be summarized with four distinct measures: 1) significant data



Figure 6. Flow chart summarizing how GRACE and compact satellite data is processed to obtain the density discrepancy metric, ΔR

expansion, 2) increasing the time window of effective density integration to 8 days, 3) 493 making comparisons at narrower local time and latitude conjunctions, and 4) fitting the 494 vertical structure of observed-to-modeled density ratios to the ratio of two Bates pro-495 files in order to compare data from compact satellites at a range of altitudes. These are 496 concerted efforts to address possible sources of error in ΔR that are not C_D -related in 497 order to more convincingly make conclusions about C_D model effects. Each of these mea-498 sures will be described in the following subsections. A flow diagram summarizing how 499 ΔR is obtained from GRACE and compact satellite data is included in Fig. 6, the de-500 tails of which are further described in the following subsections. 501

502 503

3.1 Extending and Refining Aerodynamic Drag Observations during Conjunctions

To expand the datasets, this analysis covers the full range of available GRACE accelerometer-504 derived density data (Sutton et al., 2007; P. M. Mehta et al., 2017) from 2002-2010, in-505 clusive. This is a significant increase in time covered, as in Bernstein et al. (2020), only 506 four weeks of data were used. The empirical Naval Research Laboratory's Mass Spec-507 trometer and Incoherent Scatter Extended (MSIS) (Picone et al., 2002) atmospheric model 508 has been used to obtain composition and neutral density for this time period. This ex-509 panded time period covers the deep solar minimum years from 2008-2010 and thus sam-510 ples atmospheric conditions with more helium. 511

TLE data has decreased signal-to-noise during deep solar minimum times. This can 512 be addressed using insights from the TLE variability and error analysis described in Bernstein 513 et al. (2021). The time window of integration in the TLE mean mean motion process-514 ing is selected to be a moving window of 8 days. This means that the effective densities 515 derived from TLE data for the compact satellites are densities that have been averaged 516 over 8 days. A multi-day time window was chosen since relative error between TLE-derived 517 and accelerometer-derived densities decreases as the integration time interval increases 518 (Bernstein et al., 2021). However, integrating between 8 and 20 days produced very sim-519 ilar relative error profiles across density bins. Choosing the smallest time interval in this 520 range is preferable in order to avoid over-averaging. Choosing to integrate over 8 days 521 is also advantageous for avoiding certain short-term, solar-related periodicities, like the 522 13.5 day period associated with active solar longitudes and the tilted dipole structure 523 which is observed in the solar wind and interplanetary magnetic field (Nayar, 2006; Bouwer, 524 1992). It is relevant to note here that Emmert (2009) used 3-6 days as the time integral 525 when processing TLEs to derive long-term density trends in the thermosphere. 526

The GRACE accelerometer densities have been averaged over the same 8-day moving windows to ensure reliable comparisons with the compact satellites' TLE data. Consequently, many local or short-term density signatures have been averaged over in favor of increasing the signal-to-noise in the TLE densities. This highlights the trade-off between resolution and noise as a result of using TLE data. In addition to increasing the

NORAD ID	Common Name	Perigee h (km)	$i \ (deg)$	e	A/m estimate (m ² /kg)
00045	TRANSIT 2A	600	66.7	0.025	0.00673
00046	SOLRAD 1	590	66.7	0.023	0.0107
00060	EXPLORER 8	370	49.9	0.035	0.011
00932	EXPLORER 25	530	81.4	0.115	0.0091
02826	SURCAL NRL 160/CALSPHERE 3	725	69.9	0.0005	0.0815
02909	SURCAL NRL 150B/CALSPHERE 4	685	69.9	0.0003	0.0834
06073	VENUS LANDER	200	52.0	0.127	0.00181
07337	COSMOS 660	390	83.0	0.054	0.00534
12138	COSMOS 1238	415	83.0	0.080	0.00531
14483	COSMOS 1508	395	83.0	0.082	0.00537
20774	COSMOS 2098	405	83.0	0.090	0.00546
23278	COSMOS 2292	400	83.0	0.095	0.00531

Table 2.	High-Inclination	Compact	Satellites
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time window of integration, only leading GRACE satellite density data has been used in these comparisons. This is due to the identification of the density bias associated with trailing GRACE satellite data in Bernstein et al. (2021). GRACE A is the leading satellite from 2002 to late 2005, and GRACE B is the leading satellite from late 2005 onward.

Additional measures to make the GRACE and compact satellite density compar-536 isons robust include finding conjunction times for the satellite orbits in the comparison. 537 This involves selecting comparison times where the local time and latitude of the two 538 satellite orbits match within a reasonable range, thereby further ensuring that they should 539 be sampling similar orbital and density conditions. This means constraining the avail-540 able compact satellites to only those objects with high inclinations $(>\sim 50^{\circ})$ because GRACE 541 inclination is approximately 87°. Table 2 lists these high-inclination compact satellites. 542 For the density comparisons between GRACE and the compact satellites, both local time 543 and latitude conjunctions are found. Local time matching is done by comparing the lo-544 cal time at perigee for the compact object with the local time of the ascending or de-545 scending node for GRACE near the middle of the 8-day window. For windows in which 546 these local times match within ± 1 hour, latitude filtering is employed for the GRACE 547 data within the window. GRACE data in the window for which GRACE latitude is within 548 $\pm 20^{\circ}$ of the latitude at perigee for the compact satellite is used to compute the GRACE 549 effective densities for those conjunction windows. Finding conjunctions in local time and 550 latitude between GRACE and the compact objects helps to justify that the satellites should 551 be sampling similar observed and modeled atmospheric density conditions in the same 552 8-day window. 553

554 555

3.2 Density Comparisons at GRACE Altitude Through Bates Profile Fitting

It is clear from Table 2 that the set of compact satellites has a wide range of perigee 556 altitudes spanning ~ 200 - 700 km. GRACE, in contrast, orbits at altitudes between ~ 475 -557 525 km in the analysis time period prior to 2010. These significant altitude differences 558 between GRACE and some of the compact satellites can lead to differences in derived 559 densities as well as measured-to-modeled density ratios. These altitude-dependent dif-560 ferences might interfere with the inconsistencies that can be attributed to C_D model er-561 ror. For this reason, the authors needed to find a way to mitigate the altitude differences 562 among the objects. Simply removing all of the compact objects with perigees that ex-563 tended beyond a reasonable threshold, like ± 50 km from GRACE, was not a desirable 564 option because this would significantly cut the amount of usable data in the analysis. 565

Instead, a fitting scheme was implemented to fit the altitude structure of measuredto-modeled density ratios for the compact satellites to the ratio of two Bates profiles (Bates,

GRACE Avg. He/O	C_1	C_2	C_3	RMSE
0.04	0.7153	1.140	1015.2	0.0779
0.1	0.6902	1.1774	942.9	0.0595
0.23	0.6168	1.1804	793.7	0.0597

Table 3. Sentman Diffuse Density Ratio Altitude Fit Parameters

⁵⁶⁸ 1959). This fitting scheme is presented in detail in Appendix A and is based on the as-⁵⁶⁹ sumption that exospheric temperature differences are primarily causing the offset between ⁵⁷⁰ the observed and modeled densities. The purpose of obtaining a physically-informed ver-⁵⁷¹ tical profile of measured-to-modeled density ratio for the compact satellites was to be ⁵⁷² able to attain a value for $\frac{\rho_{obs}}{\rho_{rnod}}$ for a spherical satellite at the precise location of GRACE. ⁵⁷³ This value comes from the fitted density ratio profile, which is constructed using data ⁵⁷⁴ from each of the compact satellites.

Following the procedure in Appendix A, a parametric system of equations for the ratio of two mass density profiles was obtained (see Eq. (A6)). A simplified version of this expression is shown in Eq. (5), where an 'observed' density profile is signified by the superscript A and a 'modeled' density profile is signified by the superscript B.

