Seasonal and solar cycle dependence of energy transfer rates in the Auroral E-region

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Abstract

We report one of the first comprehensive ground-based investigations of energy transfer rates in the E-region ionosphere compared relative to geomagnetic activity, seasonal effects, and solar activity level using nearly continuously sampled data collected with the Poker Flat Incoherent Scatter Radar (PFISR) between 2010-2019. We quantified the integrated electromagnetic (EM) energy transfer rate and the integrated Joule heating rate in the E-region between 90-130 km, which includes the contributions from the neutral winds. We find that (1) the median Joule heating rate and electromagnetic (EM) energy transfer rate in the evening sector is larger in the winter versus the summer and have similar magnitudes in the spring and fall for the same solar activity and geomagnetic conditions. (2) The seasonal dependence of the energy transfer rates is closely associated with the seasonal variations of the electric fields. Our analysis shows that the larger EM energy transfer and Joule heating rates in disturbed conditions in the winter versus the summer are associated with the combined effects of both the electric field and Pedersen conductance with the electric field playing a dominant role. Given that the Pedersen conductance in the evening sector is closely related to the particle precipitation and field-aligned currents in the auroral region, this study provides complementary ionospheric evidence of the winter-summer asymmetry of the intensity and density of field-aligned currents (e.g. Ohtani et al., 2009). (3) The geomagnetic activity level has the most significant impacts on the magnitude of the energy transfer rates, followed by seasonal variations, and last the solar activity level.

Seasonal and solar cycle dependence of energy transfer rates in the Auroral E-region

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Key	Points:
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7	• Geomagnetic activity, seasonal, and solar activity variability of the integrated
8	energy transfer rates are quantified in the E-region between 90-130 km span-
9	ning 2010-2019 using PFISR observations for the first time.
10	• The integrated Joule heating and EM energy transfer rates in the evening sec-
11	tor are larger in winter versus summer and have similar magnitudes in spring
12	and fall equinoxes.
13	• The larger energy transfer rates in winter relative to summer in disturbed

The larger energy transfer rates in winter relative to summer in disturbed
 conditions are associated with a combination of electric field and Pedersen
 conductance.

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16 Abstract

We report one of the first comprehensive ground-based investigations of energy 17 transfer rates in the E-region ionosphere compared relative to geomagnetic activity, 18 seasonal effects, and solar activity level using nearly continuously sampled data col-19 lected with the Poker Flat Incoherent Scatter Radar (PFISR) between 2010–2019. 20 We quantified the integrated electromagnetic (EM) energy transfer rate and the in-21 tegrated Joule heating rate in the E-region between 90-130 km, which includes the 22 contribution from the neutral winds. We find that (1) the median Joule heating rate 23 24 and electromagnetic (EM) energy transfer rate in the evening sector are larger in the winter versus the summer and have similar magnitudes in the spring and fall for 25 the same solar activity and geomagnetic conditions. (2) The seasonal dependence 26 of the energy transfer rates are closely associated with the seasonal variations of 27 the electric fields. Our analysis shows that the larger EM energy transfer and Joule 28 heating rates in disturbed conditions in the winter versus the summer are associated 29 with the combined effects of both the electric field and Pedersen conductance with 30 the electric field playing a dominant role. Given that the Pedersen conductance in 31 the evening sector is closely related to the particle precipitation and field aligned 32 currents in the auroral region, this study provides complementary ionospheric evi-33 dence of the winter-summer asymmetry of the intensity and density of field-aligned 34 currents (e.g. Ohtani et al., 2009). (3) The geomagnetic activity level has the most 35 significant impact on the magnitude of the energy transfer rates, followed by sea-36 sonal variations, and last the solar activity level. 37

³⁸ Plain Language Summary

We reported a comprehensive statistical study of energy transfer rates (EM 39 transfer rate and Joule heating rate) in the auroral E region in aspects of season, 40 geomagnetic activity, and solar activity. This study is done by collecting ground-41 based 10 years of measurements from Poker Flat Incoherent Scatter Radar. We find 42 a consistent seasonal variation of energy transfer rates for different geomagnetic and 43 solar activity levels. That is the EM energy transfer and Joule heating experience 44 larger enhancements in the evening sector in the winter versus the summer and show 45 similar magnitudes in the spring and fall equinoxes. We further find that the sea-46 sonal variation of the enhancement is mainly associated with the seasonal variation 47 of the electric field while the seasonal variation of the E region conductance could 48 also play a smaller role during disturbed conditions. We also compared the relative 49 importance of season, geomagnetic, and solar activity levels in impacting the energy 50 transfer rates and find that the geomagnetic activity has the largest impact, followed 51 by season and solar activity. 52

53 1 Introduction

Joule heating is an important mechanism by which electromagnetic (EM) 54 energy input from the magnetosphere is dissipated in the high latitude ionosphere-55 thermosphere (Thayer & Semeter, 2004). Therefore, the energy deposited through 56 Joule heating is an important source in the ionosphere-thermosphere (IT) system, 57 and can lead to temperature and density changes of the neutrals (Thayer & Seme-58 ter, 2004; Barth et al., 2009; Barth, 2010). Joule heating can further induce gravity 59 waves which impact the IT system on a large scale (Brekke, 1979; Sofko & Huang, 60 2000; Yuan et al., 2005). Many factors that modulate the transfer of EM energy into 61 the IT system have been investigated, although primarily through modeling inves-62 tigations. These factors include solar activity (e.g. Zhang et al., 2005; Bjoland et 63 al., 2015), Interplanetary Magnetic Field (IMF) clock angles and magnitudes (e.g. 64 McHarg et al., 2005; Zhang et al., 2005; Bjoland et al., 2015; Cai et al., 2016), the 65

solar wind (e.g. Zhang et al., 2005; Cai et al., 2014), geomagnetic activity (Fujii
et al., 1999; Aikio et al., 2012; Weimer, 2005; Zhang et al., 2005, etc.) and season
(dipole tilt angle) (e.g. Foster et al., 1983; Weimer, 2005; Zhang et al., 2005). IMF
has also showed effects on the E-region neutral wind through ion drag (Richmond et
al., 2003), which could potentially affect the distribution of EM energy.

Observational investigations of seasonal and solar cycle dependence of the 71 EM energy transfer and Joule heating rates are sparse, especially in the E-region. 72 Foster et al. (1983) presents one of the only observational studies that examined 73 the seasonal dependence of Joule heating using satellite observations during solar 74 minimum conditions. The solar cycle dependence of the F-region Joule heating was 75 investigated by Bjoland et al. (2015) using ion convection measurements from the 76 Super Dual Auroral Radar Network (SuperDARN) of high-frequency (HF) over-the-77 horizon radars, and neutral wind measurements from the CHAMP satellite during 78 2001-2009, but no similar study was conducted in the E-region. The purpose of this 79 paper is to quantify the seasonal and solar cycle dependence of the integrated EM 80 energy transfer rates and Joule heating by utilizing nearly continuously sampled 81 observations obtained with the Poker Flat Incoherent Scatter Radar (PFISR), which 82 also includes the contribution to the EM transfer rates from the neutral winds. This 83 investigation provides a unique opportunity to quantify the seasonal and solar cy-84 cle effects on the EM transfer rates with relatively high temporal resolution, thus 85 enabling providing a new perspective of Joule heating, although spatially limited to 86 one geographic location. 87

The previous investigations have demonstrated that Joule heating exhibits 88 89 different characteristics with respect to season during different conditions. Foster et al. (1983) used Atmosphere Explorer C (AE-C) satellite measurements and found 90 that the Joule heating input is 50% larger in summer than in the winter during solar 91 minimum, on a global scale. This result was attributed to the larger conductance 92 caused by solar illumination in the summer hemisphere. Weimer (2005) used nu-93 merical simulations to show that the total Joule heating in the northern hemisphere 94 doubles as the Earth's dipole tilt angle increases from -30° to $+10^{\circ}$ and decreases 95 when the angle is above $+10^{\circ}$ during fixed southward IMF conditions and fixed solar 96 wind velocity. These results indicate larger Joule heating in the summer and smaller 97 Joule heating during the winter. However, another simulation study by Zhang et al. 98 (2005) showed that for a fixed IMF angle the Joule heating pattern and intensity 99 do not change significantly as the dipole tilt angle increases. These contradictory 100 simulation results can be reconciled by utilizing a large dataset of observations cov-101 ering different solar cycles and seasons to investigate the seasonal and solar cycle 102 dependence of Joule heating. 103

Moreover, there are not observational investigations that compared the relative 104 importance of solar activity levels versus geomagnetic activity level. Geomagnetic 105 and solar activity have been shown to be associated with an increase in Joule heat-106 ing (Fujii et al., 1999; Aikio et al., 2012; Zhang et al., 2005; Bjoland et al., 2015), 107 but the relative importance of these two mechanisms on the variability of Joule 108 heating is not clear. Zhang et al. (2005) showed that the intensification of solar 109 110 EUV radiation, indicated by F10.7, is associated with a significant increase of Joule heating for a fixed geomagnetic activity level (AL and Kp), while the geomagnetic 111 activity level can change the Joule heating rate mainly in the postmidnight sector 112 when the solar activity level is fixed. Their relative importance is not compared 113 specifically. 114

We hypothesize that there should be seasonal variation in Joule heating due to seasonal variations in both the Pedersen conductivity and electric field. For example, using sparse observations from the Sondrestrom ISR spanning 5 years near solar minimum, de la Beaujardiere et al. (1991) showed that the polar cap potential

drop is largest in fall, followed by winter, then spring and was smallest in summer. 119 However, the seasonal dependence of the conductivity is not well resolved. The con-120 ductivity in the ionosphere in different seasons is closely related to the solar EUV ra-121 diation, therefore the conductivity is larger in summer and smaller in winter. When the particle precipitation at night in winter is considered, especially in the E-region 123 under disturbed condition, the conclusion is not intuitive. Ohtani et al. (2009) 124 showed that the absence of the solar EUV radiation can be often overcompensated 125 by more intense and energetic electron precipitation in the dark hemisphere, which 126 leads to the corresponding larger Pedersen conductivity. Using Constellation Ob-127 serving System for Meteorology, Ionosphere, and Climate (COSMIC) satellites ob-128 servations at high latitudes during 2008-2011, Sheng et al. (2014) showed that the 129 ratio of the E-region Pedersen conductance to F-region Pedersen conductance in the 130 nighttime reaches minimum at local summer and maximum at local winter, which 131 is interpreted as the seasonal variations of the solar irridance and auroral activity. 132 In addition, the neutral wind is also believed to play a role to reduce or enhance 133 Joule heating depending on the MLT sector and range (e.g. Thayer, 1998a; Thayer 134 & Semeter, 2004). However, the extent to which the neutral wind influence the Joule 135 heating varies among studies (Thayer, 1998a; Lu et al., 1995). 136

In this investigation, we use nearly continuous PFISR observations of electron 137 density and neutral winds in the E-region, and the F-region electric field to quantify 138 the seasonal and solar cycle variations of energy transfer rates in the E-region. We 139 use a dataset of energy transfer rates between 2010-2019 derived from nearly contin-140 uous observations of the local E- and F-region ionosphere obtained with PFISR. A 141 recent study by Zhan et al. (2021) using the same dataset analyzed the energy trans-142 fer rates in Fall 2015. This investigation expands upon the previous results by Zhan 143 et al. (2021) by understanding seasonal, solar cycle, and geomagnetic variability of 144 the Joule heating rate. Therefore, the dataset of energy transfer rates is divided into 145 subgroups according to seasons, solar, and geomagnetic activity levels. We present 146 the climatology of energy transfer rates for different seasons under different solar 147 and geomagnetic activity levels. To first order, we find that the seasonal variability 148 of the electric field has a significant contribute to the variability of the EM transfer 149 rate. 150

In the next section, we present a brief introduction of the data and the methods used to estimate energy transfer rates. The results section will show the seasonal variation of the energy transfer rates during low, medium, and high solar flux conditions for quiet, moderate and active geomagnetic activity. The discussion section will examine the relation between the seasonal variation of the Joule heating enhancement and the seasonal variation of the electric field for different solar and geomagnetic activity levels; the main findings are summarized in the last section.

158

2 Data and Methodology

We use observations collected with PFISR (65.13° N, 147.47°W, MLAT: 65.4° 159 N) from 2010–2019, including the E-region altitude resolved electron densities and 160 line-of-sight (LOS) velocities in the International Polar Year (IPY) operational 161 $mode(\sim 1\% duty cycle)$ and other high duty cycle operational modes (>1% duty 162 cycle). For a detailed description of PFISR, please refer to Heinselman and Nicolls 163 (2008) and the IPY mode are described in Sojka et al. (2009), and Makarevich et al. 164 (2013). The majority of the data we use for this investigation were collected from 165 the IPY mode, which covers 1° of geomagnetic latitude in the E-region (please see 166 Figure 1 in Makarevich et al. (2013). The LOS velocities of the ions drifts in the 167 E-region are used to derive the neutral winds and F-region ion drifts are used to es-168 timate the electric fields using a Bayesian inversion method described by Heinselman 169 and Nicolls (2008). This technique assumes that the ion velocity and wind fields are 170

spatially uniform (i.e. Thayer, 1998a; Heinselman & Nicolls, 2008). The Pedersen
and Hall conductivities were calculated using the E-region electron densities combined with the neutral densities from NRLMSISE-00 (Picone et al., 2002), which are
used to calculate the ion-neutral collision frequency.