 $\frac{\rho^{A}}{\rho^{B}} = f(C_{1}, C_{2}, C_{3})$ $\begin{cases}
C_{1} = \frac{n_{O,\ell}^{A}}{n_{O,\ell}^{B}} \\
C_{2} = \frac{n_{He,\ell}^{A}}{n_{He,\ell}^{B}} \\
C_{3} = T_{\infty}^{A}
\end{cases}$ (5)

579

In Eq. (5), T_{∞} represents exospheric temperature, and the subscript ℓ indicates the lower boundary altitude of the diffusive equilibrium region. Eq. (5) has been derived from Bates number density profiles and leverages model output from MSIS in order to reduce the number of unknown parameters. The three parameters to fit are represented by C_1 , C_2 , and C_3 . To maintain a reasonable number of parameters to fit, the thermosphere is assumed to be composed only of O and He at the satellite altitudes of interest.

Non-linear least squares fitting of Eq. (5) to the $\frac{\rho_{obs}}{\rho_{mod}}$ data for the compact satel-lites is performed. Note that GRACE data is not used for this fitting process, since the 586 587 goal of the fitting scheme is to get an altitude profile of measured-to-modeled density 588 ratios for the compact satellites from which a fitted value for the compact satellites at 589 GRACE's location can be obtained. The density ratio fits include C_D effects, since the 590 observed density data used to construct the fit depends on the C_D model used. Thus, 591 altitude profiles were fit to the density ratios for each C_D model used in the analysis. Ad-592 ditionally, the data used to construct the fit profiles were averaged into three bins. The 593 bins were sorted by average atmospheric helium-to-oxygen mass ratio along the GRACE 594 orbit. Helium-to-oxygen mass ratio is computed using MSIS. Fitted altitude profiles of 595 observed-to-modeled density ratios for the compact satellites are shown for the three av-596 erage He/O bins for two C_D model cases in Fig. 7. Fits for the compact satellite observed-597 to-modeled density ratios are shown for the Sentman diffuse C_D model in Figs. 7a, 7b, 598 and 7c as well as for the Schamberg specular model in Figs. 7d, 7e, and 7f. Tables 3 and 599 4 include fit parameter estimates and root mean squared error (RMSE) for the selected 600 Sentman diffuse and Schamberg specular C_D model cases. 601



Figure 7. Altitude fits of the observed-to-modeled density ratio for the compact satellites are shown by the red curves. Observed densities from panels (a), (b), and (c) have been derived using the Sentman diffuse C_D model. Observed densities from panels (d), (e), and (f) have been derived using the Schamberg specular C_D model. Compact satellite observed-to-modeled density ratios are marked by the round blue markers. The average GRACE density ratio is marked by the black star, and the interpolated geopotential altitude location of GRACE is marked by the dashed black vertical line.

GRACE Avg. He/O	C_1	C_2	C_3	RMSE
0.04	0.7740	1.232	1013.9	0.0773
0.1	0.7553	1.330	939.6	0.0596
0.23	0.6732	1.317	793.2	0.0636

Table 4. Schamberg Specular Density Ratio Altitude Fit Parameters

In Fig. 7, the panels show the data and fits for the three distinct He/O bins, where 602 the average He/O mass fraction increases in the atmosphere from the left-most panel to 603 the right-most panel. These He/O ratios are averages that have been weighted by drag 604 along the time intervals of integration (8 day periods) for the effective density, which cov-605 ers many orbits of the satellite. In other words, these He/O ratios are multi-day weighted 606 averages along the orbit of GRACE. Compact satellite observed-to-modeled density ra-607 tios are marked by the round blue markers. The density ratio altitude profile fits based 608 on the ratio of two Bates profiles are included as the red curves. The interpolated geopo-609 tential altitude location of GRACE is marked by the dashed black vertical line. The av-610 erage GRACE observed-to-modeled density ratio is included as the black star. Note that 611 this value is computed as the average GRACE effective density corresponding to con-612 junctions with each individual compact satellite, described in Section 3.1. 613

Note that the vertical structure of the density ratio is influenced by the C_D model. 614 The values of the fitted density ratio profile for the compact satellites are higher for the 615 specular C_D model case than for the diffuse C_D model case. This suggests that, for the 616 compact satellites, the observed densities derived with specularly modeled C_D are larger 617 than the observed densities derived with diffusely modeled C_D . This is because specu-618 lar C_D is smaller than diffuse C_D for spherical satellites, and the derived densities are 619 scaled by $\frac{1}{C_D}$. Fig. 7 also helps to visualize how GRACE density ratios compare to the 620 compact satellite density ratio fits. In the diffuse C_D model case, the average GRACE 621 observed-to-modeled density ratio, represented by the black star, is too high compared 622 to the red compact satellite density ratio fit line, especially in the left-most panel (Fig. 623 7a) which shows data for the smallest He/O bin. This suggests that GRACE densities 624 are too large compared to the compact satellite densities in these He/O conditions due 625 to C_D mismodeling by the Sentman diffuse model at these altitudes. Looking at the spec-626 ular C_D -based density ratios in Figs. 7d, 7e, and 7f, GRACE average density ratios are 627 more significantly offset from the compact satellite density ratio fit line for all three He/O 628 bins. This is consistent with the finding that specular C_D for GRACE is likely too high 629 compared to diffuse C_D , so specular C_D modeling yields densities for GRACE that are 630 too low. This is an important finding, as it shows that modeling the drag coefficient with 631 specular scattering yields significant density inconsistencies between the GRACE and 632 compact satellites. Specular scattering drives the C_D up for GRACE due to the front-633 plate scattering dominating the drag force. Specular scattering on the front plate yields 634 the maximum momentum exchange (more than diffuse) in GRACE gas-surface interac-635 tions, as is illustrated in Section 2.1. 636

For both the diffuse and specular C_D cases, and for all three He/O bins, the fits suggest that the MSIS model overpredicts the atomic oxygen number density and underpredicts the helium number density. This is shown in Tables 3 and 4, where C_1 , the estimated observed-to-modeled atomic oxygen number density ratio at the lower boundary, is less than 1 for all cases. The estimated fit parameter C_2 , the observed-to-modeled helium number density ratio at the lower boundary, is greater than 1 for all cases.

To compute ΔR , the value of the fit profile at GRACE's altitude (where the fit line intersects with the vertical black dashed line in Fig. 7) is subtracted from the average observed-to-modeled density ratio for GRACE at conjunctions. Recall that the purpose of doing the fitting was to resolve differences in perigee altitudes of the compact satellites while still making use of all of the compact satellite density data. This procedure to compute ΔR is implemented for a variety of C_D model cases and for each of the three GRACE He/O mass fraction bins.

650 4 Results

⁶⁵¹ Density discrepancies between GRACE and high-inclination compact satellites for ⁶⁵² the years 2002-2010 are shown in Fig. 8. The ΔR values are sorted and binned accord-



Figure 8. Binned average density discrepancies between the leading GRACE satellite and compact satellites. Each panel include ΔR results for one of the three He/O bins labeled by the panel title. Different colors represent different C_D models, and different markers are used to show that some C_D models share similar GSI assumptions. Vertical error bars represent the RMSE associated with the density ratio altitude profile fits.

ing to the He/O mass fraction along the GRACE satellite orbit. Each panel in Fig. 8 653 shows the ΔR results for one of the three He/O bins indicated by the panel title. The 654 left-most panel includes density discrepancy data in low-helium atmospheric conditions, 655 while the right-most panel includes data in helium-rich atmospheric conditions. Differ-656 ent colors represent different C_D models used to derive the effective density ratios. Each 657 of the colored data points are separated by a small amount in the horizontal direction 658 in order to distinguish C_D model differences. Vertical error bars represent the root mean 659 squared error of the compact satellite density ratios compared to the best-fit density ra-660 tio altitude profile. 661

The C_D models used are numbered and labeled in the legend of Fig. 8. The C_D 662 model labels in general include a combination of model name (Sentman, Schamberg, or 663 CLL), reemission characteristics (diffuse, quasi-specular (QS), or specular), and accom-664 modation coefficient parameters (α , σ_n , and σ_t). For model #2: Sentman Diffuse with 665 SESAM α , a fixed value for energy accommodation is not included because α varies based 666 on atomic oxygen pressure for this model (Pilinski et al., 2013). Different marker shapes 667 are used to help differentiate the model names: triangles for Schamberg models, circles 668 for Sentman models, and diamonds for CLL models. 669

A horizontal black dashed line is included at $\Delta R = 0$ in each of the panels of Fig. 8. Proximity to this zero-line is one indicator that the inter-satellite density comparisons are consistent for a given C_D model. As ΔR shifts away from 0, the orbital evidence may be indicating that the corresponding C_D model is not appropriate in the atmospheric conditions examined.