The integrated energy transfer rates (Joule heating rate, passive energy depo-175 sition rate, mechanical energy transfer rate, and EM transfer rate) between 90-130 176 km in the E region are estimated using the equations from previous investigations 177 (Aikio et al., 2012) and described further in Zhan et al. (2021)(accepted by JGR-178 Space Physics). We summarize these equations and the corresponding derivations in 179 Table 1. For more details on how to calculate the energy transfer rates using radar 180 measurements, please refer to Zhan et al. (2021). Significant additional conductivity 181 and Joule heating (the integrated Pedersen conductance between 130-150 km could 182 be $27 \sim 38\%$ of the integrated Pedersen conductance between 90-150km depending on 183 the geomagnetic activity level) does exist above 130 km (Zhan et al., 2021); however, 184 this study will only consider the E-region Joule heating below 130 since this is the 185 altitude region where PFISR can reliably estimate the neutral winds. For conve-186 nience, the different energy transfer rates mentioned in this paper are all integrated 187 with respect to altitude between 90-130 km. Measurements with SNR \geq -20 dB are 188 used in this investigation to perform the statistical study with detectable electron 189 densities. We use a 1 hour running median filter with a 15 minutes time step and 190 obtain the median energy transfer rates. Each 15-minute time interval must satisfy 191 a threshold for the number of measurements (18). The raw data has a resolution of 192 around 10 minutes. A timestep of 15 minutes will ensure small temporal variations 193 are not smeared out. More details about the PFISR measurements and the proce-194 dure to derive the energy transfer rates as well as the limitations of this dataset can 195 be found in Zhan et al. (2021). 196

We use the regional local time version of the Supermag Auroral Electrojet in-197 dex (SMEr) (Newell & Gjerloev, 2014; Gjerloev, 2012) for the geomagnetic activity 198 index and divide the data set into three categories corresponding to quiet (SMEr <199 100 nT), moderate ($100 \leq \text{SMEr} < 200$) and active ($\text{SMEr} \geq 200 \text{ nT}$) geomagnetic 200 activity levels. This local index can better characterize the variations of Joule heat-201 ing from local observations versus global indices (Thayer, 2000). We use F10.7A (81) 202 day average of F10.7 index) as the solar activity index and divide the data set into 203 three categories corresponding to low (F10.7A < 95 SFU, SFU: solar flux unit, 1 204 $SFU = 10^{-22}W \cdot m^{-2} \cdot Hz^{-1}$), medium (95 $SFU \le F10.7A < 130 SFU$) and high 205 (F10.7A > 130 SFU) solar activity levels. 206

207

208 **3 Results**

The seasonal variations of the integrated energy transfer rates between 90-209 130 km during quiet, moderate and active conditions are presented in Figures 1, 210 2, and 3, respectively. In each figure, the results for different seasons are presented 211 as columns (from left to right: spring, summer, fall, and winter) and during differ-212 ent solar activity conditions as rows (from top to bottom: All, low, medium, and 213 high solar activity). For completeness, the overall seasonal variation of the energy 214 transfer rates during 2010-2019 are presented in the first row. In each subplot, the 215 green, blue, black, and red curves correspond to the median Joule heating rate Q_i , 216 passive energy deposition rate Q_j^E , mechanical energy transfer rate Q_m , and EM 217 energy transfer rate Q_{EM} . The associated shaded light green, light blue, gray and 218 pink areas correspond to the lower and upper quartiles for Q_j , Q_i^E , Q_m , and Q_{EM} , 219 respectively. The total number of measurements (count of the 15-min intervals) used 220

Parameter	Symbol	Derivation
Electron density	n_e	ion-line spectra in the E region
Ion velocity	$\mathbf{V_i}$	ion-line spectra in the E region
Electron velocity	$\mathbf{V_e} = \mathbf{V_{i_{Fregion}}}$	ion-line spectra in the F region
Neutral wind	U_n	line-of-sight ion velocity ^{a} in the E region
Electric field	${f E}$	$-\mathbf{V_{i_{Fregion}}} imes \mathbf{B}$
Pedersen conductance	${\Sigma^E_P} {f J}$	$\int_{90km}^{130km} \frac{en_e}{B} [\Omega_i v_{in} / (\Omega_i^2 + v_{in}^2)]^b dz \text{ (z, altitude)}$
Currents	J	$en_e(\mathbf{V_i} - \mathbf{V_e})$
Passive energy deposition rate	Q_j^E	$\Sigma^E_P {f E^2}$
Joule heating rate	$ec{Q_j}$	$\Sigma_P^E (\mathbf{E} + \mathbf{U_n} imes \mathbf{B})^{2}$
Mechanical energy transfer rate	Q_m	$\mathbf{U_n} \cdot (\mathbf{J} \times \mathbf{B})$
Electromagnetic energy transfer rate	Q_{EM}	$\mathbf{J}\cdot\mathbf{E}$

 Table 1.
 Electrodynamics Parameters, Symbols and the Derivation through PFISR Measurements.

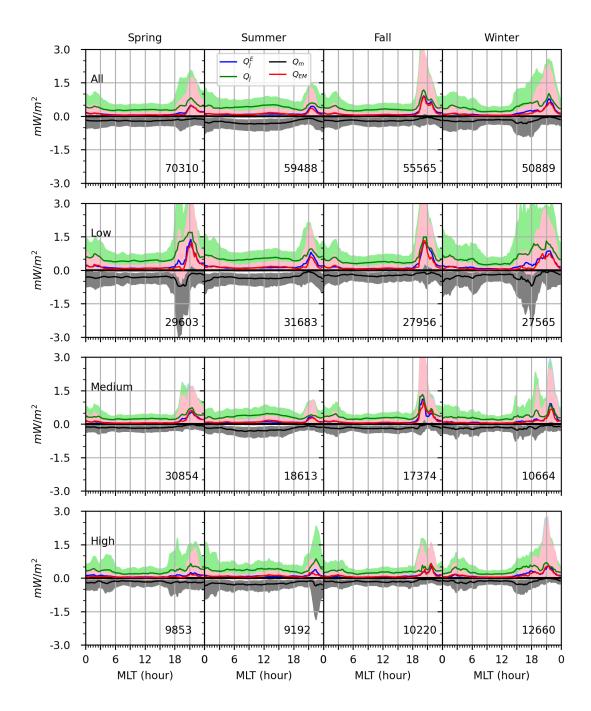
^aDetailed derivation of the neutral wind can be found in Heinselman and Nicolls (2008).

^bMagnetic field: **B**, gyro-frequency of ion: $\Omega_i = \frac{eB}{m}$, e: amount of elementary charge, ion-neutral frequency: v_{in} , see the definition in Schunk and Nagy (2009). The electron density needed is from the measurements, and the neutral densities are from the NRLMSISE-00 model (Picone et al., 2002).

in each subplot is labeled. Note the different scales of the vertical axes in Figures 1-3.