A few key trends can be discerned from the ΔR results in Fig. 8. Recall that scat-675 tering in gas-surface interactions can generally be described as diffuse, specular, or quasi-676 specular based on reemitted velocity. Diffuse and specular scenarios cover opposite ends 677 of the range of possible scattering dynamics, while the quasi-specular descriptor is more 678 broad and can be used to describe scattering that falls in between the diffuse and spec-679 ular cases, characteristic of incomplete momentum accommodation. In Fig. 8, all of the 680 C_D models which implement diffuse scattering (models #1-6) produce average ΔR val-681 ues that are greater than zero across the atmospheric conditions examined. The com-682 pletely specular C_D model (model #11) yields ΔR values that are less than zero in all 683 three He/O bins. The quasi-specular C_D models (models #7-10) produce ΔR values that 684

fall between those associated with the diffuse and specular models, where ΔR decreases as the QS models get increasingly closer to specular.

To interpret these trends, it is helpful to return to the discussion in Section 2.1 on 687 the effects of tangential momentum accommodation and incident angle for elongated satel-688 lites. The inter-satellite density comparisons, as well as the relative force contributions 689 of the side panels to the front panel for GRACE shown in Fig. 2, indicate that the scat-690 tering on the front plate dominates the drag force on GRACE. As the amount of helium 691 increases in the atmosphere, the effect of the assumed side panel GSI dynamics on the 692 total drag force is greater. Since $\Delta R > 0$ for the diffuse C_D models (models #1-6) in 693 Fig. 8, the inter-satellite density comparisons are indicating that the observed-to-modeled 694 density ratio for GRACE is too high compared to the sphere density ratio. Given that 695 the observed densities are inversely proportional to the drag coefficient, this indicates 696 that diffuse C_D is too small for GRACE with respect to the C_D response for the sphere. 697 Diffuse reemission would produce less drag than specular reemission on the front panel 698 of GRACE due to the low incident angle of the GSIs (see Fig. 1), which explains why 699 diffuse C_D may be too low for GRACE at these atmospheric conditions resulting in pos-700 itive density discrepancies between GRACE and the spheres. It is worth noting here that 701 models #1 (Schamberg diffuse) and #3 (Sentman diffuse) yield significantly different ΔR 702 values even though they both assume diffuse scattering with complete accommodation. 703 This may be because Sentman and Schamberg implement different formulations of dif-704 fuse scattering, which causes the two models to give different C_D values for a sphere and 705 even more so for an elongated satellite like GRACE. For one thing, Sentman (1961) and 706 Schamberg (1959) use different expressions for most probable velocity. Additionally, dif-707 fuse scattering in Sentman's model is represented with a cosine angular distribution, while 708 Schamberg's diffuse scattering distribution is modeled with a cone shape (Sentman, 1961; 709 Schamberg, 1959). As such, these two diffuse GSI models do not produce identical C_D 710 values, especially for elongated shapes, leading to significant differences in ΔR associ-711 ated with each model. 712

In contrast, Fig. 8 shows that ΔR is too low when a specular C_D model (model 713 #11) is used. This is because the density ratio for GRACE is too low in this case, as was 714 shown in Fig. 7. This indicates that the specular C_D for the elongated GRACE satel-715 lite is too high across the range of atmospheric He/O conditions presented here. As dis-716 cussed in Sections 2.1 and 3.2, specular scattering increases the C_D for GRACE because 717 the large change in momentum associated with specular scattering on the front plate dom-718 inates the drag force on GRACE. A similar argument can be made to discredit model 719 #10, a CLL QS model with $\sigma_n = 0.9$ and $\sigma_t = 0.1$ which yields scattering that is close 720 to specular, based on the significantly negative ΔR results associated with this C_D model. 721

The observed inter-satellite density comparisons in Fig. 8 also clearly show the ef-722 fect of decreasing energy accommodation in diffuse C_D models on density consistency. 723 As α decreases in the fixed DRIA models (models #3-6), average ΔR in all three He/O 724 bins decreases and gets closer to zero. This follows the discussion on C_D sensitivity to 725 model parameters in Section 2.2, where it was shown in Fig. 3 that diffuse C_D gets larger 726 as α decreases. Since GRACE C_D increases as α decreases, the observed-to-modeled den-727 sity ratios for GRACE get smaller as C_D DRIA models with less energy accommoda-728 tion are used. This results in smaller ΔR values associated with diffuse C_D models which 729 assume less energy accommodation. 730

In comparing the three panels of Fig. 8 from left to right, it is clear that the ΔR values associated with each C_D model shift down as helium increases in the atmosphere. The magnitude of this shift depends on C_D model assumptions. This trend shows how helium drives the speed ratio. As helium increases in the atmosphere, the speed ratio decreases causing C_D for the elongated GRACE satellite to increase. As the GRACE C_D increases with more helium, GRACE derived densities decrease resulting in a measurable decrease in ΔR . This is the case for all C_D models shown in Fig. 8, however the diffuse C_D models yield the greatest decrease in ΔR from the left bin to the right bin. This also is aligned with the discussion in Section 2.1, since more helium in the atmosphere leading to increased diffuse scattering at high incidence angles, like on the side panels of GRACE, causes C_D to increase more than if the side scattering were specular in nature. Thus, the behavior of the observed ΔR as helium increases in the atmosphere follows the theoretical examination of increased drag at high incidence angles from GSI models that assume full tangential momentum accommodation.

Examination of models #7-10 in Fig. 8 reveals that a quasi-specular C_D model with 745 incomplete normal and tangential momentum accommodation is more appropriate than 746 a diffuse or specular model at the altitudes and atmospheric conditions sampled in this 747 analysis. Since some of the QS models produce ΔR values that fall above the $\Delta R = 0$ 748 line and other QS models produce ΔR values below the $\Delta R = 0$ line, it follows that 749 there should be an optimal set of GSI accommodation parameters which yield $\Delta R =$ 750 0. Based on ΔR proximity to zero in all three bins in Fig. 8, model #8: CLL QS with 751 $\sigma_n = 0.5$ and $\sigma_t = 0.9$ appears to be a reasonable candidate for an appropriate GSI 752 model at these altitudes and He/O conditions. It is also possible that helium may scat-753 ter differently than atomic oxygen when interacting with a satellite surface, which could 754 mean that the optimal GSI assumptions in a helium-rich atmosphere might not be the 755 same as the optimal GSI assumptions in an atomic-oxygen rich atmosphere. Further anal-756 ysis which samples altitudes with a greater helium concentration is needed to confirm 757 this, though this idea is supported by the result that there is a C_D model that appears 758 to work best in only one of the He/O bins in Fig. 8. Model #7, CLL QS with $\sigma_n = 0.8$ 759 and $\sigma_t = 0.8$, appears to produce consistent inter-satellite density comparisons in the 760 highest average He/O conditions in the right panel of Fig. 8. However in the left panel 761 of Fig. 8 representing the lowest average He/O conditions, model #7 appears to be in-762 appropriate based on its corresponding ΔR value that is greater than zero. This may 763 suggest that the optimal C_D model assumptions may shift as helium increases in the at-764 mosphere. 765

To better visualize the angular distributions of the diffuse and QS models used in 766 this study, the CLL scattering kernels which correspond to the GSI parameters in mod-767 els #3 and #7-9 can be examined. It is worth noting that model #3, Sentman diffuse 768 with complete energy accommodation, can be replicated with a CLL model assuming 769 $\sigma_n = 1$ and $\sigma_t = 1$. Scattering kernels for the CLL models are given through proba-770 bility distribution functions for normal and tangential velocity components. Scattering 771 angular probability distributions for a set of accommodation parameters can be obtained 772 through Direct Simulation Monte Carlo (DSMC) simulations (Lord, 1991; Turansky, 2012). 773 It should be noted that the DSMC implementation of the CLL model depends on two 774 input parameters: α_n and σ_t (Lord, 1991). In contrast, the analytical CLL C_D expres-775 sions depend on a different set of accommodation parameters: σ_n and σ_t . In order to 776 represent the assumed GSI of CLL C_D models with DSMC scattering kernels, the au-777 thors make use of a fitted relationship between σ_n and α_n derived by Walker et al. (2014): 778

$$\sigma_n = 1 - \sqrt{1 - \alpha_n}$$

779

(6)

Eq. (6) was derived through iterative validation and Latin hypercube sampling with C_D values computed by NASA's DSMC Analysis Code. Eq. (6) is not the full fitted relationship between σ_n and α_n derived by Walker et al. (2014) (which includes two equations), however, Eq. (6) is all that is needed for the purposes of this study since for a given value of σ_n , there is only one possible solution for α_n between 0 and 1.