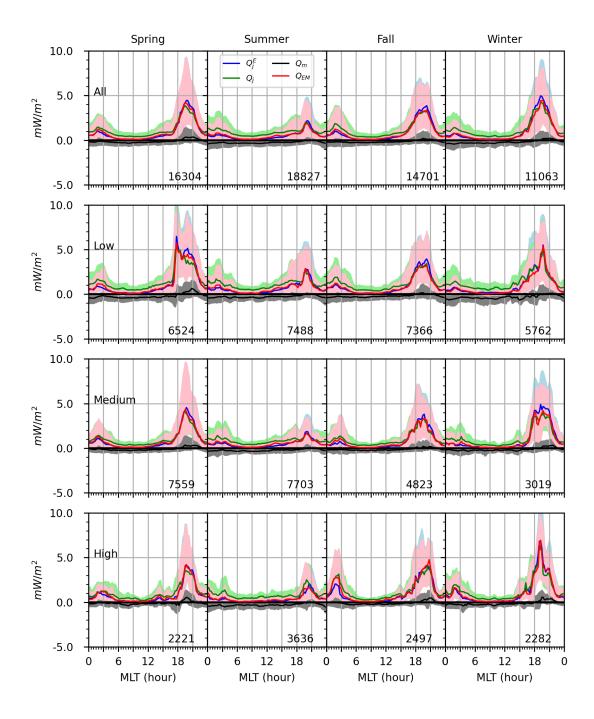
Figure 1 shows the results in geomagnetic quiet conditions. The results in the 223 first row for all years show that the passive energy deposition rate, Q_i^E , and EM 224 energy transfer rate, Q_{EM} , are very small ($<0.1 \, mW/m^2$) in most MLT sectors but 225 there are weak enhancements ($0.36 \sim 0.92 \, mW/m^2$) in the evening sector (1500 -226 2400 MLT) for all seasons. Here the enhancement refers to the increase of energy 227 transfer rates relative to the small daytime (0600-1500 MLT) values. The magni-228 tudes of the Joule heating rate, Q_j , are larger (maximum of Q_j : $1.2 \, mW/m^2$) than the passive energy deposition rate, Q_j^E , and the electromagnetic energy transfer 229 230 rate, Q_{EM} , in all seasons. These larger magnitudes of Q_i are mainly due to the 231 neutral wind-associated mechanical energy. Under quiet conditions the ion drifts 232 are frequently opposing the neutral winds such that the neutral gas is experienc-233 ing a net loss of kinetic energy (negative Q_m), and the frictional (Joule) heating 234 rates are higher than the passive rates would be ignoring the neutral winds (i.e. 235 $|u_i - u_n|^2 > |u_i|^2$). 236

Comparing the results for different seasons, we find that the enhancements 237 of the median Q_j^E , Q_j and Q_{EM} in the evening sectors are weakest in the summer $(0.45 \, mW/m^2, 0.60 \, mW/m^2, 0.36 \, mW/m^2)$ and have slightly larger magnitudes in 238 239 other seasons $(0.50 \sim 0.92 \, mW/m^2, 0.82 \sim 1.18 \, mW/m^2, 0.47 \sim 0.90 \, mW/m^2)$. When 240 comparing the results for different solar activity levels in the following rows, the 241 maxima of Q_j^E , Q_j , and Q_{EM} in the same season show a descending trend with the solar activity level. When comparing the results for different seasons, the maxima 242 243 of Q_i^E , Q_j , and Q_{EM} are smallest in the summer for low and medium solar activity 244 levels. For high solar activity level, the maxima of Q_j^E , Q_j , and Q_{EM} show similar 245 magnitudes in spring and summer and larger values in fall and winter. The mechani-246 cal energy transfer rates show enhanced magnitudes in the spring and winter for low 247 solar activity level and smaller magnitudes during medium and high solar activity 248 conditions. Detailed comparisons are included in Tables S1 in the Supplemental 249 Information SI. 250



Integrated Energy Transfer Rates (Quiet)

Figure 1. Energy transfer rates (green: Q_j , blue: Q_j^E , black: Q_m , red: Q_{EM}) during quiet condition in different seasons and different solar activity levels. From left to right: spring, summer, fall and winter. From Top to bottom: All years, low solar activity, medium solar activity and high solar activity. Shaded areas corresponds to the regions bounded by the upper and lower quartiles of the energy transfer rates (light green: Q_j , light blue: Q_j^E , gray: Q_m , pink: Q_{EM}).



Integrated Energy Transfer Rates (Moderate)

Figure 2. The same as Figure 1 but for results during moderate condition.

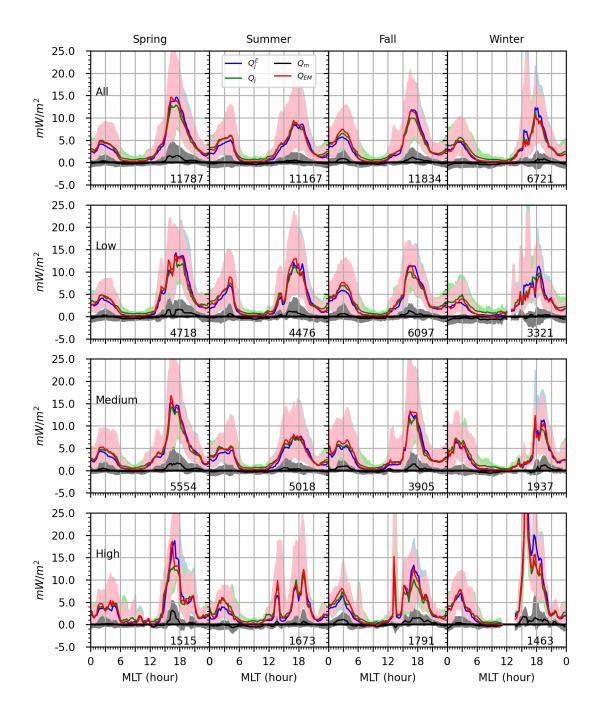
Figure 2 shows the results during moderate conditions that are plot-251 ted in a similar format to Figure 1. The results for all years in the first 252 row show enhancements of the median Q_j^E , Q_j and Q_{EM} in the morn-253 ing and evening sectors, with larger enhancements in the evening sector 254 for all seasons. Further comparison shows that the magnitudes of the en-255 hancements of Q_j^E , Q_j , and Q_{EM} in the evening sector have the follow-256 ing order: smallest in summer $(2.18 \, mW/m^2, 1.98 \, mW/m^2, 1.85 \, mW/m^2)$ 257 followed by fall $(3.86 \, mW/m^2, 3.27 \, mW/m^2, 3.40 \, mW/m^2)$, spring 258 $(4.45 \, mW/m^2, 3.87 \, mW/m^2, 4.21 \, mW/m^2)$, and slightly larger in winter 259 $(4.97 \, mW/m^2, 4.28 \, mW/m^2, 4.45 \, mW/m^2)$. However, there are relatively small 260 differences in the morning sector. The smallest magnitude of the energy transfer 261 rate in the summer relative to other seasons is consistent across solar activity levels, 262 as shown in the second through fourth rows. The solar activity level has a small 263 impact on the enhancements of energy transfer, especially in the summer and fall. 264 The differences in the same season among different solar activity levels are less than 265 $1 \, mW/m^2$ in summer and fall. 266