Fig. 9 includes DSMC CLL scattering angular probability distributions generated from 500,000 sample particles. The incident particles are monoenergetic beams of atomic oxygen with a fixed incident velocity of 7.5 km/s at a specified incident angle from the surface normal. Figs. 9a and 9b show the DSMC-generated reemission angular probability distributions assuming CLL GSI with $\sigma_n = \alpha_n = 1.0$ and $\sigma_t = 1.0$ at an incident ⁷⁹⁰ angle of 0° and an incident angle of 85° , respectively. At both low and high incident an-⁷⁹¹ gles, these model assumptions exhibit diffusive scattering with a cosine distribution. As ⁷⁹² a reminder, these diffusive scattering assumptions are adopted by all C_D models with ⁷⁹³ the description 'diffuse' in their label (models #1-6).

In contrast, the changes associated with decreasing normal and tangential momen-794 tum accommodation in the CLL model can be examined. Figs. 9c and 9d show the low 795 incident angle and high incident angle scattering kernels associated with model #7, re-796 spectively. Note that $\sigma_n = 0.8$ for model #7, which corresponds to $\alpha_n = 0.96$ based 797 on Eq. (6). It is clear that model #7 produces more quasi-specular scattering given that the associated scattering angular distributions are narrower than for diffuse scattering. 799 At high incidence angles, model #7 produces sub-specular scattering as shown in Fig. 800 9d, where the angle of reflection is less than the angle of incidence. It is also helpful to 801 consider the average energy accommodation assumed with model #7. For the CLL model, 802 tangential energy accommodation is related to tangential momentum accommodation 803 through the following expression (Cercignani & Lampis, 1971): 804

$$\alpha_t = \sigma_t (2 - \sigma_t) \tag{7}$$

805

Thus, for model #7, $\alpha_t = 0.96$. Since α_t and α_n are equivalent for this model, the average energy accommodation is nearly complete at 0.96. Model #7, a quasi-specular CLL model with incomplete momentum accommodation and near complete energy accommodation, appears to work well in helium-rich conditions based on ΔR proximity to zero in the right-most bin of Fig. 8.

The scattering kernels for the GSI assumptions in model #8 at low and high in-811 cident angles are shown in Figs. 9e and 9f, respectively. Model #8 produces nearly con-812 sistent inter-satellite density comparisons across the atmospheric He/O conditions sam-813 pled in this study. The scattering kernel differences from model #7 to model #8 include: 814 1) the low-incident angle reemission distribution narrows, and 2) the high-incident an-815 gle reemission distribution widens. This means that, for an elongated satellite like GRACE, 816 scattering on the front plate is closer to specular while scattering on the side plates is 817 closer to diffuse for model #8 in comparison to model #7. The normal and tangential 818 momentum accommodation parameters in model #8 correspond to energy accommoda-819 tion components of $\alpha_n = 0.75$ and $\alpha_t = 0.99$. For GRACE, this means that energy 820 accommodation on the front plate is 0.75, while energy accommodation on the side plates 821 is 0.99. To obtain an estimate of the average energy accommodation in model #8 for GRACE, 822 the authors considered the force contribution of the side panels relative to the front panel 823 in the same manner as was computed in Section 2.1 to obtain Fig. 2. Given that, on av-824 erage, $\frac{C_D A_{sides}}{C_D A_{sides}}$ for GRACE using model #8 is ~0.35, the average energy accommoda-825 tion for GRACE can then be estimated as $\alpha = 0.35\alpha_t + 0.65\alpha_n = 0.834$. Thus, aver-826 age energy accommodation is smaller for model #8 than for model #7. 827

Figs. 9g and 9h show low and high incident angle scattering kernels, respectively, 828 for model #9 which assumes significantly lower momentum accommodation. Normal and 829 tangential energy accommodation for model #9 are both equal to 0.75, meaning the av-830 erage energy accommodation for this model is 0.75. For model #9, scattering is closer 831 to specular with a narrower angular distribution for both low and high incident angle 832 interactions as compared to models #7 and #8. It makes sense, then, that model #9833 produces ΔR results that approach those produced by a fully specular C_D model (model 834 #11). The density discrepancy results then indicate that model #9 assumes energy and 835 momentum accommodation that is too low, meaning that it is not an appropriate model 836 at GRACE altitudes. 837

Some limitations should be discussed regarding the interpretation of ΔR across atmospheric He/O conditions. The examination of Fig. 8 reveals that it is harder to distinguish differences between C_D models at high He/O than at low He/O. Differences in ΔR between models #1-11 are smaller in the right-most bin of Fig. 8. This is one rea-



Figure 9. CLL angular probability distributions of reemitted particles assuming the GSI characteristics of model #3 at an incident angle of 0° (a), model #3 at an incident angle of 85° (b), model #7 at an incident angle of 0° (c), model #7 at an incident angle of 85° (d), model #8 at an incident angle of 0° (e), model #8 at an incident angle of 85° (f), model #9 at an incident angle of 0° (g), and model #9 at an incident angle of 85° (h)

son why it is important to consider how to quantify the uncertainty associated with each 842 of these measurements. Table 5 compares different error metrics for each of the C_D mod-843 els for the largest He/O bin alongside the measured ΔR range across the atmospheric 844 He/O bins. The column titled ' ΔR measurement uncertainty from TLEs' includes ΔR 845 uncertainty from TLE error estimates for the compact satellites. The TLE error esti-846 mates come from the TLE error model presented by Bernstein et al. (2021), which is a 847 function of the satellite observed density. These TLE error estimates are then linearly 848 propagated to the ΔR function. The column titled 'Robustness Range' alternatively presents 849 the range in measured ΔR associated with removing any single compact satellite from 850 the analysis. These uncertainty estimates provide a sense of how much the ΔR measure-851 ments could change given any of the compact satellites are removed from the analysis. 852 The column titled 'RMSE' indicates the root mean squared error associated with the den-853 sity ratio altitude fits based on the binned compact satellite observed densities for each 854 C_D model. These RMSE estimates are the vertical error bars shown for each ΔR mea-855 surement in Fig. 8. In comparison to each of these ΔR error metrics, the range in ΔR 856 measured for each C_D model across the atmospheric He/O bins is also shown in Table 857 5. These values are of similar magnitude compared to each of the error metrics, which 858 makes it difficult to comprehensively quantify the effect of increased helium in the at-859 mosphere. 860

Table 5. ΔR Error Metrics

C_D model	ΔR measurement uncertainty from TLEs	Robustness Range	RMSE	ΔR Range
Schamberg Diffuse, $\alpha = 1.0, \nu \to \infty, \phi_0 = 90^{\circ}$	0.037	0.021	0.060	0.063
Sentman Diffuse, SESAM α	0.031	0.022	0.045	0.061
Sentman Diffuse, $\alpha = 1.0$	0.037	0.024	0.060	0.044
Sentman Diffuse, $\alpha = 0.93$	0.036	0.027	0.055	0.041
Sentman Diffuse, $\alpha = 0.85$	0.035	0.027	0.053	0.040
Sentman Diffuse, $\alpha = 0.6$	0.033	0.027	0.049	0.038
CLL QS, $\sigma_n=0.8$, $\sigma_t=0.8$	0.037	0.025	0.060	0.038
CLL QS, $\sigma_n = 0.5$, $\sigma_t = 0.9$	0.034	0.023	0.051	0.031
CLL QS, $\sigma_n = 0.5$, $\sigma_t = 0.5$	0.038	0.027	0.061	0.030
CLL QS, $\sigma_n=0.9, \sigma_t=0.1$	0.050	0.038	0.099	0.032
Schamberg Specular, $\alpha = 0.0, \nu = 1, \phi_0 = 0^{\circ}$	0.039	0.029	0.064	0.020

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4.1 Drag Coefficient, Density, and Orbit Propagation Error Exercise

Standard, widely-used C_D models in research and operational applications typi-862 cally assume diffuse gas-surface interactions with complete or incomplete accommoda-863 tion. The density discrepancy results shown in Fig. 8 indicate that diffuse and DRIA 864 models are insufficient at GRACE altitudes near the oxygen-to-helium transition region. 865 A quasi-specular C_D model with incomplete energy and momentum accommodation, like 866 model #8, may be more suitable in this orbital regime given that it produces more con-867 sistent inter-satellite density comparisons. Considering model #8 to be a reasonable truth 868 model in this orbital regime, errors in C_D and consequently derived density caused by 869 making the standard assumption of diffuse scattering can be examined. 870

Fig. 10 compares C_D model #3 (Sentman diffuse with $\alpha = 1.0$) and C_D model #8 (CLL QS with $\sigma_n = 0.5$ and $\sigma_t = 0.9$) for GRACE over the course of ~5 hours during the 120th day of the year in 2003. During this time period, using the Sentman diffuse C_D model instead of the quasi-specular CLL C_D model would result in derived density errors of ~25%. In 2009 during solar minimum conditions, implementing the Sentman diffuse C_D model would yield similar derived density errors of ~27%.



Figure 10. GRACE C_D (1st panel), derived density (2nd panel), altitude (3rd panel), latitude (4th panel), and local time (5th panel) are shown over the course of ~5 hours on day 120 in the year 2003.

Table 6. Example Orbit Propagation Maximum In-Track Differences

Date	24hr In-Track Difference (m)	72hr In-Track Difference (m)
2009, DOY 150	31	64
2003, DOY 120	341	380

It is also helpful to consider potential orbit propagation effects associated with us-877 ing different C_D models. The authors performed a simple exercise to examine the C_D 878 effect on orbit propagation for GRACE in solar maximum and solar minimum conditions. 879 GRACE position and velocity were integrated and MSIS model density was used to prop-880 agate the orbit of GRACE over 24-hour and 72-hour time periods, using each of the 11 881 C_D models examined in this study. At the end of the 24-hour and 72-hour integration 882 time periods, the authors examined the in-track differences for GRACE. Results of this 883 exercise are summarized in Table 6. In solar minimum conditions (day 150, year 2009), 884 the maximum in-track difference across the 24-hour orbit propagation runs for each of 885 the 11 different C_D models was 31 meters. This value approximately doubled to 64 me-886 ters after 72-hours. In solar maximum conditions (day 120, year 2003), the in-track dif-887 ferences increased by an order of magnitude to $\sim 300-400$ meters for both the 24-hour 888 and 72-hour orbit integration runs. It should be noted that solar maximum in-track dif-889 ferences were more variable and could change significantly if the authors chose a differ-890 ent starting day. However, maximum in-track differences typically were on the order of 891 hundreds of meters in 2003. This exercise suggests that the divergence of C_D models at 892 GRACE altitudes could lead to orbital in-track errors of tens to hundreds of meters de-893 pending on the level of solar activity. 894

$_{895}$ 5 Discussion

The results presented in this paper provide evidence to support the implementation of a gas-surface interaction model which assumes quasi-specular scattering with incomplete energy and momentum accommodation at altitudes near ~500 km. This is the orbital regime where the oxygen-to-helium transition region emerges. This result marks a significant shift from the standard assumption of diffuse scattering in LEO orbital conditions, an assumption that is rooted in measurements from lower altitudes (<350 km).

Based on the ΔR analysis, the GSI models that implement quasi-specular gas-surface 902 interactions produce the most consistent densities between GRACE and the compact satel-903 lites. These results are novel for on-orbit GSI in the context of widely used DRIA mod-904 els. For the elongated GRACE satellite, diffuse models produce C_D values that are too 905 low, while specular or near-specular models produce C_D values that are too high. Mod-906 els #7 and #8 yield ΔR values closest to zero depending on the average atmospheric He/O, though the optimal model implementation could take a variety of different forms. 908 Both low-incident angle and high-incident angle gas-surface interactions drive the den-909 sity consistency results. As helium increases in the atmosphere, high-incident angle in-910 teractions become more important. 911

It is important to discuss potential limitations of this study that may impact the 912 density consistency results. Most of the compact satellites used in this study have mod-913 erately elliptical orbits in contrast to the nearly circular orbit of GRACE. It is possible 914 that there may be a suborbital time dependence or perigee velocity dependence in the 915 physical drag coefficient for the compact satellites. In this scenario, optimal GSI mod-916 els may need to incorporate differences in scattering based on incident velocity. There 917 may be different adsorption characteristics associated with compact object circular or-918 bits compared to elliptical orbits. Objects with elliptical orbits have greater perigee ve-919 locities than objects with circular orbits, which means that atmospheric molecules strike 920 the elliptical orbit satellites with greater kinetic energy. This may mean that energy ac-921 commodation is smaller, and atomic oxygen molecules are less likely to adsorb to the sur-922 face of an elliptical orbit object at perigee (K. Moe & Moe, 2005; Pilinski et al., 2013). 923 Examination of this potential effect on ΔR is beyond the scope of this paper. 924

An additional limitation of this study is the use of MSIS for atmospheric inputs 925 to the C_D models. Helium-rich atmospheric conditions and helium-related uncertainties 926 in the C_D models at upper thermospheric altitudes are of particular interest to the au-927 thors. These are the regimes that tend to be the most uncertain in empirical models like 928 MSIS. This highlights the circular nature of this problem, wherein the authors are at-929 tempting to validate appropriate physical C_D assumptions in order to estimate and mit-930 igate C_D -related biases at upper thermospheric altitudes, but the C_D models themselves 931 depend on modeled upper atmospheric constituents which may carry the same uncer-932 tainties and biases. If helium is wrong in MSIS, the C_D model will also assume the wrong 933 amount of helium. 934

Additionally, it is important to note that there are multiple theoretical solutions which may exist that could yield inter-satellite density consistency. The authors have discussed two potential CLL models that work well in varying atmospheric He/O conditions based on ΔR proximity to zero. To better constrain all possible solutions for optimal GSI parameters, the authors plan to leverage laboratory scattering data. One way to do this might be to fit CLL scattering parameters to laboratory scattering data for He and O using DSMC-generated scattering kernels, like the ones shown in Fig. 9.

The density discrepancy analysis could also be enriched by incorporating additional
datasets in this analysis. Including densities from the Swarm satellites would be a way
to sample higher altitude orbital conditions. Swarm B has an orbital altitude of ~530
km, which is slightly higher that the orbital altitude of GRACE, and Swarm densities

have been derived from GPS data instead of accelerometers (March, Doornbos, & Visser,
2019). With higher altitudes, Swarm B is likely to sample more helium-rich atmospheric
conditions than GRACE. Additionally, the Swarm satellites were launched in 2013, which
means they provide data covering the most recent solar cycle. The GRACE density data
analyzed in this study covers the years 2002 - 2010. Sampling years post-2013 also opens
the possibility of including more compact satellites with more recent launch dates.

This analysis might also be expanded by including comparisons between satellites 952 with different orientations for which the appropriate scattering scenarios might be more 953 clearly emphasized. This could include data where Swarm A and C, orbiting at ~ 470 954 km, have different attitudes. Planet Labs Flock 1-C CubeSats could also provide data 955 for these comparisons, as these satellites have low-drag and high-drag control modes (Foster 956 et al., 2015). These modes correspond to orientations parallel to the flow and perpen-957 dicular to the flow, respectively. Scattering comparisons between satellites with differ-958 ent attitudes would be helpful for evaluting C_D model assumptions related to incident 959 angle, given that such satellites fly through the same atmosphere but may have very dif-960 ferent levels of low-incidence and high-incidence scattering. 961

The inter-satellite density consistency results presented here provide a novel contribution to GSI modeling efforts in the upper thermosphere. This orbital regime is host to a growing number of resident space objects and is where physical C_D models tend to diverge due to uncertainty in model assumptions. With this work the authors have been able to constrain GSI assumptions at ~500 km to quasi-specular with incomplete energy and momentum accommodation, providing evidence that both diffuse and specular scattering assumptions are inappropriate at these altitudes.