When we compare these results with the results in Figure 1, we see that the increase of the peak energy transfer rates associated with geomagnetic activity is a factor of five larger relative to the solar activity level for the same season. In addition, while the mechanical energy transfer rates are still small, we see a short interval (1900 - 2200 MLT) of positive values in the spring for all solar activity levels and in the fall and winter mainly for high solar activity level. Positive Q_m implies that the neutral atmosphere is accelerated by the plasma forcing.

In Figure 3, during active conditions the results for all years in the first row 274 show that the enhancements of the peak Q_j^E , Q_j and Q_{EM} in the evening sector are larger than that in the morning sector for all seasons and the peaks are a factor 275 276 of two larger relative to moderate geomagnetic conditions. The mechanical energy 277 transfer rate is still close to zero in most of the MLT sectors and all seasons ex-278 cept for a short interval with positive values in the morning (0200 - 0400 MLT) 279 and evening (1600 - 2000 MLT) sectors in the spring, fall and winter. Similarly, the 280 enhancements of the energy transfer rates in the morning sector have similar mag-281 nitudes for all seasons, while the enhancements in the evening sector are smallest in 282 the summer relative to other seasons. 283

The results for different solar activity levels during active geomagnetic condi-284 tions do not show consistent features. The results for low solar activity level in the 285 evening sector show very small differences with respect to season. The results for 286 medium solar activity level show the smallest enhancements of energy transfer in 287 the summer which is consistent with the results for quiet and moderate conditions. 288 The results for high solar activity level show abnormal large values and less smooth 289 curves compared to low and medium levels. This could be partially explained by 290 fewer observations for this solar activity level. For the same reason, the results for 291 spring, fall and winter show small changes from low to medium solar activity, while 292 high solar activity level shows large differences compared to low and medium solar 293 activity levels. 294

From the results in Figures 1, 2, and 3, we see that the geomagnetic activity 295 has the strongest impact on the energy transfer rates, followed by seasonal varia-296 tions, and last solar activity level variations. We also find that the duration of the 297 enhancements of energy transfer rates in the morning and evening sectors become 298 longer and the MLT of the peak shifts towards noontime sector with the increase of 299 geomagnetic activity similar to what were found by Aikio et al. (2012). We find that 300 the median energy transfer rates in the evening sector show the smallest enhance-301 ment in the summer during most solar and geomagnetic activity conditions, which is 302 different from observational and modeling results (Foster et al., 1983; Weimer, 2005; 303



Integrated Energy Transfer Rates (Active)

Figure 3. The same as Figure 1 but for results during active condition.

³⁰⁴ Zhang et al., 2005). A detailed summary of the maximum enhancement of energy

transfer rates in different seasons and different solar activity levels is provided in

Tables S1 in the Supporting Information.

307 4 Discussion

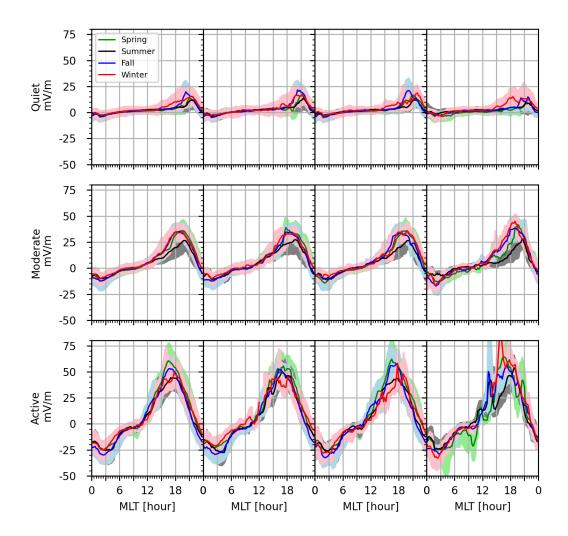
We find that the geomagnetic activity level has the most significant contribution to the MLT dependence of the auroral E-region EM energy transfer rate and Joule heating rate, followed by the seasonal variability, and finally solar activity level at Poker Flat latitude. We seek to understand which terms in the Joule heating equations are responsible, to first order, for the variability that we observe. The electric field, neutral wind and conductivity are the main factors in the energy transfer rates.

It is well-established that the electric field and particle precipitation- related 315 conductivity are closely related to the geomagnetic activity, therefore, we expected 316 to observe stronger energy transfer with an increase of geomagnetic activity. How-317 ever, it is not well understood what are the seasonal dependencies of the conduc-318 tivities and how those seasonal variations impact the EM energy transfer rate. In 319 particular, the differences of the energy transfer between winter and summer could 320 be associated with the differences of the particle precipitation-related Pedersen con-321 ductivity in the dark and sunlit hemisphere (Ohtani et al., 2009). 322

Our analysis will focus primarily on variations of the electric field and Peder-323 sen conductance. The integrated mechanical energy transfer rates shown above are 324 comparatively small during moderate and active geomagnetic conditions, therefore. 325 we will not investigate the integrated contribution from the neutral winds in this 326 study. However, the altitudinal structure of the neutral winds are known to have a 327 significant role on the altitude distribution of the energy transfer rates (Thaver & 328 Vickrey, 1992; Lu et al., 1995; Thayer, 1998b; Fujii et al., 1999; Thayer, 2000; Cai et 329 al., 2013). The seasonal variations during geomagnetic quiet intervals (AP < 16) of 330 the neutral winds has been investigated by Nozawa and Brekke (1999). 331