⁹⁶⁹ Appendix A Bates Profile Fitting

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Altitude differences between GRACE and some of the compact satellites can lead 970 to differences in measured-to-modeled density ratios. The authors implemented a least-971 squares fitting scheme to fit the altitude structure of measured-to-modeled density ra-972 tios for the compact satellites to the ratio of two Bates profiles (Bates, 1959). The pur-973 pose of obtaining a physically-informed vertical profile of measured-to-modeled density 974 ratio for the compact satellites was to be able to attain a value for $\frac{\rho_{obs}}{\rho_{obs}}$ for a spherical 975 satellite at the precise location of GRACE. This value comes from the fitted density ra-976 tio profile, which is constructed using data from each of the compact satellites. 977

The observed-to-modeled density ratio fits are constructed based on the assump-978 tion that exospheric temperature differences are causing the offset between the observed 979 and modeled densities. As such, the ratio of two Bates profiles is fit to the observed-to-980 modeled density ratios for the compact satellites, where the two profiles have differ-981 ent exospheric temperatures. The ratio of two Bates profiles is chosen as the paramet-982 ric equation to fit to the density ratio data because the Bates temperature profile is con-983 sidered to be a physically consistent representation of temperature in the middle and up-984 per thermosphere, including above the turbopause in diffusive equilibrium. Also, the MSIS 985 model is constructed based on Bates profiles (Picone et al., 2002). In this region of dif-986 fusive equilibrium, where each of the satellites in this analysis reside, integration of the 987 species vertical hydrostatic balance equation for a Bates temperature profile gives the 988 number density of a constituent in diffusive equilibrium (Chamberlain & Hunten, 1987; 989 Picone et al., 2013): 990

$$n_i(\zeta) = n_{i,\ell} \left[\frac{T_\ell}{T(\zeta)} \right]^{1+\gamma_i + \alpha_i} \exp\left(-\gamma_i \sigma(\zeta - \zeta_\ell)\right)$$
(A1)

where ζ is the geopotential altitude, *i* is the index for each species constituent, and the subscript ℓ indicates the lower boundary of the diffusive equilibrium region. The geometric altitude of this lower boundary is often chosen to be 120 km. Picone et al. (2016) ⁹⁹⁵ point out that the region between ~ 100 km and ~ 200 km should be considered a tran-⁹⁹⁶ sition region such that ~ 200 km marks the lower boundary of the diffusive equilibrium ⁹⁹⁷ region. For this study, the authors choose 150 km as the geometric altitude of the lower ⁹⁹⁸ boundary of the diffusive equilibrium region. The inverse temperature scale height σ is ⁹⁹⁹ given as:

$$\sigma = \frac{T_{\ell}'}{T_{\infty} - T_{\ell}} \tag{A2}$$

where T'_{ℓ} is the vertical temperature gradient at the reference lower boundary altitude. The ratio γ_i of temperature and species scale heights is given as:

$$\gamma_i = \frac{1}{\sigma H_{i,\infty}} \tag{A3}$$

where *H* signifies scale height. The species thermal diffusion coefficient α_i is given as (Picone et al., 2002, 2013):

$$\alpha_i = \begin{cases} -0.38 & \text{for He and H} \\ 0.0 & \text{for O, O_2, and N_2} \end{cases}$$
(A4)

Then a parametric fitting equation can be written for the ratio of two Bates profiles; an 'observed' profile signified by the superscript A and a 'modeled' profile signified by the superscript B:

$$n_{1010} \qquad \frac{n_{i}^{A}}{n_{i}^{B}} = \frac{n_{i,\ell}^{A}}{n_{i,\ell}^{B}} T_{\ell}^{\left(\frac{1}{\sigma^{A}H_{i,\infty}^{A}} - \frac{1}{\sigma^{B}H_{i,\infty}^{B}}\right)} \frac{T_{\infty}^{B^{\left(1 + \frac{1}{\sigma^{B}H_{i,\infty}^{B}} + \alpha_{i}\right)}}{T_{\infty}^{A^{\left(1 + \frac{1}{\sigma^{A}H_{i,\infty}^{A}} + \alpha_{i}\right)}} \exp\left[\left(\frac{1}{H_{i,\infty}^{B}} - \frac{1}{H_{i,\infty}^{A}}\right)(\zeta - \zeta_{\ell})\right]$$
(A5)

In Eq. (A5), exospheric temperature T_{∞} has replaced $T(\zeta)$ due to the isothermal nature of this region, and Eq. (A3) has been plugged into Eq. (A1) for both A and B profiles. Additionally, the temperature at the lower boundary T_{ℓ} is assumed to be the same for both the A and B profiles.

Eq. (A5) applies to each individual thermospheric constituent species. To maintain a reasonable number of parameters to fit, the authors assume the thermosphere is composed primarily of O and He at the satellite altitudes of interest. Then the system of equations to fit to the data becomes:

$$\begin{cases} \frac{n_{O}^{A}}{n_{O}^{B}} = \frac{n_{O,\ell}^{A}}{n_{O,\ell}^{B}} T_{\ell}^{\left(\frac{1}{\sigma^{A}H_{O,\infty}^{A}} - \frac{1}{\sigma^{B}H_{O,\infty}^{B}}\right)} \frac{T_{\infty}^{B^{(1+\frac{1}{\sigma^{B}H_{O,\infty}^{B}} + \alpha_{O})}}{T_{\infty}^{A^{(1+\frac{1}{\sigma^{A}H_{O,\infty}^{A}} + \alpha_{O})}} \exp\left[\left(\frac{1}{H_{O,\infty}^{B}} - \frac{1}{H_{O,\infty}^{A}}\right)(\zeta - \zeta_{\ell})\right] \\ \frac{n_{He}^{A}}{n_{He}^{B}} = \frac{n_{He,\ell}^{A}}{n_{He,\ell}^{B}} T_{\ell}^{\left(\frac{1}{\sigma^{A}H_{He,\infty}^{A}} - \frac{1}{\sigma^{B}H_{He,\infty}^{B}}\right)} \frac{T_{\infty}^{B^{(1+\frac{1}{\sigma^{A}H_{He,\infty}^{A}} + \alpha_{O})}}{T_{\infty}^{(1+\frac{1}{\sigma^{B}H_{He,\infty}^{B}} + \alpha_{He})}} \exp\left[\left(\frac{1}{H_{He,\infty}^{B}} - \frac{1}{H_{He,\infty}^{A}}\right)(\zeta - \zeta_{\ell})\right] \\ \frac{\rho^{A}}{\rho^{B}} = \sum_{\Sigma m_{i}n_{i}^{A}} = \frac{\frac{n_{O}^{A}}{n_{O}^{B}}(n_{O}^{B}m_{O}) + \frac{n_{Ae}^{A}}{n_{He}^{B}}(n_{He}^{B}m_{He})}{n_{O}^{B}m_{O} + n_{He}^{B}m_{He}}}$$

$$(A6)$$

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Model output from MSIS is leveraged to reduce the number of unknown parameters in 1020 Eq. (A6). Since B represents 'modeled' values, and the temperature at the lower bound-1021 ary is assumed to be the same for the 'observed' and 'modeled' profiles, MSIS is used 1022 to obtain values for n_O^B , n_{He}^B , T_ℓ , T'_ℓ , and T^B_∞ . The MSIS runs used to obtain these val-1023 ues were selected to represent the observed atmospheric conditions based on the F10.7 1024 and Ap inputs used to initiate the model runs. Then, plugging in the MSIS-modeled pa-1025 rameters and known constants (including the masses of atomic oxygen and helium), this 1026 1027 leaves three unknown parameters to fit:

$$\begin{cases} C_1 = \frac{n_{O,\ell}^A}{n_{O,\ell}^B} \\ C_2 = \frac{n_{He,\ell}^A}{n_{He,\ell}^B} \\ C_3 = T_{\infty}^A \end{cases}$$
(A7)

The authors perform non-linear least squares fitting of Eq. (A6) to the $\frac{\rho_{obs}}{\rho_{mod}}$ data for the compact satellites.

Note that the parametric expression in Eq. (A6) is a function of geopotential altitude. In order to make the density ratio data a function of geopotential altitude, derived densities were converted into corresponding geopotential altitudes by interpolating to the MSIS profiles used to obtain some of the parameter estimates in Eq. (A6).

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The Two-Line Element data used to compute integrated densities from orbital changes 1036 for the selected set of compact satellites are available at Space-Track.org via https:// 1037 www.space-track.org/#gp with account registration required. The accelerometer-derived 1038 density datasets for the GRACE A and B satellites are provided by P. M. Mehta and 1039 E. K. Sutton at http://tinyurl.com/densitysets, courtesy of P. M. Mehta et al. (2017). 1040 MSIS simulations used in this work are obtained from the model downloaded from https:// 1041 gitlab.com/afedynitch/Python-NRLMSISE-00. Drag coefficient simulations in this study 1042 were completed using the Vehicle Environment Coupling and TrajectOry Response (VEC-1043 TOR) software, which can be downloaded at https://github.com/SWxTREC/vector-code. 1044

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1048 **References**

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1060

1061

1062

- 1049Bates, D. R. (1959). Some problems concerning the terrestrial atmosphere above1050about the 100 km level.Proceedings of the Royal Society of London. Se-1051ries A. Mathematical and Physical Sciences, 253(1275), 451–462.doi:105210.1098/rspa.1959.0207
- Beletskii, V. (1970). Interaction of the aerodynamic stream with a satellite according to an analysis of the motion of "proton-2" about its center of mass. Cosmic Research, 8, 189.
 - Bernstein, V., Pilinski, M., & Knipp, D. (2020). Evidence for drag coefficient modeling errors near and above the oxygen-to-helium transition. Journal of Spacecraft and Rockets, 57(6), 1246–1263. doi: 10.2514/1.A34740
 - Bernstein, V., Pilinski, M., & Sutton, E. K. (2021). Assessing thermospheric densities derived from orbital drag data. In *Proceedings of the 31st aas/aiaa space flight mechanics meeting.* AAS Publications Office.
 - Bouwer, S. (1992). Periodicities of solar irradiance and solar activity indices, ii. Solar Physics, 142(2), 365–389. doi: 10.1007/BF00151460
- Bowman, B., & Moe, K. (2005). Drag coefficient variability at 175-500km from
 the orbit decay analyses of spheres. In *Aas/aiaa astrodynamics specialist con- ference.*
- 1067Carter, V. L., Ching, B. K., & Elliott, D. D. (1969). Atmospheric density above1068158 kilometers inferred from magnetron and drag data from the satellite ov1-106915 (1968-059a). Journal of Geophysical Research, 74 (21), 5083-5091. doi:107010.1029/JA074i021p05083
- Cercignani, C., & Lampis, M. (1971). Kinetic models for gas-surface inter actions. Transport Theory and Statistical Physics, 1(2), 101–114. doi:
 10.1080/00411457108231440
- Cercignani, C., & Lampis, M. (1997). New scattering kernel for gas-surface interac tion. AIAA journal, 35(6), 1000–1011. doi: 10.2514/2.209
- Chamberlain, T. P., & Hunten, D. M. (1987). Theory of planetary atmospheres: an introduction to their physics and chemistry. Academic Press, Inc., New York.

- 1078Ching, B., Hickman, D., & Straus, J.(1977). Effects of atmospheric winds and1079aerodynamic lift on the inclination of the orbit of the s3-1 satellite.Jour-1080nal of Geophysical Research, 82(10), 1474–1480.Retrieved from https://1081agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/JA082i010p014741082doi: 10.1029/JA082i010p01474
- ¹⁰⁸³ Cook, G. (1965). Satellite drag coefficients. *Planetary and Space Science*, 13(10), ¹⁰⁸⁴ 929–946. doi: 10.1016/0032-0633(65)90150-9
- 1085Crisp, N., Roberts, P., Livadiotti, S., Rojas, A. M., Oiko, V., Edmondson, S., ...1086others (2021). In-orbit aerodynamic coefficient measurements using soar1087(satellite for orbital aerodynamics research). Acta Astronautica, 180, 85–99.1088Retrieved from https://www.sciencedirect.com/science/article/pii/1089S00945765203075911081doi: https://doi.org/10.1016/j.actaastro.2020.12.024
- 1090Doornbos, E. (2012). Thermospheric density and wind determination from satellite1091dynamics. Springer Science & Business Media, Heidelberg, Germany. doi: 101092.1007/978-3-642-25129-0
 - Emmert, J. (2009). A long-term data set of globally averaged thermospheric total mass density. Journal of Geophysical Research: Space Physics, 114(A6). doi: 10.1029/2009JA014102

1093

1094

1095

1118

1119

- Foster, C., Hallam, H., & Mason, J. (2015). Orbit determination and differential drag control of planet labs cubesat constellations. In *Proceedings of the aiaa astrodynamics specialist conference.*
- Gaposchkin, E. M. (1994). Calculation of satellite drag coefficients (Tech. Rep.
 No. 998). Massachusetts Institute Of Technology Lincoln Laboratory.
- Goodman, F. O., & Wachman, H. Y. (1967). Formula for thermal accommodation
 coefficients. The Journal of Chemical Physics, 46(6), 2376-2386. doi: 10.1063/
 1.1841046
- Gregory, J. C., & Peters, P. N. (June 1986). A measurement of the angular distribution of 5 ev atomic oxygen scattered off a solid surface in earth orbit. In
 Proceedings of the 15th international symposium on rarefied gas dynamics (pp. 644–654). Stuttgart: B. G. Teubner, Germany.
- Harrison, I., & Swinerd, G. (1996). A free molecule aerodynamic investigation using
 multiple satellite analysis. *Planetary and space science*, 44 (2), 171–180. doi:
 10.1016/0032-0633(95)00077-1
- Hedin, A. E., Hinton, B. B., & Schmitt, G. A. (1973). Role of gas-surface
 interactions in the reduction of ogo 6 neutral particle mass spectrometer data. *Journal of Geophysical Research*, 78(22), 4651–4668. doi:
 10.1029/JA078i022p04651
- 1115Imbro, D. R., Moe, M. M., & Moe, K. (1975). On fundamental problems in the
deduction of atmospheric densities from satellite drag. Journal of Geophysical
Research, 80(22), 3077–3086. doi: 10.1029/JA080i022p03077
 - Liu, S.-M., Sharma, F. K., & Knuth, E. L. (1979). Satellite drag coefficients calculated from measured distributions of reflected helium atoms. AIAA Journal, 17(12), 1314–1319. doi: 10.2514/3.7629
- Lord, R. (1991). Some extensions to the cercignani-lampis gas-surface scattering kernel. *Physics of Fluids A: Fluid Dynamics*, 3(4), 706–710. doi: 10.1063/1 .858076
- March, G., Doornbos, E. N., & Visser, P. (2019). High-fidelity geometry models for improving the consistency of champ, grace, goce and swarm thermospheric density data sets. Advances in Space Research, 63(1), 213–238. doi:
 10.1016/j.asr.2018.07.009
- March, G., van den IJssel, J., Siemes, C., Visser, P. N., Doornbos, E. N., & Pilinski,
 M. (2021). Gas-surface interactions modelling influence on satellite aerody namics and thermosphere mass density. *Journal of Space Weather and Space Climate*, 11, 54. doi: 10.1051/swsc/2021035
- ¹¹³² March, G., Visser, T., Visser, P., & Doornbos, E. N. (2019). Champ and goce