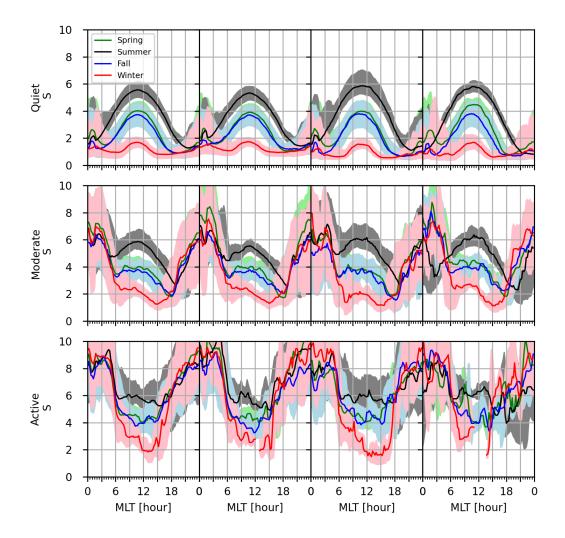
To compare the relative importance of the seasonal variations of the electric 332 field and the Pedersen conductance on the auroral E-region energy transfer rates, 333 we present the results of the median electric field and median Pedersen conduc-334 tance in Figures 4 and 5, respectively, and then compare their relative importance 335 in leading to the seasonal and solar activity dependence of the energy transfer rates. 336 Given that the magnitudes of the zonal electric field are small relative to that of the 337 meridional component, we only present the meridional electric field. In Figure 4, the 338 variations of the meridional electric field are plotted in green, black, blue, and red 339 curves corresponding to spring, summer, fall, and winter, respectively. The shaded 340 areas correspond to their lower and upper quartiles. The rows from top to bottom 341 correspond to quiet, moderate, and active conditions, respectively and the columns 342 correspond to all, low, medium, and high solar activities, respectively. We present 343 the variations of Pedersen conductance in Figure 5 in the same way. 344

We first summarize the key features shown in Figures 4 and 5. The magnitude 345 of the median meridional electric field in the evening sector is smallest in summer 346 while it has similar values in spring, fall, and winter during quiet and moderate ge-347 omagnetic activity conditions for all solar activity levels. This is consistent with the 348 smallest enhancements of energy transfer rates in summer relative to other seasons 349 shown above. The large electric field in the spring and fall equinox can be explained 350 by the semi-annual variation of the geomagnetic activity (Russell & McPherron, 351 1973; Lockwood et al., 2020). The electric field is closely related with the geomag-352 netic activity. In the evening sector it varies from between 20-25 mV/m during quiet 353



Electric Field

Figure 4. Variations of meridional electric field under quiet, moderate and active conditions (from top to bottom) in different seasons:spring (green), summer (black), fall (blue) and winter (red) and solar activity levels (from left to right: all (2010-2019), low, medium and high).



Pedersen Conductance

Figure 5. The same as Figure 4 but for Pedersen conductance.

conditions, to between 25-40 mV/m during moderate conditions, and to between 40-60 mV/m during active conditions.

The Pedersen conductance shows large seasonal variations during daytime while the solar and geomagnetic activity show small impacts on daytime conductance, which could be explained by the ISR being at subauroral magnetic latitudes during the day. The nighttime conductance shows a large dependence on geomagnetic activity while the solar activity and seasonal effects are generally small. The Pedersen conductance in the evening sector increases from below 2 S during quiet conditions, to between 2-6 S during moderate conditions, and to between 6-10 S during active conditions.

During quiet conditions, the conductance around 2100 MLT in the evening 364 sector is generally larger in the summer. The slightly larger conductance during 365 summer night could be partially attributed to weak solar EUV illumination, which 366 is a more pronounced effect during quiet intervals. To maintain current continuity a larger electric field in the winter/dark hemisphere is needed relative to the sum-368 mer/sunlit hemisphere. However, the smaller conductance in winter is not large 369 enough to offset the larger contribution from E^2 (see equation for Q_i^E in Table 1). 370 Therefore, we still see larger EM energy transfer and Joule heating rates in the win-371 ter. During moderate conditions, the conductance in the evening sector shown in 372 Figure 5 there is slightly larger in the winter than in the summer, due to the auro-373 ral precipitation. Thus there is larger EM energy transfer and Joule heating in the 374 winter than in the summer during moderate conditions which is the result of both 375 the electric field and the enhanced conductance. During active conditions, we do 376 not find a distinction of the electric field and conductance for different solar activity 377 levels; however, we can still see that the electric field and conductance in the evening 378 sector are slightly larger in the winter than in the summer from the data for all 379 years in the first column of Figures 4 and 5. 380

A detailed calculation of the ratios of E^2 and the Pedersen conductance be-381 tween winter and summer are presented in the Figure 1S in Supplement Information. 382 These results show that E^2 mostly dominates the Joule heating term during energy 383 transfer enhancements in the evening sector. Our observations of a larger electric 384 field in winter than summer is also consistent with the results from Foster et al. 385 (1983) and de la Beaujardiere et al. (1991). In addition, Foster et al. (1983) showed 386 that the Joule heating input is 50% greater in summer than in winter, primarily due 387 to conductivity enhancements caused by the solar production. For the case where 388 conductance from particle precipitation is used, Foster et al. (1983) showed that the 389 summer peaks at dawn and dusk are greatly reduced. The study by Foster et al. 390 (1983) corresponds to global scales and covers both the E- and F-region ionosphere, 391 and the E-region conductance has been shown to be less sensitive to solar activity 392 than the F-region (Sheng et al., 2014). Therefore, it is reasonable to conclude that 393 the variability of the E-region Pedersen conductance during disturbed conditions in 394 the evening sector in different seasons is mainly due to particle precipitation. Thus, 395 the larger energy transfer rates in the evening sector in the winter relative to sum-396 mer is consistent with the results excluding the conductance associated with solar 397 398 production (Foster et al., 1983).

Joule heating is a result of the closure of field-aligned currents (FACs) in the E-region, the summer-winter asymmetry of energy transfer can be associated with the summer-winter asymmetry of the FACs. Ohtani et al. (2009) used a large dataset of Defense Meteorological Satellite Program (DMSP) satellite observations to show that the Region 1 current density is larger in the dark hemisphere because the absence of solar illumination is often overcompensated by more intense and energetic electron precipitation thereby causing larger Pedersen conductance. This larger conductance in the dark hemisphere is consistent with the slightly larger conduc-tance in winter in this study.

Through a comparison between Figure 4 and Figure 5, we find that the geomagnetic dependence of the energy transfer enhancement is a result of the combination of enhanced electric field and conductance because geomagnetic activity is strongly correlated with large electric fields and Pedersen conductance. These geomagnetic dependence of energy transfer rates is consistent with the study by Fujii et al. (1999) and Aikio et al. (2012).

In addition, there are larger energy transfer rates in the evening sector than in 414 the morning sector for each season for the same geomagnetic and solar activity levels 415 which is driven by the larger electric field in the evening sector. Figure 4 shows that 416 the electric field in the evening sector is much stronger than in the morning sector 417 418 in all seasons for all geomagnetic and solar activity levels. Figure 5 shows that during all geomagnetic and solar activity conditions, though the median conductance 419 in the morning sector is slightly larger than in the evening sector, it cannot offset 420 the larger contribution from the electric field. This morning-evening asymmetry has 421 been investigated in detail in Zhan et al. (2021) and also reported in Thayer (2000). 422

423 5 Conclusion

We present one of the first comprehensive investigations of the Joule heating 424 and EM energy transfer rates in the high latitude E-region between 90-130 km as a 425 function of geomagnetic activity, season, and solar activity level. These results are 426 possible given the unique, nearly continuously sampled incoherent scatter radar data 427 obtained with PFISR. For this investigation we analyze observations from 2010-2019, 428 which nearly covers a solar cycle. The results we present also include the contribu-429 tion from the neutral winds to the Joule heating rate, making this one of the first 430 investigations to quantify the neutral wind contribution on the energy transfer rate 431 as a function of season and solar activity level. We summarize our main findings: 432

The median Joule heating and EM energy transfer enhancements in the 433 evening MLT sector show an asymmetry with respect to season; the heating rates 434 are smaller in summer versus winter and have similar magnitudes in the spring and 435 fall. The larger energy transfer rates in the evening sector in the winter versus sum-436 mer show different characteristics relative to global scale studies by satellites and 437 numerical simulations (Foster et al., 1983; Weimer, 2005; Zhang et al., 2005). We 438 find this result during quiet and moderate geomagnetic activity levels, and for all 439 solar activity levels. 440

We have demonstrated that the seasonal dependence of the energy transfer 441 enhancements, to first order, are associated with the seasonal variation of the elec-442 tric field. Further analysis shows that during quiet conditions, while the Pedersen 443 conductance is smaller in the winter than summer, the contribution from the electric 444 field is much larger, which leads to the larger Joule heating in the evening sector 445 in the winter. During moderate and active conditions, both the Pedersen conduc-446 tance and the electric field contribute to the larger magnitude of Joule heating in 447 the evening sector in the winter, although the contribution from the electric field is 448 generally larger than the conductance. We find the conductance to be larger in the 449 winter than in summer, which is consistent with the results in previous studies on 450 the dependence of FACs and particle precipitation (Ohtani et al., 2009). 451

⁴⁵² We also compared the relative importance of geomagnetic activity and solar ⁴⁵³ activity levels on the energy transfer rates. Geomagnetic activity has a larger im-⁴⁵⁴ pact than solar activity. Our results show that the maximum energy transfer rates ⁴⁵⁵ increases by a factor of ~ 5 from geomagnetic quiet to moderate condition and by

factor of 3 from geomagnetic moderate to active conditions for the same season and 456 solar activity level. However, the change of the energy transfer rates due to the solar 457 activity for the same season and a fixed geomagnetic activity level is much smaller. 458 Generally, the change due to solar activity varies from below 1 mW/m^2 , to around 3 459 mW/m^2 , to around 5 mW/m^2 for geomagnetic quiet, moderate, and active intervals, 460 respectively. We find that the geomagnetic activity has the most significant impact 461 on the EM energy transfer and Joule heating rates, followed by seasonal variability, 462 and by solar activity variations. 463

This large dataset of energy transfer rates in the high latitude E-region has for the first time provided resolved observations of the energy transfer rates which show seasonal and solar activity dependencies during different geomagnetic activity levels. These results provide a climatological perspective of the energy transfer rates that can be used in ionosphere-thermosphere model development.

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470 The raw data of PFISR measurements used in this study can be obtained from

the Madrigal database (https://isr.sri.com/madrigal). The processed data

472 of neutral winds, electric fields, conductivities and energy transfer rates can

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480 References

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481	Aikio, A. T., Ca	ai, L., & Nygrén,	T. (2012). Statistica	al distribution of height-
482	integrated	energy exchange	e rates in the	ionosphere.	Journal of Geophysical
483	Research:	Space Physics, 1	17(10), 1-14.	doi: 10.1029/2	2012JA018078
484	Barth, C. A.	(2010).	Joule heatin	g and nitric ox	ide in the thermosphere,

2. Journal of Geophysical Research: Space Physics, 115(10), 1–7. doi: 10.1029/2010JA015565

- Barth, C. A., Lu, G., & Roble, R. G. (2009). Joule heating and nitric oxide in the thermosphere. Journal of Geophysical Research: Space Physics, 114(5), 1–9. doi: 10.1029/2008JA013765
- Bjoland, L. M., Chen, X., Jin, Y., Reimer, A. S., Skjæveland, A., Wessel, M. R., ...
 McWilliams, K. A. (2015). Interplanetary magnetic field and solar cycle dependence of Northern Hemisphere F region joule heating. *Journal of Geophysical Research: Space Physics*, 120(2), 1478–1487. doi: 10.1002/2014JA020586
- Brekke, A. (1979). On the relative importance of Joule heating and the Lorentz force in generating atmospheric gravity waves and infrasound waves in the auroral electrojets. *Journal of Atmospheric and Terrestrial Physics*, 41(5), 475–479. doi: 10.1016/0021-9169(79)90072-2
- Cai, L., Aikio, A. T., & Milan, S. E. (2016, jul). Joule heating hot spot at high lat itudes in the afternoon sector. J. Geophys. Res. Sp. Phys., 121(7), 7135–7152.
 doi: 10.1002/2016JA022432
- Cai, L., Aikio, A. T., & Nygrén, T. (2013). Height-dependent energy exchange
 rates in the high-latitude e region ionosphere. Journal of Geophysical Research:
 Space Physics, 118(11), 7369–7383. doi: 10.1002/2013JA019195
- Cai, L., Aikio, A. T., & Nygrén, T. (2014, dec). Solar wind effect on Joule heating
 in the high-latitude ionosphere. J. Geophys. Res. Sp. Phys., 119(12), 440–455.
 doi: 10.1002/2014JA020269