1133	thermospheric wind characterization with improved gas-surface interac-
1134	tions modelling. Advances in Space Research, $64(6)$, $1225-1242$. doi:
1135	10.1016/j.asr.2019.06.023
1136	Mehta, N. A., Murray, V. J., Xu, C., Levin, D. A., & Minton, T. K. (2018). Non-
1137	reactive scattering of n2 from layered graphene using molecular beam ex-
1138	periments and molecular dynamics. The Journal of Physical Chemistry C ,
1139	122(18), 9859-9874. doi: $10.1021/acs.jpcc.7b11721$
1140	Mehta, P. M., Walker, A., Lawrence, E., Linares, R., Higdon, D., & Koller, J.
1141	(2014). Modeling satellite drag coefficients with response surfaces. Advances in
1142	Space Research, $54(8)$, 1590–1607. doi: 10.1016 /j.asr.2014.06.033
1143	Mehta, P. M., Walker, A. C., Sutton, E. K., & Godinez, H. C. (2017). New den-
1144	sity estimates derived using accelerometers on board the champ and grace
1145	satellites. Space Weather, 15(4), 558–576. doi: 10.1002/2016SW001562
1146	Minton, T. K., Tagawa, M., & Nathanson, G. M. (2004). Energy accommodation
1147	in hyperthermal gas-surface collisions: Aerobraking in planetary atmospheres.
1148	Journal of Spacecraft and Rockets, $41(3)$, 389–396. doi: $10.2514/1.10724$
1149	Moe, K. (1966). Absolute atmospheric densities determined from the spin and or-
1150	bital decays of explorer vi. Planetary and Space Science, $14(11)$, 1065–1075.
1151	doi: 10.1016/0032-0633(66)90022-5
1152	Moe, K., & Bowman, B. (2005). The effects of surface composition and treatment on
1153	drag coefficients of spherical satellites. In Proceedings of the aas/aiaa astrody-
1154	namics specialists conference. American Astronautical Society Publications Of-
1155	fice, San Diego, CA.
1156	Moe, K., & Moe, M. M. (1992). The deduction of in-track winds from satellite mea-
1157	surements of density and composition. Geophysical Research Letters, 19(13),
1158	1343–1346. doi: 10.1029/92GL01114
1159	Moe, K., & Moe, M. M. (2005). Gas-surface interactions and satellite drag coeffi-
1160	cients. Planetary and Space Science, $53(8)$, $793-801$. doi: $10.1016/J.pss.2005.03$
1161	
1162	Moe, K., Moe, M. M., & Yelaca, N. W. (1972). Effect of surface neterogeneity on
1163 1164	search, 77(22), 4242–4247. doi: 10.1029/JA077i022p04242
1165	Moe, M. M., Wallace, S. D., & Moe, K. (1993). Refinements in determining satellite
1166 1167	drag coefficients: Method for resolving density discrepancies. Journal of Guid- ance, Control, and Dynamics, 16(3), 441–445. doi: 10.2514/3.21029
1168	Moe, M. M., Wallace, S. D., & Moe, K. (1995). Recommended drag coefficients for
1169	aeronomic satellites. The Upper Mesosphere and Lower Thermosphere: A Re-
1170	view of Experiment and Theory, Geophysical Monograph Series, 87, 349–356.
1171	doi: 10.1029/GM087p0349
1172	Murray, V. J., Pilinski, M. D., Smoll Jr, E. J., Qian, M., Minton, T. K., Madzunkov,
1173	S. M., & Darrach, M. R. (2017). Gas-surface scattering dynamics ap-
1174	plied to concentration of gases for mass spectrometry in tenuous atmo-
1175	spheres. The Journal of Physical Chemistry C , $121(14)$, $7903-7922$. doi:
1176	10.1021/acs.jpcc.7b00456
1177	Nayar, S. P. (2006). Periodicities in solar activity and their signature in the terres-
1178	trial environment. In N. Gopalswamy & A. Bhattacharyya (Eds.), Proceedings
1179	of the ilws workshop "the solar influence on the heliosphere and earth's envi-
1180	ronment: Recent progress and prospects", goa (pp. 170–178).
1181	Pardini, C., Anselmo, L., Moe, K., & Moe, M. M. (2009). Drag and energy accom-
1182	modation coefficients during sunspot maximum. Advances in Space Research,
1183	45(5), 638–650. doi: 10.1016/j.asr.2009.08.034
1184	Picone, J., Emmert, J., & Drob, D. (2016). Consistent static models of local thermo-
1185	spheric composition profiles. arXiv preprint arXiv:1607.03370 [physics.space-
1186	ph/.
1187	Picone, J., Emmert, J., & Lean, J. (2005). Thermospheric densities derived from

1188	spacecraft orbits: Accurate processing of two-line element sets. Journal of Geo-
1189	physical Research: Space Physics, $110(A3)$. doi: $10.1029/2004JA010585$
1190	Picone, J., Hedin, A., Drob, D. P., & Aikin, A. (2002). Nrlmsise-00 empiri-
1191	cal model of the atmosphere: Statistical comparisons and scientific issues.
1192	Journal of Geophysical Research: Space Physics, 107(A12), 1-16. doi:
1193	10.1029/2002JA009430
1194	Picone, J., Meier, R., & Emmert, J. (2013). Theoretical tools for studies of low-
1195	frequency thermospheric variability. Journal of Geophysical Research: Space
1196	<i>Physics</i> , 118(9), 5853–5873. doi: 10.1002/jgra.50472
1197	Pilinski, M. D., Argrow, B. M., & Palo, S. E. (2011). Drag coefficients of satel-
1198	lites with concave geometries: Comparing models and observations. Journal of
1199	Spacecraft and Rockets, 48(2), 312–325. doi: 10.2514/1.50915
1200	Pilinski, M. D., Argrow, B. M., Palo, S. E., & Bowman, B. R. (2013). Semi-
1201	empirical satellite accommodation model for spherical and randomly tum-
1202	bling objects. Journal of Spacecraft and Rockets, $50(3)$, $556-571$. doi:
1203	10.2514/1.A32348
1204	Poovathingal, S. J., Xu, C., Murray, V., Minton, T. K., & Schwartzentruber, T. E.
1205	(2019). Gas-surface model in dsmc for molecules passing through a funnel-type
1206	gas concentrator. In Aiaa scitech 2019 forum. doi: $10.2514/6.2019-1281$
1207	Qian, L., Burns, A. G., Emery, B. A., Foster, B., Lu, G., Maute, A., Wang,
1208	W. (2014). The near tie-gem: A community model of the coupled thermo-
1209	sphere/ionosphere system. Modeling the Ionosphere-Thermosphere System,
1210	201, 73–83. doi: 10.1002/9781118704417.ch7
1211	Schaaf, S., & Chambre, P. (1958). Flow of rarefied gases, high speed aerodynam-
1212	ics and jet propulsion. In H. W. Emmons (Ed.), Fundamentals of gas dynamics
1213	(Vol. 3, pp. 687–741). Princeton University Press, Princeton, NJ.
1214	Schamberg, R. (1959). A new analytic representation of surface interaction for
1215	hyperthermal free molecular flow with applications to neutral-particle drag es-
1216	timates of satellites (Tech. Rep. No. RM-2313). Rand Corporation, Santa
1217	Monica, CA.
1218	Seidl, M., & Steinheil, E. (1974). Measurements of momentum accommodation
1219	coefficients on surfaces characterized by auger spectroscopy, sims and leed. In
1220	M. Becker & M. Fieldg (Eds.), Proceedings of the ninth international sympo-
1221	sium on rarefiea gas aynamics. DF VLR-Press, Porz-whan.
1222	Sentinan, L. H. (1901). Free molecule flow theory and its application to the determi-
1223	rition of devolgnamic forces (Tech. Rep. No. LMSC-448514). Lockneed Mis-
1224	shes and Space Co. Inc., Sunnyvale, CA.
1225	(Tash Dan No. TD 2002(2207) 1) The Assessment Comparation Vehicle Star
1226	(Tech. Rep. No. 1R-2005(5597)-1). The Aerospace Corporation, vehicle Sys-
1227	Sutton F K Norom P S & Forbes I M (2007) Density and winds in the then
1228	sutton, E. K., Nereni, R. S., & Fordes, J. M. (2007). Density and winds in the ther-
1229	//(6) 1210 1210 doi: 10.2514/1.28641
1230	44(0), 1210–1219. doi: 10.2014/1.20041 Turanely C P (2012) High fidelity dynamic modeling of engagement in the
1231	continuum-rarefied transition regime (Unpublished doctoral dissortation)
1232	University of Colorado at Boulder
1233	Virgili I & Roberts P C (2013) $\delta dsat = ab50$ cubesat mission to study rarefied-
1234	gas drag modelling Acta Astronautica 89 130–138 doi: 10.1016/j.acta.astro
1235	2013 04 006
1227	Wachman H V (1962) The thermal accommodation coefficient: Λ critical survey
1237	American Rocket Society Journal 39(1) 2-12 doi: 10.2514/8.5030
1238	Walker A Mehta P & Koller I (2014) Drag coefficient model using the
1240	cercignani-lampis-lord gas-surface interaction model <u>Journal of Spacecraft</u>
1240	and Rockets 51(5) 1544–1563 doi: 10.2514/1.A32677
****	www.ives/veres, of (9), forf 1000, doi: 10.2011/1.102011