- de la Beaujardiere, O., Alcayde, D., Fontanari, J., & Leger, C. (1991).Seasonal 507 dependence of high-latitude electric fields. Journal of Geophysical Research, 508 96(A4), 5723. doi: 10.1029/90JA01987 509 Foster, J. C., St.-Maurice, J.-P., & Abreu, V. J. (1983). Joule heating at high lat-510 Journal of Geophysical Research, 88(A6), 4885–4896. itudes. doi: 10.1029/ 511 JA088iA06p04885 512 Fujii, R., Nozawa, S., Buchert, S. C., & Brekke, A. (1999).Statistical charac-513 teristics of electromagnetic energy transfer between the magnetosphere, the 514 ionosphere, and the thermosphere. Journal of Geophysical Research: Space 515 *Physics*, 104 (A2), 2357–2365. doi: 10.1029/98ja02750 516 Gjerloev, J. W. (2012). The SuperMAG data processing technique. J. Geophys. Res. 517 Sp. Phys., 117(9), 2018. doi: 10.1029/2012JA017683 518 Heinselman, C. J., & Nicolls, M. J. (2008).A Bayesian approach to electric 519 field and E -region neutral wind estimation with the Poker Flat Advanced 520 Modular Incoherent Scatter Radar . Radio Science, 43, RS5013. doi: 521 10.1029/2007rs003805 522 Lockwood, M., Owens, M. J., Barnard, L. A., Haines, C., Scott, C. J., McWilliams, 523 K. A., & Coxon, J. C. (2020). Semi-annual, annual and Universal Time varia-524 tions in the magnetosphere and in geomagnetic activity: 1. Geomagnetic data. 525 Journal of Space Weather and Space Climate, 10. doi: 10.1051/swsc/2020023 526 Lu, G., Richmond, A. D., Emery, B. A., & Roble, R. G. (1995).Magnetosphere-527 ionosphere-thermosphere coupling: Effect of neutral winds on energy transfer 528 and field-aligned current. Journal of Geophysical Research: Space Physics, 529 100(A10), 19643-19659. doi: 10.1029/95JA00766 530 Makarevich, R. A., Koustov, A. V., & Nicolls, M. J. (2013). Poker Flat Incoherent 531 Scatter Radar observations of anomalous electron heating in the e region. Ann. 532 Geophys., 31(7), 1163–1176. doi: 10.5194/angeo-31-1163-2013 533 McHarg, M., Chun, F., Knipp, D., Lu, G., Emery, B., & Ridley, A. (2005, aug).534 High-latitude Joule heating response to IMF inputs. Journal of Geophysical 535 Research: Space Physics, 110(A8), 1–9. doi: 10.1029/2004JA010949 536 Newell, P. T., & Gjerloev, J. W. (2014).Local geomagnetic indices and the pre-537 diction of auroral power. Journal of Geophysical Research: Space Physics, 538 119(12), 9790-9803. doi: 10.1002/2014JA020524 539 Nozawa, S., & Brekke, A. (1999). Seasonal variation of the auroral e-region neutral 540 wind for different solar activities. Journal of Atmospheric and Solar-Terrestrial 541
- Physics, 61(8), 585-605. doi: 10.1016/S1364-6826(99)00016-4
 Ohtani, S., Wing, S., Ueno, G., & Higuchi, T. (2009, dec). Dependence of premidnight field-aligned currents and particle precipitation on solar illumination. J. *Geophys. Res. Sp. Phys.*, 114 (A12). doi: 10.1029/2009JA014115
- Picone, J. M., Hedin, A. E., Drob, D. P., & Aikin, A. C. (2002, dec). NRLMSISE 00 empirical model of the atmosphere: Statistical comparisons and scientific
 issues. Journal of Geophysical Research: Space Physics, 107(A12), 1648. doi:
 10.1029/2002JA009430
- Richmond, A. D., Lathuillère, C., & Vennerstroem, S. (2003, feb). Winds in the
 high-latitude lower thermosphere: Dependence on the interplanetary magnetic
 field. J. Geophys. Res. Sp. Phys., 108(A2), 1–14. doi: 10.1029/2002JA009493
- Russell, C. T., & McPherron, R. L. (1973, jan). Semiannual variation of geomagnetic activity. Journal of Geophysical Research, 78(1), 92–108. doi: 10.1029/ JA078i001p00092
- Schunk, R., & Nagy, A. (2009). Ionospheres: Physics, Plasma Physics, and
 Chemistry. Cambridge: Cambridge University Press. doi: 10.1017/
 CBO9780511635342
- Sheng, C., Deng, Y., Yue, X., & Huang, Y. (2014). Height-integrated Pedersen conductivity in both E and F regions from COSMIC observations. J. Atmos. Solar-Terrestrial Phys., 115-116, 79–86. doi: 10.1016/j.jastp.2013.12.013

562	Sofko, G. J., & Huang, C. S. (2000). SuperDARN observations of medium-scale
563	gravity wave pairs generated by Joule heating in the auroral zone. Geophysical
564	Research Letters, 27(4), 485–488. doi: 10.1029/1999GL003692
565	Sojka, J. J., Nicolls, M. J., Heinselman, C. J., & Kelly, J. D. (2009). The PFISR
566	IPY observations of ionospheric climate and weather. J. Atmos. Solar-
567	Terrestrial Phys., 71(6-7), 771–785. doi: 10.1016/j.jastp.2009.01.001
568	Thayer, J. P. (1998a, jan). Height-resolved Joule heating rates in the high-latitude
569	E region and the influence of neutral winds. Journal of Geophysical Research:
570	Space Physics, 103(A1), 471–487. doi: 10.1029/97JA02536
571	Thayer, J. P. (1998b). Radar measurements of the electromagnetic energy rates
572	associated with the dynamic ionospheric load/generator. Geophysical Research
573	Letters, $25(4)$, 469–472. doi: 10.1029/97GL03660
574	Thayer, J. P. (2000). High-latitude currents and their energy exchange with the
575	ionosphere-thermosphere system. Journal of Geophysical Research: Space
576	Physics, 105(A10), 23015-23024.doi: 10.1029/1999ja000409
577	Thayer, J. P., & Semeter, J. (2004). The convergence of magnetospheric energy
578	flux in the polar atmosphere. Journal of Atmospheric and Solar-Terrestrial
579	<i>Physics</i> , $66(10)$, 807–824. doi: 10.1016/j.jastp.2004.01.035
580	Thayer, J. P., & Vickrey, J. F. (1992). On the contribution of the thermospheric
581	neutral wind to high-latitude energetics. Geophysical Research Letters, $19(3)$,
582	265-268. doi: $10.1029/91$ GL02868
583	Weimer, D. R. (2005). Improved ionospheric electrodynamic models and application
584	to calculating Joule heating rates. J. Geophys. Res., $110(A5)$, A05306. doi: 10
585	.1029/2004JA010884
586	Yuan, Z., Fujii, R., Nozawa, S., & Ogawa, Y. (2005). Statistical height-dependent
587	relative importance of the Lorentz force and Joule heating in generating at-
588	mospheric gravity waves in the auroral electrojets. Journal of Geophysical
589	Research, 110(A12), A12303. doi: 10.1029/2005JA011315
590	Zhan, W., Kaeppler, S. R., Larsen, M. F., Reimer, A., & Varney, R. (2021). An
591	investigation of auroral e region energy exchange using poker flat incoherent
592	scatter radar observations during fall equinox conditions. Earth and Space
593	Science Open Archive, 25. doi: 10.1002/essoar.10508005.1
594	Zhang, X. X., Wang, C., Chen, T., Wang, Y. L., Tan, A., Wu, T. S., Wang, W.
595	(2005). Global patterns of Joule heating in the high-latitude ionosphere. J .

⁵⁹⁶ Geophys. Res., 110(A12), A12208. doi: 10.1029/2005JA011222